

# Characterizing the impact of fractured caprock heterogeneity on supercritical CO<sub>2</sub> injection

Jeffrey D. Hyman, Joaquin Jiménez-Martínez, Carl W. Gable, Philip H. Stauffer, Rajesh J. Pawar

Received: date / Accepted: date

**Abstract** We present a set of multiphase flow simulations where supercritical CO<sub>2</sub> (scCO<sub>2</sub>) displaces water at hydrostatic conditions within three-dimensional discrete fracture networks that represent paths for potential leakage through caprock above CO<sub>2</sub> storage reservoirs. The simulations are performed to characterize and compare the relative impact of hydraulic and structural heterogeneity in fractured media on the initial movement of scCO<sub>2</sub> through these caprock formations. In one scenario, intrinsic fracture permeabilities are varied stochastically within a fixed network structure. In another scenario, we generate multiple independent, identically distributed network realizations with varying fracture network densities to explore a wide range of geometric and topological configurations. Analysis of the simulations indicates that network structure, specifically connectivity and the presence of hanging fractures, plays a larger role in controlling the displacement of water by scCO<sub>2</sub> than variations in local hydraulic properties. We identify active surface area of the network as a single-phase feature that could provide a lower bound on the percentage of the network surface area reached by scCO<sub>2</sub>.

J. D. Hyman  
Computational Earth Science (EES-16), Earth and Environmental Sciences Division, Los Alamos National Laboratory, Los Alamos, New Mexico, USA 87505  
E-mail: jhyman@lanl.gov

J. Jiménez-Martínez  
Department Water Resources and Drinking Water, EAWAG, 8600 Dubendorf, Switzerland  
Department of Civil, Environmental and Geomatic Engineering, ETH Zurich, 8093 Zurich, Switzerland  
Earth and Environmental Sciences Division, Los Alamos National Laboratory, Los Alamos, NM, USA 87505

C. W. Gable  
Computational Earth Science (EES-16), Earth and Environmental Sciences Division, Los Alamos National Laboratory, Los Alamos, New Mexico, USA 87505

P. H. Stauffer  
Computational Earth Science (EES-16), Earth and Environmental Sciences Division, Los Alamos National Laboratory, Los Alamos, New Mexico, USA 87505

R. J. Pawar  
Computational Earth Science (EES-16), Earth and Environmental Sciences Division, Los Alamos National Laboratory, Los Alamos, New Mexico, USA 87505

This document is the accepted manuscript version of the following article:  
Hyman, J. D., Jiménez-Martínez, J., Gable, C. W., Stauffer, P. H., & Pawar, R. J. (2020). Characterizing the impact of fractured caprock heterogeneity on supercritical CO<sub>2</sub> injection. *Transport in Porous Media*, 131(3), 935-955. <https://doi.org/10.1007/s11242-019-01372-1>

21 **Keywords** Discrete Fracture Networks · Multiphase Flow · Active Surface Area · Carbon  
22 Sequestration · Fractured Porous Media

## 23 1 Introduction

24 One of the primary risks associated with the long term sequestration of carbon dioxide ( $\text{CO}_2$ )  
25 in geological formations is leakage through the heterogeneous caprock formations that are  
26 intended to act as a barrier between storage reservoirs where supercritical  $\text{CO}_2$  ( $\text{scCO}_2$ ) is  
27 injected and drinking water aquifers residing above them [2, 3, 6, 43]. For geologic seques-  
28 tration to be effective, the caprock must remain an impermeable barrier that prevents  $\text{scCO}_2$   
29 from exiting the storage reservoir for thousands of years [46]. There are wide variety of  
30 processes that could jeopardize caprock integrity including the reactivation of faults that  
31 connect to the reservoir, changes in chemical properties due to  $\text{CO}_2$ -brine-mineral inter-  
32 actions, and changes in the stress field [42]. The initial migration of  $\text{scCO}_2$  through these  
33 systems will primarily occur within the fractures embedded in the caprock matrix, which are  
34 the primary pathways for flow and transport in the low permeability caprocks [49]. These  
35 fractures form interconnected networks where the range of relevant length scales can span  
36 multiple orders of magnitude [9]. At the smallest scale, the roughness of individual frac-  
37 tures can lead to irregular flow fields, which can result in variable precipitation/dissolution  
38 and further increase the variability of the flow field [14]. At the largest scale, the network  
39 structure and hydraulic gradient determines the overall organization of the flow field [18, 28,  
40 25]. In between these two extremes there is the effective permeability of each fracture in the  
41 network. However, across all these length scales there is a high degree of uncertainty that  
42 must be reduced for the assurance of structural containment and caprock integrity. While the  
43 interplay between these length scales is becoming better understood for non-reactive single  
44 phase flow, little is known with regard to how reactive transport and multiphase flow are  
45 influenced by properties of these fracture networks.

46 In order to characterize the relative importance of these larger length scale features  
47 (network structure and fracture permeability) on multiphase flow behavior and the initial  
48 displacement of water by  $\text{scCO}_2$  in caprock formations, we perform three-dimensional sim-  
49 ulations using a high-fidelity discrete fracture network (DFN) simulator where individual  
50 fractures are explicitly represented. The choice to use a DFN, rather than a continuum model  
51 (dual-permeability/porosity or stochastic continuum) stems from the ability to directly ob-  
52 serve how properties of the network impact the organization of the multiphase flow field.  
53 We assume that the non-fractured caprock has zero permeability, is non-reactive, and that  
54 fractures are the only pathways for fluid flow through the caprock.  $\text{scCO}_2$  is injected at a  
55 constant rate for a 10 years into each network and displaces water in the system to simu-  
56 late the initial injection and sweep through the networks. We study multiphase fluid flow  
57 behavior under two scenarios. In the first, we fix the network geometry and vary hydraulic  
58 properties and in the second, we vary network densities and connectivity.

## 59 2 Computational Methods

60 We simulate the displacement of water by  $\text{scCO}_2$  in fractured media using the high-fidelity  
61 three-dimensional discrete fracture network modeling suite `DFNWORKS` [29]. `DFNWORKS`  
62 combines the feature rejection algorithm for meshing (`FRAM`) [26] and the `LAGRIT` mesh-  
63 ing toolbox [33] to generate three-dimensional fracture networks and create a computational

64 mesh representation of the DFN in parallel; each fracture is meshed independently on a dif-  
 65 ferent processor. The set of meshed fractures is assembled into a single mesh that is used in  
 66 the flow simulations. Multiphase flow simulations are performed using the code FEHM [53],  
 67 which is integrated into DFNWORKS. FEHM is a porous media multiphase flow simulator  
 68 that can solve fully coupled heat, mass, and stress balance equations. FEHM uses an unstruc-  
 69 tured control volume finite element discretization approach based on the dual mesh of the  
 70 conforming Deluanay triangulation of the DFN generated by FRAM. The discretized balance  
 71 equations are solved using a pre-conditioned Krylov-space solver (GMRES or BCGSTAB)  
 72 with an Newton-Raphson iteration scheme. FEHM has been used for numerous studies ad-  
 73 dressing CO<sub>2</sub> sequestration [11, 15, 32, 40, 41]. DFNWORKS has been previously used to  
 74 study flow and transport in fractured media in a variety of other physical scenarios including  
 75 hydraulic fracturing operations [31, 34, 23], inversion of micro-seismicity data for charac-  
 76 terization of fracture properties [38], and the long term storage of spent civilian nuclear  
 77 fuel [22], but this study marks the first time DFNWORKS has been used in problems related  
 78 to CO<sub>2</sub> sequestration.

## 79 2.1 Flow: Supercritical CO<sub>2</sub> Simulations

80 Each multiphase fluid flow simulation begins at a hydrostatic initial condition, i.e., gravity  
 81 equilibrium of saturated water in the domain. scCO<sub>2</sub> is injected uniformly into the bottom  
 82 of the fracture network at a constant rate for 10 years to simulate the migration of scCO<sub>2</sub>  
 83 into a caprock formation via the caprock-storage reservoir interface that results from the  
 84 injection of scCO<sub>2</sub> into the storage reservoir residing below. The top of the domain is open  
 85 so both water and scCO<sub>2</sub> can exit through the outlet boundary. All other boundaries are  
 86 closed; no-flow boundary conditions are applied along all lateral boundaries. We model  
 87 the displacement of water by scCO<sub>2</sub> as two mobile phases. In this context, the governing  
 88 equations for mass conservation of water are

$$\frac{\partial}{\partial t} \phi [\rho_{\text{CO}_2} S_{\text{CO}_2} Y_w + \rho_w S_w X_w] + \nabla \cdot [Y_w \rho_{\text{CO}_2} \mathbf{v}_{\text{CO}_2} + X_w \rho_w \mathbf{v}_w] = 0 \quad (1)$$

89 where  $\phi$  is the porosity,  $S$  is the saturation,  $\rho$  is the density,  $X$  is the fraction of total mass  
 90 of given component within the water phase,  $Y$  is the fraction of the total mass of a given  
 91 component within the scCO<sub>2</sub> gas phase, and  $\mathbf{v}$  is the volumetric flux. The subscript CO<sub>2</sub>  
 92 represents the scCO<sub>2</sub> and  $w$  denotes water. We do not account for the diffusion of scCO<sub>2</sub> into  
 93 the water because our primary focus is the initial displacement of water by scCO<sub>2</sub> due to  
 94 advection and buoyancy rather than the long term behavior of the system. A straightforward  
 95 calculation of the distance traveled by scCO<sub>2</sub> under steady conditions due solely to diffusion  
 96 is 2.5 m over 10 years, which compared to the total domain size (100 m) is negligible. We  
 97 revisit this choice in the discussion and describe some of the implications.

98 Similar to (1), the conservation of mass for CO<sub>2</sub> is given by

$$\frac{\partial}{\partial t} \phi [\rho_{\text{CO}_2} S_{\text{CO}_2} Y_{\text{CO}_2} + \rho_w S_w X_{\text{CO}_2}] + \nabla \cdot [Y_{\text{CO}_2} \rho_{\text{CO}_2} \mathbf{v}_{\text{CO}_2} + X_{\text{CO}_2} \rho_w \mathbf{v}_w] = 0. \quad (2)$$

99 We also apply the following constrains in every computational cell

$$S_{\text{CO}_2} + S_w = 1 \quad (3a)$$

$$X_{\text{CO}_2} + X_w = 1 \quad (3b)$$

101

$$Y_{\text{CO}_2} + Y_w = 1. \quad (3c)$$

102 We complete the governing equations by assuming that Darcy's law applies for the momen-  
103 tum of each phase

$$\mathbf{v}_i = \frac{kk_i}{\mu_i} (\nabla P_i - \rho_i \mathbf{g}) \quad (4)$$

104 where  $k$  is the intrinsic permeability,  $k_i$  is the relative permeability,  $\mu_i$  is the viscosity,  $P$   
105 is the pressure,  $\mathbf{g}$  is the gravitational vector, and the subscript  $i = \text{CO}_2, w$ . We use a linear  
106 relative permeability relationship and set capillary pressure to zero due to limited availability  
107 of constitutive relationships within fractured media backed by experiments and/or direct  
108 numerical simulations of multiphase flow, cf. Berre et. al [5] for a discussion. We revisit the  
109 implications of adopting these conventional relationships from porous media for fractured  
110 systems in the discussion.

111 The difference in density between the two fluids at our assumed pressure and tempera-  
112 ture, scCO<sub>2</sub> is less dense than the water (brine), leads to buoyancy effects where the scCO<sub>2</sub>  
113 rises and then floats above the water. A lower density scCO<sub>2</sub> bubble sitting below higher  
114 density water leads to the water having a higher pressure at the bottom of the scCO<sub>2</sub> than  
115 the top and this difference creates a force gradient that pushes the scCO<sub>2</sub> upwards. Numer-  
116 ous field and numerical studies have shown that buoyancy forces can play a critical role in  
117 the vertical flow of scCO<sub>2</sub> [7, 10, 32, 39].

118 Each simulation is run for 3650 days (10 years). We compare simulations by measuring  
119 the portion of the DFN that is occupied by either water or scCO<sub>2</sub>. Let

$$\widehat{S}_{\text{CO}_2}(t) = \int d\Omega S_{\text{CO}_2}(t) \quad (5)$$

120 and

$$\widehat{S}_w(t) = \int d\Omega S_w(t) \quad (6)$$

121 denote the total volume of scCO<sub>2</sub> and water in the domain, denoted as  $\Omega$ , at a time  $t$ . Then,  
122 the relative mass fraction (bulk saturation) occupied by scCO<sub>2</sub> at a given time  $t$  is given by

$$S'_{\text{CO}_2}(t) = \frac{\widehat{S}_{\text{CO}_2}(t)}{\widehat{S}_{\text{CO}_2}(t) + \widehat{S}_w(t)} \quad (7)$$

123 and the relative mass fraction occupied by water is

$$S'_w(t) = \frac{\widehat{S}_w(t)}{\widehat{S}_{\text{CO}_2}(t) + \widehat{S}_w(t)} \quad (8)$$

124 Note that  $S'_w(t) = 1 - S'_{\text{CO}_2}(t)$ . We also compute the maximum percentage of the network  
125 surface area where scCO<sub>2</sub> is present over the entire simulation.

## 126 2.2 Discrete Fracture Networks

127 We construct multiple sets of fracture networks to assess how their hydraulic and structural  
 128 properties influence the behavior of multiphase fluid flow within the system. In the first set  
 129 of numerical simulations, we fix the network structure and vary the fracture permeabilities.  
 130 In the second set, we vary the network density and hold intrinsic permeability constant.

131 In the discrete fracture network modeling methodology, individual fractures are repre-  
 132 sented as planar  $N - 1$  dimensional objects embedded within an  $N$  dimensional space [5];  
 133 here  $N = 3$ . Due to the inherent uncertainty surrounding fracture attributes in the subsurface,  
 134 e.g., size, shape, aperture, and location, DFN models are constructed stochastically. Fracture  
 135 shapes, locations, and orientations are sampled from distributions whose parameters are de-  
 136 termined by geologic characterization of the formation of interest, cf. [47] for an example.  
 137 The fractures form a network embedded within an impermeable porous medium. Interac-  
 138 tions with the matrix, via diffusion and other mass and energy processes, are not included in  
 139 these simulations.

140 Each DFN is constructed in a cubic domain with sides of length 100 meters. Fracture  
 141 centers are uniformly distributed throughout the domain. During the generation stage, the  
 142 domain size is slightly increased in all directions to avoid boundary effects and ensure uni-  
 143 form density through the entire domain. The final simulation domain is truncated to a 100  
 144 meter cube once the generate phase has completed. The networks are composed of two fami-  
 145 lies of mono-disperse disc-shaped fractures that have a radius of 25 meters. The orientations  
 146 of the two fracture families are given by the Fisher distribution,

$$147 \quad f(\mathbf{x}; \mathbf{a}, \kappa) = \frac{\kappa \exp(\kappa \mathbf{a}^T \mathbf{x})}{4\pi \sinh(\kappa)}, \quad (9)$$

148 sampled using the algorithm provided by Wood [52]. In (9),  $\mathbf{a}$  is the mean direction vector  
 149 and  $\kappa \geq 0$  is the concentration parameter that determines the degree of clustering around  
 150 the mean direction. Values of  $\kappa$  approaching zero represent a uniform distribution on the  
 151 sphere while larger values generate small average deviations from the mean direction. The  
 152 first family has parameters  $\mathbf{a} = (0, 0, 1)$  and  $\kappa = 10$  and the second  $\mathbf{a} = (0, 1, 0)$  and  $\kappa = 10$ .  
 153 Thus, the mean orientation of the first family is horizontal in our coordinate systems, which  
 154 will be perpendicular to the primary direction of flow, while the second family is vertical,  
 155 which will be parallel to the primary direction of flow. These parameters create a DFN with  
 156 a block-like structure. We use uniformly sized fractures rather than a distribution of fracture  
 157 sizes to isolate the effects of variability in hydraulic and structural properties from geometric  
 158 variations induced by a range of fracture sizes.

159 An example of one DFN composed of 52 fractures is shown in Fig. 1. Fractures from  
 160 one family are colored orange, the other family is colored blue, and the conforming Delau-  
 161 nay triangulation is shown in black. The mesh is composed of 122012 triangles with 60505  
 162 nodes. This results in 60505 control volumes in the Voronoi tessellation (the dual of the  
 163 Delaunay triangulation shown here) on which FEHM performs the flow simulation. These  
 164 Voronoi control volumes are geometrically two-dimensional objects but are computationally  
 165 treated as three-dimensional volumes by FEHM where the height of the control volumes is  
 166 the aperture of the fracture. These Voronoi tessellations are in a sense optimal for two-point  
 167 flux finite volume codes [19] and allow us to use existing finite volume codes for multi-  
 168 phase flow simulations, rather than having to develop numerical methods specifically for  
 169 DFN modeling, which would otherwise be required. The mesh resolution close to fracture  
 170 intersections is 0.5 m and gradually coarsens away from the intersection using the algorithm

171 presented in Ushijima-Mwesigwa et. al., [48] where the mesh resolution is a piecewise linear  
 172 function of distance from intersections on each fracture plane. Gradients in the flow  
 173 field are higher closer to fracture intersections, which is why we refine the mesh in these  
 174 regions. We performed an initial set of simulations using a uniform mesh resolution of 0.5  
 175 m, i.e., equal sized triangles throughout the DFN, so that numerical diffusion was constant  
 176 throughout the domain, for comparison with the variable mesh. Differences in our primary  
 177 quantities of interest, relative mass fraction of the domain occupied by water/scCO<sub>2</sub>(8), were  
 178 not substantial in the comparison between mesh resolutions, although there were some local  
 179 discrepancies in the coarser mesh regions. Therefore, we selected to use a variable mesh  
 180 resolution to perform more simulations at the same computational cost.

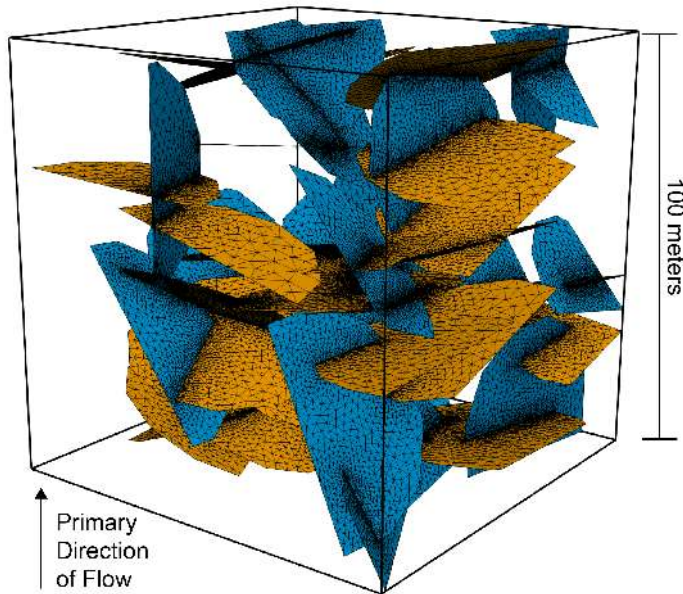


Fig. 1: A DFN composed of 52 fractures drawn from two families of fractures (orange and blue) in a 100 meter cube. The conforming Delaunay triangulation of the DFN is shown in black. In our simulations, the network is initially saturated with water that is displaced by scCO<sub>2</sub> injected into the DFN at the bottom of the domain. Thus, the primary direction of flow is from bottom to top.

### 181 2.2.1 Variable Hydraulic Properties

182 The first objective of this study is to compare the relative influence of variations in intrinsic  
 183 permeability between fractures on the displacement of water by injected scCO<sub>2</sub>. We use the  
 184 network composed of 52 fractures shown in Fig. 1. We consider one case where the intrinsic  
 185 permeability of all fractures is the same and three additional cases where the intrinsic perme-  
 186 ability varies between fractures. In the case with uniform permeability, we set  $k = 8.3 \cdot 10^{-12}$   
 187 m<sup>2</sup>. This value corresponds to a hydraulic fracture aperture of  $b = 10^{-5}$  m, assuming that the

Table 1: Static Network Information.  $\ln(\sigma_k)$ : log variance of intrinsic permeability.  $P_{32}$ : network intensity.  $d_Q$ : flow channeling density indicator.  $d_Q/P_{32}$ : portion of the total surface area where there is a significant flow.

$\ln(\sigma_k)$	$P_{32}$ [m <sup>-1</sup> ]	$d_Q$ [m <sup>-1</sup> ]	$d_Q/P_{32}$ [-]
0.1	0.13	$0.10 \pm 2.63 \cdot 10^{-3}$	$0.73 \pm 1.93 \cdot 10^{-2}$
0.5	0.13	$0.09 \pm 7.97 \cdot 10^{-3}$	$0.67 \pm 5.85 \cdot 10^{-2}$
1.0	0.13	$0.09 \pm 9.45 \cdot 10^{-3}$	$0.63 \pm 6.94 \cdot 10^{-2}$

188 hydraulic aperture and permeability are related via the cubic law  $k = b^2/12$  [51]. In the other  
 189 three cases, values of the intrinsic permeability are assigned as an independent, identically  
 190 distributed random variable sampled from a lognormal distribution. The use of a lognormal  
 191 distribution is motivated by the observation that conductivity values in many natural media  
 192 are described by a log-normal distribution [44]. We consider three log variances of intrinsic  
 193 permeability,  $\ln(\sigma_k) = 0.1, 0.5, 1.0$ , which are small to moderate levels of heterogeneity, and  
 194 hold the mean value fixed at  $8.3 \cdot 10^{-12}$  m<sup>2</sup>. When a range of fracture sizes are present, it  
 195 is typical to correlate larger fractures with wider apertures and less resistance to flow [8, 17,  
 196 24, 30, 50]. We do not consider in-fracture aperture variability in these simulations, as was  
 197 done for single phase simulations by [18, 21, 36].

198 A useful measure of fracture networks is the fracture intensity (total fracture surface  
 199 area per unit volume), which is commonly referred to as  $P_{32}$  [16],

$$200 \quad P_{32} = \frac{\sum_f S_f}{V}, \quad (10)$$

201 where  $S_f$  is the fracture surface area and  $V$  is the total size of the domain. For this network,  
 202  $P_{32} \approx 0.13$  m<sup>-1</sup>. The variations in aperture do not significantly influence  $P_{32}$ ; they all remain  
 203  $\approx 0.13$  m<sup>-1</sup>. However, the influence of these variations can be observed in the distribution  
 204 of flow at the beginning of the simulations. Recall that all simulations begin fully saturated  
 205 with water at hydrostatic conditions. We investigate how much of the domain is actively  
 206 flowing at this point using the flow channeling density indicator  $d_Q$  presented by Maillot et  
 207 al. [35]:

$$208 \quad d_Q = \frac{1}{V} \cdot \frac{(\sum_f S_f \cdot Q_f)^2}{(\sum_f S_f \cdot Q_f^2)}. \quad (11)$$

209 This definition is inspired by the participation ratio developed in solid state physics [4, 20]  
 210 and has been adapted for use in the geosciences as well [1, 12, 13]. In (11),  $Q_f$  is one-half  
 211 the absolute value of the total volume of fluid exchanged by a fracture  $f$  with its neighbors.  
 212 Comparing (10) with (11) indicates that  $d_Q$  is a measure of *active*  $P_{32}$  or *flowing*  $P_{32}$  and note  
 213 that if  $Q_f$  is uniform across all fractures  $d_Q = P_{32}$ . Thus, the flow channeling indicator is a  
 214 measure of the portion of the total surface area where there is significant flow, which can be  
 215 quantified using the ratio  $d_Q/P_{32}$ . Mean and standard deviations for  $d_Q$  and  $d_Q/P_{32}$  for the  
 216 three cases of log variances are provided in Table 1. In the homogeneous network,  $d_Q = 0.10$   
 217 m<sup>-1</sup>, so  $d_Q/P_{32} = 0.75$ , which implies that there is significant flow in 75% of the network.  
 218 Although the structure of the network does not change, the active surface area exhibits a  
 219 dependence on  $\ln(\sigma_k)$ . As  $\ln(\sigma_k)$  increases, the portion of the domain with a significant  
 220 amount of flow decreases.

### 2.2.2 Network Structure

The second objective of this study is characterizing the influence of variations in network density on the displacement of water by injected  $\text{scCO}_2$  in the fracture network. We adopt a graph-based representation of the fracture network to quantify the network structure [27] where each vertex in the graph corresponds to a fracture in the network and there is an edge between two vertices if the corresponding fractures intersect. We augment the graph to include source and target nodes that represent the inflow (bottom of the domain) and the outflow (top of the domain). For any fracture that intersects the bottom/top of the domain, an edge between the corresponding node and the source/target node is added to the graph. We focus on the distribution of fracture degrees, which is the number of other fractures that a fracture intersects, a principal topological trait of the network. In practice, we measure the degree of the corresponding node in the graph, i.e., number of edges attached to that node.

We consider thirty network realizations at three different densities with uniform fracture apertures. In the lowest density, generation of the network is stopped once 100 fractures are accepted into network. Note, that this is not the final number of fractures in the simulation because isolated fractures, those that are not in a cluster that connects inflow to outflow boundaries, are removed because they do not contribute to flow. In the second network set, generation of the network is stopped once 200 fractures are accepted into network. In the third network set, generation of the network is stopped once 400 fractures are accepted into network. The final number of fractures in each set are reported in Table 2. The intrinsic permeability of all fractures is the same and equal to  $k = 8.3 \cdot 10^{-12} \text{ m}^2$ .

The selected generation parameters lead to differences in the network density. There are a variety of definitions and measurements for network density, and we adopt the dimensionless form of density  $\chi'$  defined by Mourzenko et al. [37], which is the mean of the distribution of fracture degrees (values and standard deviations reported in Table 2). Note, however, that the reported values are for the DFN used in the flow simulations where isolated clusters of fractures have been removed. As the number of fractures increases, so does  $\chi'$ , but the variations are not drastic. Figure 2 shows the distribution of fracture degrees and highlights the disparity between the networks, even though  $\chi'$  is roughly the same. In the lowest density networks, almost all fractures have low degree and the maximum degree is seven. In contrast, the high density network is also primarily composed of fractures with low degree, but there are a significant number of fractures with higher degree. Thus, in general as the number of fractures in the domain increases, the networks are better connected. In the lowest density network 28% of fractures in the network are dead-end fractures, in the medium density, 22% of fractures are dead-ends, and in the highest density only 9% of the fractures are dead ends. Here we define the dead-ends as being fractures who are the complement of the network 2-core, which is the maximal subgraph where all nodes have degree 2 or larger [45], and is formal definition of the natural definition of dead-ends.

We also measure the local node connectivity  $n_c$  with respect to the inflow and outflow boundaries. Local node connectivity is a scalar value of a graph that corresponds to the fewest number of vertices that have to be removed so that two other vertices are disconnected. Physically, it corresponds to the fewest number of fractures that need to be removed from the DFN so that the bottom and top are disconnected and can be thought of as a measure of robustness in the network and structurally imposed channelization. In the low density networks, the mean value of  $n_c$  is close to 1, which indicates a single connected path of fractures from inflow to outflow. In the high density networks, the mean value of  $n_c$  is close to 6, thereby indicating that there are numerous pathways between inflow to outflow.



Table 2: Network Information.  $n_f$ : number of fractures.  $\chi'$ : mean of fracture degree.  $n_c$ : local node connectivity.  $P_{32}$ : network intensity.  $d_Q$ : flow channeling density indicator.  $d_Q/P_{32}$ : portion of the total surface area where there is a significant flow.

Set	$n_f$	$\chi'$	$n_c$	$P_{32}$ [m <sup>-1</sup> ]	$d_Q$ [m <sup>-1</sup> ]	$d_Q/P_{32}$ [-]
1 (lowest density)	38.23 ± 14.97	2.39 ± 0.17	1.47 ± 0.72	0.09 ± 0.04	0.05 ± 0.02	0.50 ± 0.14
2 (medium density)	118.96 ± 13.57	2.89 ± 0.13	3.03 ± 1.08	0.25 ± 0.03	0.11 ± 0.03	0.44 ± 0.09
3 (highest density)	202.00 ± 15.12	3.29 ± 0.11	5.79 ± 1.42	0.38 ± 0.02	0.18 ± 0.02	0.47 ± 0.05

268 Mean and standard deviations of  $P_{32}$ ,  $d_Q$ , and  $d_Q/P_{32}$  are also reported in Table 2.  $P_{32}$   
 269 and  $d_Q$  increase with network density. There is more variation between realizations at lower  
 270 density as well. At the lowest density, the flow tends to be more channelized due to the  
 271 lower connectivity of the networks. As the density rises the flow becomes more disperse  
 272 and homogeneous. This homogenization of the flow fields between realizations is the result  
 273 of better-connected networks. These two processes balance one another and the mean of  
 274 the ratio  $d_Q/P_{32}$  does not change significantly with density. Note that  $d_Q/P_{32}$  is a ratio of  
 275 flow rates as well, so all things are relative the network at hand. Additionally, there is more  
 276 disparity amongst the low density networks than the higher density ones.

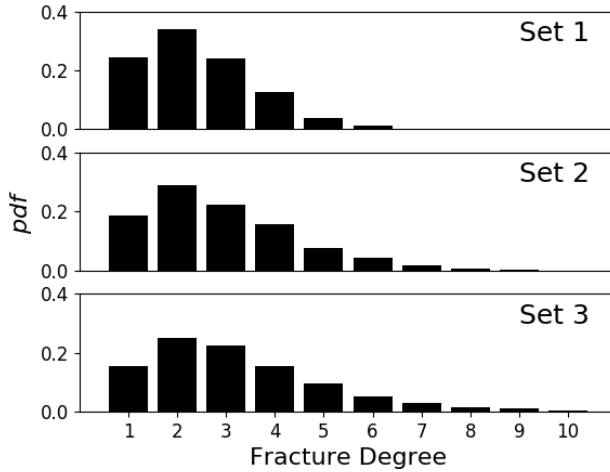


Fig. 2: Fracture degree distributions. Top: set 1, Middle: set 2, Bottom: set 3. The degree of a fracture is the number of other fractures that it intersects. In the low density networks (top) almost all fractures have low degree. The highest density networks (bottom) are also primarily composed of fractures with low degree, but there are a significant number of fractures with rather high degree. Thus, in general, as the number of fractures in the domain increases, the networks become more connected.

### 277 3 Results

278 We begin with a discussion of flow in the single network with homogeneous properties,  
 279 then proceed to the variable hydraulic properties, and conclude with simulations in multiple  
 280 networks with various densities.

#### 281 3.1 Single Network

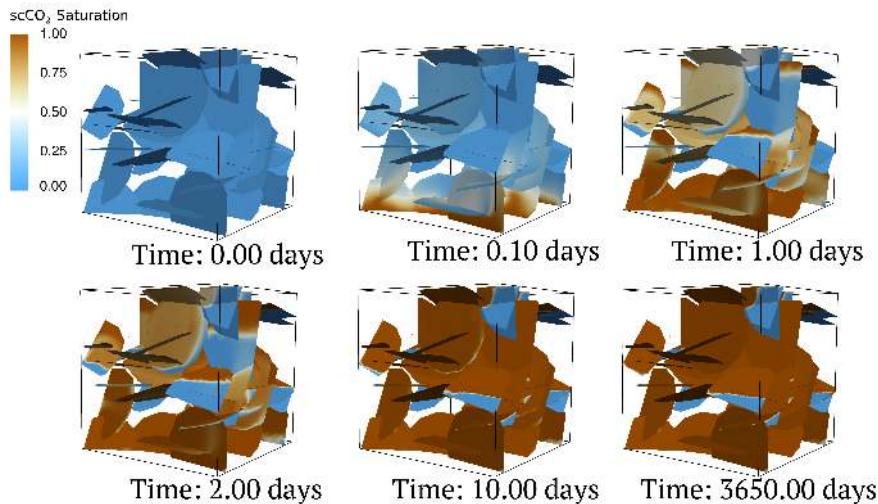


Fig. 3: Snapshots of  $\text{scCO}_2$  displacing water in a DFN with constant permeability.  $\text{scCO}_2$  is injected at a constant rate into the domain through fractures connected to the bottom. The  $\text{scCO}_2$  rises through the network displacing water. Quasi-steady state conditions are reached after  $\approx 10$  days in the 10 year simulation. At this point, the majority of the domain is fully saturated with  $\text{scCO}_2$  but there are some regions (hanging fractures) that remain fully saturated with water.

282 Six images of the simulation results at different times in the fixed geometry with uniform  
 283 permeability are shown in Fig. 3. The DFN is colored by  $\text{scCO}_2$  saturation which ranges  
 284 between  $[0,1]$ . Light blue indicates a region is fully saturated with water and brown is fully  
 285 saturated with  $\text{scCO}_2$ . The initial condition, fully saturated with water, is shown in the top  
 286 left subfigure. After  $\approx 0.1$  days (top middle subfigure), the  $\text{scCO}_2$  being injected into the  
 287 domain through fractures connected to the bottom becomes visible and starts to displace  
 288 water. At  $\approx 1$  days (top right subfigure),  $\text{scCO}_2$  has entered more of the domain, in part due  
 289 to buoyancy and in part due to the constant injection at the bottom. At  $\approx 2$  days (bottom left  
 290 subfigure)  $\text{scCO}_2$  has reached the top of the domain and nearly fully saturated the bottom  
 291 of the domain. The system has reached a quasi steady-state by  $\approx 10$  days (bottom middle  
 292 subfigure); there is little change between this time and the end of the simulation (bottom  
 293 right subfigure). Nearly the entire DFN is fully saturated with  $\text{scCO}_2$ . However, there are

294 some regions where the scCO<sub>2</sub> has not displaced the water. These regions tend to be below  
 295 (with respect to the  $z$  axis) lines of intersection with a fracture.

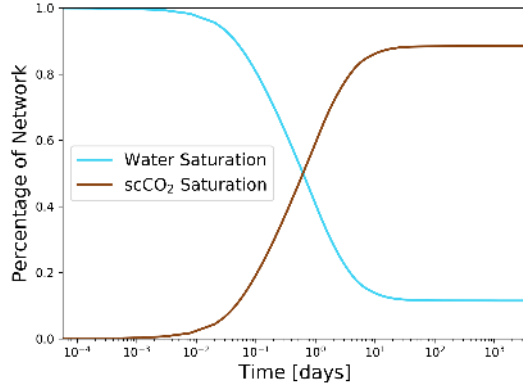


Fig. 4: Bulk saturation in the network. scCO<sub>2</sub> (brown) displacing water (blue) in the DFN shown in Fig. 3. At the end of the simulation, scCO<sub>2</sub> occupies 89% of the DFN.

296 Figure 4 reports the bulk saturation in the network, (7) and (8), for the simulation shown  
 297 in Fig. 3 (blue-water, and brown-scCO<sub>2</sub>). After  $\approx 1$  day, there is more scCO<sub>2</sub> in the DFN  
 298 than water. After  $\approx 10$  days, the system has reached a quasi-steady state, indicated by the  
 299 stabilization of the saturation volumes. Notice that the domain does not become fully satu-  
 300 rated with scCO<sub>2</sub>. Rather, only 89% of the DFN is filled with scCO<sub>2</sub> with water remaining  
 301 in the rest of the domain. The regions that are not flushed by scCO<sub>2</sub> are *hanging fractures*,  
 302 fractures whose intersections with others in the domain are at a higher elevation than other  
 303 portions of the fracture. We explore this phenomenon in the following subsection using a  
 304 simple network.

### 305 3.1.1 Hanging Fractures

306 To highlight the influence of hanging fractures on sweep efficiency, the volume of water  
 307 pushed out of the network or equivalently final percentage of the network occupied by  
 308 scCO<sub>2</sub>, we construct the five fracture network shown in Fig. 5. All fractures have the same  
 309 permeability ( $k = 8.3 \cdot 10^{-12} \text{ m}^2$ ). The network has one primary fracture (labeled-1) con-  
 310 necting the bottom of the domain (inflow) to the top of the domain (outflow). Fracture 1 is  
 311 intersected by two fractures (labeled 2 and 4) that intersect two additional fractures (labeled  
 312 3 and 5). In this network, portions of fractures 3, 4, and 5 are hanging fractures. Note, as  
 313 well, that all of the secondary fractures are *dead-end* subnetworks, they do not connect back  
 314 to the primary network.

315 We perform the same numerical simulations in this network as in the more complicated  
 316 ones to illustrate the interplay between dead-end fractures, hanging fractures, and the move-  
 317 ment of scCO<sub>2</sub>. A snapshot at an early time in the simulation is shown in the left sub-figure.  
 318 scCO<sub>2</sub> is injected at the bottom of fracture 1 and colors are the same as in Fig. 3. Water in  
 319 the primary fracture is quickly displaced. The middle sub-figure shows a later time in the

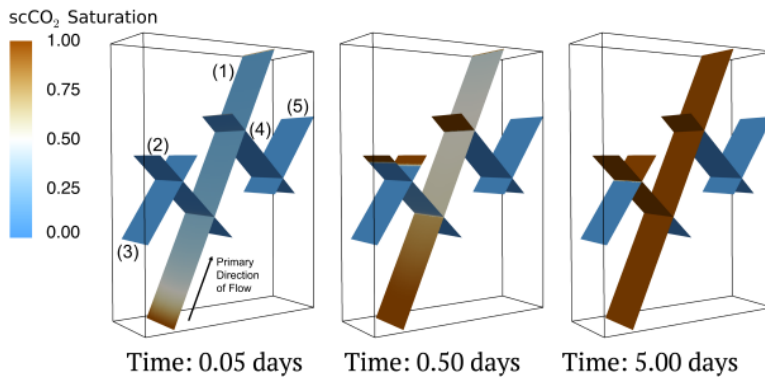


Fig. 5: Simulation snapshots of  $\text{scCO}_2$  displacing water in a DFN composed of 5 fractures. The primary fracture (1) is connected to two sets of hanging fractures (2-3) and (4-5). At the end of the simulation, the primary fracture and regions of subnetworks are fully saturated with  $\text{scCO}_2$  but the hanging portions of fractures remain fully saturated with water, having never been accessible to  $\text{scCO}_2$  due to buoyancy effects.

320 simulation where the non-hanging regions of fractures 2, 3, and 4 are fully saturated with  
 321  $\text{scCO}_2$  but other regions of fractures 3 and 4 remain fully saturated with water. The accessi-  
 322 bility of some regions and not others is due to buoyancy of  $\text{scCO}_2$  that allows  $\text{scCO}_2$  to rise  
 323 into these regions. Regions above the lines of intersection are accessible while the regions  
 324 below them are not. Note that all of these regions are portions of dead-end subnetworks that  
 325 conventional wisdom suggests to be no-flow regions. However, that is clearly not the case  
 326 here due to (a) multiphase flow effects, specifically buoyancy, and (b) that the DFN is three-  
 327 dimensional. At the end of the simulation (right sub-figure) fracture 1, 2, and pieces of 3 and  
 328 4 are fully saturated with  $\text{scCO}_2$  but the rest of the domain is saturated with water. In particu-  
 329 lar, fracture 5, while not technically a hanging fracture, remains fully saturated with water  
 330 because fracture 4, which connects to the primary fracture is a hanging fracture, and  $\text{scCO}_2$   
 331 cannot enter the lower region. This example highlights that the geometric and topological  
 332 structure of the fracture network plays a principal role in initial sweep efficiency, even in  
 333 the absence of hydraulic variability. At longer times, diffusion of  $\text{scCO}_2$  into the water will  
 334 occur, and the hanging fractures will contain a mixture of water with dissolved  $\text{scCO}_2$  at  
 335 equilibrium saturation, which depends on the pressure and temperature of the system.

336 While the conceptual definition of a hanging fracture is rather straightforward, the iden-  
 337 tification of these regions within a complex DFN is not. Unless an intersection crosses enti-  
 338 rely through a fracture thereby partitioning it into two disjoint regions, the hanging regions  
 339 are geometrically ambiguous. Moreover, multiple intersections on a fracture obfuscate their  
 340 location more so. Unlike the five-fracture example presented in this subsection, where hang-  
 341 ing fractures are obvious and clearly defined, their identification in more complicate DFNs  
 342 are difficult to rigorously define.

### 3.2 Variable Hydraulic Properties

Next we characterize the impact of variable hydraulic properties on how scCO<sub>2</sub> displaces water. In these simulations, the network geometry is fixed and we vary permeability between fractures; permeability within each fracture is constant but varies between fractures. We consider three log variances of permeability (0.1, 0.5, 1.0) and thirty realizations of the permeability field at each log variance.

Figure 6 shows snapshots from four simulations 1 day into the simulation. The top-left subfigure is the constant permeability scenario, which is used as a control case for comparison. At this time in the simulation, 57.1% of the network is saturated with scCO<sub>2</sub>. The other subfigures are realizations where (top-right)  $\ln(\sigma_k) = 0.1$  is 58% saturated with scCO<sub>2</sub>, (bottom-left)  $\ln(\sigma_k) = 0.5$  is 56.5% saturated with scCO<sub>2</sub>, and (bottom-right)  $\ln(\sigma_k) = 1.0$  is 53.2% saturated with scCO<sub>2</sub>. The influence of fracture to fracture variability subtle but present. The realization with  $\ln(\sigma_k) = 0.1$  has slightly more scCO<sub>2</sub> than the control case, while the realization with  $\ln(\sigma_k) = 1.0$  has less. In this particular realization, the case with  $\ln(\sigma_k) = 0.5$  has about the same amount as the control case. Thus, the amount of scCO<sub>2</sub> in the domain relative to the control case is realization dependent, a feature that we explore next.

The variations between realizations at a value of  $\ln(\sigma_k)$  is demonstrated in Fig. 7, which shows the bulk saturation in the networks; (left)  $\ln(\sigma_k) = 0.1$ , (middle)  $\ln(\sigma_k) = 0.5$ , and (right)  $\ln(\sigma_k) = 1.0$ . The bulk saturation in the constant permeability case is included in each sub-figure as a thick line while the individual profiles from each realization is a semi-transparent line. In the case of  $\ln(\sigma_k) = 0.1$ , the simulations center around the control case. In some realizations there is slightly more scCO<sub>2</sub> saturation at the end of the simulation than in the control case and there are others where there is slightly less. The mean of the final values, along with standard deviation, is provided in Table 3. There are larger variations when  $\ln(\sigma_k) = 0.5$  (middle sub-figure) than  $\ln(\sigma_k) = 0.1$  and even larger variations in  $\ln(\sigma_k) = 1.0$ . For  $\ln(\sigma_k) = 1.0$ , there are a few realizations where the final fracture of the domain filled with scCO<sub>2</sub> is higher than the control case but most conclude with a lower fracture of the domain being saturated with scCO<sub>2</sub>. Note, that there are also deviations from the control case at early times as well as late times. This observation of limiting behavior can be attributed to the geometric and topological constraints imposed on the flow field by the network structure and, in particular, the hanging fractures previously discussed. Specifically, if a hanging fracture is assigned a lower permeability, it will act as a bottleneck making more of the network less accessible to scCO<sub>2</sub>. Homogeneity in the distribution of scCO<sub>2</sub> in each network can be measured using the coefficient of variation  $C_v$  (standard deviation over the mean) of the relative mass fraction occupied by scCO<sub>2</sub> on each fracture across each the network at the end of the simulation; higher values indicate more variation across fractures within a network. As the log variance increases, so does  $C_v$  indicating that larger variations in the hydraulic heterogeneity result in more heterogeneity in the distribution of scCO<sub>2</sub> within the networks.

Figure 8 shows the maximum percentage of the network surface area containing scCO<sub>2</sub> over time for all realizations plotted as a function of  $d_Q/P_{32}$ : the percentage of the network at the beginning of the simulation where there is significant flow. The constant permeability network  $\ln(\sigma_k) = 0.0$  is represented by a black circle,  $\ln(\sigma_k) = 0.1$  : blue diamonds,  $\ln(\sigma_k) = 0.5$  : green triangles,  $\ln(\sigma_k) = 1.0$  : red squares. In all cases the marker are above the 1-1 line, indicating that  $d_Q/P_{32}$  is a lower bound on the surface area reached by scCO<sub>2</sub> in this network regardless of the intrinsic permeabilities we consider. There is more spread in  $d_Q/P_{32}$  compared to the final scCO<sub>2</sub> percentage across the ensemble of samples. A possible explanation for this observation is that the multiphase flow features that are not present in

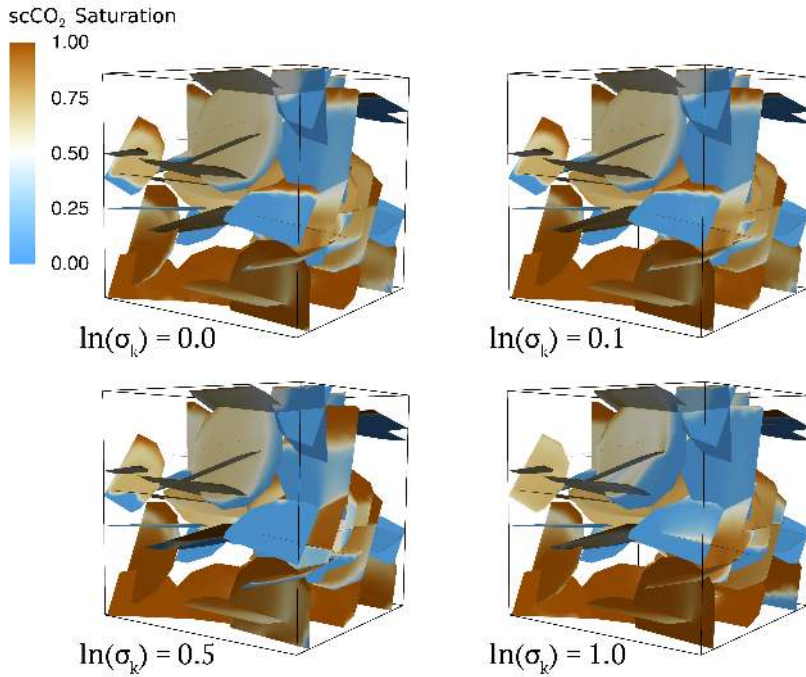


Fig. 6: Simulation of scCO<sub>2</sub> displacing water 1 day into the simulation in the same DFN with varying levels of hydraulic heterogeneity. Fracture permeabilities are sampled from a log normal distribution with log variance of (top-left)  $\ln(\sigma_k) = 0.0$  (top-right)  $\ln(\sigma_k) = 0.1$  (bottom-left)  $\ln(\sigma_k) = 0.5$  (bottom-right)  $\ln(\sigma_k) = 1.0$ . The variability in permeability can both enhance and inhibit the rate at which scCO<sub>2</sub> displaces water.

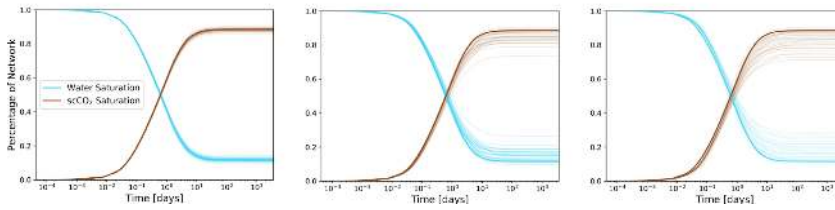


Fig. 7: Relative saturation curves in simulations with varying fracture permeability in the same network geometry. Different log-variances are shown in each sub-figure (left)  $\ln(\sigma_k) = 0.1$ , (middle)  $\ln(\sigma_k) = 0.5$ , (right)  $\ln(\sigma_k) = 1.0$ . The bulk saturation in the constant permeability case is included in each sub-figure as a thick line while the individual profiles from each of the thirty realizations are semi-transparent lines. As the log-variance increases, there is more variability between realizations are larger variants from the constant case.

Table 3: Simulation results for a fixed network geometry and varying levels of hydraulic heterogeneity. Final scCO<sub>2</sub>: fraction of the domain filled scCO<sub>2</sub> after 10 years.  $C_v$ : Coefficient of variation

$\ln(\sigma_k)$	Final scCO <sub>2</sub>			$C_v$
	mean & std. dev	max	min	
0.0	0.89 ± 0.00	0.89	0.89	0.32 ± 0.00
0.1	0.88 ± 0.01	0.90	0.86	0.32 ± 0.01
0.5	0.85 ± 0.03	0.90	0.74	0.35 ± 0.04
1.0	0.82 ± 0.05	0.90	0.74	0.36 ± 0.06

391 single phase, e.g., buoyancy, allow for the scCO<sub>2</sub> to reach areas that are more or less no-flow  
 392 regions in single phase and homogenize the structural heterogeneity of the network.

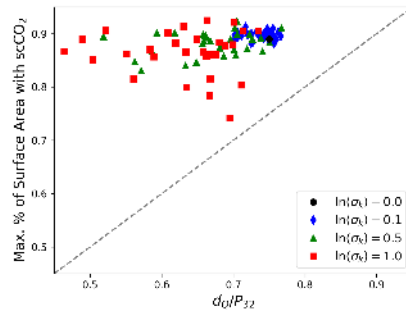


Fig. 8: Maximum percentage of the network surface area containing scCO<sub>2</sub> plotted against  $d_Q/P_{32}$  (active surface area in the single phase simulations in the same domain). The values of  $d_Q/P_{32}$  provide a lower bound on the percentage of the network surface area reached by scCO<sub>2</sub>.

### 393 3.3 Variable Structural Properties

394 In this section, we compare simulation results for networks with different densities but con-  
 395 stant permeability throughout the network. For each density, we generate thirty independent  
 396 network realizations. Figure 9 shows snapshots of the simulations in three networks, one  
 397 from each of the densities considered: (left) set 1: lowest density, (middle) set 2: middle  
 398 density, (right) set 3: highest density, 1 day into the simulation. The sample from set 1 (left)  
 399 is the network with constant permeability used in section 3.1. At this time in the simulation,  
 400 there fractures in the system that contain very little scCO<sub>2</sub> and others that are nearly fully  
 401 saturated, i.e., there is heterogeneous distribution of scCO<sub>2</sub> due to the connectivity of the  
 402 fracture network. As the density increases, this heterogeneity decreases and scCO<sub>2</sub> is more  
 403 uniformly distributed throughout the system. This homogeneity can be measured using the  
 404 coefficient of variation (standard deviation over the mean) of the relative mass fraction occu-  
 405 pied by scCO<sub>2</sub> on each fracture across each the network at the end of the simulation; higher

406 values indicate more variation across fractures within a network. In the lowest density case  
 407 the mean (with standard deviation) of the coefficient of variation across the ensemble of net-  
 408 works is  $0.56 \pm 0.34$ , for the medium it is  $0.52 \pm 0.13$ , and the highest density is  $0.39 \pm 0.09$ .  
 409 These values indicate that not only is the distribution of  $\text{scCO}_2$  more uniform across the  
 410 higher density networks, but that higher density networks are more similar to one another.  
 411 A principal reason for this homogenization is that fractures are better connected so there  
 412 are more pathways for  $\text{scCO}_2$  to enter and exit fractures. In the highest density sample, the  
 413 distribution of  $\text{scCO}_2$  is fairly uniformly, with the exception of edge effects (both fracture  
 414 and domain).

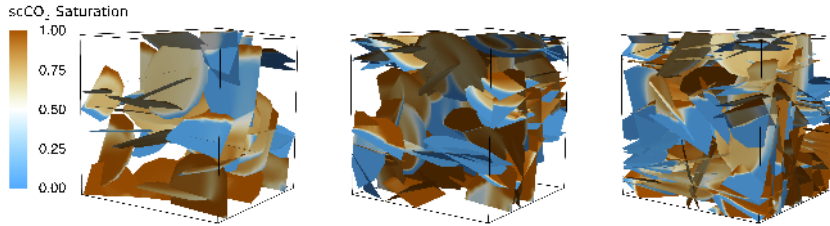


Fig. 9: Simulation snapshot after 1 day for three networks with various densities: low (left) medium (middle) and high (right) high.  $\text{scCO}_2$  is more uniformly distributed at higher densities.

415 The bulk saturation in the networks in these simulations are provided in Fig. 10; (left)  
 416 lowest density (middle) middle density, (right) highest density. The mean of the final values  
 417 of  $\text{scCO}_2$ , along with standard deviation, are provided in Table 4 along with minimum and  
 418 maximum values. There is a large amount of variation between realizations in set 1- the low-  
 419 est density. There are some networks where the final  $\text{scCO}_2$  is less than 40%, while others  
 420 are almost at 100%. In this set of networks, the sparsity results in the flow of  $\text{scCO}_2$  being  
 421 primarily controlled by the structure of the network. Many of the fractures that make up  
 422 these networks have few connections and therefore the options for inflow and outflow con-  
 423 figurations within these fractures are limited. Moreover, many fractures in these networks  
 424 are either hanging fractures or have portions of fractures that are hanging and thus not acces-  
 425 sible to  $\text{scCO}_2$  at the short time scales considered. In set 2, the bulk saturations profiles are  
 426 more uniform than in the low density case, but there is still some variation between samples.  
 427 Simulations in the set 3 (highest density) are more or less similar, especially when compared  
 428 to the low density networks. In sets 2 and 3, there are fewer hanging fractures / hanging por-  
 429 tion of the fractures. The regions that are hanging tend to be close to the domain boundaries  
 430 and the lack of accessibility appears to be the consequence of boundary effects. If we con-  
 431 sidered open lateral boundaries, it would allow more  $\text{scCO}_2$  to flow through the domain and  
 432 delay the arrival at quasi-steady state and these small numerical boundary effects would be  
 433 eliminated, but the final flow field would be relatively unaffected.

434 Figure 11 shows the maximum percentage of the network surface area occupied by  
 435  $\text{scCO}_2$  plotted as a function of  $d_Q/P_{32}$ : the percentage of the network at the beginning of  
 436 the simulation where there is significant flow, cf. Fig. 8. Set 1 is indicated by red square  
 437 markers, set 2- blue diamonds, and set 3- green triangles. There is the highest amount of



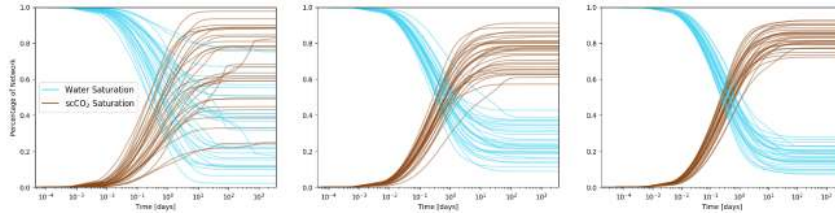


Fig. 10: Bulk saturation curves for variable density networks low (left) medium (middle) and high (right) high. As the density of the networks increases, there is less variability between network realizations.

Table 4: Simulation results for a multiple network geometries with varying densities (1-low density, 2-medium density, 3-high density) and fixed hydraulic properties. Final scCO<sub>2</sub>: amount of the domain filled scCO<sub>2</sub> after 10 years.  $C_v$ : Coefficient of variation

Network Set	Final scCO <sub>2</sub>			
	mean & std. dev	max	min	$C_v$
1	$0.67 \pm 0.21$	0.98	0.25	$0.56 \pm 0.34$
2	$0.75 \pm 0.09$	0.91	0.57	$0.52 \pm 0.13$
3	$0.83 \pm 0.05$	0.93	0.72	$0.39 \pm 0.09$

438 scatter in set 1, which is the lowest density. In all but one network, markers are above the  
 439 1-1 line, indicating that  $d_Q/P_{32}$  is less than maximum percentage of surface area occupied  
 440 by scCO<sub>2</sub>. For this set, there is 1 out of 30 where  $d_Q/P_{32}$  is greater than the final percentage  
 441 of the network saturated by scCO<sub>2</sub>; 0.61 compared to 0.54. For both sets 2 and 3,  $d_Q/P_{32}$   
 442 is less than the maximum percentage of surface area occupied by scCO<sub>2</sub> in all simulations.  
 443 Similar to what was observed in the case of varying hydraulic properties,  $d_Q/P_{32}$  appears  
 444 to provide a lower bound on the final percentage of the network surface area where scCO<sub>2</sub>  
 445 is present. Note that, however, this bound appears to depend on the network density. At  
 446 higher densities, the bound holds for all networks, but not at lowest density, where there is  
 447 an out-lier, which is discussed below. Moreover, we again observe that there is more spread  
 448 in  $d_Q/P_{32}$  compared to the final scCO<sub>2</sub> percentage across the ensemble of samples due to  
 449 multiphase flow features.

450 The one network out of ninety where the lower bound did not hold presents some in-  
 451 teresting features that could provide insights into the conditions where this apparent bound  
 452 breaks down. First, the network is poorly connected having only a single path from inlet to  
 453 outlet. Second, the area of the inlet region is consists of a single fracture with a short region  
 454 connecting to the inlet plane. The results is fairly small surface area where scCO<sub>2</sub> travel  
 455 into the domain and through. The coefficient of variation for the distribution of scCO<sub>2</sub> in  
 456 this network is 0.86, which indicates a higher degree of heterogeneity in final flow state on  
 457 this particular network than others at this density. In the case of single phase flow, the flow  
 458 spreads out across the fractures and the flow path in terms of fractures taken by scCO<sub>2</sub> is the  
 459 same as where there is activity in single phase.

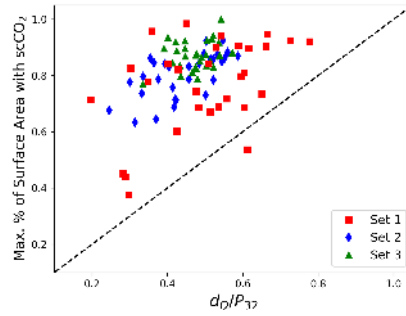


Fig. 11: Maximum percentage of the network surface area occupied by  $\text{scCO}_2$  plotted against  $d_Q/P_{32}$  (active surface area in the single phase simulations in the same domain). For sets 2 and 3, the values of  $d_Q/P_{32}$  provide a lower bound on the percentage of the network fully saturated by  $\text{scCO}_2$ . In the case of set 1, this bound holds with one exception.

#### 460 4 Discussion & Conclusions

461 We presented a set of numerical simulations performed to study the displacement of water at  
 462 hydrostatic conditions by  $\text{scCO}_2$  injected into three-dimensional discrete fracture networks.  
 463 We considered two physical scenarios. In the first, the network geometry is held constant and  
 464 the hydraulic properties of the network are varied stochastically. In the second, independent,  
 465 identically distributed network realizations are generated at various densities with constant  
 466 permeability. The results demonstrate that variations in hydraulic properties, e.g., fracture  
 467 permeability, and structural, e.g., network density, can both impact displacement of water  
 468 by  $\text{scCO}_2$  in a fracture network. In particular, the following relationships were observed:

- 469 1. At long times, once the simulations reach a quasi steady-state, the regions of the DFN  
 470 fully saturated by  $\text{scCO}_2$  are determined by network structure and fluid properties.  
 471 Specifically, hanging fractures and portions of fractures that are below fracture intersec-  
 472 tions, which is a geometric and topological property of the network, are not accessible  
 473 to  $\text{scCO}_2$  due to buoyancy, which is a property of the fluid. However, at longer times,  
 474 diffusion of  $\text{scCO}_2$  into the water will occur, and the hanging fractures will contain a  
 475 mixture of water with dissolved  $\text{scCO}_2$  at equilibrium saturation, which depends on the  
 476 pressure and temperature of the system. Also, note that these simulations were primar-  
 477 ily vertically upwards flow and if the primary flow direction was changed, the effects of  
 478 buoyancy could be different.
- 479 2. For a fixed network geometry, variations in intrinsic permeability between fractures can  
 480 both enhance and inhibit the displacement of water. As the variance of the permeability  
 481 distribution increases, the variations in bulk saturation of the networks and final dis-  
 482 tribution of the phases between realizations is more pronounced. Moreover, variations  
 483 in the volume of the domain where both phases are present is influenced in a similar  
 484 manner. The residual water in the system is typically larger with higher log variance.  
 485 This final point is due to lower permeabilities being assigned to choke points, which are  
 486 fractures that control the accessibility to hanging fractures and dead-end subnetworks at  
 487 early times.

- 488 3. Changes in the network density have a significant impact on the displacement of water.  
489 At low densities, the networks are more heterogeneous across the ensemble and unique  
490 attributes of the fracture network are reflected in the saturation profiles observed in the  
491 networks. In this set of networks, the highest variation between realizations was ob-  
492 served. These networks have numerous hanging fractures, which are not accessible to  
493 scCO<sub>2</sub>, and dead-end subnetworks, which are less accessible but not entirely inaccessi-  
494 ble. As the density of the networks increases, so does the connectivity between fractures.  
495 Thus, dead-end subnetworks and hanging fractures are eliminated, and more of the DFN  
496 is accessible to scCO<sub>2</sub>. In turn there is more uniformity between DFN realizations and  
497 structural aspects of the individual fractures become less important.
- 498 4. In comparison, variations in the network structure, such as density and topology, have a  
499 more pronounced impact on how water is displaced in fracture networks than variations  
500 in fracture permeability for the moderate values of hydraulic heterogeneity we consid-  
501 ered. These variations appear to primarily be the result of hanging fractures, which can-  
502 not become significantly more accessible to higher permeability, but can become better  
503 connected as the network density increases.
- 504 5. Our results indicate a possible attribute of single phase simulations, namely the active  
505 surface area, that could provide a lower bound on the maximum percentage of the net-  
506 work surface area that scCO<sub>2</sub> reaches. A possible explanation for this relationship is  
507 that the combination of injection and buoyancy allows scCO<sub>2</sub> to reach areas in the net-  
508 work that are more or less no-flow regions in single phase. This bound is observed for  
509 all cases of variable permeability and in the higher density networks (sets 2 and 3 in  
510 the variable DFN case). However, the bound does not hold for all networks at the lowest  
511 densities. The one exception is in a rather sparse network with a single pathways from  
512 inlet to outlet where many of the fractures are poorly connected. In this network, the  
513 flow of scCO<sub>2</sub> is highly channelized and does not disperse through the network evenly.  
514 While the bound is not firm, this observation does lend to providing a starting point to  
515 estimate what percentage of the network surface will be accessed by injected scCO<sub>2</sub>.  
516 However, we have only considered three densities of uniformly sized fractures. Addi-  
517 tional network structures with a range of geological attributes, e.g., variability length  
518 distributions along with lower and higher densities, are needed to further explore the  
519 causes of this relationship and its limitations.

520 We designed these simulations to isolate the effects of two hypothesized principal con-  
521 trols on multiphase flow through fractured media, hydraulic and structural heterogeneity.  
522 There are numerous extensions to these simulations that will build off of the presented re-  
523 sults to identify and rank additional features and their impact on multiphase flow in fracture  
524 networks. At large scales, performing simulations with fractures of variable sizes, i.e., radii  
525 following a powerlaw or lognormal distribution, is critical to advance our understanding  
526 of the interplay of geometry and multiphase flow behavior. At the opposite scale, the in-  
527 fluence of in-fracture aperture variability on capillary pressures and relative permeability  
528 remains unresolved. Additionally, we do not account for the diffusion of scCO<sub>2</sub> into the  
529 water, which is important at long time scales and could lead to more of the network being  
530 expose to scCO<sub>2</sub> and result in geochemical interactions between scCO<sub>2</sub>, water and fracture  
531 surface. Additional factors to take into consideration include the adopted functional form  
532 of relative permeability and transfer with the matrix surrounding the fracture network. The  
533 currently adopted boundary conditions (closed lateral boundaries) and open top boundary  
534 slightly influence the structure of the flow field near the lateral boundaries of the domain.  
535 If we considered open lateral boundaries, it would allow more scCO<sub>2</sub> to flow through the

536 domain and delay the arrival at quasi-steady state. These small numerical boundary effects  
 537 would be eliminated, but the key results of the study, that network structure dominates over  
 538 internal hydraulic properties, would remain the same. All of these scenarios deserve in-  
 539 dependent studies to better characterize their impact on our conceptual understanding of  
 540 multiphase flow in fracture networks.

#### 541 Acknowledgments

542 J.D.H., R.P. and P.S. are thankful for support from the US Department of Energy through the  
 543 Los Alamos National Laboratory. Los Alamos National Laboratory is operated by Triad Na-  
 544 tional Security, LLC, for the National Nuclear Security Administration of U.S. Department  
 545 of Energy (Contract No. 89233218CNA000001). Specifically, J.D.H acknowledges support  
 546 through the Laboratory-Directed Research and Development Program grants 20180621ECR.  
 547 This work was completed as part of the National Risk Assessment Partnership (NRAP)  
 548 project. Support for this project came from the U.S. Department of Energy's (DOE) Office  
 549 of Fossil Energy's Coal Research program. The authors wish to acknowledge Traci Rodosta  
 550 and Mary K. Underwood (NETL Strategic Center for Coal) and Mark Ackiewicz and Darin  
 551 Damiani (DOE Office of Fossil Energy) for programmatic guidance, direction, and support.  
 552 J.D.H. also thanks the partial support of DOE's Office of Science Basic Energy Sciences  
 553 E3W1. J.J.M. acknowledges discretionary funding from EAWAG and Guest Scientist status  
 554 from LANL. LA-UR-19-25316

#### 555 References

- 556 1. Andrade, J.S., Costa, U.M.S., Almeida, M.P., Makse, H.A., Stanley, H.E.: Inertial effects on fluid flow  
 557 through disordered porous media. *Phys. Rev. Lett.* **82**, 5249–5252 (1999)
- 558 2. Bachu, S.: Sequestration of CO<sub>2</sub> in geological media: criteria and approach for site selection in response  
 559 to climate change. *Energ. Convers. Manage.* **41**(9), 953–970 (2000)
- 560 3. Bachu, S., Adams, J.: Sequestration of CO<sub>2</sub> in geological media in response to climate change: capacity  
 561 of deep saline aquifers to sequester CO<sub>2</sub> in solution. *Energ. Convers. Manage.* **44**(20), 3151–3175 (2003)
- 562 4. Bell, R., Dean, P., Hibbins-Butler, D.: Localization of normal modes in vitreous silica, germania and  
 563 beryllium fluoride. *J. Phys. C Solid State* **3**(10), 2111 (1970)
- 564 5. Berre, I., Doster, F., Keilegavlen, E.: Flow in fractured porous media: A review of conceptual models  
 565 and discretization approaches. *Transport Porous Med.* pp. 1–22 (2018)
- 566 6. Bickle, M.J.: Geological carbon storage. *Nat. Geoscience* **2**(12), 815 (2009)
- 567 7. Birkholzer, J.T., Zhou, Q., Tsang, C.F.: Large-scale impact of CO<sub>2</sub> storage in deep saline aquifers: A  
 568 sensitivity study on pressure response in stratified systems. *Int. J. Greenh. Gas. Con.* **3**(2), 181–194  
 569 (2009)
- 570 8. Bogdanov, I., Mourzenko, V., Thovert, J.F., Adler, P.: Effective permeability of fractured porous media  
 571 with power-law distribution of fracture sizes. *Phys. Rev. E* **76**(3), 036309 (2007)
- 572 9. Bonnet, E., Bour, O., Odling, N.E., Davy, P., Main, I., Cowie, P., Berkowitz, B.: Scaling of fracture  
 573 systems in geological media. *Rev. Geophys.* **39**(3), 347–383 (2001)
- 574 10. Chaudhuri, A., Rajaram, H., Viswanathan, H., Zyvoloski, G., Stauffer, P.: Buoyant convection resulting  
 575 from dissolution and permeability growth in vertical limestone fractures. *Geophys. Res. Lett.* **36**(3)  
 576 (2009)
- 577 11. Class, H., Ebigbo, A., Helmig, R., Dahle, H.K., Nordbotten, J.M., Celia, M.A., Audigane, P., Darcis,  
 578 M., Ennis-King, J., Fan, Y., et al.: A benchmark study on problems related to CO<sub>2</sub> storage in geologic  
 579 formations. *Computat. Geosci.* **13**(4), 409 (2009)
- 580 12. Davy, P., Hansen, A., Bonnet, E., Zhang, S.Z.: Localization and fault growth in layered brittle-ductile  
 581 systems: Implications for deformations of the continental lithosphere. *J. Geophys. Res.-Sol. Ea.* **100**(B4),  
 582 6281–6294 (1995)
- 583 13. Davy, P., Le Goc, R., Darcel, C., Bour, O., De Dreuzy, J.R., Munier, R.: A likely universal model of  
 584 fracture scaling and its consequence for crustal hydromechanics. *J. Geophys. Res.-Sol. Ea.* **115**(B10)  
 585 (2010)

- 586 14. Deng, H., Molins, S., Trebotich, D., Steefel, C., DePaolo, D.: Pore-scale numerical investigation of the  
587 impacts of surface roughness: Upscaling of reaction rates in rough fractures. *Geochim. Cosmochim. Ac.*  
588 **239**, 374–389 (2018)
- 589 15. Deng, H., Stauffer, P.H., Dai, Z., Jiao, Z., Surdam, R.C.: Simulation of industrial-scale CO<sub>2</sub> storage:  
590 Multi-scale heterogeneity and its impacts on storage capacity, injectivity and leakage. *Int. J. Greenh.*  
591 *Gas. Con.* **10**, 397–418 (2012)
- 592 16. Dershowitz, W.S., Herda, H.H.: Interpretation of fracture spacing and intensity. In: *The 33th US Sym-*  
593 *posium on Rock Mechanics (USRMS)*. American Rock Mechanics Association (1992)
- 594 17. de Dreuzy, J.R., Davy, P., Bour, O.: Hydraulic properties of two-dimensional random fracture networks  
595 following power law distributions of length and aperture. *Water Resour. Res.* **38**(12) (2002)
- 596 18. de Dreuzy, J.R., Méheust, Y., Pichot, G.: Influence of fracture scale heterogeneity on the flow properties  
597 of three-dimensional discrete fracture networks. *J. Geophys. Res.-Sol. Ea.* **117**(B11) (2012)
- 598 19. Du, Q., Faber, V., Gunzburger, M.: Centroidal voronoi tessellations: Applications and algorithms. *SIAM*  
599 *Rev.* **41**(4), 637–676 (1999)
- 600 20. Edwards, J., Thouless, D.: Numerical studies of localization in disordered systems. *J. Phys C Solid State*  
601 **5**(8), 807 (1972)
- 602 21. Frampton, A., Hyman, J., Zou, L.: Advective transport in discrete fracture networks with connected and  
603 disconnected textures representing internal aperture variability. *Water Resour. Res.* **55**(7), 5487–5501  
604 (2019)
- 605 22. Hadgu, T., Karra, S., Kalinina, E., Makedonska, N., Hyman, J.D., Klise, K., Viswanathan, H.S., Wang,  
606 Y.: A comparative study of discrete fracture network and equivalent continuum models for simulating  
607 flow and transport in the far field of a hypothetical nuclear waste repository in crystalline host rock. *J.*  
608 *Hydrol.* **553**, 59–70 (2017)
- 609 23. Hyman, J.D., Karra, S., Carey, J.W., Gable, C.W., Viswanathan, H., Rougier, E., Lei, Z.: Discontinuities  
610 in effective permeability due to fracture percolation. *Mech. Mater.* **119**, 25 – 33 (2018)
- 611 24. Hyman, J.D., Aldrich, G., Viswanathan, H., Makedonska, N., Karra, S.: Fracture size and transmis-  
612 sivity correlations: Implications for transport simulations in sparse three-dimensional discrete fracture  
613 networks following a truncated power law distribution of fracture size. *Water Resour. Res.* **52**(8), 6472–  
614 6489 (2016). DOI 10.1002/2016WR018806
- 615 25. Hyman, J.D., Dentz, M., Hagberg, A., Kang, P.: Linking structural and transport properties in three-  
616 dimensional fracture networks. *J. Geophys. Res.-Sol. Ea.* (2019)
- 617 26. Hyman, J.D., Gable, C.W., Painter, S.L., Makedonska, N.: Conforming Delaunay triangulation of  
618 stochastically generated three dimensional discrete fracture networks: A feature rejection algorithm for  
619 meshing strategy. *SIAM J. Sci. Comput.* **36**(4), A1871–A1894 (2014)
- 620 27. Hyman, J.D., Hagberg, A., Osthus, D., Srinivasan, S., Viswanathan, H., Srinivasan, G.: Identifying back-  
621 bones in three-dimensional discrete fracture networks: A bipartite graph-based approach. *Multiscale*  
622 *Model. Sim.* **16**(4), 1948–1968 (2018)
- 623 28. Hyman, J.D., Jiménez-Martínez, J.: Dispersion and mixing in three-dimensional discrete fracture net-  
624 works: Nonlinear interplay between structural and hydraulic heterogeneity. *Water Resour. Res.* **54**(5),  
625 3243–3258
- 626 29. Hyman, J.D., Karra, S., Makedonska, N., Gable, C.W., Painter, S.L., Viswanathan, H.S.: dfnWorks: A  
627 discrete fracture network framework for modeling subsurface flow and transport. *Comput. Geosci.* **84**,  
628 10–19 (2015)
- 629 30. Joyce, S., Hartley, L., Applegate, D., Hoek, J., Jackson, P.: Multi-scale groundwater flow modeling dur-  
630 ing temperate climate conditions for the safety assessment of the proposed high-level nuclear waste  
631 repository site at Forsmark, Sweden. *Hydrogeol. J.* **22**(6), 1233–1249 (2014)
- 632 31. Karra, S., Makedonska, N., Viswanathan, H., Painter, S., Hyman, J.: Effect of advective flow in fractures  
633 and matrix diffusion on natural gas production. *Water Resour. Res.* **51**(10), 8646–8657 (2015)
- 634 32. Keating, E.H., Newell, D.L., Viswanathan, H., Carey, J., Zvoloski, G., Pawar, R.: CO<sub>2</sub>/brine transport  
635 into shallow aquifers along fault zones. *Environ. Sci. Technol.* **47**(1), 290–297 (2012)
- 636 33. LaGriT: Los Alamos Grid Toolbox, (LaGriT) Los Alamos National Laboratory. <http://lagrit.lanl.gov>  
637 (2013)
- 638 34. Lovell, A.E., Srinivasan, S., Karra, S., O'Malley, D., Makedonska, N., Viswanathan, H.S., Srinivasan,  
639 G., Carey, J.W., Frash, L.P.: Extracting hydrocarbon from shale: An investigation of the factors that  
640 influence the decline and the tail of the production curve. *Water Resour. Res.* (2018)
- 641 35. Maillot, J., Davy, P., Le Goc, R., Darcel, C., De Dreuzy, J.R.: Connectivity, permeability, and channeling  
642 in randomly distributed and kinematically defined discrete fracture network models. *Water Resour. Res.*  
643 **52**(11), 8526–8545 (2016)
- 644 36. Makedonska, N., Hyman, J.D., Karra, S., Painter, S.L., Gable, C.W.W., Viswanathan, H.S.: Evaluating  
645 the effect of internal aperture variability on transport in kilometer scale discrete fracture networks. *Adv.*  
646 *Water Resour.* **94**, 486–497 (2016)

- 647 37. Mourzenko, V., Thovert, J.F., Adler, P.: Percolation of three-dimensional fracture networks with power-  
648 law size distribution. *Phys. Rev. E* **72**(3), 036103 (2005)
- 649 38. Mudunuru, M.K., Karra, S., Makedonska, N., Chen, T.: Sequential geophysical and flow inversion to  
650 characterize fracture networks in subsurface systems. *Stat. Anal. Data. Min.* **10**(5), 326–342 (2017)
- 651 39. Oh, J., Kim, K.Y., Han, W.S., Park, E., Kim, J.C.: Migration behavior of supercritical and liquid CO<sub>2</sub>  
652 in a stratified system: Experiments and numerical simulations. *Water Resour. Res.* **51**(10), 7937–7958  
653 (2015)
- 654 40. Porter, M.L., Plampin, M., Pawar, R., Illangasekare, T.: CO<sub>2</sub> leakage in shallow aquifers: A benchmark  
655 modeling study of CO<sub>2</sub> gas evolution in heterogeneous porous media. *Int. J. Greenh. Gas. Con.* **39**,  
656 51–61 (2015)
- 657 41. Raduha, S., Butler, D., Mozley, P., Person, M., Evans, J., Heath, J., Dewers, T., Stauffer, P., Gable, C.,  
658 Kelkar, S.: Potential seal bypass and caprock storage produced by deformation-band-to-opening-mode-  
659 fracture transition at the reservoir/caprock interface. *Geofluids* **16**(4), 752–768 (2016)
- 660 42. Rutqvist, J., Tsang, C.F.: A study of caprock hydromechanical changes associated with CO<sub>2</sub>-injection  
661 into a brine formation. *Environ. Geol.* **42**(2-3), 296–305 (2002)
- 662 43. Saadatpoor, E., Bryant, S.L., Sepehrnoori, K.: New trapping mechanism in carbon sequestration. *Trans-  
663 port Porous Med.* **82**(1), 3–17 (2010)
- 664 44. Sanchez-Vila, X., Guadagnini, A., Carrera, J.: Representative hydraulic conductivities in saturated  
665 groundwater flow. *Rev. Geophys.* **44**(3) (2006)
- 666 45. Seidman, S.B.: Network structure and minimum degree. *Soc. Networks* **5**(3), 269 – 287 (1983)
- 667 46. Shukla, R., Ranjith, P., Haque, A., Choi, X.: A review of studies on CO<sub>2</sub> sequestration and caprock  
668 integrity. *Fuel* **89**(10), 2651–2664 (2010)
- 669 47. Svensk Kärnbränslehantering AB: Data report for the safety assessment SR-site (TR-10-52). Tech. rep.,  
670 Svensk Kärnbränslehantering AB (2010)
- 671 48. Ushijima-Mwesigwa, H., Hyman, J.D., Hagberg, A., Safro, I., Karra, S., Gable, C.W., Srinivasan, G.:  
672 Multilevel graph partitioning for three-dimensional discrete fracture network flow simulations. arXiv  
673 preprint arXiv:1902.08029 (2019)
- 674 49. Watson, F.E., Mathias, S.A., van Hunen, J., Daniels, S.E., Jones, R.R.: Dissolution of CO<sub>2</sub> from leaking  
675 fractures in saline formations. *Transport Porous Med.* **94**(3), 729–745 (2012)
- 676 50. Wellman, T.P., Shapiro, A.M., Hill, M.C.: Effects of simplifying fracture network representation on inert  
677 chemical migration in fracture-controlled aquifers. *Water Resour. Res.* **45**(1) (2009)
- 678 51. Witherspoon, P.A., Wang, J., Iwai, K., Gale, J.: Validity of cubic law for fluid flow in a deformable rock  
679 fracture. *Water Resour. Res.* **16**(6), 1016–1024 (1980)
- 680 52. Wood, A.T.: Simulation of the von Mises Fisher distribution. *Commun. Stat. Simulat.* **23**(1), 157–164  
681 (1994)
- 682 53. Zyvoloski, G.: FEHM: A control volume finite element code for simulating subsurface multi-phase  
683 multi-fluid heat and mass transfer. Los Alamos Unclassified Report LA-UR-07-3359 (2007)