- 1 Aggregate stability and size distribution of red soils under different land uses
- 2 integrally regulated by soil organic matter, and iron and aluminum oxides
- Jinsong Zhao^{a, b}, Shan Chen^a, Ronggui Hu^{a, c, *}, Yayu Li^a
- 4 a College of Resources and Environment, Huazhong Agricultural University, Wuhan 430070,
- 5 PR China
- 6 b Department of Ecology, Evolution and Marine Biology, University of California at Santa
- 7 Barbara, Santa Barbara, CA 93106, USA
- 8 CKey Laboratory of Arable Land Conservation (Middle and Lower Reaches of Yangtse River),
- 9 Ministry of Agriculture, Wuhan 430070, P R China
- * Corresponding author
- 12 Phone: +86-27-87282152
- 13 Fax: +86-27-87282137
- 14 Email: rghu@mail.hzau.edu.cn

Abstract

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The stability and size distribution of soil aggregatesis profoundly affected by land use, but the influencing mechanisms of land use are not clear. Astudy wascarried out to investigate and attempted to interpret the effects of land use on soil aggregates from types of land use and soil properties in soil samples and size fractions of soil aggregates. Soil samples were taken from 9 sites under paddy, forest, and upland in southern China. Thewet-sievingmethod was used to obtain 6 size fractions of soil aggregates: >5, $5\sim2$, $2\sim1$, $1\sim0.5$, $0.5\sim0.25$, and <0.25mm. The stability and size distribution of soil aggregates was measured as mean weight diameter (MWD), the percentage of water-stable aggregate (WSA) and the percentage of each size fraction (PSA). The quantities of soil organic carbon (SOC), humic substances, dithionite-citrate-bicarbonate (DCB) and oxalate extractable iron (Fe) and aluminum (Al) oxides were also measured. The results showed that types of land use solely explained66.6% variation of soil aggregates; SOC, DCB-extractable Fe and Al oxides, and oxalate-extractable Al oxide caused 84.3% variation, in which SOC contributed 29.0%, Fe and Al oxides contributed 33.8%, and their interactions contributed 21.4%. The multiple linear regression and partial correlation analysis showed that soil organic matterandFe and Al oxides had significant effects but played different roles on the stability and size distribution of soil aggregates. The study suggests that land use affects the stability and size distribution of soil aggregates through the integration of soil organic matter and Fe and Al oxides.

- **Keywords:** soil aggregates; stability; size fraction; land use; soil organic matter;
- 40 iron and aluminum oxides

1. Introduction

43	Soil structure has important influences on edaphic conditions and the
44	environment. The structure is often measured by the stability of soil
45	aggregates(Six et al., 2000; Bronick and Lal, 2005). Soil aggregation sustains soil
46	fertility because it reduces erosionand mediates soil aerationas well as water
47	infiltration and retention. Furthermore, soil aggregation protects soil organic
48	matter(SOM) from mineralizing because it physically reduces the accessibility of
49	organic compounds for microorganisms, extracellular enzymes, and oxygen
50	(Oades, 1984; Six et al., 2002a; von Lützow et al., 2006; Spohn and Giani, 2010).
51	The size distribution and stability of soil aggregates are under the control of
52	various mechanisms. Generally, the size distribution and stability of soil
53	aggregates positively correlates with the main cementing agents, such as SOM,
54	clay minerals, multi-valent cations and their complex in soil aggregates. According
55	to the hierarchical theory of aggregate formation, these materials may be
56	distributed unevenly in different size fractions of soil aggregates, andthey may
57	have close relationships with the stability of soil aggregates(Tisdall and Oades,
58	1982; Six et al., 2004).
59	Land use and associated management, such as crop sequencing,
60	fertilization, soil conditioning, drainage and irrigation, are the most important and
61	direct ways to affect soil structure and properties,throughitsimpact on
62	destruction forces and aggregate forming processes(Haynes et al., 1991; Besnard
63	et al., 1996; Jastrow, 1996; John et al., 2005; Ashagrie et al., 2007; Lehrsch et al.,

2012). However, the extent of the impact and the associated mechanisms of land use on soil aggregates remains unclear. Studies on the stability of soil aggregates under different land use rarely focus on the effects of various soil organic components, Fe and Al oxides in different size fractions of soil aggregates, and their interactions.

In southern China, the red soils cover approximately 1.14 million km². Because of intensive weathering and leaching, the soil is rich in Fe and Al oxides, and the dominant clay mineral is kaolinite(He et al., 2004). In recent years, soil structure has degraded rapidly, due to strong human intervention and high cultivating intensity(Iqbal et al., 2009), and the sustainability of the agricultural ecosystem is under serious threat.

Therefore, for soilsthat are rich in Fe and Al oxides, we hypothesize that the stability and size distribution of soil aggregates is controlled by SOM, which is heavily affected by the land use, whileFe and Al oxides and their interactions withSOMmanifest another important regulating mechanism for the stability and size distribution of soil aggregates. In order to test this hypothesis, in this study,the stability and size distribution of soil aggregates of red soils in southern Chinawere investigated from different types of land use, which were treated as a factor that represents the comprehensive effects brought about by the land use and its associated management, and the soil properties in soil samples and size fractions of soil aggregates.

2. Materials and Methods

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2.1 Study sites and soil samplings

Nine sites located in Jinxian, Jiangxi Province, Changsha and Taoyuan, Hunan Province, and Xianning, Hubei Province, southern China, were chosen, taking into account the variability of soil type (Inceptisols and Ultisols), in order to compare how the land use (paddy, forest, and upland) and the associated soil properties affected the stability and size distribution of soilaggregates. All sampling sites are in the subtropics and share similar climate. Those regionsare humid and have a monsoon climate, with an annual meantemperature above 15°C and annual mean precipitation above 1300 mm. The paddy sites have been used for planting rice for more than ten years. The forest sites are secondary forests, with predominance of *Pinus massoniana*in Xianning (F₁), *Cunninghamia* lanceolatain Changsha (F₃), and a mixed forest in Changsha (F₂). The uplands are used for grain or cash crops such as maize (*Zea mays*), rapeseed (*Brassica napus*), soybean (Glycine max), and sweet potato (Ipomoea batatas). Paddy and upland sites are managed with conventional tillage. The same type of land use shares similar agricultural practices.

At each site, 5 soil monoliths with diameter 10 cm were randomly taken from 0-15cm surfacebetweenOctober and December 2010. All samples were transported to the laboratory in their intact formand broken into small pieces along the natural crackby hand during the air-drying process. The soils from paddies are Endoaquepts, and those from uplands and forests are Plinthudults,

according to Soil Survey Staff(2014). All soil developed out ofmiddle

Pleistocene(Q2) red clay. The basicsoil properties, parent, land use, and soil type

are shown in Table 1 and Table 2.

2.2 Water-stable aggregates

To obtain different size fractions of water-stable aggregates, for the 5 samples from each site, 50 gof soilwereplaced in the top of a set of sieves with mesh sizes of5 mm, 2 mm, 1 mm, 0.5 mm, and 0.25 mm from top to bottom. The sieve set was placed on the shock rack of a Yoder aggregates analyzer (Institute of Soil Science, Chinese Academy of Sciences, Nanjing, China), submerged in water, and shaken with anamplitude of 3 cm and a frequency of 30 min⁻¹ for a duration of 30 mins. The size fraction at each sieve was then washed into a beaker of known mass and then dried and weighed.

Mean weight diameter (MWD) was calculated as

$$MWD = \frac{\sum_{i=1}^{6} d_i w_i}{w};$$

the percentage of water-stable aggregate (WSA) was calculated as

WSA =
$$100 \times \frac{\sum_{i=1}^{5} w_i}{w}$$
;

and the percentage of each size fraction of soil aggregate (PSA)was given as

$$PSA_i = 100 \times \frac{w_i}{w}$$
;

where d_i is the average diameter of i^{th} size fraction of the aggregate, w_i is the mass of i^{th} size fraction of the aggregate, andw is the total mass of all size fractions of the aggregate. i=1,2,...,6 represent the aggregate size >5 mm, $5\sim2$ mm, $2\sim1$ mm,

2.3 Carbon content of soil organic components

Humus (HF) in the soil sample or a size fraction of soil aggregateswas extractedusing the method proposed by the Institute of Soil Science, Chinese Academy of Sciences(1978). For each sample, 5g of air-dried soilwere placed into a 250-mL Erlenmeyer flask. Then 100 mL 0.1 mol L^{-1} sodium pyrophosphate-sodium hydroxide were added. The mixture was shaken for 30 mins and then left to stand for $13\sim14$ h. Depending on the color depth, $2\sim15$ mL supernatant were drawn into the digestive tract and neutralized with 0.5 mol $L^{-1}H_2SO_4$. The supernatant was evaporated to dryness on a boiling water bath, after which the carbon content of HF was determined using the dichromate oxidization method(Bao, 2000).

Fifty millilitersof the previous supernatant were drawn into a 100-mL Erlenmeyer flask. The pH was adjusted to 3 or less with 0.5 mol L-1 H₂SO₄ at $60\sim70$ °C. It was then kept at 80 °C for 30 mins with a water bath. After cooling for about 8h, it was filtered with a fine filter paper (Grade 44, Whatman). The precipitation on the filter paper was washed 3 times using 0.025 mol L-1 H₂SO₄ and then solubilized and washed into a 50-mL flask with 0.05 mol L-1NaOH at $40\sim50$ °C. Then, $10\sim25$ mLsolution weredrawn into the digestive tract and neutralized with 0.5 mol L-1H₂SO₄. The solution wasevaporated to dryness on a boiling water bath, after which the carbon content of humic acid (HA)was

determined using the dichromate oxidization method(Bao, 2000). The carbon content of fulvic acid (FA) was obtained by subtracting the carbon content of HA from HF(Institute of Soil Science, Chinese Academy of Sciences, 1978).

2.4 Organic carbon content in soil and size fractions

The soil sample or a size fraction of soil aggregateswasgrounded to pass through a sieve with a mesh size of 0.149 mm. For each sample, 0.3 g soil wasplaced in a small crucible to measure the total content of soil organic carbon (SOC) with an elemental analyzer (Vario MAX C/N, Elementar Analysensysteme GmbH, Hanau, Germany). A reference samplewas added after each 15 soil samples.

2.5 Free-form Fe and Al oxides

Free-formFe and Al oxides were extracted using the dithionite-citrate-bicarbonate (DCB) method(Li, 1997). For each sample, 0.3 g air-dried soil was placed into a 50-mL centrifuge tube and then 20 mL 0.3 mol L-1 sodium citrate and 2.5 mL 1 mol L-1 sodium bicarbonatewere added. When the temperature of the water bath reached 80 °C, 0.5 g sodium dithionite was added. The mixture was stirred constantly for 15 mins. After cooling and centrifugation, the supernatant was transferred into a 250 mL volumetric flask. The sediment in the centrifuge tube was treated with the above process again. Finally, all material in the tube was washed into the same flask with 1 mol L-1 NaCl. The Fe and Al in

the solution were determined with an inductively coupled plasma (ICP) analyzer (Vista-MPX, Varian, Inc., Palo Alto, CA, USA).

2.6 Amorphous Fe and Al oxide

Amorphous Fe and Al oxide was extracted using the ammonium oxalate method(Li, 1997). For each sample, 0.5 g air-dried soilwas placed intoa 50-mL centrifuge tube and then 25 mL 0.2 mol L⁻¹ ammonium oxalate was added (soil:solution = 1:50). The mixture was shakenfor 2h in the dark and then centrifugedimmediately with rotation speed 4000 rmin⁻¹ to obtain supernatant. The supernatantwas diluted 5 times and the Fe and Al content in the solution were determined withan ICP analyzer (Vista-MPX, Varian, Inc., Palo Alto, CA, USA).

2.7 Statistical analysis

All experimentally determined data are expressed as mean \pm SE (standard error of mean). For the soil properties in each site, the mean and SE that were calculated from the measurements of 5 samples (n = 5), while for the soil properties in a specific type of land use, they were calculated from 3 averages (n = 3). The differences of MWD, WSA, and PSA of soil aggregates under different types of land use were tested by one-way analysis of variance (ANOVA). If the results were statistically significant, then the post hoc multiple comparisons were performed with least significant difference (LSD).

Redundancy analysis (RDA) was used to explore the relationship between response variables (i.e., MWD, WSA, and PSA) and explanatory factors such astypes of land use, and the contents of Fe and Al oxides, FA, HA, humus, HA/FA ratio, and SOC in soil and in various size fractions of soil aggregates. The RDA model with up to 6 explanatory variables was built through best subset searching. The multicollinearity between explanatory variables was checked with variance inflation factors (VIF), and any model containing a variable with VIF greater than 10 was discarded. The Akaike information criterion (AIC) was used to select an optimal model. The validation of the RDA model and the significance of each variable on MWD and WSA were checked by permutation tests. The permutation was repeated 1999 times. The significant level was set to 0.05. The variation of response variables with respect to SOC and Fe and Al oxides were also partitioned, using the coefficient of determination of the RDA model (Borcard et al., 1992).

The relationship between individual response variables and soil properties in size fractions of soil aggregates was explored with multiple linear regression, based on best subset searching. The same procedures as in RDA were used to check the multicollinearity and to choose an optimal model. Partial correlation analysis was used to explore the correlation between the individual response variable and the corresponding soil properties in the optimal model.

All statistical analysis and plotting were carried out using the statistical and computing environment R version 3.3.0(R Core Team, 2016). RDA, variance

partition, and permutation testswere provided by vegan version 2.3(Oksanen et al., 2016), an add-on package for R.

3. Results

3.1 Soil properties and land use

The basicsoil properties, such as pH, bulk density, SOC, soil total nitrogen (STN), C/N ratio and texture of the 9 sites, are presented in Table 2. Although the soils were developed from the same parent material, the properties varied significantly among the different types of land use. Paddy sites had the lowest bulk density $(1.09\pm0.07~{\rm g~cm^{-3}})$, highest pH (5.0 ± 0.14) , and highest STN $(1.56\pm0.04~{\rm gN~kg^{-1}})$, which were significantly different from those in upland or forest sites. The contents of silt, sand, and clay were not significantly different among the types of land use.

Other soil properties, such as SOCand its components, revealed significant differences among different types of land use (Table 3). The differences between paddy and forestwerenot significant, but in both types of land use, SOC, HF, and FA were significantly higher than those in uplands (p<0.05), while HA and HA/FA ratios were not significantly different. The Fe and Al oxides did not vary significantly between any types of land use.

3.2Stability and size distribution of soil aggregates

Under different types of land use, the average MWD and WSA appearedin

order as paddy>forest >upland (Fig. 1).Paddy has the highest MWD, followed by forest. Upland has the lowest MWD,with half of this region being under forest land use. PSA₁in paddy washighest, while the rest of the size fractions weresmaller. PSA₆washighest in upland. The differences amongMWD, WSA, PSA₁, PSA₄,and PSA₆ under different types of land use weresignificant (Fig. 1). MWD, WSA, and PSA₆showedsimilar patternsandno significant differences were found among upland and the other 2 types. While PSA₁and PSA₄showed similar patternsofsignificant differences between paddy and other types of land use, there was no significant difference between forest and upland.PSA₂,PSA₃, and PSA₅were not significantly different among the 3 types of land use.

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The effects of land usetypeonthe stability and size distribution of soil aggregates (MWD, WSA, and PSA)were explored by the RDA. The effects of land use accounted for 66.6% of variation of the stability and size distribution of soil aggregates. The first axis of RDA accounted for 45.3%, mainly contributed by upland and paddy, while the second axis of RDA accounted for 21.3%, which was mainly contributed by forest (Fig. 2). MWD, WSA and PSA presented a clear pattern, where MWD, WSA and PSA₁were grouped and closely related with each other. Similarly, the middle-sized fractions (PSA₂, PSA₃, and PSA₄), and small-sized fractions of soil aggregates (PSA₅, and PSA₆)were grouped separately (Fig. 2). Paddy, forest and uplandhada positive contribution togroup 3, respectively. Permutation tests showed that the land use types had significant

effects(p<0.05) on the stability and size distribution of soil aggregates.

The relationships between soil properties and the stability and size distribution of soil aggregateswere also explored by RDA, and showed thatthe DCB-extractable Fe and Al oxides, oxalate-extractable Al oxide, and SOC couldexplain 84.3% of the variation in the stability and size distribution of soil aggregates under different types of land use. The first RDA axis accounted for 57.4%, mainly contributed by SOC, while the second RDA axis accounted for 25.0%,mainly contributed by the oxalate-extractable Al oxide (Fig. 3). The DCB-extractable Fe and Al oxides contributed equally to both ordination axes. All soil properties in the RDA model had significant positive correlations with MWD, WSA, and PSA₁, but negative with PSA₅ and PSA₆. For PSA₂, PSA₃ and PSA₄, Fe and Al oxides were positive contributors, while SOC presented a negative contribution (Fig. 3). Permutation tests showed that all soil properties significantly affected soil aggregates (*p*<0.05).

The variation of the stability and size distribution of soil aggregates with respect to SOC and Fe and Al oxides was partitioned (Fig. 4). The contribution of SOC was 29.0%, while the contribution of Al and Fe oxides was 33.8%. The interactions between SOC and Fe and Al oxides contributed 21.4% to the variation of the stability and size distribution of soil aggregates.

The pattern of site scores presented in the ordination plot (Fig. 2 and Fig. 3) was same for both RDA models. It clearly showed that the first RDA axis indicated SOC in different sites increasing to left, while the second RDA axis

indicated the vegetation increasing to top or anthropogenic disturbance increasing to bottom.

3.3 Soil properties of different aggregatesize fractions

Different types of land use had a significant impact on soil properties of different size fractions of soil aggregates (Table 4). HA in $0.5\sim0.25$ mm aggregate (HA₅), organic carbon in <5 mm aggregate (SOC₂₋₆), and FA in >5mm aggregate (FA₁)were significantly different under different types of land use. In the <0.25 mm aggregate size fraction,the oxalate-extractable Al oxide(Al₆^{Oxa}) and HA/FA ratio (HA₆/FA₆) were also significantly different under different types of land use.

The stability of soil aggregates (MWD and WSA)were closely related to the corresponding soil properties in various size fractions of soil aggregates (Table 5). The oxalate-extractable Al oxide in <0.25 mm aggregates (Al_6^{0xa}) and the organic matter in >2mm aggregate (HA_1 and FA_2) positively contributed to MWD and WSA, while the oxalate-extractable Fe (Fe_4^{0xa}) and Al oxides (Al_2^{0xa}) and humus (HF_3) in mid-sized aggregates negatively contributed to MWD and WSA.

Except for PSA₂, percentages of all size fractions of soil aggregates resulted insignificantly linear relationships with their corresponding soil properties. The organic matter (including SOC, HA, and FA) was negatively related to the percentage of smaller aggregates (<1 mm) but positively correlated with the percentage of bigger aggregates (>1 mm). However, the correlations of Fe and Al

oxides werenot consistent within different size fractions of soil aggregates.

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4. Discussion

4.1 Effects of land use on soil aggregates

The effects of land use on soil were profound and comprehensive. The stability and size distribution of soil aggregateswere significantly different in terms of types of land use (Fig. 1). The macro-aggregates, which generally had anorder offorest>rangeland/grassland>arable land, have been observed todecrease or disappear with sustained planting (Haynes et al., 1991; Jastrow, 1996; John et al., 2005; Ashagrie et al., 2007). However, the percentage of macro-aggregates in the paddy washighest among the 3 types of land use. Although the paddy is one kind of arable land that involves long-term anthropogenic disturbances, its special water management systems (Kögel-Knabner et al., 2010) resulted in highest SOC, HF, and HA(Table 3). The soil properties, especially SOM related properties, were significantly different in terms of land use type (Table 3). On the one hand, SOM is significantly affected by land use(John et al., 2005; Don et al., 2011; Li et al., 2015); on the other hand, SOM playsprofoundly important roles in the stability and size distribution of soil aggregates (Tisdall and Oades, 1982; Barral et al., 1998; Abiven et al., 2009; Alagöz and Yilmaz, 2009), it is suspected that the effects of land use on soil aggregates are driven by the SOM or its components.

Additional,66.6% variations of soil aggregates were explained by land use

(Fig. 2), implying that the driving mechanism for the stability and size distribution of soil aggregates should include other factors, such as Fe and Al oxides, although they were not influenced by land use.

4.2 Effects of soil properties on soil aggregates

The 84.3% variations of soil aggregates interpreted by organic matter and Fe and Al oxides (Fig. 3) mean that SOC and Fe and Al oxides and their interactions may be an important mechanism for the stability and size distribution of soil aggregates. It is also explicitly demonstrated by the partition of the variation of soil aggregates (Fig. 4).

SOM participates in various processes of the formation and stability of soil aggregates (Tisdall and Oades, 1982; Six et al., 2004). Furthermore, Fe and Al oxides and SOM can form stable organo-mineral complexes that can enhance the tensile strength and stability of the aggregate (Barral et al., 1998). Edwards and Bremner (1967) suggested that the formation of soil aggregates is a specific process that involves the absorption of organic matter into clay through the bridge of the multivalent metal cation; this is also known as a composite process of soil organic-inorganic complexes (Bronick and Lal, 2005).

Amézketa (1999) argued that Fe and Al oxides, which bond with organic and inorganic compounds, or aggregates through a cation bridge to improve soil structure, are the most active components in soil.

4.3 Role of soil properties in different size fractionsofsoil aggregates

The soil properties do not differ significantly in terms of fraction size, which is inconsistent with the hierarchical theory of soil aggregates (Tisdall and Oades, 1982; Puget et al., 2000; Six et al., 2000; John et al., 2005). It may result from kaolinite being the dominant clay, which as a 1:1 structure. Six et al. (2000; 2002b) suggested that soil aggregates were nother archical in this kind of soil, and the bonding material between aggregates included not only organic materials but also Fe and Al oxides and other cementing materials..

The soil properties in each size fraction played different roles in the stability of soil aggregates. The oxalate-extractable Al oxide in the smalleraggregate (Al_6^{Oxa}) and the soil organic components in the bigger aggregates (HA_1 , FA_2)contributed positively to the MWD and WSA (Table 5). However, the oxalate-extractable Al oxides in $5{\sim}2$ mm aggregate (Al_2^{Oxa}), the humus in $2{\sim}1$ mm aggregate (HF_3), and the oxalate-extractable Fe oxides in $1{\sim}0.5$ mm aggregate (Fe_4^{Oxa}) had negative contributions to the MWD and WSA.

In the smalleraggregates (<0.5mm), the role of organic components(SOC₅, HF₆, and HA₆)to the size distribution of soil aggregates were consistent. When organic components increased, the percentage of the smalleraggregates was significantly reduced, which indicates that the smalleraggregates were combining to form larger aggregates. This is also consistent with the theory that macro-aggregates are formed from micro-aggregates.

Meanwhile, the rolesof different forms of Fe and Al oxides in those processes were inconsistent. In the smaller aggregates (0.5~0.25 mm and <0.25 mm), the oxalate-extractable Al oxide (Al $_5^{0xa}$ and Al $_6^{0xa}$) was negatively correlated with the corresponding PSA, indicating that it was an important cementing substance beneficial to the formation of large aggregates, while the Fe oxide (Fe $_5^{0xa}$, Fe $_5^{DCB}$, and Fe $_6^{0xa}$) was a basic component of the aggregates due to a positive correlation with the corresponding PSA. In the >0.5 mmaggregates, the oxalate-extractable Fe oxide (Fe $_3^{0xa}$ and Fe $_4^{0xa}$) acted as a cementing substance.

The roles of soil properties in the middle size fractionson the stability and size distribution of soil aggregates are worthy of further study. On the one hand, these middle-sized aggregates can be combined through the cementing agents, such as Fe and Al oxides, and organic matter, to formlarger aggregates. In this case, the soil properties were positively related to the MWD and WSA and negatively related to the corresponding PSA. On the other hand, middle-sized aggregates can breakdown into smaller components, after which the corresponding variables should have the completely opposite pattern. Both processes of soil aggregates are homeostatic; they are indistinguishable unless dynamically monitored. In addition, because there was no conversion among the three types of land use, the relationships of transformations among the various size fractions of soil aggregates could not be determined.

5. Conclusions

The effects of land use onthe stability and size distribution of soil aggregateswere investigated based on soil properties and size fractions of soil aggregates in red soils in southern China with three types of land use, i.e., paddy, forest, and upland. The results suggested that land use significantly affected the stability and size distribution of soil aggregates through changing the SOM in soil and size fractions by high cultivating intensity or management related hydro-regime. Additionally, the Fe and Al oxides, which were not affected by land use, also hadremarkable influence. The interaction of Fe and Al oxides and SOM, and the differential roles of Fe and Al oxides in formation or break-down of soil aggregates with middle size, werekey to the stability and size distribution of soil aggregates in red soil in southern China.

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References

Abiven, S., Menasseri, S., Chenu, C., 2009. The effects of organic inputs over time on soil

413	aggregate stability – A literature analysis. Soil Biology and Biochemistry 41, 1–12.
414	doi:10.1016/j.soilbio.2008.09.015
415	Alagöz, Z., Yilmaz, E., 2009. Effects of different sources of organic matter on soil aggregate
416	formation and stability: A laboratory study on a Lithic Rhodoxeralf from Turkey. Soil
417	and Tillage Research 103, 419–424. doi:10.1016/j.still.2008.12.006
418	Amézketa, E., 1999. Soil Aggregate Stability: A Review. Journal of Sustainable Agriculture 14,
419	83-151. doi:10.1300/J064v14n02_08
420	Ashagrie, Y., Zech, W., Guggenberger, G., Mamo, T., 2007. Soil aggregation, and total and
421	particulate organic matter following conversion of native forests to continuous
422	cultivation in Ethiopia. Soil and Tillage Research 94, 101–108.
423	doi:10.1016/j.still.2006.07.005
424	Bao, S.D., 2000. Soil and agricultural chemistry analysis, 3rd ed. China Agriculture Press,
425	Beijing.
426	Barral, M.T., Arias, M., Guérif, J., 1998. Effects of iron and organic matter on the porosity and
427	structural stability of soil aggregates. Soil and Tillage Research 46, 261–272.
428	doi:10.1016/S0167-1987(98)00092-0
429	Besnard, E., Chenu, C., Balesdent, J., Puget, P., Arrouays, D., 1996. Fate of particulate organic
430	matter in soil aggregates during cultivation. European Journal of Soil Science 47,
431	495-503. doi:10.1111/j.1365-2389.1996.tb01849.x
432	Borcard, D., Legendre, P., Drapeau, P., 1992. Partialling out the spatial component of
433	ecological variation. Ecology 73, 1045–1055. doi:10.2307/1940179
434	Bronick, C.J., Lal, R., 2005. Soil structure and management: A review. Geoderma 124, 3–22.

435	doi:10.1016/j.geoderma.2004.03.005
436	Don, A., Schumacher, J., Freibauer, A., 2011. Impact of tropical land-use change on soil organic
437	carbon stocks – a meta-analysis. Global Change Biology 17, 1658–1670.
438	doi:10.1111/j.1365-2486.2010.02336.x
439	Edwards, A.P., Bremner, J.M., 1967. Microaggregates in Soils. Journal of Soil Science 18, 64–73.
440	doi:10.1111/j.1365-2389.1967.tb01488.x
441	Haynes, R.J., Swift, R.S., Stephen, R.C., 1991. Influence of mixed cropping rotations
442	(pasture-arable) on organic matter content, water stable aggregation and clod
443	porosity in a group of soils. Soil and Tillage Research 19, 77–87.
444	doi:10.1016/0167-1987(91)90111-A
445	He, Z.L., Zhang, M.K., Wilson, M.J., 2004. Distribution and classification of red soils in China, in:
446	Wilson, M.J., He, Z.L., Yang, X.E. (Eds.), The Red Soils of China: Their Nature,
447	Management and Utilization. Kluwer Academic Publishers, Dordrecht, pp. 29–33.
448	Institute of Soil Science, Chinese Academy of Sciences, 1978. The Physical and Chemical
449	Analysis of Soil. Shanghai Science and Technology Press, Shanghai.
450	Iqbal, J., Hu, R., Lin, S., Ahamadou, B., Feng, M., 2009. Carbon dioxide emissions from Ultisol
451	under different land uses in mid-subtropical China. Geoderma 152, 63–73.
452	doi:10.1016/j.geoderma.2009.05.011
453	Jastrow, J.D., 1996. Soil aggregate formation and the accrual of particulate and
454	mineral-associated organic matter. Soil Biology and Biochemistry 28, 665–676.
455	doi:10.1016/0038-0717(95)00159-X
456	John, B., Yamashita, T., Ludwig, B., Flessa, H., 2005. Storage of organic carbon in aggregate and

457	density fractions of silty soils under different types of land use. Geoderma 128,
458	63-79. doi:10.1016/j.geoderma.2004.12.013
459	Kögel-Knabner, I., Amelung, W., Cao, Z., Fiedler, S., Frenzel, P., Jahn, R., Kalbitz, K., Kölbl, A.,
460	Schloter, M., 2010. Biogeochemistry of paddy soils. Geoderma 157, 1–14.
461	doi:10.1016/j.geoderma.2010.03.009
462	Lehrsch, G.A., Sojka, R.E., Koehn, A.C., 2012. Surfactant effects on soil aggregate tensile
463	strength. Geoderma 189–190, 199–206. doi:10.1016/j.geoderma.2012.06.015
464	Li, H., Han, X., You, M., Xing, B., 2015. Organic matter associated with soil aggregate fractions
465	of a black soil in Northeast China: Impacts of land-use change and long-term
466	fertilization. Communications in Soil Science and Plant Analysis 46, 405–423.
467	doi:10.1080/00103624.2014.956887
468	Li X.Y., 1997. Soil chemistry and experimental guidelines. China Agriculture Press, Beijing.
469	Oades, J.M., 1984. Soil organic matter and structural stability: mechanisms and implications
470	for management. Plant Soil 76, 319-337. doi:10.1007/BF02205590
471	Oksanen, J., Blanchet, F.G., Kindt, R., Legendre, P., Minchin, P.R., O'Hara, R.B., Simpson, G.L.,
472	Solymos, P., Stevens, M.H.H., Wagner, H., 2016. vegan: Community Ecology Package.R
473	package version 2.3-5. https://CRAN.R-project.org/package=vegan
474	Puget, P., Chenu, C., Balesdent, J., 2000. Dynamics of soil organic matter associated with
475	particle-size fractions of water-stable aggregates. European Journal of Soil Science
476	51, 595-605. doi:10.1111/j.1365-2389.2000.00353.x
477	R Core Team, 2016. R: A language and environment for statistical computing. R Foundation
478	for Statistical Computing, Vienna, Austria.https://www.R-project.org/

479	Six, J., Bossuyt, H., Degryze, S., Denef, K., 2004. A history of research on the link between
480	(micro)aggregates, soil biota, and soil organic matter dynamics. Soil and Tillage
481	Research 79, 7–31. doi:10.1016/j.still.2004.03.008
482	Six, J., Conant, R.T., Paul, E.A., Paustian, K., 2002a. Stabilization mechanisms of soil organic
483	matter: Implications for C-saturation of soils. Plant and Soil 241, 155–176.
484	doi:10.1023/A:1016125726789
485	Six, J., Elliott, E.T., Paustian, K., 2000. Soil structure and soil organic matter II. A normalized
486	stability index and the effect of mineralogy. Soil Science Society of America Journal
487	64, 1042–1049. doi:10.2136/sssaj2000.6431042x
488	Six, J., Feller, C., Denef, K., Ogle, S.M., de Moraes, J.C., Albrecht, A., 2002b. Soil organic matter,
489	biota and aggregation in temperate and tropical soils - Effects of no-tillage.
490	Agronomie 22, 755–775. doi:10.1051/agro:2002043
491	Soil Survey Staff, 2014. Keys to soil taxonomy, 12th ed. USDA-Natural Resources
492	Conservation Service, Washington, DC.
493	Spohn, M., Giani, L., 2010. Water-stable aggregates, glomalin-related soil protein, and
494	carbohydrates in a chronosequence of sandy hydromorphic soils. Soil Biology and
495	Biochemistry 42, 1505–1511. doi:10.1016/j.soilbio.2010.05.015
496	Tisdall, J.M., Oades, J.M., 1982. Organic matter and water-stable aggregates in soils. Journal of
497	Soil Science 33, 141–163. doi:10.1111/j.1365-2389.1982.tb01755.x
498	von Lützow, M., Kögel-Knabner, I., Ekschmitt, K., Matzner, E., Guggenberger, G., Marschner, B.,
499	Flessa, H., 2006. Stabilization of organic matter in temperate soils: mechanisms and
500	their relevance under different soil conditions – a review. European Journal of Soil

Table 1 Location, land use, annual mean precipitation, parent matherial and soil type of study sites

Location	Code	Coordinates	Land use	Site description	AMP (mm)	Parent material	Soil type
Jinxian	P1	N28°21′31.9″	Paddy	>50 yrs paddy field	1580	Redeposited	Endoaquepts
		E116°10′26.8″				Q2 red clay	
Jinxian	P2	N28°21′02.4″	Paddy	>20 yrs paddy field	1580	Redeposited	Endoaquepts
		E116°10′30.9″				Q2 red clay	
Changsha	Р3	N28°13′31.1″	Paddy	>10 yrs paddy field	1410	Q2 red clay	Endoaquepts
		E113°09′94.2″					
Xianning	F1	N30°00′03.6″	Forest	>40 yrs horsetail pine forest	1547	Q2 red clay	Plinthudults
		E114°22′38.7″					
Changsha	F2	N28°13′31.9″	Forest	>10 yrs mixed forest	1319	Q2 red clay	Plinthudults
		E113°09′45.7″					
Changsha	F3	N28°15′40.3″	Forest	>30 yrs Cunninghamia	1319	Q2 red clay	Plinthudults
		E113°11′10.6″		lanceolata forest			
Taoyuan	U1	N29°13′43.4″	Upland	>3 yrs corn with straw return	1510	Q2 red clay	Plinthudults
		E111°31′17.8″					
Xianning	U2	N30°01′18.9″	Upland	>3 yrs rapeseed and soybean	1547	Q2 red clay	Plinthudults
		E114°21′44.4″					
Changsha	U3	N28°16′06.4″	Upland >	>3 yrs soybean and sweet potato	1410	Q2 red clay	Plinthudults
		E113°14′25.9″					

Note: Q2 means Middle Pleistocene, the second part of Quaternary.

Table 2 Basic soil properties, texture and structure of study sites

Code	BD	рН	SOC	STN	C/N	Silt	Clay	Sand	Texture	Structure
	(g cm ⁻³)		(gC kg ⁻¹)	(gN kg ⁻¹)		(%)	(%)	(%)		
P1	0.96±0.01	4.9±0.04	18.6±0.35	1.71±0.02	10.8±0.17	66.6±0.12	30.5±0.29	2.9±0.17	silty clay loam	Subangular blocky
P2	1.14±0.03	4.8±0.05	14.9±0.39	1.51±0.06	9.9±0.15	67.9±2.67	25.8±1.06	6.3±3.73	silty loam	Subangular blocky
Р3	1.18±0.06	5.2±0.06	13.9±0.12	1.47±0.01	9.5±0.09	46.9±3.56	41.9±3.50	11.2±0.06	silty clay	Subangular blocky
F1	1.31±0.12	4.5±0.06	15.5±0.51	1.34±0.04	11.6±0.43	52.2±0.08	41.8±0.85	6.1±0.77	silty clay	Subangular blocky
F2	1.45±0.08	4.3±0.05	11.5±0.15	1.15±0.02	10.0±0.08	46.7±0.71	37.7±1.71	15.6±1.00	silty clay loam	Crumb
F3	1.39±0.13	4.2±0.06	11.9±0.22	1.28±0.01	9.3±0.14	42.9±0.47	52.7±0.71	4.4±0.24	silty clay	Crumb
U1	1.25±0.12	4.9±0.05	8.2±0.32	1.30±0.05	6.4±0.24	53.4±1.50	24.4±0.88	22.2±0.62	silty loam	Subangular blocky
U2	1.35±0.05	4.6 ± 0.05	10.0±0.60	1.42±0.04	7.1±0.45	57.9±0.62	34.0±0.54	8.2±1.16	silty clay loam	Subangular blocky
U3	1.32±0.15	4.7±0.02	6.0±0.23	0.99±0.03	6.1±0.18	45.5±1.12	48.2±0.87	6.3±1.99	silty clay	Crumb

Note: All data are expressed in mean±SE, which are calculated based on 5 samples. BD is bulk density; SOC is soil organic carbon; and STN is soil total nitrogen; C/N is the ratio of SOC to STN. pH was measured in a 2.5:1 mixture of water:soil.

Table 3 Soil properties under different types of land use

						_	_		
	Fe ^{0xa}	Fe ^{DCB}	Al ^{0xa}	Aldcb	HF	FA	НА	SOC	HA/FA
		g k	g-1			g C k	g-1		
Paddy	5.68±1.59	16.29±2.08	1.54±0.25	1.43±0.25	5.07±0.66a	3.87±0.42a	1.20±0.50	15.8±1.41a	0.32±0.12
Forest	1.93±0.58	17.49±2.55	2.34±0.40	1.52±0.17	4.53±0.31ab	3.39±0.08a	1.15±0.23	13.0±1.28a	0.33±0.06
Upland	3.73±0.82	14.20±1.01	1.41±0.33	1.10±0.20	2.93±0.42b	2.06±0.27 b	0.87±0.15	8.1±1.16b	0.42±0.02

Note: All data is expressed as mean±SE, where sample size is 3, i.e., the number of sites for a specific land use. Different letters in a column indicates the significant difference of the specific property under the different types of land use based on the post hoc multiple comparison by least significant difference with significant level 0.05; Fe^{0xa} and Al^{0xa} are the oxalate extractable Fe and Al oxides; Fe^{DCB} and Al^{DCB} are the DCB extractable Fe and Al oxides; HF, FA, HA, SOC are the carbon content of humus, fulvic acid, humic aid and soil organic matter, respectively.

Table 4 Soil properties in size fractions of soil aggregates under different types of land use

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	Al ₆ ^{Oxa}	HA ₅	FA ₁	SOC ₂	SOC ₃	SOC ₄	SOC ₅	SOC ₆	(HA/FA) ₆	
Paddy	1.77±0.16a	1.22±0.17a	3.47±0.27a	17.2±1.55a	17.8±1.70a	16.5±0.94a	15.6±1.35a	14.6±1.94a	0.32±0.06b	
Forest	1.66±0.08a	0.52±0.02b	3.37±0.30a	11.4±0.50ab	10.0±0.36b	9.0±0.26b	8.4±0.32b	8.6±0.79b	0.24±0.04b	
Upland	1.20±0.11b	0.86±0.14ab	2.10±0.41b	9.3±2.60b	9.9±2.82b	8.7±2.18b	7.8±1.68b	7.1±1.50b	0.61±0.08a	

Note: Only the properties that significantly affected by land use are show; All data is expressed as mean \pm SE, where sample size is 3, i.e., the number of sites for a specific land use. Different letters in a column indicates the significant difference of the specific property in terms of types of land use based on the post hoc multiple comparison by least significant difference with significant level 0.05; Al^{0xa} is the oxalate extractable Al oxides; FA, HA, and SOC are the carbon content of humus, fulvic acid, humic aid and soil organic matter, respectively; The subscript of each properties indicates the fraction size of the soil aggregate, i.e., 1, 2, ..., 6 represents the size of >5 mm, 5 \sim 2 mm, 2 \sim 1 mm, 1 \sim 0.5 mm, 0.5 \sim 0.25mm and <0.25mm.

Table 5 Multiple linear regression and partial correlation analysis between stability and size distribution and soil properties in the different size fraction of soil aggregates under different types of land use

Aggregate properties	Linear regression model with partial correlation coefficient	R ²
MWD	$-2.785 + 2.792 \text{ Al}_{6}^{\text{Oxa}} (0.999) - 0.400 \text{ Al}_{2}^{\text{Oxa}} (-0.974) - 0.044 \text{ Fe}_{4}^{\text{Oxa}} (-0.985) + 1.144 \text{ HA}_{1} (0.999)$	0.9997
WSA	$-34.344 + 50.763 \text{ Al}_{6}^{\text{Oxa}} (0.9998) - 10.233 \text{ HF}_{3} (-0.998) + 32.434 \text{ HA}_{1} (0.9997) + 7.513 \text{ FA}_{2} (0.995)$	0.999 <u>6</u>
PSA ₁	$-19.104 - 3.150 \text{ Fe}_{1}^{\text{DCB}} (-0.825) - 17.367 \text{ Al}_{1}^{\text{Oxa}} (-0.919) + 56.646 \text{ Al}_{1}^{\text{DCB}} (0.948) + 15.731 \text{ FA}_{1} (0.956)$	0.952
PSA ₃	$18.556 - 0.613 \frac{\text{Fe}_3^{\text{Oxa}}}{1.0000} (-0.560) - 1.212 \frac{\text{Fe}_3^{\text{DCB}}}{1.0000} (-0.657) + 9.730 \text{ Al}_3^{\text{Oxa}} (0.762)$	0.797
PSA ₄	$13.508-0.286 \text{ Fe}_{4}^{\text{Oxa}} (-0.890)+0.311 \text{ Fe}_{4}^{\text{DCB}} (0.897)-2.233 \text{ Al}_{4}^{\text{DCB}} (-0.874)-1.523 \text{ HF}_{4} (-0.979)$	0.975
PSA ₅	$9.995 + 0.651 \text{ Fe}_{5}^{\text{Oxa}} (0.985) + 0.151 \text{ Fe}_{5}^{\text{DCB}} (0.913) - 1.624 \text{ Al}_{5}^{\text{Oxa}} (-0.937) - 0.443 \text{ SOC}_{5} (-0.992)$	0.987
PSA ₆	$126.066 + 7.673 \text{ Fe}_{6}^{0\text{xa}} (0.966) - 40.242 \text{ Al}_{6}^{0\text{xa}} (-0.978) - 6.175 \text{ HF}_{6} (-0.937) - 32.253 \text{ HA}_{6} (-0.892)$	0.977

Note: All variables are partially correlated with MWD, WSA and PSA with p < 0.05, except for the ones underlined; The partial correlation coefficient is shown in parentheses. The significance test of the partial correlation is based on a t distribution with degree of freedom: n-2-nv, where n is sample size, i.e., 9, nv is the number of the controlled variables. Fe^{0xa} and Al^{0xa} are the oxalate extractable Fe and Al oxides; Fe^{DCB} and Al^{DCB} are the DCB extractable Fe and Al oxides; HF, FA, HA, SOC are the carbon content of humus, fulvic acid, humic aid and soil organic matter, respectively; The subscript of each properties indicates the fraction size of the soil aggregate, i.e., 1, 2, ..., 6 represents the size of >5 mm, $5\sim2$ mm, $2\sim1$ mm, $1\sim0.5$ mm, $0.5\sim0.25$ mm and <0.25mm.

Figure captions:

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- Fig. 1. Mean weight diameter (MWD), water-stable aggregate (WSA), and size
- 4 distribution of soil aggregates under 3 types of land use. PSA_i is the percentage of
- 5 the ith size fraction of soil aggregates, with i = 1, 2, ..., 6 representing the size
- of > 5 mm, $5\sim2$ mm, $2\sim1$ mm, $1\sim0.5$ mm, $0.5\sim0.25$ mm and <0.25mm. Different
- 7 letters on the top of the error bars indicate the significant differences in terms of
- 8 types of land use. The error bar indicates the standard error (n = 3) of a
- 9 measurement.

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- 11 **Fig. 2.** Ordination plot of redundancy analysis (RDA) for the stability and size
- distribution of soil aggregates with types of land use as the constraining variable.
- The solid points indicate scores of sampling sites (P for paddy, F for forest, and U for
- upland). The squares are centroids of types of land use. The ellipses indicate the
- standard errors (n = 3) of types of land use with a 95% confidence levels. MWD is
- mean weight diameter; WSA is water-stable aggregate (WSA), PSA_i is the percentage
- of the i^{th} size fraction of soil aggregates, with i = 1, 2, ..., 6 representing the size
- of > 5 mm, $5\sim2$ mm, $2\sim1$ mm, $1\sim0.5$ mm, $0.5\sim0.25$ mm and <0.25mm.

- Fig. 3. Ordination plot of redundancy analysis (RDA) for the stability and size
- 21 distribution with soil properties as constraining variables. The solid points indicate

- scores of sampling sites (P for paddy, F for forest, and U for upland). MWD is mean
- weight diameter; WSA is water-stable aggregate (WSA), PSA_i is the percentage of the
- ith size fraction of soil aggregates, with i = 1, 2, ..., 6 representing the size of > 5 mm,
- $5\sim2$ mm, $2\sim1$ mm, $1\sim0.5$ mm, $0.5\sim0.25$ mm and <0.25mm. Al^{0xa} is the oxalate
- extractable Al oxides; Al^{DCB} and Fe^{DCB} are the DCB extractable Al and Fe oxides; SOC
- is soil organic carbon.
- 28
- Fig. 4. Diagrams describing the partitions of variation of the stability and size
- distribution of soil aggregates by soil organic carbon (SOC), and Fe and Al oxides.







