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First Reprocessing of Southern Hemisphere Additional Ozonesondes (SHADOZ) Ozone Profiles (1998-2016). 2. Comparisons with Satellites and Ground-based Instruments

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KEY POINTS:

Data from 14 long-term SHADOZ station have been reprocessed according to best practices to correct for ozonesonde instrument variability.

Comparisons of total column ozone among reprocessed sonde data, satellite and ground-based instruments agree within 2% for 12 of 14 stations.

In the tropics the 10-station data bias among stratospheric ozone profiles has been markedly reduced and a tropospheric wave-one remains.

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Abstract. The SHADOZ network was assembled to validate a new generation of ozone-monitoring satellites and to better characterize the vertical structure of tropical ozone in the troposphere and stratosphere. Beginning with nine stations in 1998, more than 7000 ozone and P-T-U profiles are available from 14 SHADOZ sites that have operated continuously for at least a decade. We analyze ozone profiles from the recently reprocessed SHADOZ dataset that is based on adjustments for inconsistencies caused by varying ozonesonde instruments and operating techniques. First, sonde-derived total ozone column amounts are compared to the overpasses from the EP/TOMS, OMI and OMPS satellites that cover 1998-2016. Second, characteristics of the stratospheric and tropospheric columns are examined along with ozone structure in the tropical tropopause layer (TTL). We find that: (1) Relative to our earlier evaluations of SHADOZ data, in 2003, 2007 and 2012, sonde-satellite total ozone column offsets at 12 stations are 2% or less, a significant improvement. (2) As in prior studies, the ten tropical SHADOZ stations, defined as within $\pm$19 degrees latitude, display statistically uniform stratospheric column ozone, 229 $\pm$ 3.9 DU, and a tropospheric zonal wave-one pattern with a 14 DU mean amplitude. (3) The TTL ozone column, which is also zonally uniform, masks complex vertical structure; this argues against using satellites for lower stratospheric ozone trends. (4) Reprocessing has led to more uniform stratospheric column amounts across sites and reduced bias in stratospheric profiles. As a consequence the variability in total column ozone now averages 5%.

1. INTRODUCTION

1.1. Design and Features of the SHADOZ Network

The Southern Hemisphere ADditional OZonesonde network (SHADOZ; refer also to *Acronym List* below) was initiated in 1998 as an international partnership with both technological and scientific goals that required augmenting the number of tropical ozone soundings in the troposphere and stratosphere [*Thompson et al.,* 2003a; 2011*a,b;* 2012]. The ozonesonde data are collected from electrochemical concentration cell (ECC) type [*Komhyr,* 1969] sensors launched with a standard radiosonde. Details of ozonesonde-radiosonde pairings used at SHADOZ stations are given in the archival papers [*Thompson et al.,* 2003*a,* hereafter referred to as T03; *Thompson et al.,* 2003*b*; *Thompson et al.,* 2007, hereafter referred to as T07; *Thompson et al.,* 2012, hereafter referred to as T12] and in the companion paper to this one [*Witte et al.,* 2017a; referred to hereafter as Witte17a].

The original spatial coverage of SHADOZ was determined by two requirements: (1) that the network consist of existing stations; and (2) full zonal coverage to resolve an equatorial "wave-one" feature observed in satellite total ozone [*Fishman and Larsen,* 1987; *Shiotani*, 1992; *Kim et al.,* 1996; *Thompson et al.,* 2000; T03; *Thompson et al.*, 2003*b;* *Sauvage et al.,* 2006]. At the initiation of SHADOZ in 1998 there were nine stations meeting these criteria, all in the southern hemisphere, hence the name of the network [T03]. Stations north of the equator joined SHADOZ as follows: Kuala Lumpur (in 1999; *Yonemura et al.,* [2002]); Paramaribo (1999 [*Peters et al.,* 2004; *Fortuin et al.,* 2007]); Costa Rica (in 2005; *Selkirk et al.,* 2010); Cotonou (operated 2004-2007; *Thouret et al.,* 2009); Hanoi, where soundings began in 2004 [*Ogino et al.,* 2013]. Hilo, Hawaii, with a record extending back to the 1980s, joined SHADOZ in 2009. The 14 stations that have operated at least a decade during SHADOZ and that are covered in this paper, appear in **Figure 1.** More than 7000 sets of ozone and pressure-temperature-humidity (PTU) profiles from a total of 17 stations that have been part of SHADOZ and are available at the website: <https://tropo.gsfc.nasa.gov/shadoz>.

For the first 10-15 years of SHADOZ the data were used principally in three ways. First the profiles were used to create climatologies for satellite algorithms [*McPeters et al.,*

2007; *McPeters and Labow,* 2012; *Labow et al*. 2015], to evaluate chemical-transport models [*Martin et al.* 2002; *Stevenson et al.,* 2006; *Kaminski et al.* 2008] and as a reference for coupled chemistry-climate models [*Eyring et al.,* 2005] in intercomparison exercises that support UNEP/WMO Ozone [*WMO,* 2007; 2011; 2014] and the IPCC 2007 Assessment. Statistical approaches have been used to create more geographically coherent climatologies [T12; *Tilmes et al.,* 2012; *G. Liu et al.,* 2013; *J. Liu et al.,* 2013]. Self-organizing maps of sondes, in particular, capture meteorological and chemical impacts on profile structure [*Jensen et al.* 2012; *Stauffer et al.*, 2016].

Second, ozone structure near the tropopause has been studied, in the so-called "tropopause transition layer" or "tropical tropopause layer" (TTL; *Folkins et al.,* 2002; *Takashima and Shiotani,* 2007), with ozone-water vapor relationships captured in the relatively high-accuracy hygrometer data for H_2O [*Fujiwara et al.,* 2001; *Vömel et al*., 2002]. Signatures of convection, waves, and climate oscillations (ENSO, QBO) dominate ozone interannual variability in the free troposphere (FT) and lower stratosphere (LS). *Selkirk et al.* [2010] and *Thompson et al.* [2010] quantified wave impacts on Costa Rican sondes launched during the Tropical Composition, Clouds and Chemical Coupling (TC^4) campaign. At Ascension, even though convection is less prevalent than over the western Pacific and Indian Oceans, profiles with the lowest mean tropospheric ozone mixing ratio display an S-shape year-round [T12]. This is presumably a combination of the regional subsidence along with convection in some months (Figures 4 and 8 in T12; Figure 4 in *Jensen et al.,* [2012]). In *Thompson et al.* [2011a] gravity waves were identified through ozone laminae within the upper troposphere UT) and TTL. A gravity-wave index constructed for 12 SHADOZ stations captured responses to ENSO episodes.

Third, satellites use SHADOZ profiles for validation of tropospheric and/or stratospheric data (e.g., Aura validation; Special Issue of *J. Geophys. Res.,* 2007-2008) and newer sensors like OMPS on Suomi-NPP, IASI and GOME-2 (**Figure 2)**. SHADOZ data have also been used to derive ozone trends [*Heue et al.,* 2016]. In *Randel and Thompson* [2011] a composite SAGE II and SHADOZ dataset (1984-2009) for ozone over eight equatorial stations, defined a lower stratosphere (LS) trend of -(2-4)%/decade. Two subtropical sounding stations with data prior to the start of SHADOZ, Irene and La Réunion, displayed large positive trends in FT ozone, not in summer or spring, when biomass fires are

widespread, but in winter [*Thompson et al.,* 2014]. In *Gebhardt et al.* [2014], combined SCIAMACHY and SHADOZ data for 2002 to 2012 show flat to slightly increasing ozone in the equatorial LS. *Hubert et al.* [2016] employed sonde profiles to evaluate satellite drift over the past 10-15 years. To date, it appears that most operational satellites (14 limb sounders were evaluated by *Hubert et al*., [2016]) are quite stable in the lower and middle stratosphere. To be used as references for satellite drift evaluation, the sonde precision needs to be 3-5%, somewhat better than the level achieved in the past 10-15 years.

1.2 Quality Assurance: Technological Aspects of SHADOZ

An essential technological goal of SHADOZ is to adopt, when appropriate, recommendations for sonde technique and data processing when researchers have used laboratory and field test results to produce consensus-based recommendations for data handling. The beginning of SHADOZ coincided with a series of JOSIE (Jülich [Germany] Ozonesonde Intercomparison Experiments) in test chambers [*Smit and Kley,* 1998; *Smit and Straeter,* 2004], as well as laboratory studies elsewhere [*Johnson et al.*, 2002]. ECC ozonesonde manufacturers modify materials from time to time [*Komhyr,* 1986; *Komhyr et al.,* 1995] and several variations of the potassium iodide sensing solution that reacts with ozone molecules are widely used [Witte17a]. Furthermore, the motor-driven Teflon piston pump that draws air into the cells of the ECC instrument drops in efficiency as the sonde ascends into the stratosphere. Several groups have employed different formulae to correct for this effect, in a "pump efficiency correction factor" or PCF. Thus, different combinations of instrument manufacturer, sensing solution type (SST) and PCF can lead to divergent values of ozone partial pressure. These variations prompted the World Meteorological Organization (WMO) to support the formation of the World Ozonesonde Calibration Chamber System [*Smit and Kley,* 1998] in Jülich, and, starting in 1996, to sponsor the JOSIE series of chamber tests. In JOSIE experiments to date, different types of sondes, not all of them ECC, measured ozone introduced into the chamber at changing temperatures and pressures that simulate a balloon ascent to 10 hPa. A UV photometer [*Profitt and McLaughlin,* 1983] supplies the reference ozone measurement to which the sonde reading is compared.

The JOSIE-2000 campaign [*Smit et al.,* 2007] accommodated SHADOZ by testing combinations of instrument type and SST in chamber simulations that followed the

temperature and typical ozone trace of a tropical sounding. Clear biases from differences in the instrument manufacture (there are basically two ECC hardware types) and SST were apparent. In 2004 the BESOS field campaign, conducted in Wyoming USA, with a standard UV photometer and a gondola of 18 ozonesondes on a large balloon, displayed similar biases among sonde types [*Deshler et al*., 2008]. These intercomparison activities led to recommendations of certain ozonesonde instrument-SST combinations for more consistent sonde results and served as the basis for WMO/GAW standard operating procedures [*Smit and ASOPOS,* 2014]. Several SHADOZ stations modified their technique and/or reprocessed data during the period 2000-2005, improving agreement between total ozone values from the sondes and co-located ground-based instruments and total ozone satellite instruments [T07; T12] relative to the first comparisons in T03. JOSIE and related activities have reduced the total uncertainty of the sonde measurement from ~15-20% to 5-10% [*WMO,* 2007; 2011]. Nonetheless, SHADOZ total ozone column amounts referenced to satellite (TOMS or OMI) columns as well as stratospheric ozone columns from the sondes indicated station-to-station biases [T07; T12]. Some of the biases followed discrepancies associated with variations in SST and instrument that were identified in JOSIE and BESOS.

1.3 Ozonesonde Reprocessing: Goals of this Study

Since 2010, workshops of the SI2N (SPARC-IO3C-IGACO-NDACC) activity and a series of Ozonesonde Experts Meetings have led to recommendations on how to reprocess soundings to compensate for the range of instrument type, SST, and various choices of background current and PCF [*Smit and ASOPOS,* 2012]. An important element of reprocessing data from different SST and ozonesonde instruments is to homogenize data with a "transfer function." *Deshler et al.* [2017] supply transfer functions for the major combinations of instrument type and SST. At present 6-8 providers of ozonesonde data, representing ~30 stations globally, have adopted the guidelines. Papers describing the reprocessing of Canadian network data [*Tarasick et al.,* 2016] and two European records [*Van Malderen et al.,* 2016] have appeared. In the past two years we have been reprocessing the SHADOZ profiles for 14 stations that have records of at least 10 years within the period 1998-2016. The basis for the reprocessing is *Smit and ASOPOS* [2012] with some modification for several stations as described in Section 2. Details for the reprocessing of data from the first seven stations appear in Witte17a. Additional

reprocessing, including for an 8th station, are given, along with ozone profile and column uncertainties in J. C. Witte et al., First reprocessing of Southern Hemisphere ADditional Ozonesondes (SHADOZ) profile records: 3. Uncertainty in ozone profile and total column, *J. Geophys. Res. Atmos.,* doi: 10.1002/2017JD027791, submitted, hereafter referred to as Witte17b. The complete reprocessed SHADOZ dataset will be released as SHADOZ v6.0. The present study evaluates the reprocessed data to date and addresses the following:

- How do total column ozone (TCO) amounts for the SHADOZ stations agree with satellite TCO in the period 1998-2016? How do the sonde ozone columns compare to co-located Dobson, Brewer, and SAOZ instruments at nine SHADOZ stations?
- How do stratospheric, tropospheric and TTL column ozone amounts for the tropical stations (defined as within ±19 degrees of the equator), calculated from the reprocessed data, compare to one another? Are the stratospheric and TTL ozone columns zonally uniform? What is the variability in TCO after reprocessing?
- What are tropospheric column ozone amounts at each station with the reprocessed data? What is the magnitude of the wave-one pattern in tropospheric ozone?
- Using a mean reference profile from the tropical SHADOZ stations, are there station biases in stratospheric ozone? Have biases changed from T07 and T12?

We summarize operating characteristics of all SHADOZ stations (**Section 2**) and present the results in **Section 3. Section 4** is a summary.

2. Data and Methods of Analysis

2.1 SHADOZ and other Ozone Instrumentation

The 14 SHADOZ sites (**Figure 1)** launch 2-4 times/month mostly between 0800 and 1400 Local Time. **Table 1** presents station location and numbers of soundings that are used in our analyses, along with independent total ozone instrumentation. The station labeled as Costa Rica represents sonde launches near San Jose, the capital, that have changed location several times since late 2005.

SHADOZ archives ozone profiles with records of temperature, pressure, and relative humidity from standard radiosondes. Over the 19-year record covered here, Vaisala RS-80,

RS-92, MW-41 and Intermet (iMet) radiosondes have been used at most stations (**Table 2**). Exceptions are Réunion, where a Meteo Modem radiosonde has been employed, and Kuala Lumpur where various ozonesonde-radiosonde combinations appear with four different radiosonde manufacturers. The Lockheed-Martin Sippican radiosonde, a successor to Viz, is used at Natal and was employed at Ascension from 1998 through mid-2010. When Ascension launches resumed in March 2016, EnSci ozonesondes replaced Science Pump Corporation (SPC) instruments and an iMet radiosonde was used. Details of each station's equipment are found at <https://tropo.gsfc.nasa.gov/shadoz>; every record's meta-data lists instrument type and usually serial number for ozonesonde and radiosonde. Details of the SHADOZ radiosonde/ozonesonde systems for seven stations appear in Witte17a. In the case of Watukosek, only data from mid-2001, taken using Strato software with EnSci ECC and RS80 sondes have been reprocessed (reprocessing of 1999-2001 EnSci/RS-80 data with a V03 data system is pending). Hilo, Fiji, Samoa and San Cristóbal sondes, reprocessed by C. Sterling and B. J. Johnson (*Personal Communication,* 2017; reprocessing method referred to as Sterling17 in **Table 1**), switched from Vaisala RS-80 radiosondes to iMet or RS-92 after 2010. **Table 2** shows the time period of each radiosonde at the nine stations where there have been changes. For comparisons with satellite and ground-based total ozone instruments, data compromised by balloon bursts lower than 15 hPa are not used. Above 10 hPa or burst, extrapolation to total column is made with the climatology of *McPeters and Labow* [2012]. No normalization to the total ozone reading of a satellite or ground-based instrument is made. Ground-based data are taken from the WOUDC and NDACC public archives.

Sonde total ozone comparisons are made with satellite measurements from three instruments of similar design that operated in the 1998-2016 period (**Figure 2**). These are: NASA's Earth Probe/TOMS (Total Ozone Mapping Spectrometer) instrument, covering sonde launches from January 1998-September 2004; the Dutch/Finnish OMI (Ozone Monitoring Instrument) on board NASA's Aura spacecraft (September 2004 to the present); OMPS (Ozone Mapping Profiler Suite) on the NASA-NOAA Suomi-National Polar-orbiting Partnership from February 2012 to the present. In the latter period both OMI and OMPS overpass total ozone are used; there are more OMPS measurements because of post-2008 degradation in the OMI detector array. Not every sonde record is evaluated with the

satellite because overpass data are screened for cloudiness greater than 60% and the satellite distance is limited to < 200 km.

2.2 Reprocessed Ozonesonde Data

A summary and reference for the reprocessing of each station appears in **Table 1**. Details for seven stations that were reprocessed at NASA/Goddard appear in Witte17a: Natal, Ascension, Irene, La Réunion, Kuala Lumpur, Hanoi, and Watukosek (**Figure 1**). In Witte17a, comparisons of sonde total ozone for 1998-2015 with TOMS and OMI satellites are made with original (SHADOZ version 5.2) and reprocessed data to demonstrate improvements in the satellite-sonde offset. Also in Witte17a mid-stratospheric ozone from the sonde profiles before and after reprocessing is compared to coincident Aura/MLS data from September 2004 through 2015. Notable improvements appear between 65 hPa and 10 hPa for six of the seven stations. Since Witte17a was completed, we added all the 2016 reprocessed SHADOZ data to our seven-station reprocessing and completed reprocessing of Nairobi data [Witte17b]. Witte17a has since reprocessed Costa Rica data in the same way. A transfer function, developed by *Deshler et al.* [2017], was applied to the La Réunion data after 2007 to correct for a solution change. These data are now homogenized to an EnSci/0.5%, half buffer SST [Witte17b].

The four stations with data reprocessed by the NOAA ESRL/GMD Ozone group (Hilo, Samoa, San Cristóbal, Fiji) used a 2.0% unbuffered SST with EnSci's for the 1998-2005 period, after which they switched to a 1.0% 1/10th buffer SST (Sterling17, **Table 1**). The 1998-2005 data have been reprocessed in the same way as the Watokusek data that are based on the EnSci/2.0% unbuffered SST [Witte17a]. A transfer function based on *Deshler et al.* [2017] is applied to the San Cristóbal 1998 - 2006/11/02 data to convert ozone measurements made with a SPC/6A sonde to an EnSci/Z equivalent. This gives roughly a 4% increase in the ozone profile measurements.

Paramaribo data have been reprocessed by KNMI (M. Allaart and A. Piters, *Personal Communication*, 2017) according to the O3S-DQA [*Smit and ASOPOS,* 2012] guidelines; where background currents required adjustment, the guideline of *Newton et al.* [2016] was followed.

2.3 Analyses. Tropopause and TTL Definitions.

In the analyses that follow, we refer to the TTL and the tropopause, terms that sometimes vary from one study to another. The TTL is defined as a region in which both tropospheric and stratospheric properties are found in terms of constituent mixing ratios and temperature gradients, wave activity, radiative heating rates and other thermodynamic quantities [*Gettelman and Forster,* 2002; *Fueglistaler et al.,* 2009]. We adopt 15-18.5 km for the TTL (**Figure 3**) that encompass the minimum and maximum tropopause values for tropical sites within 19 degrees of the equator**.** This range is ~1 km higher than *Fueglistaler et al.* [2009]. However, our TTL definition encompasses the locations of the tropical cold-point tropopause (CPT, *Selkirk et al.,* [2010]), the thermal lapse-rate tropopause and the ozonopause, i.e., an ozone tropopause defined as the height at which there is a sharp gradient in ozone concentration at the base of the stratosphere [*Bethan et al.,* 1996; *Sivakumar et al.,* 2011; T12]. We use the *Bethan et al.* [1996] definition for the tropopause. The values in **Figure 3** are virtually identical for 10 stations, averaging 17.1 km, and there are few tropopause heights below 14 km.

3. Results and Discussion

3.1 Sonde Total Ozone Columns Compared to Ground-based Instrument and Satellite Overpass Total Ozone

For the 14 stations **Figure 4** shows a time-series of integrated total ozone from the sonde (red) and the daily TOMS/OMI/OMPS overpass total. The order of the individual panels is west to east by longitude, starting at 180W longitude. The satellite total ozone from TOMS appears in dark gray, OMI total ozone is in black and OMPS total ozone is in silver. For nine stations (**Table 1**) total column ozone from a co-located instrument, a Brewer Spectrometer, a Dobson or SAOZ, is also shown in **Figure 4** (blue). The lower panel of each time-series summarizes the ratio of sonde total ozone to the corresponding satellite overpass or ground-based measurement. At the right in each figure there is a histogram of offsets between the sonde total from the reprocessed data and the satellite (yellow). There are five stations without ground-based total ozone instruments: Ascension, Hilo, San Cristóbal, Costa Rica, and Fiji. For the nine stations with ground-based instruments, a histogram of the sonde total and the ground-based instrument ozone total (hatched in **Figure 4**) also appears in the right panel. Comparisons of total ozone from the ground-

based instrument and the concurrent satellites appear in **Figure 5**. In **Figure 4** sonde and satellite comparisons display a high fraction of near-zero differences for every station except two datasets where the distribution peaks are skewed closer to +5% (Paramaribo) and -5% (Kuala Lumpur). All but two of the sonde-satellite discrepancies show the sonde total to be lower than the satellite (**Table 3**), similar to prior SHADOZ satellite-sonde total ozone comparisons [T03; T07; T12; Witte17a].

Among the 14 stations all but Kuala Lumpur and Paramaribo have an average agreement with the satellite overpasses within 2% absolute or better (**Table 3,** column 4). This represents a significant improvement over the sonde data currently archived, as described in Witte17a (Table 4) and Sterling17. Overall, sonde-satellite agreement is improved relative to the total ozone comparisons in T07 (EP/TOMS satellite reference, for sondes launched in 1998-2004) and in T12 (Figure 11, for OMI comparisons, 2005-2009).

3.1.1 Sonde-Satellite-Ground-based Time-series at Individual Stations

Samoa is one station for which the total ozone agreement has improved considerably since the T03 evaluation where the sonde-TOMS ozone offset, based on 1998-2000 soundings, was ~10%. For 1998-2004, with some instrument and sensing solution changes after JOSIE-2000 [*Smit et al.,* 2007], satellite and Dobson comparisons implied that the sondes were 5-7% too low [T07]. Reprocessing improves the satellite agreement on average, from 1998-2016, to 0.9% mean difference (column 4 in **Table 3**). The Dobson total ozone also agrees well with the sonde TCO. This is seen in the histograms in the right panel of **Figure 4a** that display a relatively tight and Gaussian distribution of sonde-satellite (yellow) and sonde-Dobson (hatched) total ozone.

Referring to the lower panel in **Figure 4a,** the sonde/TOMS ratio (1998-late 2004) averages 0.97-0.98. There are other features to note with the Samoa data (**Figure 4a**). First, the sonde/Dobson ratio for 1998-late 2004 is slightly higher than sonde/TOMS. This implies that TOMS total ozone is greater than the Dobson; that is consistent with the Dobson/TOMS ratio peaking at less than 1.0 in **Figure 5a** (yellow shading). Second, from 2006 through 2010 in the OMI period, the sonde/Dobson ratio and sonde/OMI are both closer to 1.0 (**Figure 4a**) than the 1998-2004 sonde-Dobson-TOMS comparisons (T07). This too agrees with the peaking of Dobson/OMI total ozone ratio peaking near unity in **Figure 5a** (blue hatched). Third, the close sonde-satellite-Dobson agreement continues

into mid-2014, when most of the sonde-satellite comparisons are based on OMPS. The ratios, both sonde/OMI and sonde/OMPS, declined on average from late 2014 through 2016. During OMPS there are many Dobson comparisons with the satellite (**Figure 4a)** but few sonde comparisons with the Dobson. The Dobson/OMPS (green hatched in **Figure 5a**) peaks slightly lower than the Dobson/OMI, but it is higher than Dobson/TOMS.

In **Figure 4b** for Hilo (elevation 10m), only satellite comparisons are displayed because the nearby Dobson spectrometer at Mauna Loa is 3.4 km above mean sea level. Hilo displays one of the best homogenized sonde records throughout the SHADOZ period. Because of its northern sub-tropical location, there is a large variation in total ozone, close to 100 DU in a highly regular cycle (see also Figure 10e in T12). Maximum total ozone over Hilo from February to May [*Oltmans et al.,* 2004; T12] during which the UT/LS displays extratropical ozone-rich air and a lower tropopause (Figure 3b in T12). In contrast, July-September are characterized by lower ozone throughout the free troposphere with a mid-troposphere minimum that resembles boundary-layer Hilo ozone (Figure 3b in T12), suggestive of convective mixing. The Hilo-satellite comparisons (**Figure 4b**) show similar patterns to the Samoa (**Figure 4a**) sonde ratios. For the TOMS to early OMI period, 1998-2005, the Hilo ratios average an equivalent of 1-2% lower ozone than the satellites (**Figure 4b**). After that, up through 2013, the sonde/satellite ratio clusters about 1.0, with nearly all data within 0.95-1.05. In the first two years of the OMI-OMPS overlap the sonde-satellite ratio resembles the 2004-2011 record. In mid-2014, as with Samoa, Hilo sonde ratios decline several percent with a number of values falling below 0.9. This pattern is reflected in the histogram of satellite to sonde offsets (yellow at right of **Figure 4b**). The tail of negative values, corresponding to sonde lower than satellite, extends almost to -20%. However, the overall sonde-satellite difference is -1.2% (**Table 3,** sonde lower).

In **Figure 4c** the San Cristóbal time-series displays several data gaps, e.g., no launches in 2009-2011, with a drop in launch frequency in 2012-2015. The sonde/satellite ratios for the TOMS and early OMI era (1998-2005) averages 0.93. As with Samoa and Hilo (**Figures 4a,b**), the sonde/OMI ratio for 2006-2008 exceeds 1.0 but in 2012-2015, the ratio drops to ~0.95. The San Cristóbal histogram (right in **Figure 4c**) skews toward -20%, but 80% of the differences are close to zero, averaging a -0.9% offset, i.e., sondes have lower ozone.

The Costa Rican sondes that started in 2005 (**Figure 4d**) span the OMI and OMPS periods. They suggest a pattern somewhat different from the three Pacific stations illustrated in **Figures 4a-c.** In 2005 through most of 2007 the ratio is frequently greater than 1.0. From late 2007 through early 2015 ratios span 0.9 to 1.07, during which there seem to be two distinct phases. From late 2007 through early 2014 the ratios are predominantly less than 1.0, followed by many values > 1.0 after mid-2012 (note: more statistics partly result from frequent OMPS overpasses). After 2014, although the percentage of sonde/satellite ratios greater than 1.0 increases, the distribution shifts to more values < 0.9. In 2016 there are many sondes with ratios 0.80-0.90. The Costa Rica histogram (right in **Figure 4d**) displays negative values to -20% but the overall mean relative to OMI and OMPS is -0.2% (column 4, **Table 3**).

The sondes at Paramaribo (**Figure 4e**) represent the only reprocessed dataset for which the sondes display markedly higher ozone than the satellites and the ground-based instrument for most the period 1999-2016. Paramaribo is the SHADOZ station with the second largest ground-based dataset; the Brewer/satellite ozone ratios in **Figure 5b** are based on more than 5000 coincidences**.** Comparisons of Brewer total ozone with the three satellites agree well; as for Samoa (**Figure 5a**), the ozone from the ground-based instrument relative to the satellite varies among the three satellites at Paramaribo. TOMS has highest ozone (lowest ratio, yellow in **Figure 5b**) with sonde/OMI somewhat higher than sonde/OMPS. In **Figure 4e**, between 2011 and 2014 there is an almost monotonic increase in the sonde/satellite ratios that disappears by 2015. These persistent positive ratios, interpreted in the past as signifying a bias at Paramaribo relative to profiles of the other 9 tropical stations (Figure 13b in T12), disappear after 2015. Then the sonde/ satellite (OMI and OMPS) and sonde/Brewer ratios more closely resemble the four stations discussed above. Around 2015 there was a change in the sonde sensing solution; the data are being reprocessed again (A. Piters and M. Allaart*, Personal Communication,* 2017).

The Natal record (**Figure 4f**) displays several data gaps; the hiatus from 2011 to early 2014 was due to equipment malfunction, then obsolescence. As with the Pacific and South American stations (**Figures 4a-e)**, agreement among sonde, satellites and Dobson total ozone overall at Natal is very good (**Table 3**). The histogram of ratios (as % offset), right frame in **Figure 4f,** averages -1.7% for the composite satellite record (yellow); the Dobson

appears to peak slightly lower (hatched in **Figure 4f)**, skewing more negative than the satellite-referenced offsets. Similar to the five stations discussed so far**,** there is a 2014-2016 drop-off in sonde ratios, consistent among OMI/OMPS and Dobson, at Natal (**Figure 4f**), even though the instrumentation used at Natal differs from the other stations. The Natal time-series during the TOMS period is unique, with many of the ratios starting out in 1999 and 2000 at 10% lower than the satellite and Dobson [T03, T07]. From late 2000 until early 2002, the sonde/TOMS corresponds to 3-4% positive; the sonde also increases relative to the Dobson but mostly remains below 1.0. After early 2002 the sonde-Dobson and TOMS ratios are nearly all below 1.0 again, values that persist to late 2004 to 2006 as the satellite data transition from TOMS to OMI. In mid-2007 and throughout 2008 the sonde ratios relative to both OMI and Dobson become largely >1.0. Natal has > 3300 Dobson-satellite coincidences (**Figure 5c**) with all three satellite ratios peaking near 1.0.

The instrumentation, preparation technique and data-processing for Ascension (**Figure 4g**) were the same as at Natal from 1998-2010 when the Ascension launches were suspended. Data from 2016 onward are provided by an Ensci/iMet combination instead of LMS with SPC (**Table 2**). The sonde data at Ascension are fairly noisy and much of the distribution of ratios is skewed to values < 0.90 (more negative than 10% in the histogram in **Figure 4g).** Low ratios are prominent in the beginning of the record, from 1998 to late 2001. After that, from the late TOMS to early OMI period, 2002-2008, there is an increase in the factors, to a majority at > 1.0. A short gap follows, after which the values drop to an average ~0.95. After the longer gap, there is a tight clustering between 0.90 and 1.0 (histogram average ~-5%). The mean sonde-satellite offset is -2.4%.

There was a 4-plus year data gap in Irene (**Figure 4h**) largely due to equipment failure. The sonde/satellite and sonde/Dobson ratios are similar; this is also seen in the histograms at the right in **Figure 4h**. Most satellite offsets (yellow in the histogram), range from -2% to +4% with the mean satellite-sonde offset 2.0% (sonde higher, **Table 3,** column 4). As with many sonde-satellite comparisons, the ratios in 1998-2001 in the TOMS period are mostly < 1.0; this changes to mostly > 1.0 in 2002, a value that applies for OMI comparisons in the remainder of the first time-series and continues with the OMI and OMPS comparisons with sonde and Dobson from late 2012 through 2016. These variations are consistent with the statistics for satellite and Dobson coincidences at Irene (**Figure 5d**).

However, the Dobson/satellite differences are not large, similar to Natal (**Figure 5c**), although at Irene, the ratio is slightly smaller for Dobson/TOMS than it is at Natal.

The Nairobi sondes (**Figure 4i**) are in excellent agreement with the three satellites throughout the 1998-2016 period. The Nairobi histogram (right in **Figure 4i**) has the highest absolute number of comparisons between sonde and satellites compared to the other sites (scale up to 225). The peak is close to 1.0; the mean offset (**Table 3**) is -1.7%. During the earliest part of the SHADOZ record, Nairobi (**Figure 4i**) sonde/satellite and sonde/Dobson ratios are close to unity, which is higher than most of the other stations examined so far. These ratios dip in 2002-2003, then increase in 2004 and even more in the early OMI era. A preponderance of factors > 1.0 persists into mid-2010 when there is a short gap. From late 2010 throughout mid-2014 the sonde ratio relative to OMI and OMPS is close to unity. The Dobson/satellite statistics are summarized in **Figure 5e,** where the ratios cluster near 1.0; there is less data than for the satellites because of Dobson gaps in 2000-2004 and 2012-2016. The lowest ratios in **Figure 5e** (blue shading) correspond to the early OMI period (2005-2011) and are consistent with OMI total ozone being greater than TOMS, for which the Dobson ratio peaks slightly higher than 1.0, like Natal and Irene (**Figures 5 c,d)**.

As described in **Section 2,** the ozone record at La Réunion (**Figure 4j**) represents further reprocessing since Witte17a. Agreement among the sondes, TOMS and co-located SAOZ is consistent from 1998-2004 with ratios ~0.95 or a mean sonde offset of -5%. During the OMI period from 2005 to 2010, most of the sonde ratios increase to an average ~1.0. In 2011-2012, most factors are > 1.0. After 2012 there is a slight drop-off, with a few ratios below 0.9. These variations, taken across the 1998-2016 period, lead to a mean offset of -0.9% for sonde-satellite (yellow in histogram, **Figure 4j**) and sonde-SAOZ (hatched in histogram). In **Figure 5f** La Réunion SAOZ/satellite comparisons have a very high sample number, with ratios for all three satellites peaking near 1.0.

For Kuala Lumpur (**Figure 4k**) from 1998 through 2012, fewer than 10 sonde/satellite ratios were > 1.0, but in 2013, there is a jump to many more values > 1.0. This is followed by a noisier record with factors both greater and less than 1.0 from late 2013 to early 2016. If it is assumed that the Brewer is a stable reference, the satellite-Brewer total ozone comparisons (**Figure 5g**) explain these changes. The Brewer/TOMS ratio indicates a

higher satellite reading in 1998-2004 than the Brewer so the sonde/satellite ratios (gray in **Figure 4k**) are lower than for Brewer. For the OMI period (Brewer is lower than the satellite, **Figure 5g**), higher sonde/satellite ratios are expected. They increase slightly from 2005 through 2009 (**Figure 4k)** and much more after the re-start of launches in 2012. Most sonde/satellite ratios after 2012 are with OMPS ozone which the Brewer comparison (**Figure 5g**) suggests measures more ozone than OMI so many ratios fall to < 1.0 (**Figure 4k)**. Sonde offsets from the three satellites average -5.5% (yellow histogram, **Figure 4k**); the Brewer histogram (hatched in **Figure 4k**) follows the satellite but with fewer samples.

At Hanoi (**Figure 4l,** note the late 2004 start) there have been two Brewer instruments (**Table 1**). The first one spanned the beginning of the sondes in 2004 until 2012 but the noisy data are so divergent from the satellites that we have not used them. Only measurements from the second Brewer, with data from 2012 to 2014, are compared to the sondes; we note that the Brewer/satellite ratios (**Figure 5h**) are much higher than OMI and OMPS and are outside the guideline for ideal Brewer-satellite agreement (*Fioletov et al.,* 2008). The lower left panel in **Figure 4l,** as well as the small sample number in the histogram (hatched, at right), show that sonde-Brewer coincidences are also infrequent. Sonde-satellite offsets average -1.2%. The Hanoi sonde-satellite data across the OMI and OMPS period until 2014 are fairly consistent (not surprising with similar statistics relative to the Brewer, **Figure 5h**). From 2014 to 2016 the sonde/satellite ratios are nearly all <1.0. The sonde-satellite histogram in **Figure 4l** has a long negative tail.

The SHADOZ data at Watukosek (**Figure 4m**) are compared to a Brewer at Bandung, which is several hundred km to the west and at elevation 740 m (Watukosek is 50 m). Watukosek is also frequently polluted near the surface by local pollution [T03; T12]. Many Brewer readings have been removed in the comparisons, e.g., from 2009-2011, because they diverged > 10% absolute from the satellite readings. However, those that remain give ratios to OMI and OMPS that are well-centered about 1.0 (**Figure 5i**). The sonde/satellite ratios (lower frame in **Figure 4m**) also average close to 1.0. There is a mean -1.3% offset in the sonde-satellite histogram (yellow at right, **Figure 4m; Table 3**).

Fiji has no ground total ozone instrument (**Figure 4n**) and there are two year-plus gaps in the measurement record. For the TOMS and early OMI period, from 1998-2005 when the first gap began, sonde/satellite ratios typically fall in the range 0.95-1.0. As for

most of the other SHADOZ stations, the sondes record more total ozone than the satellites during the mid-OMI period, 2007-2013. In 2015 the ratios trend lower for both OMI and OMPS, particularly in 2016. The histogram for the satellite comparisons (right, in yellow) skews slightly negative, averaging 0.9% lower for the sondes (**Table 3**).

3.1.2 Residual Discontinuities in Total-column Ozone Time-series

The discussion of ratios in **Figure 4** shows that for nearly all of the SHADOZ stations reprocessing has not completely resolved the systematic variability exhibited in the original data [Witte17a]. In some cases, quite a bit of noise persists even though the mean TCO is closer to both satellite total ozone and ground-based instruments and is presumed to be more accurate than the original. **Section 3.1.1,** with satellite-ground-based-sonde comparisons over three different satellites, suggests that variations among the satellite sensors (discussed in Witte17b) and perhaps unevenly calibrated Dobsons or Brewers, is one reason for persistent discontinuities in sonde ratios after reprocessing.

Several additional factors may also contribute. First, there is inherent uncertainty in the ECC sonde measurement, which varies with altitude. This has been evaluated in Witte17b for reprocessed SHADOZ ozone profiles and for the column amounts. Secondly, the reprocessing prescribed in *Smit and ASOPOS* [2012] and *Deshler et al.* [2017] and applied in Witte17a, Witte 17b and Sterling17 corrects only for variations in the ECC sondes and their operating characteristics. The impact of the many ECC sonde-radiosonde combinations (**Table 2**) used in SHADOZ has only partially been addressed. For example, additional corrections were made by Witte17a and Sterling17 to compensate for what are assumed to be incorrect pressure readings from iMet radiosondes [*Stauffer et al.,* 2014] at five stations (**Table 2).** Vertical dashed lines in **Figures 4a,b,d, g** and **n** mark the first deployment of the iMet radiosondes at each station. However, these first-order corrections are not based on rigorous lab or field tests with ECC sondes. Looking at distinct changes in the sonde ratios after 2013, their coincidence with the iMet changeover is only straightforward for two datasets. One of those is for Fiji (**Figure 4n**). The other discontinuity that coincides with iMet is at Ascension (**Figure 4g**), although the last part of the LMS-SPC instrument combination (2009-2010) also displays sharply lower sonde ratios relative to the record up to 2009. At Samoa, there is a sharp dropoff in ratios at the start of the iMet period (2014 in **Figure 4a**) but there are clusters with values > 1.0 during

2015 and the latter part of 2016. At Hilo (**Figure 4b**) the falloff in the ratio lags the iMet introduction by about a year; a similar delay is observed for Costa Rica (**Figure 4d**).

An additional factor in the residual discontinuity of sonde ratios could be the ECC ozonesonde itself. For the six stations using the iMet radiosonde, the coupling is to EnSci/DMT ECC ozonesondes. This instrument may have been inadvertently modified in the 2011-2016 period as the manufacturer changed, first in 2012 when the original EnSci Corporation became part of Droplet Measurement Technology (DMT), followed by a shift to a new version of the EnSci Corporation in 2016. We tested this hypothesis for four stations that use the Ensci sonde and where the sonde ratio dropped after 2013: Samoa, Hilo, Costa Rica, Fiji. Several operating characteristics of the sondes, e.g., flow rate, background current, total column ozone, were also examined as a function of the sonde date since 2006, (using the 2Z serial number to track the ECC). Most parameters appeared to be independent of time, except for a lower TCO amount at two stations (**Figure 6**). Total ozone and the sonde/OMI (OMPS) ratio have dropped off ~15% at Costa Rica since late 2015 (**Figure 4d**); the relationship to serial number (**Figure 6, left**) suggests something might have changed after No. 25000 but there is no clear cause when meta-data are examined (this also agrees with findings by H. Vömel, *Private Communication,* 2017). At Fiji (**Figure 6, right**) there have also been a few anomalously low TCO values in 2016, after the 2Z #25000 period. However, analysis of instruments used at Samoa and Hilo, where sonde ratio falloffs after 2014 are observed (**Figures 4a,b**), does not show similar behavior. No chamber tests have been carried out over the time of the evolving EnSci instrument. Field tests with the EnSci sondes and the current suite of radiosondes (iMet, RS92, RS41, LMS) need to be conducted. Moreover, the residual discontinuity in sonde ratios with this first set of reprocessed SHADOZ data is a reminder that the current reprocessing guidelines do not systematically account for radiosonde instrument changes at long-term stations.

3.1.3 Summary of Reprocessed data and Satellite Total Ozone

Figure 7 summarizes mean total column ozone from TOMS-OMI-OMPS over each SHADOZ station in black (with $\pm$ 1σ denoted by the bars). The mean sonde total column ozone value is in red, with the $\pm$ 1σ range represented by the pink shading. **Table 3** summarizes the corresponding mean sonde-satellite offsets over the 1998-2016 period, except for Costa Rica and Hanoi, where the comparisons are for late 2004 or 2006 to 2016,

493 respectively. For all SHADOZ stations, reprocessing has improved the agreement of the
494 sondes with both the satellites and ground-instruments (Figure 11 in T12). With the
495 reprocessed data only two stations display a mean offset > (+)2% (**Figures 4, 7; Table 3)**.
496 By contrast, in T07 (Figure 8), sondes for five of the ten stations were offset 5-10% from
497 OMI. Witte17a showed that the greatest improvements in satellite-sonde agreement are
498 for Hanoi and Watukosek (**Figures 4l,m**); this now applies to reprocessed Réunion, San
499 Cristóbal and Costa Rica data as well.

500 **3.2 Partial Column Ozone Comparisons, TTL Ozone Structure and the**
501 **Tropospheric Wave-One.**

502 In prior analyses of SHADOZ column ozone amounts, we computed ozone column
503 amounts within the stratosphere, troposphere and TTL, along with the structure of the
504 wave-one pattern in tropospheric ozone. Updates to these analyses appear in **Figures 8-**
505 **10**. In some of the station-to-station comparisons, as in T12, we focus on the 10 stations
506 designated as "tropical"**.** The latter sites are distinct from four stations that are subtropical
507 due to seasonally low tropopauses (*Baray et al.,* 1998; *Diab et al.*, 2004; *Ogino et al*., 2013):
508 Hilo, Irene, Réunion, Hanoi. The tropopause heights in km for the tropical sites are shown
509 in **Figure 3**, along with the TTL, here taken as 15-18.5 km or 130-70 hPa.

510 The stratospheric ozone column, depicted in **Figure 8a,** obtained by subtracting
511 tropospheric column ozone from total ozone (**Table 3**), includes the climatological add-on
512 amount above balloon burst. Except for Paramaribo, the other nine tropical stations
513 display a mean of 229 ±3.9 DU. The relatively small 1σ values shown in **Figure 8a**
514 represent a significant improvement in data consistency for several stations compared to
515 earlier SHADOZ climatologies [T03; T07; T12]. For example, in the first compilation of
516 stratospheric ozone (Table 3, header 7, in T03) the Natal 1σ amount was 9% and in T12
517 (Table 4) it was 7.5%. This has been reduced to 5% after reprocessing (**Figure 8a)**. There
518 are also absolute percentage changes of this order from T07 (Table 4) compared to the
519 reprocessed data for San Cristóbal, Ascension, Watukosek and Fiji. In summary,
520 reprocessing has reduced stratospheric column ozone variability (i.e, standard deviation)
521 to 5%. The TCO variability has improved accordingly. Inspection of the ranges and sonde
522 averages (pink edge and red circles in **Figure 7**) shows that they vary from 4-6%.

523 The TTL column for all 14 SHADOZ stations (**Figure 8b**) only shows two prominent
524 outliers, Hilo and Irene (numbered 2 and 8, respectively), not surprising because of their
525 subtropical character and expected higher ozone content. They share some of the highest
526 outliers as does the subtropical Réunion station (labeled 10 in **Figure 8b**). However,
527 Hanoi, also subtropical, does not stand out from the other tropical stations.
528 A major finding from the SHADOZ data, as described in *Thompson et al.* [*2003b*;
529 2001*b*] was the isolation of the "wave-one" pattern in total ozone first described by
530 *Fishman and Larsen* [1987] and *Shiotani* [1992] (cf Figure 4 of *Kim et al*., 1996). For
531 tropical latitudes, the wave-one refers to 10-20 Dobson Units (DU) more column ozone
532 over the Atlantic and eastern Africa, a maximum relative to the central-western Pacific
533 ozone minimum. The SHADOZ observations from 1998-2000 [*Thompson et al.,* 2003*b*],
534 demonstrated that because there are no statistically significant differences among column-
535 integrated stratospheric ozone over the individual tropical stations (as in **Figure 8a, Table**
536 **3**), the additional ozone must reside in the troposphere. These findings apply to the
537 reprocessed data from 1998-2016 where detailed statistics, in the form of a box-and-
538 whisker plot (**Figure 8c**) for the ten tropical stations, are presented. Although **Figure 8c**
539 does not display station locations realistically**,** noting from **Figure 7** that the wave-one
540 maximum lies between Natal and Ascension, the wave magnitude can be approximated.
541 Based on the data in **Figure 8c** the mean tropospheric ozone maximum is ~38 DU. The
542 eastern Indian Ocean-to-Pacific minimum (**Figure 8c)** extends across four stations from
543 Kuala Lumpur to Samoa, where the range of tropospheric column ozone is 22-26 DU.
544 Taking the mean tropospheric column ozone of 24 DU to represent the zonal minimum, the
545 tropical wave magnitude is estimated as 14 DU. This is the same as estimated in T12.
546 A cross-section of mean TTL mixing ratios, from 15 to 18.5 km, based on 0.25 km
547 averages (**Figure 9a**), captures nearly uniform ozone mixing ratios above the nominal
548 tropopause (> 17 km) with distinct zonal variations below 16 km. Viewed on a seasonal
549 basis (**Figures 9b,c**) the zonal gradients are more pronounced. In March-April-May (MAM
550 in **Figure 9b**), a highly convective time of year at most tropical sites, the zonal gradients
551 across 15 km are greater than 50%. The same holds when subsidence and pollution
552 impacts in the UT are strongest between Paramaribo and Nairobi (eastern South America
553 to east Africa) in September-October-November (SON in **Figure 9c**).

554 Seasonal variations in tropospheric ozone, as in the TTL, cause the wave-one
555 magnitude to vary over the year (**Figure 10)**. As in *Thompson et al.* [2003b] **Figure 10**
556 indicates that the wave-one is a minimum when most tropical stations are convectively
557 active (MAM). The maximum wave amplitude is observed in SON where there is greater
558 subsidence and pollution sources in the mid-troposphere from southern hemisphere
559 biomass fires [*Thompson et al.,* 1996; T12].

560 **3.3 Tropical Station-to-station Bias**

561 Our prior studies of station-to-station bias focused on the stratosphere because, as
562 **Figure 8a** implies, the total stratospheric ozone column for the tropical stations is uniform
563 within statistical significance. A mean tropical ozone profile from surface to 10 hPa has
564 been calculated from the ten most equatorward stations (seasonal means in **Figure 11**).
565 Note that in **Figure 11** the variability of two seasons, DJF (December-January-February)
566 and JJA (June-July-August), is greater than MAM and SON (larger 1σ standard deviations,
567 denoted by gray shading) in the troposphere and also near the stratospheric peak. The
568 high variability suggests that it may be difficult to compute reliable trends in the TTL
569 except in SON when ozone is a maximum and in MAM when ozone is a minimum [T03].

570 When the mean profile for each tropical station is referenced to an all-tropical profile
571 average based on **Figure 11**, the offsets illustrated in **Figure 12** are obtained. Results for
572 the three regions, eastern Indian Ocean to Pacific (labeled as W. Pacific in **Figure 12a**),
573 eastern Pacific (**Figure 12b**) and Atlantic plus Nairobi (**Figure 12c**) are shown separately.
574 It is assumed that above the region marked TTL, stratospheric ozone is zonally invariant
575 and that deviations from zero signify bias. All the offsets in **Figure 12** except for Fiji (blue
576 in **Figure 12a**) and Ascension near the TTL (red in **Figure 12c**) are 10% absolute or less.
577 The higher ozone over Fiji below 40 hPa likely owes to subtropical characteristics, perhaps
578 mixing of tropical and mid-latitude air, in the South Pacific Convergence Zone (the site is at
579 18°S, *Pickering et al.,* 2001). A nearly identical deviation from the other tropical stations
580 based on 1998-2009 Fiji data appears in Figure 13a of T12. Changes in Watukosek and
581 Kuala Lumpur offsets (**Figure 12a**) compared to T12 are mostly in the TTL. The three east
582 Pacific stations (Costa Rica grouped with the island sites, **Figure 12b**) are all close to the
583 mean and show almost no variation throughout the stratosphere. Compared to their T12
584 offsets, based on 1998-2009 data, Samoa and San Cristóbal in **Figure 12b** have increased

585 relative to the other stations. Samoa now has a slightly positive offset and San Cristóbal is
586 close to the zero line. These patterns for Samoa and San Cristóbal are also close to those
587 shown in T07 (Figure 5a) that were based on 1998-2004 observations.
588 In **Figure 12c** Nairobi is nearly on the zero line whereas it had a 5% high-bias in T12.
589 This is because reprocessing increases stratospheric ozone at many stations [Witte17a].
590 Paramaribo is consistently high (green in **Figure 12c**) as in T12. Natal is nearly unchanged
591 from T12 and Ascension is slightly more negative (red and blue, respectively, **Figure 12c**).
592
593 **4. Summary**
594 The first set of reprocessed data from 12 SHADOZ stations (1998-2016) and 2 stations
595 (2004-2016) with records covering at least a decade are evaluated in three ways. First,
596 total ozone from the sondes is compared to satellite overpass total ozone from the three
597 buv-type instrument series that cover the SHADOZ period, EP/TOMS up through 2004, and
598 the operational OMI (since September 2004) and OMPS, since February 2012. Total ozone
599 column coincident with the sondes are also compared to co-located ground-based
600 instruments at nine stations. The main findings are:
601 • Reprocessing leads to more consistent total ozone records within a given
602 station except for Paramaribo, which is currently being reprocessed again.
603 • Offsets between sonde total ozone and ozone from the satellite series are
604 greatly improved relative to T07 and T12. Two stations register mean absolute
605 sonde-satellite disagreements of 5% absolute; all others are within $\pm$2%. The
606 same applies to sonde TCO agreement with nine co-located ground-based
607 instruments.
608 • Although TCO from the sondes, satellites and ground-based instruments agree
609 better after reprocessing there is still unexplained noise in most time-series of
610 ratios and clear discontinuities when the sonde TCO is compared to the satellite
611 record. The two prominent features are: (1) the sonde/satellite TCO record
612 seems to vary among the satellites with TOMS v8 TCO running higher (lowest
613 sonde ratio) than the first 5-7 years of OMI (higher sonde ratio) before the
614 sonde ratio drops during the OMI-OMPS era. The pattern is confirmed through

comparison of the satellite and ground-based instrument TCO. (2) The second discontinuity is observed after 2012-2014 at eight stations with the sonde ratio dropping relative to OMI, OMPS and, where applicable, the local ground-based instrument. This suggests that something is shifting in the sondes. Radiosonde changes appear to play a role at several stations and changes in the ECC sonde can be detected at Costa Rica and Fiji. However, the statistics for the post-2012 drop in sonde ratio are limited compared to a 19-year record for most SHADOZ stations. These patterns are being investigated further. Possible ECC sonde issues will be addressed in JOSIE-type chamber tests in the near future.

The second analysis of the reprocessed data examined three partial column amounts for ten tropical stations, defined as having at latitude < 19 degrees: stratospheric ozone, tropospheric ozone, ozone in the tropical tropopause layer (TTL). As in our prior comparisons [T03; T07; T12] the stratospheric column was zonally invariant except for slightly elevated amounts over Paramaribo. In addition:

- The TTL column amount is also invariant across the ten stations (9 $\pm$ 1 DU) and two subtropical stations, Réunion and Hanoi, display nearly the same value. However, there is considerable zonal structure below 17 km at all times of year, with strong vertical gradients over the western Pacific and eastern Indian Ocean. These features are not well-resolved by profiling satellites which argues for caution in trends based on satellite data below 70 hPa.
- The tropospheric wave-one feature first characterized with SHADOZ data from nine stations [*Thompson et al.,* 2003b] is better delineated with the addition of Costa Rican data. The mean amplitude of the wave is 14 DU.
- Reprocessing has reduced the standard deviation of stratospheric and total column ozone at SHADOZ stations to 5%.

Third, using mean profiles, we examined mid-stratospheric bias given that the stratospheric column appears to be zonally invariant among the 10 tropical sites. At all but two stations the bias above 70 hPa has improved relative to the T07 and T12 analyses; the mean for all stations less than 10% absolute. This demonstrates that key elements in reprocessing procedures, using appropriate pump corrections, applying transfer functions,

screening for anomalous background currents, and, in general, applying corrections that account for known biases in the ozonesonde system, indeed homogenize SHADOZ data. The results underscore the need for complete meta-data in archived sonde records.

Satellite-based tropospheric ozone products and models suitable for exploring complex variability in the tropics are still in development. Given the improvements quantified in this study, we expect reprocessed SHADOZ data to become a standard reference for evaluating new satellites, emerging tropospheric ozone products, assessment model simulations and for detecting satellite drift. In addition, the seasonal and longitudinal variations in TTL ozone structure observed in SHADOZ data, as well as our classification of tropospheric profiles by self-organizing maps [*Jensen et al.* 2012; *Stauffer et al.,* 2016] suggest that more attention be given to tropical satellite retrievals below the stratosphere. In the TTL critical interactions take place among ozone, water vapor, temperature and dynamics; in the free troposphere ozone mediation of OH determines atmospheric lifetimes for myriad species.

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Acronyms

BESOS = Balloon Experiment on Standards for Ozone Sondes
BUV = Back-scattered Ultraviolet
CPT = Cold-point Tropopause
DJF = December-January-February
DU = Dobson Unit; 1 DU=2.69 x10^{16} molec cm^{-2}
ECC = Electrochemical Concentration Cell
ENSO = El Niño Southern Oscillation
FT = Free Troposphere
GOME = Global Ozone Monitoring Experiment (ERS-2 GOME, 1995-2011; GOME II, 2003-, 2012-)
IASI = Infrared Atmospheric Sounder Instrument
IO3C = International Ozone Commission
IGACO = Integrated Global Atmospheric Chemistry Observations
IPCC = Intergovernmental Panel for Climate Change
JJA = June-July-August
JOSIE = Jülich Ozonesonde Intercomparison Experiment; (http://www.fz-juelich.de/icg/icg-2/josie)

LRT = Lapse-Rate Tropopause
LS = Lower Stratosphere
MAM = March-April-May
MLS = Microwave Limb Sounder (on UARS, 1991-2005; on Aura, 2004-)
NASA = National Aeronautics and Space Administration
NDACC = Network for Detection of Atmospheric Composition Change (http://www.ndacc.org)
NOAA = National Oceanic and Atmospheric Administration
NPP = National Polar-Orbiting Partnership
OMI = Ozone Monitoring Instrument (on Aura, 2004-)
OMPS = Ozone Mapper Profiler Suite (on S-NPP, October 2011-)
PTU = Pressure-Temperature-Humidity (data from radiosonde)
QBO = Quasi-biennial Oscillation
SAOZ = System d'Analyse par Observation Zenitale (http://gosic.org/gcos/SAOZ-prog-overview.html)
SCIAMACHY = Scanning Imaging Absorption SpectroMeter for Atmospheric ChartographY (2002-2012)
SHADOZ = Southern Hemisphere Additional Ozonesondes; (http://tropo.gsfc.nasa. gov/shadoz)
TC⁴ = Tropical Composition, Clouds and Climate Coupling; (http://www.espo.nasa.gov/tc4/ (2007)
TOMS = Total Ozone Mapping Spectrometer (in SHADOZ era, Earth-Probe/TOMS, 1996-2005)
TTL = Tropical Tropopause Layer
UT/LS = Upper Troposphere/Lower Stratosphere
WMO = World Meteorological Organization
WOUDC = World Ozone and Ultraviolet Data Centre; (http://woudc.org)

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FIGURE CAPTIONS 3 Oct

Figure 1 Map of SHADOZ stations for which reprocessed data are analyzed in this paper, within the years 1998-2016. Period of operations and sample numbers are in **Table 1.** Technical details of the reprocessing appear in Witte17a and Sterling17**.** Details for SHADOZ sondes used at each site during 1998-2004, the Earth-Probe/TOMS period, appear in *Thompson et al.* [2003a, 2007].

Figure 2 Timeline of satellite ozone instruments that have used SHADOZ for validation. Beyond 2017, expected lifetime for operational sensors is indicated with satellites scheduled for launch by US and international partnerships in red.

Figure 3 Box-and-whisker plot of tropopause heights are shown for the 10 SHADOZ stations within $\pm$19 degrees of the equator, designated as "tropical". Yellow boxes show the interquartile of the 25th, median, and 75th percentiles. Outliers are displayed as blue circles. The TTL region is bounded by the dashed lines.

Figure 4 Time-series of TOMS-OMI-OMPS overpass total column ozone (O_3, gray-black-silver lines, respectively) with ozonesonde total column O_3 (red circles) values that reached 10 hPa. *McPeters and Labow* (2012) O_3 climatology is added from 10 hPa to the top-of-the-atmosphere to yield sonde total column amounts. After February 2012 OMI and OMPS operated simultaneously. Where available, total column O_3 from co-located Dobson, Brewer or SAOZ instruments (blue cross) is given. The lower panel includes sonde/satellite and sonde/ground instrument ratios. Here too, after February 2012, sonde/satellite ratios include both OMI (black circles) and OMPS (silver circles). Histograms of the percent difference with respect to sondes are also given for each station (yellow). Where available, the percent difference is also shown between the sondes and ground-based total column O_3 (blue hashed). (a) American Samoa; (b) Hilo, Hawaii; (c) San Cristóbal; (d) Costa Rica; (e) Paramaribo; (f) Natal; (g) Ascension; (h) Irene; (i) Nairobi; (j) Réunion; (k) Kuala Lumpur; (l) Hanoi; (m) Watukosek; (n) Fiji. Satellite overpasses are filtered for reflectivity greater than 0.6 and distance from the station location greater than 200 km.

Figure 5 Histograms of the ratio of total column O_3 from the ground-based instruments at nine stations and coincident satellite measurements from TOMS (yellow-solid), OMI (blue-hashed), OMPS (green-hashed). (a) American Samoa; (b) Paramaribo; (c) Natal; (d) Irene; (e) Nairobi; (f) Réunion; (g) Kuala Lumpur; (h) Hanoi; (i) Watukosek.

Figure 6 ECC sensor serial number (ENSCI 2Z type divided by 1000) versus total column O_3 at Costa Rica (left panel) and Fiji (right panel) since 2005.

Figure 7 Longitudinal cross-section of 1998-2016 reprocessed sonde total column O_3 (red circles with $\pm\underline{1}\sigma$ light red shading) and the collective mean TOMS-OMI-OMPS overpass columns (black with $\pm\underline{1}\sigma$ black bars). Comparison of total column amounts from the satellites and reprocessed sonde O_3, on average, appear in **Table 3**, column 4.

Figure 8 For the 10 tropical SHADOZ stations within $\pm$19 degrees of the equator, box-and-whisker plots of (a) integrated column amounts of stratospheric O_3 (in DU) based on 1998-2016 reprocessed data, (b) TTL column amounts from 15-18.5 km (130-70 hPa) for all SHADOZ stations, and (c) tropospheric column amounts integrated from the surface to the

974 tropopause showing the characteristic wave-one pattern. Yellow boxes denote the
975 interquartile ranges.
976
977 **Figure 9** (a) Longitudinal cross-section of O_3 mixing ratios (units of ppm), computed from
978 0.25 km averages of 1998-2016 reprocessed data, for the 10 tropical stations, labeled by
979 country, within the TTL region. (b, c) same as (a) except for seasons MAM and SON,
980 respectively, and labeled by station. Labeling by Comparing to Figure 8b, where the column
981 amounts are statistically identical, distinctions in O_3 structure in the cross-section are
982 pronounced.
983
984 **Figure 10** Same as **Figure 9,** except contours of O_3 mixing ratio in the troposphere up to
985 the lower region of the TTL where the wave-one pattern appears, as in T03 and T12. Four
986 seasons illustrated: (a) SON; (b) DJF; (c) MAM; (d) JJA.
987
988 **Figure 11** 1998-2016 seasonally averaged O_3 partial pressure profiles (units of nbar) for
989 the ten reprocessed tropical stations. (a) SON; (b) DJF; (c) MAM; (d) JJA.
990
991 **Figure 12** Deviations from the 1998-2016 mean climatological profile from 100 to 10 hPa.
992 TTL upper boundary is marked by the dashed line. Above the TTL, in the stratosphere,
993 station offset differences are more likely to indicate bias. (a) western Pacific and eastern
994 Indian Ocean stations; (b) sites in equatorial Americas; (c) eastern South America,
995 Ascension, Nairobi. Station mean profiles are from 1998-2016, except for Costa Rica and
996 Hanoi, where launches started in 2004 and 2005, respectively.
997

Station Name	Location	Local Launch Times*	Ground Total Ozone Instr.	Profile #	Reprocessing Reference
Am. Samoa	14.23S, 170.56W	0800 – 1200, 1300 - 1800	Dobson #80	480	Sterling17
Hilo, Hawaii	19.40N, 155.0W	0800 - 1300	---	600	Sterling17
San Cristóbal, Galapagos (Ecuador)	0.92S, 89.60W	0500 - 1700	---	387	Sterling17
Costa Rica (various sites)**	9.94N, 84.04W	0600 - 1200	---	440	Witte17, This study
Paramaribo, Surinam	5.81N, 55.21W	0900 - 1400	Brewer #159	476	Allaart et al., in prep., *AMT*
Ascension Is., UK	7.98S, 14.42W	1300 - 1400	---	538	Witte17a
Natal, Brazil	5.42S, 35.38E	1200 - 1400	Dobson #93	517	Witte17a
Irene, South Africa	25.9S, 28.22E	1000 - 1300	Dobson #89	258	Witte17a
Nairobi, Kenya	1.27S, 36.80E	0700 - 1800	Dobson #18	719	Witte17b
La Réunion, France	21.1S, 55.48E	0800 - 1300	SAOZ#20	521	Witte17a,b with modification described in Sec 2.2.
Kuala Lumpur, Malaysia	2.73N, 101.7E	1000 - 1200	Brewer #90	271	Witte17a
Hanoi, Vietnam	21.02N, 105.80E	1300 - 1400	Brewer #208	148	Witte17a
Watukosek-Java, Indonesia	7.57S, 112.65E	1200 - 1300	Brewer #116 in Bandung	146	Witte17a
Suva, Fiji	18.13S, 112.65E	0800 – 1100, 2100 - 2300	--	222	Sterling17

*On average times.
** Current site is San Pedro. Alajuela and Heredia historic sites are within 0.5 degrees of San Pedro.

Table 1. SHADOZ stations with co-located ground-based total ozone instrument (where applicable) and sample numbers (total reprocessed records with bursts between 10-15 hPa). Reference for reprocessing is also given.

Table 2. Radiosondes used at SHADOZ stations

	RS80 or RS92	Pressure Offset Correction	iMet	Pressure Offset Correction
Samoa	RS80: 1998 - 2013	No	2014/01/10 – present	Yes
Fiji	RS80: 1998 - 2013	Intermittent	2015 – present	Yes
Hilo	RS80: 1998 – 2012/09/05	Intermittent – yes after 2009/04/08	2012/09/12 – 2012/11/28	No
			2012/12/05 - present	Yes
San Cristöbal	RS80: 1998 – 2008	No	---	---
	RS92: 2012/05 – present	Yes		
Costa Rica*	RS80: 2005 - 2013/08/16	Yes – Vömel method**	2010/08/19 - present	Yes
Paramaribo	RS80: 1999 – 2005/09/17	No	---	---
	RS92: 2005/10/03 – present			
Nairobi	RS80: 1998 – 2010/03/03	No	---	---
	RS92: 2010/06/09 - present			
Ascension			2016 – present***	Yes

*Overlap of iMet and RS80 between 2010/08 – 2013/08
** Using rise rate
***LMS: 1998 – 2010 (refer to Witte17a)

Table 3. The mean and 1-sigma standard deviations of TCO, stratospheric, and tropospheric O_3, the latter from the sondes. Values are reported in units of DU. TCO are reported from satellite, ground-based, and ozonesonde instruments. Column 4 includes the percent difference of sonde and satellite mean TCO with respect to sondes. Stratospheric O_3 includes the residual climatology. Sites within ±19 degree of the equator itemized first from west to east to show the tropospheric wave-one amplitude (column 6). Sub-tropical sites follow.

Site	**Satellite TCO***	**Ground-based TCO**	**Sonde TCO (mean Sat. Δ, %)**	**Sonde Strat. O3**	**Sonde Trop O3**	**Tropopause [hPa]**
Tropical Sites within ±19° of Equator						
Samoa	251.7±11.0	248.0±10.1	249.4±12.1 (-0.9)	227.9±9.7	21.6±6.3	94.5±13.2
San Cristobal	260.4±12.1	--	258.2±15.0 (-0.9)	231.2±11.9	27.5±10.8	90.6±16.3
Costa Rica	255.7±14.34	--	255.3±19.2 (-0.2)	229.4±17.2	26.0±5.4	94.8±27.4
Paramaribo	262.1±13.8	260.3±12.7	273.0±19.0 (+4.0)	242.1±15.2	30.1±6.8	96.7±22.1
Natal	265.4±11.3	268.8±12.7	260.9±15.1 (-1.7)	226.6±10.1	34.3±8.1	92.9±9.9
Ascension	266.7±10.4	--	260.4±17.9 (-2.4)	221.5±12.5	38.6±8.4	92.8±15.8
Nairobi	260.1±12.4	257.9±13.7	255.8±11.3 (-1.7)	227.2±9.4	28.6±4.9	90.0±14.7
Kuala Lumpur	259.5±13.4	256.6±22.3	246.0±13.5 (-5.5)	220.1±12.8	26.0±5.8	93.3±10.3
Watukosek	257.5±9.5	250.4±10.4	254.3±10.9 (-1.3)	226.9±8.7	27.2±6.8	88.4±8.2
Fiji	258.4±14.5	--	256.0±16.0 (-0.9)	231.0±11.2	24.6±7.9	93.9±11.9
Sub-Tropical Sites						
Hilo	273.7±18.4	--	270.4±20.3 (-1.2)	238.8±17.7	31.9±8.0	117.5±41.2
Irene	268.1±15.9	274.6±17.3	273.6±20.0 (+2.0)	237.4±17.3	36.2±6.7	124.5±35.8
La Reunion	267.1±15.1	266.6±14.1	264.8±18.0 (-0.9)	229.6±13.2	35.5±7.8	103.3±25.8
Hanoi	264.7±17.6	276.2±17.6	261.7±21.1 (-1.2)	220.3±18.2	41.2±7.7	91.7±10.8

*Includes TOMS, OMI, OMPS TCO overpasses.

Figure 1.


Hilo, HI
Costa Rica (various sites)
San Cristobal, Ecuador
Pago Pago, Am. Samoa
Paramaribo, Surinam
Natal, Brazil
Ascension Is.
Irene, South Africa
Nairobi, Kenya
La Reunion Is., France
Hanoi, Vietnam
Kuala Lumpur, Malaysia
Watukosek–Java
Suva, Fiji

Figure 2.

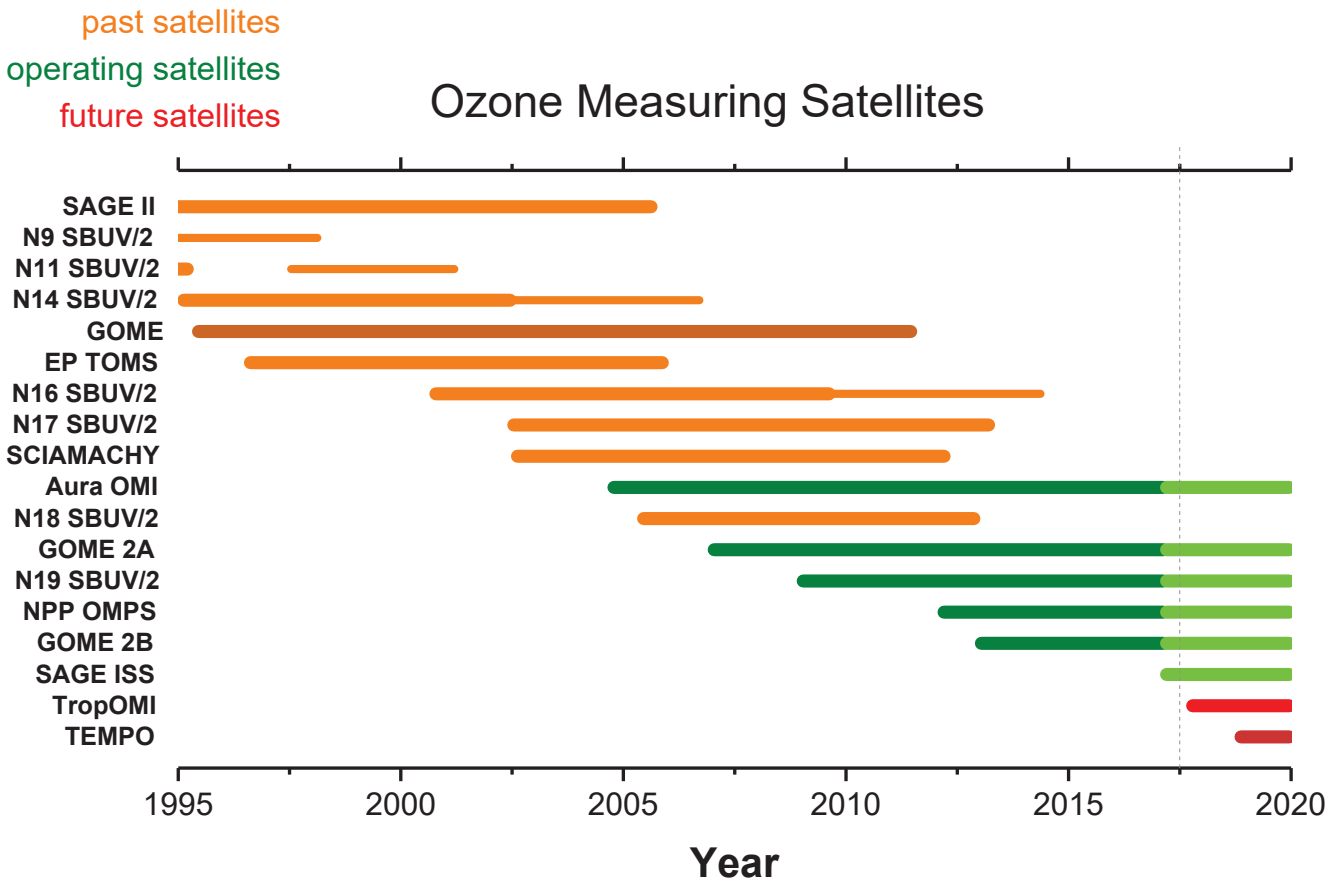

past satellites
operating satellites
future satellites
Ozone Measuring Satellites
SAGE II
N9 SBUV/2
N11 SBUV/2
N14 SBUV/2
GOME
EP TOMS
N16 SBUV/2
N17 SBUV/2
SCIAMACHY
Aura OMI
N18 SBUV/2
GOME 2A
N19 SBUV/2
NPP OMPS
GOME 2B
SAGE ISS
TropOMI
TEMPO
1995
2000
2005
2010
2015
2020
Year

Figure 3.

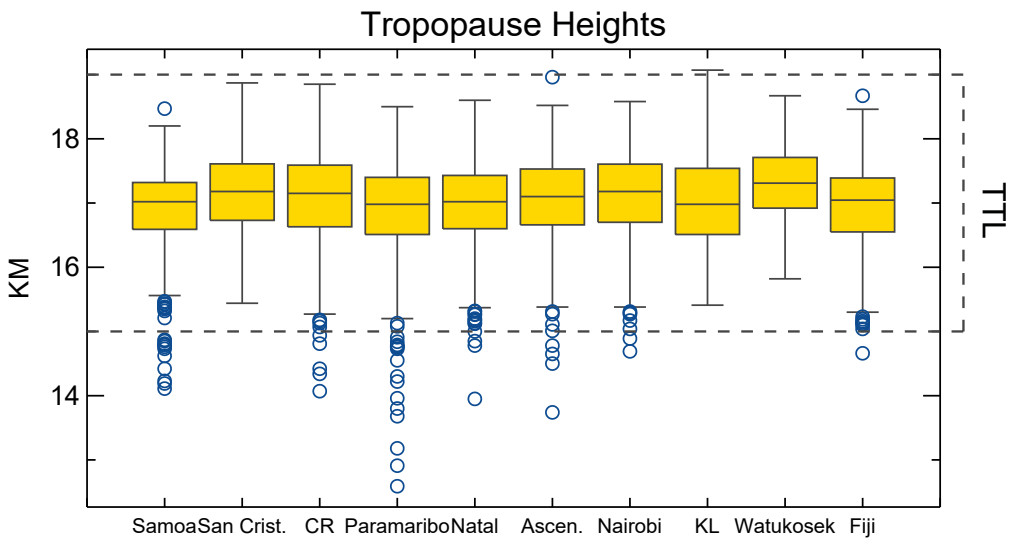
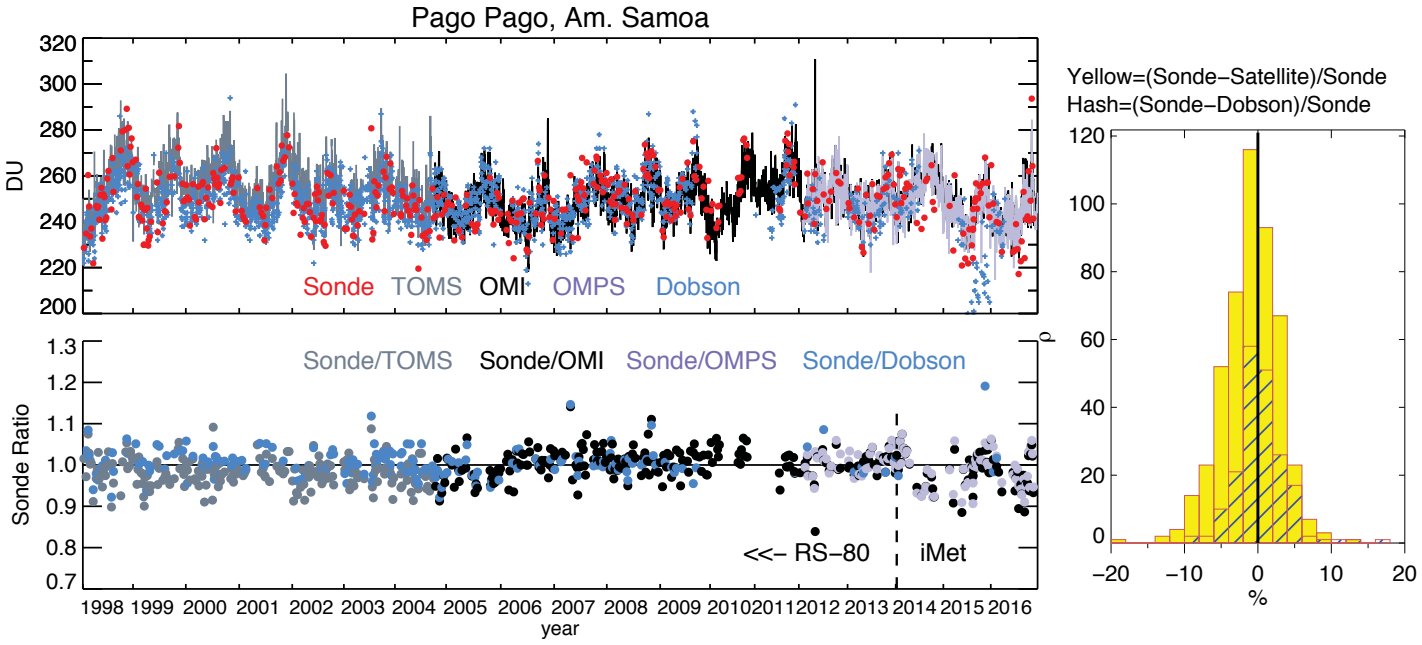
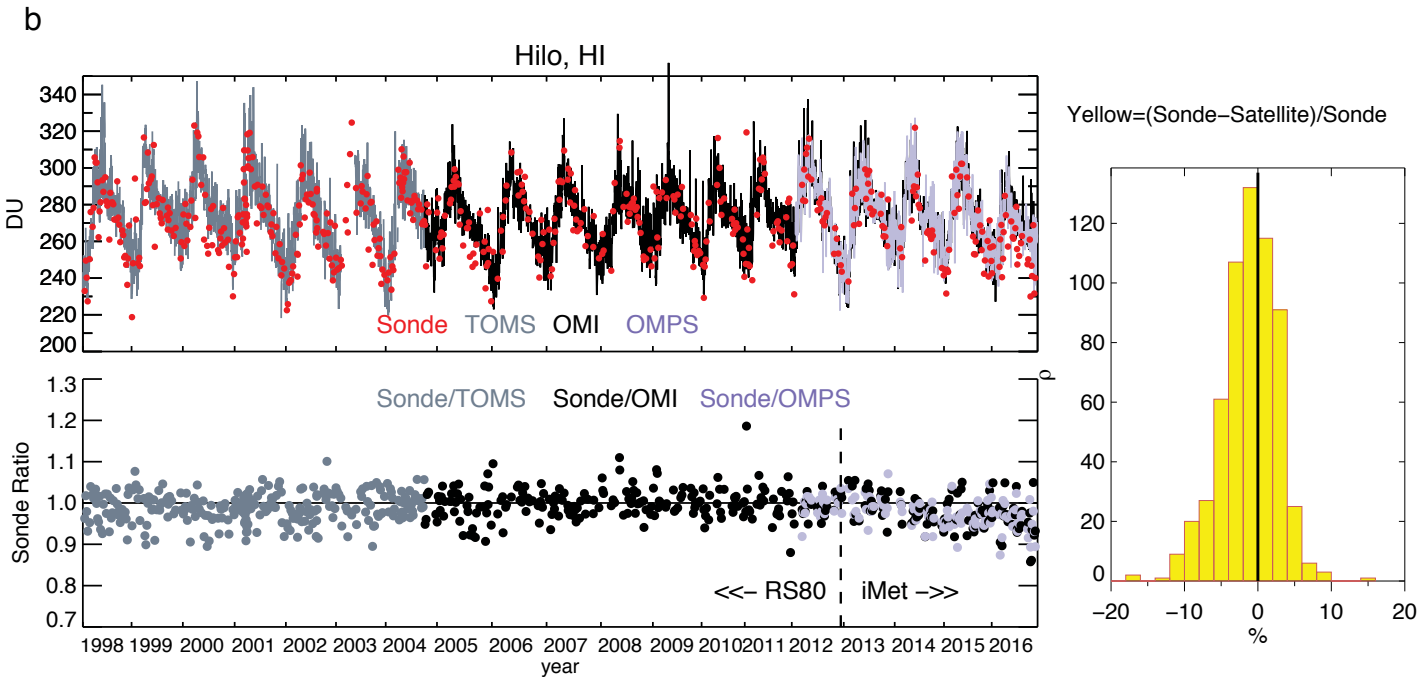
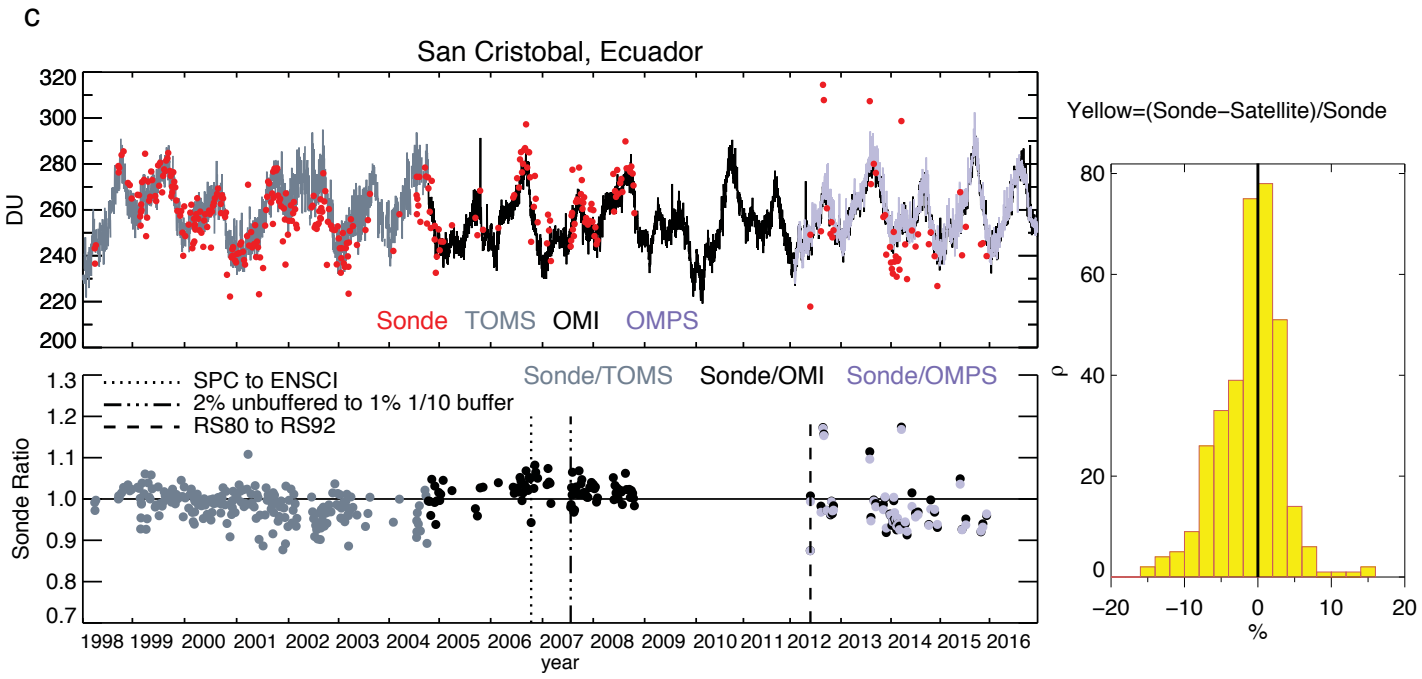

Tropopause Heights
KM
18
16
14
TTL
Samoa
San Crist.
CR
Paramaribo
Natal
Ascen.
Nairobi
KL
Watukosek
Fiji

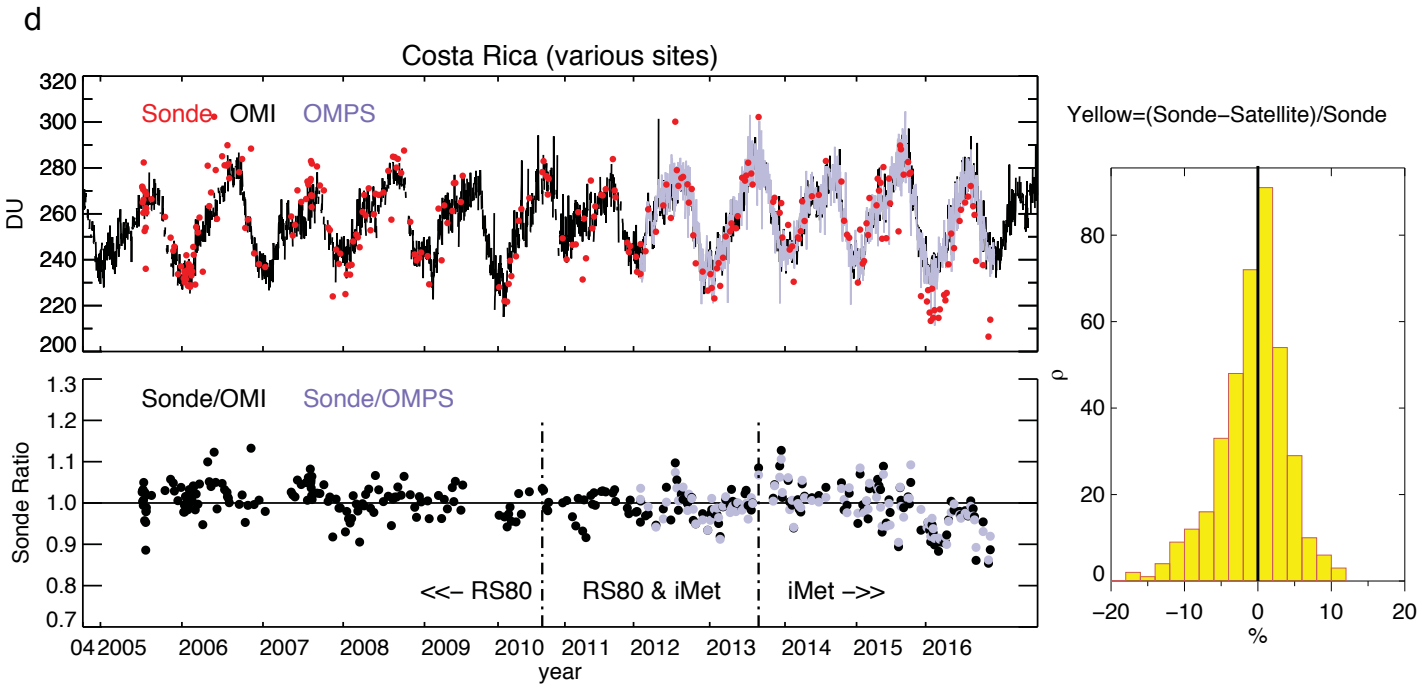
Figure 4.

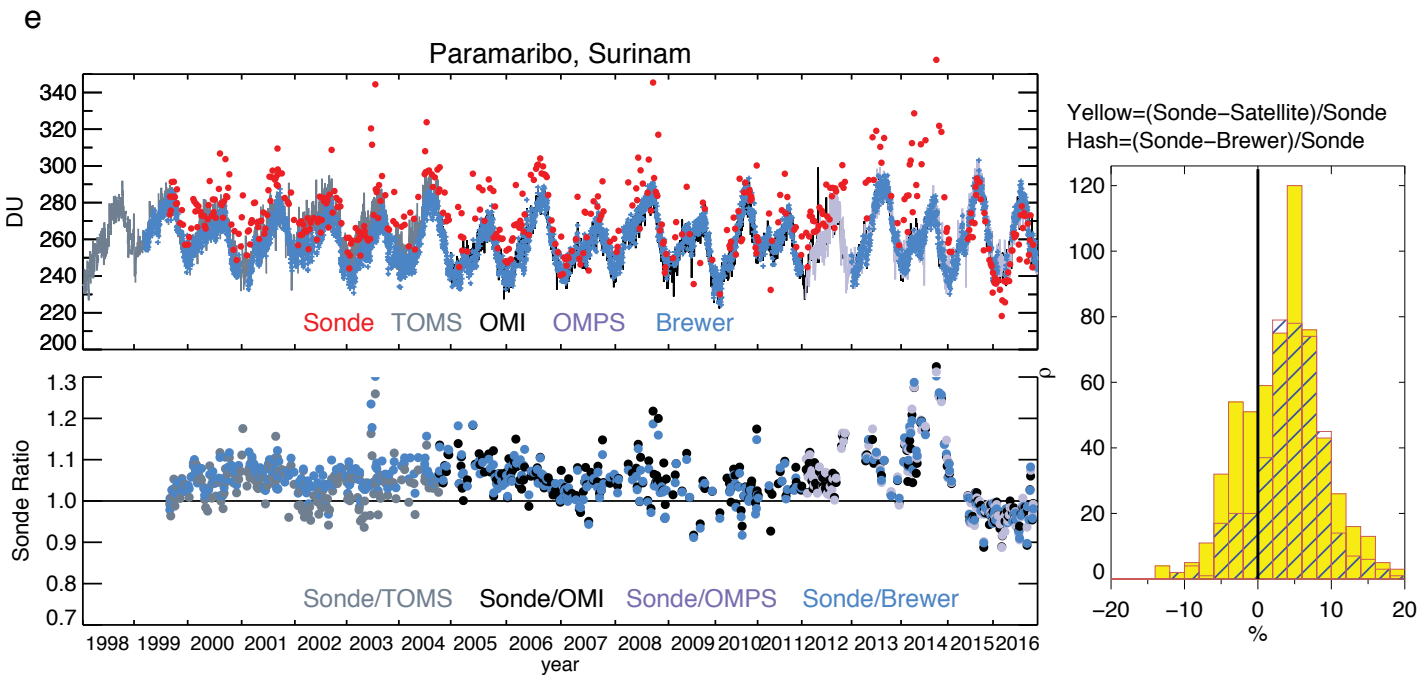
a


Pago Pago, Am. Samoa
DU
Sonde
TOMS
OMI
OMPS
Dobson
Sonde Ratio
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde/Dobson
<<– RS–80
iMet
year
Yellow=(Sonde–Satellite)/Sonde
Hash=(Sonde–Dobson)/Sonde
%

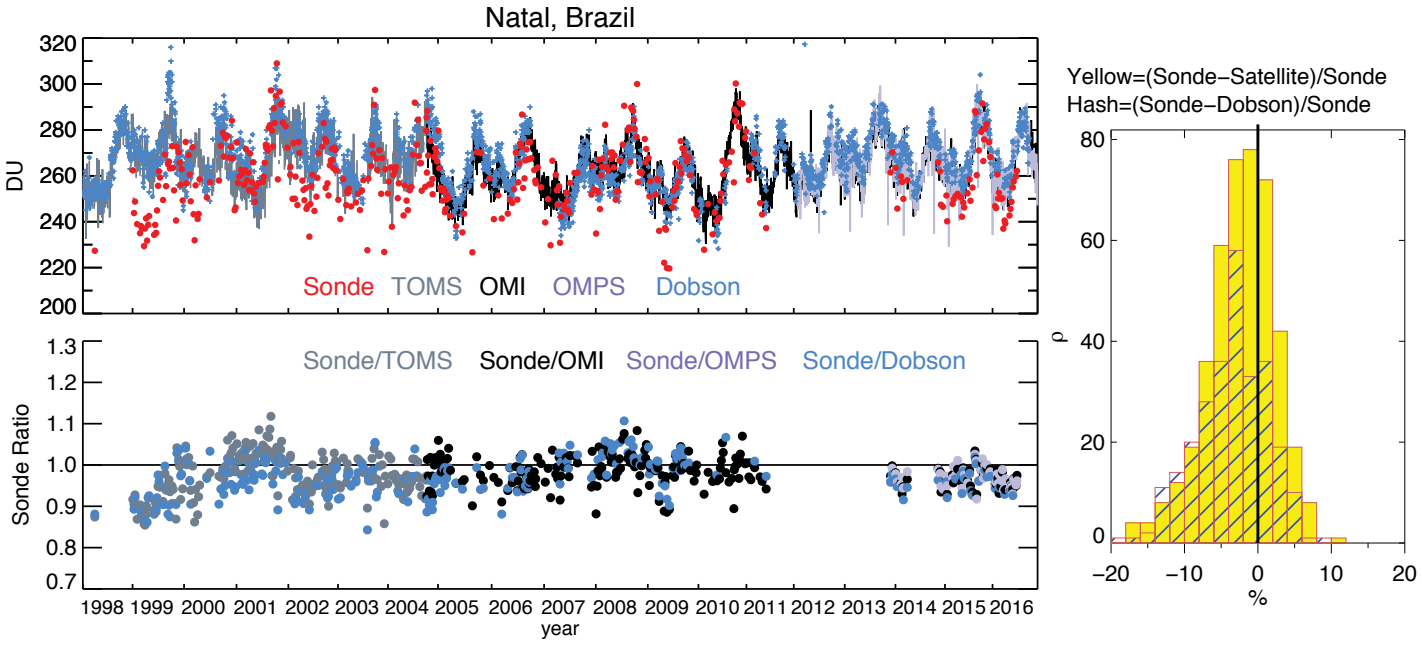

b
Hilo, HI
DU
340
320
300
280
260
240
220
200
Sonde
TOMS
OMI
OMPS
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
<<– RS80
iMet –>>
1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013 2014 2015 2016
year
Yellow=(Sonde–Satellite)/Sonde
120
100
80
60
40
20
0
–20
–10
0
10
20
%


c
San Cristobal, Ecuador
DU
Sonde
TOMS
OMI
OMPS
SPC to ENSCI
2% unbuffered to 1% 1/10 buffer
RS80 to RS92
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde Ratio
year
Yellow=(Sonde−Satellite)/Sonde
%

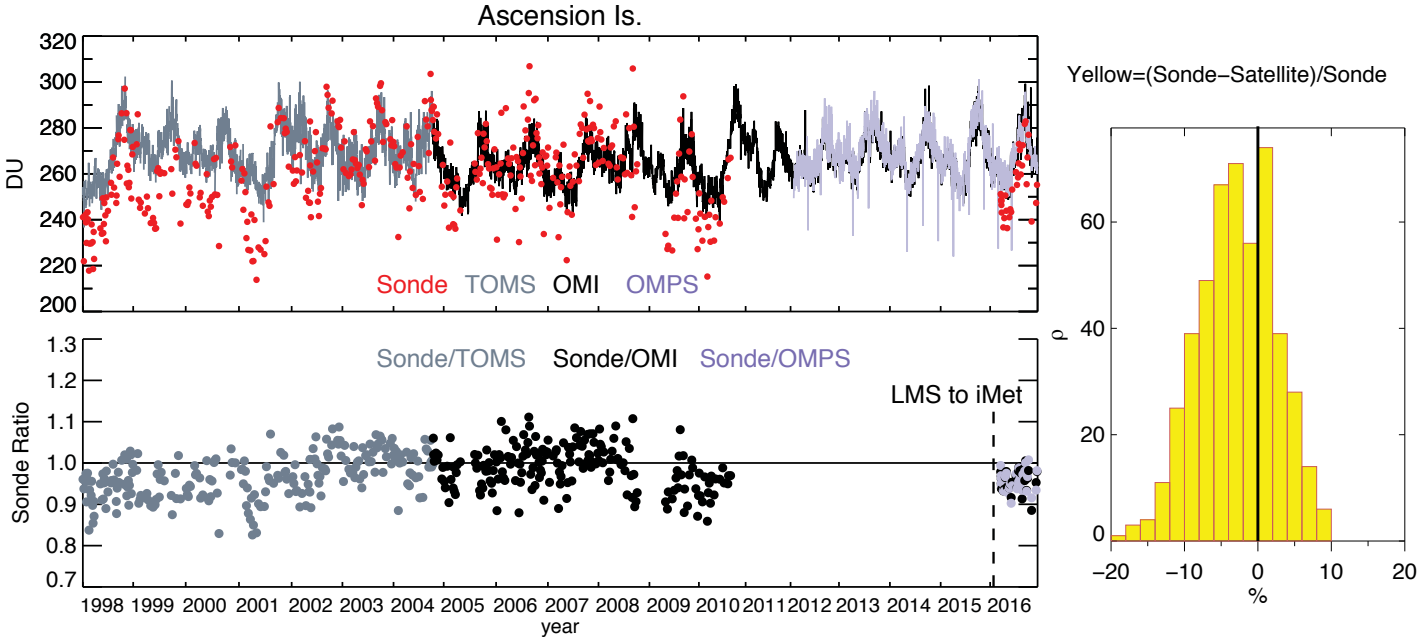

d
Costa Rica (various sites)
Sonde
OMI
OMPS
DU
320
300
280
260
240
220
200
Sonde/OMI
Sonde/OMPS
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
<<– RS80
RS80 & iMet
iMet –>>
04 2005
2006
2007
2008
2009
2010
2011
2012
2013
2014
2015
2016
year
Yellow=(Sonde–Satellite)/Sonde
ρ
80
60
40
20
0
–20
–10
0
10
20
%


e
Paramaribo, Surinam
DU
340
320
300
280
260
240
220
200
Sonde
TOMS
OMI
OMPS
Brewer
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde/Brewer
1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013 2014 2015 2016
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Brewer)/Sonde
120
100
80
60
40
20
0
−20
−10
0
10
20
%

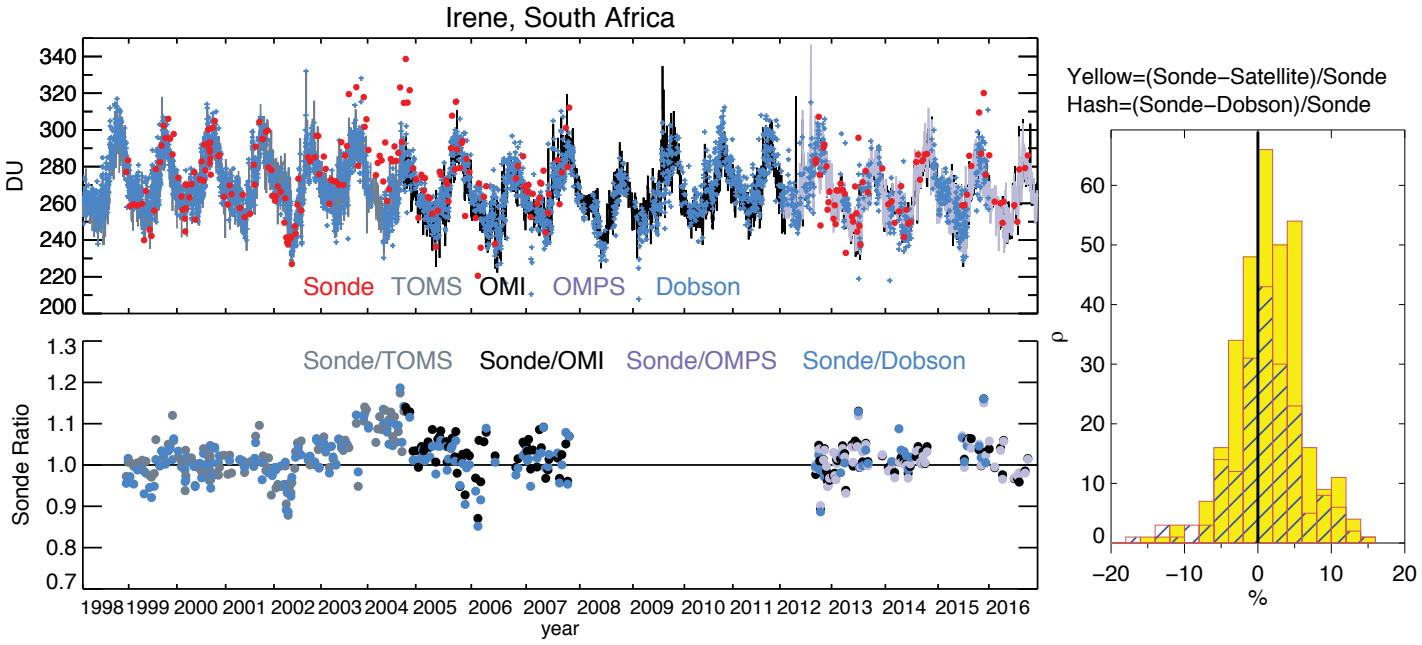
f


Natal, Brazil
DU
Sonde TOMS OMI OMPS Dobson
Sonde Ratio
Sonde/TOMS Sonde/OMI Sonde/OMPS Sonde/Dobson
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Dobson)/Sonde
ρ
%

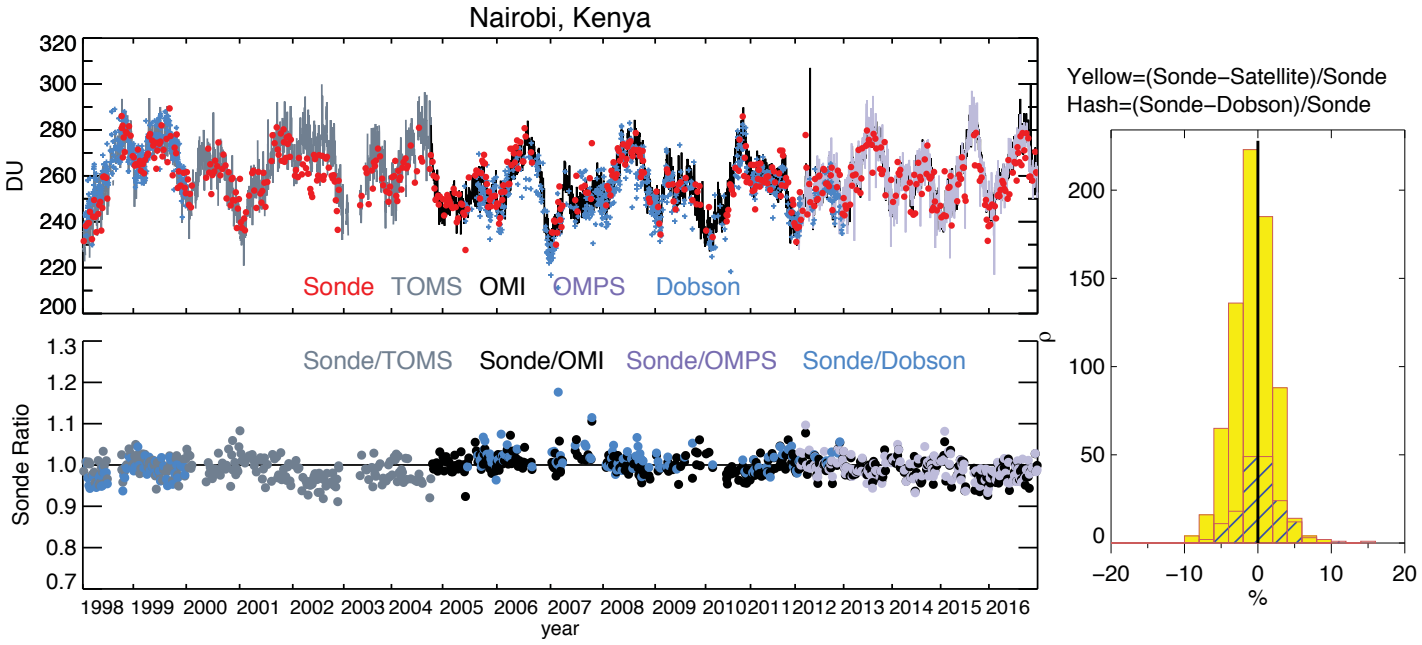
g


Ascension Is.
DU
Sonde TOMS OMI OMPS
Sonde/TOMS Sonde/OMI Sonde/OMPS
LMS to iMet
Sonde Ratio
year
Yellow=(Sonde−Satellite)/Sonde
%

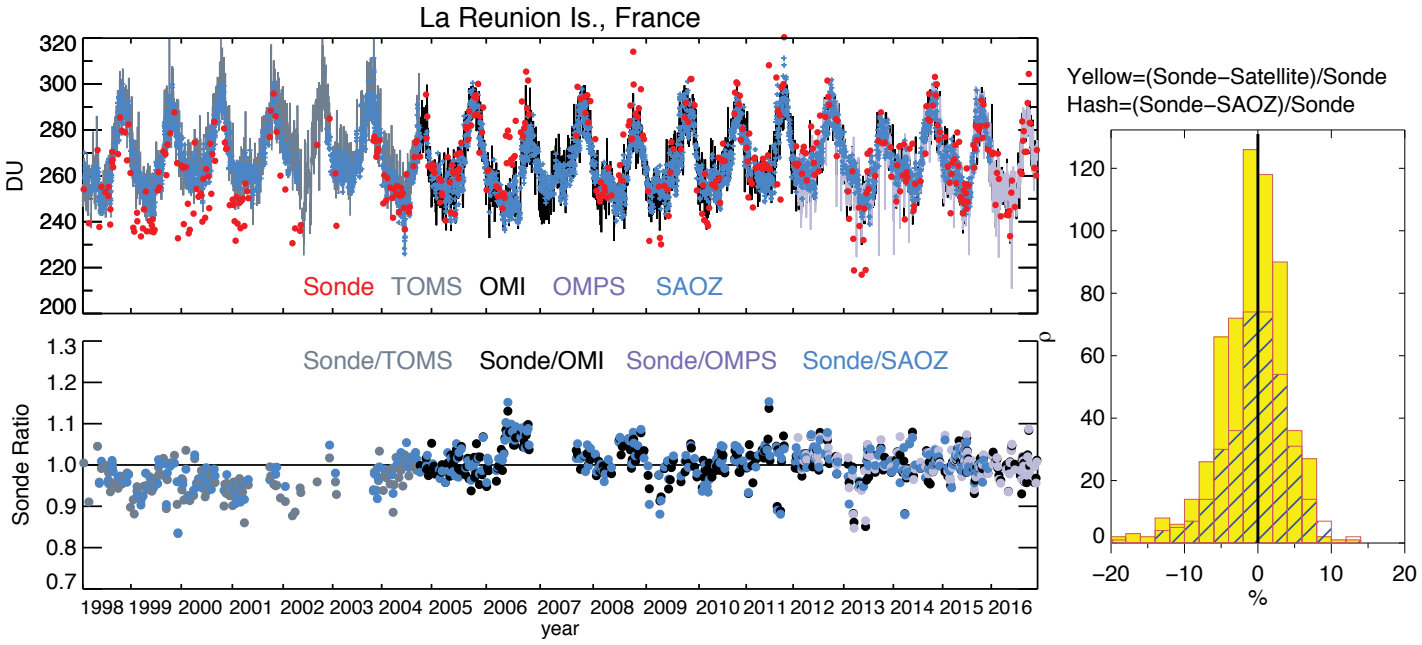
h


Irene, South Africa
DU
Sonde TOMS OMI OMPS Dobson
Sonde Ratio
Sonde/TOMS Sonde/OMI Sonde/OMPS Sonde/Dobson
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Dobson)/Sonde
ρ
%

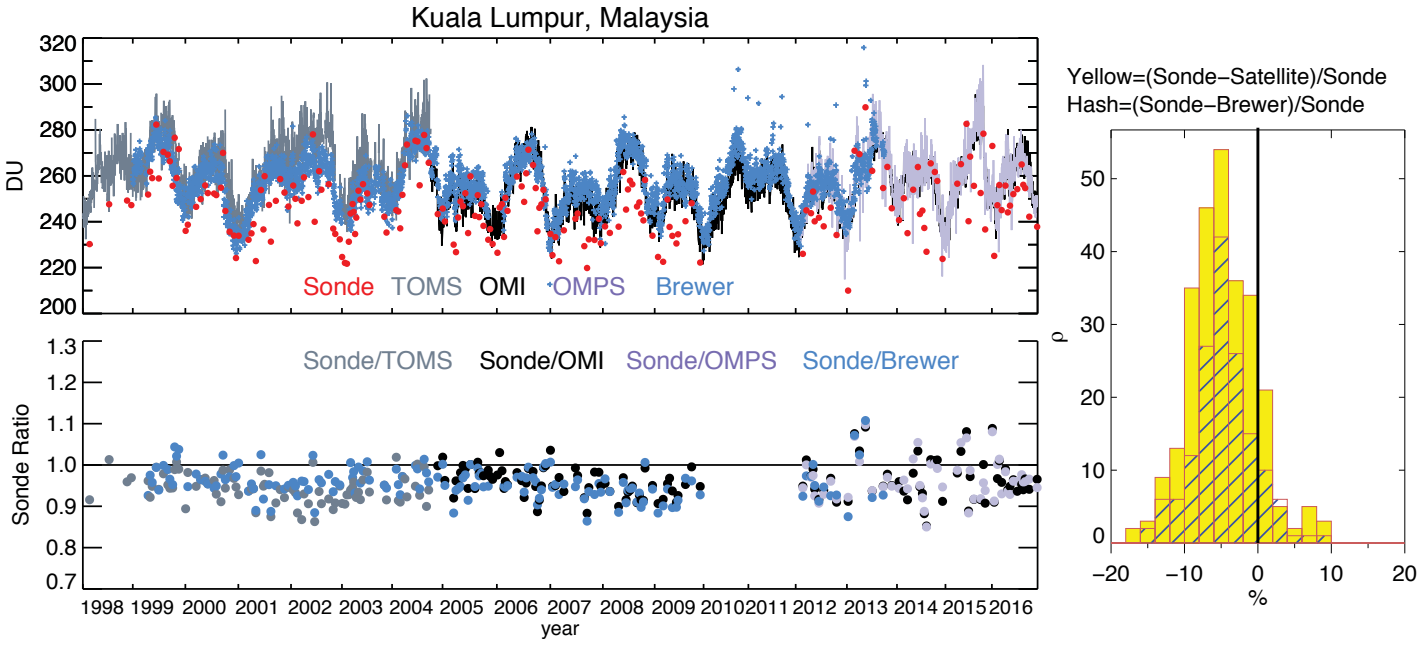
i


Nairobi, Kenya
DU
Sonde TOMS OMI OMPS Dobson
Sonde/TOMS Sonde/OMI Sonde/OMPS Sonde/Dobson
Sonde Ratio
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Dobson)/Sonde
%

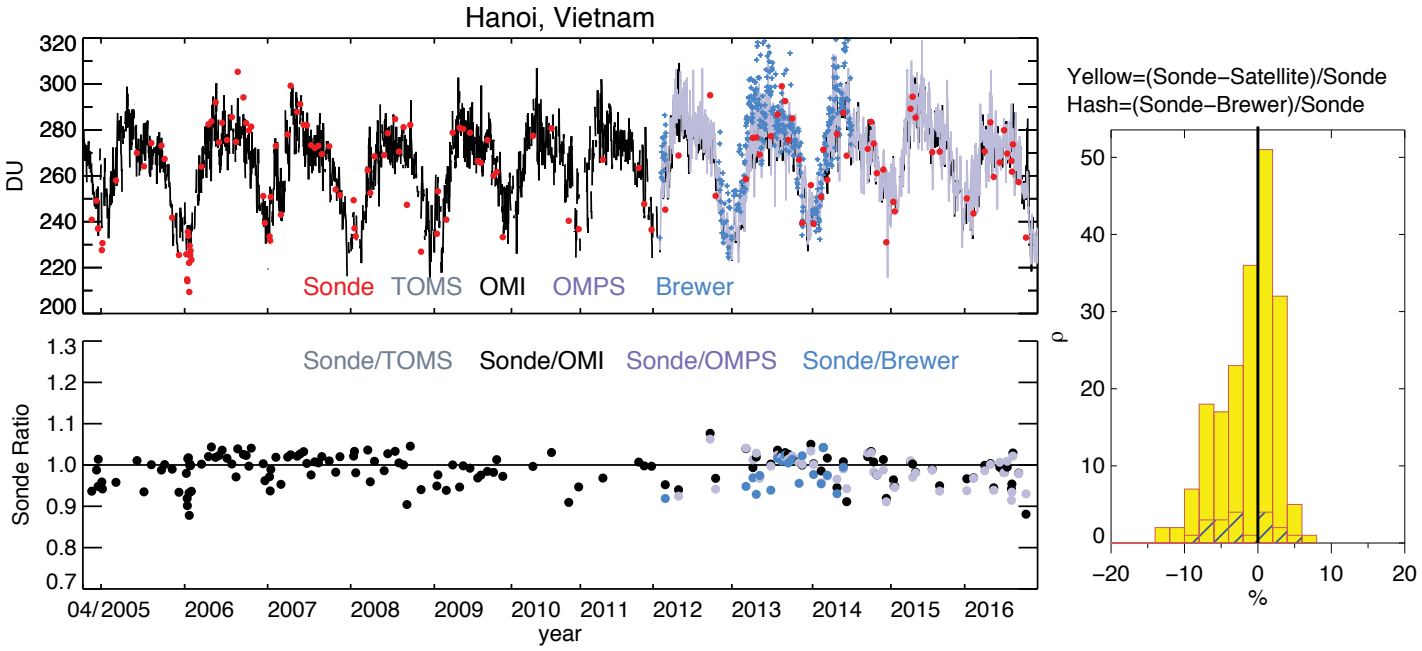
j


La Reunion Is., France
DU
Sonde TOMS OMI OMPS SAOZ
Sonde Ratio
Sonde/TOMS Sonde/OMI Sonde/OMPS Sonde/SAOZ
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−SAOZ)/Sonde
%

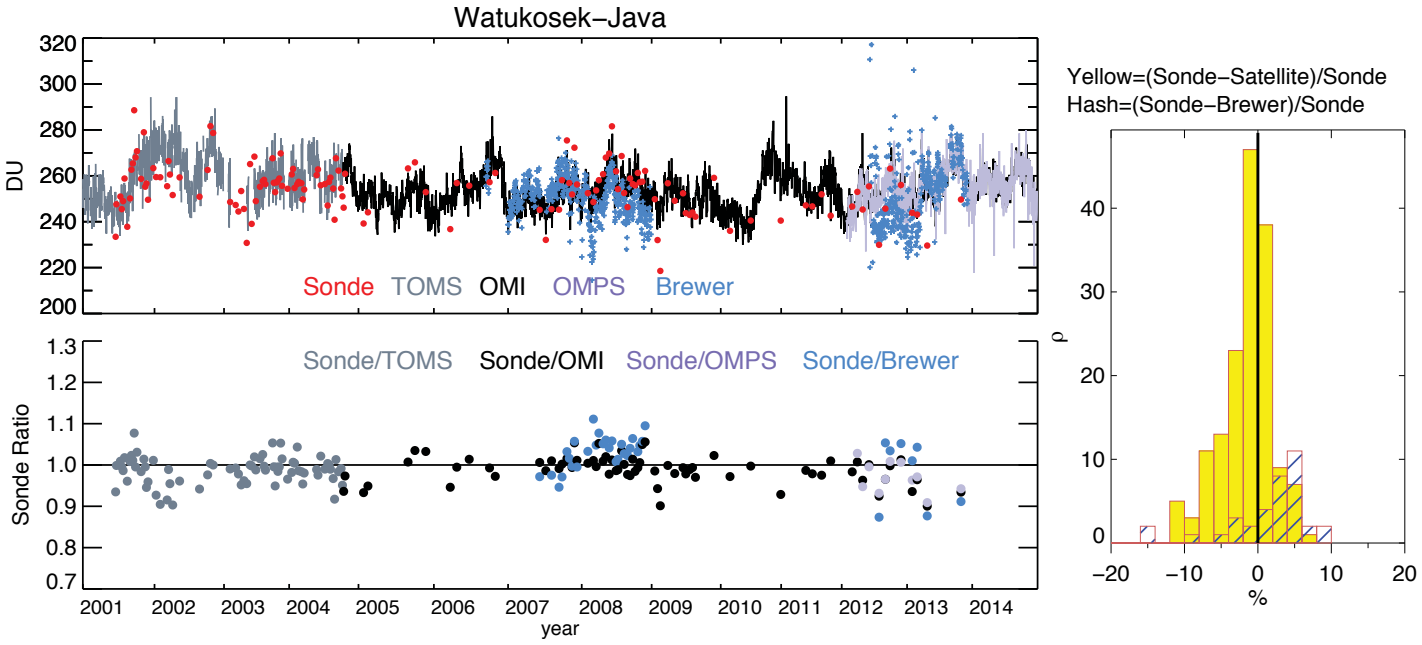
k


Kuala Lumpur, Malaysia
DU
Sonde TOMS OMI OMPS Brewer
Sonde Ratio
Sonde/TOMS Sonde/OMI Sonde/OMPS Sonde/Brewer
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Brewer)/Sonde
ρ
%

I


Hanoi, Vietnam
DU
320
300
280
260
240
220
200
Sonde
TOMS
OMI
OMPS
Brewer
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde/Brewer
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
04/2005
2006
2007
2008
2009
2010
2011
2012
2013
2014
2015
2016
year
Yellow=(Sonde−Satellite)/Sonde
Hash=(Sonde−Brewer)/Sonde
ρ
50
40
30
20
10
0
−20
−10
0
10
20
%

m


Watukosek–Java
DU
320
300
280
260
240
220
200
Sonde
TOMS
OMI
OMPS
Brewer
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde/Brewer
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
2001
2002
2003
2004
2005
2006
2007
2008
2009
2010
2011
2012
2013
2014
year
Yellow=(Sonde–Satellite)/Sonde
Hash=(Sonde–Brewer)/Sonde
ρ
40
30
20
10
0
–20
–10
0
10
20
%

n

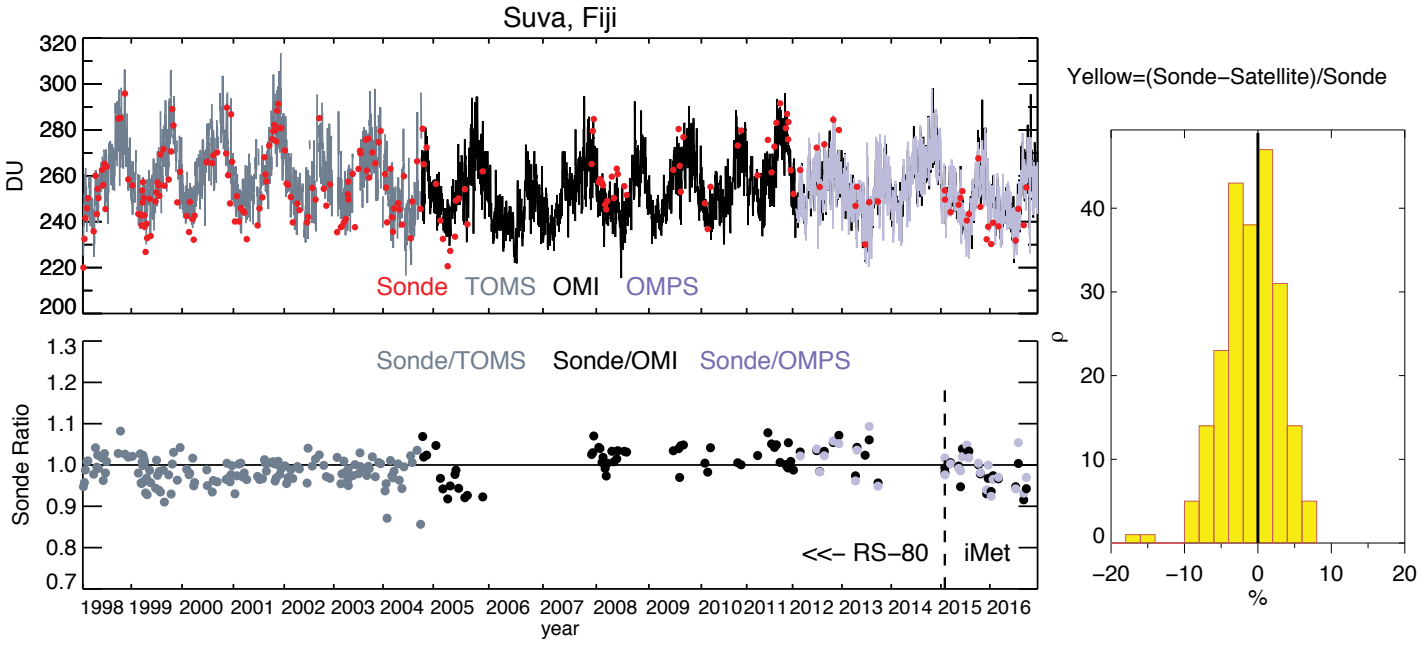
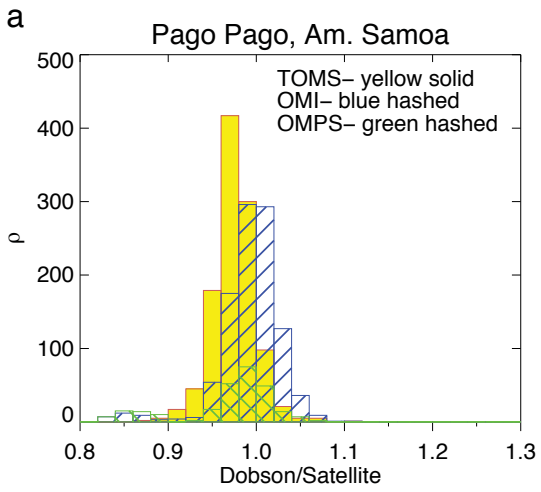
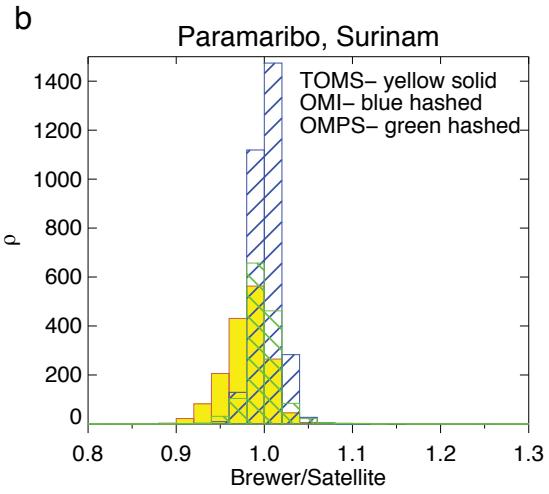
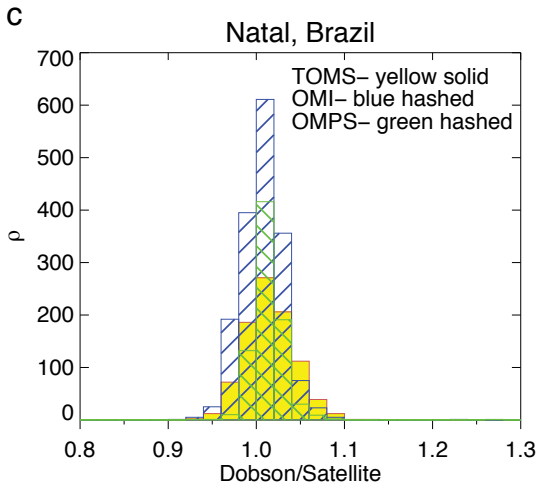

Suva, Fiji
DU
320
300
280
260
240
220
200
Sonde
TOMS
OMI
OMPS
Sonde/TOMS
Sonde/OMI
Sonde/OMPS
Sonde Ratio
1.3
1.2
1.1
1.0
0.9
0.8
0.7
<<– RS–80
iMet
1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013 2014 2015 2016
year
Yellow=(Sonde–Satellite)/Sonde
ρ
40
30
20
10
0
–20
–10
0
10
20
%

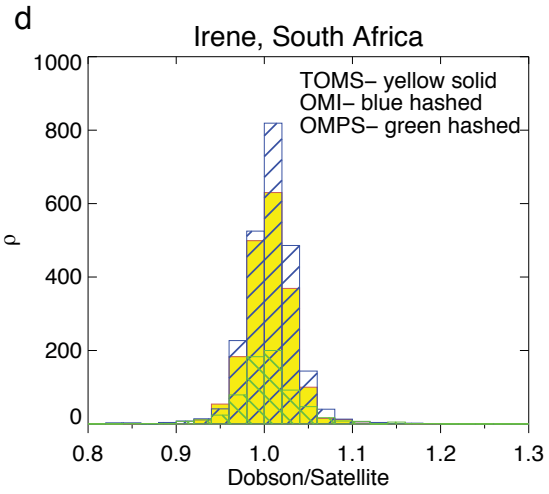
Figure 5.

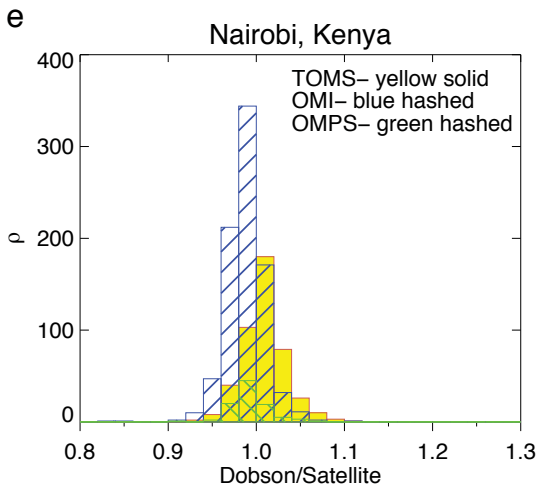

a
Pago Pago, Am. Samoa
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
500
400
300
200
100
0
0.8
0.9
1.0
1.1
1.2
1.3
Dobson/Satellite


b
Paramaribo, Surinam
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
1400
1200
1000
800
600
400
200
0
0.8
0.9
1.0
1.1
1.2
1.3
Brewer/Satellite

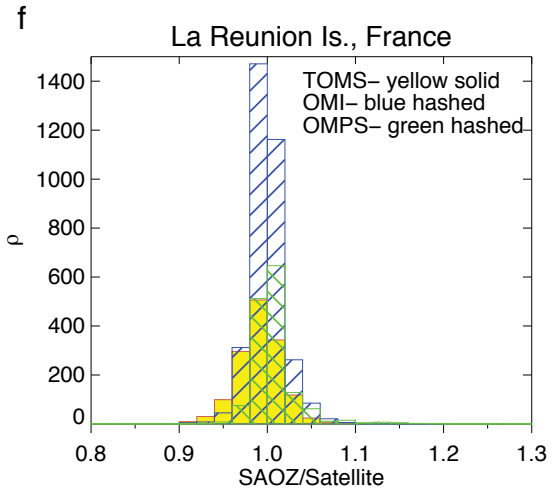
c

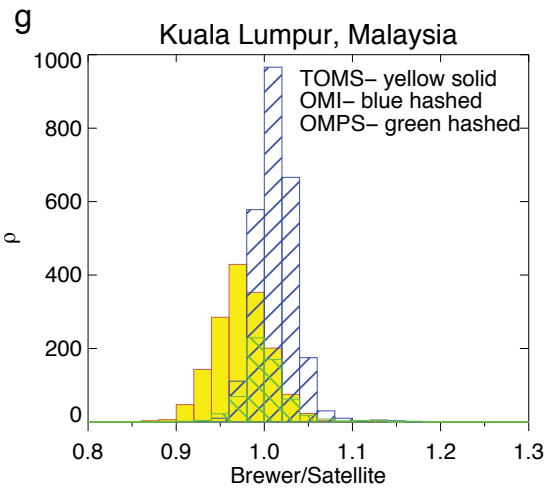

Natal, Brazil
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
700
600
500
400
300
200
100
0
0.8
0.9
1.0
1.1
1.2
1.3
Dobson/Satellite


d
Irene, South Africa
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
1000
800
600
400
200
0
0.8
0.9
1.0
1.1
1.2
1.3
Dobson/Satellite

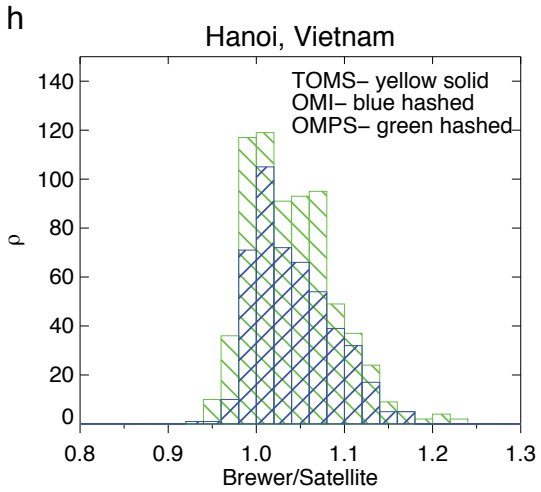

e
Nairobi, Kenya
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
400
300
200
100
0
0.8
0.9
1.0
1.1
1.2
1.3
Dobson/Satellite

f


La Reunion Is., France
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
1400
1200
1000
800
600
400
200
0
0.8
0.9
1.0
1.1
1.2
1.3
SAOZ/Satellite


g
Kuala Lumpur, Malaysia
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
1000
800
600
400
200
0
0.8
0.9
1.0
1.1
1.2
1.3
Brewer/Satellite

h


Hanoi, Vietnam
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
140
120
100
80
60
40
20
0
0.8
0.9
1.0
1.1
1.2
1.3
Brewer/Satellite

i

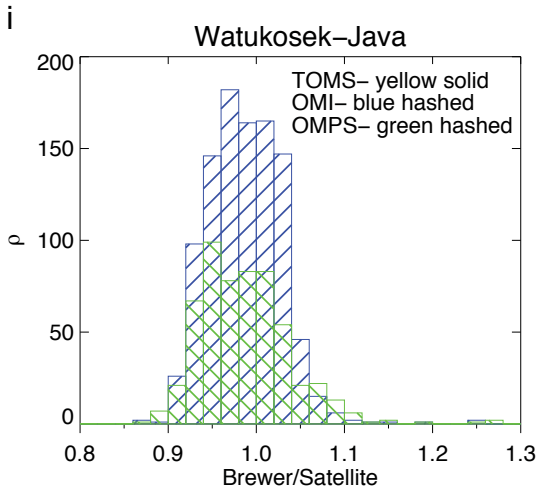
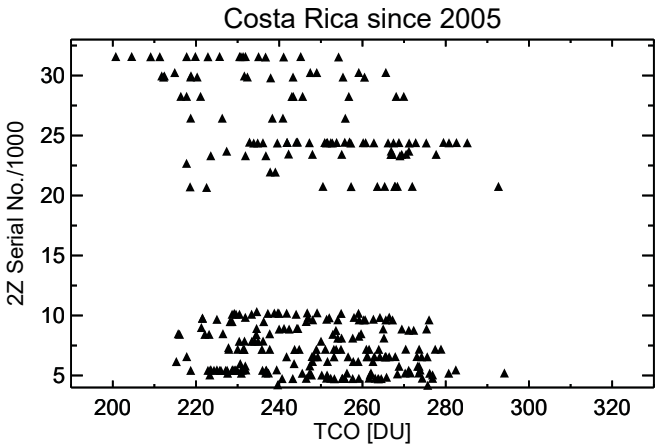

Watukosek–Java
TOMS– yellow solid
OMI– blue hashed
OMPS– green hashed
ρ
200
150
100
50
0
0.8
0.9
1.0
1.1
1.2
1.3
Brewer/Satellite

Figure 6.


Costa Rica since 2005
2Z Serial No./1000
30
25
20
15
10
5
200
220
240
260
280
300
320
TCO [DU]

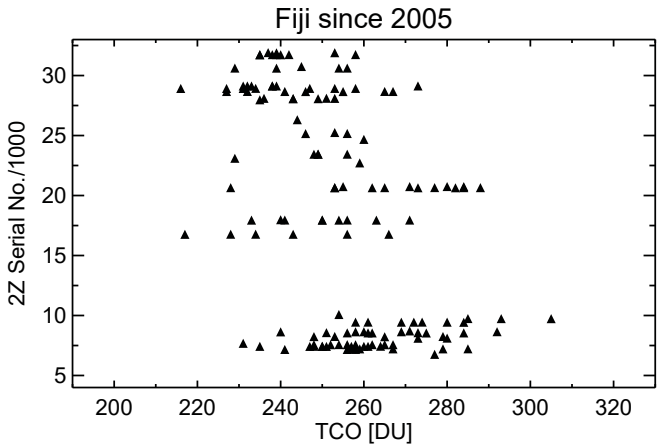

Fiji since 2005
2Z Serial No./1000
30
25
20
15
10
5
200
220
240
260
280
300
320
TCO [DU]

Figure 7.

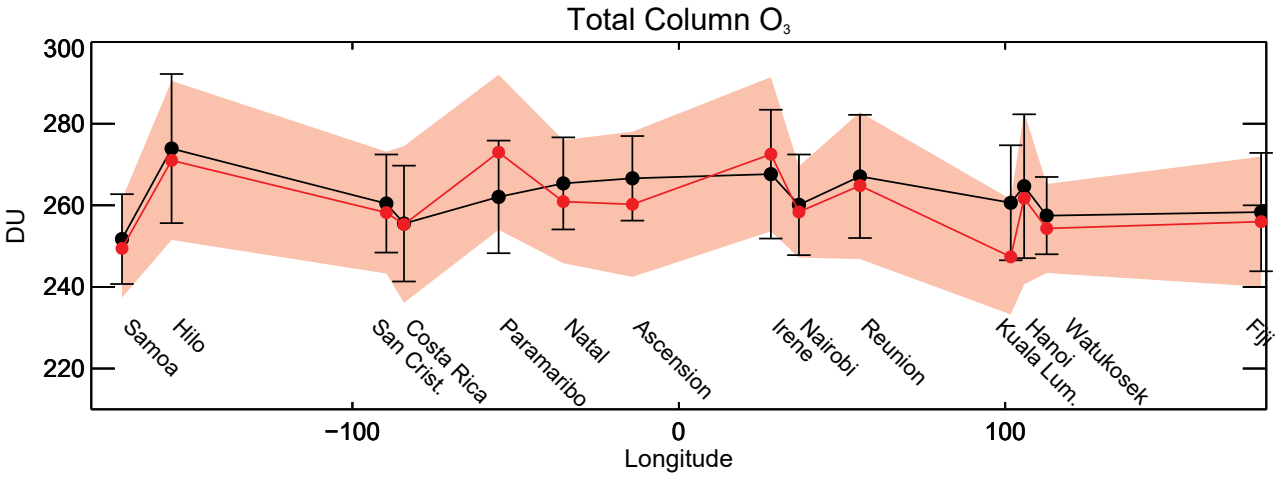
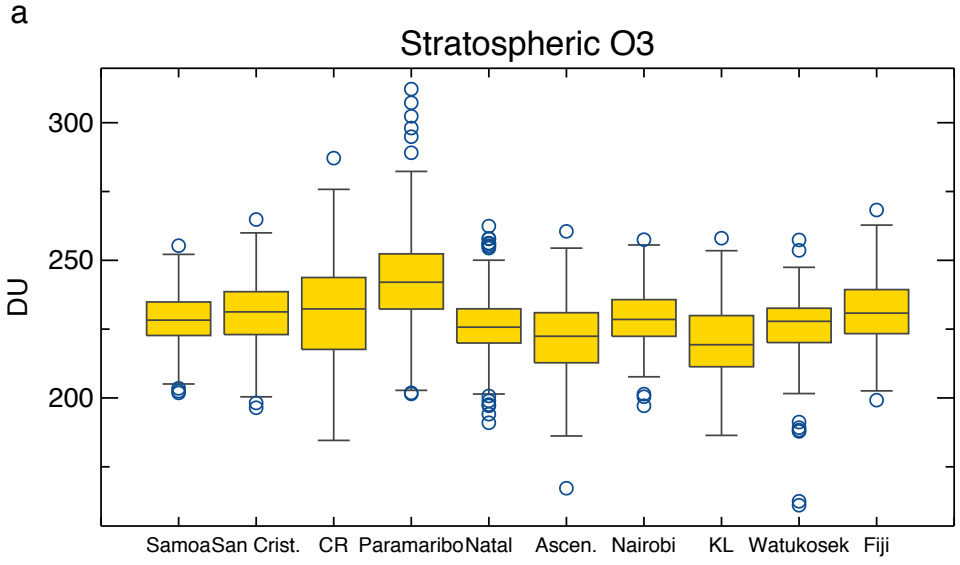
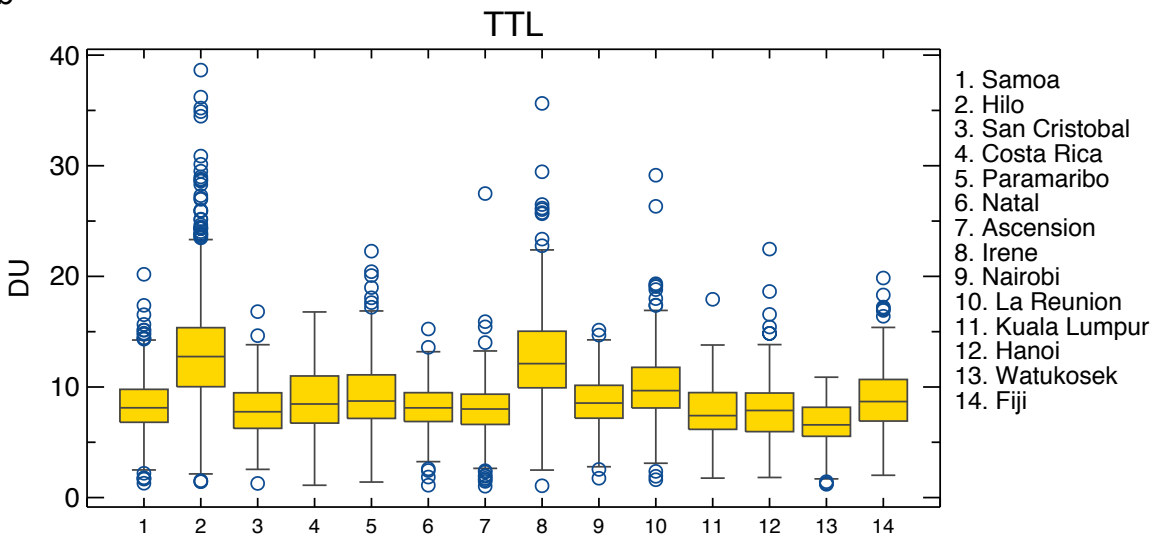

Total Column O3
300
280
260
240
220
DU
Samoa
Hilo
San Crist.
Costa Rica
Paramaribo
Natal
Ascension
Irene
Nairobi
Reunion
Kuala Lum.
Hanoi
Watukosek
Fiji
−100
0
100
Longitude

Figure 8.

a


Stratospheric O3
DU
300
250
200
Samoa
San Crist.
CR
Paramaribo
Natal
Ascen.
Nairobi
KL
Watukosek
Fiji

b


TTL
DU
40
30
20
10
0
1
2
3
4
5
6
7
8
9
10
11
12
13
14
1. Samoa
2. Hilo
3. San Cristobal
4. Costa Rica
5. Paramaribo
6. Natal
7. Ascension
8. Irene
9. Nairobi
10. La Reunion
11. Kuala Lumpur
12. Hanoi
13. Watukosek
14. Fiji

c

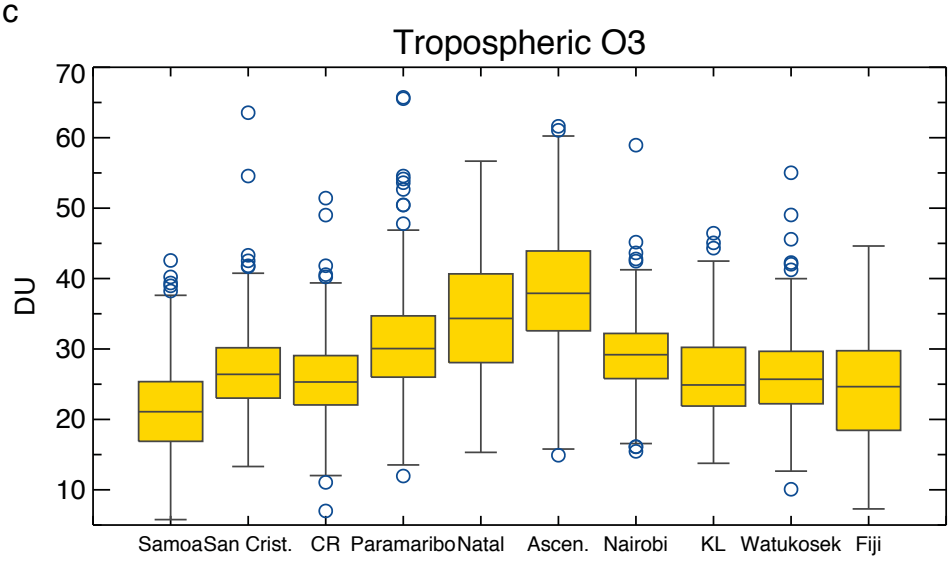
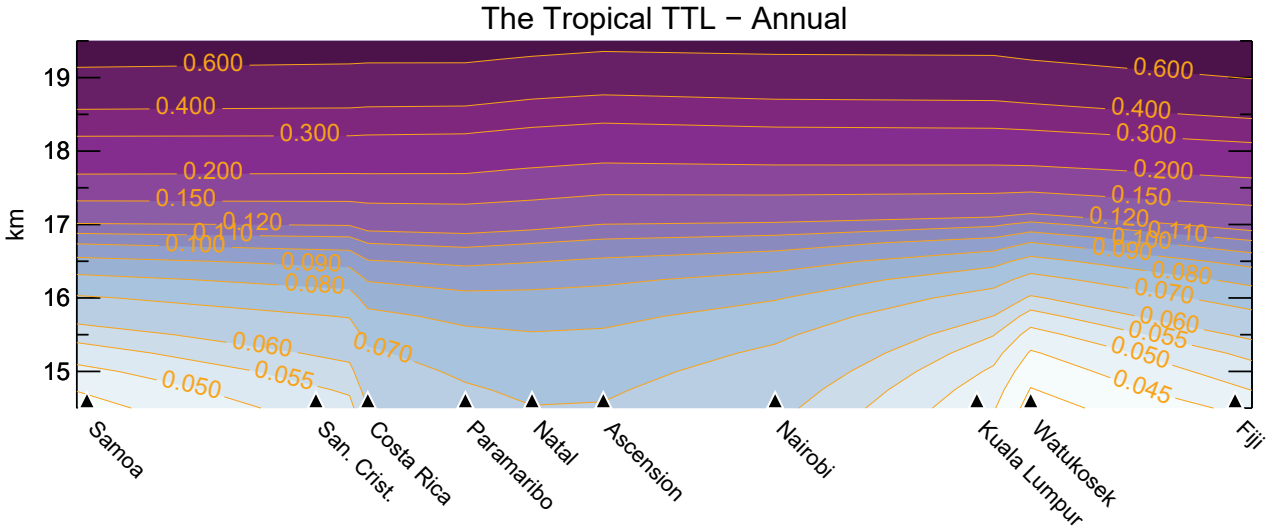
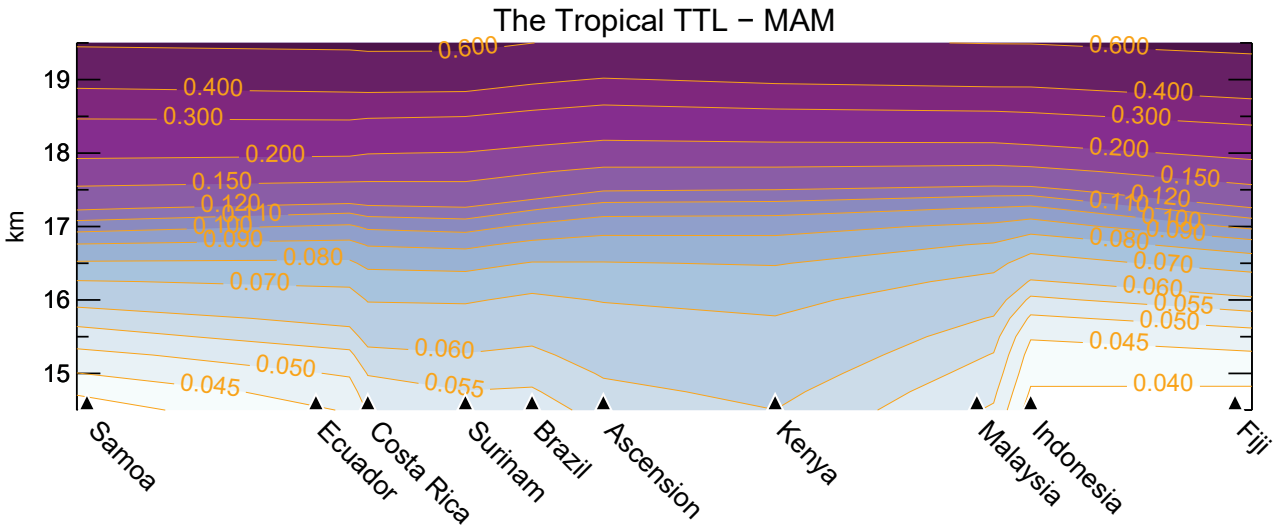

Tropospheric O3
DU
70
60
50
40
30
20
10
Samoa
San Crist.
CR
Paramaribo
Natal
Ascen.
Nairobi
KL
Watukosek
Fiji

Figure 9.

a


The Tropical TTL − Annual
km
19
18
17
16
15
0.600
0.400
0.300
0.200
0.150
0.120
0.110
0.100
0.090
0.080
0.070
0.060
0.055
0.050
0.045
Samoa
San. Crist.
Costa Rica
Paramaribo
Natal
Ascension
Nairobi
Kuala Lumpur
Watukosek
Fiji

b


The Tropical TTL − MAM
km
19
18
17
16
15
0.600
0.400
0.300
0.200
0.150
0.120
0.110
0.100
0.090
0.080
0.070
0.060
0.055
0.050
0.045
0.040
Samoa
Ecuador
Costa Rica
Surinam
Brazil
Ascension
Kenya
Malaysia
Indonesia
Fiji

c

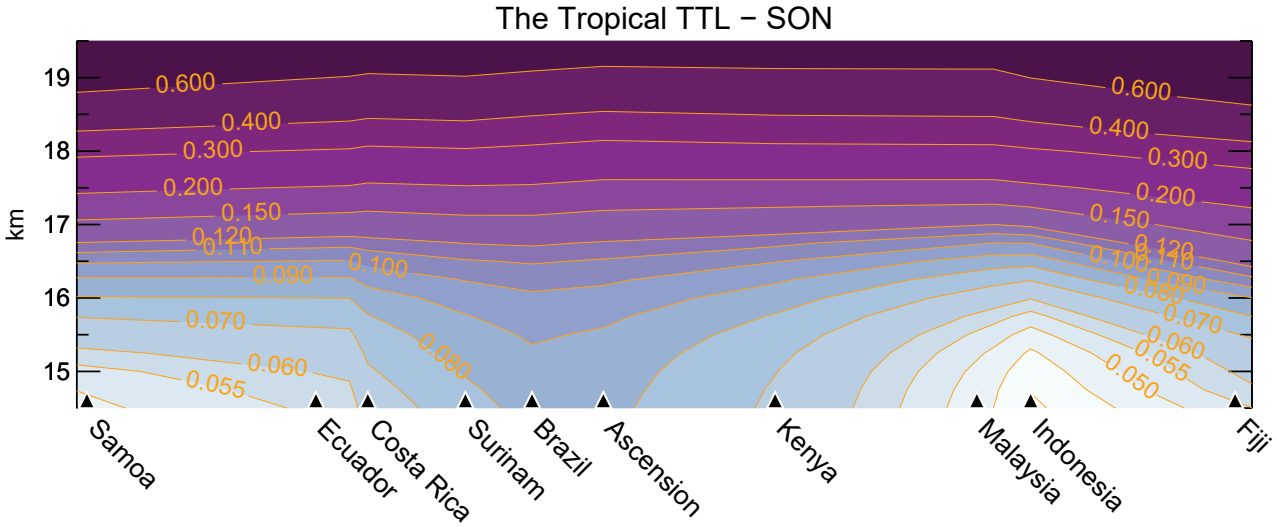

The Tropical TTL − SON
km
19
18
17
16
15
0.600
0.400
0.300
0.200
0.150
0.120
0.110
0.090
0.100
0.070
0.080
0.060
0.055
0.050
Samoa
Ecuador
Costa Rica
Surinam
Brazil
Ascension
Kenya
Malaysia
Indonesia
Fiji

Figure 10.

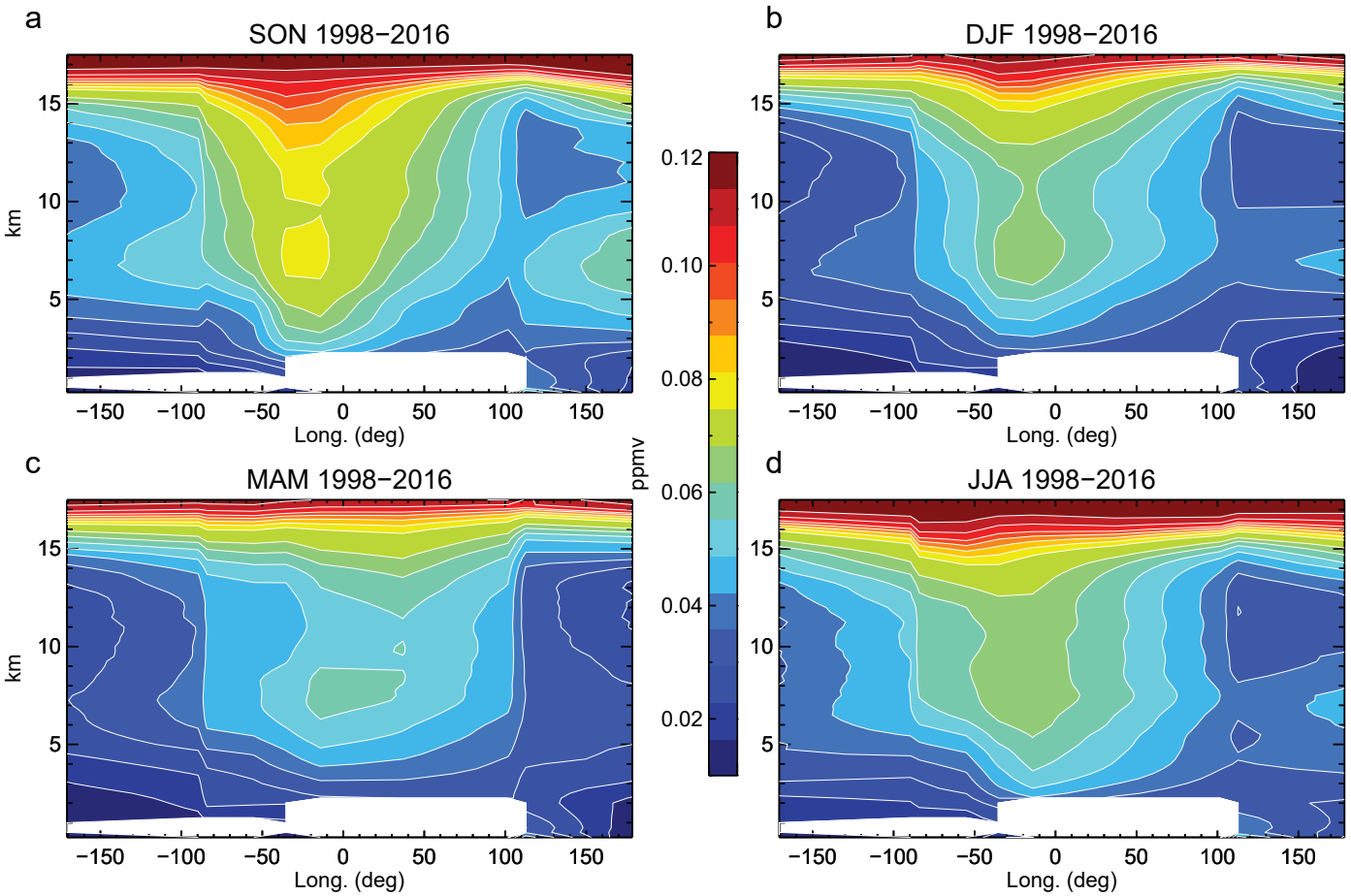

a
SON 1998–2016
b
DJF 1998–2016
c
MAM 1998–2016
d
JJA 1998–2016
km
Long. (deg)
ppmv
0.12
0.10
0.08
0.06
0.04
0.02
15
10
5
−150
−100
−50
0
50
100
150

Figure 11.

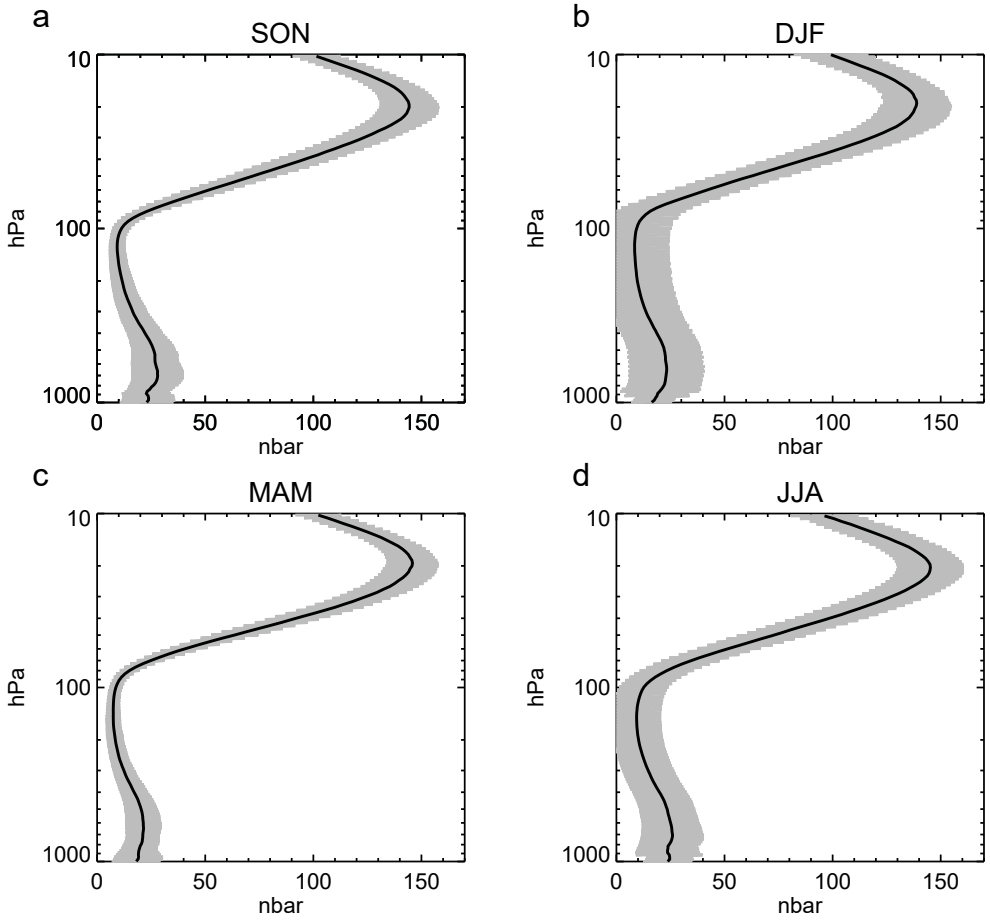
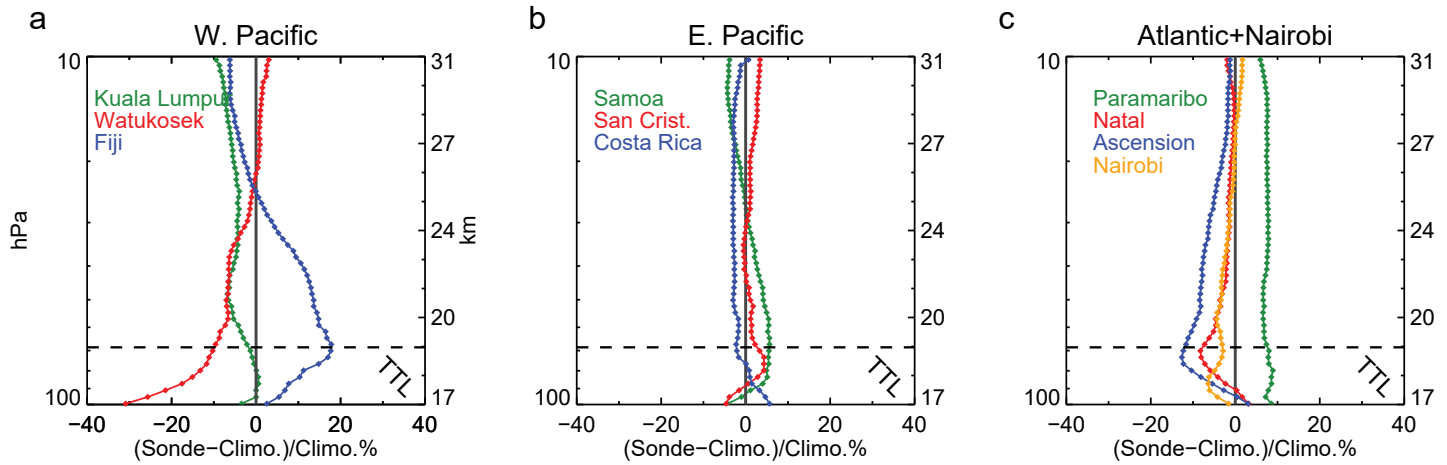

a
SON
b
DJF
c
MAM
d
JJA
hPa
10
100
1000
0
50
100
150
nbar

Figure 12.


a
W. Pacific
Kuala Lumpur
Watukosek
Fiji
hPa
km
10
100
31
27
24
20
17
TTL
−40
−20
0
20
40
(Sonde−Climo.)/Climo.%
b
E. Pacific
Samoa
San Crist.
Costa Rica
c
Atlantic+Nairobi
Paramaribo
Natal
Ascension
Nairobi