

# AMERICAN METEOROLOGICAL SOCIETY

Journal of Atmospheric and Oceanic Technology

## EARLY ONLINE RELEASE

This is a preliminary PDF of the author-produced manuscript that has been peer-reviewed and accepted for publication. Since it is being posted so soon after acceptance, it has not yet been copyedited, formatted, or processed by AMS Publications. This preliminary version of the manuscript may be downloaded, distributed, and cited, but please be aware that there will be visual differences and possibly some content differences between this version and the final published version.

The DOI for this manuscript is doi: 10.1175/JTECH-D-13-00083.1

The final published version of this manuscript will replace the preliminary version at the above DOI once it is available.

If you would like to cite this EOR in a separate work, please use the following full citation:

Fietzek, P., B. Fiedler, T. Steinhoff, and A. Körtzinger, 2013: In situ quality assessment of a novel underwater pCO2 sensor based on membrane equilibration and NDIR spectrometry. J. Atmos. Oceanic Technol. doi:10.1175/JTECH-D-13-00083.1, in press.

© 2013 American Meteorological Society



<ul> <li><i>p</i>CO<sub>2</sub> sensor based on membrane equilibration and NDIR spectrometry</li> <li><i>p</i>er Fietzek<sup>1</sup></li> <li></li></ul>
<ul> <li>and NDIR spectrometry</li> <li>and NDIR spectrometry</li> <li>Peer Fietzek<sup>1</sup></li> <li>GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany</li> <li>and CONTROS Systems &amp; Solutions GmbH Kiel, Germany</li> </ul>
4 5 6 7 Peer Fietzek <sup>1</sup> 8 GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany 9 and CONTROS Systems & Solutions GmbH Kiel, Germany
5 6 7 Peer Fietzek <sup>1</sup> 8 GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany 9 and CONTROS Systems & Solutions GmbH Kiel, Germany
<ul> <li>6</li> <li>7 Peer Fietzek<sup>1</sup></li> <li>8 GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany</li> <li>9 and CONTROS Systems &amp; Solutions GmbH Kiel, Germany</li> </ul>
<ul> <li>Peer Fietzek<sup>1</sup></li> <li>GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany</li> <li>and CONTROS Systems &amp; Solutions GmbH Kiel, Germany</li> </ul>
8 GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, Germany and CONTROS Systems & Solutions GmbH Kiel, Germany
0 and CONTROS Systems & Solutions (Impher Kiel (Iermany
<sup>3</sup> and CONTROS Systems & Solutions Oniori, Rici, Ocimany
10 11 Biörn Fiedler, Tobias Steinhoff and Arne Körtzinger
12 GEOMAR Helmholtz Centre for Ocean Research Kiel Kiel Germany
13
PRELIMIAR

 <sup>1</sup> Corresponding author address: Peer Fietzek, GEOMAR Helmholtz Centre for Ocean Research Kiel, Düsternbrooker Weg 20, 24105 Kiel, Germany.
 E-mail: <u>pfietzek@geomar.de</u>

14 15 16	Abstract	
16	We present a detailed quality assessment of a novel underwater sensor for the	
18	measurement of $CO_2$ partial pressure ( $pCO_2$ ) based on surface water field deployments	
19	carried out between 2008 and 2011. The commercially available sensor, which is based	
20	on membrane equilibration and NDIR spectrometry is small and can be integrated into	
21	mobile platforms. It is calibrated in water against a proven flow-through $pCO_2$ instrument	
22	within a custom-built calibration setup. The aspect of highest concern with respect to	
23	achievable data quality of the sensor is the compensation for signal drift inevitably	
24	connected to absorption measurements. We use three means to correct for drift effects: (i)	
25	a filter correlation or dual-beam setup, (ii) regular zero gas measurements realized	
26	automatically within the sensor and (iii) a zero-based transformation of two sensor	
27	calibrations flanking the time of sensor deployment.	
28	Three sensors were tested against an underway $pCO_2$ system during two major	
29	research cruises providing an in situ temperature range from 7.4 to $30.1^{\circ}$ C and $p$ CO <sub>2</sub>	
30	values between 289 and 445 $\mu$ atm. The average difference between sensor and reference	
31	$pCO_2$ was found to be -0.6 ± 3 µatm with a RMSE of 3.7 µatm.	
32		

#### 33 1. Introduction

34

valuable for a large number of scientific, industrial as well as socio-economic issues. 35 36 Major scientific interest is related to the anthropogenic increase of atmospheric  $CO_2$ 37 concentrations and the resulting oceanic uptake of this most important anthropogenic 38 greenhouse gas (Sabine et al. 2004; Rogner et al. 2007). The exchange of  $CO_2$  across the 39 air-sea interface and the dynamics and trends of the carbon cycle in coastal and open 40 ocean regions (Takahashi et al. 2009; Gruber et al. 2010) as well as in the interior ocean 41 are key aspects of current marine carbon cycle research. 42 Dissolved  $CO_2$  reacts with water to form carbonic acid ( $H_2CO_3$ ), which rapidly dissociates into hydrogen ( $H^+$ ), bicarbonate ( $HCO_3^-$ ), and carbonate ions ( $CO_3^{2^-}$ ). The 43 44 exact speciation within the marine  $CO_2$  system, i.e. between the above species of the 45 carbonic acid dissociation system, strongly affects the pH of seawater (Millero 2007) and 46 is of major interest. An increasing amount of dissolved  $CO_2$  therefore leads to a 47 decreasing pH, a process also referred to as ocean acidification in the context of the 48 anthropogenic  $CO_2$  transient. Its impact on calcifying organisms as well as on the 49 physiology and reproduction of other marine species is presently not well understood 50 (Doney et al. 2009). The potential long-term influence of acidification on fisheries but 51 also carbon capture and underwater storage scenarios (IPCC 2005) show the socio-52 economic relevance of CO<sub>2</sub> measurements and highlight the demand for baseline 53 monitoring of  $CO_2$  parameters in the ocean. 54 Since the thermodynamic constants of the marine carbonate systems are known 55 rather precisely (e.g., Millero 2007), the marine  $CO_2$  system can be fully determined by

The measurement of dissolved carbon dioxide  $(CO_2)$  in seawater is important and

56	measurement of any two of the following variables: dissolved inorganic carbon (DIC),
57	total alkalinity (TA), pH and CO <sub>2</sub> fugacity ( $f$ CO <sub>2</sub> ) or CO <sub>2</sub> partial pressure ( $p$ CO <sub>2</sub> ; Millero
58	2007). Although the sole measurement of $pCO_2$ is not sufficient to fully characterize the
59	marine $CO_2$ system, $pCO_2$ is still a useful parameter on its own: It is the determining
60	factor for air-sea gas exchange and responds sensitively to biogeochemical processes
61	such as photosynthesis and respiration. Hence $pCO_2$ is both, a valuable stand-alone
62	measurement parameter and a useful measured variable within multi-parameter
63	measurements for the determination of the $CO_2$ system. So far only $pCO_2$ and pH can be
64	measured autonomously with commercial underwater sensors (DeGrandpre et al. 1995;
65	Seidel et al. 2008; Martz et al. 2010). Sensors for autonomous measurements of the other
66	carbonate system parameters TA and DIC (Byrne et al. 2002; Wang et al. 2007) and even
67	for direct measurements of carbonate ions are under development (Byrne and Yao 2008).
68	Due to a lack of a commercially available underwater $pCO_2$ sensor with a
69	sufficient accuracy and platform integratability expressed by adequate dimensions, an
70	appropriately short response time at surface and at depth as well as by the ability for
71	continuous measurements, we have advanced the development of an autonomous and
72	commercially produced underwater $pCO_2$ sensor (HydroC-CO <sub>2</sub> , CONTROS GmbH, Kiel,
73	Germany). Its versatility and specifications allow for integration into various platforms
74	and provide a suitable tool for $pCO_2$ measurements at an improved spatial and temporal
75	resolution in the water column. Here, we present a comprehensive overview of the
76	sensor's measurement principle, its key components and its calibration. We also assess
77	the achievable data quality by means of a detailed analysis of data from surface water
78	field applications. Measurements against a proven flow-through system only represent

one application of the sensor, but at the same time enable the clearest assessment of the
sensors data quality. Many further applications in the water column are possible and best
require a dedicated discussions such as the measurements on a profiling float presented in
Fiedler et al. (2013).

83 a. Suitable sensor platforms

84 Beside classical mobile sensor platforms on the one hand, such as voluntary observing 85 ships (VOS) or research vessels, and stationary buoys and moorings on the other hand, 86 innovative mobile platforms are receiving growing interest (Fietzek et al. 2011). These 87 platforms, e.g. autonomous underwater vehicles (AUV), profiling floats and gliders, 88 provide an increased autonomy, mobility and versatility as opposed to the classical 89 carriers. They allow for a more cost-efficient data collection on so far largely unexplored 90 temporal and spatial scales. These newly accessible scales are of high interest for the 91 investigation of various biogeochemical processes, making these modern mobile 92 platforms desirable carriers for  $pCO_2$  sensors (Gruber et al. 2010). By equipping 93 autonomous platforms as a regional or global sensor array with high spatiotemporal 94 resolution, such as the prominent Argo project, a high scientific potential can be achieved 95 (Roemmich et al. 2009; Johnson et al. 2009; Fiedler et al. 2013). However, in order to be 96 easily installed on modern mobile platforms, new sensors need to cope with the demands 97 of the more complex carriers and fulfill the more stringent requirements with respect to 98 payload capacity, power availability, response time, etc. A more detailed discussion of 99 these facts and circumstances can be found in Fietzek et al. (2011).

101 The first measurements of dissolved  $CO_2$  were developed for medical applications in the 102 1950's. These were based on wet-chemical pH-determination behind a Teflon<sup>™</sup> 103 membrane (Stow et al. 1957; Severinghaus and Bradley 1958). Today optical CO<sub>2</sub> 104 measuring techniques are prevailing. One method is to detect and quantify  $CO_2$  molecules 105 within an equilibrated gas stream by means of direct absorption in the infra red (IR) 106 region of the electromagnetic spectrum. Another technique is an indirect measurement 107 making use of the pH affecting property of CO<sub>2</sub> by applying spectrophotometry within an 108 equilibrated pH-sensitive dye solution of known characteristics (DeGrandpre et al. 1995; 109 Lefèvre et al. 1993). An overview of current sensor techniques for carbonate system 110 species can be found in Byrne et al. (2010). A technical evaluation of  $pCO_2$  sensors that 111 also includes two sensors of the type discussed here is presented in Tamburri et al. 112 (2011). Between the evaluation and the measurements discussed herein the status of the 113 sensors was mainly improved through an optimized calibration process and data 114 processing. 115 The usage of underway flow-through instruments to measure  $pCO_2$  both in the 116 oceanic surface layer and in the atmosphere dates back to the 1960s (Takahashi 1961). 117 While in the beginning the application of these systems was restricted to research vessels, 118 current, improved systems are suitable also for application on unattended platforms such 119 as VOS (Watson et al. 2009; Pierrot et al. 2009; Pfeil et al. 2013). Continuous 120 optimization of the overall setup and the components used lead to some standardized 121 design that is nowadays also commercially available. The key components of such a 122 flow-through instrument are the air-seawater equilibrator and the IR gas analyzer. The

126 seawater  $pCO_2$  measurements (Pierrot et al. 2009). The flow-through system data are

127 commonly reported as  $pCO_2$  although the  $fCO_2$  is the value suggested for most accurate

128 carbonate system calculations. Based on approximated expressions the *f*CO<sub>2</sub>, which

129 considers the slightly non-ideal behavior of CO<sub>2</sub> in the gas phase, can be calculated from

130 temperature and  $pCO_2$ , which presumes ideal gas behavior.

### 131 2. Sensor principle and description

#### 132 a. Development and design

123

124

125

133 The development aims of the  $pCO_2$  sensor among others were to obtain (i) a versatile and 134 autonomous sensor that could be deployed on a profiling float with (ii) a response time of 135 less than 5 min, (iii) an accuracy better 5 µatm, and (iv) a stability and reliability that 136 would allow for long-term deployments of several months.

137 Since the developed sensor is based on the same measuring principle as proven 138 flow-through systems, it has the same key components: an equilibrator and an IR CO<sub>2</sub> 139 detector. A planar, semi-permeable membrane with a silicone active layer is installed in 140 the head of the sensor. It acts as an equilibrator as well as a phase separator between the 141 ambient water and an internal headspace. The sensor is commonly equipped with a water 142 pump that provides a continuous seawater flow to the membrane and thus reduces the 143 thickness of the static boundary layer in front of the membrane. By that, the response 144 time is effectively shortened and made independent of a relative movement between the

145 membrane and the surrounding water. In order to withstand high hydrostatic pressures the 146 membrane is mechanically supported from behind with a sintered metal disc. A gas pump 147 continuously circulates air between the membrane equilibrator and a non-dispersive IR 148 detector (NDIR). Figure 1 shows a model of the sensor and provides a schematic 149 overview of its setup. The gas tightness of the internal gas stream as well as of the 150 integrated values is checked thoroughly prior to calibration. Opposed to  $pCO_2$  underway 151 systems in which gas stream leakages are a major source for measuring errors (Pierrot et 152 al. 2009), the biggest "leak" within the gas stream of this sensor remains the equilibration 153 membrane; the high gas permeability of the membrane related to the volume of the gas 154 stream compensates for possible influences caused by small leaks. If bigger leakage 155 occurs within the gas stream, the entire, much larger internal gas volume of the sensor 156 will be equilibrated, leading to noticeably slower response times but not necessarily 157 biases. The gas circuit also features a specially developed gas heater upstream of the 158 NDIR detector whose heating control system is also used to stabilize the temperature of 159 the IR detector. The gas heater buffers seawater temperature gradients ( $\Delta T_{\text{in-situ}}$ ) in such a 160 way that large  $\Delta T_{\text{in-situ}}$  are damped to a much smaller gas temperature gradient. We 161 choose to set the control temperature just high enough for the heating control circuit to 162 keep the controlled temperature stable even at the maximum in situ temperature expected 163 during deployment. When the control temperature is set higher than necessary, avoidably 164 high power consumption is the consequence and the abundant absolute temperature 165 differences between the surrounding water and the internal gas becomes larger. In 166 addition to the gas heater and the temperature stabilization, the sensor's pressure housing 167 is thermally insulated and temperature sensitive components are separately protected.

168 Within the gas stream sensors for pressure, temperature and relative humidity 169 (RH) are installed to determine the conditions within the NDIR detector as well as behind 170 the membrane. Their exact position was chosen upon laboratory tests to be most 171 beneficial for their consideration within the IR sensor data analysis. The quality and a 172 deep understanding of the NDIR detector are crucial for the data quality of both, 173 underway instruments as well as the new underwater sensor. All additional components 174 within the underway instrument's and the underwater sensor's gas circuit beside the 175 equilibrator and NDIR detector, such as the additional sensors mentioned above, are 176 required for accurate and precise absorption measurements and allow for preferably long 177 deployments. 178 In contrast to common practice of underway instruments the absorption

measurement within the sensor is carried out in wet air and without interrupting the gas flow for measurement. In the underwater sensor regular zero gas measurements can be carried out. Therefore valves are included into the circuit that lead the pumped air through a soda lime cartridge instead of the membrane equilibrator at desired intervals (see Fig. 1). In the presence of water vapor soda lime scrubs the  $CO_2$  binding it as calcium carbonate (CaCO<sub>3</sub>) thus creating a zero gas with respect to  $CO_2$ .

The sensor operates by consecutively switching through different intervals, the durations of which can be individually set. As soon as the sensor is powered it starts with a warm-up interval followed by continuous repetition of measuring cycles. One measuring cycle consists of three intervals: zero, flush and measure. The warm-up interval is only passed through once after the sensor has been turned on. The required warm-up time depends on the water temperature and the supplied voltage (c.f. Table 1).

191 During the warm-up, the water pump is disabled and data are neither transmitted nor 192 stored in the internal logger. During zero intervals, a zero CO<sub>2</sub> gas is created as described 193 above and the sensor provides the current zero reading used later for drift correction. The 194 zeroing typically lasts a few minutes and repetition is recommended at least every 12 h. 195 Data stored during that time are flagged for easy handling during analysis. Flush intervals 196 are only used to flag data acquired during the signal recovery from the zero value to the 197 ambient  $pCO_2$  reading. Technically the measuring process of the sensor does not differ 198 between the flush and the subsequent measuring interval. The time the sensor needs for 199 full equilibration depends on the sensor configuration and the environmental conditions, 200 mainly the water temperature. The warmer the water, that faster the response time and 201 thus the shorter the required flush time. Laboratory and field tests for the response time 202 determination of the sensor as configured within the deployments presented here indicate 203 a linear dependency of response time on water temperature at a slope of the order of -1 s 204 per 1  $^{\circ}$ C (data not shown here). Sensor response to a step input can be well described by 205 first order kinetics and a corresponding exponential fit. Any response time given here 206 hence represents a time constant or a  $t_{63\%}$  depending on which formulation is favored. 207 Isothermal pressure vessel experiments up to 200 bar have neither shown indications for a 208 pressure hysteresis, nor could a significant pressure influence on the response time be 209 identified (data not shown here). The actual response time of the sensor can be derived 210 based on the course of the signal recovery during the flush interval as applied by Fiedler 211 et al. (2013). It is typically faster than the 2 min response time of the flow-through system 212 (Pierrot et al. 2009) and allows for measurements on moving platforms; especially when 213 a response time correction is applied to derive the "true" ambient  $pCO_2$  from the time214 lagged sensor signal (Fiedler et al. 2013). A more detailed analysis of the sensor's 215 response characteristics is in preparation. During flush and measurement intervals, the 216 water pump is active and data are recorded as configured. The interval settings simplify 217 the data processing and provide the means to generate measurement data from a fully 218 equilibrated and internally temperature stabilized sensor. Figure 2 shows the sensor signal 219 during calibration with the zero, flush and measurement intervals indicated.

220 The comparatively fast response time, small size and operability of the sensor 221 allow for deployments on various platforms. Due to its design and compared to classical 222 flow-through systems the maintenance intensity and the risk for leaks in the gas stream 223 are low. If used in situ, the sensor lacks the demand for an accurate water temperature 224 probe as it is crucial for typical flow-through systems (Körtzinger et al. 2000). In order to 225 derive the actual amount of dissolved  $CO_2$  from the measured partial pressure, the sensor 226 is commonly deployed together with a CTD probe, as the solubility of CO<sub>2</sub> depends on 227 temperature and salinity (Weiss 1974).

228 b. Specifications

Table 1 lists the specifications of the developed sensor as applicable for the

230 measurements discussed in this work. Different data communication options are feasible.

An internal data logger is optional that can either be used as a stand-alone memory (e.g.

232 Saderne et al. 2013) or for backup storage (e.g. Fiedler et al. 2013). A sleep mode

233 function further facilitates autonomous installations. The sensor development also

234 comprised a surface water flow-through version of the sensor which is not further

addressed here.

237 The properties of the NDIR detector are relevant for the overall performance of the 238 sensor. Drift, cross-sensitivities and the signal-to-noise ratio (S/N) of the NDIR detector 239 directly affect the data quality. 240 NDIR spectrometry in general is a proven direct measuring technique for mole 241 fractions of gases absorbing in the IR. It is non-destructive and traceable to standards. 242 NDIR detectors for CO<sub>2</sub> have small dimensions (several cm) and moderate power 243 consumption (here: about 0.5 W) that allow for easy integration. In addition they provide 244 good mechanical strength and are unaffected by vibrations if realized without moving 245 parts. Due to their high selectivity and limited cross sensitivity NDIR detectors are well 246 suited for qualitative analysis. NDIR detectors are composed of three main components: light source, absorption/beam path and detector. Each of them has different influences on 247 248 the final sensor signal and depends differently on environmental variables, e.g. 249 temperature. This may lead to complex overall sensor properties. Their choice defines the 250 S/N, the sensitivity and the measurement range (detection limit and upper range value) of 251 an NDIR unit. The basic idea is to get enough light energy at the desired wavelength to 252 the detector and to make the absorption path sufficiently long such that changes in 253 absorbed light intensity can be clearly resolved by the detector and amplifying 254 electronics. Various options exist to optimize and dovetail these components. 255 The linear relation between the transmitted light intensity, *I*, and molecule 256 concentration, c, as described by the Beer–Lambert Law makes the technique suitable for 257 quantitative analysis:

 $I = I_0 10^{-\varepsilon cl} , \qquad (1)$ 

with  $I_0$  being the initial light intensity, l the distance the light travels through the absorbing medium and  $\varepsilon$  the molar absorption or extinction coefficient of the target molecule to be detected. According to (1) the extinction or absorbance, A, is defined to be directly proportional to the molecule concentration in the medium:

263 
$$A = \log\left(\frac{I_0}{I}\right) = \varepsilon c l \,. \tag{2}$$

In reality there is a small non-linearity between the absorbance measured by the NDIR detector and molecule concentration due to the fact that the Beer–Lambert law is only defined for a single wavelength of infinitesimal small width, while in applications spectral dependencies occur (Wiegleb et al. 2001).

268 The NDIR unit used features a black body radiator as a broad band IR light source 269 and a pyro-electrical IR detector which requires pulsed operation of the IR emitter. 270 Interference filters in front of the detectors select the desired measurement as well as 271 reference wavelength and together with electronics for control and evaluation complete 272 the single beam dual wavelength NDIR detector.  $CO_2$  is typically sensed around 4.26  $\mu$ m 273 where it shows its highest absorption due to its fundamental asymmetric stretching vibration (2349.1 cm<sup>-1</sup>) and the reference wavelength is chosen to be placed in the water 274 275 vapor window at around 4.0 µm.

Within an NDIR detector, several temperature influences exist that either require a compensation, calibration or stabilization. The spectral properties of the filters are temperature dependent: The central filter wavelength can shift as well as the transmission width can vary. A temperature influence on the sensor can occur in the form of thermal noise, a thermal background signal and changes in sensitivity. In the case of a dual beam setup the influences might even be different for both channels. The emission properties 282 and emitted intensities of the light source show a temperature influence as well. Thermal 283 expansion of the cuvette or other mechanical deformations of the absorption path may 284 have an additional effect on the measurements. Finally temperature dependencies of the 285 analog electrical components directly behind the detector need to be considered. The 286 overall effect caused by temperature changes is hence both variable in magnitude and 287 sign for measurement channels of two similar products as well as for a measurement and 288 reference channel within the same instrument with separate filters, detectors and 289 electronics. For this reason and in order to enhance the measurement quality, the entire 290 NDIR detector is temperature stabilized within the sensor. An active temperature 291 stabilization furthermore helps to reduce the required warm-up time present in any NDIR 292 detector due to self heating effects. A separation of the light source and the detector from 293 the gas stream by windows enhances the temperature stabilization capability and protects 294 the sensitive detector with the filter. A gas heater further reduces the temperature gradient 295 within the cuvette and simultaneously reduces the risk of condensation within the optical 296 components of the sensor.

Any sensor based on an absorption principle such as an NDIR detector senses the highest raw signal in the case of a complete absence of the target molecules in the beam path as in that case no absorption occurs and the maximum radiation intensity reaches the detector (see (1)). Therefore the regular determination of the sensor signal of a zero gas, the *zeroing*, is essential to account for drift effects that alter the light intensity with time and that otherwise would be erroneously interpreted as changes in target gas concentration. Typical effects are:

304

(i) Intensity variations or spectral shifts of the light source over time,

305 (ii) Contamination of any component within the beam path that might cause
306 shadowing or growing reflectivity losses in the cuvette,

307 (iii) Aging effects that alter the detector sensitivity over time,

308 (iv) Changes in the pre-amplifier gain of the detector.

309 Within the sensor the zeroing does not only account for long-term drift influences

310 but also for changes of the measurement conditions such as large changes in water

311 temperature that cause internal temperature gradients and different water vapor

312 concentrations within the gas stream.

313 Drift compensation by means of a differential setup in our case realized in the 314 form of filter correlation is referred to as a two-beam/two-wavelength method. It is 315 supposed to compensate any unwanted influences that cause signal drift of both channels 316 in the same manner (aspects (i) and (ii) above) as the measured signal is continuously 317 referenced. Any effects that cause changes in the detector signals and that are not caused 318 by actual concentration changes within the cuvette shall be compensated in real time and 319 parallel to measurements. In reality, this technique has its limitations in accounting for 320 influences resulting from the usage of the two different channels with their own filters 321 and detectors. Theoretically the zero-point of a two-beam instrument should not be 322 affected by the above mentioned drift reasons ((i) through (iv)). But as we measure at two 323 different wavelengths, spectral differences as well as effects related to the two physically 324 different detectors still affect the zero signal of the "two-beam corrected" signal. 325 Therefore we combine the zeroing and the two-beam drift correction means within our 326 sensor. The latter provides a continuous correction applied parallel to measurements, 327 while the zeroing discontinuously further enhances the drift correction capabilities by

correcting for effects that differently affect the measured intensities at both detectors.
Related to the origin of NDIR detector drift in combination with the dual wavelength
setup it should be pointed out that changes in the zero concentration measurements can
(i) occur erratically especially after transportation of storage, (ii) are not necessarily linear
in time with (iii) the slope commonly decreasing over running time or (iv) even changing
its sign.

The zero correction of a two-beam sensor signal does not cover concentration dependent effects that equate to changes in the characteristics of the NDIR sensor's calibration polynomial. Hence for achieving best accuracies particularly with two-beam NDIR sensors, the sensor in addition to regular zeroings needs to be re-calibrated after deployment at different concentrations.

#### 339 d. Membrane equilibration

340 The solution-diffusion model can be used to describe gas transport through a dense, semi-341 permeable membrane. Assuming a partial pressure gradient between the water phase and 342 the internal gas stream, the first step within the transport process is adsorption of the 343 molecule at the membrane surface. In the case of a sufficiently small gas concentration 344 the dissolution process into the membrane can be explained by the Henry-Dalton Law, 345 which states the linearity between the partial pressure of a gas and the concentration of 346 that gas in solution connected by a temperature, pressure and, in the case of sea water, 347 salinity dependent constant. The dissolution process is followed by diffusion based 348 transport of the molecules within the membrane along the concentration gradient. 349 Outgasing into the headspace again follows the Henry-Dalton Law. This process 350 continues within the membrane of the sensor until partial pressure equilibrium with

351 respect to every single gas component is achieved. Hence semi-permeable membranes 352 can be used as equilibrators for dissolved gas measurements (McNeil et al. 2006). They 353 enable the application of gas phase based measuring techniques in an aquatic medium. 354 The time limiting and overall response time dominating step within the entire transport 355 process is the diffusion within the membrane material as well as within a boundary layer 356 in front of the membrane. The permeability of a material is a parameter integrating both, 357 the solubility of a gas within the membrane substance as well as its diffusion constant 358 within the membrane according to Fick's Law. Like the solubility and the diffusion 359 coefficient, the permeability is also temperature and pressure dependent as well as 360 concentration independent. The permeabilities of different gases for a given membrane 361 material differ (Robb 1968; Merkel et al. 2000) and hence determine the time constant for 362 the corresponding partial pressure equilibration. The direction of the transport process is 363 defined by the orientation of the individual partial pressure gradient. Silicone, 364 polydimethylsiloxane (PDMS), was chosen as the membrane material due to its high 365 permeability for  $CO_2$  (Merkel et al. 2000). 366 In addition to all the dissolved gases, water vapor also permeates through the

367 membrane. Its amount within the gas stream is related to the temperature and salinity

368 driven water vapor pressure. The risk of condensation within the headspace and

369 especially within the NDIR absorption path induced by steep gradients of warm to cold

370 water is minimized by the heating of the gas on its way to the NDIR detector.

371 Furthermore the permeability of water vapor is more than a magnitude higher than for

372 CO<sub>2</sub> (Robb 1968), leading to time constants for water vapor equilibration of about 10 s at

373 the given  $CO_2$  response time of about 70 s for this sensor.

374 Typically silicone layer thicknesses of around 10  $\mu$ m are used. The thickness is 375 determined during membrane production by permeability measurements. Pure silicone 376 monolayer membranes of this thickness could not be easily handled. Therefore we use 377 thin film composite membranes consisting of the dense silicone layer on top of 378 supporting substructures. In the case of no or minor fouling the membranes can be 379 deployed for several months to years. Cleaning of the membranes with e.g. diluted 380 sulfuric or oxalic acid at pH 2 has successfully been tested. To avoid physical damage of 381 the thin silicone layer mechanical cleaning of the membrane surface should be avoided. 382 Instead the membrane should and can be changed even in the field. The response time of 383 the sensor determined from the flush interval data can also be used to identify organic 384 ongrowth, since heavy fouling slows down the membrane permeation process or the 385 volume rate of pumped water, which both leads to a reduced response time of the sensor.

### 386 3. In-water calibration setup

387 Along with the sensor development we established a laboratory calibration setup for 388 direct underwater  $pCO_2$  calibration (Fig. 3) that can hold up to three sensors 389 simultaneously. An early version of the setup was successfully used in Friedrichs et al. 390 (2010). The setup includes a 120 L insulated and temperature stabilized water tank. The 391 water temperature can be controlled over the temperature range of 0°C to 30°C to within 392  $\pm 0.02$  °C. It is filled with de-ionized water. Sodium carbonate and bicarbonate are added 393 in the required quantities in order to mimic the CO<sub>2</sub> buffer system of seawater (DIC:TA 394 ratio) and thereby allow for a better  $pCO_2$  level control. Silver nitrate is added as an anti-395 foulant. Water is continuously pumped through a reference flow-through  $pCO_2$  system 396 (Körtzinger et al. 1996), which was slightly modified to suit the laboratory conditions.

397	Additionally it was equipped with drying components to facilitate continuous reference
398	measurements in dry gas. Special care was taken that the return flow of the water from
399	the combined bubble-type/laminar-flow equilibrator into the main tank occurs without
400	flow restrictions and thus without altering the pressure conditions within the equilibrator.
401	Temperature probes within the equilibrator and the main tank are regularly calibrated
402	against a reference probe with an accuracy of $\pm 0.02$ °C. The flow-through system is
403	equipped with a bench top IR analyzer (LI-6262 or LI-7000, LI-COR Inc, Lincoln,
404	Nebraska, USA), which is calibrated against 3 primary (certified National Oceanic and
405	Atmospheric Administration (NOAA) standards) or secondary (referenced to NOAA
406	standards) CO <sub>2</sub> -in-natural-air standards in the beginning and in the end of each
407	calibration run. Processing of the flow-through system data is carried out according to the
408	procedures described in Dickson et al. (2007) and Pierrot et al. (2009), leading to $pCO_2$
409	reference values referred to the water temperature in the tank with an accuracy of 2-
410	3 $\mu$ atm. During a full calibration run, the <i>p</i> CO <sub>2</sub> of the tank water is altered by pH
411	variation through injection of NaOH or HCl solutions. A new concentration in the tank
412	water is set and resolved by the reference system with a time constant of approx. 150 s.
413	By application of the pH-varying technique a wide $pCO_2$ -range can be realized and
414	calibration steps can be set as desired. Since the overall setup cannot be entirely
415	encapsulated from the surrounding air a small drift of the partial pressures in the tank of
416	typically around 3 $\mu$ atm hr <sup>-1</sup> can be observed. Magnitude and timescale of this drift as
417	well as the fact that both, the flow-through system and the underwater sensor, detect this
418	$pCO_2$ change make this effect negligible with respect to the assumed accuracy of the
419	calibration process. The adjustable range is limited by the measurement range of the IR-

420	analyzer of the flow-through system, which is 3000 $\mu$ mol mol <sup>-1</sup> for the dry CO <sub>2</sub> mole
421	fraction in the equilibrated gas stream ( $x$ CO <sub>2</sub> ). It was found that four calibration steps are
422	sufficient for a sensor calibration in the range of 200 to 1000 µatm. The course of a
423	calibration is depicted in Fig. 2.
424	There are several reasons to calibrate the sensor in water against a proven
425	underway system as opposed to a mere dry gas calibration of the IR detector. The fact
426	that some of the following influences are already considered in data processing
427	corrections or their minimization was addressed in sensor design, does not debilitate the
428	following compensatory advantages of an in-water calibration. Firstly, the temperature
429	stabilization including the gas heater used in the sensor does not completely avert the
430	presence of temperature gradients within the instrument's housing, components and gas
431	stream. Hence the temperature stabilization can only minimize the above mentioned
432	possible temperature effects on NDIR sensors. An in-water calibration at a temperature as
433	close as possible to the expected temperature in the field helps to further reduce these
434	signal influences. Secondly, an in-water calibration as described compensates for all
435	effects related to the large absolute humidity present in the sensor's gas stream. These
436	effects are gas-gas-interactions causing band broadening, potential cross sensitivities of
437	the NDIR signal against H <sub>2</sub> O due to minor H <sub>2</sub> O absorption at the transmitted
438	wavelengths, or $H_2O$ molecule interaction with the cuvette's surface. For a sensor
439	calibration at only one temperature it is deliberately accepted that the humidity related
440	compensations are only entirely compensated for a deployment at a water temperature
441	equal to the calibration temperature. Thirdly, an in-water calibration compares the fully
442	processed signal of the instrument with a reference value. Hence it characterizes the

overall instrument including the entire membrane equilibration process of the headspace
as identified to be important by Byrne et al. (2010). Any not otherwise considered effects
occurring in the sensor's gas stream and behind the membrane are taken care of by an inwater calibration.

#### 447 4. Data processing

448 The dual-beam NDIR detector provides two signals. The raw signal,  $S_{raw}$ , corresponds to

the transmitted light intensity around the wavelength at which CO<sub>2</sub> efficiently absorbs

450 (4.26  $\mu$ m) and the reference signal,  $S_{ref}$ , expresses the intensity at around 4  $\mu$ m where

451 practically no relevant absorption occurs. Water vapor is weak absorber at both

452 wavelengths. A continuously referenced sensor signal, the "two-beam signal", is:

$$S_{2\text{beam}} = \frac{S_{\text{raw}}}{S_{\text{ref}}} \,. \tag{3}$$

453 As a result of the regular zeroing (Z), we obtain "two-beam zero signals" at 454 discrete points in time:

$$S_{2\text{beam},Z} = \frac{S_{\text{raw},Z}}{S_{\text{ref},Z}}.$$
(4)

455 Linear interpolation in time between two adjacent  $S_{2\text{beam},Z}$  provides zero reference 456 signals for every point in time:  $S_{2\text{beam},Z}(t)$ .

457 An NDIR signal that is improved by both drift correction (DC) means, dual beam 458 and zeroing, is derived as follows:

$$S_{\rm DC}(t) = \frac{S_{\rm 2beam}(t)}{S_{\rm 2beam,Z}(t)}$$
(5)

459 The final, drift corrected NDIR signal, which is assumed to be directly

460 proportional to the amount of target molecules in the beam path is related to  $S_{DC}(t)$ .

461	During calibration the flow-through system provides a reference $pCO_2$ for every
462	calibration step. The $pCO_2$ equilibrium is assumed to be established in the membrane
463	equilibrator of the sensor at that time. Using data of the peripheral sensors in the gas
464	stream, the $xCO_2$ in moist air present at the NDIR detector is derived. As the NDIR signal
465	is proportional to the number of molecules in the beam path and not to $xCO_2$ , the
466	reference $xCO_2$ needs to be density corrected by using data of the additional temperature
467	and pressure sensors built into the gas stream of the $pCO_2$ sensor. A polynomial of rank 3
468	with a forced zero crossing is then used to calibrate the individual sensor characteristics.
469	It correlates the absorbance signal of the NDIR detector with the corresponding and
470	density corrected $xCO_2$ in the gas stream (c.f. Fig. 4). Now, all required dependencies are
471	known and the sensor provides the $pCO_2$ based on the absorbance signal of its NDIR
472	detector in combination with the data of the peripheral sensors in its gas stream. Beside
473	the density correction no other NDIR signal correction addressing a band broadening
474	effect as a consequence of $CO_2$ -H <sub>2</sub> O molecule interactions or any other H <sub>2</sub> O cross
475	sensitivity is explicitly included in the sensor sided data processing at this point. These
476	aspects are considered through the in-water calibration as mentioned in section 3.
477	Since the consideration of the zeroings requires an interpolation in time, this
478	calculation step is best applied during post processing of field data to obtain a smooth
479	behavior. In order to achieve the best measurement accuracy, changes in the sensor
480	characteristics should also be included into the processing. Therefore an interpolation
481	between the polynomial of a pre-deployment calibration and the polynomial of a post-
482	deployment calibration over the course of the deployment is conducted. We apply an
483	interpolation that is not linear with time, but instead linear with the actual value of the

484 zero signal throughout a deployment,  $S_{2\text{beam},Z}(t)$ . The pre-deployment polynomial is 485 transformed to the post-deployment polynomial by proportionately using the coefficients 486 of the two polynomials according to the actual zero signal. This approach assumes a 487 causal relationship between the temporally often non-linear change in the zero signal and 488 the change in concentration dependent sensor response.

489 The entire calibration calculations as well as the post processing are accomplished
490 with custom-designed LabVIEW routines (National Instruments, Austin, Texas).

491 5. Field evaluation

492 Field evaluations of the new pCO<sub>2</sub> sensor were carried out in April/May 2010 during a 6-493 week cruise in the North and South Atlantic (R/V Polarstern, ANT-XXVI/4) and in 494 June/July 2011 during a 4-week cruise in the eastern tropical Atlantic (R/V Maria S. 495 *Merian*, MSM-18/3; Fig. 5). During these cruises, oligotrophic (i.e., subtropical gyres) as 496 well as mesotrophic regions (e.g., continental shelfs, equatorial upwelling) provided a 497 reasonably wide range in  $pCO_2$  (295 to 430  $\mu$  atm; c.f. Fig 6. and Fig 7. top panel) and 498 temperature (7.4 to 30.3 °C, Fig. 5). Furthermore, strong  $pCO_2$  and temperature gradients 499 were found near hydrographic fronts. Thus, the conditions were ideal for a thorough 500 evaluation of the sensor, which was tested during both cruises in nearly identical 501 underway setups: Seawater, either supplied by the ship's clean seawater supply systems (ANT-XXVI/4) or by a submersible pump installed in the moon pool near the ship's keel 502 503 (MSM-18/3), was pumped to the laboratory into a thermally insulated flow-through box 504 (80 L volume) at a flow rate of approx. 12 L min<sup>-1</sup>. Sea surface temperature (SST) and 505 sea surface salinity (SSS) were determined for both cruises at the seawater intake. A 506 sensor package containing the  $pCO_2$  sensor (two units during ANT-XXVI/4, HC1 and

507	HC2; one unit on MSM-18/3, HC3) and an oxygen optode (Model 3830 or 4330,
508	Aanderaa Data Instruments AS, Bergen, Norway), which also provided the water
509	temperature in the flow-through box with an accuracy of $\pm 0.05^{\circ}$ C, were placed in the
510	underway box. Data were binned into 1-min intervals. Since the $pCO_2$ sensor data were
511	initially stored at 5-s intervals the transformation to the 1-min intervals represents an
512	averaging of typically 12 spot values. Zeroings were carried out every 12 hours. The
513	membrane interface of the $pCO_2$ sensor was supplied with a constant seawater flow by a
514	SBE 5T pump (Sea-Bird Electronics Inc., Bellevue, Washington). A fully automated
515	pCO <sub>2</sub> underway instrument (GO, General Oceanics, Miami, USA; Pierrot et al. 2009)
516	based on a spray head equilibrator and a LI-7000 CO <sub>2</sub> analyzer was operated in parallel.
517	Throughout the expeditions and beside the deployment in the flow-through box the
518	sensors were additionally used for measurements on a CTD rosette system (HC1 and
519	HC3; part of the data shown in Fiedler et al. 2013) as well as on a surface drifter (HC3;
520	unpublished data).
521	The two sensors HC1 and HC2 were calibrated at 19.67°C before and after the
522	deployment. HC3 was pre- and post-calibrated at 27.00°C and 26.00°C, respectively.
523	The averaged sensor data were prepared for processing by filtering for obvious
524	outlier zero values as, for example, caused by improper sensor warm-ups, as well other
525	outliers and data biased through insufficient water supply to the flow-through box. All
526	pCO <sub>2</sub> sensor data recorded during flush and zero intervals were excluded from the
527	comparison with the reference. The data of HC2 obtained between $28^{th}$ of April and $4^{th}$ of
528	May were excluded from the comparison as well, as the excessively high water
529	temperature did not allow for temperature stabilization at the calibration control

- temperature anymore. Although the control temperature within HC1 also temporarily
- 531 exceeded the set point, no data were removed here, as the unit seemed to be more robust
- 532 in this respect as compared to HC2.
- Figures 6 and 7 show the  $pCO_2$  measured by the flow-through reference system as well as the  $pCO_2$  data of the HC1-3 sensors with all datasets corrected to SST for direct
- 535 comparison (Takahashi et al. 1993). In addition, the  $pCO_2$  differences ( $\Delta pCO_2$ ) between
- 536 the sensor and the reference  $(pCO_{2,GO})$  are plotted vs. time. The three  $\Delta pCO_2$  curves for
- 537 every sensor are related to different drift compensation mechanisms:
- 538

- (i) pre-calibration, no zeroings
- 539 (ii) pre-calibration + zeroings
- 540 (iii) pre-calibration + zeroings + post-calibration

541 In the case of (i) only the dual-beam correction is applied and the polynomial of 542 the pre-deployment calibration is used for the entire data set. Since in this case data are 543 processed by referencing all signals to the zeroing carried out during shore-based 544 calibration, large offsets can already occur at the beginning of a deployment due to a 545 sensor drift that occurred during storage and transport or due to measurement conditions 546 that strongly deviate from calibration conditions. Also more or less strong drift behavior 547 is observed over time. When the regular zero correction is carried out as described above 548 (section 4), both initial offsets and drift over the course of the deployment are strongly 549 reduced demonstrating the effectiveness of this first correction. However, even smaller 550 residuals were obtained when both zeroings as well as pre- and the post-deployment 551 calibration polynomials are considered in the post-processing (section 4). Table 2 shows 552 the statistics of the  $pCO_2$  residuals as obtained by this optimized processing routine. In all 553 three sensor deployments the mean  $pCO_2$  offset to the reference system is within 554  $\pm 3 \mu$  atm. With an average pCO<sub>2</sub> residual over all three sensors of  $-0.6 \pm 3.0 \mu$  atm 555 (RMSE=3.7  $\mu$  atm) no systematic offset between pCO<sub>2</sub> sensor and reference system could 556 be found. This indicates that the agreement between sensor and flow-through system 557  $pCO_2$  is of similar magnitude as the accuracy of the flow-through system which was 558 estimated at 2 µatm (Pierrot et al. 2009). We use the mean RMSE from all three 559 deployments of 3.7 µatm as a conservative estimate of the sensor accuracy. This result is 560 very promising as the sensors were only calibrated at a single water temperature and 561 experienced a large temperature range during deployment (> $20^{\circ}$ C). We note that during 562 these field tests the sensors ran autonomously and without maintenance but were also 563 used in other tests (e.g., deployments of CTD rosette casts; part of the data shown in 564 Fiedler et al., 2013). For this purpose the instruments had to be restarted several times. 565 This appears to not have affected sensor performance negatively, since sensor HC3 was 566 restarted most frequently but shows the smallest overall offset. 567 The compensation routines applied, which exclusively rely on data measured by 568 the sensors themselves and the information obtained from calibrations, account well for 569 the signal drift for all 3 sensors. After processing the data do not exhibit significant

570 unaccounted drift behavior. This is even the case for the most strongly drifting sensor

571 HC1, whose signal change over time is also reflected by the change in its calibration

572 polynomials (Fig. 4). Although an NDIR detector drifting as strongly as in the case of

573 HC1 would not pass current quality controls within the manufacturer's production, it is

574 still a good example to demonstrate the effectiveness of the described processing

575 algorithms. We would like to note that sensor HC1 also participated in a different sensor

576 evaluation project (Tamburri et al. 2011). At that time, the sensor also showed a strong 577 drift that could not be adequately compensated for due to lack of the full understanding of 578 the required post-processing steps that is presented here. Although the drift of HC1 could 579 be well corrected for, in the end it still shows a slightly larger mean value and RMSE 580 compared to HC2 and HC3. NDIR detectors that show a smaller zero drift, typically 581 show a smaller concentration dependent signal change over time as well. This conclusion 582 is further corroborated by the observation that a transformation of the pre-deployment 583 into the post-deployment calibration polynomial based on the course of the zero values 584 finally provides a better correction as opposed to a transformation assumed to occur 585 linear in time (data not shown here).

586 In order to identify any remaining issues in the drift corrected  $pCO_2$  sensor data, 587 the  $\Delta pCO_2$  residuals were plotted against  $pCO_{2,GO}$ , SST and  $pH_2O$  (Fig. 8). All sensors show a weak correlation with all three parameters ( $0.0 < R^2 < 0.6$ ), which is most clear for 588 sensor HC2 (0.5 < R<sup>2</sup> < 0.6). Since pCO<sub>2 GO</sub>, SST and pH<sub>2</sub>O are strongly correlated in the 589 590 field data, the cause of these remaining residual correlations cannot be clearly discerned. 591 Nevertheless, there is indication for a weak NDIR signal dependency on  $pH_2O$ . In fact, 592 CO<sub>2</sub>-NDIR detectors over-estimate in the presence of water vapor due to pressure 593 broadening effects (McDermitt et al. 1993). Furthermore, the magnitude of this effect 594 increases with  $pCO_2$  and  $pH_2O$ . Since the sensor calibrations were only conducted at one 595 temperature and band broadening effects due to varying water vapor concentrations are 596 not considered within the sensor's data processing, this might explain at least part of the 597 dependencies. The fact that HC2 and HC3 show the smallest  $\Delta p CO_2$  around the water 598 temperature at which they were calibrated, support this observation. The limitations of

the present data set do not allow to further investigate this issue and additional tests have to be carried out to assess the potential for further improvement, e.g., by performing laboratory tests with and calibrations of the  $pCO_2$  sensor at more than one temperature.

#### 602 6. Summary and outlook

603 The development of a new underwater  $pCO_2$  sensor based on membrane equilibration and 604 NDIR spectrometry was described. Special emphasis was put on compensation measures 605 for NDIR sensor drift as well as on the in-water calibration of the sensor. The 606 performance of the  $pCO_2$  sensor was assessed based on surface water field data obtained 607 during two cruises both lasting at least one month and covering a wide range in  $pCO_2$ 608 (289 to 445  $\mu$ atm) and temperature (7.4 to 30.1°C). A wet gas stream within a small 609 underwater sensor represents a demanding environment for NDIR detectors. Against this 610 background the observed mean offset of  $-0.6 \pm 3.0 \mu$  atm with a RMSE of 3.7  $\mu$  atm to the 611  $pCO_2$  reference instrument as obtained through application of a pre- and post-deployment 612 calibration in combination with regular zeroings is a very promising result, especially 613 since the sensors were also used for various other measurements including profiling 614 applications in the water column during the deployments (data not part of this 615 assessment). At a  $pCO_2$  of 400 µatm the observed mean  $pCO_2$  difference corresponds to 616 about 0.2 % and the RMSE to less than 1 %. This favorable result underlines the 617 efficiency of the applied processing algorithms. The acquisition of high-quality field data 618 by the new  $pCO_2$  sensor calls for regular checks of the sensor parameters gas 619 temperature, pressure, RH and control temperature to guarantee optimal functioning of 620 the sensor. In addition, regular zero gas measurements need to be carried out at least 621 every 12 h under deployment conditions as a drift correction means beside the inherent

622	single-beam dual wavelength setup of the sensor. To properly apply the zero information
623	during post processing measurement data should always be flanked by zeroings. Thirdly,
624	in order to also account for the concentration dependant effects on zero and dual beam
625	corrected signals, the sensor needs to be re-calibrated at different $pCO_2$ levels on a time
626	scale of several months to a year to achieve the highest accuracies through data post
627	processing. A water calibration as presented in this paper at a temperature close to the
628	expected water temperatures in the field is beneficial. Nevertheless, the RMSE found
629	within this work is based on field data obtained under conditions where water
630	temperatures deviated by $\pm 10^{\circ}$ C from the calibration temperature.
631	The assessment given here represents an important milestone for the development
632	of the sensor. The procedures discussed are planned to be further automated and
633	implemented into data processing routines. The next development steps include
634	investigation of potential improvements with respect to NDIR data processing and the
635	laboratory calibration routines to identify and compensate for minor signal dependencies
636	on water vapor and on changes in gas matrix composition as caused by e.g. strongly
637	varying oxygen concentrations. Moreover the long-term stability during deployments on
638	moorings and profiling buoys in the water column will be investigated as well as the
639	sensor performance on different moving platforms analyzed. Especially the latter has
640	been simplified by a recently released smaller and faster version of the HydroC.
641	

The authors would like to thank all members of the "(C)O<sub>2</sub>" working group at the 643 644 GEOMAR as well as Gernot Friedrichs from the CAU for fruitful discussions and helpful 645 advice. Furthermore, we thank the CONTROS team for their support in this development 646 project. We explicitly acknowledge Christian Rettich for his assistance especially during 647 the construction of the first prototypes as well as Claus Hinz and Matthias Lunge for 648 valuable help with laboratory experiments. Janna Bohlen and Christoph Kirbach are 649 thanked for commenting on the manuscript. We would also like to thank captains and 650 crews of R/V Polarstern and of Maria S. Merian for assistance. This work was partly 651 funded by the SAW project OCEANET of the Leibniz Association (2008-2010) and by 652 the SOPRAN project of the German Federal Ministry of Education and Research (grants 653 03F0462A and 03F0611A).

## 655 **References**

656 657 658	Byrne, R. H., and W. Yao, 2008: Procedures for measurement of carbonate ion concentrations in seawater by direct spectrophotometric observations of Pb(II) complexation. <i>Mar. Chem.</i> , <b>112</b> , 128–135, doi:10.1016/j.marchem.2008.07.009.
659 660 661	Byrne, R. H., X. Liu, and E. A. Kaltenbacher, 2002: Spectrophotometric measurement of total inorganic carbon in aqueous solutions using a liquid core waveguide. <i>Anal. Chim. Acta</i> , <b>451</b> , 221–229, doi:10.1016/S0003-2670(01)01423-4.
662	Byrne, R. H. and Coauthors, 2010: Sensors and Systems for in situ Observations of
663	Marine Carbon Dioxide System Variables. <i>Proceedings of OceanObs'09: Sustained</i>
664	<i>Ocean Observations and Information for Society</i> , J. Hall, D.E. Harrison, and D.
665	Stammer, Eds., Venice, Italy, ESA Publication WPP-306.
666	DeGrandpre, M. D., T. R. Hammar, S. P. Smith, and F. L. Sayles, 1995: In-Situ
667	Measurements of Seawater pCO <sub>2</sub> . <i>Limnol. Oceanogr.</i> , <b>40</b> , 969–975.
668 669	Dickson, A. G., C. L. Sabine, and J. R. Christian, eds., 2007: <i>Guide to best practices for ocean CO</i> <sub>2</sub> <i>measurements</i> . PICES Special Publication 3.
670	Doney, S. C., V. J. Fabry, R. A. Feely, and J. A. Kleypas, 2009: Ocean Acidification: The
671	Other CO <sub>2</sub> Problem. Ann. Rev. Mar. Sci., 1, 169–192,
672	doi:10.1146/annurev.marine.010908.163834.
673	Fiedler, B., P. Fietzek, N. Vieira, P. Silva, H. C. Bittig, and A. Körtzinger, 2013: In situ
674	CO <sub>2</sub> and O <sub>2</sub> measurements on a profiling float. <i>J. Atmos. Oceanic Technol.</i> , <b>30</b> , 112–
675	126.
676 677 678 679	Fietzek, P., S. Kramer, and D. Esser, 2011: Deployment of the HydroC <sup>TM</sup> (CO <sub>2</sub> /CH <sub>2</sub> ) on stationary and mobile platforms - Merging the trends in the field of platform and sensor development. <i>Proceedings of Oceans 11 MTS/IEEE Conference</i> , <i>19-22 Sept. 2011</i> , Kona, Hawaii, USA, 1–9.
680	Friedrichs, G., J. Bock, F. Temps, P. Fietzek, A. Körtzinger, and D. W. R. Wallace, 2010:
681	Toward continuous monitoring of seawater ${}^{13}CO_2/{}^{12}CO_2$ isotope ratio and $pCO_2$ :
682	Performance of cavity ringdown spectroscopy and gas matrix effects. <i>Limnol.</i>
683	<i>Oceanogr.: Methods</i> , <b>8</b> , 539–551, doi:10.4319/lom.2010.8.539.
684	Gruber, N. and Coauthors, 2010: Toward an Integrated Observing System for Ocean
685	Carbon and Biogeochemistry at a Time of Change. <i>Proceedings of OceanObs'09:</i>
686	<i>Sustained Ocean Observations and Information for Society</i> , J. Hall, D.E. Harrison,
687	and D. Stammer, Eds., Venice, Italy, ESA Publication WPP-306.

- IPCC, 2005: *IPCC special report on carbon dioxide capture and storage*. B. Metz, O.
  Davidson, H. De Coninck, M. Loos, and L. Meyer, Eds. Cambridge University
  Press.
- Johnson, K. S. and Coauthors, 2009: Observing biogeochemical cycles at global scales
  with profiling floats and gliders: prospects for a global array. *Oceanography*, 22,
  216–224.
- Körtzinger, A., H. Thomas, B. Schneider, N. Gronau, L. Mintrop, and J. C. Duinker,
   1996: At-sea intercomparison of two newly designed underway *p*CO<sub>2</sub> systems encouraging results. *Mar. Chem.*, **52**, 133–145.
- Körtzinger, A. and Coauthors, 2000: The international at-sea intercomparison of fCO<sub>2</sub>
  systems during the RV Meteor Cruise 36/1 in the North Atlantic Ocean. *Mar. Chem.*, 72, 171–192, doi:10.1016/S0304-4203(00)00080-3.
- Lefèvre, N., J. P. Ciabrini, G. Michard, B. Brient, M. DuChaffaut, and L. Merlivat, 1993:
  A new optical sensor for *p*CO<sub>2</sub> measurements in seawater. *Mar. Chem.*, 42, 189–
  198, doi:10.1016/0304-4203(93)90011-C.
- Martz, T. R., J. G. Connery, and K. S. Johnson, 2010: Testing the Honeywell Durafet for
   seawater pH applications. *Limnology And Oceanography Methods*, 8, 172–184,
   doi:10.4319/lom.2010.8.172.
- McDermitt, D. K., J. M. Welles, and R. D. Eckles, 1993: Effects of temperature, pressure and water vapor on gas phase infrared absorption by CO<sub>2</sub>. *LI-COR, Inc, Lincoln, USA.*, 1–5.
- McNeil, C., E. D'Asaro, B. Johnson, and M. Horn, 2006: A Gas Tension Device with
  Response Times of Minutes. *J. Atmos. Oceanic Technol.*, 23, 1539,
  doi:10.1175/JTECH1974.1.
- Merkel, T. C., V. I. Bondar, K. Nagai, B. D. Freeman, and I. Pinnau, 2000: Gas sorption,
  diffusion, and permeation in poly(dimethylsiloxane). *J. Polym. Sci., Part B: Polym. Phys.*, **38**, 415–434, doi:10.1002/(SICI)1099-0488(20000201)38:3<415::AID-</li>
  POLB8>3.0.CO;2-Z.
- Millero, F. J., 2007: The marine inorganic carbon cycle. *Chem. Rev.*, **107**, 308–341,
   doi:10.1021/cr0503557.
- Pfeil, B. and Coauthors, 2013: A uniform, quality controlled Surface Ocean CO<sub>2</sub> Atlas
  (SOCAT). *Earth Syst. Sci. Data*, 125–143, doi:10.5194/essd-5-125-2013.
- Pierrot, D. and Coauthors, 2009: Recommendations for autonomous underway pCO<sub>2</sub>
   measuring systems and data-reduction routines. *Deep-Sea Res.*, *Part II*, 56, 512–
   522, doi:10.1016/j.dsr2.2008.12.005.

- Robb, W. L., 1968: Thin silicone membranes Their permeation properties and some applications. *Annals New York Academy of Science*, 119–137.
- Roemmich, D., G. C. Johnson, S. Riser, R. Davis, and J. Gilson, 2009: The Argo
  Program: observing the global ocean with profiling floats. *Oceanography*, 22, 34–
  43.
- Rogner, H. H., D. Zhou, R. Bradley, P. Crabbé, O. Edenhofer, B. Hare (Australia), L.
  Kuijpers, and M. Yamaguchi, 2007: *Introduction. In Climate Change 2007: Mitigation. Contribution of Working Group III to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change.* Cambridge University Press.
- Sabine, C. L. and Coauthors, 2004: The Oceanic Sink for Anthropogenic CO<sub>2</sub>. *Science*,
  305, 367–371.
- Saderne, V., P. Fietzek, and P. M. J. Herman, 2013: Extreme variations of *p*CO<sub>2</sub> and pH
  in a macrophyte meadow of the Baltic Sea in summer: evidence of the effect of
  photosynthesis and local upwelling. *PLoS ONE*, 8, e62689,
  doi:10.1371/journal.pone.0062689.
- Seidel, M. P., M. D. DeGrandpre, and A. G. Dickson, 2008: A sensor for in situ
  indicator-based measurements of seawater pH. *Mar. Chem.*, 109, 18–28,
  doi:10.1016/j.marchem.2007.11.013.
- Severinghaus, J. W., and A. F. Bradley, 1958: Electrodes for Blood pO<sub>2</sub> and pCO<sub>2</sub>
  Determination. J. Appl. Physiol., 13, 515–520.
- Stow, R. W., R. F. Baer, and B. F. Randall, 1957: Rapid measurement of the tension of
  carbon dioxide in blood. *Archives of physical medicine and rehabilitation*, 38, 646–
  650.
- Takahashi, T., 1961: Carbon dioxide in the atmosphere and in Atlantic Ocean water. *J. Geophys. Res.*, 66, 477–494.
- Takahashi, T., J. Olafsson, J. G. Goddard, D. W. Chipman, and S. C. Sutherland, 1993:
  Seasonal variation of CO<sub>2</sub> and nutrients in the high-latitude surface oceans: A
  comparative study. *Global Biogeochem. Cycles*, 7, 843–878.
- Takahashi, T. and Coauthors, 2009: Climatological mean and decadal change in surface
   ocean *p*CO<sub>2</sub>, and net sea–air CO<sub>2</sub> flux over the global oceans. *Deep-Sea Res., Part II*, 56, 554–577, doi:10.1016/j.dsr2.2008.12.009.
- Tamburri, M. N. and Coauthors, 2011: Alliance for Coastal Technologies: Advancing
  Moored *p*CO<sub>2</sub> Instruments in Coastal Waters. *Mar. Technol. Soc. J.*, 45, 43–51,
  doi:10.4031/MTSJ.45.1.4.

- Wang, Z. A., X. Liu, R. H. Byrne, R. Wanninkhof, R. E. Bernstein, E. A. Kaltenbacher,
  and J. Patten, 2007: Simultaneous spectrophotometric flow-through measurements
  of pH, carbon dioxide fugacity, and total inorganic carbon in seawater. *Anal. Chim. Acta*, **596**, 23–36.
- Watson, A. J. and Coauthors, 2009: Tracking the variable North Atlantic sink for
   atmospheric CO<sub>2</sub>. *Science*, **326**, 1391–1393, doi:10.1126/science.1177394.
- Weiss, R. F., 1974: Carbon dioxide in water and seawater: the solubility of a non-ideal
  gas. *Mar. Chem.*, 2, 203–215.
- Wiegleb, G. and Coauthors, 2001: Industrielle Gasssensorik: Messverfahren Signalverarbeitung Anwendungstechnik Prüfkriterien. G. Wiegleb, Ed. expert Verlag, Renningen.
- 768

### 770 List of Figures

771 TABLE 1. Specifications of the developed  $pCO_2$  sensor as used during the deployments 772 discussed in this paper. The specifications of the currently available sensor model differ 773 from these values with respect to size and power consumption. The power required for 774 the temperature stabilization as well as the warm-up duration depend on the actual water 775 temperature, the chosen control temperature as well as on the thickness of the insulation 776 material. The given warm-up times correspond to 24 V supply voltage in 20°C water and 777 to 12 V supply voltage in 3°C water. Please refer to the text for further details regarding 778 the warm-up and the zeroing interval. The response times refer to the usage of two 779 different pump models at 20°C water temperature. The pump SBE 5T has a flow rate of approx. 105 mL s<sup>-1</sup>, while the smaller model, SBE 5M, provides a flow rate of approx. 780  $35 \text{ mL s}^{-1}$ . 781

782

TABLE 2. Statistics of the  $\Delta p CO_2$  residuals for all three field deployments of the  $p CO_2$ sensor with mean, standard deviation,  $\sigma$ , and root mean square error, RMSE, for a total of *n* observations. Also shown is the mean of all three deployments.

786

FIG. 1. Drawing of the *p*CO<sub>2</sub> sensor as used in the present study (upper panel). The sensor
is equipped with a water pump (SBE 5T) and a flow-head. The lower panel shows a
schematic drawing of the sensor. Partial pressure equilibration occurs at the planar, semipermeable membrane separating the water from the internal head space of approx. 20 mL.
A pump continuously circulates the gas between the membrane equilibrator, a heater and
the NDIR detector. Valves can be toggled to realize a zero gas measurement by guiding

the gas stream through a soda lime cartridge instead of through the membrane

794 equilibrator.

796 FIG. 2. Course of a calibration carried out at a constant water temperature within the 797 calibration setup described in the text. The absorbance NDIR signal shown in arbitrary 798 units is dual-beam and zero corrected. Also shown is the  $CO_2$  mole fraction ( $xCO_2$ ) 799 measured by the reference flow-through system. The numbers 1-4 indicate the 800 calibrations steps and "Z", "F" and "M" mark the sensor intervals: zero, flush and 801 measure. 802 803 FIG. 3. In-water  $pCO_2$  sensor calibration setup. One water pump is used to allow for 804 sufficient mixing of the water in the tank, dispersion of the injected acid and base as well 805 as to provide water to the reference flow-through system, which is installed in a bypass 806 and whose NDIR unit is regularly calibrated with reference gases. 807 808 FIG. 4. Calibration polynomials of sensor HC1 before and after the deployment on R/V 809 *Polarstern.* The absorbance NDIR signal is calibrated against the  $xCO_2$  within the 810 sensor's gas stream obtained according to the text. The polynomials match well the sensor's response characteristics ( $R^2$  in both cases >0.999). A concentration dependent 811 812 change in the sensor characteristics between the different points in time (i.e., before and 813 after deployment) of the calibrations is clearly visible in this example for a strongly 814 drifting NDIR sensor. 815

817	Merian cruise MSM-18/3 (left) as well as the encountered sea surface salinity and
818	temperature (right panel). R/V Polarstern sailed from Germany to Chile and R/V Maria
819	S. Merian from Cape Verde to Gabon.
820	
821	FIG. 6. $pCO_2$ data obtained by the GO-system as well as two HydroC sensors, HC1 and
822	HC2, during ANT-XXVI/4. The top panel shows the absolute values, while the middle
823	and bottom panel show the differences between the HydroC- $p$ CO <sub>2</sub> and the reference

FIG. 5. The cruise track of R/V Polarstern cruise ANT-XXVI/4 and of R/V Maria S.

824  $(\Delta p CO_2)$  over time for three different processing methods in order to visualize the

825 potential of different drift compensation mechanisms (refer to text). The black curve of

HC1 and HC2 represent the final drift corrected and post-processed signal. Refer to textfor further details on the different processing methods as well as on the peculiarities of

828 strongly drifting HC1.

829

816

```
FIG. 7. The pCO_2 trace during MSM-18/3 measured by the GO reference pCO_2 system,
```

and a HydroC sensor, HC3 (top panel). The lower panel shows the  $pCO_2$  residuals

between the HydroC and the reference ( $\Delta p CO_2$ ) for three different drift compensation

scenarios. The black curve in the lower panel represents the final drift corrected and post-processed signal.

835

836 FIG. 8. In the top row the residuals of HC1 are plotted over the reference  $pCO_2$  (left), the

837 SST (middle) and over the  $pH_2O$  as measured within the internal gas stream close to the

838 NDIR detector (right). The middle and the bottom row show the same parameter plots for

- 840 positive correlation with all three parameters can be made out and is further discussed in
- the text.
- 842

843	TABLE 1. Specifications of the developed $pCO_2$ sensor as used during the deployments
844	discussed in this paper. The specifications of the currently available sensor model differ
845	from these values with respect to size and power consumption. The power required for
846	the temperature stabilization as well as the warm-up duration depend on the actual water
847	temperature, the chosen control temperature as well as on the thickness of the insulation
848	material. The given warm-up times correspond to 24 V supply voltage in 20°C water and
849	to 12 V supply voltage in 3°C water. Please refer to the text for further details regarding
850	the warm-up and the zeroing interval. The response times refer to the usage of two
851	different pump models at 20°C water temperature. The pump SBE 5T has a flow rate of
852	approx. 105 mL s <sup>-1</sup> , while the smaller model, SBE 5M, provides a flow rate of approx.
853	$35 \text{ mL s}^{-1}$ .

Measurement principle	IR absorption measurement in a membrane equilibrated
	headspace
Detector	Single beam dual wavelength NDIR detector; zeroings at
	desired intervals
Housing, dimensions	Cylindrical titanium housing, 90 x 530 mm (without connector)
Depth capability	2000 m (standard)
Weight	5.5 kg in air, approx. 2.6 kg in water
Operating temperature	$3 - 30^{\circ}\mathrm{C}$
Supply voltage	11-24 VDC
Power requirements	<3 W for the detector and all the electronics,
	+ <1  W - 3.5  W for temperature stabilization,
	+ <6 W during warm-up for $2 - 30$ min,

	+ 4 W during zeroing,			
	+ water pump: 7 W or 1.5 W (pump SBE 5T and 5M,			
	respectively)			
Sampling rate	$\leq 1 \text{ Hz}$			
Response time ( $t_{63\%}$ )	Approx. 70 s or 130 s (pump SBE 5T and 5M, respectively)			
Measurement range	200 µatm – 1000 µatm (standard)			

855 TABLE 2. Statistics of the  $\Delta p CO_2$  residuals for all three field deployments of the  $p CO_2$ 

856 sensor with mean, standard deviation,  $\sigma$ , and root mean square error, RMSE, for a total of

857	<i>n</i> observations.	Also shown	is the mean	of all three	deployments.
					1 2

Sensor	Mean (µatm)	$\sigma(\mu atm)$	RMSE (µatm)	n
HC1	-3.1	2.9	4.2	24791
HC2	1.8	3.4	3.9	24163
HC3	-0.7	2.8	2.8	12770
Overall mean	-0.6	3.0	3.7	



860

861 FIG. 1. Drawing of the  $pCO_2$  sensor as used in the present study (upper panel). The sensor 862 is equipped with a water pump (SBE 5T) and a flow-head. The lower panel shows a 863 schematic drawing of the sensor. Partial pressure equilibration occurs at the planar, semi-864 permeable membrane separating the water from the internal head space of approx. 20 mL. 865 A pump continuously circulates the gas between the membrane equilibrator, a heater and 866 the NDIR detector. Valves can be toggled to realize a zero gas measurement by guiding 867 the gas stream through a soda lime cartridge instead of through the membrane 868 equilibrator.



FIG. 2. Course of a calibration carried out at a constant water temperature within the calibration setup described in the text. The absorbance NDIR signal shown in arbitrary units is dual-beam and zero corrected. Also shown is the  $CO_2$  mole fraction ( $xCO_2$ ) measured by the reference flow-through system. The numbers 1 - 4 indicate the calibrations steps and "Z", "F" and "M" mark the sensor intervals: zero, flush and measure.



FIG. 3. In-water  $pCO_2$  sensor calibration setup. One water pump is used to allow for

879 sufficient mixing of the water in the tank, dispersion of the injected acid and base as well

as to provide water to the reference flow-through system, which is installed in a bypass

and whose NDIR unit is regularly calibrated with reference gases.



FIG. 4. Calibration polynomials of sensor HC1 before and after the deployment on R/V *Polarstern*. The absorbance NDIR signal is calibrated against the xCO<sub>2</sub> within the sensor's gas stream obtained according to the text. The polynomials match well the sensor's response characteristics (R<sup>2</sup> in both cases >0.999). A concentration dependent change in the sensor characteristics between the different points in time (i.e., before and after deployment) of the calibrations is clearly visible in this example for a strongly drifting NDIR sensor.





*Merian* cruise MSM-18/3 (left) as well as the encountered sea surface salinity and

- 896 temperature (right panel). R/V Polarstern sailed from Germany to Chile and R/V Maria
- 897 S. Merian from Cape Verde to Gabon.



900 FIG. 6.  $pCO_2$  data obtained by the GO-system as well as two HydroC sensors, HC1 and 901 HC2, during ANT-XXVI/4. The top panel shows the absolute values, while the middle 902 and bottom panel show the differences between the HydroC-pCO<sub>2</sub> and the reference 903  $(\Delta p CO_2)$  over time for three different processing methods in order to visualize the 904 potential of different drift compensation mechanisms. The black curve of HC1 and HC2 905 represent the final drift corrected and post-processed signal. Refer to text for further 906 details on the different processing methods as well as on the peculiarities of strongly 907 drifting HC1.





911 FIG. 7. The  $pCO_2$  trace during MSM-18/3 measured by the GO reference  $pCO_2$  system,

912 and a HydroC sensor, HC3 (top panel). The lower panel shows the pCO<sub>2</sub> residuals

913 between the HydroC and the reference ( $\Delta p CO_2$ ) for three different drift compensation

914 scenarios. The black curve in the lower panel represents the final drift corrected and post-

<sup>915</sup> processed signal.



FIG. 8. In the top row the residuals of HC1 are plotted over the reference  $pCO_2$  (left), the SST (middle) and over the  $pH_2O$  as measured within the internal gas stream close to the NDIR detector (right). The middle and the bottom row show the same parameter plots for HC2 and HC3, respectively. Although not statistically significant in all cases, a general positive correlation with all three parameters can be made out and is further discussed in the text.