1	Electron microprobe technique for the determination of iron
2	oxidation state in silicate glasses
3	Chao Zhang ¹ , Renat R. Almeev ^{1,*} , Ery C. Hughes ² , Alexander A. Borisov ³ , Eric P. Wolff ¹ ,
4	Heidi E. Höfer ⁴ , Roman E. Botcharnikov ⁵ and Jürgen Koepke ¹
5	1 Leibniz Universität Hannover, Institut für Mineralogie, Callinstrasse 3, D-30167, Hannover, Germany
6	2 School of Earth Sciences, University of Bristol, Bristol BS8 1RJ, UK
7 8	3 Institute of Geology of Ore Deposits, Petrography, Mineralogy, and Geochemistry, Russian Academy of Sciences, Staromonetny 35, 109017 Moscow, Russia
9 10	4 Institut für Geowissenschaften, Mineralogie, Johann Wolfgang Goethe-Universität, Altenhöferallee 1, D-60438 Frankfurt am Main, Germany
11 12	5 Institute für Geowissenschaften, Johannes Gutenberg Universität Mainz, J-J-Becher-Weg 21, D-55128 Mainz, Germany
13	* Corresponding author. Email: r.almeev@mineralogie.uni-hannover.de
14	ABSTRACT
15	We present a new calibration for the determination of the iron oxidation state in silicate
16	glasses by electron probe micro-analysis (EPMA) with the "flank method". This method is based
17	on the changes in both intensity and wavelength of the FeL α and FeL β X-ray emission lines with
18	iron oxidation state. The flank method utilizes the maximum difference for the FeLa and FeL β
19	spectra observed at the peak flanks between different standard materials, which quantitatively
20	correlates with the Fe ²⁺ content. Provided that this correlation is calibrated on reference materials,
21	the Fe ²⁺ / Σ Fe ratio can be determined for samples with known total Fe content. Two synthetic Fe-
22	rich ferric and ferrous garnet endmembers, i.e. andradite and almandine, were used to identify
23	the FeLa and FeL β flank method measuring positions that were then applied to the measurement
24	of a variety of silicate glasses with known $Fe^{2+}/\Sigma Fe$ ratio (ranging from 0.2 to 1.0). The
25	measured intensity ratio of FeL β over FeL α at these flank positions ($L\beta/L\alpha$) is a linear function of

26 the Fe^{2+} content (in wt%). A single linear trend can be established for both garnets and silicate

27 glasses with 4–18 wt% FeO_T (total iron expressed as FeO). In glasses with up to 18 wt% FeO_T 28 and 15 wt% TiO₂, no systematic compositional (matrix) effects were observed. A possible influence of Ti on the Fe^{2+} determination has only been observed in one high-Ti glass with ~25 29 30 wt% TiO₂, a content that is not typical for natural terrestrial silicate melts. The accuracy of the $Fe^{2+}/\Sigma Fe$ determination, which depends on both the Fe^{2+} content determined with the flank 31 method and on the total Fe content, is estimated to be within ± 0.1 for silicate glasses with 32 33 $FeO_T > 5$ wt% and within ± 0.3 for silicate glasses with low $FeO_T < 5$ wt%. The application of the flank method on silicate glasses requires minimization of the EPMA beam damage which can be 34 35 successfully achieved by continuous movement of the sample stage under the electron beam 36 during analysis, e.g. with a speed of 2 μ m/s.

37

INTRODUCTION

38 Fe is the most abundant transition metal in magmatic systems of the Earth. Depending on the redox condition. Fe can be present in different oxidation states (Fe^{3+} , Fe^{2+} , and Fe^{0}). The 39 40 oxidation state of Fe in natural silicate glasses is an important parameter that reflects the redox 41 conditions prevailing during magma generation and/or crystallization (e.g. Christie et al., 1986; 42 Bézos and Humler, 2005; Cottrell and Kelley, 2011; Kelley and Cottrell, 2009). It varies as a 43 complex function of oxygen fugacity, temperature, pressure, and melt composition (e.g., Sack et 44 al., 1981; Borisov and Shapkin, 1990; Kress and Carmichael, 1991; Nikolaev et al., 1996; 45 Moretti, 2005; Schuessler et al., 2008; Borisov et al., 2015). Due to the influence of ferrous and 46 ferric Fe on the local structure of silicate melt, the oxidation state of Fe can significantly 47 influence physical and chemical properties of silicate melts (e.g., viscosity, density, heat capacity, 48 degree of polymerization and phase equilibrium, see review by Wilke, 2005).

49 Both bulk and *in-situ* techniques are available to determine the oxidation state of Fe in geological samples, which is usually expressed as $Fe^{2+}/\Sigma Fe$ or $Fe^{3+}/\Sigma Fe$. The wet-chemistry 50 51 colorimetric method of Wilson (1960) has been used as the most popular bulk analytical method 52 providing a high accuracy (e.g., Schuessler et al., 2008). For the purpose of non-destructive 53 and/or local high-resolution analysis, several *in-situ* techniques have been developed, such as 54 micro-Mössbauer spectroscopy (McCammon, 1991; Potapkin et al., 2012), X-ray absorption near 55 edge structure (XANES) spectroscopy (Wilke, 2002), electron energy loss spectroscopy (EELS) 56 (van Aken et al., 1998; van Aken and Liebscher, 2002) and micro-Raman spectroscopy (Di Muro 57 et al., 2009). Electron probe micro-analysis (EPMA) has also been utilized to determine the $Fe^{2+}/\Sigma Fe$ ratio in geological samples, such as iron oxides (Höfer et al., 2000), garnets (Höfer and 58 59 Brey, 2007), olivines (Ejima et al., 2011), amphiboles (Enders et al., 2000; Lamb et al., 2012) 60 and silicate glasses (Fialin et al., 2001; 2004; 2011). Despite methodological challenges observed 61 so far, such as low sensitivity in some analytical protocols and lack of standard materials, the 62 easy access and low costs of EPMA compared to other methods keep it as a promising routine 63 method for measuring the oxidation state of iron in various geological samples including silicate 64 glasses.

In this paper, we present a new analytical technique for measuring the Fe oxidation state of silicate glasses with the EPMA flank method (Höfer and Brey, 2007). Our tests performed on a number of silicate glasses show that the method can provide determination of $Fe^{2+}/\Sigma Fe$ with an accuracy of ±0.1 for glasses containing 5–18 wt% FeO_T, and up to ±0.3 for glasses containing FeO_T ≤5 wt%.

70

THE FLANK METHOD

71 The application of EPMA to determine the oxidation state of Fe is based on the peak shift 72 and energy difference of the FeL α and FeL β emission lines for divalent and trivalent iron, which are induced by different electron energies of different bonding associated with Fe^{2+} and Fe^{3+} and 73 74 their different self-absorption (see details in Fischer, 1965; Tossell et al., 1974; Höfer et al., 1994). Changing from Fe^{2+} to Fe^{3+} , the FeLa and FeL β lines are both shifted to a higher energy, 75 76 and the intensity of L β peak is reduced preferentially to the L α peak (Höfer et al., 1994). To date, 77 two quantification techniques have been proposed: the "peak-shift method" and the "flank 78 method".

The *peak-shift method* utilizes the correlation between the peak positions of the FeLa line and Fe³⁺/ Σ Fe ratio (Kimura and Akasaka, 1999; Fialin et al., 2001; Fialin et al., 2004; Fialin et al., 2011). The peak-shift method requires accurate peak searches of the FeLa line for all materials under investigation (both standards and unknowns), and may have large uncertainties for samples with low total Fe. This method does not consider the changes in intensity between the FeLa and FeL β emission lines.

The *flank method* exploits both the peak shift and the intensity change of the FeL α and FeL β lines with ferric iron content by measuring the intensities at specific positions on the flanks of FeL α and FeL β peaks, respectively. Therefore, the flank method demonstrates higher sensitivity and better accuracy when compared to the peak-shift method (Höfer et al., 1994; Höfer and Brey, 2007). So far, no application of the EPMA flank method for silicate glasses has been reported in the literature.

For the flank method, the optimal FeL α flank and FeL β flank positions can be determined by the difference spectrum for a pair of materials with similar crystal structure and/or Fe coordination polyhedra but contrasting Fe oxidation states, such as wüstite-hematite (Höfer et al.,

1994; Höfer et al., 2000) and andradite-almandine (Figure 1; Höfer, 2003; Höfer and Brey, 94 2007). The andradite $(Ca^{2+}{}_{3}Fe^{3+}{}_{2}Si_{3}O_{12})$ and almandine $(Fe^{2+}{}_{3}Al^{3+}{}_{2}Si_{3}O_{12})$ used by Höfer (2003) 95 and Höfer and Brey (2007) are synthetic garnet endmembers containing Fe^{3+} and Fe^{2+} . 96 97 respectively. As shown by Höfer and Brey (2007), the flank positions determined by this 98 "mineral-difference method" are consistent with self-absorption spectra calculated from X-ray 99 emission spectra at different accelerating voltages. The ratio of intensities measured at the FeL β and FeLa flank positions, expressed in this paper as $L\beta/L\alpha$, is a function of Fe²⁺ content. After 100 some earlier attempts to correlate $L\beta/L\alpha$ with Fe³⁺/ Σ Fe or Fe³⁺ content with the flank method 101 (Höfer et al., 1994; Enders et al., 2000; Höfer, 2002), the unambiguous and accurate 102 quantification of $\text{Fe}^{3+}/\Sigma\text{Fe}$ in garnet was demonstrated by Höfer and Brey (2007). They also 103 104 found that for different mineral groups (e.g., garnet, olivine, spinel, wüstite, etc), the slopes of the regression lines of $L\beta/L\alpha$ versus Fe^{2+} content may differ significantly, implying that the 105 correlation between $L\beta/L\alpha$ and Fe^{2+} content might be a function of coordination number of Fe^{2+} . 106 107 Therefore, to achieve the high precision and accuracy of measured Fe oxidation state as it is now achieved in garnets (Höfer and Brey, 2007), it is necessary to calibrate the flank method for each 108 mineral group and glass, i.e. for each crystal or non-crystal structure. While Fe^{2+} in garnet is 8-109 fold coordinated, the coordination number of Fe^{2+} in silicate glasses is variable (4, 5 or 6, see 110 111 Wilke et al., 2007). This difference in coordination between garnet and silicate glass needs to be 112 examined before using garnets as standard materials for determining the Fe oxidation state of 113 silicate glasses. As indicated by our tests (see below), the two garnet references (almandine and and radite) and a number of silicate glasses show a consistent correlation between $L\beta/L\alpha$ and Fe²⁺ 114 115 content. Therefore, we propose that garnets can be used as standards for measuring the Fe 116 oxidation state of silicate glasses using the EPMA flank method.

117

SAMPLE SELECTION

Two end-member garnets (Höfer and Brey, 2007) with ferric (and radite, $Ca^{2+}{}_{3}Fe^{3+}{}_{2}Si_{3}O_{12}$) 118 and ferrous (almandine, $Fe^{2+}_{3}Al^{3+}_{2}Si_{3}O_{12}$) iron were used to calibrate the flank method in this 119 120 study. Forty-five silicate glasses belonging to five different glass groups were measured (Table 121 1), including twenty-nine Na- and K-free synthetic glasses (Borisov et al., 2004; Borisov, 2007; 122 Borisov et al., 2015), seven synthetic alkali-bearing glasses (ferrobasalts and basaltic andesite), 123 two synthetic hydrous glasses, four re-melted natural glasses (MORB and basanite), and three 124 natural basaltic glass references from the Smithsonian Microbeam Standards collection (USNM 125 111240/52 VG-2, USNM 113498/1 VG-A99 and USNM 113716, Jarosewich et al., 1980). The 126 synthesized or re-melted glasses were treated experimentally under controlled oxygen fugacity 127 (see Table 1 for experimental conditions), and these glasses cover a wide range of FeO_T content (4–18 wt%) and Fe²⁺/ Σ Fe ratio (0.2–1.0). 128

129

WET CHEMISTRY ANALYSIS OF FE OXIDATION STATE

130 The oxidation state of Fe in all selected experimental glasses has been analyzed using a 131 wet chemistry technique based on the colorimetric method of Wilson (1960) that was modified 132 following the procedure given by Schuessler et al. (2008). The sample powders were first placed 133 in an ammonium vanadate solution, which was then mixed with sulfuric acid. With additional 134 HF, the mixed solution was sealed and kept overnight at room temperature. In this technique Fe^{2+} is oxidized to Fe^{3+} due to the simultaneous reduction of V^{5+} to V^{4+} . Afterwards, the excess 135 136 HF in the solution was neutralized by adding saturated boric acid solution. The resultant solution 137 was then mixed with quantified ammonium acetate solution, 2:2' bipyridyl solution and distilled 138 water. The pH value in the solution was adjusted to ~ 5 as buffered by ammonium acetate. The complex of Fe^{2+} with 2:2' bipyridyl shows an intensive absorption band at ~523 nm, which 139

allows quantification of Fe²⁺ by UV spectrometer. We used a Shimadzu UV-1800 spectrometer 140 on the same solution to measure Fe^{2+} and total Fe before and after adding hydroxylamine 141 hydrochloride solution (this reducing agent forces total Fe as Fe^{2+}). This method ensures that the 142 uncertainty in measured Fe²⁺/ Σ Fe is exclusively sourced from the spectrometric measurement but 143 144 not related to weighing and dilution errors. An in-house standard andesite PU-3 (with known $Fe^{2+}/\Sigma Fe = 0.39 \pm 0.03$; Schuessler et al., 2008) and USGS basaltic standard BHVO-1 ($Fe^{2+}/\Sigma Fe$ 145 146 = 0.77 ± 0.03) were measured over all analytical sessions, and the results were identical within the 147 error.

148

ELECTRON PROBE MICRO-ANALYSIS

149 Measurements of Fe oxidation state in silicate glasses using the flank method as well as 150 the major element analyses have been performed with a Cameca SX100 electron microprobe 151 equipped with five spectrometers and "PeakSight" operation software at the Institute of 152 Mineralogy, Leibniz Universität Hannover, Germany. All standards and samples were coated with a thin carbon layer with a thickness of ca. 200 Å. The major elements (including total Fe as 153 154 FeO_T) were measured using calibration standards of synthetic oxides (Al₂O₃, Fe₂O₃, Mn₃O₄, 155 MgO and TiO₂), natural wollastonite (for Si and Ca), orthoclase (for K), jadeiite (for Na) and 156 fluorapatite (for P). The quantifications of all major elements were based on $K\alpha$ intensities, and 157 raw data were corrected using the standard PAP procedure (Pouchou and Pichoir, 1991). The 158 accelerating voltage was set at 15 kV for measuring both the major elements and $L\beta/L\alpha$, as 159 recommended by Höfer and Brey (2007). For alkali-free glasses, major elements were measured 160 with a focused 15 nA beam (Borisov et al. 2004; Borisov 2007; Borisov et al. 2015). For alkali-161 bearing glasses, we used a defocused beam (10 µm diameter) and a lower current (10 nA) to 162 minimize the loss of alkalis during electron bombardment of the sample surface. At least ten

points were measured on fresh surface of each sample (i.e., locations where not previouslybombarded) to obtain averages and standard deviations for elemental analyses.

165 For flank method measurements, we first collected FeL α and FeL β spectra in garnets. The 166 settings of the TAP spectrometer were optimized to measure the FeL lines according to the 167 recipe given by Höfer and Brey (2007). This includes the optimization of the pulse-height 168 analysis (PHA) setting for the FeL α line and the use of the "differential mode" for the X-ray counter. The differential mode was used to diminish high-energy X-ray lines (such as the 9th 169 170 order of FeK α X-Ray emission line) that are common when using the integral mode (Figure 1a). 171 A beam current of 200 nA and 10 µm diameter was used to increase the intensity of the signal, 172 and the sample stage was moved during analysis to diminish beam damage (see below).

173 As the *first step* of the method, the optimal positions of FeLa and FeL β flanks were determined by collecting FeL X-ray emission spectra of andradite and almandine. Figure 1b 174 175 shows the results indicating that the relative positions and intensities of the FeL α and FeL β peaks are displaced for both Fe^{2+} and Fe^{3+} endmembers. Before subtracting the spectra to obtain the 176 177 difference spectrum as described in Höfer and Brey (2007), we normalized the spectra to equal 178 total Fe concentration (i.e., spectra intensity divided by mineral total Fe content) to compensate 179 for the difference in bulk Fe contents between andradite and almandine. The resulting difference 180 spectrum demonstrates minima and maxima (Figure 1c). The most prominent minimum and 181 maximum have been selected for the FeL β and FeL α flank positions, respectively (vertical lines 182 in Figure 1c). The above difference spectra calculation has been measured with a relatively short 183 acquisition time (1000 points, 5 accumulations, 100 ms dwell time). Therefore, the data points of 184 the difference spectrum are scattered resulting in poorly defined flank positions.

185 In a *second step*, for achieving a better accuracy in defining the flank positions, we 186 performed a flank position adjustment by measuring intensities along a shorter spectral range (from -60 to +60 $10^5 \times \sin\theta$ relative to the approximate flank positions determined in the first step) 187 with a longer acquisition time (120 s). Figure 2 demonstrates that this procedure allows one to 188 189 specify a peak position based on a more smoothed spectral pattern compared to the raw spectral 190 scan data. In addition, this second-step adjustment shows that the new re-constrained 191 minimum/maximum positions can be different from the approximate flank positions determined 192 in the first step. As demonstrated by the tests on garnets of Höfer et al. (2000), slight changes in 193 spectrometer position for measuring positions at the flanks would introduce significant variations 194 in measured ratio $L\beta/L\alpha$ between sessions. Due to a variety of potential factors (such as drift of 195 machine conditions, major changes in laboratory conditions, see Höfer and Brey, 2007), the 196 optimal flank positions vary between different analytical sessions, and therefore such flank 197 position adjustment must be performed for each session independently. In addition, to avoid 198 potential problems with oxidation-reduction induced by electron beam bombardment, standards 199 (garnets and glasses) need to be re-polished and carbon-coated before each session (see below).

200 We acquired the spectral intensities of FeLa and FeLB at the re-constrained flank 201 positions for both garnet standards and unknown silicate glasses using a beam current of 200 nA 202 and a counting time of 120 s. This high beam current immediately poses the question whether 203 beam damage is significant. Beam damage is well known to be a problem for analyzing alkali-204 bearing glasses (Morgan VI and London, 1996). For example, in several publications, Fialin and 205 co-authors thoroughly discussed the role of beam-induced Fe oxidation or/and reduction caused 206 by electromigration of alkalis during EPMA analysis (Fialin et al., 2004; Fialin et al., 2001; 207 Fialin and Wagner, 2012). Surprisingly, the same authors reported, "neither oxidation- nor

208 reduction-induced peak shifts" during measurements of dry and hydrous glasses utilized for the 209 calibration of their peak shift method (Fialin et al., 2011; operating conditions were: 15 kV 210 accelerating potential, 250 nA beam current, 20 µm beam diameter and counting time 240 s). 211 Using static sample stage (conventional analysis, when the same analytical spot is exposed to the 212 beam for the whole acquisition time), we applied the Fialin's et al. (2011) protocol of peak-shift 213 method to our set of experimental glasses, however we failed to observe a robust correlation 214 between the shift of FeLa peak position and Fe oxidation state, which suggests that the 215 calibration of Fialin et al. (2011) should be revised on a more extensive dataset (see 216 Supplementary Figure 1). Thus, our first test measurements clearly demonstrated that beam-217 induced oxidation/reduction needs to be seriously considered. In this study, to minimize the 218 beam damage, we suggest moving the sample stage with a rate of 2 µm/s during acquisition (see also discussion below). Three independent measurements on different areas ($\sim 240*10 \ \mu m^2$) have 219 220 been performed for each sample. To check the reproducibility between sessions, analyses of a 221 few samples were replicated during three different analytical sessions (Table 2) with a time gap 222 of approximately one month. The ratio of intensities, $L\beta/L\alpha$, measured at the FeL β and FeL α 223 flank positions was then calculated and used for quantifying the Fe oxidation state.

224

RESULTS

To test the flank method described above, we have measured $L\beta/L\alpha$ of the garnet standards and the five silicate glass groups with known Fe oxidation state (**Table 2**). As shown in **Figure 3**, the values of $L\beta/L\alpha$ and Fe²⁺ content vary linearly in all sessions. Moreover, all measured glasses lie closely on the trends defined by the garnet standards, indicating that wellcharacterized garnet endmembers can be used as calibrating standards to quantify the Fe

230 oxidation state in silicate glasses despite their different coordination of iron cation. The linear231 relations defined by the garnet standards for the three independent sessions are:

- 232 Fe^{2+} (wt%) = 34.20×*L* β /*L* α 14.63 (Session 1)
- 233 $Fe^{2+} (wt\%) = 33.47 \times L\beta/L\alpha 13.88 (Session 2)$
- 234 $Fe^{2+}(wt\%) = 31.14 \times L\beta/L\alpha 13.66$ (Session 3)

Using these relations and FeO_T concentrations in the glasses, the Fe^{2+} contents and 235 corresponding $Fe^{2+}/\Sigma Fe$ ratios can be calculated (Table 2). Figures 4a, 4c and 4e show that the 236 Fe²⁺ contents determined by the flank method are consistent within error with those determined 237 238 by wet chemistry in most cases, and the differences are in general less than 1 wt% for all silicate glasses. Figures 4b, 4d and 4f show that the $Fe^{2+}/\Sigma Fe$ ratios determined by the flank method are 239 240 consistent within a value of ± 0.1 with those determined by wet chemistry for samples with high 241 FeO_T contents (>5 wt%), whereas the ratio difference increase to 0.2–0.3 for samples with lower FeO_T contents. This implies that the error of the Fe²⁺/ Σ Fe determined by the flank method is 242 dominantly associated with the intensity measured at the FeL flanks; i.e. the lower the FeO_T 243 244 content, the lower the accuracy of the analysis.

245

DISCUSSION

Potential errors of determining $Fe^{2+}/\Sigma Fe$ ratio of glasses using the EPMA flank method can be related to compositional effects, which denotes self-absorption of FeL lines by Fe and variable absorption of FeL lines by other cations (Höfer et al., 1994; Fialin et al., 2001). In order to investigate potential compositional effects, we plotted the difference between Fe²⁺ measured by EPMA and wet chemistry (ΔFe^{2+}) against total Fe content in Figure 5. The data do not show any apparent correlation between the measured Fe²⁺ and FeO_T contents, therefore no systematic discrepancy between glasses with contrasting FeO_T contents. This implies that our method of

Fe²⁺ determination with the above linear equations is robust and total Fe has little effect on this calibration. As shown by Höfer and Brey (2007), $L\beta/L\alpha$ does depend on total Fe, but this effect can be split into the dependence on Fe²⁺ and Fe³⁺ (due to different self-absorption) and can be approximated by a simple linear equation at low total Fe cases (e.g. FeO_T <20 wt%).

257 We explored the potential effect of Ti on the flank method on differentially absorbing 258 FeLa and FeL β within silicate glasses, we explored it within the range of TiO₂ content between 0 259 and 25 wt%. As listed in Table 2, for silicate glasses with TiO₂ contents lower than 15 wt%, no systematic correlation is observed between ΔFe^{2+} and TiO₂ content. However, sample DAT32 260 with extremely high TiO₂ (25.09 wt%) demonstrates high ΔFe^{2+} (Figure 5b), suggesting that Ti 261 262 is indeed able to influence the absorption of FeL α and/or FeL β , but only for silicate glasses with 263 very high TiO₂ contents (at least >15 wt%). Although Fialin et al. (2001) emphasized the 264 potential effect of Cr and Ti on Fe L line emission and absorption, this problem is perhaps only 265 crucial for Cr- and/or Ti-rich phases (e.g. chromite and ilmenite). The absorption effect of Ti 266 should be extremely weak in silicate glasses with low Ti contents, as demonstrated by the data of 267 Fialin et al. (2004) involving silicate glasses with 0–1.8 wt% TiO₂. This assumption is supported by our results. 268

Potential matrix effects of other elements such as Si, Al, Ca and Mg on the flank method for glasses were not observed in this study, which is consistent with the observations of Höfer and Brey (2007) on garnets. The dataset of silicate glass in this study covers a relatively wide compositional range (**Table 1**), in terms of SiO₂ (40–56 wt%), Al₂O₃ (10–18 wt%), CaO (9–23 wt%) and MgO (4–10 wt%), and no systematic influence of these major oxides on $L\beta/L\alpha$ in the range of FeO_T (4–18 wt%) was observed. To conclude, our measurements demonstrate that Fe²⁺ in silicate glasses can be calculated from $L\beta/L\alpha$ based on the quantitative relation calibrated

against Fe^{2+} -rich and Fe^{3+} -rich garnet endmembers, and there is no significant matrix effect of other cations, except for Ti, if it is present in very high abundances.

278 Applying the peak-shift method, Fialin et al. (2004) observed both apparent oxidation and 279 reduction trends with accumulated analytical time (at a 15 kV accelerating voltage, 240 nA beam 280 current and 20 µm beam diameter). The observed variation of measured Fe oxidation state with 281 time was attributed by Fialin et al. (2004) to two factors, including (1) Na migration and 282 consequent rearrangement of oxygen atoms between bridging and non-bridging positions in the 283 close vicinity of electron beam bombardment, and (2) buildup of carbon contamination. In this 284 study, we performed additional tests on the anhydrous natural glass reference VG-2 (0.02 wt% 285 H₂O; Figure 6) and on basaltic glasses with 0, 2.8 and 5.0 wt% H₂O (Figure 7) to illustrate the 286 potential beam damage at 200 nA beam current and associated effects on measurements of $L\beta/L\alpha$, 287 in two contrasting cases with a static sample stage (points) and with a moving stage (horizontal 288 dashed line).

289 As shown in the left-side panels of Figure 6, the values of $L\beta/L\alpha$ measured with a beam 290 diameter from 5 to 20 µm show different behavior with time for the anhydrous basaltic glass 291 VG-2, with the sample stage being static for each measurement. The 5 μ m beam induces an 292 overall decrease of $L\beta/L\alpha$ after 250 s, which likely indicates oxidation of the analytical volume 293 due to electron beam bombardment. In contrast, the use of a 10 µm or 20 µm beam tends to 294 increase slightly (or does not modify) the measured ratio of $L\beta/L\alpha$ during the first 250 s. The 295 right-side panels in Figure 6 show variations of the intensities of FeKa and NaKa for the beam 296 diameters of 5, 10 and 20 µm, which can provide information on interpreting the variations of 297 $L\beta/L\alpha$. One striking observation is that the loss of the NaKa intensity occurs for all beam sizes, 298 and it is enhanced with decreasing beam size, consistent with previous studies (e.g. Morgan VI

299 and London, 1996; Fialin et al., 2004). In addition, we show that the FeK α intensity slightly 300 increases when the 5 µm and 10 µm beams were used, and it remains almost constant with the 20 301 μm beam, demonstrating the tendency of increasing relative Fe content in glass with increase of 302 beam current (probably due to alkali loss and changes of glass density), similar to what has been 303 also shown for SiKa and AlKa (Morgan VI and London, 1996; Zhang et al., 2016). Provided that $L\beta/L\alpha$ is positively correlated with Fe²⁺ content, it seems that severe beam damage (both Na 304 intensity loss and Fe intensity increase) with a small beam size (i.e. 5 µm) tends to oxidize the 305 analyzed glass volume (decreasing $Fe^{2+}/\Sigma Fe$ ratio), whereas weak beam damage (only slight Na 306 intensity loss and no Fe intensity increase) with a large beam size (i.e. 20 µm) tends to reduce (or 307 not modify) the analyzed glass volume (increasing $Fe^{2+}/\Sigma Fe$ ratio). 308

309 It is well known that the migration of Na during EPMA (i.e. loss of Na intensity) is much stronger in hydrous glasses than in dry glasses, even if the beam current is as low as 2-5 nA 310 311 (Morgan VI and London, 1996). On the other hand, water as a chemical component has almost 312 negligible effect on the ferric/ferrous ratio of silicate glasses (Botcharnikov et al., 2005). In this 313 study, we conducted a test of beam damage as a function of time on three glasses with similar 314 major element compositions but different H₂O contents (nominally dry, 2.8 and 5.0 wt% H₂O, 315 **Table 1**). The $L\beta/L\alpha$ values and FeK α and NaK α intensities have been acquired at 200 nA and 20 316 µm beam diameter over 500 s with the sample stage being static, which are compared to the 317 values obtained whilst moving the sample stage. The left-side panels of Figure 7 show the 318 variation of $L\beta/L\alpha$, and the right-side panels show the variation of intensities of FeK α and NaK α . 319 With an increase of H₂O content, the loss of Na intensity is dramatically enhanced and FeK α 320 intensity tends to increase. The value of $L\beta/L\alpha$ increases slightly during the first 100 s on 321 nominally dry glass N72, consistent with the results obtained on the VG-2 sample measured with

322 a 20 μ m diameter beam (Figure 6). In contrast, in H₂O-bearing glasses the $L\beta/L\alpha$ value decreases significantly within the same time period, indicating a decrease of Fe^{2+} content in spite 323 324 of increasing relative total Fe content in the glass as inferred from increasing FeK α intensity. As 325 shown in Figure 8, the strong decreases in NaK α intensity and $L\beta/L\alpha$ are roughly coupled for the 326 hydrous glasses, supporting the hypothesis that the migration of Na during EPMA might promote oxidization of Fe^{2+} converted to Fe^{3+} (Fialin et al., 2004). Therefore, in comparison to dry glasses, 327 328 the analyzed volume of hydrous glass is much more prone to be oxidized during EPMA as a 329 result of beam damage.

330 Besides the potential effect of Na-migration on the EPMA measurement of Fe $L\beta/L\alpha$ ratio of silicate glasses discussed above, carbon contamination or loss on C-coated sample surface 331 332 could also play a significant role. Gopon et al. (2013) showed that carbon contamination is a 333 serious problem affecting the measured stabilities of FeLa and FeLB of Fe-Si compounds, 334 especially in cases where a static high-current beam is used. Fialin et al. (2004) found buildup of 335 carbon contamination on silicate glass to be significant when measurements were performed with 336 a 240 nA beam current (20 µm diameter) on the same spot. They suggested that it might have partly resulted in the decrease of measured $Fe^{3+}/\Sigma Fe$ ratio using their peak-shift method, at least 337 338 for the initial stage of measurement time. Höfer and Brey (2007) made a similar test on an almandine sample with a 60 nA beam scanning an area of $3 \times 5 \mu m^2$, and demonstrated that 339 340 carbon contamination resulted in decrease in $L\beta/L\alpha$ ratio measured by their flank method and in overestimation of $Fe^{3+}/\Sigma Fe$ ratio. Interestingly, the effect of carbon contamination on measuring 341 342 the Fe oxidation state, observed by Fialin et al. (2004) for silicate glass and by Höfer and Brey 343 (2007) for garnet are contradicting with each other. We tested carbon contamination by 344 measuring the carbon Ka intensity on the VG-2 glass with static and moving sample stage

345 respectively. As shown in Figure 9, the CKa intensity measured on the same spot (i.e., with 346 static sample stage) decreases strongly and continuously with accumulated time up to 400 s, 347 whereas the measurements with moving sample stage demonstrate constant intensity. The 348 observed decrease of CKa intensity during beam bombardment is contradicting with Fialin et al. 349 (2004) but consistent with that observed by Gopon et al. (2013). Fialin et al. (2004) observed a 350 continuous increase of CKa intensity on a silicate glass for 15 min. However, Gopon et al. (2013) 351 made tests on carbon-coated FeSi compounds with a low-voltage high-current beam (5 kV, 100 352 nA) and found CKa intensity was firstly strongly lost in the initial 400 s but gained later on with 353 accumulated time up to 4000 s. Therefore, the effect of carbon contamination or loss seems to be 354 complicated and probably dependent on a number of factors, such as material composition, beam 355 current, time, etc. In any case, for applying the flank method described in this paper, carbon 356 contamination and loss should be avoided in order to measure glass Fe $L\beta/L\alpha$ ratios, and moving 357 sample stage is demonstrated to be a good approach.

358 Based on these results, we conclude that, if the EPMA measurements are carried out at 359 the same position for a long time on glasses, the variation of $L\beta/L\alpha$ is a consequence of the 360 combined effects of the changes in both total Fe content and Fe oxidation state of glass, 361 reflecting accumulated material damage induced by electron beam bombardment. Our tests 362 conducted with a static stage demonstrate that the values of $L\beta/L\alpha$ cannot be accurately resolved 363 for dry or hydrous glasses if a high beam current and a long acquisition time are applied. 364 However, our results show that a high accuracy in the determination of the $L\beta/L\alpha$ (and thus $Fe^{2+}/\Sigma Fe$ ratio) can be achieved when analyses are conducted with a continuously moving 365 366 sample stage (e.g. $2 \mu m/s$) during data acquisition.

367

IMPLICATIONS

When the beam damage problem is successfully resolved (e.g. by movement of the sample stage in this study), the EPMA flank method provides a promising low-cost and very simple alternative to other local non-destructive techniques, such as XANES, micro-Mössbauer spectroscopy, EELS and micro-Raman spectroscopy (see Introduction for the references). In this study, the accuracy of the Fe²⁺/ Σ Fe determination is found to be dependent both on the Fe²⁺ content determined with the flank method and on the total Fe content, and is generally within ±0.1 for silicate glasses with FeO_T >5 wt%.

In petrology, accurately determined $Fe^{2+}/\Sigma Fe$ ratio in natural glasses serves as a proxy of 375 376 the redox conditions (fO₂) prevailing in magmatic chambers (Christie et al., 1986; Bézos and 377 Humler, 2005; Cottrell and Kelley, 2011). For example, the most recent data obtained by Cottrell and Kelley (2011) by XANES for naturally-quenched pillow-rim glasses suggest that global 378 MORB Fe²⁺/ Σ Fe has a value of 0.84±0.01 (1 σ) corresponding to the favalite–magnetite–quartz 379 380 (FMQ) buffer under conditions of primary magma generation. Assuming 1 wt.% error (2σ) in the determination of Fe^{2+} by the flank EPMA method and an ideal slope of $\frac{1}{4}$ for the dependence 381 between log (Fe³⁺/Fe²⁺) and log fO_2 we provide propagated errors in the determination of fO_2 for 382 the range of $Fe^{2+}/\Sigma Fe$ values typical for natural melts (Table 3). Note, however, that additional 383 384 errors may result from the application of empirical models describing the dependence of 385 ferric/ferrous ratios on temperature, oxygen fugacity and melt composition, and also from 386 differences between real pre-eruptive temperatures of basaltic melts and the temperature of 387 1200°C typically assumed for Δ FMQ calculations (see discussion in Borisov et al. 2013). As one can see, the translated uncertainties in estimation of the oxygen fugacity for typical MORB 388 $(Fe^{2+}/\Sigma Fe \sim 0.85; FeO_T \sim 9 \text{ wt\%})$ range within only ±0.12 log units (2 σ). This high precision in 389 determination of the Fe²⁺/ Σ Fe and in turn fO_2 by the flank EPMA method also provides a new 390

6437R: Revision	1	(22.	03	.201	8)
-----------------	---	------	----	------	----

391	promising analytical tool for future experimental studies under high pressures, where controlling
392	and logging the redox conditions is usually a challenging task.
393	
394	ACKNOWLEDGMENTS
395	We thank Marius Stranghöner and Florian Pohl for the help with wet-chemistry
396	colorimetric method and Julian Feige for sample preparation. We appreciate detailed and
397	insightful comments from John Fournelle and two anonymous reviewers. RA and JK also thank
398	Anette von der Handt and Eric Hellebrandt for support of the project. This study was funded by
399	German Research Foundation (DFG project AL 1189/6-1).
400	R eferences cited
401 402 403	Almeev, R.R, Holtz, F., Ariskin, A., and Kimura, JI. (2013) Storage conditions of Bezymianny Volcano parental magmas: results of phase equilibria experiments at 100 and 700 MPa. Contributions to Mineralogy and Petrology, 166(5), 1389-1414.
404 405	Almeev, R.R, Holtz, F., Koepke, J., Parat, F., and Botcharnikov, R.E. (2007) The effect of H ₂ O on olivine crystallization in MORB: Experimental calibration at 200 MPa. American Mineralogist, 92(4), 670-674.
406 407	Bézos, A. and Humler, E. (2005) The Fe ³⁺ /ΣFe ratios of MORB glasses and their implications for mantle melting. Geochimica et Cosmochimica Acta 69 (3), 711-725.
408 409	Borisov, A. A. and Shapkin, A. I. (1990) A new empirical equation rating Fe ³⁺ /Fe ²⁺ in magmas to their composition, oxygen fugacity and temperature. Geochemistry International 27, 111–116.
410 411	Borisov, A., Behrens, H., and Holtz, F. (2013) The effect of titanium and phosphorus on ferric/ferrous ratio in silicate melts: an experimental study. Contributions to Mineralogy and Petrology, 166(6), 1577-1591.
412 413 414	Borisov, A., Behrens, H., and Holtz, F. (2015) Effects of melt composition on Fe ³⁺ /Fe ²⁺ in silicate melts: a step to model ferric/ferrous ratio in multicomponent systems. Contributions to Mineralogy and Petrology, 169(2), 1-12.
415 416	Borisov, A., Lahaye, Y., and Palme, H. (2004) The effect of TiO ₂ on Pd, Ni, and Fe solubilities in silicate melts. American Mineralogist, 89(4), 564-571.

- Borisov, A.A. (2007) Experimental study of the influence of SiO₂ on the solubility of cobalt and iron in silicate
 melts. Petrology, 15(6), 523-529.
- Botcharnikov, R.E., Almeev, R.R., Koepke, J., and Holtz, F. (2008) Phase relations and liquid lines of descent in
 hydrous ferrobasalt implications for the Skaergaard Intrusion and Columbia River Flood Basalts. Journal of
 Petrology, 49(9), 1687-1727.
- Botcharnikov, R.E., Koepke, J., Holtz, F., McCammon, C., and Wilke, M. (2005) The effect of water activity on the
 oxidation and structural state of Fe in a ferro-basaltic melt. Geochimica et Cosmochimica Acta, 69(21), 50715085.
- 425 Christie, D. M., Carmichael, I. S. E. and Langmuir, C. H. (1986) Oxidation states of mid-ocean ridge basalt glasses.
 426 Earth and Planetary Science Letters 79 (3-4), 397-411.
- 427 Cottrell, E. and K. A. Kelley (2011) The oxidation state of Fe in MORB glasses and the oxygen fugacity of the
 428 upper mantle. Earth and Planetary Science Letters 305(3-4): 270-282.
- Di Muro, A., Métrich, N., Mercier, M., Giordano, D., Massare, D., and Montagnac, G. (2009) Micro-Raman
 determination of iron redox state in dry natural glasses: Application to peralkaline rhyolites and basalts.
 Chemical Geology, 259(1), 78-88.
- 432 Ejima, T., Akasaka, M., and Ohfuji, H. (2011) Oxidation state of Fe in olivine in a lherzolite xenolith from Oku
 433 district, Oki-Dogo Island, Shimane Prefecture, Japan. Journal of Mineralogical and Petrological Sciences,
 434 106(5), 246-254.
- Enders, M., Speer, D., Maresch, W.V., and McCammon, C.A. (2000) Ferric/ferrous iron ratios in sodic amphiboles:
 Mossbauer analysis, stoichiometry-based model calculations and the high-resolution microanalytical flank
 method. Contributions to Mineralogy and Petrology, 140(2), 135-147.
- Fialin, M., Bézos, A., Wagner, C., Magnien, V., and Humler, E. (2004) Quantitative electron microprobe analysis of Fe³⁺/ Σ Fe: Basic concepts and experimental protocol for glasses. American Mineralogist, 89(4), 654-662.
- Fialin, M., Wagner, C., Métrich, N., Humler, E., Galoisy, L., and Bézos, A. (2001) Fe³⁺/Σ Fe vs. FeLα peak energy
 for minerals and glasses: Recent advances with the electron microprobe. American Mineralogist, 86(4), 456442 465.
- Fialin, M., Wagner, C., and Pascal, M.-L. (2011) Iron speciation using electron microprobe techniques: application
 to glassy melt pockets within a spinel lherzolite xenolith. Mineralogical Magazine, 75(2), 347-362.
- Fialin, M. & Wagner, C. (2012) Redox kinetics of iron in alkali silicate glasses exposed to ionizing beams:
 Examples with the electron microprobe. Journal of Non-Crystalline Solids, 358 (12–13), 1617-1623.

- Fischer, D.W. (1965) Changes in the Soft X-Ray L Emission Spectra with Oxidation of the First Series Transition
 Metals. Journal of Applied Physics, 36(6), 2048-2053.
- Fuchs, P., Almeev, R.R., and Klügel, A. (2014) Experimental constraints on the formation of basanites-phonolite
 series (Cumbre Vieja, La Palma). Goldschmidt Abstracts, 2014 744
- Gopon, P., Fournelle, J., Sobol, P.E., and Llovet, X. (2013) Low-voltage electron-probe microanalysis of Fe–Si
 compounds using soft X-rays. Microscopy and Microanalysis, 19(6), 1698-1708.
- Husen, A., Almeev, R.R., and Holtz, F. (2016) The Effect of H₂O and Pressure on Multiple Saturation and Liquid
 Lines of Descent in Basalt from the Shatsky Rise. Journal of Petrology, 57(2), 309-344.
- Höfer, H.E. (2002) Quantification of Fe²⁺/Fe³⁺ by Electron Microprobe Analysis New Developments. In P.
 Gütlich, B.W. Fitzsimmons, R. Rüffer, and H. Spiering, Eds. Mössbauer Spectroscopy: Proceedings of the
 Fifth Seeheim Workshop, held in Seeheim, Germany, 21–25 May 2002, p. 239-248. Springer Netherlands,
 Dordrecht.
- Höfer, H.E., and Brey, G.P. (2007) The iron oxidation state of garnet by electron microprobe: Its determination with
 the flank method combined with major-element analysis. American Mineralogist, 92(5-6), 873-885.
- 461 Höfer, H.E., Brey, G.P., Schulz-Dobrick, B., and Oberhaensli, R. (1994) The determination of the oxidation state of
 462 iron by the electron microprobe. European Journal of Mineralogy, 6(3), 407-418.
- 463 Höfer, H.E., Weinbruch, S., McCammon, C.A., and Brey, G.P. (2000) Comparison of two electron probe
 464 microanalysis techniques to determine ferric iron in synthetic wüstite samples. European Journal of
 465 Mineralogy, 12(1), 63-71.
- Jarosewich, E., Nelen, J.A., and Norberg, J.A. (1980) Reference Samples for Electron Microprobe Analysis.
 Geostandards Newsletter, 4(1), 43-47.
- Kelley, K. A. and E. Cottrell (2009) Water and the Oxidation State of Subduction Zone Magmas. Science 325
 (5940): 605-607.
- Kilinc, A., Carmichael, I.S.E., Rivers, M.L., and Sack, R.O. (1983) The ferric-ferrous ratio of natural silicate liquids
 equilibrated in air. Contributions to Mineralogy and Petrology, 83(1-2), 136-140.
- 472 Kimura, Y., and Akasaka, M. (1999) Estimation of Fe^{2+}/Fe^{3+} and Mn^{2+}/Mn^{3+} ratios by Electron Probe Micro 473 Analyzer. Journal of the Mineralogical Society of Japan, 28, 159-166.
- Klügel, A., Galipp, K., Hoernle, K., Hauff, F., and Groom, S. (2017) Geochemical and Volcanological Evolution of
 La Palma, Canary Islands. Journal of Petrology, 58(6), 1227-1248.

- Kress, V., and Carmichael, I. (1991) The compressibility of silicate liquids containing Fe₂O₃ and the effect of
 composition, temperature, oxygen fugacity and pressure on their redox states. Contributions to Mineralogy
 and Petrology, 108(1), 82-92.
- 479 Lamb, W.M., Guillemette, R., Popp, R.K., Fritz, S.J., and Chmiel, G.J. (2012) Determination of Fe^{3+}/Σ Fe using the 480 electron microprobe: A calibration for amphiboles. American Mineralogist, 97(5-6), 951-961.
- 481 McCammon, C., Chaskar, V., and Richards, G. (1991) A technique for spatially resolved Mossbauer spectroscopy
 482 applied to quenched metallurgical slags. Measurement Science and Technology, 2(7), 657.
- 483 Moretti, R. (2005) Polymerisation, basicity, oxidation state and their role in ionic modelling of silicate melts. Annals
 484 of Geophysics 48 (4-5), 583-608.
- 485 Morgan VI, G.B., and London, D. (1996) Optimizing the electron microprobe analysis of hydrous alkali
 486 aluminosilicate glasses. American Mineralogist, 81, 1176-1185.
- Nikolaev, G. S., Borisov, A. A. & Ariskin, A. A. (1996) Calculation of the ferric-ferrous ratio in magmatic melts:
 Testing and additional calibration of empirical equations for various magmatic series. Geochemistry
 International 34, 641-649.
- 490 Potapkin, V., Chumakov, A.I., Smirnov, G. V., Celse, J.-P., Rüffer, R., McCammon, C., and Dubrovinsky, L. (2012)
 491 The ⁵⁷ Fe Synchrotron Mössbauer Source at the ESRF. Journal of Synchrotron Radiation. 19, 559–569.
- 492 Pouchou, J.L., and Pichoir, F. (1991) Quantitative analysis of homogeneous or stratified microvolumes applying the
 493 model "PAP". In K.F.J. Heinrich, and D.E. Newbury, Eds. Electron probe quantitation, p. 31-75. Plenum
 494 Press, New York.
- 495 Sack, R., Carmichael, I., Rivers, M., and Ghiorso, M. (1981) Ferric-ferrous equilibria in natural silicate liquids at 1
 496 bar. Contributions to Mineralogy and petrology, 75(4), 369-376.
- 497 Savitzky, A., and Golay, M.J.E. (1964) Smoothing and Differentiation of Data by Simplified Least Squares
 498 Procedures. Analytical Chemistry, 36(8), 1627-1639.
- Schuessler, J.A., Botcharnikov, R.E., Behrens, H., Misiti, V., and Freda, C. (2008) Oxidation state of iron in
 hydrous phono-tephritic melts. American Mineralogist, 93(10), 1493-1504.
- Shishkina, T., Botcharnikov, R., Holtz, F., Almeev, R. R. & Portnyagin, M. (2010) Solubility of H₂O and CO₂ bearing fluids in tholeiitic basalts at pressures up to 500 MPa. Chemical Geology 277 (1-2), 115-125.
- Tossell, J., Vaughan, D., and Johnson, K. (1974) Electronic-Structure of Rutile, Wustite, and Hematite from
 Molecular-Orbital Calculations. American Mineralogist, 59(3-4), 319-334.

- van Aken, P., Liebscher, B., and Styrsa, V. (1998) Quantitative determination of iron oxidation states in minerals
 using Fe L2,3-edge electron energy-loss near-edge structure spectroscopy. Physics and Chemistry of Minerals,
 25(5), 323-327.
- van Aken, P. A. and Liebscher, B. (2002) Quantification of ferrous/ferric ratios in minerals: new evaluation schemes
 of Fe L 23 electron energy-loss near-edge spectra. Physics and Chemistry of Minerals 29 (3), 188-200.
- Wilke, M., Behrens, H., Burkhard, D.J., and Rossano, S. (2002) The oxidation state of iron in silicic melt at 500
 MPa water pressure. Chemical geology, 189(1), 55-67.
- Wilke, M. (2005) Fe in magma An overview. Annals of Geophysics 48 (4-5), 609-617.; Moretti, R. (2005).
 Polymerisation, basicity, oxidation state and their role in ionic modelling of silicate melts. Annals of Geophysics 48 (4-5), 583-608.
- Wilke, M., Farges, F., Partzsch, G.M., Schmidt, C., and Behrens, H. (2007) Speciation of Fe in silicate glasses and
 melts by in-situ XANES spectroscopy. American Mineralogist, 92(1), 44-56.
- Wilson, A. (1960) The micro-determination of ferrous iron in silicate minerals by a volumetric and a colorimetric
 method. Analyst, 85(1016), 823-827.
- Zhang, C., Koepke, J., Wang, L.-X., Wolff, P.E., Wilke, S., Stechern, A., Almeev, R., and Holtz, F. (2016) A
 Practical Method for Accurate Measurement of Trace Level Fluorine in Mg- and Fe-Bearing Minerals and
 Glasses Using Electron Probe Microanalysis. Geostandards and Geoanalytical Research, 40(3), 351-363.
- 522

523 FIGURE CAPTIONS

Figure 1. FeL X-ray emission spectra of andradite $(Ca^{2+}_{3}Fe^{3+}_{2}Si_{3}O_{12})$ and almandine $(Fe^{2+}_{3}Al^{3+}_{2}Si_{3}O_{12})$ acquired at 524 525 15 kV, beam current 200 nA and beam size 10 µm. (a) Comparison of PHA integral mode and differential mode for 526 FeL X-ray emission spectra of andradite. Note that, for integral mode, there is a small peak between the major FeL α 527 and FeL β peaks, which is the 9th order of the high-energy FeK α X-Ray emission line. (b) Spectra of andradite and 528 almandine acquired in differential mode so the high-energy FeK α X-Ray emission line was diminished. Baseline = 529 1100 mV, window = 1300 mV, beam current 200 nA, beam size 10 μ m, dwell time 0.1 s, accumulation number 5. 530 (c) Difference spectrum (original spectra were acquired with differential mode) between andradite and almandine 531 normalized to equal total Fe content (light in color). Smoothed spectrum (dark in color) is obtained by the Savitzky-532 Golay method (Savitzky and Golay, 1964). The flank positions of FeL α and FeL β are found at the maximum and 533 minimum of the smoothed difference spectrum (marked by the vertical lines).

Figure 2. Re-constrained flank positions of FeLa (a) and FeL β (b) based on andradite and almandine. The initial flank positions (relative position = 0) are determined by smoothed difference spectra (see Figure 1c). At positions of 0, ±20, ±40 and ±60 relative to the initial flank positions, andradite and almandine were analyzed again with a longer counting time (120 s each, three repeated measurements) to obtain more accurate difference spectra (spots with ±1 σ deviation and dashed curve). For comparison, the short-time scan spectra are shown in light grey. See text for details.

540 Figure 3. Plots of $L\beta/L\alpha$ versus Fe²⁺ content for garnets and glasses. Data measured in Session 1 (a), Session 2 (b). 541 and Session 3 (c). Deviation of $\pm 1\sigma$ is smaller than symbol size. See details in Table 1.

Figure 4. Comparison of glass Fe^{2+} content and $Fe^{2+}/\Sigma Fe$ ratio determined by EPMA flank method and wet chemistry. Data measured in Session 1 (a-b), Session 2 (c-d). and Session 3 (e-f). The dashed line is ± 1 wt% in the left panels and ± 0.1 in the right panels.

Figure 5. Difference of Fe²⁺ content (i.e. Δ Fe²⁺) between EPMA flank method and wet chemistry plotted against FeO_T content (see data in Table 1). Data measured in Session 1 (a), Session 2 (b) and Session 3 (c) are plotted separately. Isopleths of induced Δ Fe²⁺/ Σ Fe are also noted in the left panels. The mean standard deviation of calculated Δ Fe²⁺ is ca. 0.4 (see inserted error bar). The standard deviation of FeO_T content is smaller than symbol size.

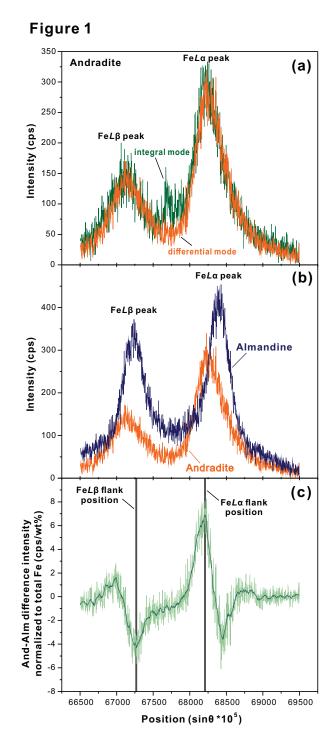
- **Figure 6.** Variation of $L\beta/L\alpha$ and count rates of FeK α and NaK α as a function of time measured with different beam size for reference glass VG-2. Beam setting is 15 kV and 200 nA for all cases. Dashed line indicates the mean value of $L\beta/L\alpha$ measured with moving sample stage, in which case beam damage is minimized.
- 553 Figure 7. Variation of $L\beta/L\alpha$ and count rates of FeK α and NaK α as a function time measured for three glasses (M72,
- 554 M6 an M11, Shishkina et al., 2010) with different H₂O contents but similar major element composition. Beam 555 current is 200 nA and beam size is 20 μ m diameter for all cases. Dashed lines indicate the mean value of *L* β /*L* α
- 556 measured with moving sample stage, in which case beam damage is minimized.
- **Figure 8.** Plots of $L\beta/L\alpha$ versus count rate of NaK α for glasses M72, M6 and M11. Beam current is 200 nA and beam size is 20 µm diameter for all cases. The values are for various times during time series measurements, with
- highest Na signal at the beginning and lowest Na signal at the ending point.

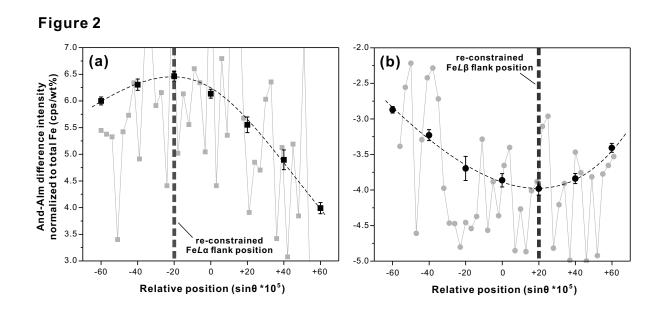
Figure 9. Count rate variations of $CK\alpha$ as a function of time measured acquired on VG-2 glass. Beam current and diameter are 200 nA and 20 μ m respectively. Note the contrasting variation trends obtained with moving and static sample stage respectively.

- 563 Supplementary Figure 1. Experimental glasses used in this study treated by the "peak-shift" calibration from Fialin 564 et al. (2011). Glass and olivine FeLa peak shift positions are scaled to that of hematite standard (measured as a 565 reference to avoid machine drift) and expressed as $\Delta \sin \theta$ versus total Fe concentration. Solid line represents the calibration for pure Fe²⁺, constructed from olivines. Diamonds are experimental glasses used in calibration of Fialin 566 567 et al. (2011); only the most reduced (filled diamonds) and most oxidized (open diamonds) are shown for clarity. 568 Colored dots are most reduced (green) and oxidized (red) glasses from this study (including some other from 569 Alexander Borisov's collection). Replicate analyses conducted for both oxidized and reduced glasses demonstrate 570 the poor reproducibility utilizing the static beam stage.
- 571
- 572

573 TABLE CAPTIONS

- 574 **Table 1.** Major element composition of garnets and silicate glasses (wt%)
- 575 **Table 2.** Fe oxidation state of garnets and silicate glasses
- 576 **Table 3.** Possible errors in fO_2 determination using the flank method for MORB glasses





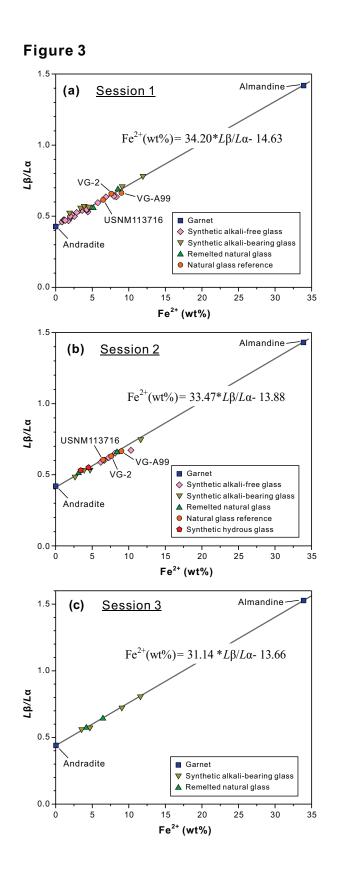
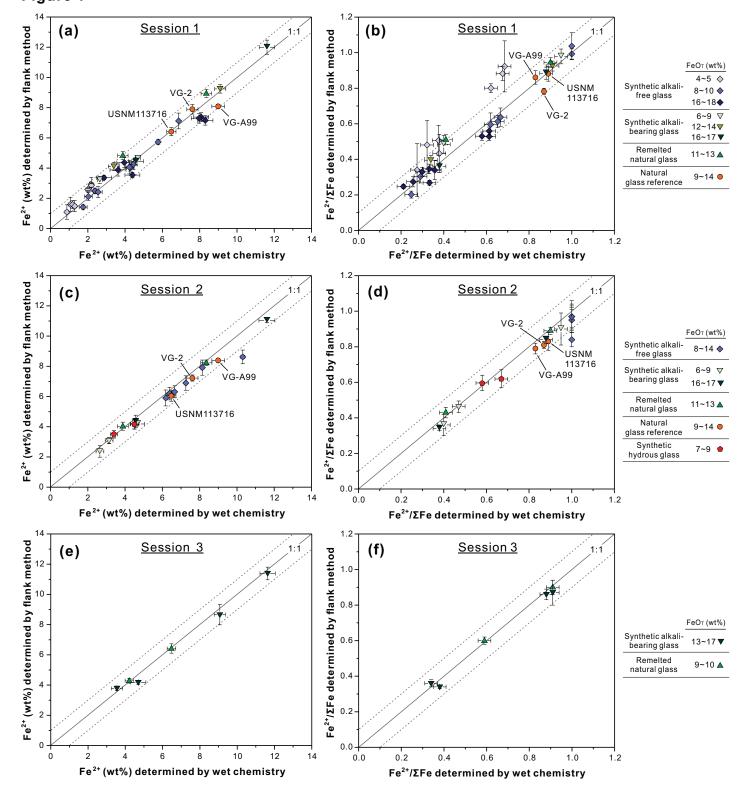


Figure 4



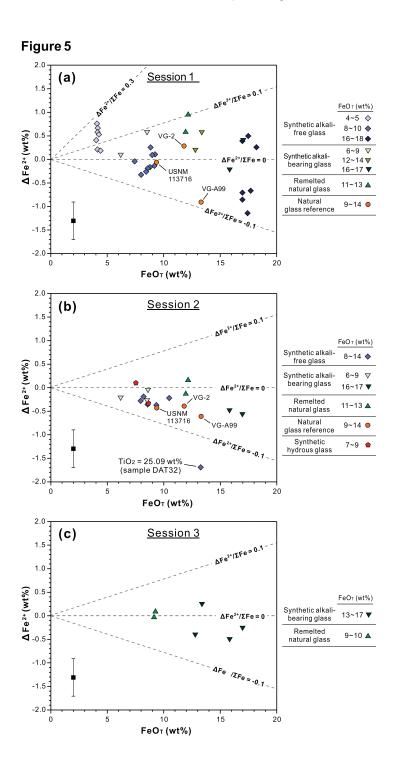
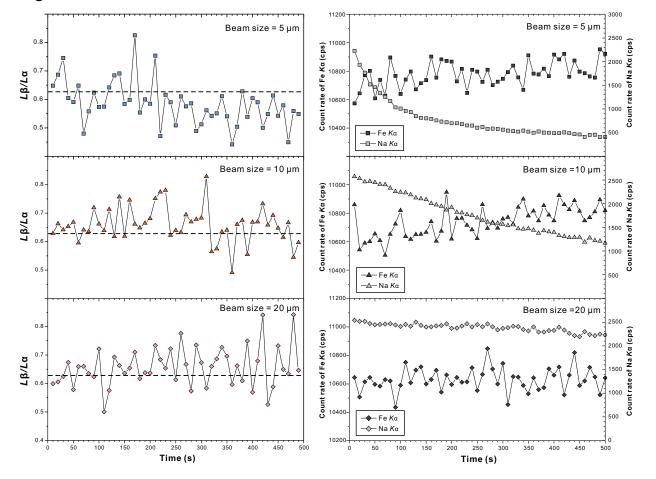


Figure 6



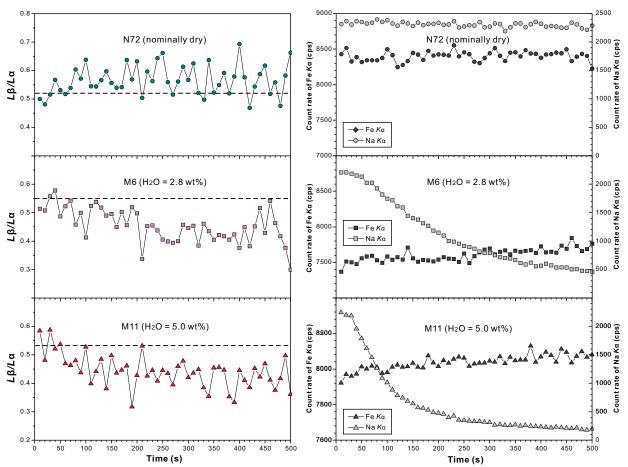


Figure 7

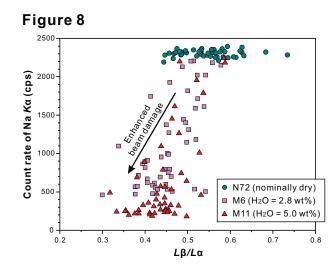
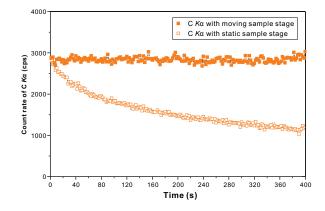
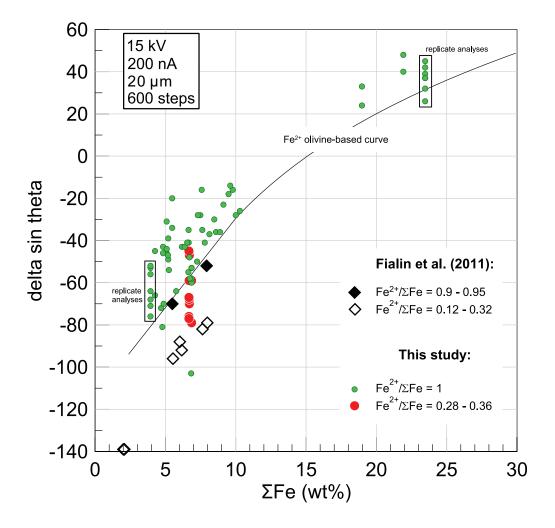


FIGURE 9





Supplementary Figure 1

Table 1. Major element composition of garnets and silicate glasses (wt%)

Group	Sample No.	SiO ₂	TiO ₂	Al ₂ O ₃	FeO _T	std	MnO	MgO	CaO	Na ₂ O	K ₂ O	Total
Garnet	Andradite	34.96	0.02	0.77	27.70	n.d.	0.09	1.01	32.20	0.01	0.03	99.87
Garriet	Almandine	35.82	0.00	20.48	43.61	n.d.	0.04	0.53	0.01	0.04	0.00	100.52
	DAFe-6	45.62	b.d.l.	14.57	8.84	0.08	b.d.l.	9.54	21.95	b.d.l.	b.d.l.	100.51
	DAFe-7	46.69	b.d.l.	14.68	7.42	0.07	b.d.l.	9.72	22.14	b.d.l.	b.d.1.	100.65
	DAF-58	45.62	b.d.l.	14.00	9.07	0.13	b.d.l.	8.84	21.25	b.d.l.	b.d.1.	98.78
	DAF5-58	48.05	b.d.l.	15.05	4.14	0.13	b.d.l.	9.57	22.68	b.d.l.	b.d.1.	99.49
	DAF20-58	40.31	b.d.l.	12.71	17.48	0.19	b.d.l.	8.11	18.94	b.d.l.	b.d.1.	97.55
	DAF-56	45.54	b.d.l.	14.04	8.98	0.07	b.d.l.	9.01	21.28	b.d.l.	b.d.1.	98.85
	DAF5-56	48.05	b.d.l.	15.02	4.26	0.11	b.d.l.	9.47	22.52	b.d.l.	b.d.1.	99.32
	DAF20-56	40.14	b.d.l.	12.57	18.2	0.19	b.d.l.	8.09	18.86	b.d.l.	b.d.1.	97.86
	DAF-57	45.55	b.d.l.	13.97	9.24	0.12	b.d.l.	9.00	21.22	b.d.l.	b.d.l.	98.98
	DAF5-57	48.13	b.d.l.	15.10	4.03	0.09	b.d.l.	9.72	22.67	b.d.l.	b.d.l.	99.65
	DAF20-57	40.53	b.d.l.	12.75	17.04	0.24	b.d.l.	8.12	18.99	b.d.l.	b.d.l.	97.43
	DAF-59	45.51	b.d.l.	14.04	9.17	0.13	b.d.l.	9.08	21.27	b.d.l.	b.d.l.	99.07
	DAF5-59	47.97	b.d.l.	14.93	4.4	0.07	b.d.l.	9.51	22.47	b.d.l.	b.d.1.	99.28
Synthetic	DAF20-59	40.69	b.d.l.	12.83	16.97	0.16	b.d.l.	8.16	19.06	b.d.l.	b.d.l.	97.71
alkali-free	DAF-83	45.97	b.d.l.	14.38	8.44	0.08	b.d.l.	9.26	21.42	b.d.l.	b.d.1.	99.47
glass	DAF5-83	48.23	b.d.l.	15.37	4.09	0.06	b.d.l.	9.81	22.59	b.d.l.	b.d.1.	100.09
8	DAF20-83	41.12	b.d.l.	13.15	16.96	0.05	b.d.l.	8.33	19.17	b.d.l.	b.d.1.	98.73
	DAF-84	46.00	b.d.l.	14.39	8.58	0.05	b.d.l.	9.23	21.36	b.d.l.	b.d.1.	99.56
	DAF5-84	48.09	b.d.l.	15.29	4.16	0.05	b.d.l.	9.74	22.49	b.d.l.	b.d.l.	99.77
	DAF20-84	40.93	b.d.l.	13.15	17.42	0.08	b.d.l.	8.35	19.09	b.d.l.	b.d.l.	98.94
	DAF-85	45.71	b.d.l.	15.00	8.78	0.08	b.d.l.	9.10	21.02	b.d.l.	b.d.l.	99.61
	DAF5-85	48.15	b.d.l.	15.40	4.15	0.06	b.d.l.	9.82	22.53	b.d.l.	b.d.l.	100.05
	DAF20-85	40.77	b.d.l.	13.09	17.7	0.10	b.d.l.	8.28	18.96	b.d.l.	b.d.l.	98.80
	DAT	46.91	b.d.l.	14.52	7.95	0.05	b.d.l.	9.81	21.76	b.d.l.	b.d.l.	100.96
	DAT3	45.47	2.17	14.26	8.21	0.07	b.d.l.	9.67	21.76	b.d.l.	b.d.l.	100.95
	DAT5	44.34	3.93	13.88	8.57	0.04	b.d.l.	9.41	20.50	b.d.l.	b.d.l.	100.63
	DAT10	41.73	8.11	12.91	9.34	0.09	b.d.l.	8.81	19.13	b.d.l.	b.d.l.	100.03
	DAT10 DAT17	37.69	14.45	11.80	10.47	0.12	b.d.l.	7.99	17.36	b.d.l.	b.d.l.	99.76
	DAT17 DAT32	31.62	25.09	9.66	13.26	0.12	b.d.l.	6.36	14.94	b.d.l.	b.d.l.	100.93
	AR39	47.53	2.89	15.05	13.39	0.38	0.04	6.25	11.11	2.45	0.31	99.02
	AR45	47.55	2.89	13.03	12.80	0.38	0.04	5.94	10.94	2.43	0.31	99.02 98.11
Synthetic	AR45 AR37	51.56	3.72	12.03	12.80	0.40	0.01	4.05	9.42	2.33	0.34	99.71
alkali-bearing	AR43	50.16	3.48	11.62	16.97	0.20	0.30	3.93	8.85	2.75	0.28	98.43
U	AR45 AR35	53.91	1.01	17.72	8.52	0.20	0.30	5.84	8.88	2.90	0.30	98.43 99.90
glass		55.63	1.01	17.72	8.32 6.19	0.18	0.20	5.84 6.00	8.88 9.21	2.80	0.99	100.02
	AR41 M72		0.92		9.43			7.05				
		50.50		18.40		0.20	0.18		11.44	2.35	0.23	100.67
D 1/1	AR36	44.73	3.84	13.35	12.16	0.51	0.21	8.93	11.82	3.49	1.65	100.37
Remelted	AR42	44.30	3.72	13.53	11.95	0.17	0.22	8.75	11.95	3.09	1.61	99.07
natural glass	AR34	49.07	0.96	15.59	9.16	0.11	0.19	9.13	12.22	2.12	0.06	98.50
	AR40	49.60	0.99	16.06	9.13	0.11	0.23	9.12	12.17	2.57	0.10	99.96
Natural glass	VG-A99	50.94	4.06	12.49	13.3	0.20	0.15	5.08	9.30	2.66	0.82	99.18
reference	VG-2	50.81	1.85	14.06	11.8	0.20	0.22	6.71	11.12	2.62	0.19	99.58
	USNM113716	51.52	1.30	15.39	9.36	0.18	0.17	8.21	11.31	2.48	0.09	99.95
Natural	M6*	47.68	0.88	17.24	8.53	0.20	0.23	6.62	10.72	2.41	0.22	94.33
hydrous glass Starting	M11*	46.44	0.86	16.72	8.19	0.20	0.20	6.39	10.46	2.22	0.23	91.70
material for M6 and M11	N72	50.07	0.90	18.36	9.35	n.d.	0.17	7.02	11.32	2.45	0.22	100.00

Note: n.d. = not determined. b.d.l. = below detection limit. Standard deviation (1σ) of Fe²⁺/ Σ Fe determined by wet chemistry is assumed as 0.03 for all experimental glasses based on estimated uncertainty of the method (see Schuessler et al. 2008). Standard deviation of FeO_T is derived from repeated EPMA measurements (n = 10 to 40). Standard deviation of Fe²⁺/ Σ Fe determined by EPMA flank method is calculated according to error propagation.

* M6 and M11 contain 2.8 and 5.0 wt% H2O, respectively (determined by Karl-Fischer Titration; Shishkina et al., 2010).

All the other glasses are nominally dry.

Synthetic alkali-free glasses:

 $DAFe_{ii}$ - series of experimental glasses produced by melting $Di_{58}An_{42}$ eutectic composition in Fe loops below IW buffer conditions (Borisov, 2007);

 DAT_{ii} - series of experimental glasses produced by melting $Di_{58}An_{42}$ composition modified with additional TiO₂ in Fe loops below IW buffer (Borisov et al., 2004);

 DAF_{ii} - series of experimental glasses produced by melting $Di_{58}An_{42}$ eutectic modified with variable additional Fe_2O_3 in air (Borisov et al. 2015).

Synthetic alkali-bearing glasses:

AR39 and AR45 are oxidized and reduced samples of the ferrobasaltic glass SC1 (Botcharnikov et al., 2008).

AR37 and AR43 are oxidized and reduced samples of the ferrobasaltic glass LS (Botcharnikov and Koepke, unpublished data).

AR35 and AR41 are oxidized and reduced samples of the basaltic andesite glass BezBA (Almeev et al., 2013).

Remelted natural glasses:

AR36 and AR42 are oxidized and reduced samples of the basanite KLA-1-6-22 (Fuchs et al., 2014; Klügel et al., 2017).

AR34 and AR40 are oxidized and reduced samples of the natural MORB glass 140ox (Almeev et al., 2007).

Experimental conditions for synthetic alkali-bearing glasses and remelted natural glasses:

The "oxidized" AR34 – AR39 glasses have been produced in a 1 atm furnace at 1600 °C, in air using Pt crucibles.

The "reduced" AR40 – AR45 glasses have been produced in internally heated pressure vessel at 1250 °C, 200 MPa under intrinsic conditions in Pt-lined graphite capsules (see details in Husen et al. 2016).

The M6 and M11 glasses are H₂O-saturated glasses have been produced in internally heated pressure vessel at 1250 °C, 100 and 200 MPa respectively, under intrinsic conditions in $Au_{80}Pd_{20}$ capsules (see details in Shishkina et al. 2010). Starting glass N72 was produced by re-melting of the island arc-tholeiitic basalt at 1atm furnace at 1600 °C, in air using Pt crucible.

Difference between EPMA Sample No. Total Fe Wet chemistry EMPA flank method and wet chemistry Fe²⁺ $\Delta \ {\rm Fe}^{2+}$ $\Delta Fe^{2+}/$ Fe²⁺ **FeO**_T std Fe²⁺/ΣFe std std Session $L\beta/L\alpha$ std std Fe²⁺/ΣFe std (wt%) (wt%) (wt%) ΣFe Garnets Andradite 27.70 n.d. 0.00 n.d. 0.00 n.d. 0.428 0.003 2 0.415 0.003 3 0.004 0.439 _ Almandine 43.61 n.d. 1.00 n.d. 33.90 n.d. 1 1.419 0.023 _ 2 1.428 0.026 3 1.527 0.024 Synthetic alkali-free glass DAFe-6 0.04 8.84 0.081.00* n.d. 6.87 0.06 0.636 0.015 7.12 0.52 1.04 0.08 0.25 1 DAFe-7 7.42 0.07 1.00* 5.77 0.05 0.595 0.005 5.73 0.16 0.99 0.03 -0.04 -0.01 n.d. DAF-58 9.07 0.13 0.25 0.03 1.76 0.21 1 0.469 0.004 1.42 0.14 0.20 0.02 -0.34 -0.05 **DAF5-58** 0.28 0.90 0.460 0.014 1.09 4.14 0.130.03 0.10 1 0.480.34 0.150.19 0.06 DAF20-58 17.48 0.19 0.21 0.03 2.85 0.41 1 0.526 0.004 3.36 0.140.25 0.01 0.51 0.04 0.29 0.490 DAF-56 8.98 0.07 0.03 2.02 0.21 1 0.007 2.12 0.26 0.30 0.04 0.10 0.01 0.474 **DAF5-56** 4.26 0.32 0.03 0.013 1.59 0.450.48 0.53 0.111.06 0.10 1 0.140.16 0.541 DAF20-56 18.2 0.190.26 0.03 3.68 0.431 0.012 3.88 0.400.27 0.03 0.20 0.01 **DAF-57** 9.24 0.12 0.33 0.03 2.37 0.22 0.500 0.006 2.49 0.19 0.35 0.03 0.12 0.02 1 DAF5-57 4.03 0.09 0.38 0.03 1.19 0.10 1 0.474 0.008 1.59 0.26 0.51 0.08 0.40 0.13 DAF20-57 17.04 0.240.30 0.03 3.97 0.400.555 0.010 4.37 0.340.33 0.03 0.400.03 1 **DAF-59** 9.17 0.13 0.36 0.03 2.57 0 22 1 0.498 0.011 2.42 0 38 0.34 0.05 -0.15-0.02DAF5-59 4.4 0.07 0.38 1.30 0.471 1.48 0.37 0.43 0.05 0.03 0.10 1 0.011 0.11 0.18 DAF20-59 16.97 0.16 0.33 0.03 4.35 0.40 0.531 0.005 3.54 0.18 0.27 0.01 -0.81 -0.06 1 4.02 -0.048.44 0.080.03 0.20 0.545 0.007 -0.24**DAF-83** 0.654.26 1 0.220.610.03DAF5-83 4.09 0.06 0.69 0.03 2.19 0.10 1 0.514 0.013 2.94 0.45 0.92 0.14 0.75 0.23 DAF20-83 0.643 7.38 -0.05 16.96 0.05 0.61 0.03 8.04 0.401 0.008 0.290.56 0.02 -0.66 DAF-84 8.58 0.05 0.67 0.03 4.47 0.20 1 0.552 0.010 4.25 0.34 0.64 0.05 -0.22 -0.03 0.511 DAF5-84 4.16 0.05 0.68 0.03 2.200.10 1 0.004 2.85 0.12 0.88 0.04 0.65 0.20 DAF20-84 17.42 8.26 0.637 0.009 0.31 -1.10 -0.08 0.08 0.61 0.03 0.41 1 7.16 0.53 0.02 DAF-85 8.78 0.08 0.62 0.03 4.23 0.21 1 0.547 0.013 4.07 0.45 0.60 0.07 -0.16 -0.02 0.503 **DAF5-85** 0.06 0.62 0.03 2.000.002 2.580.070.58 0.18 4.15 0.101 0.800.02DAF20-85 17.7 0.10 0.58 0.03 7.98 0.42 1 0.641 0.009 7.30 0.320.53 0.02 -0.68 -0.05DAT 7.95 0.05 1.00*n.d. 6.18 0.042 0.591 0.016 5.90 0.53 0.96 0.09 -0.28 -0.042 0.42 8.21 1.00*0.600 0.97 -0.19-0.03DAT3 0.07n.d. 6.38 0.05 0.012 6.19 0.07 DAT5 8.57 1.00* 2 0.603 0.012 6.31 0.40 0.95 -0.35 -0.05 0.04n.d. 6.66 0.03 0.06 1.00*2 DAT10 9.34 0.09 n.d. 7.26 0.07 0.621 0.014 6.89 0.480.95 0.07 -0.37-0.052 DAT17 10.47 0.12 1.00* n.d. 8.14 0.09 0.651 0.015 7.91 0.50 0.97 0.06 -0.23-0.03 DAT32 13.26 0.14 1.00* n.d. 10.31 0.11 2 0.672 0.014 8.62 0.45 0.84 0.04 -1.69 -0.16 Synthetic alkali-bearing glass 3.54 AR39 13.39 0.38 0.34 0.33 0.548 0.008 0.05 0.03 4.11 0.27 0.39 0.03 1 0.570.560 0.005 3.79 0.15 0.36 0.02 0.25 0.02 3 AR45 12.80 0.40 0.91 0.03 9.05 0.41 1 0.698 0.008 9.25 0.28 0.93 0.04 0.20 0.02 3 0.717 0.022 8.67 0.67 0.87 0.07 -0.38-0.040.37 AR37 0.03 15.84 0.20 0.38 4.68 1 0.558 0.008 4.46 0.260.36 0.02 -0.22-0.022 0.540 0.002 4.19 0.06 0.31 0.00 -0.49 -0.07 3 0.573 0.004 4.18 0.13 0.34 0.01 -0.50 -0.04 16.97 AR43 0.26 0.88 0.03 11.61 0.779 0.03 0.431 0.014 12.00 0.47 0.91 0.01 0.39 2 0.745 0.004 11.04 0.14 0.81 0.01 -0.57 -0.07 3 0.804 0.014 11.38 0.42 0.86 0.03 -0.23 -0.02 8.52 0.40 0.03 0.08 AR35 0.18 2.65 0.21 0.522 0.007 0.48 0.04 0.56 1 3.21 0.242 0.486 0.012 2.11 0.16 0.44 0.03 -0.540.04 AR41 6.19 0.05 0.95 0.03 4.57 0.15 1 0.564 0.006 0.20 0.97 0.04 0.08 0.02 4.65 2 0.545 0.012 4.35 0.42 0.92 0.09 -0.22 -0.03 9.43 2 M72 0.20 0.47 0.03 3.44 0.23 0.006 3.09 -0.35 -0.01 0.518 0.21 0.46 0.03 Remelted natural glass AR36 12.16 0.41 0.569 0.10 0.51 0.03 3.88 0.33 1 0.008 4.83 0.28 0.51 0.03 0.95 4.03 2 0.535 0.008 0.270.43 0.03 0.15 0.02 AR42 11.95 8.94 0.170.90 0.03 8.36 0.30 1 0.689 0.008 0.26 0.96 0.03 0.58 0.06 2 0.661 0.006 8.23 0.19 0.87 0.02 -0.13 -0.03 0.22 3 AR34 9.16 0.11 0.59 0.03 4.20 0.577 0.005 4.29 0.16 0.60 0.02 0.09 0.01

Table 2. Fe oxidation state of garnets and silicate glasses

AR40	9.13	0.11	0.91	0.03	6.46	0.23	3	0.645	0.010	6.42	0.31	0.90	0.04	-0.04	-0.01
Natural glass reference															
VG-A99	13.3	0.20	0.87**	n.d.	8.99	0.14	1	0.664	0.004	8.09	0.13	0.78	0.01	-0.90	-0.09
							2	0.665	0.004	8.39	0.12	0.81	0.02	-0.60	-0.06
VG-2	11.8	0.20	0.83**	n.d.	7.61	0.13	1	0.659	0.009	7.90	0.32	0.86	0.03	0.29	0.03
							2	0.628	0.006	7.22	0.20	0.79	0.02	-0.39	-0.04
USNM- 113716	9.36	0.18	0.89**	n.d.	6.48	0.12	1	0.615	0.008	6.41	0.27	0.88	0.04	-0.07	-0.01
							2	0605	0.004	6.05	0.36	0.83	0.05	-0.43	-0.06
Synthetic hydrous glass															
M6	8.53	0.20	0.67	0.03	4.44	0.22	2	0.550	0.009	4.17	0.32	0.62	0.05	-0.27	-0.05
M11	8.19	0.20	0.58	0.03	3.69	0.21	2	0.530	0.007	3.50	0.24	0.60	0.04	-0.19	0.02

* Assumed value. The synthetic glasses have been produced in experiments with a Fe-loop, below iron-wüstite buffer conditions (Borisov et al. 2004; Borisov 2007).

** Ratio calculated based on FeO and Fe₂O₃ contents provided in Jarosewich et al. (1980).

Table 3. Possible errors in fO_2 determination using the flank method for MORB glasses

Fe ²⁺ /ΣFe	FeO _T =	10 wt%	FeO _T =	9 wt%	$FeO_T = 8 wt\%$			
	$\Delta Fe^{2+}/\Sigma Fe^{*}$	$\Delta log fO_2^{**}$	$\Delta Fe^{2+}/\Sigma Fe$	$\Delta log fO_2$	$\Delta Fe^{2+}/\Sigma Fe$	$\Delta log fO_2$		
0.95	0.13	0.30	0.14	0.33	0.16	0.37		
0.90	0.13	0.16	0.14	0.17	0.16	0.19		
0.85	0.13	0.11	0.14	0.12	0.16	0.14		
0.80	0.13	0.09	0.14	0.10	0.16	0.11		
0.75	0.13	0.07	0.14	0.08	0.16	0.07		

* assumed $\Delta Fe^{2+} = 1$ wt% (2 σ , flank method, EPMA) and $\Delta \Sigma Fe = 1$ wt% (EPMA);

** assumed ideal slope of 0.25 for $\log(Fe^{3+}/Fe^{2+})$ versus $\log fO_2$.