



Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air

M. Ehn¹, E. Kleist², H. Junninen³, T. Petäjä³, G. Lönn³, S. Schobesberger³, M. Dal Maso³, A. Trimborn^{1,4}, M. Kulmala³, D. R. Worsnop^{3,5}, A. Wahner¹, J. Wildt², and Th. F. Mentel¹

¹Institute for Energy and Climate Research (IEK-8), Forschungszentrum Jülich, 52425 Jülich, Germany

²Institute of Bio- and Geosciences (IBG-2), Forschungszentrum Jülich, 52425 Jülich, Germany

³Department of Physics, P.O. Box 64, 00014 University of Helsinki, Finland

⁴Aeromegt GmbH, Verbindungsstraße 27, 40723 Hilden, Germany

⁵Aerodyne Research, Inc., 45 Manning Road, Billerica, MA 01821, USA

Correspondence to: M. Ehn (m.ehn@fz-juelich.de)

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Abstract. High molecular weight (300–650 Da) naturally charged negative ions have previously been observed at a boreal forest site in Hyytiälä, Finland. The long-term measurements conducted in this work showed that these ions are observed practically every night between spring and autumn in Hyytiälä. The ambient mass spectral patterns could be reproduced in striking detail during additional measurements of α -pinene ($C_{10}H_{16}$) oxidation at low-OH conditions in the Jülich Plant Atmosphere Chamber (JPAC). The ions were identified as clusters of the nitrate ion (NO_3^-) and α -pinene oxidation products reaching oxygen to carbon ratios of 0.7–1.3, while retaining most of the initial ten carbon atoms. Attributing the ions to clusters instead of single molecules was based on additional observations of the same extremely oxidized organics in clusters with HSO_4^- (Hyytiälä) and $C_3F_5O_2^-$ (JPAC). The most abundant products in the ion spectra were identified as $C_{10}H_{14}O_7$, $C_{10}H_{14}O_9$, $C_{10}H_{16}O_9$, and $C_{10}H_{14}O_{11}$. The mechanism responsible for forming these molecules is still not clear, but the initial reaction is most likely ozone attack at the double bond, as the ions are mainly observed under dark conditions. β -pinene also formed highly oxidized products under the same conditions, but less efficiently, and mainly C_9 compounds which were not observed in Hyytiälä, where β -pinene on average is 4–5 times less abundant than α -pinene. Further, to explain the high O/C together with the relatively high H/C, we propose that geminal diols and/or hydroperoxide groups may be important. We estimate that the night-time concentration of

the sum of the neutral extremely oxidized products is on the order of 0.1–1 ppt ($\sim 10^6$ – 10^7 molec cm^{-3}). This is in a similar range as the amount of gaseous H_2SO_4 in Hyytiälä during day-time. As these highly oxidized organics are roughly 3 times heavier, likely with extremely low vapor pressures, their role in the initial steps of new aerosol particle formation and growth may be important and needs to be explored in more detail in the future.

1 Introduction

Earth's atmosphere contains a large variety of different trace gases, produced both from biogenic and anthropogenic sources. Many of these trace gases impact everyday life, either through their influence on air quality (e.g. SO_2 , tropospheric ozone, NO_x) or climate. The climate-influence occurs both through the greenhouse effect from e.g. CO_2 and CH_4 , as well as through oxidation products of trace gases that can condense onto, or even form new, aerosol particles. These aerosol particles can either directly scatter incoming solar radiation, or form cloud droplets that affect radiative transfer even more efficiently. The most important inorganic precursors for aerosol formation are SO_2 , NO_x , and NH_3 , but often the organic mass fraction of sub-micron aerosol is significant or even dominant (Kanakidou et al., 2005; Zhang et al., 2007). Organic aerosol (OA) can be primary, i.e. emitted directly as aerosol particles, or secondary, formed from

a wide range of different organic trace gases. The organic trace gases are typically referred to as volatile organic compounds (VOCs), which upon oxidation form products with low enough vapor pressure to form OA. It has been estimated that 10^4 – 10^5 different atmospheric organic species have been measured, and especially biogenic terpenoids have been reported as important players in gas-phase and particulate-phase chemistry (Goldstein and Galbally, 2007).

VOCs in the atmosphere are oxidized by reactions with ozone, OH and NO_3 radicals. During daytime, OH is the most important oxidant, but it has been reported that the sum of all identified inorganic trace gases and VOCs cannot explain the observed OH reactivity (Di Carlo et al., 2004), implying that there exists groups of compounds that are yet to be identified. Holzinger et al. (2005) suggested that highly reactive biogenic precursors could at least partially explain the missing reactivity. Nevertheless, clear discrepancies still exist between models and experiments on secondary organic aerosol (SOA) formation from known organic precursors (Hallquist et al., 2009; Kroll and Seinfeld, 2008).

The mechanism of how organic molecules take part in the first steps of nucleation and growth of new particles has been the focus of several studies in recent years (Paasonen et al., 2010; Metzger et al., 2010), but no conclusive results have yet been obtained. It is widely agreed that sulfuric acid plays an important role in the initial formation step (Weber et al., 1996; Petäjä et al., 2009; Riipinen et al., 2007) whereas other compounds, most likely organics, are needed to explain the majority of the early growth (Kulmala et al., 2001; Ehn et al., 2007; Wehner et al., 2005). Some studies have shown that sulfuric acid alone could explain observed atmospheric formation rates (Sipilä et al., 2010), others that also ammonia or amines are needed (Kirkby et al., 2011; Berndt et al., 2010), and yet others that organics are important (Metzger et al., 2010). Due to high ion concentrations, Kirkby et al. (2011) could directly measure the nucleating clusters of H_2SO_4 and NH_3 , while Wang et al. (2010) suggested three different mechanisms where specific organic precursor molecules were involved in nucleation and growth up to 20 nm. Nevertheless, a conclusive determination of which organics can take part in the first steps of atmospheric new particle formation (NPF) is still incomplete, both related to the chemical mechanism and the specific organic molecules involved.

The formation of new particles can also be enhanced by ions. An additional positive or negative charge can stabilize clusters, slowing their re-evaporation, and thereby facilitating the initial growth. Frequent NPF events have been reported to take place in the boreal forest site in Hyytiälä, Finland (Dal Maso et al., 2005), and most studies on the role of ions in NPF have found that only around 10% or less of new particles are typically formed through ion-enhancement (Gagne et al., 2008; Kulmala et al., 2010). However, others have found ions to play a dominant role in NPF at the same site (e.g. Yu and Turco, 2011).

Naturally charged ions are formed through high-energy collisions of air molecules with energetic particles, mainly from galactic cosmic rays or the radio-active decay of radon (Laakso et al., 2004; Arnold, 2008; Harrison and Carslaw, 2003). The initially formed primary ions are mainly positively and negatively charged N_2 and O_2 , but as these ions quickly collide with stronger acids (negative ions) or bases (positive ions), charge transfer and/or clustering will occur. Typical long-lived negative ions comprise NO_3^- , HSO_4^- , strong organic acids, and their clusters (e.g. Eisele, 1989b; Ehn et al., 2010). In the positive ion spectrum, the dominant peaks are typically alkyl pyridines and alkyl amines (e.g. Ehn et al., 2010; Eisele, 1989a). Typical naturally charged ion concentrations are around 100–1000 cm^{-3} per polarity, but due to the complex and dynamic process forming these ions, the concentration of a certain ion can rarely be translated to a real concentration of the corresponding neutral molecule.

The aim of this work was to explore in more detail the group of high molecular weight ions (mass range roughly 300–650 Da) recently detected by Ehn et al. (2010) in Hyytiälä. These ions typically dominated the night-time negative ion spectra together with nitric acid. At that time, the authors did not have enough information available to identify these peaks, but speculated that the elemental composition suggested highly oxidized, nitrogen-containing organics. If correct, this would suggest a new, to our knowledge not previously identified group of compounds, which might contribute significantly to organic aerosol and new aerosol particle formation.

In this paper, new long-term measurements in Hyytiälä together with extensive measurements at the Jülich Plant Atmosphere Chamber (JPAC) in Jülich, Germany, are combined to provide a conclusive elemental composition for the high-mass ion species initially observed by Ehn et al. (2010). Their suggestions on the composition was confirmed, but as this group of molecules has never been observed or identified before, the emphasis of this paper is to describe the different methods used to unambiguously determine the elemental composition. Additionally, we give a lower limit estimate on the abundance of the highly-oxidized organic ions, and shortly discuss their importance and possible formation mechanisms in the atmosphere.

2 Methods

2.1 APi-TOF

The chemical characterization of the ion population was performed with the Atmospheric Pressure interface Time-Of-Flight mass spectrometer (APi-TOF, Aerodyne Research Inc. & Tofwerk AG; Junninen et al., 2010). It consists of a short APi which samples roughly 0.8 l min^{-1} through a critical orifice, and consequently guides the sampled ions towards the mass spectrometer while pumping away the gas in 3 stages.

In other words, no ionization is performed, and the instrument relies on the sampled air containing naturally charged ions. Once the ions reach the TOF they are orthogonally extracted onto a one- or three-reflection flight path, and the mass of the ions are determined based on their flight times. In this work, two APi-TOFs were used, and they will forthwith be titled the “Jülich” and the “Helsinki” APi-TOFs, respectively. As the unit for mass, we use Dalton (Da), and as the unit for mass-to-charge we will use Thomson (Th), which corresponds to Da per elementary charge. When discussing peaks at only unit mass-to-charge resolution (UMR), we will use e.g. “UMR 62 Th” to denote the location of NO_3^- (exact mass-to-charge 61.9884 Th). Additionally, we assume that all the detected ions are singly charged, as none of our results suggest otherwise. Only negative ions are discussed in this study, as this is the polarity where the highly oxidized ions were observed.

The mass accuracy of the APi-TOF is specified by the manufacturer as 20 ppm (0.002 %), although the instrument can typically be calibrated to less than 10 ppm. The typical achievable resolution is around 4000 Th/Th (ranges from 3000–5000 depending on the exact tuning) when operating in normal, single reflection mode (“V mode”). Two additional reflections inside the TOF can be used, thus producing a flight path in the shape of a W (“W mode”) to increase the resolution. However, the increased resolution is accompanied by a significant decrease in sensitivity, and thus requires longer averaging times to achieve the same signal-to-noise as in V mode. Only the Jülich APi-TOF was operated in W mode, and the resolution reached slightly above 8000 Th/Th in the region of 200–700 Th. The typical transmission of the APi-TOF in V mode is on the order of 0.1–1 % in the mass-to-charge range of 50–800 Th (Junninen et al., 2010), which was of interest in this study.

Evaporation and fragmentation may occur to some extent inside the instrument. Particularly water, which is believed to be clustered with most of the ions at ambient conditions, is typically weakly bound and evaporates from the ions before reaching the TOF region. More strongly bound clusters may also fragment inside the APi-TOF, but based on previous measurements with the instrument, clusters have been frequently detected at high concentrations, suggesting that fragmentation is not a critical issue (Junninen et al., 2010). Additionally, two low-resolution ion mobility spectrometers measuring at ambient pressures were recently found to agree within measurement accuracies with the ion distribution of the APi-TOF, further suggesting that the measured ions are not purely fragments of much larger ion clusters (Ehn et al., 2011).

2.2 Measurements

The ambient measurements were performed during May–November 2010 at the SMEAR II station in Hyytiälä, Finland (Hari and Kulmala, 2005). The site is located within the

boreal forest, with the site being mainly surrounded by Scots pine (*Pinus Sylvestris L.*). The ion measurements were performed using the Helsinki APi-TOF.

The chamber measurements were performed at the Jülich plant atmosphere chamber facility (JPAC; Mentel et al., 2009) at Forschungszentrum Jülich, Germany. The facility consists of three Borosilicate glass chambers (1641, 11501, and 14501) with Teflon floors, but only the 14501 chamber was used in this work. This reaction chamber is mounted in a climate controlled housing, and has the capability of using both real plant emissions from the other chambers, and pure compounds from diffusion sources for further oxidation reactions. For these measurements only pure α - or β -pinene was used, at concentration ranging from a few ppb to several tens of ppb. In addition to the VOCs, ozone and water vapor were constantly added to the chamber via separate inlets. The conditions used in the reported measurements were $T = 16 \pm 1$ °C, $\text{RH} = 63 \pm 3$ %, and O_3 around 85 ppb depending on the amount of added organics. These conditions were similar to the ones used by Mentel et al. (2009), who showed that particle formation events similar to those observed in the boreal forest could be attained in JPAC. VOC concentrations were measured at the inlet and the outlet air of the reaction chamber using a GC-MS (Heiden et al., 1999). The chamber is operated at a constant flow mode with in- and outflow kept constant at around 30 l min^{-1} , giving an average residence time in the chamber of roughly 50 min. The system can be kept in steady state for long periods of time, facilitating longer averages with the APi-TOF which are specifically needed for W mode measurements. These ion measurements were performed with the Jülich APi-TOF. All measurements were performed under low NO_x conditions ($\text{NO}_x < 300$ ppt), and in the absence of UV light.

2.3 Data analysis and APi-TOF mass calibration

All the data presented in this work are from naturally charged ions, therefore these mass spectra require very different interpretation than more commonly presented spectra from instruments using active ionization methods. The main reason is that in the latter instruments, the ionization is directly followed by detection, optimally resulting in a more or less linear response from neutral concentrations to detected ions. For naturally charged ions, detection often occurs several minutes after ionization, which means that the charge can be transferred countless times before detection, and therefore the process will be extremely selective, but at the same time sensitive, to molecules with high (positive ions) and low (negative ions) proton affinities. Additionally, the total ion concentration is in practice controlled by physical processes and is not affected by the molecular composition of trace gases in the air.

The APi-TOF data was analyzed using the MatLab-based tofTools package originally described by Junninen et

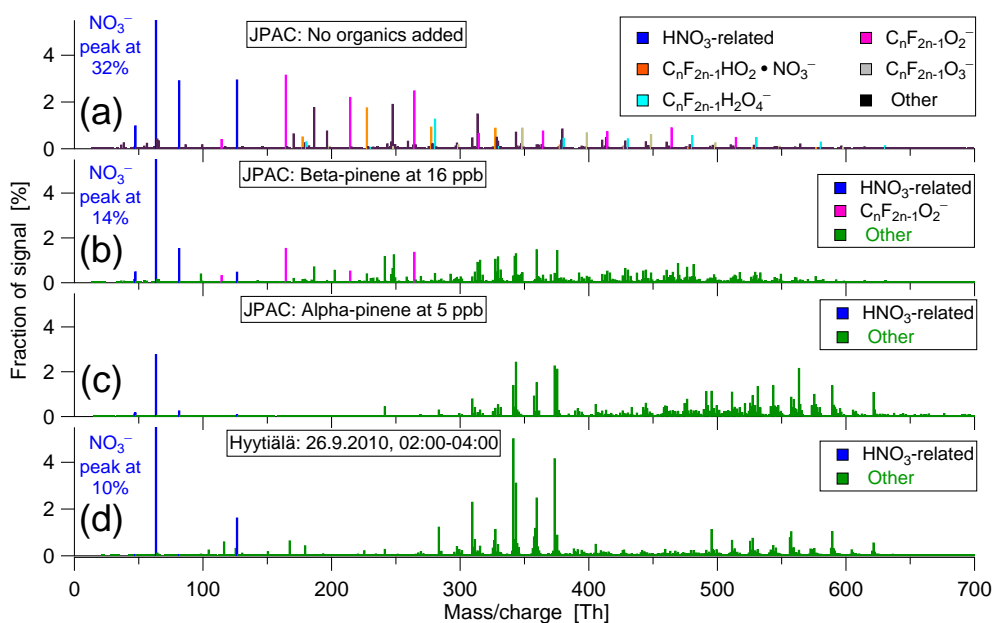


Fig. 1. Negative ion spectra from the JPAC at three different chamber conditions ((a): no addition, (b): β -pinene added, (c): α -pinene added) and from Hyytiälä (d). The NO_3^- signal is allowed to go off scale in (a), (b), and (d). In these cases the peak values are written out explicitly.

al. (2010), though the toolbox has since been extensively improved and expanded.

With the high accuracy and resolution of the APi-TOF, the possible elemental composition of a peak can be narrowed down significantly compared to unit mass-to-charge resolution (UMR) data. However, for this purpose it is extremely important to have a good conversion from time-of-flight to mass/charge (typically referred to as the “mass calibration”). For the APi-TOF, the typical 2-parameter mass calibration (Junninen et al., 2010)

$$\frac{m}{Q} = \left(\frac{t-b}{a} \right)^n, n = 2, \quad (1)$$

where m/Q is mass/charge, t is time-of-flight, and a and b are the free parameters, could not provide the desired high accuracy over the entire 50–700 Th region of interest. Allowing the exponent n in Eq. (1) to vary, greatly improved the fit over the desired range. This method with a variable exponent has been used previously in other instruments, for example in the Aerodyne aerosol mass spectrometer (AMS), which uses the same TOF mass spectrometer (DeCarlo et al., 2006).

The mass calibration is usually performed by using known ions in the spectrum for the calibration, and optimally these peaks should span the entire range to be calibrated. The APi-TOF does not have any “background” signals that are independent of the sample, and therefore the peaks used for calibration vary with the sampled air, which means that other methods need to be employed for the mass calibration. In this work, the starting point was a mass spectrum where known peaks only existed in the sub 200 Th range, and the main range of interest was in the 300–600 Th range. This

meant that peak identification was initially made in the lowest mass/charge region where the calibration was the most reliable based on previously identified peaks, such as NO_3^- , $\text{HNO}_3 \cdot \text{NO}_3^-$, $\text{C}_3\text{F}_5\text{O}_2^-$. The original calibration was used to preliminarily identify additional peaks in the 200–300 Th range, and once successful, the new peaks were used to improve the calibration. This step-wise identification allowed the entire mass range of interest to be covered. With the final 3-parameter fit acquired through this iterative method, all identified ions were well within the 20 ppm instrumental accuracy.

3 Results

3.1 Comparison of negative ion spectra from Hyytiälä and the JPAC

High molecular weight ions have previously been reported from Hyytiälä (Ehn et al., 2010), and now additional long-term measurements during May–November 2010 showed that these ions are indeed abundant almost every night. However, with frequently changing ambient conditions, the conclusive elemental identification of the ions proved a difficult task. For this reason, we tried to mimic the spectra under controlled chamber conditions at the JPAC.

To fully characterize the ion spectra in the JPAC, first the background chamber spectra were studied. Some level of contamination is inevitable in all chambers, and when only H_2O and O_3 are added to the cleaned air going into the chamber, the negative ion spectrum is dominated by the most

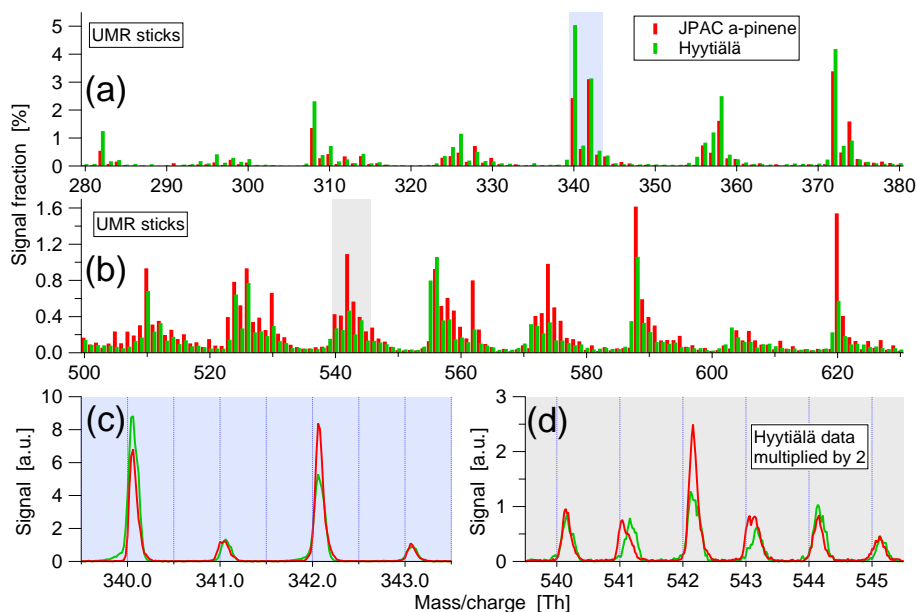


Fig. 2. UMR (a and b) and high-resolution (c and d) comparison of JPAC α -pinene (red) and Hyytiälä ambient (green) data. The bars have been offset slightly for an easier comparison. The spectra correspond to Fig. 1c and d, respectively. The blue-shaded region in plot (a) corresponds to (c), and the grey-shaded region in (b) corresponds to (d).

acidic contaminants. The dominant peak in this case is nitric acid, NO_3^- , at UMR 62 Th, as shown in Fig. 1a. The NO_3^- signal is allowed to go off-scale to better show the rest of the spectrum. NO_3^- is one of the most common negative ions also in the atmosphere, and due to its “stickiness” it is extremely hard to remove completely once introduced into a chamber. In the chamber, previous experiments with NO_x and OH have produced HNO_3 which only very slowly desorbs from the walls. The rest of the ion signal is distributed over the entire spectrum, and is mostly explained by perfluorinated organic acids. The most abundant perfluorinated ion is typically $\text{C}_3\text{F}_5\text{O}_2^-$ at UMR 163 Th. The unit mass of a CF_2 group is 50, and the homologous series is clearly visible at peaks 113, 163, 213, 263, etc. up to at least 563 Th. Several of these acids, mainly the C_8 perfluorooctanoic acid at (413 Th) is used as a surfactant in the emulsion polymerization of fluoropolymers, such as Teflon. Much of the tubing, as well as the bottom plate of the chamber is made of Teflon, and thus the presence of these molecules is not surprising, and likely a common contaminant in many chambers. However, efforts have been made to exchange as much as possible of the Teflon tubing for steel or glass equivalents.

The fluorinated acids are weaker acids than HNO_3 , and thus if the HNO_3 concentration in the chamber is increased, the ion signal from the fluorinated acids decreases, although the concentration of neutral fluorinated acids remain unchanged. Similarly, when adding organics to the chamber, the amount of neutral fluorinated compounds should not decrease, but their influence on the ion spectrum diminishes, mainly due to stronger acids forming from the oxidation

products of the organics (Fig. 1b–c). Figure 1b shows the negative ion spectrum measured in the chamber with a β -pinene concentration of 16 ppb. NO_3^- is still the largest peak, but several new peaks have appeared while many fluorinated ions have disappeared. The most clearly distinguishable fluorinated compounds are the C_3 – C_5 acids at UMR 113, 163, and 213.

Figure 1c shows the resulting spectrum with 5 ppb α -pinene in the chamber. Although using clearly lower organic precursor concentrations than in the aforementioned β -pinene experiment, many more highly acidic organic products seem to have formed, as NO_3^- has decreased dramatically and the fluorinated compounds are barely detectable. A clear pattern of peaks has formed in the 300–400 Th range, and a second pattern is visible at 500–650 Th.

Figure 1d shows a typical night-time negative ion spectrum acquired in Hyytiälä. The similarity to the JPAC α -pinene spectrum in Fig. 1c is striking. The nitric acid dimer ($\text{HNO}_3 \cdot \text{NO}_3^-$) at UMR 125 Th is much larger in Hyytiälä, reflecting that the total amount of neutral HNO_3 is higher, but the same modes in the 300–400 and 500–650 Th ranges are visible in both Hyytiälä and JPAC. The day-time negative ion spectra look very different compared to the night-time spectra in Hyytiälä, with days often dominated by H_2SO_4 -related ions, but also the amount of different organics are most likely higher due to the increased amount of OH. All-in-all, the appearance of H_2SO_4 makes the amount of negative charges available for the organics lower, and the increased amount of different organics spreads out the signal over many more ions. Therefore the ions observed at night cannot be clearly

identified during day. This does not, however, necessarily mean that the corresponding neutral molecules would not exist in the gas-phase during day-time, it only means that we cannot detect them with as high sensitivity as during night-time.

The good agreement between the JPAC α -pinene and the Hyytiälä ion spectra suggest that the ions are similar, and the identification can be done based on JPAC data. α -pinene is one of the most abundant monoterpenes in Hyytiälä, typically 4–5 times more abundant than β -pinene (Hakola et al., 2003), and based on the comparison between α -pinene and β -pinene in JPAC it seems to be especially efficient at producing molecules prone to form negative cluster ions. To make sure that the detected ions are the same, a more detailed plot of the internal patterns of the modes in Fig. 1c and d is presented in Fig. 2a and b. Both plots show unit mass-to-charge resolution data, and the spectral patterns agree extremely well, especially in the 300–400 Th region. Above 400 Th, both in Figs. 1 and 2, one can make out some differences, but the majority of the UMR patterns are largely consistent. Figure 2c and d depict some major peaks (UMR 340 and 342 Th in Fig. 2c, and 540 and 542 in Fig. 2d) in high resolution to show that also at a resolution of 3000–4000 these ions are the same in both spectra. There are peaks, like UMR 341, 541 and 543, that contain species that are visible in JPAC but not in Hyytiälä, and this will be discussed in more detail in the following sections. However, we conclude that the α -pinene negative ion spectrum in JPAC is largely representative of the typical night-time ions observed in Hyytiälä.

3.2 Identification of the elemental composition

To make a conclusive elemental identification of the α -pinene produced peaks using the maximum resolution of the instrument, the chamber was set to operate in steady state over one night, and the API-TOF was tuned to W mode with a resolution above 8000 for peaks above 300 Th. A 6h average was acquired to achieve good signal-to-noise, and 16 of the largest and “cleanest” peaks were used to calibrate the mass axis. By “clean”, we refer to the peaks which had the lowest full widths at half maximum (FWHM), in contrast to broader peaks that most likely contained several overlapping ions and were thus discarded. Gaussian peaks were fit to the 16 ions, and the mass axis was fit onto the time-of-flight axis using the 3-parameter fit described previously. The errors between the fitted Gaussians and the 3-parameter fit are plotted in Fig. 3. The dashed grey lines depict ± 10 ppm accuracy, and clearly all proposed ions are within this range. As an example, we will take a closer look at UMR 340 Th, which was used for the calibration. At this mass/charge, 10 ppm corresponds to 0.0034 Th, and the proposed ion is $C_{10}H_{14}NO_{12}$ with a mass of 340.0521. Within the range 340.0521 ± 0.0034 Th, there are only 18 different ions comprised of C, H, N and O. However, out of these 18 ions, most can be discarded

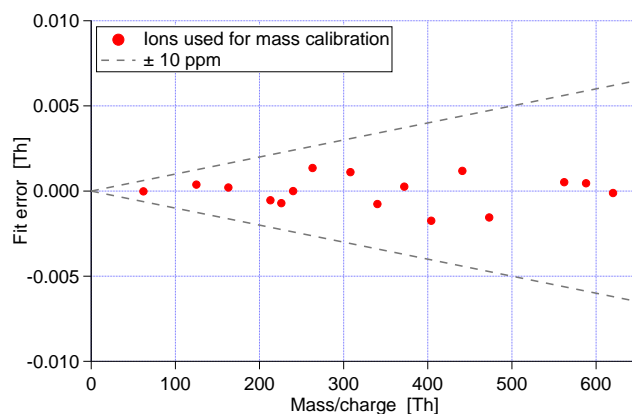


Fig. 3. Fit errors of the ions used for mass axis calibration. Gaussian fits were made to 16 peaks in the W-mode ion spectrum from JPAC. These peaks were attributed to specific elemental compositions and used for mass calibration. The selected peaks contained nitric acid, fluorinated acids, and organics to span the entire mass range of interest. The resulting errors between the fit masses and the assumed elemental compositions (red dots) were all well within ± 10 ppm accuracy (grey dashed lines), strongly suggesting that all chosen peaks were correctly identified.

as unreasonable based on their elemental compositions: 13 of the molecules have double bond equivalents higher than 8, and of the remaining 5, two have more than 3 nitrogen atoms. Discarding all the aforementioned ions, the remaining ions are $C_{10}H_{14}NO_{12}^-$, $C_7H_{16}O_{15}^-$, $H_{20}O_{20}^-$. The last two have unreasonably high amounts of oxygen, and additionally defy the “nitrogen rule” which states that a deprotonated ion at an even mass should contain an odd number of nitrogen atoms. Thus we are left with only the suggested $C_{10}H_{14}NO_{12}^-$ ion. As the precursor in the experiment was $C_{10}H_{16}$, the identification seems appropriate.

As yet another approach to test the elemental composition, one can look at the isotopic distributions of the ions. Keeping UMR 340 Th as an example, Fig. 4 shows data from Hyytiälä together with both V and W mode data from JPAC. In the left-hand graph (Fig. 4a), Gaussian fits were made to the three data sets. The peaks are scaled to not overlap, and thus the signal is in arbitrary units. However, the difference in peak resolution is still real, and one can see that the JPAC API-TOF was tuned to a slightly higher resolution (i.e. lower FWHM) than the one in Hyytiälä in V mode. The JPAC W mode resolution was still almost twice that of the JPAC V mode. The Gaussian fits provide peak areas corresponding to ion concentrations. Assuming that the ion signals at UMR 340 Th are purely from $C_{10}H_{14}NO_{12}^-$, one can calculate how much of the signal at UMR 341 Th is due to the isotopes of this ion. The measured spectra at UMR 341 Th (solid lines), together with the calculated isotopes from UMR 340 Th (dashed lines) are plotted in Fig. 4b. In Hyytiälä almost all of UMR 341 Th is explained by the isotope, and as there is some signal at almost every integer mass,

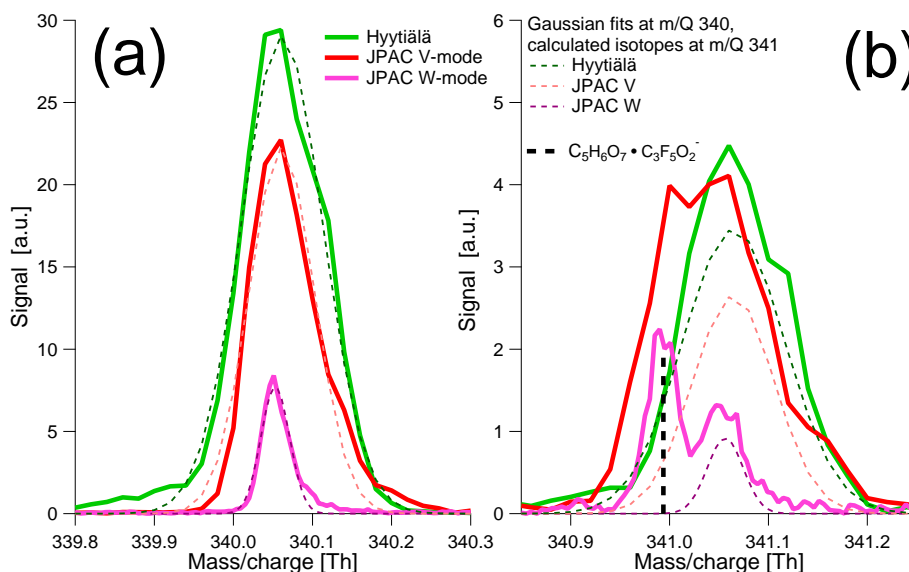


Fig. 4. Three different spectra, ambient night-time from Hyttiälä (green) and α -pinene chamber from JPAC in V-mode (red) and W-mode (purple), at UMR 340 (a) and 341 Th (b). Dashed lines in (a) depict Gaussian fits to the measured data (solid lines). Assuming that the peak at 340 Th is $C_{10}H_{14}NO_{12}^-$, the corresponding isotopes at 341 Th are plotted as dashed lines in (b). In Hyttiälä the peak at UMR 341 Th is almost completely explained by the isotope, whereas there is an interfering ion observed in JPAC. This ion was attributed to $C_5H_6O_7 \cdot C_3F_5O_2^-$.

the fact that the measured UMR 341 in Hyttiälä is larger than the isotope, is to be expected. The only way for the isotopic ratio to disprove the identification of UMR 340 Th would be if the isotopic fit would be larger than the actual signal. For the JPAC data, there is clearly some other ion at UMR 341 Th, and with the high resolution of the W mode measurement, the two peaks separate. Again, the isotopic ratio matches the signal very well. The other ion at UMR 341 has been identified as $C_5H_6O_7 \cdot C_3F_5O_2^-$, i.e. a mixed cluster of an organic molecule produced from α -pinene, and a fluorinated acid. This identification will be discussed in more detail in the next section. In the V mode spectrum, this ion causes a shift of the peak towards lower masses, which was already visible in Fig. 2c.

Based on the exact mass, isotopic distribution, and the fact that the precursor is $C_{10}H_{16}$, we now conclude that the elemental composition of UMR 340 is $C_{10}H_{14}NO_{12}^-$ as was originally speculated by Ehn et al. (2010) although this could not at that time be conclusively determined from their data. Also the other ions in the 300–400 Th range are now identified as C_{10} compounds with one N, and O ranging from 10 to 16.

3.3 Identification of the ions as clusters

Mass spectrometric methods (with the exception of MS-MS coupled systems) can typically only provide elemental identification and not any molecular information. We could, for example, not normally determine how the N atom is incorporated in the detected ions. However, in this case, more in-

formation is available. The pattern of organics around 300–400 Th in JPAC is in fact duplicated, although at lower intensity, with a shift of 101 Th in the 400–500 Th range. The mass difference corresponds to removal of NO_3^- (UMR 62) and addition of $C_3F_5O_2^-$ (UMR 163). As an example, Fig 4b showed a peak at 341 corresponding to $C_5H_6O_7 \cdot C_3F_5O_2^-$, and as can be made out in Fig. 1c, there is a lone peak at UMR 240 Th which has been identified as $C_5H_6O_7 \cdot NO_3^-$. It is believed that α -pinene in JPAC is oxidized by ozone and possibly other radicals formed from the reactions of α -pinene and ozone, followed by clustering of the highly oxidized multifunctional organic compounds (henceforth abbreviated as HOMs) with the existing NO_3^- , $C_3F_5O_2^-$, and some other fluorinated acids. Artificially increasing the amount of HNO_3 in JPAC leads to an increase of the ions from clusters of NO_3^- with the HOMs, while decreasing the corresponding clusters with $C_3F_5O_2^-$.

Additional evidence for the clustering can be found from Hyttiälä, when comparing the ion spectra during nights with high and low H_2SO_4 . As an example we will take a closer look at the nights of July 28 and 29, 2010. The first night was a “typical” summer night with northerly wind and fairly clean conditions, whereas by the second night, the wind direction had turned almost southerly, with increases in gas phase pollutants such as CO and SO_2 . The increase in SO_2 also caused an increase in gaseous H_2SO_4 , which is detectable as an increase in both HSO_4^- and $H_2SO_4 \cdot HSO_4^-$.

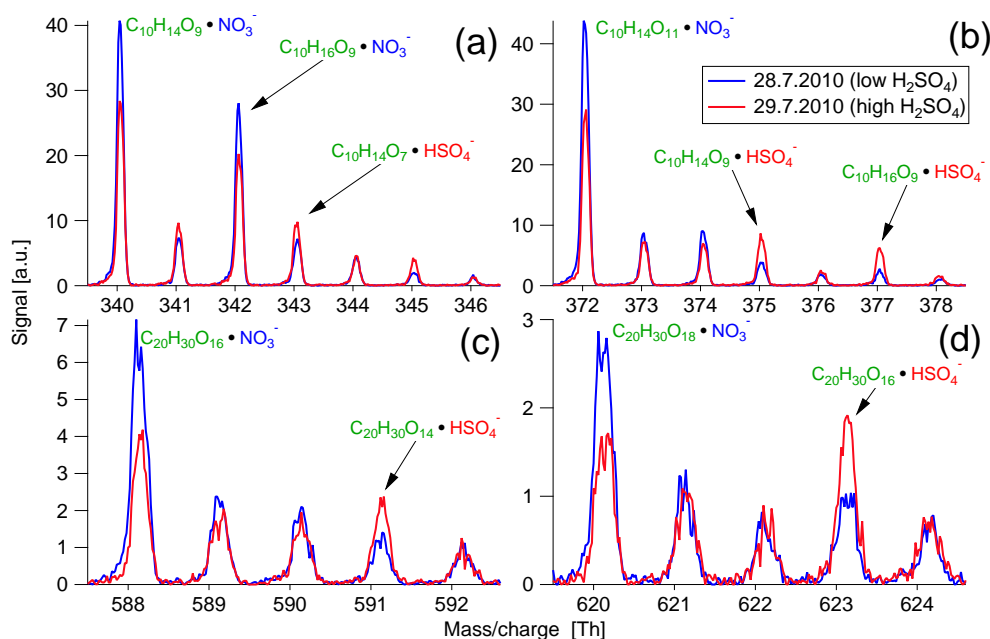


Fig. 5. Comparison of mass spectra during two consecutive nights, where the first was a “normal” night with low SO_2 and H_2SO_4 , but the second had unusually high levels of both. The four subplots show different parts of the ion spectrum, with the aim to show how $\text{HOM}\cdot\text{HSO}_4^-$ clusters increase at the expense of $\text{HOM}\cdot\text{NO}_3^-$ clusters, e.g. UMR 375 Th ($\text{C}_{10}\text{H}_{14}\text{O}_9\cdot\text{HSO}_4^-$) increases while UMR 340 Th ($\text{C}_{10}\text{H}_{14}\text{O}_9\cdot\text{NO}_3^-$) decreases.

in the APi-TOF spectra. In the 300–650 Th range, there was an accompanying change in the concentration of certain ions. The majority of this shift is believed to be due to HSO_4^- displacing NO_3^- in the clusters with the organics, causing a shift of 35 Th in the spectrum. Some example spectra from the two nights are shown in Fig. 5. In Fig. 5a the large peaks at UMR 340 and 342 Th, corresponding to $\text{C}_{10}\text{H}_{14}\text{O}_9\cdot\text{NO}_3^-$ and $\text{C}_{10}\text{H}_{16}\text{O}_9\cdot\text{NO}_3^-$, both decrease. Simultaneously UMR 343 (corresponding to $\text{C}_{10}\text{H}_{14}\text{O}_7\cdot\text{HSO}_4^-$) increases, while the corresponding cluster with NO_3^- at UMR 308 Th decreases (not shown). In Fig. 5b, UMR 372 ($\text{C}_{10}\text{H}_{14}\text{O}_{11}\cdot\text{NO}_3^-$) drops, whereas UMR 375 and 377 Th (corresponding to the ions at 340 and 342 Th in Fig. 5a with displacement of NO_3^- by HSO_4^-) increase.

Although lower signals, also the ions at 500–650 Th show this same feature. These ions typically contain 16–20 C atoms, and are thus believed to be some form of dimers of HOMs typically found in the 300–400 Th range. The dimers typically have slightly lower O/C ratios than the monomers. Ions identified as clusters with NO_3^- , such as UMR 588 and 620 Th, decrease while e.g. UMR 623 (588 + 35) Th increases (Fig. 5c and d). The signal-to-noise for these peaks is not good enough to provide conclusive identification of the ions at the highest masses, but the high-resolution data certainly do not contradict the suggested explanation. Also other ions that were not shown in Fig. 5 follow the pattern marked out above.

In summary, by using all the above-mentioned methods, a number of ions were identified. As seen already in Fig. 1, the most abundant ions are observed around 300–400 Th, but an additional group is clearly visible at 500–650 Th. Based on the elemental composition, these two groups are termed “monomers” and “dimers” and the most abundant molecules in both groups are listed in Appendix A (Tables A1 and A2).

4 Discussion

4.1 Possible formation mechanisms

The high level of oxidation of α - and β -pinene ozonolysis products that is seen in Table A1 and A2 has, to our knowledge, not been previously measured nor predicted in the gas phase. Indeed, the mechanism forming these molecules is still unclear, and we can only draw some preliminary conclusions based on the available data.

The HOMs need to be formed in the gas phase, and not through particle phase reactions, as the TSI 3025 condensation particle counter (with a d_{50} cut-off diameter of roughly 3 nm) always showed particle counts $< 10\text{ cm}^{-3}$ during the experiments in JPAC. Furthermore, the HOMs observed in this work are expected to have such low vapor pressures that they will not re-evaporate from a particle or surface, suggesting that chamber wall effects can also be discarded.

While β -pinene produced mainly C_9 compounds, α -pinene produced mostly C_{10} compounds. This is a result of

the first reaction step where ozone attacks the double bonds in the pinene molecules. In α -pinene the double bond is part of a ring, and thus merely breaking that bond will not cause fragmentation. On the other hand, when ozone attacks the exocyclic double bond in β -pinene, one C is inevitably lost. These ozone reactions also produce Criegee or OH radicals, which can play a role in the oxidation processes leading up to some of the observed HOMs.

Other HOM compounds have previously been identified in SOA samples from chamber oxidation experiments of α -pinene (Claeys et al., 2007; Warscheid and Hoffmann, 2002; Iinuma et al., 2004). Although less oxidized than the HOMs found in this study, the compounds proposed by the authors had relatively high O/C (typically C₉ and C₁₀ compounds with up to 5 or 6 O atoms). The level of oxidation was achieved mainly by carboxylic acid groups, and in one case also by ester formation (Warscheid and Hoffmann, 2002). HOMs are also formulated in the Master Chemical Mechanism MCM (Jenkin et al., 1997; Saunders et al., 2003) where they contain functional groups like hydroperoxides (-OOH), carboxyls (-COOH), peroxy-carboxyls (-C(O)OOH), and carbonyl groups (>C=O). For the observed large degree of oxygenation at high hydrogen atom contents, eventually hydroperoxide (-OOH) or hydroxyl (-OH) groups are needed, as solely increasing the number of carbonyl or carboxyl groups will increase the O/C as desired but decrease the H/C.

Aqueous reactions in aerosol particles have been shown to produce high molecular weight HOMs with O/C > 1 through oligomerization of C₂ and C₃ compounds, but these HOMs comprised a wide range of carbon numbers, also exceeding C₁₀ (Lim et al., 2010). Although fragmentation is a possibility at each oxidation step, it is highly unlikely that the gas-phase HOMs observed in this work would be clusters or oligomers of fragmentation products, as most of them contain the same amount of C as the precursor molecules. As in the case of the reactions reported by Lim et al. (2010), oligomerization would produce a variety of different carbon numbers. In our work, no molecules with 11 or more C were identified in the 300–400 Th range, though some probable fragments were detected between 200 and 300 Th. Specifically C₅H₆O₇·NO₃⁻ detected at UMR 240 Th was very common, and can be distinguished in Fig. 1c and 1d. The O/C ratio of this molecule is as high as 1.4, but the H/C is fairly low, suggesting that this molecule could be a tri-carboxylic acid. Also the detected C₇, C₈ and C₉ oxidation products from α -pinene can be formed from oxidation reactions splitting off one or two terminal C atoms, or e.g. acetone (C₃H₆O).

As the molecules are detected in the natural negative ion spectrum, they are expected to contain carboxylic groups, though the favorable cluster formation with NO₃⁻ can possibly also be due to other, unknown structures. With all the O atoms in carboxylic acid groups, the number of H atoms becomes lower than observed in the ions. Although this paper does not attempt to provide any molecular structures for

the observed HOMs, we can speculate over which functional groups that could produce the high O/C and H/C ratios. One structural element that fits the criteria is the geminal diol, formed through the addition of water to the carbonyl groups,



This reaction is known to occur in aqueous environments such as cloud droplets, but has also recently been observed to take place in the gas-phase for the cases of methylglyoxal (Axson et al., 2010) and glyoxylic acid (Plath et al., 2009). Another structural element that can add O without decreasing the H atom content is the hydroperoxide group. Both model and experimental studies have found organic peroxides to be important contributors to SOA (e.g. Reinnig et al., 2009; Chen et al., 2011). It can also not be ruled out that the HOM·NO₃⁻ clusters can contain non-covalently bound water molecules, but as mentioned previously, water molecules have not been observed at any significant amounts with any other ions. The main peaks in our spectra are typically separated by 16 Th (O) and not 18 Th (H₂O), suggesting that water clusters are not able to explain the high oxygen content of the HOMs. Additionally, if indeed water molecules clustered efficiently with the HOMs, one would still need to remove so many H₂O molecules that one would reach unreasonably low H/C ratios before achieving more commonly measured O/C ratios.

The molecules in the 500–650 Th range were termed dimers due to the number of carbon atoms in the molecules. The main peaks have 16–20 C atoms suggesting that they are products of two oxidized pinene molecules. The formation mechanism of these dimers is still unknown, however, the dimers can typically not be considered simply non-covalently bonded clusters of two of the observed monomers, as the dimer compositions can rarely be found by just adding together two monomers. Thus, the current hypothesis is that these are large molecules formed from reactions of two oxidized monomers, possibly undergoing further oxidation after the dimerization. Less oxidized HOM dimers have been reported in SOA, with an often discussed product at C₁₇H₂₆O₈ (Hoffmann et al., 1998; Yasmeen et al., 2010). As listed in Table A2, we detected C₁₇H₂₆O_n, with *n* in the range 11–16.

4.2 Estimation of neutral HOM concentrations

Naturally charged ions can only in very specific cases be used to calculate the concentration of neutral species. The absolute concentration of an ion, e.g. NO₃⁻, tells almost nothing about the concentration of the corresponding neutral molecules, e.g. HNO₃. This is because the production and loss of charges in the atmosphere depend mainly on properties that do not relate to atmospheric trace gas concentrations (e.g. cosmic radiation and radon concentration for ion production, and aerosol particle surface area acting as an ion sink). However, though NO₃⁻ does not track

the HNO_3 concentration, it was found during the measurements in Hyytiälä that the nitric acid dimer-to-monomer ratio ($\text{HNO}_3 \cdot \text{NO}_3^- / \text{NO}_3^-$) did correlate with HNO_3 , suggesting that for some species the natural ion spectrum can be used to for estimates of neutral concentration.

In the specific case of sulfuric acid, Eisele (1989b) estimated $[\text{H}_2\text{SO}_4]$ from ion spectra using a few simple assumptions. Firstly, HSO_4^- was mainly formed from $\text{NO}_3^- + \text{H}_2\text{SO}_4$, with a known rate coefficient (Viggiano et al., 1982). Secondly, the rate coefficient for the loss of HSO_4^- could be calculated as the reciprocal of the ion lifetime, which was assumed to be 100 s, since HSO_4^- is considered a “terminal” ion that will not be lost due to charge transfer to stronger acids. A similar calculation for any other, “non-terminal” ion species becomes more complicated. For example malonic acid is a stronger acid than HNO_3 in the gas phase, but still weaker than H_2SO_4 (Ehn et al., 2010). For this estimation, one would require knowledge of the rate coefficient of malonic acid with NO_3^- , but also the loss rate by reaction with H_2SO_4 needs to be known. For weaker acids the calculation becomes increasingly more complex as more and more reaction rates and concentrations need to be estimated.

With similar methods as described above we group all the HOMs observed in the ion spectra in Hyytiälä, and will attempt to estimate the total neutral concentration. A prerequisite for this is that we are in a low-OH regime, which means that the concentrations of strong acids such as H_2SO_4 , HNO_3 and malonic acid are low enough not to dominate the spectrum, but HNO_3 is high enough to form NO_3^- ions for clustering with the organics. We assume that the $\text{HOM} \cdot \text{NO}_3^-$ clusters are formed by collision of NO_3^- with the HOM. The loss of the ion clusters occurs through ion-ion recombination, coagulation with aerosol particles, and charge transfer to stronger acids. In steady-state the formation and loss are equal, and we get

$$k_1[\text{NO}_3^-][\text{HOM}] = \left(k_{\text{rec}} + k_{\text{coag}} + \sum_i k_{\text{CT}}^i \right) [\text{HOM} \cdot \text{NO}_3^-], \quad (3)$$

where k_1 is the rate coefficient of the forward reaction $\text{NO}_3^- + \text{HOM} \rightarrow \text{HOM} \cdot \text{NO}_3^-$, k_{rec} is the rate coefficient of loss by ion-ion recombination, k_{coag} by coagulation, and k_{CT}^i correspond to loss due to charge transfer to the stronger acid i . The coagulation sink is a function of the aerosol size distribution, and for this calculation it is implicitly included in k_{coag} . Here HOM refers only to the group of extremely oxidized molecules detected in the ion spectra by clustering with NO_3^- . Solving for $[\text{HOM}]$ yields

$$[\text{HOM}] = \frac{\left(k_{\text{rec}} + k_{\text{coag}} + \sum_i k_{\text{CT}}^i \right) [\text{HOM} \cdot \text{NO}_3^-]}{k_1[\text{NO}_3^-]}. \quad (4)$$

To estimate a lower limit for $[\text{HOM}]$, we will use lower limit estimates for the numerator and upper limit estimates for the

denominator in Eq. (4). To minimize the loss rate coefficients, we will assume that the loss is the reciprocal of the ion lifetime, and assume an upper range lifetime estimate of 10 min. For comparison, Eisele (1989b) used an ion lifetime of 100s for similar calculations, and Dal Maso et al. (2002) reported average condensation sink values in Hyytiälä corresponding to (neutral) vapor lifetimes of 140–250 s. Thus,

$$k_{\text{lifetime}} \equiv \left(k_{\text{rec}} + k_{\text{coag}} + \sum_i k_{\text{CO}}^i \right) \approx (10 \text{ min})^{-1} \quad (5)$$

$$= 0.0016667 \text{ s}^{-1}.$$

For k_1 we will use the collision limit for an upper limit of possible reaction rates ($3 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$).

Finally, the ratio of the clusters to pure NO_3^- is estimated from the experimental data by summing over all the identified $\text{HOM} \cdot \text{NO}_3^-$ ions and dividing by the NO_3^- signal, yielding $[\text{HOM} \cdot \text{NO}_3^-] / [\text{NO}_3^-] \approx 5$, which again is a lower range estimate. This then yields a lower limit of the concentration of the highly oxidized organics:

$$[\text{HOM}] \approx \frac{k_{\text{lifetime}} [\text{HOM} \cdot \text{NO}_3^-]}{k_1 [\text{NO}_3^-]} \approx \frac{0.0016667 \text{ s}^{-1} \cdot 5}{3 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}} \quad (6)$$

$$= 2.78 \times 10^6 \text{ cm}^{-3} \approx 0.1 \text{ ppt}.$$

A corresponding upper limit estimate is harder to calculate, mainly due to the problems in constraining the rate coefficients. Collision probabilities and steric factors depend on the exact molecules under consideration, and can have a large influence on k_1 . However, reasonable changes in k_1 , k_{lifetime} and $[\text{HOM} \cdot \text{NO}_3^-] / [\text{NO}_3^-]$ can easily yield concentrations above 1ppt. Fragmentation of $\text{HOM} \cdot \text{NO}_3^-$ clusters to pure NO_3^- inside the instrument would also lead to an underestimation of the concentration. On the other hand, the exact mass/charge dependent transmission of the instrument was not determined during the measurements, and could have an influence on the $[\text{HOM} \cdot \text{NO}_3^-] / [\text{NO}_3^-]$ ratio in either direction, though a decrease is more likely. Thus, our best estimate of the typical abundance of the extremely oxidized organics at night in the boreal forest is between 0.1 and a few ppt.

As a comparison, Ehn et al. (2010) found that the negative ion distribution in Hyytiälä was practically completely dominated by H_2SO_4 -related ions when the gaseous H_2SO_4 concentration reached 1 ppt. On the other hand, during nights when the H_2SO_4 dropped to below $\sim 10^5 \text{ cm}^{-3}$ (roughly 0.01 ppt), the HSO_4^- contribution to the ion spectrum was on the order of a percent since the ion lifetimes were not long enough to allow collisions of the ions with H_2SO_4 molecules. Replacing $\text{HOM} \cdot \text{NO}_3^-$ with HSO_4^- in all the calculations above would thus have given the correct order of magnitude estimates for $[\text{H}_2\text{SO}_4]$. As our night-time spectra were fairly dominated by the $\text{HOM} \cdot \text{NO}_3^-$ clusters, it is fair to assume that the concentrations are in the 10^6 – 10^7 cm^{-3} (~ 0.1 – 1 ppt) range, or possibly even higher.

4.3 Atmospheric relevance

Assuming the sum of the HOMs is around 1 ppt, this is higher than e.g. the typical concentration of H_2SO_4 in the boreal forest (Petäjä et al., 2009), and while the mass of the molecules are roughly 3 times larger (even more for the dimers) than that of H_2SO_4 , their role in aerosol mass formation alone could be of importance. In addition to this, the role in nucleation and initial growth of nanometer-sized clusters needs to be further investigated. The vapor pressures of the molecules, due to their size and high degree of oxidation, are expected to be extremely low. As previously stated, in the dark ozonolysis experiments, no particle formation occurred although the monomers were detected. However, the low-range cut-off of the particle counter used for particle detection was 3 nm, which does not rule out some cluster formation taking place, without the clusters having time to grow large enough before being flushed out of the chamber.

The gas-phase dimers observed in this study are probably clearly lower in abundance than the oxidized monomers, but due to their large sizes may also be very efficient at condensing onto small newly formed particles, or even taking place in the initial formation.

It is possible that these organics are present also during day-time, and this is an important point to study in the future. If the molecules are only abundant at night, they might still be relevant for SOA (mass) formation. Although rare, and with growth usually ending below 10 nm, nocturnal NPF events also occur in Hyytiälä (Junninen et al., 2008), and the highly oxidized organics may play an important role in these events. Furthermore, these compounds could possibly explain some of the night-time high concentrations of sub-3 nm particle concentrations observed with a pulse-height condensation particle counter (Lehtipalo et al., 2011; Sipilä et al., 2008).

5 Conclusions

A previously unidentified group of highly oxidized multifunctional molecules (HOMs) were observed in the naturally charged negative ion spectra in both ambient air (Hyytiälä) and during α - and β -pinene chamber ozonolysis experiments. These HOMs retained practically all of their original 10 carbon atoms, while reaching O/C ratios of 0.7–1.3. Through dedicated experiments in the Jülich plant atmosphere chamber, the elemental composition of these organics was conclusively determined. Both the exact mass and isotopic ratios in high-resolution spectra were used for the identification. Furthermore, the ions could be identified as clusters of HOMs and deprotonated inorganic acids (mainly HNO_3 , but also $\text{C}_3\text{F}_5\text{HO}_2$ and H_2SO_4).

The concentration of neutral HOMs could only be estimated using several assumptions. Based on a rough estimation, the concentration of these extremely oxidized molecules

in the night-time boreal forest is typically on the order of 0.1–1 ppt. The yield of these molecules at higher OH concentrations is yet to be determined, but this cannot be done from the ion spectra due to the signal spreading out over many more peaks, likely due to the formation of many more compounds that compete for the limited amount of charges.

Even at concentrations of 0.1–1 ppt, these molecules may have an important impact on aerosols due to their presumably low vapor pressures. Mainly the first steps of new particle formation through nucleation and growth may be influenced by these compounds, as they are present at similar concentrations as H_2SO_4 which is widely accepted to take part in nucleation of new particles, but already have roughly 3 times the mass. Furthermore, the existence of such highly oxidized species may suggest that some of the observed missing OH reactivity could be explained by VOC oxidation products reacting several times with OH.

The exact formation mechanism of the HOMs is still unknown. However, the oxidation most likely takes place in the gas phase and we hypothesize that formation of carbonyl groups, followed by water addition forming carbonyl hydrates (geminal diols), may be involved. Further work will include trying to quantitatively measure the molecules using chemical ionization techniques under different conditions. If successful, this should provide a better picture of both what drives their formation, and of their atmospheric importance.

Appendix A

Peak list

In this appendix we list the most abundant HOMs detected in this study. As all of the molecules were mainly detected in clusters with NO_3^- , this is omitted and the first column in both tables lists the elemental composition of the organic molecules that were found in the clusters. The list is not exhaustive, but does contain the most abundant detected molecules. The typically largest peaks in Hyytiälä and in JPAC α -pinene experiments are highlighted in bold in Table 1. The second column of the table shows the mass-to-charge at which the ion clusters were detected, the third the mass of the HOM in the cluster, and columns 4–6 show where this molecule has been detected.

Table A1. List of observed monomers detected in Hyytiälä and JPAC. An “X” means that the identification is fairly unambiguous, a blank means that the ion has not been detected, and “/” means that the ion might be present, but could not be identified due to insufficient signal or overlapping peaks from other compounds.

Elemental formula			Ion mass-to-charge (clustered with NO ₃ ⁻) [Th]	HOM mass [Da]	Detected in		
					Hyytiälä	JPAC α -pinene	JPAC β -pinene
C ₇	H ₁₀	O ₄	220.046	158.058		X	X
C ₅	H ₆	O ₆	224.005	162.017	X	X	X
C ₅	H ₆	O ₇	240.000	178.012	X	X	X
C ₇	H ₈	O ₈	282.010	220.022		X	\
C ₈	H ₁₂	O ₇	282.047	220.059	X	X	\
C ₈	H ₁₂	O ₈	298.042	236.054	X	X	\
C ₁₀	H ₁₄	O ₇	308.062	246.074	X	X	\
C ₉	H ₁₂	O ₈	310.042	248.054			X
C ₁₀	H ₁₆	O ₇	310.078	248.090		X	
C ₈	H ₁₂	O ₉	314.036	252.048		X	X
C ₁₀	H ₁₄	O ₈	324.057	262.069		X	
C ₁₀	H ₁₆	O ₈	326.073	264.085	\	X	
C ₉	H ₁₂	O ₉	326.036	264.048	\	X	X
C ₉	H ₁₄	O ₉	328.052	266.064		X	\
C ₁₀	H ₁₄	O ₉	340.052	278.064	X	X	
C ₉	H ₁₂	O ₁₀	342.031	280.043	\		X
C ₁₀	H ₁₆	O ₉	342.068	280.080	X	X	\
C ₁₀	H ₁₄	O ₁₀	356.047	294.059			
C ₁₀	H ₁₆	O ₁₀	358.063	296.075	X	X	
C ₉	H ₁₂	O ₁₁	358.026	296.038			X
C ₁₀	H ₁₄	O ₁₁	372.042	310.054	X	X	X
C ₁₀	H ₁₆	O ₁₁	374.058	312.070	X	X	
C ₉	H ₁₂	O ₁₂	374.021	312.033			X
C ₁₀	H ₁₄	O ₁₃	404.032	342.044	X	X	

Table A2. List of observed dimers detected in Hyytiälä and JPAC. An “X” means that the identification is fairly unambiguous, a blank means that the ion has not been detected, and “/” means that the ion might be present, but could not be identified due to insufficient signal or overlapping peaks from other compounds.

Elemental formula			Ion mass-to-charge (clustered with NO ₃ ⁻) [Th]	HOM mass [Da]	Detected in		
					Hyytiälä	JPAC α -pinene	JPAC β -pinene
C ₁₇	H ₂₆	O ₁₁	468.136	406.148			X
C ₁₈	H ₂₆	O ₁₁	480.136	418.148			X
C ₁₄	H ₂₀	O ₁₅	490.069	428.081	X	X	
C ₁₉	H ₂₈	O ₁₁	494.152	432.164	X	X	
C ₂₀	H ₃₂	O ₁₁	510.183	448.195	\	X	
C ₁₇	H ₂₆	O ₁₄	516.121	454.133			X
C ₂₀	H ₃₀	O ₁₂	524.162	462.174	X	X	
C ₁₉	H ₂₈	O ₁₃	526.141	464.153	X	\	
C ₁₈	H ₂₆	O ₁₄	528.121	466.133			X
C ₁₈	H ₂₈	O ₁₄	530.136	468.148		X	
C ₁₇	H ₂₆	O ₁₅	532.116	470.128			X
C ₂₀	H ₃₀	O ₁₃	540.157	478.169		X	
C ₂₀	H ₃₂	O ₁₃	542.173	480.185	\	X	
C ₁₇	H ₂₆	O ₁₆	548.110	486.122			X
C ₂₀	H ₃₀	O ₁₄	556.152	494.164	X	X	
C ₁₈	H ₂₈	O ₁₆	562.126	500.138		X	
C ₂₀	H ₃₀	O ₁₅	572.147	510.159		X	
C ₂₀	H ₃₂	O ₁₅	574.162	512.174	\	X	
C ₂₀	H ₃₀	O ₁₆	588.142	526.154	X	X	
C ₁₈	H ₂₈	O ₁₈	594.116	532.128		X	
C ₂₀	H ₃₀	O ₁₈	620.132	558.144	X	X	

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