

# Formal subdivision of the Holocene Series/Epoch: a summary

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## **ABSTRACT**

The Holocene Series/Epoch is the most recent series/epoch in the geological timescale, spanning the interval from 11,700 yr to the present day. Together with the subadjacent Pleistocene, it comprises the Quaternary System/Period. The Holocene record contains diverse geomorphological, biotic, climatological and archaeological evidence, within sequences that are often continuous and extremely well-preserved at decadal, annual and even seasonal resolution. As a consequence, the Holocene is perhaps the most intensively-studied series/epoch within the entire Geological Time Scale. Yet until recently little attention had been paid to a formal subdivision of the Holocene. Here we describe an initiative by the Subcommission on Quaternary Stratigraphy (SQS) of the International Commission on Stratigraphy (ICS) to develop a formal stratigraphical subdivision of the Holocene, with three new stages/ages, two underpinned by Global Boundary Stratotype Sections and Points (GSSPs) in an ice core, and a third in a speleothem. These stages/ages are defined along with their equivalent subseries/subepochs. The new stages/ages are the Greenlandian with its GSSP in the Greenland NGRIP2 ice core and dated at 11,700 yr b2k (before 2000 CE); the Northgrippian with its GSSP in the Greenland NGRIP1 ice core and dated to 8236 yr b2k; and the Meghalayan, with its GSSP in a speleothem from Mawmluh Cave, northeastern India, with a date of 4250 yr b2k. This subdivision was formally ratified by the Executive Committee of the International Union of Geological Sciences (IUGS) on 14<sup>th</sup> June 2018.

Keywords: Greenlandian, Northgrippian, Meghalayan stages/ages, Lower/Early Holocene, Middle Holocene, Upper/Late Holocene subseries/subepochs

## **INTRODUCTION**

The division of geological time is fundamental for understanding our planet's history in all its ramifications. The question of the relative sequence of events and demonstration of their equivalence beyond single localities has long been central to geological investigation. Although we may apply local criteria for the division and correlation of geological events, the nature of rock and sediment sequences demands that geologists speak a common language if we are to communicate effectively across borders, continents and oceans. Early geologists established the division of time by recognising major gaps in the sedimentary record, but it became apparent that for more accurate and precise correlations, boundaries placed at erosional or non-depositional surfaces might not be the most effective or workable means to define chronological units, as time is not recorded during the break. Today the time-units represented by geological successions are referred to as geochronological units, and their material equivalents are chronostratigraphical units, a duality in classification that dates back to the Second International Geological Congress in Bologna in 1881. By convention these

units are based on reference sections, ideally with their bases unequivocally defined and reliably and precisely correlated around the world. However, as knowledge has advanced it has become increasingly apparent that significant differences exist between the successions of events on different continents and, indeed, between different regions within continents and the ocean floors. It is obvious therefore that a coherent and globally applicable stratigraphical framework is required to establish the terminology needed for unambiguous communication.

This paper shows how such a framework has recently been developed for the Holocene, the shortest series/epoch-scale division in the International Geological Time Scale (GTS) beginning at around 11,700 years b2k (before 2000 CE) and extending to the present day (Walker et al., 2008; 2009). It summarises a proposal by a Working Group of the International Subcommittee on Quaternary Stratigraphy (ISQS) for a formal chronostratigraphic subdivision of the Holocene that was approved by the International Commission on Stratigraphy (ICS) and ratified by the Executive Committee of the International Union of Geological Sciences (IUGS) on 14th June 2018. Full details of the new scheme can be found in Walker et al. (2018).

## **INTERNATIONAL STRATIGRAPHICAL PRACTICE**

The standard approach to stratigraphy was championed after the Second World War by the American Commission on Stratigraphic Nomenclature in recognition of the desirability of uniform usage in stratigraphic classification and terminology throughout the continent of North America (American Commission on Stratigraphic Nomenclature, 1961, p. 624). This initiative was driven mainly by petroleum geologists for whom a uniform stratigraphical language or code was an essential tool. The code provided the first systematic approach to stratigraphical division and consistent definition of terminology for purposes of correlation. Subsequently the code formed the basis of the International Stratigraphic Guide published, by the International Subcommittee on Stratigraphic Classification (ISSC), originally under the editorship of Hedberg (1976) and then in revised form under the editorship of Salvador (1994). The ISSC is a constituent body of the ICS, which itself operates within the IUGS.

Founded in 1965, the ICS was established to develop a multidisciplinary global standard geological time scale that would facilitate the correlation of rock/sediment strata (Cowie et al., 1986). Correlation was based on both palaeontological and non-biological

evidence that was assembled from regions across the world. The time scale was underpinned by Global Boundary Stratotype Sections and Points (GSSPs) which were used to calibrate the stratigraphical record. Each boundary was to be indicated by a marker or 'Golden Spike' in the rock/sediment section. The first GSSP was established in 1972, and today the global geological timescale with its hierarchy of systems, series and stages is largely defined in this fashion ([www.stratigraphy.org](http://www.stratigraphy.org)) Of the 102 boundaries recognised for the Phanerozoic, 70 have now been defined by GSSPs, with work continuing to define the remaining boundaries. Based on this principle the divisions will ultimately form a continuous record of geological time to the present day. This is a remarkable achievement by the subcommissions and working groups within the ICS, and creates a global language for international stratigraphy and correlation

## **THE HOLOCENE SERIES/EPOCH**

Throughout the Quaternary System/Period there has been, understandably, an overwhelming reliance on climatic evidence to identify stratigraphical divisions. The climatic sequence of glacial and interglacial episodes, established in the late nineteenth and early twentieth centuries, still forms the basis for subdividing the record. This is because the Quaternary is so short in duration (at least relative to other systems/periods) and conventional datums, such as biostratigraphic range tops /bases and palaeomagnetic reversals, are too infrequent to provide a sufficiently refined subdivision.

This problem is even more acute in the Holocene Series/Epoch, the most recent interglacial (temperate climate) episode in the Quaternary where no global step-changes that could potentially be used as a basis for stratigraphical subdivision are manifest in the proxy climate data. On the other hand, radiometric dating (radiocarbon and uranium-series), and incremental dating methods, such as dendrochronology, varve chronology and layer-counted ice core and speleothem chronologies (Walker, 2005), mean that where distinctive climatic events can be detected, these can potentially form the basis for ultra-high-resolution (decadal or even annual) stratigraphical divisions. As the Holocene is probably the most intensively studied interval within the entire geological record and embodies a remarkable range of geomorphological, climatic, isotopic, biological and archaeological evidence, a formal subdivision of that record should be of considerable value in aiding communication between the wide range of scientists studying the present interglacial.

## **SUBDIVISION OF THE HOLOCENE**

Prior to the establishment of the Working Group on the Holocene in 2011, only limited attention had been paid to a formal stratigraphical subdivision of this time interval. The first attempt was made by Axel Blytt and Rutger Sernander who, in the early years of the twentieth century, proposed a stratigraphical scheme using plant macrofossil records from peat bogs in Scandinavia (Sernander, 1908). Their terminology, based on interpreted climatic changes comprised, in chronological order, the pre-Boreal, Boreal, Atlantic, sub-Boreal and sub-Atlantic episodes. These terms were applied to European pollen-based bizones by Lennart von Post and others (Godwin, 1975), and were subsequently incorporated into an influential paper by Mangerud et al. (1974) on the Quaternary stratigraphy of Norden (Scandinavia). This proposed that the Flandrian (regional) Stage (equivalent to the Holocene Series) should be divided into three substages with boundaries defined by the North European chronozones based on the Blytt–Sernander pollen zones and dated by radiocarbon: Early Flandrian (Preboreal and Boreal: 10,000–8,000  $^{14}\text{C}$  yr BP); Middle Flandrian (Atlantic and Subboreal: 8000–2500  $^{14}\text{C}$  yr BP); and Late Flandrian (Sub-Atlantic: post 2500  $^{14}\text{C}$  yr BP)<sup>1</sup>. But time-transgression in vegetational response to climate change, ambiguities in the use of the Blytt and Sernander classification, and problems associated with radiocarbon dating suggested that such a chronostratigraphical subdivision of the Holocene would not be applicable at anything other than the local or perhaps regional scale (Walker, 1995; Björck et al., 1998; Wanner et al., 2008; and see above).

There have since been a number of important advances in Quaternary geoscience that have encouraged a re-evaluation of the possibility of a formal chronostratigraphical subdivision. These include increasing numbers of Holocene successions resolved at annual to decadal scales, such as ice cores from Greenland, Antarctica and elsewhere; high-resolution stratigraphical records from peat deposits and lacustrine sediments; and annually-resolved tree-ring series and speleothems. These provide often highly-detailed palaeoenvironmental archives of regional, hemispherical or even global significance. In addition, refinements in numerical dating techniques offer an increasingly secure geochronological foundation, while temporal stratigraphical markers such as tephra isochrons, enhance accuracy in regional and, in some instances, extra-regional correlation. These various lines of evidence offer a stronger foundation for chronostratigraphical subdivision and correlation than was previously possible. The case for subdividing the Holocene Series/Epoch has been further strengthened by the

formal definition, following conventional chronostratigraphical procedures (Hedberg, 1976; Salvador, 1994), of the Pleistocene–Holocene boundary, with the ratified GSSP in the Greenland NGRIP2 ice core (Walker et al., 2008, 2009; see below).

One difficulty in seeking a basis for a formal subdivision of Holocene time, however, is that there is little evidence for clearly-defined and long-lasting climatic changes during the course of the present interglacial. Therefore any subdivision must be based on distinctive and globally significant events that can be identified in proxy climate records. Two such events, one occurring around 8.2 ka BP and a second at ~4.2 ka BP are recorded in many Holocene proxy climatic records (see below), and it was suggested by Walker et al. (2012) that the global signatures of these two events can form the basis for a tripartite subdivision of the Holocene, along the lines initially suggested by Mangerud et al. (1974; see above). Indeed, such a subdivision was already being widely applied, albeit in an informal sense. Assessment of the literature shows that for many years the terms ‘Early’, ‘Middle’ (‘Mid-’) and ‘Late’ Holocene have been routinely employed in a wide range of depositional contexts and environmental settings, although often in an inconsistent manner (Walker et al, 2012). Accordingly, therefore, it seemed appropriate to adopt what is current custom and practice, and to both standardise and formalize the subdivisions that are already being used by underpinning them with clearly defined marker horizons based on GSSPs.

## **THE SUBDIVISIONAL GSSPs**

Two new GSSPs have been approved to underpin the stratigraphical subdivision of the Holocene; one utilising an ice-core from Greenland and the other a cave speleothem from India. While these might seem unconventional locations for GSSPs, they were selected for strong scientific reasons. The precedent for designating an ice core as a boundary stratotype was set some ten years ago with the formal ratification of the basal Holocene GSSP in the Greenland NorthGRIP2 (NGRIP2) ice core (Walker et al., 2008; 2009). The use of a

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1. Note that ‘BP’ here, as well as below, means ‘before 1950’. This age baseline differs by 50 yrs from ice-core age estimates which sometimes are reported relative to a baseline year of 2000 CE (b2k).

speleothem however, is new. Speleothems are stratified successions that can be analysed at remarkably high resolution (sub-decadal to annual); in addition, they contain an oxygen isotope record that can be supported by an independent, high-precision chronology based on U-series dating. The stable isotope profile in Holocene speleothem calcite is known to be a highly sensitive climate proxy (e.g. Fleitmann et al., 2007; Boch et al., 2009; Cheng et al., 2009; Wu et al., 2012), and therefore changes in the isotopic signal provide a detailed and chronologically accurate record of Holocene climate change. As such, speleothems provide a remarkable archive of high-resolution climate data, and offer a basis for fixing events within very short time intervals.

## **THE SUBDIVISIONS**

As noted above, an informal tripartite subdivision of the Holocene has long been employed and describes three formally-defined chronostratigraphic and geochronological units: the Lower/Early, Middle, and Upper/Late Holocene subseries/subepochs. It should be noted, however, that while the designation of subseries/subepochs has always been acceptable under the International Stratigraphic Guide (Hedberg, 1976: table 2; Salvador 1994; table 3), divisions at this rank in the stratigraphic hierarchy had not been used previously in the GTS (Head et al., 2017). Hence, it was necessary to make an exception for the Holocene subdivisional scheme. Moreover, because the fundamental units of the GTS are stages/ages, it was necessary also to designate for each subseries/subepoch a corresponding stage/age. Accordingly, the terms Greenlandian, Northgrippian, and Meghalayan Stage/Age were proposed to represent the three divisions that correspond, respectively, to Early/Lower, Middle and Upper/Late subseries/subepochs. The terminology follows the practice of the ICS in naming stages after the geographical localities where the GSSPs are located; hence Greenland, the Greenland NorthGRIP (NGRIP) ice core and the northeast Indian state of Meghalaya where the cave that contains the speleothem is located. Since the base of the Holocene Series/Epoch was ratified in 2008 (Walker et al. 2008; 2009) albeit, somewhat unconventionally, without an accompanying Stage/Age, the Greenlandian Stage is now ratified as the lowermost stage of the Holocene and anchored by the previously-defined base-Holocene GSSP. Above the Greenlandian are the Northgrippian and Meghalayan stages/ages and their accompanying Middle and Upper/Late Holocene subseries/subepochs, both of which are underpinned by new GSSPs.

### ***Greenlandian Stage/Age; Lower/Early Holocene Subseries/Subepoch***

The lowest stage of the Holocene Series/Epoch is termed the Greenlandian Stage/Age, defined with its corresponding Lower/Early Holocene Subseries/Subepoch by the GSSP of the Holocene Series/Epoch in the NorthGRIP2 (NGRIP2) Greenland ice core (75°108'N, 42°328'W; Walker et al., 2008). The GSSP is located at 1492.45 m in the ice core, where it is marked by a shift to 'heavier'  $\delta^{18}\text{O}$  values between the preceding Greenland Stadial 1 (GS-1) and the Early Holocene; by a reduction in dust concentrations from GS-1 to modern levels and by a significant reduction in Na (sea-salt) values; and by an increase in annual ice-layer thickness (Johnsen et al., 2001; Steffensen et al., 2008). The boundary is most clearly marked, however, by an abrupt decline in deuterium (d) excess values, indicating a northward shift in ice-sheet moisture source from the middle to the northern North Atlantic as the oceanic polar front rapidly retreated. This shift occurred within a matter of a few years (Steffensen et al., 2008) and, ironically, led to a cooler (and more northerly) ice-sheet moisture source. These various data sources reflect a major change in atmospheric circulation regime and an associated temperature rise of the order of  $10 \pm 4^\circ\text{C}$ , at the onset of the Holocene (Grachev & Severinghaus, 2005; Buizert et al. 2014). The boundary is dated on the Greenland ice-core timescale (GICC05; Rasmussen et al., 2005) to 11,703 calendar yrs b2k (before 2000 CE); this corresponds to 11,653 yrs BP using the datum of the radiocarbon timescale. The maximum counting error (MCE) on the ice-core age is 99 yr, which is interpreted as broadly equating to a  $2\sigma$  uncertainty (Walker et al., 2009).

### ***Northgrippian Stage/Age; Middle Holocene Subseries/Subepoch***

The second stage/age of the Holocene is termed the Northgrippian Stage/Age and is defined with its corresponding Middle Holocene Subseries/Subepoch in the NorthGRIP1 (NGRIP1) Greenland ice core (75°108'N, 42°328'W). The GSSP is located at a depth of 1228.67m (Figure 1 upper) in an interval that shows a clear signal of climate cooling following a period of generally rising temperatures during the Early Holocene. This cooling occurs at around 8.2 ka in the NGRIP1 core and corresponds to the '8.2 ka climatic event', a short-lived near global episode that is reflected in a wide range of proxy climate records (Alley & Àgùstsdottir, 2005; Rohling & Pälike, 2005; Morrill et al., 2013). It is generally considered to result from a cessation in North Atlantic Deep Water (NADW) formation and reduction in the



associated northward heat transport following catastrophic meltwater release into the North Atlantic from glacial lakes Agassiz and Ojibway during melting of the Laurentide Ice Sheet (Barber et al., 1999; Kleiven et al., 2008; Hoffman et al., 2012), perhaps accompanied by collapse and accelerated melting of ice domes over Hudson Bay (Matero et al., 2017).

In the Greenland NGRIP1 ice core, the 8.2 ka event is marked by a significant shift in the oxygen isotope record to more negative  $\delta^{18}\text{O}$  and  $\delta\text{D}$  values; by a decline in ice-core annual layer thickness (Rasmussen et al., 2007) and deuterium excess (Masson-Delmotte et al., 2005); by a substantial sudden and short-lived minimum in atmospheric methane (a global event); and by a subsequent increase in atmospheric  $\text{CO}_2$ . The water isotope diffusion-derived temperature record indicates a cooling of  $\sim 5^\circ\text{C}$  (Gkinis et al., 2014). The GSSP is placed in the middle of a double peak in electrical conductivity measurements (Figure 1 lower). The layer also includes fluoride and is likely derived from a volcanic eruption in Iceland. Hence, while the climate signal determined from the oxygen isotope record places the GSSP within the coldest part of the 8.2 ka event in the NGRIP1 ice-core record, the independent volcanic signal enables the GSSP to be precisely located in the NGRIP1 ice core and correlated to other Greenland ice cores

The age of the GSSP is derived from the GICC05 timescale which is based on a range of physical and chemical parameters in three Greenland ice cores: DYE-3, GRIP and NGRIP (Rasmussen et al., 2006; Vinther et al., 2006). However, low accumulation rates at the NGRIP1 drill site mean that annual layers cannot easily be identified in that core. The chronology for NGRIP1 therefore derives from the DYE-3 and GRIP records where accumulation rates are higher, the cores being linked by  $\delta\text{D}$ ,  $\delta^{18}\text{O}$  fluoride and electrical conductivity measurements (ECM). In the DYE-3 ice core, the annual layer situated in the middle of the ECM double peak is dated to 8236 yr b2k with a maximum counting error of 47 yr (Vinther et al., 2006). This is the best estimate of age of the GSSP for the Northgrippian Stage/Age and Middle Holocene Subseries/Subepoch, and corresponds to 8186 yr BP using the datum of the calibrated radiocarbon time scale.

### ***Meghalayan Stage/Age: Upper/Late Holocene Subseries/Subepoch***

The uppermost subdivision of the Holocene Series/Epoch is termed the Meghalayan Stage/Age and is defined together with the corresponding Upper/Late Holocene Subseries/Subepoch by a GSSP in a speleothem (KM-A) from Mawmluh Cave in the state of Meghalaya in northeast India (25°15'N; 91°4'E). The GSSP is located at a horizon within the speleothem which shows evidence for a marked reduction in rainfall at ~4200 yrs BP, and reflects what has been referred to as the '4.2 ka climatic event' (Rousseau et al., 2018). Occurring at, or around, 4 ka BP, this climatic shift appears to involve significant reorganisations of ocean and atmosphere circulation patterns, and has also been termed the 'Holocene Turnover' (Paasche et al., 2004; Paasche & Bakke, 2009). Although the forcing factors are less clear than those for the climatic event at 8.2 ka, the 4.2 ka event nevertheless appears also to be near-global in nature, occurring in proxy records from North America and northern Europe, through the Middle East to China; and from Africa, Andean-Patagonian South America, and Antarctica (Mayewski et al., 2004; Staubwasser & Weiss, 2006; Schimpf et al., 2011; Peck et al., 2015). In higher latitudes, the event is reflected in climatic cooling and glacier readvances (Menounos et al., 2008; Gkinis et al., 2014; Balascio et al., 2015; Geirsdóttir et al., 2018), while in the mid- and low latitudes it is frequently marked by aridification (Booth et al., 2005; Wanner et al., 2015; Weiss, 2016). The 4.2 ka event caused disruption of the westerlies, the Indian Summer Monsoon, and the East Asian Summer Monsoon. It led to a ca. 250-year megadrought, that was broken perhaps only briefly, and forced synchronous societal collapse, habitat-tracking, and eventual resettlement and reorganization across Spain (Lopez-Saez 2018), Greece (Davis 2013), Egypt (Hassan et al. 2017), Palestine (Harrison 2012; Weiss 2017a), Mesopotamia (Weiss 2017b), Indus (Petrie et al. 2017) and China (Li et al. 2018).

As the 4.2 ka event is most strongly recorded in proxy climate records from mid-and low latitudes, it is appropriate that the GSSP should be located within those latitudes, and Mawmluh Cave in northeast India offers an ideal site. Formed in an area of limestone bedrock, this cave is one of the longest in the subcontinent, and high relative humidity (>90%) and minimal temperature fluctuations (18.0–18.5°C), relatively constant drip rate, and the deep location of the sampling site provide optimal conditions for calcite to form in

isotopic equilibrium with percolating precipitation. Hence variations in the  $\delta^{18}\text{O}$  of speleothem calcite closely resemble changes in the precipitation-weighted  $\delta^{18}\text{O}$ . The  $\delta^{18}\text{O}$  record from stalagmite sample KM-A (Figure 2) extends from c. 3500 to >12,000 yr BP at a resolution of  $\sim 5$  years per sample. The most significant isotopic excursion in the entire record occurs around the time of the 4.2 ka event, with an overall enrichment of  $\sim 1.5\%$  in  $\delta^{18}\text{O}$ , approximately equivalent to a 20–30% decrease in rainfall and marking a significant change in the strength of the Indian monsoon (Berkelhammer et al., 2012).

The chronology for the stable isotope record from speleothem KM-A is based on an age model that employed a Monte Carlo fitting procedure through 12 U-Th dates (Scholz & Hoffman, 2011). The analytical uncertainty on the two U-Th dates closest to the 4.2 ka event (3654 and 4112 yr BP) are 20 and 30 years respectively, with a third date at 5084 yr BP having an uncertainty of  $\pm 32$  yr (Figure 2). The KM-A record shows linear growth rates during this period, providing further confidence in the age model and hence in the timing (onset and duration) of the 4.2 ka event (Berkelhammer et al., 2012). The changes in the stable isotope record during that interval comprise a two-step sequence, with an initial enrichment at  $\sim 4300$  yrs BP and a more pronounced shift (within less than a decade) to more positive values at 4100 yrs BP. The abrupt increase in  $\delta^{18}\text{O}$  values is the primary boundary marker for the GSSP, and hence a date of 4200 yr BP, which is effectively the mid-point between these two modelled ages, can be assigned. The U-Th ages are expressed relative to a baseline date of 1950 CE and are therefore directly comparable with the calibrated radiocarbon time scale. However, in order to maintain consistency with the earlier Holocene GSSPs (the Greenlandian and Northgrippian) which are dated using the GICC05 ice-core chronology, the age of the Mawmluh speleothem GSSP is defined as 4250 yr b2k (before 2000 CE: see above).

## **SUMMARY**

This paper summarises three inter-related definitions for the formal subdivision of the Holocene Series/Epoch into the Greenlandian, Northgrippian and Meghalayan stages/ages and their corresponding Lower/Early, Middle, Upper/Late subseries/subepochs, each supported by a GSSP. The GSSP for the lowermost stage, the Greenlandian, is the previously ratified boundary for the Holocene in the NGRIP2 ice core. The GSSP for the Northgrippian is

located in the ice core from NGRIP1, Greenland, while the GSSP for the Meghalayan is placed in a speleothem from Mawmluh Cave, Meghalaya, northeast India. All three stages and their corresponding subseries/subepochs have been formally approved by the ICS and ratified by the IUGS.

While most current GSSPs are based upon biostratigraphical evidence as their primary markers, it has been recommended that future GSSPs should have physico-chemical markers as an integral part of their guiding criteria (Smith et al., 2014). Indeed, as Head & Gibbard (2015) noted, the International Stratigraphic Guide (Hedberg, 1976, p. 82; Salvador, 1994, p. 88) acknowledges that certain methods of correlation, such as climatic, palaeomagnetic and isotopic, have greater relevance for Quaternary chronostratigraphy. The definitions here follow that recommendation, as all three of the GSSPs are defined on the basis of significant physical and chemical markers. They reflect abrupt climatic events at the onset of the Holocene (~11.7 ka BP), at ~8.2 ka BP and ~4.2 ka BP, all of which are global or near-global in their expression. Stable isotope records in particular, from both Holocene ice-core and speleothem sequences, are strikingly sensitive climatic proxies, and can be dated with a very high degree of accuracy and precision. Indeed, the designated GSSPs are the best-resolved, both stratigraphically and temporally, within the entire geological time scale. They closely accord with the criteria for boundary stratotypes outlined in Remane et al. (1995), and provide stable points of reference for Holocene stages/ages and subseries/subepochs into the future.

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## Figures

**Figure 1.** *Upper:* water stable isotope ratios ( $\delta^{18}\text{O}$ ) at 55-cm resolution from the GRIP and NGRIP1 ice cores around the 8.2 ka event. The black horizontal double arrow indicates the duration of the event from ~8300 yr b2k (1234.78 m) to ~8140 yr b2k (1219.47 m). *Lower:* during the period of low  $\delta^{18}\text{O}$  values (the section marked by the grey bar in the upper panel and expanded in the lower panel), a distinct acidity double peak is reflected in electrical conductivity measurements (ECM). This layer, at 1228.67 m depth in the NGRIP1 core and 1334.04 m in the GRIP core (black dashed line), is characterized by a high fluoride content and can most likely be attributed to an Icelandic volcano. It is dated on the GICC05 timescale to 8236 yr b2k (8186 cal. yr BP), and is the primary marker for the Early–Middle Holocene boundary (after Walker et al., 2012, 2018).

**Figure 2.** The Mawmluh Cave  $\delta^{18}\text{O}$  record for speleothem KM-A after Berkelhammer et al., (2012). The continuous black line through the isotope trace is a low pass filter removing any variability with a frequency higher than 10 years. Red circles mark the U-Th dates obtained, which are shown with their  $2\sigma$  analytical uncertainty in black boxes. Age uncertainty (95% confidence interval) was assessed using a Monte Carlo fitting procedure through the U-Th dates, and is also shown by variations in colour along the trace. The envelope of the event (onset and termination) is shown by the arrowed blue lines, and the beginning of the most intensive phase of weakened monsoon is shown by a third arrowed blue line: their dates are given with uncertainty that is also assessed using the Monte Carlo fitting procedure. See Berkelhammer et al. (2012) for details on the age model calculations. The position of the GSSP, with a modelled age of 4200 yr BP (4250 b2k) is indicated by the red arrow. Note that the 8.2 ka event also registers as a significant excursion in the stable isotope record (after Walker et al., 2018).



