

1 **Cracking the code of past biodiversity responses to climate** 2 **change**

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24 **Keywords:** adaptation, dispersal, extinction, global change, natural experiments,
25 prediction

26 Ecologists are increasingly interested in understanding and predicting how
27 biological diversity will respond to climate change. In this context, past biodiversity
28 dynamics recorded in paleo-archives show a broad array of responses, including
29 tolerance, microevolution, migration, and extinction, occurring across a broad
30 range of taxonomic, temporal and spatial scales. However, this accumulated
31 knowledge is largely fragmented within disciplinary and scale boundaries, and
32 synthesis and integration across disciplines are still lacking. Significant knowledge
33 gaps center on the relative roles of evolutionary adaptation, phenotypic plasticity,
34 and dispersal for surviving climate change. Application of paleo-archives offers
35 great opportunities to understand biodiversity responses to future climate change,
36 but also poses challenges for both the paleoecological and global change
37 communities. We outline cross-disciplinary approaches that will better apprehend
38 the mechanisms allowing species to survive, facilitating prediction of future changes
39 in biological diversity.

40

41 **Does the future of biodiversity lie in the past?**

42 Atmospheric CO₂ levels may rise up to 450-500 ppm by the end of this century, driving
43 an increase in global average temperature on the order of 2 to 5 °C [1], with future
44 climate forcing potentially without precedent in the last 420 million years [2]. Projected
45 magnitudes and rates of future climate change pose major threats to biodiversity [3-8],
46 and the scientific community is struggling to fully apprehend gross and net responses of
47 biodiversity to climate change, anticipate whether species can respond quickly enough,
48 and pinpoint the various roles of life-history properties (e.g., dispersal, genetic diversity,

49 reproductive strategies, phenotypic plasticity, population growth rates) in adapting to a
50 changing environment. In this context, scientists are looking to the past, using geo-
51 historical records to learn how individuals, populations, communities and biomes have
52 responded to climatic changes [9-22]. Whether individuals and populations will adapt
53 by evolutionary change or plasticity, whether they will migrate fast enough, and
54 whether those responses will be adequate to forestall collapses of species ranges and
55 prevent widespread species extinctions can be explored using case studies from the past.
56 Indeed, past climate change, whether abrupt or gradual, and whether occurring in deep
57 time or recent history, offers a vast set of unplanned natural experiments to explore
58 biodiversity responses and test ecological and evolutionary theories. Recent years have
59 seen the accumulation of well-documented examples of the influence of climate-change
60 on persistence, adaptation and diversification, dispersal, and extinction [11-15, 23]. The
61 effect of climate change on migration rates and routes have been intensively studied by
62 biogeographers and palaeoecologists, augmented recently by molecular markers and
63 ancient DNA (aDNA). *In situ* tolerance to changing climate conditions has been
64 explored in the fossil record using functional morphology, ecophysiology, evolutionary
65 genetics, and developmental plasticity, including recent experimental approaches (e.g.,
66 ‘resurrection ecology’; [24,25]). Finally, paleoecological records of population
67 extirpations and species extinctions provide information on the nature and consequences
68 of failure of tolerance, evolution, or migration [26,27].

69

70 However, key knowledge gaps remain. The relative importance of different mechanisms
71 for persistence (e.g., evolutionary adaptive change versus phenotypic plasticity), and the
72 nature and rates of climate-driven anagenetic and cladogenetic evolution remain poorly
73 understood. The relative efficacy of *in situ* tolerance and migration under different rates

74 and magnitudes of climate change is obscure. Although much attention has been
75 devoted to paleoecological records of migration, rates and underlying controls are not
76 clear except in a few specific cases. Moreover, significant challenges remain for better
77 integrating knowledge, scales, methods and data from a variety of biological disciplines,
78 from paleoecology to genomics. In this review, we (1) synthesize the modal responses
79 of biodiversity to past climate change from deep to recent time (tolerance, migration,
80 and their simultaneous failure, resulting in extinction), (2) identify key knowledge gaps
81 concerning underlying mechanisms (which span a broad set of biological disciplines),
82 and (3) identify and advocate for new approaches that integrate multiple methods and
83 disciplines to better apprehend the strategies of life to adapt to climate change and for
84 anticipating future responses of biological diversity.

85

86 **Biodiversity responses to climate change**

87 Tolerance, Adaptive Evolution, and Diversification

88 Biotic adaptation responses to climatic and environmental changes as shown by
89 paleoarchives vary from macroevolutionary divergences (at very long (10^6 - 10^7 yr) time
90 scales to adaptive evolution (10^0 - 10^5 yr) to phenotypic adjustments in place (10^{-1} - 10^3
91 yr). Long-term climate change has been considered an important driver of high-order
92 diversification, as clades respond to new climatic regimes, although its effects vary with
93 timespan, habitat, and clade (e.g., marine bivalves [28]) and North American mammals
94 ([29]), and more generally with the ecology of individual species (e.g., Cenozoic
95 macroperforate planktonic foraminifera [30]). In shorter time spans, many individuals
96 and populations can tolerate a high degree of climate change in situ. Some long-lived
97 modular organisms (e.g., corals, plant genets) can survive in place for centuries or even

98 millennia, spanning a broad range of climate change and variation from interannual to
99 millennial [31-33]. Some species have persisted in place since before the last glacial
100 maximum 21,000 years ago; well-documented paleoecological cases include woody
101 plant species in central Europe [34,35], and in unglaciated sections of North America,
102 both eastern [33] and western [36-38].

103 The paleo genetic records indicate also that adaptive evolution can support long-term
104 persistence of clades in response to climate change [39]. Adaptations can enable
105 exploitation of new niches: Adaptive mutations in woolly mammoth haemoglobin
106 allowed the exploitation of high-latitude cold environments during the Pleistocene [40].
107 Similarly, the genome comparison of brown bears and polar bears revealed that over
108 half million years, natural selection drove major changes in polar bear genes related to
109 fat transport in the blood and fatty acid metabolism [41]. Moreover, faster traces of
110 microevolutionary changes are recorded in Przewalski's horses after they occupied
111 Yakutia some centuries ago, following the migration of the Yakut people and adapting
112 to severe cold climate [42]. In more recent periods of time, examples of
113 microevolutionary responses to climatic change include changes in the body color of
114 owls during warmer winters [43], or adaptive changes in the flowering time of Brassicas
115 in response to drought [44].

116
117 Whether adaptive evolutionary change or plasticity are the prevalent strategy to adapt in
118 situ to climate change, and at what spatial and time frame these two processes play a
119 role, can be difficult to disentangle for extant populations and even more challenging for
120 ancient extinct populations, but both have doubtless been key processes in population
121 persistence under climate change[45]. For most reported cases of climate-driven
122 phenotypic changes in the wild, it remains unclear whether they are caused by

123 microevolution or phenotypic plasticity, although recent meta-analyses [45, 46] suggest
124 that most responses to climatic change are mediated by phenotypic plasticity (see also
125 [47-50]).

126 Migration

127 Migration has been a dominant response of species to climate shifts in the past [51].
128 Past migrations are typically inferred from spatial and temporal patterns in fossil data
129 [52,53], geographic patterns in genetic markers of extant and extinct populations [54],
130 or both (55). The fossil record has indeed provided abundant evidence of (mostly)
131 poleward migration of many species following warming of the Paleocene-Eocene
132 Thermal Maximum [56], or equatorward expansion of plant cold-adapted species during
133 the late-Tertiary global cooling [57]. Likewise, both fossil and genetic approaches
134 indicate (mostly) poleward migrations since the last glacial maximum [51, 54].

135

136 Behind these general trends lay multiple different specific species trajectories when
137 migrating. The best knowledge on past migration comes from trees, particularly in
138 Europe and North America during the postglacial period, thanks to extensive fossil
139 records and phylogeographic studies (e.g. 35, 55, 58). These studies reveal variable
140 migration patterns and rates among species, and even within populations of the same
141 species [59], but typical migration rates for the last deglaciation range from a few tens
142 to a few thousand m/yr, with averages around 2.7 km/decade [60-62]. Actual rates may
143 however have been much lower after accounting for refugial populations relatively
144 close to ice sheets [63,64] –but the distribution and frequency of such ‘cryptic’ refugial
145 populations remains under debate [65]. Overall, there is evidence of both high migration
146 rates –including fast responses to abrupt climate change [66]– as well as many species

147 lagging behind climate [67], reinforcing the high specificity of migration patterns across
148 taxa.

149 There are many different mechanisms by which climate change influences range shifts.
150 First, climate change can improve suitability beyond the range limit so that species may
151 establish at formerly unsuitable areas [68]. Hence, warming would promote expansion
152 of cold-limited taxa towards higher latitudes or altitudes [3,69]. Second, climate change
153 could foster colonisation of new areas in several ways: enhanced fecundity of source
154 populations (thus increasing propagule pressure), increased propensity to disperse or
155 emigrate (in animals), or acceleration of dispersal processes [70]. Climate change can
156 also enhance establishment of propagules after arrival, both directly [71] and
157 (particularly in rapid climate change) by reducing populations of dominant species, via
158 mortality or disturbance [72]. Finally, climate change could reduce the probability of
159 extinction of leading edge populations, for instance due to extreme climatic events [73].
160 A variety of processes are involved in species migrations and range shifts, all of which
161 can be directly or indirectly (e.g. mediated by species interactions) influenced by
162 climate change [72, 74-76]. A challenge for ecologists, biogeographers, and
163 paleoecologists is to identify generalizations, and to understand the role of species-
164 specific, locale-specific, and time-specific contingencies and idiosyncracies in driving
165 patterns and rates of migration.

166

167 Extinction

168 When species cannot tolerate climate change in situ, or colonize suitable habitat
169 elsewhere quickly enough, they become extinct. In extreme cases, many high-order
170 clades can be lost in mass extinction events. The Earth has indeed witnessed five such
171 events in the last 500 million years, in connection partly with intense climatic changes,

172 both cooling (i.e., End-Ordovician) and warming (i.e., end Triassic), with estimated
173 species losses ranging from 75%, in the Devonian, to 95% (Permian) [12,77]. More
174 recently, the cooling at the inception of the Quaternary caused regional extinctions of
175 many taxa [78]. The Quaternary has also witnessed major environmental changes, with
176 intense climatic shifts between cold and warm periods occurring, on occasions, within a
177 few decades [80,81]. The last deglaciation coincided with the disappearance of many
178 large mammals, with the magnitude of climatic versus anthropic impacts still being
179 hotly debated [27,82]. In contrast there are few documented cases of global plant
180 extinctions through the Pleistocene. However, extinct taxa may leave few traces in the
181 fossil record [83], and macrofossil analysis – often needed to determine species-level
182 differentiation – reveals that plant extinctions may have been more common than
183 assumed [e.g. 26]. Similarly, recent studies document species losses of amphibians and
184 reptiles [84], and taxa in the marine realm [85], which were thought to have been
185 relatively stable during the Pleistocene (e.g. herpetofaunal stability hypothesis).
186 Evidence that historic climate shifts (which are small compared to those that occurred
187 during the glacial inter-glacial cycles of the past) drove species extinctions is limited
188 [86], with rare exceptions being synergistic functions of both 20th century human-
189 induced climate change and other proximate drivers of extinction (including infectious
190 diseases) [87]. However, anthropogenic climate disruption is predicted to soon compete
191 with habitat destruction as the most important driver of contemporary extinctions
192 [88,89].

193

194 Climate change may trigger extinctions and local extirpations by surpassing the
195 physiological limits of species, by reducing primary productivity of ecosystems and
196 thereby local population fitness across food webs, and indirectly by disrupting

197 ecological interactions via changes in species distributions or phenology. For instance,
198 coral bleaching, the loss of intracellular endosymbionts due to the increase in frequency
199 of extreme heating episodes and changes in the carbon cycle are one of the main
200 supported mechanisms behind coral extinctions during the five mass extinction events
201 [90]. Also, drier and colder climatic conditions during the LGM triggered a reduction in
202 overall primary productivity, provoking losses in genetic diversity and populations of
203 large grazers [13] , depleting lineages, for example, of bowhead whales [91], and
204 contributing to local and global extinctions [27]. These pathways to extinctions in
205 different periods of the Earth's history share some commonalities. In particular, climatic
206 changes that exceed in magnitude and speed experienced during the evolutionary
207 history of species usually trigger extinction events [92], and climate change has
208 frequently interacted with other extinction drivers.

209 **Unknowns, challenges and routes ahead**

210 Our review of the modal responses of biodiversity to past climate change unveils key
211 knowledge gaps concerning the underlying mechanisms. We identify and discuss them
212 here and propose new integrative approaches that show potential to crack the code for
213 how biodiversity responds to climate change.

214

215 Evolutionary adaptation versus plasticity?

216

217 Climate-relevant decisions and policies implemented today (e.g., levels of CO₂
218 emissions) have both short and long-term consequences for future biodiversity,
219 influencing migration, divergence, speciation, hybridization, anagenetic evolution and
220 extinction. Paleo-archives reveal that speciation, evolution and phenotypic change have
221 played roles in species responses to past environmental changes. However, the relative
222 roles of those mechanisms in different settings, for different taxa, and across different
223 timespans need clarification and exploration.

224 Comparative phylogenetics and novel macroevolutionary approaches are offering new
225 insights into speciation and phenotypic change in response to major climatic shifts [94].

226 Comparative approaches allow fitting various models of phenotypic evolution and
227 diversification to phylogenies in order to estimate evolutionary rates, including
228 speciation and extinction [14]. Recently, models that can explicitly test for the effect of
229 climatic changes on these evolutionary rates have been developed [95-98]. With
230 advances in genomics, particularly methods for treating minute and degraded samples,
231 new critical sources of information will become available. For example, full genomes
232 for thousands of bird species are being sequenced [99], which will reveal

233 macroevolutionary consequences of past climate change and dynamics of effective
234 population sizes, and link them to climatic changes over millions of years [100].
235
236 Spatial analysis of species and populations provide important insights into the
237 evolutionary processes that led to present day genetic and phenotypic diversity.
238 However, when limited to exploring current patterns, these analyses are not suitable to
239 measure past processes. Past evolutionary events can be inferred from molecular data,
240 but these events may be confounded by phylogeographic signals, complicating
241 identification of the causes of adaptive and demographic changes. Conversely, long-
242 term observational studies enable measurement of evolutionary processes by comparing
243 temporal changes in genetic and phenotypic diversity with expectations of neutral and
244 adaptive evolutionary models. Long-term studies, however, may require commitments
245 beyond the career or life spans of individual researchers. ‘Resurrection Ecology’ (see
246 Glossary) provides an alternative path to reconstructing long-term patterns of
247 evolutionary changes and unraveling mechanisms of response to climatic and other
248 environmental changes [19]. Many life forms (zooplankton, insects, algae, fungi,
249 bacteria, plants) produce resting stages as part of their life cycles in response to
250 environmental hardship. When such resting stages can be recovered from ancient
251 sediments and reared in the laboratory, they can reveal molecular targets (genes,
252 metabolites, proteins) that enable evolution and adaptation to changing climate.
253 Resurrecting individuals from such species and populations across documented
254 temporal shifts in environment uniquely permits simultaneous measurement of both
255 plastic (phenotypic and behavioural) and genetic (evolutionary) responses to climatic
256 change, using common garden or transplant experiments [101,102]. Relative fitness of
257 both historical and modern populations can be measured in response to different

258 climatic regimes, including past, present and future. Such long-term studies, replicated
259 across multiple environments and taxa, can be a powerful resource for building models
260 to forecast species persistence [103] (Figure 2).

261

262 Migrating fast enough?

263

264 Although migration is a key process underlying range shifts and the spread of native
265 and invasive species, the migration capacity of species under rapid climate change
266 remains uncertain (75). While some taxa seem unable to shift ranges under changing
267 climates (67), others seem able to migrate at a fast pace [3]. Attempts to explain
268 observed range shifts based on species traits or ecological strategies have obtained
269 modest results [104,105; but see 106]. Low predictability may be expected given the
270 large number of processes involved in range shifts, as well as the complexity and path-
271 dependence when those processes interact. The dispersal process itself is highly
272 stochastic and inherently uncertain [107]. Other important processes include size and
273 fecundity of source populations (which determine propagule pressure), gene flow, local
274 adaptation, evolution of dispersal, biotic interactions (competition, facilitation,
275 mutualisms), Allee effects, and so on, all of which are likely to be affected by climate
276 change as well [75]. Spatial heterogeneity on the landscape plays a role (e.g., dispersal-
277 target size), as does high-frequency climate variability [58, 72]. As a result, we may not
278 be able to go much farther than estimating dispersal potentials for different species or
279 populations [107]. A critical challenge is to use paleoecological and ecological data to
280 identify generalizations that can emerge from the location-specific, species-specific, and
281 event-specific particulars of detailed case studies [58,72].

282

283 Paleoecology has largely contributed to estimate how fast species migrated under past
284 climatic changes under minimum levels of pre-historic global human intervention.
285 Unfortunately, contemporaneous dispersal rates are likely to be rather different than past
286 rates due to radically different conditions: more fragmented habitats, missing and novel
287 interactions, or nearly unlimited human-mediated dispersal [75]. Hence, estimates of
288 past migrations rates, however informative, may be of limited value when attempting to
289 forecast future range shifts. Instead, a better understanding of the causes of variation in
290 migration rates may move us forward. Hence, comparative studies of migration patterns
291 among tens or even hundreds of species could throw some light into the role of
292 environmental (contingent) factors as well as intrinsic factors that make some species
293 migrate faster, slower or not at all.

294

295 Anticipating future extinctions

296

297 Revealing how the accumulative failure of tolerance and migration mechanisms under
298 climate change lead to population extirpation and ultimately species extinction is of
299 utmost importance to provide robust scenarios for future biodiversity and to enhance
300 conservation strategies. Although past extinction events have provided better
301 knowledge on extinction dynamics and their relation to climatic changes, paleo-data
302 together with current data has only recently been fully implemented in quantitative
303 assessments of future risk of extinction [108]. Moreover, correlative approaches lacking
304 key biological mechanisms have dominated the forecasting of future responses of
305 biodiversity to climate change. A large gap remains between mechanistic experiments
306 and large-scale macroecological models that forecast the distribution and persistence of
307 biological diversity under future global climate change [109].

308

309 Models that include paleo-dynamics of taxon abundance predict no continental-level
310 extinctions for the end of this century [108], but large changes in the composition of
311 ecological communities, which reflects recorded trends for plants in the Pleistocene.
312 Moreover, a paradigm shift from correlative models of different complexity to
313 mechanistic simulations will bring deeper insights on the interplay of tolerance and
314 migration to explain species range dynamics and extinctions under climate change [20].
315 Recent insights on the factors correlating with declining genetic diversity, population
316 sizes, and local and global extinctions, have been achieved for megafauna species
317 during the Late Quaternary [13]; but these still are of limited utility for identifying
318 extinction mechanisms. Stochastic process-based models have been used to pinpoint
319 mechanisms driving past changes in species ranges and biological diversity [94,110].
320 Process-based models applied to explain past recorded changes in biological diversity
321 will improve our ability to explore the role of different mechanisms and explain
322 observed patterns in paleo-archives like species distribution, abundance and genetic
323 diversity [111].

324

325 Finally, integration of the genetic level of biodiversity has been recently advocated and
326 holds great potential to include missing mechanisms such as adaptive evolution [20,
327 105]. Low quality and limited quantity of sample material has impeded application of
328 high-throughput technologies [112] to paleo records. However, recent advances in next-
329 generation sequencing technologies have at least partially mitigated these limitations,
330 allowing extraction of data from small samples and degraded materials, opening a new
331 era for the study of evolutionary processes on macroecological scales (e.g. speciation
332 and extinction of large mammals [113]), as well as at the microecological scale. These

333 technological advances will allow a truly multi-omics approach to identify the
334 molecular targets (proteins, genes and metabolites) driving evolution and speciation at
335 different temporal scales. Long term studies on paleo and neo biological archives,
336 incorporating such approaches as resurrection ecology [19] and experimental
337 macroecology [114] where possible, and combined with spatially explicit mechanistic
338 models, can provide new paths for incorporating eco-evolutionary dynamics in forecast
339 models, and for improving predictions for species survival and persistence under
340 climatic change (see Figure 2).

341

342 **Concluding remarks**

343

344 Climate change has triggered large and persistent effects on biological diversity,
345 including speciation, redistribution, local adaptations and extinction events. However, a
346 deeper mechanistic understanding of these dynamics is urgently needed (see also
347 Outstanding Questions). Until recently, lines of evidence have suggested that biotic
348 responses to climate change have been dominated by migration. It is now clear from
349 both paleoecological and ecological perspectives that *in situ* tolerance and adaptative
350 evolution are also key responses to climate change. Although adaptation is now an
351 important object of study, there remain fundamental questions to be addressed: How is
352 adaptive evolution shaped by migration? Conversely, how is migration influenced by
353 adaptive evolution? How do tolerance, adaptive evolution, and migration interact in
354 specific circumstances to reduce or amplify risk of extinction? We advocate for
355 integrating observed long-term responses, paleo experimental arrays, ecological and
356 evolutionary disciplines at macroecological scales (Figure 3). Applying multiomics
357 approaches in paleorecords applied across large spatio-temporal axes has the potential

358 to provide the unique setting to link evolution and migration across long reaches of
359 time, and for a deeper understanding of the roles of adaptation in presence of climate
360 change. Cracking the code of past biodiversity responses to climate change will increase
361 the ability to anticipate responses of biological diversity to climate change.

362

363 **Acknowledgements**

364 This review emerged from a symposium celebrated at the International Biogeography
365 Society 7th Biennial Meeting in Bayreuth, Germany. D.N.B. thanks Det Frie
366 Forskningsråd (EliteForsk) and the Center for Macroecology, Evolution and Climate.
367 F.R-S. was supported by a postdoctoral fellowship from the Spanish Ministerio de
368 Economía y Competitividad (FPD-2013-16756) and a Severo Ochoa Excellence Award
369 (SEV-2012-0262) to Estación Biológica de Doñana. HM acknowledges support from
370 the European Research Council Grant ERC 616419- PANDA.

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634 **Glossary**

635 **Adaptive evolution** –Results from the propagation of advantageous alleles in populations
636 through natural selection, driven by environmental selection pressure acting on genes
637 underlying species traits linked to fitness.

638 **Dormant propagules** – A still living but dormant seed, cyst, spore or egg preserved in ice, soil,
639 sediment, permafrost, or other media.

640 **Experimental Paleoecology:** Experimental studies to test sufficiency and necessity of
641 mechanisms (or combinations or sequences of mechanisms) invoked to explain paleoecological
642 phenomena.

643 **Migration** – Spatial displacement of organisms leading to shifts of species distributions with
644 climate change.

645 **Paleogenomics** – The study of ancient genomes to reveal functional genetic patterns through
646 time, supporting inferences concerning evolutionary adaptation, functional traits, population
647 dynamics, domestication, genetic events preceding extirpations or extinctions, and other
648 patterns of interest.

649 **Phenotypic plasticity** – Ability of individuals of a genotype to alter physiology, morphology,
650 anatomy, phenology, behaviour, or other phenotypic traits in response to environmental change.

651 **Resurrection ecology** – Study of traits and environmental responses of past populations by
652 hatching or germination of dormant propagules and culturing or cultivation of the organisms.

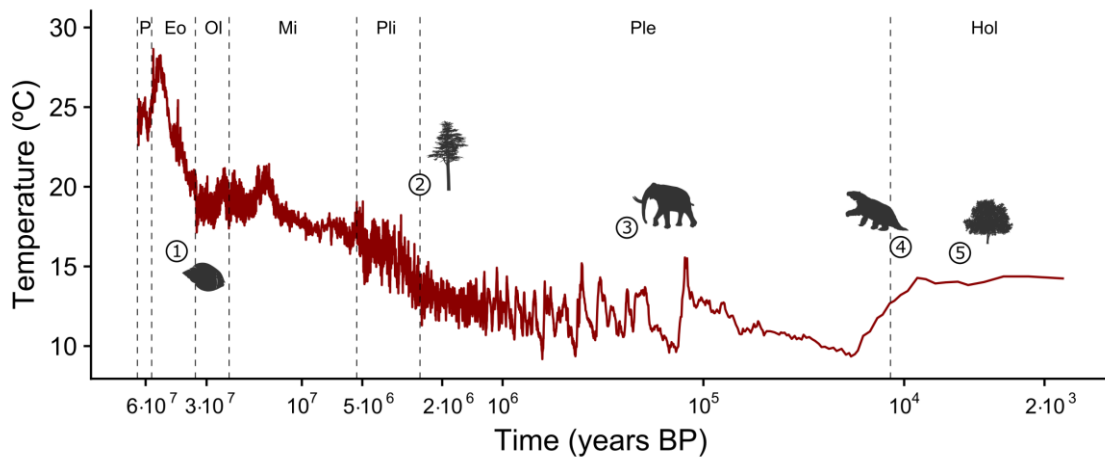
653 **Process-based models** - Spatially explicit approaches that simulate the effect of climate and
654 environmental conditions on important vital rates (including population growth, dispersal and
655 plasticity in demographic traits) to explain species distributions and their changes, including
656 range shifts and local extirpations.

657 **Tolerance** –Ability of a population to persist at a site under environmental change by adaptive
658 evolution, phenotypic plasticity, or both.

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661 **Box 1. Past responses of biological diversity under climate change.**



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663 Future climate forcing may surpass those of the previous several million years. Countless
664 individuals in thousands of species across the globe will need to tolerate climate change in situ,
665 disperse and migrate to more suitable climatic conditions, or undergo extinction. Figure 1
666 highlights a number of biodiversity responses directly or indirectly linked to climatic changes
667 along the Cenozoic (last 65 million years; temperature data from [118]). 1) Under a global cooling
668 trend, winters became $>4^{\circ}\text{C}$ colder across the Eocene/Oligocene boundary, partially driving
669 extinction of many terrestrial mammals in Europe and marine invertebrates [119]. 2) More than
670 half (52%) of the cool-temperate European tree genera did not survive to the glaciation cycles
671 starting at the end of the Pliocene [78]. 3) An adaptive mutation of haemoglobin enabled
672 mammoths to tolerate the very low temperatures at high latitudes [40]. 4) More than 70% of
673 megafauna genera in the Americas and Australia, and 40% in Eurasia, underwent extinction
674 within a relatively brief period of time (5,000-10,000 years; [12]. 5) Plants in North America
675 migrated northwards between 450 and 2200 km in less than 10,000 years under a warming of 5
676 degrees [120]. Abbreviations of geological epochs as follows: P = Palaeocene, Eo = Eocene, Ol
677 = Oligocene, Mi = Miocene, Pli = Pliocene, Ple = Pleistocene, Hol = Holocene.

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681 **Box 2. Correlations are not enough: experiments and simulations in paleobiology**

682 Much of the evidence for the impact of past climate change on biological diversity is
683 based on patterns of co-occurrence between past climatic events and biological
684 responses as speciation, migration, tolerance and extinction. For deep-time studies the
685 low temporal resolution of available dating techniques creates difficulties in aligning
686 relevant abiotic dynamics (i.e., climate change, acidification, volcanisms) with
687 biological events. Moving from correlations to causation is challenging because of the
688 co-varying climatic variables. Deep-time speciation and extinction events, for example,
689 have been shown to co-occur with both warming and cooling events but also in
690 coincidence with massive CO₂ releases and ocean acidification. Moreover, human
691 impacts during the Late Quaternary have affected biodiversity directly and indirectly,
692 interacting with climate. Enclosure ecological experiments coupled with paleo-
693 ecological data have provided evidence of widespread changes in community
694 composition and ecosystem structure and function due to large herbivore impacts [121]
695 Recently, the use of high throughput sequencing technologies on ancient DNA, the use
696 of paleogenetics, and advances in experimental approaches on living fossils (dormant
697 propagules) have provided important insights into genetic and phenotypic responses to
698 climate change [94]. The discovery of molecular targets (genes, metabolites, proteins)
699 of climate change that enable evolution and adaptation will help unveil mechanisms that
700 contribute to species and population persistence. In parallel, spatial modelling in
701 macroecology can offer more mechanistic insights by switching from correlative models
702 to process-based simulations. Stochastic simulations considering speciation, anagenetic
703 evolution, niche conservatism, range shifts and extinctions [94] are, for example, a
704 promising route to understand the role of those mechanisms in biodiversity responses
705 under climate change. We advocate for the development and further integration of

706 experiments, multi-omics and simulations on paleo-biological systems to better
707 anticipate future responses of biological diversity to climate change.

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711 **Box 3. Outstanding Questions**

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713 **-Does plasticity evolve under climate change?** The evolution of plasticity is an important
714 factor for population persistence in a variety of natural systems, but whether selection for
715 plasticity is the result of climate change acting as a main evolutionary pressure or an emergent
716 trait from selection at shorter scale needs further research. In particular, additional research on
717 the genetic basis and heritability of plasticity is needed so that we can gain a better
718 understanding of conditions under which plasticity is expected to evolve.

719

720 **-What are the long-term adaptive responses to climate change?** Temporal trajectories of
721 fitness changes in response to climatic and anthropogenic changes are paramount [19] to
722 unravel the evolutionary processes leading to population persistence (evolutionary rescue). This
723 knowledge, in turn, is critical to obtain realistic projections of species persistence in future
724 climate change scenarios. Our ability to confidently detect bottlenecks, local extirpation and
725 extinction in response to climatic or anthropogenic changes depends on ability to sample before
726 and after a drastic environmental change took place. The most direct and powerful way to gain
727 such information is by studying species producing dormant propagules (Box 1) that can be
728 sampled across temporal environmental shifts, including climate change [115].

729

730 **-Can we predict future extinctions with our current data?** Unifying the declining-species
731 paradigm and the small-population paradigm [116,117], which consider both the factors
732 contributing to the general decline of species before their populations become rare, and the
733 genetic and demographic factors promoting the extinction of small populations we can translate
734 and simulate these factors into process-based simulation models. Spatially explicit mechanistic
735 population models that include ecologically important phenotypic traits such as morphology,
736 physiology, phenology, evolutionary adaptive potential, species behaviour and species
737 interactions are a promising route. These types of models are in their infancy because both

738 environmental and empirical ecological variables are needed as prior information to correctly
739 predict future population persistence. We are therefore in much need firstly of more biological
740 and paleobiological data across large climatic and anthropogenic pressure gradients,
741 highlighting the key role field-work, expeditions, biological collections and natural history
742 museums to resolve the relevant societal challenges of the biodiversity crisis.

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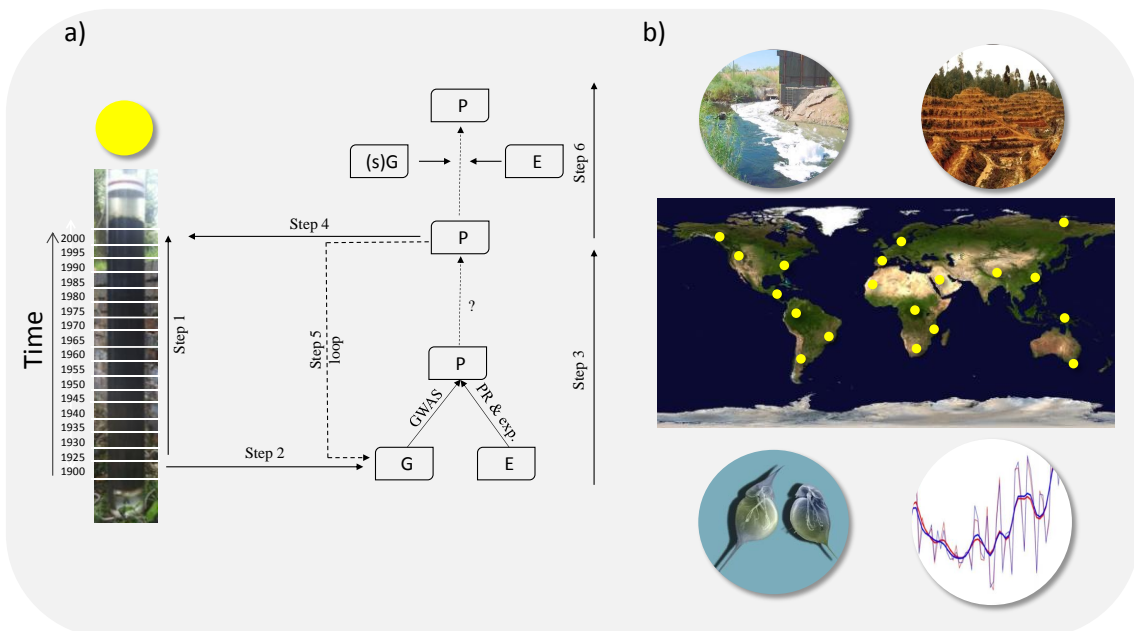
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Fig 2. Reconstructing historical patterns of evolutionary changes and unravels mechanisms of genetic and plastic response to anthropogenic environmental changes. a) Conceptual framework for the integration of resurrection ecology, paleo'omics' and macroecology. Using for example *Daphnia* as a system, we can resurrect propagules, step 1, and develop a matrix of genotypes (G) and phenotypes (P). Environmental variables (E) are also measured. The matrix of P and G is linked, step 2, via a genome wide association analysis (GWAS) and environmental variables (E) are linked to phenotypes (P) via experimental set ups and paleo records (PR). Using known E and trajectories of G over time, step 3, a model is applied to predict P. The outcome of the model is checked against empirical data, step 4. The model parameters are adjusted, step 5, to reflect real outcomes observed in empirical after several iterations). Using projected E and simulated G ((s)G) from past trajectories future P are predicted with a level of uncertainty. **b)** Scaling up the approach described in a) to a macroecological scale we can identify evolutionary and plastic responses of species to global anthropogenic pressures (within circles from upper left to bottom right: habitat degradation, land-use changes, invasive species and climate change).



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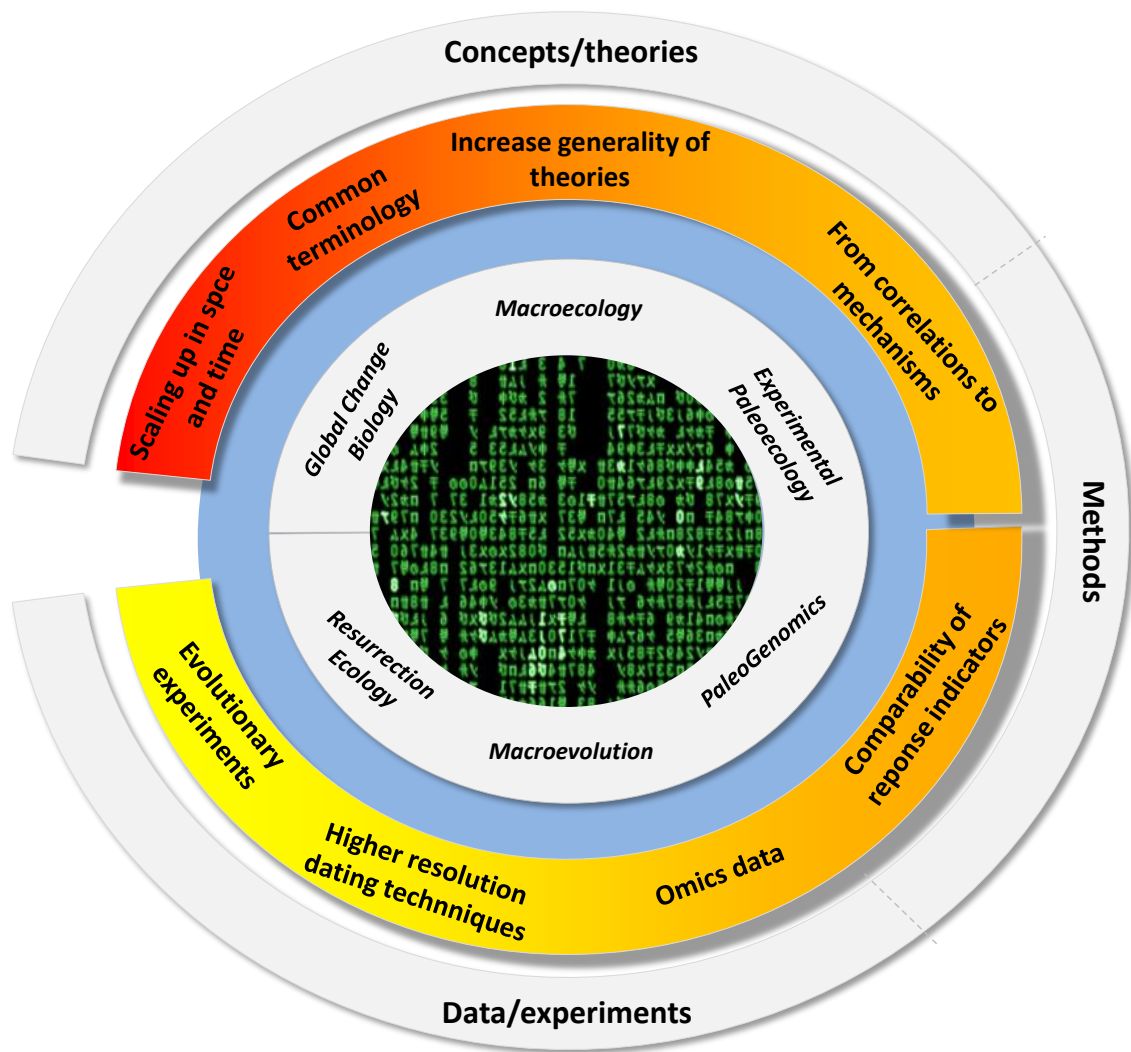
782 **Fig. 3. Cracking the code of biodiversity responses.** We identify a series of research

783 challenges, central ring, and disciplines, inner ring, to provide multiple lines of

784 evidences on the past magnitudes, rates and mechanisms of modal biodiversity

785 responses and to inform and feed future biodiversity scenarios.

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