Solid Earth, 9, 341–372, 2018
https://doi.org/10.5194/se-9-341-2018

Middle to Late Devonian–Carboniferous collapse basins on the Finnmark Platform and in the southwesternmost Nordkapp basin, SW Barents Sea

Jean-Baptiste P. Koehl[1,2], **Steffen G. Bergh**[1,2], **Tormod Henningsen**[1], **and Jan Inge Faleide**[2,3]

[1]Department of Geosciences, UiT The Arctic University of Norway in Tromsø, 9037 Tromsø, Norway
[2]Research Centre for Arctic Petroleum Exploration (ARCEx), UiT The Arctic University of Norway in Tromsø, 9037 Tromsø, Norway
[3]Department of Geosciences, University of Oslo, P.O. Box 1047 Blindern, 0316 Oslo, Norway

Correspondence: Jean-Baptiste P. Koehl (jean-baptiste.koehl@uit.no)

Received: 3 November 2017 – Discussion started: 7 November 2017
Revised: 31 January 2018 – Accepted: 9 February 2018 – Published: 28 March 2018

Abstract. The SW Barents Sea margin experienced a pulse of extensional deformation in the Middle–Late Devonian through the Carboniferous, after the Caledonian Orogeny terminated. These events marked the initial stages of formation of major offshore basins such as the Hammerfest and Nordkapp basins. We mapped and analyzed three major fault complexes, (i) the Måsøy Fault Complex, (ii) the Rolvsøya fault, and (iii) the Troms–Finnmark Fault Complex. We discuss the formation of the Måsøy Fault Complex as a possible extensional splay of an overall NE–SW-trending, NW-dipping, basement-seated Caledonian shear zone, the Sørøya-Ingøya shear zone, which was partly inverted during the collapse of the Caledonides and accommodated top–NW normal displacement in Middle to Late Devonian–Carboniferous times. The Troms–Finnmark Fault Complex displays a zigzag-shaped pattern of NNE–SSW- and ENE–WSW-trending extensional faults before it terminates to the north as a WNW–ESE-trending, NE-dipping normal fault that separates the southwesternmost Nordkapp basin in the northeast from the western Finnmark Platform and the Gjesvær Low in the southwest. The WNW–ESE-trending, margin-oblique segment of the Troms–Finnmark Fault Complex is considered to represent the offshore prolongation of a major Neoproterozoic fault complex, the Trollfjorden–Komagelva Fault Zone, which is made of WNW–ESE-trending, subvertical faults that crop out on the island of Magerøya in NW Finnmark. Our results suggest that the Trollfjorden–Komagelva Fault Zone dies out to the northwest before reaching the western Finnmark Platform. We propose an alternative model for the origin of the WNW–ESE-trending segment of the Troms–Finnmark Fault Complex as a possible hard-linked, accommodation cross fault that developed along the Sørøy–Ingøya shear zone. This brittle fault decoupled the western Finnmark Platform from the southwesternmost Nordkapp basin and merged with the Måsøy Fault Complex in Carboniferous times. Seismic data over the Gjesvær Low and southwesternmost Nordkapp basin show that the low-gravity anomaly observed in these areas may result from the presence of Middle to Upper Devonian sedimentary units resembling those in Middle Devonian, spoon-shaped, late- to post-orogenic collapse basins in western and mid-Norway. We propose a model for the formation of the southwesternmost Nordkapp basin and its counterpart Devonian basin in the Gjesvær Low by exhumation of narrow, ENE–WSW- to NE–SW-trending basement ridges along a bowed portion of the Sørøya-Ingøya shear zone in the Middle to Late Devonian–early Carboniferous. Exhumation may have involved part of a large-scale metamorphic core complex that potentially included the Lofoten Ridge, the West Troms Basement Complex and the Norsel High. Finally, we argue that the Sørøya-Ingøya shear zone truncated and decapitated the Trollfjorden–Komagelva Fault Zone during the Caledonian Orogeny and that the western continuation of the Trollfjorden–Komagelva Fault Zone was mostly eroded and potentially partly preserved in basement highs in the SW Barents Sea.

1 Introduction

The SW Barents Sea margin is located near the Iapetus suture zone that formed when Laurentia collided with Fennoscandia to produce the Caledonian Orogeny (Ramberg et al., 2008; Gernigon et al., 2014). This suture and possibly related deep-seated shear zones, which accommodated, for example, thrust nappe emplacement during the Caledonian Orogeny, are now covered by late Paleozoic to Cenozoic sedimentary basins that formed during multiple episodes of extension. These repeated extension events led to the breakup of the North Atlantic Ocean and formation of a transform plate margin at the boundary between the mid-Norwegian and SW Barents Sea margins (Faleide et al., 1993, 2008; Blystad et al., 1995; Doré et al., 1997; Bergh et al., 2007; Hansen et al., 2012; Gernigon et al., 2014). The rift margin along the SW Barents Sea, offshore western Troms and NW Finnmark (Fig. 1), consists of the Finnmark Platform and an adjacent, glacial-sediment-free strandflat and of deep offshore basins such as the Hammerfest and Nordkapp basins (Gabrielsen et al., 1990). These basins are bounded by major NE–SW-trending extensional faults such as the Troms–Finnmark Fault Complex (TFFC; Gabrielsen et al., 1990; Smelror et al., 2009; Indrevær et al., 2013), the Måsøy Fault Complex (MFC; Gabrielsen et al., 1990; Gudlaugsson et al., 1998), and potential basement-seated ductile detachments (Fig. 1). The study area also includes a deep Paleozoic basin that is located southwest of the Nordkapp Basin and east of the Hammerfest Basin and which is bounded to the southwest by the WNW–ESE-trending segment of the TFFC and to the southeast by the MFC (Fig. 1).

The SW Barents Sea margin off western Troms and NW Finnmark is segmented by margin-oblique, NNW–SSE- to WNW–ESE-trending transfer fault zones, e.g., Senja Fracture Zone and Fugløya transfer zone (Indrevær et al., 2013), which may represent analogs of the onshore, Neoproterozoic, WNW–ESE-trending Trollfjorden–Komagelva Fault Zone (TKFZ) in eastern Finnmark (Siedlecki, 1980; Herrevold et al., 2009) and to the Kokelv Fault on the Porsanger Peninsula (Fig. 1; Gayer et al., 1985; Lippard and Roberts, 1987; Rice, 2013). The TKFZ is believed to continue farther west, off the coast, where it is thought to interact with and merge into the WNW–ESE-trending segment of the TFFC (Gabrielsen, 1984; Vorren et al., 1986; Townsend, 1987b; Gabrielsen and Færseth, 1989; Gabrielsen et al., 1990; Roberts et al., 2011; Bergø, 2016; Lea, 2016). Onshore and nearshore, margin-parallel fault complexes include the Langfjorden–Vargsundet fault (LVF; Fig. 1) trending NE–SW and possibly representing an analog to the TFFC and MFC. The geometric interaction, timing, and controlling effects of the TFFC, MFC, TKFZ, LVF, and adjacent offshore basins and ridges are not yet resolved. In particular, the presence of potential Caledonian structures in the deeper portion of the Finnmark Platform, e.g., in the footwall of the TFFC (see Johansen et al., 1994; Gudlaugsson et al., 1998), is further explored in the present contribution.

The goal of this paper is to contribute to the understanding of tectonic and sedimentary processes in the Arctic in the Late Devonian–Carboniferous. To achieve this, we demonstrate the presence of an overall NE–SW-trending, NW-dipping, basement-seated, low-angle shear zone on the Finnmark Platform, the Sørøya-Ingøya shear zone (SISZ; Fig. 1), and to discuss its role played in shaping the SW Barents Sea margin during late- to post-orogenic collapse of the Caledonides in late Paleozoic times and its influence on the formation and evolution of Devonian–Carboniferous collapse basins. We mapped and analyzed basin-bounding brittle faults on the Finnmark Platform and in the southwesternmost Nordkapp basin (named the easternmost Hammerfest basin in Omosanya et al., 2015), such as the TFFC and the MFC (Fig. 1), to evaluate the impact of the SISZ on post-Caledonian brittle faults. We aim at showing the importance of structural inheritance by examining the relationship among Precambrian–Caledonian structural grains, post-Caledonian fault trends, and offshore sedimentary basin geometries. Minor Carboniferous grabens and half grabens on the Finnmark Platform (e.g., the Sørvær Basin; Fig. 1), which are thought to have formed during early stages of extension shortly after the end of the Caledonian Orogeny (Lippard and Roberts, 1987; Olesen et al., 1990; Johansen et al., 1994; Bugge et al., 1995; Gudlaugsson et al., 1998; Roberts et al., 2011), are of particular importance to the present work. We further investigate the presence of possible Devonian sedimentary deposits on the Finnmark Platform and in the southwesternmost Nordkapp basin and tentatively interpret them as potential analogs to Middle Devonian basins in western Norway (Séranne et al., 1989; Chauvet and Séranne, 1994; Osmundsen and Andresen, 2001) and mid-Norway (Braathen et al., 2000). In this context, NE–SW- to ENE–WSW-trending basement ridges in the footwall of the TFFC and on the northern flank of the southwesternmost Nordkapp basin are described and analyzed, and we compare them to adjacent basement highs such as the Norsel High (Fig. 1; Gabrielsen et al., 1990; Gudlaugsson et al., 1998), the West Troms Basement Complex (Zwaan, 1995; Bergh et al., 2010), and the Lofoten Ridge (Blystad et al., 1995; Bergh et al., 2007; Hansen et al., 2012). Finally, we propose a model of exhumation of these ENE–WSW- to NE–SW-trending basement ridges as a metamorphic core complex (see Lister and Davis, 1989) using shear zones in Lofoten–Vesterålen as onshore analogs for the SISZ (Steltenpohl et al., 2004; Osmundsen et al., 2005; Steltenpohl et al., 2011).

2 Geological setting

The bedrock geology of the SW Barents Sea margin (Fig. 1) consists of (i) an Archean to Paleoproterozoic basement suite, the West Troms Basement Complex (Zwaan, 1995;

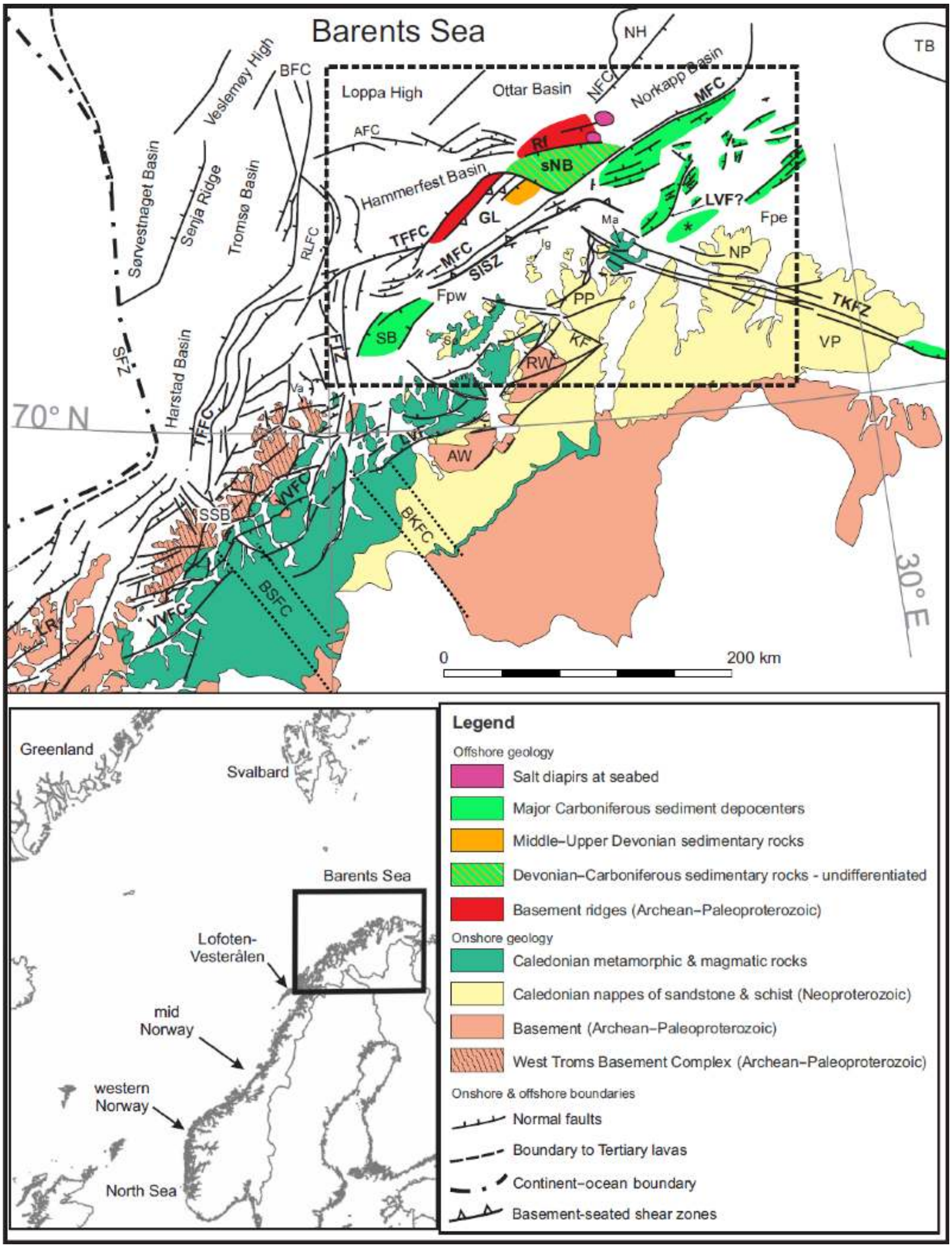

Figure 1. Regional structural map of the SW Barents Sea margin (based on Bergh et al., 2007; Faleide et al., 2008; Hansen et al., 2012; and Indrevær et al., 2013; and Koehl et al., 2018). The onshore geology is from the NGU and Ramberg et al. (2008). Dashed black frame locates Fig. 2. The black star marks the location of the speculated half-graben structure described in Bugge et al. (1995), which we reinterpret as a prograding sedimentary system unconformably resting on basement rocks. Location of the Barents Sea shown as a black frame in lower left inset map. Abbreviations are as follows: AFC: Asterias Fault Complex; AW: Alta–Kvænangen tectonic window; BFC: Bjørnøyrenna Fault Complex; BSFC: Bothnian–Senja Fault Complex; BKFC: Bothnian–Kvænangen Fault Complex; FPe: eastern Finnmark Platform; FPw: western Finnmark Platform; FTZ: Fugløya transfer zone; GL: Gjesvær Low; Ig: Ingøya; KF: Kokelv Fault; LR: Lofoten Ridge; LVF: Langfjorden–Vargsundet fault; Ma: Magerøya; MFC: Måsøy Fault Complex; NFC: Nysleppen Fault Complex; NH: Norsel High; NP: Nordkinn Peninsula; PP: Porsanger Peninsula; Rf: Rolvsøya fault; RLFC: Ringvassøya–Loppa Fault Complex; RW: Repparfjord–Komagfjord tectonic window; SB: Sørvær Basin; SFZ: Senja Fracture Zone; SISZ: Sørøya–Ingøya shear zone; sNB: southwesternmost Nordkapp basin; SSB: Senja Shear Belt; Sø: Sørøya; TB: Tiddlybanken Basin; TFFC: Troms–Finnmark Fault Complex; TKFZ: Trollfjorden–Komagelva Fault Zone; Va: Vannøya; VP: Varanger Peninsula; VVFC: Vestfjorden–Vanna fault complex.

Bergh et al., 2010), (ii) locally preserved autochthonous Neoproterozoic cover sequences (Kirkland et al., 2008), (iii) a series of Caledonian thrust nappes (Andersen, 1981; Ramsay et al., 1985; Corfu et al., 2014), and (iv) late Paleozoic to Cenozoic sedimentary units offshore (Faleide et al., 1993, 2008; Gudlaugsson et al., 1998; Worsley, 2008; Smelror et al., 2009; Fig. 1). Archean to Paleoproterozoic basement rocks are mostly exposed in major horsts and ridges in western Troms (Bergh et al., 2010; Indrevær et al., 2013; Indrevær and Bergh, 2014), whereas Neoproterozoic and Caledonian rocks dominate in the eastern part of Troms and in NW Finnmark (Kirkland et al., 2008; Corfu et al., 2014; Indrevær and Bergh, 2014; Fig. 1). In offshore areas adjacent to western Troms and NW Finnmark, extensive post-Caledonian normal faulting led to the formation of large sedimentary basins that are filled with thick, late Paleozoic to Cenozoic deposits related to the post-orogenic collapse of the Caledonides and to the opening of the NE Atlantic Ocean (Faleide et al., 1993, 2008; Gudlaugsson et al., 1998; Worsley, 2008; Smelror et al., 2009). Late Paleozoic–Cenozoic sedimentary units are missing in onshore areas of Troms and Finnmark likely due to erosion and/or nondeposition (Ramberg et al., 2008; Smelror et al., 2009).

2.1 Onshore Precambrian and Caledonian geology

2.1.1 Precambrian basement rocks

The western Troms margin is characterized by Archean to Paleoproterozoic basement rocks of the West Troms Basement Complex (Bergh et al., 2010) that are preserved and exposed in a horst block formed during post-Caledonian extension (Indrevær et al., 2013). The West Troms Basement Complex consists of tonalitic, trondhjemitic, and granitic gneisses; metasupracrustal rocks; and mafic and felsic igneous rocks (Corfu et al., 2003; Bergh et al., 2010). These rocks were deformed during the Svecofennian orogeny, which resulted in the formation of NW–SE-trending steep foliation, ductile shear zones, and upright and vertical macrofolds, which were only weakly reworked during the Caledonian Orogeny (Corfu et al., 2003; Bergh et al., 2010).

In NW Finnmark, Paleoproterozoic basement rocks occur in several tectonic windows of the Caledonides, e.g., Repparfjord-Komagfjord and Alta-Kvænangen tectonic windows (Zwaan and Gautier, 1980; Pharaoh et al., 1982, 1983; Bergh and Torske, 1988; Fig. 1), and consist of low-grade supracrustal metavolcanics and metasedimentary rocks of the Raipas Group. These Greenstone belts formed as NW–SE-trending rift basins in the Paleoproterozoic during the opening of the Kola Ocean (Bergh and Torske, 1986, 1988), although more recent studies tentatively reinterpret these rocks as foreland basin deposits derived from the Svecokarelian Orogeny (Torske and Bergh, 2004). A thin cover of Neoproterozoic to Cambrian (para-)autochthonous metasedimentary rocks occurs on top of Paleoproterozoic basement rocks in Finnmark (Siedlecki, 1980; Ramsay et al., 1985; Andresen et al., 2014; Corfu et al., 2014). Other Neoproterozoic–Ordovician units in eastern Finnmark include metasedimentary rocks of the Barents Sea and Tanafjorden–Varangerfjorden regions (Siedlecki, 1980; Siedlecka and Roberts, 1992), which are exposed on the Varanger Peninsula (Fig. 1).

The Timanian Orogeny produced major NW–SE-trending folds (Roberts and Siedlecka, 2002) and WNW–ESE-trending fault complexes like the TKFZ (Jonhson et al., 1978; Herrevold et al., 2009). The TKFZ was mapped as a narrow, single-segment fault strand all the way along the Kola Peninsula in Russia in the east, where it merges with the Sredni-Rybachi Fault Zone (Roberts et al., 1997, 2011), to the Barents shelf in the west (Gabrielsen, 1984; Gabrielsen and Færseth, 1989; Gabrielsen et al., 1990; Roberts et al., 2011). We present an alternative model in which the TKFZ splays into multiple fault segments and dies out between the Varanger Peninsula and the Barents shelf. On the Varanger Peninsula, the TKFZ is well displayed on satellite images, but is generally poorly exposed. In map view, the TKFZ is irregular, with different structural segments and branching subsidiary faults both across and along strike, locally showing duplex structures (Siedlecka and Siedlecki, 1967; Siedlecka, 1975). The TKFZ formed along the southwestern boundary of the Timanian Orogeny in the late Cryogenian–Ediacaran (Roberts and Siedlecka, 2002; Siedlecka et al., 2004) and was later reactivated as a strike-slip fault during the Caledonian Orogeny when it accommodated significant lateral displacement constrained to 200–250 km of dextral strike-slip movement (Bylund, 1994; Rice, 2013).

2.1.2 Caledonian nappes

Coastal areas of NW Finnmark are dominated by Caledonian thrust sheets of the Kalak Nappe Complex and Magerøy Nappe (Ramsay et al., 1985; Ramberg et al., 2008; Corfu et al., 2014), formed in the Neoproterozoic through Silurian (Fig. 1). The Kalak Nappe Complex is composed of amphibolite facies schists, metapsammites, and paragneisses and comprises several allochthonous thrust sheets with Proterozoic basement rocks, clastic metasedimentary rocks, and plutonic rocks of the Seiland Igneous Province (Corfu et al., 2014). A major thrust defines the contact with the underlying pre-Caledonian basement (Ramsey et al., 1985). Dominant structures include a gently NW-dipping foliation; NNE–SSW-trending, east-verging, asymmetrical recumbent folds; and low-angle thrusts that accommodated top–ESE shortening (Townsend, 1987a; Kirkland et al., 2005). The Kalak Nappe Complex was previously considered to represent an exotic terrane accreted on the Laurentian margin of Rodinia prior to the rifting of the Iapetus Ocean and to have later been thrusted over Baltica during the Caledonian Orogeny (Kirkland et al., 2008). However, paleocurrent and geochronological data suggest these rocks to be of Baltican origin (Roberts, 2007; Zhang et al., 2016).

The Seiland Igneous Province corresponds to a large, late Neoproterozoic mafic and ultramafic intrusion linked to the early–mid rifting stages of the Iapetus Ocean (Elvevold et al., 1994; Corfu et al., 2014). Recent geophysical studies by Pastore et al. (2016) show that the base of the Seiland Igneous Province defines two deep-reaching roots located below the islands of Seiland and Sørøya constraining the thickness of the Kalak Nappe Complex in this area to a maximum of 10 km. On the Porsanger and Varanger peninsulas, ENE–WSW- to NNE–SSW-trending Ediacaran metadolerite dyke swarms are particularly common, and they are associated with the rifting of the Iapetus Ocean as well (see Roberts, 1972; Siedlecka et al., 2004; Nasuti et al., 2015).

The Kalak Nappe Complex is structurally overlain by the Magerøy Nappe, which consists of Late Ordovician to early Silurian greenschist facies metasedimentary and metaplutonic rocks (Andersen, 1981, 1984; Corfu et al., 2014) that crop out on the island of Magerøya (Fig. 1). The Magerøy Nappe is characterized by asymmetrical NNE–SSW-trending, east-verging folds and low-angle, NW- and SE-dipping thrusts similar in trend to those observed within the Kalak Nappe Complex (Andersen, 1981) and is intruded by granitic and gabbroic plutons, e.g., the Silurian Honningsvåg Igneous Complex (Corfu et al., 2006) and the Finnvik Granite (Andersen, 1981). Remnants of the Magerøy Nappe thrust units are also found in northeastern Sørøya and on the Porsanger Peninsula (Kirkland et al., 2005, 2007; Corfu et al., 2014; Fig. 1).

2.2 Post-Caledonian brittle faults and basins

2.2.1 Post-Caledonian offshore basins

The SW Barents Sea margin was subjected to multiple episodes of extensional faulting after the end of the Caledonian Orogeny, starting with the collapse of the Caledonides in the Middle to Late Devonian–early Carboniferous, lasting until the early–mid Permian, although evidence of this stage is only preserved onshore in western and mid-Norway (Séranne et al., 1989; Chauvet and Séranne, 1994; Braathen et al., 2000; Osmundsen and Andresen, 2001). During this period, basement ridges in Lofoten–Vesterålen (Klein and Steltenpohl, 1999; Klein et al., 1999; Steltenpohl et al., 2004, 2011; Fig. 1) and in mid-Norway (Osmundsen et al., 2005; Fig. 1) were exhumed as metamorphic core complexes, synchronous with the development of large half-graben basins such as the Vøring and Møre basins in mid-Norway (Blystad et al., 1995) and the Hammerfest, Nordkapp, and Ottar basins in the SW Barents Sea (Gabrielsen et al., 1990; Breivik et al., 1995; Gudlaugsson et al., 1998; Indrevær et al., 2013; Fig. 1). The main rifting events occurred in the Late Jurassic and peaked in the Early Cretaceous, when major offshore basins such as the Tromsø and Harstad basins formed. The rifting ended with full breakup of the North Atlantic Ocean and formation of a transform plate margin in the SW Barents Sea at the Paleocene–Eocene transition (Faleide et al., 1993, 2008).

Off the coasts of western Troms and NW Finnmark, the SW Barents Sea margin is characterized by a relatively shallow area, the Finnmark Platform (Gabrielsen et al., 1990; Fig. 1), which is thought to have remained relatively stable since late Paleozoic times. For example, the inner part of the Finnmark Platform, here referred to as the eastern Finnmark Platform (Fig. 1), was only affected by the formation of minor Carboniferous, ENE–WSW- to NE–SW-trending half-graben and graben structures (Bugge et al., 1995; Samuelsberg et al., 2003; Rafaelsen et al., 2008; Fig. 1). In the hanging wall of the MFC, the western part of the Finnmark Platform (Fig. 1) shows a prominent gravity low, the Gjesvær Low, which was ascribed to the presence of low-density Caledonian rocks (Johansen et al., 1994; Gernigon et al., 2014). We explore and argue for an alternative explanation, i.e., the presence of Devonian collapse basin deposits draped against a low-angle extensional detachment of the SISZ, similar to the Nordfjord-Sogn Detachment Zone, a late-orogenic shear zone that bounds the Middle Devonian Hornelen, Kvamshesten, and Solund sedimentary basins onshore in western Norway (Séranne et al., 1989; Chauvet and Séranne, 1994; Wilks and Cuthbert, 1994; Osmundsen and Andersen, 2001). Ductile detachment surfaces of comparable size, showing analog kinematics and timing of activity contemporaneous with the Nordfjord-Sogn Detachment Zone are documented as far north as the Lofoten–Vesterålen Margin (Klein and Steltenpohl, 1999; Klein et al., 1999; Steltenpohl et al., 2004, 2011), but Devonian collapse basin sedimentary rocks and extensional detachments have not yet been reported along the margins of western Troms and NW Finnmark.

2.2.2 Post-Caledonian faults

Multiple studies have reported post-Caledonian brittle faults in onshore coastal areas in Lofoten–Vesterålen, western Troms, and NW Finnmark (Roberts, 1971; Worthing, 1984; Lippard and Roberts, 1987; Townsend, 1987a; Rykkelid, 1992; Lippard and Prestvik, 1997; Roberts and Lippard, 2005; Bergh et al., 2007; Hansen et al., 2012; Indrevær et al., 2013; Davids et al., 2013). A common feature is the presence of rhombic, zigzag-shaped fault trends similar in geometry to offshore basin-bounding faults. Dominant fault–fracture trends of the margin strike NNE–SSW, ENE–WSW, and NW–SE (Bergh et al., 2007; Eig, 2008; Eig and Bergh, 2011; Hansen et al., 2012; Hansen and Bergh, 2012; Indrevær et al., 2013). Typical examples are basin-bounding, NNE–SSW- and ENE–WSW-trending brittle normal faults that are part of the Vestfjorden-Vanna Fault Complex, which bounds the offshore Vestfjorden Basin southeast of the Lofoten islands and which can be traced northward to western

Troms (Indrevær et al., 2013; Fig. 1), whereas the NNW–SSE to WNW–ESE trend typically reflects margin-oblique, transform fault trends (Faleide et al., 2008). An analog to the onshore Vestfjorden–Vanna Fault Complex in NW Finnmark is the Langfjorden–Vargsundet fault (Fig. 1), described by Zwaan and Roberts (1978) and Worthing (1984) as a major NE–SW-trending, NW-dipping normal fault juxtaposing rocks from the Kalak Nappe Complex and the Seiland Igneous Province in the northwest against Precambrian basement rocks of the Repparfjord-Komagfjord and Alta-Kvænangen tectonic windows in the southeast (Fig. 1).

The NW Finnmark margin is located along the northeastward prolongation of the Lofoten–Vesterålen and western Troms segments of the Norwegian continental shelf (Fig. 1). Similar fault sets and trends as in Lofoten–Vesterålen exist in Finnmark and their interaction is thought to partly have controlled the rhombic geometry of many offshore sedimentary basins (Bergh et al., 2007; Indrevær et al., 2013). A typical example along the western Troms and NW Finnmark margins is the NW-dipping TFFC, which bounds the Harstad Basin to the east and the Hammerfest Basin to the southeast (Gabrielsen et al., 1990; Indrevær et al., 2013). The TFFC defines a system of irregular branching faults trending NNE–SSW and ENE–WSW and terminating as a WNW–ESE-trending fault zone northwest of the island of Magerøya where it merges with the NE–SW-trending, NW-dipping MFC at the southeastern boundary of the Nordkapp Basin (Gabrielsen et al., 1990) and of the triangular-shaped southwesternmost Nordkapp basin (Omosanya et al., 2015; Fig. 1). We address a possible genetic relationship and structural inheritance of the post-Caledonian MFC with the Caledonian SISZ and argue that the MFC may have initiated as an extensional splay during the reactivation of the SISZ as an extensional detachment during the late- to post-orogenic collapse of the Caledonides. Furthermore, we tentatively link basement ridges such as the Norsel High in the footwall of the Nysleppen Fault Complex (Gabrielsen et al., 1990) to bowed segments of the SISZ (Fig. 1).

2.2.3 Post-Caledonian transfer zones

The Norwegian continental shelf is segmented by transfer fault zones of which the largest is the offshore De Geer Zone (Faleide et al., 1984, 2008; Cianfarra and Salvini, 2015), the main fault segment of which is the Hornsund Fault Zone, an offshore NNW–SSE-trending fault that runs parallel to the west coast of Spitsbergen and separates the SW Barents Sea margin from the Lofoten–Vesterålen Margin (Fig. 1). In the south, the De Geer Zone proceeds through the Senja Fracture Zone and into the Senja Shear Belt on the shore of the island of Senja (Fig. 1). Olesen et al. (1993, 1997) suggested shifts of polarity of the Vestfjorden-Vanna Fault Complex along the Senja Fracture Zone, and they argued that the formation of the Senja Fracture Zone offshore was controlled by a major onshore basement weakness zone, the Bothnian-Senja Fault Complex (Fig. 1), which provided suitably oriented basement heterogeneities for the development of a transfer zone (e.g., Doré et al., 1997). Similarly, Indrevær et al. (2013) proposed the existence of a fault array termed the Fugløya transfer zone to explain offsets and shifts of polarity along the Vestfjorden-Vanna Fault Complex farther northeast in western Troms (Fig. 1). The Fugløya transfer zone trends N–S to NNW–SSE and continues on the shore of western Troms, where it merges with the NW–SE-trending Bothnian-Kvænangen Fault Complex, and offshore where it is thought to merge into the TFFC and the Ringvassøy-Loppa Fault Complex (Indrevær et al., 2013; Fig. 1).

Analogously in NW Finnmark, the WNW–ESE-trending TKFZ seems to merge into a basin-bounding fault, in this case the WNW–ESE-trending, NE-dipping fault segment of the TFFC (Gabrielsen, 1984; Gabrielsen and Færseth, 1989; Roberts et al., 2011). In nearshore areas of NW Finnmark, the TKFZ is thought to proceed offshore and seems to correlate with a large escarpment north of Magerøya and into the Barents Sea (Vorren et al., 1986; Townsend, 1987b). In the area where it terminates, it merges and links up with the TFFC to form triangular-shaped mini-basins (Gabrielsen, 1984; Gabrielsen and Færseth, 1989; Roberts et al., 2011). We explore an alternative origin for the WNW–ESE-trending fault segment of the TFFC and further examine its interaction with the onshore–nearshore TKFZ, which potentially acted as a transfer fault after the Caledonian Orogeny and contributed to offset the LVF near Magerøya and adjacent coastal areas (Koehl et al., 2018; Fig. 1). Other major WNW–ESE-trending faults exist offshore, northeast of the Varanger Peninsula, and these bound the Tiddlybanken Basin, a large WNW–ESE-trending basin that formed in Carboniferous times (Mattingsdal et al., 2015; Fig. 1).

2.2.4 Absolute age dating of post-Caledonian faulting

The absolute age of post-Caledonian brittle faults in NW Finnmark is poorly constrained, although a few contributions provide valid insights (Lippard and Prestvik, 1997; Davids et al., 2013; Torgersen et al., 2014; Koehl et al., 2016). Torgersen et al. (2014) performed K–Ar dating of brittle fault gouge in the footwall of the LVF and obtained dominantly Carboniferous to early Permian ages, as well as a subsidiary Early Cretaceous age for one of the faults. Roberts et al. (1991) and Lippard and Prestvik (1997) presented indirect evidence of early Carboniferous dolerite dykes emplaced along and sealing WNW–ESE-trending brittle fault segments of the TKFZ on Magerøya, thus providing a minimum estimate for the latest stage of faulting along this fault. These dykes produce high positive aeromagnetic anomalies (Nasuti et al., 2015) and may be used to further identify brittle faults in NW Finnmark. Late Devonian dolerite dykes emplaced along brittle faults that trend NE–SW and N–S have been identified and dated on the eastern Varanger Peninsula (Guise and Roberts, 2002) and on the Kola Peninsula (Roberts and

Onstott, 1995). By comparison, Davids et al. (2013) obtained Late Devonian–early Carboniferous ages from K–Ar dating of illite clay minerals for early extensional faulting along the Vestfjorden-Vanna Fault Complex and related faults in Lofoten–Vesterålen and western Troms.

2.3 Offshore sedimentary successions and well ties

Deep fault-bounded basins formed along the SW Barents Sea margin during successive extension events in late Paleozoic–early Cenozoic times, and these basins contain important sedimentary successions for hydrocarbon exploration. We particularly focus on the late Paleozoic succession (Fig. 3), in which sedimentary rocks were deposited on top of eroded Precambrian and Caledonian basement rocks (see Townsend, 1987a; Johansen et al., 1994; Bugge et al., 1995; Zwaan, 1995; Gudlaugsson et al., 1998; Samuelsberg et al., 2003; Bergh et al., 2010). Late Paleozoic sedimentary deposits in the study area were penetrated by only a few exploration wells, to which we tied our seismic interpretation (Fig. 2). Overlying Mesozoic to Cenozoic sedimentary units were not investigated and are better described in Omosanya et al. (2015).

The nature and age of basement rocks along the SW Barents Sea margin remain relatively complex to resolve since only a handful of wells drilled through the thick post-Caledonian sedimentary cover. Nevertheless, wells 7128/4-1 and 7128/6-1 penetrated quartzitic metasedimentary rocks on the eastern Finnmark Platform (Fig. 2) and these are believed to correlate with upper Proterozoic rocks involved in Caledonian thrusting in northern Finnmark (Røe and Roberts, 1992).

Devonian sedimentary rocks are yet to be reported in northern Norway and along the SW Barents Sea margin (Fig. 3). However, Devonian sedimentary deposits are present in western Norway (Osmundsen and Andersen, 2001) where they represent a several-kilometer-thick succession made up of clastic deposits that notably include rhythmic sandstone and coarsely grained conglomerate units. These were deposited in the hanging wall of major, low-angle extensional shear zones, e.g., the Nordfjord-Sogn Detachment Zone (Séranne et al., 1989; Wilks and Cuthbert, 1994; Osmundsen and Andersen, 2001).

Lower Carboniferous sedimentary rocks of the Billefjorden Group directly overlie basement rocks on the eastern Finnmark Platform as evidenced by exploration wells 7128/4-1 and 7128/6-1 (Larssen et al., 2002; Figs. 2 and 3). These rocks mostly correspond to fluvial clastic deposits interbedded with coal-bearing sedimentary rocks that correlate with contemporaneous deposits on Bjørnøya (Cutbill and Challinor, 1965; Gjelberg, 1981, 1984) and Spitsbergen (Cutbill and Challinor, 1965; Cutbill et al., 1976; Gjelberg, 1984). The total thickness of Billefjorden Group sedimentary deposits evidenced by exploration wells on horst blocks on the eastern Finnmark Platform ranges from 350 to 450 m. However, in the hanging wall of a minor normal fault interpreted by Bugge et al. (1995) near the coast of northern Finnmark (Fig. 2), shallow drill cores 7127/10-U-2 and 7127/10-U-3 indicate that the thickness of lower Carboniferous sedimentary rocks reaches a thickness >600 m within a NE–SW-trending mini-basin on the eastern Finnmark Platform near the coast of the Nordkinn Peninsula (see star symbol in Figs. 1 and 2). In the Serpukhovian, fluvial sediments of the Billefjorden Group were gradually replaced by shallow marine sediments of the Gipsdalen Group from which they are generally separated by a mid-Carboniferous (Serpukhovian) unconformity (Cutbill et al., 1976; Gjelberg, 1984; Bugge et al., 1995) potentially related to a global sea-level fall (Saunders and Ramsbottom, 1986).

Shallow marine sedimentary deposits of the Gipsdalen Group are widespread along the SW Barents Sea margin and have proven prolific for hydrocarbon exploration (Larssen et al., 2002; Fig. 3). Thus, this sedimentary succession benefits from a relatively high number of well penetrations and, as a result, its lateral facies and thickness variations are well-constrained (Gjelberg and Steel, 1981, 1983; Samuelsberg et al., 2003; Rafaelsen et al., 2008). The Gipsdalen Group was notably penetrated by wells 7128/4-1 and 7128/6-1 on the eastern Finnmark Platform, by well 7120/12-4 on the western Finnmark Platform, and by well 7124/3-1 on the northern flank of the southwesternmost Nordkapp basin (Larssen et al., 2002; Fig. 2). This succession consists of alluvial clastic sedimentary rocks that are progressively replaced upwards by shallow marine platform carbonates interbedded with clastic and evaporite deposits (McCann and Dallmann, 1996). In well 7124/3-1 (Fig. 2), Asselian evaporite deposits typically include thin layers of anhydrite and gypsum, but thicker, halite-rich end-members are found along the flanks of the Nordkapp Basin and southwesternmost Nordkapp basin where large pillows of upper Carboniferous–lower Permian salt were observed (Gabrielsen et al., 1992; Jensen and Sørensen, 1992; Koyi et al., 1993; Nilsen et al., 1995; Gudlaugsson et al., 1998; Koehl et al., 2017). In the Nordkapp Basin, pre-Permian deposits may in places reach a thickness of up to 7–8 km (Gudlaugsson et al., 1998). These deposits are composed of thick clastic sedimentary rocks and of upper Carboniferous to lower Permian evaporite deposits characterized by mobile salt that was involved in salt tectonism in the southwesternmost Nordkapp basin (Gudlaugsson et al., 1998; Koehl et al., 2017) and in the Nordkapp Basin (Gabrielsen et al., 1992; Jensen and Sørensen, 1992; Koyi et al., 1993; Nilsen et al., 1995).

3 Methods and databases

3.1 Seismic data and well ties

The seismic interpretation shown in this study is based on publicly available 2-D and 3-D data from the Diskos database, thus providing reasonably tight 2-D data coverage.

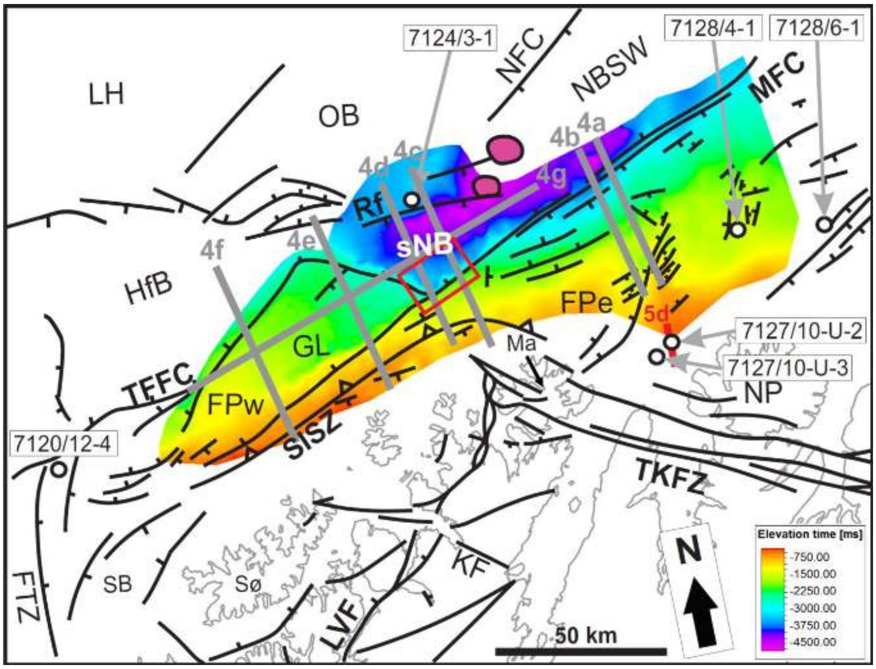

Figure 2. Regional structural map summarizing the architecture of the eastern (FPe) and western (FPw) Finnmark Platform and of the southwesternmost Nordkapp basin (sNB). The figure includes a time map of the interpreted mid-Carboniferous reflection. Grey lines show the location of seismic profiles displayed in Fig. 5a–g, the red line displays the location of the seismic section shown in Fig. 6d, and the red frame indicates the location of seismic *Z* slices described in Fig. 8. White dots show the location of exploration wells and shallow drill cores while purple blobs represent major salt diapirs in the southernmost part of the Nordkapp Basin (NBSW). See Fig. 1 for abbreviations.

However, only one seismic 3-D survey was available in the study area. The interpretation of seismic data aims at providing good constraints for the extent and geometry of offshore brittle faults and for offshore stratigraphy on the Finnmark Platform and in the southwesternmost Nordkapp basin. The present study uses ties to wells 7120/12-4, 7128/4-1, 7128/6-1, and 7124/3-1 based on publicly available well data (http://www.npd.no) and private well-tie seismograms and to shallow drill cores 7127/10-U-2 and 7127/10-U-3 from Bugge et al. (1995; Fig. 2). Seven seismic profiles from the BSS01 2-D seismic survey were used to analyze and describe offshore basin and fault geometries and provide the basis for discussion about the late Paleozoic evolution of the SW Barents Sea margin. Note that none of the seismic profiles used were depth converted. Therefore, all relevant estimates of fault offsets and stratigraphic seismic unit thicknesses will be described in seconds (s) two-way time (TWT). In addition, we analyzed two time slices from 3-D seismic survey MC3D-MFZ02 to constrain fault interaction in map view.

3.2 Aeromagnetic anomaly data

The offshore aeromagnetic data used in this study correspond to a compilation of the BASAR project of the Geological Survey of Norway (NGU) published by Gernigon and Brönner (2012) and Gernigon et al. (2014; Fig. 4). The dataset is composed of tilt derivatives of aeromagnetic data and has been used to delineate possible magmatic intrusions (dykes) emplaced along brittle faults (see Nasuti et al., 2015) and abrupt changes of lithology generally recorded across major faults, thus contributing to the mapping of post-Caledonian offshore brittle faults along the SW Barents Sea margin. However, data uncertainties arise from the fact that significantly different rock types may yield similar aeromagnetic responses. A crucial example in northern Finnmark is the similar high positive narrow aeromagnetic anomalies produced by both subvertical folded beds of metasedimentary rocks (Roberts and Siedlecka, 2012; Roberts and Williams, 2013) and dolerite dykes intruded along brittle faults (Nasuti et al., 2015; Fig. 4). In order to distinguish such features, we carefully analyzed onshore geology in coastal areas of NW Finnmark and the results of exploration wells on the Finnmark Platform and adjacent offshore basins.

4 Results

4.1 Seismic interpretation of offshore basins and faults

4.1.1 Seismic units and stratigraphy

In seismic data (Fig. 5; see the Supplement for uninterpreted seismic sections), basement rocks typically show chaotic internal reflection patterns, which complicate the task of identi-

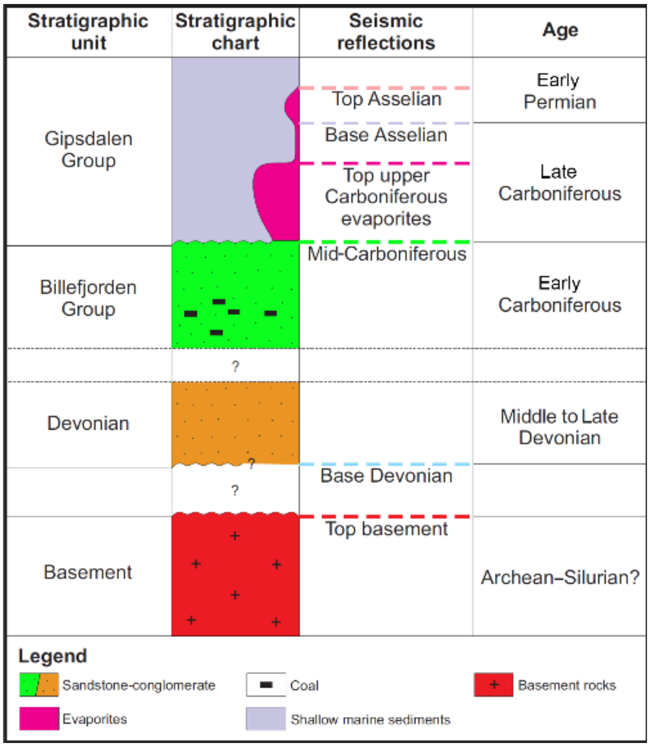

Figure 3. Simplified stratigraphic chart of late Paleozoic sedimentary successions on the Finnmark Platform and in the southwesternmost Nordkapp basin. From left to right, columns indicate the unit name, the successions' dominant lithologies and types of succession boundaries (undulating lines: erosional unconformity; straight line: conformity; dashed lines: uncertain), interpreted seismic reflections (see Figs. 4 and 5), and the units age. Lithological legend at the bottom.

fying intra-basement structures and basins, and individualize layered sedimentary sequences. However, kilometer-thick layers bearing strong basement fabrics such as widespread, gently dipping foliation or pronounced mylonitic fabric commonly found along large shear zones may turn out to be resolvable on the seismic scale (see Sect. 4.1.2.; Fountain et al., 1984; Reeve et al., 2013; Phillips et al., 2016; Fazlikhani et al., 2017). For instance, we observed a several-kilometer-thick, curved, shallow-dipping layer that is characterized by moderate-amplitude reflections, which are parallel to the layer's upper and lower boundaries (see "Sørøya-Ingøya shear zone" reflections in Fig. 5c–g). We interpret these pronounced internal fabrics as widespread mylonitic foliation separated by internal thrusts within a large-scale shear zone. Numerous smaller basement shear zones may be present below late Paleozoic–Cenozoic sedimentary rocks on the western Finnmark Platform, and these correspond to steeply to moderately dipping fabrics made of subparallel, moderate- to high-amplitude reflections (see Figs. 5b, e, f, g and 6a–c).

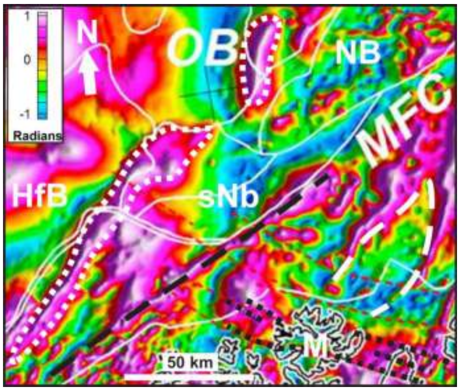

Figure 4. Enlargement of offshore tilt-derivative aeromagnetic data published by Gernigon et al. (2014). The white dashed line on the eastern Finnmark Platform represents a triangular- to rhomboid-shaped aeromagnetic low that coincides with a Carboniferous basin bounded by zigzag-shaped brittle faults (e.g., LVF). The dotted white lines on the western Finnmark Platform and on the northern flank of the southwesternmost basin represent ENE–WSW- to NE–SW-trending ridges of magnetic basement rocks. The dashed black line represents a linear, NE–SW-trending, high positive aeromagnetic anomaly that has been tied to the occurrence of the main segment of the MFC (see Indrevær et al., 2013). Dolerite dykes intruded along WNW–ESE-trending segments of the TKFZ are shown by dotted black lines. Dashed red lines are the interpretation from Gernigon et al. (2014). See Fig. 1 for abbreviations.

Potential Devonian sedimentary deposits along the SW Barents Sea are sparse and as a result their seismic character is not well constrained (Fig. 3). This sedimentary succession has not been drilled, which makes its interpretation on seismic data rather speculative. However, we believe that the best two candidates to represent Devonian sedimentary deposits analog to those in western and mid-Norway (Braathen et al., 2000; Osmundsen and Andersen, 2001; Fazlikhani et al., 2017) are located at the base of the southwesternmost Nordkapp basin and on the western Finnmark Platform near the Gjesvær Low (Fig. 1). In the southwesternmost Nordkapp basin, possible Devonian sedimentary strata are located at a deep level (below 4 s TWT) and their seismic signature is thus largely masked by overlying sedimentary successions (Fig. 5c and d). By contrast, on the western Finnmark Platform (Fig. 5e) potential Devonian sedimentary rocks are relatively shallower, which makes their seismic pattern easier to distinguish from underlying basement rocks and from overlying Carboniferous sedimentary deposits and seismic artifacts (Fig. 5e). Devonian sedimentary rocks on the western Finnmark Platform display relatively low seismic amplitudes, partly similar to analog deposits in the North Sea (see seismic facies 1 in Fazlikhani et al., 2017). The internal reflection pattern is rather chaotic apart from a few discrete, shallow-dipping, moderate-amplitude reflections that converge towards each other upwards and that we interpret as major sedimentary sequence boundaries (see dotted white

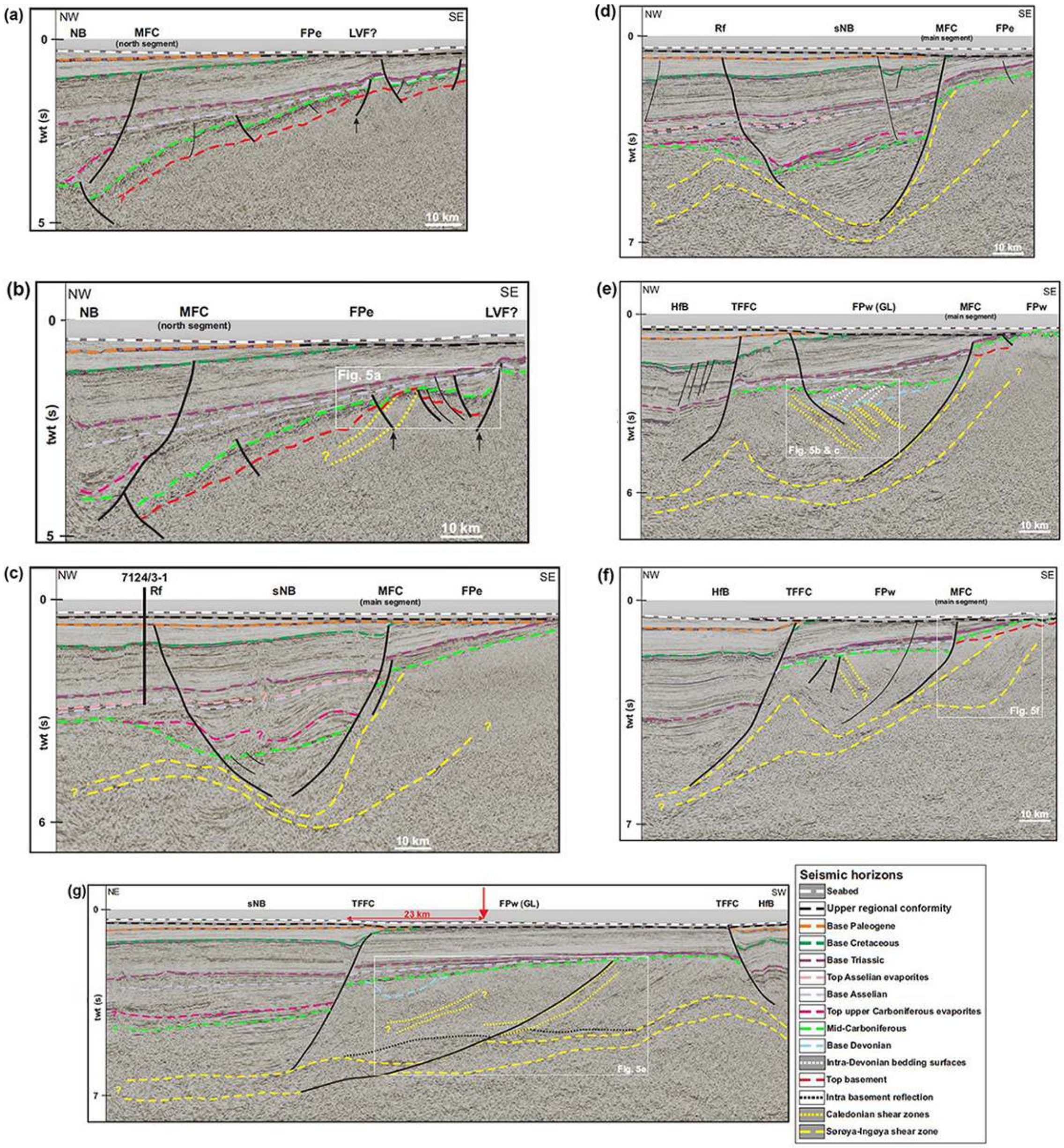

Figure 5. Examples of interpreted seismic profiles from the BSS-01 survey (2-D), the locations of which are displayed in Fig. 2. Brittle faults are shown in black and depth is in seconds (s) TWT. See Fig. 1 for abbreviations; **(a)** interpreted seismic section that shows a system of Carboniferous horst and graben structures on the eastern Finnmark Platform; **(b)** seismic profile showing increased normal displacement across the NW-dipping LVF compared with panel **(a)** and thickening of the Carboniferous sedimentary succession within the graben bounded by the LVF. Note the insignificant amount of the displacement accommodated by the northern segment of the MFC in panels **(a)** and **(b)**. Black arrows mark brittle faults that bound a triangular-shaped, negative aeromagnetic anomaly (see dashed white line in Fig. 4); **(c)** seismic profile showing a highly thickened Carboniferous succession and potential Devonian–lower Carboniferous sedimentary rocks in the southwesternmost Nordkapp basin. Note the large offset accommodated by the main segment of the MFC and the peculiar "U" shape of the southwesternmost Nordkapp basin. Also displayed is a lateral projection of exploration well 7124/3-1; **(d)** interpreted seismic section that shows the listric geometries of the main segment of the MFC and of the Rolvsøya fault; **(e)** seismic section showing potential Devonian sedimentary rocks deposited in a NE–SW-trending graben above a set of minor, SE-dipping shear zones on the western Finnmark Platform; **(f)** seismic section showing the listric geometries of the TFFC and MFC, which both seem to merge into the SISZ; **(g)** NE–SW-trending seismic cross section across the western Finnmark Platform and the southwesternmost Nordkapp basin showing the gentle dip of the SISZ to the northeast and a gradual thinning of the upper Carboniferous sedimentary succession towards the southwest. A major NNE–SSW-trending, SE-dipping brittle fault seems to offset the SISZ and an intra-basement reflection on the western Finnmark Platform before being truncated by the mid-Carboniferous reflection. The vertical red arrow shows the location of the imaginary prolongation of the TKFZ on the western Finnmark Platform as a comparison with the actual location of the WNW–ESE-trending fault segment of the TFFC, which are separated by a distance of ca. 23 km.

reflections in Figs. 5e and 6b and c). Furthermore, Devonian sedimentary deposits are likely separated from underlying basement rocks by an angular unconformity that appears as arcuate, high-amplitude seismic reflections ("base Devonian" reflection in Figs. 5e and 6b and c). We interpret these arcuate, high-amplitude seismic reflections as an erosional unconformity.

Lower Carboniferous sedimentary deposits of the Billefjorden Group, composed of thick clastic sedimentary deposits interbedded with occasional coal-bearing sedimentary rocks (Fig. 3), may produce high-amplitude seismic reflections related to their organic-rich content (Fig. 5a and b). Such sedimentary strata are present on the eastern Finnmark Platform, where they appear to thicken to the southeast near the coast of NW Finnmark (Fig. 6d), whereas they are rather sparse on the western Finnmark Platform, i.e., eroded or never deposited (Fig. 5e and f). On the eastern Finnmark Platform, the transition from basement rocks (see "Top basement" reflection in Fig. 5a and b) to lower Carboniferous sedimentary rocks is difficult to interpret on seismic sections. This is attributable to the strong similarities between high seismic amplitudes displayed locally by both basement rock fabrics such as major shear zones (see yellow dotted lines in Fig. 5b) and lower Carboniferous coal-bearing sedimentary deposits. Low-amplitude reflections also show identical chaotic patterns in both basement rocks and clastic sedimentary rocks of the Billefjorden Group (Fig. 5a and b). In the southwesternmost Nordkapp basin, lower Carboniferous sedimentary strata are believed to be present, although their seismic signature certainly appears to be affected by overlying upper Carboniferous evaporite deposits (Fig. 5c and d). The boundary between lower Carboniferous sedimentary deposits and potential underlying Devonian sedimentary rocks was not identified in the southwesternmost Nordkapp basin. Nevertheless, the maximum thickness of Billefjorden Group sedimentary strata on the eastern Finnmark Platform is ca. 600 m (Bugge et al., 1995), and this suggests that the several-kilometer-thick succession below the mid-Carboniferous reflection and above a thick shear zone in the southwesternmost Nordkapp basin is composed of lower Carboniferous sedimentary rocks probably complemented by thick Devonian sedimentary deposits (Fig. 5c and d). Alternatively, sedimentary deposits of the Billefjorden Group directly overlie basement rocks.

On the Finnmark Platform (Figs. 1 and 2), the base of the upper Carboniferous sedimentary succession is difficult to identify (see "mid-Carboniferous" reflection in Figs. 3 and 5). In places, it appears as a linear, moderate- to low-amplitude seismic reflection that separates subparallel reflections of lower and upper Carboniferous sedimentary rocks, whereas in other places the reflection is irregular and truncates high-amplitude coal-bearing sedimentary deposits of the Billefjorden Group and/or high-amplitude reflections produced by basement rocks (Fig. 6a) and/or low-amplitude reflections in Devonian sedimentary strata (Fig. 6b and c). Nevertheless, this reflection generally corresponds to an angular unconformity (e.g., Fig. 6a–c and e) and is therefore interpreted to correspond to a regional erosion surface.

In the southwesternmost Nordkapp basin, the base of upper Carboniferous sedimentary deposits (see "mid-Carboniferous" reflection in Figs. 3 and 5c and d) appears as a clear, discrete high-amplitude reflection. The strong acoustic impedance contrast producing the high seismic amplitude for the mid-Carboniferous reflection most likely arises from the presence of upper Carboniferous evaporite deposits partly composed of mobile salt (halite), which is significantly less dense than regular sedimentary rocks (see "Top upper Carboniferous evaporites" reflection in Figs. 3 and 5c and d). This evaporite succession was identified by Gudlaugsson et al. (1998) and is restricted to basinal areas located northwest of the MFC and north of the TFFC (Figs. 1 and 2). It is characterized by a highly variable thickness, which is due to the presence of lensoidal bodies bounded to the top and bottom by high-amplitude reflections on the basin edges and to the occurrence of thick bodies made of chaotic reflection patterns near the center of the basin (Fig. 5c). We interpret the lensoidal bodies on the basin edges as pillows of mobile salt and the chaotic bodies near the basin center as small salt diapirs based on similarities with large salt diapirs and evaporite deposits observed in the Nordkapp Basin (Gabrielsen et al., 1992; Jensen and Sørensen, 1992; Koyi et al., 1993; Nilsen et al., 1995). We consider that the presence of analog late Paleozoic evaporite deposits in the southwesternmost Nordkapp basin and in the Nordkapp Basin (Jensen and Sørensen, 1992; Koyi et al., 1993; Gudlaugsson et al., 1998) and the absence of such deposits in the Hammerfest Basin constitute strong arguments to justify a change of name for the "easternmost Hammerfest basin" (Omosanya et al., 2015) into the "southwesternmost Nordkapp basin". However, this basin shows a large amount of normal displacement along its southern boundary fault, the NW-dipping MFC, which is opposite to the Nordkapp Basin where basin subsidence was dominantly accommodated along the SE-dipping Nysleppen Fault Complex (Fig. 1). Hence, despite their similarities, the Nordkapp Basin and the southwesternmost Nordkapp basin should be treated as two separate basins.

Non-evaporitic, upper Carboniferous and Permian sedimentary deposits are characterized by subparallel, flat-lying to shallow-dipping, homogeneous, moderate- to low-amplitude seismic reflections (see Fig. 5). Permian deposits are relatively thin on the Finnmark Platform and are sometimes difficult to distinguish from upper Carboniferous deposits (Fig. 5a, b, e, f, and g). In the southwesternmost Nordkapp basin, however, late Paleozoic sedimentary deposits are thicker and individual units are therefore easier to identify in seismic data. Thus, we interpreted a thin unit characterized by high-amplitude reflections (see "base Asselian" and "top Asselian evaporites" reflections in Figs. 3 and 5c and d) as Asselian (earliest Permian) evaporite deposits that were evidenced by exploration well 7124/3-1 on the northern flank

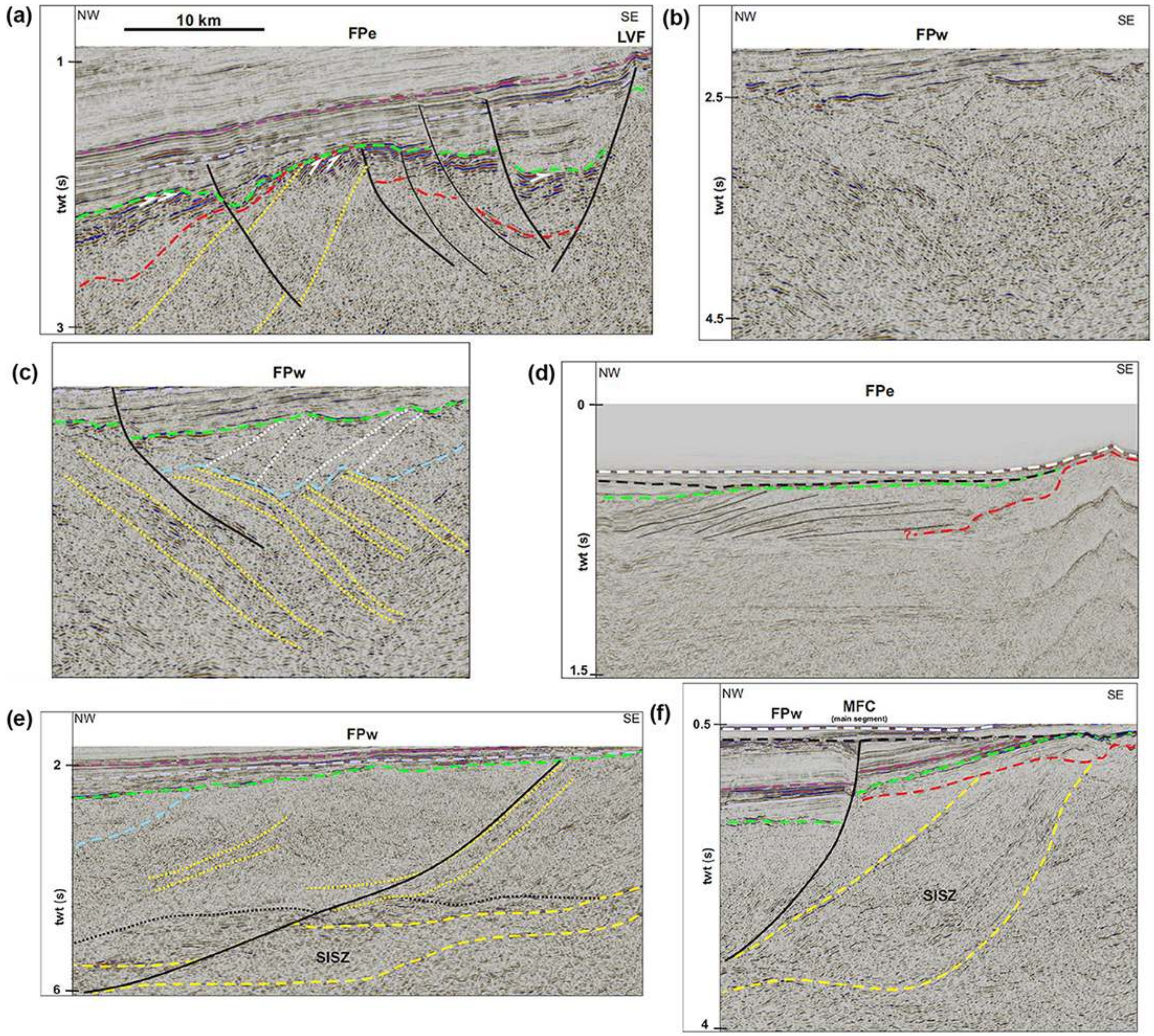

Figure 6. Enlargement of seismic sections on the eastern and western Finnmark Platform. The locations of **(a)**, **(b)**, **(c)**, **(e)**, and **(f)** are displayed as white frames in Fig. 5 and the location of **(d)** is shown as a red line in Fig. 2. See Fig. 1 for abbreviations and Fig. 5 for seismic reflection legend; **(a)** interpreted seismic section across the eastern Finnmark Platform. White arrows represent high-amplitude lower Carboniferous and basement seismic reflections that are truncated upwards (toplaps) by the mid-Carboniferous reflection. Note the contrast between low-amplitude upper Carboniferous–Permian reflections; gently dipping, high-amplitude lower Carboniferous reflections; and steeply dipping, high-amplitude basement reflections that possibly belong to a basement-seated shear zone (yellow dotted lines); **(b)** uninterpreted and **(c)** interpreted seismic zoom of a section across presumed Devonian sedimentary rocks and SE-dipping basement shear zones (yellow dotted lines) on the western Finnmark Platform; **(d)** interpreted seismic section from the IKU-87-BA (2-D) survey showing a thick lower Carboniferous succession made up of large clinoforms (thin black lines) on the eastern Finnmark Platform (location in Fig. 2). Note the presence of seismic artifacts in the southeast, including several multiples and NW-dipping diffraction rays; **(e)** interpreted seismic section across the western Finnmark Platform that displays NE-dipping basement shear zones (yellow dotted lines) including the SISZ (yellow dashed lines); **(f)** seismic zoom in the SISZ in the footwall of the main segment of the MFC on the western Finnmark Platform. The SISZ is composed of NW-dipping, moderate- to high-amplitude reflections that dip more gently than the MFC but that are steeper than basement reflections in the southeast. Note the significant thickness variations in the SISZ: thick in the footwall of the MFC and thin below the MFC.

of the southwesternmost Nordkapp basin (Figs. 2, 5c and d). Where present, this thin Asselian evaporite succession defines the base of the Permian sedimentary succession and therefore serves as an upper boundary for the Carboniferous succession (see "base Asselian" reflection in Fig. 5c and d). However, Asselian evaporites are too thin and too discontinuous to be seismically resolvable on the Finnmark Platform (Bugge et al., 1995). Occasionally, Asselian evaporites are truncated by chaotic reflections of small salt diapirs sourced from deeper upper Carboniferous evaporites in the southwesternmost Nordkapp basin (Fig. 5c).

The base Triassic reflection (see Fig. 5) defines the (near-)top of the late Paleozoic sedimentary succession and is easily interpreted through the whole Barents Sea as it corresponds to a high-amplitude reflection that represents the top of a regionally widespread carbonate unit (Bugge et al., 1995). Other important seismic reflections interpreted in the present study include the base Cretaceous; base Paleocene; the upper regional unconformity, which corresponds to a major erosional unconformity and represents the base of Quaternary sediment cover (Solheim and Kristoffersen, 1984); and the seabed reflection (Fig. 5). These reflections are penetrated by a large number of exploration wells and shallow drill cores both on the Finnmark Platform and in the southwesternmost Nordkapp basin, where they all display consistently high seismic amplitudes (Faleide et al., 1984; Bugge et al., 1995; Gudlaugsson et al., 1998; Omosanya et al., 2015).

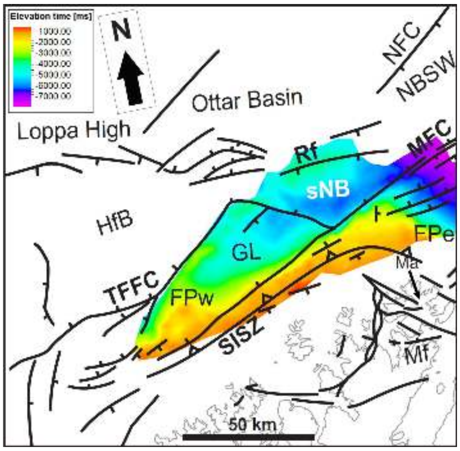

Figure 7. Time surface map of the top reflection of the SISZ and major brittle faults in the SW Barents Sea. Note the spoon-shaped depression formed by the SISZ on the western Finnmark Platform and southwesternmost Nordkapp basin, the abrupt change to a northeastward dip on the eastern Finnmark Platform, and the two narrow, NE–SW- and ENE–WSW-trending ridges in the footwall of the TFFC and of the Rolvsøya fault.

4.1.2 Structural architecture of the Finnmark Platform and of the southwesternmost Nordkapp basin

In this section, we describe the most important structural elements of the Finnmark Platform and of the southwesternmost Nordkapp basin (see Figs. 1 and 2) based on interpreted key seismic sections (Fig. 5). We also highlight the most dominant fault trends and their interactions with major structures such as the TFFC, MFC, TKFZ, and SISZ to form offshore sedimentary basins.

Faults and shear zones within basement rocks

We identified a several-kilometer-thick, curved (in cross section), shallow-dipping layer of moderate-amplitude reflections that we interpreted to represent a large-scale basement-seated shear zone, which we name the SISZ. The upper boundary surface of the SISZ (Fig. 7) appears to be relatively shallow in coastal areas. On the western Finnmark Platform, the SISZ dominantly dips to the NW but switches to a dominant dip to the northeast on the eastern Finnmark Platform. In the footwall of the MFC and in the southwestern part of the western Finnmark Platform, the SISZ occurs at a relatively shallow depth (< 1.5 s TWT). There it is believed to have been deeply eroded and is now overlain by a very thin sedimentary cover (see Figs. 5c–f and 6d). The SISZ shows significant lateral thickness variations that range from 2.0 to 2.5 s (TWT) near the coastline and in the footwall of the TFFC to 0.5 s (TWT) below the MFC and the TFFC (Fig. 5f). The SISZ deepens to the northwest towards the center of the western Finnmark Platform before bending upwards in the footwall of the TFFC (Fig. 5e and f). The SISZ then curves down where the listric TFFC merges with the shear zone at depth, thus delineating an elongated, NE–SW-trending ridge in the footwall of the TFFC (see "basement ridges" in Figs. 1 and 5e and f). A similar pattern is observed in the southwesternmost Nordkapp basin where the SISZ deepens to the northwest before curving up near the center of the basin and merging with the N-boundary fault of the southwesternmost Nordkapp basin, the Rolvsøya fault, hence giving this basin a characteristic "U" shape in cross-section (Fig. 5c and d). The SISZ also curves down in the footwall of the Rolsøya fault and defines a second elongated, ENE–WSW-trending ridge (see "basement highs" in Fig. 1). Importantly, the two basement ridges located in the footwall of the TFFC and of the Rolvsøya fault ("basement highs" in Fig. 1) are separated by a narrow trough that is bounded to the southwest by the WNW–ESE-trending segment of the TFFC (Fig. 7). Apart from this narrow trough, the attitude of the SISZ is uniform along NE–SW transects on the western Finnmark Platform and within the southwesternmost Nordkapp basin with a gentle dip to the northeast (Fig. 5g).

Notably, the spoon-shaped geometry of the SISZ, with asymmetric, NE–SW-trending, northeastward-broadening NE plunge (Fig. 7) appears to coincide with a basement gravity low on the western Finnmark Platform: the Gjesvær Low (Johansen et al., 1994; Gernigon et al., 2014; Fig. 1). The geometry of the SISZ also matches the trend and shape of the southwesternmost Nordkapp basin (Figs. 1 and 7). Farther south, along the coasts of western Troms and westwards below the Hammerfest Basin, the low quality of available seismic data did not allow us to trace the SISZ more precisely (Fig. 7). On the eastern Finnmark Platform, the SISZ bends from NE–SW into a more WNW–ESE trend and changes in dip from gentle to steep to the northeast (Fig. 7), and as a result the SISZ becomes too deep to interpret on seismic data in the northeastern part of the eastern Finnmark Platform (Fig. 7). The multiple changes of trend, dip direction, dip angle, and thickness of the SISZ gives the shear zone a spoon-shaped geometry (Fig. 7).

On the western Finnmark Platform, subsidiary, steep SE-dipping high-amplitude reflections occur in basement rocks and these are truncated by the mid-Carboniferous reflection and the base Devonian erosional unconformity in the footwall of the TFFC (see yellow dotted lines in Fig. 5e–g). Despite dipping southeast, these reflections resemble the dominant reflection pattern observed within the SISZ (Fig. 5e and f). Thus, we interpret them as SE-dipping, mylonitic shear zones (yellow dotted lines in Fig. 5e–g). The upper boundary of one of these SE-dipping shear zones coincides with an abrupt seismic facies change on the western Finnmark Platform, from moderately dipping, moderate-amplitude reflections in the west to gently dipping to subhorizontal low-amplitude seismic reflections in the east (Fig. 5g). This change also coincides with a ca. 1 s (TWT) deepening of the upper boundary of the SISZ towards the northeast (Fig. 5g) and with a small normal offset of a lensoidal, eastwards-thickening layer of subhorizontal reflections located above the SISZ (see dotted black lines in Fig. 5g). We interpret these changing attributes to be related to the presence of a NNE–SSW-trending, ESE-dipping brittle fault that flattens and merges into the SISZ and which may have developed along a preexisting, steep ductile shear zone (yellow dotted lines in Fig. 5g).

Similar NE–SW-trending but NW-dipping shear zones may exist in basement rocks on the eastern Finnmark Platform, for example in the form of steeply dipping, high-amplitude seismic reflections truncated by the mid-Carboniferous reflection (see yellow dotted lines in Figs. 5b and 6a). These reflections differ from gently dipping, high-amplitude reflections of lower Carboniferous coal-bearing sedimentary deposits (Fig. 6a) and rather resemble the SISZ reflection pattern, though these are located well above the presumed continuation of the SISZ (Fig. 5e and f). We therefore interpret these steep reflections as a NE–SW-trending, NW-dipping shear zone similar to the SISZ (Fig. 5b).

Faults within late Paleozoic sedimentary successions

Faults bounding Paleozoic sedimentary strata and basins include the major TFFC and MFC and numerous faults on the Finnmark Platform. The TFFC is made of alternating ENE–WSW- and NNE–SSW-trending, NW-dipping, listric fault segments that form a zigzag pattern and that separate the Hammerfest Basin in the northwest from the western Finnmark Platform in the southeast (Figs. 1 and 5e and f; Gabrielsen et al., 1990; Indrevær et al., 2013). Seismic data below ENE–WSW- and NNE–SSW-trending fault segments of the TFFC show that these fault segments merge with and merge into shallow-dipping reflections of the SISZ at depth (Fig. 5e and f). At the northeast termination of the Hammerfest Basin, the TFFC bends 90 degrees clockwise and continues to the southeast as a WNW–ESE-trending, NE-dipping, listric fault (Figs. 1, 2, and 5g). At depth, this fault merges with the SISZ (see Fig. 5g) near a narrow trough in the top surface of the SISZ, separating two elongated NE–SW- to ENE–WSW-trending basement ridges in the footwall of the TFFC and of the Rolvsøya fault (see "basement highs" in red in Figs. 1 and 7). In map view, the WNW–ESE-trending, NE-dipping segment of the TFFC bends anticlockwise into the main fault segment of the MFC, which corresponds to a linear, NE–SW-trending, NW-dipping fault (Figs. 1, 2, and 8a and b). The interaction of these two faults in map view gives the western Finnmark Platform and the southwesternmost Nordkapp basin triangular shapes (Figs. 2 and 8a and b). The main segment of the MFC defines the southeastern boundary of the southwesternmost Nordkapp basin (Figs. 1, 2, and 5c and d) and of a ca. 25–30 km wide graben structure on the western Finnmark Platform that is believed to be partly filled with Devonian sedimentary deposits (Figs. 1, 2, and 5e and f). Northeastwards, the main segment of the MFC (Fig. 5c–f) is replaced by several minor fault segments with limited vertical throw (Fig. 5a and b) that define the southeastern boundary of the Nordkapp Basin (Figs. 1 and 5a and b). The southwesternmost Nordkapp basin is bounded to the north by an E–W- to ENE–WSW-trending, south-dipping, listric normal fault, the Rolvsøya fault, which flattens at depth and merges into gently dipping reflections of the SISZ (Fig. 5c and d). The Rolvsøya fault separates the southwesternmost Nordkapp basin from the Ottar Basin to the northwest and from the Nordkapp Basin to the northeast (Figs. 1 and 2).

Late Paleozoic grabens on the eastern Finnmark Platform display fault patterns that are analogous to those that shape the southwesternmost Nordkapp basin and the western Finnmark Platform (Figs. 1 and 2). Numerous steeply dipping, listric normal faults made of alternating, zigzag-shaped, ENE–WSW- and NNE–SSW-trending segments bound relatively narrow, few-kilometer-wide graben and half-graben structures that are filled with wedge-shaped, late Paleozoic sedimentary successions (Figs. 2 and 5a and b). In particular, one of these zigzag-shaped faults trends NE–SW to NNE–SSW, dips to the northwest, and can be traced for about 60 km from the northern coast of Magerøya onto the eastern Finnmark Platform (Figs. 1 and 2). Southwestward, this fault roughly aligns with a similarly shaped and oriented, NW-dipping onshore and nearshore fault complex synthetic to the TFFC described as the LVF (Figs. 2 and 5a and b; Zwaan and Roberts, 1978; Lippard and Roberts, 1987; Roberts and

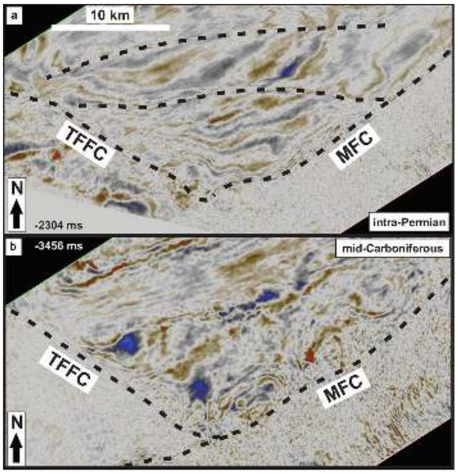

Figure 8. (a) Intra-Permian seismic time slice within 3-D seismic survey MC3D-MFZ02 in the southwesternmost Nordkapp basin. Dashed black lines correspond to interpreted brittle faults; **(b)** seismic time slice within 3-D seismic survey MC3D-MFZ02 near the interpreted mid-Carboniferous reflection in the southwesternmost Nordkapp basin. Black dashed lines represent interpreted brittle faults. See Fig. 2 for location.

Lippard, 2005; Koehl et al., 2018). We tentatively interpret the ca. 60 km long, zigzag-shaped brittle fault on the eastern Finnmark Platform, northeast of Magerøya, as the northeastward continuation of the LVF on the eastern Finnmark Platform (Figs. 1, 5a and b, and 6a).

Below the minor northern segments of the MFC, we identified a large NE–SW-trending, SE-dipping fault that is antithetic to the MFC (Fig. 5a and b). Due to the rather low quality of seismic data at large depths, the interaction of the northern segments of the MFC with the antithetic SE-dipping fault is difficult to evaluate. Our data indicate that the northern segments of the MFC crosscut the NE–SW-trending, SE-dipping in the southwest (Fig. 5b), whereas farther northeast, along strike, the northern fault segments of the MFC seem to merge and die out into upper Carboniferous evaporite deposits (Fig. 5a).

4.1.3 Fault-controlled thickness variations

In the following section, fault offsets and thickness variations in the sedimentary successions across brittle faults will be described as a basis to infer timing and sense of shear for brittle faults on the Finnmark Platform and in the southwesternmost Nordkapp basin. Regional stratigraphic thickness maps (Fig. 9a–c) show that late Paleozoic sedimentary strata on the eastern Finnmark Platform thicken from < 0.1 s (TWT) in the southeast to a maximum thickness of ca. 2 s (TWT) in the footwall of the MFC (see also Fig. 5a and b). This gradual thickness increase contrasts with the abrupt thickness increase in Devonian–Carboniferous sedimentary strata in the hanging wall of major normal faults, e.g., the WNW–ESE-trending segment of the TFFC and the main segment of the MFC (Fig. 9a–b), thus separating depositional versus tectonic thickness changes.

Intra-basement thickness changes

The dominant shear zone system within basement rocks on the western Finnmark Platform is the SISZ (Figs. 5c–g, 6b–c and e–f, and 7). A pronounced intra-basement unit made of subhorizontal, high-amplitude reflections occurs above the SISZ (Fig. 5g). The top reflection of the SISZ and the overlying intra-basement unit are offset by a NNE–SSW-trending, gently east-dipping fault, which is accompanied by a thickness increase in the intra-basement unit across the east-dipping fault (see black dotted line in Figs. 5g and 6e). This fault is interpreted to have a top–E normal sense of shear (see dotted black lines in Figs. 5g and 6e) and is itself truncated by the subhorizontal mid-Carboniferous reflection, which constrains its activity to the Middle to Late Devonian–early Carboniferous (Fig. 5g).

Fault-controlled thickness changes in Devonian–Carboniferous strata

In the southwesternmost Nordkapp basin, the Devonian–lower Carboniferous sedimentary succession (Fig. 5c and d) appears to be thickest at the intersection of the TFFC and MFC (Fig. 9a), where vertical displacement along the MFC and TFFC is estimated to be ca. 1.5 s (TWT), based on an offset of the mid-Carboniferous reflection (see Fig. 5d). The overlying upper Carboniferous succession displays a similar attitude as shown by the broad thickening of similar sedimentary strata at the intersection of the TFFC and MFC (Fig. 9b). These observations suggest that the WNW–ESE-trending segment of the TFFC and the main segment of the MFC potentially formed simultaneously in Devonian times and acted as syn-sedimentary normal faults that contributed to the thickening of Devonian–lower Carboniferous and upper Carboniferous sedimentary deposits within the southwesternmost Nordkapp basin (Fig. 5c and d). In this scenario, the Rolvsøya fault likely limits the extent of thickened Devonian–lower Carboniferous and upper Carboniferous sedimentary strata to the north. If we consider the thickness of the seismic package limited upwards by the mid-Carboniferous reflection and downwards by the top reflection of the SISZ in the footwall of the Rolvsøya fault, the maximum thickness of Devonian and lower Carboniferous sedimentary rocks on the northern flank of the basin does not exceed ca. 1 s (TWT). This thickness estimate is significantly thinner than what is observed within the southwesternmost Nordkapp basin, where the Devonian–lower Carboniferous succession reaches a maximum thickness of ca. 2–2.5 s (TWT; see Fig. 5c and d). By analogy, the thickness of upper Carboniferous sedimentary strata on the northern flank of the southwesternmost Nordkapp basin decreases from ca. 1.5 s

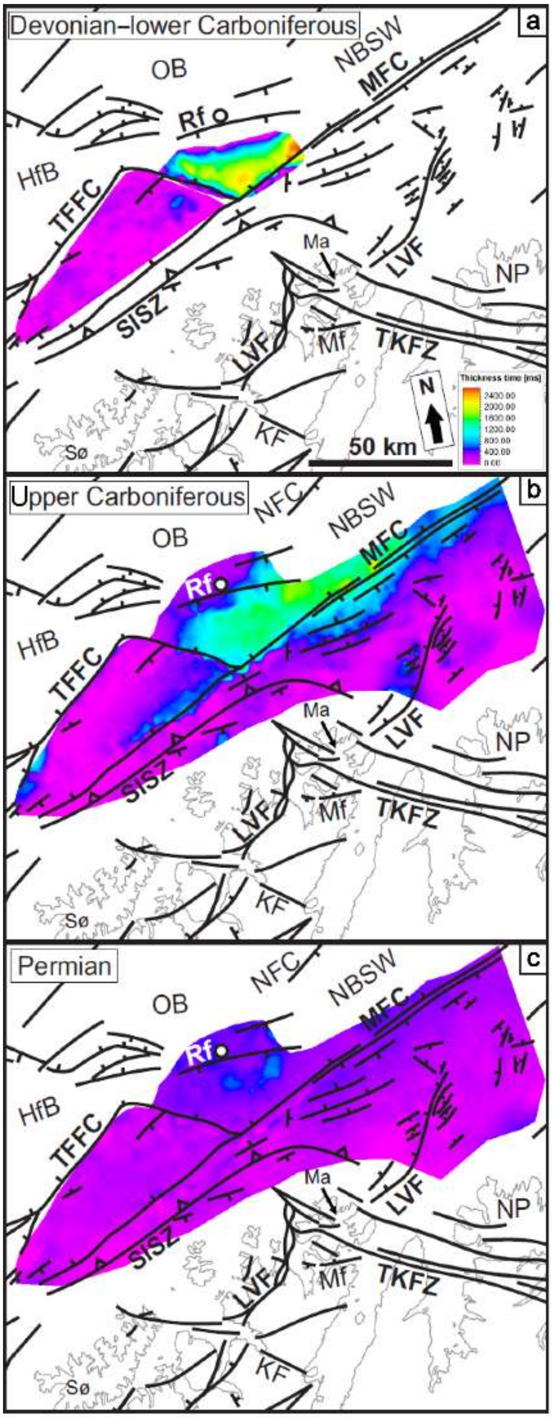

Figure 9. Thickness maps in milliseconds (ms) two-way time (TWT) of late Paleozoic sedimentary successions on the Finnmark Platform and in the southwesternmost Nordkapp basin. Color scale in panel **(a)**; **(a)** thickness map of the Devonian–lower Carboniferous succession on the western Finnmark Platform and in the southwesternmost Nordkapp basin. The succession is thickest in the southwesternmost Nordkapp basin and represents the thickest sedimentary unit of the basin. Note that in this part of the margin, the SISZ and basin-bounding faults were used as base Devonian reflections. On the western Finnmark Platform, lower Carboniferous sedimentary rocks are missing but Devonian sedimentary deposits are possibly preserved in an ENE–WSW-trending graben adjacent to the southwesternmost Nordkapp basin and bounded to the southeast by the MFC; **(b)** thickness map of the upper Carboniferous sedimentary succession showing gradual thickening of upper Carboniferous sedimentary rocks in the southwesternmost Nordkapp basin, on the western Finnmark Platform in the hanging wall of the MFC, and on the eastern Finnmark Platform in the hanging wall of the LVF and of a SE-dipping fault that parallels the MFC; **(c)** thickness map of the Permian succession showing very thin Permian sedimentary deposits and very mild thickness variations within the Permian sedimentary succession throughout the study area.

(TWT) to ca. 0.5–1 s across the Rolvsøya fault (Figs. 5c and d and 9b). Hence, the Rolvsøya fault was active and largely contributed to sediment thickening within the southwesternmost Nordkapp basin during the Middle to Late Devonian–Carboniferous.

On the western Finnmark Platform, potential Devonian sedimentary rocks are characterized by low-amplitude chaotic reflections within which we observed distinct, shallow-dipping, moderate-amplitude reflections that we interpreted as major sedimentary sequence boundaries (see white dotted lines in Figs. 5e and 6b and c). These shallow-dipping reflections diverge from each other downwards and define gently dipping, wedge-shaped layers of low-amplitude chaotic reflections that thicken downwards against arcuate, high-amplitude basement reflections that represent an erosional unconformity (see "base Devonian" reflection in Fig. 5e), and to the northwest against an ENE–WSW-trending, SE-dipping normal fault (Figs. 5e and 6b and c). We interpret these sedimentary units separated by shallow-dipping, moderate-amplitude reflections to represent growth strata deposited along an active ENE–WSW-trending, SE-dipping normal fault, which is parallel to SE-dipping basement shear zones (Figs. 5e and 6b and c). In addition, the main segment of the MFC shows a decreasing amount of vertical displacement to the southwest, accompanied by a simultaneous thickness decrease in the upper Carboniferous succession along strike (Fig. 9b), before the MFC eventually dies out on the western Finnmark Platform (Figs. 1, 2, and 5e and f). Analogously, upper Carboniferous sedimentary deposits on the western Finnmark Platform display a wedge shape that is thickest in the southeast, near the MFC, and gradually thins towards the TFFC in the northwest (Figs. 5e and f, 9b). This upper Carboniferous sedimentary wedge likely formed by syn-tectonic sedimentary growth along the main segment of the MFC.

On the eastern Finnmark Platform, the offshore portion of the LVF (see Figs. 1 and 5a and b) downthrows the mid-Carboniferous reflection by ca. 0.5 s (TWT) to the northwest (Fig. 5b) and bounds a NE–SW-trending graben structure filled with thickened lower Carboniferous and upper Carboniferous sedimentary strata (see Fig. 5a and b). In this graben structure, the lower Carboniferous and upper Carboniferous sedimentary successions thicken against the LVF (Fig. 5b), while thickness variations become negligible farther north where the LVF dies out (Figs. 1 and 5a). Consequently, similar thickness increases of lower Carboniferous and upper Carboniferous sedimentary strata elsewhere within graben and half-graben structures on the eastern Finnmark Platform suggest that syn-tectonic sediment deposition along the LVF and analog ENE–WSW- to NNE–SSW-trending faults mostly occurred in Carboniferous times. Furthermore, in the footwall of the northern segments of the MFC, we recorded anomalously thick upper Carboniferous succession (Fig. 9b) with a thickness comparable to what is observed within the southwesternmost Nordkapp basin (Fig. 9b). This succession shows a half-ellipsoid shape in map view with a NE–SW-trending major axis parallel to the MFC (Fig. 9b). We therefore argue that this thickness change on the eastern Finnmark Platform is the result of syntectonic sediment deposition in the hanging wall of a NE–SW-trending, SE-dipping fault antithetic to the MFC (Fig. 5a and b). We suggest that the half-ellipsoid shape of the thickened upper Carboniferous sedimentary deposits on the eastern Finnmark Platform reflects a large offset near the center of the SE-dipping fault and decreasing vertical throw towards the fault tips, a feature characterizing syn-sedimentary, rift-related normal faults (Fig. 9b).

By contrast, depositional sediment wedges may occur on the eastern Finnmark Platform as well, and they differ from fault-controlled thickness changes. One example is the ca. 600 m thick lower Carboniferous succession evidenced by shallow drilling between the Nordkinn Peninsula and Magerøya (see "star" symbol in Fig. 1; Bugge et al., 1995), which we reinterpreted as a prograding Carboniferous sedimentary system (Fig. 6d). The apparent thickening of the lower Carboniferous succession near the coast of NW Finnmark is more likely to be related to sedimentary processes during the formation of large clinoforms in a prograding sedimentary system (Fig. 6d) than to syn-tectonic deposition in the hanging wall of a NE–SW-trending, NW-dipping fault.

Fault-controlled thickness changes in Permian strata

In the southwesternmost Nordkapp basin and on the eastern and western Finnmark Platform, the Permian sedimentary succession is thin and shows a relatively constant thickness compared to the underlying Devonian–lower Carboniferous and upper Carboniferous successions (Figs. 5a–d and 9a–c). However, the base Asselian and base Triassic reflections marking the lower and upper boundary of the Permian succession show some offsets across the main segment of the MFC, WNW–ESE-trending segment of the TFFC and Rolvsøya fault, thus accounting for minor thickness variations in the Permian succession across these faults (Figs. 5c, d, and g and 9c). We interpret these small offsets and thickness variations as the product of minor faulting activity in the Permian and mild Mesozoic reactivation of these faults, thus suggesting that the main tectonic activity along these faults was essentially restricted to the Middle to Late Devonian–late Carboniferous (Fig. 5c, d, and g). Moreover, on the western and eastern Finnmark Platform, most brittle faults die out within the upper Carboniferous succession and only a few faults crosscut the Permian succession with a limited amount of offset (Fig. 5a, b, e, and f).

Fault-controlled thickness changes in Mesozoic–Cenozoic strata

Most faults observed within the late Paleozoic succession on the eastern and western Finnmark Platform and in the southwesternmost Nordkapp basin die out in the upper part of the succession before reaching the base Triassic reflection (Fig. 5). A few exceptions exist where the MFC and the WNW–ESE-trending segment of the TFFC show small offsets of Mesozoic sedimentary strata (Fig. 5c–g). The weak influence of these faults compared to offsets observed within late Paleozoic successions (Fig. 5c–g) suggests that at least some major faults were mildly reactivated in Mesozoic times but in general most brittle faults on the eastern and western Finnmark Platform and in the southwesternmost Nordkapp basin remained inactive after Carboniferous times.

4.2 Offshore aeromagnetic data

To better verify our 2-D interpretation of faults and basin architectures on the Finnmark Platform and in the southwesternmost Nordkapp basin, we compare and tie our results using high-resolution offshore aeromagnetic data from Gernigon et al. (2014; Fig. 4). Aeromagnetic anomalies, when combined with seismic interpretation, may provide useful results allowing the identification of brittle faults and offset patterns (see Indrevær et al., 2013).

On the eastern Finnmark Platform, offshore aeromagnetic data (Fig. 4; Gernigon et al., 2014) show multiple narrow, NNE–SSW-trending, positive aeromagnetic anomalies that bend into NW–SE and NNW–SSE orientations near the center of the Nordkapp Basin, which Gernigon et al. (2014) interpreted as arc-shaped prolongations of Caledonian nappes. A more detailed analysis of these aeromagnetic data reveals a set of triangular to rhomboidal negative aeromagnetic anomalies, the largest of which was observed northeast of the island of Magerøya (dashed white lines in Fig. 4). This highly negative anomaly is bounded to the northeast and to the northwest by narrow, linear, NNE–SSW- to NE–SW-trending, positive aeromagnetic anomalies (dashed white lines in Fig. 4). On seismic data, the locations of these linear, positive aeromagnetic anomalies coincide with a SE-dipping normal fault for the northwestern anomaly, and the NW-dipping, zigzag-shaped LVF for the southeastern anomaly (see black arrows in Fig. 5a and b). These two faults bound a triangular-shaped basin filled up with thickened Carboniferous sedimentary deposits (see Fig. 5a and b), the shape and extent of which mimic those of the triangular negative anomaly observed on aeromagnetic data northeast of Magerøya (Fig. 4). Such triangular-shaped, negative aeromagnetic anomalies may thus be indicators of offshore Carboniferous sedimentary basins.

Similarly, on the western Finnmark Platform, a large NE–SW-trending, linear positive aeromagnetic anomaly is observed in the footwall of the TFFC (dotted white lines in Fig. 4), where it extends northeastwards into the footwall of the Rolvsøya fault (Fig. 4). This NE–SW-trending, positive aeromagnetic anomaly coincides with a NE–SW-trending basement ridge in the footwall of the TFFC on the western Finnmark Platform and with the location of an ENE–WSW-trending basement ridge in the footwall of the Rolvsøya fault (Figs. 1 and 5c–f). We interpret this positive anomaly to highlight a significant compositional difference between highly magnetic basement rocks in NE–SW- and ENE–WSW-trending basement ridges and poorly magnetic, adjacent basement rocks on the western Finnmark Platform and in the southwesternmost Nordkapp basin (Figs. 1, 4, and 5c–f).

5 Discussion

Our regional and detailed seismic studies of basin-bounding faults such as the TFFC, MFC, Rolvsøya fault, and TKFZ on the Finnmark Platform and adjacent southwesternmost Nordkapp basin show multiple links and interactions. We focus the discussion on the interaction of these faults and associated minor faults on Late Devonian–Carboniferous (half-)graben basins. We specifically discuss how deep-seated ductile Caledonian shear zones, i.e., the Sørøya-Ingøya shear zone and basement ridges may have been exhumed and thus enabled to control post-Caledonian brittle faulting and formation of Late Devonian–Carboniferous basins as collapse basins. In combination, the structural architecture, timing of faulting, and fault-controlled thickness variations in the Finnmark Platform and in the southwesternmost Nordkapp basin provide the framework to discuss the evolution of the SW Barents Sea margin from the Middle to Late Devonian to the Permian.

5.1 Interaction of the main segment of the Måsøy Fault Complex with the Sørøya-Ingøya shear zone

The linear, NE–SW-trending geometry of the main segment of the MFC in map view (Figs. 1 and 2) strongly differs from the dominant ENE–WSW- to NNE–SSW-trending, zigzag pattern typically observed for post-Caledonian faults in mid-Norway (Blystad et al., 1995), Lofoten–Vesterålen (Bergh et al., 2007; Eig, 2008; Hansen et al., 2012), western Troms (Indrevær et al., 2013), and NW Finnmark (Koehl et al., 2018). Notably, the anomalously linear segment of the MFC trends fully parallel to and merges into high-amplitude, NW-dipping seismic reflections of the SISZ on the Finnmark Platform and in the southwesternmost Nordkapp basin (Fig. 5c–f). This obvious merging of the main segment of the MFC into the basement-seated SISZ (Fig. 5c–f) suggests it formed as a brittle splay fault along an inverted portion of the shear zone, likely during the collapse of the Caledonides in the Middle to Late Devonian (Gudlaugsson et al., 1998). We suggest a similar interpretation for the Rolvsøya fault, which also flattens and merges into a bowed portion of the SISZ (Fig. 5c and d), and for the northwest-boundary fault of the Devonian graben on the western Finnmark Platform merging into a minor, SE-dipping shear zone (Figs. 5e and 6b and c). These faults are thought to have remained active through the late Carboniferous as suggested by potential syn-tectonic sediment thickening within the upper Carboniferous succession (Fig. 9b) but most likely ceased before the Permian as supported by the relatively constant thickness of Permian sedimentary strata throughout the study area (Fig. 9c).

By analogy, in the North Sea, Phillips et al. (2016) successfully tied the southernmost onshore occurrence of the Karmøy Shear Zone, a major Caledonian shear zone, to a thick seismic unit made up of subparallel, high-amplitude reflections similar to those ascribed to the SISZ in the footwall of the main segment of the MFC (Fig. 5d–f). Phillips et al. (2016) argue that the Åsta Fault, a large N–S-trending, W-dipping, post-Caledonian fault in the North Sea, formed during a phase of extensional reactivation of the Karmøy Shear Zone. Similarly, in western Norway, Wilks and Cuthbert (1994) proposed that the Hornelen Basin formed along a brittle fault that splayed upwards from the Nordfjord-Sogn Detachment Zone during Middle Devonian late-orogenic extension.

5.2 Formation of the WNW–ESE-trending fault segment of the Troms–Finnmark Fault Complex as a hard-linked accommodation cross fault

Our data (Fig. 9a and b) show abrupt fault-controlled thickening of the Devonian–lower Carboniferous and upper Carboniferous sedimentary successions just northeast of the WNW–ESE-trending segment of the TFFC into the southwesternmost Nordkapp basin. On the western Finnmark Platform, potential Devonian sedimentary rocks are truncated upwards by the mid-Carboniferous reflection (Fig. 5e and f). We propose that the absence of high-amplitude, coal-bearing sedimentary deposits of the Billefjorden Group (lower Carboniferous) on the western Finnmark Platform is related to a major episode of eustatic sea-level fall in the Serpukhovian (Saunders and Ramsbottom, 1986), which may have contributed to exposing lower Carboniferous sedimentary rocks in this area to coastal erosion. Hence, part of the thickening of the Devonian–lower Carboniferous succession across the WNW–ESE-trending segment of the TFFC might be related to extensive erosion of the western Finnmark Platform in mid-Carboniferous times. In addition, the clear deepening (plunge) to the northeast of the spoon-shaped trough formed by the three-dimensionally folded and bowed geometry of the SISZ (Fig. 7) suggests that the thickening of Devonian–lower Carboniferous sedimentary strata into the southwesternmost Nordkapp basin (Fig. 9a) is also partly controlled by the shape and attitudes of the underlying SISZ. Finally, the thickened sediment depocenter observed in the southwesternmost Nordkapp basin at the intersection of the TFFC and the MFC (Fig. 9a and b) is at least partly related to syn-sedimentary normal faulting along the WNW–ESE-trending segment of the TFFC and along the main segment of the MFC. This most likely indicates that the TFFC and the MFC had already merged and acted as a single fault zone during sediment deposition in the southwesternmost Nordkapp basin from the end of the Serpukhovian and potentially from Devonian times. We propose that the WNW–ESE-trending fault segment of the TFFC acted as an accommodation cross fault, as defined in Sengör (1987), that transferred displacement between the NNE–SSW-trending segment of the TFFC and the main segment of the MFC, defining a step synthetic with the deepening direction of the spoon-shaped trough formed by the geometry of the SISZ (Fig. 7). This interpretation is based on the dominant dip-slip kinematic of the WNW–ESE-trending segment of the TFFC and on its subparallel strike to the dominant WNW–ESE-trending extension direction inferred along the SW Barents Sea margin during late Paleozoic times (Bergh et al., 2007; Eig and Bergh, 2011; Hansen and Bergh, 2012). Further, we infer that the strike and location of the WNW–ESE-trending segment of the TFFC was controlled by the geometry of the underlying SISZ (see below), which dips gently to the northeast on the western Finnmark Platform and in the southwesternmost Nordkapp basin and may therefore have favored the formation of a NE-dipping fault at this location (Figs. 5g and 7).

Alternatively, Lea (2016) proposed that the WNW–ESE-trending fault segment of the TFFC corresponds to a breached relay ramp fault between the NNE–SSW-trending fault segment of the TFFC and the MFC. However, this model implies that this portion of the TFFC would have accommodated significantly less displacement than the two faults it links (i.e., the NNE–SSW-trending segment of the TFFC and the MFC), which is clearly not the case. The offset of the mid-Carboniferous reflection and the thickness increase in both the Middle to Upper Devonian–lower Carboniferous and upper Carboniferous sedimentary successions across the WNW–ESE-trending segment of the TFFC are comparable to the offset and thickness increase observed across the main segment of the MFC (Fig. 5c, d, and g), and the TFFC and MFC seem to have evolved synchronously in the late Paleozoic.

5.3 Devonian collapse basins on the western Finnmark Platform (Gjesvær Low) and in the southwesternmost Nordkapp basin

Devonian sedimentary rocks in the SW Barents Sea may exist on the western Finnmark Platform and in the southwesternmost Nordkapp basin (Fig. 5c–g). The most probable occurrence is on the western Finnmark Platform, in the hanging wall of the main segment of the MFC (Figs. 1, 5e). The presumed Devonian seismic unit corresponds to a suite of low-amplitude reflections crosscut by a few moderate-amplitude reflections that dip gently to the northwest (Fig. 5e). The main argument for a Devonian succession is that these reflections are remarkably different from the typical seismic patterns observed for lower Carboniferous sedimentary deposits and basement rocks. Lower Carboniferous sedimentary deposits of the Billefjorden Group are characterized by high-amplitude reflections produced by coal-bearing sedimentary rocks (Figs. 5a and b and 6a), while basement rocks are mostly associated with thick packages of chaotic seismic reflections (Fig. 6a–c) and thick layers of moderate- to high-amplitude, subparallel seismic reflections that we interpreted as basement-seated shear zones (e.g., the SISZ; Fig. 6f). Another argument in favor of Devonian sedimentary deposits is the presence of a NE–SW-trending gravimetric low on the western Finnmark Platform: the Gjesvær Low (Johansen et al., 1994). Devonian sedimentary rocks in Svalbard show an average density of ca. 2.4 $\mathrm{g\,cm^{-3}}$ associated with depths of 0–8 km (i.e., average depth of 4 km; Manby and Lyberis, 1992), which is less dense than metamorphosed Caledonian rocks (2.6–3.0 $\mathrm{g\,cm^{-3}}$) and Carboniferous sedimentary rocks on the western Finnmark Platform ($<2.5\ \mathrm{g\,cm^{-3}}$; Johansen et al., 1994). However, taking into account the effect of burial up to a depth of 5–6 km on the Finnmark Platform (Johansen et al., 1994) and the resulting density increase for Devonian sedimentary deposits with an approximate rate of ca. 0.15 $\mathrm{g\,cm^{-3}\,km^{-1}}$ (see "all rocks density-depth gradient" in Table 3 in Maxant, 1980), Devonian sedimentary rocks on the Finnmark Platform may reach densities of 2.55–2.7 $\mathrm{g\,cm^{-3}}$. Thus, the occurrence of the Gjesvær low can be explained by the presence of intermediately dense Devonian sedimentary rocks below the mid-Carboniferous reflection (Figs. 5e and 6b and c). This is in accordance with density variations and related estimates of Johansen et al. (1994) in the Gjesvær Low as well.

In addition, the discrete, moderate-amplitude, NW-dipping reflections observed within the presumed Devonian sedimentary strata on the western Finnmark Platform may represent syn-tectonic sedimentary growth strata (Figs. 5e and 6b and c). These strata are located above and thickened against arcuate, high-amplitude reflections that we interpreted as a major erosional unconformity truncating SE-dipping Caledonian basement shear zones subparallel to the SISZ (dotted yellow lines in Fig. 6b and c). We consider the wedge-shaped Devonian sedimentary rocks on the western Finnmark Platform to have been deposited in late- to post-Caledonian extensional basins due to reactivation of a set of partly eroded, exhumed, SE-dipping Caledonian shear zones (dotted yellow lines in Figs. 5e and 6b and c). In mid-Norway, Braathen et al. (2000) reported a similar setting of Middle Devonian sedimentary basins located above Caledonian shear zones and folded nappe stack and proposed that these formed during extensional reactivation of the shear zones. Such a model is further supported by the geometry of Devonian sedimentary growth strata on the western Finnmark Platform, which is similar to the geometry of highly tilted sedimentary strata in Middle Devonian basins in western Norway (see Séranne and Seguret, 1987; Séranne et al., 1989; Wilks and Cuthbert, 1994; Osmundsen and Andersen, 2001). Moreover, the Devonian sedimentary basins on the western Finnmark Platform (Figs. 5e and 6b and c) and in the southwesternmost Nordkapp basin (Fig. 5c and d) define NE–SW-trending graben structures with <50 km wide sizes comparable to those of the Middle Devonian Hornelen, Kvamsheten, and Solund basins in western Norway (Séranne and Seguret, 1987; Osmundsen and Andersen, 2001).

In the southwesternmost Nordkapp basin, the presence of Middle to Upper Devonian sedimentary rocks is more speculative and is mostly based on the maximum thickness of lower Carboniferous sedimentary deposits registered in the SW Barents Sea, which is ca. 600 m thick on the eastern Finnmark Platform (Bugge et al., 1995; Fig. 6d). Assuming a seismic velocity $<6\ \mathrm{km.s^{-1}}$ for lower Carboniferous coal-bearing sedimentary deposits, a thickness of 600 m would account for only part (maximum 0.2 s) of the 2–2.5 s thick (TWT) seismic unit observed below the mid-Carboniferous reflection in the southwesternmost Nordkapp basin (Fig. 5c and d). If basement rocks were present at the base of the southwesternmost Nordkapp basin, they would most likely produce a seismic reflection pattern similar to that of the subparallel, high-amplitude reflections of the underlying SISZ (Fig. 5c and d) or potentially form an unconformity to the overlying late Paleozoic sedimentary rocks. We therefore believe that the southwesternmost Nordkapp basin, which is bounded below by the SISZ, giving the basin a peculiar "U" shape in cross-section (Fig. 5c and d), is composed of thick Middle to Upper Devonian sedimentary deposits overlain by lower Carboniferous sedimentary strata below the mid-Carboniferous reflection. Based on the brittle extensional reactivation, bowed geometry, and controlling effect of the basement-seated SISZ (Fig. 7), we suggest deposition of Devonian sedimentary rocks within a late- to post-Caledonian, spoon-shaped collapse basin formed along inverted portions of the Caledonian SISZ (Fig. 5c–f), thus representing analogs to Middle Devonian collapse basins in western and mid-Norway (Séranne et al., 1989; Wilks and Cuthbert, 1994).

5.4 Formation of NE–SW- to ENE–WSW-trending basement ridges as exhumed metamorphic core complexes

We have argued for an upward-bowed seismic geometry of the SISZ (Figs. 5c–f and 7) into which major fault complexes such as the TFFC, the MFC, and the Rolvsøya fault merge (Fig. 5c–f). In map view (Fig. 7) and cross section (Fig. 5c–f), the bowed geometry of the SISZ defines two ENE–WSW- to NE–SW-trending ridges of basement rocks on the northwestern flanks of presumed Devonian basins. These basement ridges correlate well by displaying positive gravimetric (Fig. 5 in Olesen et al., 2010, and Fig. 5 in Gernigon et al., 2014) and positive aeromagnetic anomalies (Fig. 4; Gernigon et al., 2014) that suggest these ridges are made of basement lithologies significantly different from adjacent basement rocks on the western Finnmark Platform and in the southwesternmost Nordkapp basin. These basement ridges seem to align with positive gravimetric anomalies coinciding with the NE–SW-trending Norsel High (Gabrielsen et al., 1990) along the northwestern flank of the Nordkapp Basin in the northeast. Farther southwest, these basement ridges coincide with the NE–SW-trending West Troms Basement Complex in western Troms (Bergh et al., 2010; Fig. 1) and the NE–SW-trending Lofoten Ridge in Lofoten–Vesterålen (Bergh et al., 2007; Fig. 1), among which at least the Lofoten Ridge was exhumed as a metamorphic core complex (Klein and Steltenpohl, 1999; Klein et al., 1999; Steltenpohl et al., 2004, 2011) along inverted Caledonian shear zones such as the Eidsfjord and Fiskefjord shear zones (Steltenpohl et al., 2011). By comparison, the SISZ seems to coincide with positive aeromagnetic anomalies on the western Finnmark Platform that follow the trace of the MFC (Indrevær et al., 2013) and continue past the southwestern fault tip of the MFC (Gernigon and Brönner, 2012). The aeromagnetic anomalies visible on the dataset Gernigon and Brönner (2012) appear to line up with aeromagnetic anomalies on the island of Vannøya in the northeasternmost part of the West Troms Basement Complex (Fig. 1), and these onshore anomalies correlate with NE–SW-trending, SE-dipping basement shear zones that were reactivated as extensional brittle faults (H.-K. Paulsen et al., personal communication, 2017).

These data indicate that SE-dipping portions of the SISZ propagated southwest of the MFC fault tip on the western Finnmark Platform and possibly merged with a suite of NE–SW-trending, SE-dipping shear zones on Vannøya in western Troms. As a consequence, the basement ridges observed on the western Finnmark Platform and along the northern flank of the southwesternmost Nordkapp basin may have formed as part of a large metamorphic core complex, which included the Lofoten Ridge, the West Troms Basement Complex, and possibly also the Norsel High (Fig. 1), exhumed along inverted Caledonian shear zones, e.g., SISZ on the SW Barents shelf and the analogous Eidsfjord and Fiskefjord shear zones in Lofoten–Vesterålen (Steltenpohl et al., 2011).

The timing, nature of uplift and processes of core complex exhumation can be inferred from thickness variations in the SISZ in cross section, for example, thickest in the footwall of the MFC and TFFC and thinnest below these two fault complexes (Fig. 5f). We link these thickness variations to excisement and incisement processes (see Lister and Davis, 1989) along the SISZ during core complex exhumation, after the embrittlement of the SISZ (Fig. 10). A model of Devonian late- to post-orogenic extension is proposed, when inversion of the SISZ as a low-angle, top–NW extensional detachment and extensive erosion of the Caledonides contributed to crustal thinning (Fig. 10a). Rapid thinning through extension and erosion above the upper part of the SISZ may have triggered early exhumation of basement rocks on the western Finnmark Platform and along the northern flank of the southwesternmost Nordkapp basin (Fig. 10a), causing the upper part of the SISZ to bow upwards (Fig. 10b). Continued crustal extension and continental erosion further enhanced exhumation of basement rocks below the upper part of the SISZ, leading the bowed portion of the SISZ to become unsuitably oriented to accommodate top–NW extensional displacement and thus become inactive (dashed black line in Fig. 10c). Further extension likely triggered upward splaying of the SISZ into its hanging wall, becoming suitable again to accommodate top–NW extension displacement (Fig. 10c). This upward splaying process is referred to as excisement by Lister and Davis (1989) and we tentatively apply this process to explain the observed thickening of the SISZ in the footwall of the MFC (Fig. 5f). Further extension- or erosion-related crustal thinning along the SISZ may have initiated exhumation of basement rocks along progressively deeper parts of the SISZ (Fig. 10b and c), causing bend-up of the SISZ at deeper crustal levels (Fig. 10d). Extreme bowing of lower portions of the SISZ led to an opposite top–SE transport direction on the western Finnmark Platform and in the southwesternmost Nordkapp basin (Fig. 10d), which contributed to the exhumation of NE–SW- to ENE–WSW-trending ridges of basement rocks in the footwall of the NNE–SSW-trending segment of the TFFC (Fig. 5e and f) and in the footwall of the Rolvsøya fault (Fig. 5c and d), thus forming a large spoon-shaped trough where Devonian sedimentary rocks deposited (Figs. 7 and 10d and e). Incisement (downward splaying; see Lister and Davis, 1989) may have occurred below the basement ridges during continued top–NW extension along the SISZ (Fig. 10e) and possibly contributed to thickening of the SISZ in the footwall of the TFFC (Fig. 5e and f) and of the Rolvsøya fault (Fig. 5c and d), resulting in the current geometry of the SISZ (Fig. 10f).

By comparison, in northeastern Greenland, Sartini-Rideout et al. (2006) and Hallett et al. (2014) proposed that ultra-high-pressure basement rocks were exhumed along large, mylonitic, Caledonian shear zones in Late Devonian–early Carboniferous times (ca. 370–340 Ma). The study of Sartini-Rideout et al. (2006) also shows that the last stages of exhumation were accommodated by steep, brittle normal faults that strike parallel to major Caledonian shear zones, i.e., similar to the main segment of the MFC striking parallel to the SISZ along the SW Barents Sea margin (Fig. 1). In addition, results from sediment provenance and geochronological studies by McClelland et al. (2016) in Carboniferous basins in northeastern Greenland showed that the exhumation of ultra-high-pressure basement rocks as elongated ridges could have formed a regional Serpukhovian erosional unconformity, contemporaneous with the mid-Carboniferous (Serpukhovian?) unconformity observed on the eastern Finnmark Platform (Figs. 5a and b and 6a; Bugge et al., 1995) and on the western Finnmark Platform (Figs. 5e–g and 6b and c) and in agreement with eustatic sea-level fluctuations at that time (Saunders and Ramsbottom, 1986). In late Paleozoic times, the northeastern Greenland margin was located

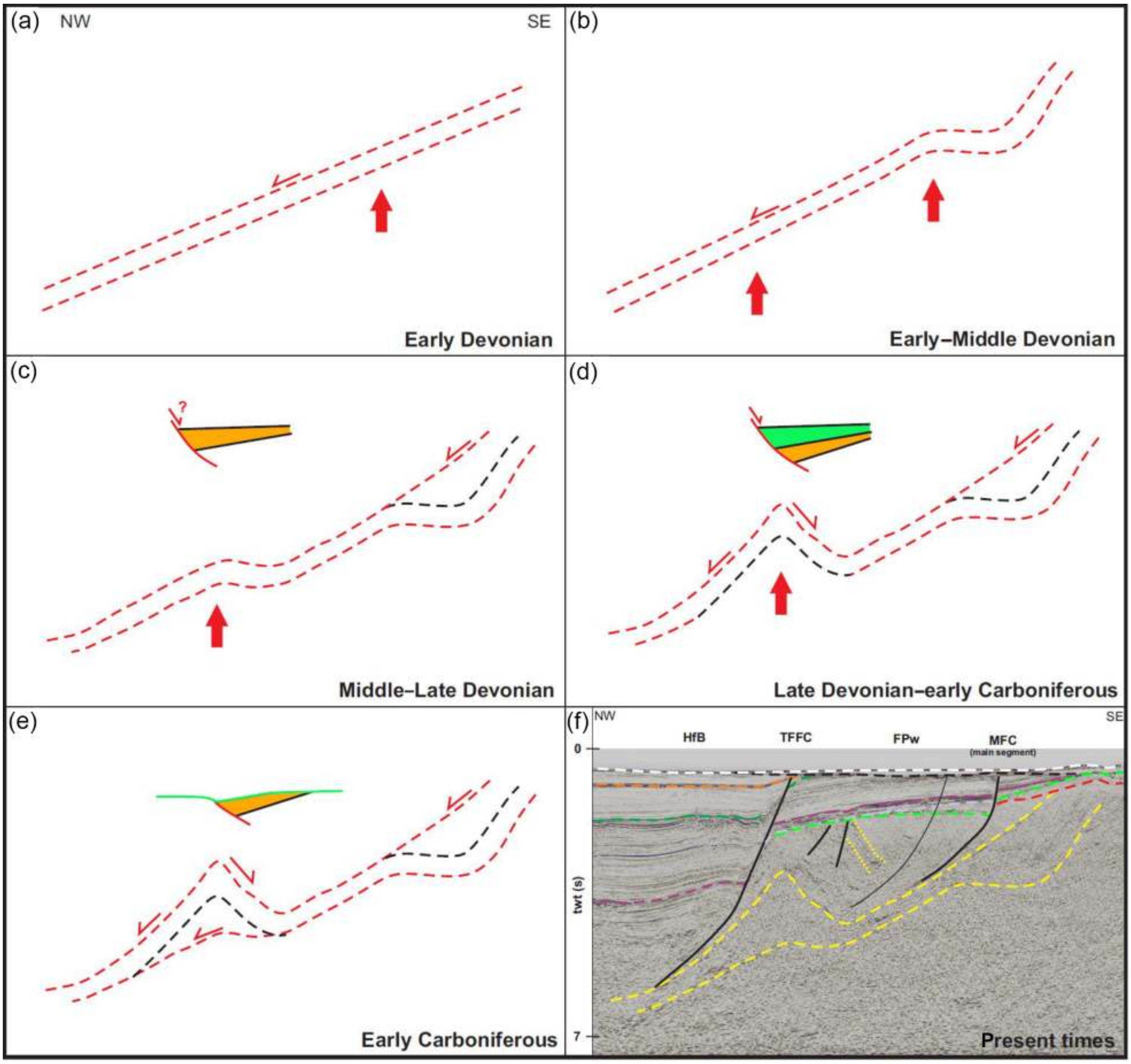

Figure 10. Evolutionary model explaining thickness variations along the SISZ. Note that the timing of **(a)** to **(e)** is tentative. Dashed red lines in panels **(a)** to **(e)** correspond to tectonically active portions of the SISZ whereas dashed black lines show inactive portions. Red lines in panels **(c)**, **(d)**, and **(e)** show presumed normal faults. Thick vertical red arrows indicate exhumation of basement rocks along the SISZ. The model is adapted to the geometry of the SISZ observed below the western Finnmark Platform (see panel **f**); **(a)** extensional reactivation (thin red arrow) of the SISZ in Early Devonian times. Rapid crustal thinning and possible erosion along the upper part of the SISZ triggers exhumation of basement rocks near the coast of NW Finnmark (thick red arrow); **(b)** in the Early–Middle Devonian, continued extension and erosion further thin the crust and exhume basement rocks in the footwall of the SISZ, leading the upper part of the SISZ to bow. Incremental crustal thinning due to continued extensional reactivation of the SISZ and continental erosion triggers exhumation of basement rocks along lower portions of the SISZ (left-hand side, thick red arrow); **(c)** in Middle to Late Devonian times, bowed portions of the SISZ become inactive and excisement (i.e., upwards splaying; see Lister and Davis, 1989) of the SISZ into its hanging wall leads to thickening of the upper portion of the SISZ. Continued extension and erosion (i.e., crustal thinning) trigger bending of the lower part of the SISZ (thick red arrow) above which brittle normal faults may have formed and localized the deposition of Devonian sedimentary deposits (orange); **(d)** further exhumation of basement rocks along lower portions of the SISZ in the Late Devonian–early Carboniferous leads to extreme bending of the SISZ, to antithetic top–SE extensional faulting, and to early Carboniferous syn-tectonic sedimentation (green); **(e)** towards the end of the early Carboniferous, the lower portion of the SISZ is thickened due to incisement (i.e., downward splaying; see Lister and Davis, 1989) of the SISZ into bow-shaped portions in its footwall. Core complex exhumation ceased in the Serpukhovian and a major sea-level fall exposed the Finnmark Platform to continental erosion (green line representing the mid-Carboniferous reflection); **(f)** present seismic expression of thickness variations along the SISZ (dashed yellow) on the western Finnmark Platform. See Fig. 5 for seismic reflections color schemes.

close to its conjugate counterpart of the SW Barents Sea margin and these two areas were most likely subjected to similar regional stresses and closely related sea-level fluctuations. Therefore, we suggest that the mid-Carboniferous unconformity reflection observed in the SW Barents Sea (see Figs. 5 and 6a–c; Bugge et al., 1995), formed as a response to major eustatic sea-level fall in the early Serpukhovian (Saunders and Ramsbottom, 1986) and due to large-scale exhumation of basement rocks in Late Devonian–early Carboniferous times. Exhumation occurred along inverted Caledonian shear zones (e.g., SISZ) and brittle splay faults such as the main segment of the MFC, the NNE–SSW-trending segment of the TFFC, the Rolvsøya fault and the NNE–SSW-trending, SE-dipping fault that bounds potential Devonian sedimentary strata on the western Finnmark Platform (Fig. 5). The timing of exhumation is constrained to the end of the Serpukhovian based on deposition of thick alluvial and shallow marine upper Carboniferous sedimentary deposits of the Gipsdalen Group (Fig. 3) on top of the mid-Carboniferous unconformity (Fig. 5) during an eustatic sea-level rise near the end of the Serpukhovian in the SW Barents Sea (Saunders and Ramsbottom, 1986).

In Lofoten–Vesterålen, Steltenpohl et al. (2004) inferred a Late Devonian age for the exhumation of metamorphic core complexes and this age was refined by recent ^{40}Ar/^{39}Ar isotopic results (Steltenpohl et al., 2011), which constrained extensional reactivation of the Eidsfjord shear zone to the Early Devonian. In the SW Barents Sea, much work is needed to better constrain the timing of late- to post-Caledonian extension and collapse basin formation. Nonetheless, we believe that the Early Devonian age obtained in Lofoten–Vesterålen (Steltenpohl et al., 2011) represents a reasonable estimate for the onset of crustal thinning in the SW Barents Sea (Fig. 10a–c). Additionally, Late Devonian–early Carboniferous timing of exhumation for basement rocks in northeastern Greenland and formation of a regional mid-Carboniferous (Serpukhovian) unconformity (Sartini-Rideout et al., 2006; Hallett et al., 2014; McClelland et al., 2016) corresponds to a realistic approximation for the final stages of late- to post Caledonian extension, ending with the formation of Devonian–Carboniferous collapse basins along exhumed NE–SW- to ENE–WSW-trending basement ridges (Fig. 10d–f).

5.5 Interaction of the Trollfjorden–Komagelva Fault Zone with the Troms–Finnmark and Måsøy fault complexes

The prolongation of the TKFZ from onshore areas in eastern Finnmark to offshore areas of the SW Barents Sea has been a matter of debate. Most studies tend to connect the onshore TKFZ with the offshore WNW–ESE-trending segment of the TFFC (Gabrielsen, 1984; Gabrielsen and Færseth, 1989; Roberts et al., 2011; Bergø, 2016; Lea, 2016). Our data, however, suggest that the TKFZ dies out near the coast of NW Finnmark (present contribution and Koehl et al., 2018), and in this section we review and discuss new evidence obtained from the interpretation of offshore seismic and aeromagnetic data.

First, the TKFZ described onshore eastern Finnmark as a major subvertical fault that accommodated dominantly strike-slip movement (Roberts, 1972; Rice, 2013). Farther west, the TKFZ crops out onshore Magerøya, where it is made of numerous, high-frequency, subparallel, subvertical, WNW–ESE-trending faults and fractures that accommodated at least small-scale post-Caledonian strike-slip to oblique-slip displacement (Koehl et al., 2018). By contrast, seismic interpretation of the WNW–ESE-trending segment of the TFFC shows that this fault exhibits a typical, high-angle (ca. 70°) normal fault geometry and accommodated significant amount of post-Caledonian normal dip-slip displacement (Fig. 5g), thus contrasting significantly with the geometry of the TKFZ. Second, the imaginary prolongation of the TKFZ from the island of Magerøya to the WNW, onto the Finnmark Platform, would crosscut the western Finnmark Platform nearly 23 km southwest of the observed trace of the WNW–ESE-trending segment of the TFFC (Fig. 5g). This represents a significant mismatch that is far too important to represent minor dextral strike-slip offset of the TKFZ across the main fault segment of the MFC, which dominantly accommodated normal dip-slip motions (Fig. 5c–f). Third, the interpretation of 3-D seismic data at the intersection of the MFC and TFFC reveals that the footwall of the MFC is largely intact and seismically unaffected by brittle faults (Fig. 8). There is no evidence of intense fracturing as typically observed along the TKFZ on Magerøya (Koehl et al., 2018). We therefore believe that the TKFZ and the WNW–ESE-trending fault segment of the TFFC represent two distinct faults. This suggests that the TKFZ dies out instead of propagating onto the Finnmark Platform. This is also supported by the absence of WNW–ESE-trending faults offshore (Figs. 1 and 2). For example, the Austhavet fault previously interpreted near the coast of Finnmark (Townsend, 1987b; Lippard and Roberts, 1987; Roberts et al., 2011) was reinterpreted as seismic artifacts related to the Djuprenna trough, a large glacial trough that trends parallel to the northeastern coast of Finnmark (Ottesen et al., 2008; Rise et al., 2015). This reinterpretation is supported by shallow drillings on the eastern Finnmark Platform, which show no sign of fault-related offset in this part of the eastern Finnmark Platform (see Fig. 4 in Bugge et al., 1995). Similarly, our mapping and regional analysis of brittle faults on the Finnmark Platform show very few occurrences of WNW–ESE-trending faults (Figs. 1 and 2).

Onshore studies (Koehl et al., 2018) show an increased number of large-scale WNW–ESE-trending fault segments and splays along the TKFZ, varying from a single-segment fault on the Varanger Peninsula in eastern Finnmark to multiple segments near Magerøya (Fig. 1), suggesting that the island of Magerøya is located near the fault-tip process zone (Shipton and Cowie, 2003; Braathen et al., 2013) of the

TKFZ and that the TKFZ therefore dies out to the west before reaching the Finnmark Platform and the southwesternmost Nordkapp basin (Fig. 1; Koehl et al., 2018). Nearby Magerøya and the Nordkinn Peninsula, high-resolution aeromagnetic data reveal the presence of highly magnetic dolerite dykes along WNW–ESE-trending fault segments of the TKFZ (Roberts et al., 1991; Nasuti et al., 2015; Koehl et al., 2018). These narrow, positive aeromagnetic anomalies also die out westwards (Gernigon and Brönner, 2012; Gernigon et al., 2014), therefore supporting that the dolerite dykes and thus the TKFZ die out before reaching the Finnmark Platform (Koehl et al., 2018).

We explore an alternative model to the fault-tip process zone of Koehl et al. (2018) in which we argue that the Precambrian orogen-parallel, WNW–ESE-trending TKFZ was not suitably oriented to be reactivated as a major thrust or strike-slip fault during the Caledonian Orogeny. Our data indicate that if the TKFZ extended farther west prior to the onset of Caledonian deformation, it was certainly truncated and decapitated by large-scale top–SE movement along the SISZ and associated NE–SW-trending Caledonian thrusts and shear zones. This is supported by the dominant top–SE transport direction inferred along Caledonian thrusts in NW Finnmark (Townsend, 1987a). Thus, we propose that the western continuation of the TKFZ below the Finnmark Platform may have been thrusted southeastwards along the SISZ and is now eroded. However, if the TKFZ ever extended westwards, portions of the fault might be preserved in offshore basement highs such as the Loppa and Veslemøy highs (Fig. 1). More work is needed on this hypothesis, but a possible insight is the recent observation of subvertical WNW–ESE-trending brittle faults analog to the TKFZ in basement rocks of the Veslemøy High (Kairanov et al., 2016).

5.6 Late Paleozoic evolution of the SW Barents Sea margin

Based on the seismic data and discussions from previous chapters we now address the tectonic evolution of the Finnmark Platform and the southwesternmost Nordkapp basin in the late Paleozoic (Fig. 11). The main structural element discussed in our model is the SISZ and we link its influence on (i) the development of the southwesternmost Nordkapp basin, (ii) the geometry of the TFFC and MFC, (iii) the deposition of Middle to Upper Devonian sedimentary rocks in the southwesternmost Nordkapp basin and on the western Finnmark Platform, (iv) transfer faults such as the TKFZ, and (v) the deposition of syn-tectonic sedimentary wedges along steep normal faults bounding triangular-shaped Carboniferous basins.

The trend and dominant northwestern dip of the SISZ (Figs. 5c–g and 6f) suggest that it formed as a large thrust that accommodated top–SE tectonic transport during the Caledonian Orogeny. The SISZ has a bow-shaped, three-dimensionally folded geometry that coincides with basement ridges in the footwall of the TFFC and of the Rolvsøya fault (Figs. 5c–f, 7, and 11). We propose that the SISZ and potential other Caledonian shear zones along the SW Barents Sea margin were inverted as low-angle extensional shear zones during late- to post-Caledonian orogenic extension and subsequent collapse. This is based on analog examples in northeastern Greenland (Sartini-Rideout et al., 2006; Hallett et al., 2014; McClelland et al., 2016), western Norway (Séranne and Seguret, 1987; Séranne et al., 1989; Wilks and Cuthbert, 1994; Osmundsen and Andersen, 2001), mid-Norway (Braathen et al., 2000), and Lofoten–Vesterålen (Klein and Steltenpohl, 1999; Klein et al., 1999; Steltenpohl et al., 2004, 2011; Osmundsen et al., 2005). Extensional reactivation of such ductile shear zones along the Barents Sea margin may have initiated in the Early Devonian, as in Lofoten–Vesterålen (Steltenpohl et al., 2011), through orogenic collapse dominated by top–NW movement along the SISZ.

Exhumation of the SISZ and underlying basement ridges as a metamorphic core complex was probably triggered by extensional reactivation of the SISZ combined with continental erosion, leading to crustal thinning. Reactivation of these exhumed basement ridges occurred by normal faulting along new, steep, brittle faults such as the main segment of the MFC and the NNE–SSW-trending fault segment of the TFFC (see Fig. 5f), likely due to incisement and excisement processes (Fig. 10; Lister and Davis, 1989). These processes also contributed to the progressive exhumation of ENE–WSW- and NE–SW-trending basement ridges along bowed portions of the SISZ (see Figs. 5c–f and 11a and b). We believe that these ridges were part of a larger-scale NE–SW-trending metamorphic core complex that included the Norsel High and the two basement ridges located in the footwall of the TFFC and the Rolvsøya fault. Farther south, this core complex may be linked to the West Troms Basement Complex (Bergh et al., 2010) and the Lofoten Ridge (Blystad et al., 1995; Fig. 11). Such a regional link is favored by the alignment of NE–SW-trending, positive gravimetric anomalies that characterize these ridges (Olesen et al., 2010; Gernigon et al., 2014). The timing of final core complex exhumation can be constrained to Middle to Late Devonian–early Carboniferous and possibly linked to the regional Serpukhovian unconformity on the Finnmark Platform (see Figs. 5a, b, e, f, and g and 6a–c; Bugge et al., 1995), in accordance with Sartini-Rideout et al. (2006) and Hallett et al. (2014) in northeastern Greenland.

The exhumation of basement ridges as metamorphic core complexes along the inverted SISZ and subsequent normal faulting along the MFC and TFFC created a deep, spoon-shaped topographic depression on the western Finnmark Platform and in the southwesternmost Nordkapp basin (Figs. 5c, d, e, and g, 6b and c, and 11a). These depressions were filled with thick Devonian clastic deposits analog to those observed in Middle Devonian collapse basins in western Norway (Séranne et al., 1989; Osmundsen and Andersen, 2001) and with lower Carboniferous coal-bearing and clas-

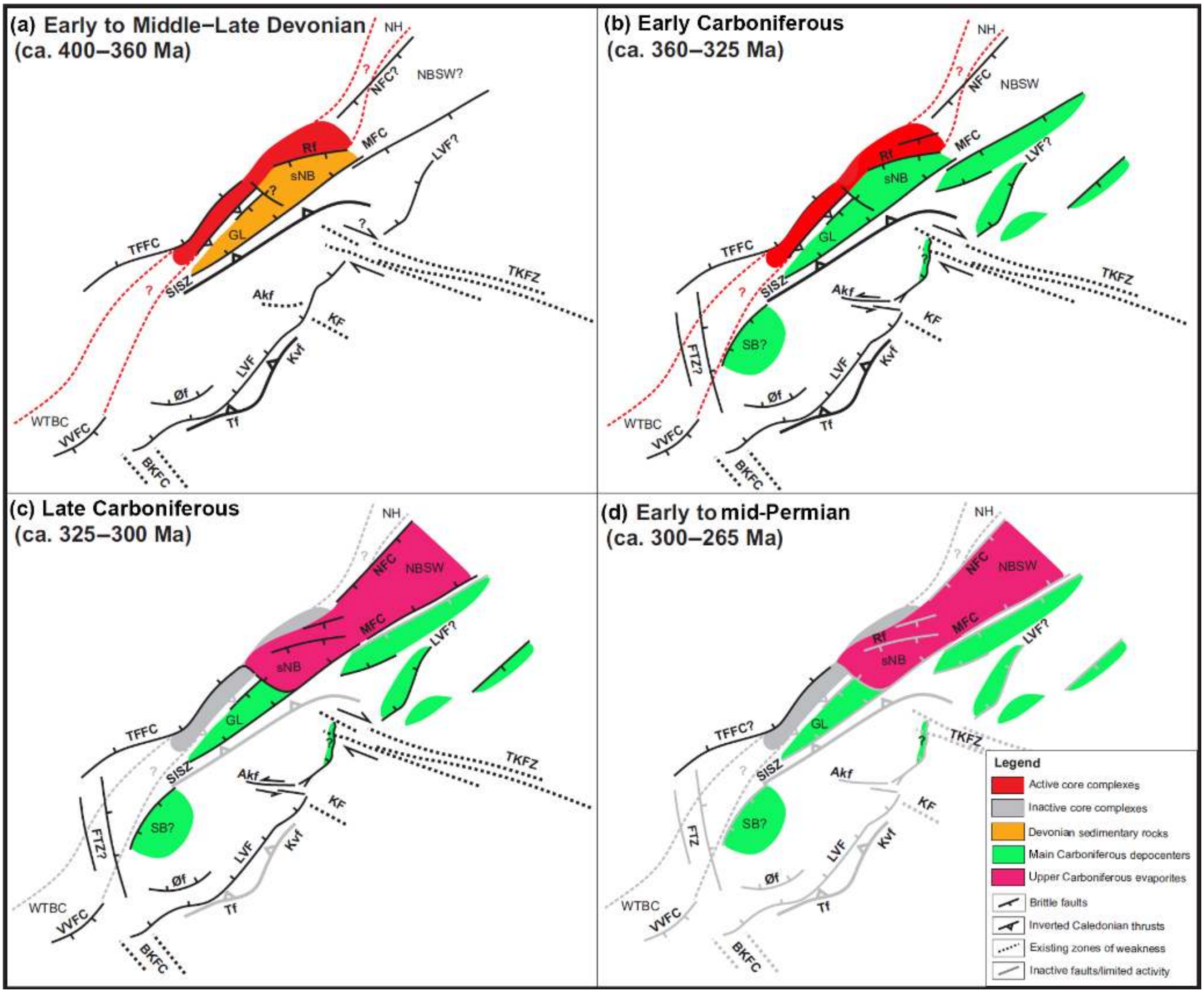

Figure 11. Map-view figures summarizing the late Paleozoic tectono-sedimentary evolution of the Finnmark Platform and southwesternmost Nordkapp basin (sNB). The tectonic evolution of onshore and nearshore faults in NW Finnmark is from Koehl et al. (2018). Abbreviations as in Fig. 1. **(a)** In the Early to Middle–Late Devonian, major Caledonian thrusts (e.g., SISZ) were inverted as low-angle extensional shear zones and exhumed metamorphic core complexes in the footwall of the TFFC and of the Rolvsøya fault. Thick Devonian sedimentary rocks were deposited within a spoon-shaped trough created by the geometry of the SISZ; **(b)** core complex exhumation continued through the early Carboniferous, though mostly accommodated by high-angle normal faults, which formed as brittle splays along Caledonian thrusts and shear zones (e.g., MFC, TFFC, Rolvsøya fault, and LVF). Core complex exhumation ceased by the end of the Serpukhovian and the WNW–ESE-trending segment of the TFFC formed as an accommodation cross fault that decoupled the western Finnmark Platform from the southwesternmost Nordkapp basin, thus contributing to the preservation of thick Devonian and lower Carboniferous sedimentary successions in the southwesternmost Nordkapp basin while these sedimentary rocks were almost completely eroded on the western Finnmark Platform. Minor graben and half-graben structures formed on the eastern Finnmark Platform. Precambrian WNW–ESE- to NNW–SSE-trending fault zones such as the TKFZ segmented the margin and acted as minor transfer faults that accommodated a limited amount of lateral displacement. Faulting along these faults ceased in the early Carboniferous; **(c)** in the late Carboniferous, inverted Caledonian thrusts and shear zones became inactive and were truncated by high-angle splay faults that accommodated the deposition of syn-tectonic sedimentary wedges on the eastern and western Finnmark Platform, and of thick, partly evaporitic deposits in the southwesternmost Nordkapp basin; **(d)** by the end of the Carboniferous, active brittle faulting came to a halt and the Finnmark Platform and the southwesternmost Nordkapp basin are believed to have remained tectonically quiet.

tic sedimentary rocks of the Billefjorden Group (Fig. 3) deposited unconformably above Devonian strata (see Fig. 11b). These collapse basins are also likely responsible for the gravimetric low observed on the western Finnmark Platform, the Gjesvær Low (Fig. 4).

On the western Finnmark Platform, the final stages of core complex exhumation and a major phase of eustatic sea-level fall in the Serpukhovian (Saunders and Ramsbottom, 1986) led to extensive erosion of Devonian and lower Carboniferous sedimentary rocks, therefore explaining the absence of lower Carboniferous sedimentary deposits and the erosional truncation of Devonian sedimentary strata along this part of the margin (Figs. 5e–g and 6b and c). On the eastern Finnmark Platform, lower Carboniferous sedimentary rocks are preserved as minor syn-tectonic sedimentary wedges within small triangular grabens and half grabens that correlate with aeromagnetic lows (dashed white line in Fig. 4). These grabens are bounded by zigzag-shaped, Late Devonian–Carboniferous normal faults such as the LVF (Figs. 5a and b and 6a; Koehl et al., 2018), which coincide with narrow, positive aeromagnetic anomalies (see Fig. 4 and black vertical arrows in Fig. 5a and b). In addition, triangular basins like the graben bounded by the LVF and the southwesternmost Nordkapp basin were partly offset and segmented by WNW–ESE-trending transfer faults that accommodated a small amount of strike-slip motion. Examples include the TKFZ in NW Finnmark, which may offset the LVF in a right-lateral fashion (Koehl et al., 2018), and accommodation cross faults (Sengör, 1987) that accommodated a large amount of orogen-parallel extension through normal dip-slip movement, e.g., the WNW–ESE-trending fault segment of the TFFC (Figs. 5g and 9a).

In the late Serpukhovian, a regional episode of eustatic sea-level rise (Saunders and Ramsbottom, 1986) flooded the eastern and western Finnmark Platform and allowed the deposition of upper Carboniferous sedimentary rocks of the Gipsdalen Group (Fig. 3). These rocks occur as syn-tectonic sedimentary wedges that thicken in the hanging wall of basin-bounding normal faults such as the LVF on the eastern Finnmark Platform and the main segment of the MFC on the western Finnmark Platform (Figs. 5a, b, e, and f and 11c). Similarly, in the southwesternmost Nordkapp basin, which may have remained flooded through the entire phase of eustatic sea-level fall and core complex exhumation, thick, partly evaporitic, upper Carboniferous sedimentary rocks were deposited in the basin and these are thickest at the intersection of the TFFC and the MFC (Figs. 5c, d, and g and 9b). Thus, the thickening of upper Carboniferous strata probably reflects significant syn-sedimentary normal faulting along these two faults, which may have acted as a single fault during the final stage of extension in the late Carboniferous (Fig. 11c).

Most faults on the eastern and western Finnmark Platform and in the southwesternmost Nordkapp basin appear to die out below the base Asselian reflection and those that propagate through this reflection show a limited amount of offset within Permian and Mesozoic–Cenozoic sedimentary strata (Fig. 5). Moreover, the Permian sedimentary succession shows a rather constant thickness through the entire study area (Figs. 5 and 9c). Thus, we argue that late- to post-Caledonian extensional faulting linked to the collapse of the Caledonides essentially took place in the Middle to Late Devonian–Carboniferous and came to a halt towards the end of this period (Fig. 11d). This presumed timing is consistent with recent K–Ar radiometric dating of brittle fault gouges in western Troms (Davids et al., 2013) and in NW Finnmark (Torgersen et al., 2014; Koehl et al., 2016), as well as with radiometric dating of dolerite dykes in NW Finnmark (Lippard and Prestvik, 1997), eastern Finnmark (Guise and Roberts, 2002), and on the Kola Peninsula in Russia (Roberts and Onstott, 1995), constraining significant extensional faulting activity in northern Norway and adjacent areas in Russia to the Late Devonian and early to mid-Permian. Minor reactivation of major fault complexes occurred in the Mesozoic–Cenozoic and are most likely associated with the rifting of the NE Atlantic (Faleide et al., 2008).

6 Conclusions

1. The atypically linear, NE–SW-trending main fault segment of the Måsøy Fault Complex formed as a brittle splay of the inverted Caledonian Sørøya-Ingøya shear zone through excisement processes during the collapse of the Caledonides in the Middle to Late Devonian–early Carboniferous and was active until the end of the late Carboniferous.

2. The WNW–ESE-trending fault segment of the Troms–Finnmark Fault Complex developed as a hard-linked, accommodation cross fault that accommodated orogen-parallel late- to post-orogen extension in the Middle to Late Devonian–Carboniferous. This fault merged with the main segment of the Måsøy Fault Complex and the two faults acted as a single fault at least during the late Carboniferous, but potentially from Devonian–early Carboniferous times, and accommodated the deposition of thick Devonian–lower Carboniferous and partly evaporitic upper Carboniferous deposits in the southwesternmost Nordkapp basin before faulting came to a halt towards the end of the late Carboniferous.

3. Low-gravity anomalies in the Gjesvær Low and southwesternmost Nordkapp basin may result from the presence of a thick Middle to Upper Devonian spoon-shaped sedimentary basin that developed along an inverted, bowed portion of the Sørøya-Ingøya shear zone during the collapse of the Caledonides and that display a geometry similar to those of Middle Devonian late- to post-orogenic collapse basins in western and mid-Norway.

4. The ENE–WSW- and NE–SW-trending basement ridges in the footwall of the Troms–Finnmark Fault Complex and on the northern flank of the southwesternmost Nordkapp basin formed through incisement processes and were exhumed along a bowed portion of the Sørøya-Ingøya shear zone during the collapse of the Caledonides in the Middle to Late Devonian–early Carboniferous. These basement ridges are thought to be part of a large-scale margin-parallel, NE–SW-trending metamorphic core complex that includes a succession of aligned basement highs such as the Lofoten Ridge, the West Troms Basement Complex, and the Norsel High. Core complex exhumation is believed to have stopped by the end of the Serpukhovian when a major eustatic sea-level rise flooded the Finnmark Platform, leading to the deposition of widespread upper Carboniferous sediments.

5. The Sørøya-Ingøya shear zone is thought to have truncated and decapitated Precambrian faults such as the Trollfjorden–Komagelva Fault Zone through top–SE thrusting during the Caledonian Orogeny and subsequent late- to post-orogenic extension. We nevertheless believe that preserved segments of these Precambrian faults might be preserved offshore on basement highs such as the Loppa and Veslemøy highs. However, more work is required in order to map and evaluate the impact of these WNW–ESE-trending, subvertical Precambrian faults on the SW Barents Sea margin.

Data availability. The seismic data analyzed in this study are part of the Diskos database and are publicly accessible from any Norwegian academic institution. Aeromagnetic data discussed in the present contribution are from Gernigon et al. (2014).

The Supplement related to this article is available online at https://doi.org/10.5194/se-9-341-2018-supplement.

Author contributions. JBPK interpreted the seismic and aeromagnetic data and is the main contributor to the writing process (workload ca. 45 %). SGB provided significant input to the "Introduction" and "Geological Setting" sections as well as detailed critical reviews of the whole paper (workload ca. 30 %). TH helped initiate the project and provided help with seismic well ties and regional seismic interpretation (workload ca. 15 %). JIF provided help with the writing process and helped improve the margin evolution model (workload ca. 10 %).

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. The present study is part of the ARCEx project (Research Centre for Arctic Petroleum Exploration), which is funded by the Research Council of Norway (grant number 228107) together with 10 academic and eight industry partners. We would like to thank all the persons from these institutions that are involved in this project. We acknowledge the NPD, the NTNU, and Schlumberger for sharing seismic data from the Diskos database, with special thanks to Dicky Harishidayat for taking care of the transfer of the data. We thank Tom Arne Rydningen from the University of Tromsø for reviewing parts of the paper. We would also like to thank Hanne-Kristin Paulsen from the University of Tromsø as well as Gwenn Péron-Pinvidic and Per-Terje Osmundsen from the NGU for fruitful discussions. Finally, we thank Thomas Phillips, Dora Marin, and David Roberts for constructive reviews.

Edited by: Mark Allen
Reviewed by: Thomas Phillips and Dora Marin

References

Andersen, T. B.: The structure of the Magerøy Nappe, Finnmark, North Norway, Nor. Geol. Unders., 363, 1–23, 1981.

Andersen, T. B.: The stratigraphy of the Magerøy Supergroup, Finnmark, north Norway, Nor. Geol. Unders., 395, 25–37, 1984.

Andresen, A., Agyei-Dwarko, N. Y., Kristoffersen, M., and Hanken, N.-M.: A Timanian foreland basin setting for the late Neoproterozoic-Early Palaeozoic cover sequences (Dividal Group) of northeastern Baltica, in: New Perspectives on the Caledonides of Scandinavia and Related Areas, edited by: Corfu, F., Gasser, D., and Chew, D. M., Geological Society, London, UK, Special Publications, 390, 157–175, 2014.

Bergh, S. G. and Torske, T.: The Proterozoic Skoadduvarri Sandstone Formation, Alta, Northern Norway: A tectonic fan-delta complex, Sediment. Geol., 47, 1–25, 1986.

Bergh, S. G. and Torske, T.: Palaeovolcanology and tectonic setting of a Proterozoic metatholeiitic sequence near the Baltic Shield Margin, northern Norway, Precambrian Res., 39, 227–246, 1988.

Bergh, S. G., Eig, K., Kløvjan, O. S., Henningsen, T., Olesen, O., and Hansen, J.-A.: The Lofoten-Vesterålen continental margin: a multiphase Mesozoic-Palaeogene rifted shelf as shown by offshore-onshore brittle fault-fracture analysis, Norw. J. Geol., 87, 29–58, 2007.

Bergh, S. G., Kullerud, K., Armitage, P. E. B., Zwaan, K. B., Corfu, F., Ravna, E. J. K., and Myhre, P. I.: Neoarchaean to Svecofennian tectono-magmatic evolution of the West Troms Basement Complex, North Norway, Norw. J. Geol., 90, 21–48, 2010.

Bergø, E.: Analyses of Paleozoic and Mesozoic brittle fractures in West-Finnmark, unpublished Master's Thesis, University of Tromsø, Tromsø, Norway, 128 pp., 2016.

Blystad, P., Brekke, H., Færseth, R. B., Larsen, B. T., Skogseid, J., and Tørudbakken, B.: Structural elements of the Norwegian continental shelf, Part II: The Norwegian Sea Region, Norwegian Petroleum Directorate Bulletin, 8, 1–45, 1995.

Braathen, A., Osmundsen, P.-T., Nordgulen, Ø., Roberts, D., and Meyer, G. B.: Orogen-parallel extension of the Caledonides in northern Central Norway: an overview, Norw. J. Geol., 82, 225–241, 2000.

Braathen, A., Osmundsen, P.-T., Hauso, H., Semshaug, S., Fredman, N., and Buckley, S. J.: Fault-induced deformation in poorly consolidated, siliciclastic growth basin: A study from the Devonian in Norway, Tectonophysics, 586, 112–129, 2013.

Breivik, A. J., Gudlaugsson, S. T., and Faleide, J. I.: Ottar Basin, SW Barents Sea: a major Upper Palaeozoic rift basin containing large volumes of deeply buried salt, Basin Res., 7, 299–312, 1995.

Bugge, T., Mangerud, G., Elvebakk, G., Mørk, A., Nilsson, I., Fanavoll, S., and Vigran, J. O.: The Upper Palaeozoic succession on the Finnmark Platform, Barents Sea, Norsk Geol. Tidsskr., 75, 3–30, 1995.

Bylund, G.: Palaeomagnetism of the Late Precambrian Vadsø and Barents Sea Groups, Varanger Peninsula, Norway, Precambrian Res., 69, 81–93, 1994.

Chauvet, A. and Séranne, M.: Extension-parallel folding in the Scandinavian Caledonides: implications for late-orogenic processes, Tectonophysics, 238, 31–54, 1994.

Cianfarra, P. and Salvini, F.: Lineament Domain of Regional Strike-Slip Corridor: Insight from the Neogene Transtentional De Geer Transform Fault in NW Spitsbergen, Pure Appl. Geophys., 172, 1185–1201, 2015.

Corfu, F., Armitage, P. E. B., Kullerud, K., and Bergh, S.G.: Preliminary U-Pb geochronology in the West Troms Basement Complex, North Norway: Archaean and Palaeoproterozoic events and younger overprints, Norg. Geol. Unders. B., 441, 61–72, 2003.

Corfu, F., Torsvik, T. H., Andersen, T. B., Ashwal, L. D., Ramsay, D. M., and Roberts, R. J.: Early Silurian mafic-ultramafic and granitic plutonism in contemporaneous flysch, Magerøy, northern Norway: U-Pb ages and regional significance, J. Geol. Soc. London, 163, 291–301, 2006.

Corfu, F., Andersen, T. B., and Gasser, D.: The Scandinavian Caledonides: main features, conceptual advances and critical questions, in: New Perspectives on the Caledonides of Scandinavia and Related Areas, edited by: Corfu, F., Gasser, D., and Chew, D. M., Geological Society, London, UK, Special Publications, 390, 9–43, 2014.

Cutbill, J. L. and Challinor, A.: Revision of the Stratigraphical Scheme for the Carboniferous and Permian of Spitsbergen and Bjørnøya, Geol. Mag., 102, 418–439, 1965.

Cutbill, J. L., Henderson, W. G., and Wright, N. J. R.: The Billefjorden Group (Early Carboniferous) of central Spitsbergen, Norsk Polarinst. Skri., 164, 57–89, 1976.

Davids, C., Wemmer, K., Zwingmann, H., Kohlmann, F., Jacobs, J., and Bergh, S. G.: K-Ar illite and apatite fission track constraints on brittle faulting and the evolution of the northern Norwegian passive margin, Tectonophysics, 608, 196–211, 2013.

Doré, A. G., Lundin, E. R., Fichler, C., and Olesen, O.: Patterns of basement structure and reactivation along the NE Atlantic margin, J. Geol. Soc. London, 154, 85–92, 1997.

Eig, K.: Onshore and offshore tectonic evolution of the Lofoten passive margin, North Norway, PhD thesis, University of Tromsø, Tromsø, Norway, 256 pp., 2008.

Eig, K. and Bergh, S. G.: Late Cretaceous-Cenozoic fracturing in Lofoten, North Norway: Tectonic significance, fracture mechanisms and controlling factors, Tectonophysics, 499, 190–215, 2011.

Elvevold, S., Reginiussen, H., Krogh, E. J., and Bjørklund, F.: Reworking of deep-seated gabbros and associated contact metamorphosed paragneisses in the south-eastern part of the Seiland Igneous Province, northern Norway, J. Metamorph. Geol., 12, 539–556, 1994.

Faleide, J. I., Gudlaugsson, S. T., and Jacquart, G.: Evolution of the western Barents Sea, Mar. Petrol. Geol., 1, 123–150, 1984.

Faleide, J. I., Vågnes, E., and Gudlaugsson, S. T.: Late Mesozoic-Cenozoic evolution of the south-western Barents Sea in a regional rift-shear tectonic setting, Mar. Petrol. Geol., 10, 186–214, 1993.

Faleide, J. I., Tsikalas, F., Breivik, A. J., Mjelde, R., Ritzmann, O., Engen, Ø., Wilson, J., and Eldholm, O.: Structure and evolution of the continental margin off Norway and the Barents Sea, Episodes, 31, 82–91, 2008.

Fazlikhani, H., Fossen, H., Gawthorpe, R. L., Faleide, J. I., and Bell, R. E.: Basement structure and its influence on the structural configuration of the northern North Sea rift, Tectonics, 36, 1–27, 2017.

Fountain, D. M., Hurich, C. A., and Smithson, S. B.: Seismic reflectivity of mylonite zones in the crust, Geology, 12, 195–198, 1984.

Gabrielsen, R. H.: Long-lived fault zones and their influence on the tectonic development of the southwestern Barents Sea, J. Geol. Soc. London, 141, 651–662, 1984.

Gabrielsen R. H. and Færseth, R. B.: The inner shelf of North Cape, Norway and its implications for the Barents Shelf-Finnmark Caledonide boundary. A comment, Norsk Geol. Tidsskr., 69, 57–62, 1989.

Gabrielsen R. H., Færseth, R. B., Jensen, L. N., Kalheim, J. E., and Riis, F.: Structural elements of the Norwegian continental shelf, Part I: The Barents Sea Region, Norwegian Petroleum Directorate Bulletin, 6, 1–33, 1990.

Gabrielsen R. H., Kløvjan, O. S., Rasmussen, A., and Stølan, T.: Interaction between halokinesis and faulting: structuring of the margins of the Nordkapp Basin, Barents Sea region, in: Structural and Tectonic Modelling and its Application to Petroleum Geology, edited by: Larsen, R. M., Brekke, H., Larsen, B. T., and Talleraas, E., NPF Special Publication, Amsterdam, 1, 121–131, 1992.

Gayer, R. A., Hayes, S. J., and Rice, A. H. N.: The structural development of the Kalak Nappe Complex of Eastern and Central Porsangerhalvøya, Finnmark, Norway, Norg. Geol. Unders. B., 400, 67–87, 1985.

Gernigon L. and Brönner, M.: Late Palaeozoic architecture and evolution of the southwestern Barents Sea: insights from a new generation of aeromagnetic data, J. Geol. Soc. London, 127, 1–11, 2012.

Gernigon, L., Brönner, M., Roberts, D., Olesen, O., Nasuti, A., and Yamasaki, T.: Crustal and basin evolution of the southwestern Barents Sea: From Caledonian orogeny to continental breakup, Tectonics, 33, 347–373, 2014.

Gjelberg, J. G.: Upper Devonian (Famennian) – Middle Carboniferous succession of Bjørnøya, a study of ancient alluvial and coastal marine sedimentation, Norsk Polarinst. Skri., 174, 1–67, 1981.

Gjelberg, J. G.: Early-Middle Carboniferous sedimentation on Svalbard. A study of ancient alluvial and coastal marine sedimentation in rift- and strike-slip basins, PhD thesis, University of Bergen, Bergen, Norway, 306 pp., 1984.

Gjelberg, J. G. and Steel, R. J.: An outline of Lower-Middle Carboniferous sedimentation on Svalbard: Effects of tectonic, climatic and sea level changes in rift basin sequences, in: Geology of the North Atlantic Borderlands, edited by: Kerr, J. W. and Ferguson, A. J., Can. Soc. Petrolm. Geol. Mem., 7, 543–561, 1981.

Gjelberg, J. G. and Steel, R. J.: Middle Carboniferous marine transgression, Bjørnøya, Svalbard: facies sequences from an interplay of sea level changes and tectonics, Geol. J., 18, 1–19, 1983.

Gudlaugsson, S. T., Faleide, J. I., Johansen, S. E., and Breivik, A. J.: Late Palaeozoic structural development of the South-western Barents Sea, Mar. Petrol. Geol., 15, 73–102, 1998.

Guise, P. G. and Roberts, D.: Devonian ages from ^{40}Ar/^{39}Ar dating of plagioclase in dolerite dykes, eastern Varanger Peninsula, North Norway, Nor. Geol. Unders., 440, 27–37, 2002.

Hallett, B. W., McClelland, W. C., and Gilotti, J. A.: The Timing of Strike-Slip Deformation Along the Storstrømmen Shear Zone, Greenland Caledonides: U-Pb Zircon and Titanite Geochronology, Geosci. Can., 41, 19–45, 2014.

Hansen, J.-A. and Bergh, S. G.: Origin and reactivation of fracture systems adjacent to the Mid-Norwegian continental margin on Hamarøya, North Norway: use of digital geological mapping and morphotectonic lineament analysis, Norw. J. Geol., 92, 391–403, 2012.

Hansen, J.-A., Bergh, S. G., and Henningsen, T.: Mesozoic rifting and basin evolution on the Lofoten and Vesterålen Margin, North-Norway; time constraints and regional implications, Norw. J. Geol., 91, 203–228, 2012.

Herrevold, T., Gabrielsen, R. H., and Roberts, D.: Structural geology of the southeastern part of the Trollfjorden-Komagelva Fault Zone, Varanger Peninsula, Finnmark, North Norway, Norw. J. Geol., 89, 305–325, 2009.

Indrevær, K. and Bergh, S. G.: Linking onshore-offshore basement rock architecture and brittle faults on the submerged strandflat along the SW Barents Sea margin, using high-resolution (5 × 5 m) bathymetry data, Norw. J. Geol., 94, 1–34, 2014.

Indrevær, K., Bergh, S. G., Koehl, J.-B., Hansen, J.-A., Schermer, E. R., and Ingebrigtsen, A.: Post-Caledonian brittle fault zones on the hyperextended SW Barents Sea margin: New insights into onshore and offshore margin architecture, Norw. J. Geol., 93, 167–188, 2013.

Jensen, L. N. and Sørensen, K.: Tectonic framework and halokinesis of the Nordkapp Basin, Barents Sea, in: Structural and Tectonic Modelling and its Application to Petroleum Geology, edited by: Larsen, R. M., Brekke, H., Larsen, B. T., and Talleraas, E., Norwegian Petroleum Society (NPF), Special Publications, Amsterdam, 1, 109–120, 1992.

Johansen, S. E., Henningsen, T., Rundhovde, E., Sæther, B. M., Fichler, C., and Rueslåtten, H. G.: Continuation of the Caledonides north of Norway: seismic reflectors within the basement beneath the southern Barents Sea, Mar. Petrol. Geol., 11, 190–201, 1994.

Johnson, H. D., Levell, B. K., and Siedlecki, S.: Late Precambrian sedimentary rocks in East Finnmark, north Norway and their relationship to the Trollfjord-Komagelv fault, J. Geol. Soc. London, 135, 517–533, 1978.

Kairanov, B., Escalona, A., Norton, I., Nadeau, P., and Lawver, L.: Early Cretaceous-Cenozoic relation between the Tromsø Basin and NE Greenland; Implications for margin evolution, in: Onshore–Offshore Relationships on the North Atlantic margins, Trondheim, 18–19 October, Norsk Geologi Forening Abstracts and Proceedings, 1, 2016.

Kirkland, C. L., Daly, J. S., and Whitehouse, M. J.: Early Silurian magmatism and the Scandian evolution of the Kalak Nappe Complex, Finnmark, Arctic Norway, J. Geol. Soc. London, 162, 985–1003, 2005.

Kirkland, C. L., Daly, J. S., and Whitehouse, M. J.: Provenance and Terrane Evolution of the Kalak Nappe Complex, Norwegian Caledonides: Implications for Neoproterozoic Paleogeography and Tectonics, J. Geol., 115, 21–41, 2007.

Kirkland, C. L., Daly, J. S., and Whitehouse, M. J.: Basement-cover relationships of the Kalak Nappe Complex, Arctic Norwegian Caledonides and constraints on Neoproterozoic terrane assembly in the North Atlantic region, Precambrian Res., 160, 245–276, 2008.

Klein, A. C. and Steltenpohl, M. G.: Basement-cover relations and late- to post-Caledonian extension in the Leknes group, west-central Vestvågøy, Lofoten, north Norway, Norsk Geol. Tidsskr., 79, 19–31, 1999.

Klein, A. C., Steltenpohl, M. G., Hames, W. E., and Andresen, A.: Ductile and brittle extension in the southern Lofoten archipelago, north Norway: implications for differences in tectonic style along an ancient collisional margin, Am. J. Sci., 299, 69–89, 1999.

Koehl, J.-B., Wemmer, K., and Bergh, S. G.: Shallow faulting and westwards migration of rifting from mid-Permian times in North Norway, evidence from K/Ar radiometric dating, in: Onshore–Offshore Relationships on the North Atlantic margins, Trondheim, 18–19 October, Norsk Geologi Forening Abstracts and Proceedings, 1, 2016.

Koehl, J.-B., Kvanli, T.-E., Rowan, M. G., and Faleide, J. I.: Late Carboniferous salt tectonics in the southwesternmost Nordkapp basin, Poster presentation, ARCEx Annual Conference 2017, 10–11 May 2017, Malangen, Norway, 2017.

Koehl, J.-B. P., Bergh, S. G., Osmundsen, P.-T., Redfield, T. F., Indrevær, K., Lea, H., and Bergø, E.: Late Devonian-Carboniferous faulting and controlling fabrics in NW Finnmark , Norw. J. Geol., submitted, 2018.

Koyi, H., Talbot, C. J., and Tørudbakken, B. O.: Salt diapirs of the southwest Nordkapp Basin analogue modelling, Tectonophysics, 228, 167–187, 1993.

Larssen, G. B., Elvebakk, G., Henriksen, S. E., Nilsson, I., Samuelsberg, T. J., Svånå, T. A., Stemmerik, L., and Worsley D.: Upper Palaeozoic lithostratigraphy of the Southern Norwegian Barents Sea, Norwegian Petroleum Directorate Bulletin, 9, 1–76, 2002.

Lea, H.: Analysis of Late plaeozoic-Mesozoic brittle faults and fractures in West-Finnmark: geometry, kinematics, fault rocks and the relationship to offshore structures on the Finnmark Platform in the SW Barents Sea, Unpublished Master's Thesis, University of Tromsø, Tromsø, Norway, 129 pp., 2016.

Lippard, S. J. and Prestvik, T.: Carboniferous dolerite dykes on Magerøy: new age determination and tectonic significance, Norsk Geol. Tidsskr., 77, 159–163, 1997.

Lippard, S. J. and Roberts, D.: Fault systems in Caledonian Finnmark and the southern Barents Sea, Norg. Geol. Unders. B., 410, 55–64, 1987.

Lister, G. S. and Davis, G. A.: The origin of metamorphic core complexes and detachment faults formed during Tertiary continental extension in the northern Coloradi River region, USA, J. Struct. Geol., 11, 65–94, 1989.

Manby, G. and Lyberis, N.: Tectonic evolution of Devonian Basin of northern Svalbard, Norsk Geol. Tidsskr., 72, 7–19, 1992.

Mattingsdal, R., Høy, T., Simonstad, E., and Brekke, H.: An updated map of structural elements in the southern Barents Sea, 31st Geological Winter Meeting, 12–14 January 2015, Norwegian Petroleum Directorate, 2015.

Maxant, J.: Variation of density with rock type, depth, and formation in the Western Canada basin from density logs, Geophysics, 45, 1061–1076, 1980.

McCann, A. J. and Dallmann, W. K.: Reactivation of the long-lived Billefjorden Fault Zone in north central Spitsbergen, Svalbard, Geol. Mag., 133, 63–84, 1996.

McClelland, W. C., Gilotti, J. A., Ramarao, T., Stemmerik, L., and Dalhoff, F.: Carboniferous basin in Holm Land records local exhumation of the North-East Greenland Caledonides: Implications for the detrital zircon signature of a collisional orogeny, Geosphere, 12, 925–947, 2016.

Nasuti, A., Roberts, D., and Gernigon, L.: Multiphase mafic dykes in the Caledonides of northern Finnmark revealed by a new high-resolution aeromagnetic dataset, Norw. J. Geol., 95, 251–263, 2015.

Nilsen, K. T., Vendeville, B. C., Johansen, J.-T., and Tulsa, O. K.: Influence of Regional Tectonics on Halokinesis in the Nordkapp Basin, Barents Sea, in: Salt tectonics a global perspective, edited by: Jackson, M. P. A., Roberts, D. G., and Snelson, S., AAPG Memoir, 65, 413–436, 1995.

Olesen, O., Roberts, D., Henkel, H., Lile, O. B., and Torsvik, T. H.: Aeromagnetic and gravimetric interpretation of regional structural features in the Caledonides of West Finnmark and North Troms, northern Norway, Norg. Geol. Unders. B., 419, 1–24, 1990.

Olesen, O., Torsvik, T. H., Tveten, E., and Zwaan, K. B.: The Lofoten-Lopphavet Project – an integrated approach to the study of a passive continental margin, Summary report, Unpublished NGU report, Trondheim, 54 pp., 1993.

Olesen, O., Torsvik, T. H., Tveten, E., Zwaan, K. B., Løseth, H., and Henningsen, T.: Basement structure of the continental margin in the Lofoten-Lopphavet area, northern Norway: constraints from potential field data, on-land structural mapping and palaeomagnetic data, Norsk Geol. Tidsskr., 77, 15–30, 1997.

Olesen, O., Brönner, M., Ebbing, J., Gellein, J., Gernigon, L., Koziel, J., Lauritsen, T., Myklebust, R., Pascal, C., Sand, M., Solheim, D., and Usov, S.: New aeromagnetic and gravity compilations from Norway and adjacent areas: methods and applications, in: From Mature Basins to New Frontiers, edited by: Vining, B. A. and Pickering, S. C., Petroleum Geology, Proceedings of the 7th Petroleum Geology Conference, London, 4 January 2011, 559–586, 2010.

Omosanya, K. O., Johansen, S. E., and Harishidayat, D.: Evolution and character of supra-salt faults in the Easternmost Hammerfest Basin, SW Barents Sea, Mar. Petrol. Geol., 66, 1013–1028, 2015.

Osmundsen, P.-T. and Andersen, T. B.: The middle Devonian basins of western Norway: sedimentary response to large-scale transtensional tectonics, Tectonophysics, 332, 51–68, 2001.

Osmundsen, P. T., Braathen, A., Sommaruga, A., Skilbrei, J. R., Nordgulen, Ø., Roberts, D., Andersen, T. B., Olesen, O., and Mosar, J.: Metamorphic core complexes along the Mid-Norwegian margin: an overview and some current ideas. In: Onshore-offshore relations on the Mid-Norwegian margin, edited by: Wandaas, B., Nystuen, J. P., Eide, E. and Gradstein, F., Norwegian Petroleum Society, Special Publications, Trondheim, Norway, 12, 29–41, 2005.

Ottesen, D., Stokes, C. R., Rise, L., and Olsen, L.: Ice-sheet dynamics and ice streaming along the coastal parts of northern Norway, Quaternary Sci. Rev., 27, 922–940, 2008.

Pastore, Z., Fichler, C., and McEnroe, S. A.: The deep crustal structure of the mafic-ultramafic Seiland Igneous Province of Norway from 3-D gravity modelling and geological implications, Geophys. J. Int., 207, 1653–1666, 2016.

Pharaoh, T. C., Macintyre, R. M., and Ramsay, D. M.: K-Ar age determinations on the Raipas suite in the Komagfjord Window, northern Norway, Norsk Geol. Tidsskr., 62, 51–57, 1982.

Pharaoh, T. C., Ramsay, D. M., and Jansen, Ø.: Stratigraphy and Structure of the Northern Part of the Repparfjord-Komagfjord Window, Finnmark, Northern Norway, Nor. Geol. Unders., 377, 1–45, 1983.

Phillips, T., Jackson, C. A.-L., Bell, R. E., Duffy, O. B., and Fossen, H.: Reactivation of intrabasement structures during rifting: A case study from offshore southern Norway, J. Struct. Geol., 91, 54–73, 2016.

Rafaelsen, B., Elvebakk, G., Andreassen, K., Stemmerik, L., Colpaert, A., and Samuelsberg, T. J.: From detached to attached carbonate buildup complexes – 3-D seismic data from the Upper Palaeozoic, Finnmark Platform, southwestern Barents Sea, Sediment. Geol., 206, 17–32, 2008.

Ramberg, I. B., Bryhni, I., Nøttvedt, A., and Rangnes , K.: The making of a land. Geology of Norway, The Norwegian geological Association, Oslo, Norway, 2008.

Ramsay, D. M., Sturt, B. A., Jansen, Ø., Andersen, T. B., and Sinha-Roy, S.: The tectonostratigraphy of western Porsangerhalvøya, Finnmark north Norwaym, in: The Caledonian OrogenScandinavia and Related Areas, edited by: Gee, D. G., and Sturt, B. A., John Wiley & Sons Ltd., 611–619, 1985.

Reeve, M. T., Bell, R. E., and Jackson, C. A. L.: Origin and significance of intra-basement seismic reflections offshore western Norway, J. Geol. Soc. London, 171, 1–4, 2013.

Rice, A. H. N.: Restoration of the External Caledonides, Finnmark, North Norway, in: New Perspectives on the Caledonides of Scandinavia and Related Areas, edited by: Corfu, F., Gasser, D., and Chew, D. M., Geological Society, London, UK, Special Publications, 390, 271–299, 2013.

Rise, L., Bellec, V. K., Chand, S., and Bøe R.: Pockmarks in the southwestern Barents Sea and Finnmark fjords, Norw. J. Geol., 94, 263–282, 2015.

Roberts, D.: Patterns of folding and fracturing in north-east Sørøy, Nor. Geol. Unders., 269, 89–95, 1971.

Roberts, D.: Tectonic deformation in the Barents Sea region of Varanger Peninsula, Finnmark, Nor. Geol. Unders., 282, 1–39, 1972.

Roberts, D.: Palaeocurrent data from the Kalak Nappe Complex northern Norway: a key element in models of terrane affiliation, Norw. J. Geol., 87, 319–328, 2007.

Roberts, D. and Lippard, S. J.: Inferred Mesozoic faulting in Finnmark: current status and offshore links, Norg. Geol. Unders. B., 443, 55–60, 2005.

Roberts, D. and Onstott, T. C.: ^{40}Ar/^{39}Ar laser microprobe analyses and geochemistry of dolerite dykes from the Rybachi and Sredni

Peninsulas, NW Kola, Russia, Nor. Geol. Unders., Special Publication, 7, 307–314, 1995.

Roberts, D. and Siedlecka, A.: Timanian orogenic deformation along the northeastern margin of Baltica, Northwest Russia and Northeast Norway, and Avalonian-Cadomian connections, Tectonophysics, 352, 169–184, 2002.

Roberts, D. and Siedlecka, A.: Provenance and sediment routing of Neoproterozoic formations on the Varanger, Nordkinn, Rybachi and Sredni peninsulas, North Norway and Northwest Russia: a review, Norg. Geol. Unders. B., 452, 1–19, 2012.

Roberts, D. and Williams, G. D.: Berggrunnskart KJØLLEFJORD 2236-4 M 1 : 50 000, foreløpig utgave, Nor. Geol. Unders., 2013.

Roberts, D., Mitchell, J. G., and Andersen, T. B.: A post-Caledonian dyke from Magerøy North Norway: age and geochemistry, Norw. J. Geol., 71, 289–294, 1991.

Roberts, D., Olesen, O., and Karpuz, M. R.: Seismo- and noetectonics in Finnmark, Kola Peninsula and the southern Barents Sea. Part 1: Geological and neotectonic framework, Tectonophysics, 270, 1–13, 1997.

Roberts, D., Chand, S., and Rise, L.: A half-graben of inferred Late Palaeozoic age in outer Varangerfjorden, Finnmark: evidence from seismic reflection profiles and multibeam bathymetry, Norw. J. Geol., 91, 191–200, 2011.

Røe, S.-L. and Roberts, D.: Basement core sandstones in well 7128/6-1, southern Barents Sea, and correlation with onshore upper Proterozoic sequences of northeastern Finnmark, Unpublished report, University of Bergen, Geological Institute, Bergen, Norway, 33 pp., 1992.

Rykkelid, E.: E69 "FATIMA" – Geologisk undersøkelse av tunnel under Magerøysundet, Veglaboratoriet, Statens Vegvesen, Oslo, Norway, 1992.

Samuelsberg, T. J., Elvebakk, G., and Stemmrik, L.: Late Paleozoic evolution of the Finnmark Platform, southern Norwegian Barents Sea, Norw. J. Geol., 83, 351–362, 2003.

Sartini-Rideout, C., Gilotti, J. A., and McClelland, W. C.: Geology and timing of dextral strike-slip shear zones in Danmarkshavn, North-East Greenland Caledonides, Geol. Mag., 143, 431–446, 2006.

Saunders, W. B. and Ramsbottom, W. H. C.: The mid-Carboniferous eustatic event, Geology, 14, 208–212, 1986.

Sengör, A. M. C.: Cross-faults and differential stretching of hanging walls in regions of low-angle normal faulting: examples from western Turkey, in: Continental Extensional Tectonics, edited by: Coward, M. P., Dewey, J. F., and Hancock P. L., Geological Society Special Publication, Oxford, Blackwell Scientific Publications, 28, 575–589, 1987.

Séranne, M., Chauvet, A., Seguret, M., and Brunel, M.: Tectonics of the Devonian collapse-basins of western Norway, Bull. Soc. Géol. Fr., 8, 489–499, 1989.

Séranne, M., and Seguret, M.: The Devonian basins of western Norway: tectonics and kinematics of an extending crust, in: Continental Extensional Tectonics, edited by: Coward, M. P., Dewey, J. F., and Hancock, P. L., Geological Society Special Publication, 28, 537–548, 1987.

Shipton, Z. K. and Cowie, P. A.: A conceptual model for the origin of fault damage zone structures in high-porosity sandstones, J. Struct. Geol., 25, 333–344, 2003.

Siedlecka, A.: Late Precambrian Stratigraphy and Structure of the North-Eastern Margin of the Fennoscandian Shield (East Finnmark – Timan Region), Nor. Geol. Unders., 316, 313–348, 1975.

Siedlecka, A. and Roberts, D.: The bedrock geology of Varanger Peninsula, Finnmark, North Norway: an excursion guide, Nor. Geol. Unders., Special Publication, 5, 1–45, 1992.

Siedlecka, A. and Siedlecki, S.: Some new aspects of the geology of Varanger peninsula (Northern Norway), Nor. Geol. Unders., 247, 288–306, 1967.

Siedlecka, A., Roberts, D., Nystuen, J. P., and Olovyanishnikov, V. G.: Northeastern and northwestern margins of Baltica in Neoproterozoic time: evidence from the Timanian and Caledonian Orogens, Geological Society, London, Memoirs, 30, 169–190, 2004.

Siedlecki, S.: Geologisk kart over Norge, berggrunnskart Vadsø – M 1 : 250 000, Nor. Geol. Unders., 1980.

Smelror, M., Petrov, O. V., Larssen, G. B., and Werner, S. C.: Geological history of the Barents Sea, Nor. Geol. Unders., Trondheim, Norway, 135 pp., 2009.

Solheim, A. and Kristoffersen, Y.: The physical environment in the Western Barents Sea, 1 : 5 000 000. Sediments above the upper regional unconformity: thickness, seismic stratigraphy and outline of the glacial history, Norsk Polarinst. Skri., 179B, 1–26, 1984.

Steltenpohl, M. G., Hames, W. E., and Andresen, A.: The Silurian to Permian history of a metamorphic core complex in Lofoten, northern Scandinavian Caledonides, Tectonics, 23, 1–23, 2004.

Steltenpohl, M. G., Moecher, D., Andresen, A., Ball, J., Mager, S., and Hames, W. E.: The Eidsfjord shear zone, Lofoten-Vesterålen, north Norway: An Early Devonian, paleoseismogenic low-angle normal fault, J. Struct. Geol., 33, 1023–1043, 2011.

Torgersen, E., Viola, G., Zwingmann, H., and Harris, C.: Structural and temporal evolution of a reactivated brittle-ductile fault – Part II: Timing of fault initiation and reactivation by K-Ar dating of synkinematic illite/muscovite, Earth Planet. Sc. Lett., 407, 221–233, 2014.

Torske, T. and Bergh, S. G.: The Caravarri Formation of the Kautokaino Greenstone Belt, Finnmark, North Norway; a Palaeoproterozoic foreland basin succession, Norg. Geol. Unders. B., 442, 5–22, 2004.

Townsend, C.: Thrust transport directions and thrust sheet restoration in the Caledonides of Finnmark, North Norway, J. Struct. Geol., 9, 345–352, 1987a.

Townsend, C.: The inner shelf of North Cape, Norway and its implications for the Baretns Shelf-Finnmark Caledonide boundary, Norsk Geol. Tidsskr., 67, 151–153, 1987b.

Vorren, T., Kristoffersen, Y., and Andreassen, K.: Geology of the inner shelf west of North Cape, Norway, Norsk Geol. Tidsskr., 66, 99–105, 1986.

Wilks, W. J. and Cuthbert, S. J.: The evolution of the Hornelen Basin detachment system western Norway: implications for the style of late orogenic extension in the southern Scandinavian Caledonides, Tectonophysics, 238, 1–30, 1994.

Worsley, D.: The post-Caledonian development of Svalbard and the western Barents Sea, Polar Res., 27, 298–317, 2008.

Worthing, M. A.: Fracture patterns on Eastern Seiland, North Norway and their possible Relationship to Regional Faulting, Norg. Geol. Unders. B., 398, 35–41, 1984.

Zhang, W. Roberts, D., and Pease, V.: Provenance of sandstones from Caledonian nappes in Finnmark, Norway: Implications for Neoproterozoic-Cambrian palaeogeography, Tectonophysics, 691, 198–205, 2016.

Zwaan, K. B.: Geology of the West Troms Basement Complex, northern Norway, with emphasis on the Senja Shear Belt: a preliminary account, Norg. Geol. Unders. B., 427, 33–36, 1995.

Zwaan, K. B. and Gautier, A. M.: Alta and Gargia. Description of the geological maps (AMS-M711) 1834 1 and 1934 IV 1 : 50 000, Nor. Geol. Unders., 357, 1–47, 1980.

Zwaan, K. B. and Roberts, D.: Tectonostratigraphic Succession and Development of the Finnmarkian Nappe Sequence, North Norway, Nor. Geol. Unders., 343, 53–71, 1978.