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Irreducible Water Saturation from Capillary Pressure and Electrical Resistivity Measurements

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Résumé — Estimation de la saturation irréductible en eau à partir de la pression capillaire et des mesures de la résistivité électrique — Cet article présente les résultats de mesures de la pression capillaire et de la résistance électrique de trois types de roches : le sable argileux (Berea), le quartz et le calcaire. Nous rapportons également des données expérimentales pour l'indice de résistivité, le facteur de formation, la tortuosité, la pression capillaire et les exposants de saturation en eau de ces roches dans le système air/saumure. Les données de résistivité électrique sont évaluées par la loi d'Archie modifiée. Les résultats des mesures de la pression capillaire et de la résistivité électrique en fonction de la saturation en saumure montrent l'existence de deux saturations irréductibles en eau correspondant à deux phénomènes physiques différents. Les deux valeurs des saturations irréductibles forment des limites hautes et basses qui peuvent être utilisées pour estimer la capacité de la production d'un milieu poreux. L'analyse détaillée des données a aussi montré un changement de valeur de l'exposant de saturation. Ce changement pourrait également être observé à la saturation irréductible en saumure obtenu lors des mesures de pression capillaire.

Abstract — Irreducible Water Saturation from Capillary Pressure and Electrical Resistivity Measurements — This paper presents the results of capillary pressure and electrical resistivity measurements for three kinds of core specimens: Berea sandstone, Quartz and Limestone. Experimental data of resistivity index, formation resistivity factor, tortuosity, capillary pressure, and water saturation exponent in the air/brine system for these cores are reported. The electrical resistivities data are evaluated using the modified Archie's law. Capillary pressure and electrical resistivity versus brine saturation results show the existence of two irreducible water saturations corresponding to two different physical phenomena. The two values of irreducible brine saturations yield upper and lower limits that can be used to estimate the production capacity of porous media. The detailed analysis of the data has also shown a change in the value of the saturation exponent. This change would be also be observed at the irreducible brine saturation as obtained with capillary pressure data.

INTRODUCTION

Capillary pressure is an important property used in the evaluation of pore fluid distributions, particularly in the transition zone between wetting and non-wetting phases (water to oil or gas in reservoir formations). This property is important in the evaluation of oil recovery processes. Also it can be used either to calculate oil reserves or to simulate reservoir behavior using computer modeling techniques. The capillary pressure curves also provide data on the irreducible water saturation of a reservoir rock and the entry pressure of the reservoir rock.

Capillary pressure curves are difficult to measure in the field. However, some of the parameters controlling the development of capillary pressure curves may be evaluated from electrical resistivity measurements. Therefore geophysical techniques can be used to estimate the electrical resistivity and the needed petrophysical properties related to wettability, interfacial tension, fluid saturation history, pore size distribution and pore geometry.

This paper presents experimental results obtained from porous plate and two electrode resistivity measurements. The electrical resistivities versus brine saturation data for all cores are evaluated using the modified Archie's law. The results of cementation factor, saturation exponent and tortuosity are related to the capillary pressure and electrical resistivity measurements.

1 PETROPHYSICAL PARAMETERS

Petrophysical parameters relate different types of flows: hydraulic, electrical, thermal, chemical, etc. As all types of flows follow similar laws, the parameter that may control one usually controls another (Wang and Anderson, 1992; Mitchell, 1993; Santamarina and Fratta, 2003). For these reasons, geophysical techniques that measure properties related to electrical flow are successfully used to evaluate hydraulic flow parameters. However, the interpretation of the results may be misleading as there are inherent physical differences on the different types of flows and these discrepancies cannot be ignored.

Conduction of electricity through porous media occurs by mainly two mechanisms: movement of ions through the bulk saturating electrolyte and by surface conduction of clay minerals. In the case of porous media with coarse grains, tube-like porous and percolating phases, the electrical resistivity of a formation ρ_c can be modeled as:

$$\frac{1}{\rho_f} = \frac{1 - \phi}{\rho_{mineral}} + \frac{\phi S_r}{\rho_{pore \ liquid}} + \frac{\phi (1 - S_r)}{\rho_{pore \ gas}} \tag{1}$$

where $\rho_{mineral}$ is the resistivity of the rock mineral, $\rho_{pore \ liquid}$ is the electrical resistivity of the pore fluid, $\rho_{pore \ gas}$ is the resistivity of the pore gas, ϕ is the porosity. S_r is the degree of pore liquid saturation. In most oil-bearing formation rocks, the rock mineral and pore gas are insulators.

One of the drawbacks of this formulation is that Equation 1 fails to model the surface conduction in a formation. The effect of surface conductance in a formation can be estimated by adding a surface conduction term to Equation 1 (Klein and Santamarina, 2003; Santamarina *et al.*, 2001). Then, the resistivity of the formation becomes:

$$\rho_f = \frac{\rho_{pore\ liquid}}{\phi S_r + (1 - \phi)\Theta \frac{\gamma_{\min}}{g} S_s \rho_{pore\ gas}}$$
(2)

where Θ is the surface conduction, γ_{\min} is the mineral unit weight, *g* is the acceleration of gravity, and *S_s* is the specific surface of the formation mineral.

Figure 1 shows the effect of porosity and surface conduction on different saturated rock formations. This figure is based in Equation 2, which in spite of its limitations yields boundaries to the evaluation of the different parameters. The proper interpretation of this figure provides a solid base for the analysis of geophysical studies. Furthermore, the electrical resistivity of a porous medium is also related to the micro-structural properties, including porosity, pore geometry and surface morphology of the mineral grains lining the pores. Archie (1942) proposes the following semi-empirical relationship:

$$\rho = a \phi^{-m} S_r^{-n} \rho_f \tag{3}$$

where a, m, and n are experimental determined factors.

The factor *a* varies between 0.5 and 2.5. The factor *m* is also known as the cementation factor and it varies between 1.3 and 2.5 for most rocks. The cementation factor is 1.3 for unconsolidated, clean glass beads, 2 for sandstone formation, and around 5 for carbonates (Telford *et al.*, 1990; Mavko *et al.*, 1998).



Figure 1

Typical electrical resistivity trends for non-conductive and saturated mineral rocks as modeled with Equation 2 (surface conduction $\Theta = 1.4 \cdot 10^{-9}$ S).

The factor n is known as the water saturation exponent and it has been related to several petrophysical parameters. Typically, its value is around 2 in water-wet formations and it is assumed constant for clean, sandy like formations (Archie, 1942; Mavko *et al.*, 1998). Lewis *et al.* (1988), Longeron *et al.* (1989a, b), Donaldson and Siddiqui (1989), and Moss *et al.* (2000) found that the saturation exponent nincreases to values as high as ten in oil wet formations. Furthermore, Diederix (1982), Anderson (1986), Longeron *et al.* (1989a, b), and Worthington and Pallat (1992) found that the saturation exponent is not always constant over the whole saturation range in disagreement with the early work by Archie (1942).

The formation resistivity factor F_R , is the ratio between the resistivity, ρ_o , of the rock at 100% saturated with conducting brine to the resistivity $\rho_{pore\ liquid}$ of this brine. This ratio depends on the lithological characteristics of the rock formation and the effective porosity. The formation resistivity factor is sometimes expressed as function of the percolation porosity ϕ_o that is the minimum porosity that connects the pores in a rock formation:

$$F_R = \left(\phi - \phi_o\right)^{-m} \tag{4}$$

The resistivity index I_R is expressed by Archie (1942) as:

$$\frac{\rho_f}{\rho_0} = I_R = S_r^{-n} \tag{5}$$

Swanson (1985) found that the resistivity-saturation plot shows a change in the slope at a saturation point where the corresponding capillary pressure curve indicate a transition from macroporosity to microporosity.

Finally, tortuosity is defined as the ratio of the ionic path L_i over the distance between the electrodes L_R . alternatively, according to Tiab and Donaldson (1996), tortuosity is defined as:

$$\tau = 1 + \frac{\phi_{trapped}}{\phi_{channel}} = 1 + \frac{\phi - \phi^m}{\phi^m}$$
(6)

where $\phi_{trapped}$ and $\phi_{channel}$ are porosities that correspond to the porosity of the trapped fluid that does not contribute to electrical conduction and to the porosity of the fluid that contributes to the conduction. In addition, Attia (2005) found that the tortuosity factor is affected by porosity, cementation factor and degree of brine saturation.

2 EXPERIMENTAL TESTS

This paper presents petrophysical properties of eight Berea sandstone, twelve Quartz and ten Limestone rock specimens. The experimental results were obtained using 5% NaCl brine solutions. The petrophysical properties of the specimens are tabulated in Tables 1 through 4. During the experimental testing program the following petrophysical parameters were determined for each specimen: pore size distribution, saturation versus capillary pressure curves, and electrical resistivity versus saturation curves. These data are obtained using two measurement techniques:

- porous plate
- impedance measurement methods

These two methodologies permit the evaluation of process of de-saturation of rock cores and the behavior of the water saturation exponent n. The description of testing procedure is as follows.

2.1 Porous Plate Method

The porous plate method was used to measure the capillary pressure-saturation curves for different types of core specimens (Berea sandstone, Quartz and Limestone) with 5% NaCl brine solution. The procedure has several steps, including:

- Porous plate saturation: the procedure indicated in the manual "Soil Moisture Capillary Pressure Cell" manual (Core Lab Instruments, 2000) is used to fully saturate the porous stone saturation.
- Core specimen preparation: the core specimens are measured, cleaned, dried, weighed, and saturated with the NaCl brine solution. After the saturation process is

TABLE 1

Geometric properties of tested Berea sandstone specimens

Specimen Number	Description	Diameter D_R (m)	Length L_R (m)	Total Volume $V_R (m^3)$	Dry mass M_d (g)	
21	Homogeneous	0.0380	0.0471	5.34.10-5	104.28	
23	Vertical layering	0.0373	0.0508	5.55·10 ⁻⁵	116.60	
24	Horizontal layering	0.0374	0.0502	5.51·10 ⁻⁵	115.53	
30	Vertical layering	0.0377	0.0384	4.29·10 ⁻⁵	88.49	
31	Vertical layering	0.0377	0.0323	3.61.10-5	74.57	
32	Vertical layering	0.0381	0.0368	4.20.10-5	83.69	
33	Vertical layering	0.0381	0.0359	4.09.10-5	82.55	
34	Vertical layering	0.0373	0.0377	4.12·10 ⁻⁵	86.79	

completed the core specimens are re-weighed to determine the porosity ϕ (*Table 1*).

- Testing procedure: with the core specimen placed on top of the porous plate, the air-brine capillary pressure procedure is ready to begin (*Fig. 2*). A certain air pressure p_a is applied to the top of the core specimen and the water within the core specimen is allowed to drain through the porous stone. The capillary pressure p_c is equal to the difference between the air pressure and the water pressure. When equilibrium is achieved, the air pressure and the top of the cell are removed. Then the core specimen is weighed and the degree saturation is calculated at the applied capillary pressure p_c . The specimen is returned to the cell and the procedure is repeated for higher values of the capillary pressure p_c .

The capillary pressure *versus* saturation curve is obtained and then used to estimate the pore size distribution.



Figure 2

Schematic of the capillary versus saturation cell (after *Core Lab Instruments*, 2000).

2.2 Impedance Measurements Method

An impedance analyzer *Hewlett Packard* 4279A LCZ meter is used to measure the electrical resistivity of the Berea sandstone (Berea as example) core specimens under varying degrees of brine saturation. The LCZ meter is capable of measuring inductance, capacitance, dissipation, quality factor, impedance magnitude, and phase angle in equivalent parallel and series circuits (*Hewlett Packard*, 1984). These measurements cover a frequency range from 100 Hz to 20 kHz. The HP 4297A model is controlled by a personal computer *via* an HP-IB card and the *Agilent* VEE Pro software. The card and the software allow the digitalization of the data for the analysis and interpretation of the measurement results.

Figure 3 presents a sketch of the testing setup, measurement system, and data collection system. The measurements include impedance magnitude and phase angle from 100 Hz to 20 kHz in steps of 100 Hz. The impedance analyzer is equipped with four coaxial BNC connectors and with an HP 16047A Test Feature. This feature combines the high and low voltage and current terminals and permits two terminal measurements of the electrical properties of the specimens. The two-terminal measurement method is a simple procedure to implement, however it has some drawbacks over competitive methods (*i.e.*, four-electrode method – see for example Santamarina *et al.*, 2001). One of the biggest drawbacks is the problem of electrode polarization. This phenomenon is explained next.

The measurement of the electrical resistivity of rock specimens involves the analysis of an equivalent electrical circuit where the rock core specimen is in series with the electrodes and the required peripheral electronics. The rock electrode interface feels the movement of ions on the rock specimen, and the movement of electrons on the electronic circuit. At the interface, where the use blocking electrodes is common (*i.e.*, electrodes that do not experience oxidation-reduction), ions accumulate at the interface causing electrode polarization.

The phenomenon of electrode polarization may be modeled by adding a thin air-gap L_{gap} at the interface between the electrode and the rock (Klein and Santamarina, 1997). Consider a parallel plate capacitor with a specimen of thickness L_R and cross sectional area A (*Fig. 4*). The rock can be represented as a resistor and a capacitor in parallel, with effective electric resistivity ρ_R (Ω m) and dielectric permittivity ε_R (F/m). Then, the electrical impedance for the equivalent circuit Z_{eq} is:

$$Z_{eq} = \frac{1}{j 2\pi f C_{gap}} + \frac{1}{\frac{1}{R_R} + j 2\pi f C_R}$$

$$= \frac{1}{A} \left[\frac{L_{gap}}{j 2\pi f \varepsilon_o} + \frac{L_R}{\frac{1}{\rho_R} + j 2\pi f \varepsilon_R} \right]$$
(7)

where $C_{gap} = \varepsilon_0 A/L_{gap}$ is the capacitance at the blocking electrodes, $R_R = \rho_R L_R / A$ and $C_R = \varepsilon_R A/L_R$ are the resistance and capacitance of the core specimen, and A is the cross sectional area of the capacitor.

The results presented in Figure 5 show that the effects of electrode polarization become negligible with smaller separation gap L_{gap} (*Fig. 5a*) and that the transition frequency is smaller with the greater rock resistivity ρ_R (*Fig. 5a* – Santamarina *et al.*, 2001). Therefore, electrode polarization effects are not expected in high-frequency laboratory



Figure 3

Sketch of the electrical resistivity measurement setup.



Figure 4

Equivalent two-electrode circuit and the effect of blocking electrodes.



Figure 5

Effect of electrode polarization on impedance measurement. a) Varying electrode L_{gap} for constant rock resistivity $\rho_R = 1 \Omega m$; b) Varying rock resistivity for constant electrode gap $L_{gap} = 10^{-6} \text{ m} (L_R = 0.0285 \text{ m}, A = 6.4 \cdot 10^{-4} \text{ m}, \varepsilon_0 = 8.85 \cdot 10^{-12} \text{ F/m} \text{ and } \varepsilon_R = 20 \varepsilon_0).$

measurements. However, due to specifications of the HP Impedance Analyzer, the range of measurement extend the low frequency range, where electrode polarization dominates the impedance measurement unless the deportation gap L_{gap} is very small or the resistivity of the measured material is high.

Two different types of electrodes are used in these measurements: Al foil and silver-based paint. The Al foil may act as a blocking electrode if low rate of chemical reaction occurs. However the problem with the Al foil electrodes is that separation gap L_{gap} is about the same order of magnitude as the sandstone grain size ($D_{50} \approx 0.1 \text{ mm}$) and this separation gap may change for different specimen and the applied normal pressure. The silver paint electrodes have been used in the past (see for example Longeron *et al.*, 1989a). The advantage of this type of electrodes is that they reduce the separation gap. Furthermore, the gap remains constant throughout the testing series and it does not depend on the normal pressure.

Figure 6 shows the measured impedance on Berea sandstone core specimen 22 using the Al foil electrodes and silver-based paint for different saturation S_r . It is clear that the Al foil electrode measurements are controlled by the effect of electrode polarization (*Fig. 6a*) while the silver-based paint electrode measurement removes most of the effect of electrode polarization in the impedance response (*Fig. 6b*). Furthermore, the transition frequency is smaller when the resistivity of the specimen is greater, as observed by comparing side-by-side the two set of results. In light of these observations the silver-based paint electrodes are selected for the electrical resistivity measurements.

2.3 Description of Electrical Resistivity Measurement Methodology

The core specimens are cleaned, dried in the oven, painted on each side, evacuated, and then saturated with 5% NaCl brine solution ($\rho_{pore\ liquid} = 0.161\ \Omega m$) by applying vacuum for 12 hours to the submerged specimens. The specimens are then removed and their surfaces are dried with tissue paper. This action renders the specimen saturated-surface dry. Then the specimens are weighed to determine the degree of saturation and placed in the cell to measure the impedance. After each measurement, the specimen is placed in a covered container to stabilize its moisture content till the next set of measurement. This technique reduces the instability of the capillary pressure when the measurements are taken.

3 TESTS RESULTS

A set of different petrophysical properties are collected including capillary pressure curves (*Fig.* 7), pore size distribution (*Fig.* 8), and electrical resistivity versus degree of saturation (*Fig.* 9) for eight Berea sandstone core specimens. Figure 7 shows that all specimens have similar capillary pressure saturation curves except for the Berea sandstone core specimen 21. All the other core specimens show that the irreducible water saturation is nearly 0.33, while specimen 21 shows it to be 0.23. This specimen is also the only one that shows no layering (*Table 1*). Furthermore, the pore size distribution also shows a marked difference in the characteristics of specimen 21 (*Fig.* 8). The average parameters for the Berea sandstone, Quartz (very fine grained) and Limestone specimens are summarized in Tables 2, 3 and 4 respectively.



Figure 6

Effect of type of electrode on the measurement of the impedance on Berea sandstone core specimen 22.

a) Al foil electrodes; b) Silver-based paint electrodes. Electrode polarization controls the impedance measurement with the Al foil electrodes.

Summary of porosity, tortuosity, ionic length, formation resistivity factor, cementation factor, and water saturation exponent for all tested Berea sandstone core specimens (Attia, 2005)

Specimen Number	Porosity φ	Tortuosity τ	Ionic length $L_i = L_R \sqrt{\tau}$ (m)	Formation resistivity factor F_R	Cementation factor	Water saturation exponent <i>n</i>
21	0.207	3.84	0.092	18.6	1.85	2.12
23	0.191	4.35	0.106	22.8	1.89	1.80
24	0.192	4.26	0.104	22.2	1.88	2.42
30	0.186	3.78	0.075	20.3	1.79	1.93
31	0.187	4.71	0.070	25.2	1.92	1.98
32	0.181	4.04	0.074	22.3	1.82	1.62
33	0.182	3.82	0.070	21.0	1.79	2.02
34	0.192	4.48	0.080	23.3	1.91	1.96

TABLE 3

Summary of porosity, quartz flour content, formation resistivity factor, and water saturation exponents for all tested Quartz sandstone core specimens (after Attia, 2005)

Specimen Number	Porosity φ	Quartz Flour (%)	Formation resistivity factor F_R	Water saturation exponent n (for $S_r < 0.32$)	Water saturation exponent <i>n</i> (for $S_r > 0.32$)
2	0.258	100	1.31	1.95	1.67
6	0.293	100	1.44	2.48	1.76
14	0.292	100	1.48	3.08	1.86
18	0.224	80	1.38	2.56	1.73
22	0.265	80	1.61	2.62	1.86
28	0.253	80	1.61	3.18	1.93
32	0.194	60	1.64	3.60	1.94
38	0.239	60	1.81	2.94	1.79
42	0.222	60	1.83	2.65	1.75
48	0.201	40	1.71	N/A	1.98
54	0.201	40	2.06	3.60	1.61
58	0.187	40	1.86	3.73	1.77

TABLE 4

Summary of porosity, tortuosity factor, formation resistivity factor, and water saturation exponents for all tested Limestone core specimens

Specimen Number	Porosity φ	Tortuosity τ	Formation resistivity factor F_R	Water saturation exponent n (for $S_r < 0.68$)	Water saturation exponent n (for $S_r > 0.68$)
1L	0.14	1.44	30.3	N/A	N/A
2L	0.145	1.44	29.7	N/A	N/A
3L	0.142	1.43	29.6	3.67	2.20
4L	0.143	1.44	30.1	3.36	1.97
5L	0.087	1.27	29.7	3.67	1.86
6L	0.143	1.43	29.4	3.70	2.14
18L	0.142	1.43	29.2	4.36	2.23
19L	0.143	1.45	31.2	4.55	2.29
20L	0.142	1.46	32.0	3.88	2.16
21L	0.158	1.49	31.2	1.41	3.01



4 DISCUSSION

The electrical resistivity and capillary pressure curves for the Berea sandstone, Quartz, and Limestone cores show similar behavior: as the brine saturation decreases, the values of the petrophysical parameters increase.

However, the point of irreducible brine saturation in the capillary pressure curve differs from the point where there is a break in the electrical conductivity path (see *Figs. 10-12*). This difference is due to the different nature of the two processes. The capillary pressure *versus* saturation curve is obtained by applying gas pressure to the brine phase until the air percolates through the porous matrix. This percolation point is the irreducible brine saturation. It is "irreducible" only in the sense that the air pressure cannot "push" the brine out beyond this point. Nevertheless, at the irreducible brine saturation point, the electrolyte paths are still active and there

are no sudden changes in the electrical conduction behavior. The change in the electrical conduction only occurs at a lower degree of saturation, when the conduction paths are broken and the resistivity rapidly increases.

The observations presented in Figures 10 through 12 can be clearly explained with the help of Figure 13. This figure shows the process of evacuation (drying) using the porous stone method. As the capillary pressure (air pressure) increases, it pushes the brine out of the porous medium decreasing the saturation and increasing the electrical resistivity of the medium (*Figs. 13a-c*). This process continues until the air paths between the top and bottom of the specimen are continuous (*Fig. 13d*). At this point, the irreducible water saturation in the "capillary pressure sense" is reached. A further increase in the capillary pressure cannot remove anymore brine and the electrical resistivity should remain constant (*Fig. 13e*). However, if the medium is allowed to dry due to evaporation the electrical resistivity would still slowly increase because the ionic paths will be strangled as the brine saturation decreases. This process continues until all the ionic paths are broken and electrical resistivity reaches its final plateau (*Fig. 13f*). From Equation 1, the theoretical electrical resistivity of the plateau is approximately equal to:

$$\rho_f = \left(\frac{1-\phi}{\rho_{mineral}} + \frac{\phi(1-S_r)}{\rho_{pore\ gas}}\right)^{-1} \tag{8}$$

The irreducible brine saturation obtained with either method can not be directly used to evaluate the production potential of a reservoir. They can only be considered as upper and lower limits of the irreducible brine saturation. These two measured levels of brine saturation yield an indication of the percolation phenomena. The true value of the irreducible water saturation will depend on the production method and the affinity of the mineral to the fluid.

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One common parameter used in the evaluation of the production potential of a formation is the saturation exponent. The saturation exponent in the measured core specimens is not constant for different degree of saturations as indicated by Archie (1942) even for clean Berea sandstone (Anderson, 1986). Figures 14a, 15a and 16a show that average saturation coefficient changes from 1.60 to 2.70, 2.00 to 3.50 and 2.09 to 3.84 for Berea sandstone, Quartz, and Limestone core specimens respectively, when the degree of saturation drops below a certain value. The break in the slope of $log(I_R)$ versus $log(S_r)$ occurs at the point that coincides with the capillary pressure-based irreducible brine saturation (see Figs. 14b, 15b and 16b) and it does help in the evaluation of the electrical resistivity-based irreducible saturation that provides the higher bound in the production potential of the formation.



Figure 8 Pore size distribution for all eight Berea sandstone core specimens.



The change in slope seems to indicate the beginning of the process that produces the rupture of electrolyte paths in the porous media. At degrees of saturation higher than the point of sudden change in the saturation exponent, the increase of electrical resistivity seems to occurs only due to the decrease in the cross sectional area of continuous electrolyte paths. Then a change in the behavior seems to occur at the point of capillary pressure-based irreducible brine saturation where the brine phase still percolates at a minimum through the surface of the porous. At this point the isolating non-wetting phase zone starts to become dominant over the conductive wetting phase.

Finally, the evaluation of tortuosity versus brine saturation is presented in Figure 17. The tortuosity is here calculated by modifying Equation 6 to include the effect of saturation:

$$\tau = 1 + \frac{\Phi_{trapped}}{\phi_{channel} \left(S_r - S_{riE} \right)} = 1 + \frac{\phi - \phi^m}{\phi^m \left(S_r - S_{riE} \right)}$$
(9)
for $S_r > S_{riE}$

where S_{riE} is the irreducible water in the electrical resistivity sense. This new equation assumes that the trapped porosity does not contribute to the electrical conductivity and the tortuosity reaches infinite value when all the brine in the channel is removed (*Fig. 17a*). The normal-normal plot (*Fig. 17a*) clearly shows the value of lower value of the irreducible water saturation, while it does not present a clear break in the log-log plot (*Fig. 17b*). The tortuosity provides an indication of the twisting of ionic path and it appears as it only changes at a slowly rate and it does not suddenly alter the conduction paths until the last brine percolation path is broken.

CONCLUSIONS

This paper presents the results of capillary pressure and electrical resistivity versus degree of saturation for three types of rocks. Berea sandstone, Quartz and Limestone core specimens were measured with the porous stone method and with the two-electrode method to evaluate the capillary



Figure 10

Capillary pressure and resistivity index *versus* brine saturation in Berea sandstone core specimens. Two irreducible water saturations are identified: one for capillary pressure and another one for the electrical properties.



Figure 11

Capillary pressure and resistivity index *versus* brine saturation in twelve Quartz (very fine-grained) sandstone core specimens. Two irreducible water saturations are identified: one for capillary pressure and another one for the electrical properties.



Figure 12

Capillary pressure and resistivity index versus brine saturation in ten Limestone core specimens. Two irreducible water saturations are identified: one for capillary pressure and another one for the electrical properties.





Effect of percolation on the capillary pressure and electrical resistivity measurements.



Figure 14

Effect of brine saturation the saturation exponent in Berea sandstone specimens. a) Log-log plot of the resistivity index *versus* brine saturation. b) Saturation exponent *versus* porosity.



Figure 15

Effect of brine saturation on the saturation exponent in Quartz sandstone specimens. a) Log-log plot of the resistivity index *versus* brine saturation; (b) Saturation exponent *versus* porosity.



Figure 16

Effect of brine saturation on the saturation exponent in Limestone specimens. (a) Log-log plot of the resistivity index *versus* brine saturation; b) Saturation exponent *versus* porosity.



Calculated tortuosity versus degree of saturation in Berea sandstone specimens. a) Normal-normal plot.(b) Log-log plot.

pressure and electrical conductivity versus brine saturation curves. In the electrical resistivity measurements, the proposed silver-based paint electrodes prevent problems with electrode polarization for the tested frequency range.

Capillary pressure and electrical resistivity versus saturation curves show that the irreducible brine saturation is a parameter that depends on the type of conduction and its corresponding value of saturation at percolation. The two values of irreducible brine saturation (Figs. 10-12) yield an upper and lower bound limits that can be used to estimate limits to the production capacity of the porous medium. The detailed analysis of the data shows that the change in the value of the saturation exponent. This change occurs at the irreducible brine saturation as obtained with capillary pressure data. The change in the slope (Figs. 14a, 15a and 16a) seems to indicate the presence of discontinuous ionic path in porous media, however this observation must be confirmed with complementary testing techniques (e.g., micro-tomography). Finally, an equation to evaluate the tortuosity versus degree of saturation is proposed. Tortuosity results provide an indication of the distortion of the ionic path. Tortuosity appears to change at slow rates and the conduction paths show no sudden changes until the last brine percolation path is broken.

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