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# Improved estimates of global terrestrial photosynthesis using information on leaf chlorophyll content

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#### **Abstract**

The terrestrial biosphere plays a critical role in mitigating climate change by absorbing anthropogenic CO<sub>2</sub> emissions through photosynthesis. The rate of photosynthesis is determined jointly by environmental variables and the intrinsic photosynthetic capacity of plants (i.e. maximum carboxylation rate;  $Vc_{max}^{25}$ ). A lack of an effective means to derive spatially and temporally explicit Vc25 has long hampered efforts towards estimating global photosynthesis accurately. Recent work suggests that leaf chlorophyll content (Chl<sub>leaf</sub>) is strongly related to  $V_{c,a}^{25}$ , since Chl<sub>leaf</sub> and  $V_{c,a}^{25}$  are both correlated with photosynthetic nitrogen content. We used medium resolution satellite images to derive spatially and temporally explicit Chliear, which we then used to parameterize  $V_{c_{max}}^{25}$  within a terrestrial biosphere model. Modelled photosynthesis estimates were evaluated against measured photosynthesis at 124 eddy covariance sites. The inclusion of Chlieaf in a terrestrial biosphere model improved the spatial and temporal variability of photosynthesis estimates, reducing biases at eddy covariance sites by 8% on average, with the largest improvements occurring for croplands (21% bias reduction) and deciduous forests (15% bias reduction). At the global scale, the inclusion of Chl<sub>leaf</sub> reduced terrestrial photosynthesis estimates by 9 PgC/ year and improved the correlations with a reconstructed solar-induced fluorescence product and a gridded photosynthesis product upscaled from tower measurements. We found positive impacts of Chl<sub>leaf</sub> on modelled photosynthesis for deciduous forests, croplands, grasslands, savannas and wetlands, but mixed impacts for shrublands and evergreen broadleaf forests and negative impacts for evergreen needleleaf forests and mixed forests. Our results highlight the potential of Chl<sub>leaf</sub> to reduce the uncertainty of global photosynthesis but identify challenges for incorporating Chl<sub>leaf</sub> in future terrestrial biosphere models.

KEYWORDS: gross primary productivity, leaf chlorophyll content, photosynthetic capacity, remote sensing, solar-induced fluorescence, terrestrial biosphere models

### 1 INTRODUCTION

The terrestrial biosphere has been a substantial carbon sink over the past half century, absorbing around one third of anthropogenic  $CO_2$  emissions (Keenan & Williams, 2018; Le Quéré et al., 2017). This net carbon sink exists due to a slight difference between carbon uptake through photosynthesis and carbon loss through ecosystem respiration and natural disturbances (i.e. wild fires). The size of individual carbon fluxes is generally one to two orders of magnitude larger than the net carbon sink (Beer et al., 2010; Bond-Lamberty & Thomson, 2010; Sitch et al., 2015), therefore, a small negative change in photosynthesis can easily influence the ability of the land sink to mitigate anthropogenic  $CO_2$  emissions.

Currently, estimates of global photosynthesis (gross primary production; GPP) from state-of-the-art terrestrial biosphere models (TBMs) vary in a wide range from 112 to 169 PgC/year (Baldocchi, Ryu, & Keenan, 2016; Piao et al., 2013), accompanied by substantial discrepancies in estimated trends and variability of GPP (Anav et al., 2015; Li, Ciais, et al., 2018). Much of the uncertainty in GPP is attributed to an inadequate constraint of the maximum leaf carboxylation rate (normalized to 25 degrees; Vc25) in TBMs (Bonan et al., 2011; Schaefer et al., 2012; Walker et al., 2017). Vc25 is related to the concentration of the photosynthetic enzyme ribulose-1,5-bisphosphate carboxylase/oxygenase (Rubisco) (Farguhar, von Caemmerer, & Berry, 1980) and is known to change over time and space under the influence of a wide range of factors, including leaf ontogeny (Croft et al., 2017), nutrient availability (Walker et al., 2014), seasonality of climatic variables (Misson, Tu, Boniello, & Goldstein, 2006), extreme weather (Zhou et al., 2014) and biodiversity and stand age (Musavi et al., 2017). However, the variability of  $Ve^{25}$  has yet to be fully considered in most currently used TBMs because of the difficulty in obtaining spatially and temporally continuous  $Ve^{25}$  for large-scale simulations from direct observations of Vc25, which require measurements of leaf photosynthesis under varying CO<sub>2</sub> concentrations (Stinziano et al., 2017; Wullschleger, 1993; Xu & Baldocchi, 2003). Therefore, current TBMs are limited to using fixed  $Vc^{25}$  values specific to plant functional types (PFTs) and empirical scaling functions to partly consider the variation in  $V_{c_{max}}^{25}$  when estimating global photosynthesis (Ryu et al., 2011; Thornton & Zimmermann, 2007).

Previous studies have attempted to derive Version from remotely sensed data, given that the spectrum of the land surface reflectance is influenced by leaf physiology and plant function. Statistical models (i.e. partial least square regression, PLSR) have been used to derive nitrogen content from the surface reflectance spectra for both leaves (Singh, Serbin, McNeil, Kingdon, & Townsend, 2015) and canopies (Dechant, Cuntz, Vohland, Schulz, &

Doktor, 2017). Some studies have adopted similar models to estimate  $Vc_{max}^{25}$ of leaves (Barnes et al., 2017; Dechant et al., 2017) and canopies (Serbin et al., 2015), considering that nitrogen content is a strong indicator of  $Ve_{max}^{25}$ (Kattge, Knorr, Raddatz, & Wirth, 2009; Walker et al., 2014). However, these empirical models are mostly only reliable in site-specific studies (Wang, Skidmore, Darvishzadeh, & Wang, 2018) and face some challenges when extrapolated to broad spatial scales because of the difficulty in separating the spectral signature of nitrogen, a nonpigment constituent, from overlapping and covarying signatures of other biochemical constituents of vegetation within different landscapes (Kokaly, Asner, Ollinger, Martin, & Wessman, 2009). Solar-induced fluorescence (SIF; 650-800 nm) is another remotely sensed indicator that has been used to derive  $V_{c,max}^{25}$  (Zhang et al., 2014; Zhang, Guanter, Joiner, Song, & Guan, 2018), considering the widely reported tight correlation between SIF and leaf photosynthesis (Yang et al., 2017) or GPP (Frankenberg et al., 2011; Guanter et al., 2014; Li, Xiao, & He, 2018). However, the derivation of  $V_{cond}^{25}$  from SIF faces the challenge of separating the impact of canopy structure on GPP from that of leaf physiological traits since the canopy-scale SIF contains information of both (Liu et al., 2018; Yang & van der Tol, 2018).

Chlorophyll is a crucial component of plant photosynthesis machinery, through harvesting photons and transporting electrons to support the production of biochemical energy for the Calvin-Benson cycle (Alton, 2017; Porcar-Castell et al., 2014). Leaf chlorophyll content (Chl<sub>leaf</sub>), defined as the total chlorophyll a and chlorophyll b content per one half of the total leaf area (all sided), can act as a proxy for the photosynthetic nitrogen pool, which is also shared by Rubisco (Evans, 1989b; Hikosaka & Terashima, 1996). Recent studies have found linear relationships between Chl<sub>leaf</sub> and  $Vc^{25}$  for several tree species (Croft et al., 2017) and agricultural crops (Houborg, Cescatti, Migliavacca, & Kustas, 2013; Houborg, McCabe, Cescatti, & Gitelson, 2015), and the integration of these Chl<sub>leaf</sub>-Ve<sup>25</sup> relationships into TBMs has been reported to reduce errors of estimated GPP at a cropland site (Houborg et al., 2013) and a forest site (Luo, Croft, et al., 2018). Since the  $Chl_{leaf}$  relationship is known to change with the light environment within a canopy (Evans, 1989a; Poorter, Kwant, Hernández, & Medina, 2000), an updated two-leaf scheme has been developed to account for the difference between the  $\mathrm{Chl}_{\mathrm{leaf}}$  relationship for sunlit leaves and that for shaded leaves in TBMs (Luo, Croft, et al., 2018). It is also argued that  $Chl_{leaf}$  serves as a better proxy for  $Vc_{max}^{25}$  than the commonly used area-based leaf nitrogen content (Narea), because N<sub>area</sub> includes the nitrogen for both photosynthetic and nonphotosynthetic components (i.e. structural development), and the nonphotosynthetic component can complicate the apparent  $N_{area}$  relationships (Croft et

al., 2017). Chl<sub>leaf</sub> effectively removes the influence of nonphotosynthetic nitrogen and only reflects the changes in the photosynthetically active nitrogen pool (Croft et al., 2017).

Recent developments in using physically based approaches to estimate Chl<sub>leaf</sub> from satellite data (Croft, Chen, Zhang, & Simic, 2013; Houborg, McCabe, Cescatti, Gao, et al., 2015; Zhang, Chen, Miller, & Noland, 2008) have made it possible to estimate globally continuous have made it possible to estimate globally continuous have continuous based approaches used established radiative transfer theories to downscale the remotely sensed canopy reflectance to leaf reflectance (Kuusk, 2018) and then use the leaf reflectance in a leaf optical model (Jacquemoud & Ustin, 2001) to estimate Chl<sub>leaf</sub>. This can remove the effect of background reflectance, multiple scattering and the bidirectional reflectance distribution function (BRDF) effects of a canopy to distil the spectral signal of Chl<sub>leaf</sub> only.

Here, we examine the potential of using satellite-derived Chl<sub>leaf</sub> to estimate temporally and spatially explicit across a diverse array of sites around the globe. To do so, we applied a two-step physically based approach (Croft et al., 2013; Simic, Chen, & Noland, 2011; Zhang et al., 2008) to derive Chl<sub>leaf</sub> at 124 eddy covariance sites (555 site-years) from the FLUXNET2015 tier 1 data set (Baldocchi, 2008; Pastorello et al., 2017), and modelled the GPP of these sites with and without Chl<sub>leaf</sub>-based (Croft et al., 2017). We examined the effects of the inclusion of Chl<sub>leaf</sub> in GPP modelling for each of the nine PFTs and further estimated global GPP. We compared the global GPP against a reconstructed SIF product (Gentine & Alemohammad, 2018), a remote

sensing-based GPP product (Zhao, Heinsch, Nemani, & Running, 2005) and a global GPP product upscaled from eddy covariance measurements (Jung et

### 2 MATERIALS AND METHODS

al., 2017).

# 2.1 Flux, meteorological and satellite data

FLUXNET2015 is the latest effort to refine the flux measurements from eddy covariance towers of different regional networks and publish the data in a standard format (Pastorello et al., 2017). It provides gap-filled GPP estimates and their concurrent meteorological records for over 200 sites. In this study, we used the FLUXNET2015 tier 1 data set released in November 2016. We selected 124 sites (555 site-years) (Supporting Information, section 1) based on the availability of the MEdium Resolution Imaging Spectrometer (MERIS) satellite data that were used for the leaf chlorophyll content (Chl<sub>leaf</sub>) derivation. We chose to use GPP partitioned from Net Ecosystem Exchange (NEE) using a night-time method (NT) with variable u\* threshold (VUT) for each year. NT uses only night-time data to parameterize a respiration model that is then applied to the whole data set to estimate ecosystem respiration and then calculate GPP (Reichstein et al., 2005). Many versions of the GPP records were created by using different percentiles of u\*. Among those GPP versions, we regarded GPP NT VUT REF as the reference GPP, which was

estimated by using the VUT and the original u\* values. The observational uncertainty of GPP was provided by the field GPP\_NT\_VUT\_SE in the data set.

Half-hourly and hourly meteorological records in the FLUXNET2015 data set were used to drive a TBM to estimate GPP. Gap-filled solar radiation (SW\_IN\_F), air temperature (TA\_F), vapour pressure deficit (VPD\_F), precipitation (P\_F) and wind speed (WS\_F) were selected as the forcing variables for the TBM, to allow the TBM to produce continuous hourly results. For GPP modelling at the global scale, we used MERRA-2 (Modern-Era Retrospective Analysis for Research and Applications, Version 2) data from Goddard Space Flight Center, NASA as the climate forcing to drive the TBM (He et al., 2018).

Leaf area index (LAI) series are required to drive the TBM and derive Chl<sub>leaf</sub>. We selected the Copernicus Global Land Service GEOV1 LAI product derived from SPOT-VEGETATION satellite, which has a global coverage of LAI from 1999 to 2014, at 10 day temporal intervals and a spatial resolution of 1 km (Baret et al., 2013). We used the Locally Adjusted Cubic-spline Capping (LACC) method (Chen, Deng, & Chen, 2006) to interpolate and smooth the discontinuous LAI into daily LAI series.

Clumping index ( $\Omega$ , or CI) describes the nonrandomness of the leaf distribution in the canopy (Chen, Rich, Gower, Norman, & Plummer, 1997). It is a key parameter to drive the canopy radiation transfer module in both the TBM and the Chl<sub>leaf</sub> derivation algorithm. Site-specific CI values were obtained from a global foliage clumping index map produced from the MODIS BRDF products (He, Chen, Pisek, Schaaf, & Strahler, 2012). The map has a resolution of 500 m, which is comparable to the size of regular tower footprints.

MEdium Resolution Imaging Spectrometer (MERIS) full resolution surface reflectance (SR) product was selected for deriving global Chl<sub>leaf</sub> maps. MERIS SR data were chosen due to the presence of chlorophyll-sensitive red-edge bands, fine temporal resolution (every 2-3 days), medium spatial resolution (300 m) and high radiometric accuracy (Curran & Steele, 2005). The SR product was produced as a 7 day temporal synthesis from images collected at the original 2-3 day revisit frequency (Rast, Bezy, & Bruzzi, 1999). The MERIS surface reflectance series were produced by a series of preprocessing steps, including radiometric, geometric and BRDF correction, pixel identification and atmospheric correction with aerosol retrieval. There are 13 bands (spectral resolution  $\sim=10$  nm) in the visible, red-edge and nearinfrared bands sampled in the reflectance data set. MERIS covers the complete years from 2003 to 2011. We extracted the surface reflectance in all bands at every flux tower site as the input for the Chl<sub>leaf</sub> derivation algorithm. In order to retain enough sampling points to detect the seasonal patterns of Chl<sub>leaf</sub>, only the site-years that have more than 10 MERIS surface reflectance records were considered in this study.

Solar-induced fluorescence (SIF). We used a recently released reconstructed SIF product (Gentine & Alemohammad, 2018) to evaluate the estimates of global photosynthesis. The reconstructed SIF (RSIF) is the product of photosynthetic active radiation (PAR) estimated by BESS (Ryu et al., 2011) and the reconstructed Global Ozone Monitoring-2 (GOME-2) SIF normalized by solar zenith angle (SZA), while the reconstructed GOME-2 SIF normalized by SZA was produced by a neural network trained on the original GOME-2 SIF normalized by SZA (Joiner et al., 2011) with the input of MODIS surface reflectance. This RSIF product exhibits much higher seasonal and interannual correlation than the original SIF when compared with eddy covariance estimates of GPP and two reference global GPP products (Gentine & Alemohammad, 2018). It has a more continuous coverage over time and space and a higher signal-to-noise ratio than the original SIF. We downloaded the monthly RSIF at 0.5°

from https://gentinelab.eee.columbia.edu/content/datasets. While using RSIF as a proxy for GPP in this study, we acknowledged that the strong correlation between SIF and GPP is largely explained by the dependence of both absorbed photosynthetic active radiation (APAR) (Zhang et al., 2016), and a recent study even suggested that SIF is more likely a proxy for APAR than for photosynthesis at a rice paddy (Yang et al., 2018).

FLUXCOM global carbon flux data set. FLUXCOM (Jung et al., 2017) produced global GPP estimates by upscaling site-level GPP from 224 flux tower sites using three machine learning algorithms: Random forests (RF), Artificial Neural Networks (ANN) and Multivariate Adaptive Regression Splines (MARS). Each machine learning algorithm was trained on daily fluxes using 11 inputs including site-level meteorological and remote sensing observations. After obtaining the trained algorithm, gridded climatic variables (CRU-NCEP v6) and satellite observations from MODIS were used to produce the gridded carbon flux estimations. We downloaded the monthly GPP estimation at 0.5° from www.bgc-jena.mpg.de/geodb/projects/Data.php.

MODIS GPP. We used MODIS GPP product (MOD17 collection 55) released by the Numerical Terradynamic Simulation Group (NTSG) at the University of Montana. The collection 55 GPP rectifies the underestimation of GPP incurred by cloud-contaminated pixels in the near-real-time MODIS GPP product (MOD17 collection 5) and was recommend for ecological studies (Zhao et al., 2005). MODIS GPP is provided at a monthly step and 0.5° resolution (http://files.ntsg.umt.edu/data/NTSG\_Products/).

Plant Functional Types (PFTs). Nine major PFTs were studied, including: croplands (CRO), deciduous broadleaf forests (DBF), evergreen broadleaf forests (EBF), evergreen needleleaf forests (ENF), grasslands (GRA), savannas and woody savannas (SAV), shrublands (SH) and wetlands (WET). For eddy covariance towers, their PFTs types were provided in metadata of the FLUXNET data set (Supporting Information, section 1). For the global-scale study, we used the PFTs classified by the MODIS Land Cover maps (Friedl et al., 2010) curated at 0.5°. For each 0.5° grid cell, we used the PFT

that was most prevalent during the period 2000–2012. FLUXNET metadata and MODIS Land Cover maps both adopted the vegetation classification protocol of the International Geosphere-Biosphere Programme (IGBP) which does not include a C4 class.

# 2.2 Derivation of global Chl<sub>leaf</sub>

Global Chl<sub>leaf</sub> maps were derived from MERIS surface reflectance, according to the procedure outlined in Croft et al. (), using a two-step process-based algorithm. The first step was to retrieve leaf reflectance spectra from satellite-derived canopy reflectance spectra through the inversion of canopy radiative transfer models. The second step was to use the retrieved leaf reflectance spectra from step 1 to estimate Chl<sub>leaf</sub> by inverting a leaf radiative transfer model. For the first step, two canopy radiative transfer models were selected, according to the structural characteristics of the vegetation present. For spatially 'clumped' vegetation types (i.e. deciduous and coniferous trees, shrubs) we selected the 4-Scale geometrical-optical model (Chen & Leblanc, 1997). The 4-Scale model simulates the BRDF based on the canopy architecture described at four scales: (a) vegetation grouping, (b) crown geometry, (c) branches and (d) foliage elements. A plant crown is, therefore, represented as a complex medium, where mutual scattering occurs between shoots or leaves. The 4-Scale model calculates canopy reflectance  $(\rho)$  as a linear summation of four components:

$$\rho = \rho_{PT}F_{PT} + \rho_{ZT}F_{ZT} + \rho_{PG}F_{PG} + \rho_{ZG}F_{ZG}(1)$$

where the sunlit vegetation  $(\rho_{PT})$ , shaded vegetation  $(\rho_{ZT})$ , sunlit ground  $(\rho_{PG})$ and shaded ground ( $\rho_{ZG}$ ), and  $F_{PT}$ ,  $F_{PG}$ ,  $F_{PG}$  and  $F_{ZT}$  represent the probability of viewing each component respectively. To derive leaf reflectance ( $\rho_L$ ) from sunlit crown reflectance ( $\rho_{PT}$ ), the enhancement of both sunlit and shaded reflectance due to multiple scattering is accounted for using a multiple scattering factor (M factor) (Croft et al., 2013; Simic et al., 2011; Zhang et al., 2008). For homogenous canopies, such as grassland and cropland species, where the distribution of foliage approaches randomness, we used the scattering by arbitrary inclined leaves (SAIL) model (Verhoef, 1984). Turbid medium models such as SAIL assume that the canopy is composed of homogeneous, horizontal layers of Lambertian scatters randomly distributed in space. SAIL is based on the following differential equations: (a) diffuse incoming flux, (b) diffuse outgoing flux, (c) direct solar flux and (d) direct radiant flux in the direction of the sensor. Both canopy reflectance models were inverted using a lookup table (LUT) approach, selected to optimize computational resources and reduce problems associated with local minima (Croft et al., ).

In the second step,  $Chl_{leaf}$  was retrieved using the PROSPECT leaf radiative transfer model (Jacquemoud & Baret, 1990) from the modelled leaf reflectance that was derived in step 1. In PROSPECT-5, leaf reflectance and transmittance (400–2,500 nm) are defined as a function of six parameters: structure parameter (N), chlorophyll (a + b) concentration ( $C_{ab}$ ), carotenoid

content (Car), brown pigment ( $C_b$ ), dry matter ( $C_m$ ) and equivalent water thickness ( $C_w$ ). This two-step approach has been successfully validated over various sites and biomes (Croft et al., 2013; Croft, Chen, Zhang, et al., 2015; Simic et al., 2011; Zhang et al., 2008). Chl<sub>leaf</sub> derived from the weekly MERIS reflectance observations was then interpolated to daily steps using the LACC method (Chen et al., 2006). Modelled Chl<sub>leaf</sub> showed a strong relationship with 248 ground measurements collected from 29 globally distributed sites. The relationship has  $R^2 = 0.47$  (p < 0.001) and RMSE = 10.8  $\mu$ g/cm² (Croft et al., ) (Supporting Information, section 2).

# 2.3 Conversion of Chl<sub>leaf</sub> into Ve<sup>25</sup><sub>max</sub>

Several linear equations were used to translate the daily  $Chl_{leaf}$  into  $Vc_{max}^{25}$  (Table 1). The  $Vc_{max}^{25}$ -Chl<sub>leaf</sub> relationships used for DBF, ENF, MF, SAV, WET were based on a recent work (Croft et al., 2017), who built a single and significant linear relationship across four deciduous broadleaf tree species through direct measurements of  $Chl_{leaf}$  and  $Vc_{max}^{25}$ . The  $Vc_{max}^{25}$ -Chl<sub>leaf</sub> relationships used for CRO were based on the approach developed by Houborg, McCabe, Cescatti, and Gitelson (2015), who used a mechanistic framework to build an equation to link  $Chl_{leaf}$  and  $Vc_{max}^{25}$  via area-based leaf total nitrogen content ( $N_{area}$ ), in which a single equation was developed for croplands. This relationship was also applied to another herbaceous type (GRA) in this study. We then applied a similar method to build  $Vc_{max}^{25}$ - $Chl_{leaf}$  equations for EBF and SH, using a comprehensive  $Vc_{max}^{25}$ - $N_{area}$  database (Kattge et al., 2009) and a constant ratio between  $Chl_{leaf}$  and  $N_{area}$  (Evans, 1989b) (Table 1).

**TABLE 1** Linear regression equations used to derive  $Vc_{max}^{25}$  from  $Chl_{leaf}$  ( $Vc_{max}^{25} = a \times Chl_{leaf} + b$ )

Plant functional types	a (μmol/m²/s per μg/cm²)	b (μmol/m²/s)	Reference
DBF, ENF, MF, SAV, WET	1.3	3.72	Croft et al., 2017
CRO, GRA	1.98	12.5	Houborg, McCabe, Cescatti, & Gitelson, 2015
SH	0.95	14.71	Kattge et al., 2009
EBF	0.66	6.99	and Evans, 1989b

CRO, croplands; DBF, deciduous broadleaf forests; ENF, evergreen needleleaf forests; GRA, grasslands; MF, mixed forests; SAV, savannas; SH, shrublands; WET, wetlands.

# 2.4 Modelling GPP using Chl<sub>leaf</sub>

The TBM used in this study is the Boreal Ecosystem Productivity Simulator (BEPS), which is a two-leaf enzyme kinetic model that has been intensively used to simulate carbon and water fluxes across different biomes (Gonsamo et al., 2013; Wang et al., 2004). It has participated in several cross-model validation studies and been proved to be one of the better performing models for GPP simulations (Grant et al., 2006; Schaefer et al., 2012).

Parameterization of BEPS was demonstrated in detail in previous studies (Chen, Liu, Cihlar, & Goulden, 1999; He et al., 2017; Luo, Chen, et al., 2018). A recent study has upgraded BEPS to integrate Chl<sub>leaf</sub> (a.k.a. BEPS-Chl<sub>leaf</sub>) in modelling fluxes in a deciduous broadleaf forest (Luo, Croft, et al., 2018).

Hourly meteorological variables (i.e. solar radiation, temperature, VPD, precipitation and wind speed) are the inputs for BEPS. The solar radiation and leaf temperature are separately calculated for sunlit and shaded leaves through a two-leaf scheme using LAI and CI (Chen et al., 1999). Leaf-level photosynthesis is then obtained from the embedded Farquhar biochemical model with the input of leaf traits.

Three cases were studied depending on the types of  $V_{\rm c}^{25}$  used in BEPS: (a) BEPS-TRY used constant  $V_{\rm c}^{25}$  provided by the TRY database (Kattge et al., 2009); (b) BEPS-Chl<sub>leaf</sub> used daily Chl<sub>leaf</sub>-based  $V_{\rm c}^{25}$ ; (c) BEPS-Chl<sub>avg</sub> used the average of daily Chl<sub>leaf</sub>-based  $V_{\rm c}^{25}$ .

After using the biochemical model to calculate leaf photosynthesis, stomatal conductance  $g_s$  is in turn derived from the Ball-Woodrow-Berry equation and used in the Penman-Monteith model to get leaf transpiration. In feedback, the changes in leaf energy budget modulate the leaf temperature and adjust the photosynthesis rate (Supporting Information, section 3).

After the leaf-level calculation, canopy-scale photosynthesis and transpiration are acquired by multiplying the leaf-level fluxes by the corresponding sunlit and shaded LAI through the two-leaf scheme. Soil texture is prescribed for each site using a global map (https://www.soilgrids.org/) to quantify the hydrological and thermal properties of soil and to simulate soil water content. We apply a linear soil stress factor derived from multilayer soil water content to simulate the stomatal response to possible droughts (Supporting Information, section 3).

Statistical analysis. We used squared Pearson correlation coefficient  $(r^2)$ , accumulated absolute bias (AAB) and root mean square error (RMSE) to evaluate the performance of models.

AAB was used to quantify the sum of daily GPP biases in a year, avoiding the potential offset of positive and negative daily biases:

$$AAB = \sum_{i=1}^{n} |A_i - B_i| \tag{2}$$

where n is the size of data population,  $A_i$  is the estimate and  $B_i$  is the observation.

#### 3 RESULTS

Our satellite-derived estimates of Chl<sub>leaf</sub> show different seasonal patterns for different PFTs (Figure 1). Chl<sub>leaf</sub> of DBF, ENF, SAV and WET mostly varied in

range from approximately 5 to 70 µg/cm² across a year with a single peak in the growing season; Chl<sub>leaf</sub> of CRO, MF and SH varied in a similar range showing two or more peaks; Chl<sub>leaf</sub> of EBF and GRA demonstrated less seasonal variability with the Chl<sub>leaf</sub> of EBF centred around 67 µg/cm² and the Chl<sub>leaf</sub> of GRA centred around 29 µg/cm². There were considerable intersite variations in Chl<sub>leaf</sub>, as we found the Chl<sub>leaf</sub> difference between the top quartile sites and the bottom quartile sites for each PFT ranged from 13 µg/cm² to 40 µg/cm² (Figure 1). We also noticed that there were temporal discrepancies shown between the development of LAI and that of Chl<sub>leaf</sub> for some PFTs (Figure 1): for DBF and CRO, the peak of Chl<sub>leaf</sub> appeared 4–5 weeks later than the peak of LAI, and LAI tended to change more dramatically than Chl<sub>leaf</sub> during a short time frame (i.e. during budburst and leaf fall); for GRA, Chl<sub>leaf</sub> was relatively stable in contrast to the clear seasonal pattern of LAI; for ENF and SH, Chl<sub>leaf</sub> showed multiple peaks in a year while LAI only peaked once.

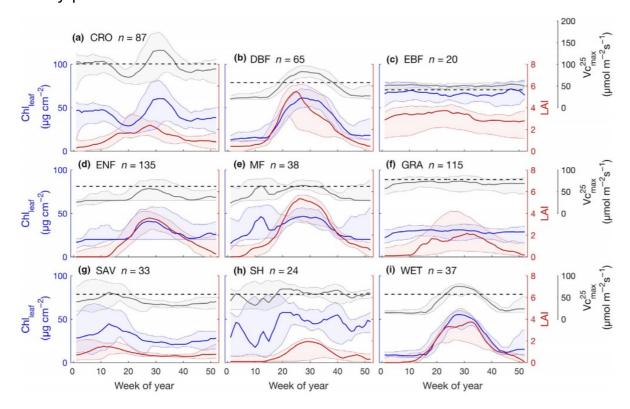


Figure 1. Seasonal dynamics of Chl<sub>leaf</sub> (blue), LAI (red) and Chl<sub>leaf</sub>-based Vc25 (solid black) for nine plant functional types (a-i). The plant functional types are croplands (CRO), deciduous broadleaf forests (DBF), evergreen broadleaf forests (EBF), evergreen needleleaf forests (ENF), mixed forests (MF), grasslands (GRA), savannas (SAV), shrublands (SH) and wetlands (WET). The dash lines (black)

indicate the constant based on the TRY database. *n* indicates the number of site-years for each plant functional type (PFT). The shadings indicate the spatial variations among site-years. In each panel, the solid line is the median of a variable for the group of site-years in a PFT, upper and lower boundaries refer to 75% and 25% percentiles of the variable respectively

We calculated  $V_{c,m,k}^{25}$  from Chl<sub>leaf</sub> using the PFT-specific equations that we compiled in Table 1. Chl<sub>leaf</sub>-based  $V_{c,m,k}^{25}$  followed the pattern of Chl<sub>leaf</sub> across the year (Figure 1). Chl<sub>leaf</sub>-based  $V_{c,m,k}^{25}$  varied seasonally and spatially in a range between 10 and 80 µmol m-2s-1 for most PFTs except for CRO where the Chl<sub>leaf</sub>-based  $V_{c,m,k}^{25}$  varied between 40 and 175 µmol m-2s-1. In comparison with  $V_{c,m,k}^{25}$  from the TRY database, we found that the TRY-based  $V_{c,m,k}^{25}$  were around the annual mean of Chl<sub>leaf</sub>-based  $V_{c,m,k}^{25}$  for CRO, DBF, SH and WET, and the TRY-based  $V_{c,m,k}^{25}$  generally were larger than their Chl<sub>leaf</sub>-based  $V_{c,m,k}^{25}$  for ENF, MF, GRA and SAV and vice versa for EBF.

We further used an established two-leaf enzyme kinetic TBM—the Boreal Ecosystem Productivity Simulator (BEPS) to estimate GPP using either Chl<sub>leaf</sub>-based (BEPS-Chl<sub>leaf</sub>) or TRY-based (BEPS-TRY) (see Material and Methods). While BEPS-TRY only used LAI to describe the variability of vegetation status, BEPS-Chl<sub>leaf</sub> considered both the variabilities of LAI and Chl<sub>leaf</sub> in the estimation of GPP. The bias of estimated GPP from BEPS-Chl<sub>leaf</sub> was much smaller, showing seasonal biases closer to zero in all weeks for all PFTs except for EBF and ENF (Figure 2). In spring and autumn, Chl<sub>leaf</sub> reduced the overestimation of GPP for CRO, DBF, GRA and WET and the underestimation for SAV; in summer, Chl<sub>leaf</sub> reduced the overestimation of GPP for CRO, GRA and SH and the underestimation for DBF, MF and WET.

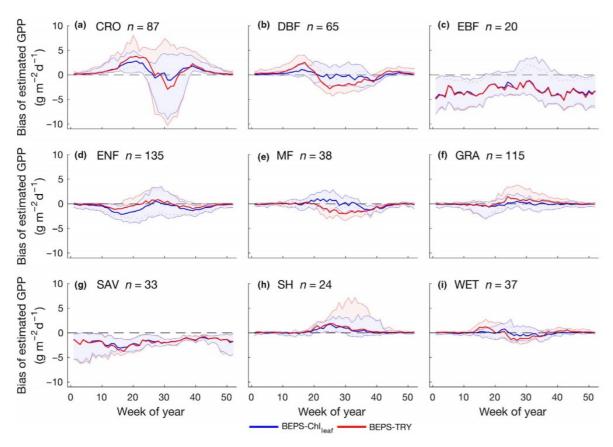


Figure 2. Temporal patterns of the bias of gross primary production estimated by BEPS-Chl $_{leaf}$  (blue) and BEPS-TRY (red) for nine plant functional types (a-i). The plant functional types are croplands (CRO), deciduous broadleaf forests (DBF), evergreen broadleaf forests (EBF), evergreen needleleaf forests (ENF), mixed forests (MF), grasslands (GRA), savannas (SAV), shrublands (SH) and wetlands (WET). n indicates the number of sites for each plant functional type. The shadings indicate the spatial variations among site-years. In each panel, the solid line is the median of a variable for the group of site-years in a PFT, upper and lower boundaries refer to 75% and 25% percentiles of the variable respectively

The results above showed that the inclusion of Chl<sub>leaf</sub> can reduce the spatial and temporal biases in GPP in some cases (Figure 2). To separate the impacts of spatial and temporal variations of Chl<sub>leaf</sub> on estimated GPP, we added the GPP estimates from BEPS-Chl<sub>avg</sub> to specifically evaluate the impact of the spatial variation of Chl<sub>leaf</sub> on GPP estimates, and then evaluated the impact of the temporal variation of Chl<sub>leaf</sub> based on the difference between BEPS-Chl<sub>leaf</sub> and BEPS-Chl<sub>avg</sub> (Figure 3).

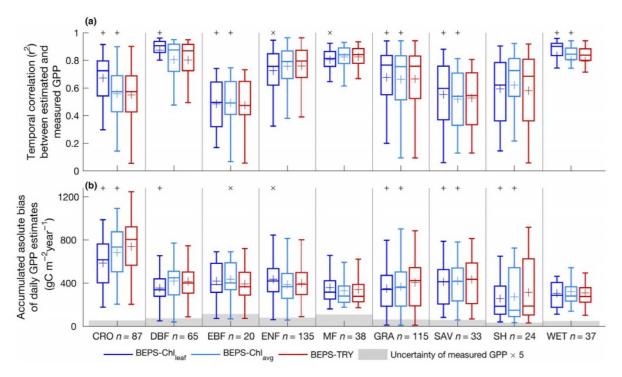


Figure 3. Comparison of daily gross primary production (GPP) estimated by BEPS-Chl<sub>leaf</sub>, BEPS-Chl<sub>lavg</sub> or BEPS-TRY and measured daily GPP for 555 site-years categorized by PFTs. (a)  $r^2$  between the estimated and the measured GPP; (b) accumulated absolute bias (AAB) of the estimated GPP. "+" indicates that BEPS-Chl<sub>leaf</sub> or BEPS-Chl<sub>avg</sub> significantly improved GPP estimates compared to those of BEPS-TRY (student's t test, p < 0.05) and "x" suggests BEPS-Chl<sub>leaf</sub> or BEPS-Chl<sub>avg</sub> results were significantly inferior to those of BEPS-TRY (student's t test, p < 0.05). Uncertainty of measured GPP (see Material and Methods) was multiplied by 5 to improve the visibility of shadings in the figure. Acronyms for PFTs are CRO (croplands), DBF (deciduous broadleaf forests), EBF (evergreen broadleaf forests), ENF (evergreen needleleaf forests), MF (mixed forests), GRA (grasslands), SAV (savannas), SH (shrublands) and WET (wetlands)

We found that the temporal correlation  $(r^2)$  between daily GPP estimates and measurements was improved from  $0.67 \pm 0.23$  to  $0.70 \pm 0.22$  after the inclusion of Chl<sub>leaf</sub> (Figure 3a). All PFTs except SH showed significant changes in  $r^2$  (student's t test, p < 0.05) after the inclusion of Chl<sub>leaf</sub>. The  $r^2$  for CRO, DBF and WET were improved the most from  $0.55 \pm 0.19$ ,  $0.80 \pm 0.17$  and  $0.79 \pm 0.17$  to  $0.67 \pm 0.17$ ,  $0.87 \pm 0.12$  and  $0.83 \pm 0.18$  respectively. In contrast, the  $r^2$  for EBF, ENF and MF decreased slightly. The inclusion of Chl<sub>avg</sub> fell short of improving the temporal correlation for most PFTs, though some PFTs like CRO, GRA, SAV and WET demonstrated significant (student's t test, p < 0.05) but slight increases in  $r^2$ .

The accumulated absolute bias (AAB) of GPP decreased by 4% from 551  $\pm$  303 to 527  $\pm$  286 gC m $^{-2}$  yr $^{-1}$  after adding Chl $_{\rm avg}$  to BEPS, and decreased by 8% to 507  $\pm$  261 gC m $^{-2}$  yr $^{-1}$  after adding Chl $_{\rm leaf}$  (Figure 3b). By using BEPS-Chl $_{\rm leaf}$ , CRO, DBF, GRA, SAV and SH demonstrated 21%, 15%, 13%, 6% and 19% of AAB reduction respectively (student's t test, p < 0.05). But we also noticed a slight 9% increase in AAB for ENF. Meanwhile, BEPS-Chl $_{\rm avg}$  demonstrated similar reductions in the AAB of GPP, only to a smaller degree than those using Chl $_{\rm leaf}$ : CRO, GRA, SAV and SH showed a significant

(student's t test, p < 0.05) decrease in AAB of 8%, 9%, 3% and 13% respectively. Considering the average AAB reductions across site-years were 4% for BEPS-Chl<sub>avg</sub> and 8% for BEPS-Chl<sub>leaf</sub>, we found that the spatial variation and temporal variation of Chl<sub>leaf</sub> contributed equally to the improvement of GPP estimates.

The reductions of daily GPP biases accumulated and led to the improvement of GPP estimates at the annual scale. We found that the spatial correlation  $(r^2)$  between estimated and measured annual GPP increased by 5% and RMSE reduced by 21% after the inclusion of  $Chl_{avg}$  in BEPS (Figure 4a,b). BEPS- $Chl_{leaf}$  further improved the estimates of GPP on top of BEPS- $Chl_{avg}$ , increasing  $r^2$  by 14% and reducing RMSE by 28% compared to BEPS-TRY (Figure 4a,c). Therefore, both the spatial and temporal variations in  $Chl_{leaf}$  improved GPP estimates. The RMSE reduction for annual GPP estimates (28%) appeared more substantial than the reduction of AAB for daily GPP estimates (8%), because the calculation of the former favoured those sites with larger changes in GPP estimates (Figure 4) while the later evaluated the percentage changes of all site-years equally (Figure 3).

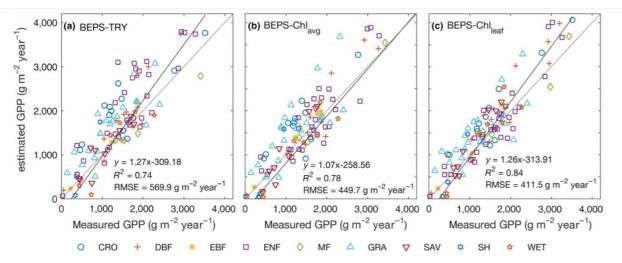


Figure 4. Correlation between measured and estimated annual gross primary production (GPP) of 124 eddy covariance sites. GPP are estimated from (a) BEPS-TRY, (b) BEPS-Chl<sub>avg</sub> and (c) BEPS-Chl<sub>leaf</sub>

The global GPP simulated by BEPS-Chl<sub>leaf</sub> and BEPS-TRY was 121 PgC and 130 PgC in 2011 respectively (Figure 5, Figure S3). The inclusion of  $Chl_{leaf}$  in BEPS brought about a 7% decrease of total global GPP, resulting from smaller GPP in the Northern Hemisphere but larger GPP in tropics and the Southern Hemisphere (Figure 5b). Around 40% of the vegetated surface GPP estimates changed more than 200 g m<sup>-2</sup> per year after the inclusion of  $Chl_{leaf}$ , and over 10% of the vegetated surface showed a GPP change larger than 500 g m<sup>-2</sup> per year.

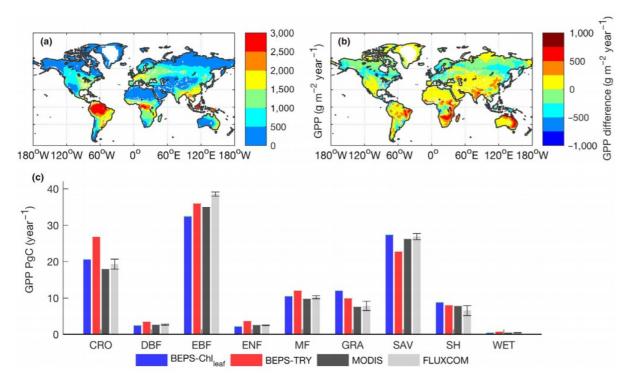


Figure 5. Global gross primary production (GPP) estimated by BEPS-Chl<sub>leaf</sub> and its comparison with other products. (a) spatial distribution of global GPP estimated by BEPS-Chl<sub>leaf</sub>; (b) difference between GPP estimated by BEPS-Chl<sub>leaf</sub> and BEPS-TRY; (c) PFT-specific GPP estimates from BEPS-Chl<sub>leaf</sub>, BEPS-TRY, MODIS and FLUXCOM. Acronyms for PFTs are CRO (croplands), DBF (deciduous broadleaf forests), EBF (evergreen broadleaf forests), ENF (evergreen needleleaf forests), MF (mixed forests), GRA (grasslands), SAV (savannas), SH (shrublands) and WET (wetlands)

Compared to BEPS-TRY, BEPS-Chl<sub>leaf</sub> reduced the annual GPP estimates by 6.1 Pg for CRO, 3.4 Pg for EBF and around 1 Pg for each of DBF, ENF and MF. Chl<sub>leaf</sub> increased the annual GPP estimates of SAV, GRA and SH by 4.7 Pg, 2.1 Pg and 0.9 Pg respectively (Figure 4c). Compared to other widely used global GPP data sets in 2011, the global GPP estimated by BEPS-Chl<sub>leaf</sub> was larger than MODIS GPP product (113 PgC) and close to the ensemble mean of GPP upscaled from eddy covariance measurements (FLUXCOM;  $122 \pm 8$  PgC).

BEPS-Chl<sub>leaf</sub> outperformed BEPS-TRY in simulating the variability of global GPP, based on the correlations between estimated annual GPP and RSIF or FLUXCOM GPP (Figure 6). The  $r^2$  between RSIF and estimated GPP increased 0.1 after using Chl<sub>leaf</sub> in BEPS, while the  $r^2$  between FLUXCOM GPP and estimated GPP increased 0.02. The RMSE between FLUXCOM GPP and estimated GPP also decreased by 5% after the inclusion of Chl<sub>leaf</sub>. There was a nonlinearity between the GPP estimated by BEPS-TRY and RSIF and FLUXCOM GPP, as BEPS-TRY tended to overestimate GPP at regions where FLUXCOM GPP was larger than 300 g m $^{-2}$  yr $^{-1}$  (Figure 6). But this overestimation of GPP by BEPS-TRY was corrected by using BEPS-Chl<sub>leaf</sub> (Figure 6).

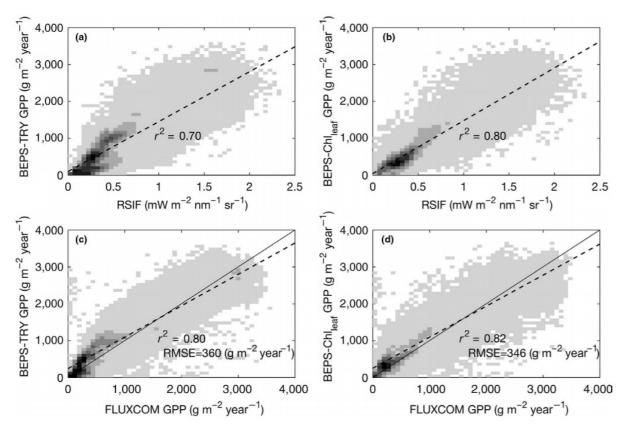


Figure 6. Validation of the spatial variability of annual gross primary production (GPP) estimated by BEPS-TRY and BEPS-Chl<sub>leaf</sub> against a reconstructed SIF product (RSIF) and the FLUXCOM gridded GPP product. (a) BEPS-TRY GPP versus RSIF; (b) BEPS-Chl<sub>leaf</sub> GPP versus RSIF; (c) BEPS-TRY GPP versus FLUXCOM ensemble mean GPP; and (d) BEPS-Chl<sub>leaf</sub> GPP versus FLUXCOM ensemble mean GPP. Density of points is indicated by grey scale with darker colour meaning higher density. Dash lines are the linear regression line and solid lines are the 1:1 reference line

Evaluation of estimated monthly GPP with RSIF and FLUXCOM GPP showed that  $Chl_{leaf}$  effectively improved GPP estimates in most months and for some PFTs (Figure 7). Based on the correlation between estimated GPP and RSIF (Figure 7a), CRO, DBF, GRA, SAV and WET were the five PFTs that improved  $r^2$  in most months, with the average  $r^2$  increased 0.06. EBF, ENF, MF and SH showed decline in  $r^2$  in most months. Meanwhile, the validation against FLUXCOM GPP (Figure 7b) confirmed the positive impacts of  $Chl_{leaf}$  in estimating GPP for CRO, DBF, GRA, SAV and WET. SH only showed increased  $r^2$  at the beginning of the year but had decreased  $r^2$  in most months. The comparison between estimated GPP and FLUXCOM GPP reaffirmed the negligible effect of  $Chl_{leaf}$  on EBF, but suggested considerable increases in  $r^2$  for ENF and MF (Figure 7b), which was not found in the validation against RSIF (Figure 7a) and eddy covariance measurements (Figure 3).

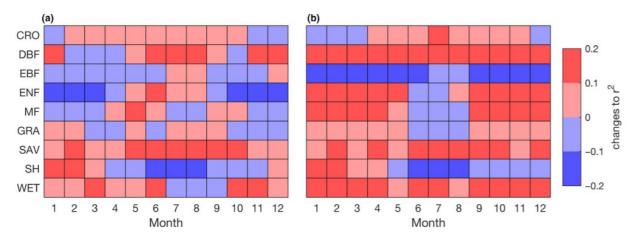


Figure 7. Evaluation of estimated monthly gross primary production (GPP) for each PFT using RSIF and FLUXCOM. (a) the increase or decrease of the correlation coefficient ( $r^2$ ) between estimated GPP and RSIF after the inclusion of Chl<sub>leaf</sub>; (b) the increase or decrease of the correlation coefficient ( $r^2$ ) between estimated GPP and FLUXCOM after the inclusion of Chl<sub>leaf</sub>. Acronyms for PFTs are CRO (croplands), DBF (deciduous broadleaf forests), EBF (evergreen broadleaf forests), ENF (evergreen needleleaf forests), MF (mixed forests), GRA (grasslands), SAV (savannas), SH (shrublands) and WET (wetlands)

#### 4 DISCUSSION

This study presents the first use of satellite-derived, temporal-spatially continuous plant physiological information (i.e. Chl<sub>leaf</sub>) to estimate GPP for globally distributed ecosystems. The inclusion of temporally and spatially explicit Chl<sub>leaf</sub> in a TBM reduced the bias of estimated daily GPP by 8% and improved the temporal correlation between estimates and observations by 5% on average for 124 eddy covariance sites (555 site-years). These improvements accumulated to a 28% reduction in RMSE of annual GPP estimates, and a 14% increase in the spatial correlation between annual estimates and observations. PFTs with strong seasonal cycles (i.e. DBF, CRO, GRA, SAV and WET) demonstrated significant improvements, with up to 21% reduction in bias and up to 22% increase in temporal  $r^2$ . We also noticed mixed impacts of Chl<sub>leaf</sub> on GPP estimates for SH and EBF and mostly negative impacts for ENF and MF. Based on the reduction of bias of daily GPP estimates, we found that the temporal and spatial variations of Chl<sub>leaf</sub> contributed equally to the improvements of GPP estimates. Global estimation of GPP was 9 PgC lower after the inclusion of Chl<sub>leaf</sub> in 2011 and better captured the variability of RSIF and FLUXCOM GPP.

# 4.1 Disparity in the temporal profiles of Chl<sub>leaf</sub> and LAI

Our results demonstrated a wide-spread disparity in the temporal profiles of  $Chl_{leaf}$  and LAI for all PFTs (Figure 1), which has been noted by several in situ studies on some tree species (Croft et al., 2017; Croft, Chen, & Zhang, 2014a; Kodani, Awaya, Tanaka, & Matsumura, 2002). This disparity highlights an asynchrony between the physiological and the structural development of plants. The reason for this disparity remains unidentified. It is possible that  $Chl_{leaf}$  and LAI develop different temporal profiles to optimally use limited resources (i.e. light), in a way similar to some leaf traits (i.e.

nitrogen, Verial) which develop a vertical profile mimicking the gradient of solar radiation in the canopy to maximize the total carbon uptake of a whole stand (Anten, Schieving, & Werger, 1995; Field, 1983; Hirose & Werger, 1987). In addition, we noticed that the degree of the asynchrony between Chl<sub>leaf</sub> and LAI varies among PFTs (Figure 1). For SH and ENF, the disparities between the temporal profiles of Chl<sub>leaf</sub> and LAI were relatively small, which may justify some modelling practices to use LAI to emulate the seasonality of Verial of ENF (Luo, Chen, et al., 2018; Ryu et al., 2011). However, for most PFTs, the temporal profiles of physiological status (i.e. Chl<sub>leaf</sub>) cannot be replaced by those of LAI in TBMs.

# 4.2 Comparison with previous studies using Chl<sub>leaf</sub> to estimate GPP

The positive impact of Chl<sub>leaf</sub> on GPP estimates found in our study is comparable to the results of previous site-level studies. Houborg et al. (2013) reported a 12% increase in temporal correlation ( $r^2$ ) and a 24% decrease in RMSE of estimated growing season GPP after using Chl<sub>leaf</sub> in the Community Land Model (Bonan et al., 2011) at a CRO site. Luo, Croft, et al. (2018) incorporated Chl<sub>leaf</sub> in BEPS at a DBF site and reported a 10% increase in temporal  $r^2$  and a 32% decrease in RMSE. In this study, we found the average improvement in the temporal  $r^2$  for 87 site-years of CRO and 65 siteyears of DBF was 22% and 9%, respectively, and their reductions in AAB, an indicator similar to RMSE, were 21% and 15% respectively (Figure 3). At the global scale, Alton (2017) used MERIS terrestrial chlorophyll index (MTCI) as a proxy for chlorophyll content to retrieve the  $V_{\rm c}^{25}$  of 296 FLUXNET sites and found the inclusion of spatially and temporally varying MTCI-based  $Vc_{max}^{25}$  in JULES-SF led to a 15% decrease in annual GPP estimates on site-average. Meanwhile, our study found a 7% decrease of annual GPP across 124 sites and a 7% decrease of global GPP. Our study took a somewhat different approach than the Alton (2017) study to incorporate Chl<sub>leaf</sub> into a TBM. The Alton (2017) study assumes a constant conversion rate from  $Chl_{leaf}$  to  $Vc^{25}$ for all leaves and PFTs, but several field-based studies have suggested that the Chl<sub>leaf</sub>-Ve<sup>25</sup> relationship is likely to change with PFTs (Croft et al., 2017; Houborg, McCabe, Cescatti, & Gitelson, 2015) and light environment (Evans, 1989a). Our result also suggests that MTCI, though widely used for carbon modelling (He et al., 2017; Yao Zhang, Xiao, et al., 2018), might not be able to capture the variability of global measured  $Chl_{leaf}(r^2 = 0.27)$  as good as our satellite-derived Chl<sub>leaf</sub> product ( $r^2 = 0.47$ ) (Supplementary Information, section 2). In addition, even though a site-level study has established a Chl<sub>leaf</sub> relationship for C4 species and reported improved photosynthesis estimates (Houborg et al., 2013), a lack of knowledge about the global distribution of C4 species in our PFT map prohibited us from applying this relationship for global GPP estimation.

4.3 Influence of light environment on Chl<sub>leaf</sub>-Vc<sup>25</sup><sub>max</sub> relationship

The ratio between Chl<sub>leaf</sub> and Vc25 can differ between sunlit and shaded leaves, with sunlit leaves showing a higher Verice to Chlieaf ratio than that of shaded leaves (Evans, 1989a). It is possible that such changes also reflect the optimal distribution of the nitrogen within the canopy to maximize the whole canopy photosynthesis (Anten et al., 1995). This difference in the  $Vc_{max}^{25}$  to  $Chl_{leaf}$  ratio between sunlit and shaded leaves is consistent with previous studies reporting Chl<sub>leaf</sub> varies little under different solar irradiance (Lambers, Chapin, & Pons, 2008; Walters, 2005) while leaf nitrogen content and  $V_{\rm c}^{25}$  both decrease proportionally from the top to the bottom of the canopy following radiation gradient (Hirose & Werger, 1987; Warren & Adams, 2001). In this study, we incorporate these observational patterns in an updated two-leaf upscaling scheme for BEPS (Luo, Croft, et al., 2018) in which we assume the sunlit and shaded leaves have Chl<sub>leaf</sub> similar to the Chl<sub>leaf</sub> of leaves on top of the canopy, and  $V_{c_{max}}^{25}$  of sunlit and shaded leave can be calculated based on a presumed nitrogen gradient (De Pury & Farquhar, 1997) and the Chl<sub>leaf</sub>-based Ve<sup>25</sup> of top leaves. The results from that site-level study have suggested that the physiological difference between sunlit and shaded leaves was effectively reproduced by this updated two-leaf scheme (Luo, Croft, et al., 2018). Although our approach assumes that leaf nitrogen content is not a robust proxy for the temporal profile of  $Ve^{25}$  (Croft et al., 2017) because of the dynamic allocation of nitrogen to photosynthetic and nonphotosynthetic components, leaf nitrogen content remains a good proxy for the vertical profile of Vertical (Wilson. Baldocchi, & Hanson, 2000).

# 4.4 Influence of leaf age on $Chl_{leaf}$ relationship

Our results demonstrated mixed impacts of Chlieaf on the estimation of GPP for EBF and negative impacts of Chl<sub>leaf</sub> for ENF and MF. These three PFTs all have evergreen species that contain leaves or needles of several age groups. Katahata, Naramoto, Kakubari, and Mukai (2007) reported that needle leaves belonging to ages of 1, 2 and 3 years showed different seasonal trends in Chl<sub>leaf</sub> and Vc<sup>25</sup>. Young needles had an increasing trend in Chl<sub>leaf</sub> while old needles showed the opposite. Warren and Adams (2001) and Ethier et al. (2006) both found reduced photosynthetic capacity for old needles while their leaf nitrogen content is almost the same as young needles. As for EBF, a study in Amazonia demonstrated that leaf age composition explains 27% of the variation in photosynthesis, and that young leaves have higher light use efficiency than the old leaves (Wu et al., 2016). These studies indicate that leaf age is potentially another factor driving the variation in  $\mathrm{Chl}_{\mathsf{leaf}}$  and  $V\mathrm{c}_{\mathsf{max}}^{25}$ within a canopy; however, the role of leaf age was not at all accounted for in our algorithms to derive Chl<sub>leaf</sub> or Chl<sub>leaf</sub> relationships. A recent study has managed to estimate leaf ages using the PLSR model driven by leaf

reflectance for some Amazonian trees (Chavana-Bryant et al., 2017), implying the possibility of deriving leaf age and its corresponding  $Chl_{leaf}$  from satellite data in the future. We suggest that a proper consideration of leaf demography and its impact on  $Chl_{leaf}$  and for evergreen species is necessary to simulate GPP with  $Chl_{leaf}$  for EBF, ENF and MF.

4.5 Influence of Chl<sub>leaf</sub> on fraction of absorbed photosynthetic active radiation (fAPAR)

Many leaf-level studies have found that absorptance of photosynthetic active radiation (PAR; 400-700 nm) of leaves increases with the increase of Chl<sub>leaf</sub> (Evans & Poorter, 2001; Poorter et al., 2000), suggesting that Chl<sub>leaf</sub> is a major factor impacting fAPAR of leaves. However, our study did not explicitly consider the impact of Chl<sub>leaf</sub> on fAPAR, due to the fact that at the canopy scale Chl<sub>leaf</sub> can only impact fAPAR to a small degree. The limited impact of Chl<sub>leaf</sub> on canopy fAPAR was suggested by observational and model results (Daughtry, 2000; Migliavacca et al., 2017). Layers of leaves in a canopy leverage reflected and scattered solar radiation within the canopy, and subsequently render the canopy fAPAR much higher and less variable than its leaf counterpart (Croft, Chen, & Zhang, 2014b). Therefore, it is the canopy structure (i.e. LAI, clumping, leaf angle distribution) that dictates canopy fAPAR (Asner, 1998; Asner & Wessman, 1997) rather than Chl<sub>leaf</sub>. A sensitivity test of our radiative transfer model shows a sevenfold increase in Chl<sub>leaf</sub> only led to around 1%-5% of increase in canopy fAPAR (Croft et al., 2014b), which is negligible compared to the impact of  $\mathrm{Chl}_{\mathrm{leaf}}$  on  $\mathrm{Vc}^{25}$  (i.e. sevenfold increase in Chl<sub>leaf</sub> means 700% increase in Vc25]). The changes in Chl<sub>leaf</sub> may impact the amount of APAR distributed between leaves but the total APAR for the canopy is likely not sensitive to the changes in Chl<sub>leaf</sub>. It explains why most studies incorporating Chl<sub>leaf</sub> in TBMs only consider the impact of Chl<sub>leaf</sub> on Vc25 (Alton, 2017; Houborg et al., 2013; Luo, Croft, et al., 2018) and light use efficiency (Croft, Chen, Froelich, Chen, & Staebler, 2015; Croft, Chen, Zhang, et al., 2015; Houborg, Anderson, Daughtry, Kustas, & Rodell, 2011) and neglect the impact of Chl<sub>leaf</sub> on APAR. However, we acknowledge that the impact of Chl<sub>leaf</sub> on canopy fAPAR, though usually small, needs to be addressed in some special cases (i.e. planophile canopies) where leaf optical properties likely impact canopy fAPAR considerably.

Chlorophyll is vital to all plants on the Earth's land surface by harvesting light and transporting electrons to support the production of biochemical energy necessary to drive photosynthesis and maintaining ecosystem services. Chl<sub>leaf</sub> is an essential vegetation trait that impacts the cycling of carbon, water and energy between the terrestrial biosphere and the atmosphere. We found that the inclusion of Chl<sub>leaf</sub> can effectively constrain the uncertainty of modelled photosynthesis for various PFTs, particularly those with strong seasonal cycles, and provide a more realistic estimation of global GPP. We

suggest that Chl<sub>leaf</sub> is a valuable leaf physiological trait to add in future TBMs to better simulate the terrestrial carbon cycle.

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