

AGES AND ORIGINS OF ROCKS OF THE KILLINGWORTH DOME, SOUTH-CENTRAL CONNECTICUT: IMPLICATIONS FOR THE TECTONIC EVOLUTION OF SOUTHERN NEW ENGLAND

JOHN N. ALEINIKOFF*, ROBERT P. WINTSCH**, RICHARD P. TOLLO***,
DANIEL M. UNRUH*, C. MARK FANNING****, and MARK D. SCHMITZ†

ABSTRACT. The Killingworth dome of south-central Connecticut occurs at the southern end of the Bronson Hill belt. It is composed of tonalitic and trondhjemitic orthogneisses (Killingworth complex) and bimodal metavolcanic rocks (Middletown complex) that display calc-alkaline affinities. Orthogneisses of the Killingworth complex (Boulder Lake gneiss, 456 ± 6 Ma; Pond Meadow gneiss, ~ 460 Ma) were emplaced at about the same time as eruption and deposition of volcanic-sedimentary rocks of the Middletown complex (Middletown Formation, 449 ± 4 Ma; Higganum gneiss, 459 ± 4 Ma). Hidden Lake gneiss (339 ± 3 Ma) occurs as a pluton in the core of the Killingworth dome, and, on the basis of geochemical and isotopic data, is included in the Killingworth complex.

Pb and Nd isotopic data suggest that the Pond Meadow, Boulder Lake, and Hidden Lake gneisses (Killingworth complex) resulted from mixing of Neoproterozoic Gander terrane sources (high $^{207}\text{Pb}/^{204}\text{Pb}$ and intermediate ϵ_{Nd}) and less radiogenic (low $^{207}\text{Pb}/^{204}\text{Pb}$ and low ϵ_{Nd}) components, whereas Middletown Formation and Higganum gneiss (Middletown complex) were derived from mixtures of Gander basement and primitive (low $^{207}\text{Pb}/^{204}\text{Pb}$ and high ϵ_{Nd}) sources. The less radiogenic component for the Killingworth complex is similar in isotopic composition to material from Laurentian (Grenville) crust. However, because published paleomagnetic and paleontologic data indicate that the Gander terrane is peri-Gondwanan in origin, the isotopic signature of Killingworth complex rocks probably was derived from Gander basement that contained detritus from non-Laurentian sources such as Amazonia, Baltica, or Oaxaquia. We suggest that the Killingworth complex formed above an east-dipping subduction zone on the west margin of the Gander terrane, whereas the Middletown complex formed to the east in a back-arc rift environment.

Subsequent shortening, associated with the assembly of Pangea in the Carboniferous, resulted in Gander cover terranes over the Avalon terrane in the west; and in the Middletown complex over the Killingworth complex in the east. Despite similarities of emplacement age, structural setting, and geographic continuity of the Killingworth dome with Oliverian domes in central and northern New England, new and published isotopic data suggest that the Killingworth and Middletown complexes were derived from Gander crust, and are not part of the Bronson Hill arc that was derived from Laurentian crust. The trace of the Ordovician Iapetan suture (the Red Indian line) between rocks of Laurentian and Ganderian origin probably extends from Southwestern New Hampshire west of the Pelham dome of northcentral Massachusetts and is covered by Mesozoic rocks of the Hartford basin.

INTRODUCTION

Tectonic models for the origins of terranes in New England have evolved to increasing complexity with the confluence of multiple techniques, including mapping (structure and stratigraphy), igneous and metamorphic petrology, geochemistry, geochronology, geophysics, and isotope geology (Tucker and others, 2001; Moench and Aleinikoff, 2003). These methods have facilitated correlation of high-grade, highly attenuated Paleozoic and Neoproterozoic rocks in southern New England with lower

*U.S. Geological Survey, Mail Stop 963, Federal Center, Denver, Colorado 80225; jaleinikoff@usgs.gov

**Department Geological Sciences, Indiana University, Bloomington, Indiana 47405

***Department of Geology, George Washington University, Washington, DC 20052

****Research School of Earth Sciences, Australian National University, Canberra ACT 0200, Australia

†Department of Geosciences, Boise State University, Boise, Idaho 83725

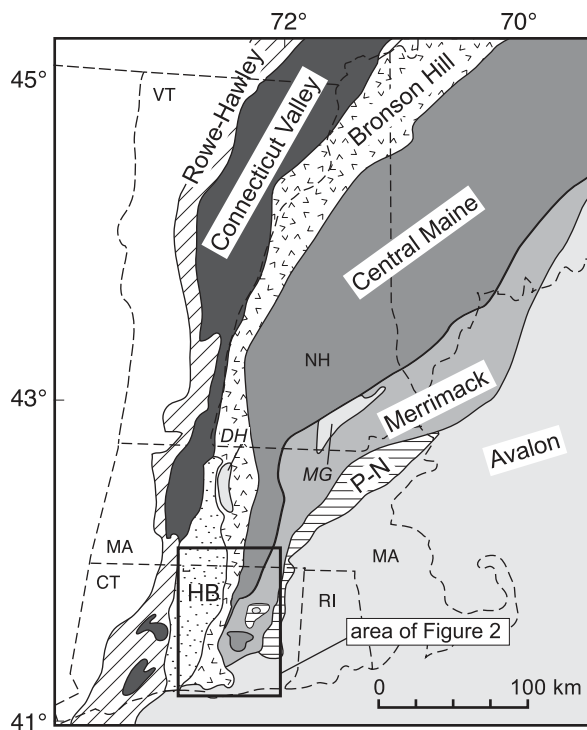


Fig. 1. Map of terranes of New England, modified from Hibbard and others (2006). Abbreviations: HB (Hartford basin), DH (Dry Hill Gneiss of Pelham dome), MG (Massabesic Gneiss), P-N (Putnam-Nashoba).

grade correlatives in northern New England and Maritime Canada (Tremblay and Pinet, 2005; Zagorevski and others, 2006; Wintsch and others, 2007). New interpretations have led to refinement of regional-scale terrane definitions and boundaries. It is now clear that several terranes (Bronson Hill, Central Maine, Merrimack, Putnam-Nashoba, Gander, and Avalon) of distinctly different origins are juxtaposed in a relatively small area of south-central New England (fig. 1). The Killingworth dome is shown here to be crucial to unraveling the accretionary history of terranes to the Laurentian margin in the middle and late Paleozoic.

The Killingworth dome (fig. 2) is the southernmost of a series of domes that expose largely calc-alkaline rocks along a narrow belt in central New England. The dome was previously interpreted as consisting of a homogeneous body of tonalitic rock (Rodgers, 1985). Webster and Wintsch (1987) used whole-rock major and trace element geochemistry to redefine units within the dome, including both intrusive and extrusive igneous rocks ranging in composition from granite to tonalite. Our new geochronologic and isotopic data show that the emplacement history of the plutonic rocks of the Killingworth dome is more complicated than previously thought. Geochemical and isotopic data indicate that this suite of igneous rocks was formed from diverse sources in both the Ordovician and Carboniferous. Moreover, in contrast to the geological record of much of New England, evidence of regional Acadian (Devonian) metamorphic overprinting is lacking. Rather, there is compelling evidence that these rocks were strongly metamorphosed, locally partially melted, and assembled by ductile faulting in the Carboniferous. The details of these interpretations are presented below to establish a new tectonic affinity for the Killingworth dome rocks in this area located

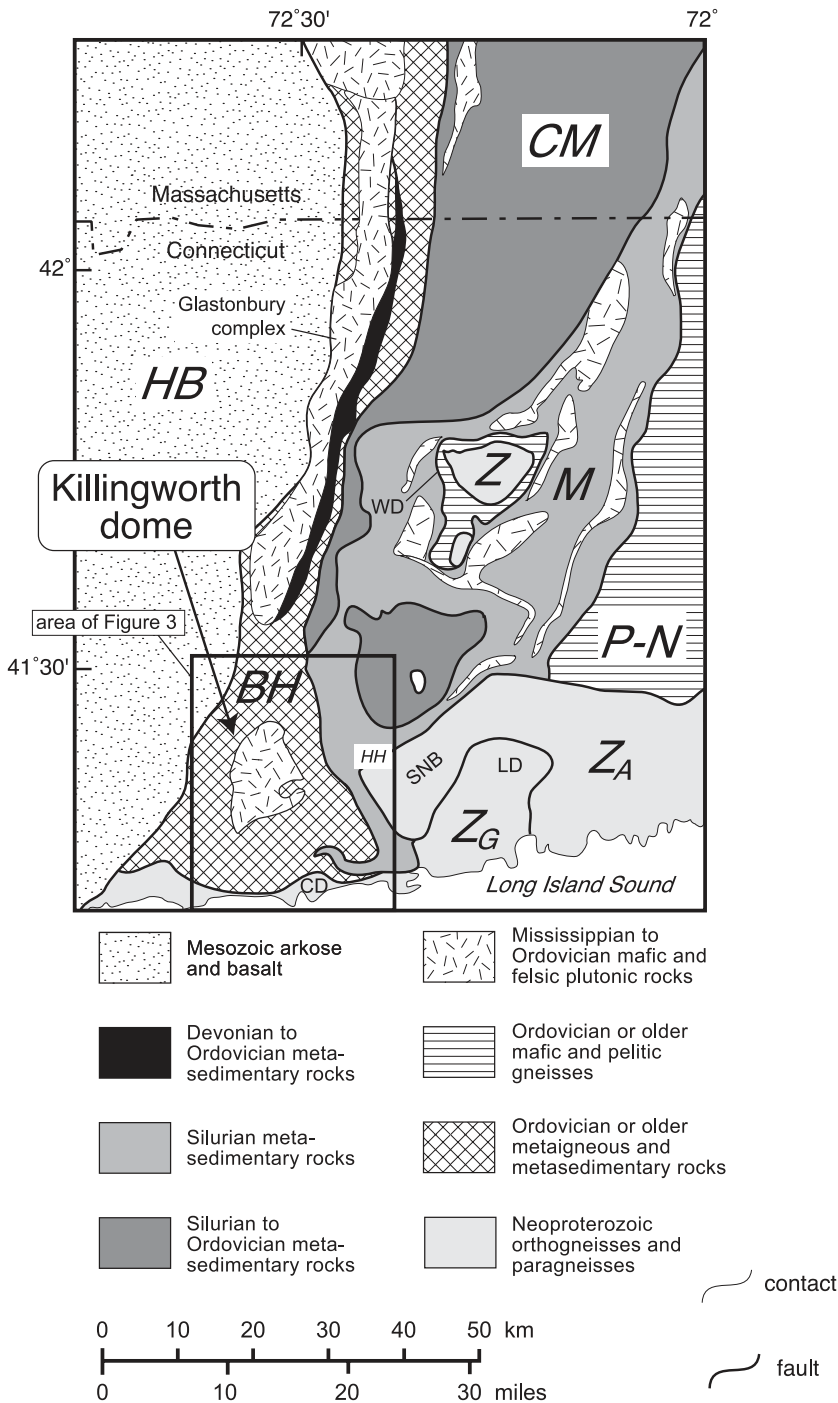


Fig. 2. Generalized geologic map of southern New England (modified from Wintch and others, 2003). Abbreviations: BH (Bronson Hill), CM (Central Maine), HB (Hartford basin), M (Merrimack), P-N (Putnam-Nashoba), Z (Neoproterozoic rocks, subdivided into Z_A—Avalon, Z_G—Gander), CD (Clinton dome), LD (Lyme dome), SNB (Selden Neck block), WD (Willimantic dome), HH (Honey Hill fault).

at the intersection of the Bronson Hill, Central Maine, Merrimack, Gander, and Avalon terranes (fig. 2).

GEOLOGIC SETTING

Orthogneisses within the Killingworth dome occur between metasedimentary rocks of the Central Maine and Merrimack terranes to the east, Neoproterozoic orthogneisses and paragneisses to the south, and Mesozoic arkoses of the Hartford basin to the west (fig. 2). The Killingworth dome traditionally has been considered to belong to the Bronson Hill Anticlinorium (Rodgers, 1985) or terrane (Wintsch and others, 2005). The orthogneisses consist primarily of plagioclase + quartz, with lesser amounts of biotite and amphibole. This limited range of modal compositions has made field and petrographic subdivision of the felsic rocks difficult (Mikami and Digman, 1957; Lundgren and Thurell, 1973). Various lithologies have been recognized, but boundaries were not identified in the field. Consequently, this lithologic assemblage has been collectively described as an undivided body of plagioclase gneiss in most previous geological investigations (for example, Percival, 1842; Rice and Gregory, 1906; Foye, 1949; Rodgers and others, 1959; Dixon and Lundgren, 1968). Distinctions between these lithologies were based on: (1) km-scale compositional zoning of plagioclase in the core region (Mikami and Digman, 1957), (2) degree of foliation development, and (3) modal content of biotite and amphibole in the Clinton quadrangle (Lundgren and Thurell, 1973). These relatively imprecise distinguishing characteristics led Rodgers (1985) to assign all varieties of these plagioclase gneisses to the Monson Gneiss. The outer margin of the Killingworth dome is occupied by metavolcanic and metasedimentary rocks of the Middletown and Collins Hill Formations.

Webster and Wintsch (1987) recognized that rocks of the Killingworth dome could be subdivided into distinct units on the basis of major- and trace-element compositions and map-scale distribution (fig. 3). These units were originally identified by reference to their geographic location around the dome (western, eastern southern, and central gneisses), but none were formally named. Our new results reinforce the interpretation that these bodies are petrologically distinct, thus justifying names for each unit. In contrast to previous studies, we regard these rocks as igneous lithodemic units rather than lithostratigraphic units. We follow the North American Stratigraphic Code (1983) in discriminating between groups and complexes, and formations and lithodemes. Use of new names supersedes Rodgers' (1985) assignment of these rocks to the Monson Gneiss because of unproven geologic continuity with Monson rocks at the type locality in Massachusetts. We follow the time scale of Gradstein and others (2004).

RESULTS

In this study, we supplement and expand on the work of Webster and Wintsch (1987). Our data confirm the designation of distinct orthogneiss bodies, and we establish ages and origins of these units. Thus, we have replaced the informal, geographically-based names used by Webster and Wintsch (1987) with the following new (informal) nomenclature: Higganum gneiss (formerly "western gneiss"), Pond Meadow gneiss (formerly "eastern gneiss"), Boulder Lake gneiss (formerly "southern gneiss;" originally called "Boulder Lake variety" of the Monson Gneiss for rocks in the adjacent Clinton 7.5-minute quadrangle by Lundgren and Thurell, 1973), and Hidden Lake gneiss (formerly "central gneiss"). We retain the name "Turkey Hill belt of Monson Gneiss" (Lundgren, 1963) for a narrow belt of rocks on the east side of the Killingworth dome that Rodgers (1985) correlated with other rocks called Monson Gneiss in north-central Connecticut and Massachusetts (Tucker and Robinson, 1990). The formal name of Clinton Granite Gneiss (Lundgren, 1964) is retained for rocks

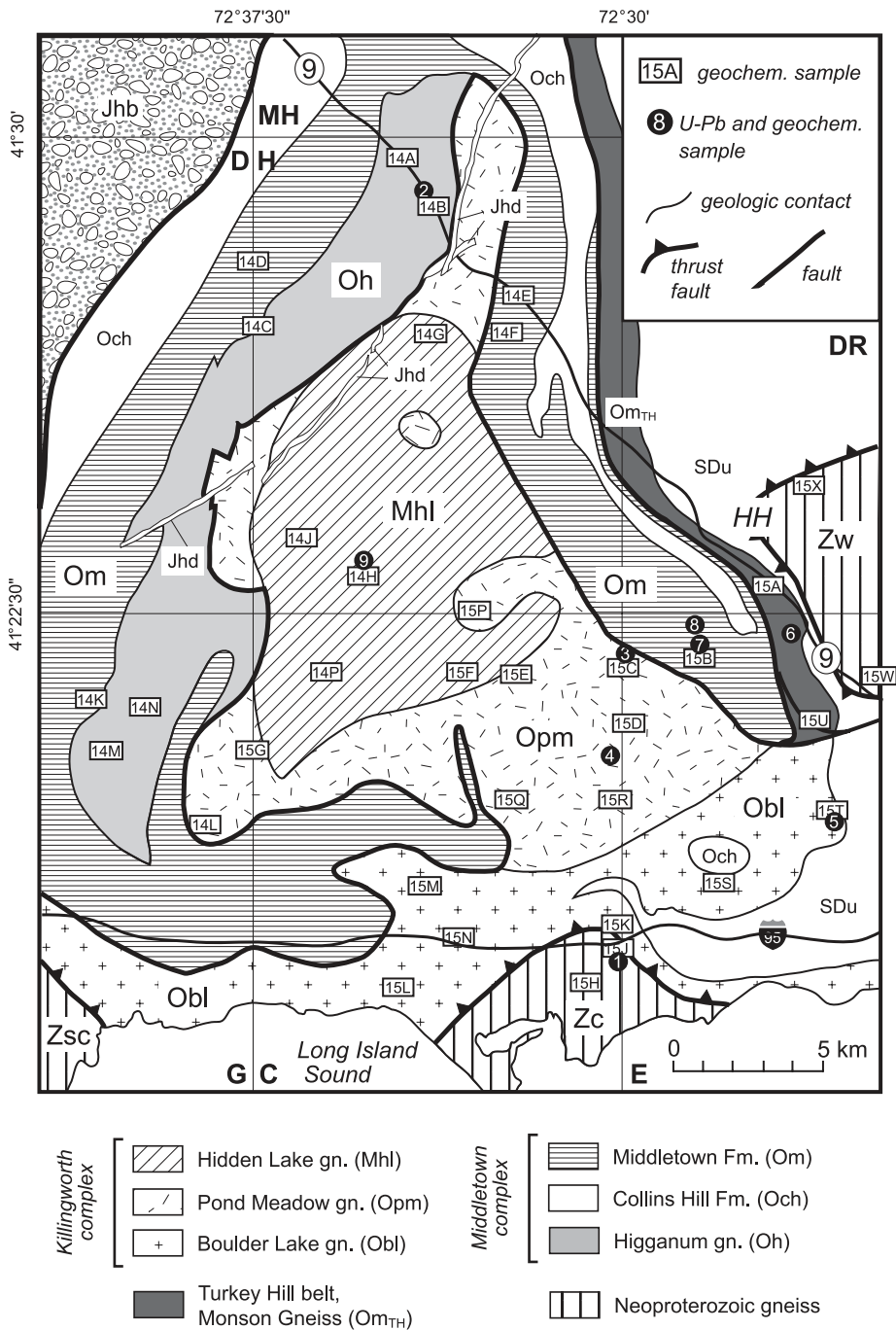


Fig. 3. Geologic map of the Killingworth dome, modified from Rodgers (1985). Lithologic subdivision of the Killingworth dome based on Webster and Wintch (1987). Quadrangle abbreviations: C (Clinton), D (Durham), DR (Deep River), E (Essex), G (Guilford), H (Haddam), MH (Middle Haddam). Abbreviations: Jhb (Jurassic clastic sediments of the Hartford basin), Jhd (Jurassic Higganum dike), Och (Collins Hill), Om_{TH} (Turkey Hill belt (Lundgren, 1963) of Monson Gneiss), SDu (Silurian and Devonian metaigneous and metasedimentary rocks, undifferentiated), Zc (Clinton Granite Gneiss), Zsc (Stony Creek Gneiss), Zw (Waterford Group), HH (Honey Hill fault). Fault contact surrounding Middletown-Higganum-Collins Hill rocks is an interpretation based on data in this paper (see Discussion section).

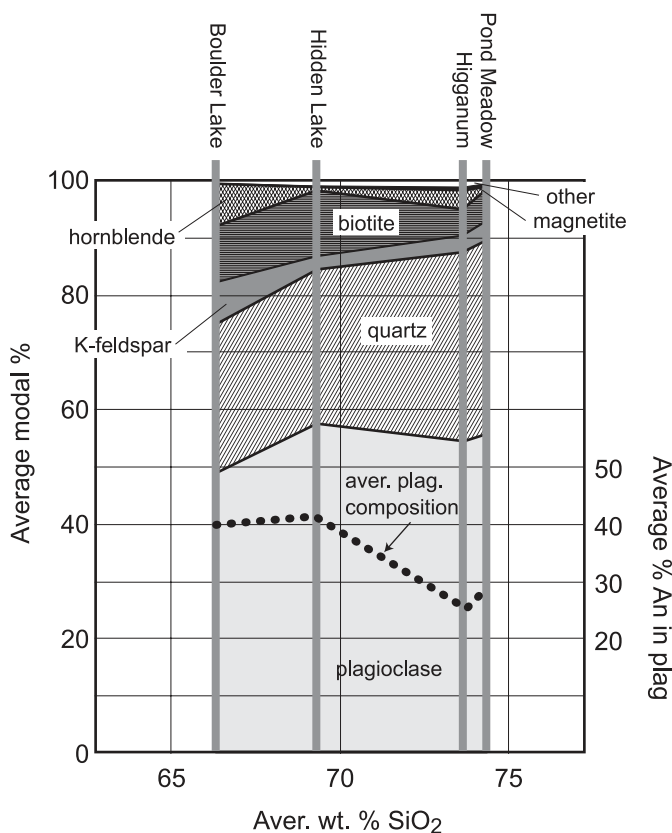


Fig. 4. Average modal compositions and average plagioclase composition (expressed as mol. % An) of lithologies in the Killingworth dome.

that crop out in the Clinton dome, south of the Killingworth dome along the Connecticut coast. The formal name Middletown Formation (Lundgren, 1963) is retained for anthophyllite-bearing quartz-plagioclase gneiss and interbedded amphibolite that occur on the outer margins of the Killingworth dome (fig. 3; Rodgers, 1985). On the basis of field occurrence, plus new geochemistry and isotopic data, we now group lithologic units that occur within the Killingworth dome into two informally-named complexes: (1) Killingworth complex (includes Boulder Lake, Pond Meadow, and Hidden Lake gneisses), and (2) Middletown complex (includes Middletown and Collins Hill Formations, and Higganum gneiss). The criteria for these groupings are discussed below.

All orthogneisses contain 50 to 60 modal percent plagioclase, 25 to 30 percent quartz, variable amounts of minor biotite, hornblende, K-feldspar, and magnetite, and trace garnet (fig. 4). For example, the hornblende/biotite ratios of the Boulder Lake and Higganum gneisses are higher than in the Hidden Lake and Pond Meadow gneisses. The minor minerals are typical, but not always diagnostic, of individual lithologic map units. In the western part of the study area, most primary mineralogical and textural features have been overprinted by moderate- to high-grade metamorphism and deformation; in the eastern part, overprinting also includes Late Paleozoic anatexis. Thus, the overall compositional similarity of most gneissic units and the pervasive regional gneissosity make it difficult to distinguish the orthogneisses in the

field. Nevertheless, our study indicates that whole-rock compositions and zircon age data largely survived high-grade metamorphism. As a result, we employ these characteristics in subdividing the lithologies and interpreting the geologic evolution of their protoliths.

Geochemistry

A total of 35 new samples from five units within the Killingworth dome (Killingworth and Middletown complexes) were collected for geochemical analysis (fig. 3, table 1). Analytical procedures for determination of concentrations of major elements and selected trace elements are described in the Appendix. This new data set supports the findings of Webster and Wintsch (1987) and offers additional perspectives because it includes analyses of trace elements not reported in the previous study. The data reported here are considered separately from the results in Webster and Wintsch (1987) because the data sets were collected by different analytical methods using different standards.

Orthogneisses of the Killingworth dome vary considerably in major-element composition, ranging from about 64 to 77 weight percent SiO_2 (table 1, figs. 5 and 6). Lithologic nomenclature for both the orthogneisses and metavolcanic rocks of the Middletown Formation is determined using standard procedures (figs. 5A and 5B, respectively). These diagrams involve use of alkali element concentrations that may have been susceptible to remobilization during metamorphism and/or anatexis. Nevertheless, we suggest that the clustering of data characterizing individual units indicates that such metasomatic effects are limited. Normative compositions indicate that the igneous protoliths of these rocks largely varied from tonalite to trondhjemitic (fig. 5A). Felsic rocks of the Middletown Formation differ from the Killingworth orthogneisses in displaying exclusively trondhjemitic compositions. The Middletown Formation also is distinctive in that it contains a bimodal lithologic association including rocks of both rhyolitic and andesitic affinity (fig. 5B), reflecting bimodality in the original magmatic compositions. The three mafic samples are widely distributed structurally and geographically around the complex, suggesting that rocks of this composition constitute an important component of the Middletown volcanic stratigraphy. As noted by Webster and Wintsch (1987), geochemical trends exhibited by most of the orthogneisses likely reflect the effects of primary magmatic processes. However, all rocks have experienced high-grade metamorphic conditions, and the southern half of the Killingworth complex is pervasively migmatitic. In some rocks, the high metamorphic grade is shown by patchy zones of tonalite with a subhedral granular (magmatic) texture interrupting an older wispy gneissic fabric (Wintsch and others, 2005). In other rocks, younger migmatitic textures are associated with an increase in modal K-feldspar. Migmatization is typical of all rocks of the southern part of the Killingworth dome. Although care was taken in collecting samples to avoid the effects of local migmatization, geochemical evidence of this anatexis is apparent from the scattered compositional data obtained from a small number of samples. This is illustrated by the relatively high normative orthoclase (Or) content of two samples (fig. 5A), decoupling of Na_2O and K_2O (table 1), and considerable variation in TiO_2 (fig. 6A) in the Hidden Lake gneiss at the high end of its silica range. Similar effects are exhibited by alkalis in some samples from the Higganum and Boulder Lake units. Such local variation notwithstanding, the preponderance of data from these samples defines coherent clusters and trends (fig. 6), supporting the compositionally-based distinction of lithologic units proposed by Webster and Wintsch (1987).

In spite of local variation in alkali contents that results in anomalous data on variation plots, the majority of geochemical data indicate that differences in aluminosity [expressed as aluminum saturation index = molar $(\text{Al}_2\text{O}_3 / (\text{CaO} + \text{Na}_2\text{O} + \text{K}_2\text{O}))$ (fig. 6D)] between the dominantly metaluminous Boulder Lake gneiss and other

TABLE 1

Geochemical analyses of samples from the Killingworth and Middletown complexes

Field No.	14G	14H	14J	14P	15F	1BW8.11.01	14L	15C	15D
Unit	Hidden Lake	Hidden Lake	Hidden Lake	Hidden Lake	Hidden Lake	Hidden Lake	Pond Meadow	Pond Meadow	Pond Meadow
SiO ₂	68.0	66.8	69.2	66.2	70.1	70.6	75.1	72.4	72.5
Al ₂ O ₃	17.5	16.6	15.4	16.4	14.8	13.8	12.7	14.1	14.3
Fe ₂ O ₃ *	2.32	3.65	3.34	3.90	3.16	2.88	2.98	3.33	3.45
MgO	0.72	1.39	1.17	1.52	1.09	0.69	0.73	0.70	0.60
CaO	4.12	4.75	4.12	5.20	3.10	1.73	2.28	3.44	3.47
Na ₂ O	4.26	3.39	3.29	3.07	3.09	2.17	3.46	3.16	3.85
K ₂ O	1.59	1.53	1.54	1.50	3.02	5.69	1.21	0.96	0.90
TiO ₂	0.21	0.38	0.29	0.40	0.34	0.51	0.23	0.17	0.19
P ₂ O ₅	0.08	0.21	0.12	0.18	0.16	0.17	0.08	0.09	0.09
MnO	0.04	0.05	0.05	0.05	0.06	0.03	0.08	0.08	0.06
LOI	0.59	0.51	0.46	0.64	0.41	0.49	0.69	0.76	0.37
Total	99.43	99.26	98.98	99.06	99.33	98.76	99.54	99.19	99.78
Rb	61	51	39	49	89	132	78	50	40
Ba	988	1098	1662	916	1589	4073	385	201	188
Sr	573	655	536	555	433	375	83	96	120
Th†	3.12	2.91	5.13	4.49	15.40	61.90	4.97	3.57	3.85
U	1.21	0.73	1.35	0.68	1.52	1.76	2.26	1.29	0.61
Zr	86	62	61	71	106	342	73	58	73
Hf	2.13	2.56	2.11	2.80	5.11	8.36	3.26	2.26	2.76
Nb	13	4	2	2	7	1	1	2	2
Ta	1.56	0.48	0.30	0.14	0.70	0.11	0.24	0.24	0.45
Ni	1	7	9	9	4	4	7	6	6
Zn	34	50	38	59	51	46	48	45	42
Cr	7	7	9	10	12	8	5	4	7
Sc	3.62	3.79	6.44	8.15	5.86	3.18	11.40	11.50	6.52
La	12.20	14.40	18.20	19.30	49.50	103.00	17.50	14.20	13.10
Ce	18.00	24.40	30.20	32.00	90.20	181.00	32.50	26.80	25.00
Nd	6.92	7.94	10.30	10.50	31.00	52.70	15.00	12.70	11.30
Sm	0.98	1.22	1.68	1.65	4.09	6.63	3.71	2.63	2.19
Eu	0.42	0.42	0.49	0.50	0.89	1.16	0.57	0.54	0.56
Gd	1.03	0.93	1.43	1.57	3.13	3.80	3.84	2.57	2.02
Tb	0.14	0.12	0.18	0.17	0.36	0.30	0.61	0.38	0.23
Ho	0.18	0.12	0.24	0.20	0.40	0.28	0.82	0.69	0.25
Tm	0.06	0.05	0.09	0.06	0.15	0.06	0.45	0.37	0.07
Yb	0.38	0.26	0.52	0.37	0.90	0.36	2.96	2.46	0.48
Lu	0.06	0.04	0.08	0.05	0.13	0.06	0.45	0.38	0.08
Y	5	4	5	6	11	10	20	16	4

peraluminous units is not a manifestation of local migmatization. Instead, these differences likely are a reflection of igneous processes, including both source-related characteristics and possible differentiation. The metaluminous Boulder Lake rocks (fig. 6D), characterized by relatively low silica contents and high concentrations of compatible major elements (figs. 6A-C, F), are the least compositionally evolved of the intermediate to felsic gneisses. In contrast, the Pond Meadow and Higganum gneisses display the most evolved compositions, with SiO₂ contents in excess of 70 weight percent. The Hidden Lake gneiss, intermediate in SiO₂ content between the Boulder Lake and Pond Meadow gneisses, plots apart from the trend defined by the other units in FeOt (total iron expressed as FeO) (fig. 6E) and the feldspar-compatible trace

TABLE 1
(continued)

Field No.	15E	15P	15R	3BS8.11.01	15G	15Q	15K	15L	15M
Unit	Pond Meadow	Pond Meadow	Pond Meadow	Pond Meadow	Pond Meadow	Pond Meadow	Boulder Lake	Boulder Lake	Boulder Lake
SiO ₂	71.9	77.0	73.2	71.1	76.4	70.4	74.3	63.8	66.6
Al ₂ O ₃	13.9	12.0	14.2	14.9	12.2	14.8	13.0	15.3	15.1
Fe ₂ O ₃ *	3.66	2.19	2.63	3.40	2.76	4.36	2.17	6.65	5.73
MgO	1.17	0.51	0.71	0.81	0.41	0.87	0.56	2.37	1.95
CaO	3.70	1.69	2.90	3.55	1.66	3.83	1.72	6.15	4.67
Na ₂ O	3.28	3.67	3.82	3.92	3.74	3.45	3.05	2.60	2.73
K ₂ O	1.20	1.67	1.27	1.08	1.17	1.14	3.56	1.55	1.68
TiO ₂	0.28	0.12	0.21	0.22	0.16	0.25	0.20	0.48	0.42
P ₂ O ₅	0.10	0.07	0.09	0.09	0.07	0.11	0.16	0.11	0.10
MnO	0.08	0.03	0.04	0.06	0.09	0.11	0.04	0.12	0.10
LOI	0.41	0.47	0.54	0.44	0.56	0.45	0.32	0.34	0.30
Total	99.68	99.42	99.61	99.57	99.22	99.77	99.08	99.47	99.38
Rb	58	52	56	44	41	55	106	63	70
Ba	173	175	242	221	254	228	1105	284	373
Sr	150	83	131	123	85	135	160	187	191
Th†	7.52	5.75	4.47	3.61	6.20	5.84	34.50	8.97	7.60
U	1.26	2.10	0.88	2.70	1.37	1.40	1.78	1.17	2.45
Zr	58	107	62	54	75	62	130	108	94
Hf	2.73	4.83	2.49	2.23	3.02	1.90	4.32	3.05	3.37
Nb	3	3	7	3	13	2	10	4	5
Ta	0.50	0.14	0.49	0.47	0.80	0.58	0.28	0.39	1.28
Ni	9	3	5	5	8	4	5	10	4
Zn	47	39	31	53	38	44	23	57	45
Cr	18	6	7	15	7	8	9	15	10
Sc	10.40	8.47	6.26	7.99	11.40	14.60	4.53	22.80	21.80
La	16.00	19.10	9.55	11.20	17.80	16.30	48.50	23.00	17.90
Ce	31.90	40.70	22.80	22.50	37.90	36.90	87.50	45.60	43.20
Nd	15.20	20.60	9.51	10.10	19.10	15.30	32.70	19.30	15.20
Sm	3.30	4.82	2.33	2.11	4.28	3.34	6.19	3.87	3.58
Eu	0.61	0.69	0.57	0.57	0.68	0.67	0.98	0.78	0.74
Gd	3.07	6.41	2.51	1.94	4.30	3.14	5.10	3.55	3.35
Tb	0.51	1.11	0.37	0.23	0.66	0.45	0.57	0.51	0.56
Ho	0.70	2.27	0.44	0.27	0.77	0.82	0.51	0.70	0.83
Tm	0.29	1.18	0.14	0.09	0.24	0.46	0.20	0.35	0.43
Yb	1.75	7.84	0.82	0.56	1.44	3.13	1.13	2.21	2.87
Lu	0.27	1.20	0.13	0.08	0.20	0.49	0.17	0.33	0.43
Y	15	49	8	6	18	20	15	19	23

elements Sr and Ba (figs. 6G and 6H, respectively), and thus is unlikely to share a common petrologic lineage. In contrast, the Higganum gneiss and felsic Middletown Formation exhibit similarities in trace-element concentrations that suggest a petrologic relationship between these units. Systematic decrease in CaO and Sr with increasing silica content for all Ordovician units (figs. 6B and 6G) points toward a petrologic origin in which differentiation of plagioclase feldspar was important.

Geochronology

With the lithologic differences found by Webster and Wintsch (1987) confirmed here, we determined the age of each unit of the Killingworth dome. In addition, we

TABLE 1
(continued)

Field No.	15N	15S	15T	4BS8.11.01	14A	14B	14M	14N	HA-1-02
Unit	Boulder Lake	Boulder Lake	Boulder Lake	Boulder Lake	Higganum	Higganum	Higganum	Higganum	Higganum
SiO ₂	66.0	67.4	64.3	67.6	71.5	75.6	72.2	72.3	71.7
Al ₂ O ₃	14.9	14.6	15.5	15.0	13.4	12.1	13.6	13.3	13.4
Fe ₂ O ₃ *	5.77	5.77	6.10	5.22	5.34	3.86	4.46	4.21	4.86
MgO	1.91	1.14	2.18	1.64	1.06	0.87	1.23	1.41	1.06
CaO	4.73	6.17	5.48	4.35	4.23	1.76	2.85	1.75	3.81
Na ₂ O	2.85	2.90	3.09	3.05	3.61	4.37	4.57	4.39	3.50
K ₂ O	2.32	0.49	1.50	1.78	0.49	0.55	0.17	1.52	0.82
TiO ₂	0.41	0.42	0.44	0.42	0.37	0.17	0.40	0.28	0.28
P ₂ O ₅	0.11	0.15	0.15	0.11	0.10	0.07	0.12	0.09	0.10
MnO	0.10	0.15	0.13	0.09	0.11	0.08	0.09	0.09	0.10
LOI	0.49	0.34	0.28	0.16	<0.01	<0.01	0.06	0.50	0.25
Total	99.59	99.53	99.15	99.42	100.21	99.43	99.75	99.84	99.88
Rb	82	4	74	87	10	10	2	113	21
Ba	537	129	325	348	90	67	81	164	113
Sr	188	273	199	185	89	99	165	100	89
Th†	8.27	2.53	12.20	9.48	1.03	0.88	3.66	0.39	1.50
U	2.62	1.15	1.98	0.86	0.81	0.86	1.65	6.77	0.85
Zr	85	58	87	95	93	98	82	74	51
Hf	3.09	1.90	2.67	3.52	3.44	3.27	2.52	2.82	2.38
Nb	2	2	4	8	3	0	2	7	0
Ta	0.42	0.17	1.03	0.70	0.20	0.17	0.21	0.76	0.25
Ni	9	6	0	6	5	2	3	4	4
Zn	36	124	41	41	67	35	30	40	56
Cr	10	4	7	18	6	16	4	4	15
Sc	19.60	24.70	22.50	18.20	35.70	8.68	16.50	22.60	17.00
La	20.80	6.50	30.20	26.40	3.91	3.05	12.50	2.17	4.58
Ce	38.80	22.30	56.00	50.20	9.35	6.58	27.80	4.18	10.70
Nd	16.90	11.80	23.40	22.80	6.74	5.15	13.60	3.00	6.91
Sm	3.79	3.95	4.76	4.64	2.86	1.56	3.81	0.99	2.46
Eu	0.78	1.04	0.93	0.83	0.67	0.65	0.81	0.36	0.55
Gd	3.54	5.10	4.40	4.08	3.99	2.02	4.18	1.11	3.19
Tb	0.46	0.83	0.63	0.53	0.75	0.39	0.72	0.22	0.61
Ho	0.75	1.28	0.88	0.70	1.27	0.72	1.14	0.34	1.10
Tm	0.41	0.68	0.39	0.28	0.60	0.40	0.51	0.16	0.58
Yb	2.57	4.53	2.40	1.57	3.90	2.69	3.46	1.01	3.66
Lu	0.39	0.67	0.36	0.22	0.61	0.45	0.54	0.14	0.58
Y	19	25	18	18	28	16	27	7	25

dated granitic gneiss from the Clinton dome (fig. 3). The geochronology results are described below in chronological order, from oldest to youngest. Isotopic data for zircon are shown on Tera-Wasserburg concordia plots; ages are calculated by determining the weighted average of selected ²⁰⁶Pb/²³⁸U ages. Isotopic data (uncorrected for common Pb) from titanite are shown on a Tera-Wasserburg concordia plot and are regressed to calculate a lower intercept age. Geochronologic techniques, including a discussion of the age of our titanite standard (determined by thermal ionization mass spectrometry), are described in the Appendix.

Clinton Granite Gneiss.—The Clinton Granite Gneiss of Lundgren (1964) and Lundgren and Thurell (1973) occurs structurally below the Killingworth dome in the

TABLE 1
(continued)

Field No.	14D	14E	14F	14K	15B	14C	2BS8.11.01	16BS11.11.02
Unit	Middle-town	Middle-town	Middle-town	Middle-town	Middle-town	Middle-town	Middle-town	Middle-town
SiO ₂	56.8	73.3	55.4	54.8	73.7	76.2	74.3	72.5
Al ₂ O ₃	14.9	12.8	15.6	15.2	12.2	12.5	12.2	13.1
Fe ₂ O ₃ *	12.3	3.73	8.40	12.6	4.40	3.48	4.47	2.55
MgO	5.38	2.17	6.62	4.13	1.79	0.48	1.24	1.48
CaO	3.02	2.23	6.65	6.71	1.96	1.41	1.60	3.24
Na ₂ O	5.64	4.28	5.54	3.98	4.62	4.68	4.63	5.24
K ₂ O	0.15	0.17	0.13	0.41	0.16	0.53	0.35	0.16
TiO ₂	0.80	0.24	0.51	0.91	0.38	0.19	0.37	0.51
P ₂ O ₅	0.22	0.10	0.12	0.10	0.12	0.08	0.10	0.17
MnO	0.27	0.05	0.16	0.18	0.06	0.12	0.05	0.06
LOI	0.12	0.42	0.55	0.12	0.09	0.22	0.08	0.14
Total	99.60	99.49	99.68	99.14	99.48	99.89	99.39	99.15
Rb	8	9	4	11	5	9	8	2
Ba	50	36	58	22	51	79	88	58
Sr	61	119	200	119	111	85	132	141
<i>Th</i> †	0.50	0.35	0.44	1.00	4.79	1.08	4.48	4.27
<i>U</i>	0.23	0.38	0.13	0.43	1.58	1.01	1.87	1.33
Zr	27	28	30	40	92	53	79	110
<i>Hf</i>	1.00	0.94	0.94	1.26	3.31	2.37	2.81	3.38
Nb	1	0	0	0	2	1	2	3
<i>Ta</i>	0.11	0.11	0.06	0.13	0.19	0.24	0.20	0.23
<i>Ni</i>	18	3	25	26	3	0	3	2
Zn	124	29	93	117	30	54	37	32
<i>Cr</i>	14	7	123	14	6	5	12	12
<i>Sc</i>	43.70	18.80	35.80	39.20	14.90	11.90	11.40	17.80
<i>La</i>	2.26	1.97	2.88	3.75	15.70	4.01	14.00	13.70
<i>Ce</i>	4.71	4.44	7.09	9.78	33.80	8.70	28.20	33.40
<i>Nd</i>	4.17	2.64	4.40	6.07	18.00	5.29	15.80	18.80
<i>Sm</i>	1.77	1.17	1.56	2.16	4.79	1.85	4.00	5.54
<i>Eu</i>	0.70	0.46	0.60	0.64	0.92	0.57	0.85	1.35
<i>Gd</i>	2.32	1.77	2.08	2.54	5.06	2.02	5.03	6.22
<i>Tb</i>	0.44	0.33	0.34	0.51	0.89	0.44	0.77	1.06
<i>Ho</i>	0.70	0.58	0.59	0.78	1.31	0.80	1.23	1.78
<i>Tm</i>	0.34	0.33	0.32	0.39	0.67	0.37	0.62	0.82
<i>Yb</i>	2.12	2.08	1.93	2.47	4.53	2.45	4.24	5.29
<i>Lu</i>	0.31	0.31	0.29	0.35	0.66	0.36	0.65	0.79
Y	18	13	13	18	30	14	21	39

* total iron expressed as Fe₂O₃

† elements in italics analyzed by instrumental neutron activation (INA) techniques; all others analyzed by X-ray fluorescence (XRF)

Clinton dome along Long Island Sound (figs. 2 and 3). It is a tan- to cream-weathering granitic gneiss, with 0.5 to 1.0 cm equant grains of K-feldspar, quartz, and plagioclase, accompanied by minor biotite. A weak gneissosity is defined by local wispy biotite-rich folia present in most outcrops. The contact between the Clinton Granite Gneiss and the structurally higher Boulder Lake gneiss to the north is a ductile fault (fig. 3). Although both rocks are shown on the Bedrock Geologic Map of Connecticut (Rodgers, 1985) as Neoproterozoic, it has long been suspected that the fault is a terrane boundary along which intrusive Ordovician rocks of the Killingworth dome are

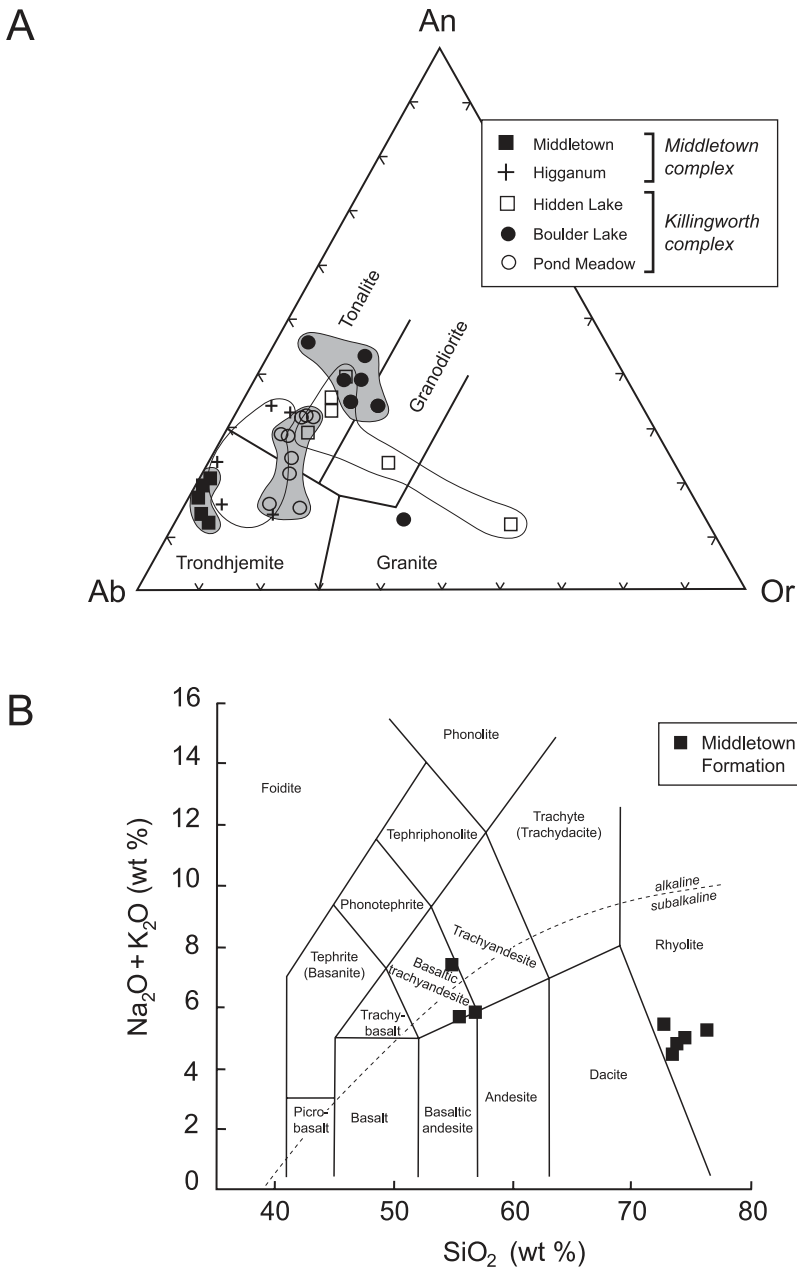


Fig. 5. Geochemical plots of whole-rock data from the Killingworth dome. (A) Plot of normative anorthite (An) vs albite (Ab) vs orthoclase (Or) for gneisses of the Killingworth and Middletown complexes. Lithologic field boundaries from Barker (1979). All normative data were calculated using $Fe^{2+}/Fe(\text{total}) = 0.9$. (B) Plot of total alkalis ($Na_2O + K_2O$) vs SiO_2 for rocks of the Middletown Formation in the study area, with field boundaries from Le Maitre and others (1989). All data are expressed in weight percent. Dashed line separating alkalic and subalkalic fields from Irvine and Baragar (1971).

juxtaposed against Neoproterozoic Gondwanan basement rocks (Wintsch and others, 2005). A new interpretation for this contact is presented in the Discussion section below.

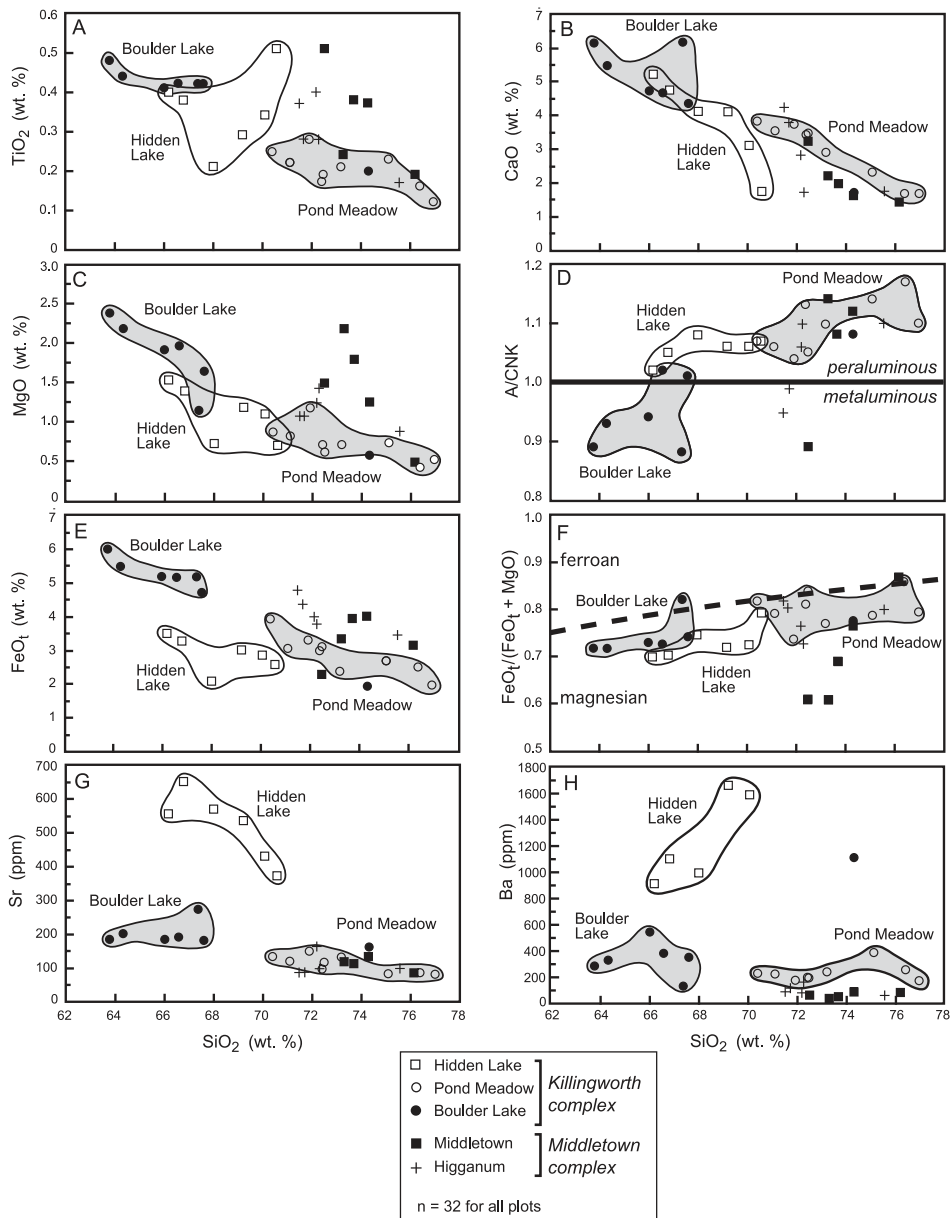


Fig. 6. Geochemical plots of whole-rock data (SiO_2 vs. various discriminants) from the Killingworth dome. (A) TiO_2 , (B) CaO , (C) MgO , (D) aluminum saturation index ($\text{A/CNK} = \text{molar } (\text{Al}_2\text{O}_3 / (\text{CaO} + \text{Na}_2\text{O} + \text{K}_2\text{O}))$), (E) FeO_t (total iron expressed as FeO), (F) $\text{FeO}_t / (\text{FeO}_t + \text{MgO})$, (G) Sr , and (H) Ba . All data are expressed in weight percent except Sr and Ba which are expressed in parts per million. Dashed line separating the ferroan (tholeiitic) and magnesian (calc-alkaline) fields in F from Frost and others (2001).

Zircons from the Clinton Granite Gneiss (sample 5BS8.11.01; locality 1, fig. 3) are subhedral to euhedral, and medium brown. Some grains are prismatic [average length-to-width ratio (l/w) = 3-4, whereas others are more stubby (l/w = 1-2)]. Most grains are fractured and contain numerous opaque inclusions. All zircons show fine, concentric, oscillatory zoning in cathodoluminescence (CL) (fig. 7A). Only a few

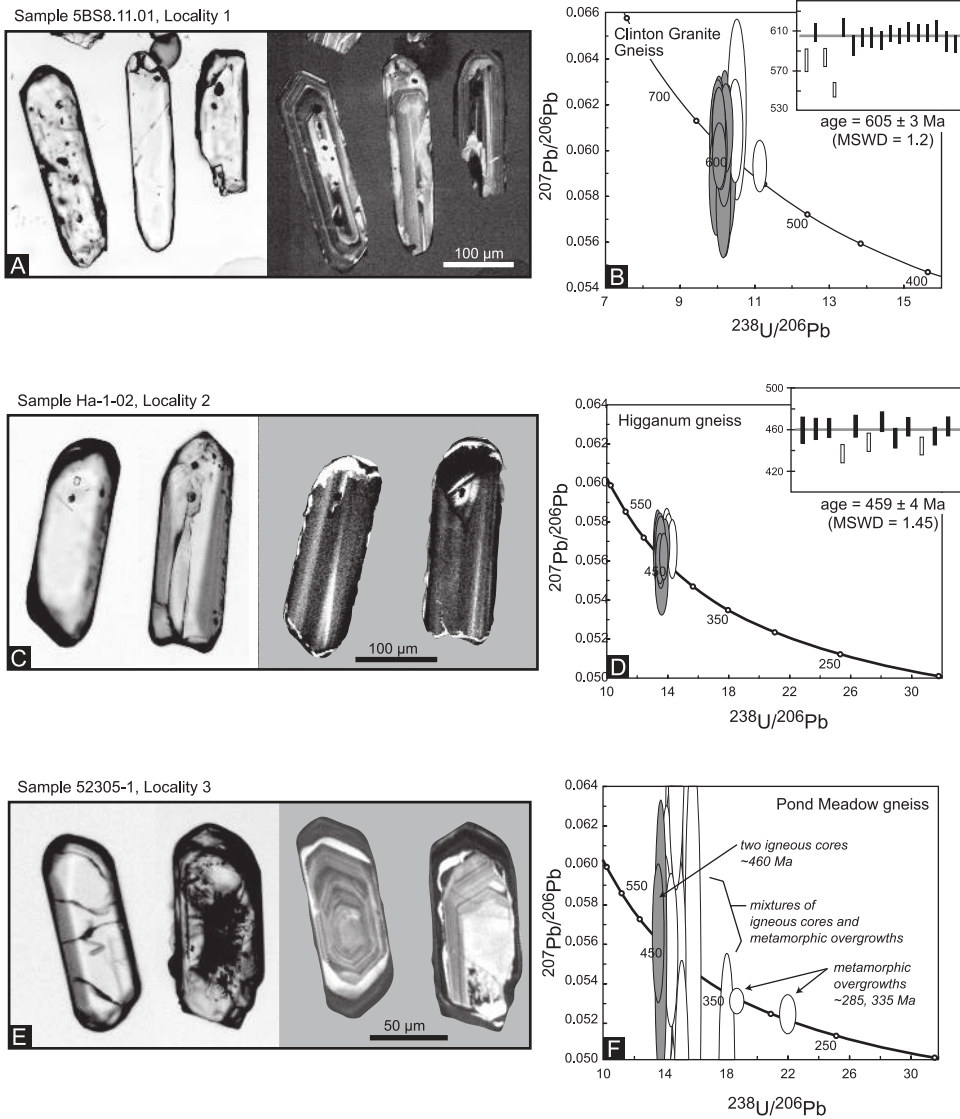


Fig. 7. Digital images (transmitted light and CL) of representative zircons and Tera-Wasserburg Concordia plots (with insets showing weighted averages of $^{206}\text{Pb}/^{238}\text{U}$ ages) for rocks of the Killingworth complex. Only gray-filled error ellipses used in weighted average calculations here and in subsequent plots. (A-B) Clinton Granite Gneiss (sample 5BS8.11.01); (C-D) Higganum gneiss (sample Ha-1-02); (E-F) Pond Meadow gneiss (sample 52305-1).

grains have very small, dark, unzoned overgrowths. Fourteen of 17 U-Pb analyses by sensitive high resolution ion microprobe (SHRIMP) yield an age of 605 ± 3 Ma (fig. 7B), interpreted as the time of crystallization of the granite protolith. Two younger analyses from cores are probably due to minor Pb loss. One analysis on the dark (in CL) outermost portion of a grain (probably an overgrowth) yields an age of about 552 Ma.

Higganum gneiss.—The Higganum gneiss (“western gneiss” of Webster and Wintsch, 1987) crops out along the western margin of the Killingworth dome (fig. 3). Rocks of

this unit are pale to medium gray-weathering calc-alkaline tonalitic and trondhjemitic gneisses (Webster and Wintsch, 1987), with subordinate amphibolite and biotite schist. The strongly gneissose fabric of the Higganum has destroyed primary structures; however, conformable layering of amphibolite and schist within the gneiss suggests an extrusive origin.

Zircons from a tonalitic sample of Higganum gneiss (sample Ha-1-02; locality 2, fig. 3) are prismatic, subhedral to euhedral, light brown to colorless, and contain numerous opaque inclusions. In CL, these grains show oscillatory zoning that is parallel to the crystallographic *c*-axis of the grains but is not concentric (fig. 7C). Also present are very small overgrowths that are black and unzoned in CL. The zircons contain relatively high concentrations of uranium (mostly 600 – 800 ppm) and have relatively high Th/U (0.2 – 0.3, table 2), indicative of magmatic origin. U-Pb SHRIMP data (9 of 12 analyses) indicate an age of 459 ± 4 Ma (fig. 7D), interpreted as the time of crystallization of the volcanic protolith. The three slightly younger ages probably are due to minor Pb loss.

Pond Meadow gneiss.—The Pond Meadow gneiss (“eastern gneiss” of Webster and Wintsch, 1987) occurs on the southern, western, and northern flanks of the Killingworth dome (fig. 3). It is a pale-gray weathering tonalitic and trondhjemitic gneiss, and is the most sodic and least mafic of all Killingworth units (table 1; Webster and Wintsch, 1987). The gneiss is poorly to moderately foliated, coarse-grained, and commonly migmatitic (Wintsch and others, 2005), suggesting that an anatectic event was superimposed on an earlier orthogneiss.

Zircons from the Pond Meadow gneiss (samples 52305-1 and 3BS8.11.01; localities 3 and 4, respectively, fig. 3) occur as two morphologic types. Both populations are prismatic and euhedral. Most grains from sample 52305-1 are medium-brown and contain numerous cracks and opaque inclusions. CL images show that these grains are composed primarily of concentric, oscillatory-zoned cores (many of which are broken), overgrown by dark, unzoned rims (fig. 7E). In contrast, most grains from sample 3BS8.11.01 are colorless and composed of very small, partially resorbed cores and broad, patchwork-zoned mantles (fig. 8A). We dated both populations, with the intent of determining: (1) the age of the igneous protolith by analyzing the oscillatory-zoned cores, and (2) the age of anatexis/metamorphism by analyzing the broad, patchwork-zoned mantles.

Two analyses of oscillatory-zoned cores from sample 52305-1 yield $^{206}\text{Pb}/^{238}\text{U}$ ages of ~456 to 462 Ma (Th/U of about 0.3, table 2), interpreted as the approximate time of igneous emplacement (fig. 7F). Two analyses of overgrowths from this sample yield ages of ~335 and ~285 Ma (Th/U < 0.02), interpreted as times of metamorphism. Several other analyses, with intermediate ages (about 418 – 448 Ma) and high, igneous-type Th/U (0.24 – 0.44), probably result either from analysis of mixed core and rim, or Pb loss from highly cracked grains. These dates have no age significance.

Patchwork-zoned mantles from zircons in sample 3BS8.11.01 have low Th/U (<0.04), indicative of metamorphic origin. Ten analyses yield a weighted average age of 335 ± 2 Ma (fig. 8B). Three analyses are slightly younger (~325 and 300 Ma) and one analysis is slightly older (~345 Ma). In addition, one oscillatory-zoned core yielded an age of about 460 Ma. We conclude from results from both samples that the Pond Meadow gneiss crystallized at about 460 Ma and was metamorphosed to high-grade, probably anatectic conditions at about 335 Ma. We see no evidence for a metamorphic event older than about 335 Ma. Limited data suggest subsequent metamorphic zircon growth as young as about 300 Ma.

Boulder Lake gneiss.—The belt of Boulder Lake gneiss, named by Lundgren and Thurell (1973) for rocks in the Clinton quadrangle, has been extended eastward into

TABLE 2
SHRIMP U-Th-Pb data for zircon from rocks of the Killingworth and Middletown complexes, Connecticut

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2.5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
1. 5BS8-11.01 (Clinton Granite Gneiss) 41°17'26" N, 72°29'55" W													
5BS8-1.1c	-----	0.0618	0.30	128	0.71	581.1	5.6	10.57	1.0	.0618	2.5		
5BS8-2.1c	0.000114	0.0621	0.24	338	0.54	608.5	4.7	10.10	0.8	.0604	1.9		
5BS8-3.1c	-----	0.0609	0.18	338	0.57	583.9	4.5	10.53	0.8	.0609	1.5		
5BS4-11r	-----	0.0594	0.10	1354	0.23	551.8	3.6	11.18	0.7	.0594	0.8		
5BS5.1c	0.000058	0.0596	-0.08	389	0.61	614.0	4.6	10.03	0.8	.0588	1.6		
5BS6.1c	0.000098	0.0611	0.16	246	0.58	595.8	5.0	10.33	0.9	.0596	2.2		
5BS7.1c	0.000059	0.0587	-0.16	395	0.81	603.7	4.5	10.21	0.8	.0578	1.6		
5BS8-8.1c	0.000068	0.0601	0.02	229	0.57	603.6	5.1	10.20	0.9	.0591	2.0		
5BS8-9.1c	-----	0.0598	-0.02	355	0.78	601.3	4.6	10.23	0.8	.0598	1.4		
5BS8-10.1c	0.000035	0.0616	0.18	621	0.45	608.2	4.2	10.10	0.7	.0611	1.1		
5BS8-11.1c	0.000167	0.0626	0.32	1120	0.54	605.5	4.0	10.15	0.7	.0602	1.2		
5BS8-12.1c	-----	0.0596	-0.07	261	0.49	609.0	4.9	10.10	0.8	.0596	1.7		
5BS8-13.1c	-----	0.0610	0.11	532	0.86	608.6	4.3	10.09	0.7	.0610	1.1		
5BS8-14.1c	-----	0.0599	-0.03	805	0.82	608.4	4.2	10.11	0.7	.0599	1.0		
5BS8-15.1c	-----	0.0605	0.04	230	0.58	610.3	5.4	10.07	0.9	.0605	1.8		
5BS8-16.1c	0.000247	0.0631	0.39	294	0.73	600.2	5.0	10.25	0.9	.0595	2.7		
5BS8-17.1c	-----	0.0610	0.15	430	0.41	597.7	4.4	10.28	0.7	.0610	1.3		
2. Ha-1-02 (Higganum gneiss) 41°29'16" N, 72°33'16" W													
Ha-1-1.1c	0.000044	0.0557	-0.09	669	0.31	468.2	4.6	13.30	1.0	.0550	1.3		
Ha-1-2.1c	-----	0.0571	0.09	843	0.33	465.8	3.3	13.33	0.7	.0571	0.9		
Ha-1-3.1c	-----	0.0561	-0.02	655	0.28	461.3	3.3	13.48	0.7	.0561	1.0		
Ha-1-4.1c	0.000029	0.0571	0.16	660	0.32	443.6	3.1	14.02	0.7	.0567	1.1		
Ha-1-5.1c	-0.000014	0.0569	0.08	696	0.32	463.6	3.8	13.40	0.8	.0571	1.0		
Ha-1-6.1c	-----	0.0574	0.17	757	0.33	451.7	3.2	13.75	0.7	.0574	0.9		
Ha-1-7.1c	-0.000058	0.0559	-0.08	472	0.26	472.5	3.5	13.15	0.8	.0567	1.4		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common ^{206}Pb	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err^3 (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err^3 (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err^3 (%)	$\frac{^{207}\text{Pb}^{2.5}}{^{206}\text{Pb}}$ (Ma)	err^3 (Ma)
2. Ha-1-02 (Higganum gneiss) 41°29'16" N, 72°33'16" W													
Ha-1-8.1c	0.000023	0.0572	0.13	779	0.31	456.5	3.5	13.62	0.8	.0569	0.9		
Ha-1-9.1c	0.000017	0.0564	0.02	743	0.23	464.4	3.2	13.39	0.7	.0562	0.9		
Ha-1-10.1c	-0.000010	0.0570	0.13	743	0.30	449.8	3.1	13.82	0.7	.0571	0.9		
Ha-1-11.1c	0.000018	0.0565	0.03	974	0.33	463.2	3.1	13.42	0.7	.0563	0.9		
Ha-1-12.1c	-0.000033	0.0560	-0.04	611	0.29	466.6	3.3	13.32	0.7	.0565	1.1		
3. 52305 (Pond Meadow gneiss) 41°19'45" N, 72°30'05" W													
52305-1.1c	0.001482	0.0754	2.57	192	0.26	397.6	5.1	15.74	1.4	.0537	9.3		
52305-2.1c	0.000124	0.0574	0.23	237	0.28	435.0	5.3	14.32	1.2	.0556	2.9		
52305-3.1c	0.000543	0.0637	1.03	255	0.27	428.6	5.2	14.54	1.2	.0558	5.7		
52305-4.1c	0.000231	0.0584	0.39	178	0.33	421.7	5.5	14.80	1.3	.0550	4.0		
52305-5.1c	0.000104	0.0580	0.22	335	0.36	461.6	5.3	13.47	1.2	.0565	2.6		
52305-6.1c	0.000759	0.0807	3.16	112	0.41	418.5	6.8	14.64	1.7	.0698	6.5		
52305-7.1c	0.000411	0.0636	0.97	426	0.34	442.9	5.0	14.03	1.2	.0577	3.8		
52305-8.1c	0.000671	0.0556	0.05	154	0.24	420.8	5.7	15.00	1.4	.0457	8.4		
52305-9.1c	0.000393	0.0616	0.68	303	0.31	456.2	5.5	13.64	1.2	.0559	5.4		
52305-10.1c	0.000429	0.0638	1.02	154	0.44	433.6	5.8	14.33	1.4	.0576	4.7		
52305-11.1c	0.000754	0.0669	1.43	448	0.37	426.0	5.3	14.63	1.3	.0559	9.5		
52305-12.1c	0.000274	0.0596	0.47	157	0.26	441.4	5.9	14.11	1.4	.0556	4.4		
52305-13.1c	0.001003	0.0639	1.28	438	0.33	351.2	4.0	17.96	1.2	.0491	5.2		
52305-14.1c	0.000283	0.0584	0.31	106	0.24	448.2	6.5	13.92	1.5	.0543	4.7		
52305-15.1c	0.000863	0.0667	1.40	182	0.26	430.3	5.9	14.51	1.7	.0541	16.4		
52305-16.1r	0.000013	0.0532	0.00	5458	0.01	337.6	3.3	18.60	1.0	.0530	0.5		
52305-17.1r	0.000040	0.0529	0.11	3428	0.00	287.2	2.9	21.94	1.0	.0523	0.8		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
4. 3BS8.11.01 (Pond Meadow gneiss) 41°20'59" N, 72°29'50" W													
3BS8-1.1r	-----	0.0534	0.06	811	0.02	324.3	2.4	19.37	0.7	.0534	1.4		
3BS8-2.1r	0.000032	0.0553	0.26	1016	0.02	334.7	2.4	18.73	0.7	.0548	1.3		
3BS8-2.2r	0.000080	0.0532	-0.02	1409	0.03	344.0	2.4	18.27	0.7	.0521	1.6		
3BS8-3.1r	0.000035	0.0526	-0.07	1136	0.03	335.2	2.3	18.76	0.7	.0521	1.3		
3BS8-4.1c	-----	0.0549	-0.16	367	0.21	461.1	3.7	13.51	0.8	.0549	1.6		
3BS8-6.1r	-----	0.0527	-0.04	958	0.03	327.7	2.3	19.18	0.7	.0527	1.2		
3BS8-7.1r	0.000025	0.0527	-0.06	1088	0.02	333.3	2.3	18.86	0.7	.0523	1.2		
3BS8-8.1r	-----	0.0542	0.12	1331	0.03	337.6	2.3	18.58	0.7	.0542	1.0		
3BS8-9.1r	0.000043	0.0535	0.04	1418	0.02	336.9	2.3	18.65	0.7	.0529	1.2		
3BS8-10.1r	-----	0.0528	-0.05	986	0.03	336.5	2.4	18.67	0.7	.0528	1.2		
3BS8-11.1r	-----	0.0532	0.00	1204	0.03	336.4	2.3	18.66	0.7	.0532	1.1		
3BS8-12.1r	-----	0.0538	0.08	1096	0.03	337.1	2.4	18.61	0.7	.0538	1.5		
3BS8-13.1r	-----	0.0542	0.14	1065	0.03	331.7	2.3	18.91	0.7	.0542	1.2		
3BS8-14.1r	0.000030	0.0536	0.06	1220	0.02	333.8	2.3	18.82	0.7	.0532	1.2		
3BS8-15.1r	0.000051	0.0522	-0.01	1441	0.02	298.2	2.0	21.15	0.7	.0514	1.4		
5. 4BS8.11.01 (Boulder Lake gneiss) 41°18'59" N, 72°26'01" W													
4BS8-1.1c	0.000110	0.0547	-0.07	238	0.28	419.8	5.4	14.90	1.3	.0531	2.6		
4BS8-2.1c	-0.000050	0.0551	-0.10	296	0.36	448.2	4.3	13.89	1.0	.0558	1.8		
4BS8-3.1c	-----	0.0564	0.08	183	0.38	441.3	4.6	14.10	1.1	.0564	1.9		
4BS8-4.1c	-0.000072	0.0563	0.08	253	0.40	436.4	4.3	14.25	1.0	.0573	2.7		
4BS8-5.1c	0.000076	0.0549	-0.17	155	0.32	461.4	5.7	13.52	1.3	.0538	2.4		
4BS8-6.1c	-0.000244	0.0542	-0.23	186	0.32	453.3	4.9	13.70	1.1	.0577	3.0		
4BS8-7.1c	-0.000088	0.0564	0.06	380	0.45	450.7	4.1	13.78	0.9	.0577	1.3		
4BS8-8.1c	0.000102	0.0551	-0.13	249	0.32	457.4	4.5	13.64	1.0	.0536	2.4		
4BS8-9.1r	0.000169	0.0527	0.00	239	0.37	315.3	3.2	20.01	1.0	.0502	3.2		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
5. 4BS8.11.01 (Boulder Lake gneiss) 41°18'59" N, 72°26'01" W													
4BS8-10.1r	0.000021	0.0567	0.03	254	0.35	471.5	4.5	13.18	1.0	.0564	1.7		
4BS8-11.1c	0.000062	0.0568	0.07	269	0.37	459.0	4.3	13.55	1.0	.0559	1.7		
4BS8-12.1c	-----	0.0565	0.05	226	0.34	454.3	4.5	13.69	1.0	.0565	1.6		
4BS8-13.1c	-----	0.0555	-0.02	346	0.36	437.8	4.0	14.23	0.9	.0555	1.4		
4BS8-14.1m	-0.000056	0.0558	-0.03	279	0.37	453.9	4.3	13.70	1.0	.0566	1.7		
4BS8-15.1m	-0.000017	0.0534	0.09	272	0.34	311.3	3.1	20.19	1.0	.0536	1.9		
4BS8-16.1c	-----	0.0545	0.25	238	0.21	309.4	3.2	20.29	1.1	.0545	2.0		
4BS8-17.1c	-----	0.0550	-0.21	190	0.36	477.6	4.9	13.03	1.1	.0550	1.9		
4BS8-17.2c	-----	0.0561	-0.01	397	0.49	461.6	4.3	13.47	1.0	.0561	1.2		
4BS8-18.1c	0.000062	0.0568	-0.03	214	0.32	494.7	5.1	12.55	1.0	.0559	2.0		
4BS8-18.2c	0.000078	0.0842	-0.17	160	0.37	1330.8	13.3	4.37	1.0	.0831	1.1		
6. 1BS8.11.01 (Turkey Hill belt of Monson Gneiss) 41°27'57" N, 72°26'27" W													
1BS8-1.1c	-0.000004	0.0563	0.16	1123	0.52	411.1	3.9	15.16	1.0	.0563	0.9		
1BS8-2.1c	0.000035	0.0556	0.01	650	0.38	433.6	3.7	14.38	0.9	.0551	1.1		
1BS8-3.1c	0.000025	0.0561	0.08	1086	0.54	430.2	3.8	14.49	0.9	.0557	0.8		
1BS8-4.1c	0.000026	0.0559	0.05	861	0.50	431.5	3.6	14.45	0.8	.0555	1.0		
1BS8-5.1c	-0.000007	0.0568	0.11	998	0.54	447.4	3.7	13.90	0.8	.0569	0.8		
1BS8-6.1c	0.000055	0.0564	0.09	430	0.35	442.1	4.0	14.09	0.9	.0556	1.4		
1BS8-7.1c	0.000008	0.0559	0.02	2106	0.55	443.3	3.7	14.05	0.8	.0558	0.6		
1BS8-8.1c	0.000017	0.0560	0.06	773	0.60	434.5	3.5	14.34	0.8	.0558	1.0		
1BS8-9.1c	-0.000004	0.0563	0.05	1190	0.59	446.6	3.8	13.93	0.9	.0563	0.7		
1BS8-10.1c	-----	0.0564	0.01	837	0.38	464.4	3.8	13.39	0.8	.0564	0.9		
1BS8-11.1c	0.000037	0.0569	0.13	1069	0.43	446.2	3.6	13.94	0.8	.0564	0.9		
1BS8-12.1c	-----	0.0563	-0.07	1047	0.64	486.1	4.2	12.78	0.9	.0563	0.8		
1BS8-13.1c	0.000050	0.0571	0.24	1196	0.69	422.1	3.4	14.75	0.8	.0564	0.8		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common ^{206}Pb	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
6. 1BS8.11.01 (Turkey Hill belt of Monson Gneiss) 41°27'57" N, 72°26'27" W													
1BS8-14.1c	0.000010	0.0559	-0.02	1233	0.60	453.6	3.6	13.72	0.8	.0557	0.7		
1BS8-15.1c	0.000017	0.0564	0.04	1083	0.60	454.8	3.7	13.68	0.8	.0562	0.8		
1BS8-16.1c	-----	0.0560	0.01	740	0.49	451.1	4.1	13.80	0.9	.0560	0.9		
1BS8-17.1c	0.000000	0.0571	0.10	748	0.50	461.1	3.8	13.47	0.8	.0571	0.9		
1BS8-18.1c	0.000031	0.0566	0.08	613	0.45	448.5	3.8	13.88	0.9	.0561	1.0		
1BS8-19.1c	0.000006	0.0567	0.15	2001	0.63	432.2	3.7	14.40	0.9	.0566	0.6		
1BS8-20.1c	0.000014	0.0555	-0.07	910	0.38	454.6	3.7	13.70	0.8	.0553	0.8		
7. 2BS8.11.01 (anthophyllite gneiss, Middletown Formation) 41°21'51" N, 72°28'07" W [session 1]													
2BS8-1.1c	0.000273	0.0605	0.66	2531	0.19	419.0	2.7	14.87	0.7	.0565	1.3		
2BS8-2.1r	0.000000	0.0531	0.00	860	0.00	335.5	2.4	18.72	0.7	.0531	1.3		
2BS8-3.1c	0.000023	0.0539	-1.10	5817	2.86	707.7	5.3	8.72	0.8	.0536	0.4		
2BS8-4.1c	0.000035	0.0539	-0.22	2149	0.38	439.5	3.0	14.21	0.7	.0534	0.8		
2BS8-5.1c	0.005911	0.1463	11.15	2118	1.28	450.7	22.5	13.73	4.9	.0606	90.1		
2BS8-6.1c	0.000090	0.0559	0.18	1823	0.07	390.6	2.6	16.01	0.7	.0546	1.1		
2BS8-7.1r	0.000136	0.0520	0.05	923	0.00	266.2	2.2	23.77	0.8	.0500	2.7		
2BS8-8.1r	0.000000	0.0544	0.12	1135	0.03	347.0	2.4	18.06	0.7	.0544	1.2		
2BS8-9.1r	0.000340	0.0592	0.74	814	0.01	340.1	2.5	18.44	0.8	.0542	2.8		
2BS8-10.1d	0.000075	0.0854	-0.18	178	0.59	1360.6	11.3	4.27	0.9	.0844	1.3	1301	13
2BS8-11.1d	0.000127	0.0850	-0.03	267	0.54	1320.2	9.8	4.41	0.8	.0832	1.3	1273	13
2BS8-12.1d	0.000063	0.0778	-0.08	173	0.27	1158.7	9.4	5.09	0.8	.0769	1.4	1118	14
2BS8-13.1d	0.000204	0.0714	-0.01	150	0.47	972.4	8.5	6.16	0.9	.0685	2.8	883	29
7. 2BS8.11.01 (anthophyllite gneiss, Middletown Formation) 41°21'51" N, 72°28'07" W [session 2]													
2BS8-1.1r	0.001770	0.0694	1.43	501	0.00	518.2	8.1	12.16	1.8	.0431	17.8		
2BS8-2.1r	0.000210	0.0553	0.06	614	0.04	405.3	5.7	15.46	1.4	.0522	2.6		
2BS8-3.1r	0.000409	0.0618	0.66	539	0.04	467.4	6.8	13.31	1.5	.0558	5.2		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	Th U	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
7. 2BS8.11.01 (anthophyllite gneiss, Middletown Formation) 41°21'51" N, 72°28'07" W [session 2]													
2BS8-4.1r	0.000970	0.0639	1.36	829	0.00	326.5	4.7	19.33	1.5	.0496	7.0		
2BS8-5.1c	0.001580	0.0830	3.37	2620	0.19	439.9	6.8	14.09	1.5	.0600	12.6		
2BS8-6.1c	0.000064	0.0545	-0.03	1424	0.07	401.9	5.6	15.57	1.4	.0536	1.5		
2BS8-7.1c	0.000029	0.0556	0.17	1514	0.08	384.3	5.4	16.26	1.4	.0552	1.3		
2BS8-8.1d	0.000000	0.0750	0.13	334	0.46	1039.3	16.6	5.71	1.6	.0750	1.5	1070	30
2BS8-9.1d	0.000463	0.0917	0.74	242	0.38	1323.6	20.6	4.39	1.6	.0852	2.9	1319	55
2BS8-10.1d	0.000000	0.0683	-0.27	110	0.39	943.8	17.3	6.36	1.9	.0683	2.5	877	52
8. 16BS11.11.02 (anthophyllite gneiss, Middletown Formation) 41°21'55" N, 72°28'08" W [session 3]													
16BS-1.1c	0.000042	0.0556	0.03	1571	0.27	426.8	3.5	14.61	0.8	.0550	1.2		
16BS-1.2r	0.000439	0.0562	0.44	370	0.12	317.7	2.9	19.87	1.0	.0498	6.9		
16BS-1.3c	0.000315	0.0592	0.40	1972	0.35	452.2	5.7	13.79	1.3	.0546	1.8		
16BS-2.1c	0.000189	0.0562	0.34	1175	0.04	350.7	2.2	17.89	0.6	.0535	1.8		
16BS-3.1c	0.000153	0.0555	0.10	1568	0.14	400.3	1.5	15.64	0.4	.0533	1.6		
16BS-3.2c	0.000099	0.0553	0.07	2107	0.70	398.2	4.1	15.71	1.1	.0538	1.2		
16BS-3.3c	0.000174	0.0554	0.18	1228	0.07	368.4	1.8	17.03	0.5	.0528	1.9		
16BS-4.1c	0.000017	0.0555	-0.52	3486	2.59	593.5	7.0	10.43	1.2	.0552	0.5		
16BS-4.2c	0.000179	0.0574	0.22	2888	1.23	438.5	3.7	14.22	0.9	.0548	1.3		
16BS-5.1c	0.000235	0.0557	0.29	1044	0.04	342.2	2.4	18.37	0.7	.0522	2.5		
16BS-6.1c	0.000042	0.0546	0.02	1917	0.06	391.2	1.5	15.99	0.4	.0540	1.1		
16BS-7.1c	0.000060	0.0559	0.27	1523	0.19	358.6	2.9	17.45	0.8	.0550	1.1		
16BS-8.1c	0.000157	0.0566	0.27	1743	0.31	391.0	2.1	16.00	0.6	.0543	1.5		
16BS-9.1c	0.000065	0.0555	0.09	2015	0.26	403.4	2.5	15.49	0.6	.0545	1.0		
16BS-10.1c	0.000278	0.0587	0.46	1181	0.51	412.7	2.7	15.13	0.7	.0547	2.4		
16BS-11.1c	0.000721	0.0650	1.52	2368	0.92	321.9	28.9	19.49	9.1	.0545	3.6		
16BS-12.1c	0.000000	0.0550	0.24	218	0.07	333.0	3.8	18.82	1.0	.0550	7.2		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common ^{206}Pb	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$ [session 4]	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
8. 16BS11.11.02 (anthophyllite gneiss, Middletown Formation) 41°21'55" N, 72°28'08" W [session 4]													
16BS-1.1c	0.000136	0.0567	0.19	1382	0.40	419.0	5.4	14.90	1.3	.0547	1.1		
16BS-2.1c	0.000060	0.0568	0.13	2539	0.39	443.0	5.6	14.05	1.3	.0559	0.7		
16BS-3.1c	0.000056	0.0556	0.19	3132	0.65	374.2	5.0	16.72	1.4	.0548	0.9		
16BS-4.1c	0.000090	0.0553	0.23	2776	0.40	346.2	5.2	18.11	1.5	.0540	0.7		
16BS-5.1c	0.000077	0.0550	0.27	1402	0.08	320.4	4.2	19.60	1.3	.0538	0.9		
16BS-6.1c	0.000057	0.0564	0.04	2917	0.21	454.1	5.8	13.71	1.3	.0556	0.6		
16BS-7.1c	-0.000013	0.0552	0.27	1324	0.25	329.7	18.5	19.00	5.7	.0554	0.6		
16BS-8.1c	0.000124	0.0580	0.25	2406	0.21	451.0	5.8	13.80	1.3	.0562	1.0		
16BS-9.1c	0.000019	0.0553	0.06	2086	0.30	405.2	5.1	15.41	1.3	.0550	0.6		
16BS-10.1c	0.000068	0.0567	0.15	2266	0.46	432.5	5.9	14.41	1.4	.0557	1.0		
16BS-11.1c	0.000004	0.0549	0.09	1744	0.23	375.6	5.0	16.65	1.4	.0548	0.8		
16BS-12.1c	0.000011	0.0545	0.01	1897	0.24	390.2	4.9	16.03	1.3	.0544	0.5		
16BS-13.1c	0.000096	0.0566	0.51	2542	1.22	307.9	4.7	20.37	1.6	.0552	0.8		
7. 2BS8.11.01 (anthophyllite gneiss, Middletown Formation) 41°21'51" N, 72°28'07" W [session 5]													
2BS8-1.1c	0.000068	0.0552	0.04	2557	0.31	408.0	4.3	15.32	1.1	.0542	1.1		
2BS8-2.1c	0.000041	0.0568	0.13	1903	0.28	444.7	4.7	14.00	1.1	.0562	1.0		
2BS8-3.1c	0.000134	0.0588	0.37	2151	0.78	447.4	4.7	13.90	1.1	.0569	1.4		
2BS8-3.2c	0.000177	0.0574	0.30	2340	0.32	412.1	4.3	15.15	1.1	.0548	1.5		
2BS8-4.1c	0.000114	0.0559	-0.19	1505	0.05	506.1	7.7	12.29	1.6	.0542	1.3		
2BS8-5.1c	0.000028	0.0568	0.04	3342	0.34	470.6	5.1	13.20	1.1	.0564	0.7		
2BS8-5.2c	0.000000	0.0574	0.15	1467	0.33	457.8	4.9	13.57	1.1	.0574	1.1		
2BS8-6.1c	0.000020	0.0564	0.08	2805	0.51	442.0	4.6	14.09	1.1	.0561	0.9		
2BS8-7.1c	-0.000003	0.0549	0.02	2174	0.20	402.7	4.3	15.51	1.1	.0550	1.1		
2BS8-8.1c	0.000038	0.0585	0.34	3279	0.18	440.6	4.7	14.10	1.1	.0579	0.8		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
7. 2BS8.11.01 (anthophyllite gneiss, Middletown Formation) 41°21'51" N, 72°28'07" W [session 5]													
2BS8-9.1c	0.000062	0.0557	0.10	2861	0.12	408.3	4.3	15.29	1.1	.0548	1.1		
2BS8-10.1c	0.001662	0.0786	2.80	4333	0.22	448.2	8.2	13.92	1.8	.0543	18.4		
2BS8-11.1c	0.000020	0.0556	-0.07	3645	0.25	461.2	4.8	13.50	1.1	.0553	0.8		
2BS8-11.2c	0.000023	0.0564	0.04	3290	0.24	456.0	4.7	13.64	1.1	.0561	0.9		
8. 16BS11.11.02 (anthophyllite gneiss, Middletown Formation) 41°21'55" N, 72°28'08" W													
16BS11m-1.1c	0.006467	0.1679	14.23	22	0.28	342.5	8.6	15.72	2.5	.1679	2.3		
16BS11m-1.2m	0.002759	0.0946	5.19	60	1.44	324.4	5.5	18.37	1.7	.0946	1.9		
16BS11m-1.3r	0.001861	0.0794	3.49	123	4.44	259.9	3.7	23.46	1.4	.0794	1.7		
16BS11m-2.1c	0.001213	0.0686	1.99	126	2.46	310.2	4.3	19.88	1.4	.0686	1.6		
16BS11m-3.1c	0.005948	0.1275	9.41	44	1.38	284.5	5.4	20.08	1.9	.1275	2.0		
16BS11m-3.2m	0.002553	0.0923	4.88	74	3.37	330.5	5.4	18.08	1.7	.0923	1.8		
16BS11m-3.3r	0.001478	0.0694	2.24	258	2.23	261.6	3.2	23.61	1.2	.0694	1.3		
16BS11m-4.1m	0.003123	0.0962	5.56	102	4.85	269.0	4.1	22.16	1.5	.0962	1.7		
16BS11m-4.2r	0.001893	0.0747	2.90	190	1.46	258.7	3.4	23.71	1.3	.0747	1.4		
16BS11m-5.1c	0.004683	0.1142	7.62	46	1.14	322.7	6.4	18.00	2.0	.1142	2.2		
16BS11m-5.2m	0.002293	0.0822	3.85	135	0.90	254.5	3.6	23.87	1.4	.0822	1.6		
16BS11m-5.3r	0.001284	0.0686	2.13	278	1.72	266.1	3.3	23.23	1.3	.0686	1.2		
16BS11m-5.4r	0.000829	0.0661	1.85	288	0.86	253.1	3.1	24.51	1.2	.0661	1.2		
16BS11m-6.1c	0.006010	0.1400	11.02	38	0.58	266.4	5.6	21.09	2.1	.1400	2.2		
16BS11m-6.2m	0.003639	0.1043	6.44	55	2.61	307.8	5.6	19.13	1.8	.1043	2.3		
16BS11m-7.1m	0.003232	0.1000	5.84	49	3.22	332.1	6.1	17.81	1.8	.1000	2.1		
16BS11m-7.2r	0.001856	0.0769	3.19	141	2.94	255.3	3.8	23.96	1.5	.0769	1.9		
16BS11m-8.1m	0.003334	0.0944	5.38	105	2.24	255.7	4.6	23.39	1.8	.0944	1.7		
16BS11m-8.2r	0.003050	0.0959	5.44	107	3.97	296.3	4.8	20.11	1.6	.0959	2.0		
16BS11m-9.1m	0.001370	0.0717	2.56	143	1.21	248.8	3.5	24.76	1.4	.0717	1.7		

TABLE 2
(continued)

sample ¹ (location)	measured $\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	% common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2+.5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
8. 16BS11.11.02 (anthophyllite gneiss, Middletown Formation) 41°21'55" N, 72°28'08" W													
16BS11m-10.1r	0.002854	0.0962	5.59	104	4.54	260.2	4.0	22.92	1.5	.0962	3.2		
16BS11m-10.2m	0.003167	0.1037	6.34	75	4.01	319.9	4.6	18.41	1.4	.1037	1.4		
G. 1BW8.11.01 (Hidden Lake gneiss) 41°27'57" N, 72°26'27" W													
1BW8-1.1c	0.000132	0.0535	0.02	400	0.68	342.4	2.8	18.37	0.8	.0515	2.8		
1BW8-2.1c	-----	0.0535	0.03	343	0.45	338.7	2.9	18.53	0.8	.0535	2.0		
1BW8-3.1c	0.000144	0.0532	-0.01	343	0.51	341.8	2.9	18.41	0.9	.0511	2.8		
1BW8-4.1c	-----	0.0527	-0.02	646	0.80	326.1	2.5	19.28	0.8	.0527	1.6		
1BW8-5.1c	-----	0.0531	-0.02	437	0.50	342.4	2.8	18.34	0.8	.0531	2.4		
1BW8-6.1c	-----	0.0565	0.41	397	0.42	339.5	2.8	18.42	0.8	.0565	1.8		
1BW8-7.1c	-----	0.0558	0.33	222	0.61	334.6	3.3	18.71	1.0	.0558	3.4		
1BW8-8.1r	-----	0.0536	0.09	613	0.16	321.1	2.5	19.56	0.8	.0536	1.6		
1BW8-9.1c	0.000089	0.0547	0.18	305	0.54	341.3	3.0	18.39	0.9	.0534	2.5		
1BW8-10.1c	-----	0.0545	0.17	246	0.43	336.1	3.2	18.65	1.0	.0545	2.4		
1BW8-11.1c	0.000089	0.0536	0.04	447	0.65	340.3	2.7	18.47	0.8	.0523	2.1		
1BW8-12.1c	-----	0.0523	-0.09	1168	0.28	330.5	2.3	19.02	0.7	.0523	1.2		
1BW8-13.1r	-----	0.0555	0.39	138	0.09	301.3	3.7	20.82	1.2	.0555	3.6		
1BW8-14.1r	-----	0.0543	0.15	401	0.23	331.3	3.0	18.93	0.9	.0543	2.3		
1BW8-15.1r	0.000168	0.0547	0.22	433	0.09	326.8	3.0	19.25	0.9	.0523	3.5		
1BW8-16.1r	-----	0.0547	0.22	351	0.13	325.1	2.8	19.29	0.9	.0547	2.0		
1BW8-17.1r	-----	0.0557	0.34	451	0.11	326.4	2.6	19.19	0.8	.0557	1.8		

TABLE 2
(continued)

sample ¹ (location)	measured		measured $\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	common $\frac{^{206}\text{Pb}}{^{206}\text{Pb}}$	U (ppm)	$\frac{\text{Th}}{\text{U}}$	$\frac{^{206}\text{Pb}^2}{^{238}\text{U}}$ (Ma)	err ³ (Ma)	$\frac{^{238}\text{U}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^4}{^{206}\text{Pb}}$	err ³ (%)	$\frac{^{207}\text{Pb}^{2,5}}{^{206}\text{Pb}}$ (Ma)	err ³ (Ma)
	$\frac{^{204}\text{Pb}}{^{206}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$												
G. IBW8.11.01 (Hidden Lake gneiss)			41°27'57" N, 72°26'27" W											
IBW8-18.1c	-----	0.0536		0.05	603	0.14	336.6	2.6	18.64	0.8	.0536	1.5		
IBW8-19.1r	0.000051	0.0527		-0.02	588	0.10	322.5	2.5	19.51	0.8	.0519	1.8		
IBW8-20.1c	0.000159	0.0547		0.20	496	0.11	335.1	2.7	18.76	0.8	.0524	2.7		
IBW8-21.1r	-----	0.0542		0.18	425	0.12	314.7	2.6	19.95	0.8	.0542	1.9		

¹ Samples with suffix "tm" are titanite; all other samples are zircon. All samples analyzed on the USGS/Stanford ion microprobe (SHRIMP-RG). Abbreviations: c (core), m (mantle), r (rim), d (detrital).

² $\frac{^{206}\text{Pb}}{^{238}\text{U}}$ ages corrected for common Pb using the ^{207}Pb -correction method; $^{207}\text{Pb}/^{206}\text{Pb}$ ages corrected for common Pb using the ^{204}Pb -correction method. Decay constants from Steiger and Jäger (1977).

³ 1-sigma errors.

⁴ radiogenic ratios, corrected for common Pb using the ^{204}Pb -correction method, based on the Stacey and Kramers (1975) model.

⁵ listed only for ages > 1.0 Ga

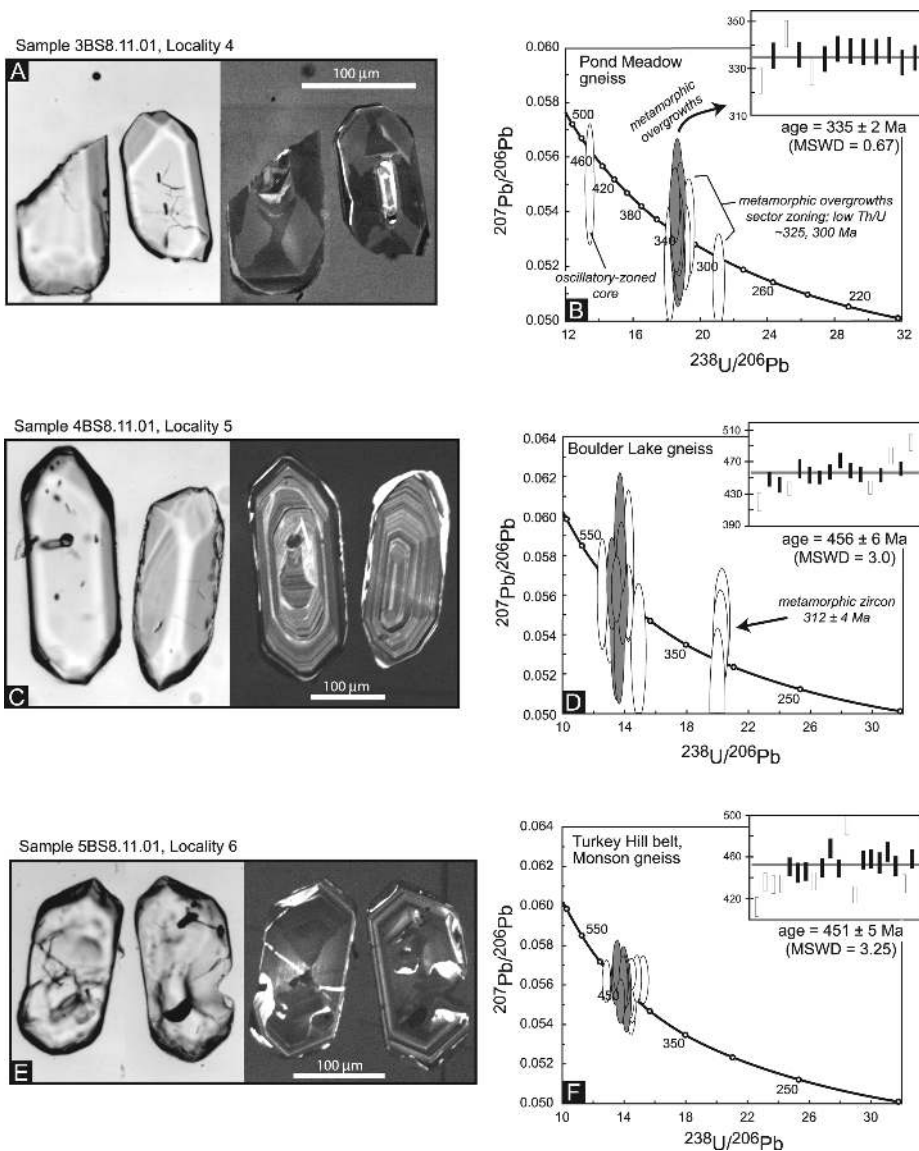


Fig. 8. Digital images (transmitted light and CL) of representative zircons and Tera-Wasserburg Concordia plots (with insets showing weighted averages of $^{206}\text{Pb}/^{238}\text{U}$ ages) for rocks of the Killingworth dome. (A-B) Pond Meadow gneiss (sample 3BS8.11.01); (C-D) Boulder Lake gneiss (4BS8.11.01); (E-F) Turkey Hill belt of Monson Gneiss (1BS8.11.01)

the Essex quadrangle (Wintsch, 1994). It occurs south of the Pond Meadow gneiss and north of the Clinton Granite Gneiss (fig. 3). The Boulder Lake is a medium- to dark-gray, weakly foliated, generally unlayered, medium-grained tonalitic pluton characterized by the common occurrence of mafic and calc-silicate inclusions (Lundgren and Thurrell, 1973). Primary igneous structures are not preserved where strain is high, such as along its southern fault contact with the Clinton Granite Gneiss (Wintsch and others, 2005). Whole-rock compositions show the protolith to be zoned in normative

anorthite content, with a more sodic interior, and more calcic margin (Wintsch and others, 1990).

Zircons from the Boulder Lake gneiss (sample 4BS8.11.01; locality 5, fig. 3) are prismatic, euhedral, and light brown. Most grains show fine concentric, oscillatory zoning in CL (fig. 8C), and some grains have narrow dark or white (in CL), unzoned overgrowths. U contents of Boulder Lake zircons are moderate (150 – 400 ppm, table 2); Th/U is 0.2 to 0.4, typical of igneous origin. Eleven of 16 analyses of the oscillatory-zoned portions of the grains yield an age of 456 ± 6 Ma, interpreted as the time of crystallization of the tonalitic protolith (fig. 8D). Two analyses are older (478 and 495 Ma). Another core, with an irregular rounded boundary and oscillatory zoning truncated at the interface with the igneous mantle, yielded an age of 1.33 Ga, indicating possible Grenville inheritance. Several ages younger than 450 Ma probably are due to Pb-loss, as these were obtained from oscillatory-zoned mantles. Three metamorphic overgrowths yield an age of ~ 312 Ma.

Turkey Hill belt of Monson Gneiss.—The Turkey Hill belt of Monson Gneiss (Lundgren, 1963) occurs in a narrow, north-northwest-trending belt of orthogneiss that crops out on the east side of the Killingworth dome in the eastern part of the study area (fig. 3). Rodgers (1985) showed this belt as a continuous unit that extends north into Massachusetts toward rocks mapped as Monson Gneiss, whereas Webster and Wintsch (1987) grouped these rocks with their “eastern gneiss.” These medium- to coarse-grained rocks include granitic gneiss, tonalitic gneiss, amphibolite, and dunite (Wintsch, 1994). The rock sampled for geochronology is a tan, strongly foliated, medium- to coarse-grained, granitic gneiss. Although originally correlated with other plagioclase gneisses in the area (Rodgers, 1985), this belt of rock was shown by Wintsch (1994) not to be part of the Killingworth complex. Consequently, we dated this sample of Monson Gneiss, but do not include this lithology in our geochemical study.

Zircons from the Turkey Hill belt of Monson Gneiss (sample 1BS8.11.01; locality 6, fig. 3) are euhedral and medium brown; most contain numerous cracks. CL images show fine oscillatory zoning in broad cores and coarse oscillatory zoning in the outer portions of the grains (fig. 8E). Many grains are invaded by wisps of white (in CL) material, suggestive of subsequent metamorphic dissolution and replacement. The zircons have moderately high U contents (600 – 2100 ppm) and Th/U of 0.3 to 0.6 (table 2). Twelve of 20 analyses form a coherent group with an age of 451 ± 5 Ma (fig. 8F), interpreted as the time of crystallization of the igneous protolith. One analysis is older (~ 486 Ma) and may reflect derivation of this gneiss from slightly older Ordovician source rocks possibly related to the Bronson Hill terrane. Younger ages (~ 410 – 435 Ma) probably are due to Pb loss from damaged, fractured grains, and thus are considered to be geologically meaningless.

Middletown Formation.—The Middletown and Collins Hill formations of Rodgers (1985) crop out on the outer margins of the Killingworth dome, except on the southeast side (fig. 3). These map units include biotite schist, plagioclase gneiss, anthophyllite gneiss, amphibolite, and calc-silicate schist (Bernold, 1976), a lithologic package suggesting a volcanic-sedimentary protolith. The Middletown Formation is considered to be primarily metavolcanic, whereas the Collins Hill Formation is mostly of sedimentary origin (Lundgren, 1979). These two units are interlayered; distinction between the two formations is based on proportion of inferred igneous or sedimentary components. On the west side of the Killingworth dome, the Middletown Formation is structurally above the Higganum gneiss with a gradational contact. The Higganum is primarily a felsic rock with lesser amounts of amphibolite and schist, whereas the Middletown is primarily an amphibolite with lesser amounts of meta-diorite and schist. The protoliths of both units probably were volcanic. Anthophyllite gneiss of the

Middletown Formation is medium-gray, strongly foliated, strongly lineated, well-layered, medium- to coarse-grained plagioclase gneiss. The occurrence within the anthophyllite gneiss of a minor population of rounded zircon grains interpreted as detrital in origin supports the suggestion of Lundgren (1979) that the protolith of the anthophyllite gneiss was an extrusive rock. Collectively, the Middletown and Collins Hill formations and the Higganum gneiss constitute our Middletown complex.

Zircons from the anthophyllite gneiss (samples 2BS8.11.01 and 16BS11.11.02; localities 7 and 8, respectively, fig. 3) have a very unusual morphology. Almost all grains are dark brown and anhedral with irregular, amoeboid shapes. In CL, they show dark, mottled textures with sparse, very dark spots irregularly overgrown by pale rims (fig. 9A). In back-scattered electron (BSE) imaging, most grains contain small, randomly distributed, very bright spots identified as thorite (using XRF energy-dispersive analysis). The white (in CL) rim material commonly invades the grains. Consequently, there are very few pristine, homogeneous areas large enough to accommodate a SHRIMP analysis without contamination by an adjacent zone. We suggest that the unusual anhedral, mottled grains are primarily igneous in origin, with discontinuous metamorphic overgrowths. A second population of zircon (less than 1% of the total) also occurs in the anthophyllite gneiss. These grains are roughly equant, rounded, light brown to colorless, and have a variety of types of oscillatory zoning, all of which are truncated at grain boundaries (fig. 9A). As shown by the U-Pb age data below, we interpret these rounded grains to be detrital in origin.

By imaging in both CL and BSE, and examining thousands of grains during five separate SHRIMP sessions, we were able to obtain sufficient U-Pb data to have confidence in assigning an age to this rock. The anhedral, mottled grains have high U contents (800 – 3600 ppm) and igneous-type Th/U ratios of 0.2 to 0.6 (table 2). The light-colored (in CL) overgrowths have lower U (mostly 800 – 1200 ppm) and very low Th/U (0 – 0.04) typical of zircons of metamorphic origin. Rounded, oscillatory-zoned grains interpreted as detrital have moderate U contents (100 – 300 ppm) and Th/U of 0.2 to 0.6 typical of igneous origin. Isotopic data from 67 analyses (5 analytical sessions) from all grain types yield a broad spectrum of ages (fig. 9B). Thirteen analyses (from 4 sessions) of homogeneous core zones result in a coherent group of data with an age of 449 ± 4 Ma (fig. 9C). Six overgrowths yield ages of about 350 to 320 Ma; one other overgrowth is about 270 Ma. Many more analyses yielded intermediate ages, and probably represent the result of mixing of Ordovician core regions with Carboniferous overgrowths. Zircons interpreted as detrital in origin have much older ages, between about 0.9 to 1.3 Ga (fig. 9B).

Titanite from the anthophyllite gneiss occurs as light-brown to yellowish, anhedral grains. BSE imaging shows that many are complexly zoned, with dark (that is, relatively low-U) cores and mantles overgrown by white to light gray (that is, relatively high-U) rims (fig. 9D). A few cores display oscillatory zoning that reflects crystal faces. Titanite cores contain about 20 to 135 ppm U, whereas the rims contain about 100 to 290 ppm. All isotopic analyses indicate the presence of a significant common Pb component (table 2). Thus, U-Pb ages are determined by calculating regressions through the isotopic data (uncorrected for common Pb) to determine concordia intercept ages of 304 ± 9 and 249 ± 4 Ma for the cores and rims, respectively (fig. 9E). We suggest that these ages record the times of titanite growth during high-grade, and locally anatectic, episodes of Alleghanian metamorphism.

Hidden Lake gneiss.—The Hidden Lake gneiss is located in the core of the Killingworth dome (fig. 3). These rocks contain relatively calcic plagioclase (Mikami and Digman, 1957), low mafic mineral content and shallow-dipping foliations (Lundgren and Thurrell, 1973), and distinctively high Al_2O_3 , Sr, and Ba contents (figs. 6G

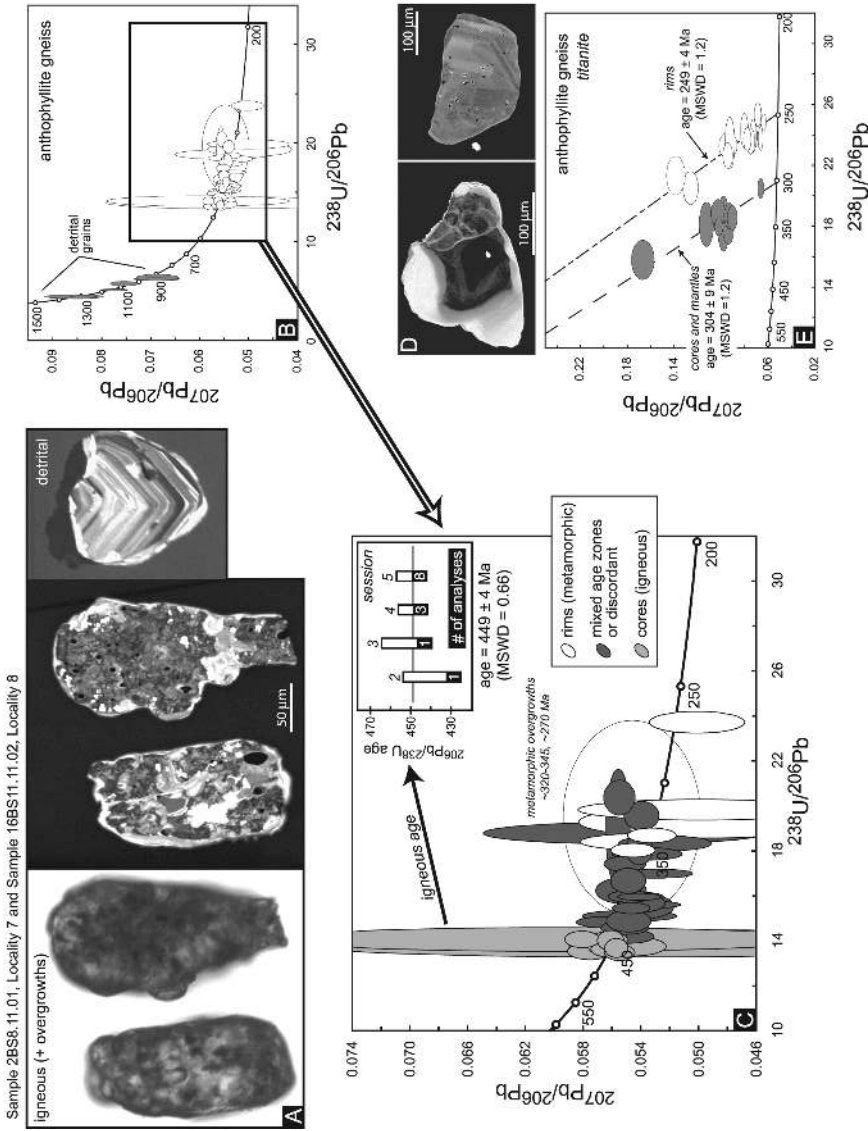


Fig. 9. Middletown gneiss (anthophyllite gneiss). (A) Transmitted light and CL images. Note truncated oscillatory zoning of detrital grain. (B) Tera-Wasserburg Concordia plot. Analyses of rounded, light colored grains (interpreted as detrital in origin) shown as light gray error ellipses. (C) Enlargement of Paleozoic data shown in B. Analyses shown as light gray error ellipses interpreted as igneous in origin. All other ellipses are younger ages of overgrowths or mixtures of cores and overgrowths. (D) Backscattered electron images of representative grains of titanite. (E) Tera-Wasserburg plot of SHRIMP U-Pb data (uncorrected for common Pb).

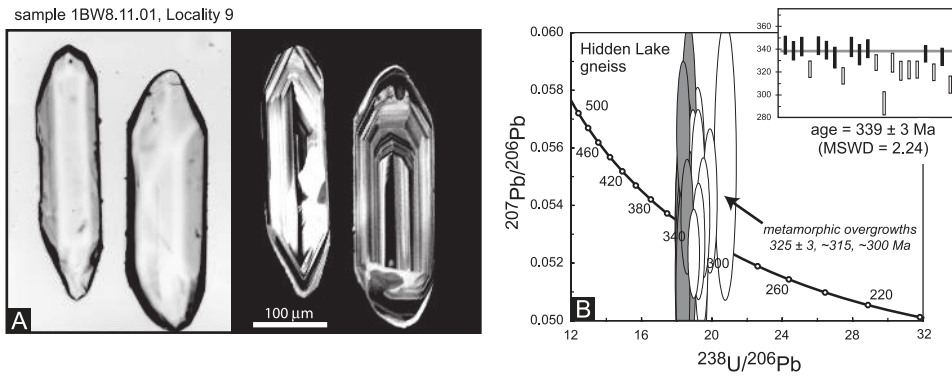


Fig. 10. Digital images (A, transmitted light and CL) of representative zircons and Tera-Wasserburg Concordia plot (B, with insets showing weighted averages of $^{206}\text{Pb}/^{238}\text{U}$ ages) for Hidden Lake gneiss (sample 1BW8.11.01).

and 6H; Webster and Wintsch, 1987). This unit is typically composed of light- to medium-gray, weakly foliated and generally unlayered, medium- to coarse-grained tonalitic gneiss and granofels. Chemical and modal analyses confirm that the Hidden Lake gneiss is a calc-alkaline meta-tonalitic pluton, zoned with an increasing normative anorthite content toward the core, and low amphibole content (fig. 4; Webster and Wintsch, 1987).

Zircons from the Hidden Lake gneiss (sample 1BW8.11.01; locality 9, fig. 3) are generally euhedral and prismatic, with few inclusions. All grains show fine concentric, oscillatory zoning in CL; many grains have dark, unzoned tips (fig. 10A) that locally invade deep into the cores. The oscillatory-zoned portions contain moderate amounts of U (about 200 – 400 ppm) and relatively high Th/U (0.5 – 0.8, table 2) typical of magmatic zircons, whereas the tips have higher U-contents (mostly 400 – 700 ppm) and have much lower Th/U (0.09 – 0.14) typical of metamorphic zircons (Hoskin and Schaltegger, 2003). Eleven analyses of the oscillatory-zoned interiors of these grains yield an age of 339 ± 3 Ma (fig. 10B), interpreted as the time of crystallization of the igneous protolith of this gneiss. Overgrowths yield a range of ages, including 325 ± 3 Ma (6 analyses) and ~ 315 and ~ 300 Ma (1 analysis each), dating times of metamorphic overprinting. This metamorphism and associated deformation produced a moderate foliation and compositional banding that dips gently north-northwest.

Summary.—Igneous protoliths of three orthogneisses in the Killingworth dome (Higganum, Pond Meadow, and Boulder Lake), crystallized in the Middle to Late Ordovician (approximately 455 – 460 Ma). Due to the strong metamorphic overprint and occurrence of interlayered amphibolite of indeterminate origin, it is unclear whether the protoliths of the Higganum and Pond Meadow gneisses were intrusive or extrusive. The homogeneous character and abundant xenoliths within Boulder Lake gneiss suggest that it is plutonic in origin. The Turkey Hill belt of Monson Gneiss (probably intrusive) and Middletown Formation (probably extrusive) are slightly younger (Late Ordovician, about 450 Ma). Hidden Lake gneiss, in the core of the Killingworth dome, was emplaced in the Mississippian (339 ± 3 Ma). The Pond Meadow gneiss was thoroughly migmatized at about this time, as shown by the occurrence of extensive zircon overgrowths dated at 335 ± 2 Ma. The Clinton Granite Gneiss, crystallized at about 605 Ma, and represents Gondwanan basement structurally below, and in fault contact with, the Boulder Lake gneiss.

Zircon overgrowths in several samples record late Paleozoic thermal events at about 325, 315, 300, and 270 Ma. Ages of titanite in the Middletown anthophyllite gneiss corroborate high-grade metamorphic conditions at about 300 Ma and 250 Ma.

Pb, Nd, and Sr Isotopic Data

Samples collected for whole-rock chemical analysis were also analyzed for Pb, Sr, and Nd isotopes (tables 3 and 4). Procedures for isotopic analysis are discussed in the Appendix. Two analyses of Sr isotopes corrected to impossibly low initial values, probably due to preferential loss of radiogenic ^{87}Sr from a high Rb/Sr phase such as biotite during late Paleozoic metamorphism. Thus, all Sr isotopic data are considered to possibly have been affected by metamorphism and are not discussed further. However, because the estimated amount of loss of radiogenic ^{87}Sr is less than 1 percent of the total Sr, concentrations of Sr used for trace element discrimination diagrams and discussions of geochemical and petrologic processes are representative of the rock chemistry.

Pb isotopic data for whole-rock samples from five units of the Killingworth and Middletown complexes plot near and above the average growth curve of Stacey and Kramers (1975) (fig. 11A). Fields of data for Pond Meadow, Boulder Lake, and Hidden Lake gneisses, and all but one analysis of the Middletown Formation mostly overlap. Note that sample 14J (Hidden Lake gneiss, table 3) has anomalously high Pb isotopic ratios and has been excluded from all plots. The field of Pb isotopic data for the Higganum gneiss is distinctly more radiogenic in $^{207}\text{Pb}/^{204}\text{Pb}$, as is one analysis of anthophyllite gneiss of the Middletown Formation. Only the intermediate to felsic rocks (SiO_2 between about 55 and 76%) within the Middletown were analyzed for geochemistry and isotopic ratios; mafic rocks (that is, amphibolites) were not analyzed.

Nd isotopic data from the five sampled lithologies have a wide range (fig. 11B). Most samples of the Ordovician Boulder Lake and Pond Meadow gneisses have relatively low initial ϵ_{Nd} values between about -5.5 and -1.6; the Mississippian Hidden Lake gneiss has similar values of -4.3 to -2.8 (table 4). These low initial ϵ_{Nd} values partly overlap data from Neoproterozoic basement rocks of the Gander Zone (about -4 to +2 at 450 Ma) and partly overlap into the field of Laurentian (Grenville) crust (about -8 to -4 at 450 Ma). Isotopic data from samples of Ordovician Middletown Formation and Higganum gneiss have much higher initial ϵ_{Nd} values of about +0.9 to +5.2 (table 4). These higher values mostly overlap the field of Nd data from the Avalon Zone (fig. 11B).

Initial ϵ_{Nd} values for Neoproterozoic gneisses of the Clinton and Lyme domes are relatively low (-2.3 to +0.9, table 4), plotting within the field of Gander zone data. Neoproterozoic gneiss samples 15W and 15X from the Selden Neck block (fig. 3) have significantly higher initial ϵ_{Nd} values of about 3.9 and 4.6, respectively, and plot within the field of data for the Avalon Zone (fig. 11B).

DISCUSSION

Geochronology and Geology

On the basis of whole-rock geochemistry, Webster and Wintsch (1987) subdivided the rocks of the Killingworth dome into four units (fig. 3). This subdivision is supported by our more recent geochemical data set. U-Pb geochronology shows that ages of three units (Pond Meadow, Boulder Lake, and Higganum gneisses) are about 455 to 460 Ma, whereas the Hidden Lake gneiss in the core of the dome crystallized

TABLE 3
Whole-rock Pb isotopic data from rocks of the Killingworth and Middletown complexes, Connecticut

Sample	Concentrations (ppm)				measured ratios				calculated initial ratios							
	U	Th	Pb	Pb	$\frac{^{238}\text{U}}{^{204}\text{Pb}}$	$\frac{^{232}\text{Th}}{^{208}\text{Pb}}$	$\frac{^{206}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	$\frac{^{208}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{206}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{208}\text{Pb}}{^{204}\text{Pb}}$	err ²	err ²		
Hidden Lake gneiss																
14G	1.05	2.69	22.44	2.97	7.89	18.804	15.637	38.288	0.83160	2.0362	18.642	0.015	15.628	0.014	38.154	0.046
14H	0.37	1.47	11.05	2.13	8.71	18.572	15.607	38.200	0.84033	2.0568	18.457	0.015	15.601	0.015	38.053	0.047
14J	1.14	4.27	18.14	4.10	15.79	20.165	15.706	38.364	0.77887	1.9025	19.943	0.017	15.694	0.014	38.096	0.047
14P	0.54	2.56	9.05	3.75	18.53	18.409	15.977	38.346	0.84728	2.0829	18.206	0.015	15.587	0.015	38.031	0.047
15F	1.09	11.20	33.57	2.07	21.98	18.605	15.624	38.543	0.83979	2.0716	18.493	0.014	15.618	0.014	38.170	0.047
1BW8	1.33	54.65	31.24	2.76	117.18	18.469	15.618	39.917	0.84565	2.1613	18.319	0.015	15.610	0.014	37.929	0.062
Pond Meadow gneiss																
14L	1.47	3.46	16.40	5.70	13.89	18.783	15.623	38.410	0.83176	2.0449	18.474	0.017	15.607	0.014	38.174	0.046
15C	1.17	4.87	13.97	5.34	22.98	18.754	15.637	38.469	0.83381	2.0512	18.465	0.017	15.622	0.014	38.079	0.047
15D	0.43	5.18	14.06	1.97	24.30	18.645	15.652	38.624	0.83943	2.0715	18.539	0.016	15.646	0.015	38.212	0.050
15E	0.75	6.26	15.11	3.17	27.42	18.801	15.636	38.727	0.83165	2.0598	18.630	0.015	15.627	0.014	38.262	0.048
15P	0.95	4.37	19.54	3.09	14.71	18.655	15.625	38.348	0.83757	2.0556	18.488	0.015	15.616	0.015	38.099	0.047
15R	0.65	4.71	13.52	3.05	23.07	18.798	15.651	38.687	0.83258	2.0581	18.632	0.015	15.642	0.014	38.295	0.047
3BS8	1.22	4.88	14.41	5.42	22.34	18.790	15.643	38.493	0.83251	2.0486	18.496	0.019	15.627	0.014	38.114	0.048
15G	1.00	4.21	25.42	2.51	10.87	18.604	15.625	38.339	0.83984	2.0607	18.469	0.015	15.618	0.014	38.154	0.046
15Q	1.27	4.65	13.45	6.07	22.98	19.197	15.661	38.602	0.81580	2.0108	18.748	0.020	15.636	0.014	38.073	0.048
Middletown Formation																
14D mafic	0.25	0.48	5.18	3.03	6.14	18.592	15.625	38.266	0.84040	2.0582	18.423	0.016	15.616	0.015	38.159	0.048
14E			6.06			18.714	15.652	38.355	0.83639	2.0496						
14F mafic	0.62	1.44	76.75	0.51	1.22	18.267	15.591	38.110	0.85353	2.0862	18.239	0.014	15.590	0.014	38.088	0.047
14K mafic	0.25	0.55	18.30	0.87	1.96	18.670	15.635	38.308	0.83741	2.0518	18.622	0.014	15.632	0.014	38.274	0.046
15B	0.47	1.51	3.83	7.98	26.44	19.270	15.693	39.010	0.81441	2.0244	18.825	0.024	15.670	0.018	38.548	0.055
14C	0.75	0.78	4.53	10.53	11.38	18.901	15.645	38.304	0.82770	2.0265	18.122	0.032	15.601	0.017	38.042	0.051
Turkey Hill belt of Monson Gneiss																
15A	0.72	2.61	19.19	2.37	8.91	18.601	15.621	38.293	0.83977	2.0586	18.430	0.015	15.611	0.014	38.092	0.047
1BS8	2.49	8.66	20.27	7.93	28.51	19.107	15.648	38.980	0.81897	2.0400	18.534	0.023	15.616	0.014	38.338	0.049
15U	0.72	5.22	11.44	4.03	30.10	18.673	15.623	38.560	0.83667	2.0650	18.381	0.024	15.607	0.020	37.882	0.061
Boulder Lake gneiss																
15K	1.40	19.98	27.03	3.35	49.48	18.603	15.658	39.711	0.84173	2.1347	18.358	0.016	15.645	0.014	38.584	0.053
15L	1.11	8.16	12.61	5.66	43.15	19.008	15.640	39.020	0.82280	2.0529	18.594	0.020	15.616	0.015	38.038	0.052
15M	1.82	5.44	9.27	12.74	39.42	19.447	15.673	39.073	0.80594	2.0092	18.516	0.033	15.621	0.015	38.176	0.052
15N	2.06	7.80	10.41	12.87	50.39	19.381	15.673	39.298	0.80865	2.0277	18.440	0.034	15.620	0.015	38.151	0.054

TABLE 3
(continued)

Sample	Concentrations (ppm)				measured ratios						calculated initial ratios						
	U	Th	Pb		$\frac{^{238}\text{U}}{^{204}\text{Pb}}$	$\frac{^{232}\text{Th}}{^{204}\text{Pb}}$	$\frac{^{206}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{208}\text{Pb}}{^{206}\text{Pb}}$	$\frac{^{207}\text{Pb}}{^{206}\text{Pb}}$	$\frac{^{208}\text{Pb}}{^{204}\text{Pb}}$	$\frac{^{206}\text{Pb}^1}{^{206}\text{Pb}}$	err^2	$\frac{^{207}\text{Pb}^1}{^{204}\text{Pb}}$	err^2	$\frac{^{208}\text{Pb}^1}{^{204}\text{Pb}}$	err^2
Boulder Lake gneiss																	
15S	1.38	2.98	20.04		4.36	9.75	18.642	15.616	38.252	0.83765	2.0519	18.323	0.017	15.598	0.014	38.030	0.046
15T	1.50	14.02	10.35		9.45	91.15	19.210	15.679	39.494	0.81616	2.0559	18.519	0.026	15.640	0.014	37.419	0.063
4B88	0.66	8.51	10.62		4.02	53.48	18.854	15.651	39.276	0.83010	2.0832	18.560	0.020	15.634	0.016	38.059	0.057
Higganum gneiss																	
14A	0.80	1.06	8.22		6.26	8.50	18.990	15.687	38.475	0.82607	2.0261	18.527	0.023	15.661	0.015	38.280	0.049
14B	0.43	0.64	32.10		0.86	1.31	18.517	15.670	38.293	0.84626	2.0680	18.454	0.014	15.667	0.014	38.263	0.046
14M	0.82	3.05	3.94		13.53	52.00	19.588	15.702	39.113	0.80161	1.9968	18.588	0.036	15.646	0.018	37.916	0.059
14N			8.11				19.469	15.694	38.457	0.80607	1.9753						
Ha-1-02	0.64	1.47	7.02		5.91	13.94	19.099	15.702	38.677	0.82212	2.0251	18.662	0.021	15.677	0.016	38.356	0.049
Neoproterozoic basement rocks																	
<i>Gander terrane</i>																	
15H	1.08	8.39	27.41		2.57	16.44	18.822	15.691	39.603	0.83367	2.1041	18.569	0.016	15.676	0.014	38.979	0.050
15J	1.95	17.27	21.47		5.90	54.14	19.158	15.675	39.516	0.81819	2.0626	18.582	0.024	15.640	0.014	37.884	0.060
5B88	2.52	29.06	25.30		6.49	77.46	19.179	15.717	39.640	0.81951	2.0669	18.546	0.024	15.679	0.015	37.306	0.069
0L-290	0.89	14.86	13.76		4.23	73.06	18.837	15.677	40.213	0.83227	2.1348	18.424	0.019	15.653	0.015	38.012	0.065
0L-293	0.78	11.20	16.12		3.10	46.03	18.316	15.642	39.266	0.85398	2.1438	18.014	0.017	15.624	0.015	37.880	0.055
<i>Avalon terrane</i>																	
15W	0.68	2.65	8.98		4.84	19.46	18.83	15.627	38.432	0.82963	2.0404	18.364	0.021	15.598	0.015	37.846	0.049
15X	0.72	4.30	9.94		4.67	28.66	18.781	15.621	38.740	0.83176	2.0627	18.326	0.020	15.594	0.014	37.876	0.050

¹ measured ratios corrected for *in situ* growth of radiogenic Pb, using the zircon U-Pb age of the unit. Decay constants from Steiger and Jäger (1977).
² 2σ uncertainties

TABLE 4
Whole-rock Sr and Nd isotopic data from rocks of the Killingworth and Middletown complexes, Connecticut

Sample	Concentrations (ppm)				measured ratios				calculated initial ratios				TDM ³	
	Rb	Sr	Sm	Nd	$\frac{87\text{Rb}}{86\text{Sr}}$	$\frac{87\text{Sr}}{86\text{Sr}}$	$\frac{147\text{Sm}}{144\text{Nd}}$	$\frac{143\text{Nd}}{144\text{Nd}}$	ϵ_{Nd}	$\frac{87\text{Sr}^1}{86\text{Sr}}$	err ²	$\frac{143\text{Nd}^1}{144\text{Nd}}$		err ²
Hidden Lake gneiss														
14G	46.6	398.8	0.71	4.80	0.338	0.707480	0.0896	0.512208	-8.35	0.705845	0.000032	0.512009	0.000023	-3.70
14H	45.7	491.0	0.65	4.85	0.270	0.706861	0.0810	0.512159	-9.30	0.705557	0.000022	0.511979	0.000014	-4.28
14J	32.8	482.5	1.24	8.04	0.197	0.706649	0.0932	0.512234	-7.84	0.705698	0.000022	0.512026	0.000014	-3.35
14P	44.8	482.5	1.07	7.25	0.269	0.707059	0.0892	0.512251	-7.51	0.705758	0.000024	0.512052	0.000015	-2.84
15F	96.8	377.3	4.64	20.44	0.743	0.709682	0.1370	0.512109	-10.28	0.706088	0.000313	0.511804	0.000016	-7.69
1BW8	113.9	299.2	5.76	48.10	1.102	0.711554	0.0723	0.512206	-8.39	0.706220	0.000090	0.512045	0.000012	-2.98
Pond Meadow gneiss														
14L	75.7	74.9	2.90	12.72	2.927	0.724167	0.1375	0.512389	-4.82	0.710002	0.000278	0.512083	0.000010	-2.24
15C	45.9	81.9	2.19	11.18	1.622	0.718663	0.1184	0.512181	-8.88	0.710812	0.000086	0.511917	0.000015	-5.48
15D	44.2	114.7	2.27	12.70	1.115	0.715686	0.1078	0.512201	-8.49	0.710291	0.000455	0.511961	0.000021	-4.63
15E	65.6	145.7	3.15	14.85	1.304	0.714884	0.1280	0.512334	-5.89	0.708571	0.000069	0.512049	0.000012	-2.91
15P	52.6	91.0	7.56	37.35	1.674	0.718549	0.1222	0.512282	-6.91	0.710450	0.000073	0.512010	0.000014	-3.67
15R	50.9	116.8	1.64	6.61	1.261	0.715159	0.1499	0.512452	-3.59	0.709058	0.000498	0.512118	0.000015	-1.55
3B88	43.8	94.9	1.44	7.36	1.336	0.716614	0.1179	0.512179	-8.91	0.710147	0.000061	0.511917	0.000016	-5.49
15G	46.5	79.5	2.51	10.88	1.693	0.717728	0.1393	0.512409	-4.43	0.709536	0.000082	0.512099	0.000027	-1.93
15Q	60.9	123.5	2.42	11.14	1.427	0.716961	0.1313	0.512227	-7.98	0.707608	0.000668	0.511832	0.000014	-4.13
Middletown Formation														
14D mafic	7.9	54.1	0.42	1.11	0.420	0.708965	0.2268	0.512905	5.25	0.706872	0.000026	0.512385	0.000016	3.91
14E	8.2	106.8	0.51	1.78	0.223	0.709207	0.1736	0.512732	1.87	0.708097	0.000022	0.512334	0.000058	2.91
14F mafic	21.4	1760.5	6.33	20.87	0.035	0.707453	0.1833	0.512872	4.60	0.707278	0.000032	0.512452	0.000027	5.21
14K mafic	14.3	95.1	1.02	3.45	0.436	0.708871	0.1775	0.512750	2.22	0.706698	0.000024	0.512343	0.000014	3.09
15B	5.7	98.0	1.51	6.07	0.169	0.709549	0.1503	0.512568	-1.33	0.708707	0.000057	0.512224	0.000013	0.76
14C	10.8	75.3	0.48	1.65	0.413	0.707113	0.1764	0.512799	3.18	0.704404	0.000028	0.512268	0.000027	4.39
Turkey Hill belt of Monson Gneiss														
15A	89.0	73.5	1.64	8.11	3.507	0.726117	0.1218	0.512361	-5.36	0.703635	0.000261	0.512002	0.000028	-1.05
1B88	178.8	2.45	11.17				0.1323	0.512456	-3.51			0.512066	0.000014	0.20
15U	50.7	95.2	1.45	6.43	1.541	0.717038	0.1363			0.707157	0.000385			
Boulder Lake gneiss														
15K	97.0	140.2	3.94	21.16	2.004	0.718903	0.1125	0.511912	-14.12	0.705916	0.000156	0.511577	0.000038	-9.24
15L	54.8	161.4	3.00	15.35	0.983	0.714352	0.1179	0.512232	-7.88	0.707981	0.000087	0.511881	0.000035	-3.30
15M	81.7	163.2	3.49	15.09	1.449	0.714870	0.1394	0.512211	-8.29	0.705477	0.000106	0.511795	0.000013	-4.96

TABLE 4
(continued)

Sample	Concentrations (ppm)				measured ratios				calculated initial ratios						
	Rb	Sr	Sm	Nd	$\frac{87\text{Rb}}{86\text{Sr}}$	$\frac{87\text{Sr}}{86\text{Sr}}$	$\frac{147\text{Sm}}{144\text{Nd}}$	$\frac{143\text{Nd}}{144\text{Nd}}$	ϵ_{Nd}	$\frac{87\text{Sr}}{86\text{Sr}}$	err^2	$\frac{143\text{Nd}}{144\text{Nd}}$	err^2	ϵ_{Nd}	TDM ³
Boulder Lake gneiss															
15N	79.8	173.7	2.75	13.83	1.331	0.716541	0.1200	0.512239	-7.74	0.707916	0.000097	0.511881	0.000027	-3.28	1.30
15S	5.3	238.8	3.67	12.03	0.064	0.707774	0.1842	0.512469	-3.26	0.707359	0.000020	0.511920	0.000015	-2.53	3.05
15T	110.9	184.7	4.14	22.37	1.738	0.714067	0.1119	0.512232	-7.88	0.702804	0.004827	0.511899	0.000015	-2.95	1.21
4BS8	89.8	157.1	3.52	19.13	1.656	0.716778	0.1112	0.512236	-7.80	0.706043	0.000130	0.511905	0.000016	-2.83	1.19
Higganum gneiss															
14A	10.2	82.5	1.76	5.04	0.359	0.708839	0.2105	0.512850	4.17	0.706487	0.000046	0.512216	0.000029	3.37	--
14B	10.8	87.1	0.53	1.75	0.357	0.709394	0.1843	0.512667	0.60	0.707052	0.000039	0.512112	0.000025	1.34	1.91
14M	1.6	127.6	1.93	8.86	0.036	0.707720	0.1317	0.512486	-2.93	0.707488	0.000033	0.512089	0.000014	0.90	1.03
14N	96.5	85.8		3.79	3.258	0.719191		0.512782	2.85	**					
HA-1-02	20.2	82.9	1.93	6.00	0.705	0.711156	0.1943	0.512755	2.32	0.706533	0.000124	0.512169	0.000017	2.47	2.24
Neoproterozoic basement rocks															
<i>Gander terrane</i>															
15H	223.0	40.4	3.64	17.24	16.158	0.837598	0.1276	0.512244	-7.65	**		0.511738	0.000027	-2.30	1.40
15J	168.4	110.8	6.48	33.97	4.410	0.740218	0.1151	0.512360	-5.38	0.702485	0.000619	0.511907	0.000033	0.88	1.05
5BS8	184.5	148.1	9.56	52.22	3.612	0.732868	0.1106	0.512337	-5.83	0.701707	0.001185	0.511899	0.000015	0.84	1.04
OL-290	111.1	121.5	12.41	80.70	2.650	0.726857	0.0929	0.512225	-8.02	0.703426	0.000313	0.511848	0.000016	0.22	1.03
OL-293	96.6	233.3	3.44	19.15	1.199	0.715975	0.1083	0.512222	-8.08	0.705377	0.000157	0.511782	0.000014	-1.07	1.18
<i>Avalon terrane</i>															
15W	28.9	281.5	3.51	19.31	0.297	0.705655	0.1099	0.512495	-2.75	0.703111	0.000026	0.512063	0.000016	3.92	0.81
15X	52.2	281.8	2.41	14.63	0.536	0.707956	0.0995	0.512490	-2.85	0.703367	0.000061	0.512099	0.000015	4.62	0.74

** Calculated initial ratios are impossibly low

¹ measured ratios corrected for *in situ* growth of radiogenic Sr and Nd, using the zircon U-Pb age of the unit. Decay constants from Steiger and Jäger (1977).² 2 σ uncertainties³ model age using the depleted mantle model of DePaolo (1981)

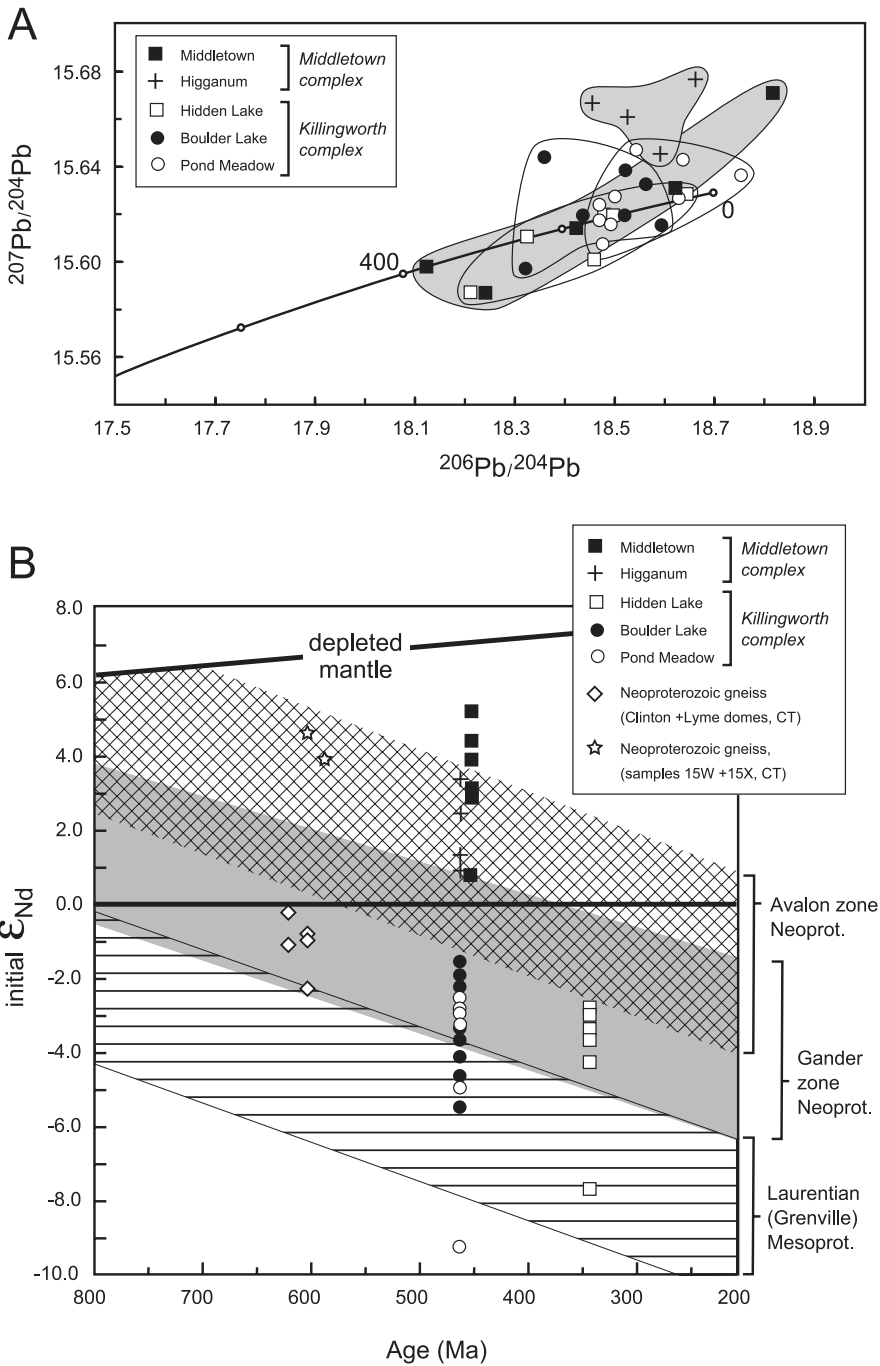


Fig. 11. Pb and Nd isotopic data from rocks of the Killingworth dome. (A) Initial Pb isotopic compositions of whole-rock samples from the Killingworth and Middletown complexes, corrected to the crystallization age of their igneous protoliths. Curve is average crustal Pb evolution of Stacey and Kramers (1975). (B) Initial ϵ_{Nd} versus age for whole-rock samples of the Killingworth and Middletown complexes, and samples of Neoproterozoic basement rocks from southern Connecticut. Fields of Avalon zone and Gander zone from Kerr and others (1995); Laurentian (Grenville) field compiled by Samson and others (2000). Depleted mantle evolution curve from DePaolo (1981).

about 340 Ma. Nd and Pb isotopic compositions suggest that the Boulder Lake, Pond Meadow and Hidden Lake magmas (now grouped as the Killingworth complex) were derived from similar crustal sources while the Higganum and Middletown magmas (now grouped as the Middletown complex) were derived from different sources.

The nature of the contact between the Killingworth and Middletown complexes is a matter of controversy. On the east side of the Killingworth dome, Rodgers (1985) shows this contact as a stratigraphic boundary, implying that the Middletown Formation was deposited on plagioclase gneisses of the dome. Mapping by Wintsch (1994) at the southeast end of the belt of Middletown Formation, in the north-central part of the Essex 7.5' quadrangle, suggests a ductile fault between Middletown and rocks now called Pond Meadow and Boulder Lake gneiss. Additional evidence for a fault contact between the Middletown and Killingworth complexes is: (1) the southeast-trending straight boundary between the Middletown and Hidden Lake gneiss in the southeastern Haddam quadrangle (fig. 3), (2) truncation of chemical zoning within the Hidden Lake gneiss by the straight contact (Webster and Wintsch, 1987), and (3) contrast in degree of metamorphic overprint between the pervasively migmatized Pond Meadow gneiss and the strongly foliated and recrystallized Higganum gneiss and Middletown Formation (only recrystallized). Because the Hidden Lake gneiss intruded these rocks at about 340 Ma, and shows no evidence of intruding into the Middletown, we consider the contact to be a post-340 Ma fault. Similarly, Lundgren (1979) mapped a thrust fault contact separating a small body of amphibolite (interpreted as Middletown Formation) from the adjacent Pond Meadow gneiss in the north-central part of the Haddam quadrangle. Thus, we conclude that the Killingworth complex and Middletown complex (including the Higganum gneiss) were juxtaposed after 340 Ma (age of zircon overgrowths in the Pond Meadow and new magmatic zircon in the Hidden Lake gneiss). The age of titanite in the Middletown anthophyllite gneiss (about 300 Ma) may date the time of faulting.

Petrogenesis and Magmatic Sources

Nearly all rocks of the Killingworth dome exhibit geochemical characteristics indicative of an origin within a magmatic arc or arc-related system. All rocks are magnesian according to the criteria of Frost and others (2001) (fig. 6F), and exhibit calc-alkaline affinity (fig. 12). Nearly all Killingworth rocks are characterized by compositions that are slightly Mg-poor relative to lavas from the mature Cascades arc (Carmichael, 1964). The Hidden Lake gneiss is distinct from other Killingworth gneisses in defining a compositional trend that is nearly coincident with the Cascade rocks. Mafic samples of the Middletown Formation plot separate from associated felsic varieties of the Middletown (fig. 12), defining a possible trend characterized by slight Mg enrichment relative to the Cascade lavas.

All felsic rocks from the Killingworth dome are characterized by a combination of relatively low Rb and Nb + Y compared to other granitoids worldwide, and thus display compositional similarity to granitoids from volcanic-arc tectonic settings (fig. 13A; Pearce and others, 1984). Island-arc affinity is also indicated by relatively low concentrations of Ta and Yb (fig. 13B). However, differences in Rb concentrations (which are consistent within units, suggesting relative immobility on the scale of individual samples) in the rocks studied define two groups. Felsic rocks of the Middletown Formation and most of the Higganum gneiss exhibit Rb depletion whereas all other gneisses display relative enrichment. Low Rb concentrations of the Higganum and Middletown rocks correspond to compositions that are characteristic of relatively immature arc suites with minimal involvement of continental crust in magma genesis, such as the Kermadec Islands in the southwest Pacific (Smith and others, 1988) and

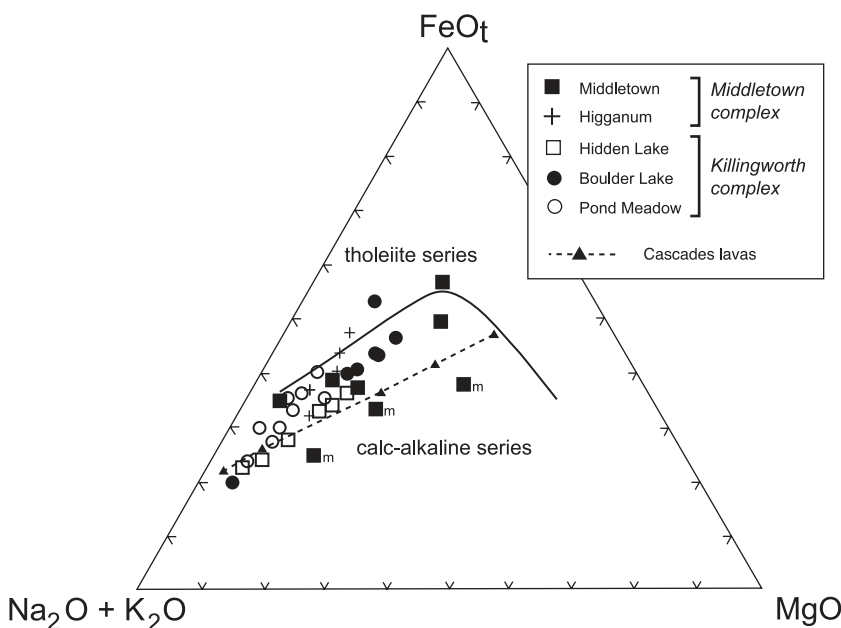


Fig. 12. AFM plot of $(\text{Na}_2\text{O} + \text{K}_2\text{O})$ vs FeO_t vs MgO for gneisses of the Killingworth dome. Coordinates are calculated using weight percent of the element oxides. Mafic samples of the Middletown Formation are labelled with "m". Solid line separating fields of tholeiitic and calc-alkaline series from Irvine and Baragar (1971). Compositions of Cascade lavas (dashed line) from Carmichael (1964).

mid-Oligocene plutonic rocks of the Cordillera de Talamanca in Costa Rica and Panama (Drummond and others, 1995). In contrast, Förster and others (1997) noted that higher Rb concentrations at comparable Y + Nb values, compositional features exhibited by felsic gneisses of the Killingworth complex (fig. 13A), are characteristic of more mature island arcs or evolving island arc-continental arc systems such as the Late Miocene Talamanca intrusive suite (Drummond and others, 1995) which, like the Killingworth gneisses, also displays modest Ta enrichment (fig. 13B). Four of the six data points for the Hidden Lake gneiss plot at lower Rb values than the moderately enriched Boulder Lake and Pond Meadow gneisses (fig. 13A) and thus were likely derived from compositionally distinct sources.

Petrologic affinity of the Killingworth rocks to arc systems is supported by moderate decoupling of large-ion lithophile (LIL) and high-field strength (HFS) elemental concentrations (Winter, 2001). Especially noteworthy is enrichment, relative to mid-ocean ridge basalt (MORB), in LIL elements K, Rb, and Ba (as well as Th), slight depletion in Ta and Nb, and MORB-like abundances of HFS elements Ce through Yb, except Ti and Y (fig. 13C). Normalized elemental concentrations of samples from the Middletown Formation and Higganum gneiss are lower than values for gneisses of the Killingworth complex, and are closer to MORB values. As a result, we suggest that the observed differences reflect petrologic processes and are not the result of alteration. Overall, the Hidden Lake gneiss is similar to some of the other gneisses but is distinguished compositionally by extreme values that fall at either the high or low end of the Killingworth range (fig. 13C). Arc affinity for the Killingworth complex rocks is further indicated by the similar pattern that is characteristic of the calc-alkaline Miocene intrusives from Talamanca (fig. 13C).

The units of the Killingworth and Middletown complexes also can be distinguished using a combination of Nd and Pb isotopic data (fig. 14). This type of plot

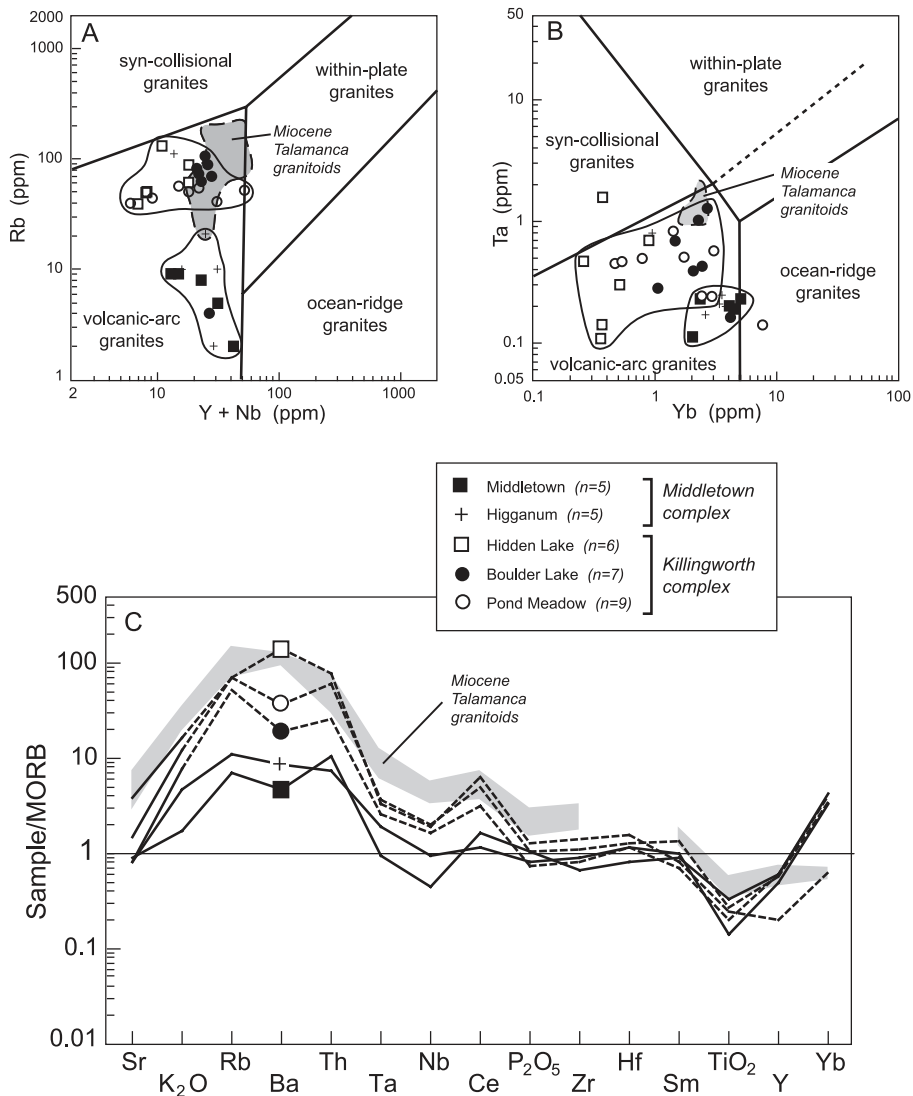


Fig. 13. Geochemical plots of whole-rock data from the Killingworth dome. (A) (Y + Nb) vs Rb. Compositional field of Late Miocene granitoids from the Cordillera de Talamanca of Costa Rica and Panama (29 samples from Drummond and others, 1995) plotted for comparison. Diagram after Pearce and others (1984). (B) Yb vs Ta. Compositional field of Late Miocene granitoids from the Cordillera de Talamanca of Costa Rica and Panama is plotted for comparison (5 samples from Drummond and others, 1995). Diagram after Pearce and others (1984). (C) Plot of average compositions of gneisses from the Killingworth complex and both felsic and mafic rocks of the Middletown Formation normalized to mid-ocean ridge basalt (MORB) using values from Bevs and others (1984) except for Sm which is from Pearce (1983). Compositional field of Late Miocene granitoids from the Cordillera de Talamanca of Costa Rica and Panama is plotted for comparison (5 samples from Drummond and others, 1995).

(initial ϵ_{Nd} versus $^{207}Pb/^{204}Pb$) is valuable because it uses two source-sensitive discriminants (Tomascak and others, 2005). Higganum gneiss and Middletown Formation (including both low and high SiO₂ lithologies, table 1) have relatively high initial ϵ_{Nd} (range of about +1 to +6), whereas Hidden Lake, Boulder Lake, and Pond Meadow gneisses have lower values of about -9 to -1. This difference in ϵ_{Nd} supports the Pb

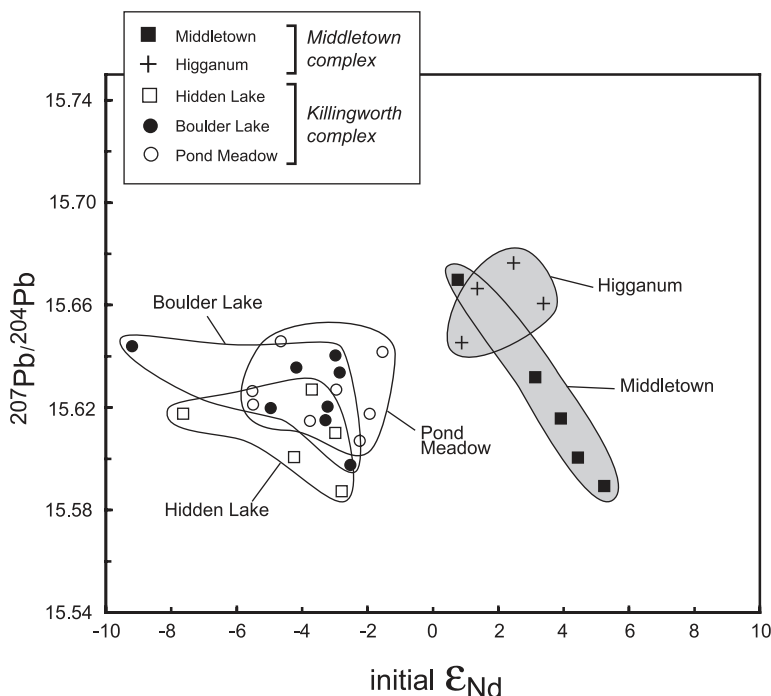


Fig. 14. Initial ϵ_{Nd} versus $^{207}\text{Pb}/^{204}\text{Pb}$ of whole-rock samples from the Killingworth and Middletown complexes.

isotopic data and several geochemical discriminants that suggest that the Higganum and Middletown magmas were derived from sources that differed from those of the Killingworth complex. In addition, the higher ϵ_{Nd} values for the Higganum and Middletown imply a more primitive source component, whereas the negative ϵ_{Nd} values for the Hidden Lake, Boulder Lake, and Pond Meadow gneisses indicate an older, more evolved crustal source.

To evaluate potential sources for the units of the Killingworth dome, we compare data from nearby samples of Neoproterozoic basement rocks, including samples of felsic orthogneiss from the Clinton dome in the Essex quadrangle (samples 15H, 15J, 5BS; table 3, fig. 3) and Lyme dome in the Old Lyme quadrangle (OL-290, and OL-293; table 3), and two samples of felsic gneisses interpreted to be of extrusive and intrusive origin, respectively (Wintsch and others, 1990), from the Selden Neck block (samples 15W and 15X; table 3, fig. 3). Pb isotopic data from samples from the Clinton and Lyme domes plot above the average growth curve (Stacey and Kramers, 1975) whereas data from samples from the Selden Neck block are distinctly less radiogenic, plotting below the average growth curve (fig. 15A). Four of five analyses from Clinton and Lyme dome rocks overlap Pb isotopic data for Gander terrane basement gneisses rocks in New Brunswick (fig. 15A; Ayuso and Bevier, 1991).¹ Samples 15W and 15X

¹This zone or belt was called “Southern Group” by Ayuso and Bevier (1991). It was originally interpreted as “Avalonian” in origin, in recognition of its very different isotopic composition compared to Laurentian (or Grenville) rocks to the west. These rocks can be traced northeast to Cape Breton Island, New Brunswick, and Newfoundland. The belt has been given a number of names, including Bras D’Or, Miramichi, Western Avalonia, Eastern Central Mobile Belt, and many local terrane names. Confusion arises when consulting older literature that refers to these rocks as Avalonian in origin. It is now known that in Newfoundland, the

have similar Pb isotopic ratios to Silurian-Devonian granitic rocks interpreted as having been derived from Neoproterozoic basement of the Avalon terrane in New Brunswick and Newfoundland. Unfortunately, there is no comparable data set for Neoproterozoic basement rocks of the Avalon terrane. The $^{207}\text{Pb}/^{204}\text{Pb}$ ratios of Paleozoic granitoids of Avalon (and samples 15W and 15X) are lower than in Paleozoic granitic rocks in the Gander terrane (Kerr, 1997).

Potential source rocks for the magmas of the Killingworth and Middletown complexes can also be evaluated on the ϵ_{Nd} versus $^{207}\text{Pb}/^{204}\text{Pb}$ plot (fig. 15B). Whole-rock data from samples of Gander terrane in southern Connecticut have relatively low ϵ_{Nd} values and relatively high $^{207}\text{Pb}/^{204}\text{Pb}$. Data from 4 (of 5) samples overlap the field of Gander basement (data from Ayuso and Bevier, 1991; Kerr and others, 1995), whereas samples 15W and 15X from the Selden Neck block have higher ϵ_{Nd} values and lower $^{207}\text{Pb}/^{204}\text{Pb}$ reflecting more primitive sources. These data confirm that gneisses of the Clinton and Lyme domes belong to the basement of the Gander terrane, and that samples 15W and 15X from the Selden Neck block belong to the Avalon terrane. The combined field of data for Boulder Lake, Pond Meadow, and Hidden Lake gneisses plots between fields of data for Gander basement and Laurentian (Grenville) crust and above isotopic data for rocks of the Bronson Hill terrane of northern New Hampshire (Lathrop and others, 1996; Moench and Aleinikoff, 2003). Higganum gneiss and Middletown Formation plot to the right of and slightly below the data field for Gander crust, on a trajectory toward model mantle at 450 Ma (fig. 15B).

We suggest that all units of the Killingworth dome have mixed sources. The two volcanic magmas (Higganum and Middletown of the Middletown complex) are the result of mixing of Gander and primitive (that is, mantle) components. The three plutonic units (Boulder Lake, Pond Meadow, and Hidden Lake of the Killingworth complex) probably were derived from a mixture of Neoproterozoic Gander rocks and a less radiogenic component. Laurentian (Grenvillian) rocks have lower $^{207}\text{Pb}/^{204}\text{Pb}$ and ϵ_{Nd} values, and now occur relatively close to the Killingworth dome. However, consideration of paleogeographic location indicates that the Gander terrane originated far from Laurentia, across the Iapetus ocean. We discuss other possible sources below in the context of Paleozoic tectonic reconstructions.

Provenance of Detrital Zircons

Isotopic evidence suggests that the Killingworth complex was derived from a mixture of Gander basement and less radiogenic material that has isotopic characteristics similar to Laurentian (Grenville) rocks. However, it is now known (on the basis of paleomagnetic data and identification of faunal realms; Harper and others, 1996; Mac Niocaill and others, 1997) that Gander formed as a peri-Gondwanan terrane. Tectonic models (van Staal and others, 2004; Valverde-Vaqueiro and others, 2006), suggest that Gander accreted to the Laurentian margin in the Silurian (~430 Ma) during the Salinic orogeny, implying that Gander would not have received Laurentian sediments until late in its accretionary history, well after creation of the Killingworth complex. In southern New England, accretion of Gander to Laurentia may have occurred somewhat earlier, perhaps beginning in the Late Ordovician (but several m.y. after emplacement of the Killingworth complex magmas), on the basis of isotopic compositions of plutonic rocks in western Connecticut (Sevigny and Hanson, 1993, 1995).

Dover fault separates Avalon *sensu stricto* (to the southeast) and Gander (to the northwest) terranes, both of which contain Neoproterozoic crystalline basement rocks (van Staal and others, 1996). This discontinuity has been traced southwest (compare, Whalen and others, 1996; Pe-Piper and Piper, 1998). Regardless of original name, the term Gander will be used throughout this paper to refer to Neoproterozoic rocks (and derivatives thereof) that have high $^{207}\text{Pb}/^{204}\text{Pb}$ and negative initial ϵ_{Nd} values. The term Avalon (Barr and White, 1996; Barr and others, 2002) will be restricted to Neoproterozoic rocks (and derivatives thereof) with low $^{207}\text{Pb}/^{204}\text{Pb}$ and positive ϵ_{Nd} values.

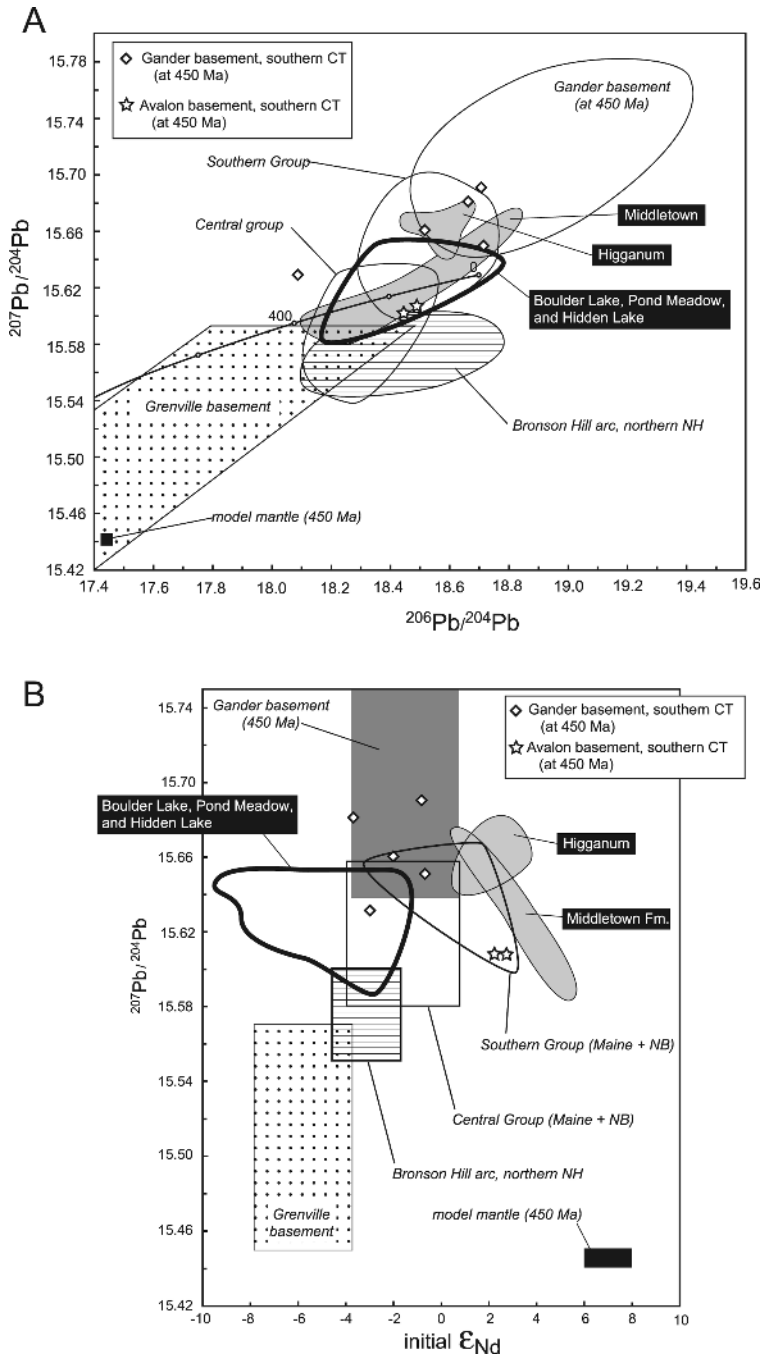


Fig. 15. Pb and Nd isotopic data from rocks of the Killingworth dome and nearby Neoproterozoic rocks. (A) Fields of initial Pb isotopic compositions of whole-rock samples of the Killingworth and Middletown complexes, plus Pb isotopic data (corrected to 450 Ma) from whole-rock samples of Neoproterozoic basement rocks of southern Connecticut. Also shown are fields of potential sources: Gander basement, corrected to 450 Ma, from New Brunswick (Ayuso and Bevier, 1991), Grenville basement (compiled in Ayuso and Bevier, 1991), granitic rocks from the northern portion of the Middle Ordovician Bronson Hill arc, NH (Moench and Aleinikoff, 2003), fields of initial ratios of Paleozoic plutons of Maine and New Brunswick (Central and Southern Groups, Ayuso and Bevier, 1991), and model mantle (Zartman and Haines, 1998).

Mesoproterozoic material is common in the Gander terrane. For example, quartzite within the Neoproterozoic metasedimentary Old Lyme gneiss of the Lyme dome (Walsh and others, 2007) contains a large population of Mesoproterozoic (about 0.9 – 2.1 Ga) detrital zircons. These ages are similar to inheritance ages in igneous zircons from the Neoproterozoic Dry Hill Gneiss, Massachusetts (Tucker and Robinson, 1990), and ages of detrital zircons from Neoproterozoic Westboro Formation of the Boston basin (Thompson and Bowring, 2000) and in our sample of anthophyllite gneiss (figs. 9A and 9B) of the Middletown Formation. Mesoproterozoic ages of detrital zircons have also been found in sedimentary rocks of Avalon and Gander terranes in New Brunswick and Nova Scotia (Barr and others, 2003; Reusch and others, 2004). Possible non-Laurentian, peri-Gondwanan sources for Mesoproterozoic zircons include Amazonia, Oaxaquia, and Baltica (Ortega-Gutierrez and others, 1995; Thompson and Bowring, 2000, and references therein). On the basis of a number of criteria (including detrital zircon age populations, fauna identification, and paleomagnetic data), Nance and Murphy (1996) and van Staal and others (1996) conclude that the most likely provenance for Mesoproterozoic material in Gander terrane rocks is Amazonia. Pb isotopic compositions from Amazonian rocks (Tohver and others, 2004; Loewy and others, 2004, and references therein) are consistent with this interpretation.

Metamorphism

One intriguing aspect of this study is the apparent lack of evidence for the Devonian Acadian orogeny in rocks of the Killingworth dome. In contrast to rocks of central and northern New England where geochronologic evidence for the Devonian Acadian orogeny is well documented (Eusden and Barreiro, 1988; Spear and Harrison, 1989; Tucker and Robinson, 1990), we find no evidence of Devonian metamorphism in the Killingworth dome. Geochronologic data from zircon and titanite from rocks of this study show that overgrowths on Ordovician magmatic zircons are all younger than Devonian, and formed nearly continuously between about 340 and 270 Ma. Titanite in anthophyllite gneiss of the Middletown Formation formed at about 300 and 250 Ma. These data are consistent with other U-Pb and argon isotopic results and modeling suggesting that while tectonic loading began in the Late Mississippian, peak metamorphic temperatures were not achieved until the Middle Permian, and were quenched to $< \sim 500^\circ\text{C}$ by early Triassic (Wintsch and others, 2003). A similar lack of evidence for Acadian overprinting was noted by Aleinikoff and others (2002) in the Glastonbury gneiss nearby to the north (fig. 2), and in Ordovician plutonic rocks of western Connecticut (Sevigny and Hanson, 1995). In contrast, Wintsch and others (2007) show that Acadian metamorphism is recorded in Gander cover terranes just east of the Killingworth dome in southern Connecticut. For example, the Canterbury and Eastford granitic gneisses were emplaced into metasedimentary rocks of the Merrimack terrane at about 414 ± 3 and 379 ± 4 Ma, respectively, and metasedimentary rocks of the Tatnic Hill Formation (Putnam-Nashoba terrane) were partially melted at ~ 400 Ma.

A detailed explanation for the lack of evidence for the Acadian remains elusive. The absence of Devonian metamorphic zircon and titanite suggests that the grade of metamorphism at that time was too low to form new mineral overgrowths. This is more

Fig. 15 (continued). (B) Fields of initial ϵ_{Nd} versus $^{207}\text{Pb}/^{204}\text{Pb}$ of whole-rock samples from the Killingworth and Middletown complexes, plus isotopic data from whole-rock samples of Neoproterozoic basement rocks of southern Connecticut. Nd data for fields of Central and Southern Groups from R. A. Ayuso (personal communication, 2006), Gander basement from Kerr and others (1995), Bronson Hill granitic rocks from Lathrop and others (1996), model mantle from DePaolo (1981). Pb data from references in figure 14A.

likely than an alternative explanation that such overgrowths were obliterated by high-grade Alleghanian metamorphism. We consider this latter scenario to be unlikely because metamorphic zircon generally contains lower concentrations of minor and trace elements than magmatic zircon (Hoskin and Schaltegger, 2003). As a result, metamorphic zircon should be even less metamict and thus less soluble or reactive than magmatic zircon during subsequent high-grade events. Nevertheless, it is possible that these rocks escaped high-grade Acadian metamorphism because they were sheltered from Acadian deformation, perhaps sequestered in a tectonic reentrant (for example, Lin and others, 1994). We speculate that the rocks of the Killingworth dome and Glastonbury gneiss were structurally high, and not buried to mid-crustal depths until the early Alleghanian orogeny in the Late Carboniferous.

Nature of the Boulder Lake gneiss-Clinton Granite Gneiss Contact

The contact of the Ordovician Boulder Lake gneiss and Neoproterozoic Clinton Granite Gneiss is a ductile fault (Wintsch and others, 2005) (fig. 3). Because the rocks of the Killingworth and Clinton domes were previously considered to belong to the Bronson Hill and Avalon terranes, respectively, this contact was interpreted as a terrane boundary. However, isotopic data presented herein result in a new interpretation suggesting that rocks of the Killingworth complex and Clinton dome both are parts of the Gander terrane. Thus, this fault separates two igneous complexes of Neoproterozoic and Late Ordovician ages. Anatectic foliation in the Boulder Lake gneiss probably formed at 312 ± 4 Ma (age of overgrowths, fig. 8D). These high grade metamorphic conditions may have persisted into the Early Permian on the basis of ages of overgrowths (as young as about 300 Ma) from related rocks of the Killingworth complex (table 1). Because high grade ($> \sim 600^\circ$ C.) foliation associated with the fault cross-cuts the anatectic foliation in the Boulder Lake gneiss (Wintsch and others, 2005), the time of juxtaposition of the Clinton Granite Gneiss and the Boulder Lake gneiss is therefore constrained to be younger than ~ 300 Ma. A minimum age of faulting is provided by cooling ages of ~ 260 Ma in amphiboles from the Boulder Lake gneiss (Wintsch and others, 2003), at which time the rocks were being uplifted and had cooled to $\sim 500^\circ$ C.

Tectonic Speculations

Regional correlations.—Ayuso and Bevier (1991) used Pb isotopic compositions of Paleozoic felsic plutons to identify three lower crustal provinces in the northern Appalachians. Their ‘northern’, ‘southern,’ and ‘central’ groups we now recognize to relate to Laurentian, Ganderian, and Mesoproterozoic+Ganderian mixtures, respectively. The Pb isotopic compositions and ϵ_{Nd} values of the Killingworth complex are similar to isotopic ratios in plutonic rocks of their “Central Group” (figs. 15A and 15B). Isotopic evidence suggests that these rocks are mixtures from Neoproterozoic Gander sources and a less radiogenic component, now thought to be from Mesoproterozoic detritus of peri-Gondwanan origin. Rocks of similar isotopic characteristics extend far to the northeast into New Brunswick (Whalen and others, 1996, 1998; Samson and others, 2000), Nova Scotia (Pe-Piper and Piper, 1998), Cape Breton Island (Barr and Hegner, 1992), and Newfoundland (Kerr and others, 1995; Whalen and others, 1997). All of these magmas are considered to have been derived from peri-Gondwanan (that is, non-Laurentian) crust (compare, O’Brien and others, 1996; Barr and White, 1996), containing Neoproterozoic basement and Paleozoic sedimentary cover and intrusive rocks.

Tonalitic gneisses of similar Ordovician age as rocks of the Killingworth and Middletown complexes occur to the north in central Massachusetts (Tucker and Robinson, 1990). Like the Killingworth rocks, their low ϵ_{Nd} values (Four Mile gneiss, Bull and Robinson, 1994) lead us to conclude that these Massachusetts magmas were

derived from Ganderian crust. Rocks of the Middletown complex are distinct from the Killingworth complex in their geochemical and isotopic signatures. However, data from R. A. Ayuso (personal communication, 2006) indicate a strong similarity between the Middletown complex and rocks of the “Southern group” of Ayuso and Bevier (1991; figs. 15A and 15B). The isotopic data suggest that the source of these Middletown rocks is a mixture of Gander basement (relatively low initial ϵ_{Nd} and relatively high $^{207}\text{Pb}/^{204}\text{Pb}$) plus a primitive (mantle) component (high initial ϵ_{Nd} and low $^{207}\text{Pb}/^{204}\text{Pb}$). A model of multiple sources is consistent with interpretations of the origins of Paleozoic granitic rocks of the Gander zone in New Brunswick and Newfoundland, considered to be derived from Neoproterozoic basement and various amounts of mantle and metasedimentary (crustal?) components (Whalen and others, 1996, 1997; Kerr, 1997).

Rocks of the Killingworth dome have previously been included within the Bronson Hill terrane (Williams, 1978; Hibbard and others, 2006), primarily due to similarities in age and structure with Oliverian domes, and apparent continuity of the belt throughout New England. Interpretations of the tectonic origin of the Bronson Hill terrane, however, remain contentious, including differing opinions about subduction zone polarity and paleogeographic location during arc formation. The arc rocks have been subdivided into two age groups: ~ 470 to 485 Ma (Shelburne Falls arc), and ~ 440 to 455 Ma (Bronson Hill arc) (Karabinos and others, 1998) with differing tectonic origins. Rocks of the Killingworth dome belong to the younger age group, similar to tonalitic rocks on strike to the north in Massachusetts (Tucker and Robinson, 1990). In the most recent regional compilation, Hibbard and others (2006), in agreement with many earlier studies (Osberg, 1978; Lyons and others, 1982; Tucker and Robinson, 1990) conclude that the Bronson Hill terrane (from Connecticut to at least northern-New Hampshire, and possibly as far northeast as New Brunswick and Newfoundland) formed as a peri-Gondwanan arc system built on Gander basement. In all these tectonic models, rocks of the Killingworth dome (including both the Killingworth and Middletown complexes) are considered part of the Bronson Hill arc, regardless of the crust (Laurentia, Gander, Avalon, or intervening Iapetan oceanic crust) upon which the arc formed.

Redefining of the Bronson Hill terrane.—An evaluation of our isotopic data and previously published data call into question the grouping of all “Bronson Hill” rocks into a single lithotectonic terrane. Middle to Late Ordovician and Devonian plutons of the Bronson Hill terrane in northern New Hampshire (fig. 16) have Pb isotopic compositions (Moench and Aleinikoff, 2003; Tomascak and others, 2005) that plot below the average growth curve (Stacey and Kramers, 1975), similar to Laurentian (Grenville) basement rocks (fig. 15A). We conclude that these magmas were derived from Laurentian crust. Lead isotope compositions of para-autochthonous mid-Ordovician intrusive rocks in western Connecticut, just east of Cameron’s Line (for example, Brookfield and Beardsley plutons; Sevigny and Hanson, 1993, 1995), suggest that they are also of Laurentian origin. We consider them to be the southern extension of the Laurentia-derived Bronson Hill magmatic arc identified in northern New Hampshire (fig. 16).

In contrast, isotopic ratios of rocks of the Killingworth complex are more radiogenic than those derived from Laurentia. Thus, along the “Bronson Hill terrane” between northern New Hampshire and southern Connecticut there appears to be a fundamental boundary separating Laurentian and Ganderian lower crustal blocks. This boundary, called the Red Indian line in Newfoundland, was originally delineated on the basis of contrasting faunal assemblages (Williams and others, 1988). It has been traced west-southwest through mainland Canada (van Staal, 1994) and into Maine (van Staal and others, 1998; Dorais and Paige, 2000). In New Hampshire, Moench and

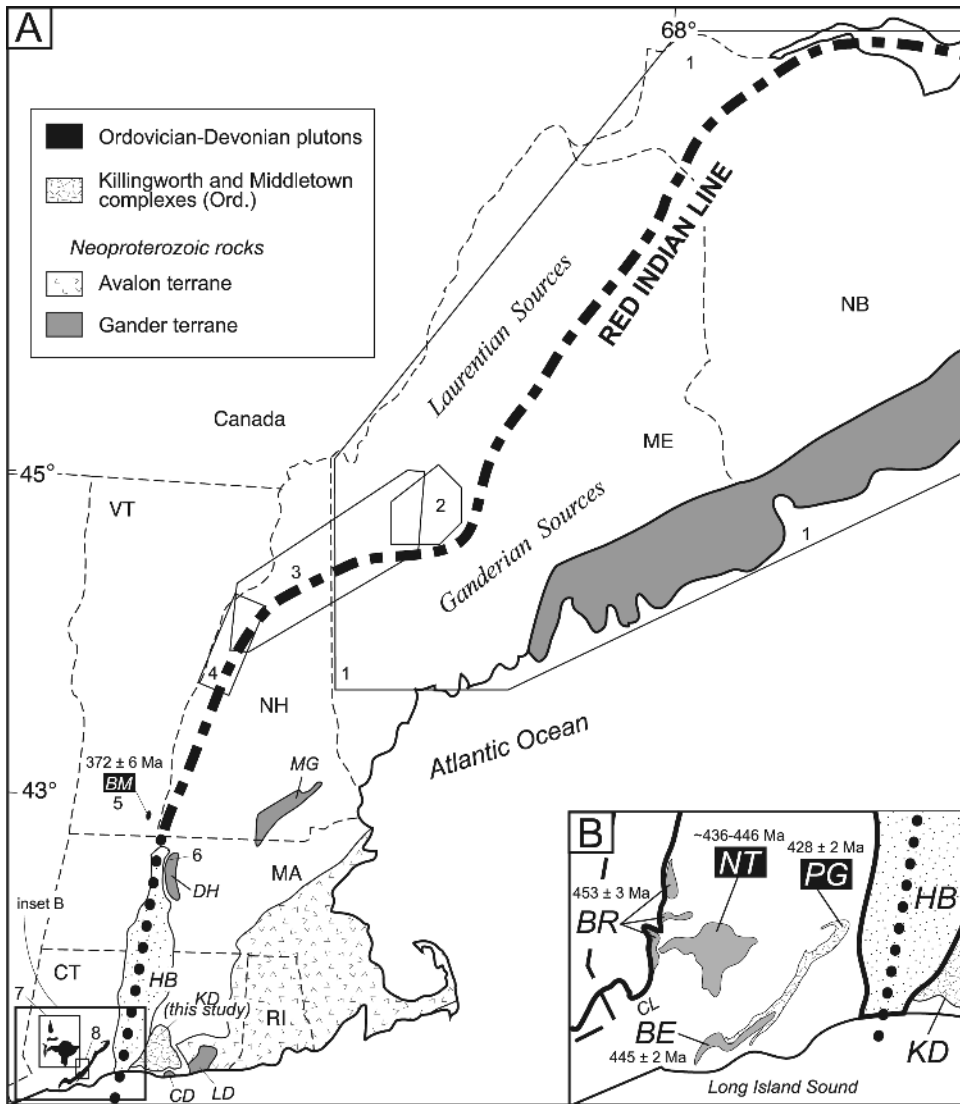


Fig. 16. New interpretation of the location of the Ordovician Iapetan suture (Red Indian line (RIL), Williams and others, 1988) between rocks of Laurentian and Ganderian affinity in southern New England. (A) Simplified map of New England and part of New Brunswick showing distribution of Gander and Avalon terranes (modified from Hibbard and others, 2006). RIL in Maine is modified from Moench and Aleinikoff (2003) to accommodate isotopic data from Tomascak and others (2005). RIL is dotted where speculative in southern New England. In this model, the Laurentia-derived Bronson Hill terrane of northern and western New Hampshire is interpreted to occur in western Connecticut. Rocks of the Killingworth dome (Killingworth and Middletown complexes) are interpreted here to be of Ganderian origin. White-lettered labels with black background in A and B are plutons with Pb isotopic values indicative of Ganderian source components that occur west of the RIL (see text for additional explanation). Abbreviations: BM (Black Mountain pluton), CD (Clinton dome), DH (Dry Hill Gneiss of the Pelham dome), HB (Hartford basin), KD (Killingworth dome), LD (Lyme dome), MG (Massabesic Gneiss). Polygons show locations of previous studies: 1—Ayuso and Bevier (1991), 2—Tomascak and others (2005), 3—Moench and Aleinikoff (2003), 4—Lathrop and others (1996), 5—J. N. Aleinikoff (unpublished data, 1990), 6—Bull and Robinson (1994), 7—Sevigny and Hanson (1995), 8—Sevigny and Hanson (1993). (B) Inset showing plutons of western Connecticut dated and analyzed for initial Pb isotopic compositions by Sevigny and Hanson (1993, 1995). Abbreviations: CL (Cameron's Line), BE (Beardsley orthogneiss), BR (plutons of the Brookville Plutonic Series), NT (Newtown gneiss), PG (Pumpkin Ground orthogneiss).

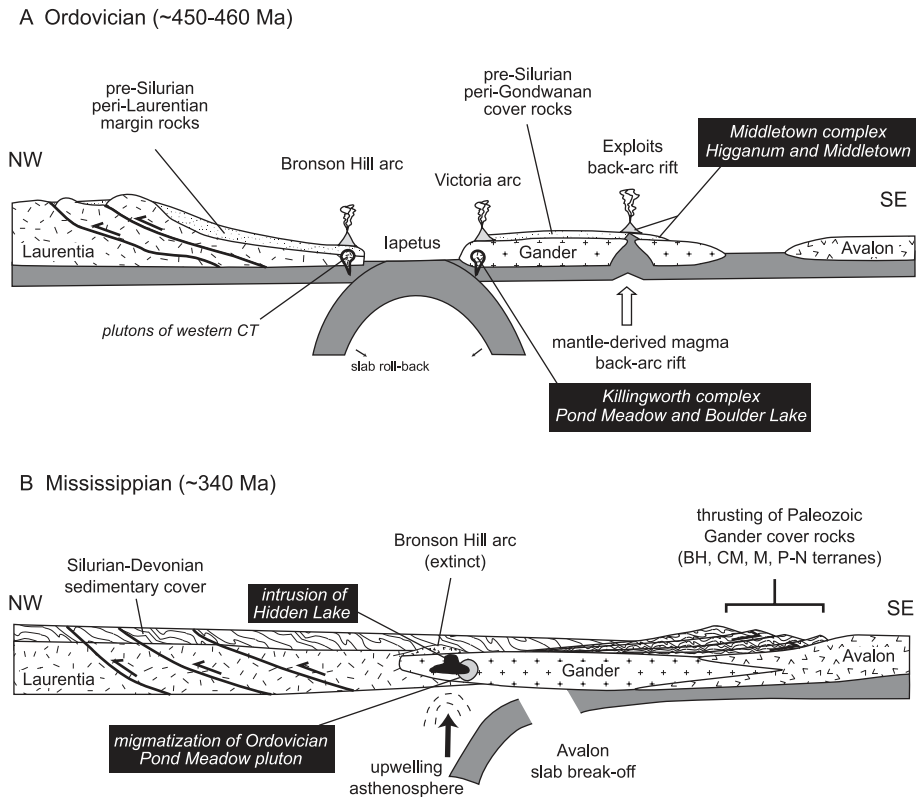


Fig. 17. Schematic cross-sections showing the tectonic evolution of the Killingworth and Middletown complexes. (A) Middle to Late Ordovician. This model assumes the interpretation that Gander accreted to Laurentia in the Silurian (Valverde-Vaquero and others, 2006). (B) Middle Carboniferous. Terrane abbreviations: BH (Bronson Hill), CM (Central Maine), M (Merrimack), P-N (Putnam-Nashoba).

Aleinikoff (2003) present two possible trajectories of the Red Indian line extending southwest into northern Massachusetts, to the east and west of the Neoproterozoic Dry Hill Gneiss of the Pelham dome. Our new interpretation suggests that the Red Indian line traces to the west, on a trend toward western Connecticut.

We conclude that the Ordovician Laurentia-Gander boundary (the Red Indian line) lies approximately under the Hartford basin between the Beardsley gneiss to the west, and Killingworth complex rocks to the east (fig. 16). If this interpretation is correct, then the southern portion of the "Bronson Hill terrane" of figure 1 requires reassignment; it would be correlative with other Ganderian margin arcs, such as the Victoria arc of Newfoundland (Valverde-Vaquero and others, 2006). In the post-Middle Ordovician, Gander wedged farther westward into Laurentia, as shown by the increasing $^{207}\text{Pb}/^{204}\text{Pb}$ ratios (that is, increasing Gander component) of Late Ordovician to Devonian plutons (Newtown and Pumpkin Ground gneisses of western Connecticut, Sevigny and Hanson, 1993, 1995; Black Mountain pluton of southeastern Vermont, J. N. Aleinikoff, unpublished data) located west of the Red Indian line (fig. 16).

Ordovician tectonic settings of the Killingworth and Middletown complexes.—We conclude that the Killingworth complex was formed in a magmatic arc environment on, or adjacent to, the western margin of the Gander terrane above an east-dipping subduction zone on the east side of Iapetus (fig. 17A). Contemporaneously, intermediate to

felsic, and mafic magmas of the Middletown complex were generated in a different tectonic setting according to trace-element and isotopic data. $^{207}\text{Pb}/^{204}\text{Pb}$ and ϵ_{Nd} isotopic values suggest that the Middletown complex is the result of mixing of Gander basement with a less radiogenic, more primitive component such as oceanic crust or mantle, without a significant older crustal component (fig. 15B). These data are consistent with an origin for the Middletown complex in a back-arc setting (fig. 17A). There is insufficient evidence at present to determine modes of origin, such as fractional derivation from more primitive magmas or partial melting of earlier underplated mantle-derived material, for these largely felsic rocks.

The proposed tectonic setting for the Killingworth and Middletown complexes is nearly identical to the model of Valverde-Vaquero and others (2006) for rocks of the Victoria arc and Exploits subzone within the Gander zone of Newfoundland. Correlative units have been traced to the British Isles and maritime Canada (Valverde-Vaquero and others, 2006, and references therein). We suggest that these zones can now be extended to southern New England. At about the same time (in the Middle to Late Ordovician), arc rocks of the Bronson Hill terrane were being formed above a west-dipping subduction zone on or adjacent to Laurentia (Sevigny and Hanson, 1995; Karabinos and others, 1998; Kim and Jacobi, 2002)

Late Paleozoic tectonic setting.—In the Mississippian, the Hidden Lake pluton was emplaced and the Pond Meadow gneiss was partially melted. We speculate that this magmatic event was brought about by renewed collision of the Avalon terrane (fig. 17B) during the assembly of Pangea. This shortening event may have led to delamination and subsequent slab break-off of Avalonian lithospheric mantle under the Gander terrane farther to the west, resulting in asthenospheric upwelling that heated the lower crust, and caused remelting of the Killingworth complex. Initially, this heat melted the root of the former arc, producing the Hidden Lake pluton. Continued heating led to partial melting of Ordovician arc rocks, as exemplified by local migmatization in the Pond Meadow gneiss. In addition, $^{40}\text{Ar}/^{39}\text{Ar}$ cooling ages of ~ 340 Ma from amphiboles of the Putnam-Nashoba terrane record the early stages of exhumation of these rocks over the Avalon terrane in the east (Wintsch and others, 1992, 1993). The Mississippian event at about 340 Ma is somewhat cryptic because it is ~ 40 m.y. earlier than final assembly of Avalon with amalgamated Laurentia+Gander in the Pennsylvanian (Wintsch and others, 2003, 2005; Walsh and others, 2007). This latter event is the beginning of the Alleghanian orogeny in this area and is recorded by titanite (~ 300 Ma) from the Middletown complex (fig. 9E); it may also represent the time of thrust emplacement of the Middletown over the Killingworth complex, which must have occurred after migmatization of the Pond Meadow gneiss and emplacement of the Hidden Lake gneiss (~ 340 Ma).

CONCLUSIONS

The Killingworth dome of southern Connecticut exposes Middle to Late Ordovician (about 455–460 Ma) and Mississippian (about 340 Ma) tonalitic to trondhjemitic orthogneisses, here named the “Killingworth complex.” These rocks are flanked by the Middle to Late Ordovician (about 450–460 Ma) “Middletown complex”, which includes the Higganum gneiss (formerly considered part of the Killingworth complex) and Middletown and Collins Hill Formations.

Geochemical and isotopic data suggest that rocks of the Killingworth complex originated in a volcanic-arc tectonic setting and were derived primarily from Gander terrane sources that likely included Neoproterozoic basement of Gander and older crustal material. Published paleomagnetic and paleontologic data indicate that this Mesoproterozoic material could not have been derived from Laurentia, but was originally derived from a peri-Gondwanan terrane such as Amazonia. Rocks of the Middletown complex were likely derived from a mixture of Gander basement and a more primitive source, probably in a back-arc setting. Although on strike with similar

rocks of the Bronson Hill terrane in northern New England, the Killingworth and Middletown complexes appear to have originated on or adjacent to Gander terrane. Rocks of the Bronson Hill terrane occur in western Connecticut.

Final juxtaposition of the Killingworth and Middletown complexes occurred after Mississippian migmatization of the Pond Meadow gneiss and crystallization of the Hidden Lake gneiss, during the Alleghanian orogeny in the late Pennsylvanian or early Permian. We have found no evidence for the Acadian orogeny in the Killingworth or Middletown complexes, suggesting that these rocks remained relatively shallow, or remote from the Laurentian margin, until the Alleghanian.

ACKNOWLEDGMENTS

This work was initiated and funded by Tom Armstrong and Bill Burton of the USGS Bedrock Regional Aquifer Systematics Study (BRASS) project and by NSF grant EAR-0510293 to R. Wintsch and M. Dorais. We thank Joe Taggart, Jim Budahn, and Brian Marshall for rock analyses, Ezra Yacob for mineral separations, and Joe Wooden and Frank Mazdab for help with SHRIMP analyses. Conversations with Robert Ayuso and Bob Scott proved illuminating. We are grateful to Janet and Byron Stone for logistical support and hospitality in the field. Conversations with John Rodgers, Larry Lundgren, and Dick Goldsmith inspired our early interest and work in this area. We greatly appreciate thoughtful technical reviews by Robert Ayuso, Paul Karabinos, Michel Malo, and Greg Walsh.

APPENDIX

ANALYTICAL METHODS

Whole-rock Analyses

Major- and trace-element concentrations were determined by X-ray fluorescence (XRF) and instrumental neutron activation analysis of whole-rock powders in the analytical facilities of the U.S. Geological Survey (USGS), Denver. Elemental oxide abundances of SiO₂, TiO₂, Al₂O₃, total Fe as Fe₂O₃, MgO, CaO, Na₂O, K₂O, P₂O₅ and MnO were determined by wavelength-dispersive x-ray fluorescence (WD-XRF). To obtain loss on ignition (LOI), a 0.8 g portion of each sample was ignited at 925° C for 45 minutes and re-weighed to measure loss. Each sample was then fused with lithium borate flux to create a homogeneous pellet for analysis.

The precision of the major- and minor-elemental determinations is typically less than or equal to 1.0 relative percent, based on replicate analysis of prepared standards, whereas accuracy is less than or equal to 2 percent. A full description of the WD-XRF technique is given in Taggart and Siems (2002). X-ray fluorescence analysis of Rb, Sr, Y, Zr, Nb, Ba, Ce, and La was performed using an energy-dispersive detector. These concentrations are precise to ±10 percent. Accuracy of all geochemical data was assured by calibration to a suite of internationally accepted standards.

Abundances of thirty-three major, minor, and trace elements, including 11 rare-earth elements (REE), were determined by instrumental neutron activation analysis (INAA). Samples of about one gram were irradiated in the USGS-TRIGA reactor at a flux of 2.5×10^{12} for eight hours. Three sequential counts at 7 days, 14 days, and 65 days after the irradiation were made on both coaxial and planar germanium detectors. A summary of the INAA procedure used at the USGS is given in Budahn and Wandless (2002). Precision and accuracy for most of the elements determined range from 1 to 5 percent, including La, Yb, Hf, Ta, Rb, Th, and U. Only five elements have precision errors of greater than 10 percent (Ho, Tm, W, Sb, and Au) based on counting statistic errors. Accuracy is based on replicate analysis of USGS standard reference materials, including BHVO-1.

SHRIMP U-Pb Geochronology of Zircon and Titanite

U-Pb geochronology of zircon was performed using the U.S. Geological Survey/Stanford University sensitive high resolution ion microprobe-reverse geometry (SHRIMP-RG). Titanite from one sample was dated using the SHRIMP II at the Research School of Earth Sciences, Australian National University. About 10 kg of rock were collected for each sample dated. Zircon (and titanite, where present) was extracted using standard mineral separations techniques, including crushing, pulverizing, Wilfley table, magnetic separator,

and heavy liquids. Individual grains were hand picked, mounted in epoxy, ground to half-thickness to expose internal zones, and polished using 6 μm and 1 μm diamond suspension. All grains were imaged digitally in transmitted and reflected light. Using the scanning electron microscope, zircons were imaged in cathodoluminescence (CL), whereas titanites were imaged by back-scattered electrons (BSE).

Following the methods of Williams (1998), SHRIMP analysis consisted of excavating a pit about 25 to 35 μm in diameter and about 1 μm in depth, using a primary oxygen beam at a current of about 4 nA. The magnet cycled through the mass stations 6 times per analysis. Raw data were reduced using Squid 1 (Ludwig, 2001) and plotted using Isoplot 3 (Ludwig, 2003). Instrument fractionation for $^{206}\text{Pb}/^{238}\text{U}$ ages was corrected using zircon standard R33 (419 ± 1 Ma; Black and others, 2004) or titanite standard BLR-1 (1047.1 ± 0.4 Ma; see below). Uranium concentrations are believed to be accurate to ± 20 percent. U-Pb data are plotted on Tera-Wasserburg concordia plots to visually identify coherent ages groups. Weighted averages of individual $^{206}\text{Pb}/^{238}\text{U}$ ages were calculated to obtain an age for each sample. For titanite, lower intercept concordia ages were calculated through data that were uncorrected for common Pb, using the Isoplot 3-D linear regression (Ludwig, 2003).

TIMS U-Pb Geochronology of Titanite

Titanite standard BLR-1 was dated by thermal ionization mass spectrometry (TIMS). A number of translucent brown fragments visually free of alteration were picked and half of the fragments abraded for 3 hours to smooth ellipsoids in an alumina-ceramic vortical abrader. Six individual fragments (three abraded, three unabraded) were ultrasonicated in double-distilled acetone, rinsed in high-purity H_2O , spiked and dissolved at 220°C for 96 hours in 29M HF and 7M HNO_3 ; samples were subsequently dried, redissolved in 6M HCl at 180°C for 24 hours and then converted to 1.1M HBr for ion exchange chemistry. Details of mass spectrometry and data reduction are stated in the footnotes to table A1.

Five of six fragments yielded essentially equivalent isotope ratios defining a precise inverse variance weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 1047.1 ± 0.4 Ma (2σ , MSWD 0.56) (fig. A1), including tracer calibration uncertainty (Schoene and others, 2006). A sixth fragment exhibits lower isotope ratios and younger ages; in retrospect, this unabraded fragment did have a slightly hackly texture on one surface, which might indicate that it was originally adjacent to a fracture, and subject to a very minor amount of surface-correlated Pb loss. For the cluster of five equivalent fragments, the slightly elevated MSWD statistics for the weighted mean $^{207}\text{Pb}/^{235}\text{U}$ (1048.0 ± 0.7 Ma; 2σ , MSWD 2.8) and $^{207}\text{Pb}/^{206}\text{Pb}$ (1049.9 ± 1.3 Ma; 2σ , MSWD 2.9) ages are due to a small dispersion in $^{207}\text{Pb}/^{235}\text{U}$ ratios. The cluster of data appears slightly discordant, although it lies within the concordia error envelope defined by the decay constant uncertainties of Jaffey and others (1971). This slight degree of discordance (0.26%) is consistent with the regular bias manifested by other high quality zircon and xenotime data sets reported by Schoene and others (2006), which was attributed to a small systematic error in one or more of the currently utilized U decay constants (most likely that of ^{235}U ; see also Mattinson, 2000). For the purposes of its use as a SHRIMP standard, BLR-1 titanite has an apparently uniform $^{206}\text{Pb}/^{238}\text{U}$ ratio of 0.17636 ± 0.000067 (2σ).

Determination of Pb, Sr, and Nd Initial Isotopic Ratios

Whole-rock fractions were analyzed for Pb, Sr, and Nd isotopes. These splits were spiked with ^{205}Pb , ^{233}U , ^{236}U , ^{230}Th , ^{87}Rb , ^{84}Sr , and ^{149}Sm - ^{150}Nd -enriched tracers and were digested in PFA-teflon screw-cap bombs with HF, and HNO_3 for a minimum of 48 hours at approximately 120°C . Lead was first isolated using anion exchange in 0.5N HBr medium. Uranium and thorium were separated together using anion exchange in 7N HNO_3 medium. Rubidium, strontium, and a rare-earth-element fraction were isolated using cation exchange in 2.5N HCl medium. Samarium and neodymium were separated from the rare-earth fraction using cation exchange in 0.2M n-methylacetic acid. Blanks for the procedure were on the order of 50 pg (Pb), 15 pg (U+Th), 40 pg (Rb), 300 pg (Sr), 50 pg (Sm), and 300 pg (Nd).

Mass spectrometry for Pb was performed using either a VG Sector 54 7-collector thermal ionization mass spectrometer or a VG54R single-collector mass spectrometer. Mass fractionation for Pb during mass spectrometry was monitored by replicate analyses of NIST standard SRM-981. Rubidium, Sr, Sm, and Nd were all analyzed using a VG 54R single-collector mass spectrometer. Rubidium was run using a triple Re filament technique. Strontium was run using a single oxidized Ta filament. Strontium isotopic data were normalized to $^{86}\text{Sr}/^{88}\text{Sr} = 0.1194$. Seventeen analyses of the NIST standard SRM-987 run during the course of this study yielded a mean $^{87}\text{Sr}/^{86}\text{Sr} = 0.710251 \pm 0.000008$ ($2\sigma_m$). Samarium and Nd were run using a triple filament technique with a Re ionizing filament and Ta sample filaments. Neodymium isotopic data were normalized to $^{146}\text{Nd}/^{144}\text{Nd} = 0.7219$. Eleven analyses of the La Jolla Nd standard run during the course of this study gave a mean $^{143}\text{Nd}/^{144}\text{Nd} = 0.511860 \pm 0.000009$ ($2\sigma_m$).

TABLE A1
U-Pb isotopic data for BLR-1 titanite

Fraction ^a	Composition				Isotopic Ratios				Dates (Ma)									
	Wt. (mg) ^b	[U] (ppm)	Th ^e (ppm)	Pb ^{*d} (pg)	$\frac{^{206}\text{Pb}^f}{^{209}\text{Pb}}$	$\frac{^{208}\text{Pb}^f}{^{206}\text{Pb}}$	$\frac{^{207}\text{Pb}^f}{^{235}\text{U}}$	% err ^e	$\frac{^{206}\text{Pb}^h}{^{238}\text{U}}$	$\frac{^{207}\text{Pb}^h}{^{235}\text{U}}$	$\frac{^{207}\text{Pb}^h}{^{206}\text{Pb}}$	$\frac{^{207}\text{Pb}^h}{^{206}\text{Pb}}$	±					
s1 aa	0.021	293	0.613	60	15.0	79	859.5	0.1868	0.17640	(.18)	1.8076	(.20)	0.074319	(.08)	0.923	1047.3	1048.2	1050.2±1.6
s2 aa	0.022	228	0.569	46	14.3	67	825.7	0.1733	0.17628	(.07)	1.8049	(.12)	0.074259	(.08)	0.689	1046.6	1047.2	1048.5±1.7
s3 aa	0.025	291	0.583	58	17.8	78	1020.7	0.1772	0.17635	(.07)	1.8056	(.10)	0.074259	(.07)	0.704	1047.0	1047.5	1048.5±1.4
s4	0.040	210	0.593	42	18.6	86	1062.7	0.1809	0.17641	(.07)	1.8086	(.10)	0.074357	(.06)	0.777	1047.3	1048.6	1051.2±1.2
s5	0.038	241	0.594	48	19.4	91	1108.2	0.1809	0.17638	(.07)	1.8080	(.10)	0.074345	(.07)	0.744	1047.2	1048.4	1050.8±1.3
s6	0.013	299	0.732	61	19.8	39	1098.9	0.2236	0.17572	(.10)	1.7996	(.13)	0.074277	(.08)	0.779	1043.6	1045.3	1049.0±1.6

^a All fractions composed of single fragments; aa indicates fragment abraded for 3 hours to a smooth ellipsoid in an alumina-ceramic vortical abrader.
^b Sample weights were estimated to within 40% using measured grain dimensions and a nominal density of 3.5 g/cm³.
^c Th contents calculated from radiogenic ²⁰⁸Pb and the ²⁰⁷Pb/²⁰⁶Pb date of the sample, assuming concordance between U-Th-Pb systems.
^d Pb* and Pbc represent radiogenic Pb and common Pb respectively.
^e Measured ratio corrected for fractionation and spike contribution; Pb fractionation estimated as 0.12 ± 0.04%/a.m.u. based on daily analysis of NBS-981.
^f Measured ratios corrected for fractionation, spike, blank, and initial common Pb; nominal Pb blank = 3.5 pg ± 50% (2σ); nominal U blank = 0.1 pg ± 50% (2σ); measured laboratory blank composition: ²⁰⁶Pb/²⁰⁹Pb = 19.10, ²⁰⁷Pb/²⁰⁹Pb = 15.72, ²⁰⁶Pb/²⁰⁹Pb = 38.65 ± 0.01 (2σ); initial common Pb composition from model of Stacy and Kramers (1972) at 1.05 Ga.
^g Numbers in parentheses are the % errors reported at the 2σ confidence interval, propagated using the algorithms of Ludwig (1980).
^h Isotopic ages calculated using the decay constants of Jaffey and others (1971): λ(²³⁵U) = 9.8485 × 10⁻¹⁰ yr⁻¹ and λ(²³⁸U) = 1.55125 × 10⁻¹⁰ yr⁻¹; error in ²⁰⁷Pb/²⁰⁶Pb date reported at the 2σ confidence interval.

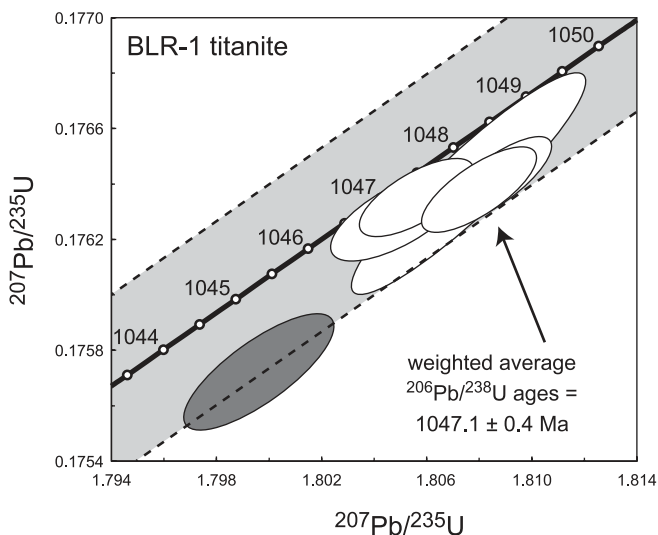


Fig. A1. Concordia plot of TIMS data for titanite SHRIMP standard BLR-1. Shaded gray band represents Concordia including uncertainties in decay constants. Discordant gray-filled error ellipse excluded from age calculation.

Initial isotopic ratios were calculated using ages determined by zircon U-Pb geochronology. This method presumes that the rocks remained closed to isotopic re-equilibration. In some cases, impossible initial Sr ratios were calculated, indicating that the presumption of isotopic closure was not achieved. For this reason, Sr data are shown in tabular form only.

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