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Chernozem (czarnoziem) – soil of the year 2019 in Poland. Origin, classification and use of chernozems in Poland

Abstract: The Soil Science Society of Poland has elected chernozem to be the Soil of the Year 2019. Although chernozems cover less than 2% of Poland, they have high importance for agriculture due to their productivity and play a specific scientific role for understanding of soil development and functioning in an environment. Chernozems are also crucial for the reconstruction of Neolithic agriculture development and human impacts on soil and landscape. This introductory paper presents (a) a specific definition of chernozem in Poland, connected with a separate distinction of black earths and colluvial chernozemic soils; (b) a review of the present and former classification schemes for chernozems in Poland and their correlation with international soil classification systems (WRB and Soil Taxonomy); (c) the spatial distribution of chernozems in Poland, their agricultural evaluation and threats for chernozems' quality and future existence related to intense land use.

Keywords: Chernozems, Phaeozems, Mollisols, soil origin, soil classification, soil functions

INTRODUCTION

Chernozem has been considered in many countries a synonym for the most fertile and productive arable soil, suitable for cultivation of demanding crops, such as sugar beet, barley, wheat and corn. Thus, chernozems¹ have been the subject of significant public concern and interest of pedologists and agronomists. In this context, it has become clear that Chernozems/Mollisols are local icons – 'state soils' – such as in Nebraska (USA), Ukraine and the Belgorod region (Russia), and have been appointed a 'soil of the year' by national pedological societies, as in Germany and Estonia (Altermann et al. 2005; Kölli and Tamm 2015; Pozniak and Havrysh 2019). Although rendzina was elected (by the Commission of Soil Genesis, Classification and Cartography of the Soil Science Society of Poland (SSSP)) the first 'soil of the year' to inaugurate the programme in Poland in 2018, due to its specific historical importance for the Polish soil science (Kabala 2018), chernozem had already been appointed the second 'soil of the year' to celebrate it in 2019. This decision was also connected with the organisation of the SSSP congress in Lu-

blin, the capital of the region known for its fertile chernozems (Borowiec 1965, 1976).

CONCEPTS OF CHERNOZEM FORMATION

Numerous early attempts were undertaken to explain the origin of soils featured by thick, black, humus-rich topsoil horizon, including their marine or alluvial formation (Łabaz et al. 2018; Vysloužilová et al. 2016). The most widely accepted until now is a theory of Dokuchaev (1883), who defined chernozem as an autogenic natural body, developed under interconnected influences of various environmental factors, including calcareous but permeable parent rock, climate, vegetation, fauna and relief. Dokuchaev confirmed, what seems today absolutely obvious, that the organic matter of chernozems originates in terrestrial plants, those species' composition being closely related to climate. The organic matter, primarily the underground remains of roots, does not decompose rapidly because of seasonal dryness and stabilization by calcium in a carbonate-rich soil substratum, mostly the loess. Moreover, the organic matter and bulk soil is vertically and horizontally translocated by burrowing animals (both earthworms and rodents), which are particularly common in a virgin continental steppe. Zooturbation results in a thickening of the humus-rich topsoil layer and an improvement of the

¹The common names of soil units (e.g., chernozems, black soils/earths and rendzina) are written in lowercase unless they refer to Reference Soil Groups of WRB or soil Taxonomy, in which case they are written in capital letters.

soil mass aggregation (Ponomarenko and Ponomarenko 2019). Because the climate-related continental steppe vegetation is zonal, the chernozem is also considered zonal soil. Dokuchaev and his followers have distinguished the common (typical) chernozems of the central, tall-grass steppe belt, the southern chernozems of the short-grass dry steppe and also the 'degraded' chernozems of the moister forest-steppe zone located north of the central (continental) steppe zone (Strzemiński 1971). The steppe origin of the common (typical) and southern chernozems generally has not been questioned because the conditions favourable for chernozem formation still exist in many sites of the world, and active chernozem development can be presently observed (Alexandrovskiy 2007; Khitrov et al. 2019; Liu et al. 2012). In contrast, the nature of 'degraded' chernozems has been continuously discussed until now (Eckmeier et al. 2007; Vysloužilová et al. 2016).

The first concept, probably the most common, assumes that the chernozems in the northern and western peripheries and isolated "islands" outside the central steppe belt have developed under a cold or temperate steppe vegetation and then have been transformed (degraded) by forest vegetation that succeeded onto former steppe areas (e.g. in Central Europe) after climate warming and moistening (Czerny and Sachsse 1965; Miklaszewski 1930). Following this concept, chernozems in northern Central Europe and northern Russia are relics, developed in the late Pleistocene or early Holocene, before the Atlantic climate optimum (Borowiec 1962; Hajdukiewicz 2010; Kołodyńska-Gawrysiak et al. 2017). Alternatively, the formation of chernozems under an open-canopy forest-steppe or meadow-forest vegetation was considered without specification of the time of development (Borowiec 1967; Strzemiński 1961; Vysloužilová et al. 2014). Following this concept, chernozems developed under a less continental (moister) climate, and transitional or variable (in space and time) vegetation cannot be considered 'degraded'. Their chernozemic features are differently expressed, respectively to conditions of their formation (Borowiec 1968). As a consequence of temporary coverage with close-canopy forest vegetation, these chernozems may be featured by clay and humus eluviation/illuviation (Kowaliński et al. 1987; Licznar 1976; Miedema et al. 1999). These two general concepts, considered contradictory, in fact complement each other and are accompanied by a few others. The concept of the relatively young age of chernozems in Central Europe (i.e. their formation in or since the Neolithic period) (Maruszczak 1998; Vysloužilová et al. 2016) has become increasingly popular. However, some researchers assume that the large scale of forest

clearance by Neolithic people has allowed reconstruction of steppe or a domination of steppe-forest/meadow-forest vegetation that created conditions supporting the formation or reconstruction of chernozems in Central Europe (Chmielewski et al. 2014; Ehwald et al. 1999; Jacomet et al. 2016; Pokorný et al. 2015; Suchodoletz et al. 2017). In contrast, other researchers assume that the large-scale forest burning by Neolithic people was a single factor that enabled chernozem formation by the production of large amounts of fine particles of black carbon (charcoal), which has strongly coloured the topsoil after mixing (Eckmeier et al. 2007). Recently, a polygenetic formation of 'extrazonal' chernozems in Central Europe has been postulated, bringing together many of the previous separate concepts (i.e. it assumes initial chernozem formation in the late Pleistocene/early Holocene under cold-steppe vegetation and their 'rejuvenation' or second phase of development in/ since the Neolithic period under human-induced forest-steppe (-meadow) vegetation developed after large-scale forest clearance (Kabała et al. 2019b; Łabaz et al. 2018). However, it is highly probable there is no one scenario of chernozem development and transformation in all Central Europe, but rather various scenarios may be true in particular sites, which differ in parent materials, relief, moisture regime, vegetation history, length and intensity of human impacts etc.

An important issue influencing the discussion on chernozem origin is the moist water regime in many chernozemic soils connected with the development of abundant redoximorphic features, typical for soils developed under humid rather than continental climates. This issue led some, even modern scientists, to the conclusions on the hydrogenic origin of chernozems (Turski 1985), or, at least, to the questioning of their steppe origin (Eckmeier et al. 2007).

DEVELOPMENT OF THE DEFINITION AND CLASSIFICATION OF CHERNOZEMS IN POLAND

The East-Slavic (Ukrainian – Russian) word 'czarnoziem' has been occasionally used in legal Polish documents referring to the evaluation of arable lands at least since the mid-19th century (Strzemiński 1980). The modern understanding of chernozems was introduced to the Polish professional literature by Miklaszewski (1906) who has good knowledge about the Dokuchaev concepts. 'Właściwy czarnoziem' (proper or common chernozem) has been defined here and in the further works of this author (Miklaszewski 1930) as a mineral soil developed from loess (at least the

humus horizon), having thick humus horizon (40–140 cm, mean 70 cm) and carbonates in the profile, sometimes even in the topsoil layer. Chernozem contains 'sweet' humus derived from a non-bog vegetation. All 'common chernozems' are originally steppe soils, featured by abundant channels of burrowing animals, presently inactive, filled with humic soil mass. Following the Russian authors, Miklaszewski (1930) considered the Polish chernozems to be 'degraded' because they contain no more than 2 to 3% of humus (on average) in topsoil layers that is 2 to 4 times less than the Russian chernozems have (Khitrov et al. 2013). The agricultural suitability of chernozems depends on their parent material. Chernozems developed from deep loess have the best physicochemical, water and air properties, and therefore the highest productivity, even if 'degraded' by lower humus content (Miklaszewski 1930). Chernozems developed from loess in topsoil and underlain by cracked limestone or sand in the subsoil have potentially good fertility, but suffer from seasonal dryness, whereas chernozems developed from loess in topsoil and underlain by heavy loam or clay, impermeable to water, are too moist, an issue which may eliminate some crops (Miklaszewski 1930). Contrary to the 'common chernozems', Miklaszewski distinguished the 'boggy chernozems' referred to by him as 'black earths' (czarne ziemie), developed from various glacial materials. The origin of these soils is closely related to former peatlands or shallow lakes because they were formed by mixing of post-bog organic matter (often a peat) with mineral subsoil after permanent and deep soil drainage. However, Miklaszewski noted the presence of transitional soils (i.e. developed from loess and having channels of burrowing animals (similar to chernozems), which are wet and contain 'acid' humus (similar to black earths). Miklaszewski (1930) also distinguished the 'podzolized chernozems' (i.e. soil with thick humus horizon, formed from loess, but leached from carbonates and having clearly higher clay content in the subsoil ('iluwjum') compared to topsoil).

Somewhat different classification of chernozems in Poland (in borders before the WWII) presented Mieczysławski (1938), who distinguished 'degraded chernozems' (czarnoziemy zdegradowane), considered to be typical for the eastern Polish territories, 'northern chernozems' (czarnoziemy północne), 'poorly developed chernozems' (czarnoziemy niedokształcone), 'colluvial chernozems' (czarnoziemy deluwialne) and the moist 'Kuyavian chernozems' (czarnoziemy kujawskie). The latter may be correlated with the 'black earths' of Miklaszewski (1930); however, Mieczysławski (1938) did not support the opinion on the post-boggy origin of these soils.

This simple distinction of chernozemic soils in Poland into the 'dry' common/degraded/podzolized chernozems and the 'moist' post-bog black earths has been maintained since the end of World War II (Przyrodniczo-genetyczna klasyfikacja gleb Polski 1956; Genetyczna klasyfikacja gleb Polski 1959) despite many pedologists reporting the presence of moist (secondary moistened) variants of chernozems and questioning the only one accepted process (a post-bog genesis) for all black earths' origination (Kowaliński 1952). Moreover, these genetically oriented classifications have omitted many specific variants of chernozemic soils, such as eroded chernozems (with thinned humus horizon) and colluvial chernozems (i.e. accumulated at the foot-slopes and local depressions) (Harasimiuk and Wicik 2010). As a result, the national system of arable soil evaluation, launched between 1956 and 1963, has applied more morphologically oriented criteria for naming and evaluating the chernozemic soils in Poland (Ministerstwo Rolnictwa 1963). The next edition of the Polish soil classification (Systematyka gleb Polski 1974) has significantly extended the divisions of chernozemic soils based on genetical criteria. Chernozems have been separated into steppe-forest and meadow-forest types to stress the presumed different conditions of their origin – more continental (drier) and more Atlantic (moister), in eastern and western Poland, respectively. Each of these two types has been further subdivided into two subtypes: common (non-degraded) and degraded, based on humus content in the topsoil layer and depth of carbonate accumulation in the soil profile. The loess-derived, moist, meadow-forest chernozems have been separated from the post-bog black earths developed from various parent materials. This classification has solved some old problems related to the genesis of chernozemic soils; however, it has also generated a number of new theoretical and practical issues. In particular, it was not clear how to classify the mosaic of chernozemic soils featured by substantial differentiation of their properties (e.g. of the moisture status) on a small area, where three different paths of chernozemic soil development seem impossible (steppe-forest in the highest, meadow-forest in intermediate and post-bog in the lowest sites). Moreover, this new classification scheme did not explain regional differences in a classification of similar soils because the criteria to distinguish the meadow-forest chernozems and some black earths were still overlapping (Łabaz and Kabala 2014). As a consequence, the next, 4th edition of the Polish Soil Classification (Systematyka Gleb Polski 1989) resigned from separating forest-steppe and meadow-forest chernozems and returned to one type of

chernozems with two subtypes (non-degraded and degraded), accompanied by black earths type (in the other soil order). Also, the concept of black earths has substantially changed: from a narrow post-bog origin to a much wider understanding of the semihydrogenic conditions of their formation that includes soils developed both in drained bogs, bottoms of drained former lakes and in moist meadows or meadow-forest environments. Also, it should be stressed that the 4th edition of the Polish Soil Classification (*Systematyka Gleb Polski* 1989) introduced the diagnostic horizons as a tool for soil type recognition, including the mollic horizon for chernozems and black earths. The 5th edition of the Polish Soil Classification (*Systematyka Gleb Polski* 2011) retained the general rules for chernozems identification based on mollic horizons and secondary carbonates but abandoned the questioned division into non-degraded and degraded ones. Instead, the following five subtypes were distinguished: typical chernozems (having a mollic horizon 30–60 cm thick and secondary carbonates present in the profile), cumulative chernozems (having a mollic horizon >60 cm thick), chernozems with cambic horizon, chernozems with argic horizon and chernozems with stagnic properties.

CURRENT DEFINITION AND CLASSIFICATION OF CHERNOZEMS IN POLAND

The 6th edition of the Polish Soil Classification (*Systematyka Gleb Polski* 2019) has defined the chernozems as soils having a thick, nearly black and structural humus horizon (mollik) and secondary carbonates present within the soil profile (i.e. no deeper than 150 cm from the soil surface). Additional, indirect requirements for Chernozems result from the position of this soil group in the key to soil types within the chernozemic soil order (Kabała et al. 2019a). Chernozems cannot be developed from alluvial and colluvial (humus-rich) materials and thus typically occur beyond river valleys and other depressions or footslopes, occupied by chernozemic alluvial or colluvial soils, respectively (Strzemiński 1961). Hence, the soils also developed from colluvial material from eroded Chernozems and located in foothills may not be classified as Chernozems if they have clearly recognisable diagnostic features of colluvial material. Chernozems may have poorly developed redoximorphic features but do not have the stagnic or gleyic properties in the upper and medium part of soil profile (<80 cm from the surface) that are diagnostic for the black earths. Moreover, Chernozems cannot be dominated by primary (lithogenic) carbo-

nates from limestones and other carbonatic rocks that are crucial for chernozemic rendzinas (Kabała 2018; Miechówka and Drewnik 2018; Smreczak et al. 2018), situated before Chernozems in the key to soil types.

Taking into account the natural conditions of Poland, Chernozems are therefore the soils developed from (initially) carbonate-bearing loess, under the climate and vegetation that allowed accumulation of organic residues and extensive activity of burrowing animals, a process which resulted in the formation of thick and base saturated mollic horizon. No strict relation to a particular kind of vegetation was stated in the definition (key) because both the 'pure' steppe formations and the open-canopy forest-steppe or meadow-forest formations may support humus accumulation and intense zooturbation, at least by earthworms (Ehwald et al. 1999; Vysloužilová et al. 2014). The long-term coverage with close-canopy forests is rather excluded due to the fact that it would not only change the kind of biomass input and eliminate a significant part of animal activity, but would also support the leaching of carbonates out of the soil profile, common in Luvisols and Retisols neighbouring to the preserved chernozems (Kabała et al. 2019b; Ostaszewska et al. 2011). Chernozems may occur in flat landscape positions but more commonly are present on slightly inclined slopes, featured by good water drainage, which eliminates the high ground water table or surface water stagnation (and stronger expression of stagnic or gleyic properties). Such a moderately dry water regime does not limit the activity of earthworms and other burrowing animals. Limited intensity of erosion on gentle slopes does not disturb the topsoil layer and thus allows the recognition of chernozem, those definitions typically include the minimum thickness requirement for humus horizon (Dobrzański and Zbysław 1956; Licznar and Drozd 1985). In contrast, the soils on steep slopes, even if originally developed as chernozems under native vegetation and a beneficial climate, may be degraded to rendzinas (Strzemiński et al. 1973; Borowiec 1967) or pararendzinas – Calcisols (Zádorová et al. 2013), when excessively eroded (Borowiec 1966).

The majority of chernozems identified in Poland are developed from thick loess. However, the definition of chernozems is not limited to loess, because some chernozems are reported to occur on a glacial till, which still contains some secondary carbonates (Borowiec 1962; Cieśla and Dąbkowska-Naskręt 1983; Cieśla et al. 1988), and other chernozems may have glaciofluvial or outwash sands, glacial tills or limestones underlying the topsoil developed from loess (Borowiec 1962; 1965; 1976; Olszewski et al. 1965).

The latter materials may significantly influence the development of soil profile and soil properties, due to either enhanced water drainage or water retention. If the underlying material influences the soil properties too strongly, the soil may no longer fulfil the criteria for chernozem; this is particularly the case for underlying sands that support deep carbonate leaching (and soil shifting to decalcified grey soils) and impermeable loams/clays that contribute to the development of redoximorphic features and soil shifting to excessively moist black earths (Kabała et al. 2015; Dudek et al. 2019). Further, too shallow presence of limestone may mask the accumulations of secondary carbonates (if present at all) and leads to soil classification as chernozemic rendzina, which, in fact, may have a similar origin to typical chernozems (Miklaszewski 1930); however, having primary instead of secondary carbonates. The presence of carbonates in loess not only stabilizes the humus, thus supporting the formation of thick humus horizon, but also limits the weathering of minerals and transformation of chernozem (Drewnik et al. 2014).

Chernozems in Poland, even if steppe vegetation was confirmed to support their development in the Neolithic period (Chmielewski et al. 2014; Mueller-Bieniek and Nalepka 2010), differ of course from the 'typical' or even of the 'normal' chernozems of the central continental steppe zone in a humus content, which in Poland commonly varies in a range of 2 to 3% (Borowiec 1962; 1965; 1968; Gašior and Partyka 1999; Łabaz 2010; Olszewski et al. 1965; Turski 1985; 1986). The structure of soil mass in topsoil layers of the Polish chernozems rarely is granular, more often it is blocky subangular or even angular due to ploughing (Chudecki and Błaszczuk 1980; Domżał et al. 1991; Łabaz et al. 2018). Such a structure visibly differs from the granular or granular/coprolite structure of native, non-arable chernozems in the steppe zone (Khitrov and Loiko 2010). However, the arable 'normal' chernozems in the Woronezh and Belgorod regions also have a blocky angular or even cloddy structure in the topsoil layers after ploughing (Kabała 2016, personal observation; Khitrov et al. 2013), thus the structure cannot be a measure of differences between the zonal and extrazonal chernozems.

The recent, 6th edition of the Polish Soil Classification (Systematyka gleb Polski 2019) has distinguished five subtypes of chernozems: (a) leached chernozems (PL: czarnoziemy wylugowane) that do not contain carbonates down to a depth of 100 cm, (b) illuvial chernozems (PL: czarnoziemy iluwialne) that have clay translocation and argik horizons starting no deeper than 100 cm from the soil surface,

(c) brown chernozems (PL: czarnoziemy zbrutniałe) that have a kambik horizon, (d) stagnogleyic chernozems (PL: czarnoziemy opadowo-glejowe) that have stagnic properties in a layer >25 cm thick, starting <80 cm from the soil surface and (e) typical chernozems (PL: czarnoziemy typowe) that have all required features of the type (coming from its position in a key to soil types) and do not have features required for the subtypes listed above. Thus, the 'typical chernozem' has a mollik horizon and accumulation of secondary carbonates starting <100 cm from the surface, does not have kambik or argik horizons and does not have clearly developed stagnic (or gleyic) properties. Comparing to the previous Polish Soil Classifications (Systematyka gleb Polski 1974; 1989), the subtype 'typical' may be correlated with 'non-degraded', the 'stagnic' subtype with 'stagnogleyed', whereas the 'leached', 'brown' and 'illuvial' subtypes were merged previously into one subtype of 'degraded' chernozems. It should be noted that not all Polish chernozems, even the 'typical' ones, fulfil the criteria of Chernozems in the WRB classification (IUSS Working Group WRB 2015), which requires the presence of secondary carbonates at a depth no larger than 50 cm below the mollic horizon. In particular, the subtype of leached chernozems may refer to the Phaeozem group (Kabała et al. 2019a). On the other hand, it must be stressed that many Polish 'black earths' fulfil the criteria of (Gleyic/Stagnic) Chernozems in the WRB (Łabaz et al. 2018). According to Soil Taxonomy (Soil Survey Staff 2014), all Polish chernozems meet the criteria of Mollisols and typically belong to Calcudolls or Hapludolls (Kabała et al. 2019a).

As stated above, the distinction between chernozems and black earths in SGP6 refers indirectly to soil genesis and directly to soil morphology and physicochemical properties. Both chernozems and black earths are featured by the presence of thick mollik horizon. Chernozems must have a layer with secondary carbonate accumulation within the profile, while black earths – may have it. Whereas, chernozems do not have strong stagnic or gleyic properties, which, in turn, are required in black earths. Thus, the present definition of black earths includes soils of various origin, including polygenetic soils, featured by at least seasonal, excessive moisture of soil profile (either from high level of ground water or stagnating rain/melting waters), evidenced by stagnic or gleyic properties. This means, also the soils originated as steppe chernozems may be classified as black earths, if presently are situated in depressions, footslopes and other positions supporting water stagnation and development of reducing conditions in soil profile.

This distinction refers to different ecological functions of "dry" chernozems and "moist" black earths" and also different agricultural and forestry management required for these soil types (Strzemiński et al. 1973; Turski 1985; Błońska and Lasota 2019).

Distribution, land use and degradation of chernozems in Poland

Soils traditionally identified in Poland as chernozems occur in isolated larger locations on the loess plateau of southern Poland, near the towns Tomaszów Lubelski-Hrubieszów, Sandomierz, Przemyśl-Jarosław, Miechów-Proszowice, and Głubczyce (Borowiec 1965, 1968; Olszewski et al. 1965; Licznar 1976; Turski 1985; Czarnecki and Lewartowska-Urbańska 1987; Gąsior and Partyka 1999). Moreover, chernozems have been documented in mosaics with black earths, for example near Wrocław, Kłodzko, Inowrocław and Pyrzyce (Borowiec 1962; Cieśla et al. 1988; Drozd et al. 2007; Łabaz et al. 2018).

Due to properties beneficial to plant growth, nearly all chernozems in Poland are arable. Typical chernozems, developed from calcareous loess and having a thick humus horizon, have the highest agricultural evaluation belonging to class I or II of the national system. Leached, illuvial and brown chernozems have a beneficial texture; however, typically these chernozems are leached of carbonates to a larger depth and often have a thinner humus horizon. Thus, they typically belong to the II or IIIa evaluation class. Potentially, some limitations to plant growth may occur in stagnogleyic chernozems, and therefore they typically belong to IIIa and IIIb evaluation classes. Chernozems of various subtypes (excluding the stagnogleyic ones) may have a lower evaluation class (IIIa or IIIb, respectively to their properties) if they are located on slopes suffering from erosion. Nevertheless, nearly all chernozems belong to the 1st and 2nd agricultural land suitability category (complex) (i.e. lands dedicated to growing the most demanding crops, such as wheat, sugar beet, corn and rape) (Strzemiński et al. 1973).

Quality of chernozems and their degradation have been the subject of numerous studies. In particular, research regarding the composition and functions of humus are frequently found in the literature. The characterization of the amounts and ratios of particular humic compounds demonstrated the similarity of Polish chernozems to Ukrainian and Russian ones but also the similarity (sometimes interpreted as similar soil origin) of chernozems and other soil types, such as black earths or smonitza (Wilk and Nowak 1977; Turski 1986; Drozd et al. 2007; Łabaz

2010). However, the composition of humus may significantly differentiate in chernozems existing at various locations, under various vegetation and cultivation practices (Kuźnicki and Skłodowski 1968; Turski and Chmielewska 1986). The quality of humus has been found to negatively change under long-term cultivation of cereals in monoculture (Turski and Flis-Bujak 1980). Wilk and Nowak (1977) reported that organic fertilization of chernozems and other soils may not enhance soil fertility and crop yields directly but is crucial for stabilizing the content and composition of humus, and thus, influences the soil productivity indirectly. Also, erosion may substantially contribute to the quality of humus in chernozems, both by selective removal/accumulation of humus complexed with fine earth particles and transformation of humic compounds after burial with colluvial material (Licznar and Drozd 1985; Licznar et al. 1993).

Chernozems in Poland generally have a high content of nutrients (Borowiec 1965; 1971; Olszewski et al. 1965; Licznar 1976; Bieganski et al. 2013; Łabaz et al. 2018) and low concentrations of chemical contaminants, including heavy metals (Skłodowski and Sapek 1977; Czarnowska 1989; Gąsior and Partyka 1999; Jeleńska et al. 2008; Domańska 2009). However, the content of cadmium and lead in topsoil layers is 3 to 5 times higher compared to buried (Neolithic) chernozems (Łabaz et al. 2019). This observation may indicate a general (large-scale) contamination of the environment with these metals since the beginning of the industrial epoch (lead) or modern agriculture (cadmium). Local contamination of chernozems due to industrial emissions, for example with salts, is rarely reported (Brogowski et al. 1975).

The common source of threats for chernozems' quality and even their existence is generated by intense agricultural land use (Witkowska-Walczak et al. 1999). Domżał and al. (1991) and Słowińska-Jurkiewicz et al. (2013) reported that the transformation of soil structure in the topsoil layer is affected by tractor wheels or long-term fertilization. Intense cultivation, including the regular ploughing and mineral fertilization, leads to a decrease of humus content in topsoil layer and transformation to 'grey soils' (Borkowski 1964; Licznar 1976; Kowaliński et al. 1987). The most serious problem for chernozems related to agriculture is the sheet erosion that may accelerate chemical denudation and nutrient loss (Jadczyński et al. 2014) and subsequent thinning of the humus layer by physical losses of surface, humus-rich material (Licznar and Drozd 1985; Licznar et al. 1993; Górka-Kostrubiec et al. 2016; Gąsior et al. 2017). Erosion of chernozems in Central Europe began with early agricultural practices, which destroyed the natural vegetation

cover and opened soil surface to water action but has critically accelerated with the introduction of ploughing (Kołodzyńska-Gawrysiak et al. 2017; Zadorova et al. 2013). Long-term erosion may lead to the disappearance of chernozem by its transformation to pararendzina/Calcisol (Dobrzański and Zbysław 1956). Therefore, conservation practices that will protect chernozems both against erosion and humus loss are urgently required (Borowiec 2002) because the true area of chernozems has clearly decreased compared to their surface, evidenced in soil-agricultural maps prepared in the 1960s (Kabała and Łabaz 2019, data unpublished).

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Czarnoziem – gleba roku 2019 w Polsce. Geneza, klasyfikacja i użytkowanie czarnoziemów w Polsce

Streszczenie: Decyzją Polskiego Towarzystwa Gleboznawczego, czarnoziem został ogłoszony glebą roku 2019. Choć czarnoziemy pokrywają mniej niż 2% powierzchni Polski, mają duże znaczenie dla rolnictwa dzięki ich wysokiej produktywności. Odgrywają też specjalną rolę w badaniach nad zrozumieniem genezy gleb oraz funkcjonowania gleb w środowisku przyrodniczym. Czarnoziemy są szczególnie użyteczne w rekonstrukcjach paleośrodowiskowych, na przykład w odtwarzaniu historii rolnictwa neolitycznego i rosnącego wpływu człowieka na gleby i krajobraz. W niniejszym artykule wstępnym (a) zaprezentowano specyficzną współczesną definicję czarnoziemów w Polsce, uwzględniającą osobną klasyfikację czarnych ziem i gleb deluwialnych czarnoziemnych, które w innych klasyfikacjach mogą również być zaliczane do czarnoziemów, (b) dokonano przeglądu historii klasyfikacji czarnoziemów w Polsce oraz ich współczesnej pozycji taksonomicznej na tle międzynarodowych systemów (WRB i Soil Taxonomy), (c) omówiono występowanie i wartość bonitacyjną czarnoziemów w Polsce, a także zagrożenia dla jakości i istnienia czarnoziemów wynikające z intensywnego użytkowania rolniczego.

Słowa kluczowe: czarnoziemy, czarne ziemie, gleby szare, geneza gleb, systematyka gleb, funkcje gleb