

This is a repository copy of A quantitative approach to fluvial facies models: Methods and example results.

White Rose Research Online URL for this paper: http://eprints.whiterose.ac.uk/80270/

Version: Accepted Version

Article:

Colombera, L orcid.org/0000-0001-9116-1800, Mountney, NP orcid.org/0000-0002-8356-9889 and McCaffrey, WD orcid.org/0000-0003-2895-3973 (2013) A quantitative approach to fluvial facies models: Methods and example results. Sedimentology, 60 (6). pp. 1526-1558. ISSN 0037-0746

https://doi.org/10.1111/sed.12050

Reuse

Items deposited in White Rose Research Online are protected by copyright, with all rights reserved unless indicated otherwise. They may be downloaded and/or printed for private study, or other acts as permitted by national copyright laws. The publisher or other rights holders may allow further reproduction and re-use of the full text version. This is indicated by the licence information on the White Rose Research Online record for the item.

Takedown

If you consider content in White Rose Research Online to be in breach of UK law, please notify us by emailing eprints@whiterose.ac.uk including the URL of the record and the reason for the withdrawal request.



A quantitative approach to fluvial facies models: methods and example results

Luca Colombera^{1,2}, Nigel P. Mountney¹, William D. McCaffrey¹

- 1 Fluvial & Eolian Research Group, School of Earth and Environment, University of Leeds, LS2 9JT, UK.
- 2 corresponding author, Email eelc@leeds.ac.uk

ABSTRACT

Traditional facies models lack quantitative information concerning sedimentological features: this significantly limits their value as references for comparison and guides to interpretation and subsurface prediction. This paper aims to demonstrate how a relational-database methodology can be used to generate quantitative facies models for fluvial depositional systems. This approach is employed to generate a range of models, comprising sets of quantitative information on proportions, geometries, spatial relationships and grain sizes of genetic units belonging to three different scales of observation (depositional elements, architectural elements and facies units). The method involves a sequential application of filters to the knowledge base that allows only database case studies that developed under appropriate boundary conditions to contribute to any particular model. Specific example facies models are presented for fluvial environmental types categorized on channel pattern, basin climatic regime and water-discharge regime; the common adoption of these environmental types allows a straightforward comparison with existing qualitative models. The models presented here relate to: (i) the large-scale architecture of single-thread and braided river systems; (ii) meandering sub-humid perennial systems; (iii) the intermediate- and small-scale architecture of dryland, braided ephemeral systems; (iv) the small-scale architecture of sandy meandering systems; (v) to individual architectural features of a specific sedimentary environment (a terminal fluvial system) and its subenvironments (architectural elements). Although the quantification of architectural properties represents the main advantage over qualitative facies models, other improvements include the capacity: (i) to model on different scales of interest; (ii) to categorize the model on a variety of environmental classes; (iii) to perform an objective synthesis of many real-world case studies; (iv) to

include variability- and knowledge-related uncertainty in the model; (v) to assess the role of preservation potential by comparing ancient- and modern-system data input to the model.

Keywords: facies models, fluvial architecture, quantitative sedimentology, channel pattern, discharge regime, basin climate.

INTRODUCTION

Background

The primary purpose of facies models is to provide a "general summary of a specific sedimentary environment" (Walker 1984), in terms of its characteristic sedimentary features. The descriptive characters of facies models are obtained by combining results from studies of both modern systems and ancient successions preserved in the rock record. The general validity of a facies model stems from the process of "distillation" by which the sedimentary features observed in many real-world examples are synthesized to develop the model; the expected generality of a facies model makes it suitable to be considered as a norm for comparison, a basis for interpretation, a guide for future observations and a predictor in new geological situations (Walker 1984).

It is commonly argued that the possible value of the facies modelling approach for the purposes claimed by Walker (1984) appears to be limited by a number of shortcomings (Hickin 1993; North 1996; Miall 1999; Reading 2001). Firstly, facies models are often based on data derived from very few or single case studies (cf. models of Miall 1996; Lunt et al. 2004; Fielding et al. 2009; Horn et al. 2012), and as such might be biased in the sense that they reflect the limited experience of individuals or research groups, whose work is often concentrated on particular geographical areas (Reading 2001). Furthermore, there exists a tendency to derive models for single field examples or for very specific categories of fluvial system such that the resultant model is excessively specialized to the extent that it is of little use as a predictive tool beyond the scope of the original study example; in such cases, the proposed model may obscure the underlying unity of the systems in order to preserve their uniqueness (Dott & Bourgeois 1983; Miall 1999). A major limitation of traditional facies models is that the degree of generality of such models in their current form is not adjustable to the particular

needs of a geologist attempting to apply the model to a new situation or dataset. Another problem relates to how the process of distillation is actually carried out: given that the process of synthesis is expected to be subjective, how can it be possible to ensure that different authors equally and objectively include the fundamental patterns and exclude accessory detail in developing their models? Also, the inclusion of some form of mechanism for the evaluation of the uncertainty ("any departure from the unachievable ideal of complete determinism" according to Walker et al. 2003) associated with developed models has not been attempted to date (Hickin 1993); it can be argued that the proliferation of categories on which facies models are classified is an endeavour to ensure that the variability between systems can be perceived. It is therefore important to devise a way to consider uncertainty (i) by measuring the variability between different systems that are classified on the basis of similar conditions and therefore represented by the same model, and (ii) by assessing the limitations and deficiencies in our knowledge of those systems. However, the most notable drawback of traditional facies models lies in their qualitative nature, as the lack of quantitative information seriously limits their predictive value (North 1996). In subsurface prediction problems it is common to combine qualitative, conceptual information about the type of sedimentary heterogeneities and their distribution with quantitative geometrical information derived from supposed outcropping analogues. Quantitative information on the geometry of sedimentary units is commonly stored in quantitative databases that serve to provide input to deterministic and stochastic subsurface models (e.g. Bryant & Flint 1993; Cuevas Gozalo & Martinius 1993; Dreyer et al. 1993; Robinson & McCabe 1997; Reynolds 1999; Eschard et al. 2002; Tye 2004); the collation of such geometrical data – as derived from a variety of case histories - combined with the classification of system parameters, permits the derivation of sets of quantitative information through a process of synthesis, as advocated by Walker (1984). One approach of this kind has been applied to fluvial systems for obtaining descriptions of channel geometries by Gibling (2006). However, facies models are not merely geometrical descriptions of a depositional system; thus, some databases have been designed to better describe spatial relationships between genetic units, for example by including summary transition statistics for deep-water genetic-unit types (Baas et al. 2005), by specifying patterns of spatial distribution for carbonate genetic-unit types (Jung & Aigner 2012), or by digitizing the spatial relationships between

individual fluvial genetic units (Colombera et al. 2012a). Also, efforts have been made to implement such systems to variably investigate the internal organization of sedimentary units (Baas et al. 2005; Colombera et al. 2012a; Jung & Aigner 2012).

Aims

The aim of this paper is to demonstrate how a database approach to the description and classification of fluvial sedimentary systems can be used to improve facies models as a benchmark for research purposes and as a tool for subsurface prediction. Whereas some techniques adopted in the study of sedimentary geology are inherently quantitative (e.g. numerical and physical modelling, sandbodygeometry quantification), facies modelling is still typically qualitative in nature. The aim is to show how the innovation in the approach lies essentially in the systematic quantification of observations and interpretations, which permits a more rigorous description and classification of architectural styles of fluvial systems. An important, broad-reaching implication for the understanding of the stratigraphic record is that the proposed approach, if used to carry out comparative studies, can be applied to deduce the relative influence of boundary conditions and potential overriding controls for given depositional contexts. Specific objectives of this paper are as follows: (i) to discuss the process of synthesis by which partial information from individual case studies is merged into a model and how this process is implemented in practical terms for different types of information, which concern the geometry, internal organization and spatial relationships and distribution of genetic units; (ii) to illustrate, through a range of example database-derived quantitative depositional models for different fluvial systems, that this database-driven quantitative approach to the development of facies models can assist in overcoming the above-mentioned problems inherent in traditional qualitative approaches.

DATABASE AND METHOD

Database structure and building blocks

Overview of FAKTS database schema

The Fluvial Architecture Knowledge Transfer System (*FAKTS*) is a database comprising field- and literature-derived quantitative and qualitative data relating to the architecture of both modern rivers

and ancient successions (Colombera et al. 2012a). Genetic units included in the database are equally recognizable in both the stratigraphic and geomorphic realms and belong to three hierarchies of observation (Fig. 1): depositional elements, architectural elements and facies units, in order of descending scale. The geometry of the genetic units is characterized by dimensional parameters describing their extent in the vertical, strike-lateral and downstream directions, relative to the channelbelt-scale (palaeo-) flow direction (thickness, width and length). The relations between genetic units are stored by recording and tracking (i) the containment of each unit within its higher-scale parent unit (e.g. facies unit within architectural elements) and (ii) the spatial relationships between genetic units at the same scale, recorded as transitions along the vertical, cross-gradient and downstream directions. Additional attributes are defined to improve the description of specific units (e.g. braiding index, sinuosity value, bank-full depth and width for channel complexes, grain-size curves for facies units), whereas accessory information (e.g. ichnological or pedological characters) can also be stored for every unit within open fields. The database also stores statistical parameters referring to genetic-unit types, as literature data is often presented in this form. Each genetic unit or set of statistical parameters belongs to a stratigraphic volume called a subset; each subset is a portion of the total dataset characterized by given attribute values, such as system controls (e.g. subsidence rate, basin type, climate type) and system-descriptive parameters (e.g. river pattern, distality relative to other subsets). For each case study of fluvial architecture, FAKTS also stores metadata describing, for example, the methods of data-acquisition employed, the chronostratigraphy of the studied interval and the geographical location. A threefold data-quality ranking system is also implemented with the purpose of rating datasets and genetic units (as A, B or C level, in order of decreasing quality). A more detailed description of the FAKTS database schema is given in Colombera et al. (2012a); for the purposes of this work, the key focus is on the adopted classifications of geological entities, described in the following paragraphs, as they are the building blocks of the quantitative facies models being developed.

Classification of bounding surfaces

The subdivision of fluvial successions into genetic packages through recognition, classification and numbering of hierarchically-ordered sets of bounding surfaces is a common sedimentological practice (Allen 1983; Miall 1988; 1996; Holbrook 2001). FAKTS permits specification of the order of bounding surfaces corresponding to the basal surface of depositional elements (highest order in case of composite surfaces) and the order of surfaces across which architectural-element or facies-unit transitions occur. FAKTS classifies bounding surfaces according to the popular hierarchical classification scheme proposed by Miall (1988; 1996), whereby surface-orders are assigned on the basis of observable characters (e.g. lateral extension, erosional or accretionary character), but are also interpretative in nature. Attribution of order (i.e. rank) to bounding surfaces is difficult in many real-world situations (Bridge 1993) and therefore has uncertainty associated with it; however, it is worthwhile to tentatively rank bounding surfaces according to a series of hierarchical orders, so as to be able to capture architectural features and changes associated to surfaces with genetic significance and often temporal and spatial relevance. Whenever observable elements on which to base the attribution of a given bounding-surface order are lacking, corresponding database fields are left undefined.

Classification of depositional elements

The general approach to the segmentation of alluvial architecture at the largest scale involves picking and indexing channel bodies, then dividing the remaining non-channelized floodplain bodies into discrete objects that are juxtaposed to the channel bodies in a spatially coherent way. Large-scale depositional elements are then classified as *channel-complexes* or *floodplain* segments on the basis of the origin of their deposits, and are distinguished on the basis of geometrical rules. The application of these rules is generally flexible, as the criteria devised for the definition of these objects may sometimes be difficult to apply due to limitations brought about by the possible lack of data of either a geometrical or geological nature (e.g. 3D channel-body geometries, recognizable internal bounding surfaces): such difficulties are recorded by data-ranking, data-type and target-scale attributes. In addition, the geometrical criteria cannot be followed altogether for cases where data are derived from published works presenting only summary results (e.g. from works presenting plots of dimensional

parameters of channelized bodies and no reproduction of the original 2D or 3D dataset from where the data were originally derived); this form of uncertainty is recorded by a data-ranking attribute.

General criteria followed for depositional-element subdivision are presented below. The choice of interpretative units at this scale is justified by the fact that the recognition of channel and floodplain segments is possible for virtually any depositional system interpreted as being fluvial in origin (cf. Miall 1996; Bridge 2006; and references therein).

Channel complex

Each stratigraphic volume that can be characterized at the depositional-element scale is firstly segmented into channel-complexes; the aforementioned set of geometrical criteria needs to be followed to distinguish individual units among channelized deposits that are complexly juxtaposed and/or interfingered with floodplain deposits. Such criteria consider geometrical change across the channel-cluster vertical extension, taking into account the interdigitation of floodplain deposits, mode and rate of change in the lateral extension of contiguous channel deposits along the vertical direction, and existence of lateral offsets where channel-bodies are vertically stacked (cf. Cuevas Gozalo & Martinius 1993). Whenever geological knowledge permits the lateral tracing of important erosional surfaces (possibly associated with high palaeo-relief), it is possible to adopt such surfaces as depositional-element bounding surfaces. When dealing with subsurface case studies, the approach is usually purely geometrical. Due to the way they are defined, channel complexes simply represent genetic bodies interpreted as having been deposited in a channelized context and encased by floodplain deposits: in geological terms they could still span a rather wide range of hierarchical orders (e.g. distributary channel-fills, channel-belts, valley-fills); definition in this way attempts to minimize interpretation, thereby still ensuring the possibility for the analysis of channel clustering in different depositional settings.

Floodplain

The subdivision of floodplain segments takes place subsequent to channel-complex assignment, such that the remainder of the stratigraphic volume is broken down into floodplain packages that are

referable as neighbouring bodies (either lateral or vertical) to each channel-complex. Thus, floodplain depositional elements simply represent geometrical genetic bodies interpreted as deposited by out-of-channel floods (cf. Miall 1996; Bridge 2006).

Classification of architectural elements

FAKTS' architectural elements are defined as components of a fluvial depositional system with characteristic facies associations that are interpretable as sub-environments. Also for these genetic units, it is not possible to separate descriptions from interpretations, as unit types are fundamentally interpretative. The attribution of a particular element type follows the criteria proposed by Miall (1985, 1996): the elements are interpreted on the basis of the characters of their bounding surfaces, their geometry, scale, and internal organization. However, FAKTS' architectural element types differ significantly from the ones included in Miall's (1985, 1996) schemes: additions and deductions strive to provide a more interpretative classification scheme containing mutually-exclusive classes that are consistent in terms of geomorphological expression, in order to make it easier to include datasets from modern rivers; an analogous attempt to define the basic geomorphic building blocks of fluvial systems was proposed by Brierley (1996). Importantly, FAKTS' architectural-element types correspond to classes of sub-environments that are commonly recognized in both the stratigraphic record and in modern rivers alike (cf. Bridge 2006), and are conveniently chosen to represent variability in sedimentary architecture.

Architectural-element types may differ from each other on just geometrical/geomorphological characters (e.g. downstream-accreting barforms from laterally-accreting barforms, crevasse splays from levees) or interpreted dominant processes (e.g. sandy aggradational floodplain from floodplain fines, abandoned channel-fill from aggradational channel-fill). The essential diagnostic characteristics of each interpretative architectural-element type are included in Table 1. In addition to the features summarized in Table 1, other characteristics concerning the geometry, internal organization, and reciprocal spatial relationships may have also been considered by the authors whose studies were incorporated into FAKTS to reach their interpretations.

Classification of facies units

According to the classification of bounding surfaces proposed by Miall (1985; 1996) and adopted in the FAKTS database, 2nd-order surfaces can be traced where a change in lithofacies or palaeocurrent are observed; on this basis, facies units represent genetic packages that are bounded by second- or higher-order bounding surfaces and are characterized by given textural and structural properties. Such genetic units are considered as corresponding to the 2nd-order units of Miall (1985; 1996) and to the microscale to mesoscale stratasets of Bridge (1993). These units are based on observable characteristics and represent more objective units than depositional and architectural elements.

As each unit is primarily classified according to the codes provided in the original works, a detailed description of grain size is optionally stored for each unit in the database field containing the original coding. The grain-size characterization given by the FAKTS' facies-unit classes is instead very limited, as the FAKTS' facies classification scheme largely follows the scheme proposed by Miall (1977; 1978; 1996), although with some additions. The adoption of this scheme has some advantages. Firstly, the use of few mutually exclusive classes is good for database use, as a more detailed description of grain size in the code could generate a high number of classes to account for all possible grain-size modalities and tails, so that description of textures that are originally less detailed (e.g. following Miall's scheme) would not be easily translated. Secondly, as many authors have adopted the Miall scheme (1977; 1978; 1996), use of this scheme (albeit in a slightly modified form) negates the requiredment to translate similar facies codes described in many case studies as they are incorporated into the database. So, although FAKTS' lithofacies coding – as well as the original facies codes of Miall (1978; 1996) - could be improved to better account for textural and structural variability, the use of a classification scheme that is well established in the scientific community is especially well-suited for database use, because for many published case examples, lithology classifications do not need to be re-coded. Nevertheless, caution must be exercised when translating original lithology data. For example, there is no consensus on the definition of matrix: the American Geological Institute defined the matrix as the "finer-grained, continuous material enclosing, or filling the interstices between the larger grains or particles of a sediment or sedimentary rock" (Gary et al. 1974). Thus, gravel-grade sediment acting as matrix could still be consistent with this definition.

However, the inclusion of clean sand- or gravel-grade deposits (cf. Shultz 1984; Sohn et al. 1999; for alluvial examples) into the definition of matrix precludes the differentiation of lithofacies associated to fundamentally different formative processes: therefore, for data entry into the FAKTS database, matrix is defined as being dominantly fine grained (clay + silt), possibly partially sandy, roughly in agreement with Miall (1996). Thus, care must be taken as the same code could be used by different authors to designate deposits that would be classified differently in the FAKTS database system.

In contrast to the approach taken to the classification of architectural elements, properties concerning the geometry or the bounding surfaces of facies units are only occasionally important for their definition (e.g. facies type Ss): facies-unit types are usually only designated on the textural and structural characteristics of the deposits. There is no scope for provision of a rich and detailed description of each facies-unit type here, as their accessory sedimentological characteristics may vary widely among the different fluvial systems included in the database. Instead, only a summary of the essential features of each of the 25 types is given, in Table 2.

Each facies-unit type may be associated with more than one genetic process, with more than one bedform type, and with variable flow regime: refer to Miall (1978; 1996) and Bridge (1993) for explanations of the genetic significance of these lithofacies types. Notably, several alternative classification schemes could be implemented into the database structure in addition to those of the original authors' and FAKTS' facies codes, possibly separating textural and structural data in different fields.

An approach to building quantitative facies models: practical considerations

As of September 2012, FAKTS comprised 111 case histories – defined as individual sedimentological studies on a particular river or succession, by specific authors – and included data referring to 4285 classified depositional elements, 3446 classified architectural elements, and 20101 classified facies units, as well as additional statistical summaries referring to architectural properties of groups of genetic units. A summary of the case studies included in the database and of the published literature considered for derivation of primary data and for system classification is given in Table S1 (see supplementary material).

Through interrogation of the database, it is possible to obtain a multi-scale quantitative characterization of the sedimentary architecture of fluvial systems primarily consisting of three types of information (Colombera et al. 2012a), respectively concerning: (i) the internal organization of genetic units and stratigraphic volumes; (ii) the geometry of genetic units; (iii) the spatial relationships between genetic units. This section discusses some issues on how to best incorporate this information within quantitative facies models by synthesizing different case studies; in particular, it is important to identify which (if any) data types might be biased, for example by under-sampling, and to specify how the integration of data from multiple scales can be achieved in practice.

At the outset, subsets should be filtered according to their suitability to given queries; this information is contained within metadata fields that specify: (i) what scales of observation (and relative orders of genetic units) each subset is focussed on; (ii) the type(s) of output that it is possible to derive from a subset (i.e. proportions and/or dimensional parameters and/or transition statistics and/or grain-size information).

A first-order description of the internal organization of genetic units or stratigraphic volumes is given by the proportion of lower-order genetic units forming them. Here, three approaches to compute such proportions are outlined.

- 1) A first approach involves computing genetic-unit-type proportions as based on the sum of all occurrences, or thicknesses, or products of dimensional parameters (e.g. thickness times width) of genetic units (cf. Fig. 2); a drawback of this approach is that case studies that have been studied more extensively for which more genetic units are recorded (e.g. datasets derived from the study of more extensive outcrops) are over-represented, resulting in a biased output that is unbalanced in favour of some case studies.
- 2) An alternative second approach is to compute genetic-unit-type proportions as based on the sum of genetic-unit percentage proportions (obtained as above) within each suitable subset, thereby obtaining corrected proportions that account for the fact that some case studies may have been studied less extensively than others (cf. Fig. 2); the principal drawback of this approach is that case studies that have been studied in only modest detail for which relatively few genetic units have been classified (e.g. datasets derived from the study of less extensive

- outcrops) are over-represented, resulting in a biased output in which some genetic-unit types are under-sampled.
- 3) In cases where the aim is to obtain unit-type proportions within genetic units that do not belong to the immediately higher scale (i.e. to derive proportions of facies-unit types composing depositional elements, or proportions of architectural-element or facies-unit types within stratigraphic volumes), it is possible to compute proportions that are weighted according to the proportions of the intermediate-scale units (cf. Fig. 2). For instance, an abundance of facies-unit types composing channel-complexes can be achieved based on a combination of facies-unit proportions forming each architectural element type with architectural-element proportions forming channel-complexes. As a specific example, if CH (aggradational channel-fill) architectural elements represent 50% of all channel-complexes and 20% of all CH elements are represented by facies unit St, it is straightforward to compute 10% as a model proportion of St within channel-complexes. Given that some case studies are focused on specific features of fluvial architecture, this approach would return more accurate proportions when scales are skipped. For example, if a case study is focussed on the facies architecture of LA (laterally-accreting barform) architectural elements, the relative facies-unit type proportions will not be an accurate description of the entire fluvial system, but of LA architecture only. Practically, constraining genetic-unit proportions to higher-scale geneticunit proportions would result in a more effective integration of observations at different scales. However, when obtaining proportions according to such an approach, it must be borne in mind that the result may be biased by not incorporating genetic relationships between different unit types. For example, if the aim is to derive the overall CS (crevasse splay) proportion for a model by integrating architectural-element-scale information from a case study in which the proportion of floodplain depositional element is 25% and in which CS elements constitute 20% of the floodplain (and therefore 5% of total volume), with depositional-element-scale information from a case study in which the proportion of floodplain is 50%, we would derive a proportion of CS within the model stratigraphic volume equal to 10%. In practical terms, this may not be realistic as the proportion of crevasse-splay

deposits may actually decrease with a decreasing proportion of channel-belt deposits, with which they are genetically related, instead of simply scaling with the proportion of floodplain depositional elements within which they are contained.

The uncertainty associated with quantitative descriptions of dimensional parameters of genetic units is partially intrinsic to the way dimensional data and metadata are stored: the width and length of a genetic unit are classified using categories of completeness of observation (*complete*, *partial*, or *unlimited*), as proposed by Geehan & Underwood (1993), whereas widths are classified as *apparent* when derived from sections oriented oblique to palaeocurrent directions; in addition, metadata qualifying the type of observations are included (e.g. outcrop extension, type of observations from which dimensional parameters are drawn). Inclusion of geometrical information in a model can lead to problems concerning over- or under-representation of specific case studies, which might also need to be confronted.

Database-informed quantitative facies models describe the spatial relationships between genetic units in each of the three directions (vertical, cross-stream, and upstream) by employing embedded transition statistics, with self-transitions (i.e. transitions between likewise-classified genetic units) considered admissible. When obtaining transition statistics, issues that are analogous to the ones related to the computation of proportions may be encountered, such as the integration of facies-unit transitions mapped from different architectural elements into a model of facies-unit transition statistics that refer to an ideal stratigraphic volume. Such problems could be tackled in a way that is entirely analogous to the approaches proposed for deriving proportions. It is also important to note that a system that allows filtering of transitions both on the bounding-surface order across which the transition occurs and on the genetic-unit type in which the transition occurs, permits the derivation of genetic-unit transitions referring to a variety of genetically-related stratigraphic packages (e.g. architectural-element transitions within channel-complexes, facies-unit transitions within 3rd-order packages contained in *LA* barforms), as envisioned by Godin (1991).

If Markov-chain analysis is attempted, two notable advantages are provided by the method the database employs to store the transition data. Firstly, because self-transitions are admissible they can be included in the Markov-chain analysis (cf. *multistory lithologies* of Carr et al. 1966), resulting in

improved independent random matrices (cf. Selley 1970; Schwarzacher 1975); this is a methodological advancement over many previously-published transition matrices containing predefined diagonal zeros (i.e. matrices that do not allow self-transitions; e.g. Gingerich 1969; Allen 1970; Miall 1973; Cant & Walker 1976), which cannot result from independent random processes (Goodman 1968; Schwarzacher 1975; Carr 1982). Secondly, the inclusion of bounding-surface information in Markov-chain analysis was advocated by Cant & Walker (1976) and Godin (1991): sorting on bounding-surface order it is possible to filter transitions on the likelihood of their genetic significance, for example by excluding erosional transitions between lithofacies (i.e. across bounding surfaces of a specified order). The necessity to incorporate variability-related uncertainty in a model can be partially tackled by quantifying the variability of architectural properties in each facies model, possibly exemplifying extreme values within the range of each property (e.g. maximum channelcomplex thickness, maximum LA proportion within any systems) by referring to real-world case studies. In addition, the implementation of a ranking system (Data Quality Index or DQI; cf. Baas et al. 2005; Colombera et al. 2012a) is employed to evaluate the quality and reliability of (i) datasets, for example by considering the type of data available; (ii) genetic-unit classification, by considering the type of observable attributes on which a class is attributed to a unit; (iii) system classification, for example by considering the reliability of proxies on which a class is attributed to a subset. Thus, uncertainty related to inadequate knowledge (rather than to the inherent variability of the system) can also be taken into account by associating to the model a measure of value that is proportional to the DQI's of the systems or units, and to the amount of data (number of systems and units) on which the model is based.

The process of synthesis (or distillation in the terminology of Walker, 1984) of the model, to which the issues presented above relate, is actually implemented only after performing the selection of the case studies or individual subsets whose parameters match with the ones chosen for the classification of the quantitative depositional model. Such a process of filtering may be performed on architectural features (e.g. choice of systems in which the thickness of gravel deposits exceed 50% of all measured thickness), descriptive-parameters (e.g. choice of systems classified as meandering), boundary conditions (e.g. choice of dryland systems), or on a combination of each (Fig. 3).

RESULTS: EXAMPLE MODELS

Large-scale architecture

The importance of including large-scale information in conceptual models of fluvial architecture has long been recognized, and such information has been included in models summarizing the distribution of channel and floodplain deposits in stratigraphic volumes (e.g. Allen 1965; Friend 1983). However, contrasting views have been expressed regarding the type of system parameters (external controls, frequency/velocity of autogenic processes, descriptive parameters) on which the categorization of the models should be based; for example, as to whether channel-pattern can actually be considered as a good predictor for large-scale organization (cf. Allen 1965; Bridge 1993). Here, large-scale models based on channel pattern are presented for single-thread and braided systems (Fig. 4). It is not the purpose of this study to assess what type of controls or control-dependent system parameters are most suitable for the categorization of models of large-scale fluvial architecture (cf. Miall 1980), but one aim is to explain how this approach can be potentially applied to solve this issue, as explained below. More generally, the main scope of this study is to show how the use of such database systems permit the generation of facies models through an objective process of synthesis, even though this does not mean that such models will necessarily be unbiased, as they will still be associated with uncertainty related to the interpretations of the systems from which the data were originally derived. These database-derived facies models describe large-scale fluvial architecture in terms of the proportions and geometries of channel-complex and floodplain depositional elements (Fig. 3 and 4). Separately computing genetic-unit type proportions for each stratigraphic volume (subset) is a sensible choice if the subset is large and few categories are included. As this is the case for subsets suitable for computing depositional-element proportions, it is then possible to quantify how proportions vary between volumes (Fig. 4a). Thus, it is possible, for example, to include information on the observed variability in channel density and geometry in the same end-member model: variability becomes part of the model, and there is no need to advocate alternative models to represent it. This also means that, ideally, the approach could be used for determining what classifications are most suitable for categorizing the models, by recognizing ensembles of categories that ensure maximum inter-type variability and minimum intra-type variability in quantities describing architectural styles.

Intermediate-scale architecture

Many traditional fluvial facies models provide a relatively detailed characterization of sedimentary architecture in terms of building blocks interpretable as sub-environments, reflecting their recognition in modern systems and the interpretation of preserved ancient facies assemblages (e.g. Galloway & Hobday 1983; Walker & Cant 1984; Miall 1985; 1996; Nadon 1994). FAKTS' architectural elements broadly match this level of detail: by querying the database, it is possible to derive quantitative information to be included in facies models describing intermediate-scale fluvial architecture in terms of the proportions, geometries and 3D spatial relationships of architectural elements (Fig. 5 and supplementary material S2). The results presented in Fig. 5 and 6 illustrate the generation of a facies model for dryland ephemeral braided systems by the application of multiple filters (based on categories of basin climate type, stream discharge regime and channel pattern type), as well as all the models resulting from intermediate filtering steps. In this case, because of the level of detail in model categorization (i.e. the number of filters), the ephemeral-river model (step 4) is built upon a limited number of systems and genetic units, thereby resulting in scant general value. Instead, the "arid to semiarid braided system" model (step 3) proposed here incorporates a far larger knowledge base, lending itself better to a discussion of its intermediate-scale architectural features. Mainly, ancient sandy systems were considered for the database-assisted creation of this model, including data from the Jurassic Kayenta Formation, USA (authors' field data; Miall 1988; Bromley 1991; Luttrell 1993; Stephens 1994; Sanabria 2001), from the Jurassic Morrison Formation, USA (Miall & Turner-Peterson 1989; Robinson & McCabe 1997; Kjemperud et al. 2008), from the Triassic Moenave Formation, USA (Olsen 1989), from the Triassic Sherwood Sandstone Group, UK (Steel & Thompson 1983; Cowan 1993), from the Miocene Vinchina Formation, Argentina (Limarino et al. 2001), from the Triassic Omingonde Formation in Namibia (Holzförster et al. 1999), and from the Permo-Triassic Balfour Formation, South Africa (Catuneanu & Elango 2001).

In agreement with other existing braided-river models (e.g. Allen 1965; Miall 1977; 1978; Cant 1982; Walker & Cant 1984; Nanson & Croke 1992), the resulting ideal braided dryland system is dominated by channel deposits because in-channel architectural elements represent over 75% by volume of the model, if only fluvial elements are considered (as in Fig. 5). As these architectural-element proportions are solely based on ancient-system data, it can be observed that the most frequently preserved product of in-channel deposition is represented by aggradational channel-fills, rather than horizontally-migrating barforms. It must be considered that this observation may not be indicative of the original geomorphic organization of channel-belts, as observed abundances may relate to channelfills having a higher preservation potential than barforms, to channel-deposit accretion directions not being discernable in all cases (for example because of inappropriate outcrop exposure and orientation, especially if surfaces dip at very low angle, cf. Bristow 1987), or to accretion surfaces not always being preserved in barform deposits (cf. Jackson 1978; Kraus & Middleton 1987) potentially resulting in deposits categorized as CH that include the product of the horizontal migration of barforms. Within the model, non-channelized deposits of high-energy sandy aggradational-floodplain elements (SF) appear to dominate over floodplain-fine elements (FF), with the former more often tending to stack on top of channel-fills and downstream-accreting barforms, and the latter more frequently developed on top of laterally-accreting barform elements. However, FF elements display the largest observed lateral extent among floodplain elements, some examples exceeding 1000 m in maximum observed width. Crevasse channels, splays, abandoned channels and levees represent only a volumetrically minor portion of the model floodplain, and the available transition statistics suggest a tendency for these elements to be associated with FF, rather than SF, floodplain elements. However, the model lacks features that are likely to be included in a qualitative model of a dryland braided system, such as dryland floodplain lakes, suggesting that the data employed to generate the model do not yet fully account for natural variability.

Small-scale architecture

Some facies models widely used for interpreting ancient systems are represented by vertical profiles summarizing fluvial styles – related to environmental categories – in terms of lithofacies occurrences,

proportions, typical thicknesses and vertical stacking (cf. Miall 1977; 1978; 1996). FAKTS permits the derivation of similar one-dimensional models, represented by proportions, thickness and vertical juxtapositional trends of facies units within system types, by performing an objective distillation of different case studies, as illustrated in Fig. S3 (see supplementary material): the inclusion of quantitative information relating to facies units may aid the interpretation of 1D subsurface data by making model comparison more objective. The approach can be generalized to include three-dimensional information: example results (Fig. 7 to 10) are again associated with the "dryland ephemeral braided system" model and with the models related to its intermediate filtering steps, to demonstrate the capability to generate multi-scale models.

As the "dryland ephemeral braided system" model currently comprises one fifth of all facies units included in the knowledge base (represented by the model at step 1), the model is richer in data than its intermediate-scale architectural-element-based counterpart, reflecting the fact that the database currently includes more data from lithofacies-scale-oriented studies than from architectural-elementscale studies, for this set of system boundary conditions. The proposed "braided dryland ephemeral" model is based on categories relying on concurrent interpretations of braiding, which requires recognition of contemporaneity in-channel activity, and of basin climate type and discharge regime, which require proxies and may refer to average conditions through time; although the quality of data and interpretations can be ranked, the possibility of including data from case studies whose environmental interpretations are incorrect increases with the number of filters applied and results must therefore be considered with care. However, the possibility to contrast this model with the ones resulting from intermediate-stage filtering serves the aim of demonstrating the capabilities of the database system in highlighting the peculiarities of the different models, in quantitative terms. For example, the "dryland ephemeral braided system" model includes case studies that collectively show a high abundance of sand-grade deposits, making this model comparable to Miall's (1985, 1996) sandy-river models 11 and 12. Compared to its intermediate-step models, the "dryland ephemeral braided system" model presented here does not show any significant increase in the proportion of Sh (horizontally bedded sandstone) and Sl (low-angle cross-bedded sandstone) lithofacies, which are often considered a diagnostic architectural feature of such systems, supposedly in relation to the influence of upper-flow regime processes associated with flash floods (Miall 1985; 1996). Instead, a comparison between the facies-unit proportions of the braided-system model (Fig. 8), and of the sandy meandering-system model (Fig. S3, see supplementary material) reveals that the proportion of *Sh* and *Sl* facies-units among sandy deposits are significantly higher in the former compared to the latter.

Facies models often contain information on individual genetic packages: models of this sort represent a tool for guiding the interpretation of lithosomes with characteristic facies associations as subenvironments, such as point bars (e.g. Allen 1970) or crevasses splays (e.g. Bridge 2003), which can be variably arranged in the rock record, thereby representing a reference to interpretations that can be flexibly applied to different fluvial environmental types. The facies architecture of lithosomes corresponding to FAKTS' depositional and architectural elements can be investigated to derive model proportions, geometries, grain-size and spatial relationships of facies units within them, as illustrated in Fig. 11 and S4 (see supplementary material). The examples shown demonstrate how basic features relating to the internal architecture of the lithosomes – such as the lack of conglomeratic beds, the dominance by flat-bedded sandstone, and the on average higher horizontal extent of the formative facies units characterizing sandy aggradational floodplain elements (Fig. S4) – can be highlighted through quantification.

Spatial and temporal evolution

Given that FAKTS stores architectural information relating to stratigraphic volumes that can be arranged in relative temporal and spatial frameworks, information on the temporal and spatial evolution of architectural features from individual case studies can be derived and included in quantitative facies models of fluvial systems. Quantitative comparative studies can be performed between different systems to investigate spatial and temporal trends with the aim being to derive models of architectural change, in terms of space and/or time. Figure 12 presents downstream changes in facies-unit proportions (cf. Miall 1977) for a modern system and an ancient system, both of which are believed to represent terminal fluvial fans, for which the identification of proximal, medial and distal fan zones is justifiable, although arbitrary rather than objective.

DISCUSSION

A database-driven method for the creation of quantitative fluvial facies models such as the one presented here has several advantages, as listed below.

- Most importantly, this approach satisfies the long-recognized need for inclusion of quantitative information in facies models (North 1996; Anderson et al. 1999; Lunt et al. 2004), improving the value of facies models as a reference for comparison, interpretation and subsurface prediction. For example, database-derived models can be used as quantitative synthetic analogues to subsurface systems with which to better inform stochastic structure-imitating simulations of sedimentary architecture (Colombera et al., 2012b).
- Although several alternative procedures can be followed for obtaining the same type of information, the process of synthesis by which information from the individual case studies is distilled into the model can be carried out objectively, and permits the preservation of local detail through incorporation of features with limited occurrence. The number of case studies and genetic units included will justify and quantify the model generality.
- Quantitative facies models generated by a database approach can be flexibly tailored on any system parameters and/or concurrent architectural properties (e.g. gravel-bed braided system), and any of the scales of observation considered can be included in the model (e.g. channel-complex distribution in an ideal alluvial basin, architectural-element distribution in a meandering-system model; lithofacies distribution in a model of a crevasse splay element), either individually or in the form of hierarchically-nested depositional products.
- As metadata concerning the quality of observations and interpretations can be stored in such a database, it is possible to include information about the uncertainty related to variability in data quality and data deficiency in the model. If all or at least all the most significant studies on the sedimentology of fluvial systems were included, the database could help identify gaps in current knowledge, in a way similar to the original intention of facies models (cf. Walker, 1984).

- The use of a database system permits inclusion of architectural variability as a character of the model, in contrast to traditional facies models. For example, Miall's models 11 and 12 (Miall 1985; 1996) are solely differentiated on the basis of architectural style, with the scope of including information on the variability of facies assemblages, despite the two model systems being categorized on non-mutually-exclusive classes. Instead, this database approach allows inclusion of information on the variability in sedimentary architecture into models classified on mutually-exclusive categories. This has implications for the recognition of the environmental categories that, by maximizing architectural variability between types and minimizing variability within types, are most suitable for facies-model classification.
- The inclusion of information that refers to interpretative system types and unit types (depositional elements, and, especially, architectural elements) permits comparison of facies associations from ancient and modern systems (cf. Fig. S5, see supplementary material), thereby providing the possibility to validate interpretations of environments or subenvironments in ancient fluvial systems. For example, the principle of comparative sedimentology can be applied to test planform-based interpretations of the rock record against observations on the facies organization of modern rivers, for which planform types are known. Additionally, as information from ancient and modern systems can be derived separately, this method overcomes the limitation of assuming that modern systems are closely analogous to ancient systems and provides the opportunity to assess the role of differential preservation potential for various types of fluvial deposits (cf. Jackson 1978; Hickin 1993; Miall 2006).

Perhaps, the most important strength of this database approach is its capability to overcome the end-member classification mentality in general; for example, the tendency to classify fluvial systems as braided or meandering – embodied by some of the example models presented herein – may tend to ignore the range of natural variability and may convey the idea that sedimentary systems must obey the ideal conditions of the end-members. A database of this kind can effectively be used to highlight the uniqueness of depositional systems, since each one is stored individually in the database and can be individually retrieved for comparison (cf. Fig. 13), thereby providing a more flexible benchmark

for reference. This system can therefore reconcile the "facies model" school-of-thought (as commonly taught, if not as originally conceived) in which there exists a discrete number of sedimentary environments, with the view that sedimentary environments tend to grade into each other (cf. Galloway & Hobday 1983; Anderton 1985; Miall 1985).

In addition, it should be apparent that, apart from generating quantitative fluvial facies models, whose scope is solely capturing patterns of sedimentary organization for environmental classes, a similar database provides the possibility to test the validity of theories concerning the genetic significance of architectural characteristics of fluvial systems and their occurrence within environmental types.

However, it must be borne in mind that the approach of utilizing a database for the generation of quantitative fluvial facies models suffers from several limitations, principally inherent in the sourceto-database workflow (cf. Saunders et al. 1995) and with the adoption of closed classification schemes, some of which include classes of purely interpretative nature: systems or genetic units may simply not fit in the existing classes, and interpretations may not be correct, may be uncertain, or may be mistakenly translated into the database system. Therefore, some precautions were taken at the database-design stage to avoid uncritical use of the system we presented. For example, to ensure consistency with original classifications and flexibility in categorization, open classification fields and multiple editable classification schemes are adopted, while the quality of interpretations and the resulting reliability of system and genetic-unit classifications is quantified by data-quality ranking (cf. Baas et al. 2005; Colombera et al. 2012a). Additionally, in cases where data do not fit in the existing classes, the relative attribute values are left undefined, signifying a lack of data or understanding on which to base the interpretation. Nevertheless, limitations in the approach must always be borne in mind and the application of such a system should never be conceived as a black-box technique. For example, creation of database-informed facies models requires that careful consideration be given to assessing uncertainty associated with the difficulty in constraining boundary conditions or system parameters for the rock record: this information could be integrated qualitatively in the model. Also, the specific database presented here could be significantly improved in the way it describes architectural styles. For example, this system currently lacks descriptors of genetic-unit shape (e.g.

wedge, sheet), descriptors of geometrical style of transition (e.g. onlap, offlap), and genetic-unit porosity and permeability data.

CONCLUSIONS

This paper demonstrates how a relational database created for the digitization of fluvial sedimentary architecture can be employed for the objective generation of facies models that are quantitative in nature and are customizable both in terms of system parameters on which they are categorized and type and scale of sedimentary units by which they are built. The type of information such models include is entirely analogous to what is traditionally presented in the form of idealized vertical logs or block diagrams, as they quantify genetic-unit abundances, geometries, spatial relationships and grain size. Data-input into the system is on-going: it is therefore still not possible to provide an exhaustive range of models spanning all environmental types and including all studied systems, and even the models presented here are only partially characterized in that they still lack information available from numerous published case studies. Yet, the example models presented herein demonstrate the value of the approach, especially in relation to its quantitative nature, its flexibility of application, and its capability to incorporate information concerning model uncertainty and variability. The proposed models may also serve as reference, as they provide insight into the sedimentary architecture of specific environmental types by quantifying the signature of basin climate regime, discharge regime and channel pattern – or of conditions conducive to the development of a channel-pattern type – on the large- to small-scale architecture of fluvial systems. Although the systems are only partially characterized in terms of their boundary conditions, future analysis of multiple case studies can be applied to the investigation of the role of a range of autogenic and allogenic controls on fluvial architecture. The method could be potentially applied to other depositional systems.

ACKNOWLEDGMENTS

We thank the Fluvial & Eolian Research Group sponsors (Areva, BHP Billiton, ConocoPhillips, Nexen, Saudi Aramco, Shell, and Woodside) for financial support to this project. Maurício Santos and Jo Venus are acknowledged for providing unpublished data. Reviewers Neil Davies, Ted Hickin and

Colin North, and Chief Editor Stephen Rice are gratefully thanked for their helpful advice, which considerably improved the scope and the form of the article.

REFERENCES

Allen, J.R.L. (1965) A review of the origin and characteristics of recent alluvial sediments. *Sedimentology*, **5**, 89-191.

Allen, J.R.L. (1970) Studies in fluviatile sedimentation: a comparison of fining upwards cyclothems, with particular reference to coarse member composition and interpretation. *J. Sed. Petrol.*, **40**, 298-323.

Allen, J.R.L. (1983) Studies in fluviatile sedimentation: bars, bar-complexes and sandstone sheets (low-sinuosity braided streams) in the Brownstones (L. Devonian), Welsh Borders. *Sed. Geol.*, **33**, 237-293.

Anderson, M.P., Aiken, J.S., Webb, E.K. and Mickelson, D.M. (1999) Sedimentology and hydrogeology of two braided stream deposits. *Sed. Geol.*, **129**, 187-199.

Anderton, R. (1985) Clastic facies models and facies analysis. In: *Sedimentology: recent developments and applied aspects* (Eds. P.J. Brenchley and B.P.J. Williams) pp. 21-48. The Geological Society, Blackwell Scientific Publications, Oxford.

Baas, J.H., McCaffrey, W.D. and Knipe, R.J. (2005) The Deep-Water Architecture Knowledge Base: towards an objective comparison of deep-marine sedimentary systems. *Petrol. Geosci.*, **11**, 309-320.

Banks, N.L. (1973) The origin and significance of some downcurrent dipping cross-stratified sets. *J. Sed. Petrol.*, **43**, 423-427.

Best, J.L. (1988) Sediment transport and bed morphology at river channel confluences. *Sedimentology*, **35**, 481-498.

Best, J.L., Ashworth, P.J., Bristow, C.S. and **Roden, J.** (2003) Three-dimensional sedimentary architecture of a large, mid-channel sand braid bar, Jamuna River, Bangladesh. *J. Sed. Res.*, **73**, 516-530.

Blair, T.C. and **McPherson, J.G.** (1992) The Trollheim alluvial fan and facies model revisited. *Geol. Soc. Am. Bull.*, **104**, 762-769.

Blair, T.C. and **McPherson, J.G.** (1994) Alluvial fans and their natural distinction from rivers based on morphology, hydraulic processes, sedimentary processes, and facies assemblages: *J. Sed. Res.*, **A64**, 450-489.

Bown, T.M. and **Kraus, M.J.** (1987) Integration of channel and floodplain suites, I. Developmental sequence and lateral relations of alluvial paleosols. *J. Sed. Petrol.*, **57**, 587-601.

Bridge, **J.S.** (1993) Description and interpretation of fluvial deposits: a critical perspective. *Sedimentology*, **40**, 801-810.

Bridge, J.S. (2003) Rivers and floodplains: forms, processes, and sedimentary record. Blackwell, Oxford, 491 pp.

- **Bridge**, **J.S.** (2006) Fluvial facies models: recent developments. In: *Facies models revisited* (Eds. H. Posamentier and R.G.Walker). *SEPM Spec. Publ.*, **84**, 85-170.
- **Brierley, G.J.** (1996) Channel morphology and element assemblages: a constructivist approach to facies modelling. In: *Advances in fluvial dynamics and stratigraphy* (Eds. P.A. Carling and M.R. Dawson), pp. 263-298. Wiley, Chichester.
- Brierley, G.J., Ferguson, R.J. and Woolfe, K.J. (1997) What is a fluvial levee? Sed. Geol., 114, 1-9.
- **Bristow, C.S.** (1987) Brahmaputra River: channel migration and deposition. In: *Recent developments in fluvial sedimentology* (Eds. E.G. Ethridge, R.M. Flores and M.D. Harvey). *SEPM Spec. Publ.*, **39**, 63-74.
- **Bristow, C.S., Best, J.L.** and **Roy, A.G.** (1993) Morphology and facies models of channel confluences. In: *Alluvial sedimentation* (Eds. M. Marzo and C. Puigdefábregas). *Int. Assoc. Sedimentol. Spec. Publ.*, **17**, 91-100.
- **Bristow, C.S., Skelly, R.L.** and **Ethridge, F.G.** (1999) Crevasse splays from the rapidly aggrading, sand-bed, braided Niobrara River, Nebraska: effect of base-level rise. *Sedimentology*, **46**, 1029-1049.
- **Bromley, M.H.** (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plateau, USA: relationship to salt tectonics in the Paradox Basin. *Sed. Geol.*, **73**, 77-99.
- **Bryant, I.D.** and **Flint, S.S.** (1993) Quantitative clastic reservoir geological modeling: problems and perspectives. In: *The geologic modelling of hydrocarbon reservoirs and outcrop analogs* (Eds. S.S. Flint and I.D. Bryant). *Int. Assoc. Sedimentol. Spec. Publ.*, **15**, 3-20.
- **Cabrera, L.I.** and **Saez, A.** (1987) Coal deposition in carbonate-rich shallow lacustrine systems: the Calaf and Mequinenza sequences (Oligocene, eastern Ebro Basin, NE Spain). *J. Geol. Soc. London*, **144**, 451-461.
- **Cain, S.A.** (2009) Sedimentology and stratigraphy of a terminal fluvial fan system: the Permian Organ Rock Formation, South East Utah. PhD dissertation, Keele University, Keele (UK), 461 pp.
- **Cain, S.A.** and **Mountney, N.P.** (2009) Spatial and temporal evolution of a terminal fluvial fan system: the Permian Organ Rock Formation, south east Utah, USA. *Sedimentology*, **56**, 1774-1800.
- **Cain, S.A.** and **Mountney, N.P.** (2011) Downstream changes and associated fluvial-aeolian interactions in an ancient terminal fluvial fan system: the Permian Organ Rock Formation, SE Utah. In: From River to Rock Record (Eds. S. Davidson, S. Leleu and C. North), *SEPM Spec. Publ.*, **97**, 165-187.
- Cant, D.J. (1978) Bedforms and bar types in the South Saskatchewan River. J. Sed. Petrol., 48, 1321-1330.
- Cant, D.J. (1982) Fluvial facies models and their application. In: Sandstone depositional environments (Eds. P.A. Scholle and D. Spearing). AAPG Mem., 31, 115-137.
- **Cant, D.J.** and **Walker, R.G.** (1976) Development of a braided-fluvial facies model for the Devonian Battery Point sandstone, Quebec. *Can. J. Earth Sci.*, **13**, 102-119.
- Cant, D.J. and Walker, R.G. (1978) Fluvial processes and facies sequences in the sandy braided South Saskatchewan River, Canada. *Sedimentology*, **25**, 625-648.

Carr, D.D., Horowitz, A., Hrabar, S.V., Ridge, K.F., Rooney, R., Straw, W.T., Webb, W. and Potter, P.E. (1966) Stratigraphic sections, bedding sequences and random processes. *Science*, **154**, 1162-1164.

Carr, T.R. (1982) Log-linear models, Markov chains and cyclic sedimentation. *J. Sed. Petrol.*, **52**, 905-912.

Catuneanu, O. and Elango, H.N. (2001) Tectonic control on fluvial styles: the Balfour Formation of the Karoo Basin, South Africa. *Sed. Geol.*, **140**, 291-313.

Coleman, J.M. (1969) Brahmaputra River: channel processes and sedimentation. *Sed. Geol.*, **3**, 129-239.

Collinson, J.D. (1996) Alluvial sediments. In: *Sedimentary environments: processes, facies and stratigraphy* (Ed. H.G. Reading) 3rd edn, pp. 37-82. Blackwell Science, Oxford.

Colombera, **L.**, **Mountney**, **N.P.** and **McCaffrey**, **W.D.** (2012a) A relational database for the digitization of fluvial architecture: concepts and example applications. *Petrol. Geosci.*, **18**, 129-140.

Colombera, L., Felletti, F., Mountney, N.P. and McCaffrey, W.D. (2012b) A database approach for constraining stochastic simulations of the sedimentary heterogeneity of fluvial reservoirs. *AAPG Bull.*, **96**, 2143-2166.

Cowan, E.J. (1991) The large-scale architecture of the fluvial Westwater Canyon Member, Morrison Formation (Jurassic), San Juan Basin, New Mexico. In: *The three-dimensional facies architecture of terrigenous clastic sediments, and its implications for hydrocarbon discovery and recovery* (Eds. A.D. Miall and N. Tyler). *SEPM Conc. Sed. Pal.*, **3**, 80-93.

Cowan, G. (1993) Identification and significance of aeolian deposits within the dominantly fluvial Sherwood Sandstone Group of the East Irish Sea Basin UK. In: *Characterization of fluvial and eolian reservoirs* (Eds. C.P. North and D.J. Prosser). *Geol. Soc. London Spec. Publ.*, **73**, 231-245.

Cuevas Gozalo, M.C. and **Martinius, A.W.** (1993) Outcrop database for the geological characterization of fluvial reservoirs: an example from distal fluvial-fan deposits in the Loranca Basin, Spain. In: *Characterization of fluvial and eolian reservoirs* (Eds. C.P. North and D.J. Prosser). *Geol. Soc. London Spec. Publ.*, **73**, 79-94.

Dott, R.H., Jr., and **Bourgeois, J.** (1983) Hummocky stratification: significance of its variable bedding sequences: reply to discussion by R.G. Walker et al. *Geol. Soc. Am. Bull.*, **94**, 1245-1251.

Dreyer, T., Fält, L., Høy, T., Knarud, R., Steel, R. and Cuevas, J.-L. (1993) Sedimentary architecture of field analogues for reservoir information (SAFARI): A case study of the fluvial Escanilla Formation, Spanish Pyrenees. In: *The geologic modelling of hydrocarbon reservoirs and outcrop analogs* (Eds. S.S. Flint and I.D. Bryant). *Int. Assoc. Sedimentol. Spec. Publ.*, **15**, 57-80.

Eschard, R., Doligez, B. and **Beucher, H.** (2002) Using quantitative outcrop databases as a guide for geological reservoir modeling. In: *Geostatistics Rio 2000 v. 1* (Eds. M. Armstrong, C. Bettini, N. Champigny, A. Galli, and A. Remacre), pp. 7-17. Dordrecht, Kluwer.

Fielding, C.R. (1984) A coal depositional model for the Durham Coal Measures of NE England. *J. Geol. Soc. London*, **141**, 919-931.

Fielding, C.R., Allen, J.P., Alexander, J. and **Gibling, M.R.** (2009) A facies model for fluvial systems in the seasonal tropics and subtropics. *Geology*, **37**, 623-626.

- **Fielding, C.R., Falkner, A.J.** and **Scott, S.G.** (1993) Fluvial response to foreland basin overfilling; the Late Permian Rangal Coal Measures in the Bowen Basin, Queensland, Australia. *Sed. Geol.*, **85**, 475-497.
- **Fillmore, D.L., Lucas, S.G.** and **Simpson E.L.** (2010) Invertebrate trace fossils in semi-arid to arid braided-ephemeral-river deposits of the Mississippian middle member of the Mauch Chunk Formation, eastern Pennsylvania, USA. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, **292**, 222-244.
- **Fisher, J.A., Krapf, C.B.E., Lang, S.C., Nichols, G.J.** and **Payenberg, T.D.** (2008) Sedimentology and architecture of the Douglas Creek terminal splay, Lake Eyre, central Australia. *Sedimentology*, **55**, 1915-1930.
- **Fisher, J.A., Nichols, G.J.** and **Waltham, D.A.** (2007) Unconfined flow deposits in distal sectors of fluvial distributary systems: examples from the Miocene Luna and Huesca Systems, northern Spain. *Sed. Geol.*, **195**, 55-73.
- **Friend, P.F.** (1983) Towards the field classification of alluvial architecture or sequence. In: *Modern and ancient fluvial systems* (Eds. J.D Collinson and J. Lewin). *Int. Assoc. Sedimentol. Spec. Publ.*, **6**, 345-354.
- **Galloway, W.E.** and **Hobday, D.K.** (1983) *Terrigenous clastic depositionalsSystems*. Springer-Verlag, New York, 423 pp.
- Gary, M., McAfee, R., Jr. and Wolf, C.L. (1974) Glossary of Geology. American Geological Institute, Washington, D.C., 805 pp.
- **Geehan, G.** and **Underwood, J.** (1993) The use of length distributions in geological modeling. In: *The geologic modelling of hydrocarbon reservoirs and outcrop analogs* (Eds. S.S. Flint and I.D. Bryant). *Int. Assoc. Sedimentol. Spec. Publ.*, **15**, 205-212.
- **Ghazi, S.** and **Mountney, N.P.** (2009) Facies and architectural element analysis of a meandering fluvial succession: the Permian Warchha Sandstone, Salt Range, Pakistan. *Sed. Geol.*, **221**, 99-126.
- **Gibling, M.R.** (2006) Width and thickness of fluvial channel bodies and valley fills in the geological record: a literature compilation and classification. *J. Sed. Res.*, **76**, 731-770.
- Gingerich, P.D. (1969) Markov analysis of cyclic alluvial sediments. J. Sed. Petrol., 39, 330-332.
- **Godin, P.D.** (1991) Fining-upward cycles in the sandy braided-river deposits of the Westwater Canyon Member (Upper Jurassic), Morrison Formation, New Mexico. *Sed. Geol.*, **70**, 61-82.
- **Goodman, L.A.** (1968) The analysis of cross-classified data: independence, quasi-independence, and interactions in contingency tables with or without missing entries. *Jour. Am. Statist. Assoc.*, **63**, 1091-1131.
- Gore, P.J. (1989) Toward a model for open-and closed-basin deposition in ancient lacustrine sequences: the Newark Supergroup (Triassic-Jurassic), eastern North America. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 70, 29-51.
- **Hampton, B.A.** and **Horton, B.K.** (2007) Sheetflow fluvial processes in a rapidly subsiding basin, Altiplano plateau, Bolivia. *Sedimentology*, **54**, 1121-1148.
- **Haszeldine**, **R.S.** (1983) Descending tabular cross-bed sets and bounding surfaces from a fluvial channel in the Upper Carboniferous coalfield of north-east England. In: J.D. Collinson and J. Lewin

(Editors), Modern and Ancient Fluvial Systems. In: *Modern and ancient fluvial Systems* (Eds. J.D Collinson and J. Lewin). *Int. Assoc. Sedimentol. Spec. Publ.*, **6**, 449-456.

Hickin, E.J. (1993) Fluvial facies models: a review of Canadian research. *Prog. Phys. Geogr.*, **17**, 205-222.

Hogg, S.E. (1982) Sheetfloods, sheetwash, sheetflow, or ... ? Earth-Sci. Rev., 18, 59-76.

Holbrook, J. (2001) Origin, genetic interrelationships, and stratigraphy over the continuum of fluvial channel-form bounding surfaces: an illustration from middle Cretaceous strata, southeastern Colorado. *Sed. Geol.*, **124**, 202-246.

Holzförster, F., Stollhofen, H. and **Stanistreet, I.G.** (1999) Lithostratigraphy and depositional environments in the Waterberg-Erongo area, central Namibia, and correlation with the main Karoo Basin, South Africa. *J. Afr. Earth. Sci.*, **29**, 105-123.

Hopkins, J.C. (1985) Channel-fill deposits formed by aggradation in deeply scoured superimposed distributaries of the Lower Kootenai Formation (Cretaceous). *J. Sed. Petrol.*, **55**, 42–52.

Horn, J.D., Fielding, C.R. and **Joeckel, R.** (2012) Revision of Platte River alluvial facies model through observations of extant channels and barforms, and subsurface alluvial valley fills. *J. Sed. Res.*, **82**, 72-91.

Horne, J.C., Ferm, J.C., Caruccio, F.T. and **Baganz, B.P.** (1978) Depositional models in coal exploration and mine planning in Appalachian region. *AAPG Bull.*, **62**, 2379-2411.

Hornung, J. and **Aigner, T.** (1999) Reservoir and aquifer characterization of fluvial architectural elements: Stubensandstein, Upper Triassic, southwest Germany. *Sed. Geol.*, **129**, 215-280.

Jackson, R.G., II (1978) Preliminary evaluation of lithofacies models for meandering alluvial streams. In: Fluvial Sedimentology (Ed. A.D. Miall) *Can. Soc. Petrol. Geol. Mem.*, **5**, 543-576.

Jones, N.S., Guion, P.D. and **Fulton, I.M.** (1995) Sedimentology and its application within the UK opencast coal mining industry. In: *European coal geology* (Eds. M.K.G. Whateley and D.A. Spears), *Geol. Soc. London Spec. Publ.*, **82**, 115-136.

Jordan, D.W. and **Pryor, W.A.** (1992) Hierarchical levels of heterogeneity in a Mississippi River meander belt and application to reservoir systems. *AAPG Bull.*, **76**, 1601-1624.

Jorgensen, P.J. and **Fielding, C.R.** (1996) Facies architecture of alluvial floodbasin deposits: three-dimensional data from the Upper Triassic Callide Coal Measures of east-central Queensland, Australia. *Sedimentology*, **43**, 479-495.

Jung, A. and **Aigner, T.** (2012) Carbonate geobodies: hierarchical classification and database – a new workflow for 3D reservoir modeling. *J. Petrol. Geol.*, **35**, 49-65.

Kirk, M. (1983) Bar developments in a fluvial sandstone (Westphalian "A"), Scotland. *Sedimentology*, **30**, 727-742.

Kirschbaum, M.A. and **McCabe, P.J.** (1992) Controls on the accumulation of coal and on the development of anastomosed fluvial systems in the Cretaceous Dakota Formation of southern Utah. *Sedimentology*, **39**, 581-598.

Kjemperud, A.V., Schomacker, E.R. and **Cross, T.A.** (2008) Architecture and stratigraphy of alluvial deposits, Morrison Formation (Upper Jurassic), Utah. *AAPG Bull.*, **92**, 1055-1076.

Kraus, M.J. and **Middleton, L.T.** (1987) Contrasting architecture of two alluvial suites in different structural settings. In: *Recent developments in fluvial sedimentology* (Eds. E.G. Ethridge, R.M. Flores and M.D. Harvey). *SEPM Spec. Publ.*, **39**, 253-262.

Lewin, J., Macklin, M.G. and **Johnstone, E.** (2005) Interpreting alluvial archives: sedimentological factors in the British Holocene fluvial record. *Quatern. Sci. Rev.*, **24**, 1873-1889.

Limarino, C., Tripaldi, A., Marenssi S., Net, L., Re, G. and **Caselli, A.** (2001) Tectonic control on the evolution of the fluvial systems of the Vinchina Formation (Miocene), northwestern Argentina. *J. S. Am. Earth Sci.*, **14**, 751-762.

Lunt, I.A., Bridge, J.S. and **Tye, R.S.** (2004) A quantitative, three-dimensional depositional model of gravelly braided rivers. *Sedimentology*, 51, 377-414.

Luttrell, P.R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. *Sed. Geol.*, **85**, 411-434.

McCabe, P.J. (1984) Depositional environments of coal and coal-bearing strata. In: *Sedimentology of coal and coal-bearing sequences* (Eds. R.A. Rahmani and R.M. Flores) *Int. Assoc. Sedimentol. Spec. Publ.*, **7**, 13-42.

McCabe, P.J. (1987) Facies studies of coal and coal-bearing strata. In: *Coal and coal-bearing strata: recent advances.* (Ed. A.C. Scott) *Geol. Soc. London Spec. Publ.*, **32**, 51-66.

Miall, A.D. (1973) Markov chain analysis applied to an ancient alluvial plain succession. *Sedimentology*, 20, 347-364.

Miall, A.D. (1977) A review of the braided river depositional environment. *Earth-Sci. Rev.*, 13, 1-62.

Miall, A.D. (1978) Lithofacies types and vertical profile models in braided river deposits: a summary. In: *Fluvial Sedimentology* (Ed. A.D. Miall). *Can. Soc. Petrol. Geol. Mem.*, **5**, 597-604.

Miall, A.D. (1979) Tertiary fluvial sediments in the Lake Hazen intermontane basin, Ellesmere Island, Arctic Canada. *Geol. Surv. Can. Pap.*, **79** -**9**, Ottawa, 25 pp.

Miall, A.D. (1980) Cyclicity and the facies model concept in fluvial deposits. *Bull. Can. Petrol. Geol.*, **28**, 59-80.

Miall, A.D. (1985) Architectural-element analysis: a new method of facies analysis applied to fluvial deposits. *Earth-Sci. Rev.*, **22**, 261-308.

Miall, A.D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Kayenta Formation (Lower Jurassic), Southwest Colorado. *Sed. Geol.*, **55**, 233-262.

Miall, A.D. (1996) The Geology of Fluvial Deposits. Springer Verlag, Berlin, 582 pp.

Miall, A.D. (1999) In defense of facies classifications and models. J. Sed. Res., 69, 2-5.

Miall, A.D. (2006) Reconstructing the architecture and sequence stratigraphy of the preserved fluvial record as a tool for reservoir development: a reality check. *AAPG Bull.*, **90**, 989-1002.

- **Miall, A.D.** and **Jones, B.G.** (2003) Fluvial architecture of the Hawkesbury Sandstone (Triassic), near Sydney, Australia. *J. Sed. Res.*, **73**, 531-545.
- **Miall, A.D.** and **Turner-Peterson, C.E.** (1989) Variations in fluvial style in the Westwater Canyon Member, Morrison Formation (Jurassic), San Juan Basin, Colorado Plateau. *Sed. Geol.*, **63**, 21-60.
- **Mjøs, R., Walderhaug, O.** and **Prestholm, E.** (1993) Crevasse splay sandstone geometries in the Middle Jurassic Ravenscar Group of Yorkshire, UK. In: *Alluvial sedimentation* (Eds. M. Marzo and C. Puigdefábregas). *Int. Assoc. Sedimentol. Spec. Publ.*, **17**, 167-184.
- Müller, R., Nystuen, J.P. and Wright, V.P. (2004) Pedogenic mud aggregates and paleosol development in ancient dryland river systems: criteria for interpreting alluvial mudrock origin and floodplain dynamics. *J. Sed. Res.*, **74**, 537-551.
- **Nadon, G.C.** (1994) The genesis and recognition of anastamosed fluvial deposits: data from the St. Mary River Formation, southwestern Alberta, Canada. *J. Sed. Res.*, **B64**, 451-463.
- **Nanson, G.C.** and **Croke, J.C.** (1992) A genetic classification of floodplains. In: *Floodplain Evolution* (Eds. G.R. Brakenridge and J. Hagedorn). *Geomorphology*, **4**, 459-486.
- **Nichols, G.J.** (2005) Sedimentary evolution of the Lower Clair Group, Devonian, West of Shetland: climate band sediment supply controls on fluvial, aeolian and lacustrine deposition. In: *Petroleum geology: north-west Europe and global perspectives, proceedings of the sixth petroleum geology conference* (Eds. A.G. Dore and B.A. Vinning), pp. 957-967. Geological Society, London.
- **Nichols, G.J.** and **Fisher, J.A.** (2007) Processes, facies and architecture of fluvial distributary system deposits. *Sed. Geol.*, **195**, 75-90.
- **North, C.P.** (1996) The prediction and modelling of subsurface fluvial stratigraphy. In: *Advances in Fluvial Dynamics and Stratigraphy* (Eds. P.A. Carling and M.R. Dawson), pp. 395-508. Wiley, Chichester.
- **North, C.P.** and **Davidson, S.K.** (2012) Unconfined alluvial flow processes: recognition and interpretation of their deposits, and the significance for palaeogeographic reconstructions. *Earth-Sci. Rev.*, **111**, 199-223.
- **Olsen, H.** (1989) Sandstone-body structures and ephemeral stream processes in the Dinosaur Canyon Member, Moenave Formation (Lower Jurassic), Utah, U.S.A. *Sed. Geol.*, **61**, 207-221.
- **Parkash, B., Awasthi, A.K.** and **Gohain K.** (1983) Lithofacies of the Markanda terminal fan, Kurukshetra district, Haryana, India In: *Modern and ancient fluvial Systems* (Eds. J.D Collinson and J. Lewin). *Int. Assoc. Sedimentol. Spec. Publ.*, **6**, 337-344.
- **Platt, N.H.** and **Keller, B.** (1992) Distal alluvial deposits in a foreland basin setting the lower freshwater Molasse (lower Miocene), Switzerland: sedimentology, architecture and palaeosols. *Sedimentology*, **39**, 545-565.
- **Reading, H.G.** (2001) Clastic facies models, a personal perspective. *Bull. Geol. Soc. Denmark*, **48**, 101-115.
- Reynolds, A.D. (1999) Dimensions of paralic sandstone bodies. AAPG Bull., 83, 211-229.
- **Robinson, J.W.** and **McCabe, P.J.** (1997) Sandstone-body and shale-body dimensions in a braided fluvial system: Salt Wash Sandstone Member (Morrison Formation), Garfield County, Utah. *AAPG Bull.*, **81**, 1267-1291.

Salter, T. (1993) Fluvial scour and incision: models for their influence on the development of realistic reservoir geometries. In: *Characterization of fluvial and eolian reservoirs* (Eds. C.P. North and D.J. Prosser). *Geol. Soc. London Spec. Publ.*, **73**, 33-51.

Sanabria, D.I. (2001) Sedimentology and sequence stratigraphy of the Lower Jurassic Kayenta Formation, Colorado Plateau, U.S.A. PhD dissertation, Rice University, Houston, 245 pp.

Sánchez-Moya, Y., Sopeña, A. and **Ramos, A.** (1996) Infill architecture of a non-marine half-graben Triassic basin (Central Spain). *J. Sed. Res.*, **66**, 1122-1136.

Saunders, M.R., Shields, J.A. and **Taylor, M.R.** (1995) Improving the value of geological data: a standardized data model for industry. In: *Geological data management* (Ed. J.R.A. Giles). *Geol. Soc. London Spec. Publ.*, **97**, 41-53.

Schumm, **S.A.** (1960) The shape of alluvial channels in relation to sediment type. Erosion and sedimentation in a semiarid environment. *US Geol. Surv. Prof. Pap.*, **352-B**, 17-30.

Schwarzacher, W. (1975) *Sedimentation models and quantitative stratigraphy.* Developments in Sedimentology, 19. Elsevier, New York, 387 pp.

Selley, R.C. (1970) Studies of sequence in sediments using a simple mathematical device. *Q. J. Geol. Soc. London*, **125**, 557-581.

Shultz, A.W. (1984) Subaerial debris-flow deposition in the Upper Paleozoic Cutler Formation, Western Colorado. *J. Sed. Petrol.*, **54**, 749-772.

Skelly, R.L., Bristow, C.S. and **Ethridge, F.G.** (2003) Architecture of channel-belt deposits in an aggrading shallow sandbed braided river: the lower Niobrara River, northeast Nebraska. *Sed. Geol.*, **158**, 249-270.

Sohn, Y.K., Rhee, C.W. and **Kim, B.C.** (1999) Debris Flow and Hyperconcentrated Flood-Flow deposits in an alluvial fan, northwestern part of the Cretaceous Yongdong basin, central Korea. *J. Geol.*, **107**, 111-132.

Steel, R.J. and **Thompson, D.B.** (1983) Structures and textures in Triassic braided stream conglomerates ('Bunter' Pebble Beds) in the Sherwood Sandstone Group, North Staffordshire, England. *Sedimentology*, **30**, 341-367.

Stephens, M. (1994) Architectural element analysis within the Kayenta Formation (Lower Jurassic) using ground-probing radar and sedimentological profiling, southwestern Colorado. *Sed. Geol.*, **90**, 179-211.

Thomas, R.G., Smith, D.G., Wood, J.M., Visser, J., Calverley-Range, E.A. and Koster, E.H. (1987) Inclined heterolithic stratification – Terminology, description, interpretation and significance. *Sed. Geol.*, **53**, 123-179.

Tye, R.S. (2004) Geomorphology: an approach to determining subsurface reservoir dimensions. *AAPG Bull.*, **88**, 1123-1147.

Wakelin-King, G.A. and **Webb, J.A.** (2007) Upper-flow-regime mud floodplains, lower-flow-regime sand channels: sediment transport and deposition in a drylands mud-aggregate river. *J. Sed. Res.*, **77**, 702-712.

Walker, R.G. (1984) General introduction: facies, facies sequences and facies models. In: *Facies Models* (Ed. R.G. Walker) 2nd edn, pp. 1-13. Geological Association of Canada Reprint Series, Toronto.

Walker, R.G. and **Cant, D.J.** (1984) Sandy fluvial systems. In: *Facies Models* (Ed. R.G. Walker) 2nd edn, pp. 71-90. Geological Association of Canada Reprint Series, Toronto.

Walker, W.E., Harremoës, P., Rotmans, J., van der Sluijs, J.P., van Asselt, M.B.A., Janssen, P. and Krayer von Krauss, M.P. (2003) Defining uncertainty: a conceptual basis for uncertainty management in model-based decision support. *Integr. Assessment*, **4**, 5-18.

Willis, B.J. (1993) Interpretation of bedding geometry within ancient point bar deposits. In: *Alluvial sedimentation* (Eds. M. Marzo and C. Puigdefábregas). *Int. Assoc. Sedimentol. Spec. Publ.*, 17, 101-114.

Wizevich, M.C. (1992) Sedimentology of Pennsylvanian quartzose sandstones of the Lee Formation, central Appalachian Basin: fluvial interpretation based on lateral profile analysis. *Sed. Geol.*, **78**, 1-47.

CAPTIONS

Table 1

Summary of the fundamental diagnostic characteristics and environmental significance of the 14 interpretative architectural-element types employed in the FAKTS database.

Table 2

Summary of the fundamental textural and structural characteristics of the 25 facies-unit types employed in the FAKTS database.

Figure 1

Representation of the main scales of observation and types of sedimentary genetic units included in the FAKTS database. Refer to Table 1 for architectural-element codes and to Table 2 for facies-unit codes (modified from Colombera et al., 2012a).

Figure 2

Example application of three different methods for computing model architectural-element proportions (see text); as no filter has been applied on either system parameters or sedimentological properties, the results refer to an ideal model of a "generic" fluvial environment derived from and constrained by the entire knowledge base.

Figure 3

Quantitative information regarding the proportion and geometry (width and thickness) of channel-complexes, constituting large-scale facies models for perennial sub-humid meandering systems and systems associated with intermediate filtering steps. In this case, as in all models presented here, the term 'basin climate type' only refers to the observed/inferred humidity-based climate class at the locus of deposition; a catchment climate classification is also stored, but it applies mostly to modern systems and may refer to average conditions.

Figure 4

Quantitative information referring to large-scale facies models for single-thread and braided river systems: a) boxplots describing the distribution of channel-complex proportions within different stratigraphic volumes (subsets) used to include information about the variability in depositional-element proportions in the models; b) log-normal probability density functions describing the distribution of channel-complex thickness; c) cross-plots of channel-complex thickness and width, classified as complete (real or apparent widths) or incomplete (partial or unlimited widths). Idealized

cross-sections comparable to traditional models and informed on such quantitative information are depicted in (d) to highlight architectural differences between the two models.

Figure 5a

Quantitative information regarding the proportion and vertical transition statistics of architectural elements, constituting intermediate-scale facies models for arid/semiarid ephemeral braided systems and systems associated with intermediate filtering steps. Idealized block-diagrams comparable to traditional models and informed on such quantitative information are depicted in the left-hand column; model architectural-element proportions, presented as pie-charts in the central column, are derived as the sum of the thickness of all elements from adequate subsets (method 1 in Fig. 2 and in the text); vertical transition statistics are presented in the right-hand column as bar charts quantifying the percentage of types of 'upper' elements (colour-coded and labelled in the bars) stacked on top of a given type of 'lower' element (labels on the vertical axis).

Figure 5b

Continuation of Fig. 5. Information on architectural-element horizontal spatial relationships, in the form of cross-gradient and up-gradient transition statistics. Results are presented in the central and right-hand column as bar charts quantifying the percentage of 'cross-gradient' or 'up-gradient' element types (colour-coded and labelled in the bars) juxtaposed to element types labelled on the vertical axis.

Figure 6a

Description of architectural-element geometries for different models. Box-plots in the right-hand column include information on the thickness of the different architectural-element types, for facies models of arid/semiarid ephemeral braided systems and systems associated with intermediate filtering steps.

Figure 6b

Continuation of Fig. 6. Cross-plots in the right-hand column include information on the relationship between width and thickness of different architectural-element types for facies models of arid/semiarid ephemeral braided systems and systems associated with intermediate filtering steps.

Figure 7

Example quantitative information that can be incorporated into a small-scale facies model referring to the entire knowledge base (no filter applied). Overall facies-unit proportions are presented as piecharts of textural classes and of 'texture + structure' facies-unit classes, and are compared with the facies organization of channel deposits, described by facies unit proportions within channel-complexes. The geometry of different facies-unit types is quantified by box-plots of their thickness distribution, summary descriptive statistics of their lateral extent, and probability density functions of the width/thickness aspect ratio of selected types. Upwards, cross-gradient and up-gradient transition statistics are presented as bar charts quantifying the percentage of types of facies units (colour-coded and labelled in the bars) juxtaposed to a given type of facies unit (labels on the vertical axis). In addition, the facies-unit-scale block diagram has been built based on database-derived information relating to the facies organization and geometry of individual architectural-element types.

Figure 8

Example quantitative information that can be incorporated into a small-scale facies model referring to braided systems, filtering the knowledge-base on the channel-pattern type. Results are presented as in Fig. 7, to render the models comparable.

Figure 9

Example quantitative information that can be incorporated into a small-scale facies model referring to dryland braided systems, filtering braided systems on the basin climate type. Results are presented as in Fig. 7 and 8, to render the models comparable.

Figure 10

Example quantitative information that can be incorporated into a small-scale facies model referring to ephemeral dryland braided systems, filtering dryland braided systems on the water-discharge regime. Results are presented as in Fig. 7, 8 and 9, to render the models comparable.

Figure 11

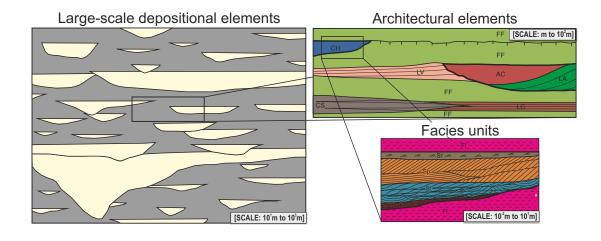
Partial quantitative information constituting a small-scale facies model of aggradational channel fills (*CH* architectural elements). The model facies association of the element is described by overall lithofacies-type proportions, presented as pie-charts of textural classes and of 'texture + structure' facies-unit classes; proportions of facies types observed at the base of channel-fills are also given. Example cumulative grain-size distributions for facies units within *CH* elements are presented for different lithofacies types; the thickness and width of classified facies units within aggradational channel fills is represented in the cross-plot; upwards, cross-gradient and up-gradient transition statistics are presented as bar charts quantifying the percentage of types of facies units (colour-coded and labelled in the bars) juxtaposed to a given type of facies unit (labels on the vertical axis) within *CH* elements. Legend and colour code are given in Fig. 10.

Figure 12

Graphs quantifying the downstream variations in the proportion of textural classes (left-hand graph) and example facies-unit types (right-hand graphs), for two different depositional systems (Parkash et al. 1983; Cain 2009, cf. Cain & Mountney 2009; 2011) classified as "terminal fans". Note that the length scales over which the variations are observed are different for the two systems, to make the results referable to a tripartite subdivision of the systems into 'proximal', 'medial' and 'distal' zones and comparable with existing models; similar results could be derived for absolute-distance scales.

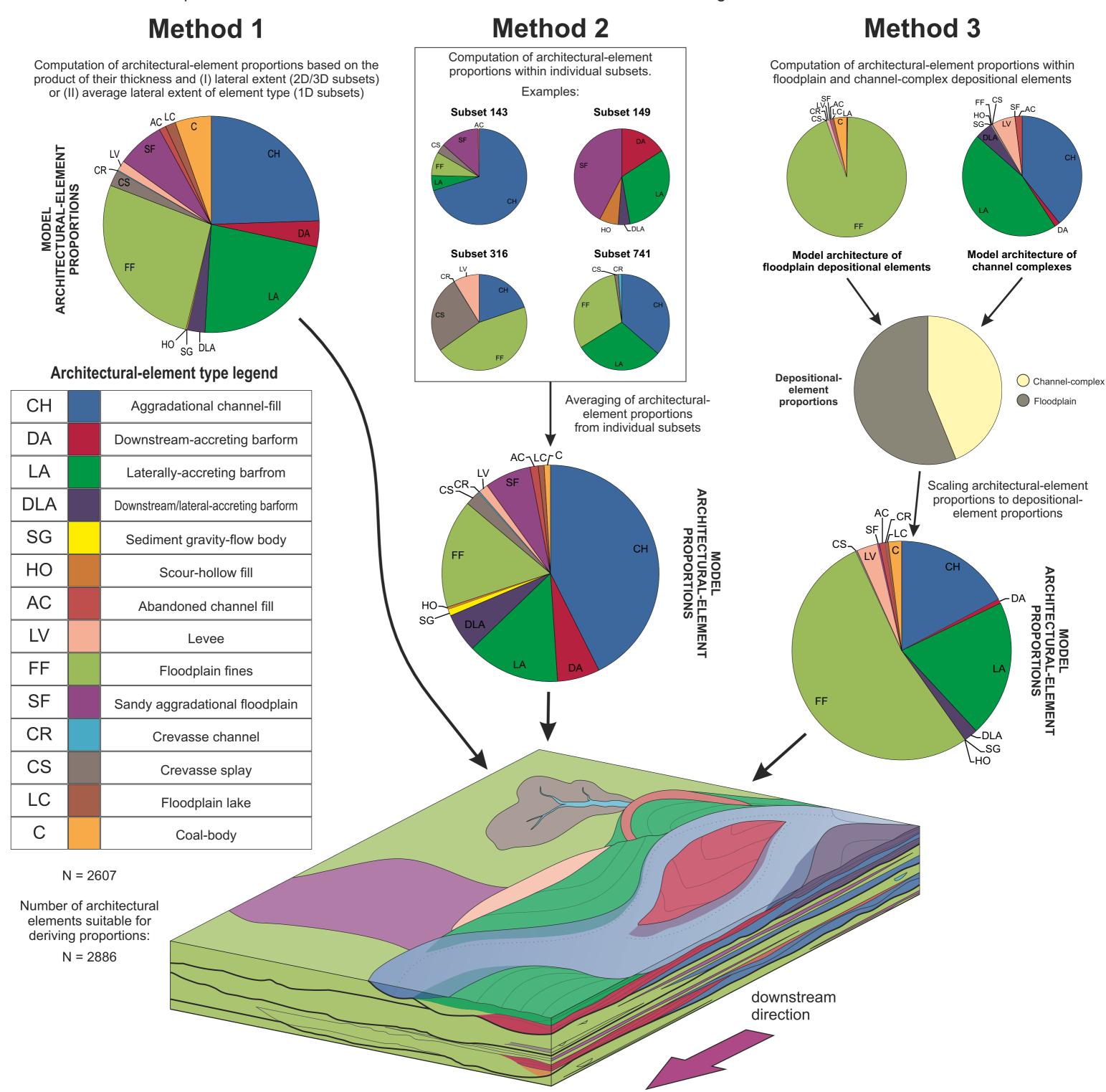
Figure 13

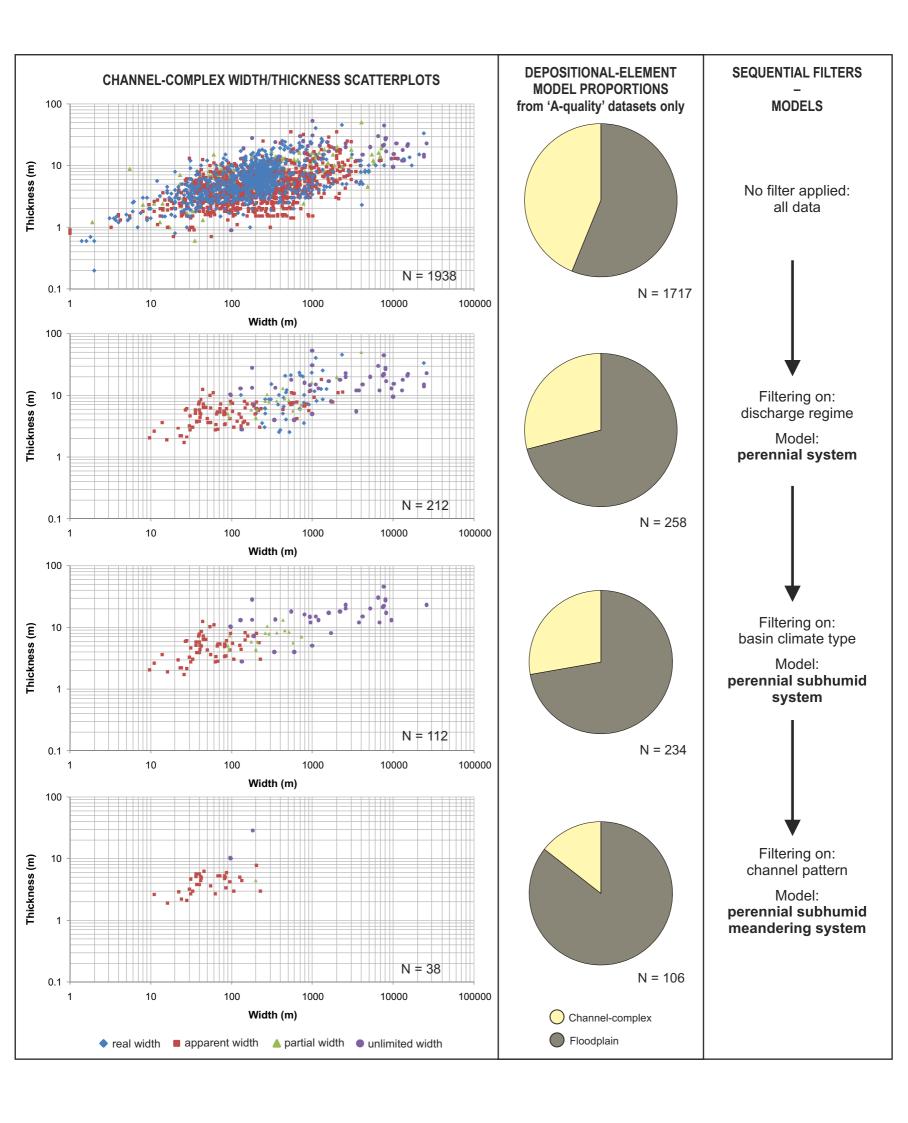
Comparison between the model facies association of 'lateral accretion barforms' (*LA* architectural elements) represented by the pie-chart, which quantifies facies-unit proportions derived as the sum of facies-unit thickness (method 1 in Fig. 2 and in the text), and the partial result of a query returning the proportion of facies-unit types within each individual *LA* architectural element, in tabulated form (e.g. 'St/0.11' means 11% of *St* facies unit with the given element). The possibility to individually store and retrieve each depositional system or genetic unit renders the FAKTS database system a reference for comparison that is richer and more flexible than traditional facies models.

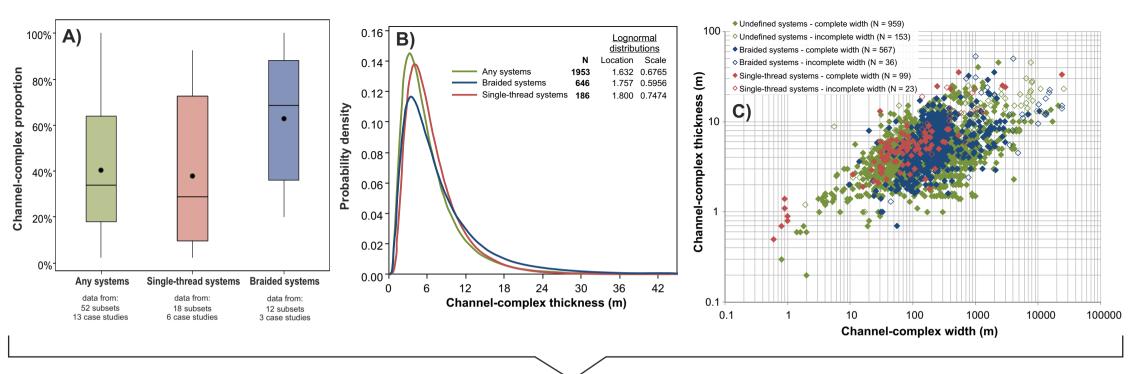


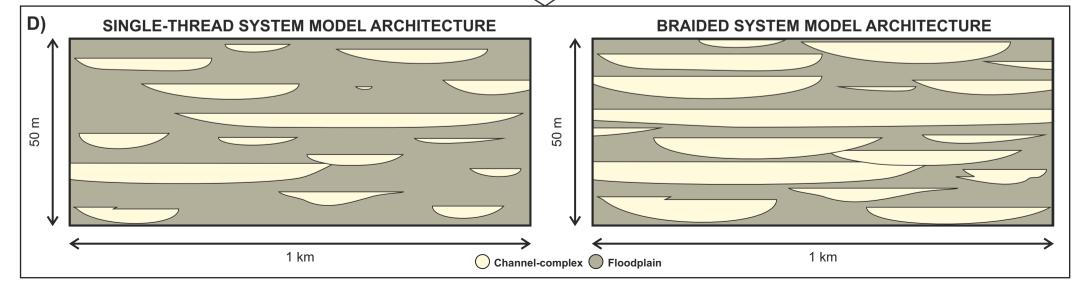
ARCHITECTURAL-ELEMENT PROPORTIONS - NO FILTER APPLIED

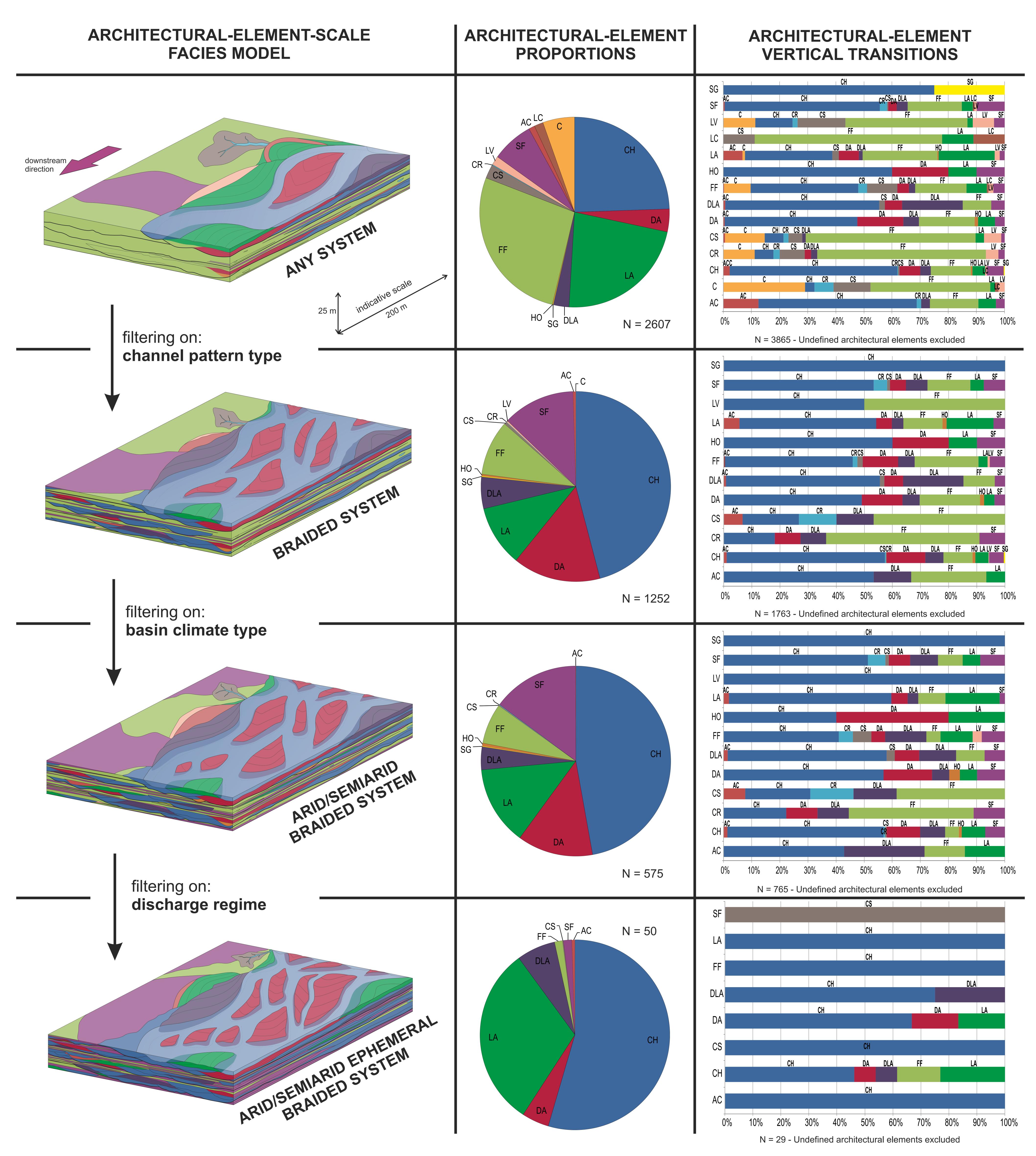
representation of the relative abundance of architectural elements among all fluvial environments

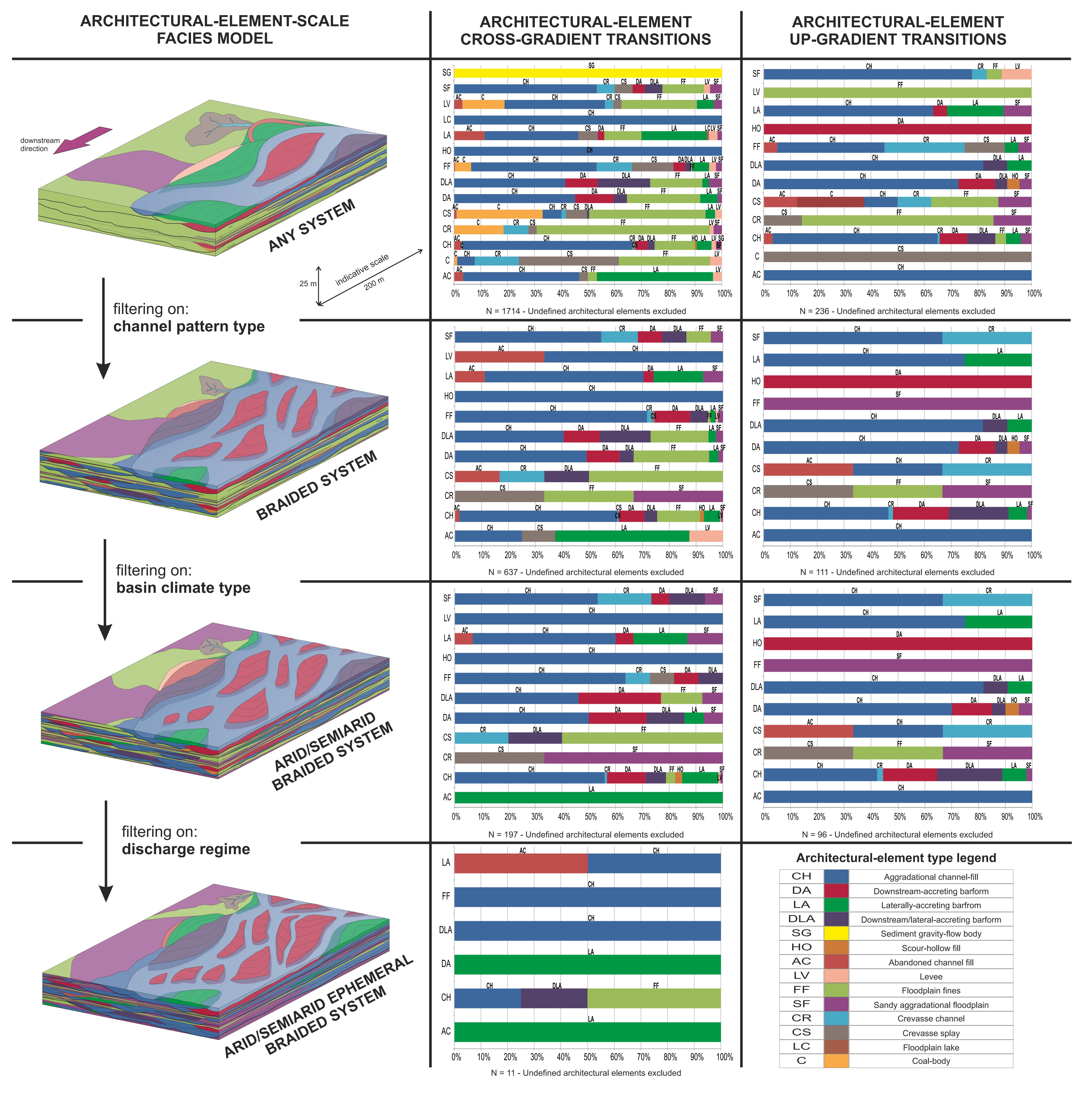


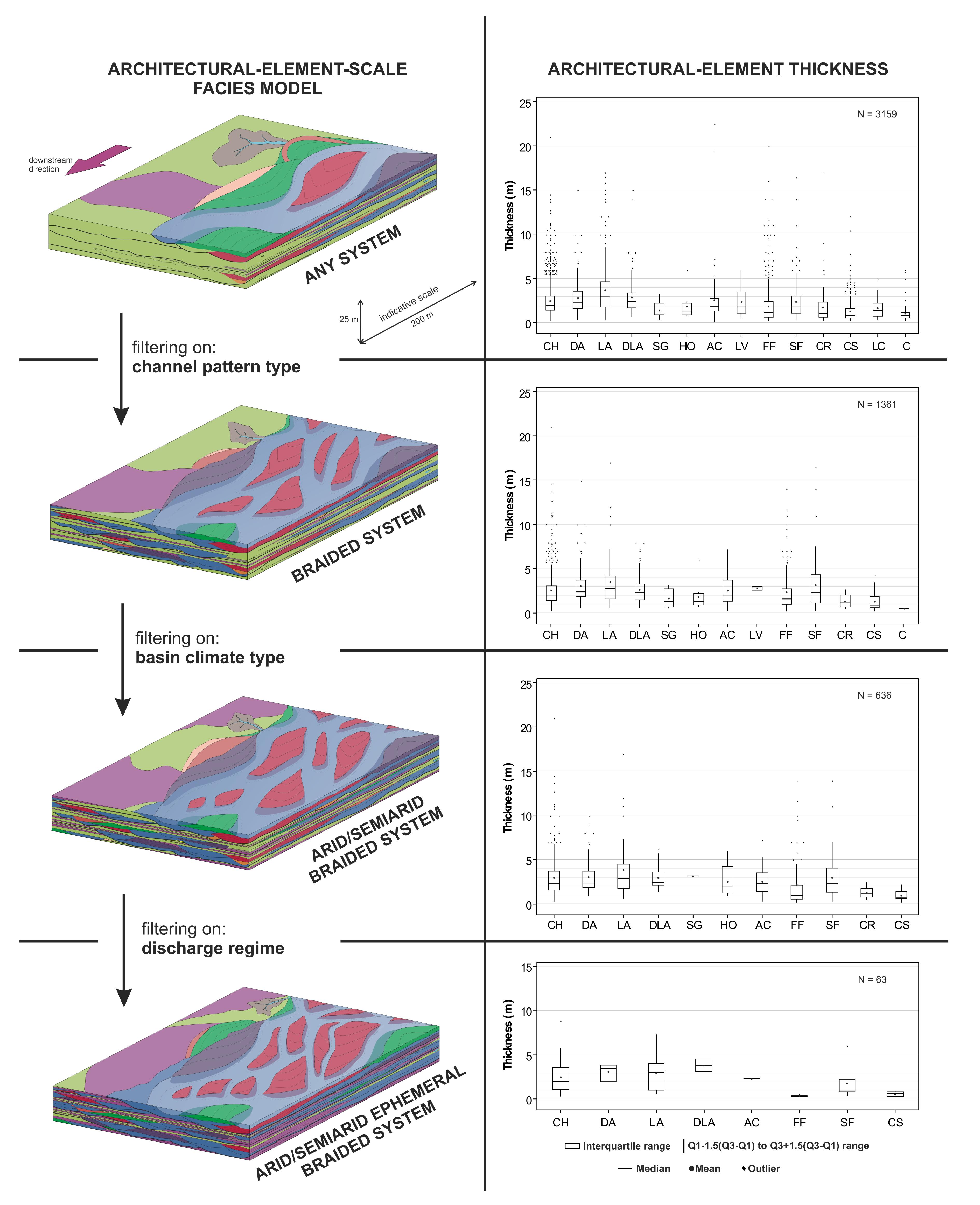


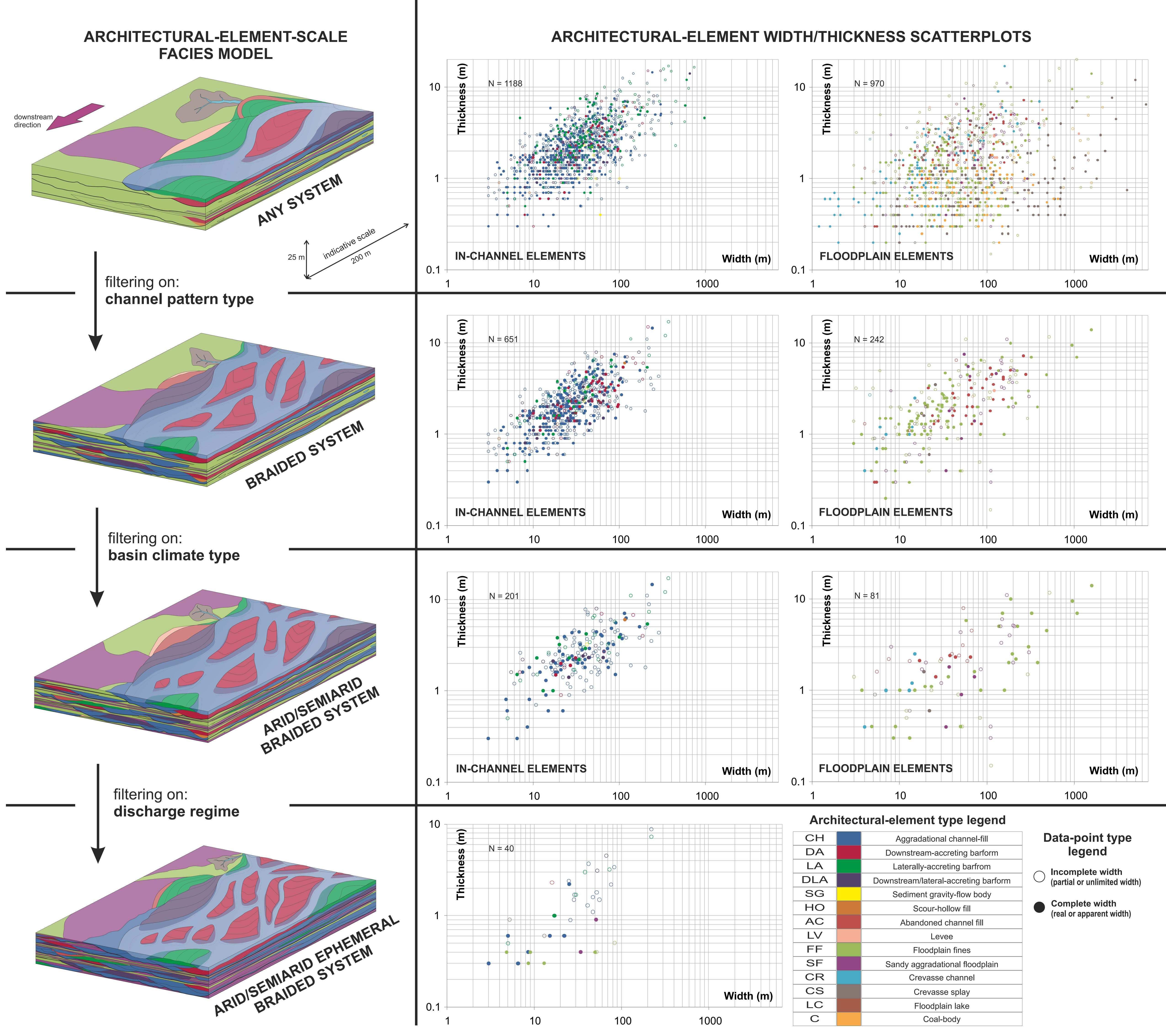


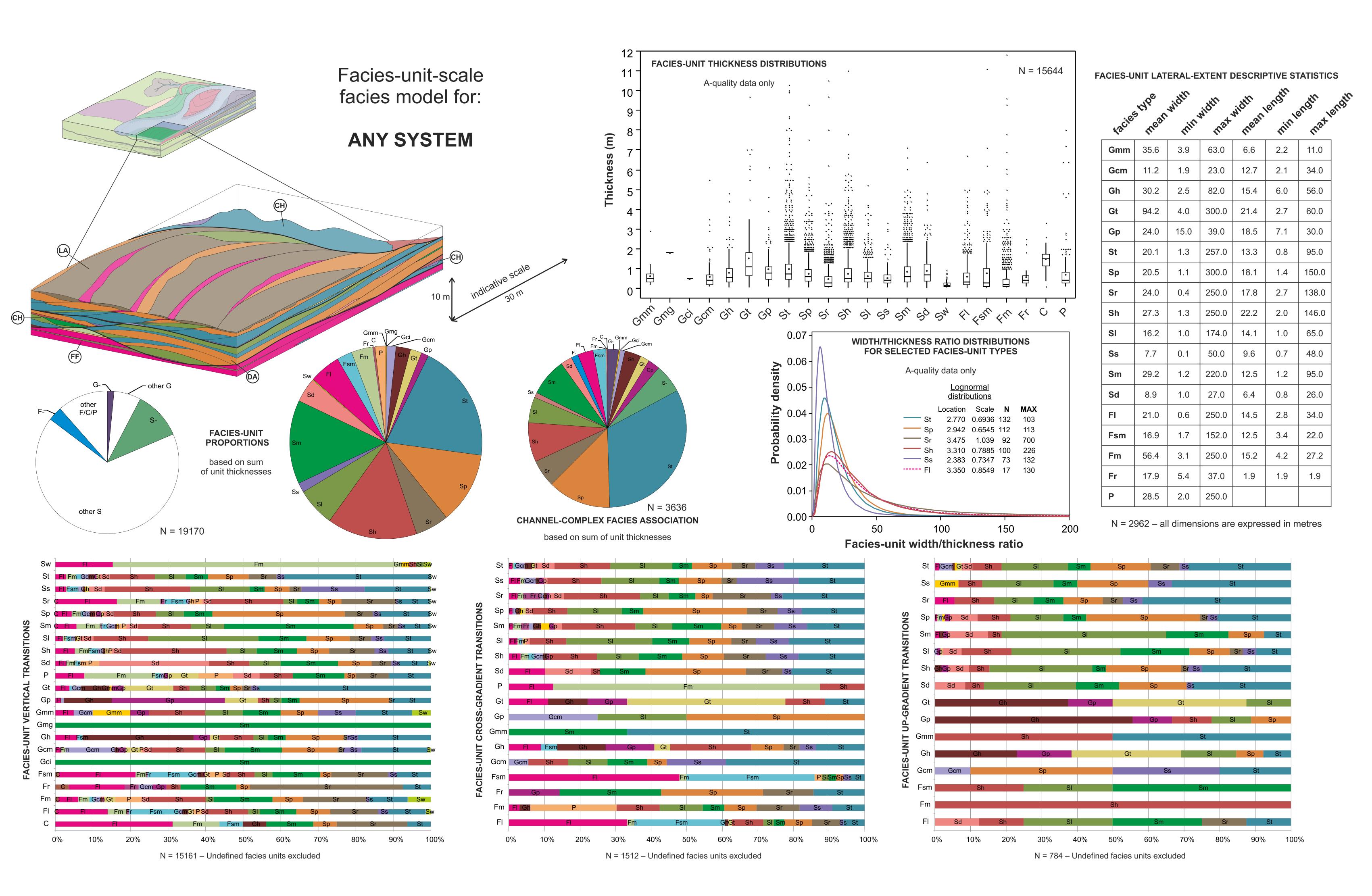


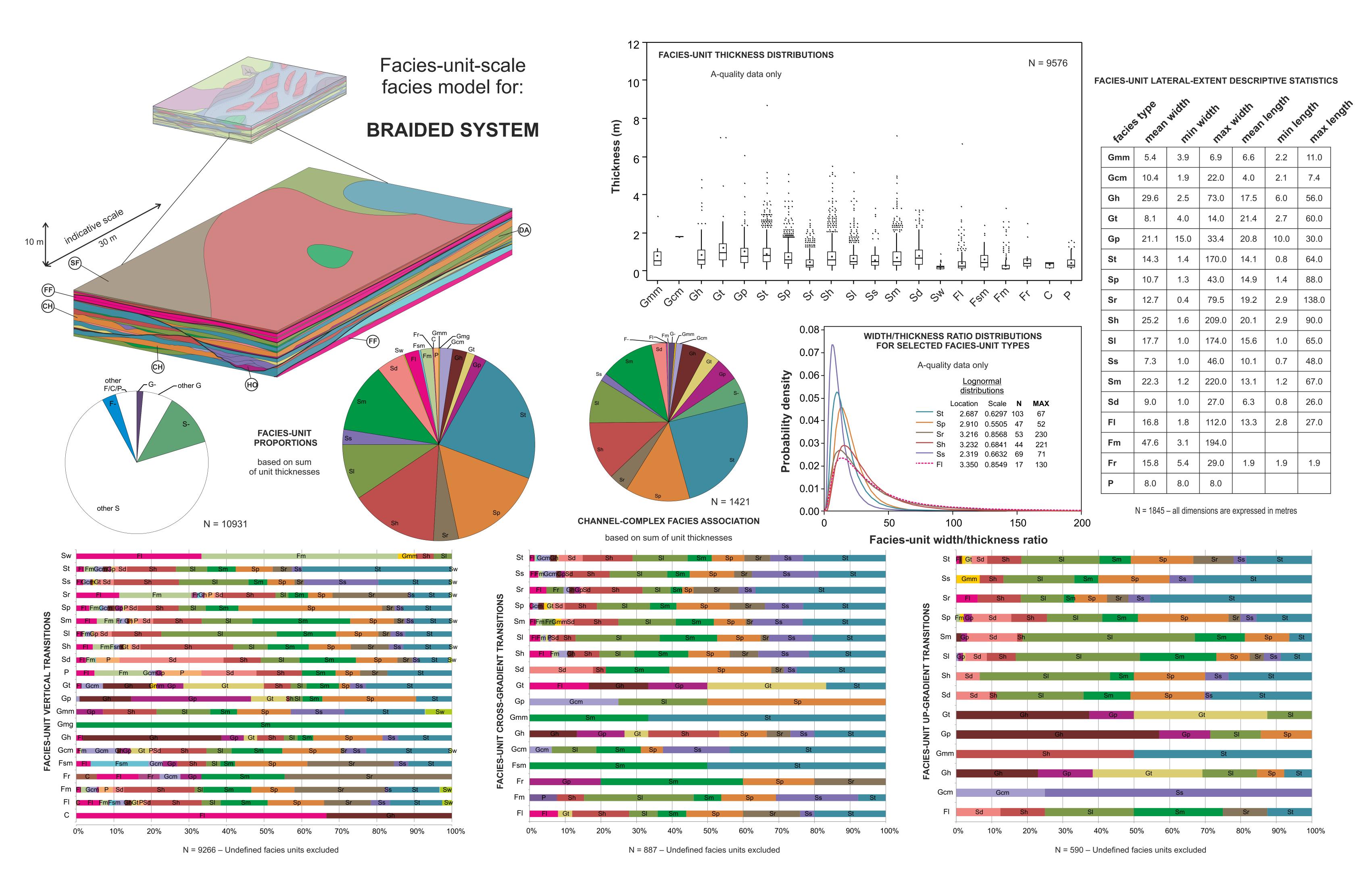


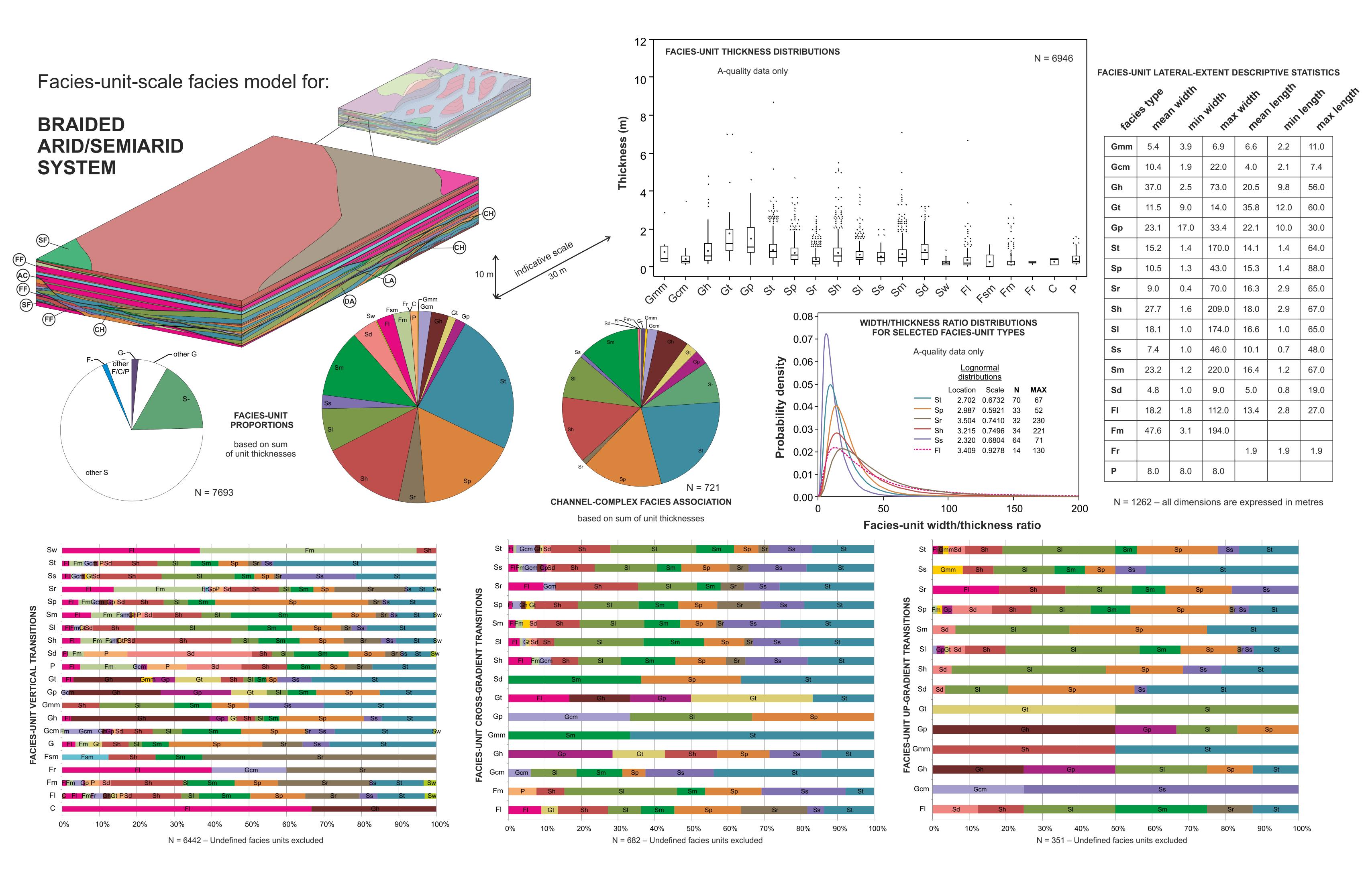


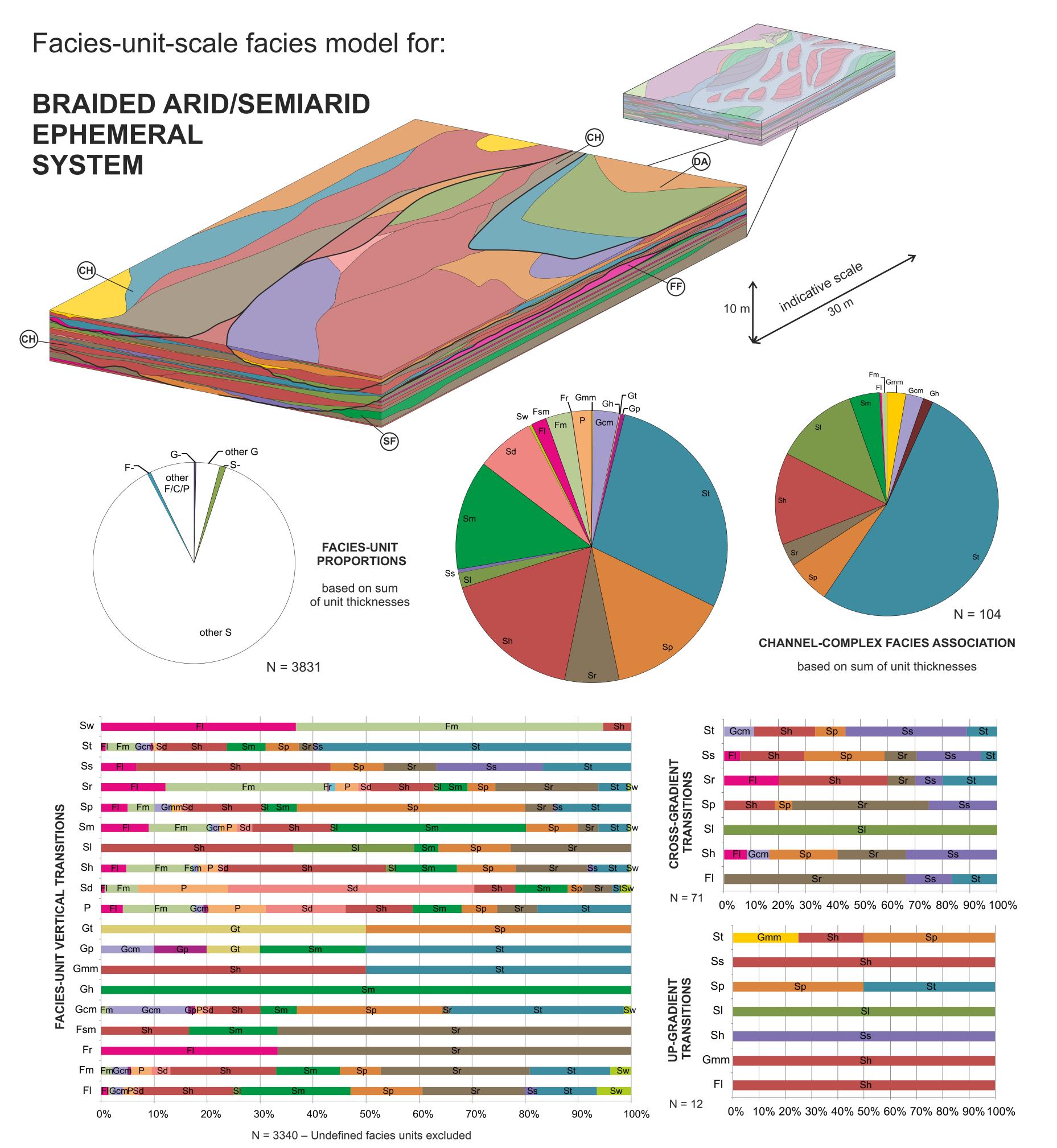


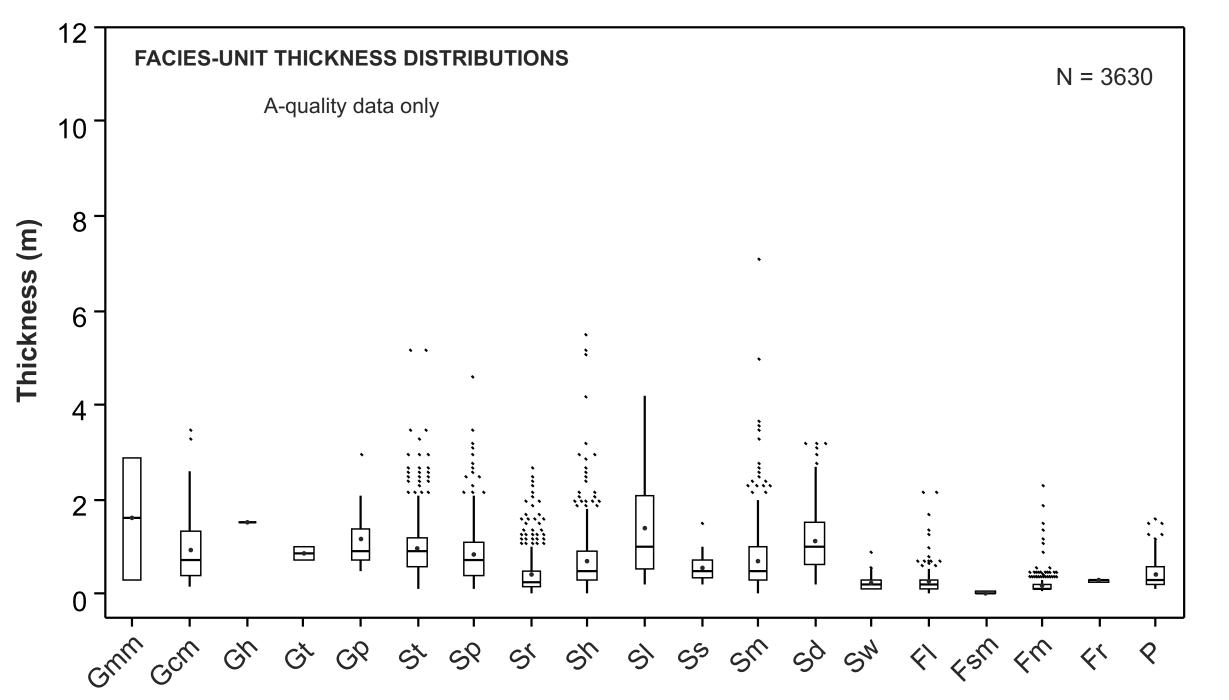


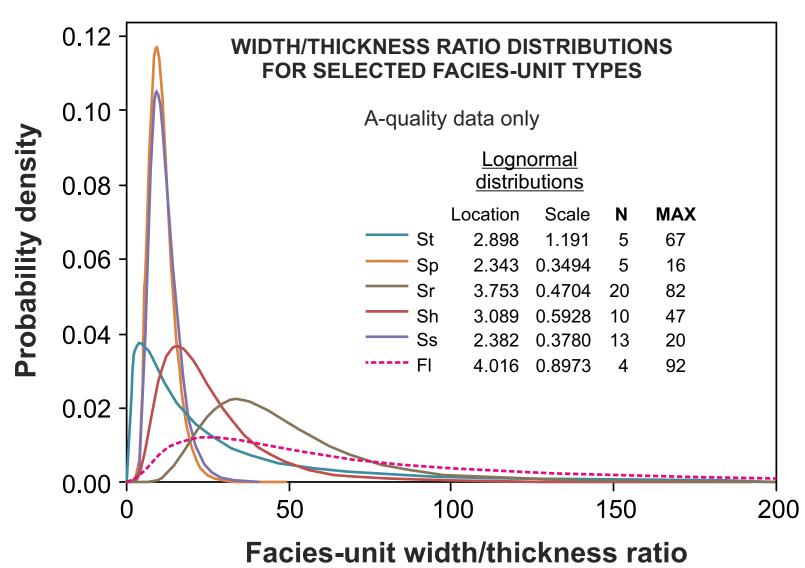








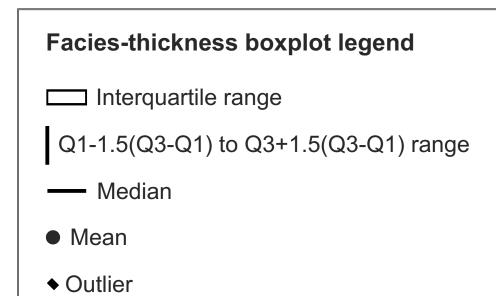


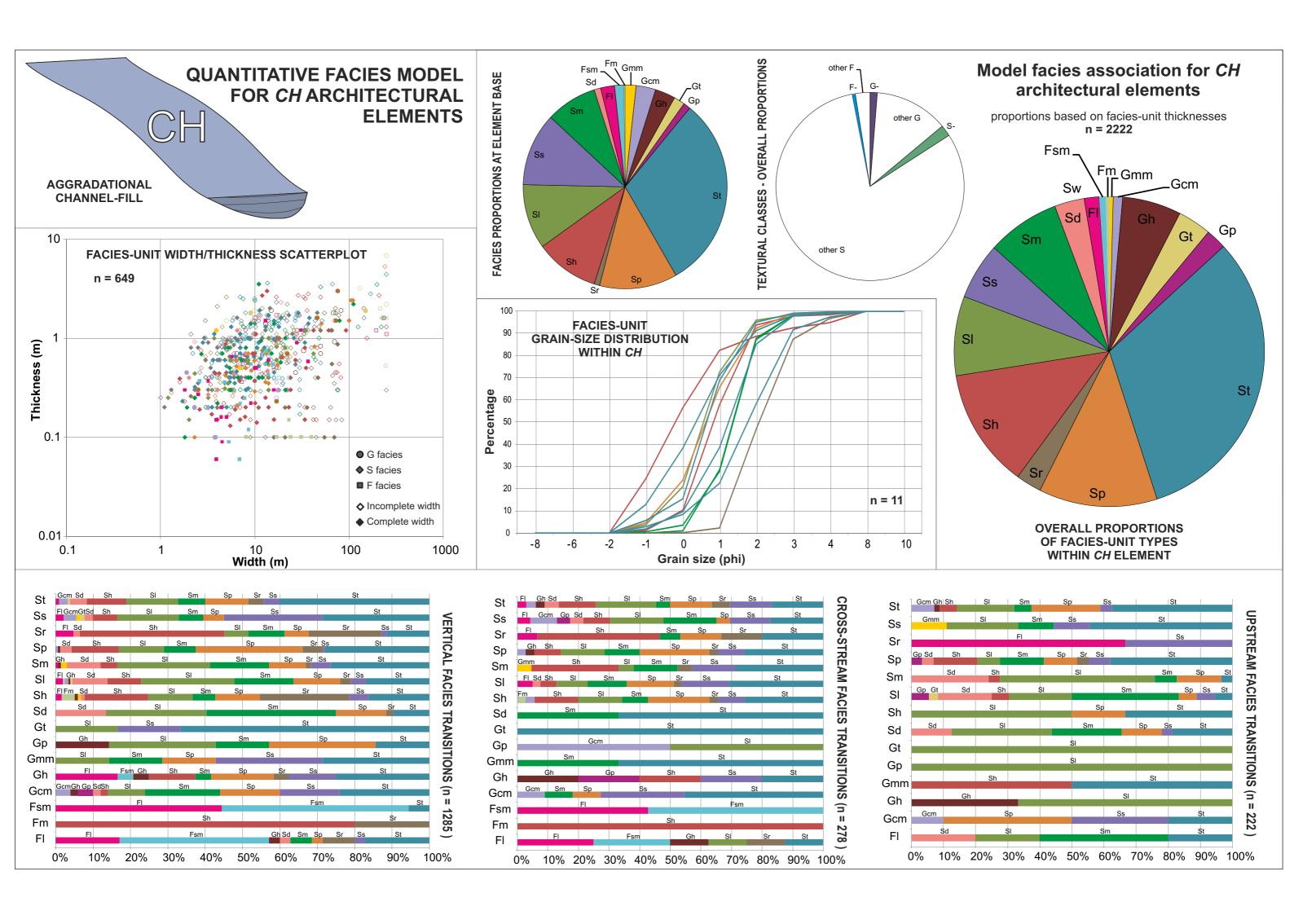


Facile	we.	nidth	idth	idth	ength	ngth
k acies	type mean	nidth rin	nidth max	vidth mean	endth le	nath
Gh	73.0	73.0	73.0			
St	12.2	2.0	22.1	13.6	1.4	38.3
Sp	7.6	1.9	25.0	14.9	10.0	33.3
Sr	6.7	1.5	22.0	15.0	10.0	20.0
Sh	21.6	2.3	209.0	15.5	10.0	33.4
SI	125.5	77.0	174.0	20.7	14.0	29.0
Ss	5.9	1.8	13.5	10.7	4.3	17.0
Sm	146.0	94.0	198.0			
FI	13.4	2.9	25.0	16.7	10.0	20.0

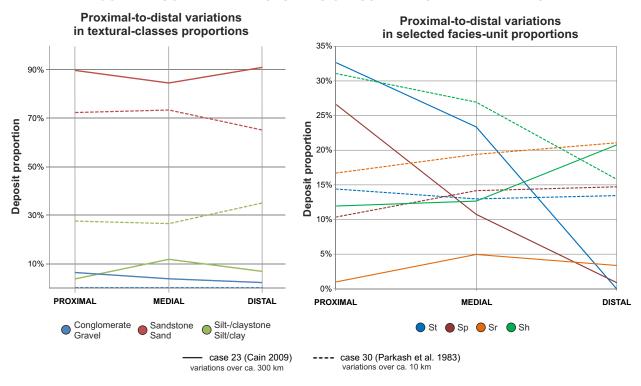
N = 128 – all dimensions are expressed in metre			N = 128 - all	dimensions	are	expressed	in	metres
---	--	--	---------------	------------	-----	-----------	----	--------

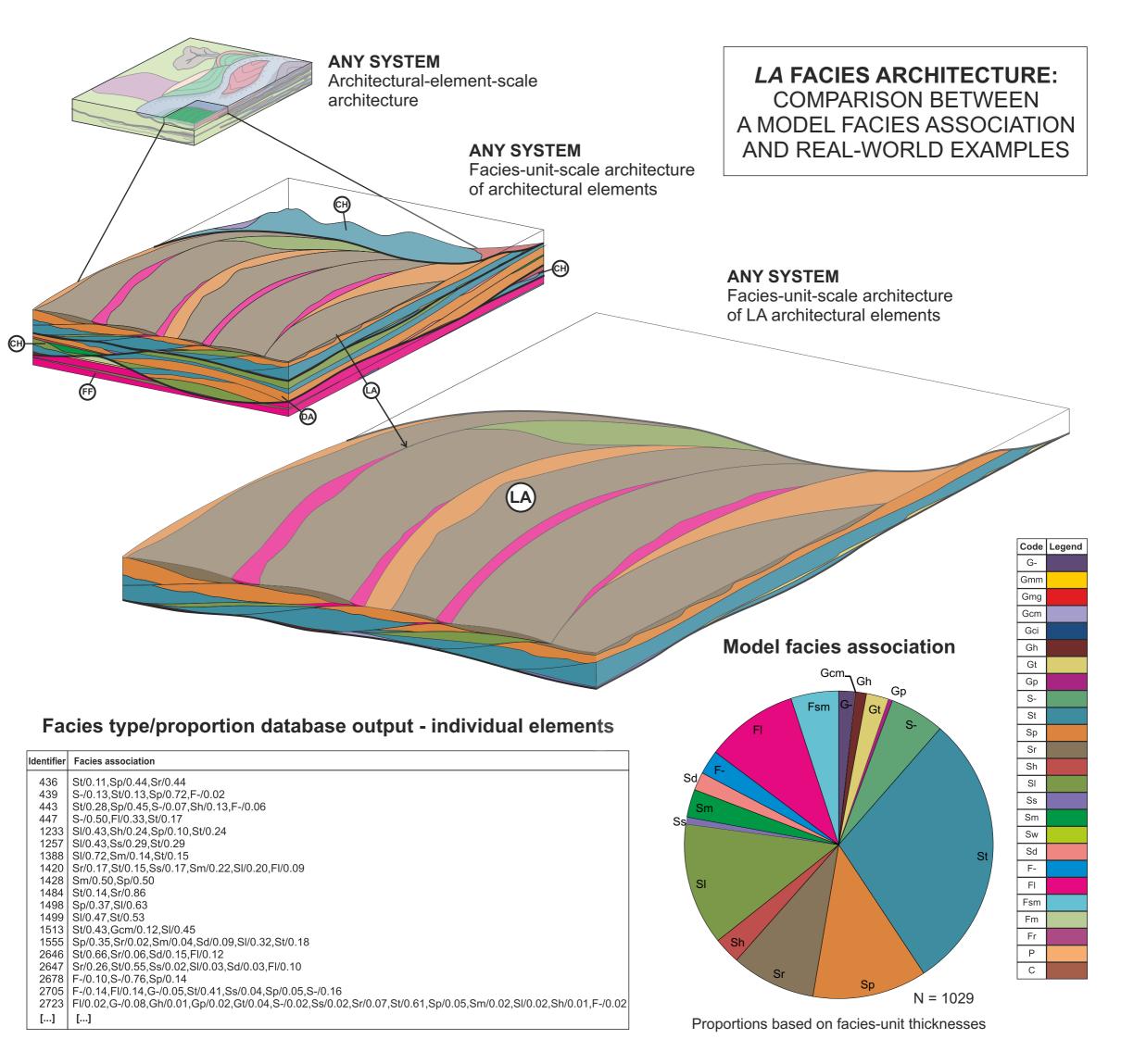
G-	Gravel to boulders - undefined structure
Gmm	Matrix-supported massive gravel
Gmg	Matrix supported graded gravel
Gcm	Clast-supported massive gravel
Gci	Clast-supported inversely-graded gravel
Gh	Horizontally-bedded or imbricated gravel
Gt	Trough cross-stratified gravel
Gp	Planar cross-stratified gravel
S-	Sand - undefined structure
St	Trough cross-stratified sand
Sp	Planar cross-stratified sand
Sr	Asymmetric-ripple cross-laminated sand
Sh	Horizontally-laminated sand
SI	Low-angle cross-bedded sand
Ss	Scour-fill sand
Sm	Massive or faintly laminated sand
Sw	Symmetric-ripple cross-laminated sand
Sd	Soft-sediment deformed sand
F-	Fines (silt, clay) - undefined structure
FI	Laminated sand, silt and clay
Fsm	Laminated to massive silt and clay
Fm	Massive clay and silt
Fr	Fine-grained root bed
Р	Paleosol carbonate
С	Coal or carbonaceous mud

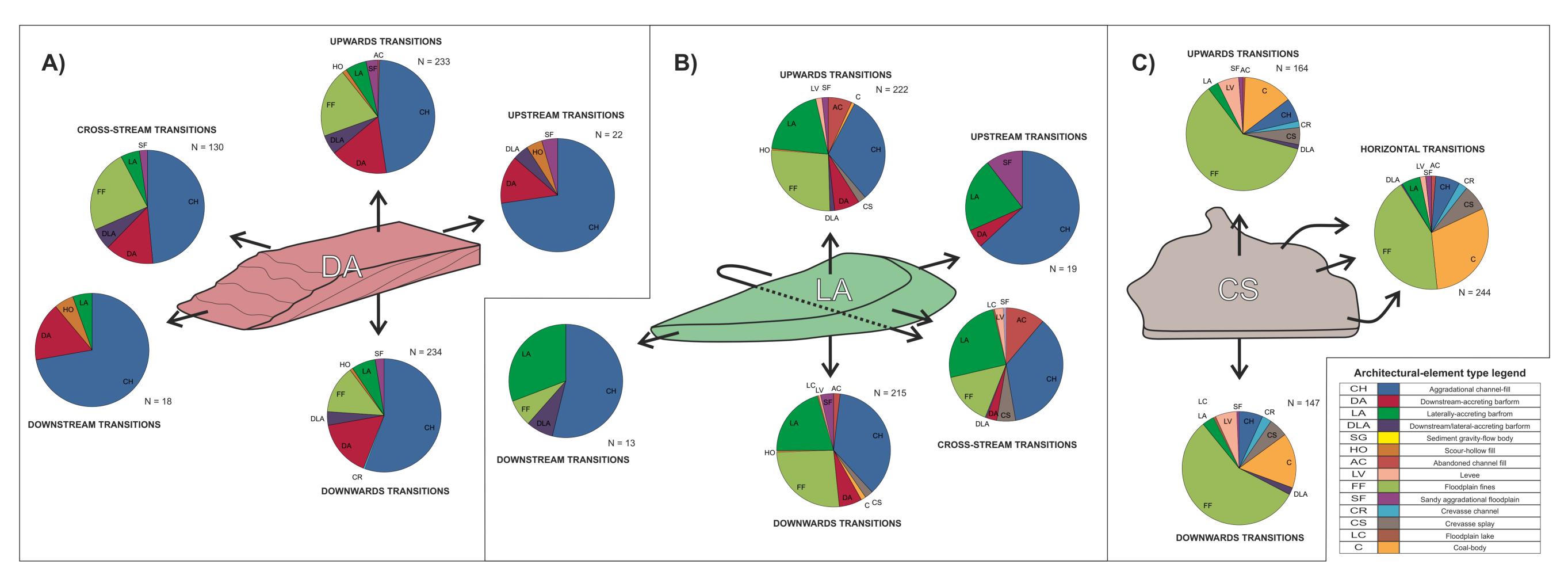




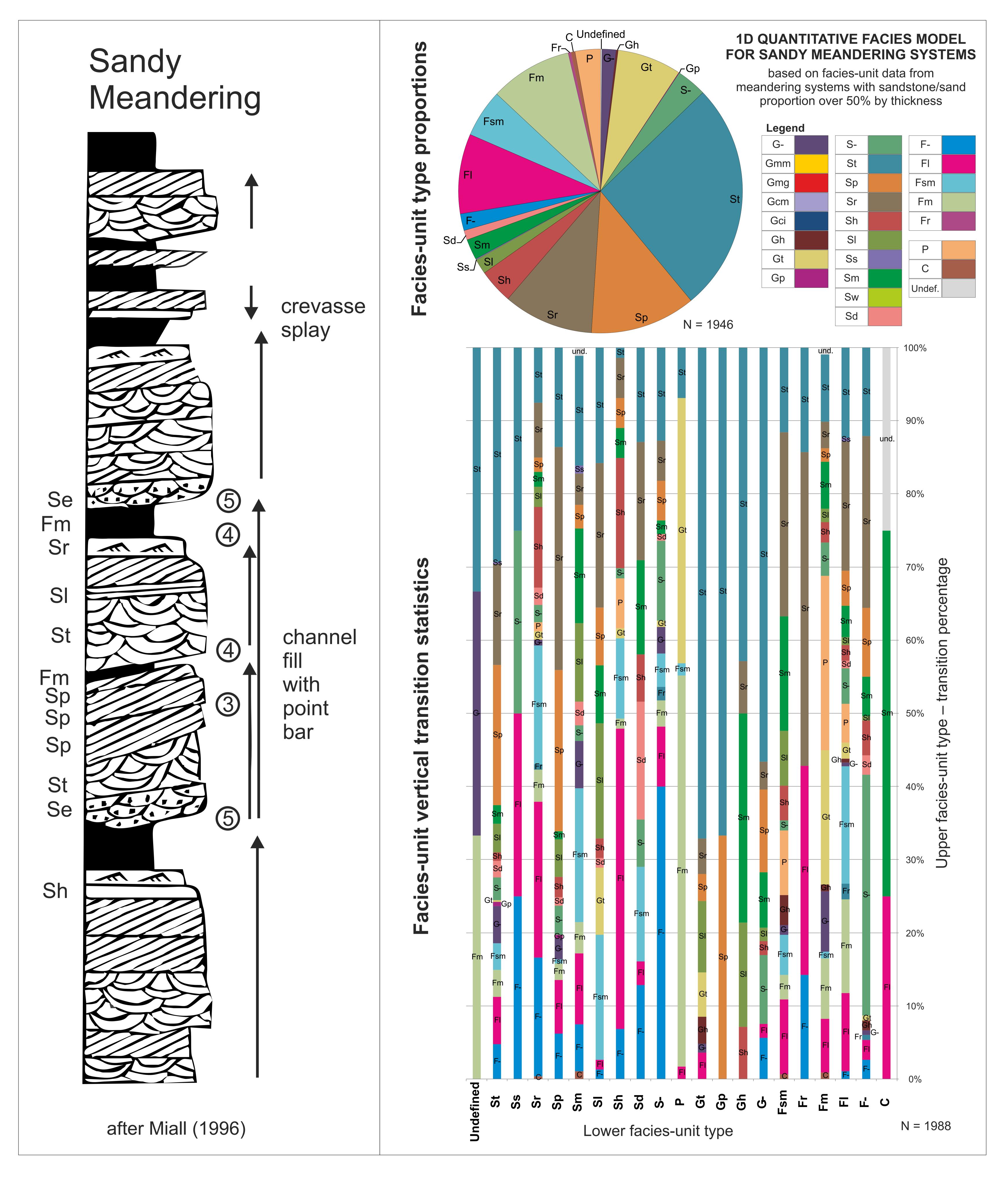
COMPARISON BETWEEN SYSTEMS CLASSIFIED AS TERMINAL FANS



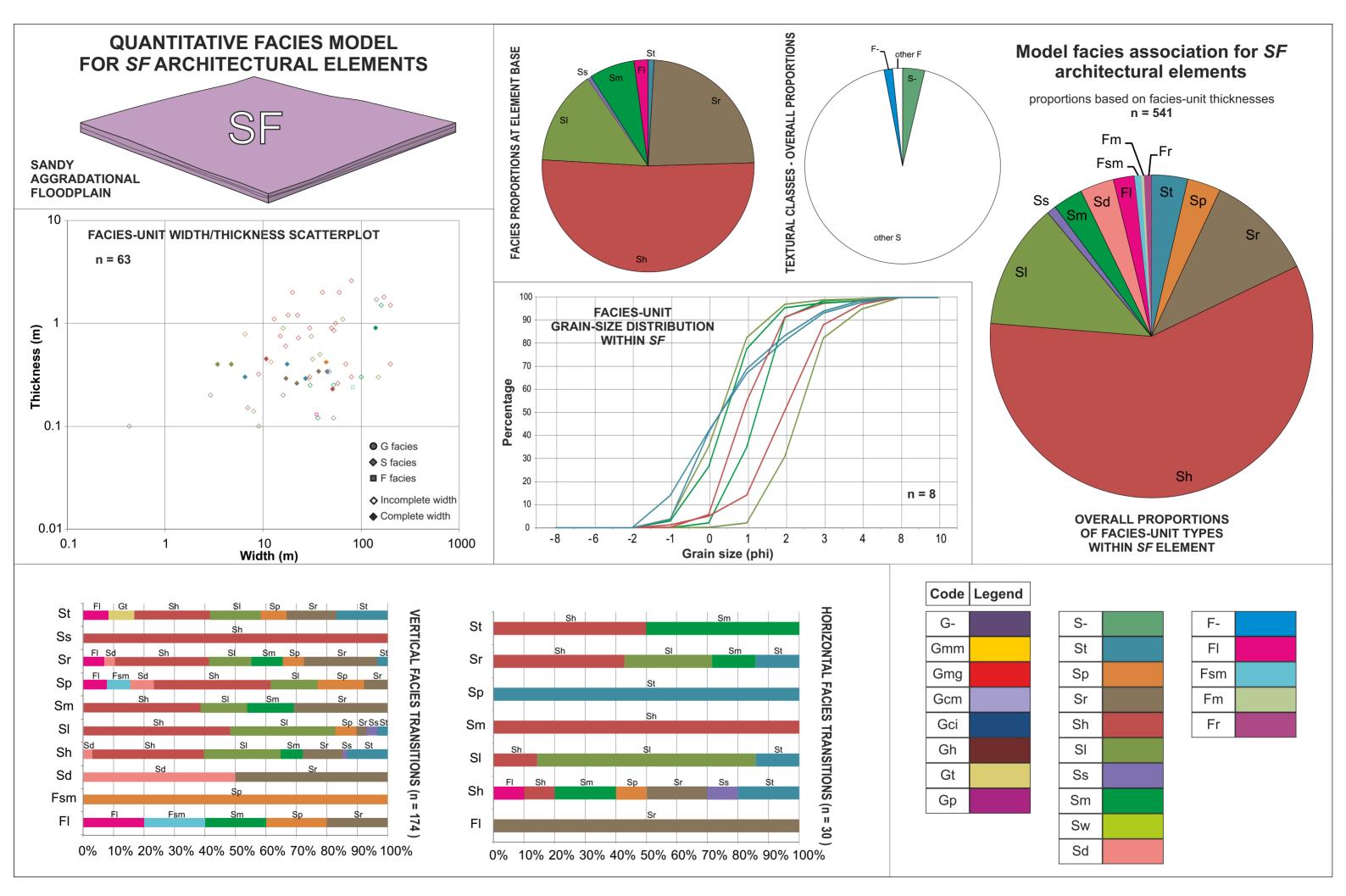




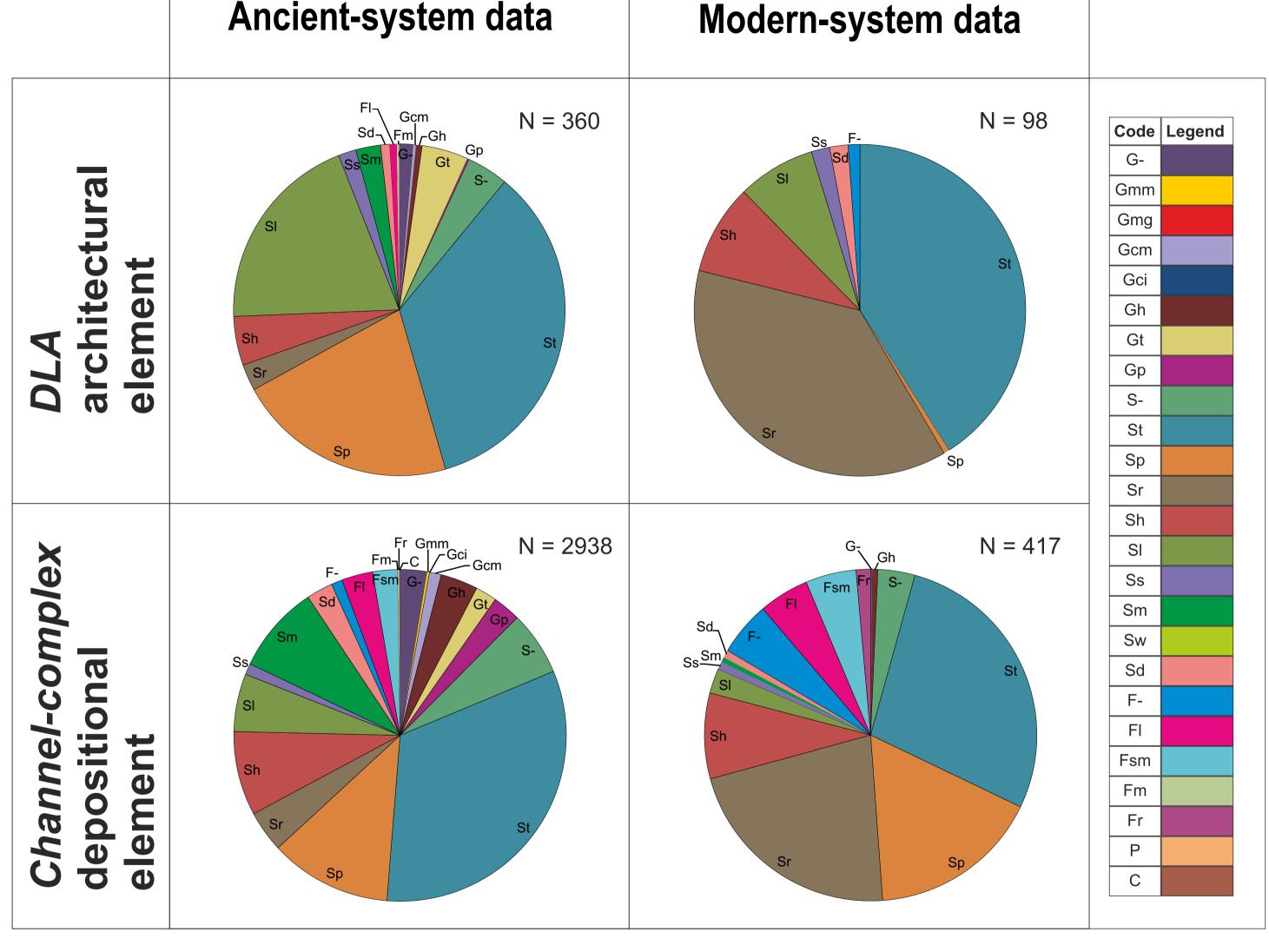
Models of architectural-element spatial relationships, in the form of pie-charts depicting transition counts between architectural-element types in the upwards, downwards, up-gradient, cross-gradient and down-gradient directions. a) transition statistics referring to downstream-accreting barforms; b) transition statistics referring to lateral-accretion barforms; cross-stream transitions conventionally refer to the right-hand direction, regardless of the dip-direction of accretion surfaces or migration direction of the barform; c) transition statistics referring to crevasse splays; lateral, upstream and downstream transitions have been grouped into horizontal transitions for convenience.



Comparison between the Miall's (1996) facies model for sandy meandering systems presented in the form of a vertical profile, on the left, and a corresponding FAKTS model, on the right. The FAKTS model has been built filtering the database on both a system parameter (meandering channel pattern) and a sedimentological feature (proportion of sandy facies units within subsets higher than 50% by thickness); lithofacies-type proportions are represented as a pie-chart, and were derived as the sum of the thickness of all facies units from adequate subsets (method 1 in Fig. 2 and in the text); vertical transition statistics are presented in the bar chart, quantifying the percentage of types of 'upper' facies units (colour-coded and labelled in the bars) stacked on top of a given type of 'lower' lithofacies (labels on the horizontal axis). In this case, results include 'undefined' lithofacies types, i.e. facies units (e.g. non-fluvial aeolian facies) that cannot be classified according to the adopted classification scheme (Table 2).



Partial quantitative information constituting a small-scale facies model of aggradational sheetflood-dominated sandy floodplain elements (SF architectural elements). As in Fig. 11, the model facies association of the element is described by overall lithofacies-type proportions, presented as pie-charts of textural classes and of 'texture + structure' facies-unit classes; proportions of facies types observed at the base of channel-fills are also given. Example cumulative grain-size distributions for facies units within SF elements are presented for different lithofacies types; the thickness and width of classified facies units within sandy aggradational floodplain elements is represented in the cross-plot; upwards and horizontal (cross-gradient + up-gradient) transition statistics are presented as bar charts quantifying the percentage of types of facies units (colour-coded and labelled in the bars) juxtaposed to a given type of facies unit (labels on the vertical axis) within SF elements.



Example facies associations for 'downstream- and lateral-accretion barforms' (DLA architectural elements) and 'channel-complex' depositional elements, as derived by separately considering data from ancient systems preserved in the rock record and modern river systems; results are presented as pie-charts quantifying facies-unit proportions derived as the sum of the thickness of all facies units from adequate subsets (method 1 in Fig. 2 and in the text).

case_ID reference_citation

lithostratigraphic_unit

- 1 Miall A. D. (1988) Architectural elements and Kayenta Fm.
- 2 Hornung J., Aigner T. (1999) Reservoir and a Middle-Upper Stubensandstein
- 3 Amorosi A., Pavesi M., Ricci Lucchi M., Sarti (-
- 4 Dalrymple M. (2001) Fluvial reservoir archite Straight Cliffs Fm.
- 5 Carter D. C. (2003) 3-D seismic geomorpholo Talang Akar Fm.
- 6 Meadows N. S. (2006) The correlation and se Ormskirk Sandstone Fm., Sherwood Sandstone Gp.
- 7 Pranter M. J., Cole R. D., Panjaitan H., Somm Lower Williams Fork Fm.
- 8 Johnson S. Y. (1984) Cyclic fluvial sedimental Bellingham Bay Mb., Chuckanut Fm.
- 9 Jones S. J., Frostick L. E., Astin T. R. (2001) Br Rio Vero Fm.
- 10 Hjellbakk A. (1997) Facies and fluvial archite Seglodden Mb., Båsnæring Fm.
- 11 Bristow C. S. (1993) Sedimentary structures (-
- 12 Robinson J. W., McCabe P. J. (1997) Sandsto Salt Wash Mb., Morrison Fm.
- 13 Tye R. S. (2004) Geomorphology: an approac-
- 14 Tye R. S. (2004) Geomorphology: an approac-
- 15 Tye R. S. (2004) Geomorphology: an approac-
- 16 Bridge J. S., Jalfin G. A., Georgieff S. M. (200(Bajo Barreal Fm.

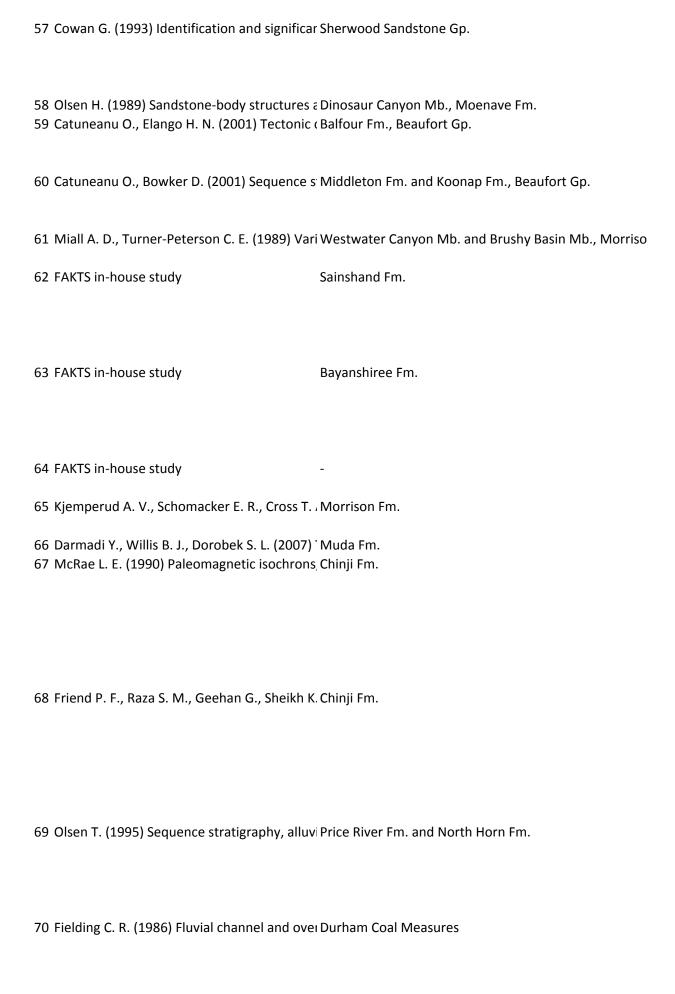
- 17 Jordan D. W., Pryor W. A. (1992) Hierarchica -
- 18 Bromley M. H. (1991) Architectural features Kayenta Fm.
- 19 Luttrell P. R. (1993) Basinwide sedimentatio Kayenta Fm.
- 20 Jo H. R. (2003) Depositional environments, a Sindong Gp.
- 21 Cuevas Gozalo M. C., Martinius A. W. (1993) Upper Unit, Tortola fluvial system
- 22 North C. P., Taylor K. S. (1996) Ephemeral-fli Kayenta Fm.
- 23 Cain S. A. (2009) Sedimentology and stratigr Organ Rock Fm.
- 24 Sanabria D. I. (2001) Sedimentology and Seq Kayenta Fm.
- 25 Stephens M. (1994) Architectural element aı Kayenta Fm.
- 26 FAKTS in-house study

- Kayenta Fm.
- 27 Abdullatif O. M. (1989) Channel-fill and shee-
- 28 Cuevas Martinez J. L., Cabrera Perez L., Marc Caspe Fm.
- 29 Fabuel-Perez I., Redfern J., Hodgetts D. (200! Oukaimeden Fm.
- 30 Parkash B., Awasthi A. K., Gohain K. (1983) L -
- 31 Fabuel-Perez I., Hodgetts D., Redfern J. (200! Oukaimeden Fm.
- 32 Tunbridge I. (1984) Facies model for a sandy Trentishoe Fm., Hangman Sandstone Gp.
- 33 Fielding C. R., Falkner A. J., Scott S. G. (1993) Rangal Coal Measures
- 34 Best J. L., Ashworth P. J., Bristow C. S., Roder-
- 35 Fielding C. R., Crane R. C. (1987) An applicati -
- 36 Friend P. F., Sinha R. (1993) Braiding and me -

- 37 Friend P. F., Sinha R. (1993) Braiding and me-38 Friend P. F., Sinha R. (1993) Braiding and me -39 FAKTS in-house study 40 Weerts H. J. T., Bierkens M. F. P. (1993) Geos-41 FAKTS in-house study Guarda Velha Fm. 42 Steel R. J., Thompson D. B. (1983) Structures Bunter Pebble Beds (Chester Pebble Beds Fm. and Canr 43 Tirsgaard H., Øxnevad I. E. I. (1998) Preserva Majût Mb., Eriksfjord Fm. 44 Pranter M. J., Ellison A. I., Cole R. D., Patters Lower Williams Fork Fm. 45 Donselaar M. E., Overeem I. (2008) Connecti Sariñena Fm. 46 Corbeanu R. M., Wizevich M. C., Bhattachar Ferron Sandstone Mb., Mancos Shale 47 Raynal J.-P., Kieffer G., Bardin G. (2004) Gark Melka Kunture Fm. 48 Tooth S., Nanson G. C. (2004) Forms and pro-49 Tooth S., Nanson G. C. (2004) Forms and pro-50 Hampton B. A., Horton B. K. (2007) Sheetflo Potoco Fm. 51 Labourdette R. (2011) Stratigraphy and stati Olson Mb., Escanilla Fm. 52 Holzförster F., Stollhofen H., Stanistreet I. G. Omingonde Fm. 53 Fillmore D. L., Lucas S. G., Simpson E. L. (201 Mauch Chunk Fm. 54 FAKTS in-house study Hawksmoor Fm. and Hollington Fm.
- 56 Brookfield M. E. (2008) Palaeoenvironments Annan Sandstone Fm., Sherwood Sandstone Gp.

Wilmslow Sandstone Fm. and Helsby Sandstone Fm., Sl

55 FAKTS in-house study



- 71 Opluštil S., Martínek K., Tasáryová Z. (2005) Kladno Fm. and Týnec Fm.
- 72 Reynolds A. D. (1999) Dimensions of paralic :-
- 73 Stewart D. J. (1983) Possible suspended-loac Wessex Fm., Wealden Gp.
- 74 Stewart D. J. (1983) Possible suspended-loac Fairlight Clay and Ashdown Beds Fm., Hastings Beds Gp
- 75 FAKTS in-house study

Undifferentiated Cutler Fm.

- 76 Hume T. M., Sherwood A. M., Campbell S. N. Hinuera Fm.
- 77 Ori G. G. (1982) Braided to meandering chan-
- 78 Martinius A. W., Nieuwenhuijs R. A. (1995) C Upper Unit, Tortola fluvial system
- 79 Martinius A. W. (2000) Labyrinthine facies as Upper Unit, Tortola fluvial system
- 80 Rygel M. C., Gibling M. R. (2006) Natural geo Joggins Fm.
- 81 Singh A., Bhardwaj B. D. (1991) Fluvial facies -
- 82 Kirk M. (1983) Bar development in a fluvial s Lower Coal Measures
- 83 Shukla U. K., Singh I. B., Sharma M., Sharma -
- 84 Ori G. G., Penney S. R. (1982) The stratigraph Templetown Fm. (Brownstown Head Mb. and Beenlea
- 85 Wood L. J. (2007) Quantitative seismic geom-
- 86 Sadler S. P., Kelly S. B. (1993) Fluvial process Gun Point Fm., Old Red Sandstone
- 87 Tunbridge I. (1981) Sandy high-energy flood Trentishoe Fm., Hangman Sandstone Gp.
- 88 Olsen H. (1987) Ancient ephemeral stream d Upper Bunter Sand, Bunter Sandstone Fm., Bacton Gp.

```
89 McKee E. D., Crosby E. J., Berryhill H. L. Jr. (1 -
```

- 90 Williams G. E. (1971) Flood deposits of the si-
- 91 Williams G. E. (1971) Flood deposits of the si-
- 92 Williams G. E. (1971) Flood deposits of the si-
- 93 Williams G. E. (1971) Flood deposits of the si-
- 94 Williams G. E. (1971) Flood deposits of the si-
- 95 Bhattacharyya A., Morad S. (1993) Proterozc Dhandraul Sandstone Fm., Kaimur Gp.
- 96 Singh I. B. (1977) Bedding structures in a channel sand bar of the Ganga River near Allahabad, Uttar
- 97 Long D. G. F. (2002) Aspects of Late Palaeopi Uairén Fm.
- 98 Sønderholm M., Tirsgaard H. (1998) Protero: Rivieradal Sandstones, Rivieradal Gp.
- 99 Dam G., Andreasen F. (1990) High-energy ep Holmestrand Fm., Ringerike Gp.
- 101 Yu X., Ma X., Qing H. (2002) Sedimentology a Yungang Fm.
- 102 Sánchez-Moya Y., Sopeña A., Ramos A. (1996 Bundsandstein
- 100 Viseras C., Soria J. M., Durán J. J., Pla S., Garr-
- 103 Limarino C., Tripaldi A., Marenssi S., Net L., F Vinchina Fm.
- 104 Ferguson R. J., Brierley G. J. (1999) Levee mc-
- 105 Ghazi S., Mountney N. P. (2009) Facies and a Warchha Sandstone Fm., Nilawahan Gp.
- 106 Mack G. H., Leeder M., Perez-Arlucea M., Ba Abo Fm.
- 107 Adams P. N., Slingerland R. L., Smith N. D. (21-
- 108 Adams P. N., Slingerland R. L., Smith N. D. (21-
- 109 Roberts E. M. (2007) Facies architecture and Kaiparowits Fm.
- 110 Kraus M. J., Middleton L. T. (1987) Contrastii Willwood Fm.
- 111 Kraus M. J., Middleton L. T. (1987) Contrastii Glenns Ferry Fm.

river -	nr_of_depositional_elements	nr_of_architectural_elements
-	31	274
-	83	38
-	241	-
-	-	-
-	16	-
-	-	-
-	-	-
-	0	37
-	2	7
Brahmaputra (Jamuna)	1	3
-	85	47
Colville	-	-
Kuparuk	-	-
Sagavanirktok -	3	0

Mississippi	1	15
-	0	22
-	0	8
-	62	30
-	0 54	72 6
		Ç
-	103	397
-	0	4
-	0	41
-	1	330
	•	
Gash -	2 85	0 23
-	20	147
Markanda	0	0
-	0	289
-	3	54
-	22	39
Duck many true (15 mm s)		
Brahmaputra (Jamuna)	0	1
	277	4
- Gandak	277 24 -	1
	- ·	

Burhi Gandak	28 -	
Baghmati	25 -	
Thomson (Cooper Creek)	3 - 14	305
- -	11 6	73 33
-	0 2	0
-	19	72
- -	1 0	7 0
Plenty	4 -	
Marshall	27 -	
-	14 297	21 15
-	28	69
-	0	0
-	0	0
-	0	0

- 5 24

		20
-	14 16	23 36
	10	30
-	4	23
-	7	51
-	0	0
-	0	0
South Saskatchewan	5 -	
-	601 -	
-	45 -	
-	203 -	
-	195	0

14

28

- 551 5

-	5	137
- -	175 4	112 21
-	14	29
-	0	0
-	0	0
Reno -	3 115	16 49
-	86	0
-	89	40
Ganges	10	0
-	2	7
Ganges -	0 21	2 34
-	31 -	
-	0	4
-	0	0
-	22	8

Bijou Creek Paralana Creek	7 1	9
The Wooldridge	1	0
Goyder Creek	1	0
Palmer Creek	1	0
The Finke	1	0
-	1	1
Ganges	1	1
-	2	0
-	0	0
-	0	0
-	7	20
-	128	64
	2	4
-	2	4
-	9	23
Tuross	2	13
-	0	15
_	15	16
	13	10
Columbia	24	12
Saskatchewan	62	31
-	148	56
-	3	22
	2	า
-	3	3

nr_or_racies_units	nr_ot_statisticai_paramete	rs
	38 -	
	463 -	
	72	
	72 -	
-	-	
-		8
-	-	
-		3
		110
-		110
	155 -	
	133 -	
	472 -	
	.,_	
	103 -	
-	-	
		6
-		6
-		5
		-
_		5
		<i>3</i>

51 -

10 -

237 -

-

57 2

5265 -

8 -

128 -

1763 -

117 -

1602 -

98 -

86 -

477 -

.

21 -

--

-

-- -

338 -280 1

229 -74 -

88 -

260 -89 -

215 -5 -

199 -132 -

298 -

934 -

35 -54 -

34 -

223 -

288 -

132 -

_

- 2

- -

77 -

300 -

681 -

28 -

-

37 -

46 -

36 -

85 -

40 -

-

109 -

65 -

1 -

7 -

2 -

2 -

2 -

70 -

27 -

47 -

130 -

254 -

_

101 -

74 -

100 -

71 -

484 -

157 -

-

372 -

_

additional literature

Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Hornung J., Aigner T. (2002) Reservoir architecture in a terminal alluvial plain: an outcrop analog stud Hornung J., Aigner T. (2002) Reservoir architecture in a terminal alluvial plain: an outcrop analog stud Bartolini C., Caputo R., Pieri M. (1996) Pliocene-Quaternary sedimentation in the Northern Apennine Carminati E., Martinelli G. (2002) Subsidence rates in the Po Plain, northern Italy: the relative impact of Carminati E., Doglioni D., Scrocca D. (2003) Apennines subduction-related subsidence of Venice (Italy) Wittmann H., Von Blanckenburg F., Kruesmann T., Norton K. P., Kubik P. W. (2007) Relation between Shanley K. W., McCabe P. J. (1993) Alluvial architecture in a sequence stratigraphic framework: a case Morley R. J. (1998) Palynological evidence for Tertiary plant dispersal in the SE Asian region in relation Tonkin P. C., Himawan R. (1999) Basement lithology and its control on sedimentation, trap formation Doust H., Sumner H. S. (2007) Petroleum systems in rift basins – a collective approach in Southeast As Meadows N. S., Beach A. (1993) Structural and climatic controls on facies distribution in a mixed fluvia Jackson D. I., Mulholland P. (1993) Tectonic and stratigraphic aspects of the East Irish Sea Basin and a Herries R. D., Cowan G. (1997) Challenging the 'sheetflood' myth: the role of water-table-controlled sa Brookfield M. E. (2008) Palaeoenvironments and palaeotectonics of the arid to hyperarid intracontine McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Europ Lorenz J. C., Heinze D. M., Clark J. A., Searls C. A. (1985) Determination of widths of meander-belt san-Gries R., Dolson J. C., Raynolds R. G. H. (1992) Structural and stratigraphic evolution and hydrocarbon Elder W. P., Kirkland J. (1993) Cretaceous paleogeography of the Colorado Plateau and adjacent areas Sommer N. K. (2007) Sandstone-body connectivity in a meandering-fluvial system: an example from t Ellison A. I. (2004) Numerical modeling of heterogeneity within a fluvial point-bar deposit using outcre Johnson S. Y. (1984) Stratigraphy, age, and paleogeography of the Eocene Chuckanut Formation, Nort Evans J., Ristow R. J. Jr. (1994) Depositional history of the southeastern outcrop belt of the Chuckanut Mustoe G. E. (2002) Eocene bird, reptile, and mammal tracks from the Chuckanut Formation, Northw Barbera X., Cabrera L., Marzo M., Pare J. M., Agusti J. (2001) A complete terrestrial Oligocene magnet Jones S. J. (2004) Tectonic controls on drainage evolution and development of terminal alluvial fans, s Hamer J. M. M., Sheldon N. D., Nichols G. J., Collinson M. E. (2007) Late Oligocene-Early Miocene pale Nystuen J. P., Andresen A., Kumpulainen R., Siedlecka A. (2008) Neoproterozoic basin evolution in Fer Røe S. L. (2003) Neoproterozoic peripheral-basin deposits in eastern Finnmark, N. Norway: stratigraph Drinkwater N. J., Pickering K. T., Siedlecka A. (1996) Deep-water fault-controlled sedimentation, Arctic Alam M. (1996) Subsidence of the Ganges-Brahmaputra delta of Bangladesh and associated drainage Allison M. A. (1998) Geologic framework and environmental status of the Ganges-Brahmaputra delta. Allison M. A., Khan S. R., Goodbred S. L. Jr., Kuehl S. A. (2003) Stratigraphic evolution of the late Holoc Tyler N., Ethridge F. G. (1983) Depositional setting of the Salt Wash Member of the Morrison Formatic Robinson J. W., McCabe P. J. (1998) Evolution of a braided river system: the Salt Wash Member of the Demko T. M., Currie B. S., Nicoll K. A. (2004) Regional paleoclimatic and stratigraphic implications of p Thornthwaite C. W. (1931) The climates of North America according to a new classification. Geog. Rev Walker H. J., Hudson P. F. (2003) Hydrologic and geomorphic processes in the Colville River delta, Alas Thornthwaite C. W. (1931) The climates of North America according to a new classification. Geog. Rev McNamara J. P., Kane D. L., Hinzman L. D. (1998) An analysis of streamflow hydrology in the Kuparuk Best H., McNamara J. P., Liberty L. (2005) Association of ice and river channel morphology determined Dery S. J., Stieglitz M., Rennermalm A. K., Wood E. F. (2005) The water budget of the Kuparuk River Ba Thornthwaite C. W. (1931) The climates of North America according to a new classification. Geog. Rev Keeley M. L., Light M. P. R. (1993) Basin evolution and prospectivity of the Argentine continental marg Rodriguez J. F. R., Littke R. (2001) Petroleum generation and accumulation in the Golfo San Jorge Basi Umazano A. M., Bellosi E. S., Visconti G., Melchor R. N. (2008) Mechanisms of aggradation in fluvial sy Umazano A. M., Bellosi E. S., Visconti G., Jalfin G. A., Melchor R. N. (2009) Sedimentary record of a Late Braile L. W., Hinze W. J., Keller G. R., Lidiak E. G., Sexton J. L. (1986) Tectonic development of the New Schweig E. S., Van Arsdale R. B. (1996) Neotectonics of the upper Mississippi embayment. Eng. Geol. 4 Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Chang K.-H., Suzuki K., Parka S.-O., Ishida K., Uno K. (2003) Recent advances in the cretaceous stratigra Lee Y. I., Lim D. H. (2008) Sandstone diagenesis of the Lower Cretaceous Sindong Group, Gyeongsang Lee Y. I. (2008) Paleogeographic reconstructions of the East Asia continental margin during the middle Martinius A. W. (2000) Labyrinthine facies architecture of the Tortola fluvial system and controls on d Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Cain S. A., Mountney N. P. (2009) Spatial and temporal evolution of a terminal fluvial fan system: the Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Miall A. D. (1988) Architectural elements and bounding surfaces in fluvial deposits: anatomy of the Ka Bromley M. H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plat Luttrell P. R. (1993) Basinwide sedimentation and the continuum of paleoflow in ancient river system: Beck M. E. Jr., Burmester R. F., Housen B. A. (2003) The red bed controversy revisited: shape analysis Davidson S. K., North C. P. (2009) Geomorphological regional curves for prediction of drainage area ar Salama R. B. (1985) Buried troughs, grabens and rifts in Sudan. J. Afr. Earth Sci. 3, 381-390. Cuevas J. L., Arbues P., Cabrera L., Marzo M. (2004) The Caspe Formation revisited (Upper Oligocene-

Cuevas J. L., Arbues P., Cabrera L., Marzo M. (2004) The Caspe Formation revisited (Upper Oligocene-Benaouiss N., Courel L., Beauchamp J. (1996) Rift-controlled fluvial/tidal transitional series in the Ouk Kumar S., Wesnousky S. G., Rockwell T., Ragona D., Thakur V. C., Seitz G. (2001) Earthquake Recurrent Benaouiss N., Courel L., Beauchamp J. (1996) Rift-controlled fluvial/tidal transitional series in the Ouk Fabuel-Perez I., Redfern J., Hodgetts D. (2009) Sedimentology of an intra-montane rift-controlled fluv Van Der Voo R. (1983) Paleomagnetic Constraints on the Assembly of the Old Red Continent. Tectono Friend P. F., Williams B. P. J., Ford M., Williams E. A. (2000) Kinematics and dynamics of Old Red Sands Michaelsen P., Henderson R. A., Crosdale P. J., Mikkelsen S. O. (2000) Facies architecture and depositi Michaelsen P. (2002) Mass extinction of peat-forming plants and the effect on fuvial styles across the Alam M. (1996) Subsidence of the Ganges–Brahmaputra delta of Bangladesh and associated drainage Allison M. A. (1998) Geologic framework and environmantal status of the Ganges-Brahmaputra delta. Allison M. A., Khan S. R., Goodbred S. L. Jr., Kuehl S. A. (2003) Stratigraphic evolution of the late Holoc

Gadgil S., Joshi N. V. (1983) Climatic clusters of the Indian Region. J. Climatol. 3, 47-63. Sinha R., Friend P. F. (1994) River systems and their sediment flux, Indo-Gangetic plains, Northern Bih

Gupta S. (1997) Himalayan drainage patterns and the origin of fluvial megafans in the Ganges foreland Gadgil S., Joshi N. V. (1983) Climatic clusters of the Indian Region. J. Climatol. 3, 47-63.

Sinha R., Friend P. F. (1994) River systems and their sediment flux, Indo-Gangetic plains, Northern Bih Gupta S. (1997) Himalayan drainage patterns and the origin of fluvial megafans in the Ganges foreland Gadgil S., Joshi N. V. (1983) Climatic clusters of the Indian Region. J. Climatol. 3, 47-63.

Sinha R., Friend P. F. (1994) River systems and their sediment flux, Indo-Gangetic plains, Northern Bih Gupta S. (1997) Himalayan drainage patterns and the origin of fluvial megafans in the Ganges foreland Nanson G. C., Rust B. R., Taylor G (1986) Coexistent mud braids and anastomosing channels in an arid Geluk M. C., Duin E. J. Th., Dusar M., Rijkers R. H. B., van den Berg M. W., van Rooijen P. (1994) Stratig Cohen K. M. (2003) Differential subsidence within a coastal prism: Late-Glacial — Holocene tectonics in Cohen K. M., Gouw M. J. P., Holten J. P. (2005) Fluvio-Deltaic Floodbasin Deposits Recording Different

Brookfield M. E. (2008) Palaeoenvironments and palaeotectonics of the arid to hyperarid intracontine McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Europe MacDonald R., Upton B. G. J. (1993) The Proterozoic Gardar rift zone, south Greenland: comparisons of Lorenz J. C., Heinze D. M., Clark J. A., Searls C. A. (1985) Determination of widths of meander-belt sanderies R., Dolson J. C., Raynolds R. G. H. (1992) Structural and stratigraphic evolution and hydrocarbon Elder W. P., Kirkland J. (1993) Cretaceous paleogeography of the Colorado Plateau and adjacent areas Sommer N. K. (2007) Sandstone-body connectivity in a meandering-fluvial system: an example from the Ellison A. I. (2004) Numerical modeling of heterogeneity within a fluvial point-bar deposit using outcrous Nichols G. J., Hirst J. P. (1998) Alluvial fans and fluvial distributary systems, Oligo-Miocene, northern S. Pérez-Rivarés F. J., Garcés M., Arenas C., Pardo G. (2002) Magnetocronología de la sucesión miocena Fisher J. A., Nichols G. J., Waltham D. A. (2007) Unconfined flow deposits in distal sectors of fluvial distramer J. M. M., Sheldon N. D., Nichols G. J. (2007) Global aridity during the Early Miocene? A terrestr Hamer J. M. M., Sheldon N. D., Nichols G. J., Collinson M. E. (2007) Late Oligocene—Early Miocene pale Bhattacharya J. P., Tye R. S. (2004) Searching for Modern Ferron Analogs and Application to Subsurfac Chavaillon J., Piperno M. (2004) Studies on the Early Paleolithic site of Melka Kunture, Ethiopia.

Piperno M., Collina C., Gallotti R., Raynal J.-P., Kieffer G., Le Bourdonnec F. X., Poupeau G., Geraads D. Morgan L. E., Renne P. R., Kieffer G., Piperno M., Gallotti R., Raynal J.-P. (2012) A chronological frame Thornthwaite C. W. (1933) The climates of the Earth. Geog. Rev. 23, 433-440.

Peel M. C., Finlayson B. L., McMahon T. A. (2007) Updated world map of the Köppen-Geiger climate c Thornthwaite C. W. (1933) The climates of the Earth. Geog. Rev. 23, 433-440.

Peel M. C., Finlayson B. L., McMahon T. A. (2007) Updated world map of the Köppen-Geiger climate c Horton B. K., DeCelles P. G. (1997) The modern foreland basin system adjacent to the Central Andes. Bentham P. A., Burbank D. W., Puigdefàbregas C. (1992) Temporal and spatial controls on the alluvial Bentham P. A., Talling P. J., Burbank D. W. (1993) Braided stream and flood-plain deposition in a rapic Kjemperud A. V., Schomacker E., Brendsdal A., Fält L.-M., Jahren J. S., Nystuen P. J., Puigdefàbregas C. Mochales T., Barnolas A., Pueyo E. L., Serra-Kiel J., Casas A. M., Samsó J. M., Ramajo J., Sanjuán J. (201 Zerfass H., Chemale F. Jr., Lavina E. (2005) Tectonic control of the Triassic Santa Maria Supersequence Smith R. M. H., Swart R. (2002) Changing fluvial environments and vertebrate taphonomy in response Hoque M. U. (1968) Sedimentologic and paleocurrent study of the Mauch Chunk sandstones (Mississi DiVenere V. J., Opdyke N. D. (1991) Magnetic polarity stratigraphy in the uppermost Mississippian Ma Meadows N. S., Beach A. (1993) Structural and climatic controls on facies distribution in a mixed fluvia Brookfield M. E. (2008) Palaeoenvironments and palaeotectonics of the arid to hyperarid intracontine McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Euro Thompson D. B. (1970) Sedimentation of the Triassic (Scythian) Red Pebbly Sandstones in the Cheshir Evans D. J., Rees J. G., Holloway S. (1993) The Permian to Jurassic stratigraphy and structural evolution Meadows N. S., Beach A. (1993) Structural and climatic controls on facies distribution in a mixed fluvia Brookfield M. E. (2008) Palaeoenvironments and palaeotectonics of the arid to hyperarid intracontine McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Euro McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Euro Meadows N. S., Beach A. (1993) Structural and climatic controls on facies distribution in a mixed fluvia Herries R. D., Cowan G. (1997) Challenging the 'sheetflood' myth: the role of water-table-controlled sa Brookfield M. E. (2008) Palaeoenvironments and palaeotectonics of the arid to hyperarid intracontine McKie T., Williams B. (2009) Triassic palaeogeography and fluvial dispersal across the northwest Europ Tanner L. H., Lucas S. G. (2007) The Moenave Formation: Sedimentologic and stratigraphic context of Hiller N., Stavrakis N. (1984) Permo-Triassic fluvial systems in the southeastern Karoo Basin, South Afr Catuneanu O., Hancox P. J., Rubidge B. S. (1998) Reciprocal flexural behaviour and contrasting stratigr Retallack G. J., Smith R. M. H., Ward P. D. (2003) Vertebrate extinction across Permian-Triassic bounds Yemane K., Kelts K. (1990) A short review of palaeoenvironments for Lower Beaufort (Upper Permian Catuneanu O., Hancox P. J., Rubidge B. S. (1998) Reciprocal flexural behaviour and contrasting stratigi Bordy E. M., Linkermann S., Prevec R. (2011) Palaeoecological aspects of some invertebrate trace foss Kowallis B. J., Christiansen E. H., Deino A. L., Peterson F., Turner C. E., Kunk M. J., Obradovich J. D. (19! Demko T. M., Currie B. S., Nicoll K. A. (2004) Regional paleoclimatic and stratigraphic implications of p Jerzykiewicz T. (1998) Okavango Oasis, Kalahari Desert: a contemporary analogue for the Late Cretace Graham S. A., Hendrix M. S., Johnson C. L., Badamgarav D., Badarch G., Amory J., Porter M., Barsbold Johnson C. L. (2004) Polyphase evolution of the East Gobi basin: sedimentary and structural records o Prost G. L. (2004) Tectonics and hydrocarbon systems of the East Gobi basin, Mongolia. AAPG Bull. 88 Davaa B.-A. (2010) Geodynamic development and hydrocarbon potential of the Tamtsag Basin, Easter Jerzykiewicz T. (1998) Okavango Oasis, Kalahari Desert: a contemporary analogue for the Late Cretace Graham S. A., Hendrix M. S., Johnson C. L., Badamgarav D., Badarch G., Amory J., Porter M., Barsbold Johnson C. L. (2004) Polyphase evolution of the East Gobi basin: sedimentary and structural records o Prost G. L. (2004) Tectonics and hydrocarbon systems of the East Gobi basin, Mongolia. AAPG Bull. 88 Davaa B.-A. (2010) Geodynamic development and hydrocarbon potential of the Tamtsag Basin, Easter Cant D. J., Walker R. G. (1978) Fluvial processes and facies sequences in the sandy braided South Sask Ashworth P. J., Sambrook Smith G. H., Best J. L., Bridge J. S., Lane S. N., Lunt I. A., Reesink A. J. H., Simi Kowallis B. J., Christiansen E. H., Deino A. L., Peterson F., Turner C. E., Kunk M. J., Obradovich J. D. (19: Demko T. M., Currie B. S., Nicoll K. A. (2004) Regional paleoclimatic and stratigraphic implications of p Tjia H. D., Liew K. K. (1996) Changes in tectonic stress field in northern Sunda Shelf basins. In: Hall R., I Willis B. J. (1993) Ancient river systems in the Himalayan foredeep, Chinji Village area, northern Pakist Willis B. J. (1993) Evolution of Miocene fluvial systems in the Himalayan foredeep through a two kilon Willis B. J., Behrensmeyer A. K. (1994) Architecture of Miocene overbank deposits in northern Pakista Badgley C., Behrensmeyer A. K. (1995) Two long geological records of continental ecosystems. Paleog Zaleha M. J. (1997) Intra- and extrabasinal controls on fluvial deposition in the Miocene Indo-Gangetic Friend P. F., Raza S. M., Geehan G., Sheikh K. A. (2001) Intra- and extrabasinal controls on fluvial depo Barry J. C., Morgan M. E., Flynn L. J., Pilbeam D., Behrensmeyer A. K., Raza S. M., Khan I. A., Badgley C. McRae L. E. (1990) Paleomagnetic isochrons, unsteadiness, and non-uniformity of sedimentation in M Willis B. J. (1993) Ancient river systems in the Himalayan foredeep, Chinji Village area, northern Pakist Willis B. J. (1993) Evolution of Miocene fluvial systems in the Himalayan foredeep through a two kilon Willis B. J., Behrensmeyer A. K. (1994) Architecture of Miocene overbank deposits in northern Pakista Badgley C., Behrensmeyer A. K. (1995) Two long geological records of continental ecosystems. Paleog Zaleha M. J. (1997) Intra- and extrabasinal controls on fluvial deposition in the Miocene Indo-Gangetic Barry J. C., Morgan M. E., Flynn L. J., Pilbeam D., Behrensmeyer A. K., Raza S. M., Khan I. A., Badgley C. Bracken B., Picard M. D. (1984) Trace fossils from Cretaceous/Tertiary North Horn Formation in centra Difley R. L., Edale A. A. (2002) Footprints of Utah's Last Dinosaurs: Track Beds in the Upper Cretaceous Horton B. K., Constenius K. N., DeCelles P. G. (2004) Tectonic control on coarse-grained foreland-basir Retallack G. J. (2005) Pedogenic carbonate proxies for amount and seasonality of precipitation in pale Roberts E. M., Deino A. L., Chan M. A. (2005) 40Ar/39Ar age of the Kaiparowits Formation, southern L Broadhurst F. M., Simpson I. M., Hardy P. G. (1980) Seasonal sedimentation in the Upper Carboniferc Haszeldine R. S., Anderton R. (1980) A braidplain facies model for the Westphalian B Coal Measures o Leeder M. R. (1982) Upper Palaeozoic basins of the British Isles-Caledonide inheritance versus Hercyn Fielding C. R. (1984) A coal depositional model for the Durham Coal Measures of NE England. J. Geol.:

Hess J. C., Lippolt H. J. (1986) 40Ar/39Ar ages of Tonstein and tuff sanidines: new calibration point for Besly B. M., Fielding C. R. (1989) Palaeosols in Westphalian Coal-bearing and red-bed sequences, cent Flint S., Aitken J., Hampson G. (1995) Application of sequence stratigraphy to coal-bearing coastal plai Birkenmajer K., Krs M., Nairn A. E. M. (1968) A Paleomagnetic Study of Upper Carboniferous Rocks frc Hess J. C., Lippolt H. J. (1986) 40Ar/39Ar ages of Tonstein and tuff sanidines: new calibration point for Bashfort A. R., Drábková J., Opluštil S., Gibling M. R., Falcon-Lang H. J. (2011) Landscape gradients and

Ruffell A. H., Batten D. J. (1990) The Barremian-Aptian arid phase in western Europe. Paleogeo., Paleo Hawkes P. W., Fraser A. J., Einchcomb C. C. G. (1998) The tectono-stratigraphic development and expl Haywood A. M., Valdes P. J., Markwick P. J. (2004) Cretaceous (Wealden) climates: a modelling perspe Brasier M., Cotton L., Yenney I. (2009) First report of amber with spider webs and microbial inclusions Sweetman S. C., Insole A. N. (2010) The plant debris beds of the Early Cretaceous (Barremian) Wessex Ruffell A. H., Batten D. J. (1990) The Barremian-Aptian arid phase in western Europe. Paleogeo., Paleo Hawkes P. W., Fraser A. J., Einchcomb C. C. G. (1998) The tectono-stratigraphic development and expl Haywood A. M., Valdes P. J., Markwick P. J. (2004) Cretaceous (Wealden) climates: a modelling perspe Brasier M., Cotton L., Yenney I. (2009) First report of amber with spider webs and microbial inclusions Sweetman S. C., Insole A. N. (2010) The plant debris beds of the Early Cretaceous (Barremian) Wessex Nuccio V. F., Condon S.M. (1996) Burial and thermal history of the Paradox Basin, Utah and Colorado, Condon S.M. (1997) Geology of the Pennsylvanian and Permian Cutler Group and Permian Kaibab Lim Schofield J. C. (1965) The Hinuera Formation and associated Quaternary events. New Zeal. J. Geol. Ge McGlone M. S., Nelson C. S., Hume T. M. (1978) Palynology, age and environmental significance of soi McGlone M. S., Nelson C. S., Todd A. J. (1984) Vegetation history and environmental significance of pr Carminati E., Doglioni D., Scrocca D. (2003) Apennines subduction-related subsidence of Venice (Italy) Cuevas Gozalo M. C., Martinius A. W. (1993) Outcrop data-base for the geological characterization of Martinius A. W. (2000) Labyrinthine facies architecture of the Tortola fluvial system and controls on d Cuevas Gozalo M. C., Martinius A. W. (1993) Outcrop data-base for the geological characterization of Martinius A. W., Nieuwenhuijs R. A. (1995) Geological description of flow units in channel sandstones Davies S. J., Gibling M. R. (2003) Architecture of coastal and alluvial deposits in an extensional basin: t Falcon-Lang H. J. (2003) Late Carboniferous tropical dryland vegetation in an alluvial-plain setting, Jog Falcon-Lang H. J. (2003) Response of Late Carboniferous tropical vegetation to transgressive-regressiv Falcon-Lang H. J., Rygel M. C., Calder J. H., Gibling M. R. (2004) An early Pennsylvanian waterhole dep Davies S. J., Gibling M. R., Rygel M. C., Calder J. H., Skilliter D. M. (2005) The Pennsylvanian Joggins For Gadgil S., Joshi N. V. (1983) Climatic clusters of the Indian Region. J. Climatol. 3, 47-63.

Singh M., Singh I. B., Müller G. (2007) Sediment characteristics and transportation dynamics of the Ga Leeder M. R. (1982) Upper Palaeozoic basins of the British Isles-Caledonide inheritance versus Hercyn Waters C. N., Browne M. A. E., Dean M. T., Powell J. H. (2007) Lithostratigraphical framework for Carb Singh I. B., Jaiswal M., Singhvi A. K., Singh B. K. (2003) Rapid subsidence of western Ganga plain during Graham J. R. (1983) Analysis of the Upper Devonian Munster Basin, an example of a fluvial distributar Van Der Voo R. (1983) Paleomagnetic constraints on the assembly of the Old Red Continent. Tectonol Williams B. P. J., Insole A. N., Bennett M. C. (1991) Geological excursion guide 8: southern County We: Williams E. A., Sergeev S. A., Stössel I., Ford M. (1997) An Eifelian U-Pb zircon date for the Enagh Tuff Galloway W. E., Ganey-Curry P. E., Li X., Buffler R. T. (2000) Cenozoic depositional history of the Gulf c Zeng H., Hentz T. F. (2004) High-frequency sequence stratigraphy from seismic sedimentology: applied Graham J. R. (1983) Analysis of the Upper Devonian Munster Basin, an example of a fluvial distributar Van Der Voo R. (1983) Paleomagnetic constraints on the assembly of the Old Red Continent. Tectonol Kelly S. B., Sadler S. P. (1995) Equilibrium and response to climatic and tectonic forcing: a study of allu Van Der Voo R. (1983) Paleomagnetic Constraints on the Assembly of the Old Red Continent. Tectono Tunbridge I. (1984) Facies model for a sandy ephemeral stream and clay playa complex; the Middle D Friend P. F., Williams B. P. J., Ford M., Williams E. A. (2000) Kinematics and dynamics of Old Red Sands Clemmensen L. B. (1985) Desert sand plain and sabkha deposits from the Bunter Sandstone Formatio Van der Zwan C. J., Spaak P. (1992) Lower to Middle Triassic sequence stratigraphy and climatology of Scheck M., Bayer U. (1999) Evolution of the Northeast German Basin – inferences from a 3D structura Weibel R., Friis H. (2004) Opaque minerals as keys for distinguishing oxidising and reducing diagenetic North C. P., Davidson S. K. (2012) Unconfined alluvial flow processes: recognition and interpretation c

-

_

Williams G. E. (1970) The Central Australian stream floods of February-March 1967. J. Hydrol. 11, 185 Bose P. K., Sarkar S., Chakrabarty S., Banerjee S. (2001) Overview of the meso- to neoproterozoic evol Kumar A., Gopalan K., Rajagopalan G. (2001) Age of the Lower Vindhyan sediments, Central India. Cur Chakraborty C. (2006) Proterozoic intracontinental basin: The Vindhyan example. J. Earth Syst. Sci. 11 Ray J. S. (2006) Age of the Vindhyan Supergroup: A review of recent findings. J. Earth Syst. Sci. 115, 14 Gadgil S., Joshi N. V. (1983) Climatic clusters of the Indian Region. J. Climatol. 3, 47-63.

Singh M., Singh I. B., Müller G. (2007) Sediment characteristics and transportation dynamics of the Ga Santos J. O. S., Potter P. E., Reis N. J., Hartmann L. A., Fletcher I. R., McNaughton N. J. (2003) Age, sour Soper N. J. (1994) Neoproterozoic sedimentation on the northeast margin of Laurentia and the openin Smith M. P., Higgins A. K., Soper N. J., Sønderholm M. (2004) The Neoproterozoic Rivieradal Group of Higgins A. K., Leslie A. G. (2008) Architecture and evolution of the East Greenland Caledonides — An in Nystuen J. P., Andresen A., Kumpulainen R. A., Siedlecka A.(2008) Neoproterozoic basin evolution in F Turner P., Whitaker V. H. McD. (1976) Petrology and provenance of late Silurian fluviatile sandstones

_

Ramos A., Sopeña A., Perez-Arlucea M. (1986) Evolution of Bundsandstein fluvial sedimentation in the Sopeña A., Ramos A., Perez-Arlucea M. (1989) Permian and Triassic Fluvial Systems in Central Spain. In López-Gómez J., Arche A. (1993) Sequence stratigraphic analysis and paleogeographic interpretation of Sopeña A., Sánchez-Moya Y. (1997) Tectonic systems tract and depositional architecture of the weste Van Wees J. D., Arche A., Beijdorff C. D., J.López-Gómez J., Cloetingh S. A. P. L. (1998) Temporal and Space López-Gómez J., Arche A., Marzo M., Durand M. (2005) Stratigraphical and palaeogeographical significal J. M., Fernández J., Viseras C. (1999) Late Miocene stratigraphy and palaeogeographic evolution Pla-Pueyo S., Gierlowski-Kordesch E. H., Viseras C., Soria J. M. (2009) Major controls on sedimentation Agustí J., Blain H.-A., Furió M., De Marfá R., Santos-Cubedo A. (2010) The early Pleistocene small verte Pla-Pueyo S., Viseras C., Soria J. M., Tent-Manclús J. E., Arribas A. (2011) A stratigraphic framework fo Tabutt K. T., Naeser C. W., Jordan T. E., Cerveny P.F. (1989) New fission track ages of Mio-Pliocene tuf Collo G., Dávila F. M., Nóbile J., Astini R. A. (2008) Burial history and estimation of ancient thermal gra Melchor R. N., Genise J. F., Farina J. L., Sánchez M. V., Sarzetti L., Visconti G. (2010) Large striated buru Ciccioli P. L., Limarino C. O., Marenssi S. A., Tedesco A. M., Tripaldi A. (2011) Tectosedimentary evolut

-

Ghazi S. (2009) Sedimentology and stratigraphic evolution of the Early Permian Warchha Sandstone, § Ghazi S., Mountney N. P. (2010) Subsurface lithofacies analysis of the fluvial Early Permian Warchha S Ghazi S., Mountney N. P. (2011) Petrography and provenance of the Early Permian Fluvial Warchha Sa Hunt A. (1983) Plant fossils and lithostratigraphy of the Abo Formation (Lower Permian) in the Socord Mack G. H., Cole D. R., Giordano T. H., Schaal W. C., Barcelos J. D. (1991) Paleoclimatic controls on sta Bensing J. P., Mozley P. S., Dunbar N. W. (2005) Importance of clay in iron transport and sediment red DiMichele W. A., Chaney D. S., Nelson W. J., Lucas S. G., Looy C. V., Quick K., Jun W. (2007) A low diver Thornthwaite C. W. (1931) The climates of North America according to a new classification. Geog. Rev Lawton T. F., Pollock S. L., Robinson R. A. J. (2003) Integrating sandstone petrology and nonmarine sec Roberts E. M., Deino A. L., Chan M. A. (2005) 40Ar/39Ar age of the Kaiparowits Formation, southern L Bown T. M., Kraus M. J. (1983) Lower Eocene alluvial paleosols (Willwood Formation, northwest Wyor Bown T. M., Kraus M. J. (1983) Ichnofossils of the alluvial Willwood Formation (lower Eocene), Bighorn Kraus M. J. (1996) Avulsion deposits in Lower Eocene alluvial rocks, Bighorn Basin, Wyoming. J. Sed. R Wood S. H., Clemens D. M. (2002) Geologic and tectonic history of the western Snake River Plain, Idal

eau, USA: relationships to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99.

: Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434.

of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335-344.

nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792."

y (Upper Triassic, southern Germany) part I: sedimentology and petrophysics. Journ. Petr. Geol. 25, 3-30 y (Upper Triassic, southern Germany) part II: ciclicity, controls and models. Journ. Petr. Geol. 25, 151-17 Foredeep and related denudation. Geol. Mag. 133, 255-273.

of natural and anthropogenic causation. Eng Geol. 66, 241-255.

. Geophys. Res. Lett. 30, 1717.

rock uplift and denudation from cosmogenic nuclides in river sediment in the Central Alps of Switzerlan history from the Upper Cretaceous of southern Utah, USA. In: Flint S. S., Bryant I. D. (eds.) The geologic to plate tectonics and climate. In: Hall R. Holloway J. D. (eds.) Biogeography and Geological Evolution o and hydrocarbon migration, Widuri-Intan oilfields, SE Sumatra. Journ. Petr. Geol. 22, 141-165. sian basins. Petr. Geosc. 13, 127-144."

al and aeolian reservoir: the Triassic Sherwood Sandstone in the Irish Sea. In: North C. P., Prosser D. J. (Edjacent areas: contrasts in their post-Carboniferous structural styles. In: Parker J. R. (ed.) Petroleum Geabkha deposits in redefining the depositional model for the Ormskirk Sandstone Formation (Lower Triasental latest Permian-late Triassic Solway basin (U.K.). Sed. Geol. 210, 27-47.

pean Basins. Geol. J. 44, 711-741."

dstone reservoirs from vertical downhole data, Mesaverde Group, Piceance Creek Basin, Colorado. AAP distribution, Rocky Mountain foreland. In: Macqueen R.W., Leckie D. A. (eds.) Foreland basins and fold in: In: Morales M. (ed.) Aspects of Mesozoic geology and paleontology of the Colorado Plateau. Museum he Williams Fork Formation, Piceance Basin, Colorado. MSc Thesis, University of Texas at Austin.

op and lidar data: Williams Fork Formation, Piceance Basin, Colorado. MSc Thesis, University of Texas at hwest Washington. Can. J. Earth Sci. 21, 92-106.

: Formation: implications for the Darrington - Devil's Mountain and Straight Creek fault zones, Washingt est Washington. Palaios 17, 403-413."

obiostratigraphy from the Ebro Basin, Spain. Earth Plan. Sci. Let. 187, 1-16.

outhern Pyrenees, Spain. Terra Nova 16, 121–127.

eosols of distal fluvial systems, Ebro Basin, Spain. Palaeogeo., Palaeoclim., Palaeoecol. 247, 220–235." noscandia, East Greenland and Svalbard. Episodes 31, 35-43.

nic revision and palaeotectonic implications. Norw. J. Geol. 83, 259-274.

c Norway and Russia: response to Late Proterozoic rifting and the opening of the lapetus Ocean. J. Geol., sedimentation, and salinity problems. In: Milliman J. D., Haq B. U. (Eds.) Sea-Level Rise and Coastal Sul J. Cost. Res.14, 826-836.

zene Ganges-Brahmaputra lower delta plain. Sed. Geol. 155, 317-342.

on, Southwest Colorado. J. Sed. Petr. 53, 67-82.

! Morrison Formation (Jurassic) in southern Utah. In: Shanley K. W., McCabe P. J. (Eds.) Relative role of € aleosols and fluvial/overbank architecture in the Morrison Formation (Upper Jurassic), Western Interio *t*. 21, 633-655.

ska. Geomorphology 56, 291-303."

i. 21, 633-655.

River Basin Arctic Alaska: a nested watershed approach. J. Hydrol. 206, 39-57.

dusing ground-penetrating radar in the Kuparuk River, Alaska. Arc. Ant. Alp. Res. 37, 157-162.

asin, Alaska. J. Hydrometeor. 6, 633-655. "

i. 21, 633-655.

zin. J. Petr. Geol. 16, 451-464.

n, Argentina: a basin modeling study. Mar. Pet. Geol. 18, 995-1028.

stems influenced by explosive volcanism: An example from the Upper Cretaceous Bajo Barreal Formatice Cretaceous volcanic arc in central Patagonia: petrography, geochemistry and provenance of fluvial vol

Madrid rift complex, Mississippi Embayment, North America. Tectonophysics 131, 1-21. 15. 185-203." yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. : Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792." yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. eau, USA: relationship to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792." aphy of Korea. J. Asian Earth Sci. 21, 937-948. Basin, southeastern Korea: Implications for compositional and paleoenvironmental controls. Island Arc e to late Mesozoic. Island Arc 17, 458-470." eposition (Late Oligocene-Early Miocene, Loranca Basin, Spain). J. Sed. Res. 70, 850-867. yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. eau, USA: relationship to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99. : Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792." Permian Organ Rock Formation, South-east Utah, USA. Sedimentology 56, 1774-1800. yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. eau, USA: relationships to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99. : Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792." yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. eau, USA: relationships to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99. : Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792." yenta Formation (Lower Jurassic), Southwest Colorado. Sed. Geol. 55, 233-262. eau, USA: relationships to salt tectonics in the Paradox Basin. Sed. Geol. 73, 77-99. : Kayenta Formation (Lower Jurassic), central portion Colorado Plateau. Sed. Geol. 85, 411-434. of Colorado Plateau units suggests long magnetization times. Tectonophysics 362, 335–344. nd screening modern analogues for rivers in the rock record. J. Sed. Res. 79, 773-792."

Lower Miocene, SE Ebro Basin, Spain). 23rd IAS Meeting of Sedimentology Abstracts Book, 19. aimeden Sandstones, High Atlas of Marrakesh (Morocco). Sed. Geol. 107, 21-36. ce and Rupture Dynamics of Himalayan Frontal Thrust, India. Science 294, 2328-2331. aimeden Sandstones, High Atlas of Marrakesh (Morocco). Sed. Geol. 107, 21-36. ial dominated succession: The Upper Triassic Oukaimeden Sandstone Formation, Central High Atlas, Mc physics 91, 271-283.

stone basins. In: Friend P. F., Williams B. P. J. (Eds.) New Perspectives on the Old Red Sandstone. Geol. S onal dynamics of the Upper Permian Rangal Coal Measures, Bowen Basin, Australia. J. Sed. Geol. 70, 87 Permian-Triassic boundary, northern Bowen Basin, Australia. Paleogeog., Paleoclim., Paleoecol. 179, 17, sedimentation, and salinity problems. In: Milliman J. D., Haq B. U. (Eds.) Sea-Level Rise and Coastal Sul J. Cost. Res.14, 826-836.

zene Ganges-Brahmaputra lower delta plain. Sed. Geol. 155, 317-342.

d basin. Geology 25, 11-14.

ar, India. Sedimentology 41, 825-845. d basin. Geology 25, 11-14.

ar, India. Sedimentology 41, 825-845.

d basin. Geology 25, 11-14.

-zone river: Cooper Creek, central Australia. Geology 14, 175-178.

graphy and tectonics of the Roer Valley Graben. Geologie en Mijnbouw 73, 129-141.

n the Rhine-Meuse delta, The Netherlands. Nederlandse Geografische Studies 316 (PhD Thesis, Universi ial Subsidence within a Coastal Prism (Central Rhine–Meuse Delta, The Netherlands). In: Blum M. D., M

ental latest Permian- late Triassic Solway basin (U.K.). Sed. Geol. 210, 27-47.

pean Basins. Geol. J. 44, 711-741."

with the East African Rift System. In: Prichard H. M., Alabaster T., Harris N. B. W., Neary C. R. (eds.) Mag dstone reservoirs from vertical downhole data, Mesaverde Group, Piceance Creek Basin, Colorado. AAP distribution, Rocky Mountain foreland. In: Macqueen R.W., Leckie D. A. (eds.) Foreland basins and fold in: In: Morales M. (ed.) Aspects of Mesozoic geology and paleontology of the Colorado Plateau. Museum he Williams Fork Formation, Piceance Basin, Colorado. MSc Thesis, University of Texas at Austin.

pp and lidar data: Williams Fork Formation, Piceance Basin, Colorado. MSc Thesis, University of Texas at pain: contrasting processes and products. J. Sed. Res. 68, 879-889.

de la Sierra de Alcubierre (sector central de la Cuenca del Ebro). Rev. Soc. Geol. España 15, 217-231. tributary systems: Examples from the Miocene Luna and Huesca Systems, northern Spain. Sed. Geol. 15 ial paleoclimate record from the Ebro Basin, Spain. J. Geol. 115, 601-608.

eosols of distal fluvial systems, Ebro Basin, Spain. Palaeogeo., Palaeoclim., Palaeoecol. 247, 220–235." te Interpretation. In: Chidsey T. C. Jr., Adams R. D., Morris T. H. (eds.) Regional to Wellbore Analog for Fl

(2009) Obsidian exploitation and utilization during the Oldowan at Melka Kunture (Ethiopia). In: Hover work for a long and persistent archaeological record: Melka Kunture, Ethiopia. J. Hum. Evol. 62, 104-115

lassification. Hydrol. Earth Syst. Sci. 11, 1633-1644."

lassification. Hydrol. Earth Syst. Sci. 11, 1633-1644." Geology 25, 895-898.

architecture of an axial drainage system: late Eocene Escanilla Formation, southern Pyrenean foreland lly aggrading basin: the Escanilla Formation, Spanish Pyrenees. In: Best J. L., Bristow C. S. (eds.) Braided (2003) The fluvial analog Escanilla Formation, Ainsa Basin, Spanish Pyrenees: Revisited. AAPG Internati L2) Chronostratigraphy of the Boltaña anticline and the Ainsa Basin (southern Pyrenees). GSA Bull. 124, of the Paraná Basin, Southernmost Brazil, and its correlation to the Waterberg Basin, Namibia. Gondw to climatic drying in a Mid-Triassic rift valley fill: the Omingonde Formation (Karoo Supergroup) of cent ppian), south-central and western Pennsylvania. AAPG Bull. 52, 246-263.

nuch Chunk Formation, Pottsville, Pennsylvania. Geology 19, 127-130. "

al and aeolian reservoir: the Triassic Sherwood Sandstone in the Irish Sea. In: North C. P., Prosser D. J. (Eental latest Permian- late Triassic Solway basin (U.K.). Sed. Geol. 210, 27-47.

pean Basins. Geol. J. 44, 711-741."

e Basin and its margins. Geol. J. 7, 183-216.

n of the central Cheshire Basin. J. Geol. Soc. 150, 857-870.

al and aeolian reservoir: the Triassic Sherwood Sandstone in the Irish Sea. In: North C. P., Prosser D. J. (ental latest Permian-late Triassic Solway basin (U.K.). Sed. Geol. 210, 27-47.

pean Basins. Geol. J. 44, 711-741."

pean Basins. Geol. J. 44, 711-741.

al and aeolian reservoir: the Triassic Sherwood Sandstone in the Irish Sea. In: North C. P., Prosser D. J. (Eabkha deposits in redefining the depositional model for the Ormskirk Sandstone Formation (Lower Triasental latest Permian-late Triassic Solway basin (U.K.). Sed. Geol. 210, 27-47.

pean Basins. Geol. J. 44, 711-741."

the Triassic–Jurassic boundary in the Four Corners area, southwestern U.S.A. Paleogeo., Paleoclim., Palica. Palaeogeog., Palaeoclim., Palaeoecol. 45, 1-21.

raphies: a new basin development model for the Karoo retroarc foreland system, South Africa. Basin Re ary in Karoo Basin, South Africa. Geol. Soc. Am. Bull. 115, 1133-1152."

) Karoo sequences from southern to central Africa: A major Gondwana Lacustrine episode. J. Afr. Earth. raphies: a new basin development model for the Karoo retroarc foreland system, South Africa. Basin Re sils from the mid- to Upper Permian Middleton Formation (Adelaide Subgroup, Beaufort Group, Karoo S 98) The age of the Morrison Formation. Mod. Geol. 22, 235-260.

aleosols and fluvial/overbank architecture in the Morrison Formation (Upper Jurassic), Western Interio eous vertebrate habitat of the Gobi Basin, Mongolia. Geoscience Canada 25, 15-26.

R., Webb L. E., Hacker B. R. (2001) Sedimentary record and tectonic implications of Mesozoic rifting in s f Mesozoic-Cenozoic intraplate deformation in Mongolia. Basin Res. 16, 79-99. , 483-513.

rn Mongolia. Unpublished PhD Thesis, Albert-Ludwigs-Universität, Freiburg."

eous vertebrate habitat of the Gobi Basin, Mongolia. Geoscience Canada 25, 15-26.

R., Webb L. E., Hacker B. R. (2001) Sedimentary record and tectonic implications of Mesozoic rifting in s f Mesozoic-Cenozoic intraplate deformation in Mongolia. Basin Res. 16, 79-99. , 483-513.

rn Mongolia. Unpublished PhD Thesis, Albert-Ludwigs-Universität, Freiburg."

atchewan River, Canada. Sedimentology 25, 625-648.

oson C. J., Thomas R. E. (2011) Evolution and sedimentology of a channel fill in the sandy braided South 98) The age of the Morrison Formation. Mod. Geol. 22, 235-260.

aleosols and fluvial/overbank architecture in the Morrison Formation (Upper Jurassic), Western Interio Blundell D. (eds.) Tectonic Evolution of Southeast Asia. Geol. Soc. Spec. Publ. 106, 291-306.

tan. Sed. Geol., 1-76.

neter-thick succession in northern Pakistan. Sed. Geol., 77-121.

n. J. Sed. Res. B64, 60-67.

eo., Paleoclim., Paleoecol. 115, 1-11.

c foreland basin, northern Pakistan. Sedimentology 44, 369-390.

sition in the Miocene Indo-Gangetic foreland basin, northern Pakistan. J. Geol. Soc. 158, 163-177.

., Hicks J., Kelley J. (2002) Faunal and environmental change in the Late Miocene Siwaliks of northern Paliocene fluvial strata of the Siwalik Group, Northern Pakistan. Jour. Geol. 98, 433-456.

tan. Sed. Geol., 1-76.

neter-thick succession in northern Pakistan. Sed. Geol., 77-121.

n. J. Sed. Res. B64, 60-67.

eo., Paleoclim., Paleoecol. 115, 1-11.

c foreland basin, northern Pakistan. Sedimentology 44, 369-390.

., Hicks J., Kelley J. (2002) Faunal and environmental change in the Late Miocene Siwaliks of northern Pa al Utah. J. Paleo. 58, 477-487.

5 (Maastrichtian) North Horn Formation of the Wasatch Plateau, Central Utah. Palaios 17, 327-346.

n sequences: An example from the Cordilleran foreland basin, Utah. Geology 32, 637-640. cosols. Geology 33, 333-336.

Jtah, and correlation of contemporaneous Campanian strata and vertebrate faunas along the margin of our of England. J. Geol. 88, 639-651.

f north-east England. Nature 284, 51-53.

ian plate margin processes. J. Geol. Soc. 139, 479-491.

Soc. 141, 919-931.

the improvement of the Upper Carboniferous Time Scale. Chem. Geol. (Isot. Geosci. Sect.) 59, 143-154 ral and northern England. Paleogeo., Paleoclim., Paleoecol. 70, 303-330.

In successions: implications for the UK Coal Measures. In: Whateley M. K. G., Spears D. A. (eds.) Europea om the Inner Sudetic Basin and the Bohemian Massif. Geol. Soc. Am. Bull. 79, 589-608.

the improvement of the Upper Carboniferous Time Scale. Chem. Geol. (Isot. Geosci. Sect.) 59, 143-154 patchiness in riparian vegetation on a Middle Pennsylvanian braided-river plain prone to flood disturbations.

clim., Paleoecol. 80, 197-212.

oration history of the Weald and Wessex basins, Southern England, UK. In: Underhill J. R. (ed.) Develop ective. Cret. Res. 25, 303-311.

from the earliest Cretaceous (c. 140 Ma) of Hastings, Sussex. J. Geol. Soc. 166, 989-997.

Formation of the Isle of Wight, southern England: their genesis and palaeontological significance. Pale Isle octim., Paleoecol. 80, 197-212.

oration history of the Weald and Wessex basins, Southern England, UK. In: Underhill J. R. (ed.) Develop ective. Cret. Res. 25, 303-311.

from the earliest Cretaceous (c. 140 Ma) of Hastings, Sussex. J. Geol. Soc. 166, 989-997.

Formation of the Isle of Wight, southern England: their genesis and palaeontological significance. Pale and petroleum potential of the Middle Pennsylvanian Paradox Formation. USGS Bull. 2000-O.

estone in the Paradox Basin, southeastern Utah and southwestern Colorado. USGS Bull. 2000-P." op. 8, 772-791.

me peat beds in the Upper Pleistocene Hinuera Formation, South Auckland, New Zealand. J. Royal Soc. Pepeat and surficial peat deposits at Ohinewai, Lower Waikato lowland. J. Royal Soc. New Zeal. 14, 233-1. Geophys. Res. Lett. 30, 1717.

fluvial reservoirs: an example from distal fluvial fan deposits in the Loranca Basin, Spain. In: North C. P., eposition (Late Oligocene-Early Miocene, Loranca Basin, Spain). J. Sed. Res. 70, 850-867. "

fluvial reservoirs: an example from distal fluvial fan deposits in the Loranca Basin, Spain. In: North C. P., in a fluvial reservoir analogue (Loranca Basin, Spain). Pet. Geosc. 1, 237-252."

he Carboniferous Joggins Formation of eastern Canada. Sedimentology 50, 415-439.

gins, Nova Scotia, Canada. Palaios 18, 197-211.

re rhythms at Joggins, Nova Scotia. J. Geol. Soc. 160, 643-648.

osit and its fossil biota in a dryland alluvial plain setting, Joggins, Nova Scotia. J. Geol. Soc. 161, 209-222 rmation of Nova Scotia: sedimentological log and stratigraphic framework of the historic fossil cliffs. Atl

nga River. Geomorphology 86, 144-175."

ian plate margin processes. J. Geol. Soc. 139, 479-491.

oniferous successions of Great Britain (Onshore). BGS Research Report RR/07/01."

3 late Pleistocene: evidence from optical dating of subsurface samples. Curr. Sci. 84, 451-454.

y sytem. In: Collinson J. D., Lewin J. (eds.) Modern and ancient fluvial systems. IAS Spec. Publ. 6, 473-48 physics 91, 271-283.

xford, Ireland. Geol. Today 7, 110-114.

Bed from the Old Red Sandstone of the Munster Basin in NW Iveragh, SW Ireland. J. Geol. Soc. 154, 189 of Mexico basin. AAPG Bull. 84, 1743-1774.

d to Miocene, Vermilion Block 50, Tiger Shoal area, offshore Louisiana. AAPG Bull. 88, 153-174."

y sytem. In: Collinson J. D., Lewin J. (eds.) Modern and ancient fluvial systems. IAS Spec. Publ. 6, 473-48 physics 91, 271-283.

ivial sequences in the Devonian Munster Basin, Ireland. In: House M. R., Gale A. S. (eds.) Orbital Forcing physics 91, 271-283.

evonian Trentishoe Formation of North Devon, UK. Sedimentology 31, 697-715.

stone basins. In: Friend P. F., Williams B. P. J. (Eds.) New Perspectives on the Old Red Sandstone. Geol. S n (L. Triassic) at the northern margin of the German Basin. Geol. Rundsch. 74, 519-536.

the Netherlands, a model. Paleogeo. Paleoclim. Paleoecol. 91, 277-290.

I model and subsidence analysis. Tectonophysics 313, 145-169.

conditions in the Lower Triassic Bunter Sandstone, North German Basin. Sed. Geol. 169, 129-149."

of their deposits, and the significance for palaeogeographic reconstruction. Earth-Sci. Rev. 111, 199-223

-200.

ution of the Vindhyan basin, central India. Sed. Geol. 141-142, 395-419.

r. Sci. 81, 806-809.

5, 3-22.

I9-160."

ınga River. Geomorphology 86, 144-175."

rce, and regional stratigraphy of the Roraima Supergroup and Roraima-like outliers in northern South A ng of Iapetus. Geol. Mag. 131, 291-299.

Kronprins Christian Land, eastern North Greenland. Geol. Surv. Den. Green. Bull. 6, 29-39.

troduction. In: Higgins A. K., Gilotti J. A., Smith M. P. (eds.) The Greenland Caledonides: evolution of the ennoscandia, East Greenland and Svalbard. Episodes 31, 35-43."

from the Ringerike Group of Norway. Sed. Geol. 16, 45-68.

e northwest Iberian Ranges (Central Spain). J. Sed. Petrol. 56, 862-875.

n: Marzo M., Puigdefábregas C. (eds.) Excursion Guidebook no. 2, 82 pp. 4th International Conference of the Buntsandstein and Muschelkalk facies (Permo-Triassic) in the SE Iberian Range, E Spain. Paleogeo rn border of the Triassic Iberian Trough (central Spain). Sed. Geol. 113, 245-267.

patial variations in tectonic subsidence in the Iberian Basin (eastern Spain): inferences from automated cance of the continental sedimentary transition across the Permian—Triassic boundary in Spain. Paleogo of the intramontane Guadix Basin (Central Betic Cordillera, Spain): implications for an Atlantic—Mediter 1 during the evolution of a continental basin: Pliocene—Pleistocene of the Guadix Basin (Betic Cordillera, 2) brate succession from the Orce region (Guadix-Baza Basin, SE Spain) and its bearing on the first human r the Pliocene-Pleistocene continental sediments of the Guadix Basin (Betic Cordillera, S. Spain). Quat. I fs in the Sierras Pampeanas and Precordillera of Argentina. Rev. Assoc. Geol. Arg. 44, 408-19.

idients in deep synorogenic foreland sequences: the Neogene Vinchina Basin, South-Central Andes. XVI rows from fluvial deposits of the Neogene Vinchina Formation, La Rioja, Argentina: A crab origin suggesion of the La Troya and Vinchina depocenters (northern Bermejo Basin, Tertiary), La Rioja, Argentina. In

Salt Range, Pakistan. Unpublished PhD Thesis, University of Leeds.

andstone, Potwar Basin, Pakistan. J. Geol. Soc. India 76, 505-517.

andstone, Salt Range, Pakistan. Sed. Geol. 233, 88-110."

o area and plant biostratigraphy of Abo red beds in New Mexico. Guideb.-New Mexico Geol. Soc. 34, 15 ble oxygen and carbon isotopes in caliche of the Abo Formation (Permian), south-central New Mexico, Idening: evidence from reduction features of the Abo Formation, New Mexico, U.S.A. J. Sed. Res. 75, 56 rsity, seasonal tropical landscape dominated by conifers and peltasperms: Early Permian Abo Formation 1, 21, 633-655.

i. 21, 633-655.

quence stratigraphy: application to the Late Cretaceous fluvial systems of southwestern Utah, U.S.A. J. § Jtah, and correlation of contemporaneous Campanian strata and vertebrate faunas along the margin of ming, U.S.A.) and their significance for paleoecology, paleoclimatology, and basin analysis. Palaeogeo. P n Basin, northwest Wyoming, U.S.A. Palaeogeo. Palaeoclim. Palaeoecol. 43, 95-128. es. 66, 354-363."

no and Oregon. In: Bonnichsen B., White C. M., McCurry M. (eds.) Tectonic and Magmatic Evolution of t

```
0.
78."
d. Journ. Geophys. Res. 112, F04010, doi:10.1029/2006JF000729."
cal modelling of hydrocarbon reservoirs and outcrop analogues. IAS Spec. Publ. 15, 21-56.
f SE Asia, 211-234.
Eds.) Characterization of Fluvial and Aeolian Reservoirs. Geol. Soc. Spec. Publ. 73, 247-264.
ology of Northwest Europe: Proceedings of the 4th Conference, 791-808.
ssic), East Irish Sea Basin. In: Meadows N. S., Trueblood S. E, Hardman M., Cowan G. (eds.) Petroleum Gi
'G Bull. 69,710-721.
belts: AAPG Memoir 55, 395-425.
of Northern Arizona Bullettin 59, 129-152.
: Austin."
ton (U.S.A.). Can. J. Earth Sci. 31, 1727-1743.
. Soc. 153, 427-436."
bsidence. Kluwer Academic Publishing, Dordrecht, Netherlands, 169-192.
eustasy, climate and tectonism in continental strata. SEPM Spec. Publ. 59, 93-107.
r, USA. Sed. Geol. 167, 115-135."
```

on, San Jorge Basin, Argentina. Sed. Geol. 203, 213-228. caniclastic deposits of the Bajo Barreal Formation, San Jorge Basin, Argentina. Cretac. Res. 30, 749-766.

17, 152-171.

procco. Sed. Geol. 218, 103-140."

ioc. Spec. Publ. 180, 29-60."

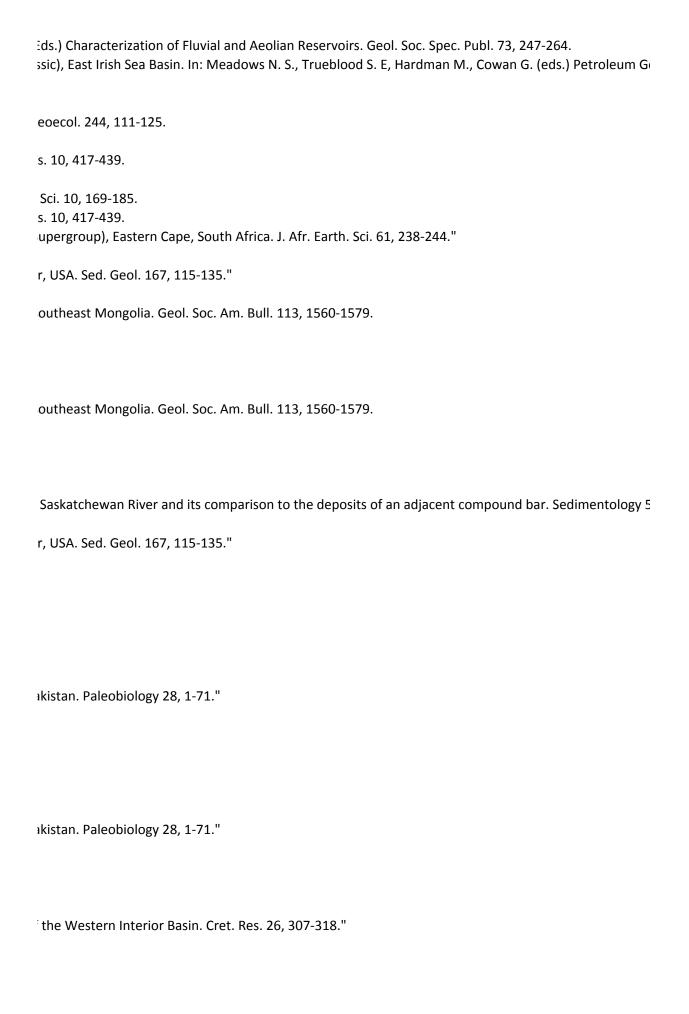
9-895.

73-188."

osidence. Kluwer Academic Publishing, Dordrecht, Netherlands, 169-192.

```
teit Utrecht).
arriott S. B., Leclair S. F. (Eds.) Fluvial Sedimentology VII. IAS Spec. Publ. 35, 295-320."
matic Processes and Plate Tectonics, Geol. Soc. Spec. Publ. 76, 427-442.
'G Bull. 69,710-721.
belts: AAPG Memoir 55, 395-425.
of Northern Arizona Bullettin 59, 129-152.
: Austin."
£35, 55-73.
luvial-Deltaic Reservoir Modeling: The Ferron Sandstone of Utah, AAPG Studies in Geology 50, 39-57.
s E., Braun D. R. (eds.) Interdisciplinary Approaches to the Oldowan. 111-128.
5."
basin, Spain. Basin Res.4, 335-352.
rivers. Geol. Soc. Spec. Publ. 75, 177-194.
onal Conference, Barcelona, Spain, September 21-24 2003.
1229-1250."
. Res. 8, 163.176.
:ral Namibia. Palaios 17, 249-267.
Eds.) Characterization of Fluvial and Aeolian Reservoirs. Geol. Soc. Spec. Publ. 73, 247-264.
```

eds.) Characterization of Fluvial and Aeolian Reservoirs. Geol. Soc. Spec. Publ. 73, 247-264.



```
an Coal Geology. Geol. Soc. Spec. Publ. 82, 1-16."
ance (Nýrany Member, Central and Western Bohemian Basin, Czech Republic). Rev. Paleobot. Palyn. 16
ment, evolution and petroleum geology of the Wessex Basin. Geol. Soc. Spec. Publ. 133, 39-65.
logeo., Paleoclim., Paleoecol. 292, 409-424."
ment, evolution and petroleum geology of the Wessex Basin. Geol. Soc. Spec. Publ. 133, 39-65.
ogeo., Paleoclim., Paleoecol. 292, 409-424."
New Zeal. 8, 385-393.
.244."
Prosser D. J. (eds.) Characterization of fluvial and aeolian reservoirs. Geol. Soc. Spec. Publ. 73, 79-94.
Prosser D. J. (eds.) Characterization of fluvial and aeolian reservoirs. Geol. Soc. Spec. Publ. 73, 79-94.
. Geol. 41, 115-142."
3.
€9-193."
3.
Timescales and Cyclostratigraphy. Geol. Soc. Spec. Publ. 85, 19-36."
ioc. Spec. Publ. 180, 29-60."
```

merica based on U-Pb geochronology. GSA Bull. 115, 331-348. northeast margin of Laurentia. GSA Mem. 202, 29-53. n Fluvial Sedimentology, Publicacions del Servei Geológic de Catalunya, Barcelona. . Paleoclim. Paleoecol. 103, 179-201. forward modelling of high-resolution stratigraphy (Permian–Mesozoic). Tectonophysics 300, 285-310. eo. Paleoclim. Paleoecol. 229, 3-23." ranean connection. Paleogeo. Paleoclim. Paleoecol. 151, 255-266. , southern Spain). Sed. Geol. 219, 97-114. occupation of Europe. Quat. Int. 223-224, 162-169. nt. 243, 16-32." I Congreso Geológico Argentino, Actas, Jujuy, 85-86. ted by neoichnology and sedimentology. Paleogeo. Paleoclim. Paleoecol. 291, 400-418. 1: Salfity J. A., Marquillas R. A. (eds.) Cenozoic geology of the Central Andes of Argentina. Salta, SCS Publ 7-163. USA. J. Sed. Petrol. 61, 458-472. 2-571.

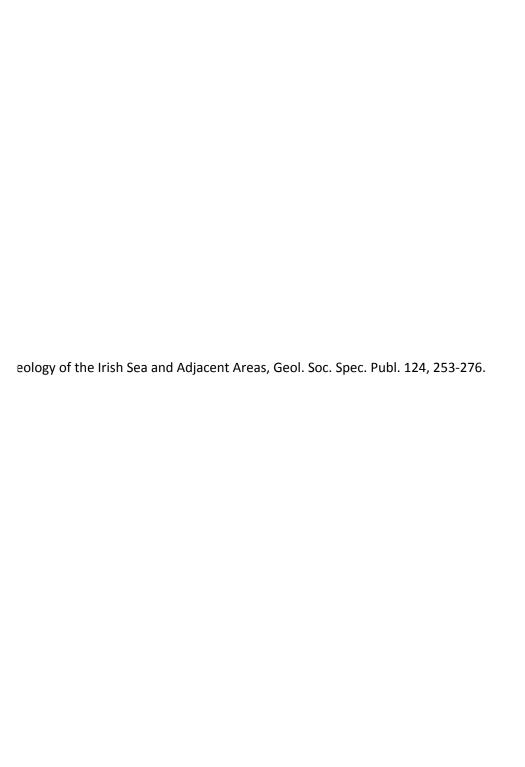
he Snake River Plain Volcanic Province. Idaho Geol. Surv. Bull. 30, 69-103.

1, New Mexico. Rev. Paleobot. Palyn. 145, 249-273."

the Western Interior Basin. Cret. Res. 26, 307-318."

Sed. Res. 73, 389-406.

'alaeoclim. Palaeoecol. 34, 1-30.



eology of the Irish Sea and Adjacent Areas. Geol. Soc. Spec. Publ. 124, 253-276.

58, 1860-1883."

isher, 91-110.