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BELOWGROUND RESPONSES TO CLIMATE CHANGE

Long-term enhancement of N availability and plant growth under elevated CO₂ in a semi-arid grassland

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Summary

1. While rising atmospheric CO_2 has the potential to enhance plant growth and biomass accumulation, rates of these processes may be constrained by soil nitrogen (N) availability. Despite much effort, it is still uncertain how elevated CO_2 affects long-term soil N dynamics.

2. We used open-top chambers to examine the effect of 5 years of elevated atmospheric CO_2 concentration (720 vs. 368 p.p.m.) on N dynamics in a semi-arid grassland ecosystem in north-eastern Colorado, USA. In the first year 0.5 g m⁻² of ammonium nitrate-N, 99.9 atom% ¹⁵N, was added to each plot. We examined the effect of elevated CO_2 on N mineralization and plant N uptake by tracking the labelled and total N in plant and soil over the following 5 years.

3. Plant growth and plant N uptake remained significantly higher under elevated than under ambient CO_2 . The fraction of labelled N (expressed per unit of total N) in above-ground biomass declined over time, and this decline was greater under elevated CO_2 . The amount and fraction of labelled N in the soil did not change with time and was unaffected by elevated CO_2 . These results suggest that with time, N released from mineralization in the soil diluted the labelled N in above-ground biomass and that this dilution effect caused by N mineralization was greater under elevated CO_2 . More of the mineralized N ended up in the above-ground biomass of *Stipa comata* and forbs (C_3) than in *Bouteloua gracilis* (C_4) under elevated CO_2 .

4. Increased soil moisture under elevated CO_2 likely supported higher rates of N mineralization, thereby reducing N constraints on plant growth. Therefore, in semi-arid systems, plant growth and species composition responses to elevated CO_2 may be more persistent than in mesic systems where N mineralization is less constrained by soil moisture.

Key-words: climate, ¹⁵N tracer, nitrogen mineralization, nitrogen uptake, soil moisture

Introduction

The global atmospheric CO₂ concentration has increased from a pre-industrial value of about 280–368 p.p.m. in 2001, and could more than double by the end of this century (Intergovernmental Panel on Climate Change 2001). Increases in atmospheric CO₂ concentration (hereafter also referred to as elevated CO₂) stimulate plant growth (Ainsworth & Long 2005), alter species composition (Smith *et al.* 2000), and may increase soil C storage, thereby slowing the rate of further increases in atmospheric CO₂ (Thompson *et al.* 2004). However, it remains uncertain if ecosystem responses to elevated CO₂ can be sustained (Dufresne *et al.* 2002; Reich *et al.* 2006a).

In particular, soil nitrogen (N) availability for plant growth may constrain the extent to which rising atmospheric CO₂ enhances plant and soil C sequestration (Hungate *et al.* 2003; Luo *et al.* 2004; de Graaff *et al.* 2006; Reich *et al.* 2006a; Van Groenigen *et al.* 2006). It has been postulated that elevated CO₂ may reduce soil N availability, also referred to as the concept of progressive N limitation (PNL, Luo *et al.* 2004). Overall, PNL develops when proportionally more of the available soil N is fixed into long-lived plant biomass and soil organic matter (SOM) under elevated CO₂. Therefore, PNL could eventually counteract the initial stimulation of plant growth

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and C sequestration in response to elevated CO_2 (Luo *et al.* 2004; Reich, Hungate & Luo 2006b).

Extensive research has been done on how elevated CO_2 affects soil N availability, but results have been inconsistent. Elevated CO_2 can increase (Hungate *et al.* 1997; Ebersberger, Niklaus & Kandeler 2003), have no effect on (Williams, Rice & Owensby 2001; Finzi & Schlesinger 2003; Zak *et al.* 2003), or decrease soil N availability under field conditions (Gill *et al.* 2002; Reich *et al.* 2006a). Moreover, most field experiments in which CO_2 has been manipulated have been relatively short in duration (3 years or less), thereby reducing their utility for understanding long-term effects of elevated CO_2 on soil N availability.

Understanding how CO₂ affects soil N availability has been further hampered by difficulties measuring soil N availability under field conditions. Soil N availability has been estimated by measuring net N mineralization in soil incubations (in the field or in the laboratory, Gill *et al.* 2002; Ebersberger *et al.* 2003; Finzi & Schlesinger 2003; Reich *et al.* 2006a) or by measuring short-term gross mineralization and immobilization reactions (up to 48 h) through additions of ¹⁵N to the soil (Hungate *et al.* 1997; Williams *et al.* 2001; Zak *et al.* 2003). However, soil incubations have severe limitations, in part because they are performed in the absence of plants (Schimel & Bennett 2004), and short-term measurements of gross N mineralization and immobilization are usually not sensitive enough to detect responses in soil N availability to elevated CO₂ (Zak *et al.* 2000; de Graaff *et al.* 2006; Reich *et al.* 2006b).

Net primary productivity and species composition responses to elevated CO_2 are predicted and have been shown to be among the largest in semi-arid and arid ecosystems (Strain & Bazzaz 1983; Melillo et al. 1993; Smith et al. 2000; Morgan et al. 2004b). While these responses may directly be related to improved water conditions, an increase in soil N availability under elevated CO_2 may be also important. When elevated CO_2 increases soil moisture and N availability, it could potentially delay the development of PNL. This effect of elevated CO₂ on soil moisture and N availability could particularly be important in semi-arid and arid ecosystems where soil moisture plays a key role in decomposition and mineralization. We hypothesize that soil N availability remains persistently higher under elevated than under ambient CO₂ in this semi-arid grassland causing sustained greater plant N uptake and plant growth during 5 years of elevated CO₂. We used a novel ¹⁵N tracer technique to obtain an integrated long-term measurement of how elevated CO2 influenced soil N availability. This new method further takes into account the direct effects of plants on N mineralization, thereby circumventing the pitfalls and shortcomings associated with soil incubation methods.

Methods

This experiment was established in 1996 on a native rangeland pasture at the USDA-ARS Central Plains Experimental Range, Colorado, USA. Mean annual precipitation is 321 mm, with the majority occurring in May, June and July, and mean air temperatures are $15.6 \,^{\circ}$ C in July and $0.6 \,^{\circ}$ C in January. Dominant species are the warm-season C4 grass Bouteloua gracilis and the cool-season C3 grasses Pascopyrum smithii and Stipa comata (together comprising c. 88% of the above-ground biomass in 1996). Other species include the sedge Carex eleocharis and the sub-shrub Artemisia frigida (each 4% of above-ground biomass in 1996). In March 1997 six open-top chambers (4.5 m diameter) were installed on the pasture, three chambers with ambient CO₂ (360 \pm 20 μ mol/mol) and three with elevated CO₂ $(720 \pm 20 \,\mu\text{mol/mol}, \text{Morgan et al. 2004a})$. Chambers were removed in late October each year after plant senescence and reinstalled in March before plant growth began. Baseline plant data were collected in 1996 (same protocol used during 1997-2001, see below), prior to the installation of the chambers. In early April, 1997, we uniformly sprayed 0.5 g N m⁻² of ammonium nitrate-N, 99.9 atom% ¹⁵N (both ammonium-N and nitrate-N were 99.9 atom% ¹⁵N, hereafter 'labelled N') as a solution (10 mm of water was applied during the N application) to the whole area of each of the plots.

Above-ground plant biomass was collected each year from 1997 to 2001 at the time of peak standing biomass (late July). A metal grid containing fifty-six 40.5×15.3 cm quadrats (total of 3.46 m²) was placed inside the plots, and vegetation in every other quadrats (28 quadrats) was clipped to the crown. The 28 unclipped quadrats were clipped at peak standing biomass the next year. By October aboveground biomass had senesced and all 56 quadrats were harvested each year, including the 28 quadrats that were not clipped that previous summer and the 28 quadrats that were clipped (regrowth biomass). In July each year, the above-ground biomass was separated by the species B. gracilis, P. smithii, S. comata and C. eleocharis, while the remaining biomass was grouped into species groups 'forbs' (mostly the sub-shrub A. frigida), 'other C₃' and 'other C₄' grasses. Each year in October two 20-cm diameter cores to 60-cm depth were removed from each plot and separated into above- and belowground plant biomass and soil. All above- and below-ground plant biomass and soil samples were dried at 60 °C and weighed. Plant and soil samples were ground and analysed for total N and ¹⁵N by combustion/isotope ratio mass spectrometry.

We calculated the labelled N expressed as a fraction of total N in plant biomass in year 'i' (N_{label,}, in mg g⁻¹) using the following equation:

$$N_{label,i} = \frac{\binom{1^5 N_{plant,i} - {}^{15} N_{plant,96}}}{\binom{1^5 N_{label} - {}^{15} N_{plant,96}}} \times 1000$$

where ${}^{15}N_{plant,i}$, ${}^{15}N_{plant,96}$, and ${}^{15}N_{label}$ are the atom% ${}^{15}N$ in plant biomass measured in year *i*, in plant biomass measured in 1996 and of the labelled N added, respectively. We calculated the fraction of labelled N in the soil similarly. We assumed that a decrease in N_{label,i} over time was due to uptake of unlabelled N from the soil, and that the rate of this dilution of labelled N with unlabelled N would be positively related to N mineralization in the soil (see discussion below). We calculated ${}^{15}N$ retention in the top 60-cm of the soil (${}^{15}N_{retention,soil}$, the amount of labelled N recovered in the soil in mg m⁻²) from the core samples harvested each year in October using the following equation:

$${}^{15}N_{\text{retentionsoil}} = N_{\text{soil},i} \times \frac{{}^{15}N_{\text{soil},i} - {}^{15}N_{\text{soil},96}}{{}^{15}N_{\text{label}} - {}^{15}N_{\text{soil},96}}$$

where $N_{soil,i}$ is the total amount of N in the soil measured in year $i \text{ (mg m}^{-2)}$, and ${}^{15}N_{soil,i}$ and ${}^{15}N_{soil,i,6}$ are the atom% ${}^{15}N$ in the soil measured in year i and in 1996, respectively. We also calculated ${}^{15}N$ retention in the whole ecosystem by summing the ${}^{15}N$ retention in the soil and in above- and below-ground biomass.

We used repeated-measures ancova to test for CO_2 treatment effects on above-ground biomass, N concentration, and total N in

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above-ground biomass (between-subject) over time (within-subject). Because of differences among plots in above-ground biomass, N concentration, and total N in above-ground biomass collected in 1996, we used these variables as covariates. There were no significant $CO_2 \times time$ interaction effects (P > 0.1). We then tested for CO_2 treatment effects on average total N in above-ground biomass (average from 1997 to 2001) using ANCOVA with above-ground biomass, N concentration, and the total N in above-ground biomass collected in 1996 as a covariate. Because above-ground biomass collected in July was separated by species and species groups, we were able to test for CO₂ treatment effects on N content in the July above-ground biomass among species and species groups over time, using repeated-measures ANCOVA (CO2 and species/species groups as main factors and above-ground N content in 1996 as a covariate). We used repeated-measures ANOVA to test for CO2 and species effects on labelled N fractions in above-ground biomass (log-transformed) and soil, and 15N retention in plant and soil over time. We also used ANOVA to test for CO2 effects on labelled N fractions in above-ground biomass in each year.

Results

Above-ground plant biomass under elevated CO₂ was significantly greater than under ambient CO₂ (P < 0.05, 5-year average) and had a significantly greater total N content by the time of senescence in October (up to 68%, P < 0.05, Fig. 1). The increase of the above-ground N pool in October under elevated CO₂ was sustained during the 5-year period (no significant time \times CO₂ interaction using repeated measures ANCOVA). The N concentration in the July above-ground biomass was significantly lower under elevated CO_2 (P < 0.001) than under ambient CO₂, which resulted in similar amounts of above-ground plant N in July under elevated and ambient CO₂. Above-ground plant biomass was reported earlier by Morgan et al. (2004a) and plant N concentration and content for the first 3 years by King et al. (2004). Elevated CO₂ had no effect on below-ground plant biomass or N content measured each year in October (data not shown).

Table 1. Repeated-measures ANCOVA results for N content (g m⁻²) in above-ground biomass collected in July by species (*B. gracilis*, *P. smithii*, *S. comata* and *C. eleocharis*) or species groups (forbs, other C₃ and other C₄ grasses)

Source of variation	d.f.	F	Р
CO ₂	1	0.11	0.74
Sp/sp gr	6	10.03	< 0·0001
Year	4	1.72	0.15
$CO_2 \times sp/sp gr$	6	2.77	0.06
$CO_2 \times year$	4	2.15	0.08
$Sp/sp gr \times year$	24	5.08	< 0·0001
$CO_2 \times sp/sp \text{ gr} \times year$	24	1.65	0.04
1996 above-ground biomass	1	61.36	< 0.0001

P-values in italics when P < 0.10, and in bold when P < 0.05.

Species and species groups responded differently to elevated CO_2 in terms of their above-ground N content collected in July ($CO_2 \times$ species/species group interaction P = 0.06), and these species/species group-specific responses significantly changed over time ($CO_2 \times$ species/species group \times year interaction P = 0.04, Table 1). Elevated CO_2 reduced above-ground N content in the C_4 grass *B. gracilis*, but increased above-ground N content in the C_3 grass *S. comata* and forbs, particularly during the last 2 years of the experiment (Fig. 2). The other species and species groups did not respond to elevated CO_2 . More than 80% of the above-ground N was tied up in the species *B. gracilis*, *P. smithii* and *S. comata* (data not shown). In the final year, above-ground N content in the forbs increased sevenfold in response to elevated CO_2 .

The ¹⁵N in plant biomass, expressed as a fraction of total plant N (labelled N fraction) clipped in July and in October, spiked in year 1, but then significantly declined over time (Fig. 3, Table 2), indicating an initial uptake of N enriched in ¹⁵N (i.e. directly after ¹⁵N application) followed by a progressive dilution of labelled N with unlabelled N. This progressive ¹⁵N



Fig. 1. (a) Average above-ground plant biomass, (b) N concentration, and (c) total N pool in above-ground biomass (5-year average) at ambient CO_2 (open bars) and elevated CO_2 (filled bars). All data were adjusted for pre-treatment (1996) values (least square means + SE from ANCOVA are shown). For October, results are shown for grids that were not harvested ('October') and for grids that were harvested ('October-regrowth') that same year in July. Significance levels shown: ns: not significant, *P < 0.05, **P < 0.01, ***P < 0.001.

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Fig. 2. The average change in above-ground biomass N content caused by elevated CO_2 (above-ground biomass N content at elevated CO_2 – above-ground biomass N content at ambient CO_2 in g m⁻²) over time for the different species and groups of species.

dilution was larger under elevated CO_2 than under ambient CO_2 (significant $CO_2 \times$ year interaction, P < 0.05). The relative difference in the labelled N fraction between the ambient and elevated CO_2 treatment steadily increased over time (Fig. 3d). The fraction of labelled N in above-ground biomass signifi-

cantly differed among species/species groups (P < 0.0001), and that difference changed with time (significant species/species group × year interaction, P < 0.0001). However, species and species groups did not respond differently to elevated CO₂ (no significant CO₂ × species/species group or CO₂ × species/ species group × year interactions, Table 2). The increased ¹⁵N dilution with elevated CO₂ occurred in all species and species groups (only shown for *B. gracilis, S. comata* and forbs, the species and species group that responded most to elevated CO₂ in terms of above-ground N content, Fig. 3e–h). The variability in the amount and fraction of ¹⁵N in below-ground biomass was large, and there was no CO₂ effect (data not shown).

The fraction of labelled N in the soil was not affected by CO_2 and did not change with time (Fig. 4, Table 2). This suggests that the dilution of labelled N in above-ground plant biomass under elevated CO_2 and with time was not caused by changes in the fraction of labelled N in the soil. Further, loss of ¹⁵N appeared not to have been affected by elevated CO_2 , since ¹⁵N retention in plant and soil during the 5 years of measurement was similar under ambient and elevated CO_2 (Fig. 5, Table 2). However, ¹⁵N retention in plant and soil significantly declined over time. The loss of ¹⁵N during the first year (only *c*. 400 mg of ¹⁵N was retained of the 500 mg applied) may have been caused by a spike in ammonia volatilization directly after the application, while removal of ¹⁵N through plant harvesting



Fig. 3. Labelled N fractions (expressed per total amount of N) in above-ground biomass over time, (a) in July, (b) in October from grids that were not harvested that same year in July, (c) in October from grids that were harvested that same year in July, and (e–g) in July for *B. gracilis*, *S. comata* and forbs. For each year CO₂ treatment effects were tested with ANOVA (ns: not significant, †P < 0.05, **P < 0.001). (d and h) The labelled N fraction in above-ground biomass from the elevated CO₂ treatment expressed as a relative percentage of the labelled N fraction in above-ground biomass from the ambient CO₂ treatment.

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Fig. 4. Labelled N fractions (expressed per total amount of N) in soil to 60-cm soil depth over time.

and gaseous N loss may have caused a further decline in 15 N retention with time. The total amount of 15 N removed through harvesting after 5 years was 67 and 66 mg m⁻² under ambient and elevated CO₂, respectively.

Discussion

Our results indicate that enhanced plant production in response to elevated CO_2 did not lead to a progressive decline in soil N availability over 5 years. On the contrary, the persistently higher N uptake and greater ¹⁵N dilution in plants suggest that soil N availability remained higher after 5 years of elevated than ambient CO_2 . Greater amounts of N in senescing aboveground biomass under elevated CO_2 by the end of the growing season (i.e. after N translocation to roots) indicate that plants under elevated CO_2 lost more N through litter-fall. However, elevated CO_2 resulted in a sustained larger N pool in above-ground biomass during the 5 years of study, suggesting that more N was taken up each year from the soil under elevated CO_2 .

The ¹⁵N fraction in plants increased directly after the ¹⁵N application (year 1), which was followed by a dilution of ¹⁵N

Table 2. Repeated-measures ANOVA results (A) for labelled N fractions in above-ground biomass and soil and (B) for 15 N retention in soil and plant + soil. Sp, species; sp gr, species group

Source of variation	d.f.	F	Р
A. Labelled N fractions in above	e-ground bio	omass and soil (1	mg ¹⁵ N g ⁻¹ N)
Above-ground biomass, Ju	ly		
CO_2	1	4.57	0.09
Year	4	558.51	< 0·0001
$CO_2 \times year$	4	3.81	0.02
Above-ground biomass, Ju	ly, by speci	es/species grou	ps
CO_2	1	5.29	0.03
Sp/sp gr	6	14.74	< 0.0001
Year	4	731.21	< 0.0001
$CO_2 \times sp/sp gr$	6	0.27	0.94
$CO_2 \times year$	4	2.23	0.07
Sp/sp gr × year	24	6.98	< 0.0001
$CO_2 \times sp/sp \ gr \times year$	24	0.38	0.99
Above-ground biomass, Oc	ctober		
CO_2	1	2.97	0.16
Year	4	161.57	< 0.0001
$CO_2 \times year$	4	2.94	0.02
Above-ground biomass, Oc	ctober, regr	owth	
CO_2	1	5.19	0.09
Year	4	310.77	< 0.0001
$CO_2 \times year$	4	2.34	0.09
Soil			
CO_2	1	0.002	0.97
Year	4	1.52	0.26
CO ₂ *year	4	2.06	0.15
B. ^{15}N retention (mg ^{15}N m $^{-2}$)			
Soil			
CO_2	1	0.08	0.79
Year	4	1.25	0.17
$CO_2 \times year$	4	1.20	0.16
Plant + soil			
CO_2	1	0.19	0.69
Year	4	7.75	0.001
$CO_2 \times vear$	4	2.20	0.12
- 2 2			

P-values in italics when P < 0.10, and in bold when P < 0.05.

in above-ground biomass over time. Most likely, progressive plant uptake of mineralized soil N that had a much lower fraction of labelled N diluted the ¹⁵N in biomass over time. The fraction of labelled N in the first 60 cm of the soil profile was much lower than the labelled N fractions observed in aboveground plant biomass (compare Fig. 4 with Fig. 3). While the fraction of labelled N of mineralized soil N may have been slightly higher than the fraction of labelled N in the total



Fig. 5. Average ¹⁵N retention + SE (a) in soil to 60-cm soil depth and (b) in plant (aboveand below-ground biomass) and soil under ambient (open bars) and elevated CO_2 (filled bars) over time. Data are from 20-cm diameter cores harvested each year in October.

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soil N pool (some of the non-labelled N in the soil may not have been accessible for microbes to mineralize), these data strongly suggest that the dilution of ¹⁵N in above-ground plant biomass with time was largely caused by uptake of mineralized N with a generally low labelled N fraction. This ¹⁵N dilution was observed to a greater degree under elevated than ambient CO₂, suggesting a greater mineralization rate under elevated CO₂. In a study with resin bags conducted during the first 3 years of the experiment we found no significant CO₂ effects on the amount of ammonium and nitrate absorbed (D.G. Milchunas, unpublished data). However, we believe that the amount of ammonium and nitrate absorbed on the bags may not only have been affected by N mineralization in the soil, but also by plant N uptake.

Atmospheric N deposition and atmospheric N fixation could potentially have diluted the labelled N fraction in the soil over time. However, the labelled N fraction in the soil did not significantly change with time (Table 2). It is unlikely that elevated CO_2 affected atmospheric N deposition, and atmospheric N fixation was most likely negligible (there were very few N-fixing plants).

It is possible that 15 N dilution was greater under elevated CO₂ because of greater exploration of non-labelled N from deep soil layers. However, there was no evidence for root growth to a greater depth to increase access of deep soil N under elevated CO₂ (LeCain *et al.* 2006) despite changes in species composition.

The greater dilution of ¹⁵N in above-ground plant biomass under elevated CO₂ probably was not caused by treatment effects on ¹⁵N loss from the ecosystem. Frequent harvesting of above-ground plant biomass and removal of ${\rm ^{15}N}$ did not result in greater dilution of plant ¹⁵N under elevated CO₂. The total amount of ¹⁵N removed after 5 years of harvesting was the same under ambient and elevated CO_2 (67 and 66 mg m⁻², respectively). The harvested biomass under elevated CO₂ had a slightly lower fraction of labelled N than under ambient CO_2 (5-year average of 12·2 and 10·2 mg ¹⁵N g⁻¹ total N for ambient and elevated CO₂, respectively), while both fractions were much greater than the soil labelled N fraction. Thus, under elevated CO₂, harvesting depleted the soil of ¹⁵N slightly less than under ambient CO2. Therefore, harvesting actually counteracted our observation of enhanced 15N dilution in aboveground plant biomass under elevated CO₂. There was no evidence for CO₂ effects on N loss through nitrification or denitrification (Mosier et al. 2002). We did not directly measure N loss through leaching, but N loss through leaching (as well as through nitrification or denitrification) was most likely small because of a tight N cycle (Mosier et al. 2002). Further, elevated CO₂ did not affect ¹⁵N retention in plant and soil during the 5 years of measurement, suggesting that there were no significant CO₂ effects on ¹⁵N loss.

It is unlikely that the greater ¹⁵N dilution was caused by reduced decomposition and mineralization of ¹⁵N enriched litter produced during exposure to elevated CO₂ because aboveground litter and root decomposition experiments revealed no significant CO₂ treatment effects (J.Y. King and E. Pendall, unpublished data). Therefore, the increased N uptake in response to elevated CO₂ and the greater dilution of plant ¹⁵N was likely caused by enhanced mineralization of relatively low-labelled soil N.

The C₃ grass S. comata and the forbs appeared to benefit more from the increase in N mineralization under elevated CO_2 than the C_4 grass *B. gracilis*. Nitrogen content in aboveground biomass of S. comata and the forbs increased in response to elevated CO₂, while it decreased in above-ground biomass of B. gracilis. Soil moisture was increased by elevated CO2 (Nelson et al. 2004), which possibly enhanced N mineralization at depth and contributed to the increased response in productivity and abundance of S. comata and the sub-shrub A. frigida, both of which are more deeply rooted compared to B. gracilis (Morgan et al. 2004a; Morgan et al. 2007). Differences in the fraction of labelled N in above-ground biomass among plant species and species groups also suggest N uptake from different soil depths by the different plant species. The ¹⁵N was added to the surface of the soil causing greater ¹⁵N enrichment at the surface (data not shown). Therefore, with a shallow root structure, B. gracilis may have taken up a greater fraction of labelled N in the above-ground biomass compared to the other species. There were no significant $CO_2 \times species/$ species group or $CO_2 \times$ species/species group \times year interaction effects on the fraction of labelled N in plant biomass, supporting the observation that elevated CO₂ did not alter rooting depth of individual species (LeCain et al. 2006).

Our results contrast with results from grassland field experiments in Texas and Minnesota in which elevated CO₂ reduced soil N availability within 2-6 years (Gill et al. 2002; Reich et al. 2006a). Elevated CO2 increased soil moisture due to decreased plant transpiration at our site (Nelson et al. 2004), which could have stimulated microbial activity and N mineralization. This CO2-induced soil moisture effect on microbial activity is likely to be more important in semi-arid grasslands than in mesic grasslands where microbial activity in the soil is less constrained by soil moisture. Although stimulation of plant growth may be directly related to increased water availability, a simultaneous increase in N mineralization may be an important 'hidden' variable controlling plant productivity in semi-arid grasslands (Burke, Lauenroth & Parton 1997). Thus, elevated CO2 may have increased soil N availability through plant control of soil moisture thereby preventing progressive plant N limitation within the first 5 years of elevated CO_2 .

Different patterns of soil N availability in response to elevated CO_2 in different studies may also stem in part from the different methods used to measure soil N availability. In the Texas and Minnesota studies, soil N availability was assessed by measuring net N mineralization in soil incubations. Because plants can interfere with these measurements through N uptake, measurements are typically done in the absence of plants. However, plant–soil interactions can significantly affect net N mineralization by enhancing SOM decomposition, known as the rhizosphere priming effect (Kuzyakov, Friedel & Stahr 2000). The rhizosphere priming effect could increase with elevated CO_2 (Cheng 1999; Hoosbeek *et al.* 2004). Thus, soil incubation assays that exclude plants could underesti-

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mate N mineralization rates, particularly under elevated CO₂. Indeed, at our site elevated CO₂ significantly increased SOM decomposition and enzyme activities (Pendall *et al.* 2003; Kandeler *et al.* 2006). Our ¹⁵N tracer method included potential rhizosphere priming effects on soil N availability. Moreover, by tracing the added ¹⁵N in plants over a 5-year period, we were able to integrate long-term effects of elevated CO₂ on soil N availability, which made our method very robust.

Our results suggest that elevated CO₂ effects on ecosystem functioning may be larger and more persistent in semiarid climates than in wetter climates. While increased plant productivity and biomass accumulation with elevated CO₂ is expected to be constrained by soil N availability (Hungate et al. 2003; de Graaff et al. 2006; Reich et al. 2006a; Van Groenigen et al. 2006), and even could result in PNL (Luo et al. 2004), our results suggest that in semi-arid climates, a CO₂-induced increase in soil moisture can reduce constraints of available N on plant productivity. Plant growth responses to elevated CO₂ in semi-arid climates, where a CO₂-induced increase in soil moisture has been shown to occur (Morgan et al. 2004b), may therefore be more pronounced and last longer than in climates with more effective precipitation and in which N mineralization is less influenced by a CO2-induced change in soil moisture. Indeed, semi-arid ecosystems have been predicted to be some of the most responsive ecosystems to elevated CO₂ (Strain & Bazzaz 1983; Melillo et al. 1993; Morgan et al. 2004b). Other ecosystem properties that are influenced by N availability, such as plant species composition and C sequestration, may also show larger and more persistent responses to elevated CO2 in semi-arid climates. For instance, growth of exotic annual grasses in the Mojave desert (NV, USA) and a weedy sub-shrub at our site showed dramatic increases in response to elevated CO₂ (Smith et al. 2000; Morgan et al. 2007). Shifts in species composition in response to elevated CO₂ could strongly impact overall forage digestibility (Morgan et al. 2004a; Milchunas et al. 2005) where grazing by domestic livestock is the primary land-use. Estimating long-term global changes in ecosystem functioning in a CO₂-rich environment requires an understanding of ecosystem responses to CO₂induced changes in water and N availability.

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