



Chemical characterization and source apportionment of PM_{2.5} in Beijing: seasonal perspective

R. Zhang^{1,*}, J. Jing^{1,2,*}, J. Tao^{3,*}, S.-C. Hsu^{4,*}, G. Wang⁵, J. Cao⁵, C. S. L. Lee⁶, L. Zhu³, Z. Chen⁷, Y. Zhao⁷, and Z. Shen⁸

¹Key Laboratory of Regional Climate-Environment Research for Temperate East Asia, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, China

²Meteorological Observation Center of CMA, Beijing, China

³South China Institute of Environmental Sciences, Ministry of Environmental Protection, Guangzhou, China

⁴Research Center for Environmental Changes, Academia Sinica, Taipei, Taiwan

⁵K LAST, SKLLQG, Institute of Earth Environment, Chinese Academy of Sciences, Xi'an, China

⁶Institute of Earth Sciences, Academia Sinica, Taipei, Taiwan

⁷ESPC, College of Environmental Sciences and Engineering, Peking University, Beijing, China

⁸Department of Environmental Science and Engineering, Xi'an Jiaotong University, Xi'an, China

*These authors contributed equally to this work.

Correspondence to: S.-C. Hsu (schsu815@rcec.sinica.edu.tw)

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Abstract. In this study, 121 daily PM_{2.5} (aerosol particle with aerodynamic diameter less than 2.5 μm) samples were collected from an urban site in Beijing in four months between April 2009 and January 2010 representing the four seasons. The samples were determined for various compositions, including elements, ions, and organic/elemental carbon. Various approaches, such as chemical mass balance, positive matrix factorization (PMF), trajectory clustering, and potential source contribution function (PSCF), were employed for characterizing aerosol speciation, identifying likely sources, and apportioning contributions from each likely source. Our results have shown distinctive seasonality for various aerosol speciations associated with PM_{2.5} in Beijing. Soil dust waxes in the spring and wanes in the summer. Regarding the secondary aerosol components, inorganic and organic species may behave in different manners. The former preferentially forms in the hot and humid summer via photochemical reactions, although their precursor gases, such as SO₂ and NO_x, are emitted much more in winter. The latter seems to favorably form in the cold and dry winter. Synoptic meteorological and climate conditions can overwhelm the emission pattern in the formation of secondary aerosols. The PMF model identified six main sources: soil

dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol. Each of these sources has an annual mean contribution of 16, 14, 13, 3, 28, and 26 %, respectively, to PM_{2.5}. However, the relative contributions of these identified sources significantly vary with changing seasons. The results of trajectory clustering and the PSCF method demonstrated that regional sources could be crucial contributors to PM pollution in Beijing. In conclusion, we have unraveled some complex aspects of the pollution sources and formation processes of PM_{2.5} in Beijing. To our knowledge, this is the first systematic study that comprehensively explores the chemical characterizations and source apportionments of PM_{2.5} aerosol speciation in Beijing by applying multiple approaches based on a completely seasonal perspective.

1 Introduction

Particulate matter (PM) is composed of various chemical components (Seinfeld, 1989). PM profoundly affects our living environments in terms of air quality (in close relation to public health), visibility, direct and indirect radiative forcing,

climate effects and ecosystems (Watson, 2003; Streets et al., 2006; Andreae and Rosenfeld, 2008; Mahowald, 2011). Numerous epidemiological studies have demonstrated that long-term exposure to pronounced PM_{2.5} increases morbidity and mortality (Dockery and Pope, 1994; Pope et al., 1995; Schwartz et al., 1996). Given its tiny size, fine-mode PM (i.e., PM_{2.5}, PM with aerodynamic diameter less than 2.5 μm) can readily penetrate the human bronchi and lungs (Pope et al., 1995; Oberdörster, 2001). Through absorption and scattering of solar radiation and serving as cloud condensation nuclei, PM_{2.5} extensively affects the global climate (Bardouki et al., 2003), and thus the hydrological cycle (Ramanathan and Feng, 2009). The diverse effects of PM_{2.5} could be a function of its complex chemical components and composition (He et al., 2009; Niwa et al., 2007; Malm et al., 2005; Eatough et al., 2006).

Due to the rapid economic and industrial developments and urbanization in the past few decades, there is an escalating increase in energy consumption and the number of motor vehicles in China, where air pollution has become ubiquitous (Chan and Yao, 2008). According to Shao et al. (2006), nearly 70 % of urban areas in China do not meet China's national ambient air quality standards, which are even much laxer than the air quality exposure standards/guidelines of the World Health Organization (WHO, 2005). The Beijing–Tianjin–Hebei region, the Yangtze River delta, and the Pearl River delta are of special concern because of their severe PM pollution, which can be explicitly shown by the spatial distribution of aerosol optical depth (AOD) retrieved by satellites (He et al., 2009; Lee et al., 2010). Three megacities that are representatives of each region, i.e., Beijing, Shanghai, and Guangzhou, are the foci, because of their dense population. Coal is the primary energy source in China, and its consumption reached up to 1528 Mtce in 2005, accounting for nearly 70 % of the total energy consumption (followed by petroleum at over 20 %). Such quantity ranks number one in the world, representing ~ 37 % of global consumption (Fang et al., 2009; Chen and Xu, 2010). The use of coal in China ranges from large power plants, industries to individual domestic households, and thus coal combustion becomes the largest contributor of air pollution (Liu and Diamond, 2005; Chan and Yao, 2008). Given the rapid growth in vehicle numbers at a rising rate of ~ 20 %, traffic has become a major urban pollutant emitter (He et al., 2002; Fang et al., 2009). Other than anthropogenic pollutants, desert and loess dust of natural origins with annual emissions of over 100 million tons also serves as important PM source in China, particularly in late winter and spring (Zhang et al., 1997; Sun et al., 2001).

Atmospheric pollutants in China are a complex mixtures of various sources, from gases to particulates, from natural to anthropogenic, from primary to secondary, and from local to regional and the term “air pollution complex” or “complex atmospheric pollution” has emerged in the last decade (He et al., 2002; Shao et al., 2006; Chan and Yao, 2008; Fang et

al., 2009). One of the major air pollutants is PM, particularly PM_{2.5}, which remains a nationwide problem despite considerable efforts for its removal (Fang et al., 2009). As the capital of China and a rapidly industrialized and typical urbanized city, Beijing has elicited much more attention domestically and internationally (Zhang et al., 2003a, 2007; Zhou et al., 2012). PM_{2.5} in Beijing is abnormally elevated, often rising to more than 100 $\mu\text{g m}^{-3}$, and characterized by multiple components and sources, ranging from inorganic to organic constituents, from anthropogenic to natural origins, from primary to secondary components, and from local to long-range transported sources, and in dynamic variability with time and/or meteorological conditions and climate regimes (He et al., 2001; Wang et al., 2005; Duan et al., 2006; Okuda et al., 2011; Song et al., 2012). In spite of many scientific research programs conducted by academic institutions and the political strategies implemented by the government, the state of air pollution in Beijing (and even across China) appears to be slowly improving. For instance, in January 2013, Beijing (and the entire inland China) suffered from the worst PM_{2.5} pollution in history, registering the highest PM_{2.5} hourly concentration of 886 $\mu\text{g m}^{-3}$ (http://www.nasa.gov/multimedia/imagegallery/image_feature2425.html). Some essential questions remain unknown, although the government has devoted itself to improving air quality and numerous studies have been conducted. Therefore, a systematically comprehensive investigation of employing multiple techniques in conjunction with chemical measurements is inevitably needed, particularly to unravel the likely contributors of PM_{2.5}.

Receptor models are used to quantitatively estimate the pollutant levels contributed by different sources through statistical interpretation of ambient measurement. The positive matrix factorization (PMF) developed by the Environmental Protection Agency of USA is a well-adopted receptor model for source apportionment analysis. A few studies have applied PMF to identify the likely dominant sources and apportion their respective contributions. For example, Wang et al. (2008) analyzed certain elements, ions, and black carbon in PM_{2.5} and PM₁₀ samples collected in Beijing in summer and winter (one month representative for each season) between 2001 and 2006. Based on the obtained data set, they performed PMF analyses and identified six main sources: soil dust, vehicular emission, coal combustion, secondary aerosol, industrial emission, and biomass burning. By applying the PMF model with only elemental data as input data, Yu et al. (2013) identified seven likely sources of PM_{2.5} in Beijing, with relative contributions following the order secondary sulfur (26.5 %), vehicle exhaust (17.1 %), fossil fuel combustion (16.0 %), road dust (12.7 %), biomass burning (11.2 %), soil dust (10.4 %), and metal processing (6.0 %). Song et al. (2007) analyzed a few elements, ions, and organic/elemental carbon (OC/EC) in PM_{2.5} collected from multiple stations in Beijing during a short period in January and August 2004, and subjected them to PMF analyses. Six potential sources were registered: coal combustion, biomass

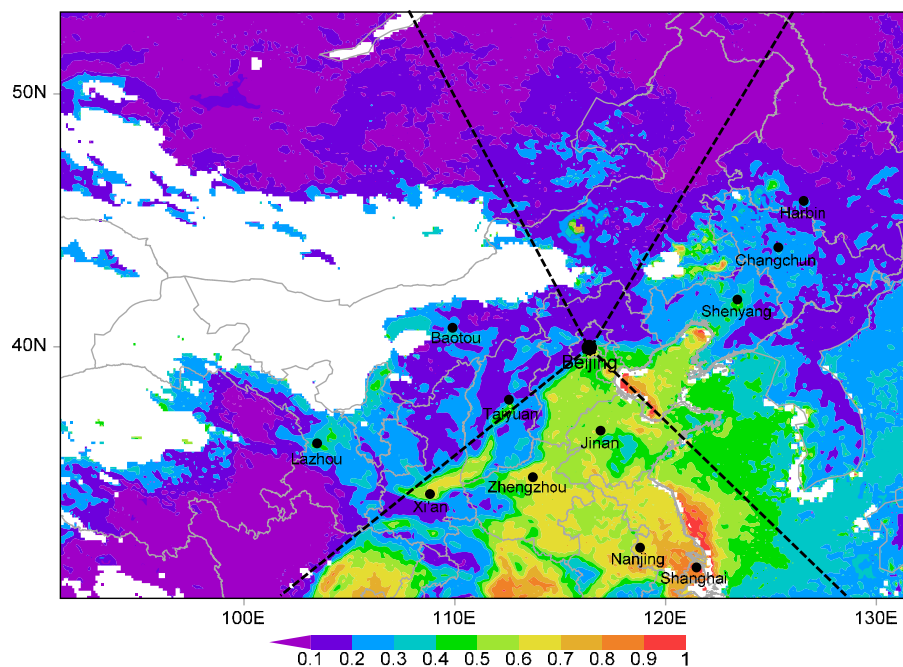


Fig. 1. Sampling location (116.30° E, 39.99° N) on a regional map superimposed with spatial distribution of annual mean fine aerosol optical depth (AOD) retrieved from MODIS satellite remote sensing in 2009. Dashed lines define four regions based on the trajectory clustering results discussed in Sect. 4.2 (seen in the text). Also shown are several major cities around Beijing.

burning, secondary sulfate, motor vehicles, secondary nitrate, and road dust, with emphasis on coal combustion in winter and secondary aerosols in summer. Xie et al. (2008) conducted PMF analyses of PM₁₀ (instead of PM_{2.5}) collected in Beijing in 10 days each in January, April, July, and October 2004 by using chemical data on metal elements, ions, and OC/EC as input data. Seven main sources were identified, including urban fugitive dust, crustal soil, coal combustion, secondary sulfate, secondary nitrate, biomass burning with municipal incineration, and vehicle emissions. All these studies were limited to a particular season and based on selected PM species.

To attain a better understanding of the chemical characteristics and sources of fine aerosols on a seasonal basis, we conducted a delicate investigation in Beijing. We continuously collected daily PM_{2.5} samples at an urban site for four months, each of which in the respective seasons (i.e., spring, summer, autumn and winter). The samples were subjected to chemical measurements of various aerosol compositions as a whole, such as a suite of crustal and anthropogenic elements, major water-soluble ions, and OC/EC. Furthermore, we identified and apportioned the main sources to PM_{2.5} by employing chemical mass closure construction and the PMF model in conjunction with trajectory cluster and potential source contribution function analyses according to the hybrid single-particle Lagrangian integrated trajectory (HYSPPLIT) model. This study will elucidate the source profile of PM_{2.5} in different seasons and the relative contribution from each

source in the complex urban airshed in Beijing and provide vital information in formulating the future air management framework to address the current alarming level of PM pollution in China which has been affecting the air quality on a vast regional scale.

2 Methodology

2.1 Sample and chemical analysis

2.1.1 Sampling site

Beijing is located on the northern edge of the North China Plain, surrounded by the Yanshan Mountains in the west, north, and northeast (Fig. 1). According to the spatial distribution of fine AOD ranging from 0.0 to 1.0 that has been retrieved from Moderate-resolution Imaging Spectrometer (MODIS) sensors on board Terra and Aqua satellites (Fig. 1), Beijing is one of the PM_{2.5} hot spots in China. The four seasons are characterized by variable meteorological conditions: spring by high-speed winds and low rainfall, summer by high temperature and frequent rain usually accounting for 75% of annual rainfall, autumn by sunny days and north-west winds, and winter by cold and dry air. The population is 17.55 million. In 2009, the number of motor vehicles increased to 4.019 million, and energy consumption was equivalent to 65.73 million tons of standard coal (2010 Beijing Statistics Yearbook). The sampling station was set up at the

roof of the Science Building in Peking University (116.30° E, 39.99° N) 26 m a.g.l. (above ground level). A few field experimental campaigns have been conducted at this urban site (He et al., 2010; Guo et al., 2012). This site is located within the educational, commercial, and residential districts, and no main pollution sources exist nearby. Thus, the observations could be typical of the general urban pollution in Beijing.

2.1.2 Sample collection

Daily PM_{2.5} samples were collected in April, July, and October 2009 and January 2010, representing spring, summer, autumn, and winter, respectively. Two collocated aerosol samplers (fmOMNI™, BGI, USA) were used to collect PM_{2.5} samples from 10:00 to 10:00 LT the next day simultaneously. The two substrates used in each sampler were 47 mm quartz filter (Whatman QM/A, England) and Teflon filter (pore size = 2 μm; Whatman PTFE, England). The flow rate was set at 5 L min⁻¹. The quartz filters were baked at 800 °C for 3 h before use. The filter samples were stored at -18 °C until pretreatment.

2.1.3 Gravimetric weighing

Before and after each sampling, the PTFE filters were conditioned at 22 °C ± 1 °C in relative humidity of 35 % ± 2 % for 24 h and then weighed in a weighing room by using an electronic balance with a detection limit of 1 μg (Sartorius, Göttingen, Germany). The corresponding PM_{2.5} mass concentration of each filter was equal to the weight difference before and after sampling divided by the sampled air volume.

2.1.4 Chemical analysis of trace elements and water-soluble ions

Prior to extraction and digestion, each aerosol-laden PTFE membrane filter was cut into two equal halves with ceramic scissors. One half was subjected to Milli-Q water extraction for ionic measurement and the other half to acid digestion for elemental measurement. For the acid digestion, the polypropylene support O-ring on half of each PTFE filter sample was carefully removed with a ceramic knife from contamination. The filter samples were digested with an acid mixture (5 mL HNO₃ + 2 mL HF) by using an ultra-high throughput microwave digestion system (MARSXpress, CEM, Matthews, NC). A blank reagent and two filter blanks were prepared in each run following the same procedure used for the samples. All the acids used in this study were of ultra-pure grade (Merck, Germany). The detailed digestion method has been published elsewhere (Hsu et al., 2008). Another half of all filter samples were used for extraction with 20 mL Milli-Q purified water (specific resistivity = 18.2 MΩ cm⁻¹; Millipore, Massachusetts, USA) for 1 h. The detailed extraction procedures have been described in Hsu et al. (2007, 2010a).

Ionic species (Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, F⁻, Cl⁻, SO₄²⁻ and NO₃⁻) in the leachate were analyzed through a Dionex model ICS-90 (for anions) and ICS-1500 (for cations) ion chromatograph equipped with a conductivity detector (ASRS-ULTRA). Trace elements in the digestion solutions, including Al, Fe, Na, Mg, K, Ca, Ba, Ti, Mn, Co, Ni, Cu, Zn, Mo, Cd, Sn, Sr, Sb, Pb, Tl, Ge, Cs, Ga, V, Cr, As, Se, and Rb, were analyzed by inductively coupled plasma-mass spectrometry (ICP-MS). Quality assurance and control of the ICP-MS was guaranteed by the analysis of a certified reference standard, NIST SRM-1648 (urban particulates). The resulting recoveries fell within ±10 % of the certified values for most elements, except for Se, As, Cs, Sb, and Rb (±15 %) (Hsu et al., 2009, 2010a).

2.1.5 OC and EC measurements

A punch of 0.526 cm² from each quartz filter was heated stepwise by a thermal/optical carbon analyzer (DRI 2001, Atmoslytic, US) in a pure helium atmosphere at 140 °C (OC1), 280 °C (OC2), 480 °C (OC3), and 580 °C (OC4), and then in 2 % O₂/98 % He atmosphere at 580 °C (EC1), 740 °C (EC2), and 840 °C (EC3) to convert any particulate carbon on the filter to CO₂. After catalyzed by MnO₂, CO₂ was reduced to CH₄, which was then directly measured. Mass concentrations of OC and EC were obtained according to the IMPROVE protocol (Chow et al., 2007), OC = OC1 + OC2 + OC3 + OC4 + OP; EC = EC1 + EC2 + EC3 - OP, where OP is the optical pyrolyzed OC. Detailed descriptions can be found in Zhang et al. (2012a).

2.2 Data analysis methods

2.2.1 Chemical mass closure

In this study, we constructed chemical mass closure (CMC) on a seasonal basis by considering mineral dust, SO₄²⁻, NO₃⁻, NH₄⁺, EC, particulate organic matter (POM), chloride salt (instead of sea salt; reason given below), trace element oxide (TEO), and biomass burning-derived K⁺. SO₄²⁻, NO₃⁻, and NH₄⁺ can be regarded as the secondary inorganic aerosols.

The aluminosilicate (i.e., soil, dust, or mineral) component is often estimated through the following formula (Malm et al., 1994; Chow et al., 1994), which includes Si.

$$[\text{Mineral}] = 2.20\text{Al} + 2.49\text{Si} + 1.63\text{Ca} + 2.42\text{Fe} + 1.94\text{Ti}$$
$$[\text{Mineral}] = 1.89\text{Al} + 2.14\text{Si} + 1.40\text{Ca} + 1.43\text{Fe}$$

However, Si is volatilized as SiF₄ in the acid digestion of aerosol samples when using HF; therefore, a few studies estimated Si from Al in the calculation of the mineral component (Hueglin et al., 2005). However, once Al is used in estimating Si concentrations, it generates relatively large uncertainty in the mineral component proportion as Si/Al mass ratios could largely vary in China's dust (Yan et al., 2012, and references

therein). Accordingly, we adopted a straightforward method conventionally used in estimating dust aerosols from Al:

$$[\text{Mineral}] = \text{Al}/0.07,$$

where 0.07 is the average Al content (7 %) reported by Zhang et al. (2003b). A similar estimation has been applied previously (Ho et al., 2006; Hsu et al., 2010b).

In estimating POM, we adopted a factor of 1.6 in converting OC to POM (Viidanoja et al., 2002), whereas a wide range of 1.4–2.2 has been utilized in previous investigations (Turpin and Lim, 2001; Andreae et al., 2008). The main determinants in selecting a conversion factor are the origin and age of the organic aerosols. The factor of 1.6 was employed in this study because the latest result shows a OM/OC ratio averaged at 1.59 ± 0.18 in PM_{2.5} over China (Xing et al., 2013). This factor was used for the PM_{2.5} of Beijing by Dan et al. (2004), who also observed a similar seasonality for EC and OC and a OC/EC ratio (2–3) close to our results.

Sea salt is usually calculated as $[\text{Sea salt}] = 1.82 \times \text{Cl}^-$ or $= 2.54 \text{Na}^+$. Given that Beijing is about 150 km away from East China's coastal oceans (i.e., Bohai Sea), sea spray-generated sea salt particles are not readily transported and are therefore insignificant to fine aerosols in Beijing. Nevertheless, dust blowing from Northern and Northwestern China is often associated with NaCl and Na₂SO₄ from salt lake sediments and saline soils (Zhang et al., 2009a). On the other hand, Cl⁻ may be essentially contributed by coal combustion in Beijing, particularly in winter (Yao et al., 2002). Thus, we considered chloride salt, instead of sea salt, as an individual component of PM_{2.5} aerosols in Beijing: $[\text{Cl salt}] = [\text{Cl}^-] + [\text{Na}^+] + [\text{ss-Mg}^{2+}]$. By considering chloride depletion in sea salt particles within the marine boundary layer because of the heterogeneous reaction, Hsu et al. (2010a) successfully evaluated such formula.

Following Landis et al. (2001), we estimated the contribution of heavy metals as metal oxides by employing the following equation:

$$\begin{aligned} \text{TEO} = & 1.3 \times [0.5 \times (\text{Sr} + \text{Ba} + \text{Mn} + \text{Co} + \text{Rb} + \text{Ni} + \text{V}) \\ & + 1.0 \times (\text{Cu} + \text{Zn} + \text{Mo} + \text{Cd} + \text{Sn} + \text{Sb} + \text{Ti} \\ & + \text{Pb} + \text{As} + \text{Se} + \text{Ge} + \text{Cs} + \text{Ga})]. \end{aligned}$$

The enrichment factor (EF) of a given element (*E*) was calculated by using the formula $\text{EF} = (E/\text{Al})_{\text{Aerosol}} / (E/\text{Al})_{\text{Crust}}$ (Hsu et al., 2010a), where $(E/\text{Al})_{\text{Aerosol}}$ is the ratio of the element to the Al mass in aerosols and $(E/\text{Al})_{\text{Crust}}$ is the ratio in the average crust (Taylor, 1964). The result of the EF is shown in Fig. S1 (Supplement). Elements with EFs of ≤ 1.0 , such as Cr and Y, were not considered, as they are of exclusive crustal origin. Elements with EFs between 1 and 5 were multiplied by a factor of 0.5, as they are possibly originated from two sources (i.e., anthropogenic and crustal sources). Elements with EFs ≥ 5.0 were multiplied by unity, as they are dominated by anthropogenic origins. Furthermore, the multiplicative factor was set at 1.3 so that metal abundance

could be converted to oxide abundance, similar to those used by Landis et al. (2001). We also considered biomass burning-derived K⁺ (K_{BB}) as an individual component, although K_{BB} salt may exist in the chemical forms of KCl and K₂SO₄ (Pósfai et al., 2004), where both Cl⁻ and SO₄²⁻ have already been considered in other components.

2.2.2 PMF model

PMF is an effective source apportionment receptor model that does not require the source profiles prior to analysis and has no limitation on source numbers (Hopke, 2003; Shen et al., 2010). The principles of PMF can be found elsewhere in detail (Han et al., 2006; Song et al., 2006; Yu et al., 2013). In the present study, PMF 3.0 was employed with the inclusion of 34 chemical species in the model computation: PM_{2.5}, Al, Fe, Na, Mg, K, Ca, Ba, Ti, Mn, Co, Ni, Cu, Zn, Mo, Cd, Sn, Sb, Pb, V, Cr, As, Se, Rb, Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Cl⁻, SO₄²⁻, NO₃⁻, OC, and EC. Six physically realistic sources were identified.

2.2.3 Air mass back trajectory cluster

We calculated 48 h air mass back trajectories arriving at the sampling site (116.30° E, 39.99° N) during our sampling period by using the National Oceanic and Atmospheric Administration (NOAA) HYSPLIT-4 model with a 1° × 1° latitude–longitude grid and the final meteorological database. The six-hourly final archive data were generated from the National Center for Environmental Prediction's Global Data Assimilation System (GDAS) wind field reanalysis. GDAS uses a spectral medium-range forecast model. More details about the HYSPLIT model can be found at <http://www.arl.noaa.gov/ready/open/hysplit4.html> (NOAA Air Resources Laboratory). The model was run four times per day at starting times of 04:00, 10:00, 16:00, and 22:00 UTC (12:00, 18:00, 00:00, and 06:00 LT – local time, respectively). The arrival level was set at 100 m a.g.l. The method used in trajectory clustering was based on the GIS-based software TrajStat (<http://www.meteothinker.com/TrajStatProduct.aspx>).

2.2.4 Potential source contribution function

The potential source contribution function (PSCF) is a method for identifying regional sources based on the HYSPLIT model. The zone of concern is divided into $i \times j$ small equal grid cells. The PSCF value in the ij -th cell is defined as m_{ij}/n_{ij} , where n_{ij} is designated as the numbers of endpoints that fall in the ij -th cell and m_{ij} denotes the numbers of “polluted” trajectory endpoints in the ij -th cell. In this analysis, average concentrations were treated as the “polluted” threshold (Hsu et al., 2003). To better reflect the uncertainty in cells with small n_{ij} values (Polissar et al., 1999), the weighting function w_{ij} was adopted:

Table 1. Statistical summary showing the means (with one standard deviation) and ranges of atmospheric concentrations for PM_{2.5} (in unit $\mu\text{g m}^{-3}$) and selected species (in unit ng m^{-3}) in the entire sampling (annual) and four-season (months) periods.

Species	Annual	Spring	Summer	Autumn	Winter
PM _{2.5}	135 ± 63 39–355	126 ± 59 39–280	138 ± 48 41–226	135 ± 55 45–251	139 ± 86 48–355
SO ₄ ²⁻	13.6 ± 12.4 0.9–52.8	14.7 ± 11.5 2.3–52.8	23.5 ± 14.5 2.5–52.0	7.9 ± 7.4 0.9–25.7	8.5 ± 8.6 1.3–34.4
NO ₃ ⁻	11.3 ± 10.8 0.3–63.8	15.5 ± 13.7 1.3–63.8	11.8 ± 8.2 1.8–31.5	10.7 ± 11.0 0.3–34.7	7.3 ± 8.1 1.6–35.5
NH ₄ ⁺	6.9 ± 7.1 0.1–39.1	7.5 ± 8.1 0.6–39.1	11.0 ± 6.9 0.5–23.9	4.7 ± 5.8 0.1–17.7	4.5 ± 5.7 0.3–23.3
Cl ⁻	1.42 ± 2.18 0.03–10.34	0.72 ± 0.81 0.04–3.74	0.30 ± 0.56 0.03–3.06	1.12 ± 0.98 0.09–3.71	3.52 ± 3.32 0.19–10.34
Na ⁺	0.46 ± 0.55 0.04–2.82	0.31 ± 0.18 0.08–0.94	0.17 ± 0.09 0.04–0.42	0.30 ± 0.22 0.05–1.06	1.08 ± 0.80 0.11–2.82
K ⁺	0.92 ± 0.75 0.03–3.66	1.08 ± 0.71 0.14–3.14	0.66 ± 0.47 0.20–2.47	1.13 ± 0.90 0.03–3.66	0.81 ± 0.77 0.05–2.53
Mg ²⁺	0.16 ± 0.13 0.02–1.04	0.24 ± 0.20 0.03–1.04	0.07 ± 0.03 0.02–0.16	0.16 ± 0.07 0.06–0.31	0.18 ± 0.09 0.06–0.45
Ca ²⁺	1.6 ± 1.5 0.2–11.3	2.6 ± 2.2 0.2–11.3	0.6 ± 0.3 0.2–1.7	1.7 ± 1.0 0.5–4.2	1.5 ± 0.9 0.5–4.0
Al	1.8 ± 1.5 0.1–6.9	2.5 ± 1.7 0.3–6.6	0.7 ± 0.4 0.1–2.0	2.0 ± 1.4 0.3–6.7	2.1 ± 1.5 0.7–6.9
OC	16.9 ± 10.0 5.9–58.6	13.7 ± 4.4 5.9–23.7	11.1 ± 1.8 7.4–16.6	17.8 ± 5.6 7.5–26.2	24.9 ± 15.6 8.5–58.6
EC	5.0 ± 4.4 0.6–28.1	2.8 ± 1.1 0.6–5.8	4.2 ± 1.2 1.5–6.8	5.3 ± 2.8 1.3–12.1	7.5 ± 7.4 1.2–28.1

$$w_{ij} = \begin{cases} 1.00 & 80 < n_{ij} \\ 0.70 & 20 < n_{ij} \leq 80 \\ 0.42 & 10 < n_{ij} \leq 20 \\ 0.05 & n_{ij} \leq 10 \end{cases}$$

The study domain was in the range of 30–60° N, 75–130° E. The resolution was 0.5° × 0.5°.

3 Results

3.1 Annual average

Table 1 provides a statistical summary of the obtained data on atmospheric concentrations for PM_{2.5}, Al (a tracer of aluminosilicate dust), water-soluble ions, OC, and EC during the sampling period. The annual mean PM_{2.5} concentration reached 135 ± 63 $\mu\text{g m}^{-3}$. This mean value is nearly three times higher than that (35 $\mu\text{g m}^{-3}$) of the interim target-1 standard for annual mean PM_{2.5} recommended by the WHO. The level of PM_{2.5} in Beijing is much higher than other mega-cities around the world. In comparison with that of domestic cities, PM_{2.5} seems to display a spatial tendency, increasing northward and decreasing southward (Zhang et al.,

2012b). Such a spatial pattern may be related to the low rainfall and high dust in northern China (Qian et al., 2002; Qian and Lin, 2005). According to Wang et al. (2008), PM_{2.5} concentrations in winter were much higher than in summer in 2001 to 2002. However, such a trend seemed to be reversed in 2005 to 2006, with rather higher concentrations in summer.

For the ionic concentrations, SO₄²⁻ ranked the highest among the water-soluble ions analyzed, with an annual mean of 13.6 ± 12.4 $\mu\text{g m}^{-3}$, followed by NO₃⁻ (11.3 ± 10.8 $\mu\text{g m}^{-3}$), NH₄⁺ (6.9 ± 7.1 $\mu\text{g m}^{-3}$), Ca²⁺ (1.6 ± 1.4 $\mu\text{g m}^{-3}$), Cl⁻ (1.4 ± 2.2 $\mu\text{g m}^{-3}$), K⁺ (0.92 ± 0.75 $\mu\text{g m}^{-3}$), Na⁺ (0.46 ± 0.55 $\mu\text{g m}^{-3}$), and Mg²⁺ (0.16 ± 0.13 $\mu\text{g m}^{-3}$). Such levels of mean concentrations are rather comparable with those measured in many Chinese cities such as Shanghai, Tianjin, Jinan, and Guangzhou (Yao et al., 2002; Tao et al., 2009; Gao et al., 2011; Gu et al., 2011). On average, the combination of SO₄²⁻, NO₃⁻, and NH₄⁺, which could be regarded as secondary inorganic aerosols, constituted the majority (88 %) of the total ionic concentrations, consistent with earlier studies (Yao et al., 2002; Duan et al., 2003). The annual mean concentrations of OC and EC reached up to 17.0 ± 10.0 and 5.0 ± 4.4 $\mu\text{g m}^{-3}$, respectively. Such levels are close to those observed for

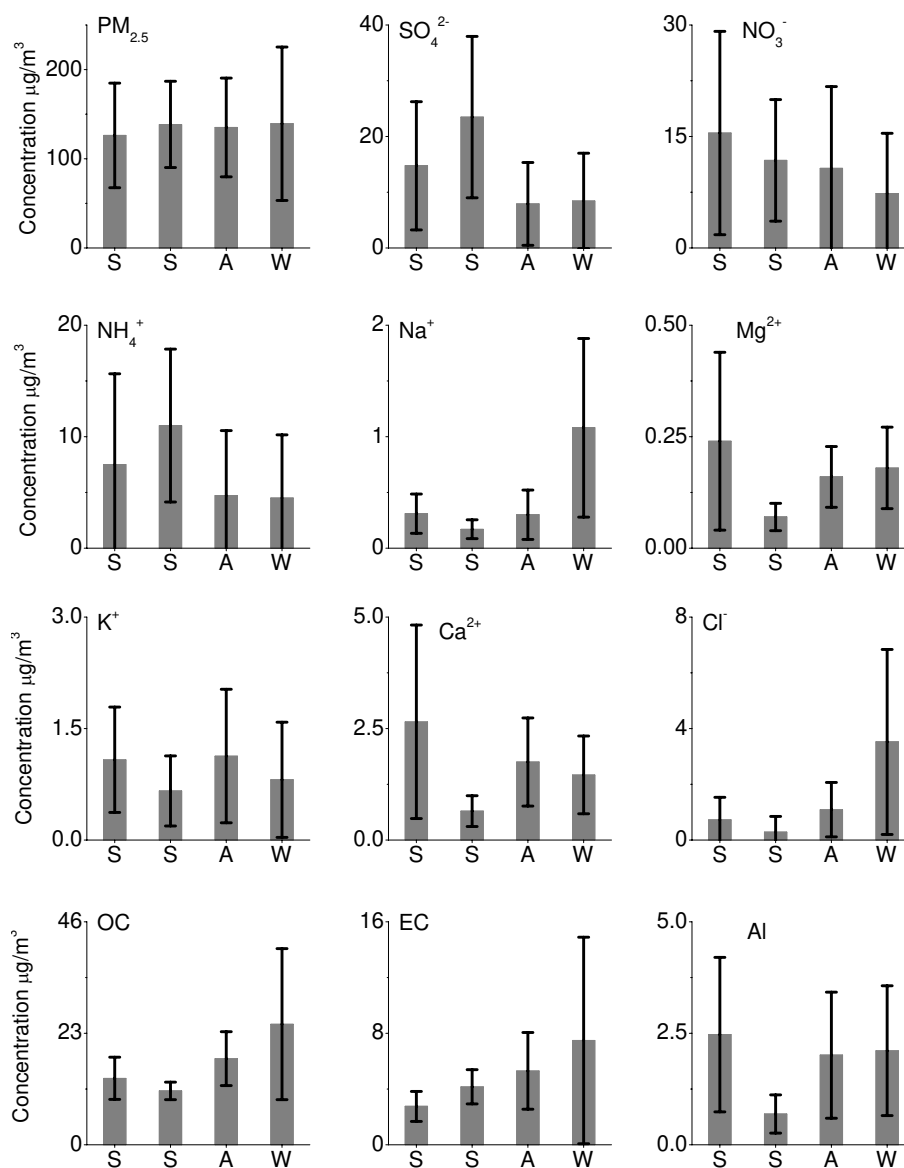


Fig. 2. Seasonal variations of PM_{2.5} mass concentration and associated species, including SO₄²⁻, NO₃⁻, NH₄⁺, Na⁺, Mg²⁺, K⁺, Ca²⁺, Cl⁻, OC, EC, and Al concentrations. Shown here are the mean and one standard deviation for each bar.

regional sites across China ($16.1 \pm 5.2 \mu\text{g m}^{-3}$ for OC and $3.6 \pm 0.93 \mu\text{g m}^{-3}$ for EC) by Zhang et al. (2008), who extensively measured carbonaceous aerosols around China; however, such concentrations are approximately half of those observed at urban sites ($33.1 \pm 9.6 \mu\text{g m}^{-3}$ for OC and $11.2 \pm 2.0 \mu\text{g m}^{-3}$ for EC) by Zhang et al. (2008).

3.2 Seasonality

As illustrated in Fig. 2, the seasonality of PM_{2.5} and these primary species were characterized by distinctive features. The seasonality of PM_{2.5} concentration was not very evident and typical, with nearly equal concentrations of around $140 \mu\text{g m}^{-3}$ in summer, autumn, and winter and a relative

minimum ($\sim 125 \mu\text{g m}^{-3}$) in spring. The minimum concentration typically occurs in summer because precipitation in Beijing is usually concentrated at that period (Fig. S2, Supplement). However, this is not the case, because the maximum concentrations of secondary sulfate and ammonium were observed in summer, arising from strong photochemistry and accounting for a large proportion ($\sim 25\%$) of PM_{2.5} in Beijing (Yao et al., 2003). This suggestion can be supported by the fact that ambient temperatures through February to October 2009 in Beijing are higher than the climatology by $\sim 1^\circ\text{C}$ or higher (Fig. S2, Supplement), while precipitation in July 2009 (i.e., 197 mm) is even slightly larger than the climatology (185 mm), demonstrating that

the photochemical effect might overwhelm the precipitation scavenging effect for fine aerosol pollutants.

In contrast to PM_{2.5}, sulfate and ammonium revealed a typical seasonality with higher concentrations in spring and summer and lower concentrations in autumn and winter, consistent with the seasonal variability of AOT (Xia et al., 2006). The summertime maximum concentrations of sulfate and ammonium were 24 and 12 $\mu\text{g m}^{-3}$, respectively, which were higher than those in Beijing before 2003 (~ 15 and $\leq 10 \mu\text{g m}^{-3}$, respectively) (He et al., 2001; Duan et al., 2006; Wang et al., 2005) but rather comparable to those observed in the last few years (Okuda et al., 2011; Song et al., 2012). By contrast, the wintertime concentration of sulfate ($8.5 \mu\text{g m}^{-3}$) was significantly reduced compared with earlier literature data (He et al., 2001; Hu et al., 2002; Wang et al., 2005). The decrease in wintertime sulfate concentration seemed to result from the effective control of SO₂ emissions over China in the recent years (Itahashi et al., 2012), particularly from coal combustion (Hao et al., 2005). High summertime sulfate concentration is ascribed to enhanced photochemistry during summer, and relatively high humidity accelerates the conversion rate of SO₂ to the particulate form (Yao et al., 2003). However, the precursor SO₂ concentrations are much higher in winter (Fig. S3, Supplement) because of higher emission at that time (Zhang et al., 2009b). One might further conclude that in Beijing, photochemistry plays a more vital role in the sulfate aerosol formation and variability than the change in precursor SO₂ emission as well as rain scavenging process. In the present study, artificial biases, particularly of nitrate and ammonium, possibly occurred during sampling because no denuder and/or back-up filter was used to trap ammonia and nitric acid (Pathak et al., 2004). The maximum concentration ($15.5 \mu\text{g m}^{-3}$) of nitrate was observed in spring rather than summer, which was different from that of sulfate (Wang et al., 2008). This observation may be ascribed to the volatility of ammonium nitrate, which is one of the main chemical forms of nitrate associations revealed by the ionic relationships. Thus, ammonium nitrate could evaporate at relatively high temperature. Besides, there are distinct emission sources for their respective precursor gases, SO₂ and NO_x. The minimum concentration of nitrate was observed in winter. Nitrate levels could be a function of various factors in terms of emissions, such as vehicular exhaust, coal combustion, and biomass burning, and complex chemical processes with respect to photochemistry, heterogeneous reaction, renoxification, and gas-aerosol equilibrium. The overall trend of NO_x emission in China is increasing primarily due to persistently increasing energy demand for industrial development and transportation, though control measures for NO_x emissions have been implemented in coal-fired power plants (Zhao et al., 2013).

Crustally derived ions and elements, such as Mg²⁺, Ca²⁺, and Al, waxed in the spring and waned in the summer, followed by significant increases toward autumn and winter. Such levels of seasonal mean concentrations and seasonality

are consistent with those observed in previous studies (Duan et al., 2006; Wang et al., 2005), which are related to dust storms and anthropogenic and fugitive dust. The seasonal concentrations of Na⁺ and Cl⁻ peaked in winter, consistent with Hu et al. (2002) and Wang et al. (2005). The seasonality of Mg²⁺ was distinguishable from that of Na⁺, demonstrating the difference in their dominant sources. K⁺ had relatively higher concentrations in both spring and autumn than in summer and winter, which was closely associated with the agricultural burning around Beijing (Zheng et al., 2005). Such seasonality was also found by some previous studies (He et al., 2001; Zheng et al., 2005), but distinct from other previous studies, in which winter often had the highest concentration (Duan et al., 2006; Wang et al., 2005). Specifically, the highest levoglucosan, which is suggested to serve as an excellent tracer of biomass burning pollutants relative to K⁺, has been measured in autumn (He et al., 2006), although no spring sample was measured in that study.

Both OC and EC had similar seasonal patterns of waxing in winter and waning in spring (for EC) or summer (for OC). Zhang et al. (2008) observed a persistently common seasonality for both OC and EC at 18 background, regional, and urban stations in China, i.e., a maximum in winter and a minimum in summer. The seasonality may be governed by the variability in emission strengths and meteorology. For instance, lower-molecular weight semi-volatile organic compounds are mostly in gaseous phase at high temperature in summer (Yassaa et al., 2001). The OC/EC mass ratio 2.0 indicates the presence of secondary organic matter (Chow et al., 1996). In this work, the OC/EC ratios mostly fell within the range of 2–5, with mean ratios of 4.8, 2.5, 3.0, and 2.7 in spring, summer, autumn, and winter, respectively (Fig. 3). These figures are very similar to those observed in previous studies (Duan et al., 2006; Zhang et al., 2008), which suggests the relative domination of secondary organic aerosols in spring but of primary sources in other seasons (Zhang et al., 2008). Another reason for the relatively higher spring-time OC/EC mass ratio may be the open biomass burning source, consistent with higher K⁺ in the spring (Fig. 2), as the aerosols from open biomass burning are generally characterized by elevated OC/EC ratios (Cao et al., 2007).

3.3 Stoichiometric analyses of cations and anions

Note that equivalent concentrations ($\mu\text{eq m}^{-3}$) are used throughout this section. Figure 4 shows the scatter plots of (a) Mg²⁺ vs. Na⁺, (b) Mg²⁺ vs. Ca²⁺, (c) Cl⁻ vs. Na⁺, and (d) Cl⁻ vs. K⁺. Figure 5 shows the scatter plots of (a) total cations vs. total anions, (b) NH₄⁺ vs. SO₄²⁻, (c) NH₄⁺ vs. [SO₄²⁻ + NO₃⁻], (d) [NH₄⁺ + Ca²⁺] vs. [SO₄²⁻ + NO₃⁻], (e) [NH₄⁺ + Ca²⁺ + Mg²⁺] vs. [SO₄²⁻ + NO₃⁻], and (f) [NH₄⁺ + Ca²⁺ + Mg²⁺] vs. [SO₄²⁻ + NO₃⁻ + Ex-Cl⁻]. Ex-Cl⁻ is the excessive Cl⁻, defined as the excessive amount of Cl⁻ relative to the amount sea salt can sustain with Na⁺

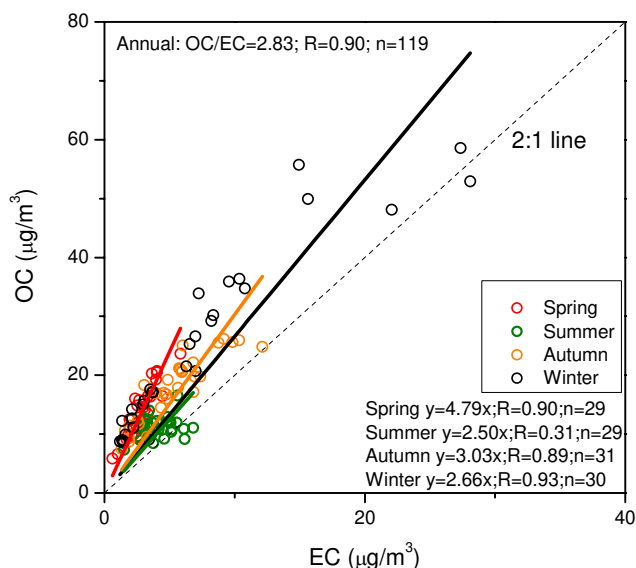


Fig. 3. Scatter plot showing the correlation between OC (y axis) and EC (x axis) in PM_{2.5} collected from Beijing. Different symbols denote the four seasons. Linear regression equations are given in the annual and seasonal cases.

as the tracer of sea salt: $Ex-Cl^- = Cl^- - [Na^+] \times 1.17$, where 1.17 is the typical Cl^-/Na^+ equivalent ratio of average seawater (Chester, 1990). If the resulting $Ex-Cl^-$ is negative, then no Cl^- excess exists. In other words, Cl^- is totally contributed by sea salt and is even depleted by heterogeneous reactions. Nevertheless, total Na^+ does not necessarily originate from sea salt alone, but could partially come from dust. The resultant biases are hence likely insignificant.

Figure 4a illustrates that Mg^{2+} mostly comes from non-sea salt sources, except in wintertime, because the regression slopes that represent the Mg^{2+}/Na^+ ratios (1.56, 0.70, 0.68, and 0.25 for spring, summer, autumn, and winter, respectively) are clearly deviated from the ratio (0.23) of average seawater (Chester, 1990). Instead, the dominant source of Mg^{2+} is mineral dust, mainly carbonate minerals (Li et al., 2007), as reflected by the good correlations (0.97, 0.83, 0.90, and 0.80 for spring, summer, autumn, and winter, respectively) between Mg^{2+} and Ca^{2+} (Fig. 4b). Similarly, most Cl^-/Na^+ ratios (1.70, 1.21, 2.51, and 2.25 for spring, summer, autumn, and winter, respectively) in PM_{2.5} are larger than the mean ratio (1.17) of seawater, except the summertime samples (Fig. 4c). This difference indicates the dominance of the non-sea salt sources, of which the most likely contributor of Cl^- is coal combustion (Yao et al., 2002), particularly in winter when Cl^- and SO_2 are maximal (Figs. 2 and S2, Supplement). In summer, air masses are dominated by the southerly monsoon from Bohai (as supported by the trajectories below), leading to a mean Cl^-/Na^+ ratio close to that of average seawater. Moreover, the correlations between K^+ and Cl^- largely varied with the seasons, with

better correlation and higher ratios in autumn and winter and moderate correlations and lower ratios (less than unity) in summer and spring (Fig. 4d). Thus, the results suggested that K^+ was not present in chemical form KCl at high temperature, but as K_2SO_4 (Pósfai et al., 2004). By contrast, low temperature in winter may favor the presence of KCl.

Furthermore, the ratio of total cation concentration to total anion concentration is averaged at near unity throughout the year (Fig. 5a), which indicates excellent charge balance in PM_{2.5} and high data quality. Figure 5b shows good correlations between NH_4^+ and SO_4^{2-} for the annual data set, with ratios (represented by the slope of the linear regression line) between 1.25 and 1.77 (all higher than unity). These good correlations reveal the dominance of $(NH_4)_2SO_4$ (Ianniello et al., 2011), rather than NH_4HSO_4 and the possible full neutralization of SO_4^{2-} by NH_4^+ throughout the year. We further considered the combination of NO_3^- and SO_4^{2-} in this charge balance analysis (Fig. 5c). The resulting NH_4^+ to $[SO_4^{2-} + NO_3^-]$ ratios ranged from 0.83 to 0.94 (all lower than unity), which demonstrates the presence of NH_4NO_3 in the fine-mode aerosols. Moreover, ratios lower than unity suggest that nitrate may be present in other chemical forms than NH_4NO_3 . Heterogeneous reactions between NO_x (and its products, such as HNO_3 and N_2O_5) and dust carbonate are often observed in northern China (Li and Shao, 2009). Accordingly, we examined the correlations of $[NH_4^+ + Ca^{2+}]$ vs. $[SO_4^{2-} + NO_3^-]$ (Fig. 5d) and of $[NH_4^+ + Ca^{2+} + Mg^{2+}]$ vs. $[SO_4^{2-} + NO_3^-]$ (Fig. 5e), given that the good correlations between Mg^{2+} and Ca^{2+} suggest the possible existence of water-soluble Mg in reacted carbonate dust. These ions are strongly correlated throughout the year, with high coefficients (all 0.99 or higher) and slopes of regression lines around unity. These correlations indicate that nitrate is partly present in $Ca(NO_3)_2$ and $Mg(NO_3)_2$, not just in NH_4NO_3 . We assumed that fine-mode sulfate is exclusively associated with ammonium and that Na^+ is present only in the associated NaCl. However, these assumptions may not always be true because Na_2SO_4 is observed in dust particles from dried lakes in northern China (Zhang et al., 2009a) and NaCl could react with nitric acid to form $NaNO_3$ via heterogeneous reaction (Hsu et al., 2007). Therefore, based on the aforementioned equivalent interrelationships and assumptions, we quantitatively estimated that the former two chemical forms ($Ca(NO_3)_2$ and $Mg(NO_3)_2$) represent $\sim 20\%$ of the total nitrate and that NH_4NO_3 is the dominant association accounting for the remaining $\sim 80\%$. The elevated Cl^-/Na^+ ratio (> 1.17) shows that excessive Cl^- seemed to be attributed to coal combustion rather than sea salt particles from dried salt lake sediment. The addition of excessive Cl^- (Fig. 5f) insignificantly changed the correlations of positive and negative charges. Nevertheless, we noted that in wintertime, the equivalent ratio improved from 1.10 to 1.00, which indicates the presence of chloride salts such as KCl, $CaCl_2$, and $MgCl_2$ other than NaCl at low ambient temperature (Ianniello et al.,

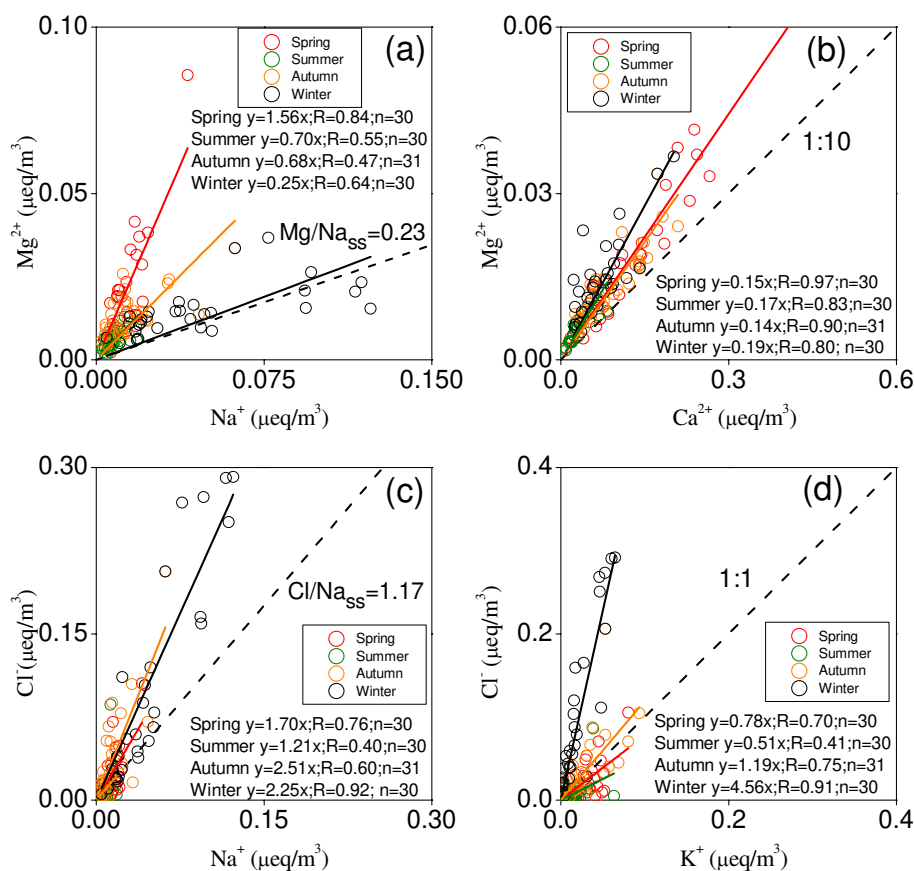


Fig. 4. Correlations between certain cations and anions: (a) Mg²⁺ versus Na⁺, (b) Mg²⁺ versus Ca²⁺, (c) Cl⁻ versus Na⁺, and (d) Cl⁻ versus K⁺. Four seasons are considered and represented by different color symbols. Linear regression equations are also given. In (a) and (c), the dashed lines indicate the Mg/Na and Cl/Na equivalent concentration ratios in seawater (sea salt), respectively.

2011). KCl may have originated from biomass burning, and CaCl₂ and MgCl₂ could have been formed through heterogeneous reactions between the dust carbonate and HCl emitted from coal combustion.

3.4 Chemical mass closure

By employing the methods in Sect. 2.6, we constructed the CMC of PM_{2.5} in Beijing on a seasonal and annual basis. The reconstructed PM_{2.5} mass concentrations were compared with the gravimetric PM_{2.5} mass concentrations, as shown in Fig. 6, which shows a good correlation with one another in each season and throughout the year. However, the ratios seasonally changed, with higher ratios of 0.82 in spring and 0.75 in winter and lower ratios of 0.59 in summer and 0.68 in autumn. The proportions of all specific components in PM_{2.5} together with the unidentified constituents as a whole are schematically illustrated by five pie charts for the four seasonal and annual cases (Fig. 7). Overall, the major components are secondary inorganic aerosols (combination of sulfate, nitrate, and ammonium), mineral dust, and POM, which account for each ~ 20 %, albeit with seasonal

variations. The minor components include EC, chloride salt, potassium salt, and TEO, each of which represents less than 5 %. Specifically, the proportions of mineral dust are maximal (33.4 %) in spring, minimal (only 8.2 %) in summer, and intermediate (23.6 and 28.9 %) in the other two seasons, consistent with the tendency of seasonal Al concentrations. The totals of secondary inorganic species (SO₄²⁻, NO₃⁻, and NH₄⁺) have the largest proportion (27 to 30 %) in spring and summer and a minimal percentage (< 15 %) in autumn and winter. However, sulfate peaks were noted in summer (15.4 %), whereas nitrate peaks were observed in spring (11.1 %). Ammonium decreased from around 5 to 7 % in spring and summer to half (2 to 3 %) in autumn and winter. The POM fractions largely varied as follows: summer (14.7 %) < spring (18.9 %) < autumn (22.1 %) < winter (28.6 %). EC and chloride salt exhibited the largest proportions (5.2 and 3.3 %, respectively) in winter. Potassium salt and TEO had slightly higher proportions in spring and autumn than in summer and winter.

Both the primary and secondary components of PM_{2.5} in Beijing are equally important, albeit with seasonal variability, which is typical of PM pollution in China (Shao et al.,

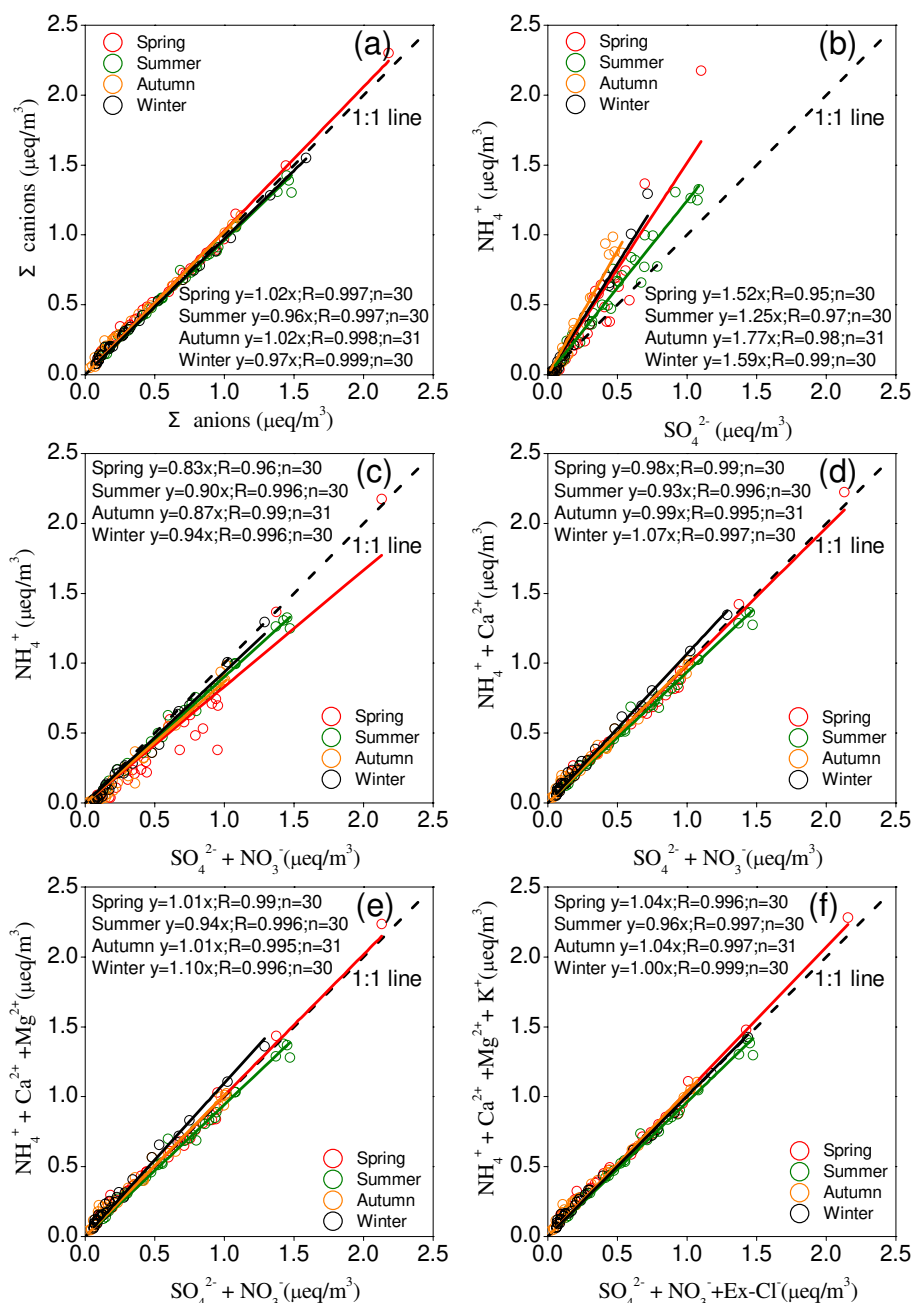


Fig. 5. Same as Fig. 4, but for (a) total cations versus total anions, (b) NH_4^+ versus SO_4^{2-} , (c) NH_4^+ versus $[\text{SO}_4^{2-} + \text{NO}_3^-]$, (d) $[\text{NH}_4^+ + \text{Ca}^{2+}]$ versus $[\text{SO}_4^{2-} + \text{NO}_3^-]$, (e) $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$ versus $[\text{SO}_4^{2-} + \text{NO}_3^-]$, and (f) $[\text{NH}_4^+ + \text{Ca}^{2+} + \text{Mg}^{2+}]$ versus $[\text{SO}_4^{2-} + \text{NO}_3^- + \text{Ex-Cl}^-]$. The detailed definition of Ex-Cl⁻ (Excessive-Cl⁻) can be found in the text. The 1 : 1 lines are given for comparison.

2006). In general, given that the seasonal variability in PM_{2.5} mass concentrations is relatively small, temporal trends in the proportions of each component of PM_{2.5} resemble the atmospheric concentrations of their corresponding chemical species. The likely factors for such seasonality are partially addressed in Sect. 3.2 and discussed in detail in the following two sections.

On average, the unidentified components reached 28.6 % of the total PM_{2.5}. They also showed seasonal variability, with the smallest (15.9 %) in spring when dust was prevalent, and the largest (42.5 %) in summer when secondary inorganic aerosol formation was favorable. Such high uncertainties in the CMCs were caused by the water absorption of water-soluble components in the weighing environment, though relative humidity was controlled (Speer et al., 1997;

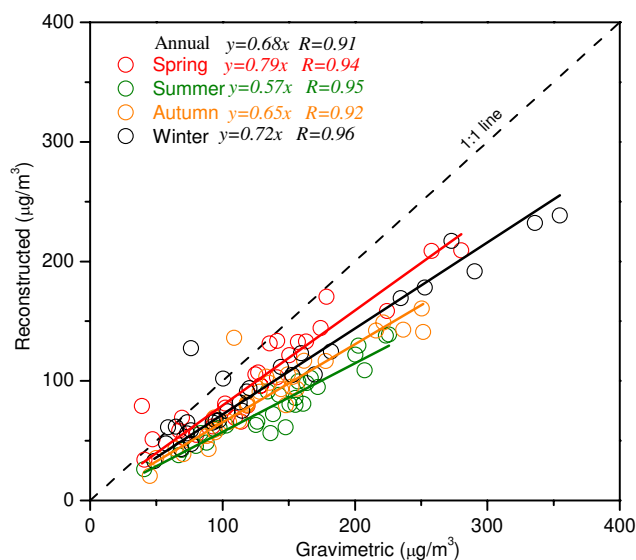


Fig. 6. Scatter plot showing the correlations between the PM_{2.5} mass concentrations reconstructed from the chemical mass balance method and obtained from gravimetric measurement, which are presented on a seasonal basis. Linear regression lines are shown with equations for the four-season period, along with the annual case.

Tsai and Kuo, 2005). The absorption likely led to positive biases in PM_{2.5} concentrations. Alternatively, such uncertainties may be partly due to the volatilization of NH₄NO₃ and organic matter, particularly in summer and autumn during the storage of the weighted samples prior to extraction, which may have resulted in negative biases in the specific components. Another likely reason for the non-match of the reconstructed and gravimetric mass concentrations is the varying factors used in transferring a given analyzed species (e.g., OC and Al) to a certain component (e.g., POM and mineral soils) (Rees et al., 2004; Hsu et al., 2010a; Yan et al., 2012). For example, a few studies adopted a factor of only 1.4 for converting OC content to organic matter (Duan et al., 2006; Song et al., 2007; Guinot et al., 2007). Another study obtained a much higher POM/OC mean ratio over China (Xing et al., 2013) of up to 1.92 ± 0.39 based on a mass balance method. If we adopt this higher ratio, the unidentified percentage would be reduced by 3%. In the present study, the EF_{crust} of Ca averages at 2.6, which shows its enrichment relative to average crust composition. In Beijing, fine-mode Ca-rich dust is partly attributed to construction activities. Therefore, we may underestimate carbonate abundance in the mineral component estimated from Al concentration alone (Guinot et al., 2007). This may have resulted in the underestimation ($\sim 2\%$) of the total mass reconstructed.

Table 2. Relative contributions from six identified sources of PM_{2.5} in Beijing within the one-year and four-season periods.

Source	Spring	Summer	Autumn	Winter	Annual
Soil dust	23 %	3 %	18 %	16 %	15 %
Coal combustion	5 %	1 %	7 %	57 %	18 %
Biomass burning	19 %	6 %	17 %	7 %	12 %
Traffic and waste incineration emission	5 %	4 %	4 %	2 %	4 %
Industrial pollution	14 %	32 %	42 %	12 %	25 %
SIA	34 %	54 %	13 %	6 %	26 %

4 Discussion

4.1 Source identification and apportionment

By utilizing the PMF model with the obtained full data set as input data, we identified six main sources: mineral dust, biomass burning, coal combustion, traffic emissions plus waste incineration, industrial pollution, and secondary inorganic aerosol. Table 2 summarizes the source apportionment results of the relative contributions from each identified source to the PM_{2.5} on both a seasonal and annual basis in Beijing. These sources have average contributions of 15, 18, 12, 4, 25, and 26% (Table 2). Figures 8 and 9 show the modeled source profiles and the time series of modeled concentrations for each identified main source. Again, the relative dominance of each identified source largely varies with changing seasons, which is roughly consistent with the CMC results. For mineral dust, only one of the six sources mostly dominated by nonvolatile substance, its proportions (e.g., annual mean $\sim 20\%$) and relative order in the four seasons are consistent with the CMC results, with the highest contribution in spring, the lowest contribution in summer, and intermediate contribution in autumn and winter. This consistency indirectly verifies the reliability of the PMF results. The other five sources all appear to be related to high-temperature activities and/or photochemical processes and involved with volatile species. We then compared the contribution percentages of the secondary inorganic aerosol (SIA) with the CMC results as this source is also identified in the CMC analyses. Apparently, the percentages of SIA in the four seasons differ from those obtained by CMC in terms of the values (e.g., 6 to 54% versus 14 to 33% for SIA), although the seasonal trends are quite similar. Thus, CMC method only offers chemical characterization instead of source apportionment. The PMF model provides real information on sources of aerosol speciation.

Moreover, we evaluated the distinctive characteristics of each modeled source profile according to well-accepted knowledge. The first component, mineral dust, is typically characterized by high crustal elements, such as Al, Ca (Ca²⁺), Fe, Mg, K, and Ti. Ca content is higher than Al, which indicates Ca-rich dust. OC content is close to $\sim 12\%$, which suggests resuspended dust because of the presence

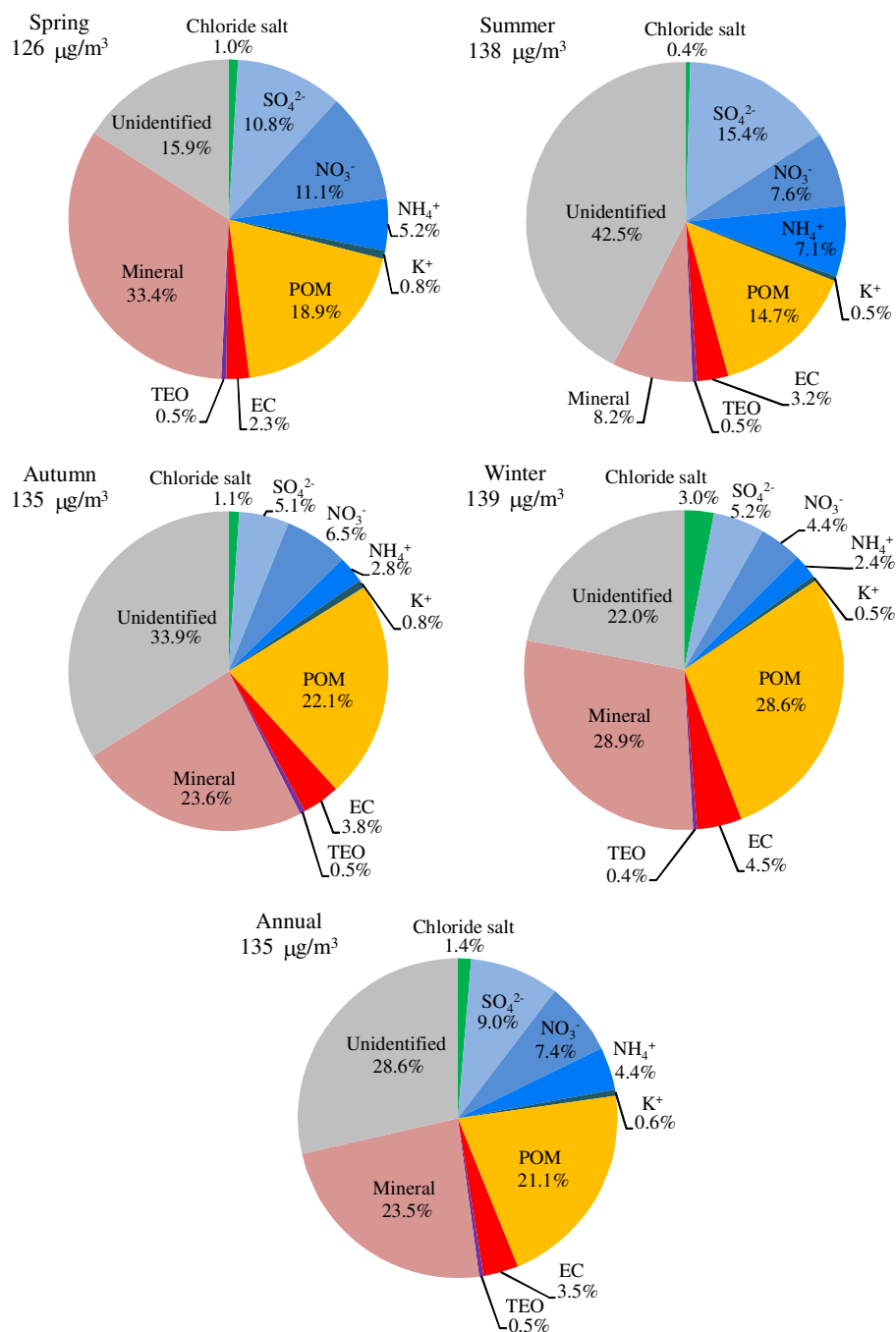


Fig. 7. Pie-charts showing the constructed chemical mass closures for PM_{2.5} in Beijing: (a) spring; (b) summer; (c) autumn; (d) winter; and (e) annual. The components include mineral dust, secondary aerosol ions (sulfate, nitrate, and ammonium), POM, EC, trace element oxides (TEO), chloride salt, and biomass burning-derived potassium. Other than the identified components, unidentified fractions comprise a significant portion of PM_{2.5}.

of high levels of OC (Watson and Chow, 2001). Thus, this source possibly mixes desert/loess dust, anthropogenic construction dust, fugitive dust, and resuspended road dust. Construction activities are prevalent in the urban cities of China, and no effective measures for dust control are implemented. Therefore, calcium is used as an indicator element

for construction dust in Beijing (Zhang and Iwasaka, 1999). The second source, coal combustion, is characterized by elevated Cl⁻ associated with high Na⁺, OC, and EC. Extraordinarily high Cl⁻ associated with fine aerosols in winter is a distinctive feature in Beijing and even around inland China, which is ascribed to coal combustion (Yao et al., 2002; Wang

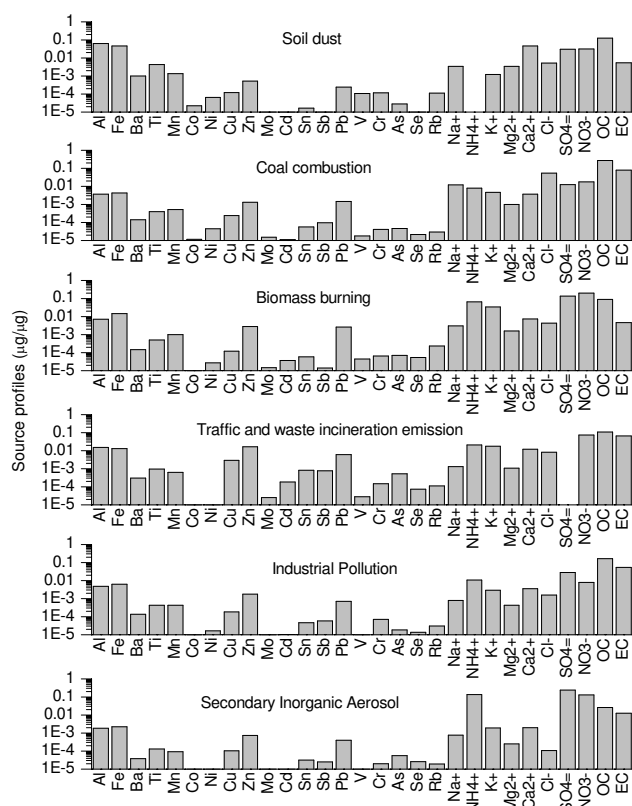


Fig. 8. Profiles of six sources identified from the PMF model, including soil dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol (from the upper to the lower panels).

et al., 2008). Coal combustion is the predominant source of fine aerosols over China (Yao et al., 2009), which has resulted in severe air pollution problem not only locally, but also regionally and globally; for instance, it alone contributes more than 10 % (268 Mg) of the global anthropogenic mercury emission (2319 Mt) annually (Pirrone et al., 2010). Besides, coal fly ash could be one of the main contributors of aerosol Pb in China as they contain abundant Pb (Zhang et al., 2009c). Sodium has also been found to be enriched in fine particulates from coal combustion (Takuwa et al., 2006). Wang et al. (2008) and Zhang et al. (2009a) attributed the observed high Na and Cl to the presence of Na₂SO₄ and NaCl that may be originated from dried lake salt sediment in Inner Mongolia, a non-local dust source. However, as discussed, salt lake aerosols alone cannot account for such strikingly high Cl⁻ in winter, which suggests that coal combustion is the most likely dominant source. Different investigations have obtained significantly different contributions of coal combustion to PM_{2.5} in Beijing, which range from 7 to 19 %. Yao et al. (2009) concluded that the likely fraction ranges between 15 and 20 %. Nevertheless, previous studies have not considered that other main identified sources, such

as secondary inorganic and organic aerosols, have contributions from coal combustion.

The third source, biomass burning, is characterized by high K (K⁺), which is an excellent tracer of biomass-burning aerosols (Cachier and Ducret, 1991; Watson and Chow, 1998), and by rich Rb, OC, and NO₃⁻. Biomass burning has higher contributions in spring and autumn than in summer and winter, consistent with cultivation in spring and harvest in autumn (Duan et al., 2004). The fourth source is a mixed source of traffic and metropolitan incineration emissions, which is characterized by high NO₃⁻, EC, Cu, Zn, Cd, Pb, Mo, Sb, and Sn. These aerosol species are all enriched in vehicular and/or waste incineration emissions (Lee et al., 1999; Alastuey et al., 2006; Birmili et al., 2006; Marani et al., 2003; Dall'Osto et al., 2013; Tian et al., 2012). For instance, Wählin et al. (2006) observed that traffic-generated aerosol particles are rich in Cu, Zn, Mo, and Sb. Christian et al. (2010) analyzed the aerosol particles emitted from garbage burning, which are rich in Zn, Cd, Sb, and Sn. Leaded gasoline was phased out in 1997 in Beijing and in 2000 in the rest of China. Coal burning was then suggested as the most important source of Pb aerosols in China (Mukai et al., 2001). However, Widory et al. (2010) argued that in Beijing, metal-refining plants are the dominant sources of aerosol Pb, followed by thermal power stations and other coal combustion sources.

The fifth source is industrial pollution, which is characterized by high contents of OC, EC, Zn, Mn, and Cr. This source may also be involved with secondary organic aerosols. Coal is the primary energy source commonly used in industries in China. Both coal combustion and vehicle emissions are the main sources of primary OC (Zhang et al., 2007; Cao et al., 2011). However, Zhang et al. (2008) estimated that secondary OC represents more than half of the measured OC at regional sites (~67 %) and urban sites (~57 %), which is higher than those reported by Cao et al. (2007) (i.e., 30 to 53 %). Therefore, industrial pollution could act as a vital source of carbonaceous aerosols, which seems to be widely ignored. Furthermore, given that Zn and Cr contents are high, this source may be relevant to smelters and metallurgical industries (Dall'Osto et al., 2013). The sixth source is relevant to secondary inorganic aerosols, which are typically characterized by remarkable SO₄²⁻, NO₃⁻, and NH₄⁺. Certain identified sources, such as biomass burning, coal combustion, vehicle exhausts, and waste incineration, can also contribute to secondary inorganic and organic aerosols through the emission of their precursor gases.

Figure 9 illustrates the time series of daily concentrations contributed by each identified source. To examine if the results are reasonable, we compared the modeled seasonality of each source with the observed seasonality of the specific chemical species that could represent respective contribution sources (Fig. 2). For instance, we compared the maximal and minimal contributions of mineral dust in spring and summer, respectively, which are consistent with the seasonality

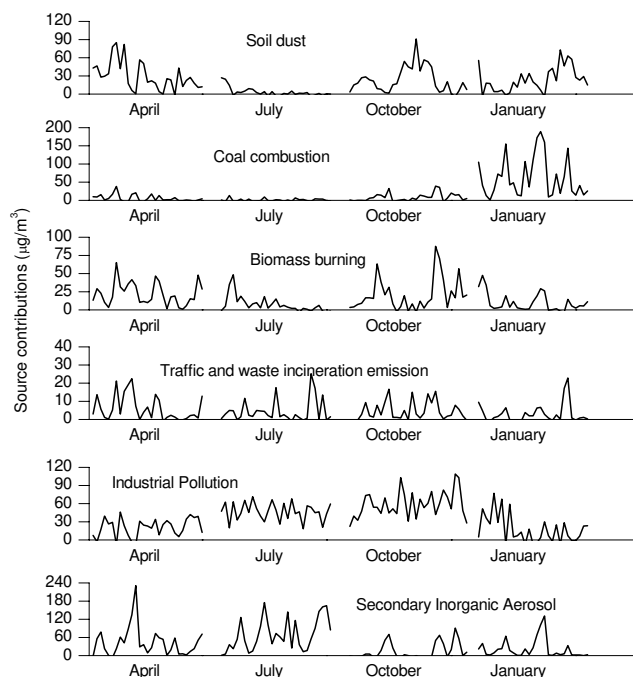


Fig. 9. Time series of daily contributions from each identified source, including soil dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and secondary inorganic aerosol (from the upper to the lower panels) during the study period between April 2009 and January 2010.

of aerosol Al. Dust storms are essentially responsible for springtime dust aerosols, whereas in autumn and winter, fugitive dust from construction and the resuspension of street dust are the main contributors. Obviously, the reconstructed time series of daily concentrations from coal combustion reveals a pronounced wintertime maximum, consistent with those of aerosol Cl⁻ (Fig. 2) and even gaseous SO₂ (Fig. S2, Supplement). Moreover, the time series of biomass burning contributions show relatively higher concentrations in spring and autumn and lower concentrations in summer and winter, consistent with the seasonality of K⁺. For traffic and waste-burning emissions, the resulting time series do not reveal evident seasonality, corresponding with the seasonality of nitrate and some trace metals, such as Pb, Cu, Sb, and Cd (Fig. S4, Supplement). Industrial pollution has higher contributions in summer and autumn, possibly corresponding with the seasonality of Zn and Cr. However, such seasonality is inconsistent with that of OC, with a wintertime maximum, because OC may be from various sources, including the former five sources identified. Coal combustion has the largest contribution in winter, and low temperature in winter facilitates the formation of secondary organic aerosols. SIA has higher contributions in summer and spring, mirroring the seasonality of sulfate, nitrate, and ammonium. This result is definitely related to the photochemistry that accounts for SIA formation. The formed SIA species may not appear in their

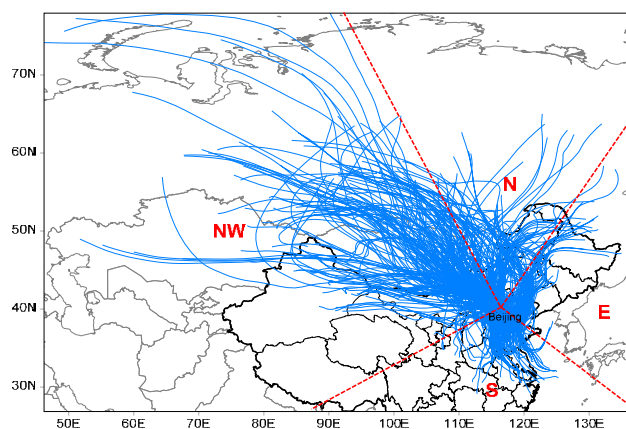


Fig. 10. Analytical results of the 48 h air mass back trajectories at 100 m elevation during the sampling periods, which were run four times per day. Four regions were defined based on the trajectory clustering results, i.e., NW, N, E, and S regions.

original emission sources (i.e., coal combustion, biomass burning, traffic exhausts, waste incineration, and industrial pollution), but in the SIA component. Based on the PMF results and chemical data in January and August 2004, Song et al. (2007) found that the most predominant sources of PM_{2.5} are coal combustion in winter and secondary aerosols in summer, along with other significant sources, such as motor vehicle emissions, road dust, and biomass burning. The PMF-modeled results seem to be promising because the corresponding time series of each source's contribution are very consistent with the observations.

4.2 Regional sources deduced from trajectory and PSCF analyses

The regional sources and transport of air pollutants exert a profound impact on local air quality in Beijing (e.g., Wang et al., 2004). To address this issue, both trajectory clustering and PSCF methods were employed. The 48 h back trajectories starting at 100 m from Beijing were computed by using the HYSPLIT model of NOAA (<http://www.arl.noaa.gov/ready.html>). Four clusters were made (Fig. 10): northwestern (including western, NW), northern (N), eastern (from north-eastern to southeastern, E), and southern (S) directions. The NW cluster was further differentiated into two types, i.e., fast (NW_f) and slow (NW_s), according to the motion speed ($\leq 7 \text{ m s}^{-1}$ for NW_s and $> 7 \text{ m s}^{-1}$ for NW_f), and distance of air parcels. The classification is consistent with the spatial distribution of fine AOD retrieved by remote sensing (Fig. 1).

Table 3 summarizes the percentages of each trajectory cluster in the total on an annual and seasonal basis and the corresponding mean concentrations of PM_{2.5} and various aerosol species associated within each trajectory cluster. Annually, the trajectory clusters are dominated by both NW and S, accounting for 44 and 34 %, respectively. The E and N

Table 3. Mean concentrations (in unit $\mu\text{g m}^{-3}$) of PM_{2.5} and selected aerosol speciations in the identified trajectory clusters within the one-year (annual) and four-season period. Also given are the percentages of each trajectory cluster classified in the one-year and four-season periods. For details on trajectory clustering, please refer to the text.

Air-mass type	Annual					Spring				Summer				Autumn			Winter	
	NW _s	NW _f	N	E	S	NW _s	N	E	S	NW _s	N	E	S	NW _f	E	S	NW _f	E
Percent (%)	11.9	32.4	7.4	14.6	33.7	8.7	29.6	17.4	44.3	9.6	8.7	8.7	73.0	49.6	26.8	23.6	88.2	11.8
PM _{2.5}	148	111	87	110	172	145	70	108	167	108	110	63	155	113	144	173	131	209
Sulfate	10.2	5	6.3	10.9	25.4	11.9	4.8	13.6	23.1	11.1	9.6	7.6	28.3	4.8	10.9	11.2	7	19.5
Nitrate	11.2	5.1	4.3	10.2	19.2	14	4.8	14.4	24.4	4.5	4	3.5	14.3	6.4	13.4	16.9	6.3	14.9
Ammonium	5.4	2.3	2.3	5.7	13.2	4.7	1.7	8	12.2	4.8	3.7	3.4	13.4	2.4	6.5	7.6	3.7	10.9
OC	22.2	18.1	10.8	13.8	16.4	15.5	10	11.5	16.7	12.4	12.2	9.4	11	15.6	18.7	21.6	23.8	33.6
EC	6.5	4.9	2.5	3.7	5.4	3.7	1.9	2.2	3.4	3.5	3.6	2.6	4.5	4	5.9	7.5	7	11.4
Mineral	31.3	36.5	26	16.4	20.8	64.8	32.8	15.3	39.1	15	21	4.8	8.5	33.3	19.4	30.4	31.3	21.8

clusters represent the rest (15 and 7 %, respectively). However, the variability is large and season dependent. For instance, the predominant clusters are N (30 %) and S (44 %) in spring, S (73 %) in summer, NW_f (50 %) in autumn, and NW_f (88 %) in winter. The resulting mean concentrations of main aerosol species seasonally vary with certain types of air masses. In winter, a few PM_{2.5} pollution cases (only 12 % of the wintertime trajectories) with mean concentration as high as $209 \mu\text{g m}^{-3}$ are associated within the E trajectories that passed over Hebei and Liaoning Provinces, where heavy industries are concentrated in certain cities (e.g., Tianjin, Tangshan, Dalian, Shenyang). However, in spring, summer, and autumn, high PM_{2.5} ($> 150 \mu\text{g m}^{-3}$) is preferentially associated with the S trajectory cluster. Overall, the general patterns agree with the spatial distribution of the MODIS-retrieved fine AOD around Beijing (Fig. 1).

Furthermore, we applied an alternative approach called PSCF to explore the likely regional sources and transport pathways of various PM_{2.5}-associated speciations, such as sulfate, nitrate, ammonium, OC, EC, and mineral dust in Beijing, as illustrated in Fig. 11. A few main features were found: (a) sulfate, nitrate, and ammonium have similar spatial patterns, with higher values in the east to the south, covering Tianjin, Shijiazhuang, and Zhengzhou; (b) both OC and EC show similar spatial distribution, with higher values in the northwest, the south, and the northeast, covering the border of Hebei and Shanxi Provinces, Inner Mongolia, the border of the Hebei, Shanxi, and Henan Provinces, and the area from Tianjin to Shenyang; (c) the higher value for mineral aerosols is localized in the northwest and the south; and (d) for these six aerosol speciations, the southern area appears to be a common hot spot. The overall PSCF results are rather consistent with the spatial distributions of fine AOD (Fig. 1) and their respective corresponding species' emissions such as SO₂, NO_x, NH₃, OC, and EC in China (Zhang et al., 2009b; Fu et al., 2012; Huang et al., 2012). In regard of dust, our PSCF result has shown that the dust transported to Beijing is primarily originated from northern/northwestern China (Wang et al., 2004) while Talimakan Desert is a very important dust source in China (Laurent et al., 2005). The

statistics obtained from the trajectory clustering (Table 3) shows that the southern air masses bring high levels of secondary inorganic and carbonaceous aerosols and the northwestern air masses are enriched in mineral dust and carbonaceous aerosols. Sun et al. (2006) and Streets et al. (2007) found that the S sector has much higher secondary species, such as sulfate, nitrate, and ammonium. During haze-fog events in Beijing, chemical constituents of secondary inorganic aerosols are also much higher when the winds blow from the south. Such high amounts of secondary fine-mode aerosols in southern air parcels may be related to high humidity (water vapor) and enhanced heterogeneous reaction in clouds/fog, aside from strong photochemistry. The association of high dust with the NW trajectories is consistent with Wang et al. (2004) and Yu et al. (2011).

4.3 Implications for atmospheric chemistry, PM control measures, and climate

Rigorous efforts exerted for air pollutant governance prior to the 2008 Beijing Summer Olympics, such as changing the energy source structure, reducing local dust emissions, controlling vehicle exhaust emissions, and relocating major industrial emitters, have achieved air quality improvement during the games. Effective control of coarse PM pollution seemed possible and the main urban air pollutants became finer PM (PM_{2.5}). The annual mean concentration of PM_{2.5} in Beijing is nearly three times and over an order of magnitude higher than the annual exposure level ($35 \mu\text{g m}^{-3}$) and air quality guideline (only $10 \mu\text{g m}^{-3}$) recommended by the WHO, which indicates that tremendous efforts of multi-pollutant alleviation measures and air quality management policies are still needed for PM abatement (Zhang et al., 2012b). The PM pollution level in Beijing is governed by the emission sources involved with natural and anthropogenic origins and particulate and gaseous phases and by synoptic meteorological conditions and atmospheric circulation systems. With sulfate/SO₂ as an example, our results demonstrate that meteorological conditions could be a crucial factor for determining fine aerosol levels and formation other than their own and/or precursor emission strengths. Accordingly,

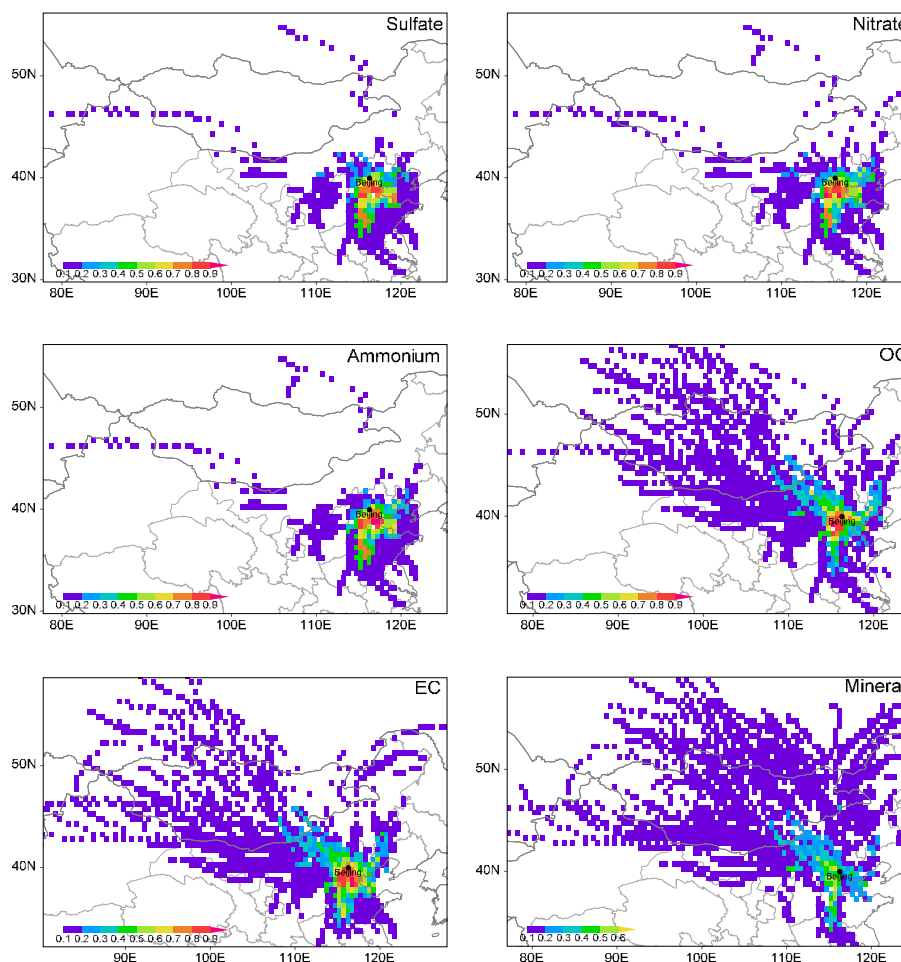


Fig. 11. The PSCF maps for sulfate, nitrate, ammonium, OC, EC, and mineral.

this study offers insights into the likely impact on atmospheric chemistry because of the changing climate. The increasingly warm climate predicted (IPCC, 2007) will enhance photochemistry in summer, which may offset the mitigation measures in China to some extent. Moreover, the decreasing wind speed forecasted in China (Chen et al., 2012) favors air quality degradation because of the likely reduction in ventilation efficiency (Zhang et al., 2007; Song et al., 2008). Complex PM pollution in terms of chemical and physical properties such as multiple sources (natural versus anthropogenic), mixing states (internal versus external), various chemical composition, size spectrum, and hygroscopicity, which are closely related to optical and direct/indirect radiative properties, complicate the modeling assessment of aerosol effects on the climate (cooling versus warming) in China and in the region.

If the contributions from the three main sources (coal combustion, industrial pollution, and SIA) are combined, fossil fuel burning-related emissions may dominate PM_{2.5} pollution in Beijing, representing two thirds (~68%) of PM_{2.5}. Rapid industrial development in provinces around Beijing,

including Liaoning, Hebei, Shandong, Shanxi, and Henan, has exacerbated regional air quality because of massive quantities of air pollutant emissions, resulting in cross-border transport. Better understanding of the pollution characteristics of PM_{2.5} in Beijing, particularly after the 2008 Summer Olympics, in terms of chemical composition and sources of PM_{2.5} from a regional and seasonal perspective, is urgently needed. Relevant air pollution control measures should be implemented locally, regionally, and nationally in China. Such measures would improve pollution abatement, public health, and climatic modeling capacity.

5 Summary

The levels of daily PM_{2.5} concentrations are still elevated in Beijing throughout the year, with an annual mean of up to $135 \pm 63 \mu\text{g m}^{-3}$, which is several times higher than the 24 h exposure standard and guideline recommended by the WHO. Seasonality is not very evident, although the highest occurred in winter and the lowest in spring. Distinctive seasonality

occurred for various aerosol species. Sulfate and ammonium peaked in summer, mainly because of photochemistry, whereas the maximum of nitrate was observed in spring, which is attributed to the heterogeneous reaction between dust carbonate and nitric acid (the product of NO_x) along with the photochemical process. Crustally derived species, such as Ca²⁺, Mg²⁺, and Al, showed typical seasonality, waxing in spring and waning in summer. The maximum of OC, BC, and Cl⁻ was observed in winter, and their minimum in summer or spring, which suggests a relation with fossil fuel (mainly coal) combustion processes. The correlations between total cation and anion equivalent concentrations were excellent with a mean ratio around unity, which indicates good charge balance. The mean NH₄⁺/SO₄²⁻ equivalent concentration ratios were higher than unity throughout the year, which reveals that NH₄⁺ could fully neutralize sulfuric acid. The correlations between NH₄⁺ and [SO₄²⁻ + NO₃⁻] suggest that ammonium could partially be present as NH₄NO₃. Furthermore, the correlations between [NH₄⁺ + Ca²⁺] and [SO₄²⁻ + NO₃⁻] demonstrate the presence of Ca(NO₃)₂ and NH₄NO₃, which was due to the heterogeneous reaction between dust carbonate and nitric acid. A few chloride salts may form, particularly in winter when low temperature is favorable for their occurrence.

CMCs were successfully constructed on a seasonal basis, although with an unidentified fraction averaged at 28.6%. The major aerosol speciation considered in this study included mineral dust, POM, and SIA (the combination of sulfate, nitrate, and ammonium), and the minor components were EC, chloride salts, trace element oxides, and biomass burning-derived potassium salts. Their respective fractions in PM_{2.5} largely varied with the season, generally corresponding to the seasonality of their concentrations. The higher unidentified fraction, particularly in summer, may be partly due to the volatilization of volatile or semi-volatile species, such as ammonium nitrate and organic species. This observation highlights the requirement for the immediate analysis of aerosol samples after collection. Through the PMF model, six potential sources were identified: mineral dust, coal combustion, biomass burning, traffic and waste incineration emission, industrial pollution, and SIA. Similarly, the relative contributions of these sources to PM_{2.5} in Beijing greatly varied with the changing seasons, which proves the “complex air pollution” in Beijing. On the average, the annual mean contributions were 15, 18, 12, 4, 25, and 26%. The highest contributions occurred in spring for soil dust (23%) and traffic and waste incineration emission (5%), in winter for coal combustion (57%), in autumn for industrial pollution (42%), in summer for SIA (54%), and in both spring and autumn for biomass burning (19 and 17%, respectively). Based on the trajectory clustering and PSCF method results, regional sources and transport pathways of various aerosol speciation associated with PM_{2.5} in Beijing were explored. The southern region may be an important source of all major

aerosol components. The eastern region is significant for anthropogenic species such as SIA and carbonaceous aerosols. The northwestern region could be the common source of naturally derived mineral dust and anthropogenically derived carbonaceous aerosols. This observation indicates that air pollution control measures must be simultaneously implemented in the surrounding provinces.

The results of our studies clearly suggest that chemical constituents and sources of PM_{2.5} can largely vary with seasons, which are characterized by variable meteorology and diverse air pollution sources. Source apportionment results do not precisely reflect the large temporal contribution variations from various sources when PMF analyses using data from the full four seasons are not considered. In addition, apart from the emission strengths of primary aerosols and gaseous precursors, the dynamically variable synoptic weather conditions and circulation patterns also have a crucial role in the anomalies of PM (both fine and coarse mode) concentrations.

Supplementary material related to this article is available online at: <http://www.atmos-chem-phys.net/13/7053/2013/acp-13-7053-2013-supplement.pdf>.

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