



Delft University of Technology

Multivariate analysis of GPS position time series of JPL second reprocessing campaign

Amiri Simkoei, Alireza; Haji Mohammadloo, Tannaz; Argus, D.F.

DOI

[10.1007/s00190-016-0991-9](https://doi.org/10.1007/s00190-016-0991-9)

Publication date

2017

Document Version

Accepted author manuscript

Published in

Journal of Geodesy

Citation (APA)

Amiri Simkoei, A., Haji Mohammadloo, T., & Argus, D. F. (2017). Multivariate analysis of GPS position time series of JPL second reprocessing campaign. *Journal of Geodesy*. <https://doi.org/10.1007/s00190-016-0991-9>

Important note

To cite this publication, please use the final published version (if applicable). Please check the document version above.

Copyright

Other than for strictly personal use, it is not permitted to download, forward or distribute the text or part of it, without the consent of the author(s) and/or copyright holder(s), unless the work is under an open content license such as Creative Commons.

Takedown policy

Please contact us and provide details if you believe this document breaches copyrights. We will remove access to the work immediately and investigate your claim.

Multivariate analysis of GPS position time series of JPL second reprocessing campaign

A. R. Amiri-Simkooei¹  · T. H. Mohammadloo² · D. F. Argus³

Received: 6 June 2016 / Accepted: 19 December 2016
© Springer-Verlag Berlin Heidelberg 2017

Abstract The second reprocessing of all GPS data gathered by the Analysis Centers of IGS was conducted in late 2013 using the latest models and methodologies. Improved models of antenna phase center variations and solar radiation pressure in JPL's reanalysis are expected to significantly reduce errors. In an earlier work, JPL estimates of position time series, termed first reprocessing campaign, were examined in terms of their spatial and temporal correlation, power spectra, and draconitic signal. Similar analyses are applied to GPS time series at 89 and 66 sites of the second reanalysis with the time span of 7 and 21 years, respectively, to study possible improvements. Our results indicate that the spatial correlations are reduced on average by a factor of 1.25. While the white and flicker noise amplitudes for all components are reduced by 29–56 %, the random walk amplitude is enlarged. The white, flicker, and random walk noise amount to rate errors of, respectively, 0.01, 0.12, and 0.09 mm/yr in the horizontal and 0.04, 0.41 and 0.3 mm/yr in the vertical. Signals reported previously, such as those with periods of 13.63, 14.76, 5.5, and $351.4/n$ for $n = 1, 2, \dots, 8$ days, are identified in multivariate spectra of both data sets. The oscillation of the draconitic signal is reduced by factors of 1.87, 1.87, and 1.68 in the east, north and up components,

respectively. Two other signals with Chandlerian period and a period of 380 days can also be detected.

Keywords GPS position time series · JPL second reprocessing campaign · Multivariate noise assessment · Multivariate power spectrum

1 Introduction

Continuous global positioning system (CGPS) time series have been widely used to study several geophysical phenomena (Segall and Davis 1997). These studies include inferring motion of the Earth's surface due to plate tectonics (Thatcher 2003; Argus et al. 2010; Kreemer et al. 2014), post-glacial rebound (Johansson et al. 2002; King et al. 2010; Peltier et al. 2015), and hydrological loading (van Dam et al. 2001; Rajner and Liwosz 2012; Argus et al. 2014). Moreover, strain accumulation (Argus et al. 2005; d'Alessio et al. 2005; Serpelloni et al. 2005; Craig and Calais 2014), sea-level variation (Wöppelmann et al. 2007), volcanic deformation (Bonforte and Puglisi 2006; Cervelli et al. 2006), and subsidence studies (Lü et al. 2008; Bock et al. 2012) can be conducted.

To effectively apply GPS time series to geophysical phenomena, appropriate functional and stochastic models are required. The functional model takes into consideration the deterministic effects—a linear trend, offsets, and potential periodicities—to name a few. The stochastic model identifies and determines the remaining unmodeled effects—white noise and power-law noise for instance. Deterministic effects, if left undetected in the functional model, may mistakenly mimic flicker noise and random walk noise (Williams et al. 2004; Amiri-Simkooei et al. 2007).

A proper stochastic model provides the best linear unbiased estimator (BLUE) of unknown parameters. It can also

✉ A. R. Amiri-Simkooei
amiri@eng.ui.ac.ir

¹ Department of Geomatics Engineering, Faculty of Civil Engineering and Transportation, University of Isfahan, Isfahan 81746-73441, Iran

² Aircraft Noise and Climate Effect Group, Faculty of Aerospace Engineering, Delft University of Technology, Kluyverweg 1, 2629 HS Delft, The Netherlands

³ Satellite Geodesy and Geodynamics, Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA

provide a realistic description of the parameters' precision. The parameter estimation in a stochastic model is referred to as variance component estimation (VCE). VCE can be conducted using various methods. The least-squares variance component estimation (LS-VCE), which was originally developed by [Teunissen \(1988\)](#), is used in the present contribution. For its geodetic and geophysical applications, we may refer to [Amiri-Simkooei et al. \(2007, 2009, 2013\)](#), [Amiri-Simkooei \(2007, 2009, 2013a, b\)](#), and [Khodabandeh et al. \(2012\)](#).

Proper analysis of GPS time series is a prerequisite for an appropriate geophysical interpretation. The VCE method based on the maximum likelihood estimation (MLE) has also been widely used to assess the noise structure of GPS time series. The differences between LS-VCE and MLE are explained by [Amiri-Simkooei et al. \(2007\)](#). [Zhang et al. \(1997\)](#) used MLE and found that the noise structure is a combination of white noise and flicker noise. Similar results have been drawn by [Bock et al. \(2000\)](#), [Calais \(1999\)](#), [Langbein and Bock \(2004\)](#), [Mao et al. \(1999\)](#), [Williams et al. \(2004\)](#). The presence of random walk noise or a combination of other noise components has been acknowledged by several scholars including [Johnson and Agnew \(2000\)](#), [King and Williams \(2009\)](#), [Langbein \(2008, 2012\)](#), [Langbein and Bock \(2004\)](#).

Cross-correlation among different series is an important issue. Errors in satellite orbits, Earth orientation parameters, and errors in daily and long-term geodetic reference frame are causes of regionally correlated errors ([Wdowski et al. 1997](#)). Moreover, large-scale atmosphere errors, receiver and satellite antenna phase center variations ([Dong et al. 2006](#)), and atmospheric and hydrospheric water loading effects ([van Dam et al. 2001](#)) are also candidates for common-mode errors (CMEs). [Williams et al. \(2004\)](#) found that in the regional GPS solutions in which CMEs have been removed, the noise is significantly lower compared to the global solutions. CMEs can be estimated with regional spatial filtering methods. We refer to the stacking approach, which was first utilized by [Wdowski et al. \(1997\)](#). [Nikolaidis \(2002\)](#) removed CMEs from daily GPS solutions by computing the daily weighted mean of residual noise from a few regional fiducial stations. [Teferle et al. \(2002\)](#) deployed a filtering technique to reduce the annual signal effect on site velocity estimates using a network of 9 stations. [Teferle et al. \(2006\)](#) used the weighted stacking method (WSM) to remove CMEs through analysis of a network consisting 6 permanent stations. Using the WSM, [Bogusz et al. \(2015\)](#) calculated CMEs for the ASG-EUPOS permanent stations.

As the regional networks expands, the magnitude of daily CMEs is reduced ([Márquez-Azúa and DeMets 2003](#)), and hence the application of the WSM becomes limited. [Dong et al. \(2006\)](#) presented a spatiotemporal filtering method based on principal component analysis (PCA) and Karhunen–Loeve expansion. Unlike the WSM, this method

allows data to reveal the spatial distribution of CMEs by disregarding the assumption of spatially uniform distribution of these errors. Because the stations we utilized are globally distributed, the concept of CMEs has lost its meaning ([Dong et al. 2006](#)). The cross-correlation (i.e. spatial correlation) among time series is thus investigated (see [Williams et al. 2004](#); [Amiri-Simkooei 2009](#)).

The GPS draconitic year (351.4 days) is the revolution period of the GPS constellation in inertial space with respect to the Sun. Harmonics of this periodic pattern have been observed in GPS-derived geodetic products. [Ray et al. \(2008\)](#) analyzed the time series of 167 IGS stations using the stacked Lomb–Scargle periodogram. They identified up to the sixth harmonic of GPS draconitic year in the east, north, and up components. [Collilieux et al. \(2007\)](#) found significant signals near the frequencies 2.08, 3.12, and 4.16 cpy in the up component. [Amiri-Simkooei et al. \(2007\)](#) computed the stacked least squares power spectra of 71 permanent GPS stations. They identified up to the eighth harmonic of the GPS draconitic signal. [Amiri-Simkooei \(2013a\)](#) identified ten harmonics of the draconitic signal by calculating the multivariate least-squares power spectrum of 350 permanent GPS stations. For more information on the harmonics of the GPS draconitic signal, we refer to the studies of [King and Watson \(2010\)](#), [Rodríguez-Solano et al. \(2012, 2014\)](#), [Ostini \(2012\)](#) and [Santamaría-Gómez et al. \(2011\)](#).

2 Second reprocessing campaign strategies

In 2008, the Analysis Centers (ACs) of the international GNSS service (IGS) initiated the reprocessing of the all GPS data gathered by the IGS global network since 1994 employing the latest methods upon that time in an entirely consistent manner. This was the first reprocessing campaign, and it was anticipated that as further analysis and improvements were made, undoubtedly, more reprocessing campaigns will be required. Thus, the 2nd reanalysis of all IGS data using the improved methods begun by the late 2013. Table 1 compares different aspects of the two processing campaigns.

Also, there are other modifications and changes in the models used within the 2nd reanalysis, which are explained in “Appendix 1”.

The new models used within the second reanalysis along with the studies conducted by [Hugentobler et al. \(2009\)](#) and [Rodríguez-Solano et al. \(2012\)](#), who emphasized the orbit mismodeling deficiencies and their effects on peculiar signals observed in GPS-derived products, motivated us to study the reprocessed daily position time series. We have suspected that since these models have been incorporated within the new reprocessing campaign, it is highly likely to observe significant improvements. Improvements expected include the reduction in the range of variations of the periodic pattern of

Table 1 Setup of the first and second reanalysis campaign

Campaign	First reprocessing	Second reprocessing
Duration	1994–2007	1994–2014
Reference frame	IGS05 (Aligned to ITRF2005)	IGb08 (Aligned to ITRF2008)
IERS convention	IERS 2003	IERS 2010
Antenna calibration	IGS05 ANTEX (absolute calibration)	IGS08 ANTEX (absolute calibration)

GPS draconitic year, the amplitude of different noise components, the spatial correlation of GPS position time series (Rodriguez-Solano et al. 2012). Reibischung et al. (2016) have recently shown that the noise characteristics of GPS position time series for JPL second reprocessing deviate from the common white plus flicker noise toward an only flicker background noise.

This contribution is a follow-up to the work carried out by Amiri-Simkooei (2013a) in which the daily position of many permanent GPS stations was analyzed. In the present contribution, the daily position time series of 66 and 89 permanent GPS stations of the length 21 and 7 years are derived from the 2nd reprocessing campaign (Fig. 1). They are referred to as data set #1 and data set #2, respectively, which are freely available in ftp://sideshow.jpl.nasa.gov/pub/JPL_GPS_Timeseries/repro2011b/post/point/. The time series with 89 GPS stations (data set #2) are also derived from the 1st reprocessing campaign to make comparisons. Therefore, for the data set #2 we have two kinds of data (Repro1 and Repro2) with the same length, time span, and time instants.

All formulas and methodologies, used in the subsequent sections, are based on those presented by Amiri-Simkooei (2013a) who used a multivariate time series analysis. This method is superior over univariate analysis because many weak signals and small noise amplitudes which cannot be detected in univariate analysis can be detected if we simultaneously analyze multiple time series. This holds, for example, when estimating the random walk amplitude, which has a high chance to be masked in the univariate analysis, but has a higher chance to be detected in the multivariate analysis. However, a drawback of this multivariate analysis is that it can only provide a kind of network-based random walk and hence such errors cannot necessarily be attributed to the individual time series. For further information, we may refer to Amiri-Simkooei (2013a).

3 Results and discussion

The multivariate method is used to study the GPS position time series of daily global solutions. These time series have been obtained using the precise point positioning (PPP) method in the GIPSY-OASIS software (Zumberge et al.

1997). The process has been carried out in an analysis center at JPL (Beutler et al. 1999).

Prior to the analysis, a multivariate offset detection method was used to identify and remove offsets in the series (Hoseini-Asl et al. 2013). Although the manual offset detection method is still more reliable than the existing methods (see Gazeaux et al. 2013), we used an automatic offset detection method having a few characteristics. This method assumes similar offsets in the three coordinate components. It also takes into account appropriate functional and stochastic models. For example, prior to offset detection, LS-VCE is applied to estimate the white and flicker noise amplitudes. Comparing the offset detection results with those in the JPL website indicates that our method detects all offsets reported by JPL. In addition, a few smaller offsets, which are likely due to other causes like small earthquakes, have been detected.

The initial functional model consists of a linear trend along with the three harmonics of the annual signal; the tri-annual signal was included because the power spectrum showed a signal near 122 days. Equation (8) in Amiri-Simkooei (2013a) is utilized to obtain the multivariate power spectrum (MPS) of multiple series. The analysis requires matrices Σ and Q , which can be estimated using a multivariate method (see Amiri-Simkooei 2009, algorithm in Fig. 1). The nonnegative least-squares variance component estimation method (NNLS-VCE) (Amiri-Simkooei 2016, algorithm in Fig. 1) has been employed to avoid nonnegative variance factors for white noise, flicker noise, and random walk noise. While the matrix Σ explains the spatial correlation among the series, Q considers the temporal correlation among observables within each series. For the flicker noise, the Hosking structure introduced by Williams (2003a) and Langbein (2004) has been employed.

Multivariate analysis requires simultaneous time series. This indicates that if there is a gap or outlier in a (couple of) series, the observations of other series should be removed to have simultaneous time instants for all series. However, if the data were available in 95% of the series (missed in 5% of the series), the observations for the gaps (missed) were reconstructed using the above functional model, and then a normally distributed noise based on the estimated stochastic model was added to reconstruct the data. For the data set # 2, 89 GPS stations were analyzed. Therefore, the total number of series is $r = 267$. Matrix Σ , which expresses the spatial

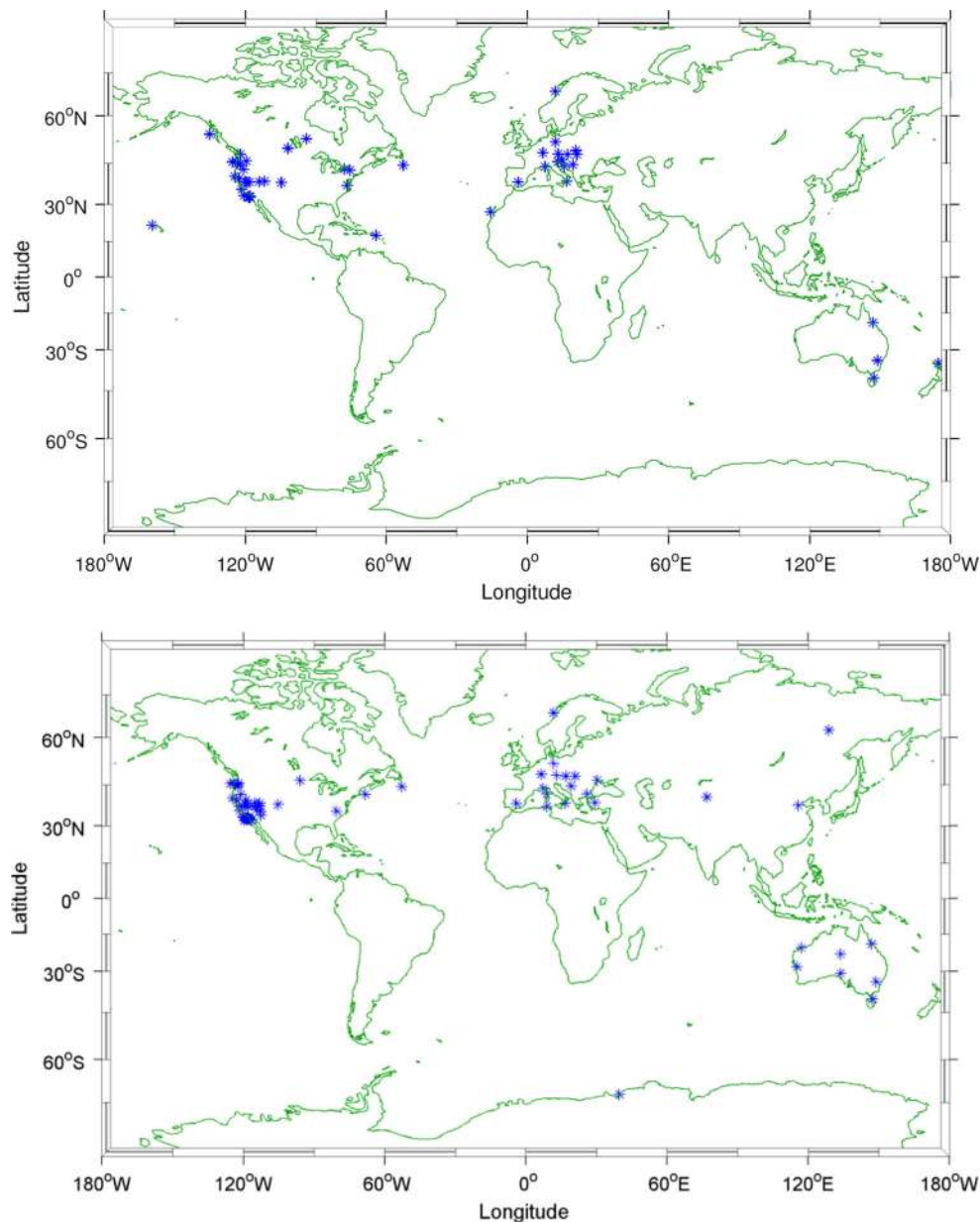


Fig. 1 World distribution of 66 GPS stations with time span of 21 years (*top*: Repro2), 89 GPS stations with time span of 7 years (*bottom*: Repro1 and Repro2)

242 correlation, is of size 267×267 . Matrix Q is of size $m \times m$,
 243 where m is the number of observables in each series; for
 244 the multivariate analysis, m is identical for all time series.
 245 While the three 89×89 block diagonals of the Σ form the
 246 spatial correlation of each coordinate component (i.e. east-
 247 east (EE), north-north (NN), and up-up (UU)), the other three
 248 89×89 off-diagonals represent the cross-correlation of the
 249 components (i.e. between north-east (NE), north-up (NU),
 250 and east-up (EU)).

251 The VCE methods can computationally be an expensive
 252 process. Some researchers have contributed to reduce the
 253 computational burden of the VCE methods. We may refer

to excellent studies by Bos et al. (2008, 2012) in which the
 computational burden of MLE is drastically reduced. One
 feature of our multivariate noise assessment method is also
 that its computational burden is similar to that of the univari-
 ate analysis (see Amiri-Simkooei 2009).

3.1 Spatial correlation

GPS position time series have been shown to have a sig-
 nificant spatial correlation (Williams et al. 2004; Amiri-
 Simkooei 2009, 2013a). The spatial (cross) correlation
 results for the data set with 89 GPS station are illustrated

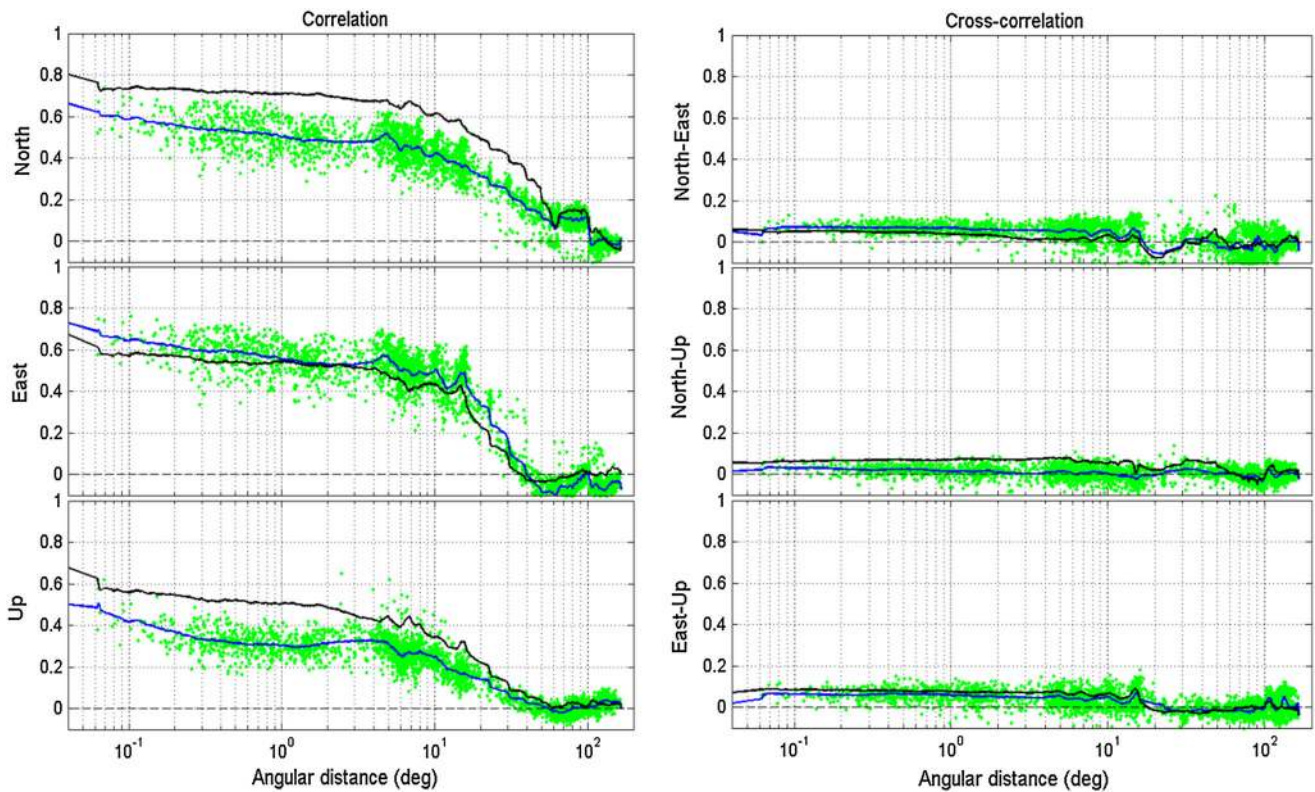


Fig. 2 Six kinds of spatial correlation estimated for position time series with the time span of 7 years as a function of angular distance (deg); (left) individual components NN, EE, and UU; (right) cross components

NE, NU, and EU. Indicated in the plots also mean correlation curves for the second (blue) and first (black) reprocessing campaigns using a moving average

264 in Fig. 2. Derived from Σ , this figure shows the spatial
 265 correlation among NN, EE, UU, NE, NU, and EU compo-
 266 nents. Significant spatial correlations for NN, EE, and UU
 267 are observed over an angular range of 0° to 30° , implying
 268 the presence of regionally correlated errors. No effort has been
 269 put forward to reduce CMEs here, and thus, as expected, the
 270 spatial correlation among stations which are close to each
 271 other (about 3000 km apart) is significant. This spatial cor-
 272 relation directly propagates into the correlation between site
 273 velocities, and hence it should be taken into consideration
 274 in the covariance matrix of the site velocities (Williams et al.
 275 2004). Over larger distances, the correlations of individual
 276 components experience a significant decline, in agreement
 277 with the findings of Amiri-Simkooei (2013a) and Williams
 278 et al. (2004). This indicates that the CME noise is significant
 279 only over nearby stations. The component EE experiences
 280 higher correlations compared with the NN and UU compo-
 281 nents.

282 The spatial cross-correlations between components (NE,
 283 NU, and EU) are negligible. The cross-correlation curve is
 284 less than 0.1 which is owing to a good GPS geometry stem-
 285 ming from simultaneous processing of all observations. To

286 fairly compare the average spatial (cross) correlations derived
 287 from the 1st and 2nd reprocessing campaign, the 1st repro-
 288 cessing campaign time series for the data set with 89 stations
 289 have been processed as well. The results are presented in
 290 Table 2. The spatial correlations of individual components
 291 have been reduced compared to those computed for the
 292 Repro1 data except for the EE component, which shows a
 293 (small) increase from 0.57 to 0.62 in the second reproces-
 294 sing. The reduction is the result of improvement in the models
 295 used within the new campaign. It could also be due to an
 296 improved alignment of the daily terrestrial frames, which
 297 makes it difficult to separate it from the impact of new mod-
 298 els used in the analysis. The spatial correlation matrix Σ ,
 299 estimated for the latest processing campaign, is to be taken
 300 into consideration in the estimation of the multivariate power
 301 spectrum.

302 In this contribution, we considered the correlation among
 303 the east, north and up components. In principle, by applying
 304 the error propagation law, these correlations can be propa-
 305 gated into the coordinate differences of X, Y, and Z in an
 306 earth-centered earth-fixed coordinate system using an appro-
 307 priate coordinate transformation.

3.2 Temporal correlation and noise assessment

The amplitudes of white noise, flicker noise, and random walk noise can be obtained using matrices Σ and Q . Noise characteristics of GPS time series have been expressed as a combination of white plus spatially correlated flicker noise (Zhang et al. 1997; Mao et al. 1999; Calais 1999; Nikolaidis et al. 2001; Williams et al. 2004; Amiri-Simkooei et al. 2007, 2009). The presence of random walk noise in GPS time series is due to monument instability (Williams et al. 2004) or the presence of nonlinear deformation behavior, for example in areas with active deformation or when the offsets remain in the data series (Williams 2003b). The presence of postseismic deformation or volcanic events could also increase the apparent amplitude of random walk noise. The reason for masking the (small) values of the random walk noise is the short time spans of the data series or the existence of dominant flicker noise (Williams et al. 2004).

The amplitudes of white noise, flicker noise, and random walk noise can simply be provided from the Kronecker struc-

Table 2 Average spatial correlation over the angular distance of 30° for the first and second reprocessing campaign using 89 GPS stations

Reprocessing campaign	Correlation			Cross-correlation		
	NN	EE	UU	NE	NU	EU
1st (Repro 1)	0.73	0.57	0.55	0.05	0.07	0.08
2nd (Repro 2)	0.56	0.62	0.37	0.07	0.03	0.06

ture $\Sigma \otimes Q$. The diagonal entries of the matrices $s_w \Sigma$, $s_f \Sigma$ and $s_{rw} \Sigma$ represent the variances of white, flicker and random walk noise for each series. To compare the amplitudes of the three noise components for the two reprocessing campaigns, the data sets with the time span of 7 years (89 GPS stations) of the two campaigns have been processed. For the second reanalysis, the time correlation results of these stations are shown in Fig. 3. The average amplitudes of white, flicker, and random walk noise components along with their estimated standard deviations for both campaigns are presented in Table 3. A few observations are highlighted.

- The amplitudes of all noise components of the vertical is larger than those of the horizontal by a factor of 3, consistent with the previously published results (Williams et al. 2004; Amiri-Simkooei 2013a; Dmitrieva et al. 2015).
- Amiri-Simkooei (2013a) published flicker noise variances for the repro1 series about 4 times smaller than those reported here. Unfortunately, there was a mistake in presenting flicker noise results in Amiri-Simkooei (2013a). There, the unit was mistakenly $\text{mm}/\text{day}^{1/4}$ (and not $\text{mm}/\text{year}^{1/4}$) for the flicker noise component. This indicates that a scaling factor of $\sqrt[4]{365.25} = 4.37$ should be applied to his flicker noise amplitudes.
- In contrast to the values obtained from the first reanalysis, the noise amplitudes of the north and east components are nearly identical in the second reanalysis.

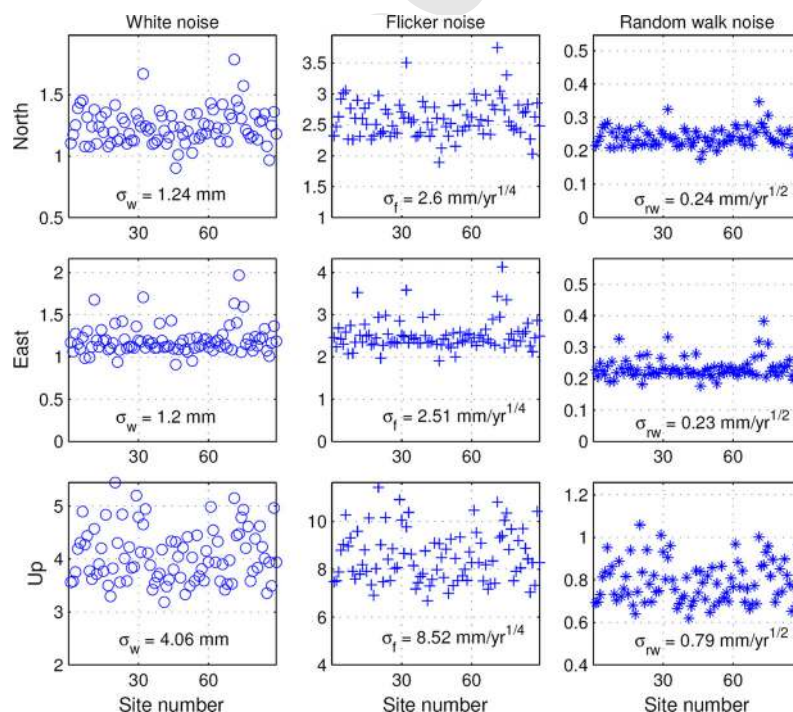


Fig. 3 Estimated amplitudes of white (left), flicker (middle), and random walk (right) noise for the data set with the time span of 7 years; top frame (north), middle frame (east), bottom frame (up)

Table 3 Average amplitudes of white noise, flicker noise, and random walk noise along with their estimated standard deviations for permanent GPS stations of 1st and 2nd processing campaigns

Processing campaign		Second	First
White noise (mm)	N	1.24 ± 0.02	2.02 ± 0.03
	E	1.20 ± 0.02	2.68 ± 0.04
	U	4.06 ± 0.06	5.69 ± 0.09
Flicker noise (mm/year ^{1/4})	N	2.60 ± 0.04	4.39 ± 0.06
	E	2.51 ± 0.04	5.80 ± 0.08
	U	8.52 ± 0.13	12.33 ± 0.18
Random walk (mm/year ^{1/2})	N	0.24 ± 0.004	0
	E	0.23 ± 0.004	0
	U	0.79 ± 0.010	0

- The amplitudes of flicker and random walk noise over different stations are multiples of the white noise amplitudes. In reality, however, this should not indicate all stations contain random walk noise, because the estimated values are an average value (over all stations) due to the special structure used (see Amiri-Simkooei et al. 2013). Therefore, the multivariate approach implemented in the present contribution can resolve only a single network-wide random walk value rather than a station specific one.
- When the values obtained from the latest reanalysis are compared to their older counterpart, the amplitudes of white and flicker noise of all components have been reduced by factors ranging from 1.40 to 2.33. This highlights that the new models used in the second reanalysis have significantly reduced the amplitude of these two noise components.
- While the amplitudes of both white and flicker noise have significantly reduced in this contribution, Rebischung et al. (2016) reported reduction in only white noise. This, however, was only speculated by explaining their power spectra and hence was not based on a real estimation of the noise amplitudes.
- The random walk noise amplitudes estimated in the second reanalysis are substantially larger than those of the first campaign. This further confirms the findings of King and Williams (2009), Dmitrieva et al. (2015) and Amiri-Simkooei et al. (2013), who identified significant random walk noise in GPS time series. As a non-stationary noise process, the variance increases over time under a random walk process. The zero amplitude of random walk in the first reprocessing campaign is likely because this noise process is being masked (or underestimated) in the ‘processing’ noise due to the lack of the new appropriate models and strategies used in the second reprocessing campaign.
- To further support the statement of the previous point, we present the detrended data (i.e. the mean residuals)

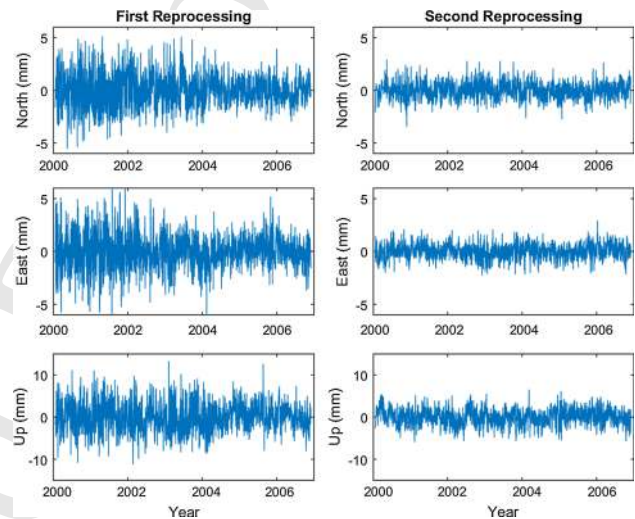


Fig. 4 Mean residuals (for the data set with the time span of 7 years) of time series for north, east, and up components after removing a linear trend, 3 harmonics of annual signal and 10 draconitic harmonics; (left) first reprocessing campaign; (right) second reprocessing campaign

of all 89 stations for these two reprocessing campaigns (Fig. 4). In contrast to the series derived from the first reanalysis, the noise of the new time series has not significantly changed over time as the latest models were used in the second reprocessing. Having a uniform ‘processing’ noise over time allows one to efficiently detect the possible non-stationary random walk noise process due to monument instability (see also Santamaría-Gómez et al. 2011).

To estimate rate errors induced by white, flicker, and random walk noise in the multivariate model, we employ a method described in “Appendix 2”. Using Eqs. (7)–(9), the rate errors of different noise structures have been estimated for the north, east, and up components (Table 4); the rate errors are determined for the data set with the time span of 7 years. It is observed that random walk rate error is larger than those of white and flicker noise. These results are in good agreement

Table 4 White flicker and random walk noise rate errors using three types of formulas; left: this contribution with white ($Q = s_w Q_w$), flicker ($Q = s_f Q_f$) and random walk noise ($Q = s_{rw} Q_{rw}$) and those obtained

using all noise components ($Q = s_w Q_w + s_f Q_f + s_{rw} Q_{rw}$); middle: Bos et al. (2008); right: Argus (2012)

Noise component	Error rates (mm/year)								
	This contribution			Bos et al. (2008)			Argus (2012)		
	N	E	U	N	E	U	N	E	U
White	0.013	0.013	0.044	0.014	0.014	0.047	0.012	0.012	0.040
Flicker	0.126	0.121	0.412	0.152	0.147	0.500	0.136	0.131	0.445
Random walk	0.092	0.090	0.301	0.097	0.093	0.316	0.091	0.088	0.298
White + Flicker + Random walk	0.160	0.155	0.525	–	–	–	–	–	–

In Argus (2012), the formula for the error in rate generated by white noise is missing a factor of $\left(\frac{12}{f}\right)^{\frac{1}{2}}$. The correct formula is $\sigma_{wh} = \left(\frac{12}{f}\right)^{\frac{1}{2}} \frac{s_{wh}}{T^{\frac{3}{2}}}$. In this table we use this corrected formula in the Argus (2012) column. The data set used consisted of 89 stations and 7 years of data ($T = 7$ years) with equal sampling frequency

with those obtained using Eqs. (30)–(31) of Bos et al. (2008) (see Table 4). We may also employ Eqs. (1)–(3) of Argus (2012), originated from Williams (2003a) and Bos et al. (2012), to calculate the rate errors (substitute $T = 7$ years and $f = 365$). Rate errors determined by employing these equations are also shown in Table 4. The last row of Table 4 presents the rate errors using the combination of all noise components.

The (large) amplitude of the random walk compared to those reported by King and Williams (2009) and Dmitrieva et al. (2015) can be explained as follows. It has been shown that white and flicker noise have nearly identical spatial correlation (Amiri-Simkooei 2009). However, random walk noise does not show such a significant correlation because this noise depends on site-related effects such as monument instability, etc. The Kronecker structure used in Amiri-Simkooei (2013a) will induce also significant spatial correlation for random walk. A sub-optimal stochastic model can bias (i.e. overestimates or underestimate) the estimated variance components (Amiri-Simkooei et al. 2009, see Eq. 33). This highlights again that the estimated random walk amplitudes of the multivariate analysis provide only a general indication of a single network-based random walk value.

3.3 Multivariate power spectrum

The multivariate power spectra (MPS), illustrated in Figs. 6, 7, 8 and the top frame of Fig. 5, are obtained using Eq. (8) of Amiri-Simkooei (2013a). The power spectrum would be flat if: (1) there were only white noise in the series, or, (2) the correct stochastic model $\Sigma \otimes Q$ were used. Both spectra shown in Fig. 5 are obtained when taking the temporal correlation of the series (estimated Q) into consideration. The spectrum at the top is derived assuming the series are spatially correlated (correct $\Sigma \otimes Q$), while the bottom frame

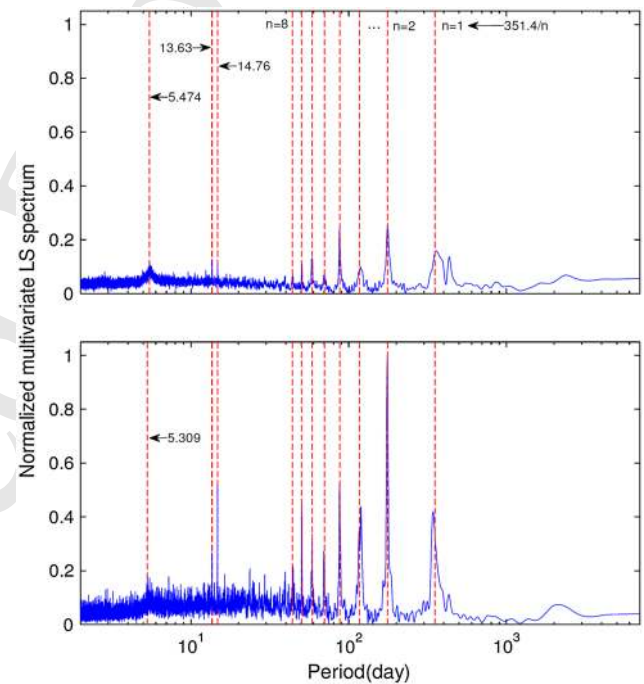


Fig. 5 Multivariate least-squares power spectrum for the data set with the time span of 7 years. Vertical axes are normalized with respect to spectral values of bottom frame to provide the maximum power of one; (top) full structure of $\Sigma \otimes Q$ is taken into consideration, and (bottom) Σ is considered to be diagonal

is derived assuming that the spatial correlation is absent, i.e. $\Sigma = \text{diag}(\sigma_{11}, \dots, \sigma_{rr})$ is a diagonal matrix. The bottom frame is similar to the weighted power spectrum in the studies of Amiri-Simkooei et al. (2007) and Ray et al. (2008, 2013), but differs in that it is based on the correct Q , rather than stationary white noise. Therefore, in contrast to their spectra, our spectra is nearly flat. This indicates that the matrix Q , which compensates for the temporal correlation of the series, affects the flatness of the spectrum, whereas the spatial correlation (matrix Σ) affects the scale of the spectrum. Therefore,

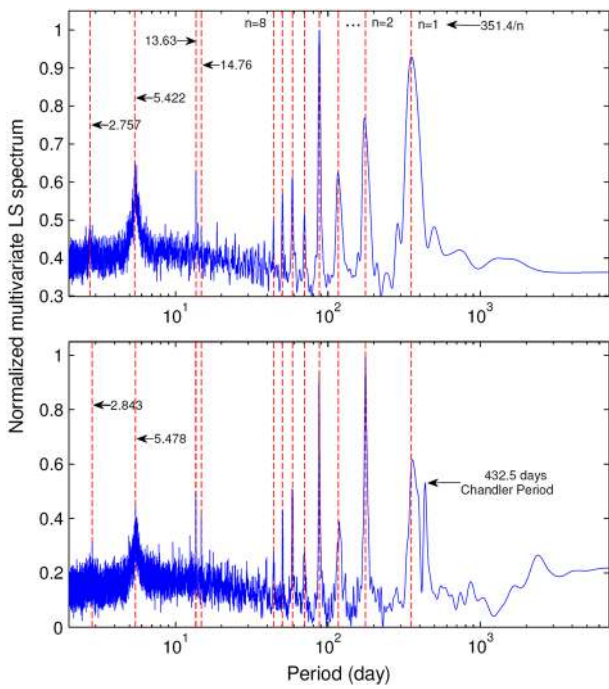


Fig. 6 Multivariate least-squares power spectrum on all coordinate components. *Vertical axes* are normalized to provide the maximum power of one; (*top frame*) data set with the time span of 7 years, (*bottom frame*) data set with the time span of 21 years

a mature stochastic model is crucial for the correct detection of signals. When employing an immature stochastic model, one takes the risk of not detecting peaks at higher frequencies (see Fig. 5); cluster of periods between 5 and 6 days, present in the top frame, are absent in the bottom frame.

The MPS in Fig. 6 shows signals with periods of 13.63 days (direct tides) and 14.76 days (direct 14.77 days tide or 24-h alias of M2). These signals are also detected in the MPS on individual components for both data sets (Fig. 7). It can be seen that the former signal is sharper in the bottom frame of Fig. 6 and the left frame of Fig. 7. The 14.76-day signal was not clearly observed in the up component of the data set with 66 stations. The signals detected for the east and north components are in good agreements with those reported by Ray et al. (2013). They, however, found that fortnightly signals are much less distinct in the up components. Our observations show that this holds indeed only for the 14.76-day signal.

The vertical dashed lines in Figs. 5, 6, 7, 8 illustrate harmonics of the GPS draconitic signal with the periods of $351.4/N$ days ($1.04N$ cpy) for $N = 1, \dots, 8$. The peaks match nearly all of the frequencies. The aliasing signal can contribute to parts of this draconitic signal. Errors in GPS satellite orbit are considered to be the origin for the harmonics because the GPS draconitic year is intrinsic to the

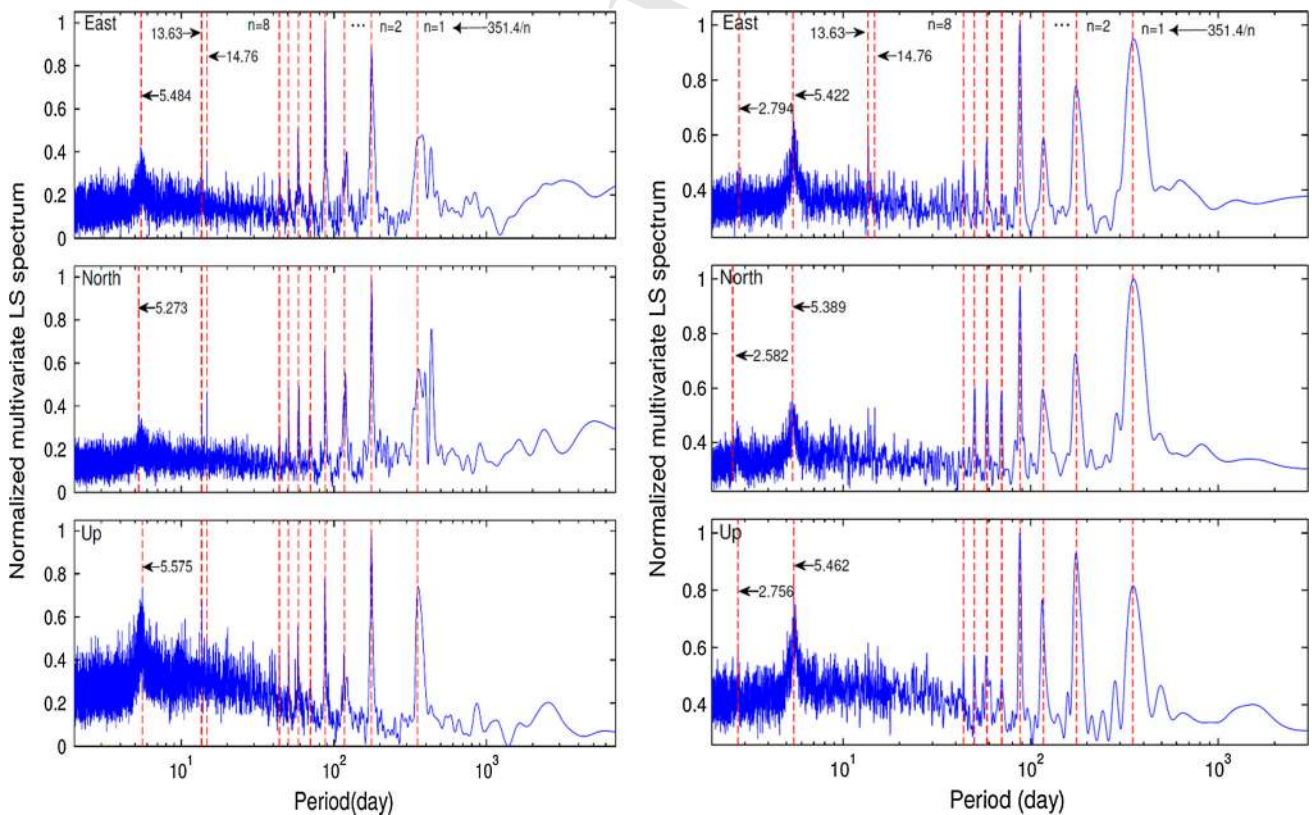


Fig. 7 Multivariate least-squares power spectrum analysis on individual components. *Vertical axes* are normalized to provide the maximum power of one; (*right frame*) data set with the time span of 7 years, (*left frame*) data set with the time span of 21 years

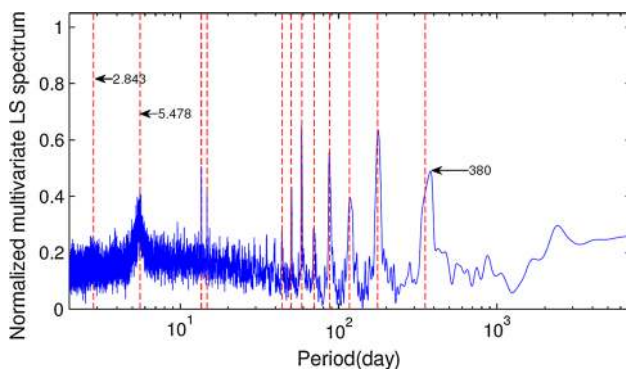


Fig. 8 Multivariate least-squares power spectrum after removing Chandlerian, annual, semiannual, tri-annual, and 8 harmonics of GPS draconitic year for the data set consisting 66 GPS stations (21 years of data). *Vertical axes* are normalized to provide the maximum power of one

satellite orbits, and hence they provide a mechanism for the generation of harmonic modulations. As an example, [Rodríguez-Solano et al. \(2011\)](#) slightly reduced the level of the sixth draconitic harmonic by taking earth albedo and thermal effects on GPS orbits into consideration. For more information, we refer to [Ray et al. \(2008\)](#), [Tregoning and Watson \(2009\)](#), [King and Watson \(2010\)](#), and [Griffiths and Ray \(2013\)](#). [Amiri-Simkooei \(2013a\)](#) shows that a similar behavior of the draconitic pattern at adjacent stations implies that the dominant draconitic effect is not likely dependent on the station-related local effects—multipath for instance. Because the GPS orbit modeling has been improved in latest reanalysis campaign using the new models for Earth radiation pressure and Earth albedo radiation, the reduction in the draconitic signal is expected. This issue will be considered later in Sect. 3.5.

[Amiri-Simkooei \(2013a\)](#) found, contrary to expected, that the first draconitic harmonic in Figs. 5, 6, 7, 8 does not have the largest and sharpest peak, owing to leakage. According to the Rayleigh criterion ([Godin 1972](#)), in order to clearly distinguish between two signals with the periods of T_1 and T_2 , the time spans of the series should be at least equal to $\frac{T_1 T_2}{T_2 - T_1}$. Applying this formula to the annual and draconitic signals with the periods of 365.25 and 351.4 days, respectively, we find that the minimum length of the time series should be equal to 25.4 years. This holds only in theory, but in reality longer time series are required because the above signals are (much) messier than the pure sinusoidal waves. If the time series are not long enough, the annual signal is leaked into the draconitic signal and prohibits it from having the largest and sharpest peak. This is, however, not the case for the higher harmonics of this periodic pattern as the length of the time series exceeds that of the minimum time span required. A sharper peak of the first harmonic in the bottom frame of Fig. 6 and the left frame of Fig. 7 in which

longer data span (21 years) have been used verifies the above statement. Compared to [Amiri-Simkooei \(2013a\)](#), the number of draconitic harmonics detected has been reduced from 10 to 8.

The multivariate analysis is applied both to the individual components (Fig. 7) and simultaneously to the three components (Fig. 6). Both spectra show a cluster of periods around 5.5 days. Using daily time series of 306 IGS stations, [Ray et al. \(2013\)](#) detected a signal with this period in the north and up residuals, but barely visible in the east residuals. We also observe a cluster of periods around 2.75 days (likely the second harmonics of 5.5 days) in the data set with 89 stations (Fig. 6, top frame and Fig. 7, right), and to a lesser extent in the data set with 66 stations (Fig. 6, bottom frame). These findings are in agreement with those of [Ray et al. \(2013\)](#). We do not offer an explanation for the origin of these two signals.

[Selle et al. \(2014\)](#) reprocessed six stations in which a large 5.5 days feature has been found. They used the same orbit, clock product and GIPSY software as the JPL GPS PPP time series, but with a different processing strategy which results in a significant reduction in the strength of the 5.5 days feature. Their result suggested that this signal is both station dependent and probably related to parts of PPP processing strategy other than orbit and clock products or the GIPSY-OASIS software. Therefore, further research is needed for investigation into the origin of the 5.5 days feature in the JPL time series.

Apart from the detected signals discussed earlier, a signal with a period of 432.5 days referred to as Chandler wobble period has been found (Fig. 6, bottom frame). The amplitude of the Chandlerian signal (averaged over 66 stations) for the east, north, and up components are 0.2, 0.2, and 0.4 mm, respectively (Table 5), and the maximum amplitude of this signal for the up and east components reaches nearly 1.2 mm. [Nikolaidis \(2002\)](#) identified a signal with a period of 439 ± 15 days in the power spectrum of the GPS position time series residuals derived from the SOPAC network. It was attributed to the unmodeled pole tide. Moreover, the amplitude of the first Chandlerian harmonics obtained by [Bogusz and Klos \(2016\)](#) was nearly 1 mm for the up component. [Collilieux et al. \(2007\)](#) identified a broad range of frequencies between 0.75 and 0.9 cpy in SLR height residuals from the ITRF2005 solution. The existence of this signal may indicate mismodeling of the Chandler period and its modulations ([Bogusz and Klos 2016](#)) on GPS time series. As the minimum time span needed for the identification of Chandlerian signal is 12 years, the signal has not been detected in the data set with the time span of 7 years. The Chandlerian signal, which is likely related to International Earth Rotation Service's (IERS) pole tide model ([Wahr 1985](#); [King and Watson 2014](#)), has not been reported in any of the IGS AC stacked spectra (including JPL) by [Rebischung et al. \(2016\)](#).

Table 5 The mean and maximum range of variations of the 3 annual harmonics, the 3 draconitic harmonics separately, the 8 draconitic harmonics, the Chandlerian signal and the signal with a period of 383 days for the north, east and up components of the data set with 66 permanent GPS stations of the second reprocessing campaign

Signal	Mean range (mm)			Maximum range (mm)		
	N	E	U	N	E	U
Annual	0.8	1.0	2.4	2.0	2.1	5.3
Semiannual	0.3	0.2	1.1	0.7	0.6	1.9
Tri-annual	0.1	0.1	0.3	0.4	0.3	0.7
Draconitic	0.3	0.4	0.7	0.7	0.9	2.7
Semi-draconitic	0.3	0.4	0.9	0.6	0.8	1.7
Tri-draconitic	0.2	0.1	0.4	0.4	0.3	1.2
All 8 draconitic	0.8	0.9	2.0	1.5	1.5	3.7
Chandlerian	0.2	0.2	0.4	0.5	1.2	1.2
383 days	0.3	0.3	0.6	1.0	0.7	3.0

3.4 Draconitic periodic pattern

This section investigates the GPS draconitic year signal. Following Amiri-Simkooei (2013a), in the linear model $y = Ax$, one can partition A and x as $[A_1 A_2]$ and $[x_1^T x_2^T]^T$, respectively, where x_1 is the unknown parameters of linear term plus annual, semiannual, and tri-annual signals and x_2 is the unknowns of the 8 harmonics of draconitic year signal. Using $y_2 = A_2 x_2$, one can investigate the signal estimated for the draconitic signal. Assume we have r time series. All estimated y_2 vectors of individual time series can be collected in an $m \times r$ matrix $Y_2 = A_2 X_2$, where m is number of observations in the time series.

An investigation on Y_2 (for the data set consisting 89 GPS stations with the time span of 7 years) indicates that the mean range of variations of the draconitic signal reaches -1.91 – 1.91 , -1.75 – 1.73 and -4.72 – 4.72 mm for the north, east, and up components, respectively. They are the amplitudes (average of all minima and maxima over all GPS stations) of the draconitic signal. Compared to the first reprocessing campaign, the mean range of variations for the north, east, and up components are reduced by factors of 1.87, 1.87, and 1.68, respectively.

This reduction stems from the combined effect of the new models used. As an example, Rodriguez-Solano et al. (2012) found that the inclusion of the Earth radiation pressure model causes a change in the north component position estimates at a submillimeter level. The effect of their proposed method has a main frequency of around six cpy, and hence a reduction of 38% occurs by applying this model. Within the latest reprocessing campaign, the UT1 libration effect has been considered, which can result in the reduction in the ranges of variations.

To clearly observe the harmonics of the draconitic signal, the 3 harmonics of the annual signal have been considered in the initial functional model. That is, the functional model consists of 8 columns (2 columns for the linear regression and 2 columns for each annual harmonics). To compare the relative oscillations of the annual and draconitic signal, we have analyzed the original data without considering the 3 annual harmonics. The investigation has been done on the time series with the time span of 21 years as in the time series with the time span of 7 years it is not possible to analyze both annual and draconitic signal (due to the shortness of the time series). The results are presented in Table 5.

The mean annual variations of the north, east, and up components are larger than those of the draconitic by factors ranging from 2.5 to 3.4. The maximum annual variations are larger than those of the semiannual by a factor ranging from 2.78 to 3.5. The annual oscillation is due to exchange of ice, snow, water, and atmosphere, mainly between the northern and southern hemispheres (Blewitt et al. 2001).

We would intuitively expect the spectrum not to show any peak around the annual signal if we were to remove 8 harmonics of the GPS draconitic year signal and the first harmonic of the Chandler wobble in addition to 3 harmonics of the annual signal. To examine our hypothesis, these signals are added to the functional model and the noise assessment was carried out and the correct matrices $\Sigma \otimes Q$ were estimated. The spectral values were then computed. Figure 8 shows the MPS for 66 stations after removing the signals mentioned above. Although the spectral values of 8 harmonics of the draconitic signals have been reduced compared to the bottom frame of Fig. 6, they are not totally removed. This indicates that the draconitic pattern is not completely of periodic nature. Moreover, a signal with a period of around 380 days has been detected, which was not previously observed. This signal is statistically significant because its spectral value (i.e. 412.56) is much larger than the critical value of $\chi_{0.99, 2 \times 66}^2 = 172.71$. We do not have an explanation for this. But it may correspond to the findings of Griffiths and Ray (2013), who computed the Doodson number 165.545 with the period of 23.9379816 h aliases into the period of 385.98 days when the 1-day sampling is used. As expected, this signal has not been observed in the data set with the time span of 7 years as the minimum length of the time series required for distinguishing between this signal and draconitic is 12.7 years (to clearly detect this signal and the annual signal at least 25.7 years of data is needed). The variations of the signal observed for the east, north, and up components of the 66 GPS stations are 0.3, 0.3, and 0.6 mm, respectively (Table 5). The variation of this signal is larger than those of the Chandlerian, tri-annual, and the third draconitic harmonics. The maximum variations of this signal for the up components is larger than those of the first draconitic and the semiannual signal (Table 5).

For further investigation of this phenomenon, two kinds of results are presented in the subsequent subsections.

3.4.1 Visual inspection

We now investigate the possible draconitic peak reduction in the data derived from the 2nd reprocessing campaign. The data sets analyzed consist of 89 GPS stations with the time span of 7 years acquired from the first and second reprocessing campaigns. Using Eq. (8) of Amiri-Simkooei (2013a), the MPS is obtained (Fig. 9). The first, fourth, sixth, and eighth draconitic peaks have been reduced by less than 15%. The third draconitic harmonic experienced a significant reduction; it has been nearly halved. The reduction in the second and fifth draconitic peaks was nearly 25%. It can thus be concluded that using new models within the second reprocessing campaign resulted in the reduction in the draconitic peaks.

To investigate the behavior of the draconitic signal on different GPS stations, we use visual inspection. Figures 10 and 11 represent typical examples on the nature of the draconitic signal for two nearby and two faraway GPS permanent stations, respectively. As expected (see Amiri-Simkooei 2013a), this signal is of similar pattern for nearby stations (<10km) (Fig. 10, compare red or black curves for each

component of stations CIT1 and OXYC). However, for two faraway stations (> 10, 000 km), this statement does not hold true (Fig. 11). The effect is thus location dependent, which originates from the CMEs. But, they are not likely station dependent, and hence multipath cannot be the main source. As expected, this periodic pattern for the 2nd reprocessing campaign (black curve) has been reduced compared to that for the first reprocessing campaign (red curve).

3.4.2 Correlation analysis

The behavior of this periodic pattern can be investigated using the correlation analysis. For this purpose, first we form a zero-mean time series by using all sinusoidal functions of the draconitic signal over one full cycle and collect them in the matrix Y of order $m \times r$. The spatial correlation induced by the matrix Y can be obtained using $\frac{Y^T Y}{m}$. Figure 12 presents the results for the data sets with 89 stations. The spatial correlation induced by the draconitic signal is significant over the angular distance ranging from 0° to 20° (2000 km). This is in agreement with the findings of Amiri-Simkooei (2013a). Therefore, this also indicates that this periodic pattern has still common-mode signatures for the adjacent stations.

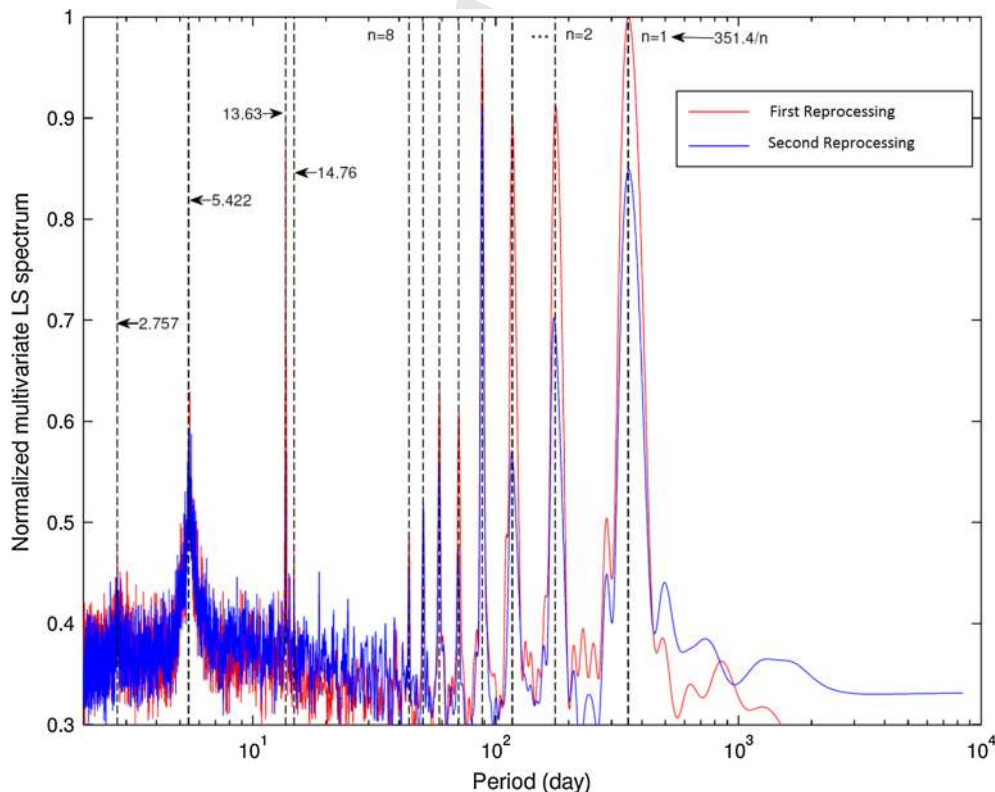


Fig. 9 Multivariate least-squares power spectrum for the data set with the time span of 7 years for first reprocessing (red) and second reprocessing (blue) campaign. Vertical axes are normalized with respect to the spectral values of the first reprocessing campaign (dashed red) to have the maximum power of one

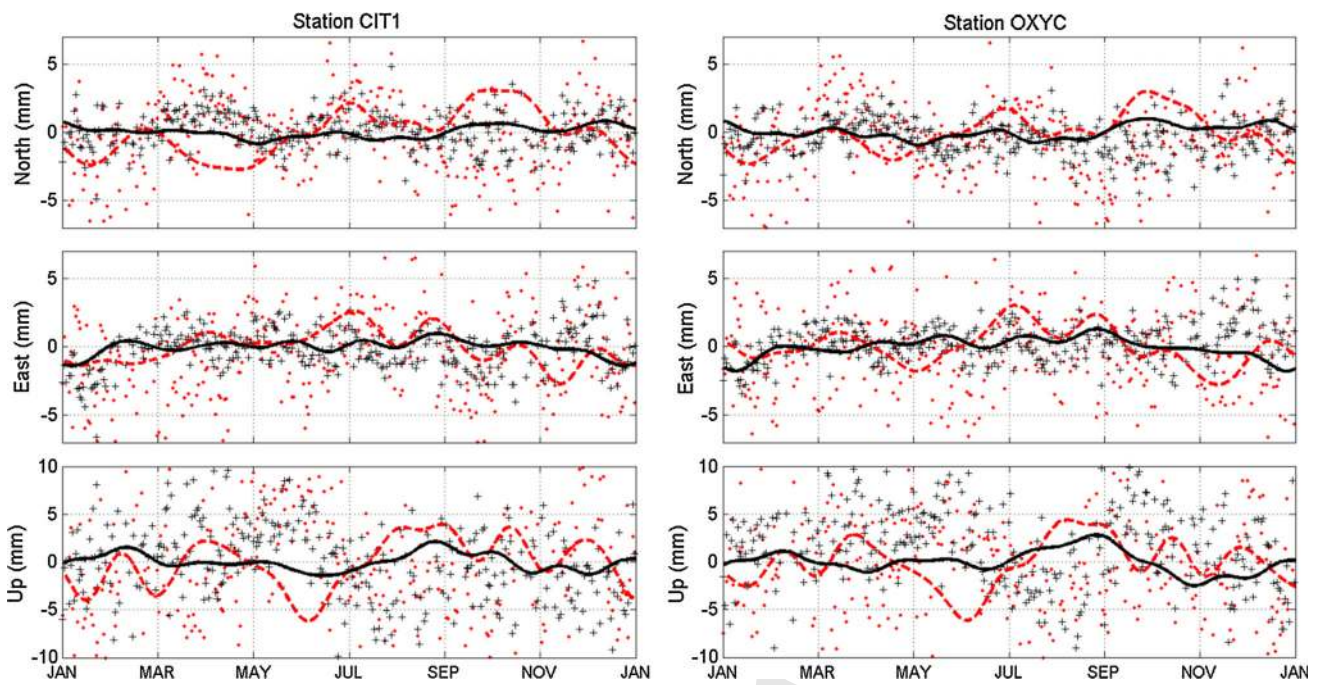


Fig. 10 Effect of periodic pattern of first reprocessing (*red*) and second reprocessing (*black*) campaign estimated for a typical example in which the stations are close to each other. CIT1 is the site at California Institute of Technology. OXYC is the site at Occidental College. OXYC and CIT1 are 7 Km apart. The *red* and *black* points denote the residual time series

after subtracting liner regression terms plus 3 harmonics of the annual signal for first and second reprocessing campaigns, respectively. The *dashed red* and *solid black lines* denote the draconitic signal estimated for the first and second reprocessing campaigns, respectively

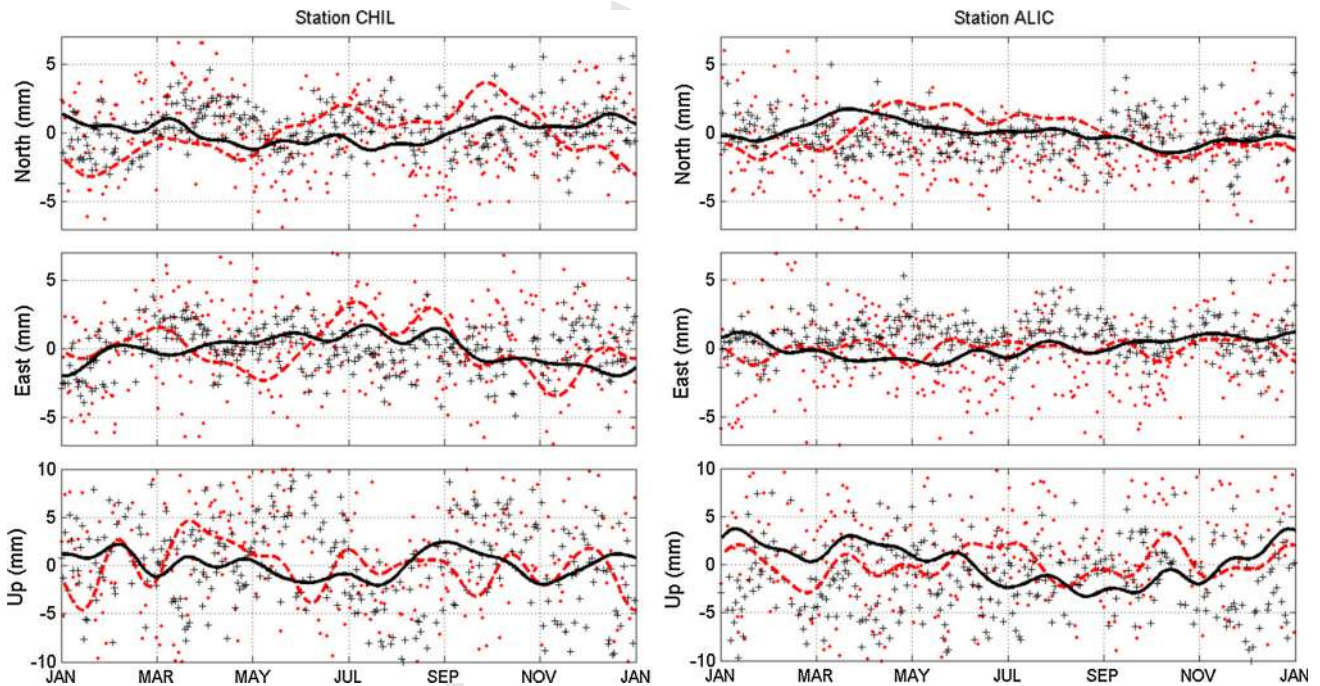


Fig. 11 Effect of periodic pattern of first reprocessing (*red*) and second reprocessing (*black*) campaign estimated for a typical example (CHIL versus ALIC) in which stations are far from each other. CHIL is the site at San Gabriel Mountains, US. ALIC is the site at Alice Springs, Australia. The two sites are 13,000 Km apart. The *red* and *black* points

denote the residual time series after subtracting liner regression terms plus 3 harmonics of the annual signal for first and second reprocessing campaigns, respectively. The *dashed red* and *solid black lines* denote the draconitic signal estimated for the first and second reprocessing campaigns, respectively

3.5 Geodetic and geophysical impact of new time series

This contribution showed improvement on both the functional and stochastic models of GPS position time series of the second reprocessing campaign. Parts of geodetic and geophysical impacts of these improvements are highlighted as follows:

- There is research ongoing in the field of Earth's elastic deformation response to ocean tidal loading (OTL) using kinematic GPS observations. [Martens et al. \(2016\)](#) estimated GPS positions at 5-min intervals using PPP. They studied the dominant astronomical tidal constituents and computed the OTL-induced surface displacements of each component. Such kinematic GPS processing can have many other geophysical applications. Precise determination of Love numbers, as dimensionless parameters characterizing the elastic deformation of Earth due to body forces and loads, is considered to be another application. Therefore, as a direct effect of the new time series, one would expect further improvements in the realization of such geophysical applications.
- GPS position time series have been widely used to study various geophysical phenomena such as plate tectonics, crustal deformation, post-glacial rebound, surface subsidence, and sea-level change ([Thatcher 2003](#); [Argus et al. 2010](#); [Kreemer et al. 2014](#); [Johansson et al. 2002](#); [King et al. 2010](#); [Peltier et al. 2015](#); [Wöppelmann et al. 2007](#); [Lü et al. 2008](#); [Bock et al. 2012](#)). Long-term homogeneous time series reanalysis using the new methods and strategies will directly affect all such phenomena—site velocities along with their uncertainties for instance. Reduction in noise components and the GPS draconitic effect allows other signals to be detected (for example signals with periods of 432.5 and 380 days). More appropriate geophysical interpretation can thus directly be expected, although many of the above references use position time series with CME filtering and hence such signals can be attenuated relative to the “global” solutions discussed in this paper.
- Strain analysis using permanent GPS networks requires proper analysis of time series in which all functional effects are taken into consideration and all stochastic effects are captured using an appropriate noise model. To investigate the effect of the normalized strain parameters on geophysical interpretation, we may recall the statistics theory on the significance of the estimated parameters. To have a statistically significant parameter, one has to compare the parameter with its standard deviation. Flicker noise is the main contributor to make these parameters insignificant ([Razeghi et al. 2015](#)). Reduction in flicker noise has thus a direct impact on the significance of the deformation parameters.

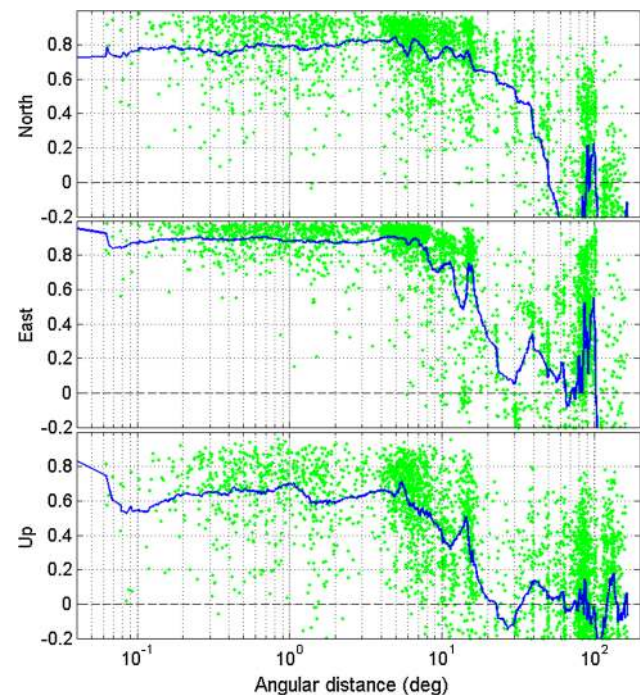


Fig. 12 Spatial correlation originated from draconitic signal of three coordinate components (*north*, *east*, and *up*) for the data set with the time span of 7 years

- Reduction in colored noise, their spatial correlation, and the GPS draconitic signal have significant benefits on the realization of International Terrestrial Reference Frame (ITRF). These improvements will significantly affect the estimation of the parameters of interest and their uncertainty ([Altamimi and Collilieux 2009](#)). They indicated that “IGS is undertaking a great effort of reprocessing the entire time span of the GPS observations with the aim to produce a long-term homogeneous time series. Preliminary analysis of some reprocessed solutions indicates a high performance of these solutions which will play a significant role in the next ITRF release”. This came true based on the results presented in this contribution.

4 Conclusions

This contribution compared the results of the processing the data derived from the first and second reanalysis campaigns to identify the areas of improvement and/or possible degradation. Daily position time series of 89 (7 years) and 66 (21 years) permanent GPS stations, obtained from the JPL second reprocessing campaign, were analyzed. The former data sets were also derived from the first reprocessing campaign to compare the possible improvements in the most realistic manner. Spatial and temporal correlations and MPS were obtained using the formulas and methodologies presented by [Amiri-Simkooei \(2013a\)](#). The following conclusions are drawn:

- 768 • Although the time series of the second reprocessing campaign showed reduction in the spatial correlation among
769 the series by a factor of 1.25, it is nevertheless significant.
770 The spatial cross-correlation also decreases; it is less than
771 0.1 for the three coordinate components.
772
- 773 • The amplitudes of white noise and flicker noise are
774 reduced by factors ranging from 1.40 to 2.33. The random
775 walk amplitudes are higher than the zero values deter-
776 mined for the first reanalysis campaign. This is likely
777 due to the new time series benefiting from a kind of
778 uniform ‘processing’ noise over time, while the noise
779 of the older series is reduced with time. As a result of
780 the revised analysis techniques, the random walk noise
781 has been detected. Further, white and flicker noise have
782 significantly reduced resulting in better detection of the
783 random walk noise amplitude. For the 89 permanent GPS
784 stations with 7 years of data, white noise, flicker noise,
785 and random walk noise rate errors are 0.01, 0.12, and
786 0.09 mm/yr, respectively, for the horizontal component.
787 The vertical rate errors are larger than those of the hori-
788 zontal by the factors ranging from 3.33 to 4.
- 789 • Unlike the results derived from the first reprocessing cam-
790 paign, the noise amplitude of the north component equals
791 that of the east. This is attributed to incorporating the new
792 model for the tropospheric delay and to taking the higher-
793 order ionospheric terms into consideration, which likely
794 improves ambiguity resolution.
- 795 • Both MPS applied to the three components and to the
796 individual components clearly show signals with periods
797 of 13.63 and 14.76 days. In addition, the spectra show a
798 cluster of periods around 5.5 days. A cluster of periods
799 around 2.75 days has been identified in the data set with
800 89 (7 years) and 66 (21 years) GPS stations. Regarding the
801 signals with lower frequencies, a significant signal with
802 period of around 351.4 days (up to its eighth harmon-
803 ics) is detected. This closely follows the GPS draconitic
804 year. Two other signals with periods of nearly 432.5 and
805 380 days have been found. While the period of the former
806 signal equals the well-known Chandler period, the latter
807 signal is not known.
- 808 • The mean range of variations (max and min) of the dra-
809 conitic pattern for the series derived from the second
810 reprocessing campaign shows a reduction of 46, 46 and
811 41% for the north, east, and up components, respectively,
812 compared to those of the first campaign. This significant
813 reduction can be a direct corollary of the improved mod-
814 els in the new campaign. While the first, fourth, sixth,
815 and eight draconitic peaks have been reduced by less
816 than 15%, the third draconitic harmonic has been nearly
817 halved. The reduction in the second and fifth draconitic
818 peaks was nearly 25%.
- 819 • Two independent measures of visual inspection and cor-
820 relation analysis were used to investigate the nature of

the draconitic pattern. While the effect of the draconitic
signal is of similar pattern for nearby stations (Fig. 10),
it differs significantly for distant stations (Fig. 11). The
periodic pattern was reduced in the second reanalysis
campaign.

- 821 • A similar behavior for the spatial correlation of the
822 time series (Fig. 2) and the periodic pattern (Fig. 12)
823 is observed. This indicates that although new models and
824 methodologies in the latest reanalysis have reduced the
825 spatial correlation among the series to an extent, the dra-
826 conitic pattern is still an error source inducing spatial
827 correlation to the time series.
828
- 829 • There are three factors that may prevent random walk to
830 be detected. The first is the dominance of flicker noise,
831 which masks random walk noise (Williams et al. 2004).
832 Flicker noise has been significantly reduced in the sec-
833 ond reprocessing. The second factor is the small length
834 of the time series. For some stations, however, there are
835 currently more than two decades of data. A few pre-
836 liminary tests confirm significant random walk noise on
837 longer time series. 3) The third factor originates from our
838 observation in this contribution, which states that second
839 reprocessing has not only reduced noise but also it shows
840 a kind of uniform processing noise over time (see Fig. 4).
841 These three factors thus indicate that random walk noise
842 can in principle be the subject of the intensive research
843 in future GPS position time series analysis.
844
845
846
847

Acknowledgements We are grateful to JPL’s GPS data analysis team
for the GPS position time series and to Christina Selle for an informal
review. Support for JPL’s time series came from NASA’s Space Geodesy
Project and MEaSURES program. D.F. Argus’s part of this research
was performed at the Jet Propulsion Laboratory, California Institute of
Technology, under a contract with the National Aeronautics and Space
Administration. We are also grateful to the associate editor, Prof. Matt
King, and three anonymous reviewers for their detailed comments which
improved the quality and presentation of this paper.

5 Appendix 1: models employed within the second IGS reanalysis campaign

5.1 Yaw attitude variations

Inconsistent yaw attitude models affects the precision of the
IGS combined clock solutions (Hesselbarth and Wanninger
2008). Therefore, the reliability of the IGS combined clocks
is impaired. To diminish the effect of the eclipsing satellites
on the IGS clock solutions, consistent modeling of attitude
changes is needed (Ray 2009). Distortions in the orientation
of the eclipsing satellites follow a simplified yaw attitude
model for Block II/IIA and Block IIR satellites (see Kouba
2009a). Attitude behavior of the Block IIF-1 (launched on
May 27, 2010) spacecraft during the eclipse has been studied

by [Dilssner \(2010\)](#). In addition, with complete modernization of the GLONASS satellites, ACs should include GLONASS observations as well. An appropriate yaw attitude modeling of these satellites may follow the model proposed by [Dilssner et al. \(2011\)](#).

5.2 Modeling of orbit dynamics

[Urschl et al. \(2007\)](#) observed anomalous pattern in the plot of GPS-SLR residuals which they attributed to the GPS orbit mismodeling. This anomalous pattern (particularly, the GPS draconitic year signal) was also identified in the geocenter Z-component ([Hugentobler et al. 2006](#)) and GPS position time series ([Ray et al. 2008](#)).

One of the potential sources for GNSS orbit mismodeling is the deficiencies in the Earth radiation pressure (ERP) model. Not all IGS ACs are yet modeling ERP. Utilizing a model for Earth radiation, proposed by [Rodriguez-Solano et al. \(2012\)](#), results in the reduction in root mean square (RMS) of orbit's height component by about 1–2 cm and smaller perturbations of other components related to the orbit. [Rodriguez-Solano et al. \(2012\)](#) showed that the model can compensate the SLR residual bias observed.

The GPS orbit perturbations due to ERP depend on the relative position of Sun, Earth, and satellite. Parts of the observed periodic patterns in GPS time series may stem from failure to correctly model ERP ([Rodriguez-Solano et al. 2012](#)). They found that the inclusion of the ERP model results in the reduction in the sixth draconitic signal for the north component at a submillimeter level (equal to reduction of around 38%). [Ray \(2009\)](#) also suggested taking ERP into consideration. Hence, the model proposed by [Rodriguez-Solano et al. \(2011\)](#) has been used within the IGS in the operational reprocessing.

Earth albedo radiation (EAR) is another source for orbit modeling deficiencies. This radiation consists of both visible reflected light and infrared emitted radiation. Most AC contributors have not taken into account the effect of EAR. The albedo acceleration may have a significant effect on the orbit of GPS satellites (a mean reduction in the orbit radial component by 1–2 cm) ([Hugentobler et al. 2009](#)). They concluded that for the high-precision GPS orbit determination, EAR and antenna thrust should be taken into consideration. However, regarding the spectra of geocenter and position time series, no significant impact has been observed when the model for EAR was used ([Hugentobler et al. 2009](#)). This indicates that there could be still unmodeled effects on the GPS orbit which can be larger than the albedo radiation.

5.3 Geopotential field

In terms of the geopotential model, a new model referred to as EGM2008 has been defined (see [Ray 2009](#)). EGM2008

exhibits significant improvements compared to its previous counterpart EGM96, thanks to the availability of CHAMP and most importantly GRACE data in the 2000s. Compared to EGM96, used for the 1st processing campaign, EGM2008 has been modified in the following aspects:

1. Its degree and order have been increased by a factor of 6.
2. Updated value for secular rate of low-degree coefficients.
3. A new model for the mean pole trajectory was proposed.
4. Model for geopotential ocean tide has been updated for FES2004.
5. A new ocean pole tide model has been introduced.

For more information, the reader is referred to IERS 2010 conventions ([Petit and Luzum 2010](#)).

5.4 Tidal effects

Tidal effects are categorized to the following two contributions. (1) Tidal displacement of station positions; (2) Tidal EOP variations. For the former, within the new processing campaign, a new model which is introduced for the mean pole trajectory IERS 2010 ([Petit and Luzum 2010](#)) has been used for the pole tide correction. Moreover, model for ocean pole tide loading presented by [Desai \(2002\)](#) should be used. For the latter, the Earth rotation axial component in terms of UT1 contains small diurnal and subdiurnal signals. Thus, the tidal gravitation effect on those features of Earth's mass distribution results in the astronomical precession-nutation of Earth rotation ([Brzeziński 2008](#)). A minor part of the astronomical variations, called libration, is a result of the tidal gravitation effect on the non-zonal terms of geopotential ([Brzeziński 2008](#)). In case of UT1, the perturbation is semidiurnal with total amplitude up to 75 μ as. [Brzeziński and Capitaine \(2009\)](#) studied the subdiurnal libration in UT1. They derived a solution for the structural model of the Earth composing of an elastic mantle and a liquid core not coupling to each other.

A key expectation in tidal EOP variations modeling compared to the 1st reprocessing campaign is the addition of the UT1 libration effect introduced by [Brzeziński and Capitaine \(2009\)](#). It is noted that the maximum effect of UT1 libration is about 105 μ as, or 13 mm at GPS altitude. It probably severely aliases into the orbit parameters.

5.5 Tropospheric propagation delay

In the second reprocessing, a new slant delay model (GPT2) was suggested. It improves its older models GPT/GMF with refined horizontal resolution, enhanced temporal coverage, and increased vertical resolution (37 isobaric levels compared to 23 ones utilized for GPT/GMF) ([Lagler et al. 2013](#)). In addition to mean value, a_0 , and annual amplitude,

966 A, estimated using the least-squares method in GPT/GMF, 1014
 967 semiannual harmonics are incorporated within GPT2. This 1015
 968 better accounts for regions with very rainy or dry peri- 1016
 969 ods. As for the temperature reduction, in contrast to the 1017
 970 GPT/GMF in which a constant $-6.5\text{ }^\circ\text{C}/\text{km}$ was assumed, 1018
 971 mean, annual, and semiannual variations of temperature lapse 1019
 972 rate are determined each grid point in GPT2. Regarding 1020
 973 the pressure reduction, unlike the GPT/GMF which utilizes 1021
 974 an exponential formula based on the standard atmosphere, 1022
 975 GPT2 deploys an exponential formula based on virtual tem- 1023
 976 perature (Lagler et al. 2013). The improved performance 1024
 977 of GPT2 compared to the previous model GPT/GMF has 1025
 978 been examined by Lagler et al. (2013). They have recom- 1026
 979 mended to replace GPT/GMF with GPT2 as an empirical 1027
 980 model. 1028

981 Because of the partial compensation of the atmospheric 1029
 982 loading by mismodeling the zenith hydrostatic delays (ZHDs) 1030
 983 (Kouba 2009b), GPT-derived ZHDs give rise to a better 1031
 984 station height repeatability compared to ECMWF ZHDs if 1032
 985 atmospheric loading is not corrected for (Steigenberger et al. 1033
 986 2009). On the other hand, if one needs to examine the coor- 1034
 987 dinates time series to reveal atmospheric loading signals, 1035
 988 application of ZHDs derived from numerical weather models 1036
 989 is a key element. 1037

990 5.6 Higher-order ionospheric terms

991 A linear combination of multi-frequency observations allows 1038
 992 for taking into consideration the first-order $\sim \frac{1}{f^2}$ ionospheric 1039
 993 term (Hofmann-Wellenhof et al. 2008). The first-order iono- 1040
 994 spheric delay is in the order of 1 to 50 meters, which depends 1041
 995 on the satellite elevation, ionospheric activities, local time, 1042
 996 season and solar cycle (Kedar et al. 2003). The higher-order 1043
 997 ionospheric terms, which are in the order of submillimeters 1044
 998 to several centimeters, are usually neglected. Kedar et al. 1045
 999 (2003) stated that the effect of second-order ionospheric term 1046
 1000 introduced by Bassiri and Hajj (1993) can likely improve the 1047
 1001 position repeatability and reduce the small biases in geocenter 1048
 1002 estimates. Fritsche et al. (2005) and Hernández-Pajares 1049
 1003 et al. (2007) showed that the second-order ionospheric term 1050
 1004 affects the geocenter Z-component estimates. Fritsche et al. 1051
 1005 (2005) processed the double difference phase observation of a 1052
 1006 global network and compared solutions with and without the 1053
 1007 higher-order ionospheric terms. They concluded that apply- 1054
 1008 ing these higher terms will become a standard part of precise 1055
 1009 GPS applications. IERS 2010 conventions (Petit and Luzum 1056
 1010 2010) suggested that while the first- and second-order iono- 1057
 1011 spheric terms are to be considered for GNSS applications, 1058
 1012 the third order is at the limited significance and the fourth 1059
 1013 order can be neglected. 1060

1014 5.7 Analysis constraints

1015 Ferland (2010) found that high-frequency smoothing may be 1016
 1017 due to unremovable continuity constraints for some ACs. Ray 1018
 1019 (2009) suggested that, for the 2nd reprocessing campaign, 1020
 1021 ACs constraints and procedures should be reconsidered 1022
 1023 from the following aspects: (1) Reviewing the necessity of 1024
 1025 applying constraints, (2) Paying particular attention to the 1026
 1027 constraint on the orbit and UT1/LOD, (3) Elimination and 1028
 1029 minimization of the constraints as many as possible, and (4) 1029
 1030 Better understanding of the impacts of constraints retained 1030
 1031 is necessary. Accordingly, in the IGS2008 recommendations 1031
 1032 (<http://igs.org/overview/pubs/IGSWorkshop2008/>), all ACs 1032
 1033 should report their a-priori constraints. Although remov- 1033
 1034 able constraints are acceptable, unconstrained solutions are 1034
 1035 preferred. Inner constraints (origin, orientation, scale) are 1035
 1036 acceptable. 1036

1030 6 Appendix 2: rate errors in multivariate model

1031 Having r time series available, a multivariate linear model is 1031
 1032 of the form (Koch 1999) 1032

$$1033 E(\text{vec}(Y)) = (I_r \otimes A) \text{vec}(X), D(\text{vec}(Y)) = Q_{\text{vec}(Y)} \quad 1033$$

$$1034 \quad (1) \quad 1034$$

1035 where vec is the vector operator and \otimes is the Kronecker prod- 1035
 1036 uct. I_r is the identity matrix of size r . X and Y are the matrices 1036
 1037 of the sizes $n \times r$ and $m \times r$ collecting unknown param- 1037
 1038 eters and observations from r number of series, respectively. 1038
 1039 A and $Q_{\text{vec}(Y)}$ are, respectively, the functional and stochastic 1039
 1040 models describing all deterministic effects and statistical 1040
 1041 characteristics of the observables. E indicates the expectation 1041
 1042 operator, and D is the dispersion operator. 1042

1043 The following structure for the stochastic model, referred 1043
 1044 to as the more practical model, is used (Amiri-Simkooei 1044
 1045 2009) 1045

$$1046 D(\text{vec}(Y)) = \Sigma \otimes Q = \Sigma \otimes \sum_{k=1}^p s_k Q_k \quad 1046$$

$$1047 \quad (2) \quad 1047$$

1048 where Q_k 's are the known cofactor matrices of size $m \times m$. 1048
 1049 The matrix Σ and the unknown factors s_k are to be estimated 1049
 1050 using LS-VCE. 1050

1051 The least-squares estimate of X reads then (Koch 1999) 1051

$$1052 \hat{X} = (A^T Q^{-1} A)^{-1} A^T Q^{-1} Y \quad 1052$$

$$1053 \quad (3) \quad 1053$$

1054 The covariance matrix of the nr -vector $\text{vec}(\hat{X})$ is 1054

$$1055 Q_{\text{vec}(\hat{X})} = \Sigma \otimes (A^T Q^{-1} A)^{-1} = \Sigma \otimes N^{-1} \quad 1055$$

$$1056 \quad (4) \quad 1056$$

where $N = A^T Q^{-1} A$ is the normal matrix. Here, we assume that the functional model contains two columns for the linear regression terms plus two columns for each of the annual, semiannual, and tri-annual signal. A is thus of size $m \times 8$. Its i th row at the time instant t_i is

$$\begin{bmatrix} 1 & t_i & \cos 2\pi t_i & \sin 2\pi t_i & \cos 4\pi t_i & \sin 4\pi t_i & \cos 6\pi t_i & \sin 6\pi t_i \end{bmatrix} \quad (5)$$

Therefore, the unknown parameters are the intercept, rate, and amplitudes of the annual, semiannual, and tri-annual signals. The covariance matrix of the slopes (for all series) is given as $Q_r = \Sigma \times (N^{-1})_{22}$, where $(N^{-1})_{22}$ is the second diagonal element of N^{-1} . It is further assumed that Q matrix has the form

$$Q = s_w I + s_f Q_f + s_{rw} Q_{rw} \quad (6)$$

where s_w , s_f , s_{rw} are the white, flicker, and random walk noise amplitudes, respectively. Q_f and Q_{rw} are the flicker and random walk noise cofactor matrices, respectively. LS-VCE has been employed to estimate s_w , s_f , s_{rw} , and Σ . As the three coordinate components of all stations have been processed simultaneously, Σ is of the size $r \times r$. Its corresponding, north, east, and up components are referred to as Σ_N , Σ_E , and Σ_U , respectively (block diagonals). To compute the white, flicker, and random walk noise rate errors for the east components, matrix Q in Eq. (4) is substituted with $Q_w = s_w I$, $Q_f = s_f Q_f$ or $Q_{rw} = s_{rw} Q_{rw}$, respectively. Matrices N_w , N_f , N_{rw} are then obtained. The rate errors of the east component read

$$\sigma_r^w = \sqrt{\text{diag} \left(\Sigma_E N_w^{-1} (2, 2) \right)} \quad (7)$$

$$\sigma_r^f = \sqrt{\text{diag} \left(\Sigma_E N_f^{-1} (2, 2) \right)} \quad (8)$$

$$\sigma_r^{rw} = \sqrt{\text{diag} \left(\Sigma_E N_{rw}^{-1} (2, 2) \right)} \quad (9)$$

where σ_r^w , σ_r^f and σ_r^{rw} are the vector of rate errors for the east component of all stations. Their mean indicate the average error rates over all stations. The corresponding values for the north and up components can accordingly be obtained.

References

- Altamimi Z, Collilieux X (2009) IGS contribution to the ITRF. *J Geod* 83(3):375–383
- Amiri-Simkooei AR (2007) Least-squares variance component estimation: theory and GPS applications (Ph.D. thesis). Delft University of Technology. Publication on Geodesy. 64. Netherlands Geodetic Commission, Delft
- Amiri-Simkooei AR (2009) Noise in multivariate GPS position time-series. *J Geod* 83(2):175–187

- Amiri-Simkooei AR (2013a) On the nature of GPS draconitic year periodic pattern in multivariate position time series. *J Geophys Res: Solid Earth* 118(5):2500–2511
- Amiri-Simkooei AR (2013b) Application of least-squares variance component estimation to errors-in-variable models. *J Geod* 87(10):935–944
- Amiri-Simkooei AR (2016) Non-negative least squares variance component estimation with application to GPS time series. *J Geod*. doi:10.1007/s00190-016-0886-9
- Amiri-Simkooei AR, Tiberius CCJM, Teunissen PJG (2007) Assessment of noise in GPS coordinate time series: methodology and results. *J Geophys Res: Solid Earth* 112:B07413. doi:10.1029/2006JB004913
- Amiri-Simkooei AR, Teunissen PJG, Tiberius CCJM (2009) Application of least-squares variance component estimation to GPS observables. *J Surv Eng* 135(4):149–160
- Amiri-Simkooei AR, Zangeneh-Nejad F, Asgari J (2013) Least-squares variance component estimation applied to GPS geometry-based observation model. *J Surv Eng* 139(4):176–187
- Argus DF (2012) Uncertainty in the velocity between the mass center and surface of Earth. *J Geophys Res: Solid Earth* 117:B10405. doi:10.1029/2012JB009196
- Argus DF, Fu Y, Landerer FW (2014) GPS as a high resolution technique for evaluating water resources in California. *Res. Lett, Geophys*. doi:10.1002/2014GL059570
- Argus DF, Heflin MB, Peltzer G, Webb FH, Crampe F (2005) Interseismic strain accumulation and anthropogenic motion in metropolitan Los Angeles. *J Geophys Res* 101:B04401. doi:10.1029/2003JB002934
- Argus DF, Gordon RG, Heflin MB, Ma C, Eanes RJ, Willis P, Peltier WR, Owen SE (2010) The angular velocities of the plates and the velocity of Earth's center from space geodesy. *Geophys J Int* 180:913–960. doi:10.1111/j.1365-246X.2009.04463.x
- Bassiri S, Hajj G (1993) High-order ionospheric effects of the global positioning system observables and means of modeling them. *Manuscr Geod* 18:280–289
- Beutler G, Rothacher M, Schaer S, Springer TA, Kouba J, Neilan RE (1999) The International GPS Service (IGS): an interdisciplinary service in support of Earth sciences. *J Adv Space Res* 23(4):631–635
- Blewitt G, Lavallée D, Clarke P, Nurutdinov K (2001) A new global mode of Earth deformation: seasonal cycle detected. *Science* 294:2342–2345
- Bock Y, Nikolaidis RM, de Jonge PJ, Bevis M (2000) Instantaneous geodetic positioning at medium distances with the Global Positioning System. *J Geophys Res: Solid Earth* 105(B12):28223–28253
- Bock Y, Wdowinski S, Ferretti A, Novali F, Fumagalli A (2012) Recent subsidence of the Venice Lagoon from continuous GPS and interferometric synthetic aperture radar. *Geochem Geophys Geosyst* 13(Q03023):2011G. doi:10.1029/C003976
- Bogusz J, Klos A (2016) On the significance of periodic signals in noise analysis of GPS station coordinate time series. *GPS Solut*. 20(4):655–664. doi:10.1007/s10291-015-0478-9
- Bogusz J, Gruszczynski M, Figurski M, Klos A (2015) Spatio-temporal filtering for determination of common mode error in regional GNSS networks. *Open Geosci* 7(1):140–148
- Bonforte A, Puglisi G (2006) Dynamics of the eastern flank of Mt. Etna volcano (Italy) investigated by a dense GPS network. *J Volcanol Geotherm Res* 153(3):357–369
- Bos MS, Fernandes RMS, Williams SDP, Bastos L (2008) Fast error analysis of continuous GPS observations. *J Geod* 82(3):157–166
- Bos MS, Fernandes RMS, Williams SDP, Bastos L (2012) Fast error analysis of continuous GNSS observations with missing data. *J Geod* 87(4):351–360
- Brzeziński A (2008) Recent advances in theoretical modeling and observation of Earth rotation at daily and subdaily periods.

- In: Soffel M, Capitaine N (eds) Proceedings of the Journées 2008 “Systèmes de référence spatio-temporels” & X. Lohrmann-Kolloquium: Astrometry, Geodynamics and Astronomical Reference Systems, TU Dresden, Germany, 22–24 September 2008. Lohrmann-Observatorium and Observatoire de Paris
- Brzeziński A, Capitaine N (2009) Semidiurnal signal in UT1 due to the influence of tidal gravitation on the triaxial structure of the Earth. In: Proceedings of XXVII General Assembly of the International Astronomical Union, 3–14 August. Rio de Janeiro, Brazil
- Calais E (1999) Continuous GPS measurements across the Western Alps, 1996–1998. *Geophys J Int* 138(1):221–230
- Cervelli PF, Fournier T, Freymueller J, Power JA (2006) Ground deformation associated with the precursory unrest and early phases of the January 2006 eruption of Augustine Volcano. *Geophys Res Lett*, Alaska. doi:[10.1029/2006GL027219](https://doi.org/10.1029/2006GL027219)
- Collilieux X, Altamimi Z, Coulot D, Ray J, Sillard P (2007) Comparison of very long baseline interferometry, GPS, and satellite laser ranging height residuals from ITRF2005 using spectral and correlation methods. *J Geophys Res: Solid Earth*. 112(B12403). doi:[10.1029/2007JB004933](https://doi.org/10.1029/2007JB004933)
- Craig TJ, Calais E (2014) Strain accumulation in the New Madrid and Wabash Valley seismic zones from 14 years of continuous GPS observation. *J Geophys Res: Solid Earth* 119(12):9110–9129
- d’Alessio MA, Johanson IA, Bürgmann R, Schmidt DA, Murray MH (2005) Slicing up the San Francisco Bay Area: block kinematics and fault slip rates from GPS-derived surface velocities. *J Geophys Res* 110:B06403. doi:[10.1029/2004JB003496](https://doi.org/10.1029/2004JB003496)
- Desai SD (2002) Observing the pole tide with satellite altimetry. *J Geophys Res: Solid Earth* 107(C11): doi:[10.1029/2001JC001224](https://doi.org/10.1029/2001JC001224)
- Dietrich R, Rulke A, Scheinert M (2005) Present-day vertical crustal deformations in west Greenland from repeated GPS observations. *Geophys J Int* 163(3):865–874
- Dilssner F (2010) GPS IIF-1 satellite: antenna phase and attitude modeling. *Inside GNSS* 5:59–64
- Dilssner F, Springer T, Gienger G, Dow J (2011) The GLONASS-M satellite yaw-attitude model. *J Adv Space Res* 47(1):160–171
- Dmitrieva K, Segal P, DeMets C (2015) Network-based estimation of time-dependent noise in GPS position time series. *J Geod* 89(6):591–606
- Dong D, Fang P, Bock Y, Webb F, Prawirodirdjo L, Kedar S, Jason P (2006) Spatiotemporal filtering using principal component analysis and Karhunen–Loeve expansion approaches for regional GPS network analysis. *J Geophys Res: Solid Earth* 111(B03405). doi:[10.1029/2005JB003806](https://doi.org/10.1029/2005JB003806)
- Ferland R (2010) Combination of the reprocessed IGS Analysis Center SINEX solutions. IGS10. Newcastle upon Tyne, England. 28 June–2 July (http://acc.igs.org/repro1/repro1-combo_IGSW10.pdf)
- Fritsche M, Dietrich R, Knöfel C, Rülke A, Vey S, Rothacher M, Steigenberger P (2005) Impact of higher-order ionospheric terms on GPS estimates. *Geophys Res Lett*. doi:[10.1029/2005GL024342](https://doi.org/10.1029/2005GL024342)
- Gazeaux J, Williams SDP, King MA, Bos M, Dach R, Deo M, Moore AW, Ostini L, Petrie E, Roggero M, Teferle FN, Olivares G, Webb FH (2013) Detecting offsets in GPS time series: first results from the detection of offsets in GPS experiment. *J Geophys Res Solid Earth* 118(5):2397–2407
- Godin G (1972) The analysis of tides. University of Toronto Press, Toronto
- Griffiths J, Ray JR (2013) Sub-daily alias and draconitic errors in the IGS orbit. *GPS Solut* 17(3):413–422
- Hesselbarth A, Wanninger L (2008) Short-term stability of GNSS satellite clocks and its effects on precise point positioning. ION GNSS2008. Savannah, Georgia, USA, September 16–19. pp 1855–1863
- Hernández-Pajares M, Juan JM, Sanz J, Orús R (2007) Second-order ionospheric term in GPS: implementation and impact on geodetic estimates. *J Geophys Res: Solid Earth* 112(B08417): doi:[10.1029/2006JB004707](https://doi.org/10.1029/2006JB004707)
- Hofmann-Wellenhof B, Lichtenegger H, Walse E (2008) GNSS-Global navigation satellite system. Springer Vienna, Vienna
- Hoseini-Asl M, Amiri-Simkooei AR, Asgari J (2013) Offset detection in simulated time-series using multivariate analysis. *J Geom Sci Technol* 3(1):75–86
- Hugentobler U, van der Marel H, Springer T (2006) Identification and mitigation of GNSS errors. In: Springer T, Gendt G, Dow JM (eds) The International GNSS Service (IGS): perspectives and visions for 2010 and beyond, IGS Workshop 2006. ESOC. Darmstadt, Germany. May 6–11
- Hugentobler U, Rodríguez-Solano CJ, Steigenberger P, Dach R, Lutz S (2009) Impact of albedo modelling on GNSS satellite orbits and geodetic time series. In: Eos Transactions American Geophysical Union, 90(52), Fall Meeting. San Francisco, California, USA. December 14–19
- Johansson JM, Koivula H, Vermeer M (2002) Continuous GPS measurements of postglacial adjustment in Fennoscandia, 1, Geodetic results. *J Geophys Res* 107(B8)
- Johnson HO, Agnew DC (2000) Correlated noise in geodetic time series. U.S. Geology Surveying. Final Technical Report, FTR-1434-HQ-97-GR-03155
- Kedar S, Hajj GA, Wilson BD, Heflin MB (2003) The effect of the second order GPS ionospheric correction on receiver positions. *Geophys Res Lett* 30(16)
- King MA, Williams SDP (2009) Apparent stability of GPS monumentation from short-baseline time series. *J Geophys Res: Solid Earth* 114:B10403. doi:[10.1029/2009JB006319](https://doi.org/10.1029/2009JB006319)
- King MA, Watson CS (2010) Long GPS coordinate time series: multipath and geometry effects. *J Geophys Res: Solid Earth* 115:B04403. doi:[10.1029/2009JB006543](https://doi.org/10.1029/2009JB006543)
- King MA, Watson CS (2014) Geodetic vertical velocities affected by recent rapid changes in polar motion. *Geophys J Int* 199(2):1161–1165
- King MA, Altamimi Z, Boehm J, Bos M, Dach R, Elosegui P, Fund F, Hernández-Pajares M, Lavallee D, Cervera PJM et al (2010) Improved constraints on models of glacial isostatic adjustment: a review of the contribution of ground-based geodetic observations. *Surv Geophys* 31(5):465–507
- Khodabandeh A, Amiri-Simkooei AR, Sharifi MA (2012) GPS position time-series analysis based on asymptotic normality of M-estimation. *J Geod* 86(1):15–33
- Koch K-R (1999) Parameter estimation and hypothesis testing in linear model, 2nd edn. Springer, Berlin 333p
- Kouba J (2009a) A simplified yaw-attitude model for eclipsing GPS satellites. *J GPS Solut* 13(1):1–12
- Kouba J (2009b) Testing of global pressure/temperature (GPT) model and global mapping function (GMF) in GPS analyses. *J Geod* 83(3–4):199–208
- Kreemer C, Holt WE, Haines AJ (2003) An integrated global model of present-day plate motions and plate boundary deformation. *Geophys J Int* 154(1):8–34
- Kreemer C, Blewitt G, Klein EC (2014) A geodetic plate motion and Global Strain Rate Model. *Geochem Geophys Geosyst* 15:3849–3889. doi:[10.1002/2014GC005407](https://doi.org/10.1002/2014GC005407)
- Lagler K, Schindelegger M, Böhm J, Krásná H, Nilsson T (2013) GPT2: empirical slant delay model for radio space geodetic techniques. *J Geophys Res Lett* 40(6):1069–1073
- Langbein J (2004) Noise in two-color electronic distance meter measurements revisited. *J Geophys Res: Solid Earth* 109(B04406). doi:[10.1029/2003JB002819](https://doi.org/10.1029/2003JB002819)
- Langbein J (2008) Noise in GPS displacement measurements from southern California and southern Nevada. *J Geophys Res: Solid Earth* 113(B05405). doi:[10.1029/2007JB005247](https://doi.org/10.1029/2007JB005247)

- Langbein J (2012) Estimating rate uncertainty with maximum likelihood: differences between power-law and flicker-random-walk models. *J Geod* 86(9):775–783. doi:[10.1007/s00190-012-0556-5](https://doi.org/10.1007/s00190-012-0556-5)
- Langbein J, Bock Y (2004) High-rate real-time GPS network at Parkfield: Utility for detecting fault slip and seismic displacements. *Geophys Res Lett*. doi:[10.1029/2003GL019408](https://doi.org/10.1029/2003GL019408)
- Lü WC, Cheng SG, Yang HS, Liu DP (2008) Application of GPS technology to build a mine-subsidence observation station. *J China Univ Mine Technol* 18(3):377–380
- Mao A, Harrison CGA, Dixon TH (1999) Noise in GPS coordinate time series. *J Geophys Res: Solid Earth* 104(B2):2797–2816
- Márquez-Azúa B, DeMets C (2003) Crustal velocity field of Mexico from continuous GPS measurements, 1993 to June 2001: implications for the neotectonics of Mexico. *J Geophys Res: Solid Earth* 108(B9):2450. doi:[10.1029/2002JB002241](https://doi.org/10.1029/2002JB002241)
- Martens HR, Simons M, Owen S, Rivera L (2016) Observations of ocean tidal load response in South America from subdaily GPS positions. *Geophys J Int* 205:1637–1664
- Nikolaidis RM (2002) Observation of geodetic and seismic deformation with the global positioning system. Ph.D Thesis, University of California, San Diego
- Nikolaidis RM, Bock Y, de Jonge PJ, Shearer P, Agnew DC, Domselaar M (2001) Seismic wave observations with the global positioning system. *J Geophys Res* 106(B10):21897–21916
- Ostini L (2012) Analysis and quality assessment of GNSS-derived parameter time series. PhD thesis, Astronomisches Institut der Universität Bern. http://www.bernese.unibe.ch/publist/2012/phd/diss_lo_4print.pdf
- Peltier WR, Argus DF, Drummond R (2015) Space geodesy constrains ice age terminal deglaciation: The global ICE-6G_C (VM5a) model. *J Geophys Res: Solid Earth* 119. doi:[10.1002/2014JB011176](https://doi.org/10.1002/2014JB011176)
- Petit G, Luzum B (2010) IERS conventions. IERS technical note 36, Verlag des Bundesamts für Kartographie und Geodäsie, 2010, ISBN 3-89888-989-6. Frankfurt am Main, Germany
- Rajner M, Liwosz T (2012) Studies of crustal deformation due to hydrological loading on GPS height estimates. *Geod Cartog* 60(2):135–144
- Ray J (2009) Preparations for the 2nd IGS reprocessing campaign. American Geophysical Union Fall Meeting. San Francisco, California, USA. December 14–19. Poster G11B-0642
- Ray J, Altamimi Z, Collilieux X, van Dam T (2008) Anomalous harmonics in the spectra of GPS position estimates. *GPS Solut* 12(1):55–64
- Ray J, Griffiths J, Collilieux X, Reibschung P (2013) Subseasonal GNSS positioning errors. *J Geophys Res Lett* 40(22):5854–5860
- Razeghi SM, Amiri-Simkooei AR, Sharifi MA (2015) Coloured noise effects on deformation parameters of permanent GPS networks. *Geophys J Int* 204(3):1843–1857
- Reibschung P, Altamimi Z, Ray J, Garayt B (2016) The IGS contribution to ITRF2014. *J Geod* 90(7):611–630
- Rodriguez-Solano CJ, Hugentobler U, Steigenberger P (2011) Impact of albedo radiation on GPS satellites In: *Geodesy for Planet Earth*. Proceeding of the 2009 IAG Symposium, Buenos Aires, Argentina, 31 August– 4 September, pp 113–119
- Rodriguez-Solano CJ, Hugentobler U, Steigenberger P, Lutz S (2012) Impact of Earth radiation pressure on GPS position estimates. *J Geod* 86(5):309–317
- Rodriguez-Solano CJ, Hugentobler U, Steigenberger P, Bloßfeld M, Fritsche M (2014) Reducing the draconitic errors in GNSS geodetic products. *J Geod* 88(6):559–574
- Santamaría-Gómez A, Bouin MN, Collilieux X, Wöppelmann G (2011) Correlated errors in GPS position time series: Implications for velocity estimates. *J Geophys Res: Solid Earth* 116(B01405). doi:[10.1029/2010JB007701](https://doi.org/10.1029/2010JB007701)
- Segall P, Davis JL (1997) GPS applications for geodynamics and earthquake studies. *Annu Rev Earth Planet Sci* 25(1):301–336
- Selle C, Desai S, Garcia Fernandez M, Sibos A (2014) Spectral analysis of GPS-based station positioning time series from PPP solutions. In: 2014 IGS workshop. Pasadena, CA, USA
- Serpelloni E, Anzidei M, Baldi P, Casula G, Galvani A (2005) Crustal velocity and strain-rate fields in Italy and surrounding regions: new results from the analysis of permanent and non-permanent GPS networks. *Geophys J Int* 161(3):861–880
- Steigenberger P, Boehm J, Tesmer V (2009) Comparison of GMF/GPT with VMF1/ECMWF and implications for atmospheric loading. *J Geod* 83(10):943–951
- Thatcher W (2003) GPS constraints on the kinematics of continental deformation. *Int Geol Rev* 45(3):191–212
- Teferle FN, Bingley RM, Dodson AH, Penna NT, Baker TF (2002) Using GPS to separate crustal movements and sea level changes at tide gauges in the UK. In: Drewes H et al (eds) *Vertical reference systems*. Springer-Verlag, Berlin, pp 264–269
- Teferle FN, Bingley RM, Williams SDP, Baker TF, Dodson AH (2006) Using continuous GPS and absolute gravity to separate vertical land movements and changes in sea level at tide gauges in the UK. *Philos Trans Royal Soc A* 364:917–930
- Teunissen PJG (1988) Towards a least-squares framework for adjusting and testing of both functional and stochastic models. In: *Internal research memo, Geodetic Computing Centre, Delft*. A reprint of original 1988 report is also available in 2004, *Series on mathematical Geodesy and Positioning*, No. 26 <http://saegnss1.curtin.edu.au/Publications/2004/Teunissen2004Towards.pdf>
- Teunissen PJG, Amiri-Simkooei AR (2008) Least-squares variance component estimation. *J Geod* 82(2):65–82
- Tregoning P, Watson C (2009) Atmospheric effects and spurious signals in GPS analyses. *J Geophys Res: Solid Earth* 114:B09403. doi:[10.1029/B006344](https://doi.org/10.1029/B006344)
- Urschl C, Beutler G, Gurtner W, Hugentobler U, Schaer S (2007) Contribution of SLR tracking data to GNSS orbit determination. *J Adv Space Res* 39(10):1515–1523
- van Dam T, Wahr J, Milly PCD, Shmakin AB, Blewitt G, Lavalée D, Larson KM (2001) Crustal displacements due to continental water loading. *Geophys Res Lett* 28(4):651–654
- Wahr JM (1985) Deformation induced by polar motion. *J Geophys Res* 90(B11):9363–9368. doi:[10.1029/JB090iB11p09363](https://doi.org/10.1029/JB090iB11p09363)
- Wdowinski S, Bock Y, Zhang J, Fang P, Genrich J (1997) Southern California permanent GPS geodetic array: spatial filtering of daily positions for estimating coseismic and postseismic displacements induced by the 1992 landers earthquake. *J Geophys Res: Solid Earth* 102(B8):18057–18070
- Williams SDP (2003a) The effect of colored noise on the uncertainties of rates estimated from geodetic time series. *J Geod* 76(9–10):483–494
- Williams SDP (2003b) Offsets in global positioning system time series. *J Geophys Res: Solid Earth* 108(B6):2310. doi:[10.1029/2002JB002156](https://doi.org/10.1029/2002JB002156)
- Williams SDP, Bock Y, Fang P, Jamason P, Nikolaidis RM, Prawirodirdjo L, Miller M, Johnson DJ (2004) Error analysis of continuous GPS position time series. *J Geophys Res: Solid Earth* 109:B03412. doi:[10.1029/2003JB002741](https://doi.org/10.1029/2003JB002741)
- Wöppelmann G, Martin Miguez B, Bouin MN, Altamimi Z (2007) Geocentric sea-level trend estimates from GPS analyses at relevant tide gauges world-wide. *Glob Planet Change* 57(3):396–406
- Zhang J, Bock Y, Johnson H, Fang P, Williams S, Genrich J, Wdowinski S, Behr J (1997) Southern California permanent GPS geodetic array: error analysis of daily position estimates and site velocities. *J Geophys Res: Solid Earth* 102(B8):18035–18055
- Zumberge JF, Heflin MB, Jefferson DC, Watkins MM, Webb FH (1997) Precise point positioning for the efficient and robust analysis of GPS data from large networks. *J Geophys Res: Solid Earth* 102(B3):5005–5017