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1	Forest resilience and tipping points at different spatio-temporal scales:
2	approaches and challenges
3	
4	Running head: Forest resilience to environmental change
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# 39 Summary

40	1.	Anthropogenic global change compromises forest resilience, with profound
41		impacts to ecosystem functions and services. This synthesis paper reflects on the
42		current understanding of forest resilience and potential tipping points under
43		environmental change, and explores challenges to assessing responses using
44		experiments, observations and models.
45	2.	Forests are changing over a wide range of spatio-temporal scales, but it is often
46		unclear if these changes reduce resilience or represent a tipping point. Tipping
47		points may arise from interactions across scales, as processes such as climate
48		change, land-use change, invasive species, or deforestation gradually erode
49		resilience and increase vulnerability to extreme events. Studies covering
50		interactions across different spatio-temporal scales are needed to further our
51		understanding.
52	3.	Combinations of experiments, observations, and process-based models could
53		improve our ability to project forest resilience and tipping points under global
54		change. We discuss uncertainties in changing CO <sub>2</sub> concentration and quantifying
55		tree mortality as examples.
56	4.	Synthesis: As forests change at various scales, it is increasingly important to
57		understand whether and how such changes lead to reduced resilience and potential
58		tipping points. Understanding the mechanisms underlying forest resilience and
59		tipping points would help in assessing risks to ecosystems, and presents
60		opportunities for ecosystem restoration and sustainable forest management.

- 61 Keywords: climate change, CO<sub>2</sub>, drought, environmental change, mortality, regime shifts,
- 62 review, spatio-temporal scales, vegetation–climate interactions, vulnerability

## 63 Introduction

64 Global environmental changes in, for instance, atmospheric CO<sub>2</sub> concentration, 65 climate, nitrogen deposition, and their interactions are affecting forests around the world 66 (Hyvönen et al. 2007; Phillips et al. 2009; Lindner et al. 2010; Dietze & Moorcroft 2011; 67 Brouwers et al. 2012; Poulter et al. 2013). These changes occur in combination with 68 direct local anthropogenic drivers of forest change such as land conversion and intensive 69 forest management (e.g., Malhi et al. 2008; Barona et al. 2010; DeFries et al. 2010; 70 Gibbs et al. 2010; Kurz 2010; Vayreda et al. 2012; Bryan et al. 2013). The structural 71 (e.g., reduction in biomass or canopy cover) and functional changes (e.g., changes in 72 community structure) in forests in response to these global and local drivers may 73 compromise the capacity of forest ecosystems to recover from natural or human-induced 74 perturbations (Laurance 2004; Lenton et al. 2008). Resilience theory suggests that 75 ecosystems exhibit threshold responses to changing environmental drivers, and move 76 towards a new regime when a critical threshold is exceeded (Scheffer *et al.* 2001; 77 Scheffer *et al.* 2003). However, while evidence for the effects of gradual environmental 78 change on forests exists, the theory of regime shifts has rarely been tested in forest ecosystems and evidence for tipping points at which feedbacks cause ecosystems to enter 79 80 alternative stable states remains sparse (Hirota et al. 2011). 81 In the scope of the special feature on "Forest Resilience, Tipping Points and Global 82 Change Processes", we reflect here on the current understanding of forest resilience and 83 potential tipping points under environmental change at different spatial and temporal 84 scales, and the challenges to quantify these ecosystem attributes with experiments, 85 observations and models.

## 86 **Concepts and definitions**

87 We use the term forest resilience in the sense of Scheffer (2009) as "the ability of a 88 forest to absorb disturbances and re-organize under change to maintain similar functioning and structure". This definition is wider than what is understood as 89 90 "ecological resilience" (sensu Holling 1973) to account for the fact that systems rarely 91 recover to exactly the same pre-disturbance conditions (Scheffer 2009). Furthermore this 92 definition allows resilience to be framed as the rate of recovery after a disturbance and as 93 the maximum disturbance a forest can absorb before switching to a different ecosystem 94 type (Gunderson 2000; Scheffer 2009) – both being important features of resilience in the 95 context of global change. In systems with multiple stable states, reduced resilience can 96 lead to regime shifts (Scheffer et al. 2001). Scheffer (2009) defines a regime shift as a 97 "relatively sharp change from one regime to a contrasting one, where a regime is a 98 dynamic 'state' of a system with its characteristic stochastic fluctuations and/or cycles". 99 A tipping point describes a threshold in conditions at which a small change in conditions 100 leads to a strong change in the state of a system (cf. Brook et al. 2013). 101 A typical example to illustrate these concepts is the so-called "Amazon forest dieback" (White et al. 1999, Cox et al. 2004, Malhi et al. 2008; Rammig et al. 2010). 102 103 Model simulations including changing precipitation and/or ongoing anthropogenic 104 deforestation in the Amazon show reduced evapotranspiration and hence lower amounts 105 of water recycled for regional precipitation (e.g. Cox et al. 2004, Sampaio et al. 2007; 106 Costa & Pires 2010). Such an altered hydrological cycle may lead to further tree 107 mortality and reduces forest resilience to recover from fire and drought. Repeated 108 occurrences of strong drought/fire years, for example in the context of El-Nino events,

109 could drive the Amazon past a tipping point into a savannah-like state. Sampaio *et al.* 

110 (2007) suggested that at 40% deforestation, such a regime shift could occur in the

111 Amazon basin. Hirota et al. (2011) predicted that the south-eastern parts of the Amazon

112 are the least resilient and most likely to turn into savannah or a treeless state.

Another example to illustrate resilience and tipping points in forests at a different spatio-temporal scale is the analysis of drought-induced mortality by Camarero *et al.* (2015). They suggest a framework in which already stressed trees with reduced resilience are further affected by drought up to a point where hydraulic failure occurs or carbon reserves are depleted and additional drought stress can lead to widespread tree mortality

and a tipping point.

119 The spatio-temporal scales at which forest resilience and potential tipping points 120 can be assessed are of critical importance in understanding and interpreting ecological 121 processes (Wiens 1989; Levin 1992). While sophisticated concepts and methods for 122 delineating scales have been discussed recently (Angeler *et al.* 2011; Chave 2013; Allen 123 et al. 2014; Nash et al. 2014; Soranno et al. 2014), we here use the following scale 124 classification to group studies that are relevant to address our main objective of 125 understanding forest resilience and tipping points (Table 1). We refer to 'local' scale as 126 the spatial extent from hundreds of m<sup>2</sup> up to 10 km<sup>2</sup>, 'regional' scale as larger areas 127 spanning >10 km<sup>2</sup> to water catchment, ecoregion, country or continental-level, and the 128 'global' scale. For temporal scales, we define 'short-term' as periods of up to 10 years 129 capturing processes such as tree mortality, and 'long-term' as decadal to millennial time 130 periods capturing longer-term processes such as species replacement (see Chapin et al. 131 2004; Hughes et al. 2013a).

# 132 At which temporal and spatial scales can forest resilience and potential tipping

133 points be assessed?

134 Local scale responses

135	At the local scale, short-term and long-term responses to environmental change
136	have been well documented in different types of forests (Table 1). For example,
137	Mediterranean ecosystems have been profoundly transformed by long human use and
138	their response is well studied across different continents. The capacity of Mediterranean
139	ecosystems to respond to disturbances strongly interacts with climate condition.
140	Vegetation changes in the driest regions tend to be persistent and sometimes very
141	difficult to reverse because of positive feedbacks that maintain alternative vegetation
142	states (Holmgren & Scheffer 2001; Acacio et al. 2007, 2009, Kefi et al. 2007).
143	Mediterranean climate regions are likely to become drier and warmer over the decades to
144	come (Stocker et al. 2013), which will likely impact on their resilience as is already
145	occurring in southwest Western Australia (e.g. Brouwers et al. 2012). Here, forest tree
146	species are increasingly showing severe dieback and mortality in response to extreme
147	climatic conditions (Bader et al. 2014, Matusick et al. 2012, Matusick et al. 2013) but
148	whether this constitutes a tipping point is unclear.
149	Allen et al. (2010) concluded that in already dry regions, forests are most
150	vulnerable to further drought possibly indicating reduced resilience. It is precisely these
151	dry regions that have become drier during the second half of the last century (Bates et al.
152	2008; Steinkamp & Hickler 2015). Relative to the small spatial extent of forests with a
153	dry climate, many of the observed drought-induced forest mortality events in Allen et al.
154	(2010) occurred in drought-prone forests.

## *Regional scale responses*

157	At the regional scale, short and long-term responses to environmental change have
158	been documented in different forest types (Table 1). Allen et al. (2010) provide examples
159	of reduced resilience due to drought-induced tree mortality over both short and long
160	periods of time. Their findings are confirmed by a recent risk analysis of ecosystem
161	vulnerability to drought by Van Oijen et al. (2013) indicating that drought vulnerability is
162	expected to increase in Southern Europe.
163	Inferring resilience at regional scales is challenging. Hirota et al. (2011) and Staver
164	et al. (2011) inferred tropical forest resilience from remotely-sensed tree cover
165	distribution. They studied tree cover classes over different precipitation levels that
166	suggest instabilities of forest cover. The "resilience maps" of Hirota et al. (2011) show
167	the probability of finding tropical forests, savannas and treeless conditions in a specific
168	location given the mean annual precipitation. These maps highlight which parts of Africa,
169	Australia and South America are more or less likely to persist in certain vegetation state
170	if disturbed pointing towards possible tipping points. This approach has also been used to
171	infer the resilience of boreal ecosystems (Scheffer et al. 2012a). However, Hanan et al.
172	(2014) recently pointed out that such tree cover datasets may be unsuitable for the
173	diagnosis of alternative stable states due to their inherent error distribution. Moreover,
174	remotely-sensed data are two-dimensional, time-delayed and often not available as time
175	series.
176	

For the Amazon forests, Cox *et al.* (2004) have projected a tipping point based onclimate-change induced rainfall reductions leading to vegetation-atmosphere feedbacks

178 and further drying using coupled carbon cycle-climate models. However, in more recent 179 studies, Cox et al. (2013) and Huntingford et al. (2013) show that the Amazon is 180 probably more resilient than projected in the original Cox et al. (2004) study. This change 181 in perspective is due to a better understanding of model uncertainties. Uncertainties with 182 regard to parameterization and structural representation of physiological processes such 183 as CO<sub>2</sub>-fertilization or responses to elevated temperatures have been highlighted as being 184 more important than uncertainties related to emission scenarios and climate projections 185 (Huntingford et al. 2013). However, another study, using an Earth System Model, found 186 projected changes in dry-season length to be the dominant single factor explaining 187 differences in forest dieback projections (Good et al. 2013). 188

189 Global scale responses

190 Recently, the potential for an anthropogenically-driven global tipping point of the 191 biosphere has been discussed (Barnosky et al. 2012; Brook et al. 2013; Hughes et al. 192 2013b; Lenton & Williams 2013). Barnosky et al. (2012) argue that in the past, state 193 shifts of the totality of the global ecosystem have occurred and that a tipping point of the 194 Earth's biosphere is plausible given the strong and enduring alterations of natural 195 ecosystems by human interventions such as land-use change. Brook et al. (2013) 196 highlighted three criteria that need to be fulfilled for such a state shift to occur globally: 197 1) a spatially homogenous response towards changing environmental conditions, 2) an 198 interconnection of responses, and 3) a spatial homogeneity of drivers of environmental 199 change. They conclude that spatial heterogeneity in drivers of, and responses to, 200 environmental change as well as the lack of continental connectivity are smoothing

201 global-scale responses and will therefore unlikely result in a global tipping point (Brook 202 et al. 2013). Hughes et al. (2013b) argue that the speed of change does not determine the 203 presence or absence of a global tipping point but rather the nonlinear relation between the 204 driving force of ecosystem change and the state of the ecosystem. Lenton & Williams 205 (2013) add that it is the existence of possible feedbacks (and the balance between positive 206 and negative feedbacks) between different components of the Earth system that 207 ultimately determines whether a global tipping of the biosphere exists. They also stress 208 that a clear definition of the "biosphere" is needed to discuss the possibility of a global 209 tipping point. Lenton & Williams (2013) argue that although a strong global ecological 210 response may be perceived as a tipping point of the biosphere, it is not necessarily 211 associated with a tipping point caused by a feedback within the biosphere itself. They 212 agree with Brook et al. (2013) in that interactions of smaller-scale tipping points that 213 influence and trigger each other are more likely than a global tipping point (Lenton & 214 Williams 2013). However, Lenton & Williams (2013) also point out that under certain 215 circumstances a cascade of tipping points may lead to a tipping point with a global reach. 216

## 217 Interactions across scales

Besides the individual impacts of natural and anthropogenic pressures, it is perhaps primarily the interaction of processes across scales that determine forest resilience and tipping points (cf. Gunderson & Holling 2002). For example, the interaction of changes in longer-term mean climate with changes in short-term climate variability determines plant responses to climate change (Reyer *et al.* 2013). In regions where recent climate change has had a significant impact, tipping points may occur when extreme climatic

224 events take place. Hence, continuous slow changes in climatic conditions alter the 225 resilience of forests until extreme events or disturbances, such as prolonged drought, push 226 a specific forest ecosystem across a tipping point into a new state. We therefore argue 227 that, in regions where gradual and persistent long-term drying and warming trends are 228 evident, forest resilience and tree health has likely decreased (e.g., van Mantgem et al. 229 2009; Brouwers *et al.* 2012). This decrease maybe a consequence of idiosyncratic 230 interactions of increasing physiological stress, altered natural disturbance regimes and 231 comparably rapid climatic changes, which makes forests vulnerable to regime shift in the 232 future. Similar gradual erosion of resilience through land-use change, invasive species 233 and deforestation may increase vulnerability to extreme events pushing forests towards a 234 tipping point.

235 Interaction across scales may however also increase resilience at larger scales. 236 Higgins & Scheiter (2012) for example, show that although potentially large areas across 237 Africa are threatened by regime shifts in vegetation cover, asynchrony in their timing 238 may dampen their effect on the Earth system globally. Thus, although abrupt transitions 239 can occur locally with potentially devastating effects, they may smooth out on lager 240 spatial scales and prevent strong regional to global scale tipping points (Higgins & 241 Scheiter 2012). Similarly, it is important to consider that over larger spatial or temporal 242 scales, there is also a larger potential for adaptation of species and communities to 243 changing conditions, which may increase resilience and further dampen the possibility for 244 a tipping point.

# What are the challenges for understanding forest resilience and tipping pointsunder global environmental change?

247 This section focuses on the challenges in capturing different environmental 248 changes in experiments, observations and models to assess the effects of changing 249 environmental conditions on forests and how this limits our understanding of resilience 250 and tipping points (Table 2). There are, in theory, a number of early warning signals 251 when approaching a regime shift. For instance, these include indicators related to the 252 'critical slowing down' of a system under stress, a flickering in time-series of relevant 253 indicators (e.g. population abundance, nutrient flows), or simply approaching a state 254 where the system has not been observed under current conditions (see Scheffer et al. 255 (2012b) and Camarero *et al.* (2015)). However, the ability to predict a tipping point will 256 remain limited due to the inherent stochasticity of the mechanisms involved (Scheffer et 257 al. 2012b).

258 Another challenge for understanding forest resilience and tipping points is that 259 background conditions are changing, for example CO<sub>2</sub>-concentration and temperature. 260 Changes in human-driven pressures such as alterations of fire regimes, deforestation, 261 management and their interactions, also need to be considered. Consequently, to deepen 262 our understanding of forest resilience and tipping points integrating insights from 263 experiments, observations and modelling is needed. Experiments may be helpful to 264 determine potential thresholds of ecosystem variables that influence tipping point 265 behaviour, particularly when manipulating whole ecosystems (e.g., Nepstad *et al.* 2007, 266 Brando *et al.* 2008). Observations show conditions under which tipping of a forest 267 ecosystem is possible or has already happened (e.g., Leadley et al. 2010). Taken together,

this understanding of resilience and tipping points may help to improve process-basedmodels.

270 Process-based models can be used as diagnostic tools to understand system 271 components that are complicated or costly to measure or as predictive tools that allow 272 using different scenarios to project forest responses to changing environmental conditions 273 (e.g., Reyer *et al.* 2014). Improving process-based models is crucial, because they tend to 274 be very complex and often include mechanisms that are still poorly understood, 275 particularly under changing environmental conditions. With regard to the application of 276 such models to assess resilience and detect tipping points, there is prospect for 277 implementing observations and experimentally- or theoretically-derived mechanisms in 278 models. For example, integrating species-specific rooting profiles into a forest model has 279 been shown to affect competition in mixed forests (Reyer et al. 2010). Similar 280 implementations in dynamic global vegetation models could help to better explain forest-281 savannah interactions as a response to changes in extreme precipitation (Kulmatiski & 282 Beard 2013). 283 Nonetheless, because models are simplified representations of reality, they should 284 always be considered as generators of hypotheses or projectors of possible pathways, and 285 therefore further checked against experimental and independent data on system 286 responses. Most current forest and vegetation models were not meant to simulate tipping 287 points and may not be able to do so because of underpinning assumption like the 288 inclusion of equilibrium pools representing an average steady state response, and the lack 289 of feedbacks between vegetation and climate. Moreover, most models still do not contain 290 disturbance processes (e.g. pathogens/insects, wind-throw), which are crucial for

291 simulating mortality of trees and abrupt vegetation shifts (Steinkamp & Hickler 2015). A 292 future challenge is to incorporate these processes in forest models to better simulate 293 resilience and tipping points. In a first attempt 'tipping point' behaviour could be 294 implemented into models by introducing them explicitly, depending on threshold values. 295 However, in that case the model would depend totally on the parameter or threshold 296 value imposed. The larger challenge is to build models that do not have these explicit 297 thresholds hard-coded, but have tipping point as an emerging property as a result of the 298 combination of the underlying processes included in the model. 299 Although most forest models are still limited for exploring tipping points and 300

regime shifts, they are usually able to identify indicators of reduced resilience, for 301 example by predicting changes in biogeochemical cycles. A metric developed by Heyder 302 et al. (2011) indicates severe ecosystem change highlighting loss of resilience and the 303

potential approach of tipping points (Gerten et al. 2013; Warszawski et al. 2013).

304 Recently, also more simple ecological models have been developed that are explicitly

305 designed to simulate regime shifts (e.g., Staver & Levin 2012; van Nes et al. 2013). In

306 the following sections we take the example of  $CO_2$ -effects and mortality as two key

307 challenges for understanding forest resilience and tipping points under changing

308 environmental conditions in process-based forest models.

309

#### 310 *Effects of changing CO*<sub>2</sub>*-concentrations*

311 From a plant's perspective, it is not the role of  $CO_2$  in atmospheric forcing that is 312 most important, but its direct effect on photosynthesis and gas exchange. Since rising 313  $CO_2$  tends to increase productivity and reduce water use, it may increase forest resilience

314	or lead to greening of non-forested ecosystems (e.g. Higgins & Scheiter 2012). However,
315	there are counter-examples where rising CO <sub>2</sub> has decreased resilience. In some cases,
316	elevated $CO_2$ has increased plant leaf area thus increasing the impact of drought when it
317	occurs (Ghannoum & Way 2011). Also, since elevated CO2 reduces transpiration, leaf
318	temperatures tend to increase, which may increase the impact of extremely high
319	temperatures (Warren et al. 2011) and decrease resilience to frost events (Barker et al.
320	2005). Increased productivity under elevated CO <sub>2</sub> may also increase litter and reduce
321	litter decomposition rates (Hyvönen et al. 2007) leading to increased fuel and higher
322	vulnerability to fire.
323	Although the effects of elevated $CO_2$ can be observed close to natural sources of
324	CO <sub>2</sub> (Haettenschwiler et al. 1997), derived from time-series of measured forest functions
325	under the recent increase in CO <sub>2</sub> (e.g., Keenan et al. 2013) or inferred from
326	paleoecological data, experiments and models are most commonly used to study how
327	trees and forests respond to CO <sub>2</sub> . The logistics involved in experiments manipulating
328	atmospheric CO <sub>2</sub> levels within existing forests are tremendous and explain why, to date,
329	very few of these experiments exist in natural or semi-natural forests (Norby et al. 2010;
330	Leuzinger & Bader 2012; Bader et al. 2013). Such experiments remain largely artificial,
331	because the atmospheric CO <sub>2</sub> increase has to occur stepwise instead of gradually as in
332	reality. Furthermore, because transpiration is affected through stomatal closure, forests
333	under elevated atmospheric CO <sub>2</sub> affect local climate (e.g. air humidity). These changes,
334	however, do not feed back on the canopy under study, as the surrounding non-treated
335	trees dominate the climate feedbacks (see Leakey 2009). Tipping points are not to be
336	expected via the direct CO <sub>2</sub> response, but rather via the much more important indirect

transpiration response (Holtum & Winter 2010; Hartmann 2011). For example, the
rainfall amount and pattern will determine when the CO<sub>2</sub>-induced water response starts to
become important (Leuzinger & Körner 2010; Hovenden *et al.* 2014). This CO<sub>2</sub>-induced
water response is likely to be the main process that affects resilience and determines
whether a forest tips into an alternate state.

342 In models, CO<sub>2</sub> effects on forest productivity are key uncertainties. Crucial 343 assumptions relate to whether photosynthesis and stomatal conductance are proportional 344 to CO<sub>2</sub> levels and how this translates to carbon sequestration, how leaf-level processes 345 are coupled to the atmosphere, and how drought affects stomatal conductance or how 346 rainfall is intercepted (Medlyn et al. 2011). For example, decreasing productivity under 347 climate change may only be observed in those models that are assuming no interaction 348 between changing temperatures and CO<sub>2</sub>-concentrations (Medlyn *et al.* 2011; Rever *et al.* 349 2014). Models that include a nitrogen cycle and interactions with CO<sub>2</sub> provide a better 350 representation of CO<sub>2</sub> effects on forest productivity (Grant 2013; Zaehle *et al.* 2014).

351

352 *Mortality* 

Mortality of individual trees does not mean that resilience of the forest is impaired or a tipping point may occur. However, improving our understanding of mortality will enhance our capacity to simulate forest resilience and tipping points. The mechanistic controls over tree mortality remain a major uncertainty in modelling forest resilience to drought-induced mortality (McDowell *et al.* 2011; Adams *et al.* 2013). Many tree species across the globe operate with small 'safety margins' for xylem embolism and cavitation (Choat *et al.* 2012) and vulnerability to drought is a key factor of tree mortality (Allen *et* 

*al.* 2010; Table 1). Mortality is particularly challenging to predict because tree death can
occur from a variety of causes, including resource limitation, environmental stress, or
large-scale natural disturbance. These causes are also strongly interdependent, as reduced
carbon or water stores impair a tree's capacity to defend against pests, or to withstand
drought. Predicting the onset of mortality thus requires an integrated framework for
quantifying critical thresholds and feedbacks in tree responses to physiological stress
(McDowell *et al.* 2011).

367 Regional-scale analyses of tree mortality from forest inventory or forest monitoring 368 plot networks offer outstanding opportunities for quantifying mortality patterns. Long-369 term, plot-based records have been used to reveal increases in mortality rates over the 370 past 50 years both in the western US and Canada (van Mantgem et al. 2009; Peng et al. 371 2011). Such observations can help inform and calibrate the mechanisms used to predict 372 mortality in process-based models, or can serve as a basis for hybrid approaches that 373 combine mechanistic and phenomenological elements to derive predictions of mortality 374 and its role in defining forest resilience or tipping points.

375

### 376 *Combination of methods*

A combination of experiments, observations and models may help to better understand scale-dependent responses (Fig. 1). Leaf-level CO<sub>2</sub>-effects are wellunderstood in the laboratory but integrating this mechanistic knowledge with experiments and observations of larger-scale CO<sub>2</sub>-effects in models in conjunction with other ecophysiological and demographic processes is necessary to paint a more realistic picture of forest responses to elevated CO<sub>2</sub> (Field *et al.* 1995).

383	Advances have also been made in experiments and observations with regard to the
384	mechanisms underpinning forest growth. While there is empirical evidence that net
385	primary production is at least partly sink- rather than source- driven (Körner 2003;
386	Muller et al. 2011; Fatichi et al. 2014), current modelling approaches still rely on
387	photosynthesis-driven (i.e., source-driven) plant growth (e.g., Sitch et al. 2003; Bonan et
388	al. 2003; Krinner et al. 2005; Sitch et al. 2008; Bonan et al. 2011). This is highly relevant
389	because photosynthesis often correlates with productivity and there is a danger of
390	obtaining the right answer for the wrong reason. Sensitivity analyses of vegetation
391	models have shown an unrealistic importance of mostly photosynthesis-related
392	parameters (Pappas et al. 2013). However, if tipping points are to be simulated
393	accurately, it is imperative that the causal mechanisms in modelling plant growth are
394	correct, which may involve the reconsideration of the core architecture of vegetation
395	models (Fatichi et al. 2014).
396	Besides combinations of experiments, observation and models, the integration of
397	different model types can help to better understand forest resilience. For example,
398	Schelhaas et al. (submitted) combined simulations of a process-based forest growth
399	model by Reyer et al. (2014) with species distribution maps from Hanewinkel et al.
400	(2013). By doing this, they could assess where tree species would need to be changed by
401	forest managers as a consequence of changing species distributions as well as changing
402	productivity under climate change to increase forest resilience.
403	

## 404 **Concluding remarks**

405 This paper shows that forests are changing over a wide range of spatial and 406 temporal scales. There seems to be more evidence for ongoing changes at smaller spatial 407 and shorter temporal scales and in particular areas such as the Mediterranean. But also 408 larger and longer spatio-temporal scales as well as other forest types face strong changes 409 (Table 1). However, in many cases it is not clear if these changes reduce resilience or a 410 tipping point is about to be reached. Our synthesis indicates that global-scale disruptions 411 of forest ecosystems are unlikely. We argue that studies covering interactions across a 412 range of spatio-temporal scales are needed to further our understanding of forest 413 resilience and tipping points. Changing CO<sub>2</sub> concentration and tree mortality are 414 processes that require an integration of results from experiments, observations and 415 models to improve the ability of process-based models to simulate forest resilience and 416 potential tipping points.

417 While acknowledging the important role of gradual changes in affecting resilience 418 of forests under global change and the lack of evidence for mechanisms leading to tipping 419 points, we stress that tipping points are important to investigate from a risk assessment 420 perspective. This means that even though a low likelihood is attributed to a specific high-421 impact event such as a tipping point, it is still important to discuss if we are willing to 422 take the risk of such events to occur. Especially because reduced resilience and potential 423 tipping points may have a profound effect on ecosystem functions and services. Such an 424 approach requires assessments of how society values ecosystem functions and services 425 and would benefit from framing forests as social-ecological systems affected by press and 426 pulse dynamics (cf. Collins et al. 2011). At the same time, a better understanding of

- 427 resilience and tipping points may actually highlight opportunities for ecosystem
- 428 restoration (e.g. Holmgren *et al.* 2013) or sustainable forest management (Puettmann *et*
- 429 *al.* 2009; Rist & Moen 2013) which may provide strong co-benefits for improving human
- 430 livelihoods.

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## 831 Figures



833 Figure 1: Temporal and spatial scales at which different research approaches 834 (experiments, observations and models) are available to assess forest resilience and 835 potential tipping points. Drought, pests, fire and other processes that structure the forests 836 such as heatwaves or also deforestation can occur across different spatio-temporal scales. 837 Precipitation, temperature and CO<sub>2</sub> changes cover all spatial scales but only long time 838 scales as we refer to them occurring within the context of climate change and not as part 839 of short-term meteorological variability. The figure is adapted from Gunderson & Holling 840 (2002).

## 841 **Tables**

842 **Table 1**. Examples of changes in forests that are possibly related to reduced resilience and tipping points in different types of forest

843 ecosystems across different temporal and spatial scales as defined in the text. Observed examples in **bold**, projected or theoretical

## 844 examples in normal font

	Local (<~10km <sup>2</sup> )	Regional (10km <sup>2</sup> -continental scale)	Global
Short-term	• <u>Mediterranean:</u>	• <u>Mediterranean</u>	<ul> <li>Catastrophic geophysical</li> </ul>
$(<\sim 10 \text{ years})$	-Drought and heat-induced tree	-Drought and heat induced patchy forest die-back and	events with wide-spread
( • 10 j • • 15)	dieback in southwest Australia	mortality across a Mediterranean-climate ecoregion in	biological consequences
	(Matusick et al. 2012; Bader et	southwest Australia (Brouwers et al. 2013; Matusick et al.	(e.g. asteroid strike,
	<i>al.</i> 2014)	2013 Brouwers <i>et al.</i> in review)	supervolcanoes, cosmic
	-Drought and fire have induced	-Gradual long-term declines in rainfall and increasing	radiation from a nearby
	transition from forest and	temperatures affecting tree health in southwest Australia	supernova (e.g. Benton &
	savannas to shrublands in south	(Brouwers <i>et al.</i> 2012)	Twitchett 2003)
	Portugal (Acacio et al. 2007,	• Tropical	
	2009; Acacio and Holmgren	-Resilience of three alternative stable states of tree covers	
	2014)	depends on precipitation in Africa, Australia and South	
	-Drought induced increase in	America (Hirota et al. 2011)	
	tree mortality in Spain	-Large parts of Amazon and Congo forest could be either	
	(Camarero et al. 2015)	forest or savannah, depending on fire regime (Staver et al.	
	• Different forest types	2011)	
	-Drought-induced forest die-	• Boreal	
	back around the world (Allen et	-Bark beetle-induced tree mortality in western North	
	<i>al.</i> <b>2010</b> )	America (Kurz et al. 2008; Raffa et al. 2008)	
	-Low safety margin of many	-Drought-induced dieback of aspen ecotone in midcontinental	
	species worldwide to drought	North America (Michaelian <i>et al.</i> 2011)	
	stress (Choat et al. 2012)	• Different forest types	
		-Drought-induced forest die-back around the world (Allen <i>et</i>	
		al. 2010)	
Long-term	• <u>Tropical</u>	• <u>Mediterranean</u>	• State-shift of the global
<u> </u>	-Drought decreases biomass in a	-Increasing drought stress in Southern Europe (van Oijen et al.	biosphere (Barnosky et al.

(10-1,000 years)	Monodominant Amazon forest	2013)	2012; Hughes et al. 2013b;
	(Nascimento et al. 2014)	• <u>Tropical</u>	Brook et al. 2013; Lenton
	<ul> <li><u>Different forest types</u></li> </ul>	-Climate-change induced dieback of the Amazon (Cox et al.	& Williams 2013)
	-Drought-induced forest die-	2004; Malhi et al. 2009 but see also Cox et al. 2013; Huntingford	
	back around the world (Allen et	<i>et al.</i> 2013)	
	al. 2010)	• <u>Boreal</u>	
	-Drought-induced mortality	-Boreal biome transitions (Scheffer et al. 2012a)	
	(Steinkamp & Hickler 2015)	<u>Different forest types</u>	
		-Drought-induced forest die-back around the world (Allen et	
		al. 2010)	

Table 2. Examples of changing environmental conditions that are relevant for assessing forest resilience and tipping points and 845

Changing conditions	Observations	Experiments	Models
CO <sub>2</sub>	<ul> <li>Long time series available only close to natural CO<sub>2</sub>-sources which usually represent very specialized environments (e.g., Haettenschwiler <i>et al.</i> 1997)</li> <li>Europe fluxtower records are generally not yet long enough to provide long-term average references (Ciais <i>et al.</i> 2005)</li> </ul>	<ul> <li>Only small spatial and temporal coverage of experiments, costly, few in mature stands (Norby <i>et al.</i> 2010; Bader <i>et al.</i> 2013), no tropical forests but new FACE experiment starting in Amazon</li> <li>CO<sub>2</sub> increased stepwise vs. gradual change in reality</li> </ul>	<ul> <li>CO<sub>2</sub> assumptions in models: sink- vs. source-driven models (Cox <i>et al.</i> 2013; Fatichi <i>et al.</i> 2014)</li> <li>Assessment of CO<sub>2</sub> response requires N availability to be included in model (Grant 2013)</li> </ul>
Temperature	• Time series are easy to observe, but the problem is mostly the disentanglement by concurrent changes such as precipitation, CO <sub>2</sub> and nitrogen	<ul> <li>Lack of whole tree warming experiments</li> <li>Mostly small spatial and temporal coverage, typically greenhouse experiments or open-top chambers</li> </ul>	<ul> <li>Temperature responses in models are poorly represented</li> <li>Acclimation of model parameters is poorly represented</li> </ul>
Drought	<ul> <li>Drought and heat often act together, and their interactions need to be considered (Allan <i>et al.</i> 2010; Matusick <i>et al.</i> 2013; Brouwers <i>et al.</i> 2013)</li> <li>Drought impact responses of forest vegetation are highly variable depending on the positioning within the landscape (Brouwers <i>et al.</i> 2013, Brouwers <i>et al.</i> 2013, Brouwers <i>et al.</i> in</li> </ul>	<ul> <li>Expensive, challenging and laborious rainfall exclusion experiments (REE) (Nepstad <i>et al.</i> 2007; Phillips <i>et al.</i> 2010; Misson <i>et al.</i> 2011)</li> <li>Lateral flow of water not completely controlled in REEs (Fisher <i>et al.</i> 2006)</li> <li>REEs limited because they alter only rainfall, and not temperature, VPD or radiation (Fisher <i>et al.</i> 2007)</li> </ul>	• Mechanisms of drought-induced mortality such as embolism and cavitation not well covered in many models (Choat et al. 2012; Steinkamp & Hickler 2015)
	review)	• Few actual tropical drought experiments exist (Phillips <i>et al.</i> 2009; da Costa <i>et al.</i> 2010)	

846 challenges in understanding them in different research approaches