## SUPPORTING INFORMATION

## Iron-Mediated Anaerobic Oxidation of Methane in

## Brackish Coastal Sediments

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The 8 pages of supporting information consist of 2 Figures (Figure S1: correction of AVS dissolution during Fe-oxide extraction; Figure S2: determination of Fe -AOM rate) and 3 Tables (Table S1: porewater data; Table S2: solid-phase data; Table S3: model parameters).

## POREWATER CALCULATIONS

The diffusive flux of dissolved $\mathrm{Fe}^{2+}$ into the SMTZ can be used to estimate the rate of Fe reduction below the SMTZ. Thus, assuming that all of the Fe-reduction below the SMTZ is due to $\mathrm{Fe}-\mathrm{AOM}$ and limited removal of dissolved $\mathrm{Fe}^{2+}$ below the SMTZ, one can estimate the contribution of $\mathrm{CH}_{4}$ removal by $\mathrm{Fe}-\mathrm{AOM}$ to the total amount of $\mathrm{CH}_{4}$ removed in the sediment.

Porewater gradients were used to calculate diffusive fluxes of $\mathrm{CH}_{4}$ and $\mathrm{Fe}^{2+}$ into the SMTZ applying Fick's first law ${ }^{1}$ :
$J=\phi D_{s} \frac{d C_{i}}{d x}$
where J equals the diffusive flux in mol me ${ }^{-2}$ year $^{-1}, \varphi$ is the measured porosity, $D_{s}$ is the sediment diffusion coefficient in $\mathrm{m}^{2}$ year ${ }^{-1}$, C is the concentration of molecule $i$ in $\mathrm{mol} \mathrm{m}^{-3}$, and x is the depth in m . Sediment diffusion coefficients were calculated as follows:
$D_{s}=\frac{D_{i, s w}}{1-\ln \phi^{2}}$
where $\mathrm{D}_{\mathrm{i}, \mathrm{sw}}$ (diffusive flux of $i$ in seawater) was corrected for in-situ temperature and salinity according to ref ${ }^{2}$. Based on the porewater profiles, our estimated diffusive fluxes for $\mathrm{CH}_{4}$ and $\mathrm{Fe}^{2+}$ are $\mathrm{J}_{\mathrm{CH} 4}=60 \mu \mathrm{~mol} \mathrm{~cm}{ }^{-2}$ year $^{-1}$ and $\mathrm{J}_{\mathrm{Fe} 2+}=13 \mu \mathrm{~mol} \mathrm{~cm}{ }^{-2}$ year $^{-1}$, respectively. Thus, taking into account that only one molecule of $\mathrm{CH}_{4}$ is needed to reduce $8 \mathrm{Fe}(\mathrm{OH})_{3}, \mathrm{CH}_{4}$ removal by Fe-AOM is estimated to account for about $3 \%$ of total $\mathrm{CH}_{4}$ removal. If part of the $\mathrm{Fe}^{2+}$ production below the SMTZ were still due to organoclastic Fe reduction, the contribution of the $\mathrm{Fe}-\mathrm{AOM}$ to the total removal of $\mathrm{CH}_{4}$ would even be lower. Thus, the $3 \%$ should be seen as an upper estimate.

## DIFFUSION MODEL PARAMETERS

Implemented rate constants for $\mathrm{SO}_{4}-\mathrm{AOM}$ (32.2 year ${ }^{-1}$ ) and $\mathrm{Fe}-\mathrm{AOM}\left(0.64\right.$ year ${ }^{-1}$ ) are based on porewater profiles of $\mathrm{CH}_{4}$ and $\mathrm{Fe}^{2+}$. $\mathrm{CH}_{4}$ production rates were chosen in each scenario such that the model best represents measured values of dissolved $\mathrm{CH}_{4}$. Diffusive mixing reduces porewater gradients and therefore will also result in a reduction of isotope differences along a gradient, lowering the isotopic fractionation implemented in the model ('effective AOM fractionation') to the one derived from Figure 4 ('apparent AOM fractionation', with $\varepsilon_{\mathrm{H}}$
$=98 \%$ and $\left.\varepsilon_{\mathrm{C}}=9 \%\right)$. Effective fractionation factors for $\mathrm{SO}_{4}-\mathrm{AOM}\left(\varepsilon_{\mathrm{H}}=138 \%\right.$ and $\varepsilon_{\mathrm{C}}=13$ $\%$ ) were therefore estimated by fitting the model with the observations in the top 8.5 cm , i.e. the zone with $\mathrm{SO}_{4}-\mathrm{AOM}$ as the main mechanism effecting porewater $\mathrm{CH}_{4}$. As an approximation, isotope fractionation for $\mathrm{Fe}-\mathrm{AOM}$ was assumed to be the same as for $\mathrm{SO}_{4}{ }^{-}$ AOM. In a last step, $\delta^{13} \mathrm{D}$ and $\delta^{13} \mathrm{C}$ source signatures for methanogenesis were defined.

## SUPPLEMENTARY REFERENCES

(1) Berner, R. A. Early diagenesis: A theoretical approach; Princeton University Press, 1980.
(2) Boudreau, B. P. Diagenetic models and their implementation: Modelling transport and reactions in aquatic sediments; Springer Verlag, 1997.
(3) Poulton, S.; Canfield, D. Development of a sequential extraction procedure for iron: implications for iron partitioning in continentally derived particulates. Chem. Geol. 2005, 214, 209-221.

## SUPPLEMENTARY FIGURES



Figure S1. Results for the $\mathrm{Fe}\left(\mathrm{ox}_{1}\right)$ and Fe (carb) fractions suggest that most of the AVS (Acid Volatile Sulfide) present in the SMTZ was extracted during those two initial extraction steps. Thus, AVS was subtracted from the sum of the $\mathrm{Fe}\left(\mathrm{ox}_{1}\right)+\mathrm{Fe}(\mathrm{carb})$ within the SMTZ (Feoxides). The $\mathrm{Fe}\left(\mathrm{ox}_{1}\right)$ fraction includes ferrihydrite and lepidocrocite. $\mathrm{Fe}\left(\mathrm{ox}_{2}\right)$ is assumed to represent Fe associated with goethite, hematite and akagenéite. Fe (carb) is an estimate for Fe associated with siderite and ankerite. Fe (magn) represents Fe in recalcitrant oxides (mostly magnetite) ${ }^{3}$.


Figure S2. Determination of Fe-AOM rate based on linear intervals (marked gray) of ${ }^{13} \mathrm{CO}_{2}$ production after the initial (0 days; (A)) and second (55 days; (B)) Fe(III) addition. In total, 8 different slopes were calculated from slurry incubations of sediments from $20-30 \mathrm{~cm}$ and $30-$ 35 cm depth. Headspace ${ }^{13} \mathrm{CO}_{2}$ concentrations are corrected for dissolved ${ }^{13} \mathrm{CO}_{2}$ and changes in the slurry volume during the experiment.

## SUPPLEMENTARY TABLES

Table S1. Measured porewater concentrations at site US5B in 2012. $\delta \mathrm{D}-\mathrm{CH}_{4}$ and $\delta^{13} \mathrm{C}_{-} \mathrm{CH}_{4}$ are in \% vs. SMOW (Standard Mean Ocean Water) and vs. VPDB (Vienna Pee Dee Belemnite), respectively. $\sum \mathrm{H}_{2} \mathrm{~S}$ equals the sum of the different sulfide species $\left(\mathrm{H}_{2} \mathrm{~S}+\mathrm{HS}^{-}+\right.$ $\mathrm{S}^{2-}$ ) and "n.d." $=$ no data available. $*\left[\mathrm{SO}_{4}{ }^{2-}\right]=$ values below detection limit of $75 \mu \mathrm{M}$.

| Depth | $\mathrm{SO}_{4}{ }^{\mathbf{2 -}}$ | $\sum \mathrm{H}_{2} \mathrm{~S}$ | $\mathrm{Fe}^{2+}$ | Depth | $\mathrm{CH}_{4}$ | $\delta \mathrm{D}-\mathrm{CH}_{4}$ | $\delta^{13} \mathrm{C}-\mathrm{CH}_{4}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| [cm] | [ $\mu \mathrm{mol} / \mathrm{l}$ ] | [ $\mu \mathrm{mol} / \mathrm{l}$ ] | [ $\mu \mathrm{mol} / \mathrm{l}$ ] | [cm] | [mmol/l] | [\%o] | [\%o] |
| 0.0 | n.d. | 0.0 | 0.0 | 0.0 | 0.0 | n.d. | n.d. |
| 0.5 | 4870 | 0.0 | 0.0 | 1.0 | 0.0 | -23.9 | n.d. |
| 1.5 | 4269 | 0.0 | 5.4 | 3.5 | 0.1 | -29.1 | -52.5 |
| 2.5 | 3781 | 0.0 | 19.7 | 6.0 | 0.3 | -100.1 | -60.5 |
| 3.5 | 2801 | 0.3 | 9.0 | 8.5 | 0.8 | -184.9 | -67.6 |
| 4.5 | 2118 | 16.8 | 1.8 | 11.0 | 1.7 | -223.5 | -68.7 |
| 5.5 | 1389 | 111.1 | 1.8 | 13.5 | 1.9 | -230.2 | -68.7 |
| 6.5 | 751 | 186.7 | 1.8 | 16.0 | 2.5 | -235.8 | -68.3 |
| 7.5 | 115 | 95.4 | 1.8 | 18.5 | 3.0 | -238.5 | -68.5 |
| 8.5 | 304 | 123.7 | 1.8 | 21.0 | 3.3 | -239.1 | -68.3 |
| 9.5 | *19 | 1.3 | 275.8 | 23.5 | 3.8 | -245.8 | -68.0 |
| 11.0 | *25 | 1.8 | 275.8 | 26.0 | 3.7 | -243.6 | -68.4 |
| 13.0 | *52 | 0.7 | 551.5 | 28.5 | 4.3 | -240.5 | -68.0 |
| 15.0 | *15 | 0.9 | 827.3 | 31.0 | 4.5 | -243.2 | -67.7 |
| 17.0 | *15 | 0.8 | 1002.8 | 33.5 | 5.0 | -248.6 | -67.9 |
| 19.0 | *41 | 0.0 | 1178.3 | 36.0 | 5.2 | -244.8 | -67.5 |
| 21.0 | *19 | 0.2 | 1281.0 | 38.5 | 5.2 | -245.0 | -67.6 |
| 23.0 | *6 | 2.2 | 1371.3 | 41.0 | 5.4 | -246.9 | -67.8 |
| 25.0 | *8 | 0.2 | 1451.5 | 43.5 | 5.2 | -239.8 | -67.5 |
| 27.0 | *17 | 0.3 | 1594.4 | 46.0 | 4.9 | -237.3 | -67.5 |
| 29.0 | *7 | 0.2 | 1586.9 | 48.5 | 5.2 | -243.5 | -67.6 |
| 31.0 | *8 | 0.0 | 1684.7 | 51.0 | 5.5 | -248.8 | -67.6 |
| 33.0 | *8 | 0.2 | 1764.9 | 53.5 | 4.8 | -238.3 | -67.4 |
| 35.0 | *6 | 0.4 | 1757.4 |  |  |  |  |
| 37.0 | *8 | 0.0 | 1664.6 |  |  |  |  |
| 39.0 | *5 | 0.0 | 1810.0 |  |  |  |  |
| 41.0 | *7 | 0.0 | 1717.3 |  |  |  |  |
| 43.0 | *4 | 0.6 | 1617.0 |  |  |  |  |
| 45.0 | *6 | 0.0 | 1830.1 |  |  |  |  |
| 47.0 | *6 | 0.0 | 1827.6 |  |  |  |  |
| 49.0 | *10 | 0.0 | 1722.3 |  |  |  |  |
| 51.0 | *6 | 1.5 | 1762.4 |  |  |  |  |

Table S2. Measured solid phase sediment concentrations of different elements at site US5B in 2012. $\mathrm{Fe}_{\text {carb }}=$ carbonate associated Fe , including siderite and ankerite. $\mathrm{Fe}_{\mathrm{ox} 1}=$ easily reducible (amorphous) oxides, such as ferrihydrite and lepidocrocite. $\mathrm{Fe}_{\mathrm{ox} 2}=$ reducible (crystalline) oxides, including goethite, hematite and akagenéite. $\mathrm{Fe}_{\text {magn }}=\mathrm{Fe}$ in recalcitrant oxides (mostly magnetite). AVS $=$ Acid Volatile Sulfide (Fe-monosulfide, FeS) and CRS = Chromium Reducible Sulfide (pyrite, $\mathrm{FeS}_{2}$ ). All values are per g dry weight sediment. "n.d." = no data available.

| Depth | $\mathrm{S}_{\text {tot }}$ | $\mathrm{Fe}_{\text {tot }}$ | $\mathbf{F e}_{\text {carb }}$ | $\mathrm{Fe}_{\text {ox1 }}$ | $\mathrm{Fe}_{\text {ox } 2}$ | $\mathrm{Fe}_{\text {magn }}$ | AVS | CRS | Fe/Al | Porosity |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| [cm] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mu \mathrm{mol} / \mathrm{g}$ ] | [ $\mathrm{mol} / \mathrm{mol}$ ] |  |
| 0.5 | 52.1 | 1073.4 | 76.1 | 278.9 | 56.2 | 45.9 | 7.8 | 9.1 | 0.40 | 0.94 |
| 1.5 | 49.4 | 1239.3 | 101.0 | 235.4 | 61.2 | 46.5 | 8.5 | 4.8 | 0.40 | 0.92 |
| 2.5 | 82.4 | 1194.0 | 168.3 | 127.4 | 55.2 | 43.3 | 21.9 | 11.2 | 0.38 | 0.91 |
| 3.5 | 324.6 | 1240.6 | 172.2 | 201.0 | 36.2 | 34.1 | 166.2 | 21.2 | 0.39 | 0.91 |
| 4.5 | 479.0 | 1295.0 | 140.0 | 246.9 | 30.8 | 29.8 | 237.2 | 22.3 | 0.40 | 0.91 |
| 5.5 | 442.4 | 1132.6 | 140.3 | 231.0 | 30.2 | 29.9 | 260.0 | 25.2 | 0.39 | 0.91 |
| 6.5 | 604.5 | 1344.6 | 189.8 | 248.3 | 28.6 | 29.7 | 283.4 | 21.6 | 0.43 | 0.90 |
| 7.5 | 468.3 | 1198.8 | 172.7 | 233.9 | 27.8 | 28.2 | 188.7 | 15.2 | 0.40 | 0.89 |
| 8.5 | 694.2 | 1433.0 | 234.8 | 292.6 | 27.1 | 27.4 | 227.6 | 15.4 | 0.45 | 0.89 |
| 9.5 | 339.4 | 1203.5 | 193.1 | 110.7 | 33.0 | 31.5 | 67.1 | 27.7 | 0.39 | 0.90 |
| 11.0 | 86.4 | 1189.6 | 136.6 | 99.0 | 44.7 | 38.2 | 27.7 | 13.7 | 0.35 | 0.90 |
| 13.0 | 43.9 | 1212.2 | 139.2 | 103.7 | 46.6 | 38.1 | 18.6 | 163.5 | 0.35 | 0.90 |
| 15.0 | 40.1 | 1099.4 | 149.9 | 106.4 | 44.0 | 36.0 | 18.4 | 1.6 | 0.35 | 0.90 |
| 17.0 | 30.1 | 967.4 | 156.8 | 100.9 | 42.9 | 34.5 | 25.0 | 1.4 | 0.34 | 0.89 |
| 19.0 | 37.2 | 1129.0 | 158.7 | 105.2 | 42.4 | 35.4 | 21.8 | 31.9 | 0.35 | 0.89 |
| 21.0 | 45.4 | 1165.7 | 163.3 | 103.3 | 45.4 | 36.7 | 25.1 | 25.7 | 0.36 | 0.89 |
| 23.0 | 45.9 | 1284.3 | 164.7 | 102.4 | 47.3 | 38.4 | 21.4 | n.d. | 0.37 | 0.89 |
| 25.0 | 54.7 | 1171.0 | 164.5 | 103.4 | 47.2 | 37.8 | 11.8 | 19.5 | 0.37 | 0.89 |
| 27.0 | 63.2 | 1143.4 | 169.5 | 105.5 | 48.3 | 40.0 | 11.5 | 6.8 | 0.37 | 0.89 |
| 29.0 | 92.4 | 1267.3 | 182.7 | 100.0 | 48.5 | 39.0 | 24.4 | 13.8 | 0.38 | 0.89 |
| 31.0 | 132.1 | 1203.8 | 222.2 | 87.1 | 42.7 | 35.8 | 23.4 | 22.8 | 0.39 | 0.90 |
| 33.0 | 157.6 | 1182.3 | 249.8 | 98.8 | 48.9 | 41.3 | 46.3 | 24.9 | 0.40 | 0.90 |
| 35.0 | 117.8 | 1135.0 | 214.1 | 87.6 | 43.7 | 35.8 | 35.4 | 10.5 | 0.38 | 0.90 |
| 37.0 | 98.0 | 1160.1 | 185.9 | 110.1 | 46.3 | 37.5 | 20.8 | 13.9 | 0.39 | 0.90 |
| 39.0 | 92.4 | 1292.9 | 183.3 | 99.0 | 42.9 | 35.4 | 21.9 | 20.0 | 0.37 | 0.90 |
| 41.0 | 92.0 | 1120.3 | 183.4 | 108.4 | 46.0 | 38.2 | 21.4 | 3.3 | 0.37 | 0.89 |
| 43.0 | 100.7 | 1134.8 | 207.5 | 102.9 | 45.3 | 36.8 | 28.2 | 14.6 | 0.38 | 0.89 |
| 45.0 | 140.5 | 1296.0 | 251.8 | 92.1 | 45.0 | 35.9 | 43.1 | 22.9 | 0.40 | 0.90 |
| 47.0 | 146.9 | 1228.0 | 219.4 | 105.0 | 47.5 | 38.0 | 23.6 | 22.8 | 0.39 | 0.90 |
| 49.0 | 99.0 | 1157.3 | 228.6 | 84.8 | 41.9 | 35.3 | 35.3 | 10.1 | 0.38 | 0.89 |
| 51.0 | 121.5 | 1232.8 | 212.5 | 107.4 | 43.9 | 35.6 | 38.8 | 18.3 | 0.39 | 0.89 |

Table S3. Model parameters used for the column diffusion model simulation of the total concentration and isotopic composition of porewater $\mathrm{CH}_{4}$ at station US5B. Sim $1=$ Simulation 1 with $\mathrm{SO}_{4}$-AOM at top ( $0-8.5 \mathrm{~cm}$ ), Fe-AOM and methanogenesis below (8.5-36 $\mathrm{cm})$. Sim $2=$ Simulation 2, with $\mathrm{SO}_{4}-\mathrm{AOM}$ at top $(0-8.5 \mathrm{~cm})$, but only methanogenesis below (8.5-36 cm). Sim $3=$ Simulation 3, with $\mathrm{SO}_{4}$-AOM at top $(0-8.5 \mathrm{~cm})$, but no $\mathrm{Fe}-\mathrm{AOM}$ nor methanogenesis below ( $8.5-36 \mathrm{~cm}$ ).

| Parameter | Unit | Sim 1: With Fe-AOM | Sim 2: No Fe-AOM | Sim 3: Diffusion only |
| :--- | :---: | ---: | ---: | ---: |
| $\mathrm{SO}_{4}$-AOM rate constant | $1 /$ year | 32.2 | 32.2 | 32.2 |
| Fe-AOM rate constant | $1 /$ year | 0.64 | - | - |
| $\mathrm{CH}_{4}$ production rate constant | $1 /$ year | 3.6 | 1.6 | - |
|  |  |  |  | -138 |
| $\varepsilon(\mathrm{H}) \mathrm{AOM}$ | -138 | -13 | -138 |  |
| $\varepsilon$ (C) AOM | $\%$ | -13 | 3 | -13 |
| $\varepsilon$ (Diffusion) | $\%$ | 3 |  | 3 |
|  | $\% 0$ |  | -255 | - |
| $\delta$ D source |  | -305 | -75.7 | - |
| $\delta^{13} \mathrm{C}$ source | -78.7 |  |  |  |

