

Assessing the oscillation of channel geometry and meander migration cardinality of Bhagirathi River, West Bengal, India

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Abstract: Lateral migration of the Bhagirathi River temporally creates unavoidable geomorphic hazards in West Bengal, India. The Bhagirathi River flows SW for ~67.30 km between the confluence point of Ajay and Jalangi rivers in East Burdwan and Nadia districts of West Bengal. The course of Bhagirathi is notably migration prone and cultivates problematic changes along its course over time. In the study, we have looked into its migration tendency and unpredictability for past 238 years and then predicted the lateral shifting of river center-line using temporal satellite imageries – Landsat-5 (TM) of 1987 (8, December), 1995 (28, January), 2005 (7, January) and LISS-IV satellite imagery (2017, 5 January); SOI Toposheet – 1968–1969 (79A/2, 79A/3, 79A/6 and 79A/7) and Rennell's map of 1779. Other highlights are the quest of fluvial features, oxbow lakes, mid-channel bars, channel migration rate, meander geometry, channel sinuosity in different parts of river course and the parts that experience intensive bank erosion. The entire river course has been subdivided into three segments; viz. reaches A, B and C. Investigation displays that degree of sinuosity decreases from its anterior course (1968) to the existing course (2017). Reach-specific outputs display that reach B is highly sinuous (SI value 1.94 in 2017) and SI increases temporarily, whereas for reaches A and C it decreases with time. The rate of migration is higher in reach B than that in reaches C and A. The study displays a notably decreasing trend of migration in comparison with its previous lateral migration and shows that the migration nature over time is intensively inconsistent and unpredictable except very few portions of the river course. The nature of deposition within the river channel shows an unstable behavior during the entire period of the study. Meander geometry depicts a rapid change of river course innate to meander bends and shows a higher rate of migration by meander loop cut-off rather than lateral migration that reflects the inconsistency, erosion vulnerability and unpredictable nature of migration. The present work offers a valuable source to comprehend channel changes in Bhagirathi River and serve as an efficacious base for river-bank migration and erosion hazard planning and management.

Keywords: channel geometry; sinuosity index; braiding index; meander geometry; radius of curvature; bend tightness index; linear regression; GIS

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1 Introduction

The Bhagirathi–Hooghly River is the western most major distributary in the Ganga delta. In the east, other distributaries are Bhairab, Sialmari, Jalangi, Mathabhanga–Churni–Ichamati rivers (commonly known as “Nadia Group of Rivers”) up to the Bangladesh border. In general, the Bhagirathi–Hooghly River is considered as the western limit of the Ganga delta (Mondal *et al.*, 2016). “*River and river processes are considered as the most significant geomorphic systems which are active on the earth’s surface. They have shaped much of the royal scenery. The earth’s surface is a dynamic surface and the forces of gravity, heat, orogeny, and isostasy is continuously modifying earth materials and forms. These changes impact the action of water flowing over the surface*” (Morisawa *et al.*, 1985). Meandering is a natural process in a floodplain and alluvial river system which causes gradual shifting of the river’s course by depositing sediment in a convex bank and at the same time eroding the concave bank of a meander. Most of the rivers on the Earth are subject to meandering due to natural and human activities and shift their course across the valley bottom (Ayman *et al.* 2009). Floodplain and braided river system are frequently affected by bank erosion and channel migration hazards and these are caused by numerous agents, *visa-a-viz.*, flow characteristics, bank material composition, climate, vegetation, geology and anthropogenic agencies (Knighton, 1998; Panda and Bandyopadyay, 2011). In agreement with this traditional view, some river scientists have paid more importance to some factors than the others. Among them, Lacey *et al.* (1939, 1958), and Luna and Thomas (1953) discovered that the bank-full discharge was more vulnerable to calculate the stable channel geometry and river width, depth and flow velocity vary with mean annual discharge. Recently, Rudra *et al.* (1981, 2008, 2011, 2014), Bandyopadhyay *et al.* (1996) and Basu *et al.* (2005) have depicted channel dynamics and meander geometry.

The present study investigates the relationship between channel geometry, lateral migration of the river course and bank erosion hazards for Bhagirathi River up to the non-tidal part from the confluence of Ajay River. Human activities along the river course including urbanization, settlement, agriculture, industries exert excessive stress on river stability and its dynamic processes. Continuous and haphazardly meandering lead to channel migration and bank erosion problem and causes significant loss of crops, land property, livestock and human lives. This work has been carried out methodically by RS and GIS techniques and will provide a blue print in the sustainable management of bank erosion and river migration hazards. The major objectives of this study are to investigate the trend of river course migration and its unpredictability over time and find out major changes of Bhagirathi River in terms of bank line positions, channel geometry and its pattern since historical past to present.

2 Study area

The Bhagirathi River take-off on the right bank of the Ganga in the district of Murshidabad and after traversing a distance of 564 km falls into the Bay of Bengal. The non-tidal part of the Bhagirathi up to Nabadwip is known as Bhagirathi, whereas the tidal part from Nabadwip to the Bay is called the Hooghly (Garrett, 1910; Willcocks, 1930; Chatterjee, 2014). The channel of Bhagirathi River within the study region is excessively dynamic and characterized by irregular meandering and lateral migration. Consequently, the adjacent area is susceptible to continuous erosion and channel migration hazards, which may cause substantial

losses of crops, land property, livestock and human lives. The right bank tributary of Bhagirathi is Ajay; which falls into the Bhagirathi near the Katwa town. This tributary causes huge sedimentation in the Bhagirathi-Hooghly River and creates the biggest navigation problem in the Kolkata port (Bose, 1972). The left bank tributary Jalangi joins the Bhagirathi opposite of Nabadwip town.

The study area lies in east-central West Bengal in East Burdwan and Nadia districts between 23°23'22.55"N to 23°40'30.92"N latitude and 88°6'34.03"E to 88°25'20.33"E longitude and comprises an area of ~1000.44 km² (Figure 1). It is a part of Bengal basin (Mukherjee,

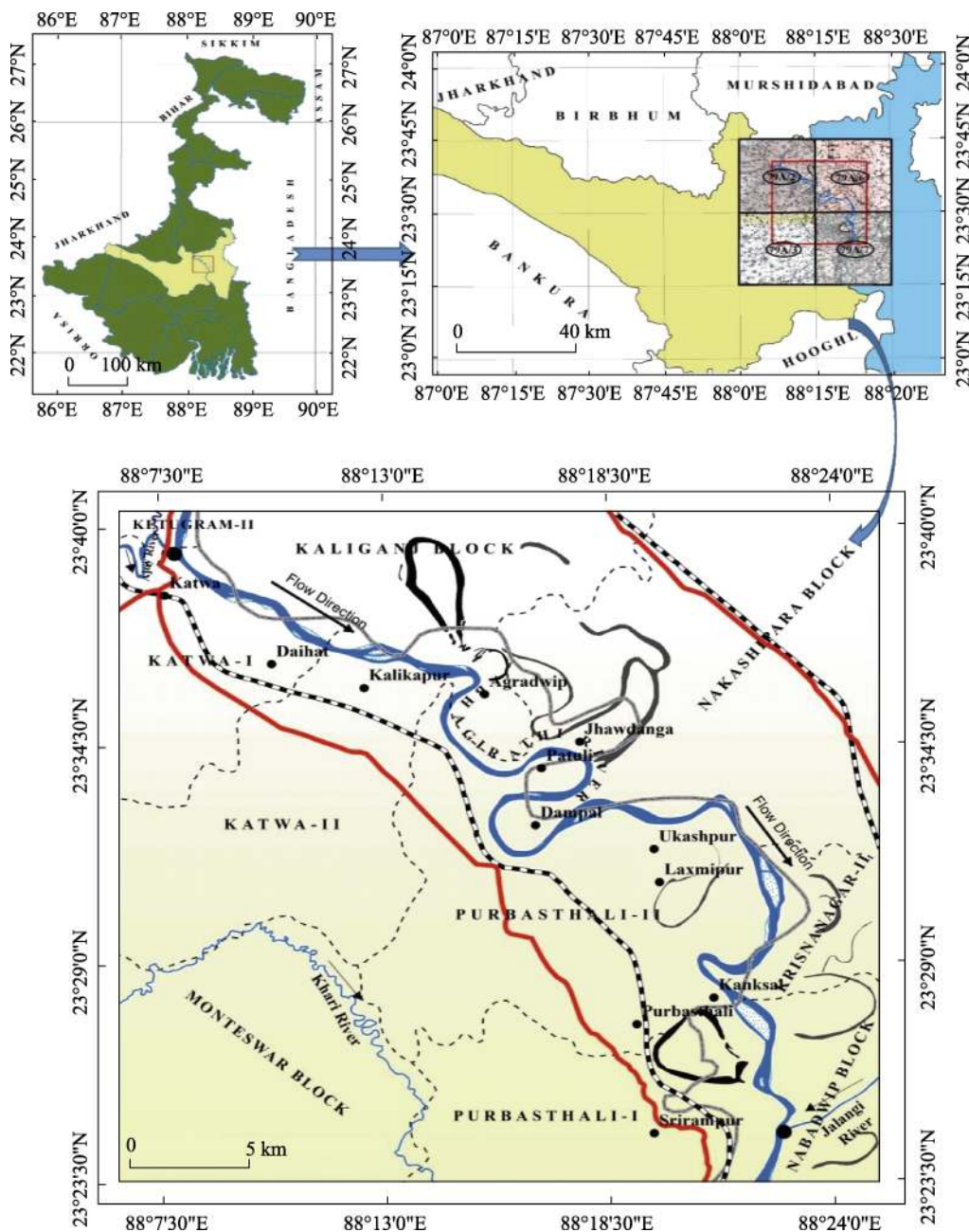


Figure 1 Location map of the study area

1938; Mukherjee, 2012), deliberated under the fluvial landscape dominated by the active and older floodplain prepared by Bhagirathi and Jalangi rivers (Bagchi *et al.*, 1944, 1972). Gentle slope, very low elevation (<5 m to 35 m), migrating river channel, oxbow lakes, natural leaves, meander scrolls, floodplain formation etc. are common features. The area is characterized by the tropical monsoon climate (Am type). Average annual rainfall is approximately 120 mm and 80% of this concentrated during the summer monsoon (June to August month). Arsenic pollution has been reported at a few places in this region (Mukherjee, 2005). The soil of this area is of new alluvial origin, brought down by the Ganga and its distributaries like Bhagirathi, Jalangi and western tributaries of Bhagirathi like Khari, Ajay. The soil type of the Bhagirathi river embankment is sandy-clay where cohesive clay layer overlies non-cohesive sandy layer, which enhances the bank erosion hazards.

3 Materials and methods

The time period for this study is 238 years, which has been separated into two, qualitative (1779 to 1968) and quantitative (1968 to 2017). All quantitative measurements have been done from 1968 to 2017. The topographical map from the Survey of India (SOI) has been considered as the base map. The river course has been identified from the orthorectified satellite imagery of Landsat-5 (TM) of 1987 (8, December), 1995 (28, January), 2005 (7, January) and LISS-IV satellite imagery (2017, 5 January) along with SOI Toposheet of 1968–1969 (79A/2, 79A/3, 79A/6 and 79A/7) and Rennell's map (1779) and digitized using ArcGIS 10.3 software and subsequently saved as a Geodatabase files (Table 1).

Table 1 Data used for the study and analysis

Data types	Description of data	Date of data (dd/mm/yy)	Spatial resolution/scale	Source
Remote sensing images	Landsat-5 (TM) imagery	08.12.1987	30 m	USGS global visualization viewer
	Landsat-5 (TM) imagery	28.01.1995	30 m	USGS global visualization viewer
	Landsat-5 (TM) imagery	07.01.2005	30 m	USGS global visualization viewer
	LISS-IV imagery	05.01.2017	5.8 m	www.nrsc.gov.in
Collateral data	Topographical maps (79A/2, 6, 3, 7)	1968–1969 (Surveyed)	1:50,000	Survey of India (SOI)
	Rennell's map	1779	1:250,000	Rennell (1788), collected from National Library, Kolkata

The UTM projection with zone 45 North and WGS-84 datum is used for all the spatial datasets. Resampling and geometric correction has been done for all spatial datasets using ERDAS IMAGINE 2014 software to bring them to the same scale. Channel centerlines were obtained from the temporal vector layer of river courses. Fourteen cross-sections have been created across different sections of the river course for measuring river width, lateral movement of bank line and channel centerline. The length of lateral migration of river course at each cross-section has been measured and migration rates for different points were calculated. For the prediction of channel centerline, the linear regression method was used. Deb and Ferrieria (2014), and Das (2016) had extremely used this method for the prediction of

bank line position. The LR model shows that among the 14 cross-sections, only five cross-sections strongly trail a linear migrating trend, which has been selected by the test of significance of r-values. Then the cross-sections wise river centerline position was measured for 2030. Lastly, for validating of predicted values of 2030, the known values of 2017 and predicted values of 2017 have been compared and standardized.

4 Results and discussion

4.1 Channel geometry

4.1.1 Channel width

The width of a channel at any point represents the cross-sectional distance of a channel and is the key indicator for showing the stage and evolution characteristics of channel. The absolute channel width of Bhagirathi River has been measured based on fourteen fixed cross-sections at different locations of the channel (Figure 2). The study period since 1968 to 2017 shows that the width of the channel gradually decreases. During the assessment

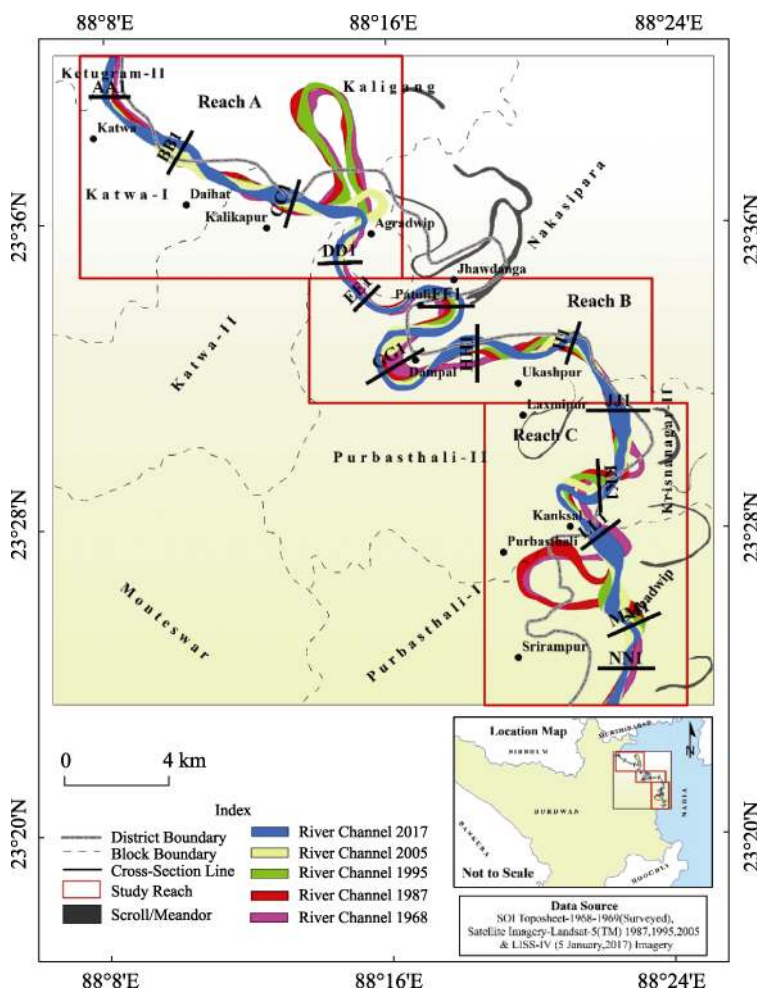


Figure 2 Showing the temporal river course with location of cross-sections and study reaches

period of 1968–1987, maximum channel widening was recorded at cross-sections AA1, BB1, HH1, JJ1 and MM1 (333.39–825.79 m) and maximum channel narrowing was recorded at cross-sections FF1, GG1 (1155.38–507.3 m) and KK1. The period 1987–1995 was documented for overall channel narrowing trend except BB1, DD1, JJ1 (589.10–742.84 m), MM1 and LL1 cross-sections. The assessment period during 1995–2005 noted channel widening trends and at cross-sections NN1, FF1, KK1 showed significant channel narrowing. Subsequently, the final assessment period (2005–2017) listed channel narrowing trend. Extreme channel narrowing trend was recorded at cross-sections BB1, CC1, DD1, MM1 and channel widening was recorded at cross-sections JJ1 and LL1. On the whole, the entire period of study (1968–2017) actually shows channel narrowing trend. Maximum channel narrowing is noted at cross-sections GG1 (1155.38–272.07 m), FF1 (1063.62–278.34 m) while widening was observed at cross-sections JJ1 (336.53–995.52 m), III, HH1 and BB1. Figure 3 shows that channel width is changing at a higher rate; which may result in maximum sedimentation, higher discharge and intensive bank erosion.

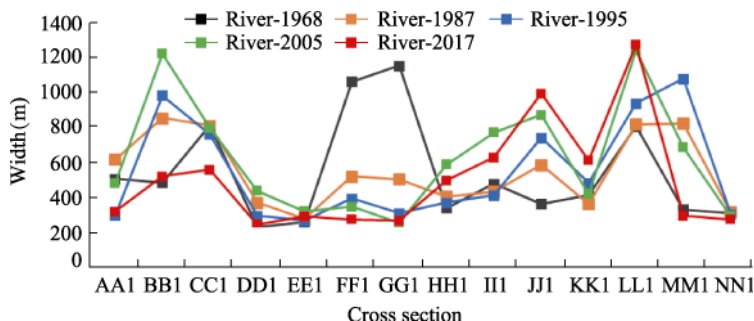


Figure 3 Temporal changes of river width at various cross-sections

4.1.2 Channel length

Temporal channel length of Bhagirathi River between the meeting point of Ajay and Jalangi Rivers has been measured along the midway between the left and right bank of the channel polygon. During the entire period of the study, the maximum channel length was recorded in 1987 (84.58 km) and minimum recorded in 2017 (67.30 km). From 1987 to 1995 channel length was reduced by 8.06 km due to neck cutoff which occurred in Reach C. Figure 4 points out that the overall length of the river course reduced by 17.18 km from 1987 to 2005 and by 12.28 km from 1987 to 2017. Also, the channel length decreased by 9.12 km during 1995–2005 due to the initiation of neck cutoff along Reach A, near Agradwip. Throughout

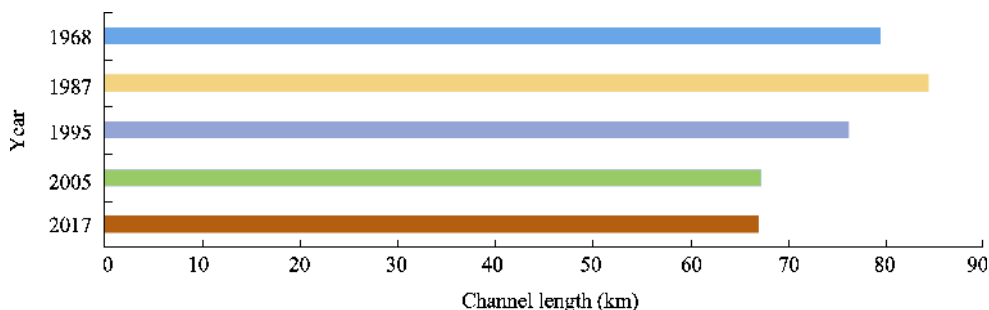


Figure 4 Temporal change of the river length from 1968 to 2017

1968 to 1987, channel length increased by 4.85 km, which highlights that the length of Bhagirathi is changing at a higher rate by the meander loop cutoff rather than lateral migration of the river course.

4.1.3 Meander geometry

Within the study area, Bhagirathi River strongly preserves a meandering course. Historical analysis of meander geometry has shown that meandering tendency of the Bhagirathi River is very complex as it varies spatially. A two-folded curve near Patuli and Dampal (Figure 5) and a three-folded curve near Purbasthali have been selected for assessment of the meander

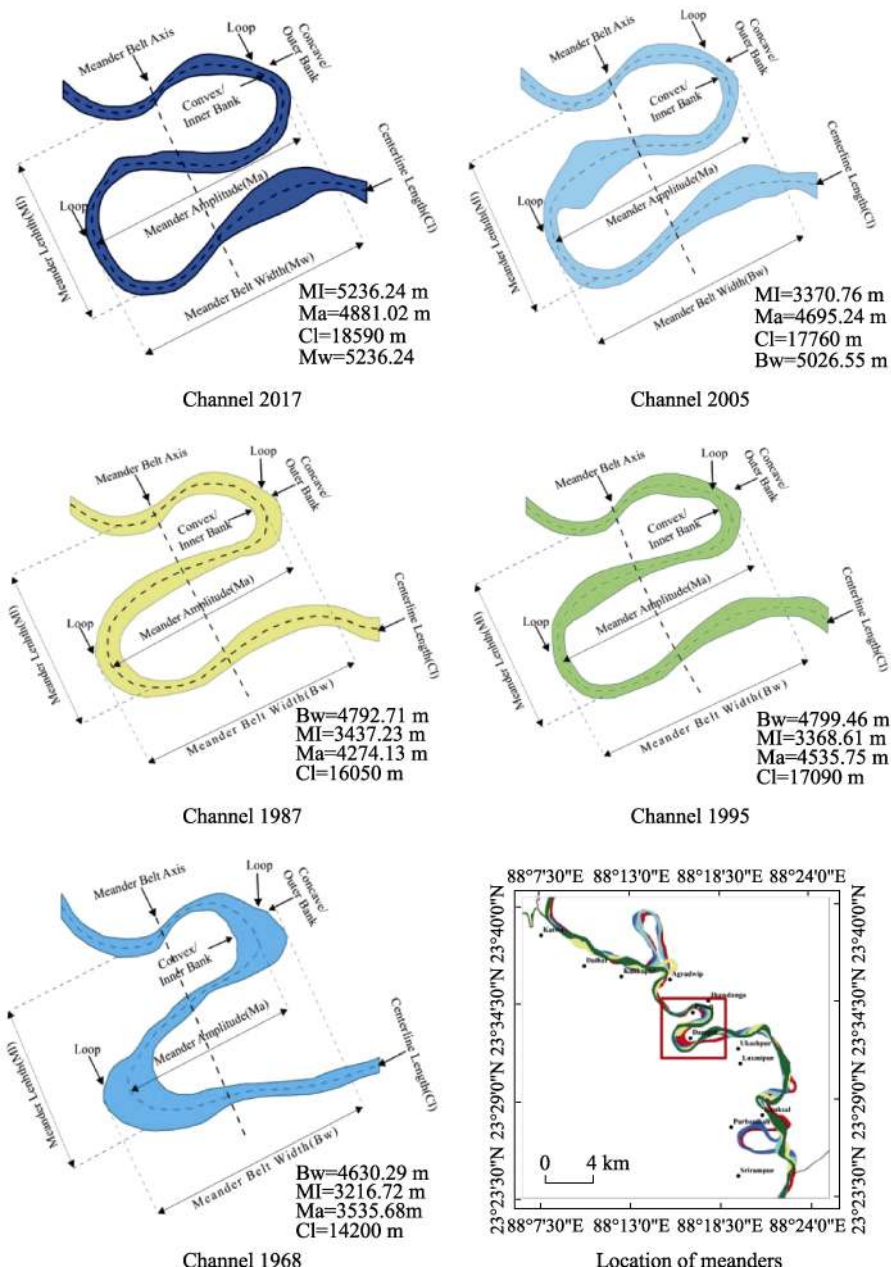


Figure 5 Temporal measurement of meander geometry near the Patuli and Dampal bends

geometry. Cautious measurement of older datasets near Purbasthali reveals an older meandering course of Bhagirathi than the present-day. This bend listed decreasing trend in its radius of curvature, meander amplitude and meander belt width from 1987 to 2017, due to neck cutoff around the Purbasthali bend (Figure 6) although the entire study period shows an increasing trend of meandering near the Patuli and Dampa bends (Figure 7).

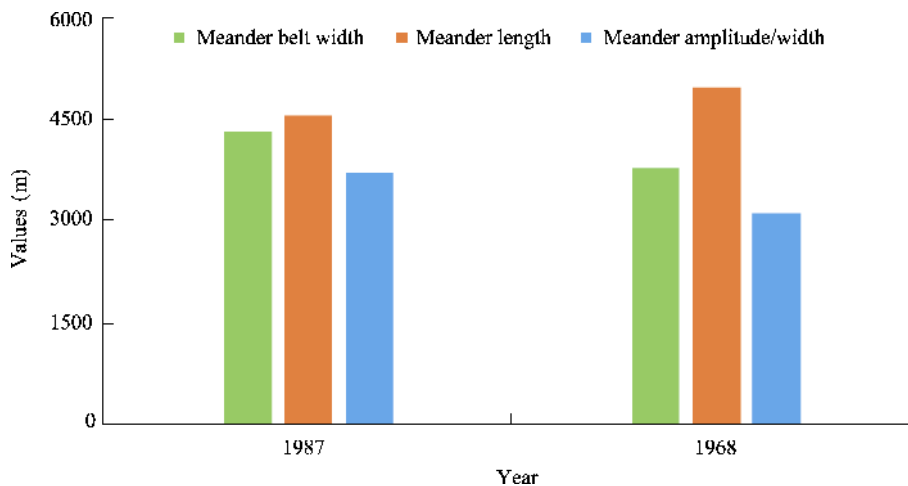


Figure 6 The temporal changes of meander geometry near the Purbasthali bend

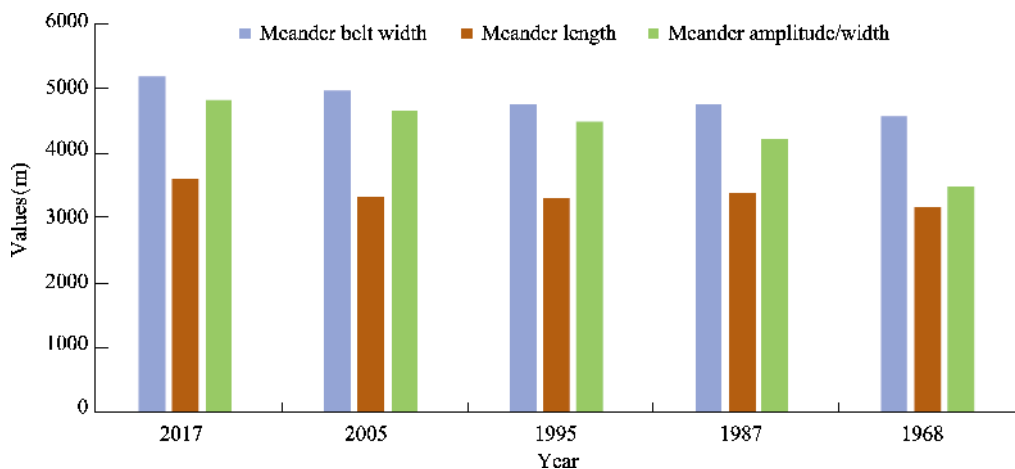


Figure 7 The temporal changes of meander geometry near the Patuli and Dampal bends

(1) Radius of curvature (R_c)

Radius of curvature (R_c) is a measure of the ‘tightness’ of a single meander bend and is inversely correlated with sinuosity. The R_c is measured by the radius of a circle that fits the meander arc from the outside of the bank-full channel to the crossing point of two lines that perpendicularly bisect the tangent lines of each meander departure point (Figures 8 and 9). The R_c near Patuli bend was 915 m in 2017 and was 597 m in 1968 (Figure 10 and Table 2) increasing by 318 m from 1968 to 2017, while near Dampal bend maximum R_c listed in 2017 was 1270 m and it increased by 650 m (Figure 11 and Table 3) from 1968 (620 m) to 2017 (1270 m). Near Purbasthali bend, the R_c increased from 1035 m to 1295 m during the period of 1968–1987 (Figure 12).

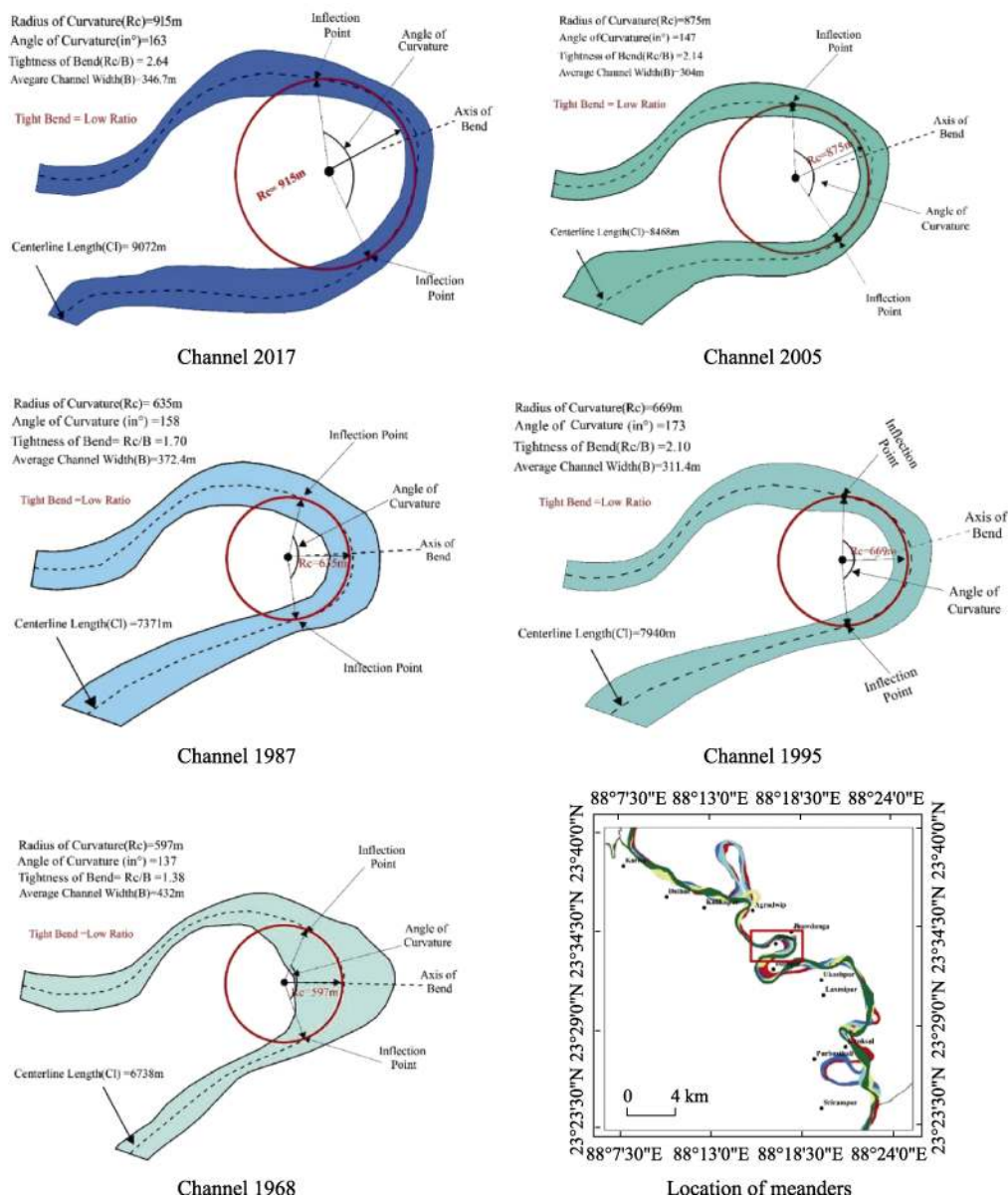


Figure 8 Temporal measurement of radius of curvature near the Patuli bend

(2) Bend Tightness Index (BTI)

The Bend Tightness Index (BTI) is a ratio between the radius of curvature and mean channel width near a bend. BTI value indicates the tightness and looseness of the bend. Higher BTI values indicate the loose bends and lower values specify the tight bends. BTI values near the Patuli bend was 1.38 in 1968 and 2.64 in 2017 (Figure 10 and Table 2), whereas, near the Dampal bend, it was 1.37 in 1968 and 4.16 in 2017 (Figure 11 and Table 3). During the period of 1968–1987, BTI values near the Purbasthali bend increased from 3.17 to 3.7 (Figure 12). Study shows that the migration rate and tightness index are positively correlated. Higher tightness values associated with higher rates of bend migration.

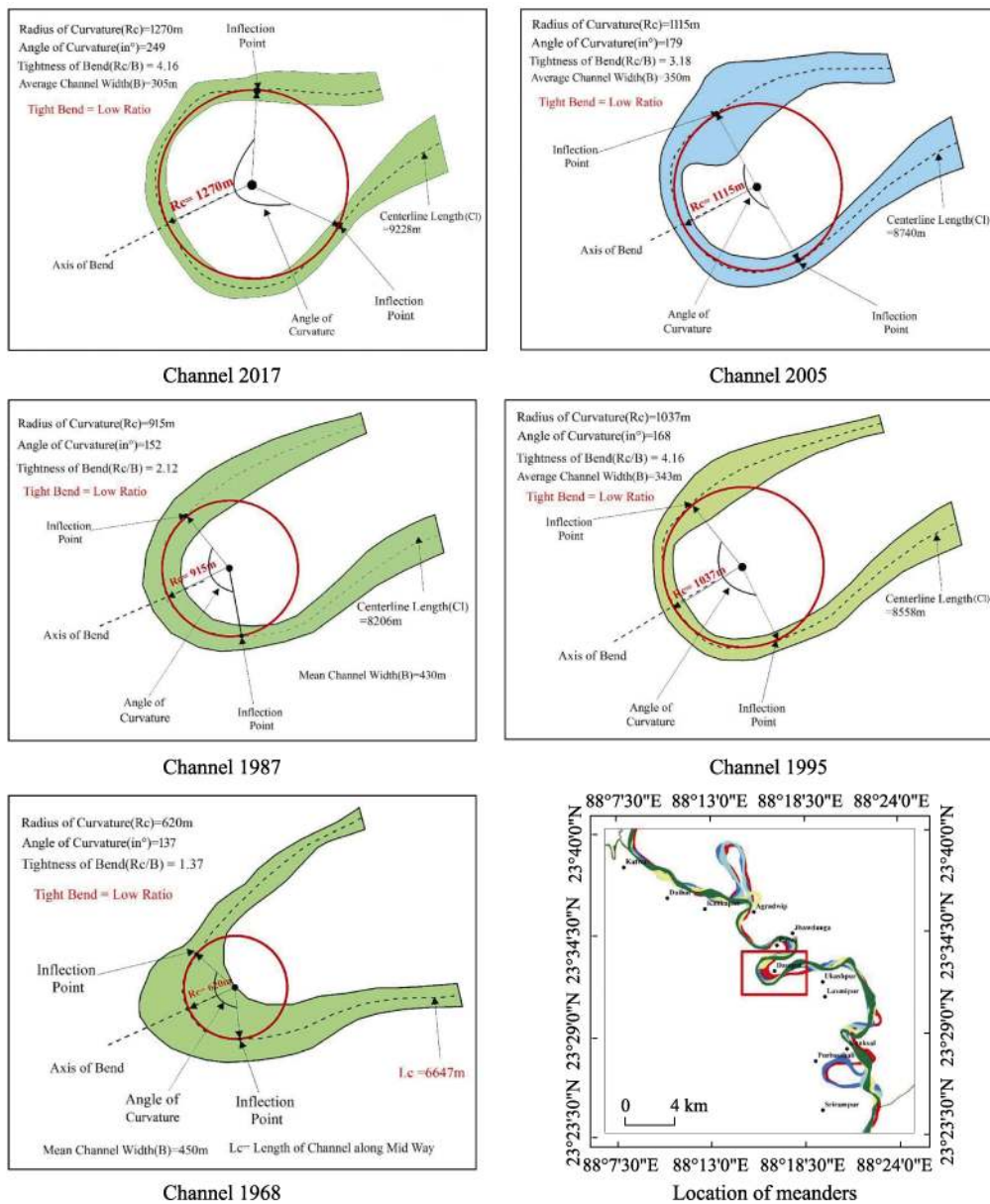


Figure 9 Temporal measurement of radius of curvature near the Dampal bends

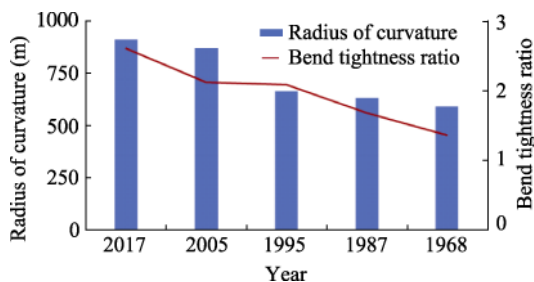


Figure 10 The temporal changes of radius of curvature and bend tightness ratio near the Patuli bend

Table 2 Radius of curvature near the Patuli bend

Year	Radius of curvature (m)	Angle of curvature (°)	Mean channel width (m)	Tightness index
2017	1250	249	305	4.16
2005	1115	179	350	3.18
1995	1037	168	343	3
1987	915	152	430	2.12
1968	620	137	450	1.37

Table 3 Radius of curvature near the Dampal bend

Year	Radius of curvature (m)	Angle of curvature (°)	Mean channel width (m)	Tightness index
2017	915	163°	346.7	2.62
2005	875	147	304	2.14
1995	669	173	312	2.1
1987	637	158	372	1.7
1968	597	137	432	1.38

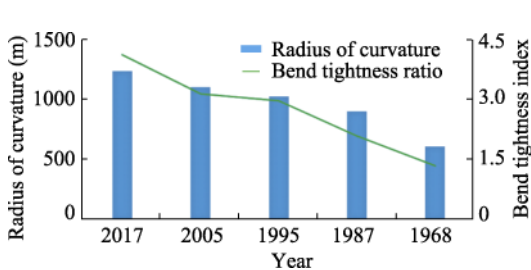


Figure 11 The temporal changes of radius of curvature and bend tightness ratio near the Dampal bend

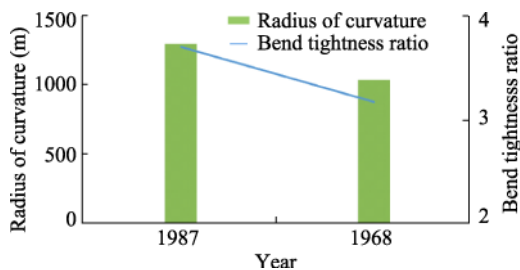


Figure 12 The temporal changes of radius of curvature and bend tightness ratio near the Purbasthali bend

4.2 Channel pattern

4.2.1 Sinuosity index (SI)

There are several methods to compute SI of a river channel (Schumm, 1963; Leopold *et al.*, 1964; Brice 1964; Leopold and Langbein, 1966; Mueller, 1968). For the present analysis, the method of Muller’s (1968) has been adopted as it can compute SI of the individual meander loops easily. Computation of SI by this method is done by the following formula

$$SI = \text{Channel Length} / \text{Air Length}$$

The meandering nature of Bhagirathi can be quantified by the temporal variation of its SI magnitude. SI ranges from 1.70 to 2.14 during the period of 1968–2017. The sinuosity value was notably higher in 1987 (2.14) and 1968 (2.02), while the lowest sinuosity value was recorded in 2017 (1.7). The general trend of sinuosity values shows that sinuosity of Bhagirathi River decreases through the entire assessment period of the study (Figure 13). These give a picture that the straight pattern of Bhagirathi River has increased gradually. In the study area, there are abundant oxbow lakes and paleo-channels that unfold migration nature of Bhagirathi River, which was previously part of the present river course. The overall sinuosity values of Bhagirathi River since past to present shows a decreasing trend, but

reach wise sinuosity values expressed a complex trend. Reaches A and C follow a reducing trend of SI, however for the reach B it follows an increasing trend from 1968 to 2017. In 2017, the SI value for reach B was 1.94 while it was 1.58 in 1968. Whereas the SI value for reach A was 2.01 in 1968 and 1.30 in 2017. For reach C, it was 1.82 in 1968 and 1.35. Figures 14 and 15 display reach wise increasing and decreasing trend of sinuosity values during the whole period of the study.

4.2.2 Braiding index (BI)

BI of Bhagirathi River within the study section has been estimated using the Brice (1964) formula to recognize the nature of sediment deposition and its consistency in the riverbed. According to Brice (1964), $BI = 2 \left(\frac{\text{Sum of the length of islands or bars}}{\text{Length of the river}} \right)$

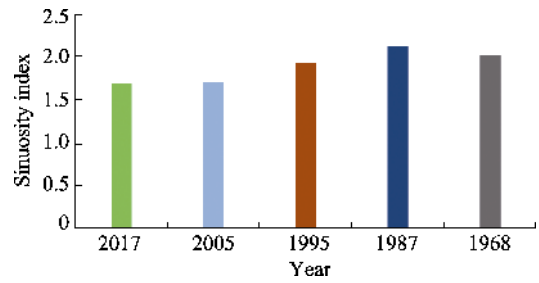


Figure 13 Temporal change of sinuosity index values of Bhagirathi River

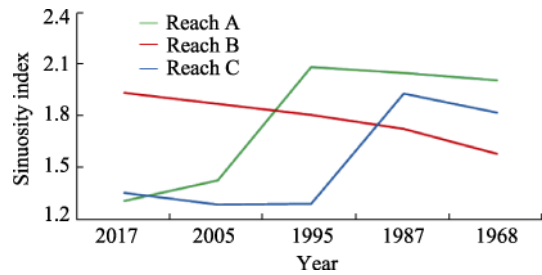


Figure 14 Reach wise temporal change of sinuosity index values from 1968 to 2017

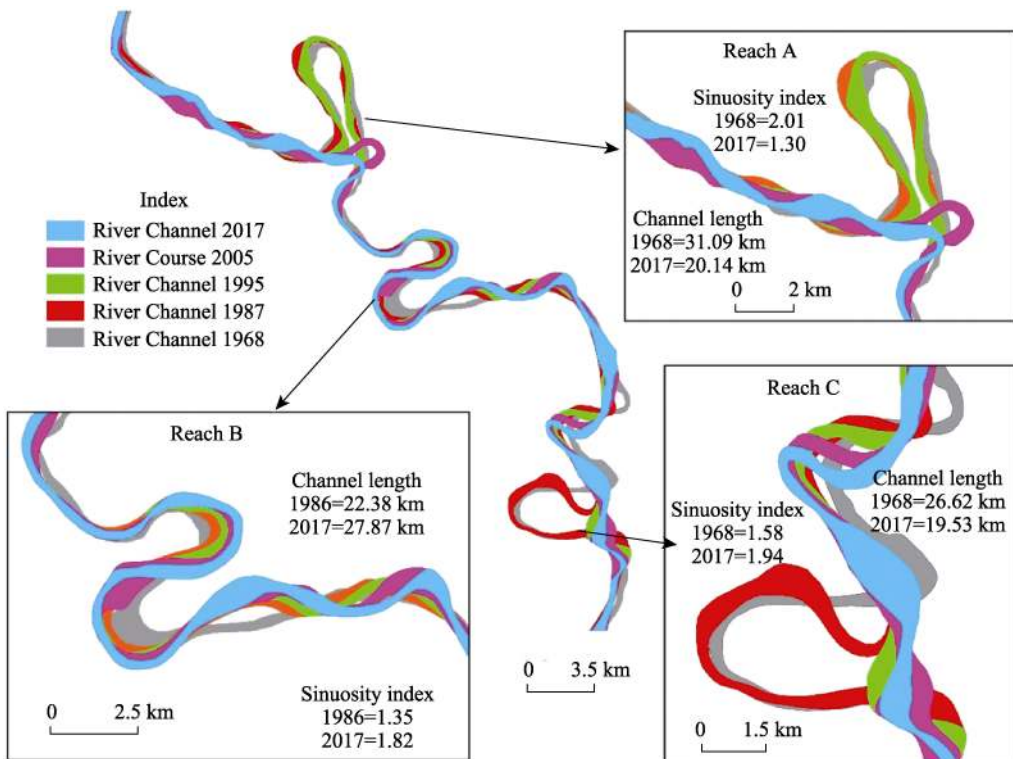


Figure 15 The reach wise change of channel length and sinuosity index values from 1968 to 2017

Figure 16 displays the sediment deposition pattern of the Bhagirathi through BI values from 1968 to 2017. BI ranges from 0.08 to 0.37. The maximum BI value has been registered in 2017(0.37) while the minimum BI value computed in 1987 (0.08). The general trend of BI values shows that the sediment load and deposition nature of Bhagirathi River have increased gradually from 1987 to 2017. Remarkably, the BI values vary reach to reach (Figure 17). Reach C recorded highest BI value (0.53) in 2017 and recorded the lowest 0.06 in 1987. Reach C follows an increasing trend of BI values from 1987 to 2017, while reach A shows a complex trend of braiding in the different years, which indicate that sediment deposition and bar formation tendencies of Bhagirathi River are highly unstable in reach A. Reach B demonstrated a more stable nature of braiding.

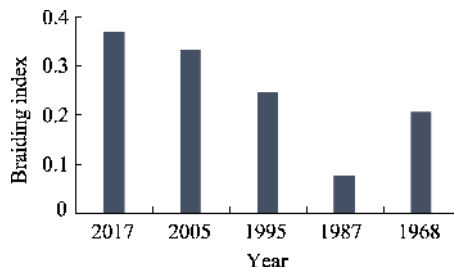


Figure 16 Temporal change of braiding index of Bhagirathi River

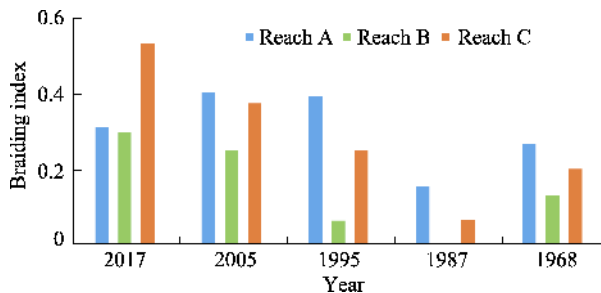


Figure 17 Reach wise temporal change of braiding index

4.3 History of temporal change of river morphometry

4.3.1 River course (18th century, 1779) in the time of Rennell’s survey

The river course has been identified and analysed from Rennell’s map (Figure 18). Rennell’s map (Rennell J J, 1781; Rudra K, 2010) shows that Bhagirathi River used to flow towards SE with a highly sinuous course (SI value 1.75). The length of the river channel along the midway between the left and right bank was 69.57 km and SI value was 1.75. Rennell’s map presented that near Agradwip, Patuli, Dampal and Purbasthali, the river course has been highly sinuous from historical time to present. These parts are very complex and dynamic in terms of meander geometry and lateral migration of bend. From Katwa to the Bagnalpur and Baladanga, river course was comparatively less sinuous (SI value 1.17) than another parts (SI value 2.14 between Bagnalpur and Patuli). This part of the river has pursued a straight to low sinuous incline still now.

Data source: Rennell’s map obtained from the book ‘Rivers of Bengal’ collected from the National Library, Kolkata. River course 1968, 1987, 1995, 2005 and 2017, identified from SOI Toposheet (1968-1969), Landsat-5 (TM) satellite imagery of 1987,1995, 2005 and LISS-IV satellite imagery of 2017.

4.3.2 Present situation (1968–2017)

1968 river course

The river course has been digitized from the SOI Toposheet of 1968–1969. During this

period, the river course has been highly meandering (SI value 2.02). The length of the river was 79.72 km. The SI and BI values for the entire river course was 2.02 and 0.21, respectively indicating a meandering river course with less sedimentation. The average width of the river was 542.37 m.

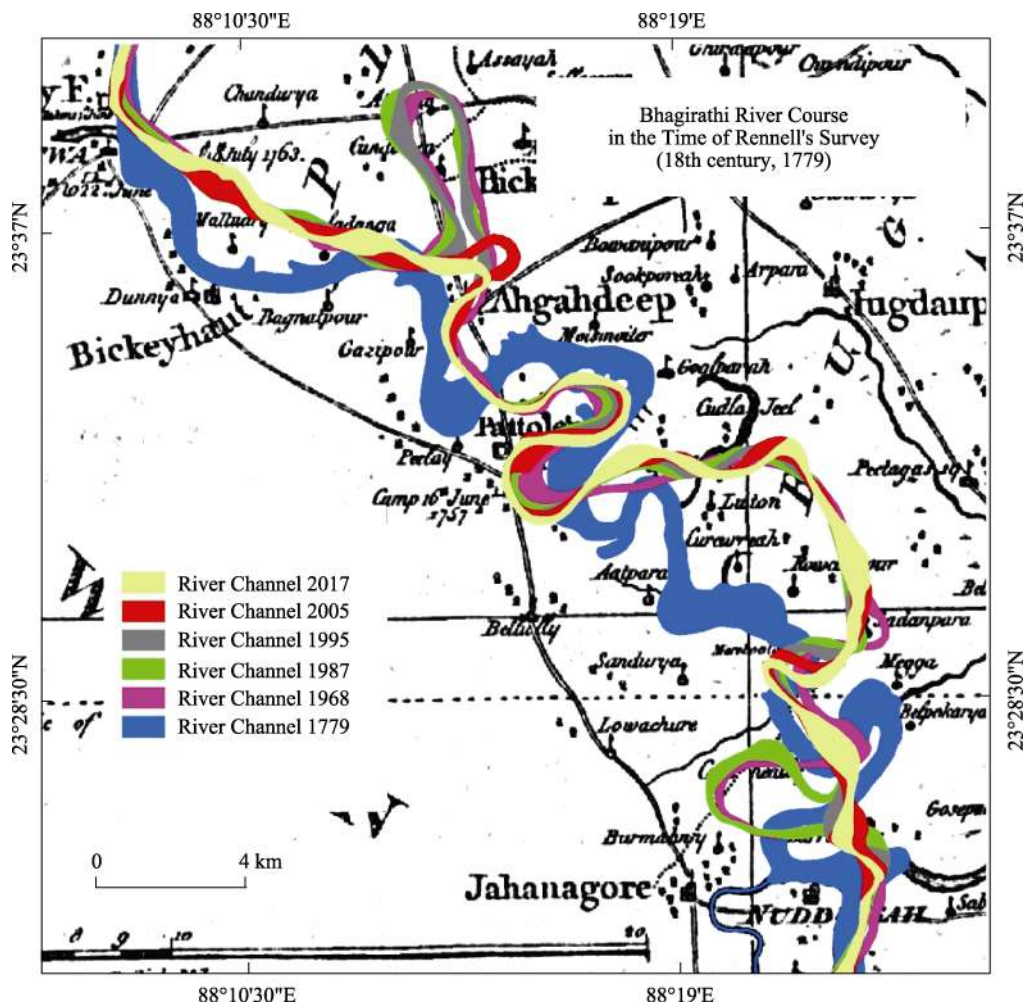


Figure 18 Historical river course of Bhagirathi River since 1779 to 2017

1987 river course

The river course has been obtained from Landsat-5 (TM) satellite image of 8 December 1987. In this year, the bends of river course have increased and the bent zone persisted for more distance (84.58 km) and the length has increased by 4.79 km from 1968 to 1987. The SI became maximum at 2.14. The BI value was 0.08 indicating that BI value has decreased from 1968 to 1987. It also indicates that the sediment load was low and lateral erosion might be high. The average width of the river was 553.45 m.

1995 river course

The river course has been digitized from orthorectified Landsat-5 (TM) satellite image of 28 January 1995. In this year, the length of rivers was noted to be 76.51 km and during 1987 to 1995, the length has decreased by 8.07 km. The SI value for this year was 1.94. The BI

value was 0.25. During the period of 1987 to 1995, a cut-off meander occurred near the Purbasthali river course, which reduced the length of the river during this period. The average width of the river was 548.71 m.

2005 river course

The river channel was identified and analyzed from Landsat-5 (TM) satellite imagery of 7 January 2005. The length of the river was 67.40 km and the length has reduced by 9.11 km during the period of 1995 to 2005, due to neck cutoff that occurred near the Agradwip bend. Similarly, SI value (1.72) has decreased in this year compared to 1995 whereas; the BI value (0.34) has increased during 1995 to 2005. The average width of river course was 627.19 m.

2017 river course

The present river channel has been obtained from the LISS-IV satellite imagery of 5 January 2017. The length of river is 67.30 km and SI value listed 1.69. Both length and SI values have decreased from 2005 to 2017. The BI in this year recorded 0.37, indicating an increasing trend from 2005 to 2017 (5 January). Only due to a small neck cutoff in the upper segment of river course near the Agradwip, the river length has decreased by 95.42 m during 2005–2017. At present, Bhagirathi River represents a less sinuous course with maximum sediment load. The average width of river course is 501.99 m.

4.4 Quantitative measurement of lateral dynamics

4.4.1 Bank line migration from 1968 to 2017

The present study exposed the actual shifting of left and right bank during the 49-year time span (Figure 19). During all the assessment periods, the maximum shifting was recorded by the left bank (Figure 20). It was 1112.78 m at cross-section KK1 towards the SW and the right bank shifted 82.14 m towards S W at cross-section MM1. Higher left bank shifting was registered at cross-sections HH1 (1048.4 m, NE), GG1 (1093.59 m, SW), LL1 (560.42 m, W) and right bank shifting found at cross-sections LL1 (1021.84 m, W), KK1 (910.54 m, SW), HH1 (988.23 m, NE), FF1 (1031.93 m, E) and CC1 (595.6 m, NE). The mean shifting rate of the left bank towards the left side and right side are 10.63 m/year and 10.71 m/year, respectively and for the right bank these are 14.28 m/year and 8.18 m/year.

4.4.2 Channel centerline migration from 1968 to 2017

Centerline is the mid-line between the left and right bank, obtained from the temporal river courses. Overall centerline migration length during 1968 to 2017 and cross-sections wise migration rate have been assessed. Chronologically superimposing of channel centerline depicts an uncertain picture of the migration pattern of Bhagirathi River (Figure 21). During the period of 1968 to 2017, the centerline shifted ~4.49 km towards E in the lower river course due to neck cutoff occurring near the Purbasthali. Whereas the period of 1995–2005, channel centerline shifted ~5.8 km towards S in the upper river course due to another cutoff occurring near the Agradwip. While centerline followed a moderate shifting rate in the middle river reach.

Figure 22 shows the cross-section wise migration of centerline during the period of 1968 to 2017. The maximum shifting of centerline was 1047.77 m towards NW and the minimum was 9.69 m towards W at cross-sections HH1 and JJ1, respectively. Whereas,

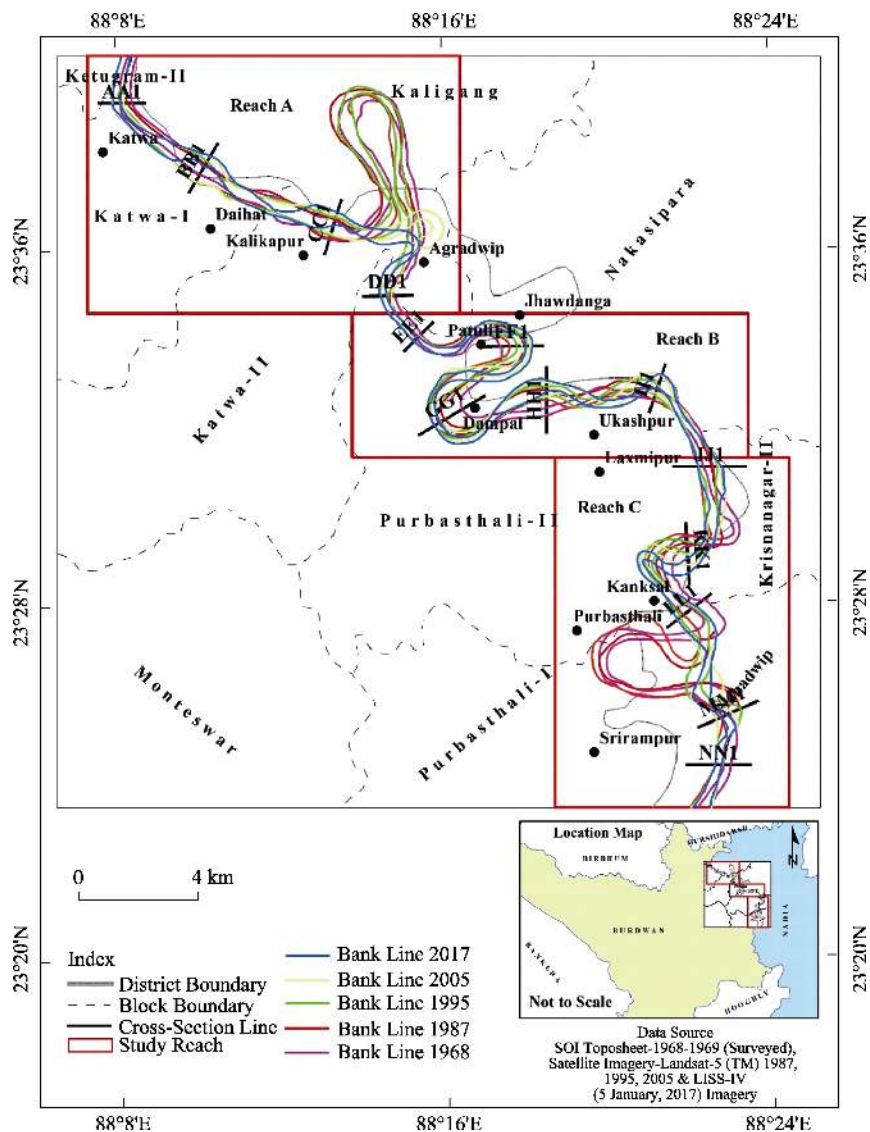


Figure 19 Measurement of bank line shifting at different section lines from 1968 to 2017

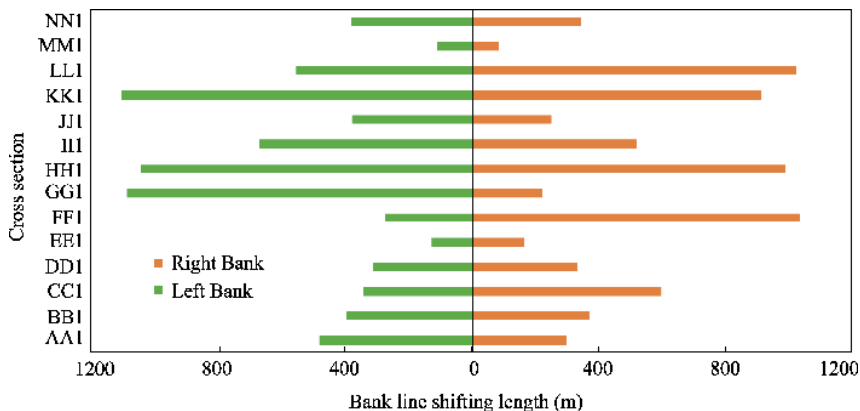


Figure 20 Actual shifting of left and right bank during 1968-2017

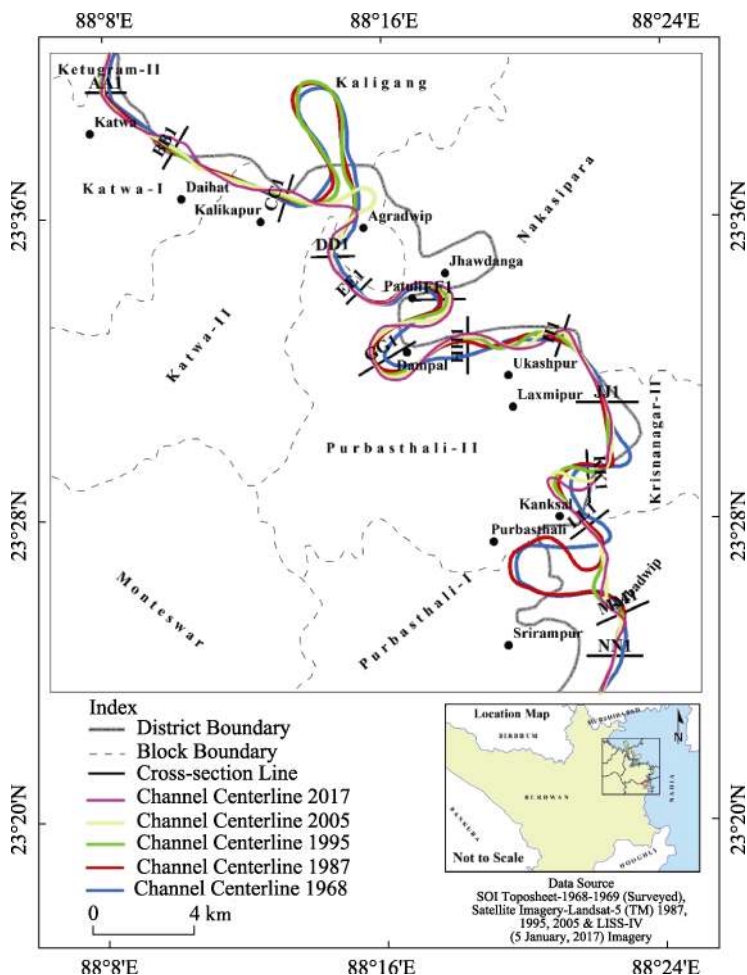


Figure 21 Measurement of channel center line shifting at different section lines from 1968 to 2017

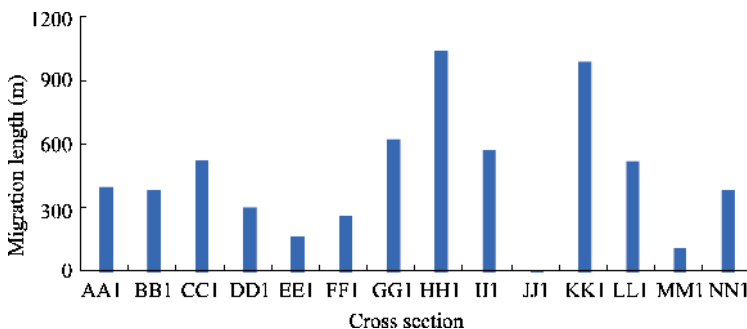


Figure 22 Actual shifting of river centerline during 1968 to 2017

cross-sections KK1 (1001.46 m, SW), LL1 (859.05 m, W), II1 (583.59 m, NE), GG1 (631.19 m, SW) and CC1 (533.5 m, NE) listed comparatively higher amount of centerline migration length during the entire time span of the study.

Figure 23 represents cross-sections wise centerline migration rate for the entire evaluation period of the study. At the cross-section HH1 near Uday Chandrapur the maximum (21.38

m/year) centerline migration rate was observed and near Rukunpur at cross-section JJ1 the minimum (0.20 m/year) shifting rate was seen. Relatively higher rate of centerline migration was recorded at cross-sections KK1 (20.44 m/year, near Gopipur), LL1 (17.53 m/year, near Kashthali Char), GG1 (12.88 m/year, near Dam Pal), III (11.91 m/year, near Chandanpur), CC1 (10.89 m/year, near Balidanga Char) and lower rate of migration was observed at cross-sections MM1 (2.51 m/year, near Mayapur), EE1 (3.53 m/year, near Simuldanda), FF1 (5.56 m/year, near Patuli) and DD1 (6.35 m/year, near Gazipur).

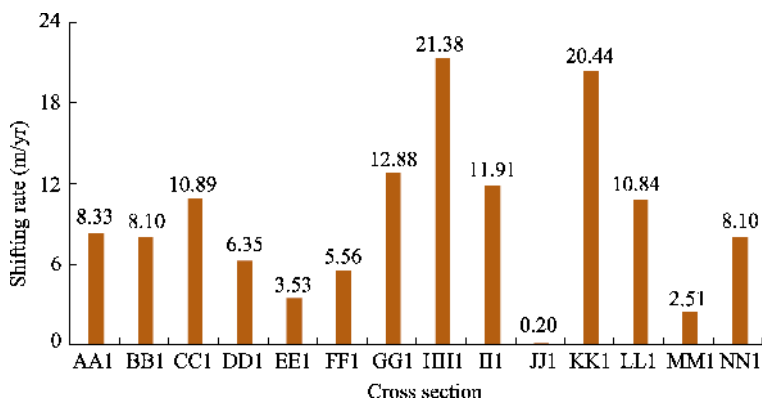


Figure 23 The cross-section wise centerline shifting rate (m/year) during the entire time period of study (1968-2017)

4.5 Channel centerline prediction using linear regression (LR) model

Migration pattern of the Bhagirathi River constantly did not retain a linear tendency. All through the assessment period of 49 years, it was observed that, cross-sections DD1, EE1, FF1 are very uneven as they change abruptly to the left and right direction and does not follow a linear migration trend. Centerline of the river is migrating linearly at the cross-sections BB1, CC1, HH1, II1, and KK1. At cross-sections AA1, GG1, JJ1, LL1, MM1, and NN1, river centerline has maintained a linear trend with moderate regression values. First cross-sections wise shifting length and direction were extracted from the base year 1968 to 1987, 1995, 2005 and 2017. Then, using the LR method, cross-sections wise regression values (r^2) were calculated for assessing the accuracy of the trend line and finally predicted values for 2017 and 2030 were generated (Figure 24 and Table 4). All through the evaluation, it could be observed that the actual values and the projected values of 2017 are exactly similar, which point out that the LR model is highly acceptable for projecting the centerline of the Bhagirathi River. Figure 25 shows a predicted centerline and it demonstrates the future Potentiality of Cutoff along the Bhagirathi River course near the Patuli, Dampal, and Kanksal bends.

5 Discussion and conclusions

This study has been attempted to analyze the lateral migration and changes of the Bhagirathi river course in details for the last 238 years using the remote sensing and GIS techniques. The previous work in this area (Pal, *et al.* 2016; Islam and Guchhait 2015) had underlined the problem of riverbank migration and its socioeconomic effects on the

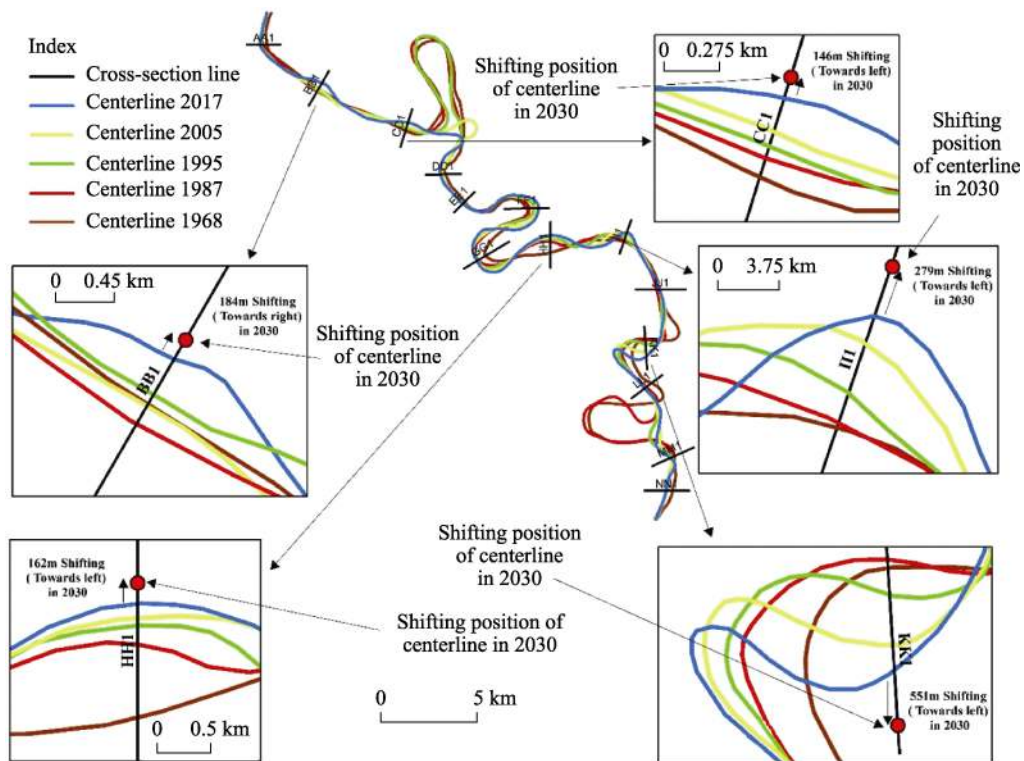


Figure 24 The future position of channel centerline of Bhagirathi in 2030 at selected cross-sections. The maximum shifting of centerline was recorded in 2030, at cross-section KK1 near the Gopipur (551 m, towards left side), while the minimum at cross-section CC1 near the Balidanga Char (146 m, towards left side). Whereas at cross-sections HH1 near the Uday Chandrapur (146 m, towards left side), BB1 near the Rajnagar (184 m, towards right side), III near the Jaganathpur and Chandanpur (279 m, towards left side) a moderate shifting trend of channel centerline was recorded.

Table 4 Cross-sections wise channel centerline shifting prediction using linear regression method, (-) value indicates the shifting to the left side and (+) value indicates the shifting to the right side. Table values represents in m.

Cross-section	1968–1987	1968–1995	1968–2005	1968–2017	Predicted values (1968–2017)	1968–2030	
						Shift (m)	R ²
BB1	172.55	112.56	15.43	-396.82	-360.8	-563.7	0.88
CC1	-126.4	-199.19	-299.57	-533.5	-522.58	-679.19	0.96
HH1	-666.78	-833.46	-901.79	-1047.77	-1054.07	-1209.75	0.96
III	-25.48	-166.68	-435.08	-583.59	-601.38	-862.15	0.97
KK1	87.67	-270.6	-744.28	-1001.46	-1020.77	-1552.64	0.97

neighboring community. However, in this study the focus has been devoted to investigating major morphometric changes of river course during the period of study. Similarly, we focused on retaining the accuracy of the study. This work is based on dry seasons images (December to January), which eliminate the river regime effects on quantitative measurement. However, due to lack of high-resolution satellite imagery up to the assessment period 2005, there may be small error in quantitative outputs and riverbank position. Yet, we used

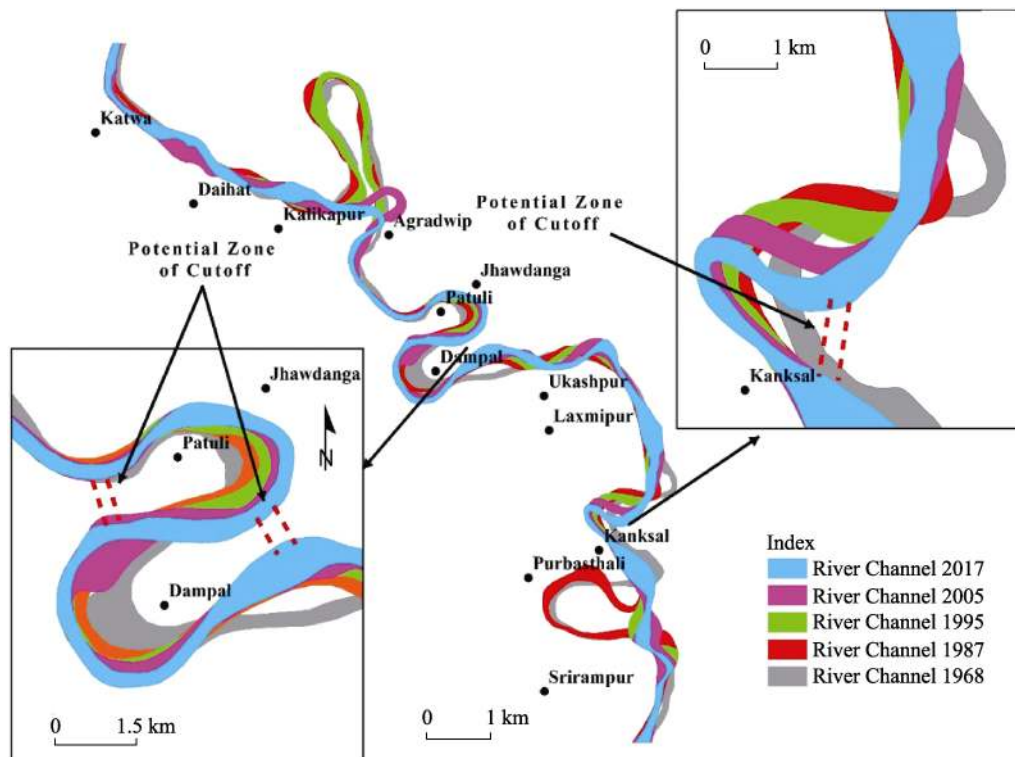


Figure 25 The meander loops of Bhagirathi River near Patuli, Dampal and Kanksal showing the potentiality of cutoff by broken red line

high-resolution LISS-IV satellite imagery for precise measurement of bank line position and channel geometry in case of final evaluation (2017). Apart from this limitation, the study follows a precise and accepted methodology, specifically for the measurement of major changes of river geometry.

In the study area, the Bhagirathi River is too dynamic and causes severe bank erosion. The river course near Patuli, Purbasthali, and Agradwip is notorious for severe bank erosion and avulsion hazards. A number of the previous studies in this area, particularly by Mukherjee (2010); Islam and Guchhait (2017) show that the fluctuation of the river discharge (through the feeder canal) due to the construction of Farakka Barrage and composition of bank materials (cohesive clay layer overlies non-cohesive sandy layer) may cause bank erosion mostly during the monsoon period. Frequently wetting and drying can contribute to cracking which reduces the shear stress of bank materials and increases its vulnerability. During the time of bank-full discharge, a huge amount of water enters into the side of bank and the return water creates pressure on the bank in post-monsoon period causing bank erosion (Mukherjee 2010). One cause that is more contemporary for bank erosion is the collection of alluvial soil from and near the river bank to provide for brick kilns.

The study shows that the Bhagirathi River shifted to both bank side and follows a linear trend of migration and also displays an irregular migration pattern. The braiding nature of the Bhagirathi is highly unstable and inconsistent. Meander geometry represents a rapid change of river course close to the meander bends and shows a higher rate of mi-

gration by meander loop cutoff rather than lateral migration. The degree of sinuosity follows a declining trend over time, whereas BI has gradually increased with time. Finally, remote sensing and GIS technology shows huge potentiality and capabilities for measurement of channel migration, erosion zone mapping and related various valuations.

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