


Research Article

Gamma Ray and Radon Anomalies in Northern Taiwan as a Possible Preearthquake Indicator around the Plate Boundary

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Taiwan is tectonically situated in an oblique collision zone between the Philippine Sea Plate (PSP) and the Eurasian Plate (EP). Continuous observations of gamma rays at the Yangmingshan (YMSG) station and soil radon at the Tapingti (TPT) station were recorded in the volcanic area and around a major fault zone, respectively, in Taiwan for seismic studies. A number of anomalous high gamma ray counts and radon concentrations at certain times were found. It is noted that significant increases of soil radon concentrations were observed and followed by the increase in gamma rays a few days to a few weeks before earthquakes that occurred in northeastern Taiwan. Earthquakes such as these are usually related to the subduction of the PSP beneath the EP to the north along the subduction zone in northern Taiwan (e.g., $M_L = 6.4$, April 20, 2015). It is suggested that the preseismic activity may be associated with slow geodynamic processes at the subduction interface, leading to the PSP movement triggering radon enhancements at the TPT station. Furthermore, the further movement of the PSP might be blocked by the EP, with the accumulated elastic stress resulting in the increase of gamma rays due to the increase in porosity and fractures below the YMSG station. The continuous monitoring of the multiple parameters can improve the understanding of the relationship between the observed radon and gamma ray variations and the regional crustal stress/strain in north and northeastern Taiwan.

1. Introduction

The island of Taiwan is the result of an oblique collision zone, with a convergence rate of about 8 cm/yr in the direction of 306° [1], between the Philippine Sea Plate (PSP) and the Eurasian Plate (EP) (Figure 1). In the south, the EP underthrusts the PSP to the east along the Manila trench, while to the north, the PSP subducts beneath the EP along the Ryukyu Arc and induces spreading of the Okinawa Trough. The interaction of these two plates generates many earthquakes. The Plio-Pleistocene volcanism of the Tatun Volcano Group

(TVG) is situated both at the northern tip of Taiwan and the western tip of the Ryukyu Arc, which might be associated with the opening of the Okinawa Trough and the postcollisional extension [2–4]. According to previous radiometric dating investigations, two major eruptions of the TVG occurred at 2.5–2.8 Ma and 0.8–0.2 Ma [5–10]. Juang [11] considered the TVG extinct because the last major volcanism occurred from 0.8 to 0.2 BP. Also, Chen and Shen [12] reviewed some records of historical eruptions in northern Taiwan and suggested that volcanism of the TVG has ended.

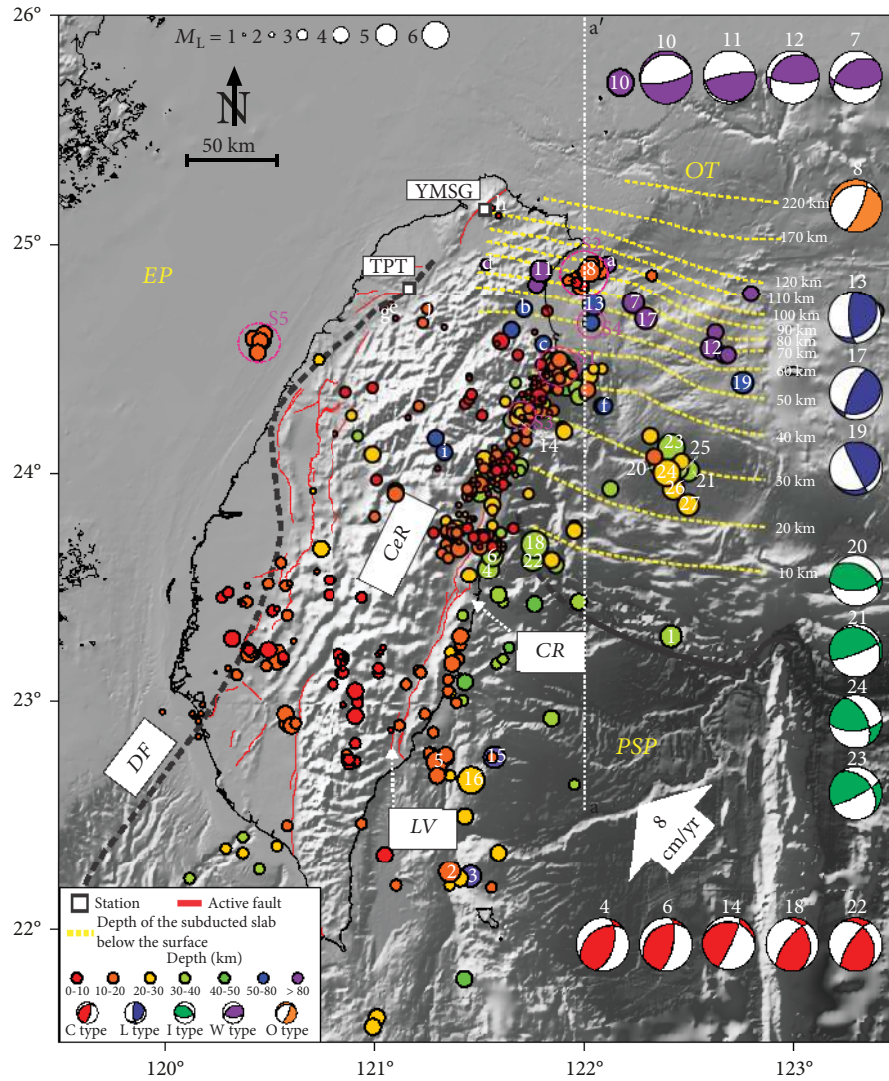


FIGURE 1: The simplified sketch of the regional map and principal tectonic structures around Taiwan shows that the Philippine Sea Plate (PSP) is moving northwest at 8 cm/yr towards the Eurasia Plate (EP). The circle symbols represent the earthquake epicenter with the focal depths corresponding to different colors from July 1, 2014, to June 1, 2015. The open square represents the location of the YMSG gamma ray station and the TPT radon station. The red lines indicate the active fault proposed by the Central Geological Survey of Taiwan [66]. Seismic events with magnitudes larger than five and some specific earthquakes are labeled and listed in Table 2. The yellow dashed lines indicate the isodepth contours between 10 and 220 km. OT = Okinawa Trough; CeR = central range; CR = costal range; LV = Longitudinal Valley; DF = deformation front.

However, Yang et al. [13, 14] and Ohba et al. [15] suggested that a magma chamber might still exist beneath the TVG area based on a high $^3\text{He}/^4\text{He}$ ratio from fumarolic gas and the high $\text{CO}_2/\text{H}_2\text{O}$ ratio of magmatic components. Konstantinou et al. [16] and Wen et al. [17] proposed that magma chambers might exist at depths of 7 to 20 km as indicated by the seismicity characteristics and attenuation structure. Lin [18] proposed the presence of a deep magma reservoir in the lower crust using S-wave shadow and P-wave delay. In recent years, there have been an increasing number of studies in the TVG area focusing on geochemical investigations and monitoring of the gas compositions, soil flux, and hot spring water [19–27] and seismological and geological research, such as seismicity, seismic variations, and ground deformation [16, 28–33], which has provided

more evidences that the TVG could still be active. As the TVG is located close to the Taipei metropolis and two nuclear power plants, long-term monitoring is necessary for providing possible early warnings of volcano-related hazards and major earthquakes.

Radon (^{222}Rn) is the major radioactive gas in this volcanic area. To investigate volcanic activity in the volcanic areas, the monitoring of radon level variations has been considered as a useful tool [20, 27, 34–38]. This research provides another way to record gamma rays related to radon gas emission from underground towards the surface. The anomalous gamma rays may be associated with the radioactive material released during crustal fracturing due to plate motions and geodynamic processes [39]. The objective of this paper is to assess the feasibility of using the gamma ray method to

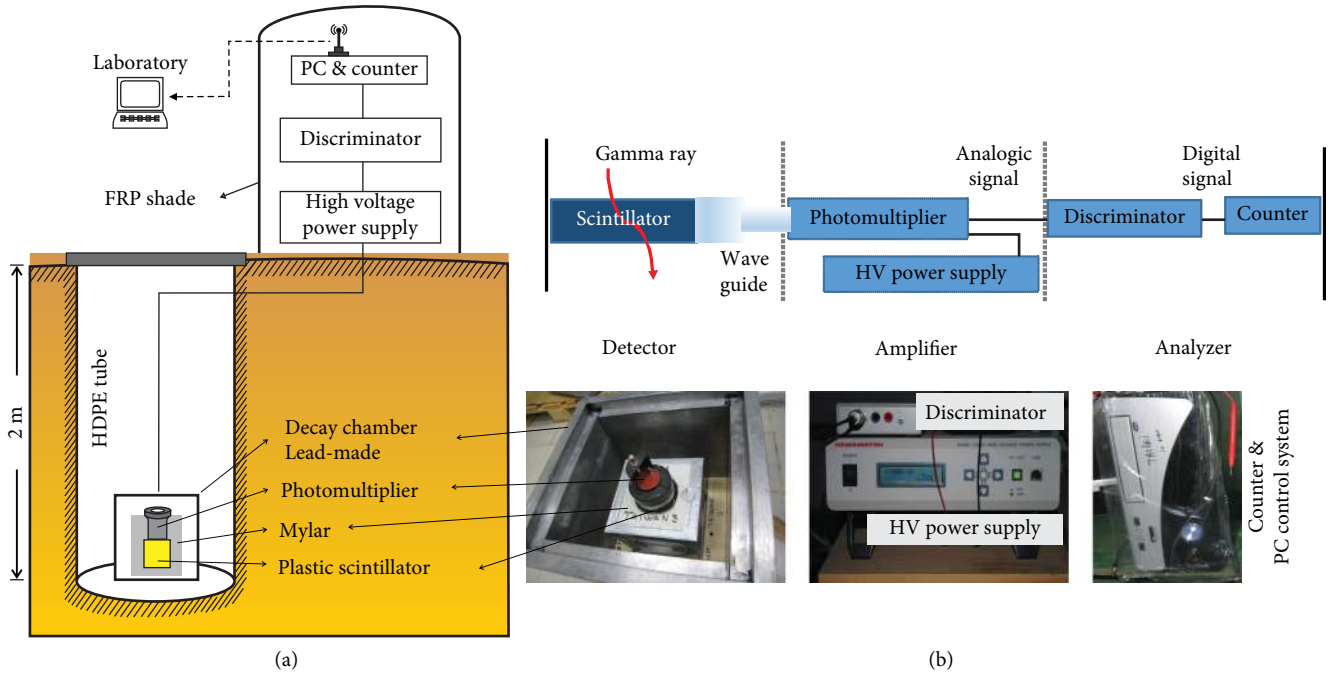


FIGURE 2: (a) Schematic of the gamma ray sensor installation in the YMSG monitoring station. (b) The block diagram and pictures of the gamma sensor assembly.

monitor seismic and volcanic activities in northern Taiwan. The location of the gamma ray station in the Yangmingshan (YMSG) area near the Taiwan Volcano Observatory at Tatun (TVO) was chosen for the first gamma ray observations in a volcanic region of Taiwan. In this study, gamma rays were continually monitored from July 1, 2014, to June 1, 2015. The potential for using a gamma ray anomaly as the possible precursor of earthquake and/or volcanic activity in northern Taiwan was further evaluated on this basis. In order to evaluate the accuracy of possible anomalous gamma ray changes, the radon variations at the Tapingti (TPT) station were chosen for comparison, being located about 70 km southwest of the YMSG gamma ray station.

2. Methodology

Radon is a naturally radioactive noble gas, widely distributed in various places throughout the Earth's crust. Radon is generated in the uranium decay series. The emission of energy from radon can be in the form of alpha or beta particles and gamma rays. Thus, gamma radiation can be used for indirectly determining the radon concentration.

The YMSG monitoring station is equipped with a gamma ray spectrometer with a scintillation counter inside a container made of 7 cm thick lead, inserted in a High-Density Polyethylene (HDPE) tube at a depth of 2 meters (Figure 2). The NE110 gamma ray is a plastic scintillator with a size of 400 mm × 400 mm × 400 mm, connected to photomultipliers with a diameter of 7.6 cm. Any cosmic gamma rays from atmospheric sources can be prevented from reaching the scintillation counter by using the lead box as a decay chamber to provide passive shielding. When radon gas migrates into the lead box, specific gamma rays may be emitted from

^{214}Pb and ^{214}Bi as decay products in the uranium decay series after radon. The energy of the window channel is limited within the range of 250 to 700 keV, which includes the response of gamma rays at 351 keV for ^{214}Pb and 609 keV for ^{214}Bi . The pulse of the chosen energy signal is converted to generate electron currents by the photomultipliers, and then, they are recorded by the discriminator and computer. The description of a similar operation and setting was described in detail by Giuliani et al. [40] and Fu et al. [39]. Meteorological records (e.g., atmospheric pressure, temperature, humidity, and hourly precipitation) are available from a meteorological station of the Central Weather Bureau (CWB) of Taiwan, which is located at a distance of 1 km from the YMSG gamma ray station. Measurements are recorded once every hour and immediately wirelessly transmitted to the server at the Institute of Earth Sciences, Academia Sinica, from the gamma ray and meteorological stations.

The methodology of the operation at the TPT radon station follows that as described in detail by Fu et al. [41]. Seismic data were acquired from the earthquake catalogs of the CWB. In total, there were 482 seismic events with magnitudes (M_L) ranging from 1.4 to 6.7 during the monitoring period in the Taiwan area.

3. Results

The gamma ray results along with meteorological data recorded on an hourly basis at the YMSG station were presented in Figure 3. The gamma counting rate was observed to vary from 589 to 25,020 cph (counts per hour) with an average value of 3,172 cph. During the same period, atmospheric pressure (P_{atm}) varies from 924 to 958 mbar, atmospheric humidity (H_{atm}) from 31 to 100%, and atmospheric

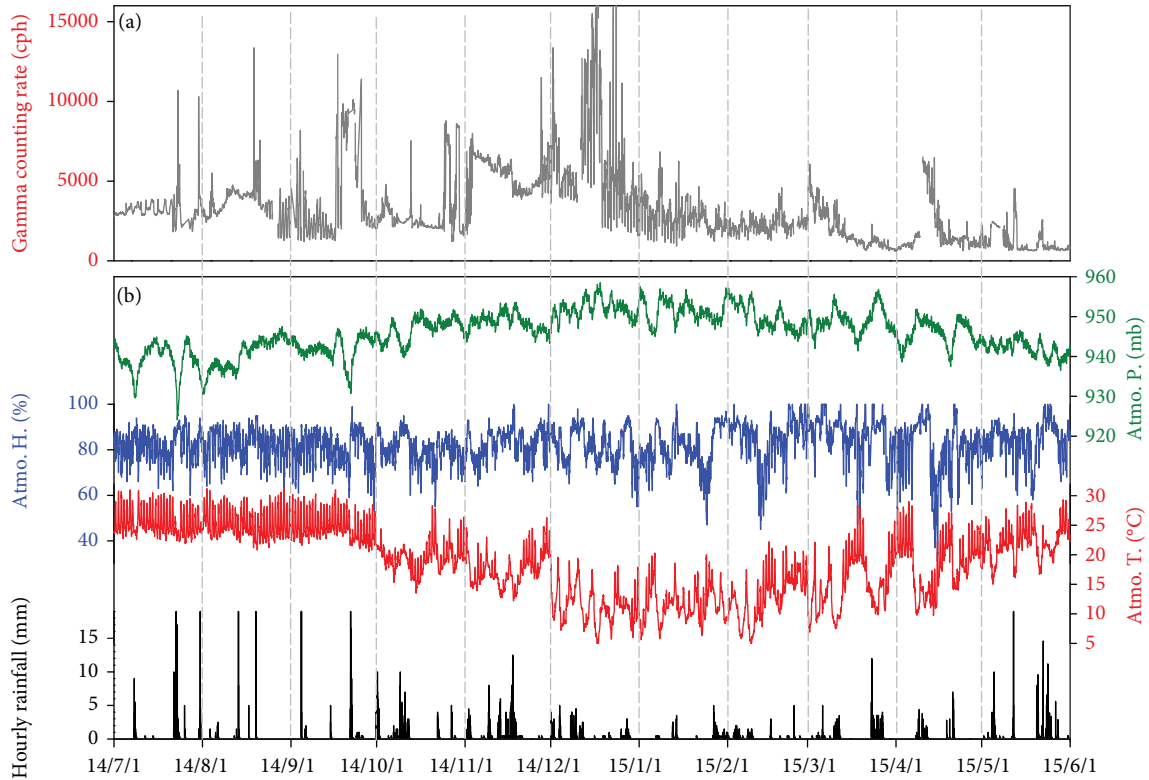


FIGURE 3: Temporal variations of gamma rays at the YMSG station (a) and (b) atmospheric data, including atmospheric pressure (green), humidity (blue), temperature (red), and hourly rainfall (black) (b) from July 1, 2014, to June 1, 2015.

temperature (T_{atm}) from 4.4 to 31.4°C, with the mean value of 946 mbar, 82.3%, and 18.5°C, respectively. Temporal variations of the meteorological factors are not evident. After the initial adjustment for the detector, the gamma counting rate decreases to a low background value of about 1,100 cph for the period from March 15 to June 1, 2015. Substantial increases in the gamma counting rate were observed periodically over the observation period. A wide range of radon gas concentrations in the soil at the TPT station from 7,337 to 31,931 Bq/m³ with an average value of 15,589 Bq/m³ was observed and is shown in Figure 4.

4. Discussion

4.1. The Influence of Meteorological Parameters. Some earlier investigations have revealed that changes in gamma rays may be affected by meteorological factors, such as atmospheric pressure, humidity, temperature, and precipitation [39, 42–46]. Therefore, it is necessary for meteorological effects to be taken into consideration when analyzing gamma ray results. Fu et al. [39] proposed that the air temperature had significant influence on the gamma ray measurements in eastern Taiwan due to a high correlation coefficient between the air temperature and the presence of the gamma ray. However, the influences of meteorological pressure, humidity, and temperature on gamma ray variations were relatively small due to the extremely low correlation coefficients of 0.0178, 0.0022, and 0.0139, respectively, for the study period (Table 1). These low correlation coefficients indicated that

the gamma variations at the YMSG station are not mainly controlled by meteorological parameters.

Fu et al. [39] presumed that the significant decrease in temperature cause reduced sensitivity of electrical sensors. In this study, the marginal effect of atmospheric temperature at the YMSG station, which is situated in a mountainous area (approximately 700 m a.s.l.), may be associated with the relatively small fluctuation of underground temperatures because of the tree shade. To evaluate this possibility, the in situ temperature measurement with a gamma sensor in the monitoring station should be considered in the future.

The temporary, abrupt increase in gamma ray was observed periodically, which may be related to heavy rainfall or precipitation accumulation. This sudden increase in the gamma dose was usually followed by a rapid return to background levels within a few hours after the rain stopped. These observations suggest that short-term (few hours) fluctuations in the rise of gamma rays were partially influenced by a heavy precipitation event, as has been similarly described before [43, 47–49].

For the TPT station, Fu et al. [41] have pointed out that the effects of meteorological parameters on soil radon concentration were insignificant. Therefore, meteorological effects were not taken into consideration during the duration of soil radon concentration anomalies at the TPT station.

4.2. The Relationship between Gamma Ray and Radon Variations and Seismicity. The time series gamma ray data from YMSG was plotted with seismic events, rainfall records, and the radon concentrations of TPT for a period from

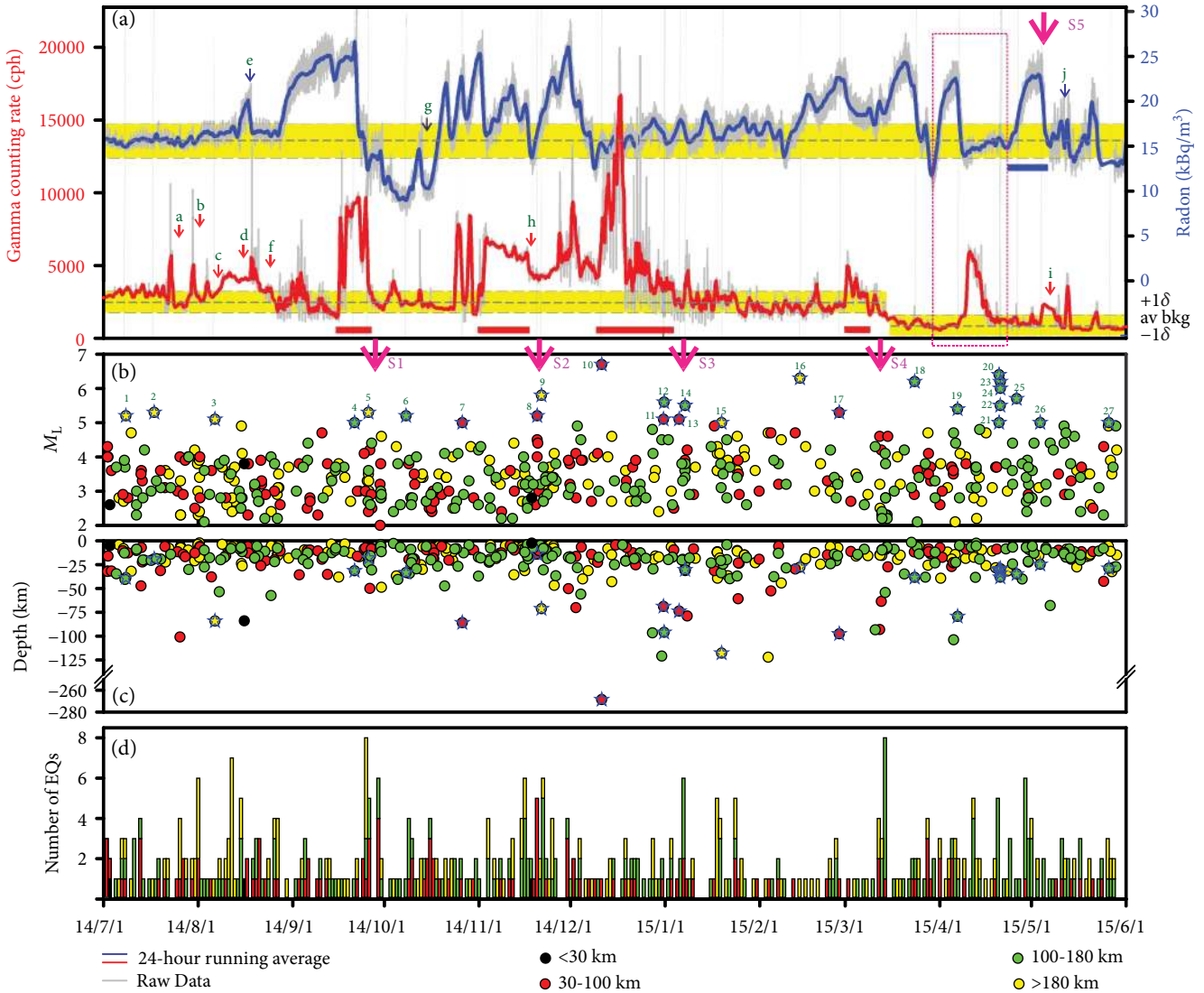


FIGURE 4: The temporal variations of (a) radon and gamma rays from July 1, 2014, to June 1, 2015. Blue and red lines indicate the 24-hour running averages of radon and gamma ray data, respectively; grey lines represent the raw data. Horizontal dashed lines indicate the annual averages of the observed results, and the anomalous threshold value ($\pm 1\sigma$) is included with the yellow bar. The magnitude and depth of the earthquakes are plotted in (b) and (c), respectively. The circle symbols in different colors indicate that the epicenter distance of the earthquake to YMSG station was <30 km (black), 30 to 100 km (red), 100 to 180 km (green), and >180 km (yellow). Red and blue arrows indicate the specific seismic events. The accumulated number of earthquakes at different distances is also shown as a histogram (d). One specific period of time marked by the dashed rectangle is enlarged in Figure 5.

June 1, 2014, to June 1, 2015, in Figure 4. Many substantial increases in gamma ray and radon concentrations were recorded periodically, and slight diurnal variations were also recorded as shown by the grey line in Figure 4(a). The 24-hour running average of the gamma counting rate and radon concentrations are shown as red and blue lines, respectively, in Figure 4(a) to distinguish their background values. The threshold values were determined using the mean value plus one standard deviation as shown by the grey dashed lines.

The magnitude and depth of the earthquakes were plotted in Figures 4(b) and 4(c), respectively. The seismic event data included the magnitude and depth of the earthquakes, which are sorted by the distances between the YMSG station and the epicenter of the earthquakes for discussing the spatial

distribution of the seismic events. The earthquakes at different distances are plotted in different colors as black, red, green, and yellow circles for distances less than 30 km, 30 to 100 km, 100 to 180 km, and greater than 180 km, respectively. To show the intensity of seismic activity, the accumulated numbers of earthquakes at different distances are also included (Figure 4(d)). The distribution of some specific seismic events is marked in Figure 1, including earthquakes having a magnitude ≥ 5 , an earthquake swarm, and a local earthquake with an epicentral distance less than 30 km to the monitoring station, and the catalog of those earthquakes is listed in Table 2.

The temporal variation of gamma rays and radon shows similar patterns, whereas some high gamma ray

TABLE 1: Correlations between gamma ray variations and meteorological parameters, including atmospheric pressure (P_{atm}), temperature (T_{atm}), and humidity (H_{atm}) during different time periods.

| | γ -Rays & P_{atm} | γ -Rays & T_{atm} | γ -Rays & H_{atm} |
|---------|-----------------------------------|-----------------------------------|-----------------------------------|
| 2014/07 | 0.2306* | 0.0339 | 0.0064 |
| 2014/08 | 0.1056 | 0.0585 | 0.0504 |
| 2014/09 | 0.1105 | 0.0253 | 0.0091 |
| 2014/10 | 0.021 | 0.0095 | 0.0041 |
| 2014/11 | 0.0316 | 0.0124 | 0.0267 |
| 2014/12 | 0.249 | 0.0171 | 0.1124 |
| 2015/01 | 0.1932 | 0.0285 | 0.3428 |
| 2015/02 | 0.0082 | 0.0053 | 0.0405 |
| 2015/03 | 0.0004 | 0.1215 | 0.3099 |
| 2015/04 | 0.2202 | 0.0326 | 0.3606 |
| 2015/05 | 0.1838 | 0.02 | 0.3988 |

*R-squared values.

and radon concentration peaks in the entire spectrum can also be observed. In addition, it is noted that the increase of soil radon concentrations usually occurred before the change in gamma ray counts. For calculating the anomalies, the anomalous thresholds are found using the average value plus one to three standard deviations [50–52]. In this study, the precursory anomalies were determined as when the running average exceeds one standard deviation, which then can be used to correlate with seismic events (Figure 4 and Table 2).

A comparison of gamma ray data and earthquakes detected by the CWB over the monitoring period indicates that the presence of gamma ray anomalies was quite often accompanied by seismic activity, except for fourteen of twenty-seven marked earthquakes with magnitudes ≥ 5 (1, 2, 3, 4, 5, 9, 14, 15, 16, 18, 19, 22, 25, and 26 in Figure 1 and Table 2). The Events 14 and 22 are considered to be triggered by an earlier earthquake series due to different focal mechanisms and will not be further discussed. Some short-term impulsive gamma ray increases were also recorded with peak values greater than the threshold value before the Events a, b, c, d, f, h, and i. It is noted that the magnitude of those events is smaller than five with a hypocenter depth deeper than 40 km or an epicenter distance of <30 km to the monitoring station. In addition, four significant long-term increases of the gamma ray were also observed from September 15 to 25, 2014, from November 2 to 17, 2014, from December 11, 2014, to January 3, 2015, and from February 28 to March 10, 2015. These anomalies could not be correlated with any significant earthquake, except for the third period. However, the accumulated number of earthquakes displays unusually high seismic activity, such as S1, S2, S3, and S4, which may be related to gamma ray anomaly periods as shown by the red bars in Figure 4(a). These four seismic swarms occurred at a depth of 4 to 50 km from September 26 to 29, 2014, with magnitude 2.0 to 4.4 for S1; at a depth of 4.9 to 71 km from November 20 to 22, 2014, with magnitude 2.7 to 5.2 for S2; at a depth of 23 to 31 km

on January 7, 2015, with magnitude 3.4 to 5.5 for S3; and at a depth of 6 to 93 km from March 12 to 14, 2015, with magnitude 2.2 to 4.6 for S4, respectively. The distance between the epicenters of these four seismic swarms and the YMSG station was approximately 50 to 130 km.

Similarly, the presence of radon anomalies at the TPT station could also be connected with marked seismic events with a magnitude ≥ 5 , except for the Events 1, 2, 3, 5, 9, 11, 12, 13, 14, 15, 16, 19, 22, 25, and 26 in Figure 1 and two local earthquakes (Events e and g). It is noted that the soil radon obviously increased during April 24 to May 3, 2015, and there was a sudden drop in radon concentrations on May 3 followed by an earthquake swarm (S5), which occurred at depths of 11–14 km from May 1 to 15, 2015, with magnitudes 4.2 to 4.9. The distance was approximately 70 km from the epicenter of the earthquake swarm (S5) to the TPT station as shown in Figure 1.

Each gamma ray and radon anomaly is assigned to a definite earthquake and could only be correlated with the precursory anomalous increase a few days to a few weeks before the event occurrence. Each precursory time was illustrated in Figure 5, and the precursory catalog of each seismic event is listed in Table 2.

Figure 5 shows an example for the preearthquake anomalous period of gamma ray and radon results from March 29 to April 22, 2015, before Event 20. Here, ΔT_{Gamma} and ΔT_{Radon} are the time differences between the start time of an anomaly and the earthquake event for the gamma ray and radon anomalies, respectively. In Figure 5, the radon at TPT rose progressively until April 7, 2015, with a peak value of about 23 kBq/m³. A temporal decrease in radon occurred from April 7 to 9, 2015, and the value returned to the background value with a slight variation before Events 20 to 24. Meanwhile, a significant increase in gamma rays at the YMSG was observed from April 7 to 11, 2015, with a peak value of 6,000 cph. Then, a progressive decrease happened from April 11 to 19, 2015, when the gamma counting rates returned to the background value, followed by Events 20 to 24, which occurred at depths of 29–38 km with magnitudes of 5.0 to 6.4 on April 20, 2015. The distance was approximately 150 km from the epicenter of the largest earthquake ($M_L = 6.4$) to the YMSG station. In this case, the significant changes in radon and gamma ray were recorded as precursors about 21 days and 12 days, respectively, before Event 20, which occurred on April 20, 2015.

4.3. Spatial Variation of Related Earthquakes. Most large events could be linked to anomalous changes in gamma rays and radon at both stations (Table 2), and few events were only associated with one of the two anomalies. Anomalies in gamma and/or radon accompanying the relevant earthquake, which can be called the precursory event (PE), and the distribution of those events are shown in Figure 1. For YMSG, two groups of earthquakes may be recognized based on the epicenter distance.

Group A included the PE with an epicenter distance of <30 km and showed anomalous gamma ray changes over a period of a few days (e.g., Event h). The PE with epicenter distances of 30 to 180 km showed anomalies in gamma rays

TABLE 2: Catalog of gamma ray and radon anomalies and related earthquakes occurring in Taiwan from July 1, 2014, to June 1, 2015.

| No. ¹ | Date | Long. (°E) | Lat. (°N) | Mag. (M_L) | Depth (km) | Distance (km) | $\Delta T_{\text{Radon}}^2$ (day) | $\Delta T_{\text{Gamma}}^2$ (day) | Type |
|------------------|------------------|------------|-----------|----------------|------------|---------------|-----------------------------------|-----------------------------------|------|
| 1 | 2014/7/8 07:06 | 122.41 | 23.29 | 5.2 | 38.8 | 224 | None | None | — |
| 2 | 2014/7/17 14:11 | 121.36 | 22.25 | 5.3 | 19.1 | 322 | None | None | — |
| 3 | 2014/8/6 11:46 | 121.46 | 22.23 | 5.1 | 84.2 | 324 | None | None | — |
| 4 | 2014/9/21 05:14 | 121.54 | 23.59 | 5 | 31.5 | 173 | 21 | None | C |
| 5 | 2014/9/25 18:35 | 121.3 | 22.74 | 5.3 | 18 | 269 | None | None | — |
| 6 | 2014/10/8 02:08 | 121.56 | 23.65 | 5.2 | 33.4 | 167 | 15 | 10 | C |
| 7 | 2014/10/26 12:47 | 122.23 | 24.75 | 5 | 86.0 | 81 | 8 | 3 | W |
| 8 | 2014/11/20 01:46 | 122.02 | 24.89 | 5.2 | 13.9 | 55 | 17 | 17 | O |
| 9 | 2014/11/21 11:29 | 120.13 | 20.65 | 5.8 | 71.3 | 520 | None | None | — |
| 10 | 2014/12/11 05:03 | 122.17 | 25.7 | 6.7 | 268.6 | 86 | 23 | 14 | W |
| 11 | 2014/12/31 11:06 | 121.79 | 24.89 | 5.1 | 68.9 | 37 | None | 20 | W |
| 12 | 2014/12/31 15:54 | 122.6 | 24.55 | 5.6 | 96.1 | 125 | None | 20 | W |
| 13 | 2015/1/5 13:53 | 122.04 | 24.74 | 5.1 | 73.8 | 67 | None | 20 | L |
| 14 | 2015/1/7 12:48 | 121.7 | 24.26 | 5.5 | 30.4 | 100 | None | Unknown | C |
| 15 | 2015/1/19 11:48 | 121.57 | 22.76 | 5 | 117.9 | 265 | None | None | — |
| 16 | 2015/2/14 04:06 | 121.46 | 22.66 | 6.3 | 27.8 | 276 | None | None | — |
| 17 | 2015/2/27 00:50 | 122.29 | 24.68 | 5.3 | 97.6 | 90 | 13 | 11 | L |
| 18 | 2015/3/23 18:13 | 121.76 | 23.7 | 6.2 | 38.4 | 162 | 9 | None | C |
| 19 | 2015/4/6 19:37 | 122.75 | 24.4 | 5.4 | 79.4 | 146 | None | None | — |
| 20 | 2015/4/20 09:42 | 122.37 | 24.05 | 6.4 | 30.6 | 147 | 21 | 12 | I |
| 21 | 2015/4/20 09:49 | 122.49 | 24.02 | 5 | 30.7 | 157 | 21 | 12 | I |
| 22 | 2015/4/20 19:20 | 121.75 | 23.63 | 5.5 | 38.3 | 170 | Unknown | Unknown | C |
| 23 | 2015/4/20 19:45 | 122.41 | 24.12 | 6.2 | 33.9 | 143 | 21 | 12 | I |
| 24 | 2015/4/20 19:59 | 122.39 | 24.01 | 6 | 29.4 | 152 | 21 | 12 | I |
| 25 | 2015/4/26 04:01 | 122.44 | 24.03 | 5.7 | 34.8 | 153 | Unknown | Unknown | — |
| 26 | 2015/5/3 21:09 | 122.42 | 23.94 | 5 | 24.8 | 160 | None | None | — |
| 27 | 2015/5/26 08:56 | 122.49 | 23.87 | 5 | 28.9 | 171 | 7 | 6 | I |
| a | 2014/7/26 00:43 | 122.11 | 24.92 | 4 | 101 | 61 | None | 4 | — |
| b | 2014/8/1 12:38 | 121.6 | 24.58 | 4 | 6.1 | 64 | None | 3 | — |
| c | 2014/8/5 08:34 | 121.67 | 24.4 | 3.6 | 38 | 84 | None | 3 | — |
| d | 2014/8/16 02:51 | 121.53 | 24.92 | 3.8 | 84 | 26 | None | 7 | — |
| e | 2014/8/17 06:52 | 121.04 | 24.72 | 2.3 | 6.2 | 71 | 3 | None | — |
| f | 2014/8/24 20:39 | 122.09 | 24.3 | 4 | 57.5 | 109 | None | 6 | — |
| g | 2014/10/19 17:50 | 121.1 | 24.68 | 2.9 | 6 | 70 | 4 | None | — |
| h | 2014/11/18 07:22 | 121.56 | 25.16 | 2.8 | 2.5 | 1 | None | 2 | — |
| i | 2015/5/7 04:04 | 121.33 | 24.1 | 4.4 | 68 | 119 | None | 4 | — |
| j | 2015/5/9 20:32 | 121.23 | 24.66 | 3.5 | 10 | 64 | 1 | None | — |

¹The label of relevant earthquakes marked in Figures 1 and 4. ²The time difference between the start time of an anomaly and the earthquake event. “Unknown” indicates that the relationship between an anomaly and the earthquake cannot be identified.

over several days to weeks (e.g., Events 10, 17, and 20) which was identified as Group B. It was noted that no anomaly was found when the distance of the earthquake epicenter was farther than 180 km, such as Events 1, 2, 3, 5, 9, 15, and 16.

A diagram showing the relationship between the depth of the PE and latitude location and the seismic velocity structure along the a-a' line (Figure 1) is presented in Figure 6. The seismic V_p structure was found by using a local earthquake tomography model developed by Wu et al. [53]. The yellow dashed line indicates the possible location of the subduction zone as highlighted by contoured V_p/V_s

values higher than 1.78 [54]. It is noteworthy that a series of the PE located along the subduction zone progressively deepen to the north from ~28 to 268 km deep, except for Event 8. This suggests that an anomaly of gamma rays at the YMSG station may be connected with the geodynamics process within the subduction zone in north and north-eastern Taiwan.

For the TPT, Fu et al. [41] proposed that radon anomalies at the TPT station could be found when a small earthquake has hypocenters shallower than 15 km within 30 km of the monitoring station or when a large earthquake has deeper

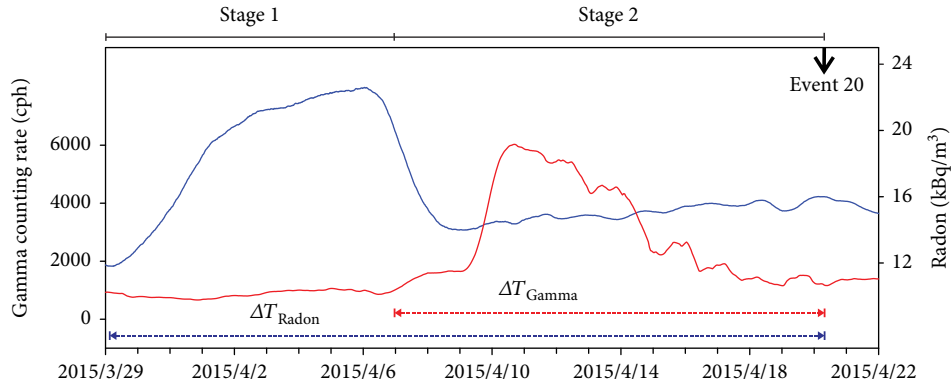


FIGURE 5: Variations of gamma ray and radon data during the period from March 29 to April 22, 2015. The preearthquake anomalous time from the gamma ray and radon results is represented by ΔT_{Gamma} and ΔT_{Radon} , respectively, showing that a radon anomaly occurs earlier than a gamma ray anomaly with respect to an earthquake event.

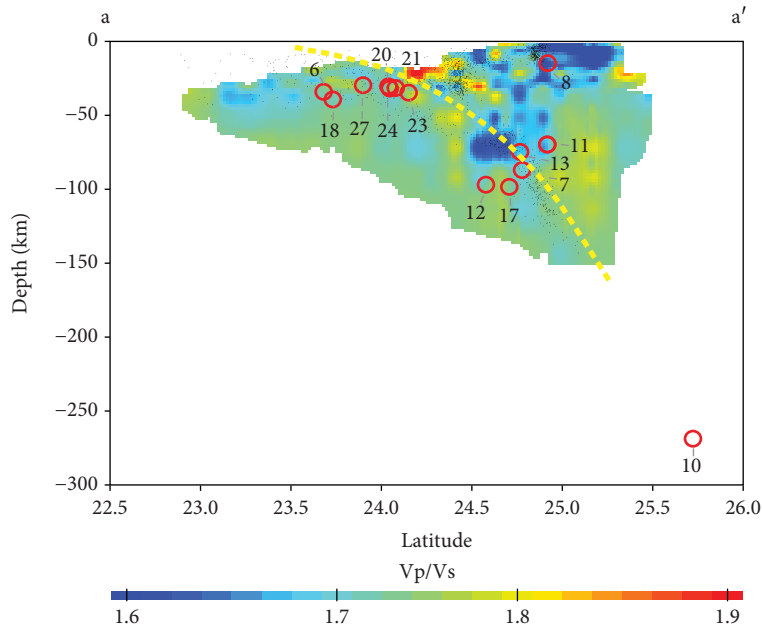


FIGURE 6: The seismic velocity structure transecting northern Taiwan along the a-a' line of Figure 1 is derived based on the local earthquake tomography model by Wu et al. [53]. The dots show the relocated events for $M > 3$ within ± 5 km of the profile. The red open circle means that the projected earthquake with a preanomalous signal is listed in Table 2. Note that the hypocenters get deeper towards the north. The yellow dashed line indicates the subduction boundary between the PSP and the EP.

hypocenters (>20 km) at a greater distance (>45 km). In this paper, the similar anomalies in radon are also confirmed before the near- and far-field earthquakes, such as Events 6, 7, 8, 10, 17, 18, 20, 21, 23, 24, 27, e, g, and j. The preearthquake activities of the far-field events may be transmitted to the region below the TPT site along the major decollement (e.g., [55, 56]) and can facilitate the release of soil gases when squeezed. Hence, radon anomalies could be detected at the TPT station away from the large event [41].

Some fault plane solution diagrams of the studied earthquakes were plotted in Figure 1, and five predominant patterns can be classified. Similarly, Kao et al. [57] proposed five major seismogenic structures in the southernmost Ryukyu Arc-Taiwan region. They were (a) the Collision Seismic

Zone (CSZ, C type), associated with the relative plate convergence between the EP and the PSP along the Longitudinal Valley (e.g., Event 6), (b) the Interface Seismic Zone (ISZ, I type), related to a plate boundary between the EP and the PSP with low-angle thrust faulting to the north with a depth range of 10 to 30 km (e.g., Events 20, 21, 23, 24, and 27), (c) the Wadati-Benioff Seismic Zone (WBSZ, W type), indicating that the earthquakes deeper than ~ 60 km occurred within the Wadati-Benioff zone striking east to west and extending westward beneath NE Taiwan (e.g., Events 7, 10, 11, and 12), (d) the Lateral Compression Seismic Zone (LCSZ, L type), related to a mixture of thrust and oblique strike-slip with P -axes in approximately an east-west direction, roughly parallel to the local strike of the trench-arc system thrust or

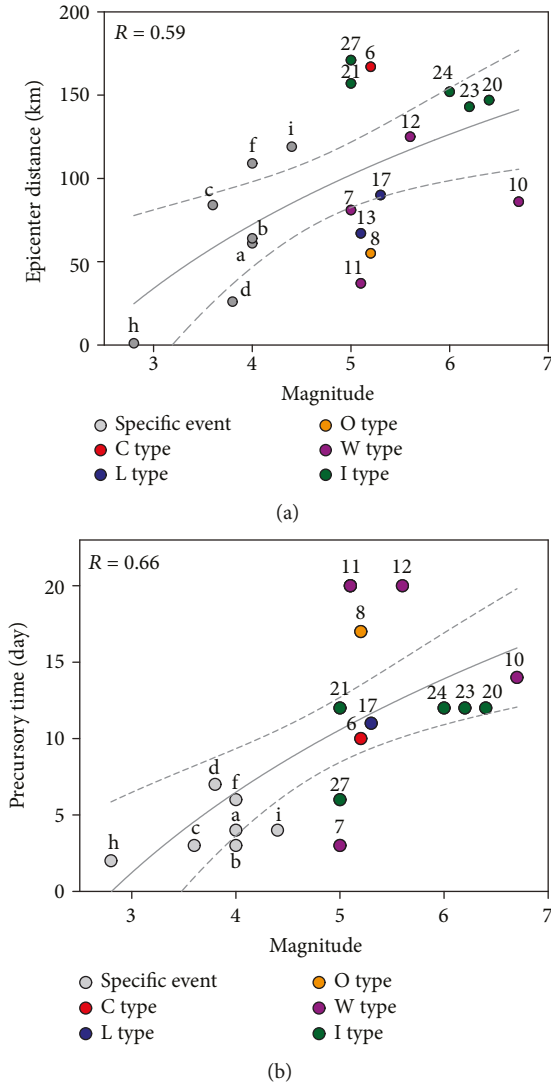


FIGURE 7: Variation diagrams showing the relationship between earthquake magnitude and (a) epicenter distance to the YMSG monitoring site and (b) distribution of the precursory time. Fitted curves are shown as solid lines. Dashed lines indicate the 95% confidence interval of the data set.

oblique-thrust faulting (e.g., Events 13 and 17), and (e) the Okinawa Seismic Zone (OSZ, O type), reflecting the gradual changing state of strain associated with the crustal deformation of the Okinawa opening with a characteristic of shallow earthquakes occurring (e.g., Event 8).

The dependent relationship between the epicenter distance to the YMSG site and the magnitude of the PE was shown in Figure 7(a). A good correlation ($R = 0.66$) was obtained between these variables. The best-fit line may be used to estimate the magnitude of an earthquake with a given epicenter distance. In other words, a larger magnitude of the PE might be distributed in greater epicenter distances of the YMSG station. As mentioned above, when the distance of the earthquake epicenter was farther than 180 km, no anomaly was detected. And the distributions of the PE, including I type, L type, O type, W type, and specific events, get closer towards the YMSG station.

The good correlation between the magnitude of PE and precursory time (ΔT_{Gamma}) was presented in Figure 7(b). The moderate correlation ($R = 0.59$) indicates that it is possible to forecast the magnitude of an impending earthquake. A longer precursory time of gamma ray anomalies appears to be associated with a larger upcoming earthquake, especially the occurrence of I type and W type with high probability. We also found that an anomaly of C type was recorded in only one (Event 6) out of five events, suggesting that C type may be not sensitive for the YMSG station but valuable for the TPT station (Table 2).

The YMSG station is located at the western end of the subduction system in the northern Taiwan area, where the PSP is subducting northward beneath the EP. Furthermore, this region is still seismically active due to continuous convergence and volcanic activity, with changes in high-pressure fluid affected by a deeper pump system [58]. Some microcracks may be also developed during this process, causing the short duration gamma ray anomalies. Similarly, Padrón et al. [59] found that crustal deformation and fracturing during high seismic activity enhance the release of radiogenic helium produced in the crust. Lee et al. [23] reported a short-term negative anomaly of $^3\text{He}/^4\text{He}$ ratios, which indicated that the radiogenic gas release of crustal components affected the degassing system for the short term.

It may be suggested that due to continuous convergence, the preseismic slow slip may occur around the subduction zone before the event. The volumetric expansions of the EP around northern Taiwan then released the additional radiogenic gas from the crustal component. The increase in gamma rays can be attributed to the expansion around the YMSG that produced new fractures for gas and fluid migrations, potentially associated with the slow slip (e.g., Events 6, 7, 8, 10, 17, 20, 21, 23, 24, and 27). A similar explanation of the preseismic slow slip before the Tohoku earthquake was illustrated by Orihara et al. [60].

Hence, based upon the above-mentioned points, the preliminary outline of anomalous data regarding relevant earthquakes can be concluded, which may be associated with the dynamics of the subduction process.

4.4. Relationship between the Tectonic Setting and the Mechanism for Anomalous Signals. The geodynamic setting of the arc-continent collision in Taiwan is well defined by the oblique collision between the PSP and the EP, with the former plate moving about 8 cm/year [1]. To the north of the island lies the Ryukyu subduction zone, which results from the subduction of the PSP beneath the EP. A schematic diagram of the proposed physical mechanism based on the anomalies in radon at TPT and changes in gamma rays at YMSG is shown in Figure 8. Stages 1 to 2 are considered to have occurred successively before the significant earthquake, which occurred along the subduction region. During Stage 1, the prestress caused by the northwestern movement of the PSP may be transmitted by collision with the Kuanyin High along the major decollement to trigger radon enhancements due to the development of microcracks below the TPT station, while no significant changes at the YMSG station occur during this stage. The description of the similar model for

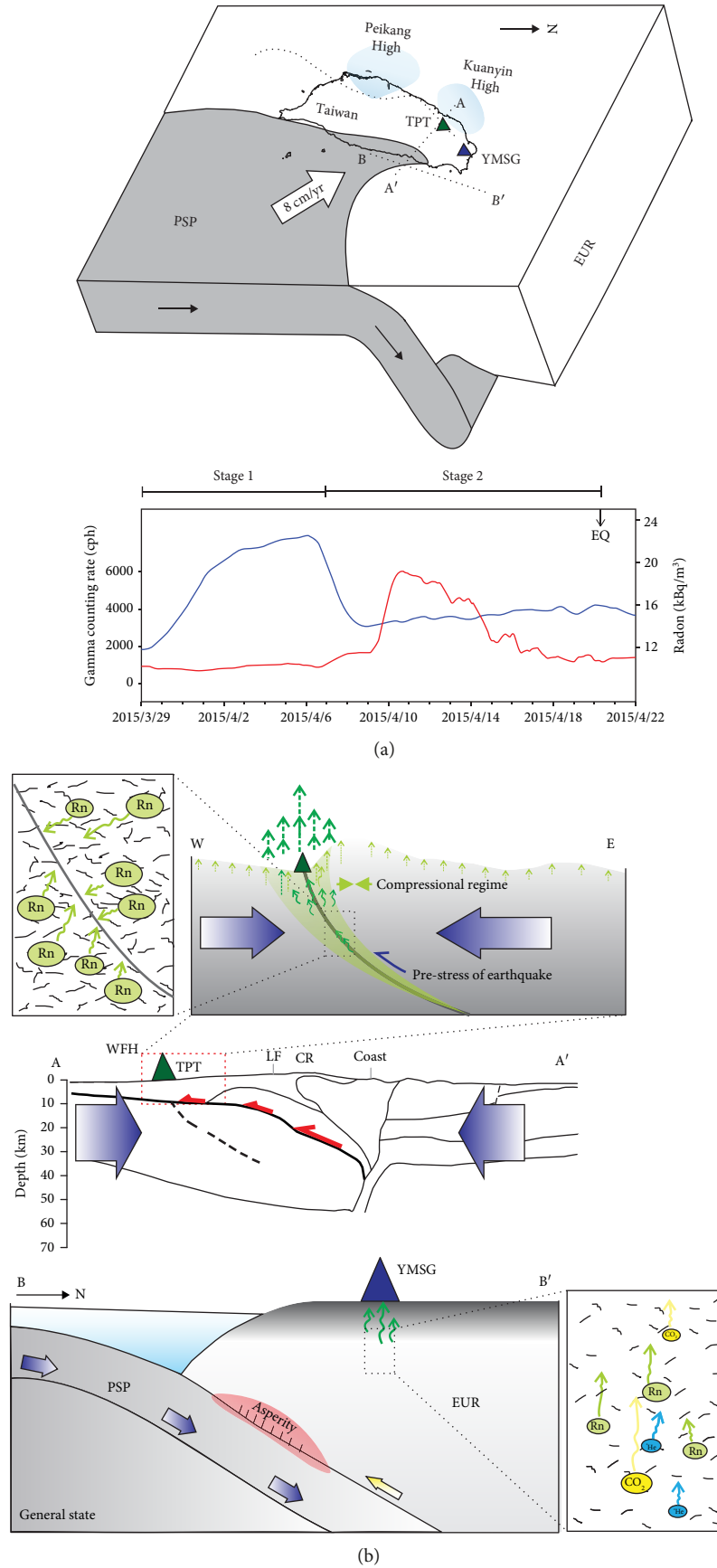


FIGURE 8: Continued.

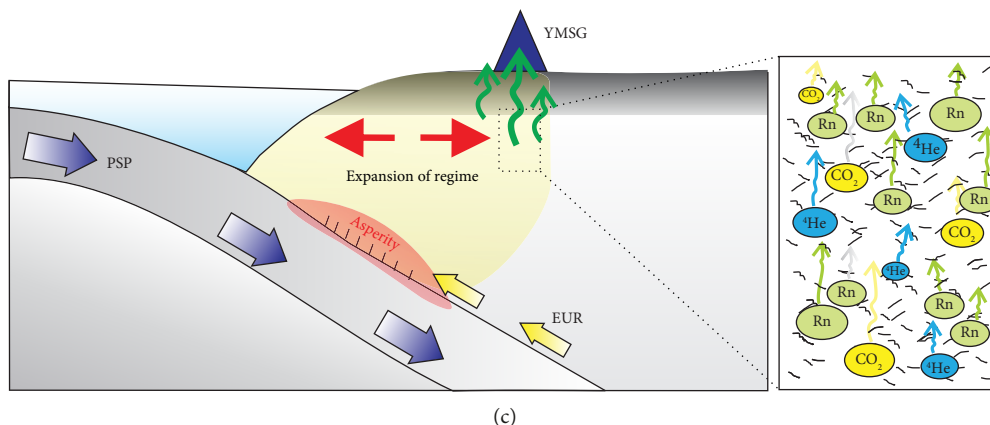


FIGURE 8: (a) A schematic diagram of the proposed physical mechanism based on the gamma ray anomaly of YMSG and radon anomaly at TPT. Stages 1 (b) and 2 (c) are considered to have occurred in order before the significant earthquakes. The schematic diagram showing a crustal-scale cross-section along the geophysical transect (AA' in (a)) modified by Fu et al. [41] is shown in (b).

TPT was described in detail by Fu et al. [41]. At Stage 2, further movement of the PSP may be locked by the EP and accumulated elastic stress results in the increase of radiogenic materials, such as gamma rays and radon gas, due to an increase in the porosity and fractures below the YMSG station.

Based on GPS observations, Rau et al. [61] and Ching et al. [62] concluded that crustal deformations showed the transition of surface strain in the type of tectonic activity from an arc-continent collision to a subduction/back-arc opening in northern Taiwan. Continental collision (around the TPT region) to extension (around the YMSG region) may be associated with the similar geodynamic process of stage 1 to stage 2, respectively, at different time scales.

Therefore, anomalous signals from gamma rays and radon that are attributed to changes in pore pressure by the Coulomb stress changes may be attributed to near/far-field earthquakes or strain-induced changes in permeability within the preparation zone of the earthquake [63–65]. During the build-up of stress by preseismic activity, some of the existing features and the permeability of the fault zone are ruptured. The highly porous and permeable fault zones are then filled with gas and fluid. These closing and opening processes within the fault system are induced by the tectonic stress, which may be responsible for the episodic and/or periodic seismicity in northern and northeastern Taiwan.

5. Conclusions

The major findings of this study are listed below:

- (1) Gamma ray variations are monitored at the YMSG monitoring station, located in the volcanic area of north Taiwan. Results of the temporal gamma ray variations are compared with meteorological records. The marginal effect of atmospheric parameters is observed in gamma ray measurements at the YMSG station
- (2) Under continuous monitoring, many anomalously high radon values at TPT and gamma rays at YMSG

were observed a few days to a few weeks before the seismic events

- (3) Two groups of earthquakes can be identified at the YMSG station based on relevant earthquakes: earthquakes with shallower hypocenters and that are concentrated within 30 km of the monitoring station with a smaller magnitude and the earthquakes with deeper hypocenters and that are distributed at greater distances (~30 to 180 km) from the monitoring site
- (4) The precursory changes in soil radon at the TPT station and gamma rays at the YMSG station may represent the preparation stage of an earthquake. When the observed precursory signals from these two stations with time delay occur, the possible impending large earthquake around the plate boundary in northeastern Taiwan can then be expected, which is located approximately between latitudes 23.9°N and 26.7°N and longitudes 121.3°E and 122.6°E
- (5) Both radon and gamma ray measurements provide a useful tool and act as a good indicator for exploring earthquake precursors, especially in north and northeastern Taiwan

Data Availability

The data used to support the findings of this study are available from the corresponding author upon request.

Disclosure

A preliminary study has been presented as a conference abstract in the 2005 AGU-Fall Meeting.

Conflicts of Interest

The authors declare that they have no conflicts of interest.

Acknowledgments

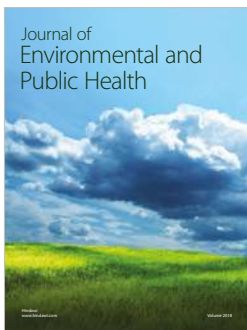
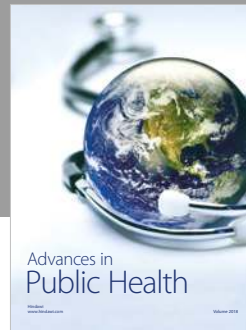
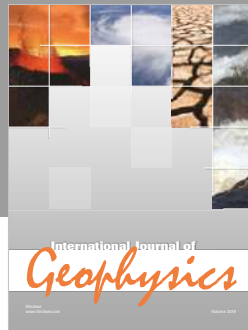
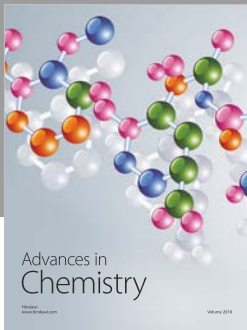
We dedicate this paper to the late Prof. Tsanyao Frank Yang (TFY), who passed away on the 12th of March 2015. He had worked as a geochemist at the Department of Geosciences, National Taiwan University (NTU), Taiwan, for more than 20 years. His passion for research included setting new standards for relevant measurements, which enables us to present our work here. We thank Messrs. Shin-Jung Lin and Kou-Wei Wu for helping in the operation and maintenance of the monitoring station. The research is supported by grants MOST 107-2116-M-002-019 through the Ministry of Science and Technology (MOST) and MOTC-CWB-106-E-01 from the Central Weather Bureau (CWB), Taiwan. The TEC contribution number for this article is 00151. We would like to thank Uni-edit (<http://www.uni-edit.net>) for editing and proofreading this manuscript. We are also thankful to the editor Yunpeng Wang and the anonymous referee for their valuable comments and suggestions, which helped us to improve the manuscript.

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