

1 Clouds, solar irradiance and mean surface temperature over
2 the last century

3 ¹ A.D.Erlykin^(1,2), T.Sloan⁽³⁾, A.W.Wolfendale⁽²⁾

4 (1) *P. N. Lebedev Physical Institute, Moscow, Russia*

5 (2) *Dept. of Physics, University of Durham, Durham, UK*

6 (3) *Dept. of Physics, Lancaster University, Lancaster, UK*

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8 **1 Introduction**

9 The undoubted ‘global warming’ over the past half century or so has focused at-
10 tention on the role of changes in solar irradiance (and the solar wind) on a variety
11 of timescales and the relevance of cloud cover (CC). As is well known, the effect of
12 ‘solar forcing’ on the Earth’s climate is not fully understood (eg Foukal et al, 2004;
13 2006). In particular, the observed temperature changes are greater than would have
14 been expected, so this is one reason for yet another examination of the problem. An-
15 other is our analysis of the maps of Voiculescu et al (2006) in which we were unable
16 to find a good meteorological reason for the observed geographical pattern of the
17 regions having strong cloud cover (CC), solar irradiance (denoted SI) correlations.
18 As is well known, (eg Kristjansson and Kristiansen, 2000 and Erlykin et al 2009a),
19 the observed CC, Sunspot Number (SSN) - which can be taken in first order as a
20 proxy for SI - correlation is most unlikely to be due to cosmic ray variations, as
21 proposed by a number of authors, and SI variations are favoured. In what follows,

¹ Corresponding author at P.N.Lebedev Physical Institute, Moscow, Russia
E-mail address: erlykin@sci.lebedev.ru (A.D.Erlykin)

22 we refer to ‘SI’ but are mindful that the closely related solar wind may be the op-
23 erative agent, instead. The possible distinction is taken up later. Here, we examine
24 the variations over the 20th Century of both temperature and SI (via SSN) and
25 CC. Data on ‘temperature versus time’ are available for most of the 20th Century
26 for many points on the Earth’s surface (eg Hegerl et al, 2007) and sunspot records
27 are readily available over an even longer period. A complementary study is made,
28 albeit with less rigour, of the last thousand years and, in view of the contemporary
29 significance of the extended period of low (zero) sunspot numbers, the last 20 years
30 is examined in some detail.

31 The biggest problem relates to cloud cover. Clouds are inevitably connected with
32 temperature change and their relationship to surface temperature is cloud-height
33 dependent. Satellite data have been available since 1983, only, and even here there
34 are calibration uncertainties (eg Norris, 2000). Cloud data over a longer period,
35 post-1952, are, however, available from the synoptic reports summarised by Norris
36 (1998, 2004). These relate to ‘upper’ and ‘low’ clouds and cover specific latitude
37 ranges : 30° S - 30° N (ocean) and 30° N - 60° N (ocean), and these data are used
38 here.

39 The main thrust of the paper is to study the correlation of changes in solar irra-
40 diance with ‘climate’ (temperature and clouds) for the various time periods from
41 the standpoint of both the 11-year and 22-year cycles. We aim to check that the
42 temperature variations are in fact excessive and, in particular, to study the reason
43 for the 22-year cycle being so much stronger than that for the 11-year cycle (as
44 observed already by Miyahara et al, 2008).

45 We then go on to determine a best estimate of the contribution of ‘natural’ (SI)
46 effects to the well known increase of Global temperature since 1900. It is appreci-
47 ated that others have also examined this topic but an independent study is clearly
48 desirable.

49 2 Temperature versus Solar Irradiance

50 2.1 An Overview

51 Figure 1 shows a collection of estimates of Global average surface temperature
52 changes (ΔT) versus estimated changes of SI (ΔSI) from (mainly) previous work,
53 in order to ‘set the scene’. Details are given in Table 1. The data here are all from
54 the analyses of others, except for 7 and 3. Concerning ‘7’ this is a straightforward
55 plot of the variation of SI over the year for latitude bands and the corresponding
56 summer/winter temperature difference. Although a naive approach it has most of
57 the ingredients for a realistic expectation for the rate of change of ΔT with ΔSI . It
58 is extrapolated linearly downwards. Another prediction is the line indicated ‘model’.
59 This was derived using the relation $\frac{\Delta T}{T} = \frac{1}{4} \frac{\Delta SI}{SI}$, which fits the data for almost all
60 planets and is as expected from evaluation of the equilibrium temperature of an
61 insulated black surface normal to the Sun. It is appreciated that, understandably,
62 the planets do not satisfy the condition of an ‘insulated black surface’ but there is
63 just a systematic off-set of this equilibrium temperature and the actual (bright-side)
64 temperature. The off-set for most planets is a reduction of actual below equilibrium
65 temperature of about 30%. The exception is Mercury, having no atmosphere. In
66 the case of the Earth the ratio is a little higher, the equilibrium temperature being
67 about 394°K and the actual 295°K (ie a reduction of about 25%).

68 In Figure 1, we use $T = 294^\circ\text{K}$, the equilibrium temperature, so the line is somewhat
69 of an upper limit. The line is approximately $\Delta T(^{\circ}\text{C}) = \frac{\Delta SI}{SI}$ (in %) ie at our datum
70 $\Delta T = 0.1^{\circ}\text{C}$, $\frac{\Delta SI}{SI} = 0.1\%$. The line is clearly just a datum in that feedback effects
71 (positive or negative) can cause differences. This topic is taken up again later.

72 The point marked 3 is from our recent paper (Erlykin et al, 2009a). In this paper
73 we pointed out that the change in temperature since 1956 about the smooth trend
74 follows closely the change in SI (both averaged over the 11 year solar cycles). The
75 change in the cosmic ray rate was also observed to follow the trend but is delayed
76 by $\sim 2 - 4$ years. We assume that the observed change in temperature is caused by

77 the change in SI to give point 3 on Figure 1.

78 Point 6 differs from the others in that it is from calculation of the effect of a change
79 of distance of the Earth from the sun; the others all relate to measured, or inferred,
80 temperature changes (ΔT) and associated changes in sunspot number ($\Delta SI/SI$).

81 Inspection of the Figure shows that with the exception of ‘7’, the Global seasonal
82 variations, the ΔT values are higher than would have been expected from ‘expec-
83 tation’, a result of importance in view of the need to know the magnitude of the
84 solar forcing at the $\Delta SI \simeq 0.1\%$ level as a help to understanding fully the cause of
85 temperature changes in general and Global Warming in particular.

86 One of the objectives of the present work is to attempt to clarify the situation of
87 these small irradiance changes and to confirm, or otherwise, the ‘excess’ values of
88 ΔT . Another is to endeavour to identify the actual cause of the temperature changes
89 : SI as such or another phenomenon connected with the solar wind and to go on
90 to determine the ‘natural’ contribution to the 0.7°C Global Warming since 1900.
91 Another objective is to examine the role of clouds, particularly from the standpoint
92 of the 22-y cycle.

93 *2.2 The last half-century - our own earlier work*

94 As remarked, we have already examined temperature, SI and CR data since 1956
95 (Erlykin et al (2009a)). It was concluded that the CR time variations did not fit the
96 temperature variations and this was yet another reason to disbelieve a significant
97 role for CR in generating clouds and thereby affecting temperature (other work
98 includes Sloan and Wolfendale, 2008). The conclusions in the post-1956 data related
99 largely to the information from the large dip in ΔT in the region of 1970, which
100 coincided with the well known low sunspot maximum (for Cycle 20) in that year in
101 comparison with the neighbouring Cycles (Figure 2). Next, we go further back to
102 the beginning of the Century and see whether there is confirmation of the conversion
103 from SI change to temperature change derived there is confirmed (point 3 in Figure

104 1).

105 *2.3 The temperature record for the last century*

106 *2.3.1 The mean over the Globe*

107 A disconcerting feature is the fact that the profile of surface temperature versus time
108 is not unique but varies from place to place, not only in absolute magnitude but
109 also in shape. Figure 3a shows the average surface temperature over land (LAN),
110 over ocean (OCE) and averaged for the whole earth (GLO). (The results are from
111 Hegerl et al, 2007). Inspection of the profiles of surface temperature vs year, from
112 1910 to 2000 for the 22 regions distributed over the Globe (each having an area of
113 $\sim 10^7$ km²) shows not only differences in the temperature rise from place to place,
114 but other differences, too. Most pronounced is the movement of the peak at 1940 in
115 Figure 3a which, although stable between ocean and land, is variable from one region
116 to another. Specifically, there are 12 regions with peaks between 1938 and 1942 and
117 8 regions with peaks between 1948 and 1952. The cause of the dichotomy is probably
118 the phenomenon encountered in the Altai region where the identification by Eichler
119 et al (2009), of a 10–30 year lag between solar forcing and temperature response, led
120 the authors to postulate an indirect sun-climate mechanism involving ocean-induced
121 changes in atmospheric circulation. Having said that, it must be remarked that the
122 3 regions nearest to Altai (‘NAS’, ‘CAS’ and ‘TIB’) all had peaks in 1940, rather
123 than being delayed by 10 years to 1950. In any event, the stability of the profiles in
124 Figure 3a for land (LAN) and ocean (OCE) is plain to see.

125 The role of the oceans in comparison with that of the land can be seen by means of
126 the overall increase in temperature over land (LAN) being significantly bigger than
127 that over the oceans (OCE). Having said that, it is not clear why the ‘structure’ in
128 OCE is greater than that in LAN; although the peak at 1940 and the dips at 1950
129 - 1970 are close in time, they are sharper in OCE. Presumably the answer lies in
130 the fact that there is a greater homogeneity of ‘ocean’ than ‘land’, the latter having
131 a wide variety of terrains : industrial areas, farm lands, lakes, deserts, etc, all with

132 different albedos and other properties.

133 In this connection it is necessary to consider further the possible time lags that can
134 occur between SI changes and subsequent changes in ΔT . For a start, differences
135 might be expected between ocean and land, in view of the different specific heats
136 (5:1) and thermal conductivities. However, inspection of Figure 3a shows that the
137 differences between LAN and OCE are only in amplitude and not in temporal po-
138 sition as mentioned already. Thus, the strong 1940s peaks are in the same year, as
139 are the minima at ~ 1950 and 1969. That the near-surface air temperature has a
140 very short time lag is evident from a night/day comparison; it is invariably ‘cold at
141 night’. Our earlier work (Erlykin et al, 2009a) for the 11-year averages showed a lag
142 less than a year (indeed the best estimate would appear to be about minus one year,
143 with respect to the SSN!).

144 In what follows for ‘mean Global temperature’ we use the values from Hansen et al
145 (2006), with an 11-year smoothing to eliminate the first order solar-cycle variation,
146 as adopted in our earlier work (Erlykin et al, 2009a).

147 *2.3.2 The 1940 peak*

148 It has been remarked already that there is a consistent peak in temperature in 1940,
149 for the Global averages, for all situations : GLO, LAN and OCE.

150 Inspection of Figure 2, the sunspot number versus time for the last century, surpris-
151 ingly shows no expectation of a maximum for 1940, rather the envelope of the SS
152 number leads to an expectation of a maximum ΔT in about 1959 if there is a causal
153 connection between solar activity and temperature. Here, there is, indeed, a small
154 peak in OCE ($\simeq 0.04$ ° C) and GLO ($\simeq 0.03$ ° C) but this is dwarfed by the 1940
155 peak.

156 The surprising peak in 1940 has been commented on by a number of workers (eg
157 Hegerl et al, 2007). In the period 1930 - 1960, volcanic aerosols are thought to
158 have had a negligible effect on Global temperatures (Lean et al, 2005) but this

159 is not the case with the ENSO tropical temperature index, which has a peak-to-
160 peak temperature excursion of about 0.15°C for 1 y binning (and the well-known
161 2 - 3-y oscillation). There was, in fact, a particularly strong El Nino in the period
162 1940 – 42 but the geographical distribution of the ‘1940-peak’, which not only varies
163 by about 10 years, does not accord with expectation. The detailed correction of the
164 surface temperature for ENSO, volcanic aerosols, greenhouse gases and tropospheric
165 aerosols by Lean et al (2005), although including a ‘peak’ of about 0.05° , and a width
166 of half height of ~ 6 years, still leaves an excess ΔT of $\sim 0.2^\circ\text{C}$ over nearly 10 years
167 unaccounted for.

168 Another explanation put forward is that the 1940 ‘peak’ was due to a post 1940
169 ‘dip’ caused by the effects of bio-mass burning (Nagashima et al, 2006).

170 The detailed discussion of the ‘1940-peak’ is to draw attention to the hazards in-
171 volved in separating out a (small) solar forcing signal in the presence of other forc-
172 ings, of inevitably uncertain magnitude.

173 *2.3.3 Temperature, SSN correlation over the whole Century*

174 In an attempt to apply a consistent analysis to the data we inspect the temperature
175 record in Figure 3b, (the upper line) which is decadal averaged and corrected for the
176 long-term trend (as in Erlykin et al, 2009a) and endeavour to correlate the patterns
177 of ΔT and SSN. In view of the 11-year smoothing of the SSN there are minima at
178 the even Solar Cycle numbers indicated in the Figure. It is evident that with the
179 exception of the ‘1940-problem’ there is a generally good correlation between the
180 patterns and that ΔT and SSN are causally correlated. In order to derive a value
181 to add to Figure 1 we examine the magnitudes of the dips in both ΔT and SSN
182 assuming that the two are strongly correlated. Times are identified which are near
183 the peaks of the (smoothed) SSN - identified by small vertical arrows - and the SSN
184 dip determined for each of the 5 regions (by ‘dip’ we mean the mid point value with
185 respect to the mean of the two end points). The same end points were taken for ΔT
186 and the temperature dips found in a similar fashion. The prominent dip marked by
187 the 3rd arrow in Figure 2, ie Cycle 20, is that studied in our earlier work (Erlykin

188 et al, 2009a).

189 We have taken the data from Figure 3b and measured off 3 points in each of the
190 5 dips, equally spaced. In each case ΔT and ΔSSN are determined with respect to
191 the chord joining the peaks in ΔT corr. Figure 4 shows the values so derived. It will
192 be observed that there is a very approximate linear dependence of ΔT on ΔSSN .
193 The slopes can be converted to ΔT vs $\Delta SI/SI$ using the conversion from Sloan and
194 Wolfendale (2008) that $\Delta SSN = 150$ corresponds to the $\Delta SI/SI = -0.07\%$ from Lean
195 et al (2005). The result is that $\Delta T = (0.1 \pm 0.03)^\circ C$ corresponds to $\Delta SI/SI = (0.014$
196 $\pm 0.004)\%$. Eliminating points ‘30’ and ‘50’ in Figure 4, ie the data containing the
197 dips in 1930 and 1950, which might be justifiable because of problems with the 1940
198 peak, can be seen to have no effect on the overall slope. The point is plotted as ‘10’
199 in Figure 1.

200 It should be remarked that SI (taken from the work of Wang et al (2005)) follows
201 SSN rather closely from 1956 to 1992 but there is divergence thereafter; this is
202 another indicator that Cycle 23 is anomalous.

203 As a check, we have used an alternative set of Global temperature data, that over
204 land alone and for the two Hemispheres separately (Peixoto and Oort, 1992). The
205 averages to 1985 are: $\langle \Delta T \rangle = 0.15 \pm 0.05^\circ C$ for the N hemisphere, and $\langle \Delta T \rangle$
206 $= 0.05 \pm 0.02^\circ C$ for the S hemisphere.

207 The Global mean ($0.06 \pm 0.02^\circ C$) is consistent with our $0.05 \pm 0.02^\circ C$ just derived.

208 The fact that the mean value of ΔT over land is greater than that over the oceans
209 is understandable in view of the higher thermal inertia of the oceans (see 2.3.1 and
210 Figure 3a). The value of ΔT over land in the S Hemisphere is presumably lower
211 than that in the North for reasons of the greater proximity to water in the South.

212 We are mindful of the contributions of other effects to changes in surface temperature
213 : ENSO, volcanoes, ozone and greenhouse gases. Concerning the last mentioned, the
214 smoothed contribution is included in the slow systematic change in ΔT with time; it
215 is the shorter term variable part which is of concern here. Estimates of the necessary

216 corrections have been made by us using data from Lean et al (2005) and Mechl et
217 al (2004). The values were smoothed over 5 year intervals and have been applied to
218 the upper line in Figure 3b to give the middle curve, labelled ΔT_{corr} . The median
219 spread in corrections for the 5-yearly smoothing is $\pm 0.05^\circ \text{C}$. Our estimate for the
220 1-year smoothing is $\pm 0.09^\circ \text{C}$; this value will be needed later.

221 Repeating the analysis to give the mean values of ΔT and SI for the ‘corrected’ time
222 dependence in Figure 3b gives $\Delta T = 0.01^\circ \text{C} \pm 0.02^\circ \text{C}$, for $\frac{\Delta SI}{SI} = 0.013\% \pm 0.003\%$ ie
223 very similar to the earlier value. It is reassuring that the perturbing factors (ENSO,
224 volcanoes, etc) do not invalidate our analysis.

225 The conversion value for 1956-2001 derived by us in Erlykin et al (2009a), of $\frac{\Delta SI}{SI} =$
226 0.015% for $\Delta T = 0.1^\circ \text{C}$, is therefore confirmed within the uncertainties.

227 *2.3.4 Temperature, SSN correlations over the last 1,000 years*

228 Although there are no direct SSN measurements over the whole of the millennia,
229 proxy indicators of the SI have been used. For example, Crowley (2000) has used
230 cosmogenic isotopes, specifically ^{10}Be in ice cores, residual ^{14}C from tree ring records
231 and an estimate of ^{14}C from ^{10}Be fluctuations. The same author derived a tempera-
232 ture record for the Northern Hemisphere, using instrumental data after 1860 and a
233 proxy record prior to this date, the proxy being tree rings, corals and ice cores. It is
234 appreciated that there are many uncertainties for a time period of such length but
235 we would contend that an analysis of correlations in the proxy data has some value.

236 Crowley gives the resulting ‘observed’ temperature variation (we call it ΔT (ob-
237 served) and the expected, from the inferred ΔSI temperature relation derived above,
238 (we call it ΔT (predicted) versus time from the year 1000 to 2000. The data have
239 been used by us to study the correlation for each century: 1000 - 1100... to 1800-
240 1900 with the results for the correlation parameter, p , and the slope of the line for
241 ΔT (observed) vs ΔT (predicted) shown in Table 2. It is interesting to note that
242 the p -values are very small (ie the correlation is very significant) for the periods for
243 which volcanoes contributed significantly to the ΔT -value (1200-1300; 1400-1500

244 and 1800-1900), ie the corrections must have validity. Taking all the 1000 year data
245 together and plotting ΔT (obs) vs ΔT (pred), we find a straight line fit with slope
246 0.77 ± 0.05 , a correlation coefficient $r = 0.98$ and a correlation probability $p < 0.001$.

247 The overall situation regarding the probabilities is satisfactory and supports the
248 contention that there is a good correlation over 1000 years.

249 The near-proportionality of ΔT to the solar forcing (ΔSI) used in the calculations,
250 with the small volcanic forcing correction adds validity to the arguments put forward
251 in section 2.3.3 for the last 50 years. It remains to examine reason for the difference
252 in the 11-year and 22-year temperature - SI relations. That there are dependences
253 with both 11-year and 22 year components is evident from many workers for the
254 1-year averaged direct solar cycle (eg Lean et al, 2005).

255 Concerning the 22-year (approximately) variation over the century, this follows di-
256 rectly from the present work (Figure 3b) for the last 100 years. It is also present
257 for the same period in (Erlykin et al, 2009b). In that work we found for the Fourier
258 frequency spectrum, peaks at 0.0039 and 0.0072 month⁻¹, i.e. 21- and 11.6 - years.

259 Independent analyses have been made by others (eg Vecchio and Nanni, 1994, for
260 the last century) show peaks in the 'relative variance' at ~ 20 years (the Hale cycle).

261 Turning to the last millennium, inspection of the ΔT , time profile shows the presence
262 of some 50 peaks in this period, ie a mean separation of ~ 20 years. This is borne
263 out by the Fourier analysis which shows a peak at about 20 years.

264 At the level of examination here there is no evidence for significant phase lag between
265 the changes in SI and temperature, except, perhaps for the periods 1000 - 1100 and
266 1600 -1700, where the correlation probabilities in Table 2 are poor.

267 *2.4 Discussion of the temperature versus solar irradiance results*

268 *2.4.1 General Remarks*

269 Despite the fact that most of the ‘points’ in Figure 1 are ‘high’ - it does appear there
270 is evidence in their favour. Additional evidence comes from the fact that there is an
271 upward progression of ΔT with length of time over which the averaging is made, at
272 least for those observations for which we have made the analysis, viz points: 7, 3, 10
273 and the point from the exhaustive study by Lean et al, (2005): 2. Of the others, 6
274 relates to a first order calculation of the expected effect of changes in the sun-earth
275 distance and the others are approximate.

276 The ‘preferred’ points are thus 7 (1-year), 2 (11-year), 3 and 10 (22-year). There are
277 now two questions: is the progression ‘reasonable’? and, what is the reason for the
278 excess values of ΔT over the Model expectation? Such a behaviour is not completely
279 understandable. Thermal inertia per se appears to be ruled out. The datum line
280 (denoted ‘7’) relates to a yearly period and it seems that some 20-30 year is needed
281 for this inertia to be largely overcome. The inertia must be ‘resistive’ in character
282 (in part, at least); a ‘capacitive’ component would give a phase-lag, which seems
283 not to have been observed for the earth as a whole. It is not self-evident that such
284 a time is correct particularly because the ‘Model’ line should be an equilibrium
285 value and it is ‘low’;presumably some form of positive feedback is operative. The
286 comments of Shindell et al (1999), with regard to the ‘disproportionate effects of
287 UV’ changes on the upper wind patterns are relevant, as in the model of Haigh
288 (2007) and elsewhere, as will now be described. Many workers have found increased
289 values of ΔT at heights well above ground level. Specifically, Hood (1997), Hood
290 and Soukharev (2000) and Gray et al (2009) have derived ΔT values higher than
291 the 0.1°C ground level value, by a factor ~ 4 at a pressure of 10hPa rising to $\sim 20\times$
292 at a pressure of 0.1hPa. These high values arise from the large 11-year cycle in solar
293 UV and include positive feedback effects due to ozone. At 0.1 hPa (65 km) the ΔT
294 change is seen to be some 50 times the Model prediction of Figure 1.

295 Referring to Figure 3b, it appears that the dips at 1910, 1930, 1970 and 1991 (ie

296 the ‘22-year cycle’), are reflections of the fact that the peak SSN values alternate
297 from Cycle to Cycle. A useful factor is the ratio of the sunspot numbers (monthly
298 averages) for an even-numbered cycle to the mean of the adjacent ones. The ratios
299 are: Cycle 14 : 0.82, Cycle 16 : 0.78, Cycle 18: 0.97, Cycle 20: 0.79 and Cycle 22 :
300 1.03. This last-mentioned arises because the peak at 2002 (cycle 23) is anomalously
301 low; indeed cycle 23 is anomalous in many ways (see Section 5). This aspect will
302 now be examined.

303 2.4.2 ‘Solar irradiance’

304 Although in first order, the change in Solar Irradiance, ‘SI’, is proportional to sunspot
305 number, ‘SSN’, there are subtleties. These arise from the spectral shape of the solar
306 radiation. It is well known that, although the fraction of the energy content of SI
307 falls with increasing frequency, its 11-year cycle increases in amplitude. Thus, there
308 is the possibility that the pattern in Figure 3b is a consequence of UV as distinct
309 from visible radiation. Inspection of available data indicates typical radiance changes
310 over the 11-year cycle of 0.15 Wm^{-2} in the range 200 - 300 nm which, at the high
311 altitudes at which this radiation is absorbed (above 20 km) is very large, considering
312 the very low air density involved (the total, for all wavelengths, is only $\sim 1 \text{ Wm}^{-2}$).

313 If the odd-even Cycle differences found for SSN (and other indicators) are present in
314 the UV, too, then UV irradiance is a good candidate for the observations in Figure
315 3b. This aspect can be considered further.

316 UV data are only available from 1978 (Viereck and Puga, 2005; Deland and Cebula,
317 2008), but can be extrapolated back just two years to the SSN minimum in 1976,
318 yielding reasonable results for all 3 Cycles: 21, 22 and 23. Integration over the UV
319 intensity vs time for each Cycle gives $(\text{UV}(22)/\langle \text{UV}(21,23) \rangle = 0.78$, without doubt
320 less than the (anomalous) ratio for SSN (1.03).

321 If the feature is common, viz that the UV (Even Cycle) is always significantly less
322 than the mean of its neighbours, to a greater extent than is usually true for SSN,
323 then we would have a ready explanation of the ΔT value for the 22-year mean being

324 proportionately greater than for the 11-year cycle (Figure 1, ie the black summary
325 point from Figure 3b being ‘higher’, with respect to either line than the 11-year
326 points 1 and 2).

327 The role of UV in probably explaining the high value of ΔT with respect to the
328 Model prediction has already been referred to in 2.4.1. It has relevance, too, to the
329 22-year cf 11-year difference. In the work of Haigh (2007) the energy (excess) in
330 the 11-year Cycle is deposited at about 50 hPa, ie 20 km altitude where 300 nm
331 radiation is absorbed (this is one optical depth from the top of the atmosphere). The
332 air density here is $\sim 10^{-4}$ of that at ground level so that it is not surprising that
333 the predicted temperature effect at this altitude is so great. The model involves a
334 perturbation of the atmospheric air circulation system so that the lower troposphere
335 and the Global surface are affected, specifically by $\Delta T \sim 0.6^\circ\text{C}$ at 20 km and $\sim 0.2^\circ\text{C}$
336 at ground level.

337 Mention should also be made of the work of Mohakumar (1988) who examined
338 the middle atmosphere (65-70 km altitude) temperature associated with the 11-
339 year solar cycle. This worker argued that enhanced solar emission of Lyman Alpha
340 (121.6 nm) plays a major role in the physico-chemical processes involving minor
341 constituents. Lyman Alpha is absorbed mainly at the heights mentioned. The 10.7
342 cm radio emission, which is generally regarded as a proxy of the UV flux, and for
343 which there is data back to 1947 (UKSSDC, 2009), is also very relevant to this height
344 region. We find that, for Cycles 19, 20 and 21, the Odd/Even Cycle maximum of
345 $\langle 19, 21 \rangle$ to that of 20 in the Solar Radio Flux is 1.8 compared with 1.54 for the
346 peaks of the SSNs. As remarked already, the hard UV has a bigger peak to peak
347 variation than the sunspots so that if the ‘1.8’ factor itself increases as wavelength
348 increases then this will help with the 11-year, 22-year problem. It also helps with
349 the positive feedback suggestion for the ‘high’ ΔT -values in Figure 3b.

350 That this is probably not the whole story, however, comes from the linearity in the
351 plot of ΔT (obs) vs ΔT (pred) - predicted on the basis of SSN, referred to in 2.3.4
352 (and having a correlation probability $p = 0.000$). One would have needed a concavity
353 in the plot, ie low SSNs (even numbered cycles) to yield lower UV fluxes (per SSN)

354 than high for the 22-y, 11-y contrast to be due entirely to UV. It seems likely that
355 there is a big component from the inertia effect already referred to.

356 At this stage, the dependence of ‘efficiency of change in temperature for change in
357 SI’ as a function of relevant time interval can be extended to include annual and
358 daily variations. Summarising, we have the following $\Delta\text{SI}/\text{SI}$ percentages needed for
359 a change in temperature of 0.1°C :

360 1 day $\sim 0.6\%$, 1 year $\sim 0.6\%$, together with the values discussed already 11-year \sim
361 $(0.10 \pm 0.02)\%$ and 22-year $\sim (0.014 \pm 0.004)\%$.

362 The trend seems physically sound.

363 *2.5 The extent to which the solar irradiance change accounts for Global Warming*

364 By confirming the conversion factor for changes in SI to changes in temperature we
365 have confirmed the conclusions of our earlier work (Erlykin et al, 2009a) that less
366 than 14% of the temperature increase (of 0.5°C) between 1956 and 2001 is due to
367 the change in solar irradiance. Applying the same conversion factor for $\Delta\text{SI}/\text{SI}$ to
368 ΔT from the second half of the Century to the first half, we find that the observed
369 temperature increase (of 0.25°C) can be compared with our prediction of $0.3 \pm 0.1^\circ\text{C}$.
370 There is thus no evidence for any excess warming over and above ‘natural causes’
371 in this period.

372 **3 Cloud cover, solar irradiance correlations**

373 *3.1 Post-1984 results : the 11-year cycle*

374 There is a wealth of literature on the relationship between the mean cloud cover,
375 as deduced from satellite observations (ISCCP) and the SSN or the closely related
376 cosmic ray (CR) intensity. Much of it relates to ‘low’ clouds (LCC : heights less than

377 3.2 km), eg Marsh and Svensmark (2000), where, over the 11-y cycle, the peak-to-
378 peak range of LCC is ~ 2 %. For higher clouds, the medium clouds (MCC : 3.2 -
379 6.5 km) and high clouds (HCC : > 6.5 km) there are the maps, already referred to,
380 of Voiculescu et al (2006). An analysis has also be given by ourselves (Erlykin et
381 al, 2009c) in which the CC, SSN correlations have been examined as a function of
382 latitude. The two analyses are consistent in the sense that there are variations with
383 latitude and the sign of the correlation depends on height.

384 Specifically, the LCC, SSN (or UV) correlation is negative for LCC (this is why
385 the apparent LCC, CR correlation is positive - since CR and SSN are inversely
386 correlated). The MCC, SSN correlation is positive, and the HCC, SSN correlation
387 has equal (small) areas of positive and negative correlation (Voiculescu et al, 2006).
388 However there is a bigger area having a negative HCC, CR correlation and this may
389 have arisen from a mis-identification of SSN and CR, (actually UV and CR)), in
390 which case the HCC, SSN correlation is also positive like that for MCC, SSN.

391 *3.2 Longer period correlations : post-1952*

392 As remarked earlier, synoptic cloud data are available for the last half-century and
393 these are very useful for the present analysis. They comprise measurements from a
394 very large number of sites distributed over the oceans. Some 60 million observations
395 were involved in the period 1954 - 2000. Every effort was taken to ensure that the
396 same criteria were adopted by the observers. Unfortunately, the observations are
397 divided only into ‘high clouds’ and ‘low clouds’ and thus a comparison with the
398 HCC, MCC and LCC is difficult; however, it can and will be done.

399 Figure 5 shows the results for the CC magnitudes, as a function of time, for the
400 two latitude regions, both over the oceans. The SSN data are averaged over 11-year
401 cycles as in Erlykin et al (2009a), but the CC are not(an unimportant fact), but
402 rather to means over a 5 year bin. The random error on each point is about $\pm 0.15\%$.
403 There is seen to be a strong anti-correlation of CC with SSN for the latitude range
404 $30^\circ - 60^\circ$ N and a lesser one for 30° S - 30° N, both for low clouds. These results are

405 in the same sense as for our own work on LCC, where, as with Marsh and Svensmark
406 (2000), we found a positive correlation of LCC with CR - in the form of a negative
407 correlation of LCC with SSN (CR and SSN being inversely correlated). This follows
408 because we found that the positive LCC, CR averaged over all latitudes was greater
409 than the negative MCC, CR correlation averaged in the same way. Thus LCC plus
410 a fraction of MCC will have a positive CC, CR correlation. There is also a marked
411 upward trend in the CC with time.

412 For high clouds, the correlation is in the same sense as we found for the 11-year
413 analysis of HCC and MCC. Here, for high clouds there is a modest downward trend
414 in CC with time. The opposite trend with time for ‘low’ and ‘high’ clouds is in the
415 spirit of our suggestion referred to earlier of the inverse correlation in general of
416 LCC and MCC.

417 It is interesting that the ‘1970 dip’ so strongly marked in the ΔT plot (Figure 3b), is
418 associated with the peaks in low cloud and the minima in high cloud, both in Figure
419 5. This again agrees with the fact that there is a strong negative correlation between
420 LCC and MCC; indeed our explanation of the apparent LCC, CR correlation itself
421 is that it is SI (rather than CR) that is responsible, the mechanism being the heating
422 of the earth’s surface causing a change in mean cloud height (by about 40m from
423 1985 - 2005), which caused the LCC to fall and the MCC to rise.

424 Concerning the 11-year cycle in the synoptic cloud data, although none is visible
425 in the low, high and latitude-divided data in Figure 5 it is readily apparent in the
426 work of Pallé Bagó and Butler (2000), who combined the data of Norris (1999). The
427 authors found a peak-to-peak variation for a three-year running mean of the yearly
428 daytime total cloud cover over the ocean of $0.7 \pm 0.2\%$.

429 A slow rise in low cloud cover over the 50-year period with a somewhat smaller fall
430 in high cloud cover over the same period is a prominent feature of Figure 5. We
431 estimate an overall mean increase of $3.0 \pm 0.5\%$ over both cloud height ranges and
432 both latitude ranges. Pallé Bagó and Butler’s value is $2.5 \pm 0.5\%$. Some confirmation
433 comes from the summary of cloud data by Bryant (1997) who finds increases of $\sim 2\%$

434 over each of the N and S oceans and 3% over Europe. Interestingly, it is found that
435 there is a peak in cloud cover in 1945 for the N and S oceans, Europe and Canada,
436 with a peak in 1950 for the USA (but none for Australia and India).

437 Presumably they are connected with ‘our’ temperature peaks in 1940 and 1950 (see
438 Section 2.3.2).

439 It can be added that Pallé Bagó and Butler (2000) find support for the cloud cover
440 trend from sunshine records over 4 sites in Ireland, and provide arguments why such
441 records should have wider application.

442 *3.3 Discussion of the Cloud Data*

443 Concerning the 11-year cycle there is consistency in the sense that the sign of the
444 correlation of CC with SSN changes phase with increasing altitude for both the maps
445 of Voiculescu et al (2006) and our own analysis (Erlykin et al, 2009b). The reason
446 for the ‘sign’ reversal Haigh (2007), for temperature, at least, in which the Solar
447 UV causes changes in the atmospheric air circulation system and thereby mean air
448 temperature versus height. In this model the magnitude of the 11-year peak-to-peak
449 temperature cycle increases with height systematically, unlike in the cloud case,
450 where the CC ‘amplitude’ falls with height, as well as oscillating in sign. The answer
451 might be the role of ice crystals, which have different contributions as a function
452 of height. The same situation may be responsible for the difference between the
453 observed and expected latitude variation of CC.

454 Turning to the ‘22-y oscillation’, the observation of the same pattern in both the
455 30° - 60° N region and that for 30° S - 30° N, albeit with reduced magnitude, is
456 reassuring. Also reassuring is the accuracy with which the minimum SSN in 1970 and
457 maximum in 1985 is reproduced for the High Cloud (30° - 60° N) and the inverses for
458 the Low Cloud (30° - 60° N). Explanation of the cloud pattern in terms of systematic
459 errors is surely ruled out. Turbulent effects associated with the equatorial region can
460 be invoked to explain the reduced peak-to-peak variation, in the near Equatorial

461 regions.

462 The long-term drift visible in both Figures 5a and 5b can be considered next. Norris
463 (2000) has argued that the trend may be spurious but, in our view, the consistency
464 between the two latitude ranges and the analyses by other authors, argues against
465 this likelihood. Comparison of the trends can be made with the ISCCP data. The
466 latter show, for LCC, a distinct fall of mean cloud cover with time from 1984 onwards
467 and this is opposite to the trend found for the synoptic cloud results for ‘low’ clouds.
468 It would be surprising if the difference in the ‘low’ definitions had such a big effect.
469 More relevant is to examine the trend for the whole cloud amount, ie ‘low’ and ‘high’
470 (synoptic) with LCC + MCC + HCC. The former is still positive (Figures 5a and
471 5b) at the rate of about +3% per 30 years. For the ISCCP data the overall trend is
472 -3% per 20 years. There is a clear discrepancy. The situation of CC with respect to
473 temperature will be discussed next.

474 **4 The relationship of surface temperature and cloud cover**

475 It is generally accepted that low clouds are negatively correlated with surface tem-
476 perature because of their degree of absorption of incoming radiation and high clouds
477 have a positive correlation because of their ‘greenhouse effect’, ie reflection of ter-
478 restrial infra-red radiation.

479 Comparison of Figures 3b and 5a, b show that for the 22 y ‘cycle’ this is indeed the
480 case; the temperature dip in 1970 corresponds to the high value of low cloud cover
481 in Figure 5a and the low value of high cloud cover in Figure 5b. This is a satisfying
482 result.

483 Less satisfying is the situation with the long term trends in CC in Figures 5a and 5b.
484 It will be remembered that the upward temperature trend has been removed from
485 Figure 3b. Thus, we have an overall temperature rise associated with a rising low
486 cloud cover and a slightly falling high cloud cover, ie opposite to the situation for the
487 22-year modulation. Although by no means certain, it is likely that this situation

488 arises because the ‘Global Warming’ is due to anthropogenic materials which do not
 489 result in changes to the cloud cover in the same sense as that for SI changes. The
 490 manner in which the magnitude of the CC change depends on the ‘time interval’
 491 can be considered, in a similar way to that for temperature (Section 2.4.2).

492 Summarising, we have the following $\Delta SI/SI$ changes needed to change the CC by
 493 2%.

	1 day	29% (night/day, typically)	29%
	1 year	(latitude variation of ISCCP data)	6%
494	11 year	(LCC)	0.1%
	22 year	(50y analysis, here)	0.01%

495 As with the temperature changes, the trend seems reasonable.

496 **5 The last 20 years**

497 The SSN has been anomalously low for several years and currently shows no sign
 498 of increasing! The associated CR intensity may have at last reached its maximum,
 499 at least there is a flattening in the contemporary neutron monitor rates at some
 500 locations, McMurdo and Inuvik, sites in the Arctic region where particle of the
 501 lowest energy - and which are most affected by the solar wind - are able to reach the
 502 Earth along the magnetic field lines (Bartol neutron monitors, 2009). However other
 503 sites show a continuing rise. There is agreement, however, that the CR intensity is
 504 higher than has ever been previously recorded (eg Calgary neutron monitor, 2009).

505 The CR peak is interesting in that normally the SSN would have started to rise
 506 up to a year previously (the well known hysteresis effect caused by CR diffusion)
 507 but it did not. A further anomaly has been the unusually low peak SSNo for Cycle
 508 23 - the peaks for odd numbered cycles have usually been higher than the mean of
 509 their neighbours, as can be seen in Figure 2. Interestingly, the UV profile for Cycle
 510 23 was normal. The CR time-profile for the Cycle was also anomalous (Ahluwalia

511 and Ygbuhay; 2008) in the sense that there was a ‘shoulder’ in the neutron monitor
512 count rate in 2004 and 2005.

513 All these facts make us aware of a possible anomaly in the temperature record
514 (Figure 6) - there was a dramatic drop in ΔT for 2008. Specifically, ΔT fell by
515 0.20°C whereas an increase of $\sim 0.1^{\circ}\text{C}$ might have been expected, since anthro-
516 pogenic heating is supposed to be gaining ground. It might be thought conceivable
517 that all the anomalies were related and that the 0.2°C cooling was indicative of a
518 mechanism related to the prolonged very low SSN, or the forecast by Ermakov et
519 al (2009) involving the frequency distribution of past temperature periodicities and
520 the involvement of periodicities for the arrival of cosmic dust - which should have
521 implications for climate.

522 It is relevant at this stage to point out that there have been forecasts of ‘Global
523 Cooling’. For example, Landscheidt (2000) found a correlation of Global temperature
524 with the strength of the solar wind and used the fall in wind strength (as indicated
525 by the ‘aa’ geomagnetic parameter) after 1990 to predict a cooling, after a phase - lag
526 of some 4-6 years. Clearly, such a cooling did not start as predicted (in 1994-1996)
527 but an open mind should be kept about the cooling prospect.

528 The statistical significance of the 2008 reduction is not high, however, for the fol-
529 lowing reasons:

- 530 (i) The dispersion of yearly values about the smoothed temperature variation with
531 time (shown dashed) is as indicated ‘obs’ in Figure 6. If many factors contribute
532 to the temperature variation in a random way (a rather uncertain assumption,
533 although there are, in fact, several independent contributors) the probability of
534 the 2008 deviation can be estimated. It is at the 2 standard deviations level, ie a
535 2.5% probability.
- 536 (ii) A 0.2°C or more dip has occurred on 16 occasions in the period of the observations,
537 128 y, ie a probability per year, if random, of 12.5%
- 538 (iii) In fact, there are corrections to be applied to the temperature record for ENSO,
539 volcanoes, ozone and greenhouse gas fluctuations, as described earlier. The stan-

540 dard deviations for the 5-y and 1-y corrections (none of which have been applied
541 to the data) are also shown in Figure 6. If added to the (obs.) fluctuations in
542 quadrature, the significance of the 2008 dip falls to about one standard deviation,
543 ie $\sim 16\%$, assuming a Gaussian distribution.

544(iv) The ‘noise’ correction (ENSO + etc) is not, in fact, completely random, largely
545 because of the near-periodicity in ENSO. This feature is presumably responsible
546 for the fact that most of the minima in Figure 6 are separated by 2y. In a sense,
547 a reduction in 2008 might have been predicted from known mechanisms.

548 Taking all the factors above into account, no case can be made for the 2008 minimum
549 having an anomalous origin.

550 It can be remarked that the shallow convexity in the dashed curve is consistent with
551 the ‘standard’ SI effect.

552 **6 Conclusions**

553 The relation between changes in solar irradiance and changes in mean surface tem-
554 perature and cloud cover at various levels has been examined. The comparatively
555 high temperature changes associated with changes in the solar irradiance (as evinced
556 by change in sunspot numbers) are confirmed; changes in UV as distinct from longer
557 wavelengths are a strong candidate. Positive feedback may also be a contributory
558 process, such as occurs in the Arctic where an increased surface temperature melts
559 ice and reduces the albedo so that the temperature rise is enhanced. A similar situa-
560 tion pertains in the stratosphere. The increase in ΔT with respect to expectation as
561 a function of ‘integration time’ - 1y, 11y and 22y - points to an inertial effect, ie the
562 time-constant of the atmosphere/Earth’s surface temperature system, but it must
563 be said that the difference between the 22-year and 11-year temperature variations
564 - a factor ~ 5 - is rather dramatic. Our result is bigger than the factor 1.7 found by
565 Miyahara et al (2008) (for 26-year and 12-year cycles) from tree ring data over the
566 last 500 years.

567 It might be thought that the 22-year cycle temperature change was due to a solar
568 wind effect, in view of the well-known 22-year cycle of the solar magnetic field
569 direction. Indeed, a cosmic ray origin might even be postulated. However, inspection
570 of the CR induced atmospheric ionization versus time (eg Bazilevskaya et al, 2008)
571 does not show an adequate 22-year modulation. Instead, we prefer an intrinsic-to-the
572 sun origin, involving ultra-violet radiation together with positive feedback.

573 The changes in cloud cover correlate appropriately with the temperature changes on
574 the 20-30 y scale but the mechanism is still unclear. The longer term slow increase in
575 low cloud cover (and small reduction in high cloud cover) is of the opposite ‘sign’ to
576 expectation. An explanation in terms of anthropogenic causes for the temperature
577 rise seems likely, although it cannot be ruled out that the slow cloud cover changes
578 are an artifact.

579 The lack of an explanation for the actual geographical pattern of the strong correla-
580 tion of low cloud cover with solar irradiance (negative)(see Section 1), is still present.
581 An explanation in terms of the changes to the atmospheric circulation ‘geography’
582 not being as predicted by the model of Haigh (2007) is a distinct possibility, although
583 the magnitude of the effect may well be as has been estimated.

584 The mechanism responsible for the temperature and Cloud Cover changes is clearly
585 ‘solar’ but whether the initiating energy is supplied by radiation (UV, as described)
586 or whether it is the solar wind is not yet clear. However, in view of the energy in
587 the solar wind being only of order one millionth of that in sunlight, the solar wind
588 hypothesis has difficulties: a ‘positive feedback’ of the magnitude required would
589 appear to be very unlikely.

590 Our estimate of ‘less than 14% for the period 1956-2002 is confirmed. For the pre-
591 vious 50 years, changes in solar irradiance appear to be responsible.

592 Thus, extra forcing (presumably anthropogenic) started to become important only
593 in the 1950s. This conclusion confirms that of Lean et al (2005), and others (notably
594 that of the IPCC).

595 **7 Acknowledgement**

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597 support of this project. Dr Joan Kenworthy is thanked for helpful comments on the
598 manuscript.

599 **Table 1. Sources of the data for Figure 1 : temperature change, ΔT , ver-**
600 **sus change in solar irradiance, ΔSI**

601

‘Number’	Remarks	References	Time period
1	‘The sensitivity of the Earth’s surface temperature to changes in irradiance’	Lean et al (2005) from IPCC	≈ 10 y
2	The 11-year temperature anomaly	Lean et al (2005); Haigh (2007)	11 y
3	The long scale 1970 dip	Erlykin et al (2009a)	20 - 30 y
4	Maunder Minimum (variable over the Earth)	Lean et al (2005), mean of range for ΔSI	~ 100 y
5	Decadal variations of sea surface temperatures (very variable)	White et al (1997) Van Loon et al (2007)	10 y
6	Orbital changes (calculations for $65^\circ N$, for periods with no magnetic field changes)	Rusov et al (2008)	10^5 y
7	Global seasonal - summer, winter temperature differences vs latitude	ΔT data from Allen (1973)	1 y
8	Changes over the last 10^4 y	IPCC (1990)	10^4 y
9	Change since the last Ice Age	Lean et al (2005)	$2 \cdot 10^4$ y
10	The average from Figure 3b	The present work	20-30 y

602

603 **Table 2. Correlation probabilities and slopes of the line, ΔT (obs.) vs ΔT**
 604 **(predicted) for the period 1000-1900AD. The ‘correlation probability’ is**
 605 **the chance of a random correlation giving the observed value or greater;**
 606 **the number of standard deviations from zero for the slope is another in-**
 607 **dicator of the significance of the correlation.**

608

Period	Correl. prob.	Slope
1000 - 1100	0.327	0.36 ± 0.20
1100 - 1200	0.167	0.63 ± 0.28
1200 - 1300	0.000	1.55 ± 0.15
609 1300 - 1400	0.037	0.399 ± 0.092
1400 - 1500	0.013	1.06 ± 0.21
1500 - 1600	0.013	0.883 ± 0.25
1600 - 1700	0.248	0.379 ± 0.37
1700 - 1800	0.021	0.922 ± 0.19
1800 - 1900	0.07	1.11 ± 0.18

610 8 References

- 611 1. Ahluwalia,H.S. and Ygbuhay,R.C., 2008. Observed Galactic Cosmic Ray 11-year
612 modulation for Cycle 23, 30th Int. Cosmic Ray Conf., Merida, Mexico, 493.
- 613 2. Allen,C.W., 1973. Astrophysical Quantities, Athlone Press, Univ. of London.
- 614 3. Bartol Neutron Monitors: (2009) <<http://neutron.bartol.udel.edu.html>>
- 615 4. Bazilevskaya,G.A. et al (12 authors), 2008. Ionization Processes in Planetary At-
616 mospheres Cosmic Ray Induced Ion Production in the Atmosphere, Space Science
617 Reviews 137, 1.
- 618 5. Bryant,E., 1997. Climate processes and Change, Cambridge University Press.
- 619 6. Calgary Neutron Monitor: (2009)
620 <<ftp://ftp.ngdc.noaa.gov/STP/SOLAR-DATA/COSMIC-RAYS/calgary.tab>>
- 621 7. Crowley,T.J., 2000. Causes of Climate Change over the past 1000 years, Science,
622 289, 270.
- 623 8. Deland,M.T. and Cebula,R.P., 2008. Creation of a composite solar ultraviolet
624 irradiance data set, submitted to Journal of Geophysical Research available from
625 <<http://lasp.colorado.edu/lisird/delandcomposite/index.html>>
- 626 9. Eichler,E., Olivier,S., Henderson,K., Laube,A., Beer,L., Papina,T., Gaggeler,H.W.,
627 and Schwikowski,M., 2009. Temperature response in the Altai region lags forcing,
628 Geophysical Research Letters, 36, L01808.
- 629 10. Erlykin,A.D., Sloan,T. and Wolfendale,A.W., 2009a. Solar Activity and the Mean
630 Global Temperature, Environmental Research Letters, 4, 014006.
- 631 11. Erlykin,A.D., Gyalai,G., Kudela,K, Sloan,T. and Wolfendale,A.W., 2009b. ‘On
632 the correlation between cosmic ray intensity and cloud cover’, Journal of Atmo-
633 spheric and Solar-Terrestrial Physics, DOI: 10.1016/j.jastp.2009.06.012; arXiv:0906.4442
- 634 12. Erlykin,A.D., Sloan,T. and Wolfendale,A.W., 2009c. ‘Correlations of clouds, cos-
635 mic rays and solar irradiation over the earth’, Solar Physics (submitted for pub-
636 lication)
- 637 13. Foukal,P., North,G. and Wigley,T., 2004. A stellar view on solar variations and
638 climate. Science 306, 68.
- 639 14. Foukal,P., Frohlich,C., Spruit,H., and Wigley,T.M.L., 2006. Variations in solar
640 luminosity and their effect on the Earth’s climate. Nature, 443, 161.

- 641 15. Gray,L.J., Rumbold, S.T. and Shine, K.P., 2009. Stratospheric temperature and
642 radiative forcing response to 11-year solar cycle change in irradiance and ozone.
643 DOI:10.1175/2009JAS 2866.1.
- 644 16. Haigh,J.D., 2007. The Sun and the Earth's Climate, Living Rev. Solar Physics 4,
645 2.
- 646 17. Hansen,J.E., Ruedy,R., Sato,M., Imhoff,M., Lawrence,W., Easterling,D., Peter-
647 son,T., and Karl,T., 2001. A closer look at United States and global surface tem-
648 perature change, Journal of Geophysical Research 106, 23947, updated by GISS,
649 NASA 2009-01-09.
- 650 18. Hegerl,G.C. et al, 2007. 'Understanding and Attributing Climate Change 2007',
651 CUP, UK and NY.
- 652 19. Hood,L.L., 1997. The solar cycle variation of total ozone : dynamical forcing in
653 the lower stratosphere. Journal of Geophysical Research 102, 1355.
- 654 20. Hood,L.L. and Soukharev,B.E., 2000. The solar component of long-term strato-
655 spheric variability: observations, model comparisons and possible mechanisms.
656 Proc. SPARC 2nd General Assembly November 6-10.
- 657 21. IPCC: Climate Change, The IPCC Scientific Assessment, Ed Houghton,J.D.,
658 Jenkins,G.J. and Ephraums,J.J., Cambridge University Press, (1990).
- 659 22. IPCC : Climate Change 2007: The Physical Basis. Contribution of Working Group
660 1 to the Third Assessment Report of the International Panel on Climate Change,
661 Cambridge University Press, Cambridge, UK and New York, USA.
- 662 23. The ISCCP data were obtained from the International Cloud Climatology Project
663 website : <http://isccp.giss.nasa.gov/> maintained by the ISCCP research group at
664 NASA Goddard Institute for Space Studies, New York, Rossow,W.B. and Schif-
665 fer,R.A., 1999. 'Advances in understanding clouds from the ISCCP', Bulletin of
666 the American Meteorological Society 80, 2261.
- 667 24. Khristjansson,J.E. and Kristiansen,J., 2000. Is there a cosmic ray signal in recent
668 variations in global cloudiness and cloud radiative forcing?, Journal of Geophysical
669 Research 105(11), 851.
- 670 25. Landscheidt,T., 2000. The Solar Cycle and Terrestrial Climate, ESA SP-463.
- 671 26. Lean,J., Rottman,G., Harder,J. and Kopp,G., 2005. Source contributions to new

- 672 understanding of global change and solar variability, *Solar Physics* 230, 27.
- 673 27. Marsh, N. and Svensmark, H., 2000. Low cloud properties influenced by cosmic
674 rays, *Physical Review Letters*. 85, 5004.
- 675 28. Meehl, G.A., Washington, W.M., Ammann, C.A., Arblaster, J.M., Wigley, T.L.M.
676 and Tebalid, C., 2004. Combinations of Natural and Anthropogenic Forcings in
677 the Twentieth-Century Climate, *Journal of Climate* 17, 3721.
- 678 29. Miyahara, H., Yokoyama, Y. and Masuda, K., 2008. Possible link between multi-
679 decadal climate cycles and periodic reversals of solar magnetic polarity. *Earth
680 and Planetary Science Letters* 272, 290.
- 681 30. Mohakumar, M., 1988. Response to an 11-year solar cycle on Middle Atmosphere
682 Temperature, *Physica Scripta* 37, 460.
- 683 31. Nagashima, T., et al, 2006. The effect of carbonaceous aerosols on surface tem-
684 perature in the mid twentieth century, *Geophysical Research Letters* 33, L04702.
- 685 32. Norris, J.R., 2004. Low Cloud Type over the Oceans from surface observations,
686 Part II : Geographical and Seasonal Variations, *Journal of Climate* II (1998), 383,
687 and "Change in Near-Global Cloud Cover and Reconstructed Radiation Flux since
688 1952". Private communication.
- 689 33. Norris, J.R., 2000. What can cloud observations tell us about climate variability,
690 *Space Science Reviews* 94, 375.
- 691 34. Pallé Bagó, E. and Butler, C.J., 2000. Sunshine, clouds and cosmic rays. star.arm.ac.uk/rambn/345.pdf
- 692 35. Peixoto, J.P. and Oort, A.H., 1992. *Physics of Climate*, American Institute of Physics.
- 693 36. Rusov, V. et al, 2008. Galactic Cosmic Rays - Clouds effect and bifurcation model
694 of the Earth's global climate. Private Communication (siiis@te.net.ua)
- 695 37. Shindell, D., Rind, D., Balachandran, N., Lean, J. and Lonergan, P., 1999. Solar cycle
696 variability, ozone and climate, *Science* 284, 3050-58.
- 697 38. Sloan, T. and Wolfendale, A.W., 2008. Testing the proposed causal link between
698 cosmic rays and cloud cover, *Environmental Research Letters* 3, 024001.
- 699 39. UKSSDC, 2009 : <ftp://ftp.ukssdc.ac.uk/wdc/a.w.wolfendale@durham.ac.uk/>.
- 700 40. Vecchio, G.L., and Nanni, T., 1995. The atmospheric temperature in Italy during
701 the last 1000 years and its relationship with solar output, *Theoretical and Applied
702 Climatology* 51, 159.

- 703 41. Van Loon, H., and Shea, D.J., 2007. A probable signal of the 11-year solar cycle
704 in the troposphere of the northern hemisphere, *Geophysical Research Letters* 26,
705 2893.
- 706 42. Viereck, R.A. and Puga, L.C., 1999. The NOAA MgII core-to-wing index: Con-
707 struction of a 20 year time series of chromospheric variability from multiple satel-
708 lites, *Journal of the Geophysical Research* 104, 995-10006.
- 709 43. Voiculescu, M., Usoskin, I.G. and Mursula, K., 2006. Different response of clouds to
710 solar input, *Geophysical Research Letters* 33, L21802.
- 711 44. Wang, Y-M, Lean, J.L., Sheeley Jr., N.R., 2005. 'Modeling the Sun's Magnetic Field
712 and Irradiance since 1913', *Astrophysical Journal* 625, 522-538.

713 **9 Captions to Figures**

714 **Figure 1** Changes in mean surface temperature of the Earth (ΔT) versus changes
715 in the solar irradiance (ΔSI). The sources of the data are given in Table 1. ‘Model’ is
716 expectation from the observation that for all planets the mean surface temperature
717 varies as $(SI)^{\frac{1}{4}}$.

718 **Figure 2** Sunspot number versus time since 1990. The downward arrows represent
719 periods when the long term mean was low and where (in Figure 3b) we identify
720 temperature dips. The Cycle numbers are indicated along the top.

721 **Figure 3a** Average surface temperatures versus time for the whole Globe (GLO),
722 land (LAN) and the oceans (OCE). The shaded areas represent predictions for spe-
723 cific models, the upper regions include anthropogenic changes and the lower regions
724 without. The Figures are from Hegerl et al (2007).

725 **Figure 3b** Change in temperature (ΔT) and sunspot number (SSN) smoothed (over
726 11-years) from Erlykin et al (2009a); the smooth slow increase in temperature, ie the
727 ‘Global Warming’, has been removed. We use the SSN as the proxy for SI. ΔT_{corr} is
728 the change in temperature after correction by us for the effects of volcanoes, ENSO,
729 ozone and greenhouse gases.

730 **Figure 4** ΔT vs ΔSSN for the 5 dips shown in Figure 3b, with 3 pairs of values
731 for each dip. The number (10, 30 ...90) is the year (1910, 1930...) for the minimum
732 temperature in that Figure. The dashed line shows the best fit. The probability of
733 obtaining a ‘zero-fit’ is 0.009, confirming a good correlation.

734 **Figure 5a** Changes of cloud cover (ΔCC) for low clouds for the latitude ranges
735 indicated, and associated changes in sunspot number (ΔSSN). The cloud data are
736 from Norris (2004).

737 **Figure 5b** Changes of cloud cover (ΔCC) for high clouds for the latitude ranges
738 indicated, and associated changes in sunspot number (ΔSSN). The cloud data are
739 from Norris (2004).

740 **Figure 6** ΔT versus year for the last 20 years (from Hansen et al, 2001, updated
741 by GISS, NASA, 2009) The dashed curve is the 5-year average. 'obs' denotes the
742 median spread of the yearly points about the mean line. (5) and (1) refer to the
743 median spread of the corrections (which have not been applied) for ENSO, volcanic
744 forcing, etc.

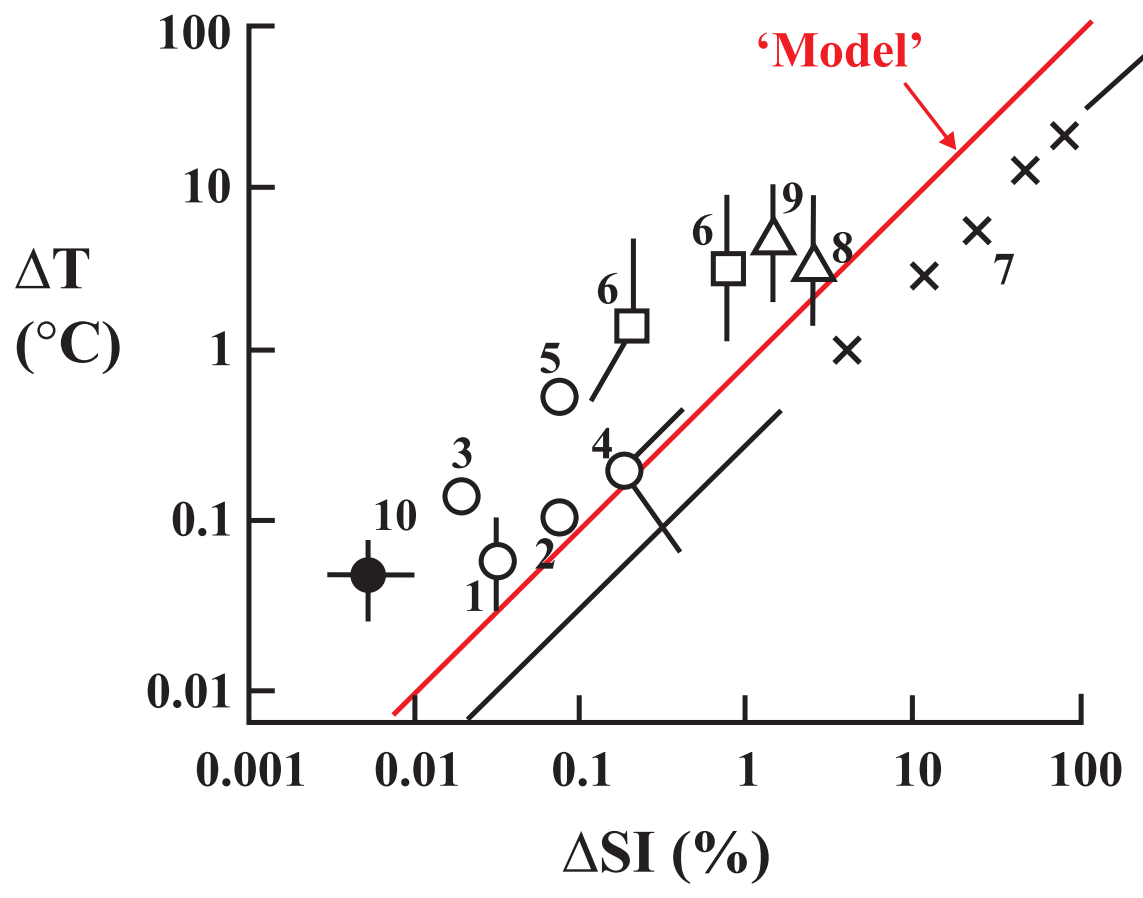


Fig. 1

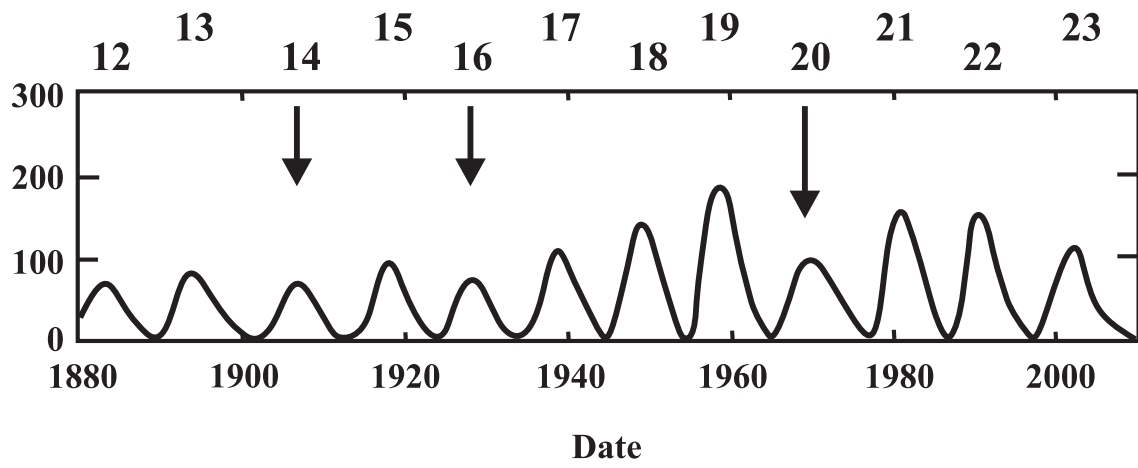


Fig. 2

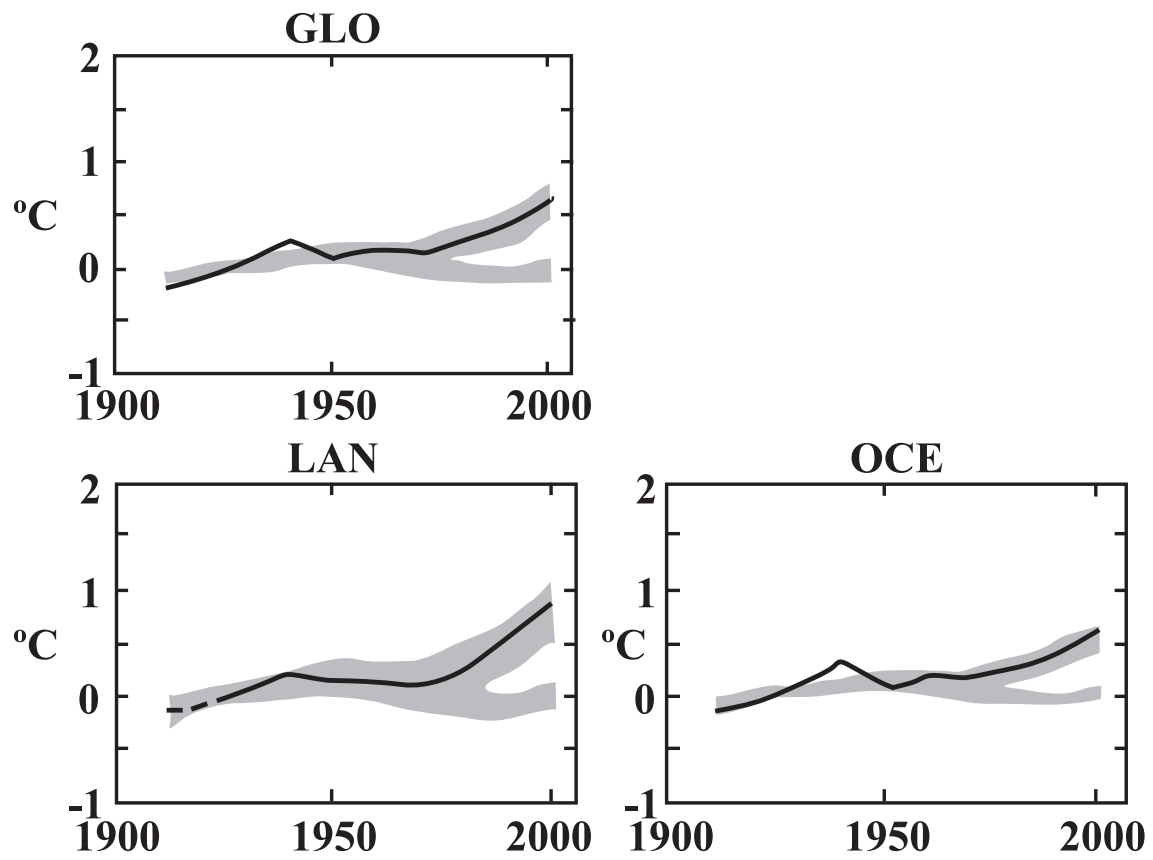


Fig. 3a

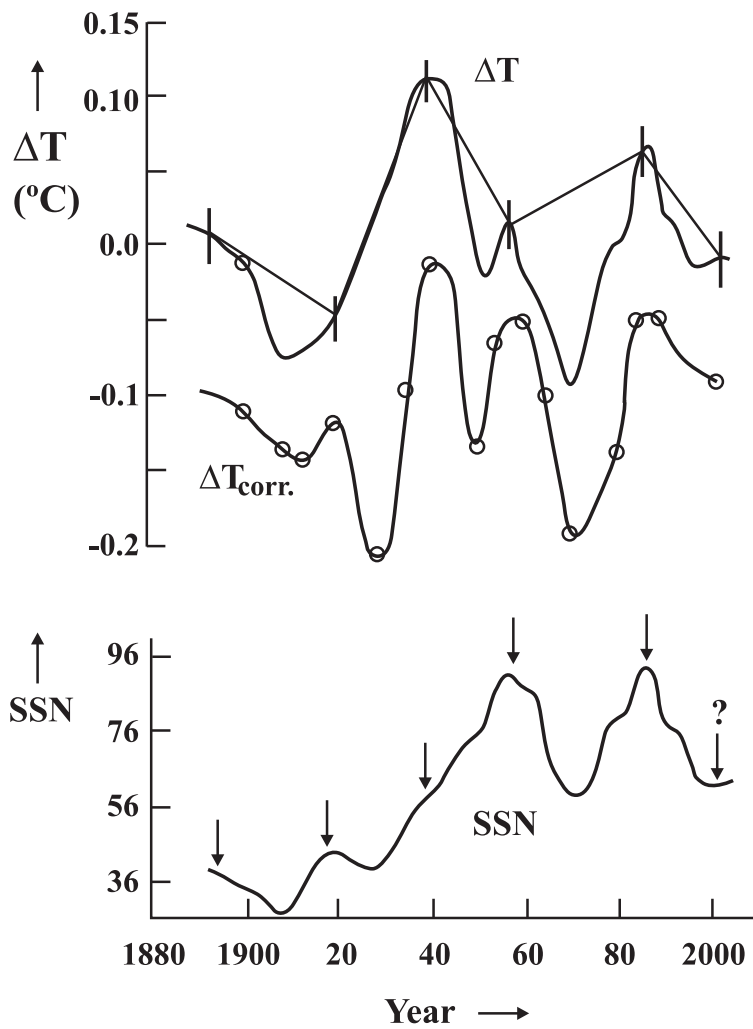


Fig. 3b

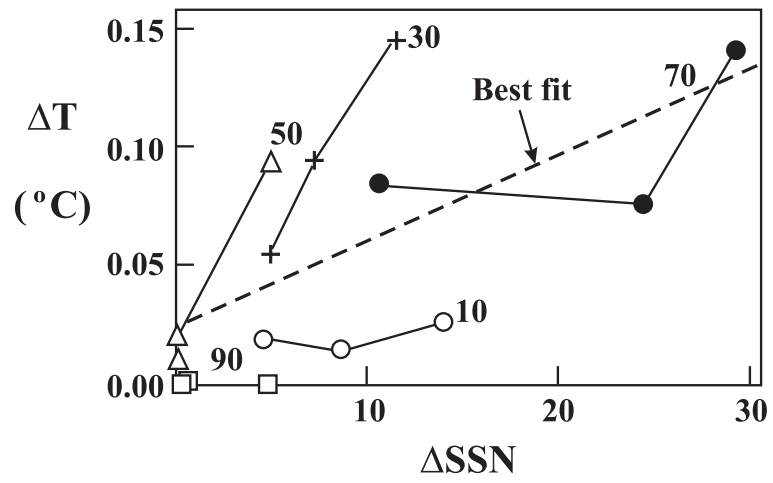


Fig. 4

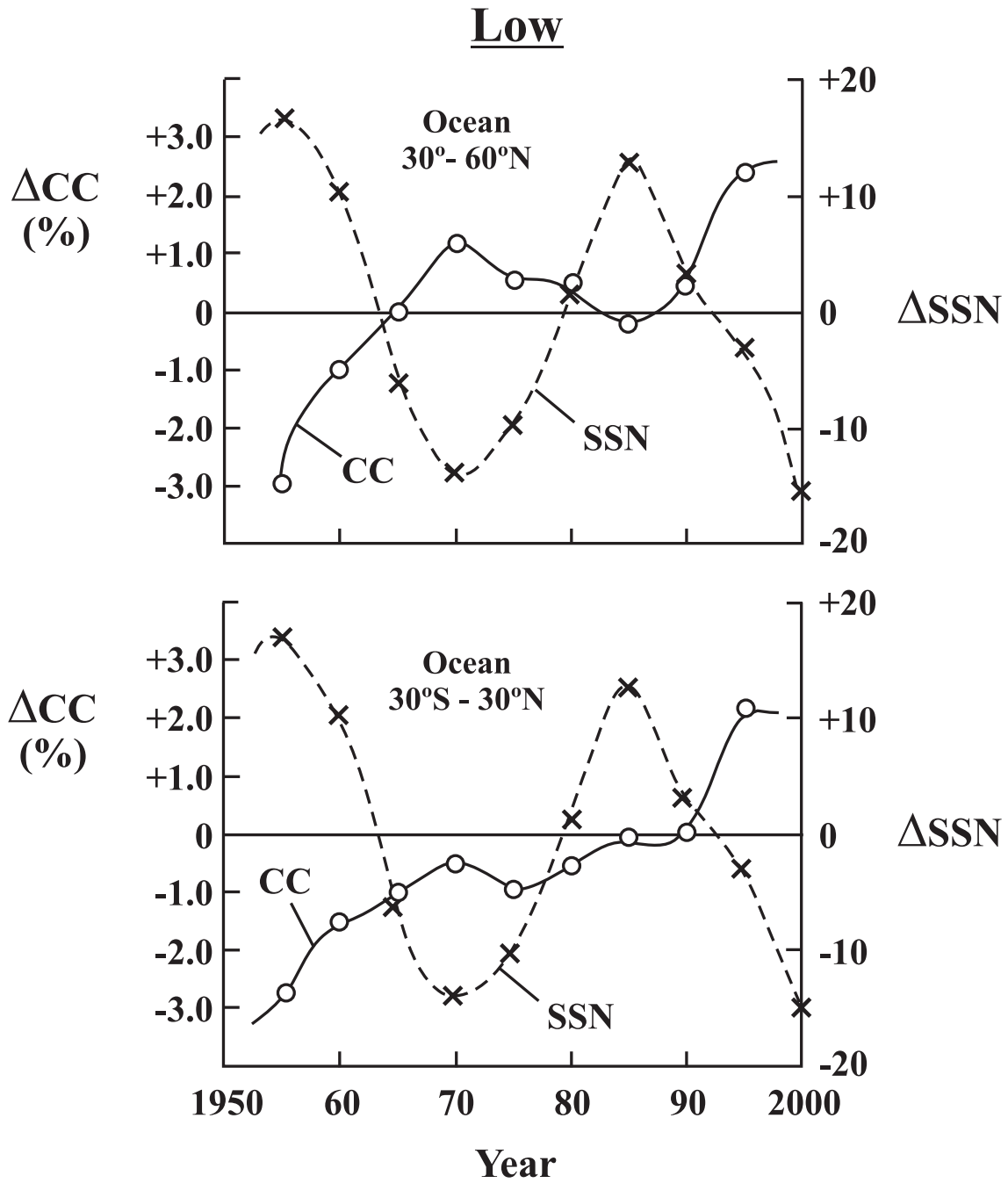


Fig. 5a

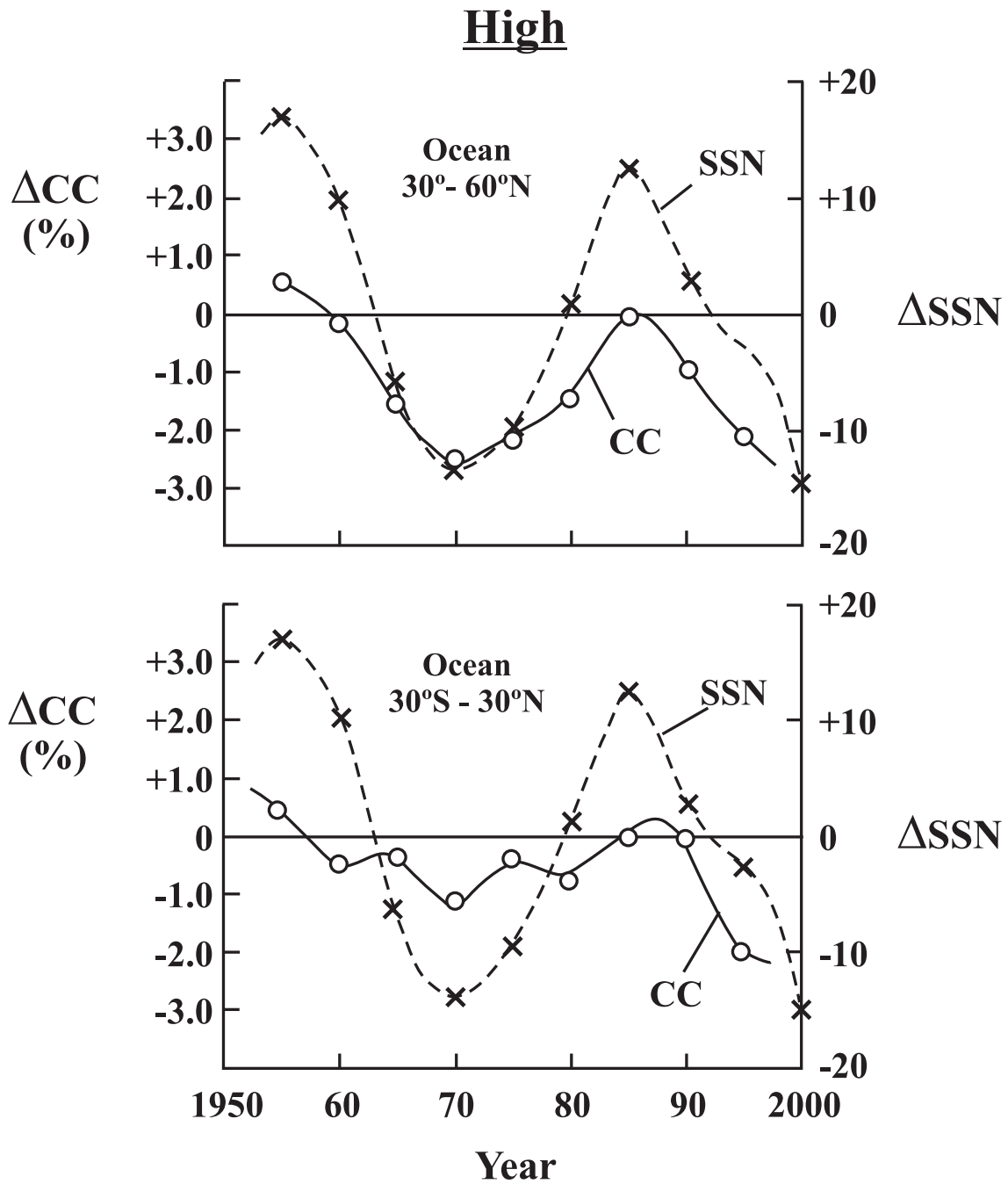


Fig. 5b

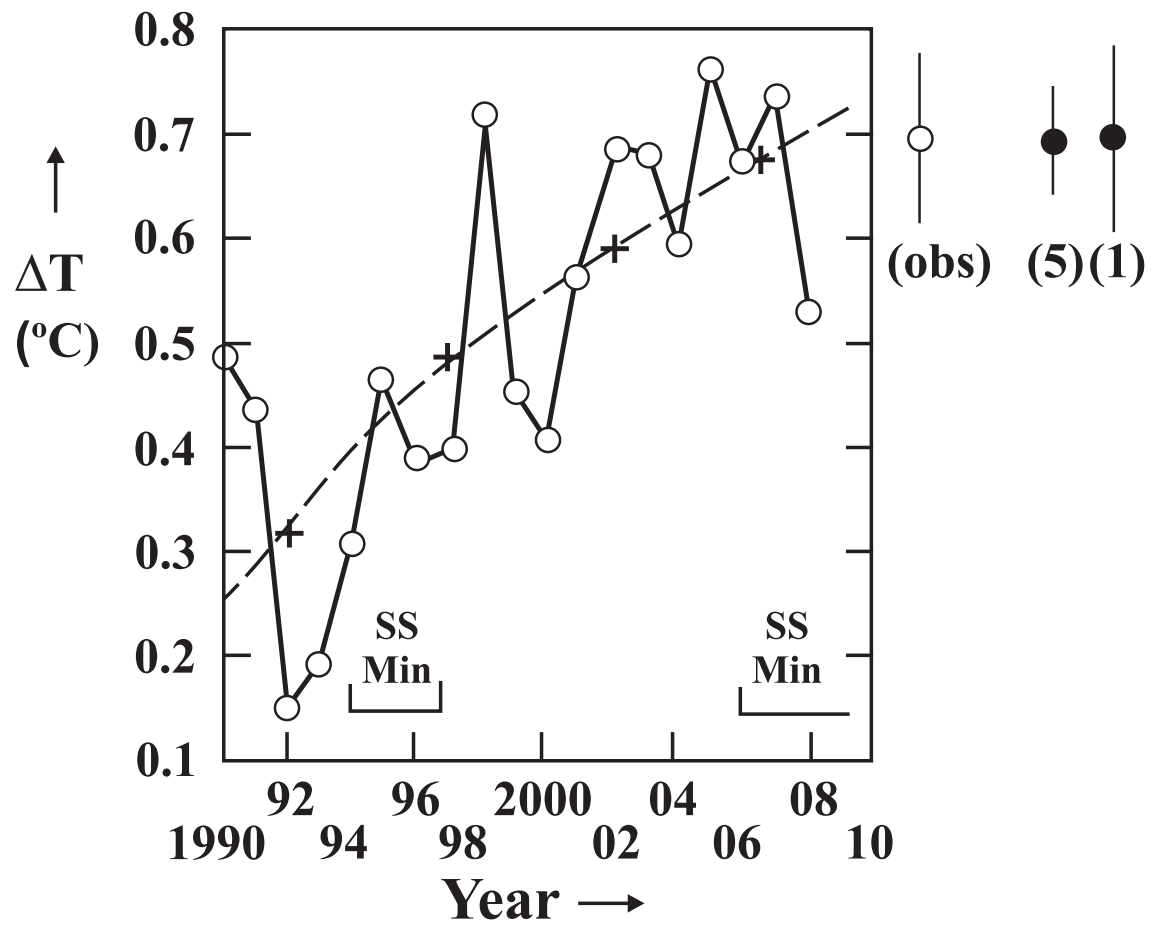


Fig. 6