

AN INTER-HEMISPHERIC VOLCANIC TIME-MARKER IN ICE CORES FROM GREENLAND AND ANTARCTICA

by

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ABSTRACT

A strong volcanic-acid signal is clearly registered, using an acidity-measuring technique, in the A.D. 1259 ice layer in four different Greenland ice cores (Camp Century, Milcent, Crête and Dye 3). This signal is similar in amplitude to the Laki (Iceland) A.D. 1783 volcanic event as recorded in the central and south Greenland ice cores. Measurement of ice layers from corresponding age levels in Antarctic ice cores (Byrd Station, South Pole and J-9 on the Ross Ice Shelf) provides similar strong acid signals. There is no historical record of a significant volcanic eruption for the period around A.D. 1260 in the Northern Hemisphere. Subsequent chemical analyses of all A.D. 1259 ice layers show similar compositions. We suggest that the A.D. 1259 signals registered in both Greenland and Antarctica were caused by the same volcanic disturbance and that its epicenter was located at the Earth's equatorial zone, which enabled global distribution of the acid gases. These results indicate that inter-hemispheric dating of ice sheets is possible by the chemical identification of major eruptive volcanic events in the equatorial zone.

INTRODUCTION

Establishing a depth-age relationship is an important first step in the fuller examination of an ice core. Until the core is dated it is not possible to interpret meaningfully the wealth of important paleo-environmental proxy data contained in the stratigraphic layers of ice sheets. Certain limitations exist in the dating of ice cores from both Greenland and Antarctica. The difficulties are related to the varying environmental conditions which existed at the original snow-deposit sites and to the post-depositional

changes occurring there and along the flow-line trajectory leading to the drilling site. In many cases it is less difficult to date Greenland ice cores than, for instance, Antarctic ice cores (Langway and others 1985). The most successful primary methods for dating long ice cores have been the oxygen-isotope and the "acidity" measurement techniques, which are most accurate when performed in parallel (Hammer and others 1986). Both methods rely on unbroken, connected ice cores and continuous measurements, and they have been used to date ice cores back 10^3 - 10^4 years B.P.

We present here results of a successful attempt to cross-correlate a single, significant volcanic event in the thirteenth century, as a chemical signal recorded and measured from the same stratigraphic time units in both polar ice sheets.

SAMPLES

Two sets of ice-core samples were used in this study. The first set was from the Greenland locations of Camp Century, Milcent, Crête and Dye 3. The other set was from the Antarctic locations of Byrd Station, South Pole and J-9 (on the Ross Ice Shelf). Table I lists details of samples and collection sites.

The *Camp Century* ice core was augered to the 500 m depth using a thermal coring device, and thereafter to bedrock—at about the 1400 m depth—using an electro-mechanical drill (Ueda and Garfield 1968). Core drilling took place between 1961 and 1966. This was the first continuous ice-core drilling operation to reach bedrock from an inland polar-ice sheet location. Although spot sampling and physical measurements were performed on the ice core in the field, it was not dated by any method until after

TABLE I. SAMPLE AND SITE IDENTIFICATION FOR GREENLAND AND ANTARCTICA.

Site	Lat.	Long.	Elevation m above sea-level	Year core drilled	Depth 1259 event m	Annual mean temp. °C	Annual mean rate of accumulation m of ice
Camp Century	77.17°N	61.13°W	1880	1963	263.6	-24	0.38
Crête	71.12°N	37.32°W	3170	1974	220.4	-30	0.28
Milcent	70.30°N	44.58°W	2410	1973	364.5	-22	0.53
Dye 3 1971	65.18°N	43.49°W	2480	1971	363.6	-20	0.54
Dye 3 1980	65.18°N	43.49°W	2480	1980	369.9	-20	0.54
Byrd	80.02°S	119.52°W	1530	1968	97.8	-28	0.11
RISP J-9	82.38°S	168.63°W	55	1976	85.4	-27	0.10
South Pole	90.00°S		2835	1978	87.8	-49	0.08

transport back to the laboratory. The core was augered by two different experimental drilling systems, which caused several sections of the core profile to be fragmented, discontinuous or lost during drilling. The physical condition of the core compounded the difficulties of dating the ice, but it was on this core that the original continuous $\delta^{18}\text{O}$ dating method for ice cores was developed (Dansgaard and others 1971).

The Holocene ice in this core was dated by reference to seasonal $\delta^{18}\text{O}$ and dust variations (Hammer and others 1978). At selected depths, d , the mean annual layer thickness λ_y (m of ice equivalent/year) was determined by counting seasonal variations over an ice-core increment of the length dy . The age t at the depth $d = H - y$ (H is the ice thickness, y is the distance from the bottom) is:

$$t = \int_y^H \frac{dy}{\lambda_y}$$

This time-scale for the Camp Century core was later refined and fixed at certain depths, such as the historically recorded large Icelandic volcanic eruptions of Laki (A.D. 1783) and Eldja (A.D. 934), by using a solid electrical-conductivity method, ECM (Hammer 1980).

The Milcent ice core was augered to the 398 m depth during the summer of 1973, using a thermal-coring device. The core was dated only by counting annual variations in the continuously measured series of $\delta^{18}\text{O}$. The Milcent core exhibits the best example of $\delta^{18}\text{O}$ seasonals yet measured. This is mainly due to the site location's relatively high rate of accumulation (53 cm of ice equivalent per year) and relatively low mean annual surface temperature (-22°C). The high accumulation rate allows the annual $\delta^{18}\text{O}$ oscillations to survive the active diffusion process which occurs during firnification, and the low surface temperature prevents extensive melting during the summer. The ice at the bottom of the core has been dated as A.D. 1176.

The Crête ice core was augered to the 404 m depth during the summer of 1974, also using a thermal-coring device. The core was dated by counting annual variations in the continuously measured series of $\delta^{18}\text{O}$ and by a later application of the ECM method. For certain sections of the core, dust measurements and nitrate-concentration levels were used to demarcate limits of annual cycles. The Crête location is almost ideal for obtaining paleo-environmental proxy data and therefore serves as our Greenland reference core for the past 1400 years. The ice at the bottom of the core has been dated as A.D. 553.

Two ice cores were obtained at Dye 3. The first was recovered in 1971, to a depth of 372 m, using a thermal-coring device. The second core was augered to bedrock, at 2037 m, during the Greenland Ice Sheet Program (GISP) 1979–81 field operations, using a newly designed and constructed Danish drill (Gundestrup and others 1984). The 1971 core was dated by continuous $\delta^{18}\text{O}$ measurements, and the 1979–81 core was dated by continuous $\delta^{18}\text{O}$ measurements, the ECM method and dust measurements. The annual $\delta^{18}\text{O}$ cycles are not as regular as those from Milcent because penetration and refreezing of melt water offsets the advantages of the high annual rate of accumulation (53 cm of ice equivalent per year) and the low mean annual surface temperature (-20°C).

The Byrd Station ice core was augered to bedrock, at 2164 m, during the 1967–68 field seasons, using the same electro-mechanical drill as at Camp Century (Ueda and Garfield 1969). This was the first ice-core drilling to reach bedrock in Antarctica.

The South Pole ice core was augered to 100 m in 1978 and the J-9 core to 100 m during the 1975–76 Ross Ice Shelf Project (RISP) field seasons, both cores were augered by electro-mechanical drills. Dating details for these cores are discussed in the next section.

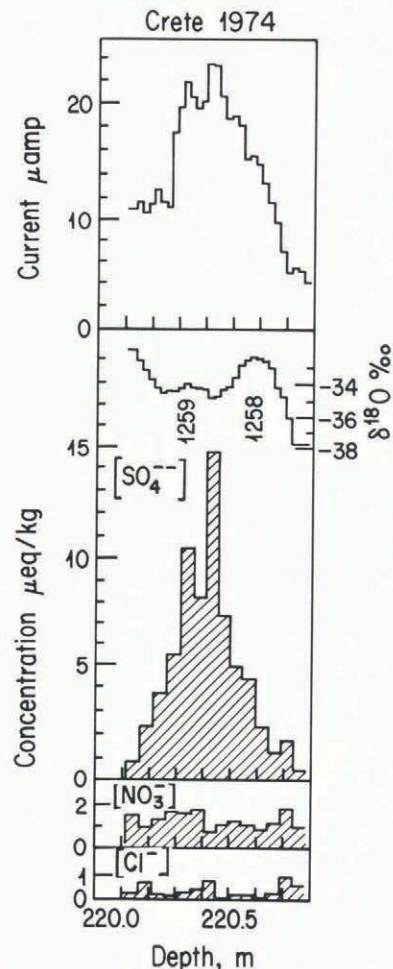
RESULTS AND DISCUSSIONS

The A.D. 1259 acid layer

Large volcanic eruptions are in general connected with the release of great amounts of strong acid gases into the atmosphere. The composition and amount of these gases vary with time and are often deposited as strong acids such

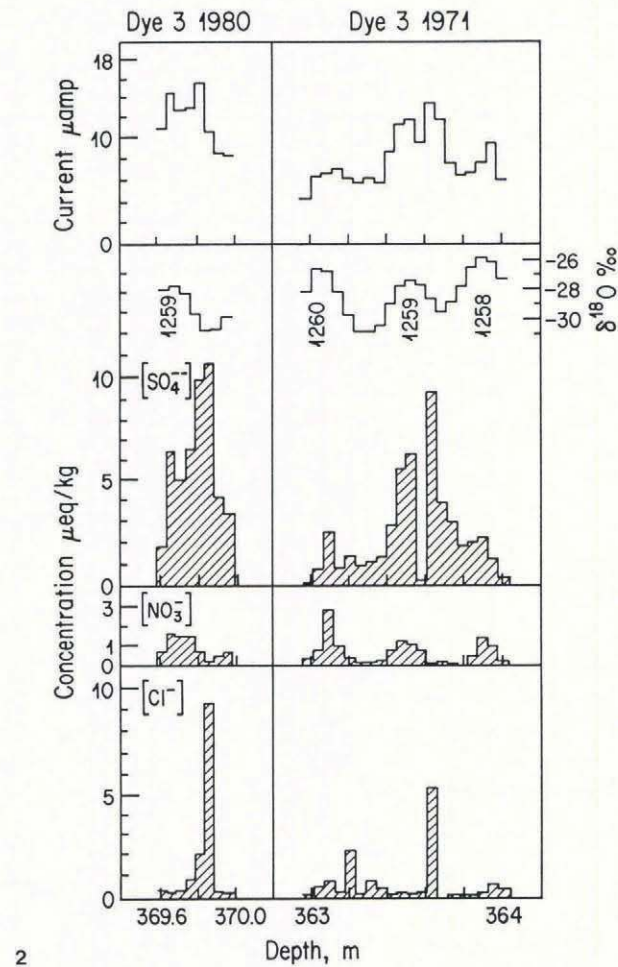
as H_2SO_4 , HCl and HF on the polar ice-sheet surfaces in an amount related to the violence and extent of the eruption. Other factors which influence the composition and amount of volcanic acids are the chemical nature of the eruption and the distance to the ice sheet. These acid layers are identified in ice cores as peaks in the electrical-conductivity records and can be detected by a solid ECM, as discussed by Hammer (1980).

Acid signals from major volcanic eruptions in Iceland and from violent eruptions south of 50°N are clearly present in snow deposits all over Greenland (Hammer 1984). In the 400 m deep Crête core, the third-largest acid signal recovered (Fig. 1) was detected in the A.D. 1258–59 layer (Hammer and others 1980). In both Dye 3 cores (Fig. 2), a large acid signal was found in the A.D. 1258–59 layers. This signal is among the largest in the Holocene Dye 3 acidity record. In the Milcent core (Fig. 3), a large acid signal was found in the layer dated A.D. 1259 by counting $\delta^{18}\text{O}$ oscillations downwards from the top. In the Camp Century core (Fig. 4), a large acid signal was detected in an ice layer dated to about A.D. 1250. This signal is among the largest found for the entire Holocene record of Camp Century.

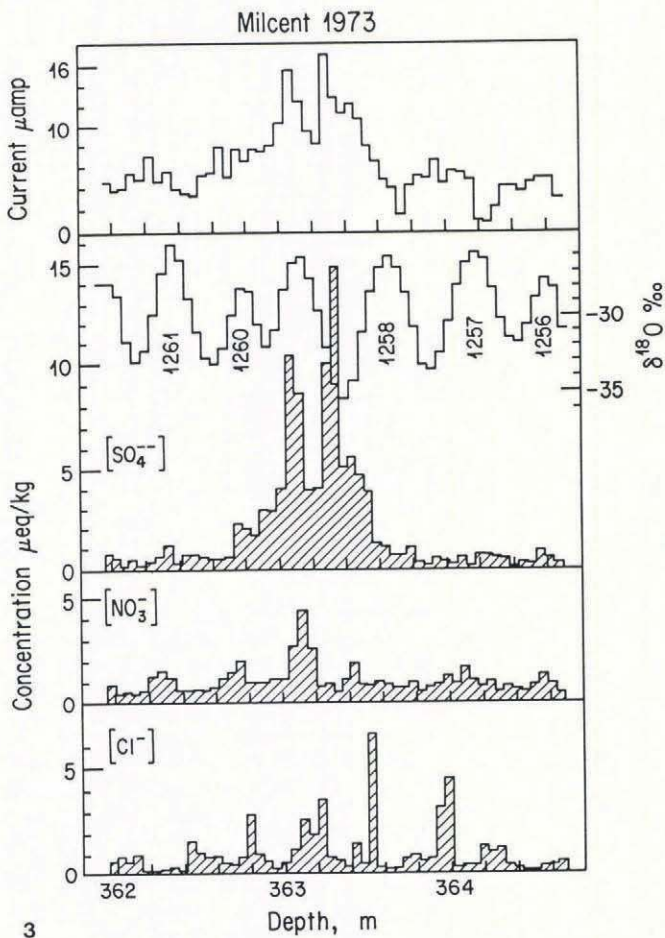


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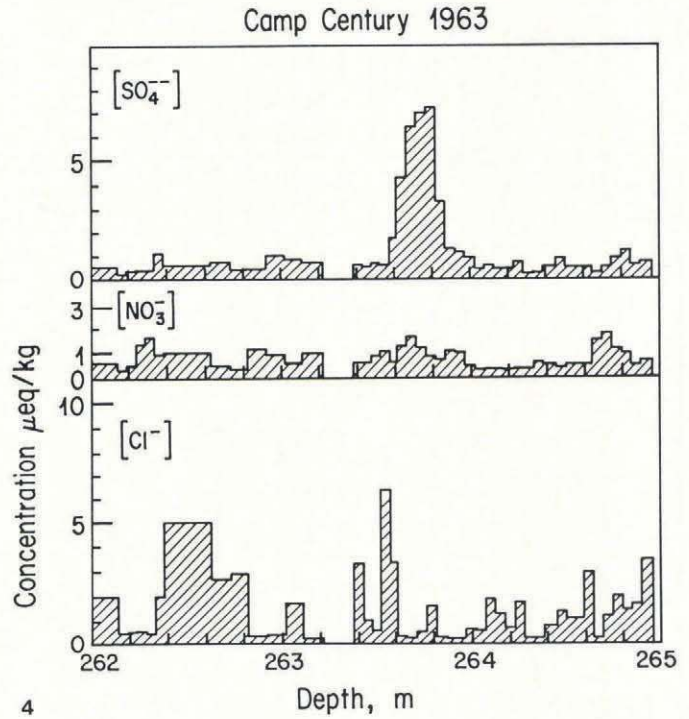
Figs. 1–9. Title at top refers to the site (Greenland and Antarctica) and the year the core was drilled. Figures 1–6 and 9 show on a depth scale in m below the surface the anion concentrations (shaded curves) of Cl^- , NO_3^- and SO_4^{2-} in $\mu\text{equivalents per kg}$ of ice ($\mu\text{eq/kg}$) from the A.D. 1259 event in the ice samples. Figures 7 and 8 exhibit the anion concentrations at South Pole for the Krakatoa (A.D. 1883) and Tambora (A.D. 1815) eruptions respectively. Figures 1–3 show the seasonal variation of $\delta^{18}\text{O}$ in per mille (scale to the right). The top curve in Figures 1–9 (except for Fig. 4, see text), shows the electrical current in μamp , determined by the ECM at -14°C . This method serves here to identify the exact position of the high volcanic-acid layer in the core.



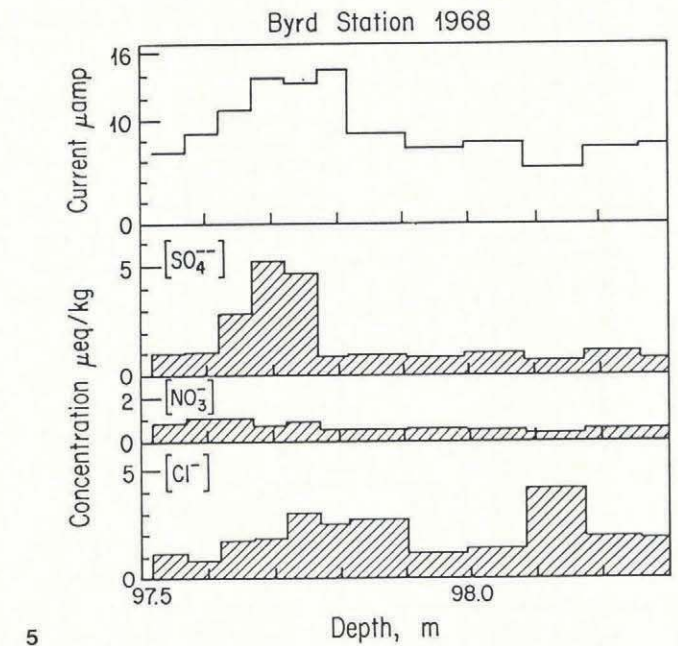
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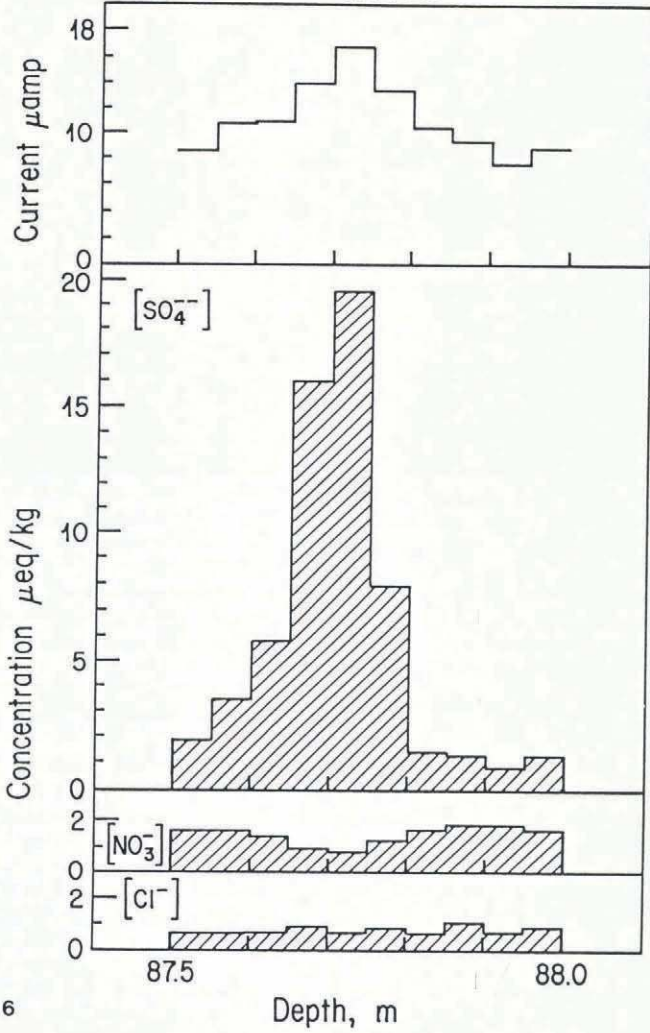
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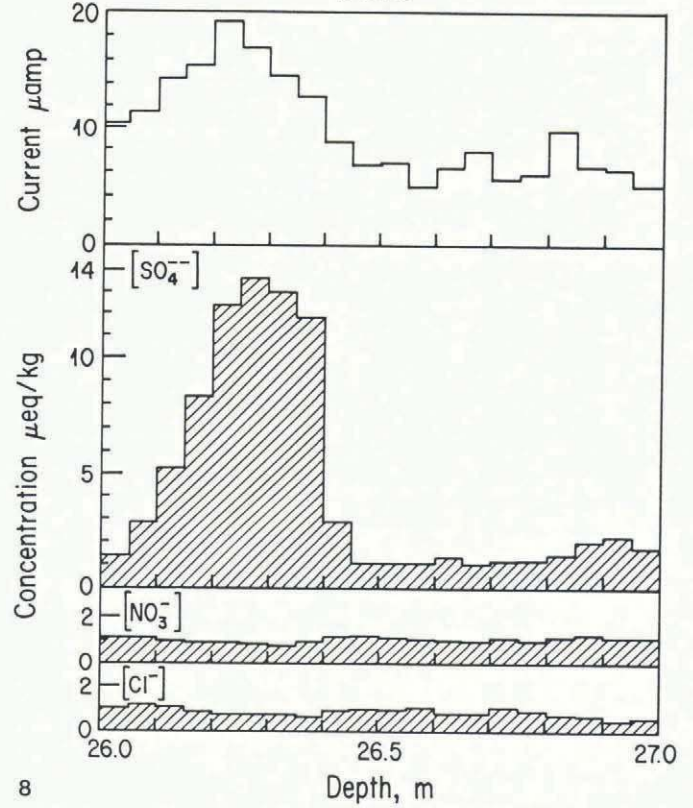
We find no record of any large volcanic eruption occurring around A.D. 1260 in any Icelandic record or other references (Simkin and others 1981). Therefore we believe that the A.D. 1259 eruption occurred farther south of 50°N. If this event took place in the equatorial zone, it would probably be recorded in the Antarctic ice. In fact, high-acidity signals were found in Antarctic ice cores from Byrd Station, South Pole and J-9 at the A.D. 1250 levels. In the *Byrd Station core*, one of the largest acid signals in the entire continuous acidity record, which covers about the last 60 ka B.P. (Hammer and others 1985), was detected at the depth of 97.8 m (Fig.5). In the *South Pole core*, a large acid signal in the continuous acidity profile was detected at the 87.8 m depth (Fig.6). Two other distinct acid signals were found at depths of 16.3 m and 26.3 m (Figs.7 and 8).

South Pole 1978
1259 AD



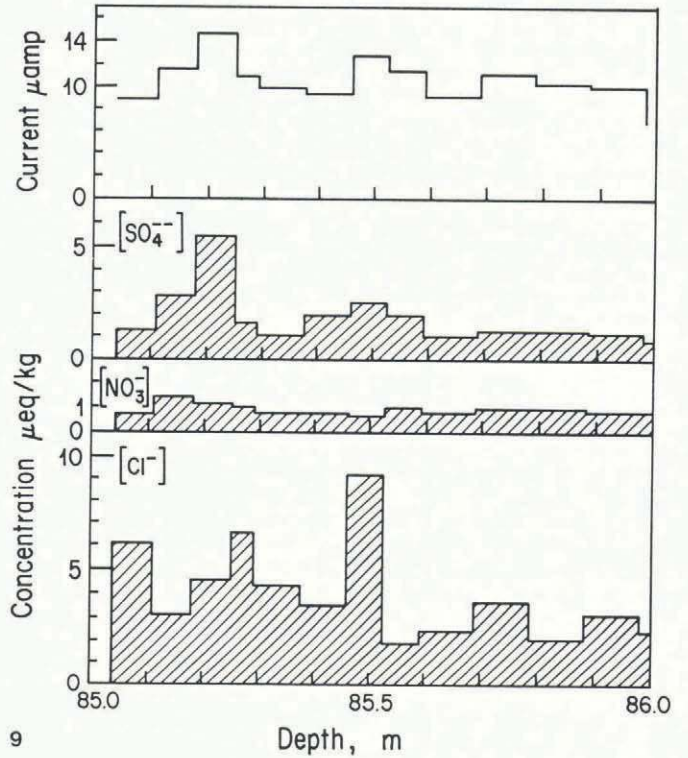
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South Pole 1978
1816 AD



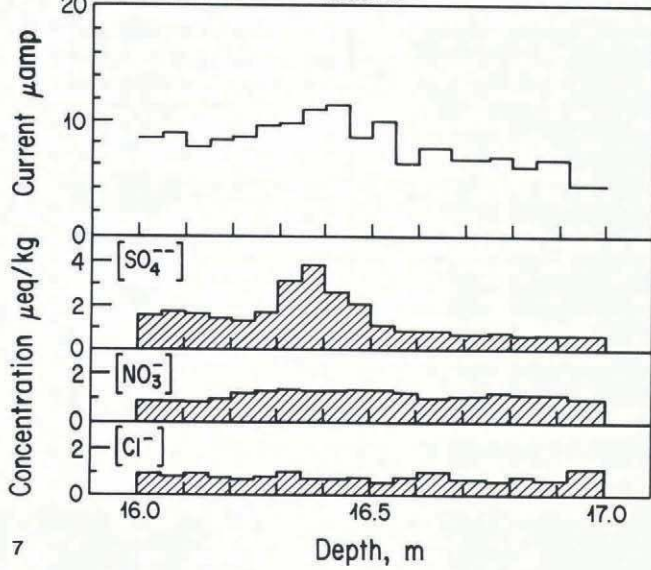
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RISP J9 1976



9

South Pole 1978
1883 AD



7

Between the depth interval 70–95 m in the *J-9 core*, the largest acidity signal measured was detected at a depth of 85.2 m (Fig.9). Due to the relatively low rate of accumulation at the Antarctic locations included in this study, seasonal $\delta^{18}\text{O}$ variations do not survive the diffusion processes during densification of the firn. Table II contains the calculated ages of the acid layers for the Antarctic

stations, assuming a uniform vertical strain-rate for all sites. Schwander (unpublished) found similar distinct acid layers at comparable depth in another ice core from the South Pole. Zanolini and others (1985) report that the largest acidity and sulfate peak found in a 200 m ice core from Terre Adélie could be due to a powerful volcanic eruption some 600 years ago. This powerful eruption is probably the

TABLE II. ACID-LAYER DEPTHS AND AGES FOR ANTARCTIC SITES.

Site	Depth of acid layer m of snow	d, Depth of acid layer m of ice equivalent		H, ice thickness m of ice equivalent	Age in years $t = \frac{H}{\lambda_H} \ln \frac{H}{H-d}$		λ_H , annual m of ice equivalent
		m of air			B.P.	A.D.	
J-9	85.2	70.1	15.1	400	786	1190	0.098
Byrd Station	97.8	78.8	19.0	2164	669	1299	0.12
					730	1238	0.11
South Pole	87.8	57.6	30.2	3000	646	1332	0.09
	26.2	13.7	12.5		153	1825	
	16.3	8.0	8.3		89	1889	

A.D. 1259 event. Fisher and Koerner (1988, this volume) found one of the largest acid signals in the Holocene acidity record of an Ellesmere Island ice core at a depth corresponding to a date of A.D. 1259.

The snow depths in Table II are converted into meters of ice equivalent, using measured density profiles from South Pole (Chiang and Langway 1978), from Byrd Station (Bender and Gow 1961) and from J-9 (Langway 1976). The annual rate of accumulation in meters of ice equivalent per year at South Pole is 0.09 m (Lambert and others 1977), at Byrd Station it is between 0.12 m (Gow and others 1972) and 0.11 m (Hammer and others, in preparation) and at J-9 it is 0.098 m (Clausen and Dansgaard 1977).

The acidity profile for the *Byrd Station core* exhibits seasonal variations thousands of years beyond the Holocene/Wisconsin transition, which is about 11 ka B.P. (Hammer and others 1985). Since the upper 88 m of the Byrd Station core are missing, we have only the calculated age for the acid layer at 97.8 m. As can be seen from Table II, a date of A.D. 1259 for this acid event agrees well with the 1238–99 range for the calculated ages.

The *South Pole* acidity profile is difficult to interpret in terms of seasonal variations, due to the low rate of accumulation at this site. From accumulation-rate studies of the Ross Ice Shelf (Clausen and Dansgaard 1977), we know it is difficult to apply the method of visible stratigraphy to dating ice cores from regions with a low annual rate of accumulation. Mosley-Thompson and Thompson (1982) dated a 104 m South Pole core (drilled in 1974) by counting microparticle oscillations, but concluded that as many as 90 years may be missing from the record, due to wind erosion and lack of precipitation. If we apply their time-scale to the three acid signals we find at the 16.3, 26.2 and 87.8 m depths, we obtain the years 1887, 1812, and 1216 respectively. Although there is a 4 year span between the drillings, and some distance between the cores, the ages we arrived at strongly indicate that these acid signals are the volcanic "footprints" of Krakatoa (A.D. 1883), Tambora (A.D. 1815) and the "A.D. 1259 event" respectively, and only about 40 years of accumulation are missing or poorly represented in the last 700 years of precipitation at the South Pole.

Using 0.098 m of ice equivalent/year for the present annual mean rate of accumulation at J-9, we calculate the 85.2 m depth as being 786 years old, or about 10% older than it should be if it represents A.D. 1259. In other words, the age may have been overestimated by 69 years, or 10%, due to changes in the accumulation rate, which would account for the difference. Since this is the largest acid signal in the 250 year depth interval we measured, we take this peak as representing A.D. 1259.

Anion composition

Increments of the ice cores at depth intervals around the acidity peaks were measured for SO_4^{2-} , NO_3^- and Cl^- concentrations by ion chromatography in the Ice Core Laboratory at the State University of New York at Buffalo. The results are shown in Figures 1–9, where the concentrations are given in μ equivalent per kg ($\mu\text{eq}/\text{kg}$). Sulfate is the dominant anion in all acid layers. The nitrate content shows an average of around 1 $\mu\text{eq}/\text{kg}$, with little

variation from year to year, even during the deposition of the volcanic acid. The chloride content of these samples shows some well-known trends of higher concentrations at coastal and lower-altitude sites (J-9, Byrd Station and Camp Century) and a decrease inland towards high-altitude sites (South Pole and Crête). These findings are consistent with other studies (Herron and others 1977, Boutron 1979, Herron and Langway 1979, Herron 1982, Rasmussen and others 1984, Finkel and others 1986, Mayewski and others 1986).

In this study, the ECM method (Hammer 1980) was merely used to identify high acid signals, and no calibration by liquid pH measurements was performed to convert the electrical currents in μamp into H^+ concentration. In Figures 1–9 (but not Fig. 4, Camp Century) the top curve shows measured currents in μamp at -14°C . For Camp Century, a very thin sample that had been stored for over 20 years was measured by ECM. Hence the current level cannot be compared directly to the current values obtained from the other cores; however, a clear increase in the current was observed in the depth interval 263.6–263.8 m. The seasonal variations in $\delta^{18}\text{O}$ are shown for the three Greenland sites also in Figures 1–3.

Volcanic H_2SO_4 deposition

On the basis of the measured sulfate-concentration levels, we calculated the deposition rate of volcanic H_2SO_4 , which is defined as the area above the general background in Figures 1–9. Table III lists the mean sulfate concentrations (column 2) over the length of ice (column 5) which represents the period of volcanic deposition, background sulfate concentrations (bg, column 3), volcanic-sulfate concentrations (column 4), deposition rates of volcanic sulfate (column 6) and deposition rates of volcanic sulphuric acid (column 7).

The H_2SO_4 deposition from the A.D. 1259 eruption was 117 and 90 kg/km^2 at Crête and South Pole (SP) respectively, and is of the same order of magnitude as given for the Tambora eruption (69 and 83 kg/km^2 respectively). For comparison, the Icelandic eruption of Laki

TABLE III. SULFATE DEPOSITION RATES.

Milcent	$\mu\text{eq}/\text{kg}$	$\mu\text{eq}/\text{kg}$	$\mu\text{eq}/\text{kg}$	m of ice	meq/m^2	H_2SO_4
1 Site	2 Mean SO_4^{2-} $\mu\text{eq}/\text{kg}$	3 SO_4^{2-} bg $\mu\text{eq}/\text{kg}$	4 SO_4^{2-} volc. $\mu\text{eq}/\text{kg}$	5 Length of event m of ice	6 Dep. SO_4^{2-} meq/m^2	7 Dep. H_2SO_4 kg/km^2
Crête	6.3	0.9	5.4	0.48	2.39	117
Dye 3 71	4.5	0.5	4.0	0.53	1.93	95
Dye 3 80	6.0	0.5	5.5	0.39	1.98	97
Milcent	6.3	0.8	5.5	0.68	3.46	170
CC	4.6	0.6	4.0	0.35	1.27	63
Byrd	4.2	0.9	3.3	0.15	0.45	22
SP 1883	2.7	1.2	1.5	0.16	0.22	11
SP 1816	8.8	1.3	7.5	0.25	1.68	83
SP 1259	11.4	1.2	10.2	0.20	1.84	90
RISP J-9	4.2	1.7	2.5	0.14	0.32	16

TABLE IV. RATIO OF DEPOSITIONS IN GREENLAND AND ANTARCTICA.

Latitude of injection	Year of injection	Greenland	Antarctica	R
75°N total β	1961-62	15.3	2.6	5.9
11°N mCi/km ²	1952, 1954	3.2	2.2	1.45
8°S H ₂ SO ₄	1815	64	83	0.77
? kg/km ²	"1259"	106	90	1.18

(A.D. 1783) deposited 137 kg/km² in the Crête region (Clausen and Hammer 1988, this volume). In the South Pole region, the Tambora eruption deposited 83 kg/km², some eight times more than that deposited by the Krakatoa eruption (11 kg/km²), which is similar to the natural background of 8 kg/km² during the time of volcanic deposition.

Latitude of 1259 eruption

To estimate the site of the A.D. 1259 eruption, we compare the H₂SO₄ deposition from known volcanic eruptions with the deposition of the total β activity from atmospheric nuclear-bomb tests. We assume a uniform global deposition pattern of volcanic-acid gases from large volcanic eruptions and of the debris produced by atmospheric nuclear tests if occurring at the same latitude. At present we know of three well-documented global events which are clearly recognized as index horizons for the entire Greenland and Antarctic ice sheets:

the Tambora eruption at 8°S in 1815,
the American bomb tests at 11°N in 1952 and 1954,
the Russian bomb tests at 75°N in 1961 and 1962 (UNSCEAR Report 1982). Table IV lists data for these events. Greenland is represented by Dye 3 and Crête, and Antarctica by South Pole, because ice cores from these regions have been analyzed both for anions from the three major volcanic eruptions ("1259", Laki and Tambora) and for total β activity from the Russian and American atmospheric bomb-test series.

The total β values are given in deposition of mCi/km², originating from a total β injection of 22.6 and 9.6 MCi (⁹⁰Sr + ¹³⁷Cs = "total β ") at 75° and 11°N respectively (Clausen and Hammer 1988, this volume). R is the ratio between the depositions in Greenland and Antarctica. In Figure 10, R is plotted on a semi-logarithmic scale as a function of the latitude of the site of injection. This figure suggests that the "1259" volcano was most likely to be located close to the Equator, probably in the Northern Hemisphere.

Total volcanic amount of H₂SO₄

Combining the deposition of the total β activity originating from the 9.6 MCi injected at low northern latitude in 1952 and 1954 (column 4, Table V) with the deposition of volcanic H₂SO₄ (column 3, Table V), we calculate in Table V the total amount of H₂SO₄ injected

TABLE V. MAGNITUDE OF THE VOLCANIC ERUPTIONS.

1 Year of max. deposit	2 Site	3 Volc. H ₂ SO ₄ kg/km ²	4 Total β mCi/km ²	5 [9.6 MCi "Total β "] × 10 ⁹	6 H ₂ SO ₄ 10 ⁶ tons
1259	Crête	117	3.36	2.86	335
	Dye 3	96	3.02	3.18	305
	Milcent	170	4.79	2.00	340
	CC	63	2.87	3.35	211
	Byrd	22	0.89	10.79	237
	J-9	16	0.58	16.55	265
	South Pole	90	2.19	4.38	394
1884-85	South Pole	11	2.19	4.38	48
1816	South Pole	83	2.19	4.38	364
	Crête	69	3.36	2.86	197
	Dye 3	59	3.02	3.18	188

into the atmosphere (column 6). Column 5 shows that the 9.6 MCi (e.g. at Crête) is equal to 2.86×10^9 times the total β deposition per km².

The total amount of H₂SO₄ in the 1259 eruption was some 300×10^6 tons, with the mean value $(298 \pm 60) \times 10^6$ and $(299 \pm 84) \times 10^6$ tons for Greenland and Antarctic sites respectively (column 6, Table V). The individual values vary by less than a factor of 2, and the coefficients of variations in the Greenland and Antarctic mean values are 20 and 30% respectively. This scatter is easily explained by drift-snow (sastrugi) noise, which in Greenland typically results in accumulation-series noise of

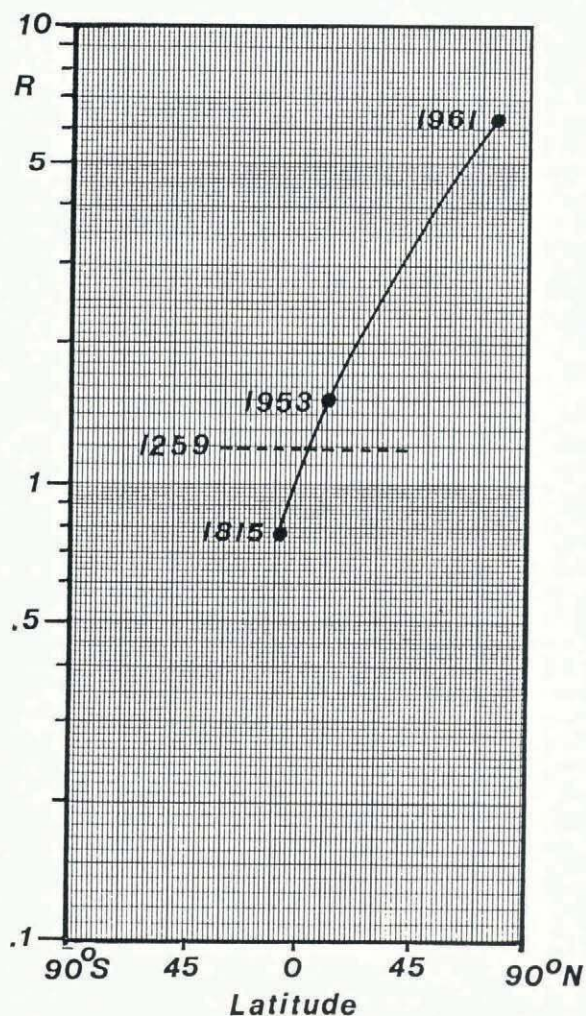


Fig.10 shows on a semi-logarithmic scale the ratio R from Table IV (between the depositions in Greenland and Antarctica) of injections (total β activity from nuclear-bomb tests and strong acids from the Tambora eruption) into the atmosphere as a function of the site of injection (degree latitude). 1961, 1953 and 1815 refer to the 1961-62, 1952 and 1954 atmospheric bomb-test series and to the 1815 Tambora eruption respectively. R of the A.D. 1259 event (dotted line) corresponds to a site of eruption close to the Equator in the Northern Hemisphere.

about 40% of the yearly value (Fisher and others 1985). In regions with a low rate of annual accumulation this kind of noise increases. The Tambora and the "1259" event are comparable in magnitude, and Krakatoa is one order of magnitude smaller. The Greenland values of the Tambora eruption are from Clausen and Hammer 1988 (this volume).

The calculation of the magnitude of the Tambora eruption at 8°S, based on the Greenland total β deposition from a low northern-latitude test site, will probably underestimate the total H_2SO_4 production, because the volcano is located south of the Equator and the total β was injected into the atmosphere north of the Equator. In a similar way the calculation based on Antarctic values has probably overestimated the production of the H_2SO_4 , suggesting a total H_2SO_4 production similar to that of the 1259 event, some 300×10^6 tons.

CONCLUSION

A major volcanic eruption in A.D. 1259, close to the Equator in the Northern Hemisphere, injected some 300×10^6 tons of sulphuric acid into the atmosphere. This volcanic eruption is one of the largest which has occurred in the last 10 000 years, of the same magnitude as the Tambora 1815 eruption, and serves as a distinct inter-hemispheric index horizon for the dating of ice cores.

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