Atmos. Chem. Phys., 16, 7969–7979, 2016 www.atmos-chem-phys.net/16/7969/2016/ doi:10.5194/acp-16-7969-2016 © Author(s) 2016. CC Attribution 3.0 License.





# Observations and implications of liquid-liquid phase separation at high relative humidities in secondary organic material produced by $\alpha$ -pinene ozonolysis without inorganic salts

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Received: 18 October 2015 - Published in Atmos. Chem. Phys. Discuss.: 26 November 2015

Revised: 14 April 2016 - Accepted: 6 May 2016 - Published: 1 July 2016

**Abstract.** Particles consisting of secondary organic material (SOM) are abundant in the atmosphere. To predict the role of these particles in climate, visibility and atmospheric chemistry, information on particle phase state (i.e., single liquid, two liquids and solid) is needed. This paper focuses on the phase state of SOM particles free of inorganic salts produced by the ozonolysis of  $\alpha$ -pinene. Phase transitions were investigated in the laboratory using optical microscopy and theoretically using a thermodynamic model at 290 K and for relative humidities ranging from < 0.5 to 100 %. In the laboratory studies, a single phase was observed from 0 to 95 % relative humidity (RH) while two liquid phases were observed above 95 % RH. For increasing RH, the mechanism of liquid-liquid phase separation (LLPS) was spinodal decomposition. The RH range over which two liquid phases were observed did not depend on the direction of RH change. In the modeling studies, the SOM took up very little water and was a single organic-rich phase at low RH values. At high RH, the SOM underwent LLPS to form an organic-rich phase and a water-rich phase, consistent with the laboratory studies. The presence of LLPS at high RH values can have consequences for the cloud condensation nuclei (CCN) activity of SOM particles. In the simulated Köhler curves for SOM particles, two local maxima were observed. Depending on the composition of the SOM, the first or second maximum can determine the critical supersaturation for activation. Recently researchers have observed inconsistencies between measured CCN properties of SOM particles and hygroscopic growth measured below water saturation (i.e., hygroscopic parameters measured below water saturation were inconsistent with hygroscopic parameters measured above water saturation). The work presented here illustrates that such inconsistencies are expected for systems with LLPS when the water uptake at subsaturated conditions represents the hygroscopicity of an organic-rich phase while the barrier for CCN activation can be determined by the second maximum in the Köhler curve when the particles are water rich.

## 1 Introduction

Particles consisting of secondary organic material (SOM) can account for 20–80% of the total submicron organic mass concentrations in the atmosphere (Zhang et al., 2007;

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Jimenez et al., 2009). SOM in the particle phase consists of the low volatility fraction of the oxidized products of biogenic or anthropogenic volatile organic compounds (Hallquist et al., 2009). To predict the role of SOM particles for climate, visibility and atmospheric chemistry, information on the phase state within individual SOM particles (e.g., one liquid, two liquids and one solid) is needed. Particle phase state influences the properties of particles such as cloud condensation nuclei (CCN) properties, optical properties and interactions with reactive and non-reactive gas-phase species (Martin et al., 2000; Raymond and Pandis, 2002; Bilde and Svenningsson, 2004; Zuend et al., 2010; Kuwata and Martin, 2012).

A possible phase transition of SOM particles during relative humidity (RH) cycling is liquid–liquid phase separation (LLPS) (Pankow et al., 2003; Petters et al., 2006). LLPS has been observed in the laboratory when SOM produced by  $\alpha$ -pinene ozonolysis was combined with ammonium sulfate and for other organic systems when mixed with inorganic salts when the average organic oxygen-to-carbon elemental ratios (O:C) were less than approximately 0.8 (Krieger et al., 2012; You et al., 2014). The presence of the ammonium sulfate causes salting-out of the organic material and the formation of two liquid phases. However, we are not aware of previous laboratory studies focusing on LLPS in SOM in the absence of inorganic salts.

This paper focuses on phase transitions of SOM produced by  $\alpha$ -pinene ozonolysis free of inorganic salts.  $\alpha$ -pinene was chosen for the precursor gas for SOM because it is an important contributor to organic particle mass in the atmosphere, especially in regions such as boreal forests (Cavalli et al., 2006). Phase transitions were investigated both in the laboratory and with a thermodynamic model over the range of <0.5 % to 100 % RH.

## 2 Methods

#### 2.1 Laboratory studies

# 2.1.1 Production and collection of secondary organic material

Particles of secondary organic material were produced by  $\alpha$ -pinene ozonolysis in a flow tube reactor by the methods described in Shrestha et al. (2013). To remove excess reactants, the aerosol in the outflow from flow tube reactor continued through a diffusion dryer charged with ozone destruction catalyst (Ozone Solutions, model ODS-2) and a carbon filter denuder (Sunset Laboratory). Particle mass concentrations in the flow tube reactor ranged from 75 to 11 000  $\mu$ g m<sup>-3</sup> (Table 1)

At the outlet of the flow tube reactor, particles were collected using one of two different methods. In the first method, after charging in a bipolar charger (TSI, model 3077), a

portion of the flow (1.5 slpm) was sampled into a Nanometer Aerosol Sampler (TSI, model 3089). The particles were collected by electrostatic precipitation (-10 kV sampler potential) onto a siliconized glass slide (Hampton Research, Canada). This method of collection resulted in submicron particles distributed evenly over the glass slide (Liu et al., 2013). In the second method, a portion (1.5 slpm) of the aerosol flow exiting the flow reactor was sampled into a single stage impactor (Prenni et al., 2009; Pöschl et al., 2010). The collection substrate was a glass slide coated with trichloro(1H,1H,2H,2H-perfluorooctyl)silane (Sigma-Aldrich, 97 % purity). The coating procedure, which was described in Knopf (2003), produced a hydrophobic surface. The size of the particles after coagulation on the glass slides ranged from 10 to 80 µm in diameter. Table 1 lists samples collected by method 1 or method 2.

For the optical microscope experiments (see Sect. 2.1.2), supermicron particles are needed, and in the case of method 1 the collected submicron particles were exposed to water supersaturation (SS) conditions to grow and coagulate the particles (Song et al., 2015). Specifically, the slides containing the submicron particles were mounted to a temperature and RH-controlled flow cell, which was coupled to a reflectance microscope, as described previously (Koop et al., 2000; Parson et al., 2004; Pant et al., 2006). The RH in the flow cell was initially set to > 100 % by decreasing the cell temperature to below the dew point temperature. At the initial RH (> 100 %) water condensed on the slides forming large (150– 300 µm) droplets by growth and coagulation. The RH was then ramped back to  $\sim 98\%$  by warming the cell back to room temperature, resulting in water evaporating from the droplets. This process of coagulation followed by evaporation resulted in SOM particles with lateral dimensions of 5-30 µm.

This hygroscopic cycling of method 1 did introduce the possibility for aqueous phase reactions to occur (e.g., simulating cloud water reactions) that might not be present under subsaturated conditions (e.g., aerosol water only). Furthermore, during the RH cycle some of the secondary organic material may have evaporated from the particles. However, the similarities in the results for two collection methods (see Table 1) suggest that neither of these possible processes, if present, changed the chemical composition enough to influence LLPS. Also, prior to collection with both methods, excess gas-phase components were removed with a carbon filter. During this process, some of the more volatile material in the SOM may have evaporated. If some evaporation of higher volatility species occurred, the SOM would likely be more similar to the chemical composition of SOM particles in the atmosphere, which are formed at lower particle mass concentrations compared to particles in the current laboratory experiments.

**Table 1.** Summary of experimental conditions for the production and collection of  $\alpha$ -pinene-derived SOM. SOM samples 2, 3, 4, 5, 7 and 8 were collected on hydrophobic substrates using a single stage impactor. SOM samples 1, 6 and 9 were collected on hydrophobic substrates using an electrostatic precipitator. The separation relative humidity (SRH) from one to two phases is listed for each SOM. The standard deviation (stdev) is derived from several cycles of RH for different deposited particles. In cases for which the SRH was only determined for one humidity cycle (SOM samples 3–5), the error represents the maximum error reported for the other SOM samples. SOM particle mass concentration refers to the concentration of organic particles suspended in the gas phase at the time of SOM production.

SOM sample	α-pinene (ppm)	O <sub>3</sub> (ppm)	SOM particle mass $(\mu g m^{-3})$ concentration	Collection time (min)	SRH (%) $\pm$ SD with decreasing RH	SRH (%) $\pm$ SD with increasing RH	Collection method
1	0.20	16	75	3120	$96.2 \pm 0.41$	$96.4 \pm 0.03$	1
2	0.35	10	85	5580	$95.8 \pm 0.18$	$95.9 \pm 0.04$	2
3	0.35	10	95	5733	$95.1 \pm 0.41$	$95.2 \pm 0.41$	2
4	0.35	10	110	2160	$94.7 \pm 0.13$	$95.0 \pm 0.41$	2
5	0.80	10	320	1590	$95.2 \pm 0.41$	$96.3 \pm 0.41$	2
6	1.00	20	1500	1440	$96.2 \pm 0.39$	$96.1 \pm 0.08$	1
7	5.00	10	2900	1520	$97.3 \pm 0.08$	$96.7 \pm 0.39$	2
8	5.00	10	2900	1472	$95.8 \pm 1.05$	$96.5 \pm 0.21$	2
9	5.00	20	11 000	330	$96.5 \pm 0.23$	$96.5 \pm 0.28$	1

# 2.1.2 Method of determining SOM phase(s)

Hydrophobic substrates containing the supermicron particles were located within a flow cell with temperature and RH control and coupled to a reflectance microscope (Zeiss, AxioTech, 50× objective) for observation (Koop et al., 2000; Parson et al., 2004; Pant et al., 2006). During experiments, the RH was changed by adjusting the moisture content of the gas flow. The RH was measured with a chilled-mirror hygrometer (General Eastern, model 1311DR), which was calibrated using the deliquescence RH of pure ammonium sulfate particles. During typical experiments, the RH was first set to  $\sim 100\,\%$ , and then the RH was ramped down at a rate of 0.1–0.3 % RH min<sup>-1</sup> and images were collected every 5–20 s. After the RH reached < 0.5 % RH, it was ramped up again at the same rate to  $\sim 100 \,\%$ . During the experiments, temperature was constant at  $290 \pm 1 \,\mathrm{K}$ . From images recorded while changing the RH, the number of phases present in the SOM was determined.

During the experiments used to determine SOM phase state the concentration of organic vapors in the flow cell was not controlled. Hence, some of the more volatile material in the SOM may have evaporated during these experiments. However, no visible change in the particle volume occurred during these experiments, suggesting evaporative loss was minimal. The SOM particle mass concentrations used when generating the SOM were similar to those used in Grayson et al. (2016), and the sample preparation methods were identical to those used in Grayson et al. (2016), who showed no visible volume change of the droplets over time periods of greater than 44 h. It is possible that condensed-phase reactions may have occurred that lowered the vapor pressure of the SOM.

## 2.2 Thermodynamic modeling studies

Liquid–liquid equilibria and water uptake were calculated with the methods developed by Zuend et al. (2008, 2010, 2011) and Zuend and Seinfeld (2012, 2013). To calculate activity coefficients of the organic species as a function of the solution composition, the thermodynamic group-contribution model AIOMFAC (Aerosol Inorganic–Organic Mixtures Functional groups Activity Coefficients) developed by Zuend et al. (2008, 2010, 2011) was utilized. To determine whether two liquid phases or a single liquid phase was the thermodynamic stable state, the Gibbs free energies of a two-liquid phase state and a one-liquid phase state were calculated (Zuend et al., 2010). If the two-liquid phase state had a lower Gibbs free energy compared to the one liquid phase state, then LLPS was predicted.

To represent SOM from the ozonolysis of  $\alpha$ -pinene, the oxidation products listed in Table 2 were used. These oxidation products were based on the calculations performed by Zuend and Seinfeld (2012), who used the Master Chemical Mechanism (Jenkin et al., 1997; Saunders et al., 2003) in combination with EVAPORATION (Compernolle et al., 2011) to establish a representative condensed-phase composition of oxidation products from the ozonolysis of  $\alpha$ -pinene. Three different mixtures of the oxidation products were used in the current study (see Tables 2 and 3). The mixtures SOMhigh and SOM-low are based on calculations by Zuend and Seinfeld (2012) carried out at 60 % RH yielding particle mass concentrations of 21.86 and  $0.81 \, \mu g \, m^{-3}$ , respectively (see Fig. 4 from Zuend and Seinfeld, 2012). The SOM-ox mixture used the same oxidation products as SOM-high and SOMlow mixtures, but the share of the more oxidized products was increased. Water uptake and CCN activation for these mixtures were simulated assuming that all oxidation products remained in the condensed phase without further gas-to-

**Table 2.** Molecular weights  $(M_w)$ , O:C elemental ratios and mole fractions of the  $\alpha$ -pinene ozonolysis products from Zuend and Seinfeld (2012) used in the thermodynamic modeling study. Three different scenarios were investigated: high SOM concentrations (SOM-high), low SOM concentration (SOM-low) and with higher shares of the more oxidized products (SOM-ox).

Name	$M_{\rm W}$ (g mol <sup>-1</sup> )	O:C	Mole fraction in mixture		
			SOM-high	SOM-low	SOM-ox
C107OOH	200.231	0.4	0.039	0.013	0.009
Pinonic acid	184.232	0.3	0.016	0.000	0.000
C97OOH	188.221	0.444	0.310	0.042	0.030
C108OOH	216.231	0.5	0.050	0.012	0.009
ALDOL_dimer	368.421	0.368	0.029	0.079	0.056
Pinic acid	186.205	0.444	0.167	0.156	0.110
C921OOH	204.220	0.556	0.138	0.271	0.192
C109OOH	200.231	0.4	0.005	0.000	0.000
C812OOH	190.194	0.625	0.128	0.277	0.245
ESTER_dimer	368.421	0.368	0.005	0.021	0.015
C811OH	158.094	0.375	0.012	0.000	0.000
Hopinonic acid	200.232	0.4	0.058	0.026	0.019
С813ООН	206.193	0.75	0.042	0.102	0.316

particle partitioning to isolate the effect of LLPS. The average O:C ratios used in the thermodynamic modeling studies are similar to those measured in environmental chambers (e.g., see Chhabra et al. (2011) and references therein).

The oxidation products and mole fractions used in the thermodynamic modeling studies were used to (1) improve our understanding of the phase state of multicomponent organic mixtures such as those generated during SOM formation from  $\alpha$ -pinene ozonolysis and (2) to explore the possible implications of liquid–liquid phase separation in multicomponent organic mixtures such as SOM. However, the oxidation products and their mole fractions were not intended to reproduce the laboratory conditions used here or atmospheric SOM.

In addition to detecting the presence of LLPS, the thermodynamic model was used to predict the hygroscopic growth factor (HGF), CCN activation and the hygroscopicity parameter ( $\kappa$ ), from calculations of hygroscopic growth ( $\kappa_{HGF}$ ) and calculations of CCN activation ( $\kappa_{CCN}$ ) (see Table 3). The hygroscopic growth factor was calculated with the following Eq. (1):

$$HGF(RH) = \frac{D(RH)}{D_0},$$
(1)

where D(RH) and  $D_0$  represent the wet and the dry diameters of the particles, respectively. The dry diameter  $D_0$  was calculated at 0% RH. The following equation was used to calculate the  $\kappa_{HGF}$  (Petters and Kreidenweis, 2007; Pajunoja et al., 2015):

$$\kappa_{\rm HGF} = 1 - {\rm HGF}^3 + \frac{{\rm HGF}^3 - 1}{\frac{{\rm RH}}{100\,\text{Ge}}} e^{\left(\frac{4\sigma\,M_{\rm W}}{RT\,\rho_{\rm W}\,D_0\,{\rm HGF}}\right)}, \tag{2}$$

where  $\sigma$  is surface tension at the particle–air interface,  $M_w$  is the molecular weight of water (18 g mol<sup>-1</sup>), R is the univer-

sal gas constant, T is temperature (298 K) and  $\rho_{\rm w}$  is the density of water (1 g cm<sup>-3</sup>). The following equation was used to calculate  $\kappa_{\rm CCN}$  (Petters and Kreidenweis, 2007; Pajunoja et al., 2015):

$$\kappa_{\rm CCN} = \frac{4A_{\rm Kelvin}^3}{27d_{\rm c}^3 \ln^2 S_{\rm c}},\tag{3}$$

where

$$A_{\text{Kelvin}} = \frac{4\sigma M_{\text{W}}}{RT\rho_{\text{W}}},\tag{4}$$

and  $d_c$  and  $S_c$  are the critical diameter and saturation ratio, respectively.

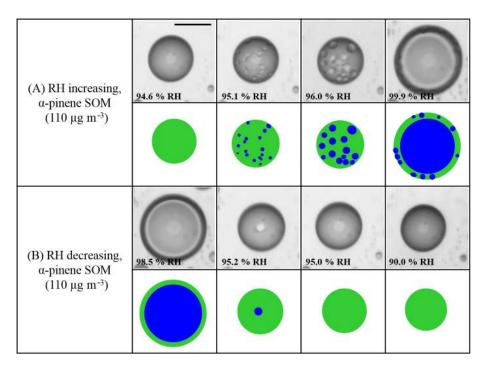
#### 3 Results and discussion

# 3.1 Observations of LLPS in α-pinene-derived SOM particles: laboratory studies

As the RH was scanned from high values ( $\sim 100\,\%$ ) to low values ( $\leq 0.5\,\%$ ) and, in reverse, LLPS in the SOM was clearly visible for RH >95 %. Example images are shown in Fig. 1. As the RH was increased from <0.5 % to  $\sim 95\,\%$ , no change in the image was observed. The particles appear as a single phase. The light-colored circle in the center of the particle in panel a is an optical effect of light scattering from a hemispherical uniform particle (Bertram et al., 2011). Above  $\sim 95\,\%$  RH, the particle underwent spinodal decomposition, resulting in two phases (see Fig. 1 and movie S1 in the Supplement). Spinodal decomposition, a phenomenon by which the phase transition occurs with essentially no free energy barrier to nucleation of a second phase, is evident from the formation of the many small inclusions of the second

Table 3. Calculated properties of the mixtures SOM-high, SOM-low and SOM-ox: average O:C elemental ratio, average molecular weight, simulated mass yields at 60% RH reported in Zuend and Seinfeld (2012), range of LLPS for a 20 μm particle in diameter, range of LLPS for a 100 nm particle in diameter, critical supersaturation SSc for a 100 nm particle with a surface tension of 72 mN m<sup>-1</sup>, critical supersaturation SSc for a 100 nm particle with a surface tension of 40 mN m<sup>-1</sup>,  $\kappa_{HGF}$  from the hygroscopic growth curve at 90 % RH for a 100 nm diameter particle and surface tension of 72 mN m<sup>-1</sup>,  $\kappa_{HGF}$  from the hygroscopic growth curve at 90 % RH for a 100 nm diameter particle and surface tension of 40 mN m<sup>-1</sup>,  $\kappa_{CCN}$  from SSc of the Köhler curve for a 100 nm particle and surface tension of 72 mN m<sup>-1</sup>,  $\kappa_{CCN}$  from SSc of the Köhler curve for a 100 nm particle and surface tension of 40 mN m<sup>-1</sup>.

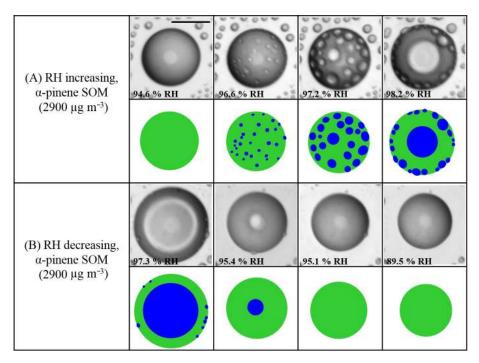
	SOM-high	SOM-low	SOM-ox
Av. O:C	0.472	0.513	0.582
Av. $M_{\rm W}$ (g mol <sup>-1</sup> )	199.5	213.5	210.6
PM mass conc. ( $\mu g  m^{-3}$ )	21.86	0.81	_
LLPS range (% RH), 20 µm	99.31-99.88	98.91–99.94	98.71-99.92
LLPS range (% RH), 100 nm	> 100 %	> 100 %	> 100 %
$SSc (\%), 72 \mathrm{mN} \mathrm{m}^{-1}$	1.206	0.668	0.432
$SSc (\%), 40 \mathrm{mN  m^{-1}}$	0.335	0.177	0.172
$\kappa_{\rm HGF}$ at 90 % RH, 72 mN m <sup>-1</sup>	0.0228	0.0278	0.0319
$\kappa_{\rm HGF}$ at 90 % RH, 40 mN m <sup>-1</sup>	0.0228	0.0278	0.0319
$\kappa_{\rm CCN}$ , 72 mN m <sup>-1</sup>	0.0093	0.0318	0.0758
$\kappa_{\rm CCN}$ , 40 mN m <sup>-1</sup>	0.0198	0.0750	0.0793



**Figure 1.** Effect of RH cycles on α-pinene-derived SOM for SOM produced at  $110 \,\mu g \, m^{-3}$ . Illustrations of the images are shown for clarity. Green: SOM-rich phase. Blue: water-rich phase. Size bar is  $20 \,\mu m$ .

phase throughout the particle (Ciobanu et al., 2009; Song et al., 2012). After phase separation by spinodal decomposition, the inclusions containing the second phase increased in size and coagulated into larger inclusions and eventually formed the inner phase of the particle as the RH was increased above  $\sim 95\,\%$  RH (see movie S1 in the Supplement). After phase separation, two liquid phases persisted until  $\sim 100\,\%$  RH.

Panel b of Fig. 1 and movie S2 in the Supplement show the same particle as panel a, but for experiments using decreasing RH starting from close to 100% RH. At  $\sim 98\%$  RH, the particle contained two liquid phases. As the RH decreased from  $\sim 98$  to  $\sim 95\%$ , the thickness of the SOM-rich phase increased, while the amount of the water-rich phase decreased. Below  $\sim 95\%$  RH, the two phases merged into a single phase. As the RH was decreased further to < 0.5% RH,



**Figure 2.** Effect of RH cycles on  $\alpha$ -pinene-derived SOM for SOM produced at 2900  $\mu$ g m<sup>-3</sup>. Illustrations of the images are shown for clarity. Green: SOM-rich phase. Blue: water-rich phase. Size bar is 20  $\mu$ m.

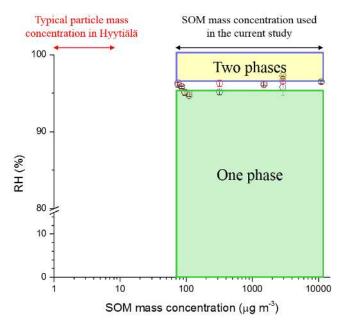
no abrupt change was observed, indicating the absence of any further phase transitions. Figure 2 and movies S3–S4 in the Supplement show similar pictures and movies as Fig. 1 and movies S1–S2 in the Supplement, except in this case the SOM was generated using a higher particle mass concentration in the flow tube reactor.

Figures 1–2 and movies S1–S4 in the Supplement show that there are differences in the process of LLPS and the resulting morphology depending on the direction of the RH change. For increasing RH, spinodal decomposition was identified as the mechanism of phase separation. For decreasing RH, disappearance of phase separation occurred by merging of the two phases.

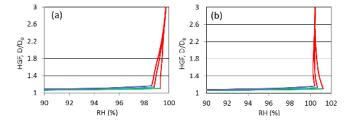
Experiments were also carried out to determine whether the lowest RH at which two phases existed depended on the direction of RH change. Values for the lowest RH at which two phases were observed when increasing and decreasing RH are listed in Table 1 and shown in Fig. 3 (black circles correspond to increasing RH and red circles correspond to decreasing RH). Table 1 and Fig. 3 illustrate that the lowest RH at which two phases were observed did not depend significantly on the direction of RH change. Figure 3 and Table 1 also show that within uncertainties of the measurements, there is no effect of the SOM particle mass concentrations in the flow tube reactor on the lowest RH at which two liquid phases were observed for the range of 75 to  $11\,000\,\mu\mathrm{g}\,\mathrm{m}^{-3}$ . Also included in Fig. 3 are typical SOM particle mass concentrations measured over a boreal forest (Raatikainen et al., 2010), where  $\alpha$ -pinene is an important

contributor to SOM (Cavalli et al., 2006). Since the SOM particle mass concentrations used in our experiments when generating the SOM were higher than typically observed in the atmosphere, additional studies are needed to confirm LLPS with SOM produced using atmospherically relevant particle mass concentrations.

The behavior observed here for SOM is consistent with bulk thermodynamics. Consider, for example, a mixture of a relatively hydrophobic organic with a less hydrophobic organic, such as a mixture containing equal mole ratios of heptanol and propanol (Stoicescu et al., 2011). Under dry conditions this mixture exists as a single phase. As water is added to the system, the mixture exists as a single (organic-rich) phase until the water content is approximately 0.3 mole fraction. At this point, the mixture separates into an organic-rich phase and a water-rich phase. As water is further added to the system, the two phases coexist until a large amount of water has been added, at which point all the organic material dissolves into the water-rich phase. The formation of two phases is due to the non-ideality of the mixture; i.e., if the mixture was ideal, LLPS would not be observed. Examples of other organic mixtures that exhibit this type of behavior include mixtures of hexanol and acetic acid (Senol et al., 2004) and mixtures of octanol and acetone (Tiryaki et al., 1994). For a long list of organic mixtures that undergo liquid–liquid phase separation when mixed with water, see Table 1 in Ganbavale et al. (2015).



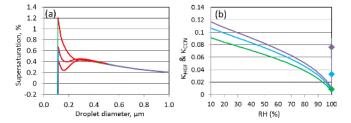
**Figure 3.** Relative humidity (RH) at which phase transition between one phase and two liquid phases were observed for  $\alpha$ -pinenederived SOM as a function of the mass concentration of SOM produced. Red circles: onset of phase separation upon moistening. Black circles: merging of the two liquid phases upon drying. The y error bars represent the standard deviation in RH determination at the phase transition. Green shaded region: one phase prevalent in  $\alpha$ -pinene-derived SOM. Yellow shaded region: two phases present. Also shown is the mass concentration observed over a representative boreal forest in Hyytiälä (Raatikainen et al., 2010).



**Figure 4.** Simulated hygroscopic growth factors HGF =  $D/D_0$  for SOM-high (green, O:C=0.472), SOM-low (blue, O:C=0.513) and SOM-ox (purple, O:C=0.582). The red segments on the lines indicate the presence of LLPS. Panel (a) corresponds to a dry diameter of  $20\,\mu\text{m}$ , which is similar in size to the particles used in the optical microscope experiments, and a surface tension of water. Panel (b) corresponds to a dry diameter of  $100\,\text{nm}$  and a surface tension of water.

# 3.2 Observations of LLPS in α-pinene-derived SOM particles: thermodynamic modeling studies

Shown in panel a of Fig. 4 are the simulated hygroscopic growth factors for the three different SOM mixtures (SOMhigh, SOM-low, SOM-ox) with a dry diameter of  $20 \,\mu m$ , which is similar in size to the particles used in the optical



**Figure 5.** Assuming the surface tension of water, Köhler curves (panel **a**) and hygroscopicity parameter  $\kappa$  (panel **b**) for a particle with a dry diameter of 100 nm for SOM-high (green, O: C = 0.472), SOM-low (blue, O: C = 0.513) and SOM-ox (purple, O: C = 0.582). The red segments on the lines in panel (**a**) indicate the presence of LLPS. In panel (**b**),  $\kappa_{\text{HGF}}$  is given as solid line as a function of RH and  $\kappa_{\text{CCN}}$  as diamond at RH = 100 %.

microscope experiments, and assuming a surface tension  $(\sigma)$ of water. At RH values < 98 % the SOM took up little water. However, when the multicomponent systems consisting of organic substances with different hydrophilicities (i.e., different O: C elemental ratios) were exposed to RH values > 98 % RH, LLPS into an organic-rich phase and a water-rich phase was observed. At the RH of LLPS, the particles took up a significant amount of water, leading to an almost vertical increase in the hygroscopic growth curve as shown in panel a of Fig. 4. For the SOM-high mixture LLPS occurred from 99.3 to 99.88 % RH as indicated by the red segment on the green line. When the share of the more hydrophilic substances is increased, as is the case for the low SOM loading (SOM-low) with a particle mass concentration of  $0.81 \,\mu\mathrm{g}\,\mathrm{m}^{-3}$ , the onset of LLPS shifted to lower RH and the RH range of LLPS was increased. In the laboratory experiments, LLPS was observed starting at 95 % RH. This lower onset may be due to more highly oxidized products produced in the laboratory compared with the oxidation products used in the thermodynamic calculations. When the range of O: C elemental ratios of the individual products used in the thermodynamic calculations is narrower than what is present in the SOM generated in the experiment, the calculated width of LLPS in terms of RH would be narrower than the measured one.

Shown in panel b of Fig. 4 are the simulated hygroscopic growth factors of a 100 nm dry particle for the three different SOM mixtures (SOM-high, SOM-low, SOM-ox), again assuming a surface tension of water. This figure illustrates that LLPS can shift to RH > 100 % in small particles due to the Kelvin effect. In 100 nm particles, the SOM took up little water at RH < 100 %, and LLPS is predicted above 100 % RH.

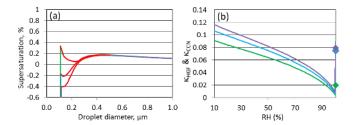
## 4 Implications

## 4.1 Cloud condensation nuclei properties

The presence of a miscibility gap at RH > 95 % has consequences for the CCN activity of particles as suggested previously (Petters et al., 2006). Shown in panel a of Fig. 5 are simulated Köhler curves for SOM particles with dry diameters of 100 nm and using the surface tension of water. The Köhler curves show a sharp increase in the equilibrium water vapor SS above the particles as the size of the particles increases from 100 to roughly 110 nm due to the Kelvin effect when they are still in their organic-rich phase (i.e., low water content state). As the particle size increases from 110 to 200 nm there is a steep decrease in SS as the particles switch from the organic-rich phase to two phases by taking up water from the gas phase. This gives rise to the first maximum in the Köhler curve, which occurs at a wet particle diameter of  $D_{\rm p} \approx 110$  nm for the SOM-high mixture. The second maximum of SS at a wet diameter of  $D_p \approx 300 \,\mathrm{nm}$ is the regular maximum of the Köhler curve, which occurs when the droplet is dilute and close to solution ideality. When the particle is composed of higher shares of the more hydrophilic substances, the first maximum decreases in height while the second maximum remains constant (see panel a of Fig. 5). For SOM-high (O:C=0.472) and SOMlow (O:C=0.513), the first maximum in the Köhler curve determines the critical SS to overcome the activation barrier. In SOM-ox (O: C = 0.582) the second maximum is higher than the first one and relevant for CCN activation. The height of the second maximum in the Köhler curve is sensitive to the molecular weight of the organic substances making up the particle (e.g., Wex et al., 2007, 2008).

Shown in panel a of Fig. 6 are simulated Köhler curves for SOM particles with dry diameters of 100 nm and using the surface tension of 40 mN m<sup>-1</sup>, which is consistent with the surface tension of aqueous mixtures of pinonic acid, pinic acid and pinonaldehyde (Tuckermann and Cammenga, 2004; Hartz et al., 2006). Panel a of Fig. 6 illustrates that a lower surface tension has a large effect on the first maximum in the Köhler curve and also lowers the barrier of the second maximum. During the activation process, the surface tension is expected to increase as the phase state changes from organic rich to water rich, but this process is not modeled here. Additional studies are needed to fully understand the effect of varying surface tension on the resulting Köhler curves (Ruehl et al., 2016).

The non-ideality of SOM also has consequences for the applicability of the single parameter  $\kappa$  representation of Köhler theory (Petters and Kreidenweis, 2007). If SOM forms an ideal mixture with water then  $\kappa_{HGF}$  is approximately constant over the whole RH range and  $\kappa_{CCN}$  corresponds well with  $\kappa_{HGF}$  for an organic particle (Petters and Kreidenweis, 2007). However, panel b of Figs. 5 and 6 show that  $\kappa_{HGF}$  for the mixtures SOM-high, SOM-low and SOM-ox with a



**Figure 6.** Assuming a surface tension of  $40 \,\mathrm{mN \,m^{-1}}$ , Köhler curves (panel **a**) and hygroscopicity parameter  $\kappa$  (panel **b**) for a particle with a dry diameter of  $100 \,\mathrm{nm}$  for SOM-high (green, O:C=0.472), SOM-low (blue, O:C=0.513) and SOM-ox (purple, O:C=0.582). The red segments on the lines in panel (**a**) indicate the presence of LLPS. In panel (**b**),  $\kappa_{\mathrm{HGF}}$  is given as solid line as a function of RH and  $\kappa_{\mathrm{CCN}}$  as diamond at RH =  $100 \,\%$ .

dry diameter of 100 nm are not constant over the whole RH range. Due to the solution non-ideality,  $\kappa_{HGF}$  decreases as the RH increases. In addition,  $\kappa_{CCN}$  strongly depends on whether the first or the second maximum in the Köhler curve is limiting CCN activation.

Recently researchers have observed inconsistencies between measured CCN properties of SOM particles and hygroscopic growth measured below water saturation. In other words, hygroscopic parameters measured below water saturation were inconsistent with hygroscopic parameters measured above water saturation. Several reasons have been put forward to explain these discrepancies (Petters et al., 2006, 2009; Prenni et al., 2007; Juranyi et al., 2009; Good et al., 2010; Massoli et al., 2010; Hersey et al., 2013; Pajunoja et al., 2015). The results shown in panel b of Figs. 5–6 illustrate that such inconsistencies are expected for systems with LLPS when the water uptake at subsaturated conditions represents the hygroscopicity of the organic-rich phase while the barrier for CCN activation is determined by the second maximum in the Köhler curve when the particles are water rich. Additional laboratory studies are needed to determine whether LLPS occurs in  $\alpha$ -pinene SOM generated with particle mass concentrations typically found in the atmosphere. Additional studies are also needed to determine if LLPS occurs in other types of SOM particles of atmospheric relevance.

# The Supplement related to this article is available online at doi:10.5194/acp-16-7969-2016-supplement.

Acknowledgements. This work was supported by the Natural Sciences and Engineering Research Council of Canada. Support from the US National Science Foundation and the US Department of Energy is also acknowledged. Claudia Marcolli acknowledges the Competence Center Environment and Sustainability of the ETH Domain (CCES) project OPTIWARES for financial support and Andreas Zuend for providing the program to perform the model

calculations. The authors would also like to thank Doug Worsnop for enthusiastic and motivating discussions related to the current manuscript.

Edited by: A. Virtanen

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