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1	Microbial uptake and utilization of low molecular weight organic substrates in
2	soil depend on carbon oxidation state
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27 Abstract

The fate of low molecular weight organic substances (LMWOS) in soil is regulated by microbial uptake. However, C oxidation state, the number of C atoms and -COOH groups in the LMWOS can affect their microbial utilization. Thus, the aim of this study was to reveal the effects of substance chemical properties on initial uptake and utilization of sugars, carboxylic and amino acids by microorganisms.

Soil solution, spiked with ¹⁴C-labelled glucose, fructose, malate, succinate, formate, alanine or
 glycine, was added to the soil and ¹⁴C was traced in the soil solution, CO₂, cytosol, and soil organic
 carbon (SOC) over 24 hours.

36 The half-life time of all LMWOS in the soil solution varied between 0.6 min (formic acid) and 37 5.0 min (sugars), indicating its dependence on C oxidation state of the substances. The half-life time of ¹⁴C in the fast mineralized pool in microorganisms, ranged between 30 (malic acid) and 80 38 39 (glycine) min and was independent on either C oxidation state, the number of C atoms, or number of 40 -COOH groups. This suggests that intercellular metabolic pathways are more important for LMWOS transformation in soil than their basic chemical properties. The portion of mineralized LMWOS 41 42 increased with their C oxidation state (20 % for sugars vs. 90% for formic acid) corresponding to the 43 decrease of C incorporated into the cytosol and SOC pools.

Concluding, the physicochemical properties of the common LMWOS allow predicting their
 microbial uptake from soil solution and subsequent partitioning of C within microbial biomass.

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- 47

48 Key words: carbon use efficiency, CUE, decomposition kinetics, dissolved organic nitrogen, organic
49 acids.

51 **1. Introduction**

52 Low molecular weight organic substances (LMWOS) in soil originate from a wide range of sources, including root and microbial exudation, animal wastes, canopy throughfall, and the 53 54 decomposition of plant and microbial necromass. Although LMWOS typically represent a small 55 proportion of the total dissolved organic carbon (DOC) pool in soil, they play a critical role in many 56 soil processes, including complexation of metal ions which increases their mobilization (e.g. 57 carboxylic acids), as an important N source (e.g. amino acids) for plants and microorganisms, and as 58 a source of C and energy for microorganisms (e.g. sugars) (Blagodatskaya and Kuzyakov, 2013; 59 Grayston et al., 1997; Hill et al., 2012). From a global perspective, LMWOS contribute significantly 60 to total soil CO₂ flux (up to 30%) (van Hees et al., 2005) and thus represent an important parameter 61 for modeling of soil organic carbon (SOC) dynamics.

62 Although LMWOS may be leached, become sorbed to the solid phase, abiotically mineralized 63 or used by plants, their uptake by the microbial communities dominates their longevity in soil 64 solution and represents the first step of their utilization (Glanville et al., 2016). The uptake of 65 LMWOS from solution depends on their properties, namely broad substrate class (e.g. sugars, 66 phenolics etc), which determines its subsequent use within cell metabolism (Gunina et al., 2014; Apostel et al., 2013), and concentration, which determines the transport systems used by 67 68 microorganisms for taking up LMWOS (Hill et al., 2008). In addition, for amino acids it has been 69 shown that substances with low C oxidation states (e.g. lysine) are taken up by microorganisms 70 slower than ones having higher C oxidation states (e.g. glycine and glutamate) (Jones and Hodge, 71 1999), while the fate of carboxylic acids in soil is dependent on their solubility and association with 72 the soil's solid phase (Gunina et al., 2014). Thus, even if the general substance class plays a major 73 role in the fate of LMWOS in soils, the physico-chemical properties of the individual compound are 74 also highly important.

The second step of LMWOS utilization by microorganisms is their incorporation into
 metabolic cycles and subsequent mineralization to CO₂ or immobilization within cellular components

77 (Apostel et al., 2013). It has also been shown that intercellular metabolism affects the fate of amino 78 and carboxylic acid derived-C in soils (Gunina et al., 2014), as each compound class enters distinct 79 metabolic cycles within the cell. The proportion of each mineralized LMWOS is also linked to the C 80 oxidation state of the substrate. Carboxyl groups (-COOH) (C oxidation state = +3.0) are mineralized 81 to CO_2 at a higher amount than methyl groups (-CH₃) (C oxidation state = -3.0) (Fischer and 82 Kuzyakov, 2010). So, the presence of a high number of reduced C atoms in LMWOS molecules can lead to low mineralization and high LMWOS-C incorporation into structural elements of the cell. At 83 84 the same time, a higher proportion of mineralized C should be observed for substances with high 85 number of oxidized C atoms (e.g. substrates rich in -COOH groups). Additionally, the standard 86 enthalpy of combustion of organic compounds seems to be dependent on substance C oxidation state: 87 for substances with "0" C oxidation state (e.g. glucose, alanine) the values of standard enthalpy of 88 combustion are in the range 1600-2800 kJ/mol, whereas for oxidized substances (C oxidation state +1 89 or +2) the values are lower: 280-1300 kJ/mol. Thus, substance physico-chemical properties can 90 directly impact the utilization processes of LMWOS within the microorganisms. In contrast, further 91 fate of C contained within LMWOS may be closely related to cell metabolite turnover, where this C 92 was incorporated during intercellular metabolisation (Glanville et al., 2016).

93 The aim of the study was to estimate the initial utilization (within 24 h of LMWOS 94 application) of three main LMWOS classes (sugars, carboxylic and amino acids) and to reveal the 95 effect of substance properties on their fate within soil. We hypothesized that: i) LMWOS half-life 96 times in soil solution will depend on substance properties, namely C oxidation state, number of -97 COOH groups and size of the molecules, ii) the half-life of LMWOS-C in microbial biomass pool 98 will depend on the properties of LMWOS and the pathway taken when entering into intercellular 99 metabolic cycles, and iii) substances with a high C oxidation state will be mineralized to a larger 100 extent than substances with a low C oxidation state.

101

103 **2. Materials and methods**

104 2.1. Site description and soil sampling

105 Soil was collected from the BangorDIVERSE long-term forest diversity experiment, located 106 in Abergwyngregyn, North Wales, UK (53°14'16" N, 4°1'1" W) (Smith et al., 2013; Ahmed et al., 107 2016). Within this experiment, soil was collected from the replicated Silver birch (Betula pendula 108 Roth.) plots. The soil is classified as a fine loamy textured Dystric Fluvic Cambisol (WRB, 2006) and 109 has a mixed glacial till parent material. The site has a mean annual soil temperature of 10.6 °C and an 110 annual rainfall of ca. 950 mm. The basic properties of the soil are presented in Table 1 and in Ahmed 111 et al. (2016). At each sampling site, surface litter (ca. 1-2 cm) was removed and the top 10 cm of the 112 mineral soil (A horizon) was collected from four independent locations within each of four replicate 113 plots and combined to make a composite soil sample. Soil samples were stored in gas-permeable 114 plastic bags at 5 °C until extraction of soil solution, which was conducted within 24 h of sample 115 collection. Substrate uptake and mineralization experiments were conducted within one week of soil 116 sample collection.

117

118 2.2. Extraction of soil solution

Soil solution was obtained by centrifugation following the technique of Glanville et al. (2012). Briefly, 100 g of fresh soil was placed into a polypropylene centrifuge tube with a perforated bottom and covered by a fine mesh (pore size 50 μ m). This was attached to a base unit which collects soil solution during centrifugation. This construction was centrifuged at 3500 g for 15 min. The extracted soil solution was subsequently passed through a 0.2 μ m cellulose acetate filter to remove microbial contaminants and stored at -20 °C prior to use in subsequent experiments.

125

126 2.3. LMWOS uptake from soil solution

127 The uptake of LMWOS by the soil microbial community was measured over 24 h for sugars 128 (glucose and fructose), carboxylic acids (malic, succinic and formic acids) and amino acids (alanine 129 and glycine). These substrates were chosen as they are either commonly found in root 130 exudates/lysates or they represents the breakdown products arising from the main organic polymers 131 entering soil (i.e. cellulose/protein). The C oxidation state of each LMWOS was calculated as sum of 132 all C oxidation states divided by the amount of C atoms in the substance (Table 2).

The ¹⁴C radiolabeled substances (<10 nM) were added separately to the extracted soil solution (see section 2.2) to obtain a total ¹⁴C specific activity of 0.83 kBq ml⁻¹ for each compound. No additional non-labeled substances were added so that we did not want to change the intrinsic concentrations of the compounds naturally present in soil solution. All LMWOS were uniformly labeled and ¹⁴C specific activities of the each initial substances were: ¹⁴C-glucose 7.4 MBq ml⁻¹, ¹⁴Cfructose 37 MBq ml⁻¹, ¹⁴C-malic acid 3.7 MBq ml⁻¹, ¹⁴C-succinic acid 3.7 MBq ml⁻¹, ¹⁴C- formic acid 35.6 MBq ml⁻¹, ¹⁴C-alanine 3.7 MBq ml⁻¹, ¹⁴C-glycine 1.8 MBq ml⁻¹.

140 To measure the depletion of the LMWOS from soil solution, fresh field-moist soil (1.2 g) was placed into a 1.5 cm³ polypropylene microcentrifuge tube and 0.3 ml of ¹⁴C-labelled soil solution was 141 142 added to the soil surface. The solution immediately infiltrated into the soil. The microcentrifuge tubes 143 were perforated at the bottom and the holes were covered with a small piece of Whatman GF/A glass 144 fiber filter paper (pore size 1.6 µm). These soil-filled microcentrifuge tubes was then placed on top of 145 another empty microcentrifuge tube and the dual-tube array was centrifuged (14,000 g, 1 min). The 146 soil solution from the upper tube passed to the lower tube where it was recovered for analysis. Soil 147 solution was obtained 1, 4, 8, 10, 20, 30, 60, 240, 960 and 1440 min after addition of the ¹⁴C-labelled solution to the surface of the soil in the upper microcentrifuge tube. ¹⁴C activity of the recovered soil 148 149 solution was measured by liquid scintillation counting (Wallac 1409 scintillation counter, Wallac 150 EG&G Ltd, Milton Keynes, UK) using Wallac Optiphase 3 scintillation cocktail (Wallac EG&G Ltd, 151 Milton Keynes, UK). This procedure was also done with sterile soil (autoclaved, 121°C, 30 min) to 152 determine the importance of abiotic losses of LMWOS from soil solution (i.e. sorption to the solid 153 phase) in the absence of the microbial activity (Hill et al., 2008). Each component of the experiment 154 was replicated four times. The uptake rate of ¹⁴C-labelled LMWOS from soil solution was calculated
155 as follows:

156
$$R = a_1 + a_2 \exp^{-kt}$$

157 where *R* is the percent of applied ¹⁴C remaining in soil solution, a_1 is an 158 asymptote to which ¹⁴C activity fells in single exponential curves, a_2 is an estimated pool size for 159 uptake, *t* is time and *k* is an uptake rate constant. The half-life times of LMWOS in soil solution ($T_{\frac{1}{2}}$ 160 solution) were calculated as $\ln(2)/k$. As the main portion (>80%) of the applied tracer was taken up 161 from soil solution within 60 min, only this period of time is presented, whereas the single first order 162 kinetic equation was fitted to all the data collected over the experimental period (24 h).

163

164 2.4. LMWOS mineralization in soil

165 To estimate the mineralization rate of each LMWOS, a similar procedure to that described above was employed except that we measured the rate of ¹⁴CO₂ evolution from the soil. Briefly, fresh 166 soil (1.2 g) was placed into a 1.5 ml microcentrifuge tube and 0.3 ml of each ¹⁴C-labeled solution 167 168 added (according to procedure described above). The microcentrifuge tubes were placed into a larger 169 50 ml polypropylene container and a 1 M NaOH trap (1 ml) added to capture evolved CO₂ in the 170 closed system. The NaOH traps were changed at 1.5, 3.5, 5.5, 8.5, 13, 22, 24, 25.5 and 27.5 h after LMWOS addition. ¹⁴C activity of the NaOH solutions was measured by liquid scintillation counting 171 172 as described above. To describe mineralization rate of each LMWOS, a double first order kinetic equation was applied to the portion of ${}^{14}C$ remaining in the soil (${}^{14}C_{SOC}$), (calculated as 173 $100^{-14}C_{CO_2}(\%)$): 174

175
$${}^{14}C_{SOC} = a \cdot \exp^{-k_a t} + b \cdot \exp^{-k_b t},$$

where *a* and *b* are pool sizes for the fast and slow mineralization phases, *t* is time and k_a and k_b are the rate constants for the fast and slow mineralization phases (Glanville et al., 2016). The $T_{\frac{1}{2}}$ for LMWOS-C of the fast and slow phases of C mineralization within the microbial community were 179 calculated as $\ln(2)/k_a$ or $\ln(2)/k_b$ and will subsequently be referred to as $T_{\frac{1}{2}-\text{fast}}$ and $T_{\frac{1}{2}-\text{slow}}$ 180 respectively.

At the end of the experiment (27.5 h), ¹⁴C activity was measured in the microbial cytosol pool 181 182 using the chloroform fumigation-extraction procedure of Wu et al. (1990). As no extraction 183 efficiency correction factor was applied to the extracted dissolved organic C pool after fumigation 184 (Glanville et al., 2016), this pool was referred to "cytosol" rather than microbial biomass. The amount of ¹⁴C remaining in the bulk soil at the end was also measured by combusting the soil at 800 °C in a 185 OX400 biological oxidiser (R.J. Harvey Instrument Corp., USA) and ¹⁴CO₂ measured by scintillation 186 counting after capture in Oxosol scintillant (National Diagnostics, Atlanta, GA, USA). To obtain ¹⁴C 187 in SOC pool (further referred to as ¹⁴C-SOC) the ¹⁴C portions in CO₂ and cytosol pools were 188 subtracted from ¹⁴C in bulk soil, and present the pool containing non-extractable microbial biomass 189 190 and microbial metabolites. Tracer incorporation into cytosol and SOC pools was presented as a percent of the total applied 14 C. 191

192 Based on the calculated ¹⁴C incorporation into CO₂ and microbial cytosol pools (for the last

193 measurement point - 27.5 h), the anabolism to catabolism ratio was calculated as:

194
$$\frac{anabolism}{catabolism} = \frac{{}^{14}C_{cytosol}}{{}^{14}C_{CO_2}},$$

which shows the proportion of ¹⁴C used for energy production relative to that incorporated into cell
components.

- 197
- 198 *2.5. Statistics*

Data on ¹⁴C in CO₂, cytosol and SOC as well as pool sizes, rate constants and $T_{\frac{1}{2}}$ were subjected to ANOVA and significant differences between the various LMWOS were tested with LSD post hoc test with *P* < 0.05. Exponential equations were fitted to the experimental results using a least squares iteration routine in Statistica 10.0 (Dell Statistica Inc., Tulsa, OK). The simple regression analysis was performed in Statistica 10.0 (Dell Statistica Inc., Tulsa, OK) with data on C oxidation state, number of C atoms, number of COOH groups vs. LMWOS $T_{\frac{1}{2} \text{ solution}}$, $T_{\frac{1}{2}\text{ - slow}}$, portion of

205 ¹⁴C in SOC, cytosol and CO₂ pools.

206 **3. Results**

207 3.1. Uptake of LMWOS from soil solution

208 The three classes of LMWOS showed a similar uptake pattern from soil solution based on the ¹⁴C depletion from the DOC pool (Fig. 1). Calculated LMWOS-C $T_{\frac{1}{2}-\text{solution}}$ changed in the order: 209 sugars > amino acids > carboxylic acids (Table 2). Glucose and fructose showed a similar $T_{\frac{1}{2}-\text{ solution}}$ 210 (3.8 min), which was 1.5 - 2 times longer than for other the substances. The lowest $T_{\frac{1}{2}$ -solution (<1 min) 211 212 was found for formic acid. Estimates of the total amount of LMWOS ascribed to modelled pool a_2 were similar for all substances (Table 2). There was a negative relationship between the $T_{\frac{1}{2}-\text{solution}}$ of 213 214 each substrate and its C oxidation state (Fig. 2 top panel) and number of -COOH groups 215 (Supplementary material; Fig. S1). Furthermore, there was a positive relationship between the $T^{1/2}$ -216 solution of all LMWOS and the number of C atoms within the individual substrates (Fig. 2, bottom 217 panel). Results for the autoclaved soil (Supplementary material; Fig. S2) showed some dilution with 218 the intrinsic soil solution and that sorption can occur for some substances (e.g. carboxylic acids and, 219 glycine). However, as shown previously (Fischer et al., 2010), biotic uptake of LMWOS out-220 competes the abiotic sorption processes, from which we predict that sorption processes will not 221 greatly influence the results in the non-autoclaved soil.

222

223 3.2. Mineralization of LMWOS in soil

224 Mineralization patterns were similar for all three LMWOS classes, namely the highest portion 225 of C was mineralized in the first 5 h, and later ¹⁴C-CO₂ reached a plateau (Fig. 3). The maximum 226 proportion of mineralized LMWOS was found for carboxylic acids, followed by amino acids and 227 sugars (Fig. 3). Overall, 15 to 80% of the applied LMWOS were decomposed to CO₂ within the first 228 mineralization phase (pool *a*, k_a) depending on substance class (Fig. 3). Constant rates for the first 229 mineralization phase were between 0.5 and 1.3 % h⁻¹ and calculated $T_{\frac{1}{2}-fast}$ values for pool *a* for each LMWOS-C were in the range of 0.52-1.34 h (30-80 min) (Table 3), with the shortest values observed for malic acid and the longest for glycine. The $T_{\frac{1}{2}-\text{fast}}$ values for each LMWOS-C were much longer than those calculated for their loss from soil solution, showing that mineralization does not occur immediately after LMWOS uptake. No significant correlation was found between the $T_{\frac{1}{2}-\text{solution}}$ values of each substrate and it subsequent mineralization during the fast utilization phase (Supplementary materials; Fig. S3).

Constant rates for the second mineralization phase (model pool *b*, k_b ; Table 3), which describes the turnover of substrate-C immobilized in the microbial biomass, were up to 3 orders of magnitude lower than for the first modeled pool (*a*, k_a). Calculated LMWOS-C $T_{\frac{1}{2}-slow}$ ranged between 25 and 290 h, with the shortest values observed for formic acid and the longest for glucose. The $T_{\frac{1}{2}-slow}$ values for each LMWOS showed relationships with C oxidation state and number of C atoms (Supplementary material; Fig. S5).

242 The partitioning of LMWOS-C between CO₂, the microbial cytosol and that remaining in 243 SOC is shown in Figure 4. The maximum proportion of mineralized substances was observed for 244 formic acid, which was followed by malic and succinic acid, amino acids and sugars. In contrast, the ¹⁴C recovered in the cytosol and remaining in SOC followed the opposite trend. The proportion of 245 mineralized LMWOS increased with substance C oxidation state, whereas the amount of ¹⁴C 246 247 incorporated into the cytosol and remaining in SOC (for all substances) followed the opposite trend 248 (Fig. 4, top panel). Additionally, the proportion of LMWOS-C incorporated into the microbial cytosol 249 increased with the number of C atoms present in the molecule and decreased with the number of -250 COOH groups (Fig. 4, bottom panel). Anabolism/catabolism ratio (Fig. 5) was the highest for the 251 sugars (both glucose and fructose) and for alanine, having zero C substance oxidation states. The 252 lowest value was found for formic acid.

253 Overall, initial utilization of LMWOS within the microbial biomass was not dependent on the 254 substance properties. In contrast, the total amount of LMWOS-C which can be utilized (including

255 mineralization to CO_2 and incorporation in to cellular compounds) within the microbial biomass was 256 clearly dependent on the physico-chemical properties of the individual substrates.

257

258 **4. Discussion**

259 In this study, the utilization of LMWOS in soil focused on: i) the initial rate of uptake from soil solution, ii) mineralization to CO₂, and iii) subsequent utilization and partitioning of C within the 260 261 microbial cells. These processes were studied within 24 h, to deduce the initial fate of LMWOS-C 262 rather than the turnover of secondary metabolites within the microbial community or the turnover of 263 the biomass itself (i.e. necromass). The fate and flux of LMWOS was studied at natural concentrations (soil solution was only labeled at trace levels for each ¹⁴C-compound), to best reflect 264 265 conditions which naturally exist in the field. This contrasts with almost all previous studies which 266 have used high substrate addition rates to investigate LMWOS turnover in the soil. Although these 267 former studies may reflect pulse additions of soluble C arising from root lysis or organic waste 268 addition, they misrepresent the much lower concentrations of LMWOS produced by the slower 269 turnover of more recalcitrant (and arguably more important) pools of soil organic matter.

270

271 4.1. Uptake of LMWOS from soil solution

272 We found that up to 90% of the applied LMWOS were taken up from soil solution within the 273 first 10 minutes (Fig. 1). Similar results have been found for glucose applied to soil in the 274 concentration range from 1 µM to 10 mM (Hill et al., 2008). The rapid removal of substrates can be 275 attributed to the rapid uptake of LMWOS by the C-limited soil microbial community, extracellular 276 enzymatic decomposition or sorption on the mineral phases. For most neutral or monovalent 277 LMWOS, microorganisms represent the dominate loss pathway from solution, particularly in 278 comparison to sorption to mineral phases (Fischer et al., 2010). In the case of di- and tri-valent 279 substrates, however, sorption can significantly suppress microbial uptake, especially in soils 280 containing large amounts of Fe and Al oxyhydroxides (Jones and Edwards, 1998), however, it was 281 not the case in our study. We attempted here to estimate the effect of abiotic sorption processes by 282 measuring the loss of LMWOS under sterile (autoclaved) and non-sterile soil. Sorption had low importance in the fate of LMWOS because larger percent of ¹⁴C was removed from soil solution in 283 284 non-sterile soil compare to sterile for the same time interval. This is the consequence of neutral pH 285 values and low contents of Fe and Al in the soil. Overall, our results are consistent with microbial 286 transformation being the dominant process. Although extracellular enzymes may exist in soil solution 287 and could extracellularly cleave our substrates (e.g. deaminases acting on alanine or glycine to 288 produce pyruvate and lactate), we expect this transformation pathway to be insignificant in 289 comparison to the direct uptake by microbial membrane transporters.

290 The fastest uptake rates from solution and subsequent $T_{\frac{1}{2}-\text{solution}}$ values (0.6-1.5 min) were 291 found for carboxylic acids while the slowest $T_{\frac{1}{2}$ -solution value was found for sugars (3.7 min) (Table 2). 292 Although the rate of depletion was very rapid for all substrates, the variation in uptake rate can be 293 attributed to differences in (i) relative diffusion speed of the substrates in solution, (ii) different 294 affinities and expression of the various transport systems within the microbial community, and (iii) 295 rate of intracellular processing of the various substrate classes which may feedback on transporter 296 activity (Hill et al., 2008; Jones and Edwards, 1998). The $T_{\frac{1}{2}-\text{solution}}$ of carboxylic and amino acids 297 decreased with the C oxidation state of substances suggesting that LMWOS with low C oxidation 298 states remain in soil solution longer than ones which are highly oxidized. At the same time, LMWOS 299 $T_{\frac{1}{2}-\text{solution}}$ values increased with the number of C atoms indicating that substances with a lower 300 molecular weight are taken up faster. For substances with a similar C oxidation state (both sugars and 301 alanine), a longer $T_{\frac{1}{2}$ -solution was found for larger molecules although more substrates would need to be 302 tested to confirm this. Overall, even if the substance class is one of the significant parameter 303 determining the fate of LMWOS in soil solution (Gunina et al., 2014), we conclude that the $T_{\frac{1}{2}$ -solution 304 of LMWOS depends also on substance C oxidation state and on molecular size. Further, the very 305 rapid uptake of all LMWOS classes from soil solution suggests that this is not the limiting step of 306 their initial utilization by microorganisms.

308 4.2. Mineralization of LMWOS

309 The $T_{\frac{1}{2}-\text{fast}}$ values, describing the initial transformation of LMWOS-C within the microbial biomass, were 30-80 times higher than the $T_{\frac{1}{2}-\text{solution}}$ values, indicating that mineralization may occur 310 311 more slowly than cellular uptake. However, we added tracer amounts of substrate to extracted soil 312 solution which was then injected to the soil to try and mimic natural C concentrations. Therefore, we 313 would expect the system to be at quasi-steady state (i.e. a stable microbial biomass) and the rate of C 314 influx into soil solution should be equal to the rate of C efflux from the microbial biomass. However, 315 it was not the case in our study and observed slow rate of C efflux and high values of $T_{\frac{1}{2}-\text{fast}}$ could be 316 due to i) dilution of the LMWOS in the labile metabolite pool within the cytosol (Hill et al., 2008), 317 and ii) passage of LMWOS through contrasting metabolic pathways which enter aerobic or anaerobic 318 respiratory cycles at different points. Additionally, natural artifacts such as release of HCO_3^- from the 319 cell, its diffusion through extracellular water films and the subsequent degassing and diffusion of CO₂ 320 through the pore network to the soil surface can effect on the temporal dynamic of captured CO₂ 321 (Boddy et al., 2007). However, due to the small amount of soil, which was used in the present 322 experiment, these artifacts should not strongly affect our results, but would need to be accounted for 323 when working with large undisturbed field samples. This highlights the intrinsic problems associated 324 with sole reliance on quantifying substrate turnover rates based on mineralization data alone, 325 especially for short-lived substrates. It also indicates that previous studies may have vastly 326 underestimated substrate turnover rates (van Hees et al., 2002).

An absence of dependence between LMWOS-C $T_{\frac{1}{2}-\text{fast}}$ and C oxidation state, number of C atoms, or number of -COOH groups of the substances (Supplementary material; Fig. S4) are likely due to incorporation of LMWOS into various metabolic pathways within the microorganisms (Gunina et al., 2014; Apostel et al., 2013; Apostel et al., 2015; Dippold and Kuzyakov, 2013; Dijkstra et al., 2011). So, calculated alanine C $T_{\frac{1}{2}-\text{fast}}$ was 1.5 times faster than glycine (Table 3). This could be explained as alanine enters the citric acid cycle as pyruvate, whereas glycine is metabolized in the

333 cells via three different pathways: i) by glycine cleavage enzyme, ii) by conversion of glycine to 334 pyruvate via serine and iii) by conversion of glycine to glyoxylate by L-amino acid oxidase or L-335 amino acid dehydrogenase (Keseler et al., 2009), thus, glycine-C can be metabolized slower than 336 alanine. In contrast, LMWOS-C $T_{\frac{1}{2}-\text{slow}}$ decreased with an increase in C oxidation state and increased 337 with the amount of C atoms in the LMWOS molecule, showing that more time is needed to oxidize the LMWOS with a low C oxidation state. Thus, the initial mineralization processes of LMWOS by 338 339 soil microorganisms are mainly connected with the point at which compounds enter into metabolic 340 cycles, whereas subsequent utilization of LMWOS-derived C can be affected by properties of the 341 substances.

342

343 *4.3. Partitioning of LMWOS-C between the CO₂, cytosol and SOC pools*

344 The amount of C mineralized followed the order: carboxylic acids > amino acids > sugars. 345 This is in agreement with some previous laboratory and field studies (Jones and Edwards, 1998), but 346 contrasts with others where no differences were observed (Gunina et al., 2014). Such contradictory 347 results are connected with i) various observation periods used during the studies, ii) the amount of 348 time elapsed between LMWOS application and the start of sampling, and iii) various half-life time of 349 cell metabolites, where LMWOS-C was incorporated. Additionally, the total amount of LMWOS 350 applied to the soil can affect the amount of substrate mineralized, especially if the amount added is 351 sufficient to stimulate microbial growth. Typically, when concentrations of LMWOS exceed 10 mM 352 the amount of C incorporated into microbial biomass compartments increases and less C is respired 353 (Hill et al., 2008). In this study, the proportion of substrate-C mineralized increased with the C 354 oxidation state of the substances (Fig. 4, top panel, Fig. 6), showing that oxidized compounds are 355 used preferentially for respiration with less C incorporated into cell metabolites.

The highest portion of ¹⁴C-LMWOS recovered from the cytosol pool was from sugars, suggesting that sugars are the universal compounds for construction of cell components (constituents of the bacterial and fungal cell membranes and cell walls, lipoteichoic and teichoic acids of Gram-

359 positive bacteria, lipopolysaccharides of Gram-negative bacteria, polysaccharides, etc) (Dippold et 360 al., 2014; Gunina and Kuzyakov, 2015; Lengeler et al., 1999). In contrast, the lowest incorporation of ¹⁴C-LMWOS found in the cytosol was from carboxylic acids, with the lowest of that group being 361 362 formic acid (Fig. 4, Supplementary material; Fig. S6). Reported ratios of mineralized-C to 363 immobilized-C for carboxylic acids is 3:2 (Jones and Edwards, 1998). A wider range of mineralized-364 to-immobilized C was reported for formic acid - 19:1 (Herlihy et al., 1987) and our results (Fig. 4, 365 Supplementary material; Fig. S6) are in accordance with these findings. Such high mineralization can 366 be explained by the fact that formic acid is a toxic substance (Herlihy et al., 1987), and thus, even if it is taken up by microorganisms it is mainly decomposed to CO₂ within the cells. The proportion of C 367 368 incorporated into the cytosol decreased with the substance C oxidation state (Fig. 4, top panel), 369 suggesting that more oxidized compounds are mainly used for respiration, whereas reduced 370 compounds are utilized for cell biomass construction. Thus, despite the initial LMWOS 371 mineralization dynamics being independent of substance properties, the final partitioning of the 372 LMWOS-C between mineralized and immobilized pools is dependent on their physiochemical 373 properties.

Anabolism/catabolism ratio (Fig. 5) declined as C oxidation state increased, suggesting that losses for respiration prevail during the assimilation of C from oxidized substances or functional groups (e.g. -COOH). This is directly connected with energy production, which microorganisms can obtain during utilization of LMWOS - with C oxidation state increases, energy content of the LMWOS decreases. Thus, it shows that substrates with high oxidation state are used primarily for energy, whereas substrates with low C oxidation state are primarily used for cell construction and maintenance.

381

382 5. Conclusions

383 Typically, the turnover of individual LMWOS in soil is estimated by measuring the rate of 384 CO₂ appearance after substrate addition to soil (i.e. substrate-induced respiration). However, this

385 approach fails to realistically capture the dynamics of LMWOS in soil. In this study, the uptake of 386 three common classes of LMWOS (sugars, carboxylic and amino acids) from soil solution and their 387 subsequent mineralization by the soil microbial community was studied over a 24 h period. While 388 previous studies have mainly focused on the effect of substance class or concentrations, in the present 389 study the main focus was on the physico-chemical properties of substances, including substance C 390 oxidation state, number of -COOH groups and C atoms. We combined the use of substrates at natural 391 abundance with repeated measurements over short time scales. This allowed us to estimate actual 392 rates of LMWOS loss from solution rather than the processing of C once it had already been 393 incorporated into cell metabolites.

394 The half-life of the LMWOS in soil solution ranged from 0.5 to 3.8 min, with the shortest for 395 carboxylic acids and the longest for sugars. Thus, the extremely fast microbial uptake of all LMWOS 396 classes from solution suggests that this is not a rate-limiting step in the utilization of LMWOS by the 397 microbial community. The $T_{\frac{1}{2}}$ of the LMWOS in solution decreased with C oxidation state. In 398 contrast, the $T_{\frac{1}{2}}$ of LMWOS in soil solution increased with the number of C atoms showing that 399 larger molecules persist longer, possibly due to their slower rate of diffusion in soil. Our data 400 suggests that the uptake of common LMWOS from soil solution by microorganisms may be possible 401 to predict from the physio-chemical properties of the substance.

The LMWOS-C $T_{\frac{1}{2}-\text{fast}}$ values ranged between 30 and 80 min and was lowest for amino acids and highest for carboxylic acids. Large differences between LMWOS $T_{\frac{1}{2}}$ values in solution and in soil shows that microbial uptake and subsequent mineralization of LMWOS are temporarily decoupled. The $T_{\frac{1}{2}-\text{fast}}$ of LMWOS-C in soil was not dependent on the properties of the substance, from which we infer that intercellular metabolism is the main factor determining initial mineralization of C derived from LMWOS.

408 The total proportion of C mineralized from each LMWOS increased with the substance's C 409 oxidation state, suggesting that oxidized compounds are mineralized to a greater degree than more 410 reduced compounds. To support this observation, the LMWOS-C $T_{\frac{1}{2}-\text{slow}}$ decreased with C oxidation

411 state increase. The portion of LMWOS-C incorporated into the cytosol and remaining in SOC 412 decreased with each substance's C oxidation state. Thus, substance properties directly affected the 413 final partitioning of LMWOS-C between mineralized and microbially utilized pools. The 414 anabolism/catabolism ratio decreased with compound C oxidation state, showing that more oxidized 415 substances are mainly mineralized, whereas less oxidized LMWOS are primarily used by 416 microorganisms for cell construction and maintenance.

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Table 1 Selected soil properties.
Table 2 Single first order kinetic coefficients describing the depletion of individual carbon substrates
 from soil solution over time. Table 3 Double first order kinetic coefficients describing the depletion of individual carbon substrates from soil over time. Figure 1. Temporal dynamics of ¹⁴C-labelled sugar, organic acid and amino acid disappearance from soil solution. Values represent means \pm SE (n = 4). Lines are the following: blue: solid - glucose, dotted - fructose; green: solid - formic acid, dashed - malic acid, dotted - succinic acid; brown: solid -glycine, dashed - alanine. Figure 2. Relationship between the half-life (min) of different LMWOS in soil solution and their C oxidation state (top panel) and number of C atoms in the molecule (bottom panel). Values represent means \pm SE (n = 4). The error bars for the half-life times of LMWOS in DOC are smaller than size of icon symbols. Figure 3. Cumulative ¹⁴C-CO₂ production from mineralization of ¹⁴C-labelled substances in soil. Values represent means \pm SE (n = 4). **Figure 4.** Relationship between ¹⁴C remaining in the cytosol, SOC and CO₂ pools and C oxidation state (top panel) and ¹⁴C remaining in the cytosol and number of C atoms and -COOH groups (bottom panel) in different LMWOS. Values represent means \pm SE (n = 4). P-values for the regression lines on the top panel figure are less than 0.002; *p*-values for the regression lines on the bottom panel figure are less than 0.004. The substance names are shown only once. Figure 5. Relationship between ${}^{14}C$ incorporated into cytosol (anabolism)/ ${}^{14}C$ incorporated into CO₂ (catabolism) and C oxidation state at the end of LMWOS mineralization experiment. Figure 6. Schematic representation showing the dependence of microbial uptake rate (red), utilization (green) and mineralization efficiency (black) of three distinct classes of LMWOS as a function of substrate C oxidation state

Tables and Figures captions