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## COMPARISON OF SUPERCRITICAL CO<sub>2</sub> WITH WATER AS GEOFLUID IN GEOTHERMAL RESERVOIRS WITH NUMERICAL INVESTIGATION USING FULLY COUPLED THERMO-HYDRO-GEOMECHANICAL MODEL

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6                   **GEOMECHANICAL MODEL**

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48 **ABSTRACT**

49 *In the present work, fully coupled dynamic thermo-hydro-mechanical (THM) model was employed to*  
50 *investigate the advantage and disadvantages of supercritical CO<sub>2</sub> (SCCO<sub>2</sub>) over water as geofluids. Low-*  
51 *temperature zone was found in both SCCO<sub>2</sub>-EGS and water-EGS systems, but spatial expansion is higher in*  
52 *water-EGS. Although, the spatial expansion of SCCO<sub>2</sub> into the rock matrix will help in the geo-sequestration.*  
53 *The expansion of stress and strain invaded zones were identified significantly in the vicinity of fracture and*  
54 *injection well. SCCO<sub>2</sub>-EGS system is giving better thermal breakthrough and geothermal life conditions*  
55 *compared to the water-EGS system. Reservoir flow impedance (RFI) and heat power are examined, and heat*  
56 *power are high in the water-EGS system. Minimum RFI is found in the SCCO<sub>2</sub>-EGS system at 45°C and 0.05*  
57 *m/s. Maximum heat power for SCCO<sub>2</sub>-EGS was observed at 35°C, 20 MPa, and 0.15 m/s. Therefore, the*  
58 *developed dynamic THM model is having greater abilities to examine behaviour of SCCO<sub>2</sub>-EGS and water-*  
59 *EGS systems effectively. The variations occur in the rock matrix and the performance indicators are*  
60 *dependent on the type of fluid, injection/production velocities, initial reservoir pressure, injection*  
61 *temperature. The advantages of SCCO<sub>2</sub>-EGS system over the water-EGS system, providing a promising result*  
62 *to the geothermal industry as geofluid.*

63 **Keywords:** *SCCO<sub>2</sub>-EGS, water-EGS, geofluid, thermo-hydro-geomechanical, reservoir flow impedance, heat*  
64 *power*

65 **1. Introduction**

66 Energy demand and carbon emissions from the fossil fuels have become critical issues in  
67 the many developed and developing countries. To meet the energy requirement  
68 geothermal is the prominent resource due to its availability and it is clean, ecofriendly,  
69 and the renewable resource [1–5]. The amount of geothermal energy resource is  
70 estimated as 40-400 M EJ ( $1\text{EJ}=10^{18}\text{J}$ ) and it is approximately 100-1000 folds than fossil  
71 fuels [6]. Enhanced geothermal systems will be producing heat from the low porosity and  
72 low permeable zones to meet the energy requirements. The carbon dioxide emissions  
73 from the fossil fuels will lead to increase global warming. Capturing and geological storing  
74 of  $\text{CO}_2$  have been major options for sequestering  $\text{CO}_2$ ; the deep oil/gas reservoirs, and  
75 deep saline aquifers are used for geological storing purpose.

76 In the recent years, the  $\text{CO}_2$  becoming most attractive as an alternative geo fluid for  
77 extracting heat from the geothermal reservoirs. Brown first proposed the advantages of  
78  $\text{CO}_2$  as geofluid compared to the water [7]. High compressibility, expansivity, and low  
79 viscosity are the advantages of  $\text{CO}_2$  over the water. These will increase the mobility in the  
80 rock matrix and fracture compared to water. This will also reduce the pumping power  
81 required to inject the geofluid into the hot rock to extract the heat. Another advantage of  
82 the fluid from the prospect of the geo sequestration, if the fluid loss occurs from the main  
83 hydraulic fractures in to rock matrix can have ability to permanently sequester the  $\text{CO}_2$   
84 [8–14]. In the geothermal reservoirs, hydraulic fractures play an important role in the heat  
85 extraction. This is because of the higher fracture permeability compared to the rock  
86 matrix [13,15–18]. The thermal energy recovered from the geothermal reservoir by  
87 injecting the relatively low temperature fluid into the hydraulic fracture from injection  
88 well and producing high temperature fluid at the production well (i.e., the heat collecting  
89 from the rock matrix) [17,19–22]. Fractures act as preferential flow channels in the  
90 geothermal reservoir for carrying cold fluid from injection location to production location  
91 and extracting heat energy from the rock matrix [10,23–26]. Geothermal reservoir  
92 performance during the thermal energy production is mainly rely on the magnitude  
93 permeable channels/fractures in the rock matrix for the movement of geofluid [27–29].  
94 Development of geofluid flow field in the porous/fracture media is highly reliant on the

95 thermal discrepancy in the rock matrix [27–32], which will be altering the temperature of  
96 surrounding rock matrix. It can generate the variations in the porous media will lead to  
97 compaction and expansion. Therefore, the type of fluid and physical properties and the  
98 fracture and rock properties will play a key role in the heat production of geothermal  
99 reservoir.

100 The advantages of super critical CO<sub>2</sub> (SCCO<sub>2</sub>) over the water as geofluid in the geothermal  
101 reservoir for heat extraction is gaining more importance. Several researchers found that  
102 the SCCO<sub>2</sub> showing promising numerical results due to the high compressibility, high  
103 expansivity, low viscosity, low density, and low chemical interaction with the rock matrix  
104 over the water [8–14,33–38]. Pan et al. [34] studied the injection of supercritical CO<sub>2</sub> into  
105 a permeable reservoir which is initially saturated with water using coupled wellbore-  
106 reservoir system. They also conducted the sensitivity analysis on the mass flow rate,  
107 injection temperature, and reservoir permeability. They found that these parameters  
108 influencing the extraction of heat from the reservoir. Biagi et al [8] proposed an injection  
109 scenarios used the TOUGH2 for the simulating the geothermal reservoir for heat  
110 production using SCCO<sub>2</sub>. They mentioned that the decline in the heat extraction rate due  
111 to the cooling of the reservoir. Liang et al [33] investigated the interactions of  
112 water/+gas(CO<sub>2</sub>) in the geothermal reservoir for the production of heat. They included  
113 the variation of porosity and permeability as functions of geochemical reactions initiated  
114 by the injection of CO<sub>2</sub> into hot reservoir. They didn't consider the variation of other  
115 properties of rock and fluid properties while injection/extraction heat. Yin et al [36] is also  
116 working on the injection of CO<sub>2</sub> into the porous media with change in porosity and  
117 permeability in a carbonate reservoir. Zhang et al [37] investigated the impact of CO<sub>2</sub>-EGS  
118 and water-EGS in terms of net power, thermal efficiency, and exergy efficiency. They  
119 found that CO<sub>2</sub>-EGS produces more power compared to the water-EGS system. Zhang  
120 et al [11] working on the CO<sub>2</sub> assisted heat recovery from the high temperature gas well  
121 using the CMG-star simulator. Most the researchers are not considered the dynamic  
122 behaviour of the reservoir and the fluid in their research.

123 From the abovementioned, the present work is focused on the comparison of super  
 124 critical CO<sub>2</sub> (SCCO<sub>2</sub>) as geofluid with water using the dynamic behaviour of rock matrix,  
 125 fracture and also the fluid properties. The evolution of rock properties, fluid properties,  
 126 and temperature with SCCO<sub>2</sub> were examined exclusively with the comparison of water.  
 127 The properties of the injection fluids are considered above the super critical condition for  
 128 both CO<sub>2</sub> and water. In additions to that the performance parameters/indicators such as  
 129 reservoir impedance, geothermal life, breakthrough time, and heat power are  
 130 determined for the evaluation and development purpose. The influence of  
 131 injection/production velocities, initial reservoir pressure and also the injection  
 132 temperature on the production temperature, reservoir flow impedance, and heat power  
 133 are examined for both SCCO<sub>2</sub> and water and also comparison between them also  
 134 presented.

## 135 2. Mathematical Equations

### 136 2.1 Governing equations

137 The mathematical equation which will govern the transfer of heat in the rock matrix is  
 138 presented in Eq. (1).

$$139 \quad (\rho C_p)_{eff} \frac{\partial T}{\partial t} + \rho_{fl} C_{pfl} u_{mat} \cdot \nabla T - \nabla \cdot (\lambda_{eff} \nabla T) = Q_{mT} + Q_{fT} \quad \dots (1)$$

$$140 \quad (\rho C_p)_{eff} = \phi_{mat} \rho_{fl} C_{pr} + (1 - \phi_{mat}) \rho_{mat} C_{pfl} \quad \dots (2)$$

$$141 \quad \lambda_{eff} = \phi_{mat} \lambda_{mat} + (1 - \phi_{mat}) \lambda_{fl} \quad \dots (3)$$

$$142 \quad u_{mat} = -\frac{\kappa_{mat}}{\mu_{fl}} \nabla p_{mat} \quad \dots (4)$$

143 The mathematical equation which will govern the transfer of heat in the fracture is given  
 144 in Eq. (5)

$$145 \quad a_{frc} \left[ (\rho C_p)_{eff} \right]_{frc} \frac{\partial T}{\partial t} + a_{frc} (\rho_{fl} C_{pw})_{frc} u_{frc} \cdot \nabla T - \nabla \cdot a_{frc} (\lambda_{eff})_{frc} \nabla T = a_{frc} (Q_{fT} + Q_{mT})$$

146 \quad \dots (5)

147 The fluid flow velocity in fracture ( $u_{frc}$ ) is given in Eq. (6)

$$148 \quad u_{frc} = -\frac{\kappa_{frc}}{\mu_{fl}} \nabla p_{frc} = \frac{-a_{frc}^2}{12\mu_{fl}} \nabla p_{frc} \quad \dots (6)$$

149 The injected cold fluid flow in the porous media and fracture are administrated by the  
 150 Darcy's law, mass conservation law, compressibility equation and the force equilibrium  
 151 equations in the porous media.

152 The mathematical equation which will govern the geofluid flow in the rock matrix is given  
 153 in Eq. (7) [39].

$$154 \quad \frac{\partial(\phi_{mat}\rho_{fl})}{\partial t} - \nabla \cdot \left[ \rho_{fl} \frac{\kappa_{mat}}{\mu_{fl}} \nabla p_{mat} \right] - q_{mat} = 0 \quad \dots (7)$$

155 The mathematical equation for the poroelastic storage model is given in Eq. (8).

$$156 \quad \frac{\partial(\rho_{fl}\phi_{mat})}{\partial t} = \frac{\rho_{fl}}{M} \frac{\partial p_{mat}}{\partial t} \quad \dots (8)$$

157 The Biot's modulus and Biot-Willis coefficient ( $\alpha_b$ ) are given in Eq. (9) and Eq. (10)[40,41].

$$158 \quad \frac{1}{M} = \frac{\phi_{mat}}{K_{fl}} + (\alpha_b - \phi_{mat}) \frac{1 - \phi_{mat}}{K_d} \quad \dots (9)$$

$$159 \quad \alpha_b = 1 - \frac{K_d}{K_{fl}} \quad \dots (10)$$

160 The mathematical equation which will govern both the geofluid flow and geomechanical  
 161 effects in the porous media is presented in Eq. (10) (i.e., after merging Eq. (7) to Eq. (10)).

$$162 \quad \left( \frac{\phi_{mat}}{K_{fl}} + \left( 1 - \frac{K_d}{K_{fl}} - \phi_{mat} \right) \frac{1 - \phi_{mat}}{K_d} \right) \frac{\partial p_{mat}}{\partial t} - \nabla \cdot \left( \frac{\kappa_{mat}}{\mu_{fl}} \nabla p_{mat} \right) + \left( 1 - \frac{K_d}{K_{fl}} \right) \frac{\partial \varepsilon_{vol}}{\partial t} = 0 \dots$$

163 (11)

164 The volumetric strain ( $\varepsilon_{vol}$ ) is given in the form of displacement vectors and presented in  
 165 Eq. (12).

$$166 \quad \varepsilon_v = \varepsilon_{11} + \varepsilon_{22} \quad : \left( \varepsilon_{ij} = 0.5 \left( \partial u_{di} / \partial x_j + \partial u_{dj} / \partial x_i \right) \right) \quad \dots (12)$$

167 The mathematical equation for the flow of geofluid in the fracture is given in Eq. (13)

$$168 \quad a_{frc} \frac{\partial}{\partial t} (\phi_{frc} \rho_{fl}) + \nabla_{Tn} \cdot (\rho_{fl} q_{frc}) = a_{frc} q_{mat} \quad \dots (13)$$

169 The mathematical equation for the flow rate of geofluid ( $q_{frc}$ ) per unit length in the  
 170 fracture is presented in Eq. (14)

$$171 \quad q_{frc} = -\frac{\kappa_{frc}}{\mu_{fl}} a_{frc} \nabla_{Tn} P_{frc} = \frac{-a_{frc}^3}{12\mu_{fl}} \nabla_{Tn} P_{frc} \quad \dots (14)$$

172 The governing equation for the force balance for the poroelastic rock matrix is presented  
 173 in Eq. (15) [3,21,42–44]

$$174 \quad \nabla \cdot \sigma + (\rho_{fl} \phi_{mat} + \rho_{mat}) = 0 \quad \text{and} \quad \sigma = \sigma_{mat} - \alpha_b P_{mat} \quad \dots (15)$$

175 The hydraulic fracture is taken as thin elastic layer. Force per unit area acting on the  
 176 fracture is represented mathematically as a function of spring constant ( $k_A$ ), damping  
 177 constant per unit area ( $d_A$ ) and fracture thickness (or fracture aperture) ( $a_{frc}$ ) is given in  
 178 the following eq. (16).

$$179 \quad F_A = -k_A (u_u - u_d - u_0) - d_A \frac{\partial (u_u - u_d - u_0)}{\partial t} - \frac{1}{2} \rho_{fl} a_{frc} \frac{\partial^2 (u_u + u_d)}{\partial t^2} \quad \dots (16)$$

180 Spring constant for unit area is given in eq. (17)

$$181 \quad k_A = k_n n \otimes n + k_s (I - n \otimes n) \quad \dots (17)$$

182 The stiffness in the normal direction, and shear stiffness are defined as a function of both  
 183 elastic modulus and Poisson's ratios of fractures are given in eq. (18) and eq. (19),  
 184 respectively.

$$185 \quad k_n = \frac{E_{frc} (1 - \nu_{frc})}{a_{frc} (1 + \nu_{frc}) (1 - 2\nu_{frc})} \quad \dots (18)$$

$$186 \quad k_s = \frac{E_{frc}}{2a_{frc} (1 + \nu_{frc})} \quad \dots (19)$$

## 187 2.2 Coupling mathematical relations

188 Variation of porosity in rock matrix is defined as a changes occur geomechanical and  
 189 thermal strains and is presented in Eq. (20) [43,45].

$$190 \quad \phi_{mat} = \frac{\phi_i + \Delta \varepsilon_v - \Delta \varepsilon_T}{1 + \Delta \varepsilon_v} = \frac{\phi_i + \Delta \varepsilon_v - [\alpha_T (1 - \phi_i) \Delta T]}{1 + \Delta \varepsilon_v} \quad \dots (20)$$

191 The elastic modulus of reservoir is considered as a function of porosity variation and  
 192 presented in Eq. (21) and it is developed by the Liu [46].

$$193 \quad \ln \left( \frac{E}{E_i} \right) = -d (\phi_{mat} - \phi_i) \quad \dots (21)$$



194 The rock matrix and fracture permeabilities are given Eq. (22) and Eq. (23), respectively  
 195 [47–49].

$$196 \quad \ln\left(\frac{\kappa}{\kappa_o}\right) = \left[\frac{(1-\phi_i)a + b\phi_i}{\phi_i}\right] \varepsilon_v = C_n \varepsilon_v \quad \dots (22)$$

$$197 \quad \kappa_{frN} = \kappa_{fr0} \exp\left(-\frac{\sigma_n}{\sigma^*}\right) \quad \dots (23)$$

198 The rock heat capacity of rock, thermal conductivity of rock are presented in Eq. (24) and  
 199 Eq.(25) [43,50] and

$$200 \quad C_{pr}(\kappa) = \begin{cases} \frac{(2.6 \log(\kappa) + 4.2) \times 10^3}{2.7 \log(\kappa) + 0.3}; & \text{if } -20 \leq \log(\kappa) \leq -11 \\ -13.0 \log(\kappa) + 699.0; & \text{if } -11 \leq \log(\kappa) \leq -2 \end{cases} \quad \dots (24)$$

$$201 \quad \lambda_r(T) = 2.6 - 0.0025(T - 293.15) \quad \dots (25)$$

202 Temperature dependent viscosity, density, heat capacity and thermal conductivity were  
 203 used in the present work. When water is used as geofluid Eq. (26) to Eq (29) [47,51] are  
 204 utilized.

$$205 \quad \mu_w(T) = \begin{cases} \left( \begin{array}{l} 1.38 - 0.028T + 1.36 \times 10^{-4}T^2 \\ -4.61 \times 10^{-7}T^3 + 8.9 \times 10^{-10}T^4 \\ -9.08 \times 10^{-13}T^5 + 3.84 \times 10^{-16}T^6 \end{array} \right); & \text{if } 0 \leq T \leq 140 \\ \left( \begin{array}{l} 0.004 - 2.11 \times 10^{-5}T + 3.86 \times 10^{-8}T^2 \\ +2.4 \times 10^{-11}T^3 \end{array} \right); & \text{if } 140 \leq T < 280 \end{cases} \quad \dots (26)$$

$$206 \quad \rho_w(T) = 838.47 + 1.40T - 0.003T^2 + 3.72 \times 10^{-7}T^3 \quad \dots (27)$$

$$207 \quad C_{pw}(T) = 12010.15 - 80.41T + 0.31T^2 - 5.38 \times 10^{-4}T^3 + 3.62 \times 10^{-7}T^4 \quad \dots (28)$$

$$208 \quad \lambda_w(T) = -0.869 + 0.009T - 1.58 \times 10^{-5}T^2 + 7.98 \times 10^{-9}T^3 \quad \dots (29)$$

209 The variation of thermophysical properties of super critical CO<sub>2</sub> such as viscosity, density,  
 210 heat capacity and thermal conductivity are given in Eq. (30) to (33). These properties are  
 211 dependent on the temperature and pressure. These equations will be applicable for a  
 212 temperature range between 273 K to 553 K and pressure range between 15 MPa to 40  
 213 MPa [35].

$$214 \quad \mu_{\text{scCO}_2}(T, p) = \left[ \begin{array}{l} 7.14 \times 10^{-9} T^2 - 5.642 \times 10^{-6} T - 5.71 \times 10^{-9} p^2 \\ + 2.186 \times 10^{-6} p + 0.0011 \end{array} \right] \quad \dots (30)$$

$$215 \quad \rho_{\text{scCO}_2}(T, p) = \left[ \begin{array}{l} 0.00036 T^3 - 0.3693 T^2 + 122 T - 0.333 p^2 \\ + 32.54 p - 12720 \end{array} \right] \quad \dots (31)$$

$$216 \quad C_{\text{scCO}_2}(T, p) = \left[ \begin{array}{l} -4.9 \times 10^{-5} T^3 + 0.084 T^2 - 49.11 T + 0.47 p^3 \\ - 42.1 p^2 + 1200 p + 276.3 \end{array} \right] \quad \dots (32)$$

$$217 \quad \lambda_{\text{scCO}_2}(T, p) = \left[ \begin{array}{l} -1.75 \times 10^{-8} T^3 + 2.29 \times 10^{-5} T^2 - 0.01 T - 1.89 \times 10^{-5} p^3 \\ + 0.0007 p^2 - 0.006 p + 1.46 \end{array} \right] \quad \dots (33)$$

### 218 3. Model Implementation in COMSOL

#### 219 3.1 Computational Model

220 In this work, we created a 2D computational domain with fracture of length 200 m as  
 221 portrayed in Fig 1. Production and Injection wells are connected with the hydraulic  
 222 fracture and is depicted in Fig 1a. it is considered as a main flow channel for the geofluid  
 223 which is injected from the injection well and fluid is extracted at the production well. The  
 224 computational porous domain is having a permeability of  $9.87 \times 10^{-16} \text{ m}^2$  with initial  
 225 porosity of 0.04. The size of the reservoir domain is considered as 500 m by 500 m and it  
 226 is sufficient to prevent the consequences from the boundaries while extraction process.  
 227 Quadrilateral meshing technique was employed in the present work with a size 0.1 m and  
 228 having a total number of 25000000 domain elements and 210000 boundary elements in  
 229 the meshed domain (Fig 1b). The properties of rock, fracture and geofluid (i.e.,  
 230 water/SCCO<sub>2</sub>) are given in Table 1.

#### 231 3.2 Initial and boundary conditions

232 The initial temperature of the geothermal reservoir is 425 K with and average reservoir  
 233 pressure is 15 MPa. The computational porous domain is suitable to restrict the effects of  
 234 boundaries during the heat recovery from the matrix while in operation. It will keep the  
 235 constant temperatures at the boundaries which is equals to the initial reservoir  
 236 temperature. The initial conditions of fluid flow, temperature and displacements field are  
 237 given in eq. (34), eq. (35), eq. (36) and eq. (37), respectively.

$$238 \quad p(x, y, t) \Big|_{t=0} = p_0 \quad \dots (34)$$

$$239 \quad T(x, y, t)|_{t=0} = T_0 \quad \dots (35)$$

$$240 \quad [u_x, u_y]^T|_{t=0} = [0, 0]^T \quad \dots (36)$$

$$241 \quad \left[ \frac{\partial u_x}{\partial t}, \frac{\partial u_y}{\partial t} \right]^T|_{t=0} = [0, 0]^T \quad \dots (37)$$

242 For the fluid flow, undrained condition was employed on the all boundaries of the  
 243 reservoir domain. In the structural module, as represented in the Fig 1b, two boundaries  
 244 were constrained with rolling boundary condition and two boundaries are constrained by  
 245 the lateral stresses. This study is attentive to the generation stress due to the thermal,  
 246 pore pressure and external stress/loads effects during the heat production. The  
 247 temperature of the geofluid (water or SCCO<sub>2</sub>) was varied from 35 °C to 45 °C with variable  
 248 injection rates between 0.05 m/s to 0.15 m/s. The initial fracture aperture is 0.5 mm with  
 249 a fracture length of 200 m, and Biot-Willie’s coefficient is 0.5 considered in the present  
 250 work. Physical fields such as pressure, temperature, and displacement are examined in  
 251 the present work during the heat extraction process by fully coupled model.

252 The boundary conditions are primarily concerned at injection well, production well and  
 253 at the boundaries of the rock matrix. The fluid flow boundary conditions at the injection  
 254 well and production well are represented in Eq. (38) and Eq. (39), respectively.

$$255 \quad \text{At injection well:} \quad \dot{m}(t)|_{inj} = u_{inj} \rho_{fl} \quad \dots (38)$$

$$256 \quad \text{At production well:} \quad \dot{m}(t)|_{prod} = u_{inj} \rho_{fl} \quad \dots (39)$$

257 Heat flux boundary condition was employed at the injection well and is given in eq. (40).

$$258 \quad \text{At injection well:} \quad q|_{inj} = (C_p)_{fl} (T_{inj} - T_0) u_{inj} \rho_{fl} \quad \dots (40)$$

### 259 3.3 Implementation in COMSOL Multiphysics

260 COMSOL Multiphysics 5.4 (Institute license version from the IIT-Madras) was used in the  
 261 present research work. Plentiful researchers and engineers utilized the COMSOL  
 262 Multiphysics for the fully coupled numerical investigations [25,44,52–57]. COMSOL inbuilt  
 263 modules are used in the present work to investigate the coupled impact of THM  
 264 interactions both in the fracture and rock matrix. The modules used in the present work

265 are heat transfer in porous media, Darcy law, and solid mechanics along with the  
266 poroelasticity and thermal expansion modules. In addition to those modules, dynamics  
267 fluid and rock/fracture properties were embedded as local variable in the component  
268 section. The comparison of the fluid properties of water and SCCO<sub>2</sub> is depicted in the Fig  
269 2. It was found that SCCO<sub>2</sub> has much less values in all the properties compared to the  
270 water in the operating conditions. In the geomechanics module (i.e., solid), hydraulic  
271 fracture is designated as a thin elastic layer's boundary element. The governing equations  
272 for the elastic layers are presented from Eq. 16 to Eq. 19. The fluid and heat flow in the  
273 fracture is employed using the fracture flow submodule was employed. The flow chart for  
274 solving the proposed THM model is presented in Fig 3. To stabilize the model, initially a  
275 stationary solver without boundary conditions was employed (i.e., with initial conditions)  
276 and then time dependent solver for 30 years.

## 277 4. Results and Discussion

### 278 4.1 Verification

279 Porous media with fracture will create complexity in dealing which will increase the  
280 intricacy in computational solving. To solve the developed fully coupled dynamic  
281 mathematical model we used the COMSOL Multiphysics, and it is a finite element tool. In  
282 the present work, hydraulic fracture is considered as the main conduit flow channel.  
283 Dynamic fluid and rock/fracture properties were explicitly employed using Eq. (20) to Eq.  
284 (33). The developed dynamic THM model and its accuracy are validated with the work of  
285 Lauwerie's [58]. Fig 4a represents the geometry with single fracture for the verification of  
286 transfer of heat. The equation for the spatiotemporal (i.e.,  $x$ , and  $t$ ) variation of  
287 temperature is given in Eq. (41)[58].

$$288 \quad T(x,t) = T_0 + (T_{inj} - T_0) \operatorname{erfc} \left[ \frac{(\lambda_m x / \rho_{fl} C_{pfl} a_{frc})}{\sqrt{\frac{\lambda_m}{\rho_m C_{pr}} (u_{fl} t - x)}} \right] U_{step} \left( t - \frac{x}{u_{fl}} \right) \quad \dots (41)$$

289 The rock/fracture and fluid parameters for the validation purpose were given in the Table  
290 2. The model results from the COMSOL Multiphysics were compared the work of  
291 Lauwerie's [58] and presented in Fig 4b and Fig 4c. It is clearly identified that; the error is

292 less than  $\pm 5\%$ . Thus, the proposed dynamic THM model can estimate the temperature of  
293 geofluid in the hydraulic fracture. Therefore, we can use the developed dynamic THM  
294 model to examine the heat production from the geothermal reservoir.

#### 295 **4.2 Evolution of temperature**

296 The spatiotemporal temperature variations in the rock matrix and fracture was examined  
297 in the present work with  $\text{SCCO}_2$  and water as geofluids and presented in Fig 5 and Fig 6.  
298 The comparison between the  $\text{SCCO}_2$  and water as geofluids also studied in the present  
299 work (i.e., Fig 6). From Fig 5, it has been found that temperature of the  $\text{SCCO}_2$  in the  
300 fracture is gradually rising (i.e., lower to high) while moving from injection well to  
301 production well. It is due to the exchange of heat from the rock matrix to the geofluid-  
302  $\text{SCCO}_2$  and it will attain to maximum before reaching to production well. It will create a  
303 temperature difference in the fracture from injection to production well (Fig 5). It was  
304 also found that the injection/production velocity is playing an important role in the  
305 expansion of low-temperature region in the vicinity of hydraulic fracture. It has been  
306 found that the expansion of low-temperature region is fast and high in the high  
307 injection/production velocity (i.e., 0.15 m/s) scenario compared to low  
308 injection/production velocity (i.e., 0.05 m/s). Similarly, the temperature in the fracture is  
309 influenced by the injection/production velocity. The comparison of water and  $\text{SCCO}_2$  as  
310 geofluids is presented in the Fig 6. At same injection/production velocity, the low  
311 temperature region is spreading comparatively faster while using water compared to  
312  $\text{SCCO}_2$ .

313 The impact of injection velocity, initial reservoir temperature, initial reservoir  
314 temperature on the production temperature is presented in Fig 7. It has been found that,  
315 the temporal decrement in the production temperature with time and injection velocity  
316 (i.e., Fig 7a). This is because of the spreading of low temperature region in the rock matrix  
317 which is near the hydraulic fracture and it is faster in the high velocity scenario compared  
318 to low velocity. It will reduce heat extraction from the rock matrix compared to the early  
319 stages of injection-production operation, furthermore it will reduce the production  
320 temperature. Alike  $\text{SCCO}_2$ , production temperature was decreased temporally and  
321 increasing velocity when water is used as geofluid (Fig 7d). Initial reservoir pressure is also

322 influencing the production temperature and extracting maximum production  
323 temperature at initial reservoir of 15 MPa and minimum at initial reservoir of 20 MPa (Fig  
324 7b). It is due to the heat capacity, and thermal conductivity of  $\text{SCCO}_2$  is a function of both  
325 temperature and pressure. Similar nature was not found when water using as a geofluid  
326 (Fig 7e). It is due to the properties of water are not a function of pressure and only  
327 dependent on the temperature. It was also found that, the injection temperature is  
328 influencing the production temperature. For higher injection temperature, production  
329 temperature is high compared to the low-injection temperature of the geofluid (Fig 7c).  
330 When water is used as a geofluid, injecting at low temperature ( $35^\circ\text{C}$ ) is more influencing  
331 compared to the higher temperatures ( $40^\circ\text{C}$  and  $45^\circ\text{C}$ ).

332 The comparison of both water and  $\text{SCCO}_2$  as geo fluid on the production temperature was  
333 depicted in Fig 7g to Fig 7i. It was found that, steep reduction of production temperature  
334 was found when using the water compared to  $\text{SCCO}_2$  with increase in  
335 injection/production velocity (Fig 7g). The initial reservoir pressure is not showing  
336 negligible impact on production temperature when using water, but it is much lesser  
337 compared to  $\text{SCCO}_2$  as geofluid at constant operating conditions (Fig 7h). Similarly,  
338 injection temperature is influencing the production temperature when using water and  
339  $\text{SCCO}_2$  as geofluids, but high production temperatures were recorded when using  $\text{SCCO}_2$ .  
340 Fig 8 depicts the production temperature at different operating conditions when using  
341  $\text{SCCO}_2$  as geofluid. It was found that, minimum production temperature was recorded  
342 when injecting/producing  $\text{SCCO}_2$  at 0.15 m/s at initial reservoir pressure of 20 MPa. Thus,  
343  $\text{SCCO}_2$  is providing the better results compared to water in the prospect of production  
344 temperature from the reservoir and also the spreading of low-temperature zone at similar  
345 operating conditions.

### 346 **4.3 Evolution of stress, and strain**

347 The variation of geomechanical properties such as stress, strain in the rock matrix and  
348 fracture are studied and presented in the present work. The Von-mises stress, Tresca  
349 stress, mechanical strain, thermal strains were examined. Von-mises stress is used to  
350 determine the fracture/yield of the rock under the load which is equal or greater than the  
351 yield strength of the rock. Fig 9 represents the spatiotemporal evolution of Von-mises

352 stress (in MPa) when using SCCO<sub>2</sub> and water at injection/production velocities of 0.05 m/s  
353 and 0.1 m/s. It was found that, von-mises stress is expanding from the fracture to the rock  
354 matrix. The stress is expanding its intensity from the injection well to fracture, then to  
355 rock matrix. The distribution is regular when using SCCO<sub>2</sub> at lower injection/production  
356 velocity, and the low stress region is increasing in the vicinity of injection well (Fig 9a<sub>1</sub> to  
357 9a<sub>5</sub>). When using the water as geofluid, the distribution of von-mises stress is irregular  
358 after the five years of injection/production operation (Fig 9b<sub>1</sub> to 9b<sub>5</sub>). The values in both  
359 the scenarios are less than the yield stress/elastic modulus of the rock. In both the cases  
360 the rock will not fail under the load due to the yield stress will not exceeds the yield  
361 strength of the rock. It was found that injection/production velocity is influencing the  
362 distribution of von-stress in the rock matrix along with the type of fluid. At higher the  
363 velocities (Fig 9), the distribution is irregular with time increases. At high velocities, the  
364 fluid may try to escape form the fracture to enter into the rock matrix. It will lead to  
365 restructuring of rock in the vicinity of fracture, furthermore it will create abnormalities in  
366 the rest of the rock. These abnormalities create the increase in the von-mises stress in the  
367 rock and lowering at the fracture (specifically near injection well). Thus, the type of fluid  
368 and the injection/production velocities are influencing the generation of von-mises stress  
369 (also distribution) in the rock.

370 The tresca stress defined as the failure of the rock occurs at a critical value of the  
371 maximum shear stress (i.e., =0.5yield strength of the rock). Fig 10 illustrates the tresca  
372 stress (in MPa) in the rock matrix and fracture with SCCO<sub>2</sub> and water. It was found that  
373 the distribution of tresca stress is different than the von-mises stress distributions. At the  
374 early stages the maximum tresca stress will be generates at the injection and production  
375 wells (i.e., Fig 10a<sub>1</sub>, 10b<sub>1</sub>, 10c<sub>1</sub>, and 10d<sub>1</sub>). With time progression the maximum tresca  
376 stress retain at the production well and lowest will be found in the vicinity of injection  
377 well. Specifically, minimum tresca stress will be found at the low-temperature zone  
378 compared to the rest of the reservoir. It was also found that the maximum tresca stress  
379 is found just away from the low temperature zone. The maximum value of the tresca  
380 stress is less than the maximum shear stress, thus the rock will not fail due to the shear

381 stress generated during the cold fluid injection and heat extraction process. The tresca  
382 stress is dependent on the type of fluid and the injection/production velocity and it was  
383 clearly observed from the numerical results (i.e., Fig 10).

384 Thermal strain, mechanical strain and combined strain are computed and illustrated in  
385 Fig 11. It was found that mechanical strain is dominated compared to the thermal strain.  
386 It was observed that negative thermal strain was playing vital role in the vicinity of  
387 production and positive thermal strain was governed in the vicinity of injection well. It  
388 was also found that thermal strain is dominated in the low-temperature zone and  
389 minimum mechanical strain is found in the same zone. The thermal and mechanical strain  
390 dominated region is also dependent on the type of fluid is using. More water is entering  
391 into the rock matrix from the injection well and fracture compared to the  $\text{SCCO}_2$ , and it  
392 was clearly identified in the form of stress-strain variation in Fig 10 and Fig 11. It was also  
393 found that the injection/production velocity is influencing the thermal and mechanical  
394 strains. Higher the injection/production velocity, the spreading of thermal strain and  
395 mechanical strain are higher specifically in the low-temperature zone. Thus, thermal,  
396 mechanical, and combined stains are highly influenced by the type of fluid, and  
397 injection/production velocity.

#### 398 **4.4 Evolution of rock and fluid properties**

399 The dynamic of fluid, rock, and fracture properties are integrated with the proposed THM  
400 model in the present work. Fig 12 depicts the variation of porosity, permeability, young's  
401 modulus, and effective thermal conductivity in the fracture and also in the matrix. It has  
402 been observed that the porosity is changed in the rock matrix, specifically in the  
403 neighborhood of fracture (Fig 12a<sub>1</sub> to 12a<sub>3</sub>). This variation is due to the expansion of grain.  
404 It will decrease the pore space in the neighborhood of injection well and fracture (Fig 12a<sub>1</sub>  
405 to 12a<sub>3</sub>) due to the expansion of low-temperature zone. The impact of type of injection  
406 fluid is also found in the variation of porosity. It was found that when using the water, the  
407 porosity disturbed zone is higher compared to  $\text{SCCO}_2$ (Fig 12b<sub>1</sub> to 12b<sub>3</sub>) Similarly, the  
408 variation of porosity will be influencing the young's modulus of the rock matrix and  
409 fracture and it is presented in the Fig 12c<sub>1</sub> to 12c<sub>2</sub> when using the  $\text{SCCO}_2$ . The variation in  
410 the young's modulus is found very negligible compared to the initial value for both water



411 and  $\text{SCCO}_2$  (Fig 12c<sub>1</sub> to 12c<sub>2</sub> and Fig 12d<sub>1</sub> to 12d<sub>3</sub>). It was found the permeability of rock  
412 matrix is reduced in the neighborhood of injection well with time (Fig 12e<sub>1</sub> to 12e<sub>3</sub>). The  
413 reduction in permeability and porosity in the low temperature invaded zone may be the  
414 cause for the pore pressure variation in the vicinity of fracture. It will be increasing with  
415 the injection/production velocities. Higher reduction of permeability zone is found while  
416 using water as geofluid (Fig 12e<sub>1</sub> to 12e<sub>3</sub> and Fig 12f<sub>1</sub> to 12f<sub>3</sub>). The maximum of effective  
417 thermal conductivity ( $\lambda_{eff}$ ) in rock matrix was found in the vicinity of fracture and  
418 decrease spatially away from the injection well (Fig 12g<sub>1</sub> to 12g<sub>3</sub> and Fig 12h<sub>1</sub> to 12h<sub>3</sub>).  
419 Maximum value of  $\lambda_{eff}$  in fracture was observed near the injection well and it will be  
420 reduced towards the production well. These higher values are due to the low temperature  
421 near the injection well. Effective thermal conductivity will influence the heat extraction  
422 capacity from rock matrix via cold fluid injection. Thus, the physical, mechanical, and  
423 thermal properties of rock/fracture were executed efficiently and observed significant  
424 variations while injecting cold fluid.

#### 425 **4.5 Performance indicators of geothermal reservoir**

426 The performance of the geothermal reservoir is studied exclusively using indicators such  
427 as thermal breakthrough time, geothermal life, reservoir impedance and the generated  
428 heat power. Thermal breakthrough is defined the time of production temperature decline  
429 was identified. In the present work we used the equation given by Rijn [58] and it is  
430 presented in eq. (42).

$$431 \quad T_{\beta T} = 0.99T_{ini} \quad \text{..... (42)}$$

432 In eq. (42),  $T_{\beta T}$  is the thermal breakthrough,  $T_{ini}$  is the initial temperature. Thermal  
433 breakthrough time is defined as the time required for the fluid to reach thermal  
434 breakthrough at the production well. Geothermal life ( $\zeta_l$ ) described as the time period  
435 from the starting of the heat production to the production temperature of the reaches to  
436 60% of the original reservoir temperature. Reservoir flow impedance is defined as the  
437 ratio of pressure difference between injection and production wells to the production  
438 flow rate [35]. Heat power represents the average extraction of heat from the geothermal

439 reservoir while producing heat. The mathematical equations for the reservoir flow  
 440 impedance and heat power are represented in eq. 43 and eq. 44, respectively [35,47,59].

$$441 \quad I_{RF} = \frac{p_{inj} - p_{prod}}{q_h(t)_{prop}} \quad \dots (43)$$

$$442 \quad \bar{W} = \frac{W_h}{\zeta_l} = \frac{\sum_{i=1}^{\zeta_l} [q_h(t)_{prop} n_{prop} h_{prod} - q_h(t)_{inj} n_{inj} h_{inj}]_i}{\zeta_l} \quad \dots (44)$$

443 In eq (43) and (44),  $p_{inj}$  is the injection pressure,  $p_{prod}$  is the production pressure,  $q_h(t)_{prop}$   
 444 is the mass flow rate at the production well,  $q_h(t)_{inj}$  is the mass flow rate at the injection  
 445 well,  $n_{prop}$  is the number of production wells,  $n_{inj}$  is the number of injection wells,  $h_{prod}$   
 446 is the enthalpy at the production well,  $h_{inj}$  is the enthalpy at the injection well,  $\zeta_l$  is the  
 447 geothermal life. The thermal breakthrough and the geothermal life are illustrated in the  
 448 Fig 7 and Fig 8. It was found the thermal breakthrough is achieved much faster when using  
 449 water as geofluid. It was also found that the geothermal life is less when using water as  
 450 geofluid. It was due to the intervention of water into the rock matrix from the injection  
 451 well/hydraulic fracture. It will create the low-temperature zone in the vicinity of fracture  
 452 and the injection well. It will reduce the heat extraction capacity of injected fluid. Thus,  
 453 the production temperature decreases rapidly compared to SCCO<sub>2</sub>-EGS. Similarly, SCCO<sub>2</sub>-  
 454 EGS system is having the better geothermal life compared to the water-EGS system.

455 Fig 13 depicts the impact of injection velocity, initial reservoir pressure and injection  
 456 temperature on the reservoir flow impedance (RFI). It was found that the RFI is increasing  
 457 with rise in injection/production velocity in both SCCO<sub>2</sub>-EGS (Fig 13a) and water-EGS (Fig  
 458 13d). It was found that the RFI is independent at the higher initial reservoir pressure (i.e.,  
 459 20 MPa and 25 MPa) and dependent at lower reservoir pressure (i.e., 15 MPa). Higher the  
 460 injection temperature and lower the RFI in SCCO<sub>2</sub>-EGS and followed ascending nature. For  
 461 water-EGS system, RFI is lower when injection temperature of 35°C and higher when  
 462 injection temperature of 40°C which is higher than the injection temperature of 45°C,  
 463 and 35°C. The comparison of RFI for the water-EGS compared to the SCCO<sub>2</sub>-EGS also  
 464 studied found higher in water-EGS system for all three scenarios (Fig 13g, 13h, 13i). Fig 14

465 depicts the RFI of the SCCO<sub>2</sub>-EGS system for the injection/production velocity, initial  
466 reservoir pressure, and injection temperature. It was found that, higher RFI is found when  
467 using the injection temperature of 35°C and lowest was found in the case of 45°C.  
468 But all the RFI values (both water-EGS and SCCO<sub>2</sub>-EGS) are showing within the approved  
469 limit of 0.2 MPa/(kg/s) which is given by Evans [60]. Thus, the SCCO<sub>2</sub>-EGS system is  
470 showing the better performance compared to the water-EGS in the prospective of RFI.  
471 Fig 15 illustrates the impact of injection velocity, initial reservoir pressure and injection  
472 temperature on the heat power generated during the operation. It was found that these  
473 three are having significant impact on heat power for SCCO<sub>2</sub>-EGS. Higher heat power was  
474 recorded for higher velocities and lower was found in low velocity (i.e., 0.05 m/s) and it  
475 was following the ascending order (Fig 15a). Lowest heat power was recorded when for a  
476 reservoir having 15 MPa initially and higher was found for 20 MPa condition (Fig 15b). It  
477 was found that when increasing injection temperature, the heat power is reducing in the  
478 SCCO<sub>2</sub>-EGS which follows the descending order (Fig 15c). For water-EGS, same trend was  
479 found like SCCO<sub>2</sub>-EGS with injection velocities but much higher than that of SCCO<sub>2</sub>-EGS  
480 (Fig 15d and Fig 15g). It is due the density, heat capacity, and thermal conductivity of  
481 SCCO<sub>2</sub> are lower than water (Fig 2). But the injection temperature and the initial reservoir  
482 pressures are showing negligible influence on heat power for water-EGS. The comparison  
483 of RFI for the water-EGS compared to the SCCO<sub>2</sub>-EGS also studied found higher in water-  
484 EGS system for all three scenarios (Fig 15g, 15h, 15i). Fig 16 depicts the heat power  
485 recorded for the SCCO<sub>2</sub>-EGS system at different injection/production velocity, initial  
486 reservoir pressure, and injection temperature. It was found that, higher heat power for  
487 SCCO<sub>2</sub>-EGS was recorded when using the injection temperature of 35°C at pressure of  
488 20 MPa (injection velocity=0.15 m/s) and lowest was found in the case of 45°C and  
489 pressure of 15 MPa (i.e., injection velocity=0.05 m/s). Thus, the SCCO<sub>2</sub>-EGS system is  
490 showing the lower performance compared to the water-EGS in the prospective of heat  
491 power. But showing better performance in the thermal breakthrough, geothermal life and  
492 also the RFI.

## 493 5. Conclusions

494 The thermo-hydro-geomechanical mathematical model was enhanced with the dynamic  
495 rock, fracture, and fluid properties in the present work. The variation of properties  
496 includes the porosity of rock/fracture, permeability of rock/fracture, young's modulus of  
497 rock/fracture, heat capacity of rock/fracture and fluid, thermal conductivity of  
498 rock/fracture and fluid, fluid viscosity, and fluid density. These variations are the functions  
499 of pressure, temperature, and stress and strains variations. This fully coupled dynamic  
500 thermo-hydro-geomechanical model was verified using the COMSOL Multiphysics 5.4 and  
501 used for the investigations of geothermal reservoir.

502 The present work is focused on the comparison of SCCO<sub>2</sub>-EGS and water-EGS examined  
503 found the merits and demerits of the SCCO<sub>2</sub>-EGS over water-EGS. Steep reduction of  
504 production temperature was identified for water-EGS system compared to SCCO<sub>2</sub>-EGS  
505 system. Enhanced production temperature was recorded for the SCCO<sub>2</sub>-EGS system at  
506 low inlet/outlet velocity (i.e., 0.05 m/s), and pressure of 20 MPa. The generation of Von-  
507 mises and tresca stresses the rock matrix are influenced by the injection/production  
508 velocities and type of geofluid used. Irregular distribution of von-mises stress in the rock  
509 matrix is found at higher inlet/outlet velocities. Lower tresca stress was recorded in the  
510 vicinity of fracture and maximum was found at the outer boundary of the low  
511 temperature zone. Thermal strain and mechanical strain were examined during the  
512 injection-production operation. Mechanical strain was dominated away from the low-  
513 temperature zone and thermal strain was dominated in the vicinity of hydraulic fracture.  
514 The expansion of thermal strain was highly dependent on the type of fluid (i.e., SCCO<sub>2</sub> or  
515 water) and injection/production velocity. The variation of the rock properties also  
516 presented in this work.

517 The performance indicator such as thermal breakthrough, geothermal life, reservoir flow  
518 impedance and heat power are determined in the present work. The SCCO<sub>2</sub>-EGS system  
519 is showing better thermal breakthrough and geothermal life compared to the water-EGS  
520 system. The RFI is found within the Evans limit for the both systems and SCCO<sub>2</sub>-EGS  
521 showing lowest compared to water-EGS system. The water-EGS system recorded best  
522 heat power compares to SCCO<sub>2</sub>-EGS system. So, these performance indicators are highly

523 dependent on the type of fluid, injection/production velocities, initial reservoir pressure,  
524 injection temperature. Moreover, the advantage of SCCO<sub>2</sub>-EGS system over the water-  
525 EGS system gives the promising results to the geothermal industry as geofluid. It will also  
526 improve the provides the geosequestration which is not at possible in the water-EGS  
527 system.

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529

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**Nomenclature:**

<b>Symbol</b>	<b>Description</b>
$\rho_{fl}$	Density of fluid
$\rho_{mat}$	Density of matrix,
$u_{mat}$	Darcy's velocity in matrix,
$u_{frc}$	Darcy's velocity in fracture.
$T$	Temperature,
$Q_{fT}$ and $Q_{mT}$	Source/sink terms fracture and matrix, respectively,
$\phi_i$	Initial porosity
$\phi_{mat}$	Porosity of the matrix,
$C_{pfl}$	Specific heat capacity of fluid,
$C_{pr}$	Specific heat capacity of matrix,
$\lambda_{eff}$	Effective thermal conductivity,
$\lambda_{mat}$	Thermal conductivity of matrix,
$\lambda_{fl}$	Thermal conductivity of fluid,
$K_{mat}$	Rock permeability,
$K_{frc}$	Permeability in fracture
$K_o$	Initial permeability of the porous media
$K_{frc0}$	Initial permeability of fracture
$\mu_{fl}$	Viscosity of the fluid
$a_{frc}$	Fracture aperture
$P_{mat}$	Pressure in matrix.
$P_{frc}$	Pressure in fracture
$\varepsilon_{vol}$	Volumetric strain
$M$	Biot's modulus.
$\alpha_b$	Biot-Wills coefficient
$K_{fl}$	Fluid bulk modules
$K_d$	Drained bulk modules
$\partial\varepsilon_v/\partial t$	Rate of change in volumetric strain of the porous matrix.
$u_{di}$ and $u_{dj}$	Displacement vectors in 'i' and 'j' directions, respectively
$\nabla_{Tn}$	Gradient is measured on the tangential plane of fracture
$q_{mat}$	Source/sink term which couple both matrix and fracture $[-\rho_{fl}\alpha_b(\partial\varepsilon_{vol}/\partial t)]$
$\sigma_s$	Effective stress tensor
$\sigma$	Deviatoric stress tensor.

$\sigma^*$	Normalizing constant (and it is considered as the initial reservoir pressure)
$\sigma_n$	Normal stress.
$\Delta \varepsilon_T$	Thermal strain
$\alpha_T$	Coefficient of thermal expansion
$\alpha_b$	Biot-Wills coefficient
$E_i$	Initial elastic modulus
$E_{frc}$	Elastic modulus of fracture
$d$	Fitting parameter (constant and equal to 1)
$a$ and $b$	Constants
$C_n$	Coefficient and it is a function of initial porosity of formation ( $C_n = 5/\phi_i$ )
$k_A$	Spring constant
$d_A$	Damping constant per unit area
$u_u$	Displacement in upside of fracture
$u_d$	Displacement in downside of fracture
$u_0$	Initial displacement of fracture
$\nu_{frc}$	Poisson's ratio of fracture
$d$	Fitting parameter (constant and equal to 1)
$U_{step}$	Unit step function
$u_{fl}$	velocity of fluid
$t$	Time
$x$	Distance in x-direction
$erfc$	Error function
$u_{di}$ and $u_{dj}$	Displacement vectors in 'i' and 'j' directions, respectively

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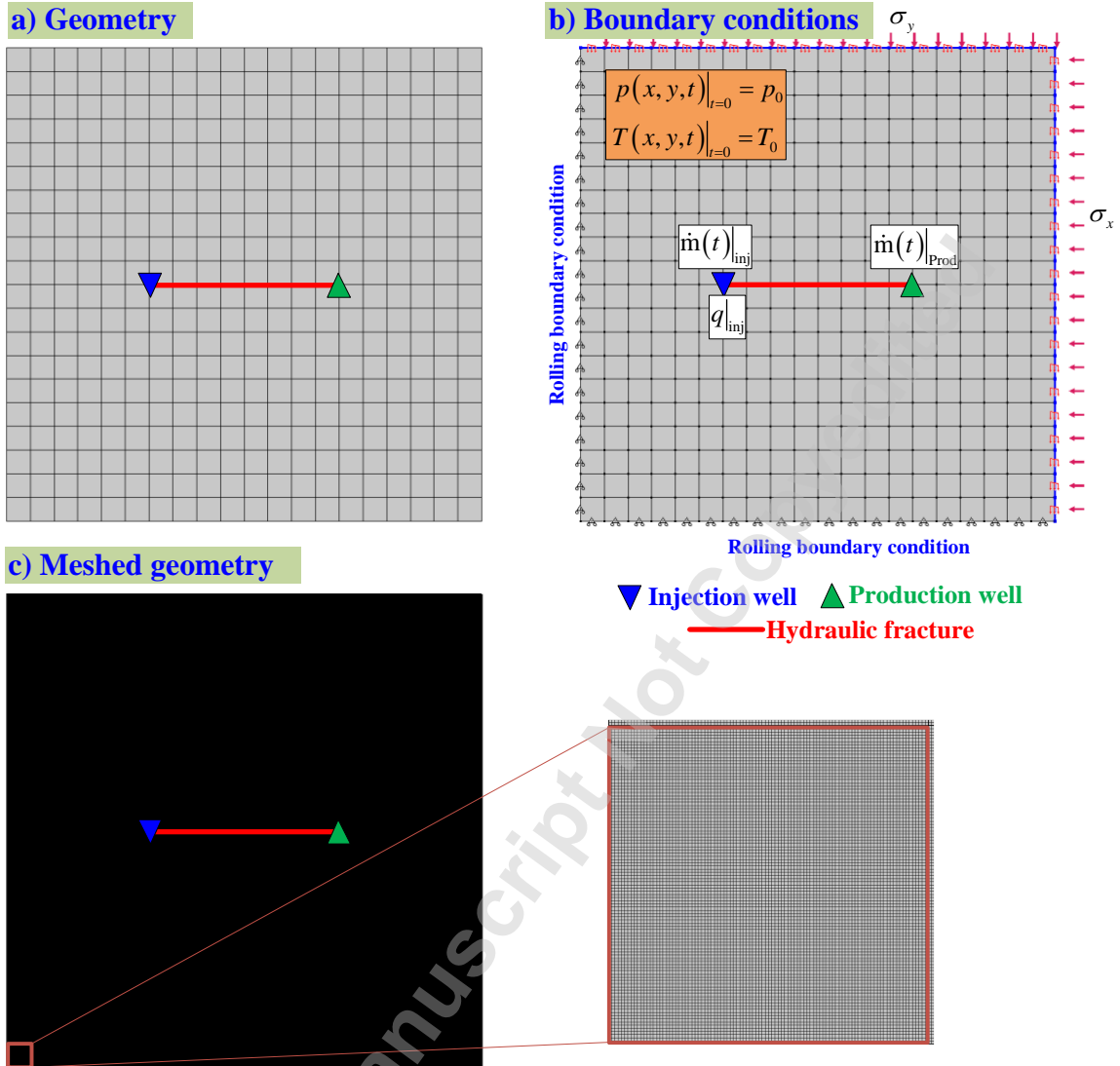
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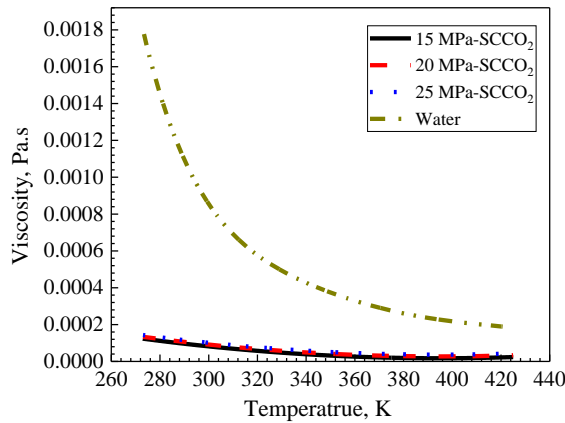
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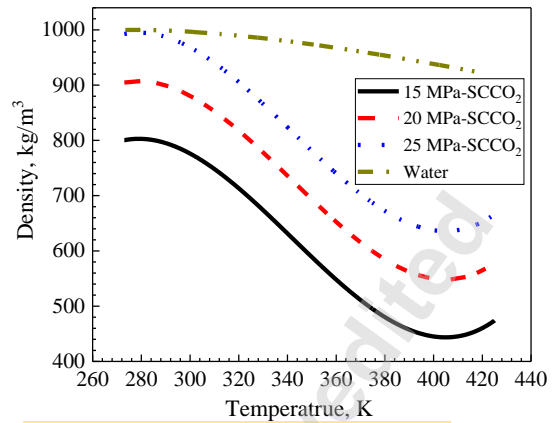
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**Fig 1** Geometry geothermal reservoir with single fracture (a), boundary conditions (b) and the meshed geometry (c).

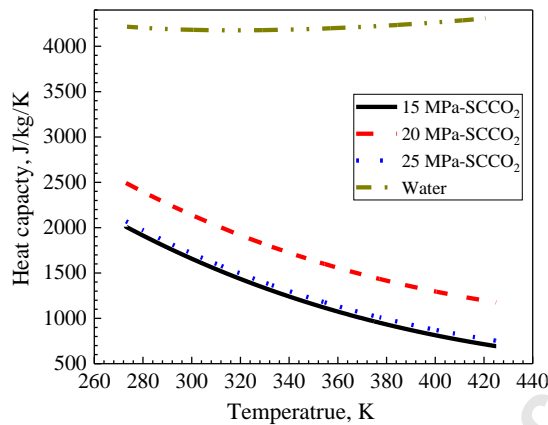
a) Viscosity, Pa.s



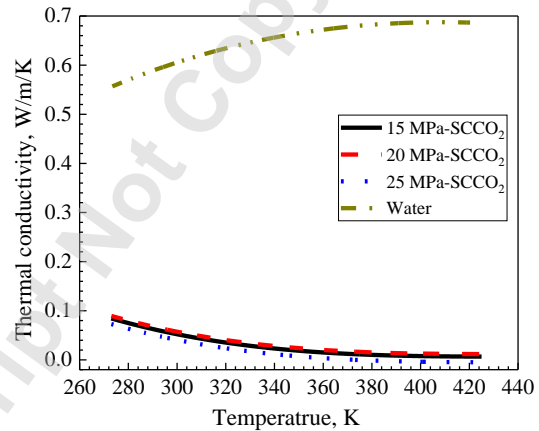
b) Density, kg/m<sup>3</sup>



c) Heat capacity, J/kg/K



d) Thermal conductivity, W/m/K

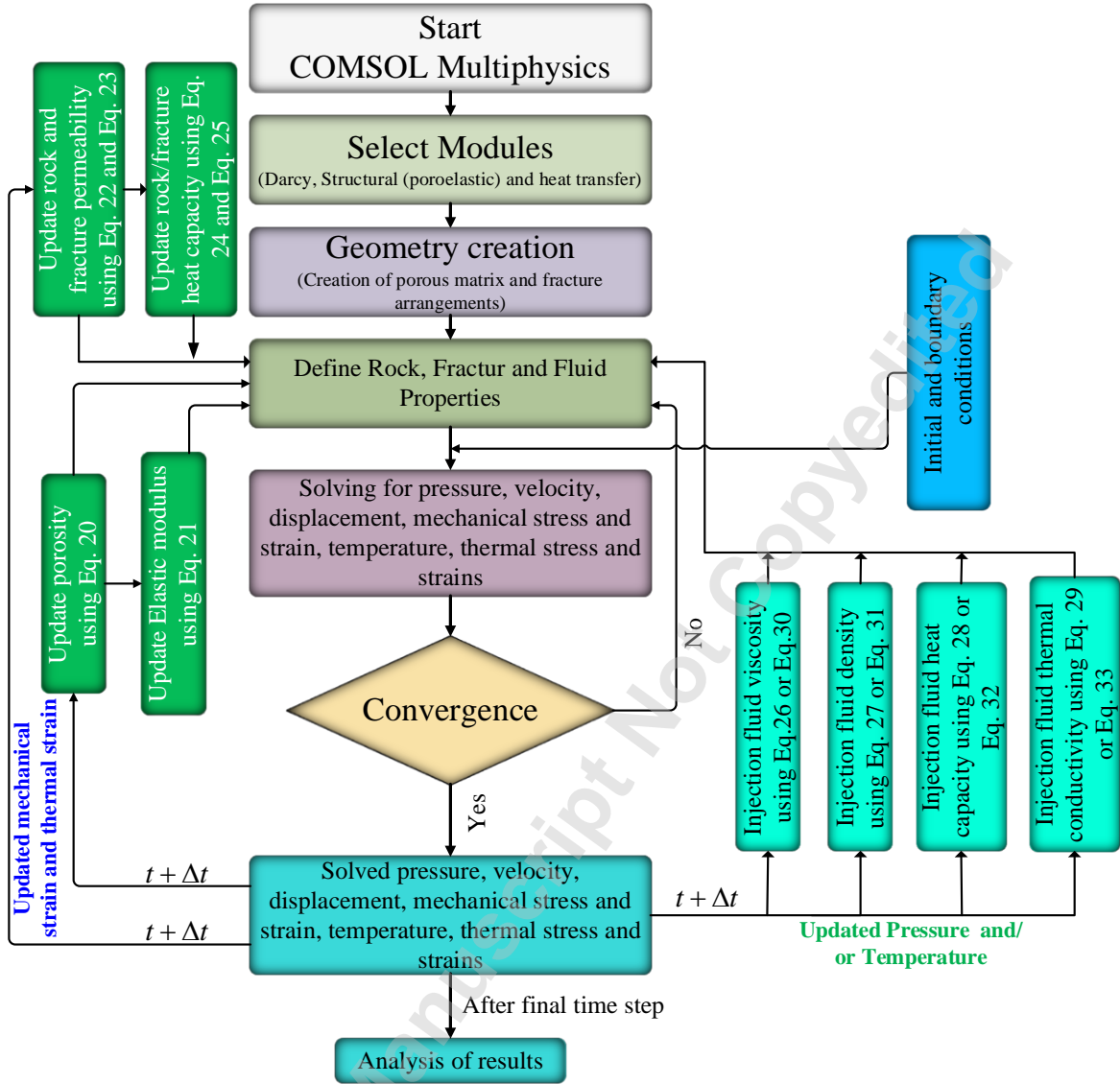


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**Fig 2** Comparison of the properties of SCCO<sub>2</sub> and water within the operating pressure and temperatures.

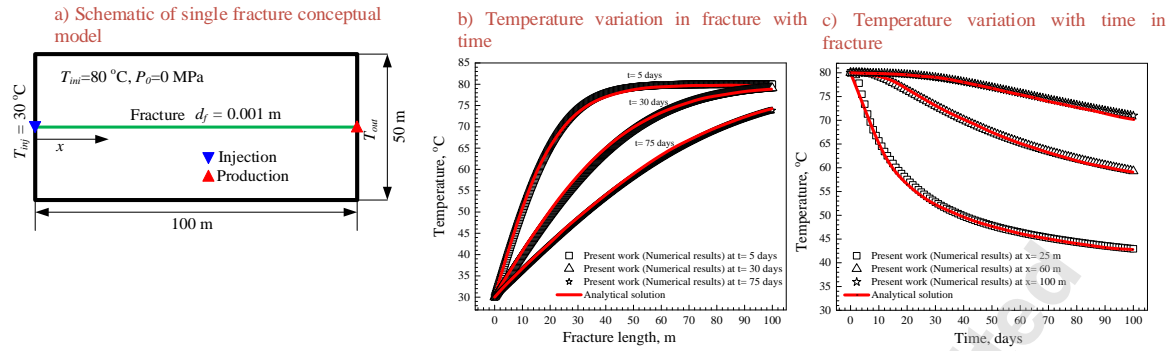
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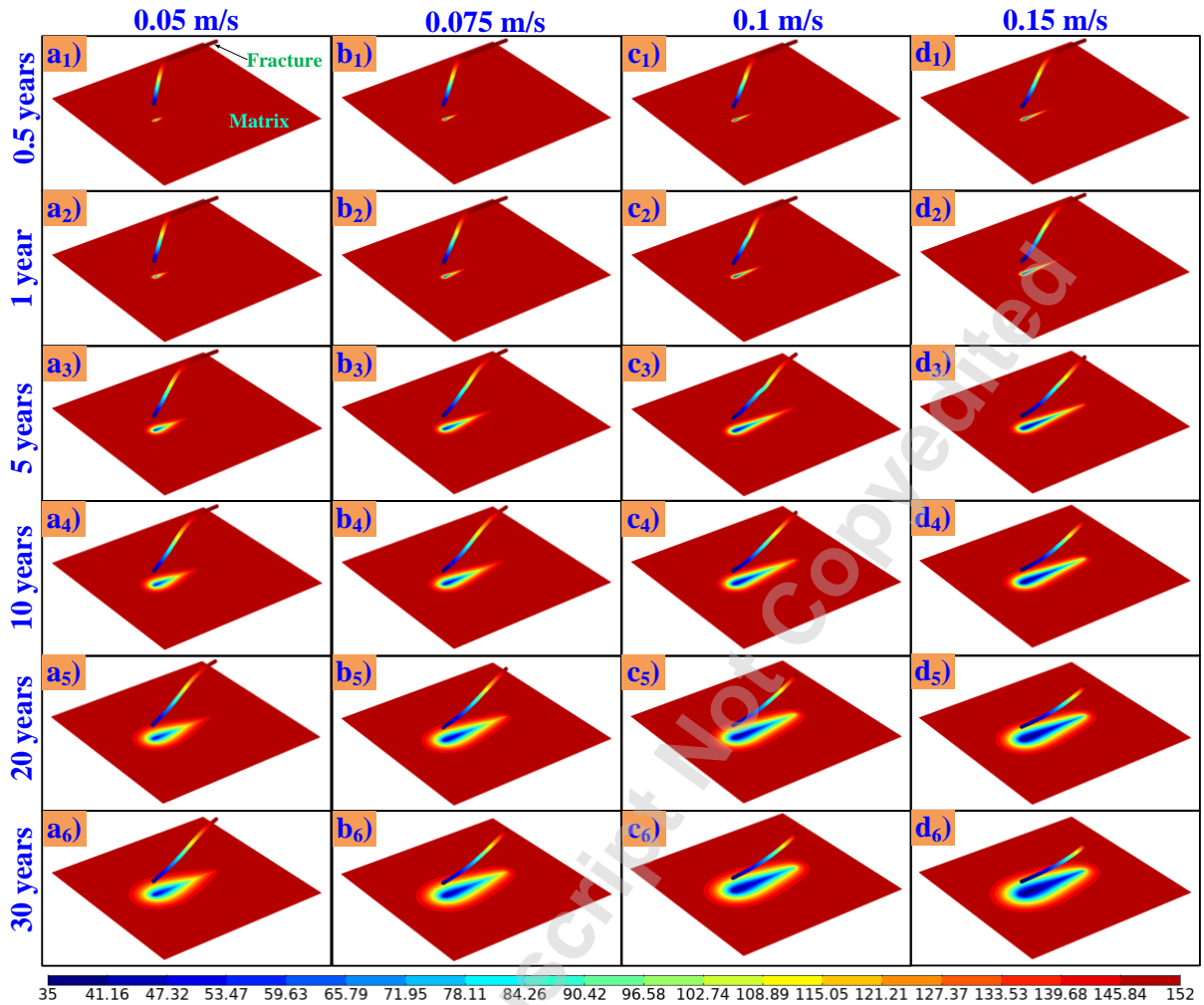
**Fig 3** Schematic of the solution process for fully coupled Thermo-hydro-geomechanical model in geothermal reservoir with fracture.



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6. Fig 4 Verification for the heat transfer in single fracture with analytical solution.

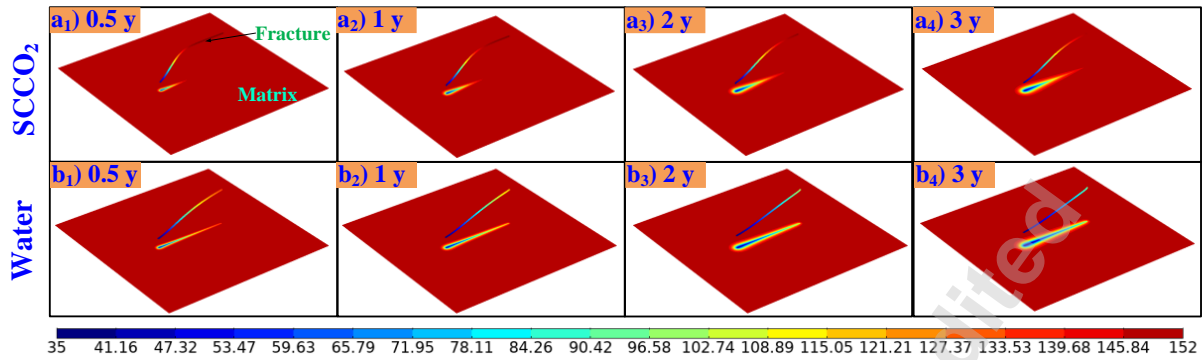
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**Fig 5** Spatiotemporal variation of temperature in the reservoir and fracture with different injection/production velocities when using SCCO<sub>2</sub> as geofluid at initial pressure of 15 MPa, and injection temperature of 35°C.

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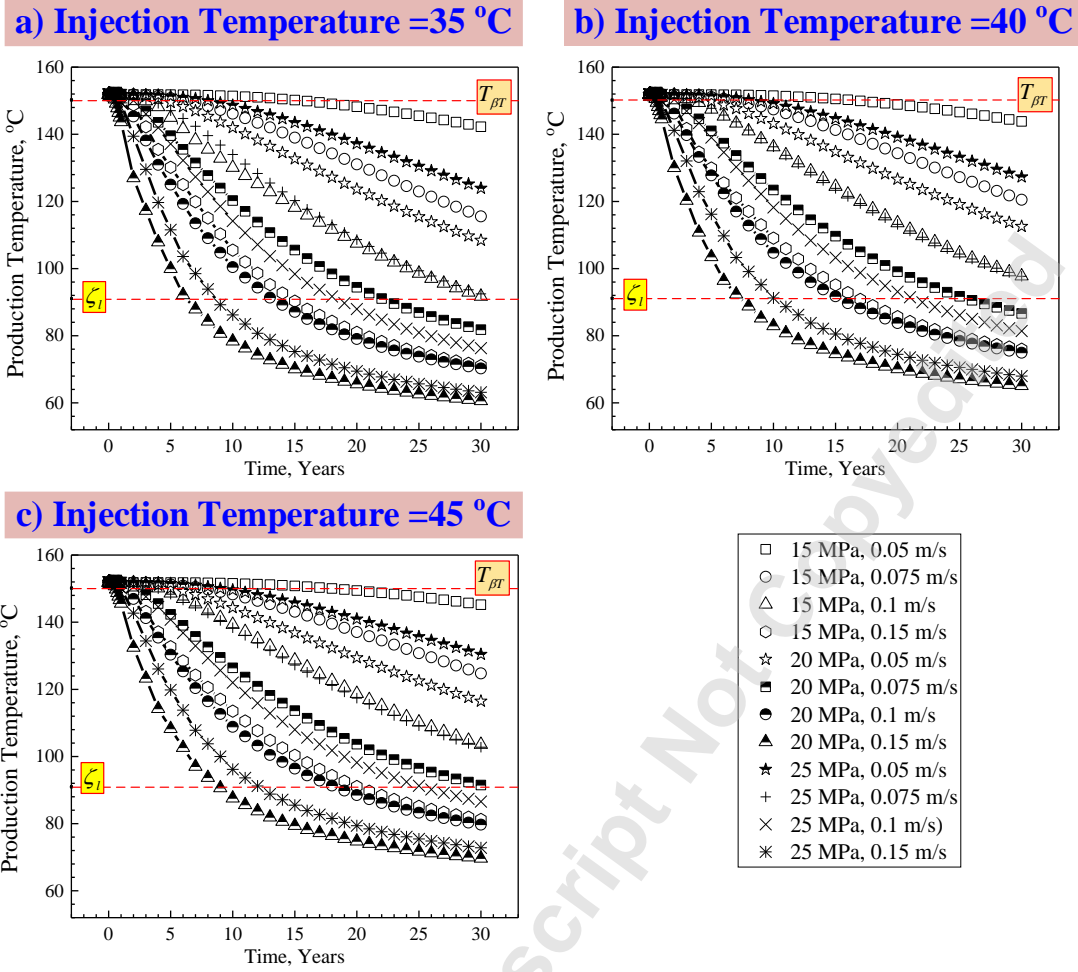
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**Fig 6** Comparison of SCCO<sub>2</sub> and water as geofluids on spatiotemporal variation of temperature in the reservoir and fracture at injection/production velocity of 0.1 m/s and initial pressure of 20 MPa, and injection temperature of 35°C.

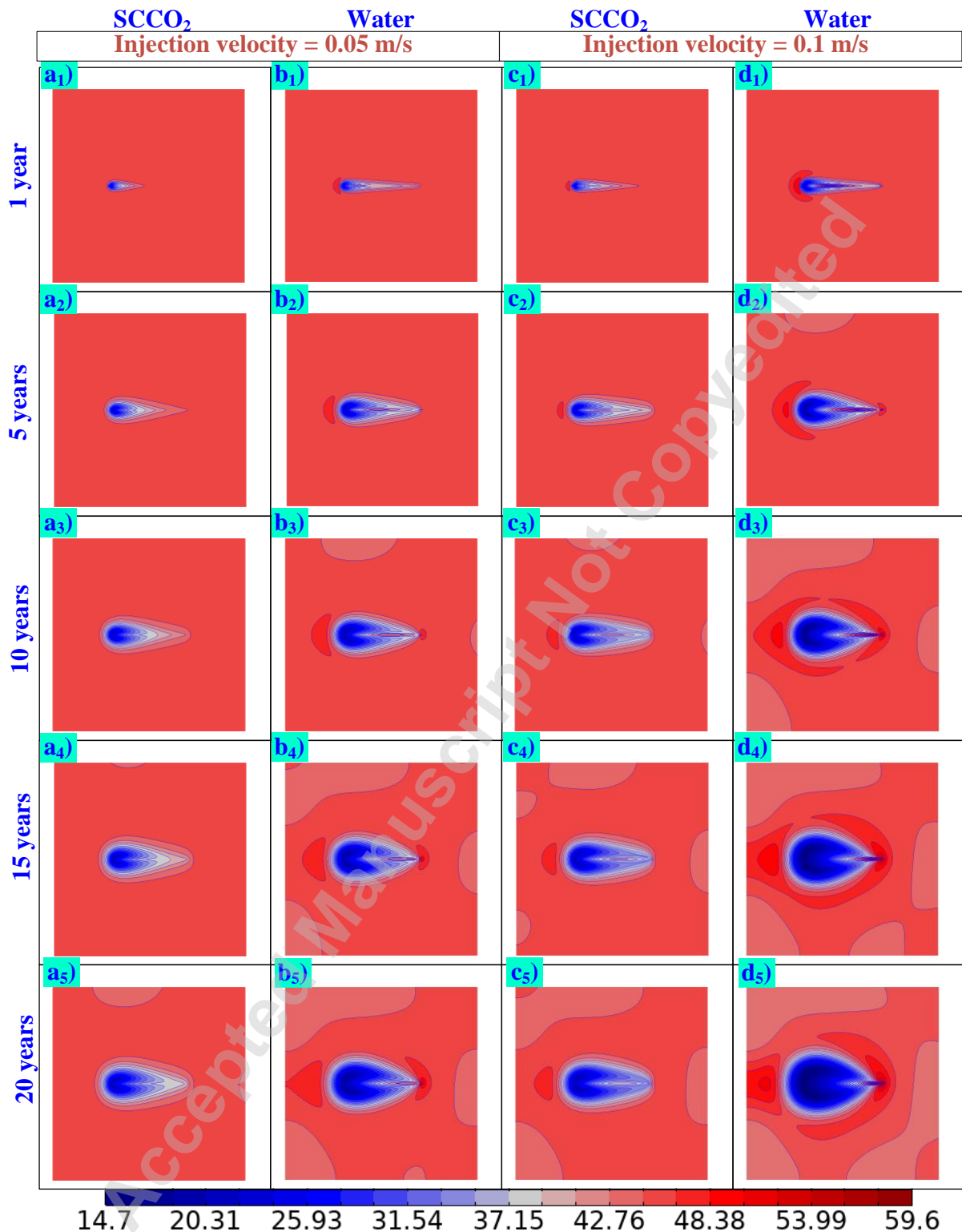
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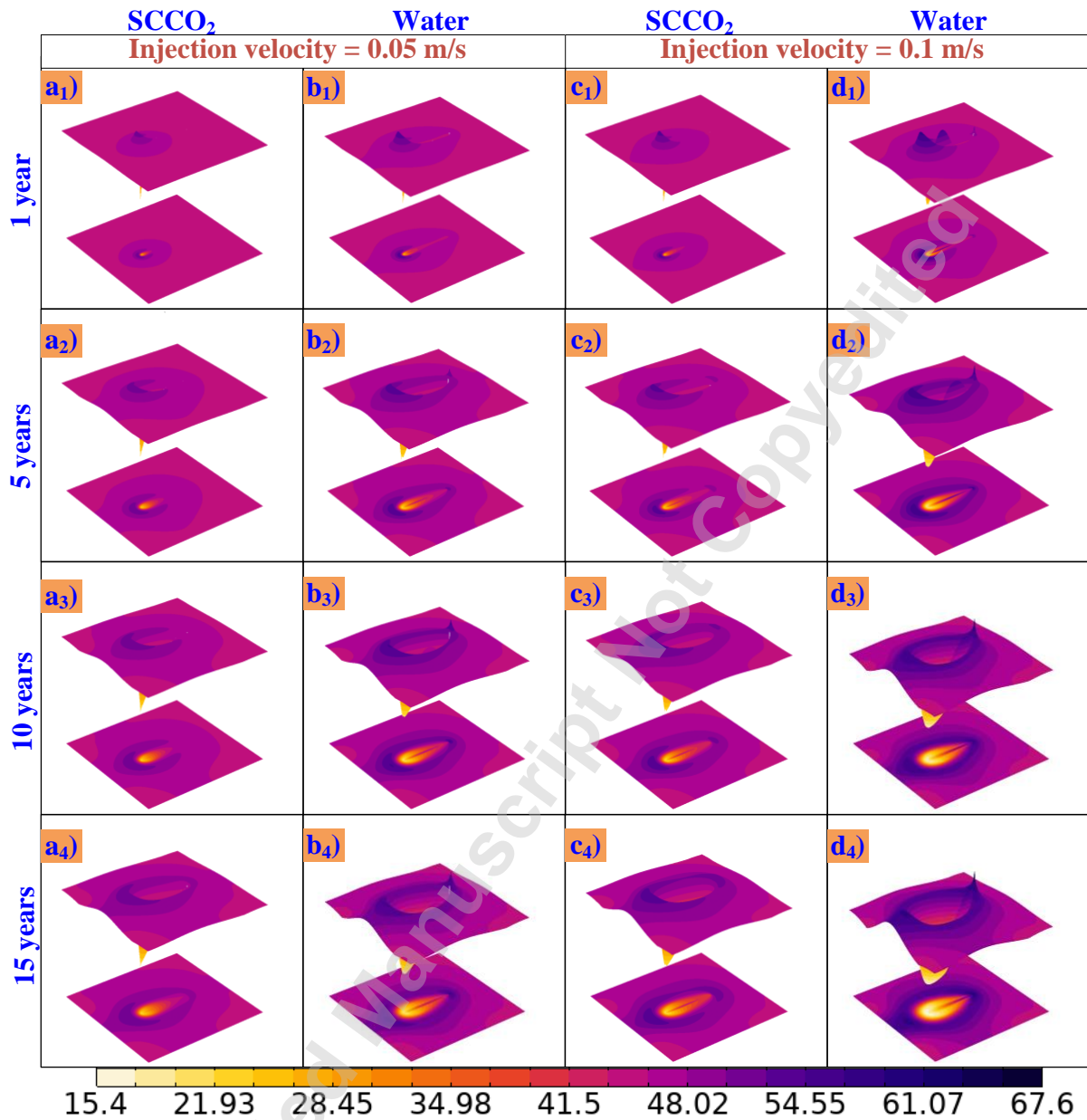
**Fig 8** Production temperature when using SCCO<sub>2</sub> as geofluid at different injection rates.



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**Fig 9** Spatiotemporal variation of Von-mises stress in MPa (aperture=0.5 mm and Biot-Willis's coefficient=0.5, initial reservoir pressure=20 MPa and injection temperature=40°C.)

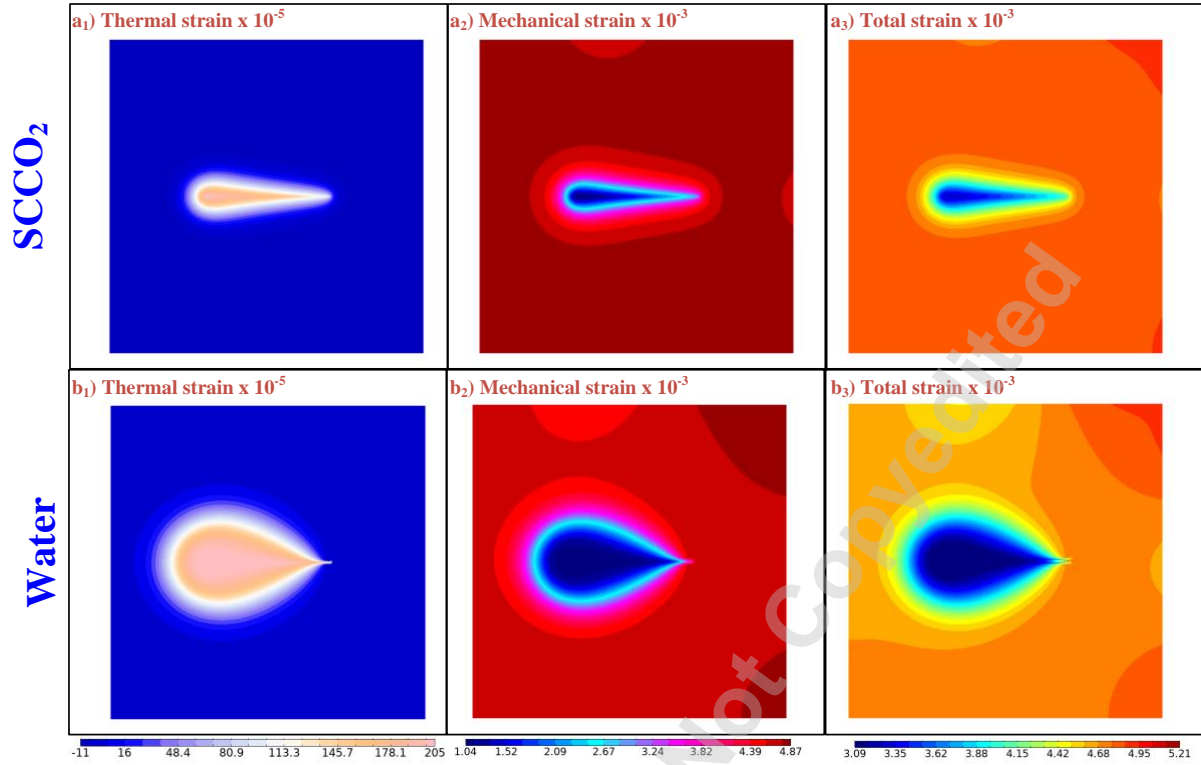
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**Fig 10** Spatiotemporal variation of tresca stress in MPa (aperture=0.5 mm and Biot-Willis coefficient=0.5, initial reservoir pressure=20 MPa and injection temperature=40°C.)

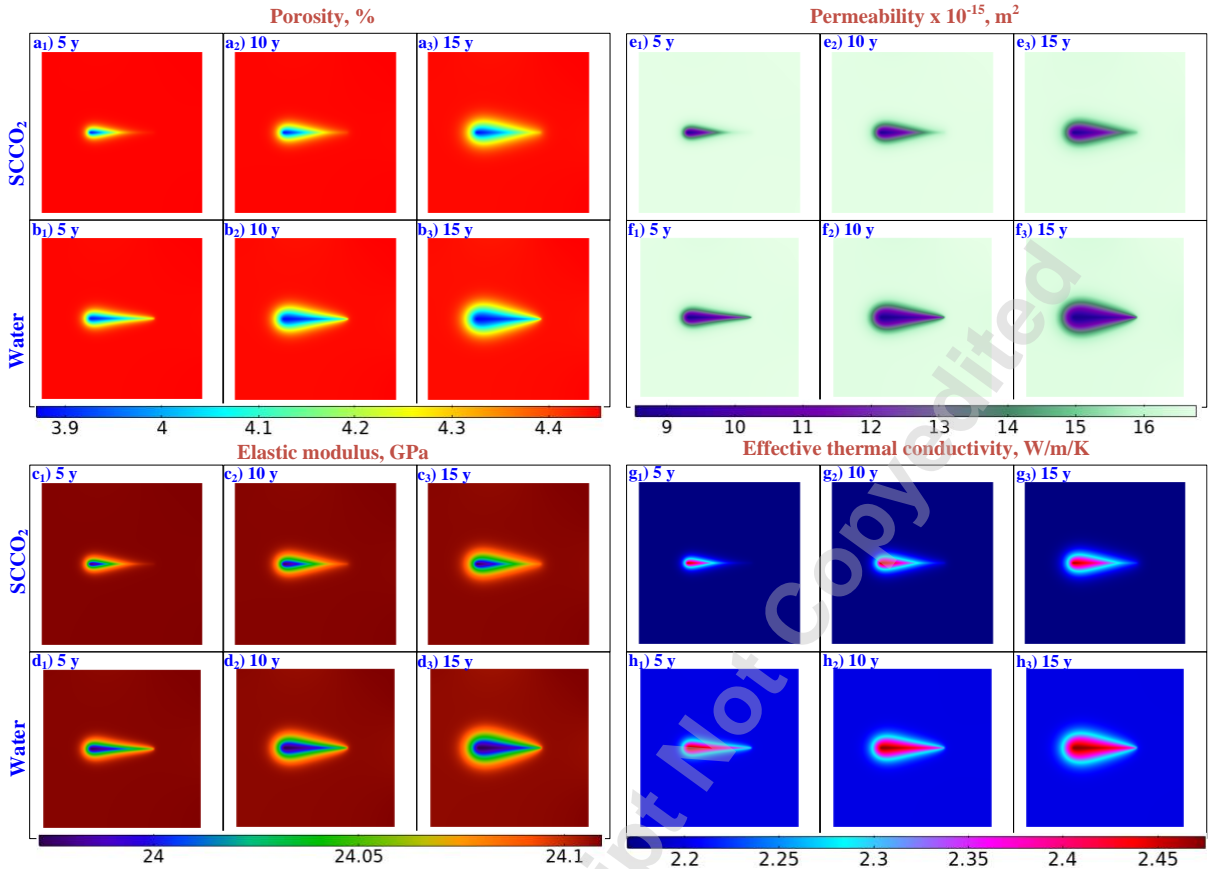




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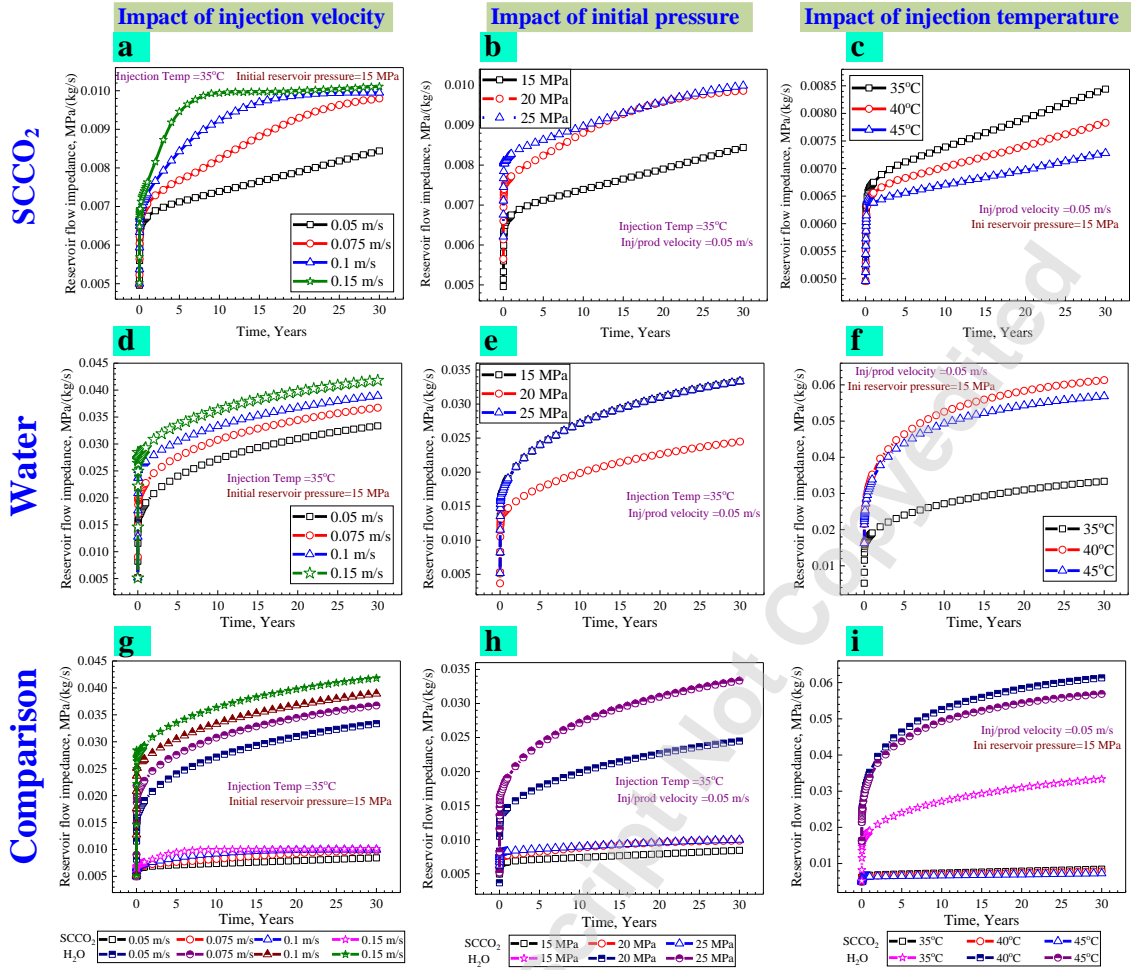
**Fig 11** Spatial variation of different variants of strains after 10 years of injection and production operation (injection/production velocity= 0.05 m/s, aperture=0.5 mm and Biot-Willis coefficient=0.5, initial reservoir pressure=25 MPa and injection temperature=45°C.)

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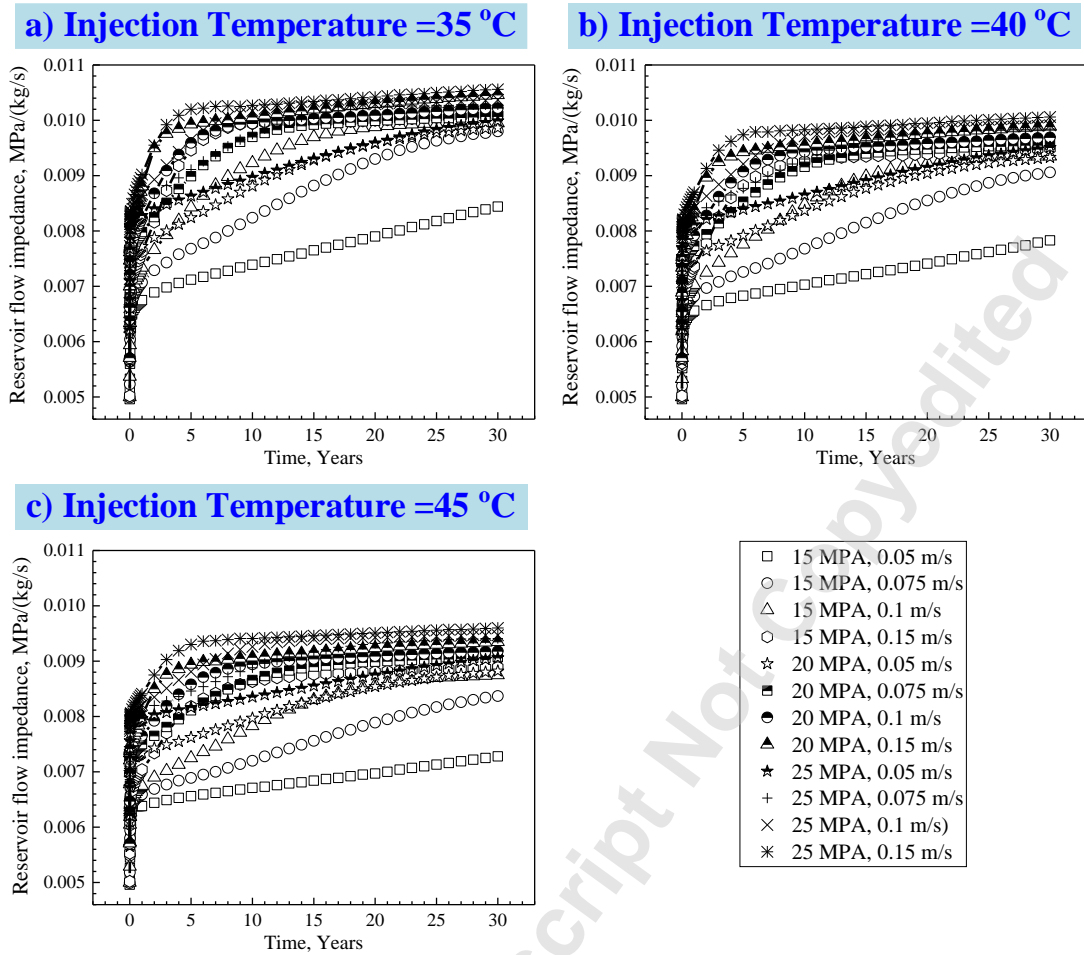
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**Fig 12** Spatiotemporal variation of permeability, porosity, young's modulus, and effective thermal conductivity (injection/production velocity= 0.05 m/s, aperture=0.5 mm and Biot-Willis coefficient=0. 5), initial reservoir pressure=20 MPa and injection temperature=40°C.



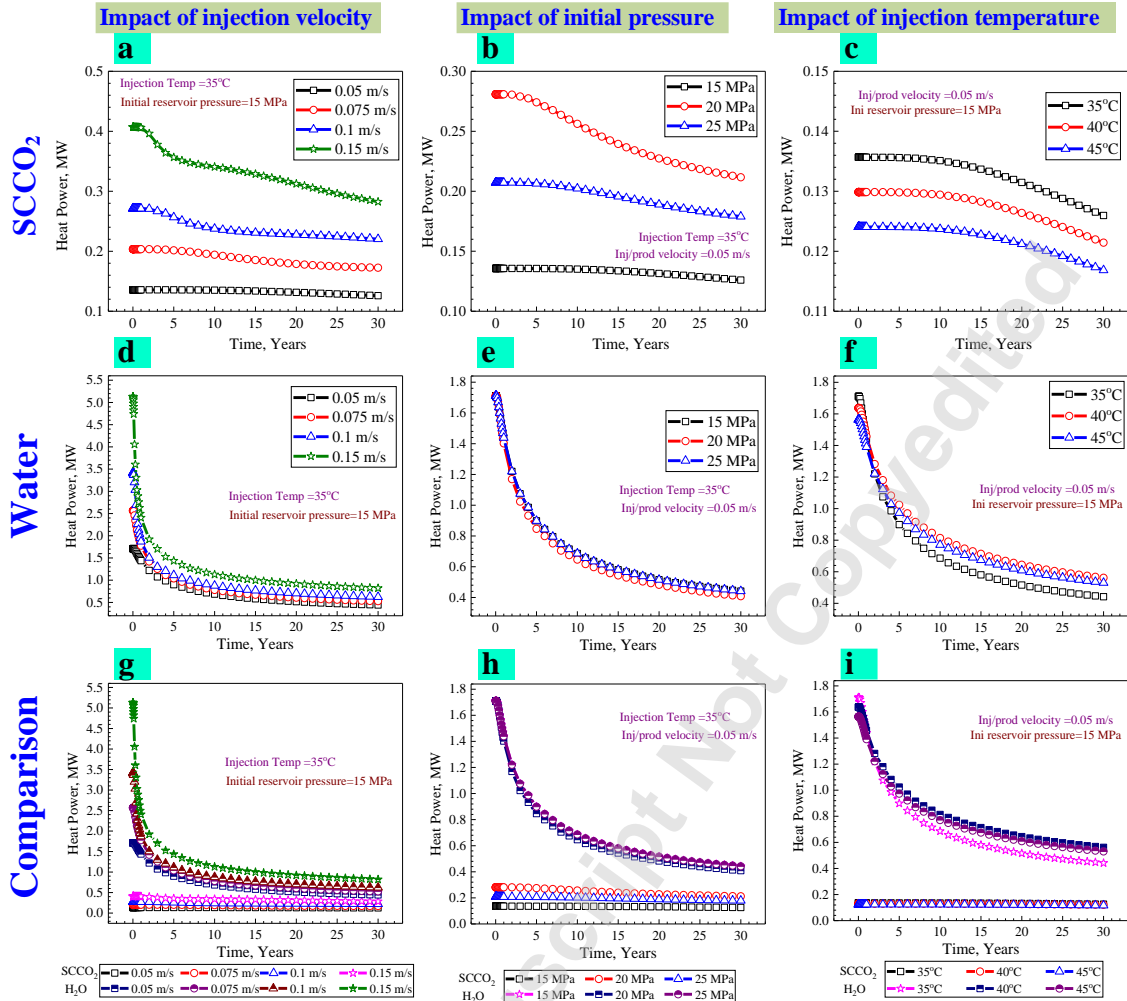
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**Fig 13** Impact of injection/production velocities, initial reservoir pressure, and injection temperature on the reservoir flow impedance when using SCCO<sub>2</sub>, and water as geofluids.



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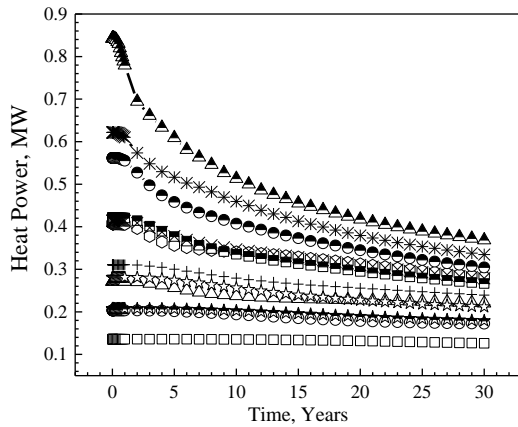
**Fig 14** Reservoir flow impedance when using SCCO<sub>2</sub> as geofluid at different injection rates.



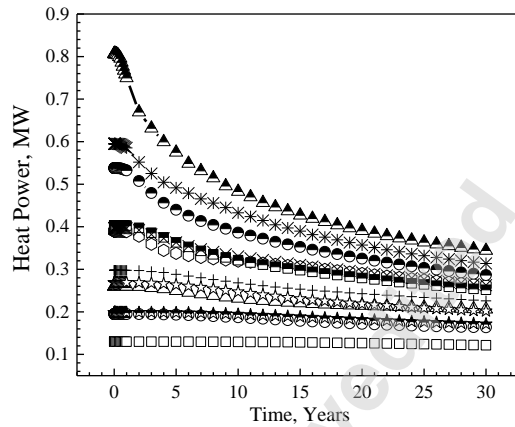
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**Fig 15** Impact of injection/production velocities, initial reservoir pressure, and injection temperature on the heat power when using SCCO<sub>2</sub>, and water as geofluids.

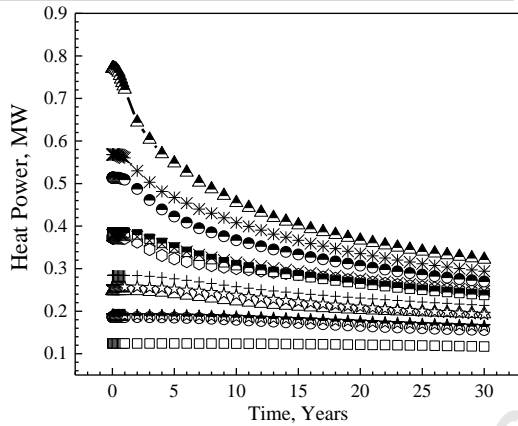
a) Injection Temperature = 35 °C



b) Injection Temperature = 40 °C



c) Injection Temperature = 45 °C



- 15 MPA, 0.05 m/s
- 15 MPA, 0.075 m/s
- △ 15 MPA, 0.1 m/s
- ◇ 15 MPA, 0.15 m/s
- ☆ 20 MPA, 0.05 m/s
- 20 MPA, 0.075 m/s
- 20 MPA, 0.1 m/s
- ▲ 20 MPA, 0.15 m/s
- ★ 25 MPA, 0.05 m/s
- + 25 MPA, 0.075 m/s
- × 25 MPA, 0.1 m/s
- \* 25 MPA, 0.15 m/s

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**Fig 16** Reservoir flow impedance when using SCCO<sub>2</sub> as geofluid at different injection rates.

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**LIST OF TABLES**

803

**Table 1** Rock matrix, fracture, and injection fluid properties

Property	Rock Properties	Fluid Properties
Density, kg/m <sup>3</sup>	2600	Eq (27) or Eq. (31)
Dynamic viscosity, Pa.s	-	Eq (26) or Eq. (30)
Thermal conductivity, W/m.K	Eq (25)	Eq (29) or Eq. (33)
Heat capacity at constant pressure, J/kg.K	Eq (24)	Eq (28) or Eq. (32)
Coefficient of Thermal expansion, K <sup>-1</sup>	2 x10 <sup>-5</sup>	-
Initial Youngs Modulus, GPa	24	-
Poisson's ratio	0.26	-
Initial Porosity	0.04	-
Initial Permeability, m <sup>2</sup>	9.8692x10 <sup>-15</sup>	-
Ration of Specific heats	-	1.0
Biot-willis coefficient	0.5	-
Fluid-injection rate, m/s	-	0.05, 0.075, 0.1, and 0.15
Fluid-production rate, m/s	-	0.05, 0.075, 0.1, and 0.15
Initial reservoir Temperature, °C	151.85 K	-
Fluid injection Temperature, °C	-	35, 40, and 45
Initial Youngs Modulus-Fracture, GPa	2.4	-
Poisson's ratio-Fracture	0.104	-
Fracture aperture ( $d_f$ ), mm	0.5	-
Fracture permeability, m <sup>2</sup>	$d_f^2/12$	-
Fracture porosity	1	-
Boundary load: x-direction, MPa	48	-
Boundary load: y-direction, MPa	48	-

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**Table 2** Values used for the validation of heat transfer in single fracture.

Parameter	Numerical Value
Rock density, kg/m <sup>3</sup>	2700
Rock Thermal conductivity, W/m·K	3.0
Rock Heat capacity at constant pressure, J/kg·K	1000
Coefficient of Thermal expansion, K <sup>-1</sup>	0.0001
Initial Youngs Modulus*, GPa	30 GPa
Poisson's ratio	0.3
Rock Initial Porosity	0.01
Initial Permeability, m <sup>2</sup>	1×10 <sup>-7</sup>
Ration of Specific heats	1
Biot-willis coefficient	1
Fluid density, kg/m <sup>3</sup>	1000
Dynamic viscosity, Pa·s	0.001
Fluid Heat capacity at constant pressure, J/kg·K	4200
Fluid Thermal conductivity, W/m·K	0.6
Flow velocity, m/s	0.02 m/s
Initial reservoir Temperature, °C	80
Fluid injection Temperature, °C	30

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**Figure Captions List**

Fig. 1	Geometry geothermal reservoir with single fracture (a), boundary conditions (b) and the meshed geometry (c).
Fig. 2	Comparison of the properties of SCCO <sub>2</sub> and water within the operating pressure and temperatures.
Fig. 3	Schematic of the solution process for fully coupled Thermo-hydro-geomechanical model in geothermal reservoir with fracture.
Fig. 4	Verification for the heat transfer in single fracture with analytical solution.
Fig. 5	Spatiotemporal variation of temperature in the reservoir and fracture with different injection/production velocities when using SCCO <sub>2</sub> as geofluid at initial pressure of 15 MPa, and injection temperature of 35°C
Fig. 6	Comparison of SCCO <sub>2</sub> and water as geofluids on spatiotemporal variation of temperature in the reservoir and fracture at injection/production velocity of 0.1 m/s and initial pressure of 20 MPa, and injection temperature of 35°C.
Fig. 17	Impact of injection/production velocities, initial reservoir pressure, and injection temperature on the production temperature when using SCCO <sub>2</sub> , and water as geofluids.
Fig. 18	Production temperature when using SCCO <sub>2</sub> as geofluid at different injection rates.
Fig. 19	Spatiotemporal variation of Von-mises stress in MPa (aperture=0.5 mm and Biot-Willis's coefficient=0.5, initial reservoir pressure=20 MPa and injection temperature=40°C.)
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Fig. 11	Spatial variation of different variants of strains after 10 years of injection and production operation (injection/production velocity= 0.05 m/s, aperture=0.5 mm and Biot-Willis coefficient=0.5, initial reservoir pressure=25 MPa and injection temperature=45°C.)
Fig. 12	Spatiotemporal variation of permeability, porosity, young's modulus, and effective thermal conductivity (injection/production velocity= 0.05 m/s, aperture=0.5 mm and Biot-Willis coefficient=0. 5), initial reservoir pressure=20 MPa and injection temperature=40°C.
Fig. 13	Impact of injection/production velocities, initial reservoir pressure, and injection temperature on the reservoir flow impedance when using SCCO <sub>2</sub> , and water as geofluids

Fig. 14	Reservoir flow impedance when using SCCO <sub>2</sub> as geofluid at different injection rates.
Fig. 15	Impact of injection/production velocities, initial reservoir pressure, and injection temperature on the heat power when using SCCO <sub>2</sub> , and water as geofluids
Fig. 20	Reservoir flow impedance when using SCCO <sub>2</sub> as geofluid at different injection rates

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**Table Caption List**

Table 1	Rock matrix, fracture, and injection fluid properties
Table 2	Values used for the validation of heat transfer in single fracture

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