

Evaluation of a present-day climate simulation with a new coupled atmosphere-ocean model GENMOM

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Abstract. We present a new, non-flux corrected AOGCM, GENMOM, that combines the GENESIS version 3 atmospheric GCM (Global Environmental and Ecological Simulation of Interactive Systems) and MOM2 (Modular Ocean Model version 2) nominally at T31 resolution. We evaluate GENMOM by comparison with reanalysis products (e.g., NCEP2) and three models used in the IPCC AR4 assessment. GENMOM produces a global temperature bias of 0.6 °C. Atmospheric features such as the jet stream structure and major semi-permanent sea level pressure centers are well simulated as is the mean planetary-scale wind structure that is needed to produce the correct position of stormtracks. Most ocean surface currents are reproduced except where they are not resolvable at T31 resolution. Overall, GENMOM captures reasonably well the observed gradients and spatial distributions of annual surface temperature and precipitation and the simulations are on par with other AOGCMs. Deficiencies in the GENMOM simulations include a warm bias in the surface temperature over the southern oceans, a split in the ITCZ and weaker-than-observed overturning circulation.

1 Introduction

We present a new non-flux corrected coupled atmosphere-ocean general circulation model (AOGCM), GENMOM, which combines GENESIS version 3 (Global ENvironmental and Ecological Simulation of Interactive Systems) and MOM2 (Modular Ocean Model version 2) general circulation models. Both models have been used widely in climate studies that demonstrate their overall ability to produce climate simulations that are in agreement both with obser-

vations and with similar models. GENESIS version 1 was developed initially in 1989 at the National Center for Atmospheric Research (NCAR) with a focus on linking terrestrial physical and biophysical processes with the atmosphere to provide a model that could be applied to simulate paleoclimate and possible future climates under global warming.

GENESIS version 1 was released in 1991 (Thompson and Pollard, 1995) and included a land-surface transfer model (LSX) and an atmospheric general circulation model derived from NCAR CCM1. GENESIS version 2 was released in 1995 and included many improvements ranging from new prognostic cloud amounts, the use of hybrid vertical coordinates, the inclusion of gravity wave drag, and additional improvements in LSX (Thompson and Pollard, 1997; Pollard and Thompson, 1997).

GENESIS version 3 expands on version 2 by including the NCAR CCM3 radiation code. The ocean can optionally be represented by fixed sea surface temperatures (SSTs), a slab or mixed layer or by the MOM2 ocean general circulation model. MOM also has a long history of use and development also spanning back to the early 1990s and is used as the ocean component in many other AOGCMs (Pacanowski, 1996). Our current version of GENMOM uses T31 ($\sim 3.75^\circ \times 3.75^\circ$ latitude and longitude) horizontal resolution for both the atmosphere and ocean to balance computational requirements needed for long simulations with the ability to simulate important features of the general circulation. A higher resolution version with $2^\circ \times 2^\circ$ ocean and LSX with T31 atmosphere is under development.

We evaluate a simulation of modern climate using observational and reanalysis data and we compare GENMOM surface temperature and precipitation with three other AOGCMs evaluated in the World Climate Research Programme (WCRP) Coupled Model Intercomparison Project phase 3 (CMIP3, Meehl et al., 2007a), a multi-model dataset that was subsequently used in the IPCC AR4 (described in



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Table 1. Three AOGCMs used in the IPCC AR4. T indicates the horizontal resolution using spectral truncation. L indicates the number of levels used in the model.

Model	Modeling Center, Country	Atmosphere Resolution	Ocean Resolution
GFDL CM 2.0	US Department of Commerce/National Oceanic and Atmospheric Administration (NOAA)/Geophysical Fluid Dynamics Laboratory (GFDL), USA	$2.0^\circ \times 2.5^\circ$ L24 GFDL GAMDT (2004)	$0.3^\circ\text{--}1.0^\circ \times 1.0^\circ$ Gnanadesikan et al. (2004)
MPI ECHAM5	Max Planck Institute for Meteorology, Germany	T63 ($\sim 1.9^\circ \times 1.9^\circ$) L31 Roeckner et al. (2003)	$1.5^\circ \times 1.5^\circ$ L40 Marsland et al. (2003)
NCAR CCSM 3.0	National Center for Atmospheric Research, USA	T85 ($1.4^\circ \times 1.4^\circ$) L26 Collins et al. (2004)	$0.3^\circ\text{--}1^\circ \times 1^\circ$ L40 Smith and Gent (2002)

Table 1, Randall et al., 2007). A full description of GENESIS and MOM2 as well as their coupling is provided in Sect. 2, atmospheric and oceanic results from a modern climate simulation are presented in Sect. 3, and concluding remarks follow in Sect. 4.

2 GENMOM description

2.1 GENESIS description

GENESIS has been developed with emphasis on terrestrial physical and biophysical processes, and suitability for paleoclimatic experiments. Earlier versions of GENESIS are described by Thompson and Pollard (1995, 1997) and Pollard and Thompson (1994, 1995, 1997), and have been applied and tested in a wide range of modern and paleoclimate applications including the Paleoclimate Modeling Intercomparison Project (e.g., Pollard et al., 1998; Joussaume et al., 1999; Pinot et al., 1999; Beckmann et al., 2005; Miller et al., 2005; Ruddiman et al., 2005; Bice et al., 2006; DeConto et al., 2006, 2008; Hostetler et al., 2006; Poulsen et al., 2006, 2007; Horton et al., 2007).

The nominal GENESIS resolution is spectral T31 ($3.75^\circ \times 3.75^\circ$) with 18 vertical sigma coordinate levels, 4 of which are above the tropopause. Spectral transform dynamics are used for mass, heat and momentum (Williamson et al., 1987). A semi-Lagrangian transport in grid space is used for water vapor (Williamson and Rasch, 1989). Convection in the free atmosphere and in the planetary boundary layer is treated using an explicit sub-grid buoyant plume model similar to, but simpler than, Kreitzberg and Perkey (1976) and Anthes (1977, Sect. 4). Stratus, convective and anvil cirrus clouds are predicted using prognostic 3-D water cloud amounts, (Smith, 1990; Senior and Mitchell, 1993) and clouds are advected by semi-Lagrangian transport and mixed vertically by convective plumes and background diffusion.

The land-surface transfer model, LSX, accounts for the physical effects of vegetation (Pollard and Thompson, 1995). Up to two vegetation layers (trees and grass) can be speci-

fied at each grid point, and the radiative and turbulent fluxes through these layers to the soil or snow surface are calculated. A six-layer soil model extends from the surface to 4.25 m depth, with layer thicknesses increasing from 5 cm at the top to 2.5 m at the bottom. Physical processes in the vertical soil column include heat diffusion, liquid water transport (Clapp and Hornberger, 1978; Dickinson, 1984), surface runoff and bottom drainage, uptake of liquid water by plant roots for transpiration, and the freezing and thawing of soil ice. A three-layer snow model, which includes fractional area cover when the snow is thin, is used for snow cover on soil, ice-sheet and sea-ice surfaces. A three-layer sea-ice model accounts for local melting, freezing, fractional sea-ice cover (Semtner, 1976; Harvey, 1988), and includes dynamics associated with wind and ocean current using the cavitating-fluid model of Flato and Hibler (1992). Version 3 of GENESIS (Zhou et al., 2008; Kump and Pollard, 2008) incorporates the NCAR CCM3 radiation code (Kiehl et al., 1998) and the ocean is represented by the MOM2 ocean general circulation model (Pacanowski, 1996).

2.2 MOM2 description

MOM2 was developed by the Geophysical Fluid Dynamics Laboratory (GFDL) in the early 1990s, but builds off previous work that began back in 1969 (Pacanowski, 1996). MOM2 is a finite difference implementation of the primitive equations of ocean circulation based on the Navier-Stokes equations with the Boussinesq, hydrostatic, and rigid lid approximations (Bryan, 1969). The Boussinesq approximation invokes constant density with depth, with the exception of terms that contain gravity, thereby reducing computational complexity. The hydrostatic approximation assumes that vertical pressure gradients are density driven. A nonlinear equation of state couples temperature and salinity to fluid velocity. An insulated lateral boundary is used such that no temperature or salinity flux is exchanged between ocean and land cells. Unlike the sigma levels used for atmospheric altitude in GENESIS, MOM uses a fixed z-axis for depth, which

simplifies the equations used in the finite difference representation. Our version of MOM2 uses 20 unevenly spaced vertical levels that become progressively thicker with depth, so that the uppermost ocean layers are well resolved. The topmost level is 25 m thick, while the bottommost level is ~ 660 m thick. A horizontal resolution of $3.75^\circ \times 3.75^\circ$ is used to match the atmospheric T31 resolution. The hybrid mixing scheme isopycnal is used with a prescribed vertical viscosity coefficient of $0.1 \text{ cm}^2 \text{ s}^{-1}$ and a prescribed vertical diffusion coefficient of $0.35 \text{ cm}^2 \text{ s}^{-1}$. Although the Gent-McWilliams mixing scheme (Gent and McWilliams, 1990) is available in MOM2, it was not used in this study.

2.3 GENMOM coupling

To simplify the coupling between the atmosphere and ocean, both the GCMs are implemented on essentially the same T31 grid. In MOM2, the latitudinal grid spacing is not exactly T31, but is adjusted with a cosine-stretching factor (Pacanowski, 1996) to closely approximate T31. GENESIS has a 30-min timestep, and MOM2 has a 6-h timestep for scalar fields. The two models interact in an essentially synchronous manner, communicating every 6 hours: 6-h averages of the surface fluxes of heat, water and momentum are passed from GENESIS to MOM2, and MOM2 is run through one 6-h scalar timestep and the updated SSTs are passed back to GENESIS and used to run it for the next 6 h. Sea ice is treated within the LSX module of GENESIS. Under sea ice, fluxes between the base and the uppermost ocean layer are passed to MOM2. Continental freshwater river runoff is globally averaged and spread over the world ocean. The T31 version of GENMOM simulates ~ 22 yr per calendar day on an 8-CPU Linux server. GENMOM coupled to MOM2 has previously been used by Zhou et al. (2008) to investigate warm Cretaceous climates; however, their study focused mainly on paleoclimates and water isotopic ratios, and not on modern climatology.

3 Simulation of the present-day climate

We analyze the annual and seasonal climatologies of the last 30 years of a 700-year GENMOM simulation. Analysis of ocean temperature indicates that spin up of the model was suitably achieved after 400 years. Over the last century of the simulation the deep ocean (> 1000 m) warmed by ~ 0.002 °C/decade, the mid layers (200 m–1000 m) warmed by ~ 0.003 °C/decade and the surface temperature was essentially free of drift. We prescribed atmospheric CO_2 concentration at 355 ppmV, near the mean value for the 1981–2005 climatology period. GENMOM was initialized with a latitudinal-dependent temperature profile and ocean salinity was uniformly prescribed at 35 PSS.

GENMOM input files for topography, bathymetry, and land-ocean mask were derived by interpolating the ICE-4G

model (Peltier, 2002) reconstruction from $1^\circ \times 1^\circ$ to T31 resolution. Ice-sheet cover and thickness is prescribed by interpolating the ICE-4G model reconstruction to T31. To maintain numerical stability, we smooth the bathymetry field over the northernmost Arctic Ocean with a 9-cell moving window. At T31 horizontal resolution the Bering Strait is closed. Modern values for the distribution of vegetation (Dorman and Sellers, 1989), soil texture (Webb et al., 1993) and freshwater lakes (Cogley, 1991) are prescribed. The use of ICE-4G orography to derive global topography, bathymetry, and ice-sheet extent is based on our goal of streamlining the configuration of GENMOM for paleoclimate applications.

3.1 Validation datasets and input files

To evaluate the GENMOM atmospheric fields we use the NOAA NCEP Reanalysis 2 data set (NCEP2, Kanamitsu, et al., 2002) for the standard climatology period of 1981–2005 unless otherwise specified. Although NCEP2 is comprised of observed and derived data it provides an internally consistent dataset with which to evaluate our model. SST data are derived from the NOAA Optimum Interpolation Sea Surface Temperature V2 (OI SST, Reynolds, et al., 2002), which is a $1^\circ \times 1^\circ$ gridded dataset based on combining in situ measurements and satellite observations. We use a climatology period of 1982–2005 for OI SST, because 1982 is the first full year for which the data are available. Subsurface ocean temperatures were obtained from The World Ocean Atlas 2005 (WOA05, Locarnini et al., 2006), which is also a $1^\circ \times 1^\circ$ gridded dataset of ocean temperature and salinity. We use the Hadley Ice and Sea Surface Temperature v1.1 (HadISST, UK Meteorological Office, 2006) for observed sea-ice extent data and to evaluate ocean surface currents and overturning, we use the German partner of the Estimating the Circulation and Climate of the Ocean dataset (GECCO), which is a 50-year (1950–2000) oceanography reanalysis (Köhl and Stammer, 2008; NCEP1; Kalnay et al., 1996).

3.2 Atmospheric fields

The zonally averaged profile of air temperature simulated by GENMOM is in overall agreement with the NCEP2 profile (Fig. 1). Seasonally, GENMOM simulates the meridional shift of peak insolation and warmest surface temperatures well although the modeled tropical warm region is slightly more compressed meridionally than that of the NCEP2 data. North of 30° N in both boreal winter and summer GENMOM produces a cold bias relative to NCEP2 from the surface extending up to the mid-troposphere. An additional cold bias in the upper atmosphere is found over the polar region during boreal winter. A cold bias south of 60° S is present during austral summer and in the upper atmosphere during winter.

The summer and winter patterns and magnitudes of the annually averaged planetary jet stream structure are well captured by GENMOM (Fig. 2). In both winter hemispheres the

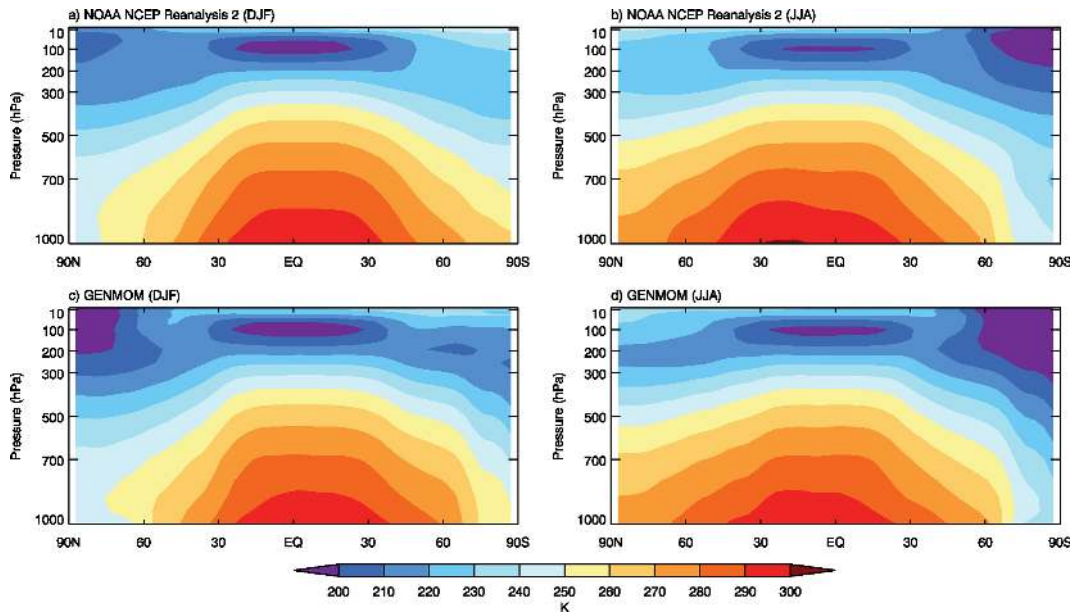


Fig. 1. Mean-annual zonal averaged atmospheric temperature profiles. (a) Observed (NCEP2, 1981–2005) December, January, February (DJF), (b) Observed June, July, and August (JJA), (c) GENMOM DJF, (d) GENMOM JJA.

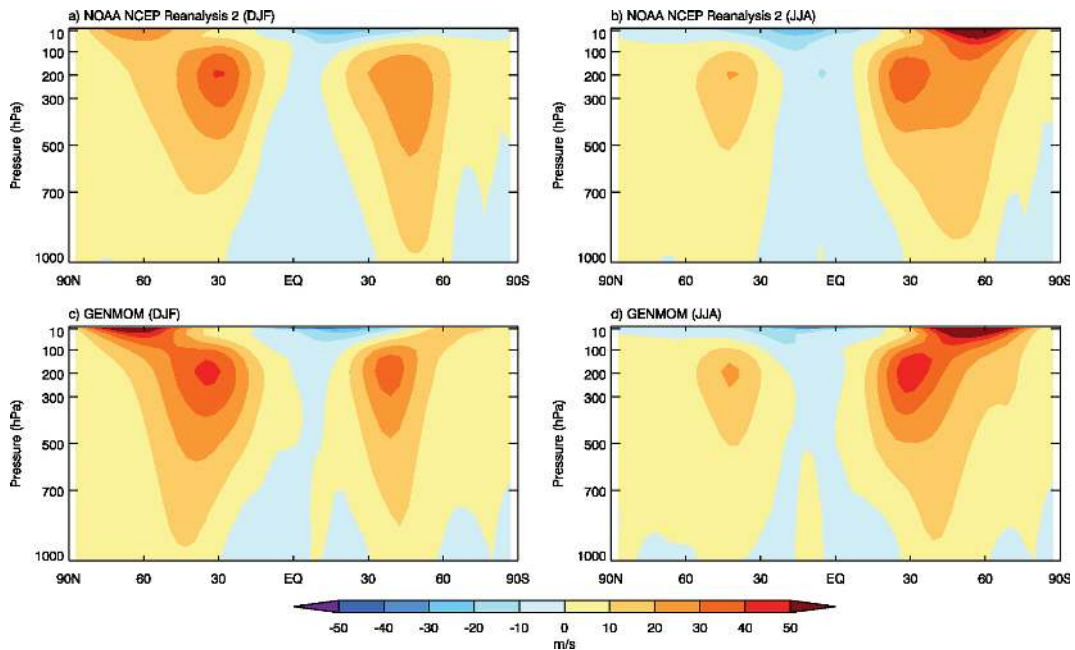


Fig. 2. Winter (DJF) and summer (JJA) zonally averaged eastward wind velocity. (a) Observed (NCEP2, 1981–2005) DJF, (b) Observed JJA, (c) GENMOM DJF, (d) GENMOM JJA.

core of the jetstream (at ~ 200 hPa) and related upper level winds (500 hPa) are slightly enhanced relative to those of the NCEP2 data. These minor mismatches notwithstanding, the overall structure of the simulated jetstream suggests that GENMOM produces a realistic mean planetary-scale wind structure that is essential to the related positioning of the stormtracks.

GENMOM simulates the positions of the seasonally persistent planetary-scale ridges and troughs and thus the resulting upper atmospheric flow and 500 hPa geopotential heights (Fig. 3a–d). During boreal winter, the ridge over western North America is shifted eastward in GENMOM relative to observations, resulting in the eastward displacement of the associated trough over northern Canada and the North

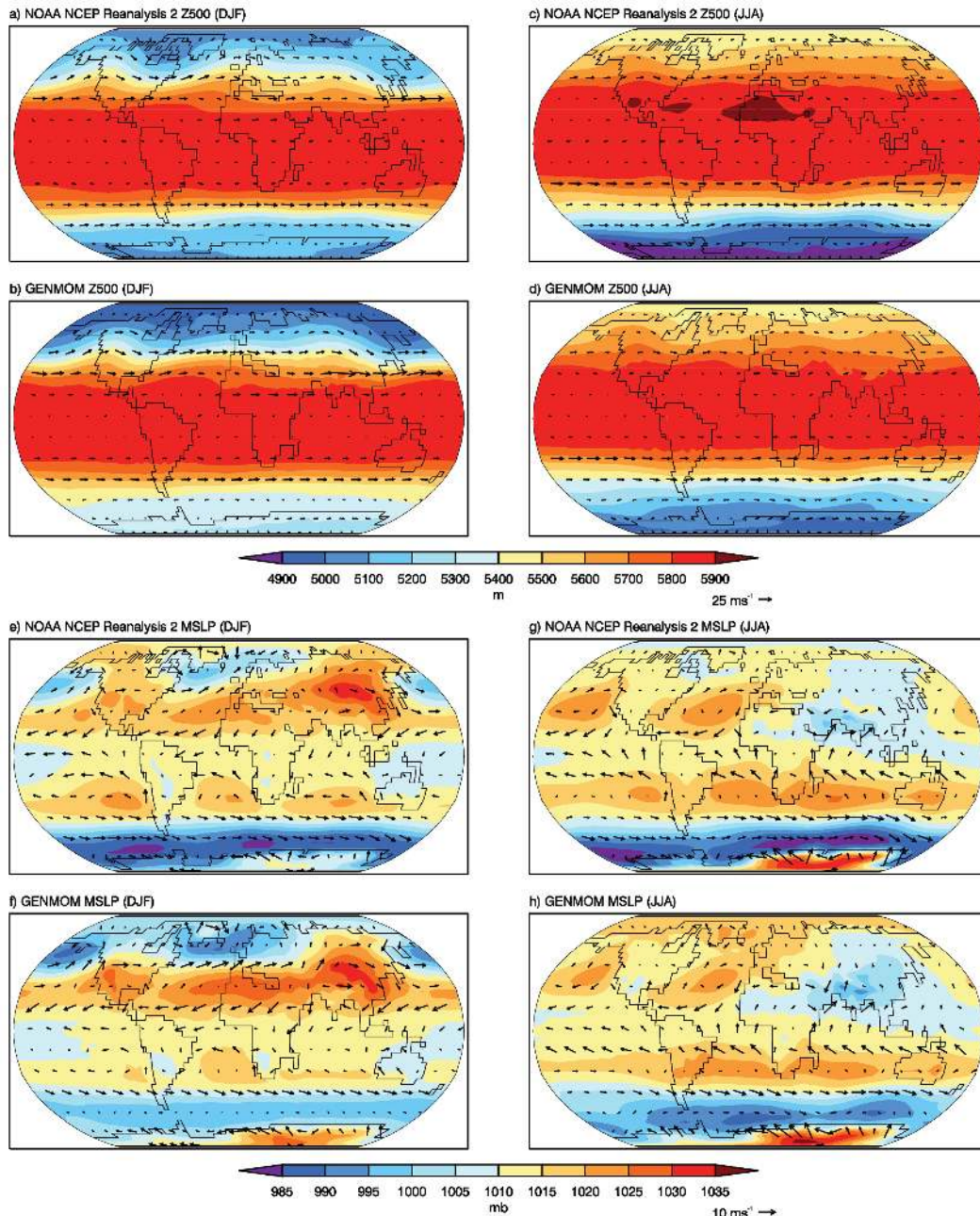


Fig. 3. 500 hPa geopotential height (Z500, a–d) and mean sea level pressure (MSLP, e–h) with wind vectors for both winter (DJF) and summer (JJA).

Atlantic and slightly more zonal flow than that of the NCEP2 data (Fig. 3b). The 500 hPa heights over North America and Eurasia are lower than those of the NCEP2 data resulting in reduced wind velocities, particularly over eastern North America and the North Atlantic. In the Southern Hemisphere (SH), austral summer 500 hPa heights are well simulated but wind velocities associated with the westerlies are somewhat weaker than NCEP2 due to the lower pressure gradient over the Southern Ocean and Antarctica and the lack of actual topographic forcing due to the resolution of the model.

During boreal summer, the ridge over western North America is correctly placed in GENMOM, but the amplitude of the ridge and the related downstream trough is greater than that of the NCEP2 data (Fig. 3c). Heights in the region extending east of the Mediterranean and across India and China appear modestly lower than observed; however, part of the apparent discrepancy stems from values that are just above or just below the color breaks in the plotting scales.

The observed spatial pattern of the semi-permanent sea level pressure (MSLP) cells is captured by GENMOM with

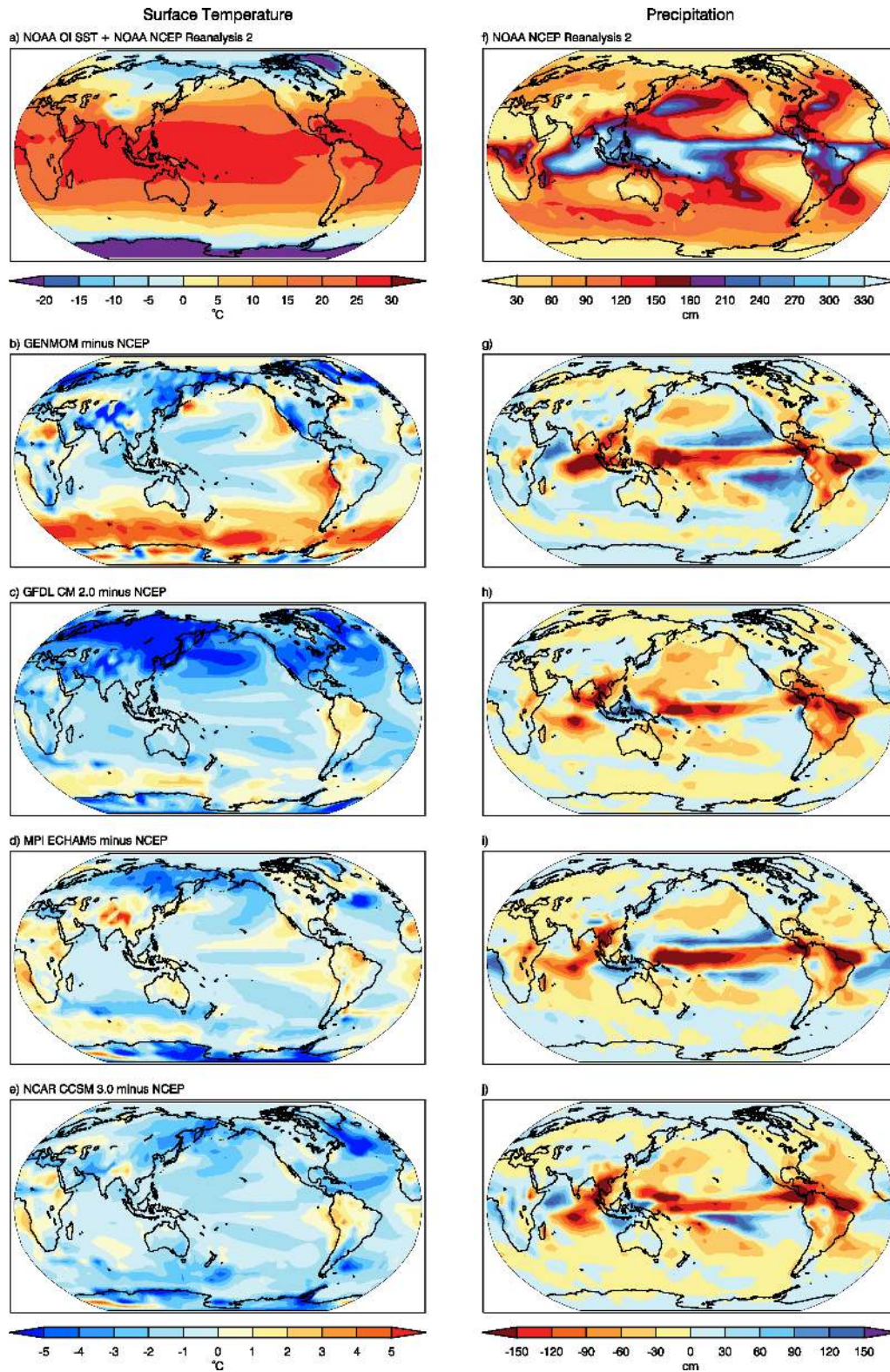


Fig. 4. Annual surface temperature (a) with model anomalies (b–e) and annual total precipitation (f) with model anomalies (g–j) from GENMOM and three AOGCMs included in the IPCC AR4. Observed data are from NOAA NCEP Reanalysis 2 (over land) and NOAA OI SST (over sea). GENMOM 2 m air temperature and SST are for model years 670–699 of the control equilibrium simulation. The three IPCC AR4 models are averaged over the last 30 years (1970–1999) of the Climate of the 20th Century experiment. All data are bi-linearly interpolated to a $5^\circ \times 5^\circ$ grid. Anomalies are calculated as simulation – observation.

some regional differences (Fig. 3e–h). During boreal winter GENMOM simulates lower-than-observed MSLP in the Aleutian and Icelandic lows. As a result, surface wind velocities are enhanced around the pressure centers and over North America. Stronger-than-observed westerly winds across southern Europe are simulated by GENMOM as a result of the enhanced Icelandic low and the cold temperature bias in the Norwegian Sea. In the SH (austral summer), surface pressure and winds are comparable with those of the NCEP2 but GENMOM simulates a weaker-than-observed subtropical high and associated anticyclonic wind flow over the eastern Pacific off the coast of South America (Fig. 3e and f). This deficiency in turn affects the magnitude of the northward wind along the coast and the strength of convergence and westerly trade winds in the Intertropical Convergence Zone (ITCZ). Over the Southern Ocean, the simulated gradient of MSLP is weaker and more diffuse than that of NCEP2 resulting in lower surface wind velocity, which affects the location and strength of circulation around Antarctica.

The boreal summer MSLP pressure patterns and wind velocities in both hemispheres are well captured by GENMOM (Fig. 3g and h). The Northern Hemisphere (NH) subtropical highs are well placed, but slightly weaker than NCEP2; the associated wind velocities are similar to those of NCEP2. MSLP and wind velocities in the tropics and the SH (austral winter) are also well simulated by GENMOM; however, as is the case with DJF, the JJA high-pressure anticyclones are somewhat weaker than observed. Again, the associated anticyclonic flow around the subtropical highs is too weak which contributes to a weakened South Pacific Gyre. The SH westerly winds are simulated to be too weak, presumably due to coarse resolution topography, which will influence ocean overturning.

3.3 Surface temperature

The simulated global mean-annual 2 m air temperature is 278.3 K, in good agreement with the NCEP2 value of 278.9 K. Over land, the simulated temperature is 1.3 K colder than observed and over the oceans simulated temperature is 0.6 K warmer than observed. GENMOM captures the observed meridional temperature and temperatures over topographic features resolved by the model such as the Rocky Mountains, the Andes and the Himalayas. These regions have temperatures that compare well with observations (Fig. 4b). The temperature bias over land and much of the ocean is similar in magnitude to other AOGCMs (Fig. 4c–d).

The cold bias over the Norwegian Sea reflects too much simulated sea-ice resulting from insufficient meridional overturning. Warm biases are found over the Southern Ocean and in the upwelling region off the western coast of South America (Fig. 4a). The Southern Ocean bias is attributed primarily to weaker-than-observed sea level pressure gradient and attendant reduced westerly wind strength (Fig. 3) which results in weakened ocean overturning. The warm bias off the

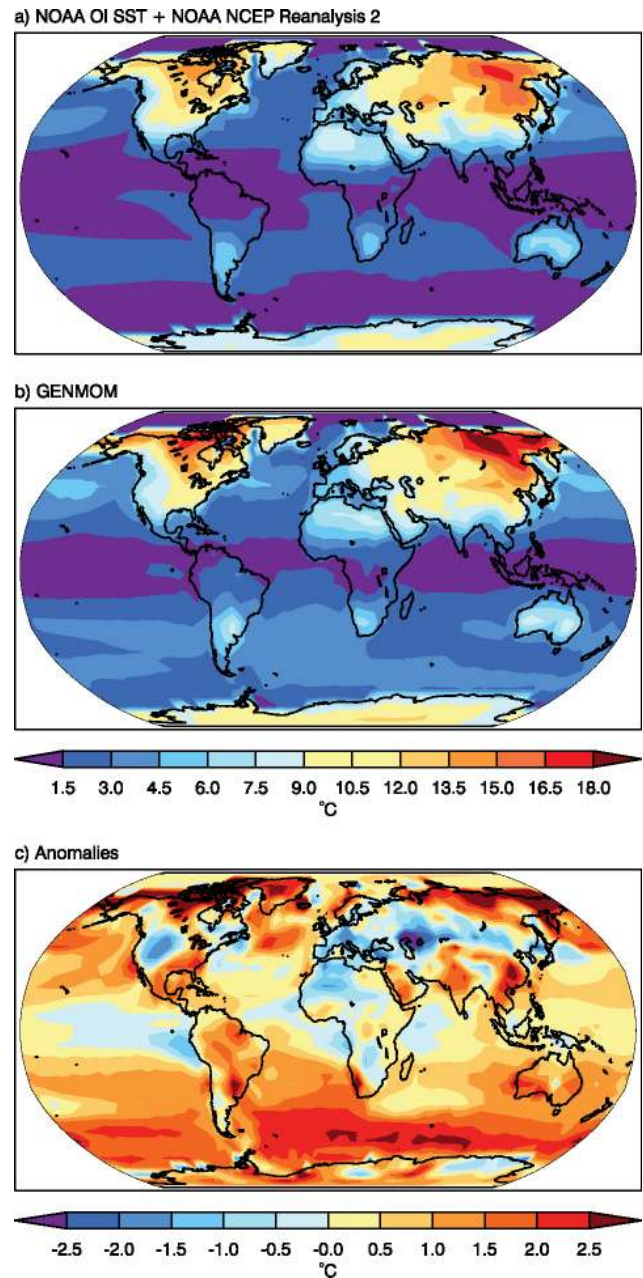


Fig. 5. Observed and modeled seasonal cycle amplitude of surface temperature and anomalies. The amplitude of the seasonal cycle is calculated as the standard deviation of the 12 climatological months.

coast of South America is in part attributed to weaker shore-parallel winds associated with the reduced subtropical high (Fig. 3) and to the lack of resolution of finer scale currents in MOM. Lack of a California Current (~ 300 km wide) is similarly attributed to the T31 resolution of the models. Further discussion of the oceanic circulation is given in Sect. 3.5.

GENMOM captures the global patterns of the seasonal cycle of temperature but overestimates the amplitude over

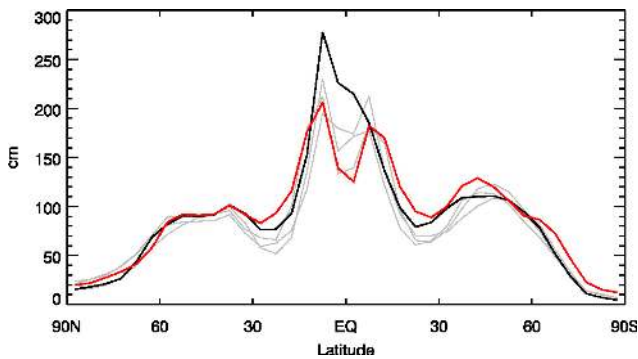


Fig. 6. Zonal averaged annual precipitation for observations (black), GENMOM (red) and three IPCC AR4 models (gray). The IPCC AR4 models are averaged over the last 30 years (1970–1999) of the Climate of the 20th Century experiment.

Greenland, South America, southeast United States and Australia and underestimates the amplitude over northern Africa, the western United States and much of Europe and Asia (Fig. 5). The model also simulates greater variability over some of the oceans, particularly in the mid latitudes.

3.4 Precipitation

GENMOM simulates global mean-annual precipitation reasonably well relative to the reanalysis data (Figs. 4g and 6). Notable exceptions are dry biases over Southeast Asia and South America. GENMOM also produces a double ITCZ in the tropical Pacific, which is characterized by displacement of precipitation maxima off the equator (Lin, 2007). The dry biases and the double ITCZ are common to other AOGCMs (Fig. 4h–j).

During DJF, the southern branch of the ITCZ simulated by GENMOM extends too far to the east, whereas during JJA, the northern branch of the ITCZ is compressed and extends too far to the north relative to observations. Lin (2007) attributes the double ITCZ to: (1) excessive tropical precipitation, (2) high sensitivity of modeled precipitation and surface air humidity to SST, (3) a lack of sensitivity of cloud amount to precipitation, and (4) a lack of sensitivity of stratus cloud formation to SST. GENMOM produces a cold SST bias in the Pacific Basin along with a confined cold tongue, both of which Lin (2007) note as factors that result in a double ITCZ. Consistent with Lin (2007), GENMOM does not produce a double ITCZ when coupled to a slab ocean. The split ITCZ problem can potentially be resolved by improving resolution of tropical ocean-atmosphere feedbacks.

3.5 Oceanic fields

The global surface and subsurface patterns of ocean potential temperatures simulated by GENMOM are consistent with the WOA05 data (Fig. 7); however, anomalies reveal biases exceeding 2°C (Fig. 7c), similar to other AOGCMs (Ran-

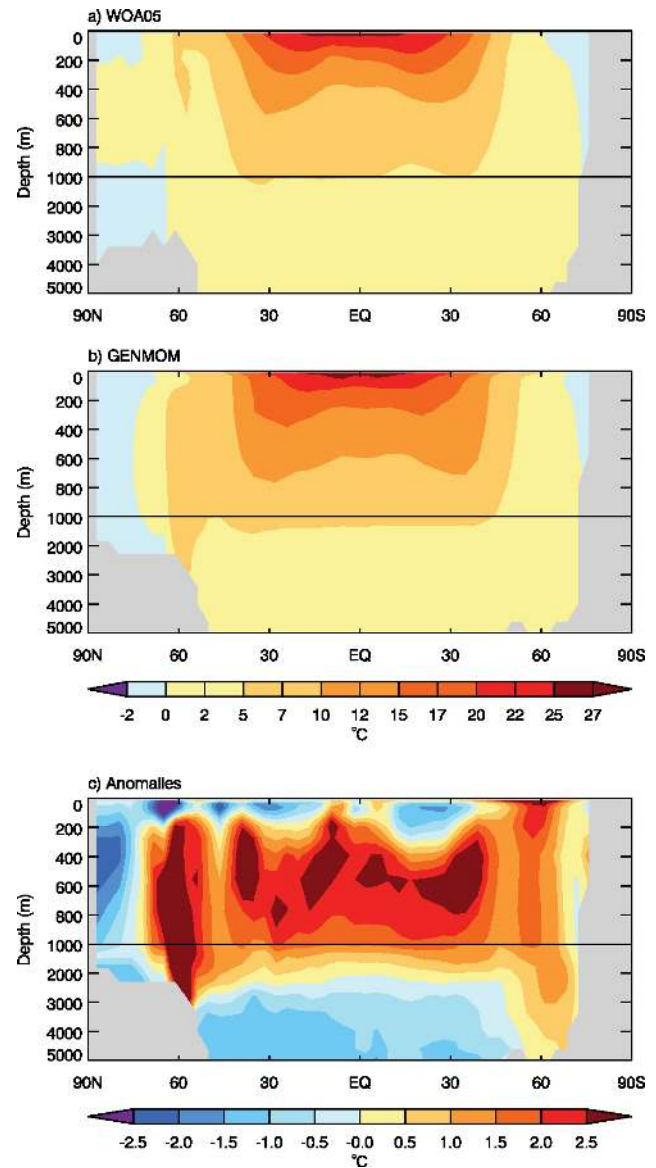


Fig. 7. Mean-annual zonally averaged ocean potential temperature profile. (a) Observed (WOA05), (b) GENMOM, (c) Anomalies, calculated as GENMOM – observed.

dall et al., 2007, Supplementary Materials, Pages 60–61). The warm bias in the near-surface of the Southern Ocean is consistent with the surface temperature anomalies shown in Fig. 4b. The Southern Ocean warm bias is attributed to weak ocean overturning and, perhaps, indicates the need for incorporating the Gent-McWilliams mixing scheme (Gent and McWilliams, 1990). The warm bias in the tropical ocean mid-depths is attributed to weakened simulated upwelling and the use of a relatively high vertical diffusion coefficient ($0.35\text{ cm}^2\text{ s}^{-1}$) that is prescribed to maintain reasonable ocean overturning; too much heat is diffused from the surface to the mid-depths.

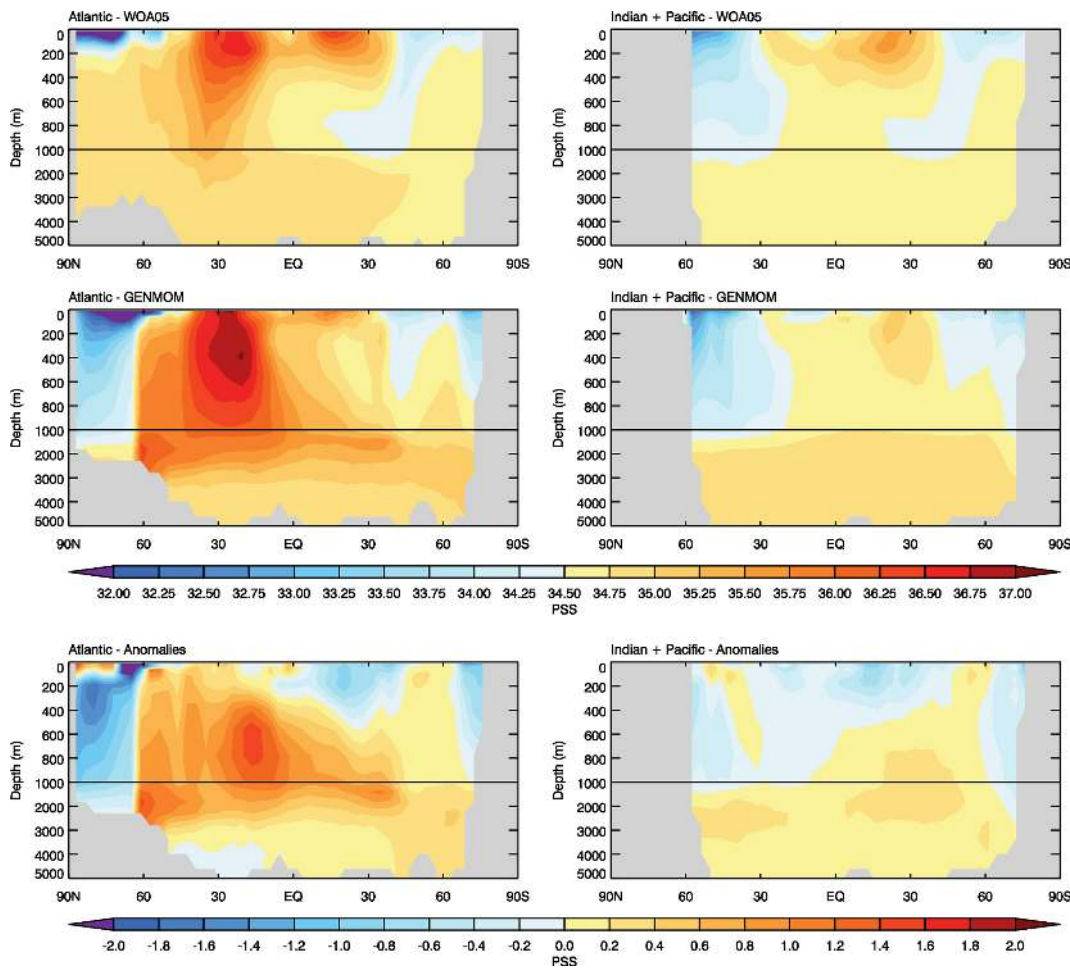


Fig. 8. Mean-annual zonally averaged ocean salinity profile for both observed (WOA05) and simulated (GENMOM) for the Atlantic Ocean (left), Indian and Pacific Oceans (right), and anomalies between observed and simulated.

GENMOM captures the zonal distributions of salinity with depth both between and within the Atlantic and Indian + Pacific basins (Fig. 8). Salinity in the Indian + Pacific Oceans matches well to observations with much of the zonal bias being less than ± 0.2 PSS. Relative to the WOA05 data, in the Atlantic, GENMOM simulates lower salinity waters at high latitudes and higher salinity waters in the northern mid latitudes; the maximum centered on 30° N exceeds observations and the maximum at 30° S underestimates observations. A $1+$ PSS salinity bias in the northern mid latitudes between 400–1000 m is attributed to a build up of salinity in the Gulf of Mexico caused by weaker-than-observed circulation associated with the coarse resolution of ocean orography. Similarly, the low salinity bias north of 60° N is associated with reduced northward penetration of the North Atlantic Drift into the Arctic, again due to the coarse resolution of ocean orography, and weaker-than-observed Atlantic Meridional Overturning Circulation (AMOC). Similar to ocean temperature, the GENMOM salinity anomalies are comparable with those of other AOGCMs (Randall et al., 2007, Supplementary Materials, Pages 66–68).

We compare simulated global and basin ocean overturning for the full 300-year GENMOM with observations (Fig. 9). The last 30 years of the simulation coincidentally displayed one of the weakest periods of overturning in the 300-yr simulation, so we use the full 300-yr record as more representative because higher frequency (decadal) variability is smoothed out in the average. Globally, GENMOM produces overturning that is similar in pattern to that of the GECCO data. The most notable shortcoming in the GENMOM simulation is that the strength and depth of the Deacon Cell, which is characterized as deep clockwise meridional circulation in the Southern Ocean driven by windstress, are poorly captured. Wind velocities across the Southern Ocean are weaker-than-observed (Fig. 3) thereby failing to produce sufficient windstress to drive deep overturning (Toggweiler and Samuels, 1995; Sijp and England, 2009). The weak westerly winds are likely due to the coarse meridional resolution (Held and Phillipps, 1993; Tibaldi et al., 1990) and may also contribute to weak AMOC biases in non-flux corrected models with coarse atmospheric resolution, as found in earlier studies (Bryan et al., 2006; Schmittner et al., 2011). Failure

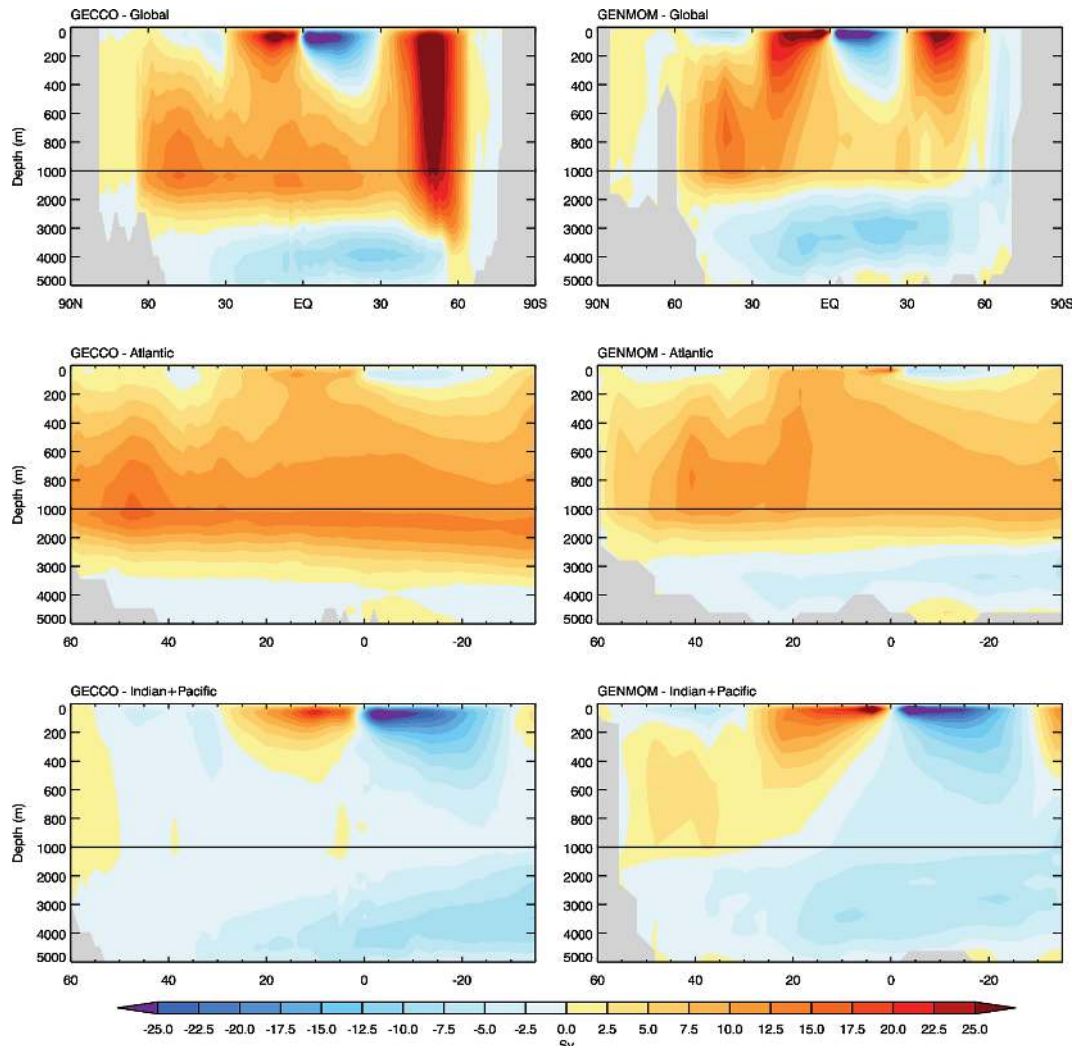


Fig. 9. Ocean overturning for both observed (GECCO) and simulated (GENMOM, full 300-year simulation) globally (top), Atlantic Ocean (middle), Indian and Pacific Oceans (bottom).

to simulate the Deacon Cell also contributes to the warm Southern Ocean temperature bias by limiting the addition of upwelled cold waters.

The simulated AMOC is somewhat weaker in strength than that of the GECCO data. The maximum AMOC strength over the 300-year simulation is 14.5 ± 0.9 Sv, which is lower than the observed 16–18 Sv range. (The models used in the IPCC AR4 generally fall between 12–20 Sv (Meehl et al., 2007b; Schmittner et al., 2005).) The combined Indian Ocean and Pacific Ocean overturning matches well with observations with the exception that GENMOM simulates deeper-than-observed clockwise overturning in the northern tropics, which may imply the vertical diffusion coefficient is too high.

The major Atlantic surface currents are well simulated in GENMOM, with the exception of the Gulf Stream, which is too weak (Fig. 10). The Antarctic Circumpolar Current flowing through the Drake Passage is well resolved, as is the con-

tinuing flow to the South Atlantic Current. In the Pacific, the equatorial currents are well simulated, but the North Equatorial Counter Current is not present and the North Equatorial Current is weaker than that of the observations. The Kuroshio Current is well placed but slightly weaker than that of the observations. The California Current is absent from the GENMOM simulated surface currents. Both the Humboldt Current and Antarctic Circumpolar Current are weaker than observed. The strength of the Antarctic Circumpolar Current through the Drake Passage is $\sim 35\%$ weaker than the 119 Sv estimated from the GECCO reanalysis. In the Indian Ocean, GENMOM simulates the Indonesian Throughflow well. The Indonesian Throughflow is 12.7 ± 0.8 Sv, compared to the observed estimates of 9.3 ± 2.5 Sv (Gordon et al., 1999) and 13.2 ± 1.8 Sv (Lumpkin and Speer, 2007).

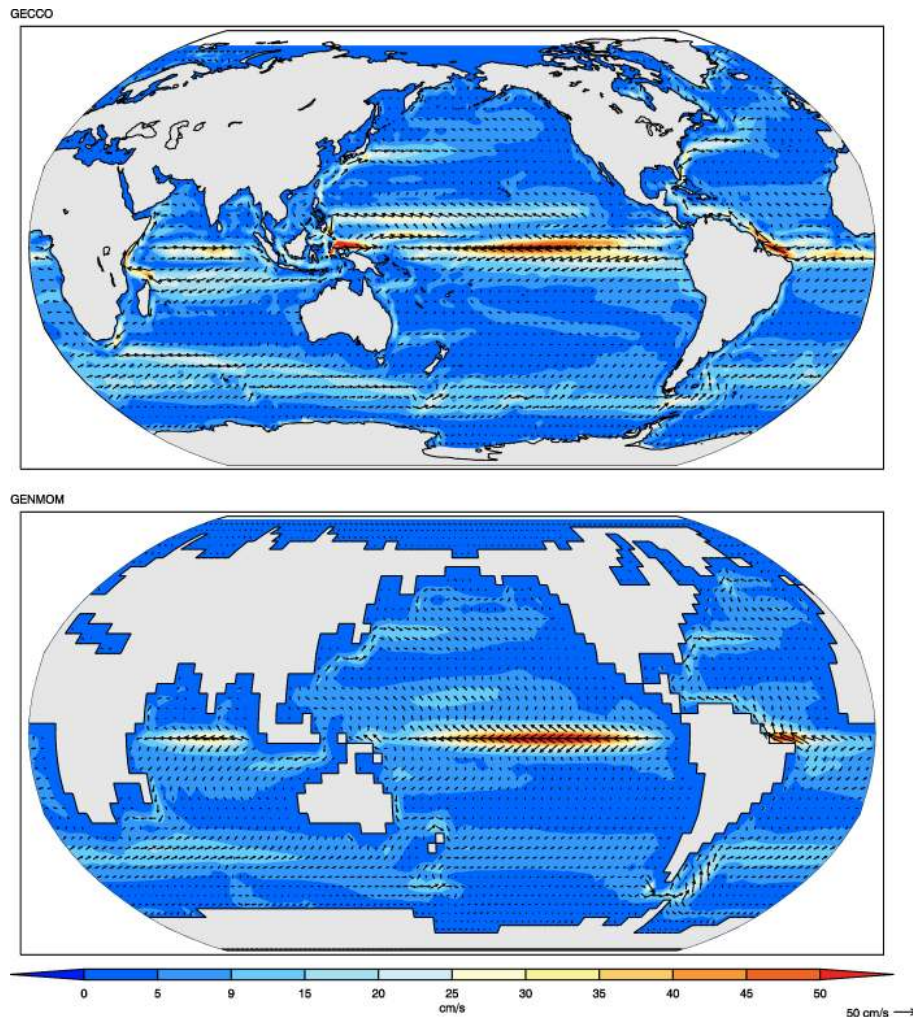


Fig. 10. Annual global surface currents for both observations (GECCO) and GENMOM simulated.

Surface currents in the northern Indian Ocean are modulated by the monsoon, where currents flow westward during winter and eastward during summer. The annually averaged surface currents in Fig. 10 show westward flow dominating in GENMOM whereas eastward flow dominates in the observations. The reversal of the surface currents is most noticeable in GENMOM during winter and spring.

GENMOM reproduces the observed NH winter and summer sea-ice extent and concentration (Fig. 11). Sea ice extends too far into the Norwegian Sea during both winter and summer and too far into Hudson Bay during winter. The excessive sea-ice in the Norwegian Sea is likely due to the weakened AMOC, which does not transport enough warm, mid-latitude water northward. Sea ice is deficient around Antarctica during both seasons. The lack of extent and concentration is a direct result of the persistent warm temperature bias in the Southern Ocean.

4 Discussion

We present the first formal evaluation of the new AOGCM GENMOM, a non-flux corrected model comprised of GENESIS 3 atmospheric model, the MOM2 ocean model and LSX land-surface model. The spectral resolution of T31 for both atmosphere and ocean is used during this evaluation. The simulated global 2 m air temperature is 0.6°C warmer over oceans and 1.3°C colder over land. The jet stream structure and major planetary features of sea level pressure are well captured by the model. GENMOM produces a realistic mean planetary-scale wind structure that is needed to produce the correct position of stormtracks. The 500 hPa ridges and troughs are well simulated, as are the seasonal surface pressure cyclones and anticyclones.

The annual surface temperature gradient, spatial distribution, and the annual distribution of precipitation compare

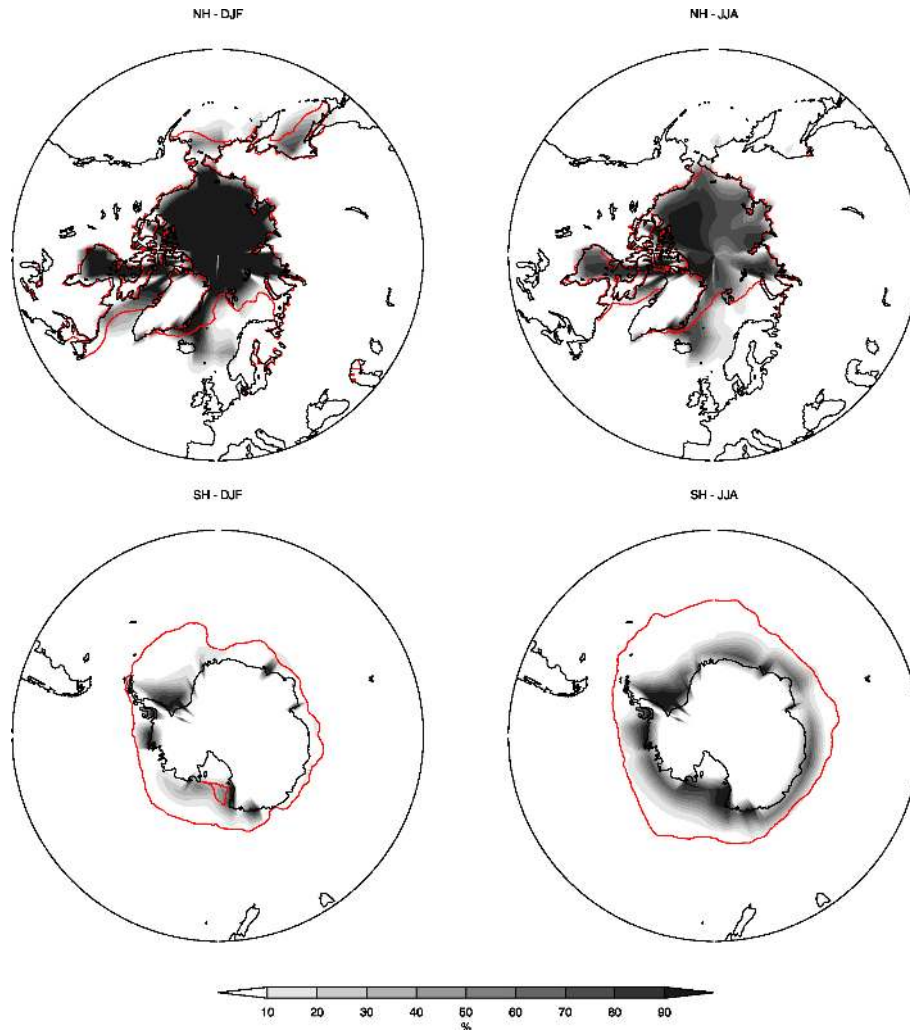


Fig. 11. Fractional sea-ice extent. HadISST v1.1 15% observed contour plotted in red.

well to observations and are on par with the three other AOGCMs. Cold SST anomalies in the Norwegian Sea are explained by excessive sea-ice in both winter and summer, which is in turn caused by weak Atlantic Ocean overturning. A warm bias in the Southern Ocean is attributed to a weak ocean overturning resulting in a poor simulation of the Deacon Cell, which suppresses associated cold water upwelling in the Southern Ocean. GENMOM fails to resolve adequately the South Pacific Gyre, which results in a warm SST bias in the eastern Pacific Ocean and weak anticyclonic atmospheric circulation around the gyre. GENMOM simulates a double ITCZ when coupled with the OGCM, which is not present when GENESIS is coupled to a slab ocean.

The global ocean temperature is generally well simulated, with the exception of a warm bias between 200–1000 m in the tropics and mid-latitudes. The warm bias is attributed to weak global overturning and the use of a high value of the vertical diffusion coefficient, which was needed to main-

tain realistic global ocean overturning. Salinity is generally well simulated, but with a fresh bias in the North Atlantic caused by underrepresentation of narrow channels (i.e., the Norwegian Sea) at T31 model resolution and a 1+ PSS salinity bias in the northern mid latitudes originating in the Gulf of Mexico. Ocean overturning is simulated with the correct spatial pattern, but is generally weaker-than-observed. We attribute the weak meridional ocean overturning to (1) weak and northwardly displaced westerly winds in the SH due to coarse topography and (2) a narrow and shallow Drake Passage also due to coarse orography.

Most ocean surface currents are well simulated by GENMOM, with the exception of narrow currents such as the Gulf Stream and the Kuroshio Current that are weaker-than-observed again due to the coarse T31 resolution. Northern Hemisphere Sea-ice is well simulated with the exception of excess sea-ice in the Norwegian Sea. However, the SH sea-ice extent is too small compared to observations. Both NH

and SH deficiencies are linked to weak ocean overturning. The use of the Gent-McWilliams mixing scheme should be considered in future GENMOM simulations with the expectation that this scheme should cool the Southern Ocean and increase SH sea-ice (Wiebe and Weaver, 1999).

The evaluation performed here demonstrates that the first generation of GENMOM produces a realistic climatology. Deficiencies in the oceanic component of the model discussed here provide guidance for improving GENMOM. The addition of a coupled ocean model allows GENMOM to be used to investigate past climates and to study phenomena such as ENSO that require dynamic ocean-atmosphere interaction.

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