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## Article

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# Global changes in oceanic mesoscale currents over the satellite altimetry record

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**Oceanic eddies play a profound role in mixing tracers such as heat, carbon, and nutrients, thereby regulating regional and global climate. Yet, it remains unclear how global oceanic eddy kinetic energy has evolved over the past few decades. Furthermore, coupled climate model predictions generally fail to resolve oceanic mesoscale dynamics, which could limit their accuracy in simulating future climate change. Here we show a global statistically significant increase of the eddy activity using two independent observational datasets of mesoscale variability, one directly measuring currents and the other from sea surface temperature. Regions characterized by different dynamical processes show distinct evolution in the eddy field. For example, eddy-rich regions such as boundary current extensions and the Antarctic Circumpolar Current show a significant increase of 2% and 5% per decade in eddy activity, respectively. In contrast, most of the regions of observed decrease are found in the tropical oceans. Because eddies play a fundamental role in the ocean transport of heat, momentum,**

21 **and carbon, our results have far-reaching implications for ocean circulation and climate, and**  
22 **the modelling platforms we use to study future climate change.**

23 Changes in the climate system over recent decades have warmed the upper ocean and modi-  
24 fied the wind stress, heat and freshwater fluxes that drive ocean circulation<sup>1,2</sup>. These changes have  
25 the capacity to modify the ocean circulation at all scales, including the overturning circulation<sup>3,4</sup>,  
26 basin-scale gyres<sup>5,6</sup>, boundary currents<sup>7,8</sup>, and the mesoscale<sup>9</sup>. The ocean's mesoscale incorpo-  
27 rates motions that occur at spatial scales from  $\sim 10$  to  $\sim 100$  km. These motions include both  
28 steady flows, such as jets and re-circulations, and time-varying flows, generally referred to as ed-  
29 dies. Mesoscale eddies are ubiquitous in the global ocean and feed back onto all scales, from  
30 regional processes<sup>10</sup> up to the meridional overturning circulation<sup>3</sup>. Moreover, these eddies act to  
31 transport and mix tracers such as heat, salt, and nutrients<sup>11,12</sup>. Thus, understanding the evolution  
32 of the mesoscale circulation is crucial to better predict our changing oceans.

33 Kinetic energy (KE) quantifies the magnitude of ocean currents<sup>9,13-15</sup>. Kinetic energy is  
34 proportional to the square of the velocity, and is commonly separated into the mean KE (MKE;  
35 computed from the time-mean velocity field) and the KE of the time-varying velocity (known as  
36 the Eddy Kinetic Energy; EKE). The EKE is dominated by mesoscale variability and is a signifi-  
37 cant fraction of the total KE<sup>16,17</sup>. A recent study has inferred a global increase of KE anomaly from  
38 ocean reanalyses and ARGO floats<sup>15</sup>. However, these reanalyses and observations do not have the  
39 spatial resolution required to resolve the mesoscale field. Satellite observations, which can resolve  
40 the mesoscale, suggest that EKE in the Southern Ocean has a robust increasing trend<sup>9,18,19</sup>. How-

41 ever, there have been no studies that have quantified the global evolution of mesoscale variability  
42 from satellite observations. This is the goal of the present study.

43 Mesoscale ocean flow has a footprint in both sea surface height (SSH) and sea surface tem-  
44 perature (SST). EKE can be directly inferred from SSH via geostrophy, and mesoscale turbulence  
45 also acts to strain and shear the temperature field, meaning that regions of high EKE are associated  
46 with strong SST gradients. Therefore, observed gradients in SST can be considered a proxy for  
47 mesoscale variability<sup>20–22</sup>.

48 In this study, we examine the mesoscale evolution of the ocean from satellite observations of  
49 sea surface height (SSH) and sea surface temperature (SST) over the available satellite altimetry  
50 record (1993 - 2019). We use two independent datasets, namely AVISO+ SSH and NOAA optimal  
51 interpolated sea surface temperature (OISST)<sup>23</sup>, to estimate EKE and SST gradients respectively.  
52 These fields are then temporally smoothed using a running average of 12 months; the trends and  
53 the significance of each field are computed using a linear regression and a modified Mann–Kendall  
54 test<sup>24</sup> (see Methods for further details). We explore the mesoscale time-series for the global ocean,  
55 for individual ocean basins, and over regions characterized by their dynamical properties.

## 56 **Surface ocean acceleration**

57 Over recent decades, sea water thermal expansion and melting of land ice have led to an  
58 increase of SSH<sup>1,25</sup> (Fig. 1a). This SSH increase can be observed in all ocean basins, but there is  
59 also regional variability (Fig. 1b). SSH gradients are proportional to the surface geostrophic flow,  
60 from which we can compute velocity anomalies and Eddy Kinetic Energy (see Methods). The time-

61 mean EKE highlights mesoscale-rich regions including boundary currents and their extensions, the  
62 Antarctic Circumpolar Current (ACC), and the equatorial band (Fig. 1d). More importantly, these  
63 oceanic mesoscale-dominated sectors show significant trends over the satellite altimetry record  
64 1993-2019 (Fig. 1f; and contours in Fig. 1d), which suggest a long-term adjustment of the global  
65 ocean mesoscale field. In particular, the surface globally-integrated EKE trend is positive  $\sim 0.56\%$   
66 per decade (95% confidence level). The spatial structure of EKE trends is highly heterogeneous,  
67 although its zonal average shows some significant net tendencies, with increasing trends polewards  
68 of  $25^\circ\text{S}$  and  $40^\circ\text{N}$  (Fig. 1e,f). A strengthening of the EKE field is a direct indication of an increase  
69 in mesoscale activity.

70 Sea Surface Temperature (SST) is an independent dataset relative to SSH, but is also influ-  
71 enced by mesoscale activity and has better temporal and spatial resolution than SSH. SST has  
72 increased on multi-decadal timescales due to climate change<sup>26,27</sup> (Fig. 2a), with the global spatial  
73 pattern also modulated by interdecadal climate variability. This increase is spatially heterogeneous,  
74 with positive trends found in western boundary current regions<sup>7</sup> and over much of the subtropical  
75 gyres, and decreasing trends in the southeast Pacific, the far north Atlantic, and parts of the South-  
76 ern Ocean (Fig. 2b). In contrast, the largest time-mean SST gradients are located in dynamic  
77 mesoscale regions, such as boundary currents, their extensions, and the ACC (Fig. 2d). Regions  
78 with large SST gradients also exhibit the largest positive SST gradient trends, while the subtropical  
79 gyres and the tropics mostly exhibit a decreasing trend (Fig. 2e-f; contours in 2d). The globally  
80 integrated SST gradient magnitude shows a significant increasing trend of  $0.42\%$  per decade ( $5.76$   
81  $\pm 2.46 \times 10^6 \text{ }^\circ\text{C m decade}^{-1}$ ; 95% confidence level) relative to the time-mean globally integrated

82 SST gradient magnitude ( $1.36 \times 10^9 \text{ }^\circ\text{C m}$ ). Moreover, mesoscale eddies enhance the strength of  
83 SST gradients on small length scales, and further analysis (Fig. S3) shows that mesoscale SST  
84 gradients, characterized by features with horizontal length scales smaller than  $3^\circ$ , dominate the  
85 observed trends. The globally integrated mesoscale SST gradient magnitude has increased at a  
86 rate of  $6.13 \pm 2.40 \times 10^6 \text{ }^\circ\text{C m decade}^{-1}$ ; significant at the 95% confidence level. The observed  
87 SST gradient signal appears to be an intrinsic response to the mesoscale field, consistent with the  
88 observed patterns of trends of the mesoscale EKE variability.

89 Eddy kinetic energy and mesoscale SST gradients show analogous spatial and temporal re-  
90 sponses in the boundary currents and their extensions, the ACC and the tropics. Note that eddy-rich  
91 regions such as the Kuroshio, the Agulhas, the Gulf Stream, and the ACC show large changes in  
92 mesoscale SST gradients co-located with some of the largest EKE increases (Fig. 3). Even though  
93 we do not expect the fields to match perfectly, the areas with same-sign trends match to a good  
94 extent (Fig. S4). The similarity between these two independent satellite products reaffirms the  
95 intrinsic response of the mesoscale processes to a changing and variable climate.

## 96 **Spatial patterns of ocean mesoscale trends**

97 Mesoscale EKE and SST gradient trends both indicate a net strengthening of the global  
98 mesoscale. However, both datasets reveal heterogeneous patterns of trends. To further understand  
99 the spatial variability, we first focus our analysis on individual area weighted regions: namely,  
100 the Southern Ocean (south of  $35^\circ\text{S}$ ), and the Pacific, Indian, and Atlantic Oceans north of  $35^\circ\text{S}$   
101 (Fig. 4d). This analysis reveals that the Southern Ocean and the Pacific Ocean are, to a large

102 extent, responsible for the globally integrated trends and variations of EKE and mesoscale SST  
103 gradients; the trends in the Indian and Atlantic Oceans are in contrast much smaller (Fig. 4a,b).  
104 The Southern Ocean shows a significant increase for both the EKE and SST gradient, however,  
105 the Pacific Ocean SST gradient decreases significantly, while the EKE signal also decreases (albeit  
106 below 95% significance level; Fig. 4c,e). The large uncertainty in the Pacific EKE trend (orange  
107 error bars in Fig. 4c) is a consequence of the pulses in the time series during 1997 and 2015, both  
108 being El Niño onset years. These large anomalous interannual signals dominate the uncertainty of  
109 the global EKE trend, so that the global trends reported here are only just statistically significant  
110 at the 95% confidence level.

111 Over the satellite altimetry record, the strongest El Niño events occurred during 1997-1998  
112 and 2015-2016<sup>28,29</sup> (gray bars in Fig. 4a,b). These El Niño events correlate with large pulses in the  
113 Pacific EKE signal, resulting in large uncertainty in the calculated global trends (Fig. 4c). El Niño  
114 events are associated with a strengthening of the North Equatorial Countercurrent and the northern  
115 branch of the South Equatorial Current particularly during extreme eastern Pacific El Niños, such  
116 as occurred during 1997-1998 and 2015-2016<sup>30</sup>. During such El Niño events, the equatorial cur-  
117 rents generate significant transient circulation anomalies extending over the equatorial band (9°N  
118 - 9°S). After a scale decomposition of the velocities, we observe that these EKE pulses correspond  
119 to features located within the equatorial band and have scales larger than the typical mesoscale  
120 eddy size (see Methods; Fig. S5a,c). Thus, we conclude that modulation of the equatorial currents  
121 during El Niño events are responsible for these distinct ~year-long EKE pulses.

122 To further investigate the effect of El Niño events on the mesoscale, we remove the equa-  
123 torial regions ( $9^{\circ}\text{S}$ – $9^{\circ}\text{N}$ ) and repeat the global trend analysis for EKE and SST gradients. The  
124 globally integrated extratropical EKE and SST gradient trends increase, while the corresponding  
125 uncertainty decreases; namely, EKE trends are now  $1.46\% \pm 0.24\%$  per decade and SST gradi-  
126 ents trends are  $1.24\% \pm 0.17\%$  per decade (see striped bars in Fig. 4c,e); both significant at the  
127 95% confidence level. It is thus clear that mesoscale activity in the Pacific, and particularly in the  
128 equatorial region, is strongly influenced by interannual variability.

129 The observed changes in EKE and SST of whole ocean basins still integrate over large het-  
130 erogeneous regions with opposing trends. For example, the Pacific Ocean aggregates the strength-  
131 ening of the equatorial currents in the equatorial Pacific Ocean during El Niño events, boundary  
132 currents, and the broader-scale oceanic gyres. These dynamical regions are not unique to the Pa-  
133 cific Ocean; the Atlantic and Indian basins also span over diverse dynamical regions. Therefore,  
134 we further decompose the ocean into dynamical regions (Fig. 5d): namely, the Antarctic Circum-  
135 polar Current (ACC), the boundary currents and their extensions, the equatorial regions, and the  
136 subtropical ocean gyres (see Methods for dynamical region definitions). The remaining regions  
137 are grouped together. We now investigate the variability and trends within each of these dynamical  
138 regions.

139 Globally, there is a significant increase of EKE and SST gradients, however, each dynami-  
140 cal region shows a different response (Fig. 5). For example, the ACC region shows a significant  
141 increase in both EKE and SST gradients, consistent with an increase in wind stress there, as demon-



142 strated in previous studies<sup>13,31</sup>. Boundary currents and their extensions collectively show a similar  
143 net response, in which EKE and SST gradients both increase. Individually, SST gradients increase  
144 in all boundary currents, however EKE in the Agulhas retroflection, the East Australian Current,  
145 the Leeuwin Current and the Malvinas Current have significantly increased, but the Gulf Stream  
146 and the Kuroshio Current do not show a significant net strengthening<sup>32,33</sup> (Fig. S6); instead, re-  
147 gions of increase and decrease tend to cancel each other out in an area-average, particularly for  
148 the Kuroshio Current (Fig. 1f). The response seen in the Gulf Stream and Kuroshio Currents is  
149 consistent with the idea that these currents have shifted polewards<sup>7,34</sup> and readjusted to climate  
150 modes<sup>35</sup>. Note that a poleward shift cannot be captured by our static climatological definition of  
151 the boundary currents (see Methods). Moreover, regions outside the dynamical regions exhibit a  
152 significant increase of EKE and SST gradients, comparable to those of the more energetic regions.  
153 Thus, the observed increase of these remaining regions could be a consequence of the poleward  
154 shift of the boundary currents, particularly in the Northern Hemisphere.

155       There are two dynamical regions that do not exhibit an increasing trend (Fig. 5). The equato-  
156 rial region is, again, dominated by interannual variability, where the largest changes correspond to  
157 El Niño events in both the EKE and mesoscale SST gradient time series. The mesoscale SST gradi-  
158 ents show a significant decrease, which suggests a reduction of the mesoscale variability. However,  
159 the equatorial EKE trends are insignificant, due to the dominance of interannual variability in the  
160 signal. Similarly, subtropical gyre mesoscale SST gradients have decreased significantly over the  
161 past few decades, while the EKE trend is negative, but not significant. This observed decrease  
162 of SST gradients across the tropics and subtropics (Fig. 2) could result from the displacement of

163 the extratropical atmospheric circulation<sup>36,37</sup> (Hadley cells) and expansion of the tropics<sup>38</sup>, which  
164 could explain the decreasing SST gradient trends in the tropics due to homogenization of the trop-  
165 ical surface SST gradients. Thus, the SST gradients induced by mesoscale stirring are expected to  
166 reduce as the surface ocean becomes more thermally homogeneous. However, a longer record is  
167 required to separate the mesoscale response from interannual-decadal climate variability.

168 We have analysed the available satellite altimetry record of SSH and SST to reveal a signifi-  
169 cant global intensification of the mesoscale field over the satellite record, where eddy-rich regions  
170 such as boundary current extensions and the Antarctic Circumpolar Current present a significant  
171 strengthening of 2% and 5% respectively. This increase is most apparent in regions where eddies  
172 are already strong, such as the Southern Ocean as well as boundary currents and their exten-  
173 sions. This mesoscale signature is observed in both eddy kinetic energy and mesoscale sea surface  
174 temperature gradients. Note that these observed trends only capture the ocean's surface. If we  
175 assume that the mesoscale flow extends vertically to around 500 m depth (a reasonable assump-  
176 tion given the vertical structure of the gravest mode<sup>39</sup>), then the EKE trends imply a significant  
177 change in the global mesoscale energy budget of 0.42% - 0.93% ( $0.05 - 0.12 \times 10^{18}$  Joules) over  
178 the satellite altimetry record<sup>40</sup>. Given the importance of mesoscale eddies for the redistribution of  
179 carbon and heat by the oceans, in addition to their capacity to control climate on regional to global  
180 scales, understanding this signal is crucial for improving future predictions of our oceans under  
181 anthropogenic forcing. Furthermore, current generation climate models used for climate projec-  
182 tions (CMIP6) do not generally resolve eddies<sup>41</sup>; thus, important climatic adjustments driven by  
183 eddy-forced changes are likely to be missing from these simulations.

184 The mesoscale evolution described here cannot be captured by coarse-resolution reanalysis  
185 products and sparse ARGO float observations, as previously used in other studies<sup>15</sup>. Here we used  
186 eddy-permitting satellite observations to reveal an insignificant decrease in EKE over the tropics,  
187 while reanalysis products suggest the tropics is where KE has recently increased the most. To  
188 better understand this discrepancy, we revisited the analysis for the oceanic KE time-anomaly, as  
189 undertaken by Hu et al.<sup>15</sup>, using our eddy-permitting satellite-derived KE. The results obtained are  
190 almost identical to the trends we presented in the analysis above (Fig. S7), yet strikingly different  
191 from those obtained by Hu et al. using non-eddy-resolving data (cf. their Fig 2a). This suggests that  
192 the inability of reanalyses and the Argo dataset to resolve the mesoscale, due to coarse-resolution  
193 and sparse sampling, leads to different estimates for regional kinetic energy trends. In addition, as  
194 we have demonstrated, the tropics are strongly influenced by interannual variability (e.g. El Niño),  
195 and yet the KE time-series from coarse reanalysis data do not detect the two extreme El Niño  
196 events observed in the satellite record. Thus, data resolution is the likely cause of the discrepancy  
197 between our eddy-permitting analysis and the results from the prior work using coarse-resolution  
198 reanalysis products<sup>15</sup>.

199 We investigated several hypotheses in order to explain the observed trends including (1)  
200 changes in winds (wind curl, wind stress), (2) changes in stratification, and (3) changes in large-  
201 scale temperature gradients. The drivers of the observed trends are likely to be a combination of  
202 (1) non-local increases of winds, (2) outcropping and tilting of isopycnals, and (3) strengthening  
203 of baroclinic instability. However, the changes of winds, isopycnals and baroclinic growth rate are  
204 small, highly variable and sparse. Thus, better observational records are required to understand

205 the increase in EKE observed from satellites. Moreover, as discussed above, coarse-resolution re-  
206 analysis products differ significantly from the mesoscale evolution we could detect from observed  
207 satellite altimetry. Thus, an in-depth analysis of the mesoscale evolution should be explored in  
208 more detail using eddy-resolving global ocean models and/or reanalyses.

209 Our study has documented major global scale reorganisation of the ocean's mesoscale kinetic  
210 energy observed over the past three decades. These observed adjustments in the mesoscale field  
211 have the potential to affect ocean circulation at all scales, and to modify the transport and redistri-  
212 bution of tracers, such as heat, carbon, and nutrients. Our findings thus have major implications  
213 for ocean readjustment to a changing climate.

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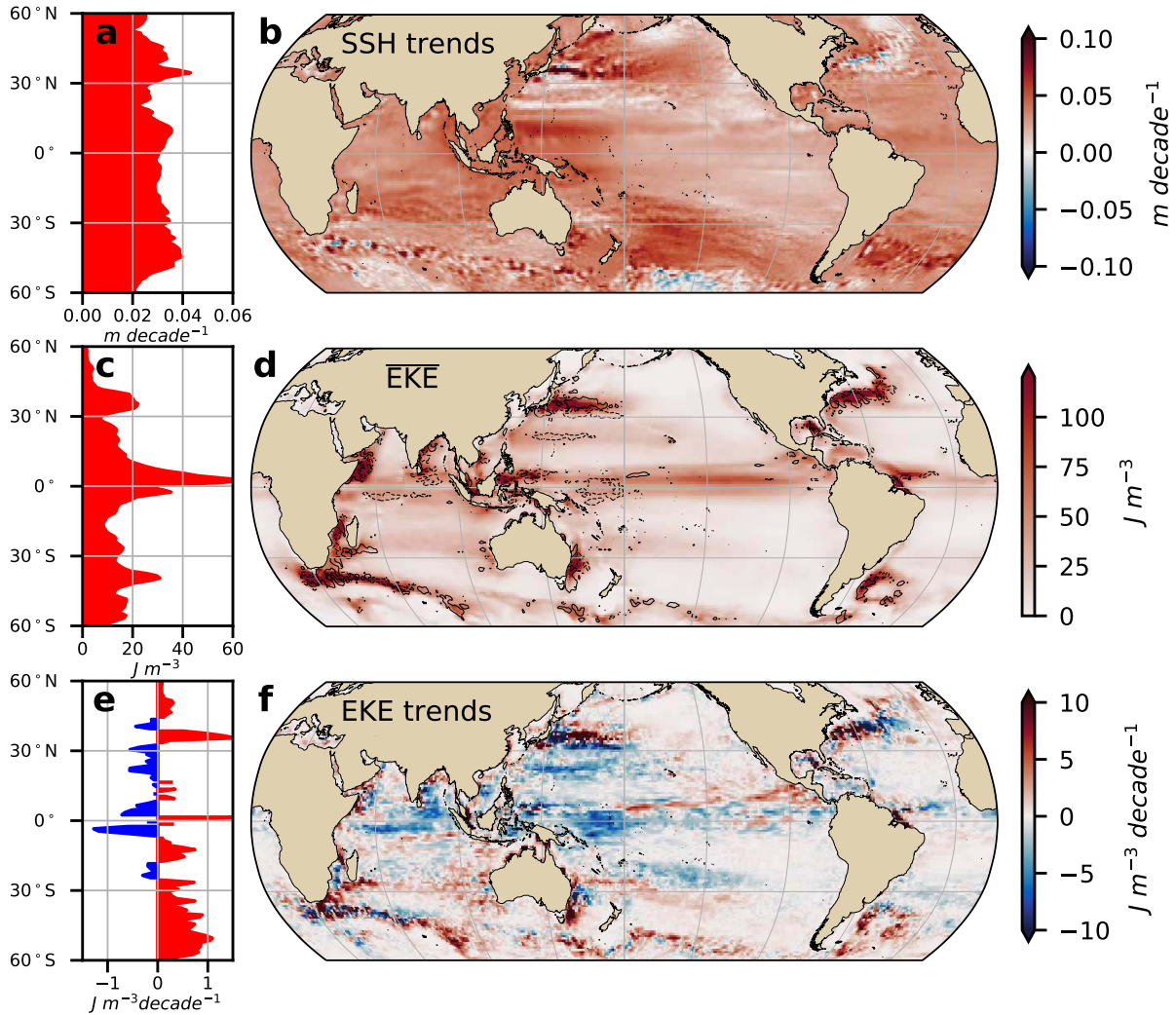


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## 312 **Acknowledgments**

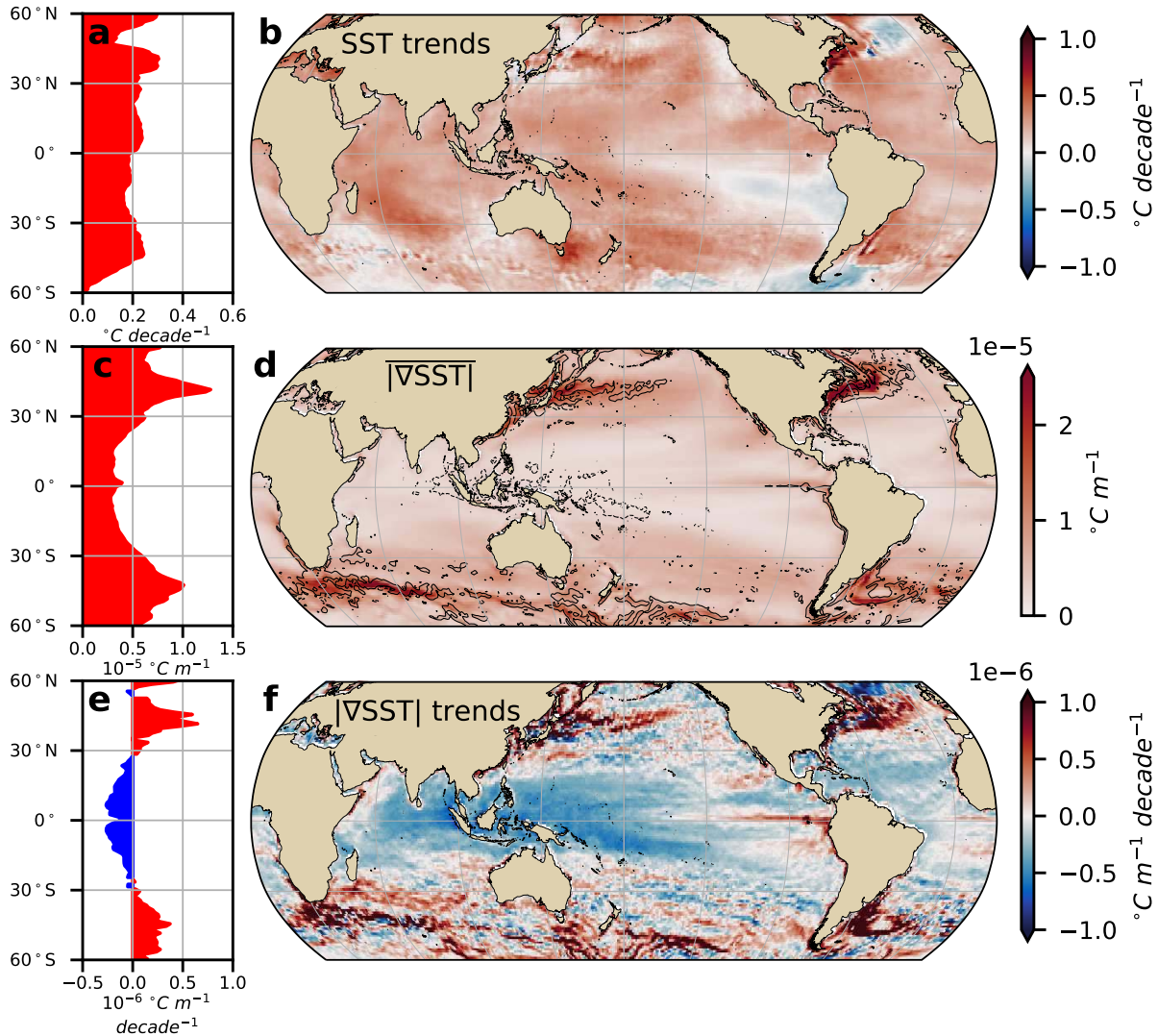
313 We thank Ryan Holmes for clarifying the equatorial response of El Niño events during  
314 1997-1998 and 2015-2016. The satellite altimetry products were produced by Ssalto/Duacs and  
315 distributed by AVISO+, with support from CNES (<https://www.aviso.altimetry.fr/>)

316 en/data/products/sea-surface-height-products/global/gridded-sea-level-height  
317 html). All datasets are available in netCDF format at [https://doi.org/10.5281/zenodo.](https://doi.org/10.5281/zenodo.3993824)  
318 3993824. Additional information and notebooks that reproduce the figures can be found at  
319 [https://github.com/josuemtzmoe/EKE\\_SST\\_trends](https://github.com/josuemtzmoe/EKE_SST_trends). J. M.-M. was supported by the  
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325 **Figure 1 Sea Surface Height trend, mean surface Eddy Kinetic Energy and surface**  
 326 **Eddy Kinetic Energy trend between 1993-2019** (a) Zonally averaged SSH trend; (b)  
 327 map of SSH trend (92.1% of global area is significant above 95% confidence level; for  
 328 spatial distribution refer to Fig. S1); (c) zonally averaged mean EKE; (d) map of mean  
 329 EKE; (e) zonally averaged EKE trend; (f) map of EKE trend (55.4% of global area is  
 330 significant above 95% confidence level). In panel (d), solid contours show positive EKE  
 331 trends and dashed contours show negative EKE trends (contours of  $\pm 5 J m^{-3} decade^{-1}$ ).  
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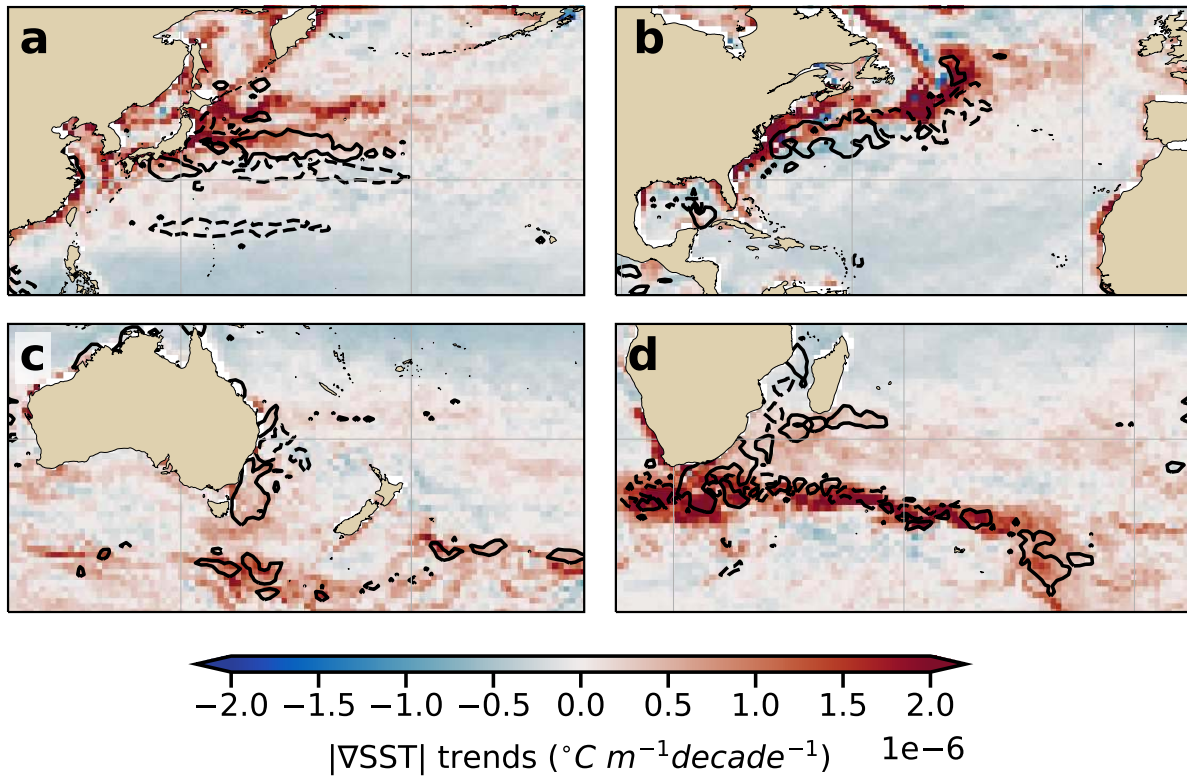
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**Figure 2 Sea Surface Temperature (SST) trends, mean SST gradient magnitude, and SST gradient magnitude trends between 1993-2019.** (a) Zonally averaged SST trend; (b) map of SST trend (76.7% of global area is significant above 95% confidence level; for the spatial distribution refer to Fig. S1); (c) zonally averaged time-mean of SST gradient magnitude; (d) map of time-mean of SST gradients magnitude; (e) zonally averaged SST gradient trend; (f) map of SST gradient trends (81.6% of global area is significant above 95% confidence level). In panel (d), solid contours show positive SST gradient trends and dashed contours show negative SST gradient trends ( $\pm 0.5 \times 10^{-6} \text{ } ^\circ\text{C m}^{-1}\text{decade}^{-1}$ ). Note that the spatial pattern of SST gradient maps is independent of the temporal extent of the SST gradient record used to compute SST gradient trends (Fig. S2).



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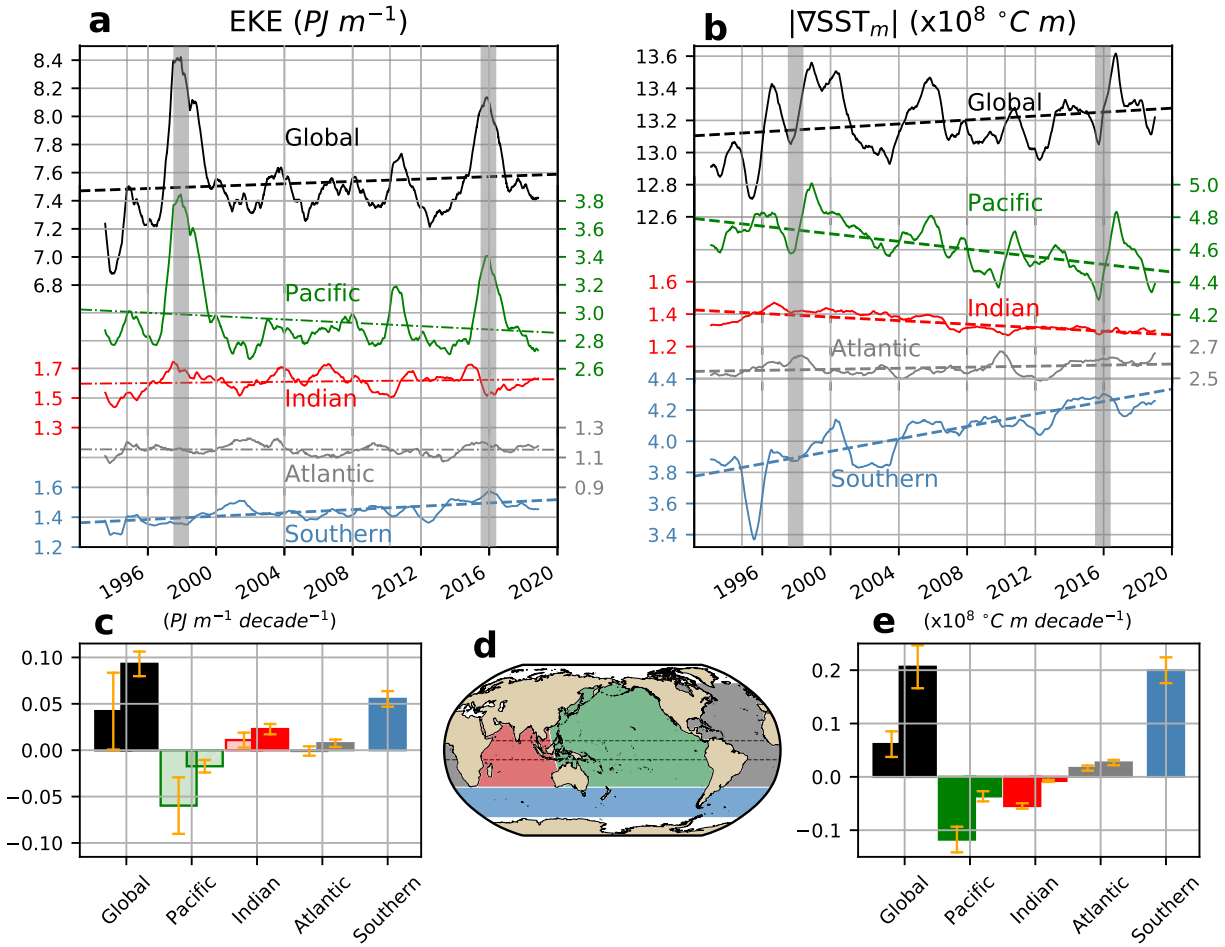
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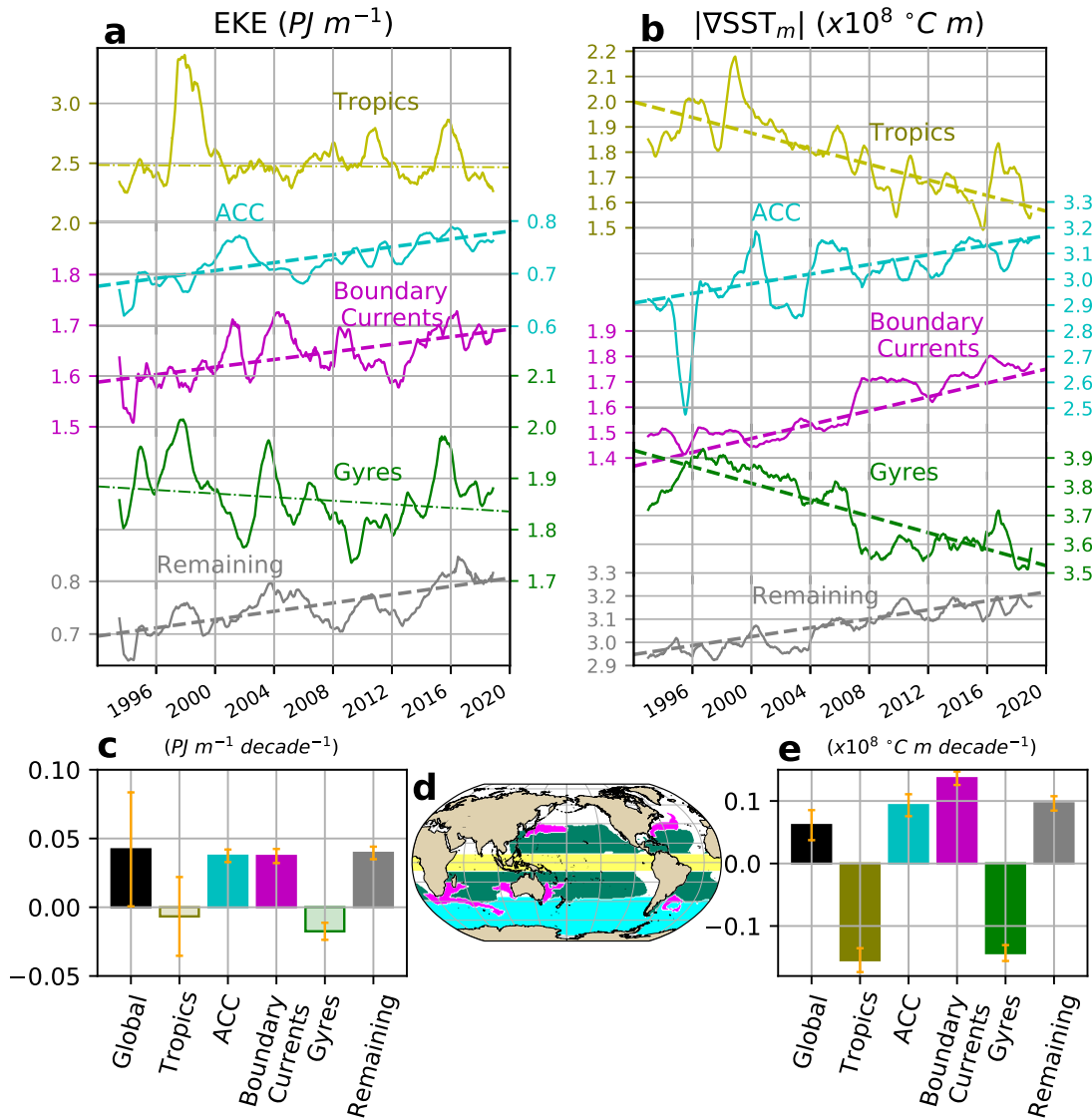
**Figure 3 Regional maps of mesoscale SST gradient trends and EKE trends.** (a) Kuroshio Current; (b) Gulf Stream; (c) East Australian Current; (d) Agulhas retroflexion. In all panels, solid contours show positive EKE trends and dashed contours show negative EKE trends (contours of  $\pm 5 J m^{-3} decade^{-1}$ ).



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**Figure 4 Time-series and linear trends of area integrated EKE and mesoscale SST gradients over various ocean basins.** Global (solid black), Southern (blue), Indian (red), Pacific (green), and Atlantic Oceans (gray) and each region separately without the equatorial region (striped bars). (a) EKE time series. (b) mesoscale SST gradient time series. In panels (a) and (b), solid curves denote 12-month running averages for each basin, dashed lines correspond to significant time-series trend, dashed-dotted lines show insignificant time-series trend, and vertical gray bars indicate El Niño events (above the 90th percentile of MEI.v2<sup>42</sup>). Note that the y axis is discontinuous in panels (a) and (b). (c) Linear EKE trends for each basin. (d) Ocean basins; equatorial region (9°S–9°N) is marked by the dashed lines. (e) Linear mesoscale SST gradient trends. In panels (c) and (e), standard errors are shown with orange bars and statistically significant trends (above 95% confidence level) denoted with solid bars while non-significant trends are translucent.





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**Figure 5 Time-series and linear trends of integrated EKE and mesoscale SST gradients over dynamical regions.** Global Ocean (black), Tropics (yellow), Antarctic Circumpolar Current (cyan), Boundary Currents and their extensions (magenta), Subtropical Ocean Gyres (green) and the rest (gray). (a) EKE time series. (b) mesoscale SST gradient time series. In panels (a) and (b), solid color curves are 12-month running averages of EKE and mesoscale SST gradients for each region, dashed lines correspond to significant time-series trend, and dashed-dotted lines show insignificant time-series trend. (c) Linear EKE trends for each dynamical region. (d) Definition of ocean regions. (e) Linear mesoscale SST gradient trends. Note that in panel (a) the top curve that corresponds to the Tropics has a different scale than the rest. In panels (c) and (e), standard errors are shown with orange bars and statistically significant trends (above 95% confidence level) are solid bars while non-significant trends are translucent.

## 377 **Methods**

### 378 **Observational products**

379 The data used in this study includes Sea Surface Height, geostrophic velocities, and Sea Sur-  
380 face Temperature. The Archiving, Validation and Interpretation of Satellite Oceanographic data  
381 (AVISO+) gridded multimission sea surface height and geostrophic velocities have a horizontal  
382 resolution of  $1/4^\circ$ . Currents within the equatorial region ( $5^\circ S$ - $5^\circ N$ ) are estimated using an equa-  
383 torial  $\beta$ -plane approximation of the geostrophic equations<sup>43</sup>. National Oceanic and Atmospheric  
384 Administration - Optimum Interpolation Sea Surface Temperature (NOAA-OISST) has a hori-  
385 zontal resolution of  $1/4^\circ$ <sup>23</sup>. This dataset is constructed by combining observations from different  
386 products (e.g. satellites, ships, buoys, and ARGO floats).

387 These datasets have a quasi-global coverage ( $65^\circ S$ - $65^\circ N$ ) and span 26 years, from January  
388 1993 to December 2019. Anomalies were computed with respect to the record's climatology.  
389 We have verified that using a different period for defining the climatology does not change the  
390 observed trends in the anomaly fields. Moreover, through wavelet analysis (edge detection) of the  
391 time-series, we have verified that the observed trends are not a product of steps in the number of  
392 satellite missions and oceanographic observations.

### 393 **Kinetic Energy decomposition**

394 Kinetic energy density is decomposed into the energy density contained by the steady flow  
395 (time-mean) and that contained by the transient flow (time-varying). In other words, the surface  
396 geostrophic velocity components are split using a Reynolds decomposition into their time-mean



397  $(\bar{u}, \bar{v})$  and time-varying components ( $u' = u - \bar{u}$ ,  $v' = v - \bar{v}$ ), with bars denoting time-averages  
 398 over the whole record. The KE is, therefore, decomposed as:

$$399 \quad \underbrace{\frac{1}{2}\rho_0 (u^2 + v^2)}_{\text{KE}} = \underbrace{\frac{1}{2}\rho_0 (u'^2 + v'^2)}_{\text{EKE}} + \underbrace{\frac{1}{2}\rho_0 (\bar{u}^2 + \bar{v}^2)}_{\text{MKE}} + \underbrace{\rho_0(\bar{u}u' + \bar{v}v')}_{\text{Cross terms}}, \quad (1)$$

400 where we approximate the density of the seawater by the constant  $\rho_0 = 1025 \text{ kg m}^{-3}$ . The  
 401 energy contained in the time-varying component of the flow is known as the eddy kinetic energy  
 402 (EKE), while the mean kinetic energy (MKE) is the energy of the time-mean flow.

403 Maps of EKE in this study correspond to the the time-mean EKE, defined as

$$404 \quad \overline{\text{EKE}}(x, y) = \overline{\frac{1}{2}\rho_0(u'^2 + v'^2)}. \quad (2)$$

405 Time-series correspond to the surface integrated EKE (globally or over specific regions),

$$406 \quad \langle \text{EKE} \rangle(t) = \iint_A \frac{1}{2}\rho_0(u'^2 + v'^2) dx dy, \quad (3)$$

407 where  $A$  refers to the area of each geographical or dynamical region,  $u'$  and  $v'$  are anomalies of  
 408 the surface geostrophic velocities provided by AVISO+, and angle brackets  $\langle \rangle$  denote the area  
 409 integral.

### 410 **Sea surface temperature gradients**

411 Sea surface temperature (SST) gradients are decomposed into mesoscale (SST gradients with  
 412 scales smaller than  $3^\circ$ ) and large-scale (SST gradients with scales larger than  $3^\circ$ ). To decompose  
 413 the SST gradients, we first compute large-scale SST by using an average spatial convolution of

414  $3^\circ \times 3^\circ$ , and a 12-months running average i.e.,

$$415 \quad \text{SST}_{ls}(x, y, t) = \frac{\iint \widetilde{\text{SST}}(x - x', y - y', t) K(x', y') dx' dy'}{\iint K(x', y') dx' dy'}, \quad (4)$$

416 where the SST field is convolved with a constant  $3^\circ \times 3^\circ$  kernel  $K$  and the tilde  $\widetilde{\phantom{x}}$  denotes a  
417 12-month running average. The mesoscale SST is then defined as

$$418 \quad \text{SST}_m = \text{SST} - \text{SST}_{ls}. \quad (5)$$

419 Small-scale EKE was computed analogously to the mesoscale SST but for the velocity components  
420 ( $u_m = u - u_{ls}$ ,  $v_m = v - v_{ls}$ ).

421 The gradients of the large-scale and mesoscale SST are computed afterwards. The SST  
422 gradient magnitude is:

$$423 \quad |\nabla \text{SST}| = \sqrt{\left(\frac{\partial \text{SST}}{\partial x}\right)^2 + \left(\frac{\partial \text{SST}}{\partial y}\right)^2}, \quad (6)$$

424 with analogous expressions for  $\text{SST}_m$  and  $\text{SST}_{ls}$ .

425 Computations of SST gradient time-series and time mean SST gradient trend maps are anal-  
426 ogous to those of EKE, namely:

$$427 \quad \langle |\nabla \text{SST}| \rangle(t) = \iint_A |\nabla \text{SST}| dx dy, \quad (7)$$

### 428 **Trends, significance & uncertainties**

429 Linear trends are calculated using a linear least-squares regression model for spatial inte-  
430 grated time-series and geographic data. The fields are first coarsened to a  $1^\circ \times 1^\circ$  grid, and then  
431 the linear trends are computed for each grid point.

432 A modified Mann-Kendall test is used to assess statistical significance of trends, while prop-  
433 erly taking into account the autocorrelation in the time-series<sup>24</sup>. The effective sample size for all  
434 the reported trends is always smaller than the actual sample size, due to autocorrelation of the  
435 time series. The significance of our observed trends for EKE and SST gradients is assessed via  
436 a t-student distribution test. If the probability of the trend is smaller than the confidence level  
437 ( $\alpha = 0.05$ ) then the null hypothesis is rejected and the estimated slope is considered significant  
438 at the 95% confidence level. The Pearson coefficient was also used to assess the significance of  
439 the trends. If the Pearson coefficient is smaller than 0.05 for a significance level of 95%, then the  
440 existing trend is considered to be statistically significant. Finally, the reported uncertainties corre-  
441 spond to the standard error using the effective sample size from the Mann-Kendall test; that is, the  
442 standard deviation of the time-series divided by the square root of the effective sample size.

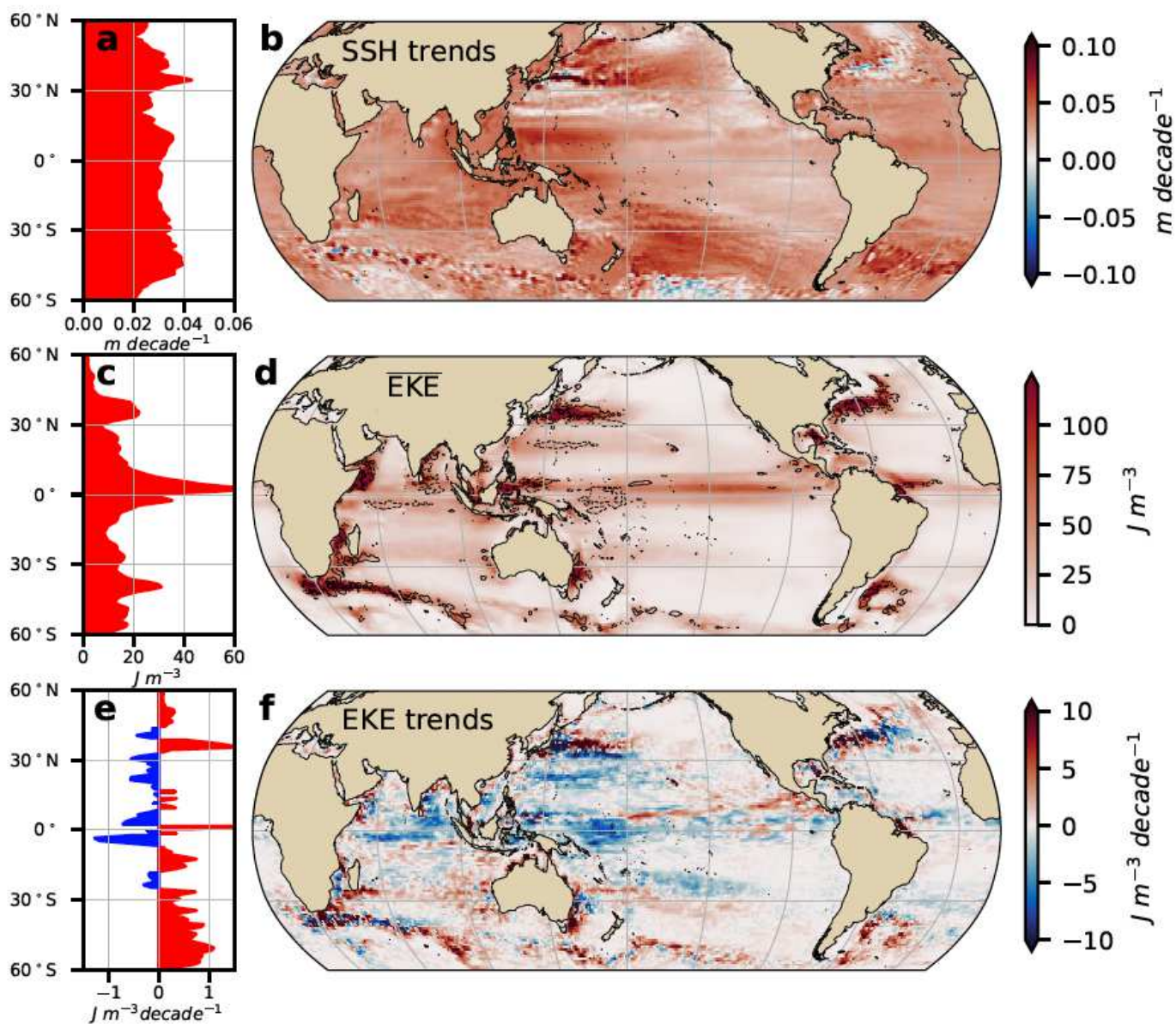
#### 443 **Geographical and dynamical regions**

444 Geographical regions consist of the following ocean basins: the Southern Ocean, the In-  
445 dian Ocean, the Pacific Ocean and the Atlantic Ocean. These ocean basins were defined to cap-  
446 ture ocean processes at all scales. The ocean basin mask can be obtained from the repository  
447 [https://github.com/josuemtzmo/EKE\\_SST\\_trends](https://github.com/josuemtzmo/EKE_SST_trends) that contains all the data used for  
448 this study (refer to acknowledgments; filename `ocean_basin_mask.nc`).

449 Dynamical regions were defined from the climatological mean SSH and the mean KE. We  
450 define a mask for each dynamical region and overlay the masks on a global mask in the follow-  
451 ing order. First, the Antarctic Circumpolar Current (ACC) is defined as all regions south of 35°S.

452 Second, the subtropical ocean gyre definitions depend on the particular ocean basin they are lo-  
453 cated within: the Pacific Ocean gyres correspond to mean SSH above the 0.65 *m* contour; the  
454 Atlantic Ocean gyres correspond to mean SSH above the 0.36 *m* contour; and the Indian Ocean  
455 gyres correspond to mean SSH above the 0.60 *m* contour. All these values were tuned to cap-  
456 ture approximately the same extension as the theoretical estimation of ocean gyres according  
457 to the Sverdrup balance. Third, the boundary currents and their extensions are defined as re-  
458 gions with mean KE above the  $\sim 99$ -th percentile ( $2.8\sigma$ ). Lastly, the equatorial region is de-  
459 fined as the region between 9°S and 9°N. The dynamical regional mask can be obtained from  
460 the repository containing all the data used for this study (refer to acknowledgments; filename  
461 `ocean_processes_mask.nc`).

# Figures

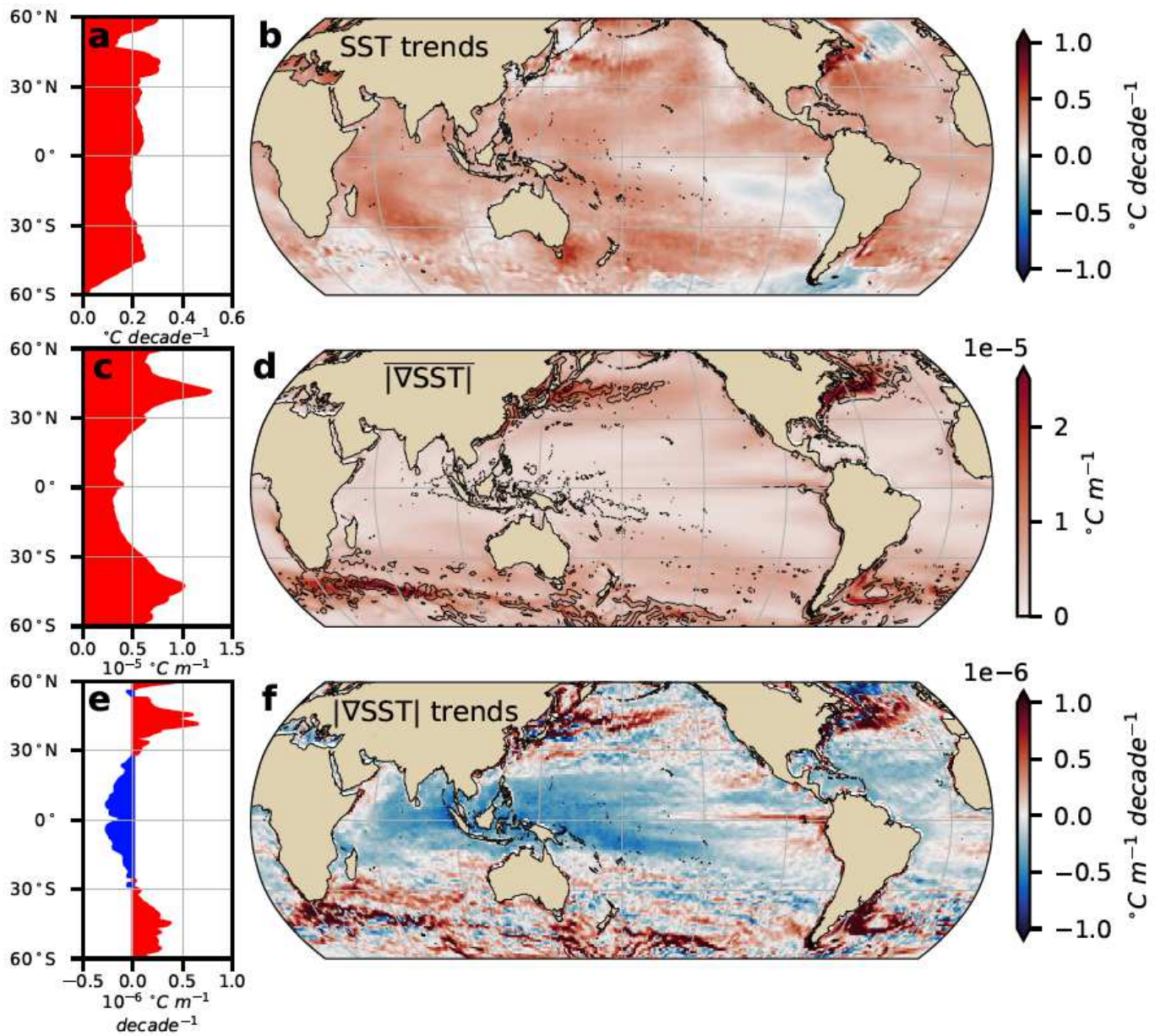


**Figure 1**

Sea Surface Height trend, mean surface Eddy Kinetic Energy and surface Eddy Kinetic Energy trend between 1993-2019 (a) Zonally averaged SSH trend; (b) map of SSH trend (92.1% of global area is significant above 95% confidence level; for spatial distribution refer to Fig. S1); (c) zonally averaged mean EKE; (d) map of mean EKE; (e) zonally averaged EKE trend; (f) map of EKE trend (55.4% of global area is significant above 95% confidence level). In panel (d), solid contours show positive EKE trends and dashed contours show negative EKE trends (contours of  $\pm 5 \text{ J m}^{-3} \text{ decade}^{-1}$ ). Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its



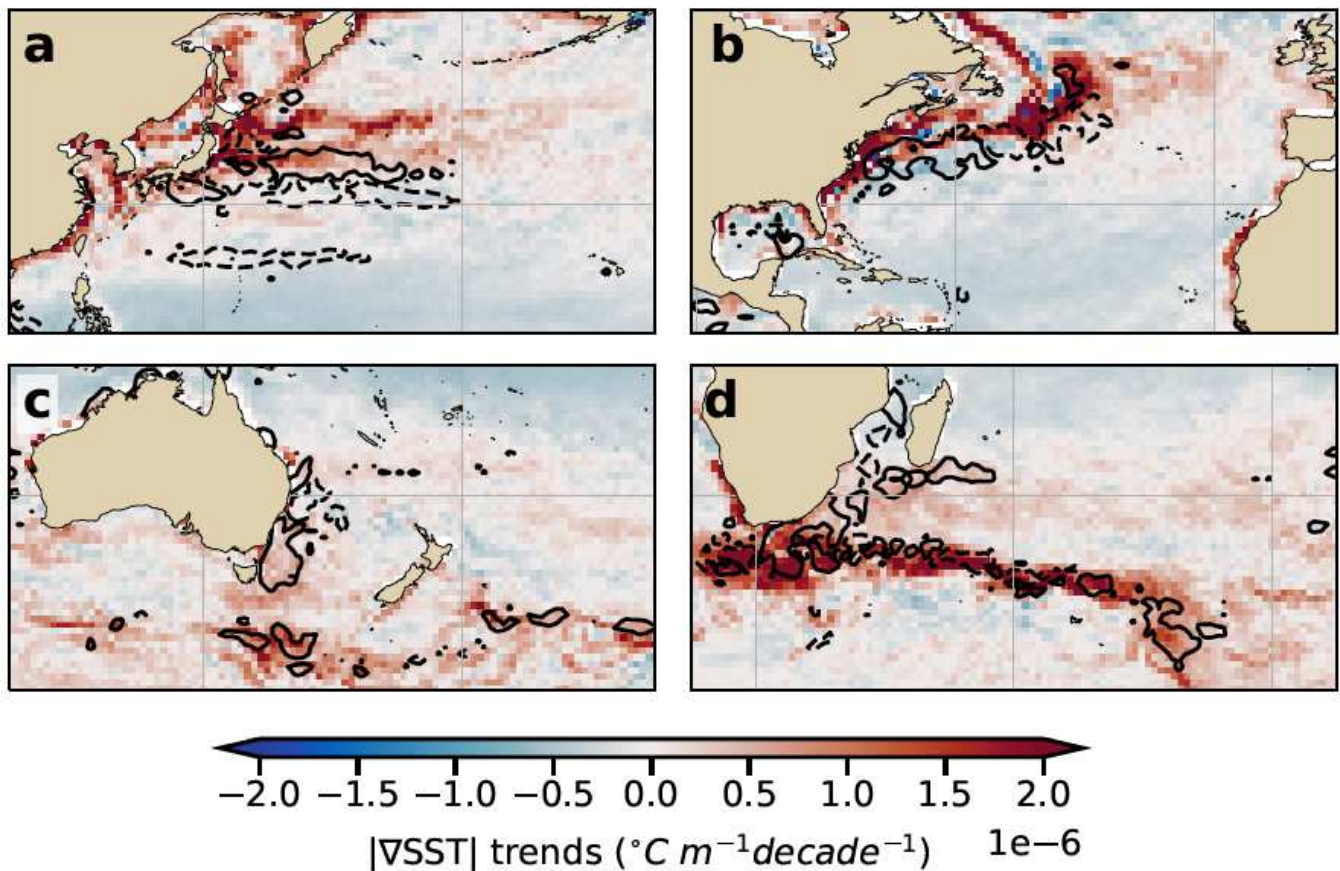
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**Figure 2**

Sea Surface Temperature (SST) trends, mean SST gradient magnitude, and SST gradient magnitude trends between 1993-2019. (a) Zonally averaged SST trend; (b) map of SST trend (76.7% of global area is significant above 95% confidence level; for the spatial distribution refer to Fig. S1); (c) zonally averaged time-mean of SST gradient magnitude; (d) map of time-mean of SST gradients magnitude; (e) zonally averaged SST gradient trend; (f) map of SST gradient trends (81.6% of global area is significant above 95% confidence level). In panel (d), solid contours show positive SST gradient trends and dashed contours show negative SST gradient trends ( $\pm 0.5 \times 10^{-6} \text{ } ^\circ\text{C m}^{-1}\text{decade}^{-1}$ ). Note that the spatial pattern of SST gradient maps is independent of the temporal extent of the SST gradient record used to compute

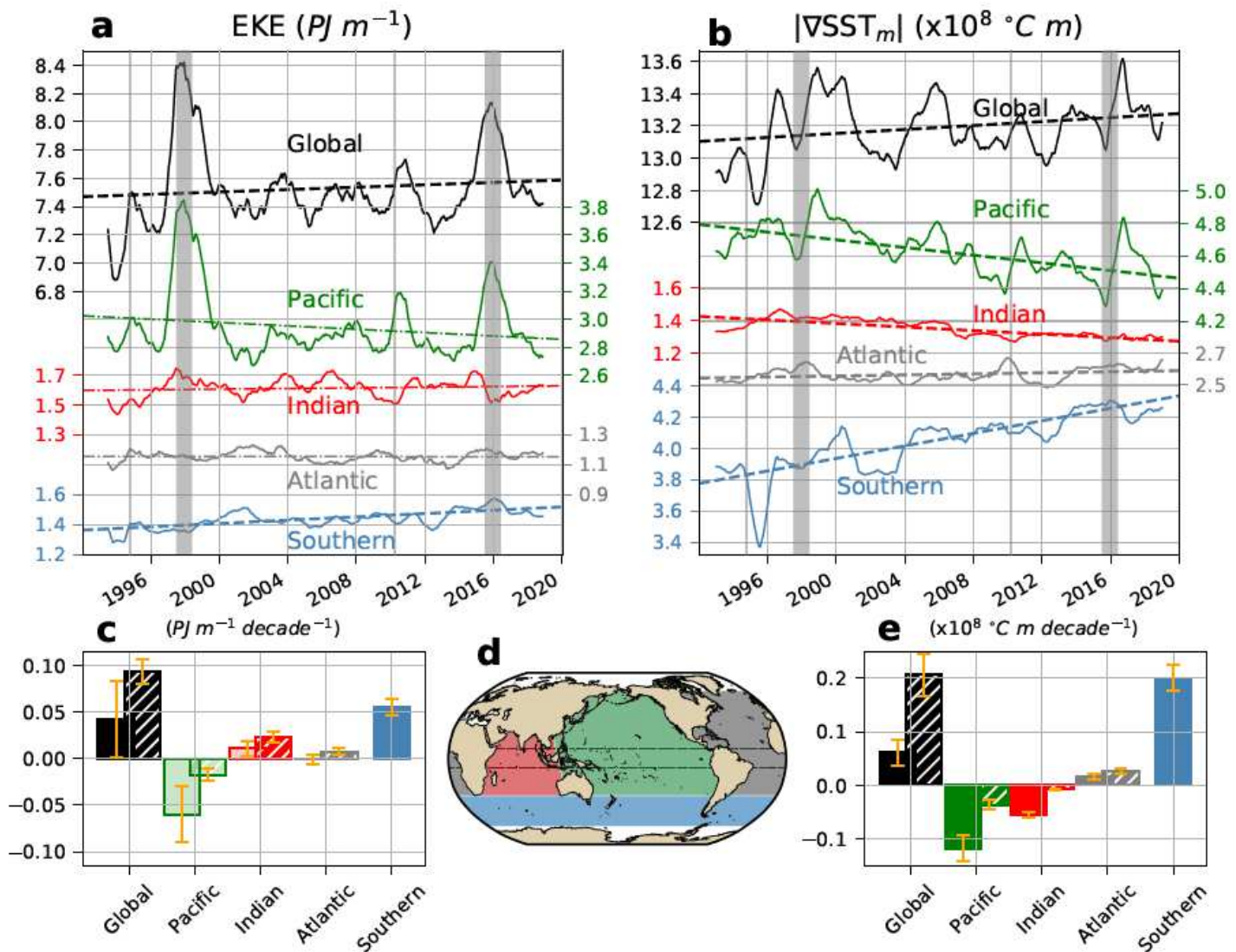
SST gradient trends (Fig. S2). Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.



**Figure 3**

Regional maps of mesoscale SST gradient trends and EKE trends. (a) Kuroshio Current; (b) Gulf Stream; (c) East Australian Current; (d) Agulhas retroflexion. In all panels, solid contours show positive EKE trends and dashed contours show negative EKE trends (contours of  $\pm 5 \text{ J m}^{-3} \text{ decade}^{-1}$ ). Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.



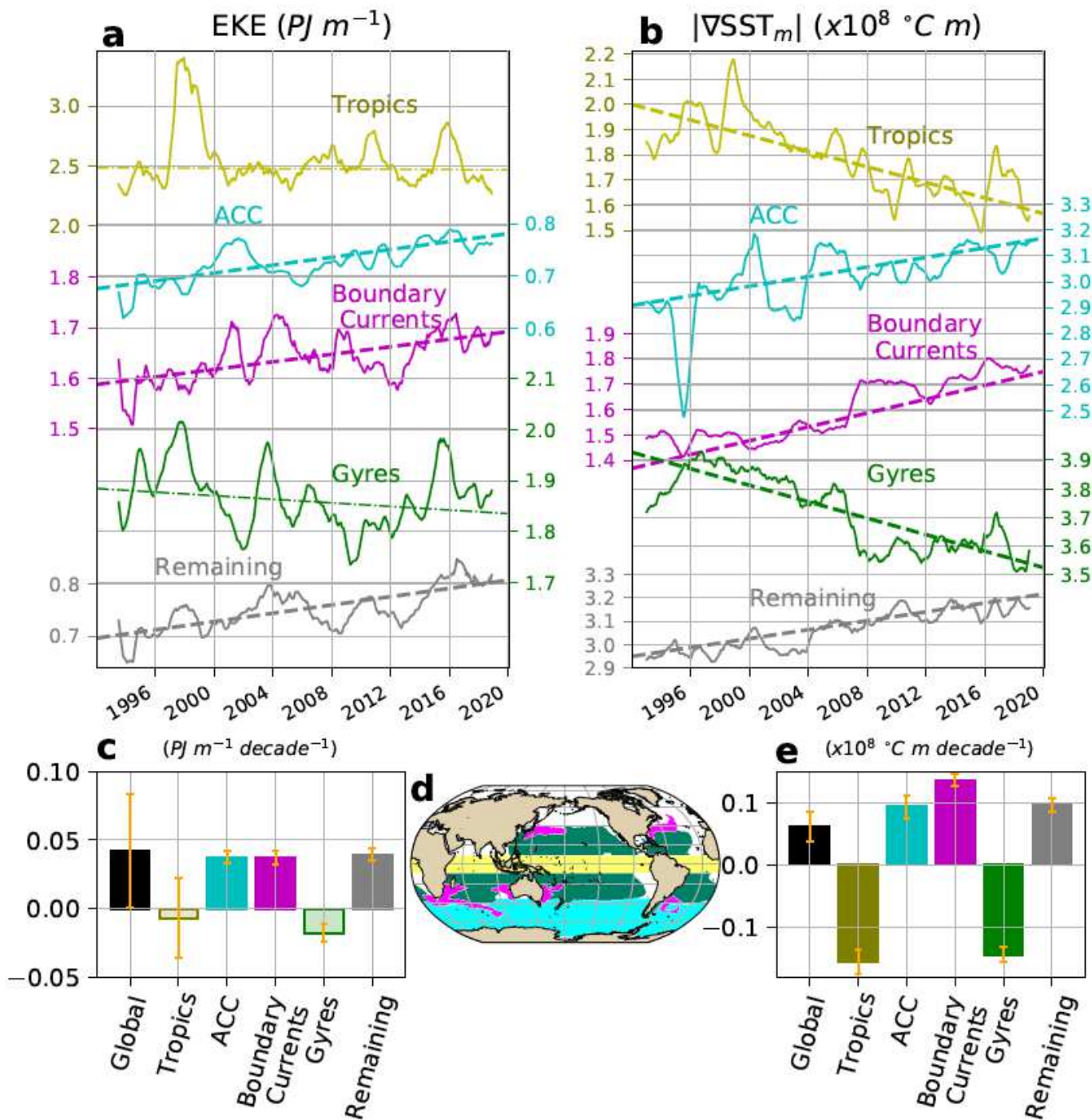


**Figure 4**

Time-series and linear trends of area integrated EKE and mesoscale SST gradients over various ocean basins. Global (solid black), Southern (blue), Indian (red), Pacific (green), and Atlantic Oceans (gray) and each region separately without the equatorial region (striped bars). (a) EKE time series. (b) mesoscale SST gradient time series. In panels (a) and (b), solid curves denote 12-month running averages for each basin, dashed lines correspond to significant time-series trend, dashed-dotted lines show insignificant time-series trend, and vertical gray bars indicate El Niño events (above the 90th percentile of MEI.v242). Note that the y axis is discontinuous in panels (a) and (b). (c) Linear EKE trends for each basin. (d) Ocean basins; equatorial region ( $9^\circ S - 9^\circ N$ ) is marked by the dashed lines. (e) Linear mesoscale SST gradient trends. In panels (c) and (e), standard errors are shown with orange bars and statistically significant trends (above 95% confidence level) denoted with solid bars while non-significant trends are translucent. Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any



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**Figure 5**

Time-series and linear trends of integrated EKE and mesoscale SST gradients over dynamical regions. Global Ocean (black), Tropics (yellow), Antarctic Circumpolar Current (cyan), Boundary Currents and their extensions (magenta), Subtropical Ocean Gyres (green) and the rest (gray). (a) EKE time series. (b) mesoscale SST gradient time series. In panels (a) and (b), solid color curves are 12-month running averages of EKE and mesoscale SST gradients for each region, dashed lines correspond to significant

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