

A 15-year record (2001–2015) of the ratio of nitrate to non-sea-salt sulfate in precipitation over East Asia

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Abstract. Acidifying species in precipitation can have severe impacts on ecosystems. The chemical composition of precipitation is directly related to the amount of precipitation; accordingly, it is difficult to identify long-term variation in chemical concentrations. The ratio of the nitrate (NO_3^-) to non-sea-salt sulfate (nss- SO_4^{2-}) concentration in precipitation on an equivalent basis (hereinafter, Ratio) is a useful index to investigate the relative contributions of these acidifying species. To identify the long-term record of acidifying species in precipitation over East Asia, the region with the highest emissions worldwide, we compiled groundbased observations of the chemical composition of precipitation over China, Korea, and Japan from 2001 to 2015 based on the Acid Deposition Monitoring Network in East Asia (EANET). The spatial coverage was limited, but additional monitoring data for Japan, southern China, and northern China around Beijing were utilized. The period of analysis was divided into three phases: Phase I (2001-2005), Phase II (2006–2010), and Phase III (2011–2015). The behaviors of NO_3^- and nss- SO_4^{2-} concentrations and hence the Ratio in precipitation were related to these precursors. The anthropogenic NO_x and SO₂ emissions and the NO_x / SO₂ emission ratio were analyzed. Further, satellite observations of the NO₂ and SO₂ column density to capture the variation in emissions were applied. We found that the longterm trend in the NO_3^- concentration in precipitation was not related to the variation in NO_x emission and the NO_2 column. In comparison, the nss- SO_4^{2-} concentration in precipitation over China, Korea, and Japan was partially connected to the changes in SO₂ emissions from China, but the trends were not significant. The long-term trends of Ratio over China, Korea, and Japan were nearly flat during Phase I, increased significantly during Phase II, and were essentially flat again during Phase III. This variation in Ratio in East Asia clearly corresponded to the NO_x / SO_2 emission ratio and the NO₂ / SO₂ column ratio in China. The initial flat trend during Phase I was due to increases in both NO_x and SO₂ emissions in China, the significantly increasing trend during Phase II was triggered by the increase in NO_x emissions and decrease in SO₂ emissions in China, and the return to a flat trend during Phase III was caused by declines in both NO_x and SO_2 emissions in China. These results suggest that emissions in China had a significant impact not only on China but also on downwind precipitation chemistry during the 15-year period of 2001-2015. In terms of wet deposition, the NO₃⁻ wet deposition over China, Korea, and Japan did not change dramatically, but the nss- SO_4^{2-} wet deposition declined over China, Korea, and Japan from Phase II to III. These declines were caused by a strong decrease in the $nss-SO_4^{2-}$ concentration in precipitation accompanied by a reduction in SO₂ emission from China, which counteracted

the increase in precipitation. These findings indicated that the acidity of precipitation shifted from sulfur to nitrogen.

1 Introduction

Accompanying the recent acceleration of anthropogenic emissions in Asia, atmospheric deposition in Asia has been a major focus worldwide (Vet et al., 2014). According to the Clean Air Status and Trends Networks (CASTNET) in the US (Sickles II and Shadwick, 2015) and the European Monitoring and Evaluation Programme (EMEP) in Europe (Tørseth et al., 2012), in Japan, which is located in the downwind region of the Asian continent, the total wet and dry deposition have surpassed those of both the US and Europe (Endo et al., 2011; Ban et al., 2016). In terms of the wet deposition of nitrogen and sulfur compounds over Japan, the influence of anthropogenic emissions from China has been determined via chemical transport model simulations (Kuribayashi et al., 2012; Kajino et al., 2011, 2013; Morino et al., 2011).

The amount of wet deposition is affected by the amount of precipitation. For this reason, the ratio of the nitrate (NO_3^-) concentration to non-sea-salt sulfate $(nss-SO_4^{2-})$ concentration in precipitation on an equivalence basis, hereafter referred to as Ratio, is useful for evaluating the relative contributions of nitrogen and sulfur to the acidity of precipitation. Ratio in Japan was about 0.41 between 1987 and 1990 (Fujita et al., 2003). Over western Japan, Ratio increased substantially by 0.09-0.17 between 1987 and 1996, reaching about 0.5 in the late 1990s. This corresponded to a large increase in NO_x emissions compared with SO_2 emissions across East Asia (Takahashi and Fujita, 2000). In our previous studies (Itahashi et al., 2014a, 2015), we showed that the Ratio trend in precipitation over Japan remained flat at around 0.5-0.6 between 2000 and 2005, with a subsequent increase to 0.6-0.7 between 2006 and 2011. During this period, the NO_x / SO_2 emission ratio in Japan constantly decreased, and changes in *Ratio* closely followed the changes in the NO_x / SO_2 emission ratio in China. A significant increase in Ratio between 2006 and 2011 was also found over both China and Korea. A correlation between Ratio in China, Korea, and Japan and the NO_x / SO_2 emission ratio in China was detected. A sensitivity simulation via a regional chemical transport model of Chinese anthropogenic emissions indicated the important contribution of NO_x emissions near the source region and the high impact of SO₂ emissions over downwind regions. Identifying the importance of acidity via nitrogen and sulfur is critical to mitigate ecosystem effects, such as soil acidification (Zhao et al., 2009) and surface water acidification (Yamashita et al., 2016). Similar studies of Ratio (or the inverse of Ratio, defined as the ratio of the nss- SO_4^{2-} to NO_3^{-} concentration in precipitation) have been conducted in northern China (Wang et al., 2012), Nanjing (Tu et al., 2005), southeastern China (Cui et al., 2014), and the Pearl River Delta region (Lu et al., 2015).

We next consider recent changes in *Ratio* in East Asia according to recent changes in anthropogenic emissions. In our previous studies (Itahashi et al., 2014a, 2015), which considered data obtained in 2000-2011, emissions from China exhibited a transition from a continuously increasing trend. NO_x emissions in China have been rising steadily (Kurokawa et al., 2013; Itahashi et al., 2014b), whereas SO₂ emissions in China peaked in 2005–2006 and subsequently declined due to the introduction of flue-gas desulfurization systems in China's 11th Five-Year Plan (2006–2010) (Kurokawa et al., 2013; Itahashi et al., 2012). Recently, it has been reported that NO_x emissions in China declined after 2011–2012 (Irie et al., 2016; Xia et al., 2016; Krotkov et al., 2016; van der A et al., 2017). To determine the result of this change, we revisited and updated an analysis of precipitation data in relation to emission variation over East Asia. We compiled precipitation observations over the 15-year period between 2001 and 2015. Considering the drastic change in emissions from China, the analysis period was divided into three parts: Phase I (2001–2005), Phase II (2006–2010), and Phase III (2011-2015). This study extends our previous studies (Itahashi et al., 2014a, 2015) in the following respects. First, this study incorporates additional data obtained over southern China and northern China around Beijing. The groundbased observation network in East Asia does not cover northern China, which is characterized by a large urban population and therefore contributes substantial anthropogenic emissions (e.g., Pan et al., 2012, 2013). Although these additional data over southern and northern China are limited to 1 year or a few years, this approach of combining and comparing datasets will further improve our understanding of precipitation chemistry over all of China. Additionally, this study uses satellite observations of NO₂ and SO₂ column density as a proxy to estimate recent emission changes. This can further enhance our knowledge of the current emission status. Analyses of long-term precipitation chemistry are still limited. For example, precipitation chemistry over periods of longer than 10 years has been evaluated for Guangzhou in 1983-2010 (Fang et al., 2013), Shenzhen in 1986-2006 (Huang et al., 2008), Lijiang in 1989-2006 (Zhang et al., 2012), Nanjing in 1992-2003 (Tu et al., 2005) and Tieshanping in 2001–2013 (Yu et al., 2017). This long-term, 15-year analysis provides a comprehensive overview of the precipitation chemistry over East Asia.

This paper is structured as follows. Section 2 introduces the dataset used in this study, which includes ground-based observations, emission inventories, and satellite observations that were applied as a proxy for emission levels. Section 3 is dedicated to the results and discussion. First, Sect. 3.1 presents the long-term trends for NO_3^- and $nss-SO_4^{2-}$ in China, Korea, and Japan. Section 3.2 describes the intensive analysis of *Ratio* and its relation to emissions. Section 3.3



Figure 1. Geographical mapping of observation sites included in this study. Detailed information about each observation site is provided in Table 1.

explores the analysis of wet deposition. Finally, Sect. 4 summarizes this research and outlines future perspectives.

2 Dataset

2.1 Ground-based observations

The ground-based observations used in this study are spatially mapped in Fig. 1. Detailed information for the observation period, including latitude (° N), longitude (° E), elevation (meters above sea level, a.s.l.), and the classification of each site, is listed in Table 1. The observation dataset for the chemical composition of precipitation, which was compiled by the Acid Deposition Monitoring Network in East Asia (EANET) program (EANET, 2017), was mainly used in this study. In China, EANET observations were obtained over three areas in southern China (Zhuhai, Xiamen, and Chongqing) and one area in central China (Xi'an), for a total of 10 sites. In Korea, EANET observations from three sites (Cheju, Imsil, and Kanghwa) were available. In Japan, data for a total of 11 sites were available on EANET. Among these 11 sites, the data obtained at Ogasawara, which is located in the northwest Pacific Ocean (27.09° N, 142.22° E, 230 m a.s.l.), includes estimates of the chemical composition of precipitation that are below the detection limit, that is, 11.8 % for NO₃⁻ and 7.7 % for nss-SO₄²⁻. Therefore, the data from Ogasawara were excluded. Samples were collected using a wet-only sampler at a daily interval, except for Banryu (sampled weekly during 2001-2015) and Ijira (sampled weekly in 2001). Concentrations of NO_3^- and SO_4^{2-} in precipitation were determined by ion chromatography and qualified via ion balance and conductivity agreement. The completeness of the data was determined from the precipitation coverage duration and total precipitation amount (EANET, 2000, 2010). In Japan, two sites (Ryori and Komae) were included in the study. Under the Global Atmosphere Watch (GAW) program of the World Meteorological Organization (WMO GAW, 2004), the Japan Meteorological Agency (JMA) has been conducting observations of the atmospheric concentration and deposition at the northern, remote Ryori site since 1976; observations of deposition at Ryori ended in 2011. Samples were collected using a wet-only sampler at a daily interval at Ryori. The Central Research Institute of Electric Power Industry (CRIEPI) conducted continuous monitoring of precipitation chemistry. The CRIEPI data obtained at Komae, which is located near Tokyo, span 1987 to the present (Fujita et al., 2000). As we have more fully described in our previous studies (Itahashi et al., 2014a, 2015), CRIEPI monitoring was also conducted at Goto Island, which is located at the western edge of Japan. However, the period of coverage was limited to February 2000 to April 2003, so we excluded the Goto data from this study. Samples were collected using a wet-only sampler at 10-day intervals in Komae. For the EANET, JMA, and CRIEPI monthly mean datasets, the nss- SO_4^{2-} concentration (in mol L^{-1}) was calculated from the conservative assumption that sodium (Na⁺) is a sea salt tracer, using the following equation:

$$nss-SO_4^{2-} = SO_4^{2-} - 0.06028 \times Na^+$$
(1)

To analyze the long-term behavior of the dataset over the 15-year period of 2001–2015, outliers in the observation of EANET, JMA, and CRIEPI were carefully examined according to method used in Itahashi et al. (2015). The volume-weighted monthly mean concentrations of NO_3^- , nss- SO_4^{2-} , and *Ratio* in precipitation were analyzed using the Smirnov–Grubbs outlier test at each site. In this study, the outliers for *Ratio* were directly checked. In this method, outliers were detected one at a time, assuming that the data likely followed

	Country	Site	Period	Latitude (° N)	Longitude (° E)	Elevation (m a.s.l.)	Classification
EANET	China	Zhuhai					
		Xiangzhou	2001-2015	22.27	113.57	40	urban
		Zhuxiandong	2001-2015	22.20	113.52	45	urban
		Xiamen		24.45	110.12	-	
		Hongwen	2001-2015	24.47	118.13	50	urban
		Changeing	2001-2015	24.85	118.03	080	remote
		Guanvingiao	2001_2007	29.57	106 52	262	urban
		Haifu	2008-2015	29.62	106.52	317	urban
		Jinvunshan	2001-2015	29.82	106.37	800	rural
		Xi'an					
		Shizhan	2001-2015	34.23	108.95	400	urban
		Weishuiyuan	2001-2006	34.37	108.85	366	rural
		Jiwozi	2001-2015	33.83	108.80	1800	remote
	Korea	Cheju	2001-2015	33.30	126.17	72	remote
		Imsil	2001-2015	35.60	127.18	-	rural
		Kanghwa	2001-2015	37.70	126.28	150	urban
	Japan	Hedo	2001–2015	26.87	128.25	60	remote
		Yusuhara	2001–2015	33.38	132.94	790	remote
		Banryu	2001-2015	34.68	131.80	53	urban
		ljira Tu	2001-2015	35.57	136.69	140	rural
		Tokyo	2007-2015	35.69	139.76	26	urban
		OKI Sadasalti	2001-2015	30.29	133.19	90	remote
		Toppi	2001-2013	56.25 41.25	136.40	130	remote
		Ochiishi	2001-2015	41.23	140.33	100	remote
		Rishiri	2003-2015	45.10	141.21	40	remote
	Ionon	Duori	2001 2013	20.02	141.21	260	Temote
	Japan	Kyon	2001-2011	25.64	141.82	200	
	Japan	Kolliae	2001-2013	55.04	139.38	27	
IMPACTS	China	Guangdong					
		Liuxihe (LXH) Guizbou	2002-2003	23.55	113.58	500	
		Leigongshan (LGS)	2002-2003	26 37	108-18	1630-1735	
		Liuchongguan (LCG)	2002 2003	26.63	106.72	1320-1400	
		Hunan	2001 2002	20100	1001/2	1020 1100	
		Caijiatang (CJT)	2001-2003	27.92	112.43	450-500	
		Chongqing					
		Tieshanping (TSP)	2001-2003	29.63	104.68	450–500	
BNU	China	Beijing					
		Beijing Normal University	2005	39.96	116.37	40	
LAPC/IAP/CAS	China	Shandong					
		Yucheng (YC)	2008-2010	36.85	116.55	23	
		Hebei					
		Luancheng (LC)	2008-2010	37.89	114.69	57	
		Cangzhou (CZ)	2008-2010	38.30	116.87	10	
		Baoding (BD)	2008-2010	38.85	115.50	21	
		Tangshan (TS)	2008-2010	39.60	118.20	24	
		Xinglong (XL)	2008-2010	40.38	117.57	872	
		Tanggu (TC)	2008 2010	20.04	117 70	0	
		Tanggu (TO) Tianiin (TI)	2008-2010	39.04 30.09	117.72	0	
		Beijing	2000-2010	39.08	117.21	0	
		Beijing (BI)	2008-2010	30.06	116 36	57	
		Yangfang (YF)	2008-2010	40.15	116.10	73	
					110.10	13	

 Table 1. Locations of observation sites included in this study.

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an approximately normal distribution. The hypothesis of no outliers is rejected when

$$G = \max_{i=1,...,N} \frac{|C_i - \overline{C}|}{s} > \frac{N-1}{\sqrt{N}} \sqrt{\frac{t_{(\alpha/2N, N-2)}^2}{N-2 + t_{(\alpha/2N, N-2)}^2}}$$
(2)

where N, \overline{C} , and s are the number, mean, and standard deviation, respectively, of NO₃⁻ or nss-SO₄²⁻ concentrations in precipitation or *Ratio* in precipitation (*C_i*). $t_{(\alpha/2N, N-2)}^2$ denotes the critical value of the t distribution with (N-2) degrees of freedom and a significance level of $(\alpha/2N)$. Outlier detection and removal were iterated until the dataset satisfied the specified significance level of 0.05. Following the above criteria, 4.7, 3.8, and 4.4 % of data points for NO₃⁻, nss- SO_4^{2-} , and *Ratio* in precipitation, respectively, were discarded from the China dataset; 5.0, 3.0, and 3.0% of these were discarded from the Korea dataset; and 2.4, 1.5, and 1.6% of these were discarded from the Japan dataset. Finally, the annual mean concentrations of NO_3^- , nss- SO_4^2 , and the annual mean Ratio in precipitation were calculated from the monthly mean data when at least 9 months of data were available for a given year at the site. For the treatment of the monthly accumulated precipitation amount, the Smirnov-Grubbs test for Ratio was applied to discard outliers. If at least 9 months were available, the annual accumulated precipitation was calculated.

These observation datasets taken from EANET, JMA, and CRIEPI were essentially the same as those used in our previous studies (Itahashi et al., 2014a, 2015). A limitation of our previous studies was a lack of spatial coverage over northern China, especially around the capital of Beijing because EANET covered only the area from southern to central China (Fig. 1). It has been recognized that anthropogenic emissions centered over this region (Kurokawa et al., 2013; Li et al., 2017), the related atmospheric concentration, and depositions are severe in China. To overcome this limitation and advance our knowledge of precipitation chemistry over the whole of China, we evaluated additional sources of data for the chemical concentration of precipitation over China.

The Integrated Monitoring Program on Acidification of the Chinese Terrestrial System (IMPACTS) was established through a Chinese–Norwegian cooperative project (Larssen et al., 2004, 2006) from 2001 to 2003. Under the IMPACTS program, atmospheric concentration, precipitation composition, and soil, water, and vegetative effects were studied at five forested sites (LXH, LGC, LCG, CJT, and TSP; refer to Fig. 1 and Table 1) over southern China. In terms of deposition, four measurements (wet-only, bulk, throughfall collected below the tree canopy, and belowground vegetation) were obtained. Wet-only sampling data were used in this study. Observations of the chemical composition of precipitation by a wet-only sampler were reported for four sites (LGS, LCG, CJT, and TSP) from 2001 to 2003 and for one site (LXH) from 2002 to 2003. Data for LGS in 2001 were not used, owing to insufficient coverage. At TSP, continuous monitoring results up to 2013 have recently been reported (Yu et al., 2017).

The precipitation chemistry over northern China has important implications. The precipitation sample at Beijing Normal University (BNU) in 2003 (Sun et al., 2004; Tang et al., 2005) and the recent work by the State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry/Institute of Atmospheric Physics/Chinese Academy of Sciences (LAPC/IAP/CAS) from December 2007 to November 2010 (Pan et al., 2012, 2013; Pan and Wang, 2015; Wang et al., 2012) were included in this study. Using the observation framework from BNU, data for a total of 53 rain events were collected at Beijing. To prevent contamination from dry deposition, the collector surface was covered with a plastic lid. A detailed description of this collection method is provided by Tang et al. (2005). In the recent work conducted by LAPC/IAP/CAS, a 3-year observation from December 2007 to November 2010 was conducted at 10 sites around Beijing. Daily rainwater samples were collected using a customized wet/dry automatic collector. The precipitation sensor opened the collection funnel of the cover device when rainfall began. In this study, December 2007 to November 2008 was regarded as the year 2008, December 2008 to November 2009 was regarded as the year 2009, and December 2009 to November 2010 was regarded as the year 2010. For the IMPACTS and LAPC/IAP/CAS datasets, the nss- SO_4^{2-} concentration was calculated using Na⁺ as a sea salt tracer based on Eq. (1). Owing to the lack of information on Na⁺ within the samples at BNU, SO_4^{2-} was used; however, it was reported that Na⁺ is not a major ion component of these samples.

2.2 Emission inventories

It is predicted that variation in NO_3^- , SO_4^{2-} , and hence *Ratio* is directly related to the emission of NO_x and SO_2 . We used the following emission inventories in this study. The Regional Emission inventory in ASia (REAS) version 2.1 (Kurokawa et al., 2013), which covers Asia from 2000 to 2008, was the main dataset used in this study to obtain NO_x and SO_2 emissions over China, Korea, and Japan. The recent status of Asian anthropogenic emission inventories with a mosaic approach named MIX (Li et al., 2017). MIX covers the years 2008 and 2010. The data from MIX were also used to acquire data for NO_x and SO_2 emissions over China, Korea, and Japan.

The latest country-level status of emissions in China is described by Xia et al. (2016). Their emissions data cover the period from 2000 to 2014. The primary case, which analyzed the contributions of advanced combustors with improved energy efficiency and air pollutant control devices with improved pollutant removal efficiency, was used in this study. In Korea, the National Institute of Environmental Research (NIER) provided estimates of the national emissions via the National Air Pollutants Emission Service, and the latest data reported cover 1999 to 2013 (NIER, 2017). In Japan, the Japan Auto-Oil Program (JATOP) provided 5-year interval emission datasets beginning in 1995 (JATOP, 2012a, b). The datasets for 2005 and 2010 were used in this study.

Using the total amounts of NO_x and SO₂ emissions estimated via inventories over China, Korea, and Japan, the NO_x / SO₂ emission ratio on a molar basis was calculated by calculating NO_x as NO₂. The behavior of the NO_x / SO₂ emissions ratio is correlated to *Ratio* in precipitation.

2.3 Satellite observations

The emission inventories had some time lags. Accordingly, satellite observations of the NO2 and SO2 vertical column density were combined to capture the recent status of NO_x and SO_2 . The NO_x / SO_2 column ratio reflects the NO_x / SO_2 emission ratio and was effective for characterizing the correspondence with Ratio in precipitation. Recently, satellite observations have been widely used as a proxy for emissions data. The NO₂ column has been used to capture NO_x emissions (e.g., Miyazaki et al., 2012; Mijling et al., 2013; Itahashi et al., 2014b; Han et al., 2015; Irie et al., 2016), and the SO_2 column has been used for SO_2 emissions (e.g., Lee et al., 2011; Li et al., 2010) and/or volcanic eruptions (e.g., Brenot et al., 2014). Several studies have indicated the importance of different technologies to control emissions (e.g., Li et al., 2010; Wang et al., 2015; Krotkov et al., 2016; van der A et al., 2017). For example, the ratio of Ozone Monitoring Instrument (OMI)-derived SO₂ / NO₂ was used to determine the effectiveness of the flue-gas desulfurization devices for power plants in China (Li et al., 2010; Wang et al., 2015).

The NO₂ and SO₂ column dataset, which was observed by OMI onboard the National Aeronautics and Space Administration (NASA) Earth Observing System Aura satellite, was used in this study (NASA, 2017). The Aura satellite was launched on 15 July 2004 in a sun-synchronous ascending polar orbit with a local equator crossing time of 13:45. During the data period, it measured sunlight backscattered from the Earth over a wide range of ultraviolet and visible wavelengths to derive abundances of ozone and other trace gases important for air quality and climate. Science-quality data operations began on 1 October 2004; hence, the data from 2005 to 2015 were used to cover our analysis period of 2001– 2015. Retrieval algorithms were based on the products provided by NASA.

In terms of the NO₂ column, we used the most recent version of level 3 daily global nitrogen dioxide product (OMNO2d) of version 3.0, which was released in August 2016 and is gridded at a resolution of $0.25^{\circ} \times 0.25^{\circ}$ (Krotkov, 2013; Krotkov et al., 2017). This product contains the total and tropospheric column for all atmospheric conditions and for sky conditions where the cloud fraction is less than 30 %.

We analyzed the tropospheric column with clouds screened on the condition of a cloud fraction of less than 30 %.

In terms of the SO₂ column, we used a level 3 daily global sulfur dioxide product (OMSO2e) of the latest version (3.0), which was released in February 2015 and is gridded at a resolution of $0.25^{\circ} \times 0.25^{\circ}$ (Krotkov et al., 2015). The dataset contains the total column of SO_2 in the planetary boundary layer (PBL). The algorithm was based on a principal component analysis, as introduced by Li et al. (2013). Cloud fraction, scene number, solar and satellite viewing angles, and row anomaly flags were provided as ancillary parameters. The data filtering of this level 3 dataset included the removal of rows with any of the following: anomaly flags, radiative cloud fraction greater than 20%, solar zenith angle greater than 70.0° , or scene number greater than 58 or less than 3. In addition, we adopted the smoothed method to average out the noise levels of the SO₂ column, following the research of Koukouli et al. (2016), who provided the anthropogenic loading of SO₂ over China as obtained from different satellite sensors. This method smoothed the SO₂ column assigned to each of the $0.25^{\circ} \times 0.25^{\circ}$ grid cells, which were weighted by the SO_2 column of the surrounding eight cells. In this process, the negative values were regarded as zero values because our aim was to construct the NO_x / SO_2 column ratio from satellite observations. Two data periods associated with volcanic activities were excluded from our analysis as follows. The Sarychev Volcano in the Kuril Islands (48.09° N, 153.20° E) had an explosive eruption that emitted a huge amount of ash and SO₂ at altitudes of 10-16 km (Brenot et al., 2014). The data from 14 to 22 June 2009 included this large amount of SO2 (over 10 D.U.) in and around the analyzed domain; accordingly, they were discarded from the calculations of monthly and annual means. An eruption of the Nabro Volcano in Eritrea (13.37° N, 41.07° E) from 12 June to 7 July 2011 was also reported. During the night of 12 June 2011, this volcano started to erupt, and on 14 June 2011, it spewed a volcanic plume across the route of many flights over east Africa and the Middle East (Brenot et al., 2014). The data from 15 June to 9 July 2011 were excluded from the calculations of monthly and annual means according to the approach detailed in van der A et al. (2017). In Japan, where many active volcanoes are located, SO₂ was continuously emitted at a level that surpassed anthropogenic emissions (e.g., Itahashi et al., 2017a). Owing to the difficulties of attempting to separate the effect of volcanic activity, the SO₂ column data for Japan, excluding the two data periods mentioned above, were used.

Based on the daily gridded data of NO₂ and SO₂ columns, monthly averages were calculated first and then the annual averages were calculated. In the calculation of annual averages, cells with monthly averaged data not available for at least 9 months were regarded as deficient cells for consistency with the criteria adopted for the ground-based observations of NO₃⁻, nss-SO₄², and *Ratio* in precipitation. The NO_x / SO_2 column ratio was obtained from the annual averaged gridded data for NO_2 and SO_2 columns.

3 Results and discussion

3.1 Long-term trend in the chemical composition of precipitation with respect to emissions

The long-term trends in the precipitation amount, NO_x emission from NO_2 columns, NO_3^- concentration in precipitation, SO_2 emission from SO_2 columns, and nss- SO_4^{2-} concentration in precipitation in 2001-2015 over China, Korea, and Japan are shown in Fig. 2. The averaged values and statistical analyses to estimate linear trends during Phase I (2001-2005), II (2006–2010), and III (2011–2015) are summarized in Table 2. Linear trends during each of the three phases were analyzed using linear regression, and significance levels were calculated using Student's t test. To support our finding of variation in the spatial distribution, satellite observations from 2005 to 2015 were used. NO₂ and SO₂ columns were mapped for 2006–2007 (the first half of Phase II), 2010–2011 (the transition from Phase II to Phase III), and 2014-2015 (the latter half of Phase III), as shown in Fig. 3. The annual variation in NO₂ and SO₂ columns, based on linear regression analyses during phases II and III, are also mapped in Fig. 4.

As shown in Fig. 2a, year-to-year variation in precipitation was found and the average values for annual precipitation accumulation were around 1300, 1100, and 1500 mm yr⁻¹ over China, Korea, and Japan, respectively. A statistical analysis revealed that, except for the increasing and decreasing trends over China and Korea during Phase III (p < 0.05), there were no clear changes in precipitation levels during the 15-year period.

The NO_x emissions and satellite observations of NO_2 columns are shown in Fig. 2b. Over China, emissions increased from 2001 to 2010, peaked in 2011-2012, and decreased after 2012, consistent with previous findings (e.g., Irie et al., 2016; Krotkov et al., 2016; van der A et al., 2017). Based on the spatial distribution, the NO₂ column above China peaked during the transition from Phase II and III (left panel of Fig. 3b) and contrasting trends, that is, increasing and decreasing trends, were revealed during phases II and III (left panel of Fig. 4a and b). Over China, the NO_x emissions of REAS and Xia et al. (2016) as well as satellite observations of NO₂ columns were highly consistent over long-term observations. Over Korea, NO_x emissions obtained from REAS and NIER showed a slight increase in 2003–2004 and a flat trend after 2008. There were some inconsistencies between NO_x emissions and the NO_2 column in 2005–2007. On the one hand, NO_x emissions showed a decreasing trend; on the other hand, the NO2 column showed a flat trend. Over Japan, both NO_x emissions and the NO_2 column exhibited a slight decrease during the 15-year period. Overall, the correlation between NO_x emissions and NO₂ columns suggests that satellite observations of the NO₂ column can serve as a proxy for NO_x emissions.

Such variation in NO_x emissions should be related to the change in the NO_3^- concentration in precipitation. The long-term trend in the NO_3^- concentration in precipitation is shown in Fig. 2c. The NO_3^- concentration in precipitation decreased in the order of China, Korea, Japan, with concentrations of approximately $50 \mu eq L^{-1}$ in China and $40 \,\mu\text{eq}\,\text{L}^{-1}$ in Korea. These levels were around 2-fold greater than that in Japan $(15-20 \mu eq L^{-1})$. The temporal variation found in the NO_3^- concentration in precipitation did not correspond to the variation in NO_x emissions (Fig. 2b). This was particularly notable for China; although NO_x emissions from 2001 to 2010 doubled, the NO_3^- concentration in precipitation did not increase substantially. The NO_3^- concentration in precipitation over Korea and Japan also did not show dramatic variation. This may be explained by changes in atmospheric composition and chemical regimes with changes in base cations and the oxidation capacity of the atmosphere, leading to changes in the lifetime of nitrogen species. Statistical analyses of the NO₃⁻ concentration in precipitation (Table 2) did not indicate a significant overall trend for phases I–III, with values remaining within -5to +3% yr⁻¹ over China, Korea, and Japan, with the exception of a negative trend of -3.4 ± 1.3 % yr⁻¹ (p < 0.05) in Japan during Phase II. When comparing EANET observations with additional datasets, IMPACTS sites around south China showed lower NO₃⁻ concentrations in precipitation, reflecting the characteristics of remote sites, while BNU and LAPC/IAP/CAS sites around Beijing showed relatively high concentrations, reflecting the emission intensity over northern China. These levels reflected the emission intensity of the particular regions and were nearly within 1 standard deviation of values obtained from the EANET observation network. For China, 10 EANET sites were classified as belonging to an urban category (six sites) and rural and remote sites (four sites) (Fig. 2c, see Table 1 for classification). Generally, concentrations were higher for urban sites than for rural and remote sites, and the correlation coefficient for mean values over urban sites and rural and remote sites was 0.21; these results can be attributed to differences in regional characteristics.

SO₂ emissions and satellite observations of the SO₂ column are shown in Fig. 2d. Over China, a steady increase was detected for 2001–2004, with a peak in 2005–2006 and a decrease after 2006; there thus seems to be a 5-year time lag prior to the decrease in NO_x emissions and NO₂ columns. The SO₂ column, which was available from 2005, showed a continuous decline over all of China (center of Fig. 3) and a decreasing trend was estimated by linear regression (center of Fig. 4a and b). In China, the SO₂ emissions of REAS and Xia et al. (2016) and SO₂ columns were highly consistent with the observed variation. Over Korea, SO₂ emissions obtained from REAS and NIER showed a slight de-

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Figure 2. Long-term temporal variation in (a) precipitation, (b) NO_x emission with the NO_2 column, (c) NO_3^- concentration in precipitation, (d) SO_2 emission with the SO_2 column, and (e) nss- SO_4^{2-} concentration in precipitation during 2001–2015 over (i) China, (ii) Korea, and (iii) Japan. Mean and 1 standard deviation across EANET observation sites (JMA and CRIEPI sites are also included in Japan) are indicated by the thick lines and shaded areas, respectively, in (a, c, e). For China, EANET observation sites are denoted with dashed lines for urban sites and dotted lines for rural and remote sites, and observation campaign data are indicated by marks corresponding to marks on Fig. 1. The numbers at the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).

crease, which was consistent with the SO_2 column reduction, but some discrepancies were found for 2007–2008. The temporal variation in the SO_2 column in Korea was similar to that observed in China. The spatial distribution data (Fig. 3) suggested that the SO_2 column over Korea is impacted by upwind regions, e.g., China. The SO_2 column obtained from SO_2 PBL products has a center of mass altitude at about 900 m, which is expected to lead to longer transport distances and longer lifetimes of SO₂. Therefore, compared with the correlation between NO_x emission and NO₂ columns, it was difficult to clearly capture the relation between SO₂ emissions and SO₂ columns. Over Japan, SO₂ emissions exhibited a slight decrease during the 15-year period, whereas a nearly flat trend was found for the SO₂ col-



Figure 3. Satellite observations of the (left) NO₂ column, (center) SO₂ column, and (right) NO₂ / SO₂ column ratio averaged over (a) 2006–2007 (the first half of Phase II), (b) 2010–2011 (transition between Phase II and Phase III), and (c) 2014–2015 (the latter half of Phase III). Dark-gray cells indicate places where annual mean could not be calculated. If either the NO₂ or SO₂ column value is less than 1.0×10^{15} molecules cm⁻², this is indicated by a light-gray cell to clarify the NO₂ / SO₂ column ratio.

umn. This is partially related to the fact that the SO_2 column includes SO_2 emissions from volcanoes. Hot spots of the SO_2 column above south Kyushu (a western island of Japan) and central Honshu (near Tokyo), shown in Fig. 3, were related to the location of volcanic activity during this period.

The nss-SO₄²⁻ concentration in precipitation also decreased in the order of China, Korea, and Japan, similar to the NO₃⁻ concentration in precipitation. The nss-SO₄²⁻ concentration in precipitation was 200–400 μ eq L⁻¹ above China, around 60 μ eq L⁻¹ above Korea, and 30 μ eq L⁻¹ above Japan. The levels over China were about 10-fold greater than those over Japan, and the levels over Korea were almost twice those over Japan (Fig. 2e). SO₂ emissions from China reached a peak in 2005–2006 (Fig. 2d),

and the nss-SO₄²⁻ concentration in precipitation above China reached approximately 393.0 µeq L⁻¹ in 2006, with a decline in 2015. The nss-SO₄²⁻ concentrations in precipitation above China were around 230 µeq L⁻¹ during Phase I and II but were 171.5 µeq L⁻¹ during Phase III (Table 2). Trends calculated by linear regression were 12.7 ± 8.3 % yr⁻¹ (not significant) during Phase I, -20.3 ± 8.8 % yr⁻¹ (not significant) during Phase I, and -13.6 ± 5.2 % yr⁻¹ (p < 0.05) during Phase III. It seems that the variation in the nss-SO₄²⁻ concentration in precipitation was partially related to SO₂ emission changes. In comparing EANET observations with other datasets, IMPACTS sites around south China showed lower nss-SO₄²⁻ concentrations in precipitation, BNU sites showed higher concentrations, and LAPC/IAP/CAS sites showed similar values. These levels were within 1 standard devia-

	Phase I (2001–2005)		Phase II (200	06–2010)	Phase III (2011–2015)				
	Mean	Trend	rend Mean Trend		Mean	Trend			
Precipit	tation (mm yr ^{-1})								
China	1330.2 (N = 35)	-4.3 ± 1.9	1390.2 (N = 31)	$+1.5 \pm 5.4$	1158.3 (N = 28)	$+6.6 \pm 2.5^{a}$			
Korea	1114.9 (N = 11)	-3.9 ± 15.4	1084.3 (N = 15)	$+6.9\pm3.0$	1015.1 (N = 15)	-11.2 ± 3.3^a			
Japan	1493.5 ($N = 44$)	-0.5 ± 3.8	1591.7 ($N = 56$)	$+0.9\pm3.2$	1555.7 (N = 56)	-0.5 ± 3.1			
$NO_3^- co$	NO_3^- concentration in precipitation (µeq L ⁻¹)								
China	51.3 (N = 31)	-1.9 ± 9.4	54.3 ($N = 31$)	$+0.6\pm5.5$	57.3 $(N = 34)$	-4.7 ± 5.0			
Korea	38.0 (N = 11)	$+3.0 \pm 4.1$	40.8 (N = 15)	-2.4 ± 4.7	42.7 (N = 15)	-1.1 ± 2.5			
Japan	17.2 (N = 44)	-0.9 ± 4.4	18.1 (N = 55)	-3.4 ± 1.3^{a}	18.0 (N = 55)	-0.6 ± 0.8			
nss-SO ₄ ²⁻ concentration in precipitation (μ eq L ⁻¹)									
China	237.0 (N = 32)	$+12.7\pm8.3$	230.6 (N = 32)	-20.3 ± 8.8	171.5 (N = 33)	-13.6 ± 5.2^a			
Korea	62.2 (N = 11)	$+10.0\pm0.8^{\rm c}$	66.9 (N = 15)	-6.8 ± 3.6	58.6 (N = 15)	-4.2 ± 2.6			
Japan	29.8 ($N = 44$)	-0.8 ± 5.2	30.4 (N = 56)	-5.4 ± 2.9	27.8 ($N = 55$)	-3.3 ± 1.5			

Table 2. Statistical analysis of average values and linear trends for precipitation, NO_3^- concentration in precipitation, and nss- SO_4^{2-} concentration in precipitation over China, Korea, and Japan during phases I, II, and III.

Note: total number of observation sites are shown in parentheses. Linear trends were estimated using linear regression and are shown as rates (% yr⁻¹). Significance levels are indicated by ^a p < 0.05, ^b p < 0.01, and ^c p < 0.001, and lack of a mark indicates a lack of significance. Analysis for Korea during Phase I was for the 2002–2005 period.

(a) Annual changes between 2006 and 2010 (Phase II)



Figure 4. Annual changes based on the linear regression results for the (left) NO₂ column, (center) SO₂ column, and (right) NO₂ / SO₂ column ratio during (a) Phase II and (b) Phase III. Dark-gray cells indicate areas where annual mean calculation could not be performed. If either the NO₂ or SO₂ column value is less than 1.0×10^{15} molecules cm⁻², this is indicated by a light-gray cell for clarity.

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Figure 5. Long-term temporal variation in the (**a**) NO_x / SO_2 emission ratio and NO_2 / SO_2 column ratio and (**b**) $NO_3^- / nss-SO_4^{2-}$ concentration in precipitation (*Ratio*) during 2001–2015 over (**i**) China, (**ii**) Korea, and (**iii**) Japan. Mean and 1 standard deviation across EANET observation sites (JMA and CRIEPI sites are also included in Japan) are indicated by the thick lines and shaded areas, respectively, in (**b**). For China, EANET observation sites are denoted with dashed lines for urban sites and dotted lines for rural and remote sites, and observation campaign data are indicated by marks corresponding to Fig. 1. The numbers at the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).

tion of the means obtained from the EANET observation network, similar to the NO_3^- concentration in precipitation. By classifying 10 EANET sites over China as urban sites or rural and remote sites (Fig. 2e; see Table 1 for classification), the mean values over urban sites and rural and remote sites were highly correlated, with a correlation coefficient of 0.78. This result implied that the nss- SO_4^{2-} concentration in precipitation was largely determined by large-scale characteristics related to SO₂ emissions. Variation in SO₂ emission from Korea and Japan was constant or declined between 2000 and 2008, and the nss-SO $_4^{2-}$ concentration in precipitation above Korea and Japan did not exhibit a clear relation with SO₂ emissions in these countries. In addition, the increasing trend of $+10.0 \pm 0.8 \% \text{ yr}^{-1}$ (p < 0.001) over Korea during Phase I was not related to the variation in SO₂ emissions in Korea. Relatively high nss- SO_4^{2-} concentrations in precipitation in Japan in 2001-2002 were partially explained by the SO₂ emissions from the Miyake-jima volcano in 2000 (Itahashi et al., 2012, 2014a, 2015). With the exception of an increase in Korea during Phase I and the peak in Japan in 2001–2002, nss-SO₄²⁻ concentrations in precipitation above Korea and Japan were high in 2005–2007 and subsequently decreased. These patterns were similar to those over China and appeared to be connected to the variation in SO_2 emissions from China, though these trends were not significant.

3.2 Long-term trend in *Ratio* in precipitation and its relationship to changes in emissions in China

The long-term variation in precipitation and NO₃⁻ and nss- SO_4^{2-} concentrations in precipitation did not indicate clear or significant trends during the 15-year period above China, Korea, and Japan, with a few exceptions. To further clarify the correlations between anthropogenic emission changes and the precipitation chemistry over East Asia, we focus on Ratio. The long-term trends in Ratio were analyzed for 2001-2015, and the results are shown in Fig. 5. The NO_x / SO₂ emission ratio and NO₂ / SO₂ column ratio were calculated using average values over each country. Over China, both the NO_x / SO_2 emission ratio and NO_2 / SO_2 column ratio were flat during Phase I, sharply increased during Phase II, and were nearly flat during Phase III. The substantial increase during Phase II was caused by an increase in NO_x emissions and a decrease in SO₂ emissions (Fig. 2b and d), as discussed in Itahashi et al. (2015). It was interesting to find that the trends in the NO_x / SO₂ emission ratio and the NO₂ / SO₂ column ratio were again flat during Phase III. This can be explained by the declines in both NO_x and SO_2 emissions. Based on the spatial mapping of the NO_2 / SO_2 column ratio (right part of Fig. 3), the NO_2 / SO_2 column ratio was substantially lower than 1.0 (red in the right panels of Fig. 3) above China during the first half of Phase II and subsequently increased above 1.0 (yellow to green in the right panels of Fig. 3) during the latter half of Phase II and into Phase III. An increase was found over all of China during Phase II, but

	Phase I (2001–2005)		Phase II (2006–2010)		Phase III (2011–2015)		
	Mean	Trend	Mean	Trend	Mean	Trend	
Ratio (µ	$(eq \mu eq^{-1})$						
China	0.29 (N = 35)	-3.3 ± 2.7	0.37 (N = 31)	$+14.8 \pm 1.9^{c}$	0.46 (N = 28)	$+10.1 \pm 3.8^{a}$	
Korea	0.62 (N = 11)	-4.9 ± 1.9^{a}	0.66 (N = 15)	$+13.6\pm4.7^a$	0.86 (N = 15)	$+3.9\pm2.6$	
Japan	0.58 (N = 44)	-0.3 ± 1.9	0.60 (N = 56)	$+3.6\pm1.3^{a}$	0.67 (N = 56)	$+2.5\pm1.3$	

Table 3. Statistical analysis of average values and linear trends for $NO_3^- / nss-SO_4^{2-}$ concentration in precipitation (*Ratio*) in precipitation over China, Korea, and Japan during phases I, II, and III.

Note: total number of observation sites are shown in parentheses. Linear trends were estimated using linear regression and results are shown as rates (% yr⁻¹). Significance levels are indicated by ^a p < 0.05, ^b p < 0.01, and ^c p < 0.001, and lack of mark indicates a lack of significance. Analysis for Korea during Phase I was for the 2002–2005 period.

only a slight increase over central China and decreases over north and south China were observed during Phase III (right panels of Fig. 4a and b). Such changes in the NO_x / SO_2 emission ratio were highly correlated with the variation in Ratio over China. Ratio was almost 0.3 during Phase I and subsequently increased to 0.5 during Phase II, with a trend of $+14.8 \pm 1.9 \% \text{ yr}^{-1}$ (p < 0.001), and to approximately 0.4– 0.6 during Phase III, with a trend of $+10.1 \pm 3.8 \% \text{ yr}^{-1}$ (p < 0.05). The analyses of different site types for EANET observations over China (Fig. 5b; see Table 1 for classification) revealed that the mean values for urban sites and rural and remote sites were highly correlated, with a correlation coefficient of 0.65. In addition to the variation in Ratio based on EANET observations for urban sites and rural and remote sites, other observations through IMPACTS, BNU, and LAPC/IAP/CAS also indicated similar variation in Ratio. Based on the IMPACTS dataset, which covered southern China with extensive data in 2001–2003, Ratio in China during Phase I was approximately 0.2, which was much lower than the current (Phase III) status. Although NO₃⁻ and nss- SO_4^{2-} concentrations in precipitation were higher than those in the EANET observation dataset, observations at BNU revealed a Ratio of 0.31, which was consistent with EANET observation results. A coordinated observation network by LAPC/IAP/CAS, which operated for 3 years around Beijing, indicated an increasing trend during Phase II from 0.39 in 2008 to 0.59 in 2010. These results obtained from multiple observations reinforce the idea that Ratio observed in the EANET network is representative of the precipitation chemistry in China.

In our previous studies (Itahashi et al., 2014a, 2015), we highlighted the impact of the NO_x / SO_2 emission ratio in China on the *Ratio* over downwind countries, including Korea and Japan. NO_x / SO_2 emissions and the NO_2 / SO_2 column ratio in Korea showed complex variation. These parameters increased during Phase I, stayed essentially flat during Phase II, and increased slightly or remained flat during Phase III. It should be noted that the NO_x / SO_2 emission ratio and NO_2 / SO_2 column ratio were not correlated in some parts of Korea. In particular, a flat trend was ob-

served for the NO_x / SO₂ emission ratio from 2006 to 2007, with a decline in the NO₂ / SO₂ column ratio; later, there was a flat trend for the NO_x / SO₂ emission ratio from 2010 to 2011, with an increase in the NO₂ / SO₂ column ratio. These discrepancies are attributed to complications from using the SO₂ column as a proxy for SO₂ emissions over Korea. *Ratio* over Korea was almost 0.6 during Phase I, with a trend of $-4.9 \pm 1.9 \% \text{ yr}^{-1}$ (p < 0.05). This subsequently increased to 0.8 during Phase II, with a trend of $+13.6 \pm 4.7 \% \text{ yr}^{-1}$ (p < 0.05), and to around 0.8–1.0 during Phase III, with a trend of $+3.9 \pm 2.6 \% \text{ yr}^{-1}$ (not significant). Considering the NO_x / SO₂ emission ratio during Phase II, the variation in *Ratio* in Korea may be connected to the variation in *Ratio* in China.

In Japan, both the NO_x / SO₂ emission ratio and NO₂ / SO₂ column ratio exhibited declining trends during the 15-year period. However, the long-term variation in *Ratio* did not show a decreasing trend. The *Ratio* over Japan was almost 0.6 during Phase I, with a trend of $-0.3 \pm 1.9 \% \text{ yr}^{-1}$ (not significant), subsequently increased to 0.7 during Phase II, with a trend of $+3.6 \pm 1.3 \% \text{ yr}^{-1}$ (p < 0.05), and remained level at approximately 0.7 during Phase III, with a trend of $+2.5 \pm 1.3 \% \text{ yr}^{-1}$ (not significant). Considering the continuous decline in the NO_x / SO₂ emission ratio during the 15-year period, the variation in *Ratio* in Japan appears to be related to the variation in *Ratio* in China.

The relationships between *Ratio* over China and *Ratio* over Korea or Japan are displayed in a scatterplot (Fig. 6). These data clearly illustrate that the *Ratio* over China and Korea and the *Ratio* over China and Japan dramatically changed over time. Along with the variation in the NO_x / SO₂ emission ratio and NO₂ / SO₂ column ratio in China, *Ratio* in precipitation over China was flat, then increased, and then returned to a flat trend during the 15-year period. The variation over China was related to patterns of variation over Korea and Japan, with correlation coefficients of 0.84 and 0.81, respectively (p < 0.001). These results revealed the impact of emission changes in China and the variation in precipitation chemistry on a regional scale in East Asia.



Figure 6. Scatterplots of (a) the $NO_3^- / nss-SO_4^{2-}$ concentration in precipitation (*Ratio*) in China and Korea and (b) those in China and Japan. Each circle indicates the annual mean *Ratio*, and the color indicates the year.



Figure 7. Long-term temporal variation in the (**a**) NO_3^- wet deposition (kg N ha⁻¹), (**b**) nss- SO_4^{2-} wet deposition (kg S ha⁻¹), and (**c**) sum of NO_3^- and nss- SO_4^{2-} wet deposition (keq ha⁻¹) with the NO_3^- fraction during 2001–2015 over (**i**) China, (**ii**) Korea, and (**iii**) Japan. One standard deviation across observation sites is indicated by the shaded area in (**a**, **b**). The numbers at the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).

3.3 Long-term trend in wet deposition

Finally, we focused on wet deposition, which was calculated by multiplying the chemical concentration in precipitation with the precipitation amount. Monthly data were checked via an outlier test for the chemical concentration, and the annual accumulated wet deposition was computed for years in which there was at least 9 months of coverage. The long-term temporal variation in NO_3^- and nss- SO_4^{2-} wet deposition over China, Korea, and Japan are shown in Fig. 7. Statistical analyses of averages values and trends during phases I, II, and III are summarized in Table 4. NO_3^- wet deposition amounts were approximately 6, 4, and 3 kg N ha⁻¹ over China, Korea, and Japan, respectively, and the year-to-year variation was high. In a statistical analysis of the NO_3^- wet deposition amount, there were no significant trends for China, Ko-

Phase I (2001–2005)		Phase I	Phase II (2006–2010)		Phase III (2011–2015)			
Mean	Trend	Mean	Trend	Mean	Trend			
NO_3^- wet deposition (kg N ha ⁻¹)								
5.88	-5.9 ± 4.6	5.90	$+4.4 \pm 4.5$	5.53	$+4.9 \pm 3.2$			
4.10	$+2.4 \pm 7.5$	3.76	$+8.2 \pm 7.3$	3.95	-8.3 ± 4.8			
2.92	-0.3 ± 2.7	3.23	-0.4 ± 1.5	3.17	$+0.2\pm0.3$			
nss-SO ₄ ²⁻ wet deposition (kg S ha ⁻¹)								
28.80	$+4.9 \pm 7.4$	24.57	-9.6 ± 3.4^a	16.90	-8.8 ± 4.3			
8.36	$+7.0 \pm 8.0$	8.26	-1.1 ± 6.3	6.01	-16.7 ± 5.6^a			
5.88	$+1.0\pm5.0$	6.32	-3.5 ± 0.9^a	5.68	-3.2 ± 2.5			
	Phase I Mean et deposi 5.88 4.10 2.92 2 ²⁻ wet de 28.80 8.36 5.88	Phase I (2001–2005)MeanTrendet deposition (kg N ha ⁻¹) $5.88 - 5.9 \pm 4.6$ $4.10 + 2.4 \pm 7.5$ $2.92 - 0.3 \pm 2.7$ $2^{2^{-}}$ wet deposition (kg S I) $28.80 + 4.9 \pm 7.4$ $8.36 + 7.0 \pm 8.0$ $5.88 + 1.0 \pm 5.0$	Phase I (2001–2005)Phase IIMeanTrendMeanet deposition (kg N ha ⁻¹) $5.88 -5.9 \pm 4.6$ 5.90 $4.10 +2.4 \pm 7.5$ 3.76 $2.92 -0.3 \pm 2.7$ 3.23 $2^{2^{-}}$ wet deposition (kg S ha ⁻¹) $28.80 +4.9 \pm 7.4$ 24.57 $8.36 +7.0 \pm 8.0$ 8.26 $5.88 +1.0 \pm 5.0$ 6.32	Phase I (2001–2005)Phase II (2006–2010)MeanTrendMeanTrendet deposition (kg N ha ⁻¹) $5.88 -5.9 \pm 4.6$ $5.90 +4.4 \pm 4.5$ $4.10 +2.4 \pm 7.5$ $3.76 +8.2 \pm 7.3$ $2.92 -0.3 \pm 2.7$ $3.23 -0.4 \pm 1.5$ $2^{2^{-}}$ wet deposition (kg S ha ⁻¹) $28.80 +4.9 \pm 7.4$ $24.57 -9.6 \pm 3.4^{a}$ $8.36 +7.0 \pm 8.0$ $8.26 -1.1 \pm 6.3$ $5.88 +1.0 \pm 5.0$ $6.32 -3.5 \pm 0.9^{a}$	Phase I (2001–2005)Phase II (2006–2010)Phase IMeanTrendMeanTrendMeanet deposition (kg N ha ⁻¹) $5.88 -5.9 \pm 4.6$ $5.90 +4.4 \pm 4.5$ 5.53 $4.10 +2.4 \pm 7.5$ $3.76 +8.2 \pm 7.3$ 3.95 $2.92 -0.3 \pm 2.7$ $3.23 -0.4 \pm 1.5$ 3.17 $2^{2^{-}}$ wet deposition (kg S ha ⁻¹) $28.80 +4.9 \pm 7.4$ $24.57 -9.6 \pm 3.4^{a}$ 16.90 $8.36 +7.0 \pm 8.0$ $8.26 -1.1 \pm 6.3$ 6.01 $5.88 +1.0 \pm 5.0$ $6.32 -3.5 \pm 0.9^{a}$ 5.68			

Table 4. Statistical analysis of average values and linear trends for NO_3^- and $nss-SO_4^{2-}$ wet deposition over China, Korea, and Japan during phases I, II, and III.

Note: linear trends were estimated using linear regression and results are shown as rates ($\% \text{ yr}^{-1}$). Significance levels are indicated by ^a p < 0.05, ^b p < 0.01, and ^c p < 0.001, and lack of a mark indicates a lack of significance. Analysis for Korea during Phase I was for the 2002–2005 period.

rea, and Japan during each phase (Fig. 7a). With respect to reactive nitrogen (Nr) deposition, the threshold value of $10 \text{ kg N} \text{ ha}^{-1}$ has been established (e.g., Bleeker et al., 2011). In the 15-year long-term analysis, the wet deposition of NO_{2}^{-} accounted for more than half of this threshold Nr deposition in China and about one-third in Korea and Japan.

The nss- SO_4^{2-} wet deposition exhibited decreasing trends during Phase II and III. The amount over China was around $30 \text{ kg S} \text{ ha}^{-1}$ during Phase I and below $20 \text{ kg S} \text{ ha}^{-1}$ during III, and the amounts over Korea and Japan decreased to about 2 and 1 kg S ha⁻¹, respectively. Based on a statistical analysis, nss- SO_4^{2-} wet deposition showed increasing (nonsignificant) trends during Phase I and decreasing trends during phases II and III. Decreasing trends over China and Japan during Phase II and over Korea during Phase III were significant (p < 0.05) (Fig. 7b). It should be emphasized that precipitation over China, Korea, and Japan during Phase II and over China during Phase III increased. Our results indicate that the decreasing trends in the nss- SO_4^{2-} wet deposition were caused by a strong decline in the nss-SO₄²⁻ concentration in precipitation, which counteracted the increase in overall precipitation. The reduction in SO₂ emissions over Korea and Japan might partially contribute to the decline in the nss- SO_4^{2-} concentration in precipitation. However, taking into account the correlation between precipitation chemistry over East Asia and emission changes in China, the SO₂ emission reduction in China after 2005-2006 strongly impacted both local deposition and downwind deposition.

To consider the impact of wet deposition from NO_3^- and nss-SO₄²⁻, the sum was calculated and is shown in Fig. 7c as an equivalent (eq) basis; additionally, the fraction of this sum attributed to NO₃⁻ wet deposition was also analyzed (expressed as a percentage). For the 15-year analysis period, it has been suggested that $nss-SO_4^{2-}$ wet deposition is reduced

in East Asia. The results indicated that the fraction of NO_{3}^{-} wet deposition increased gradually during the 15-year period over East Asia, indicating that further attention should be paid to the deposition of NO_3^- and related Nr species. These results are consistent with those of previous analyses of several cities in China, including Nanjing (Tu et al., 2005), Shenzhen (Huang et al., 2008), Guangzhou (Fang et al., 2013), the Pearl River Delta region (Lu et al., 2015), Tieshanping, which is one of the IMPACTS sites (Yu et al., 2017), and Beijing (Wang et al., 2012). Our study reinforced these inferences on a larger scale over East Asia. Excess Nr deposition will result in eutrophication, which normally reduces biodiversity. Transboundary nitrogen air pollution and deposition have been found over East Asia (Itahashi et al., 2016, 2017b). To reduce their impact, further understanding of NH₃ emissions will be important to develop a multipollutant control strategy (Zhao et al., 2009). In China, the pH of precipitation has not been lower than those of Korea and Japan due to substantial buffering species (Duan et al., 2016); therefore, base cations should be included in forthcoming analyses.

4 Summary and future perspectives

In this study, the chemistry of precipitation in East Asia in 2001–2015 was evaluated, with a focus on the behavior of *Ratio* (defined as the NO_3^- / nss- SO_4^{2-} concentration in precipitation). The monitoring networks over East Asia were initially used, and the study was augmented by intense observation campaigns over southern and northern China. NO2 and SO₂ column density satellite observations were used as proxies for NO_x and SO_2 emissions, and this was a key update from our previous studies (Itahashi et al., 2014a, 2015). The results for the NO₃⁻ concentration in precipitation suggested

that there is no clear relation between NO_x emissions and corresponding trends in China, Korea, and Japan. In comparison, the nss-SO₄²⁻ concentration in precipitation was partially correlated to the SO₂ emissions in China, and a corresponding (but nonsignificant) decline in Korea and Japan was also found. The analysis of Ratio clarified the trend (i.e., flat, followed by an increase and subsequent flattening) during the 15-year period in China, Korea, and Japan, and this variation was correlated to the NO_x / SO_2 emission ratio and the NO_2 / SO_2 column ratio in China. The initial flat trend was due to the increases in both NO_x and SO_2 emission in China, the subsequent increase was caused by the increase in NO_x emissions and the decrease in SO_2 emission in China, and the eventual flat trend was due to decreases in both NO_x and SO_2 emission in China. Due to the impact of upwind regions, it is difficult to use the SO₂ column as an accurate proxy of SO₂ emissions in Korea. Despite the increasing trends in the precipitation amount, decreasing trends for nss- SO_4^{2-} wet deposition over China were observed after 2005– 2006 and matched reductions in SO₂ emissions from China. This decrease in China triggered the decreases in $nss-SO_4^{2-}$ wet deposition over Korea and Japan. Both nitrogen and sulfur compounds pose acidification risks in ecosystems via atmospheric deposition processes. During the 15-year period analyzed in this study, sulfur wet deposition in East Asia decreased. This reduction seems to be important for future patterns, and additional studies focused on nitrogen are needed. Although this study was focused on wet deposition, a synergetic analysis of dry deposition processes is also needed to understand the impact of total wet and dry deposition on East Asian ecosystems.

Data availability. All of the data used in this study is publicly available. The data set of the ground-based and satellite observations and emission inventories are available from each reference as described in Sect. 2. For further information, please contact Syuichi Itahashi at the Central Research Institute of Electric Power Industry via email (isyuichi@criepi.denken.or.jp).

Author contributions. SI performed the analysis of observation, emission inventory, and satellite data and prepared the manuscript with contributions from all coauthors. KY and IU conducted the satellite data analysis and discussed the results. HH and SF measured precipitation at the Komae site and performed the analysis. YP and YW measured precipitation at 10 observation sites around Beijing and performed the analysis.

Competing interests. The authors declare that they have no conflict of interest.

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