



Analysis of ozone and nitric acid in spring and summer Arctic pollution using aircraft, ground-based, satellite observations and MOZART-4 model: source attribution and partitioning

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Abstract. In this paper, we analyze tropospheric O₃ together with HNO₃ during the POLARCAT (Polar Study using Aircraft, Remote Sensing, Surface Measurements and Models, of Climate, Chemistry, Aerosols, and Transport) program, combining observations and model results. Aircraft observations from the NASA ARCTAS (Arctic Research of the Composition of the Troposphere from Aircraft and Satellites) and NOAA ARCPAC (Aerosol, Radiation and Cloud Processes affecting Arctic Climate) campaigns during spring and summer of 2008 are used together with the Model for Ozone and Related Chemical Tracers, version 4 (MOZART-4) to assist in the interpretation of the observations in terms of the source attribution and transport of O₃ and HNO₃ into the Arctic (north of 60° N). The MOZART-4 simulations reproduce the aircraft observations generally well (within 15%), but some discrepancies in the model are identified and discussed. The observed correlation of O₃ with HNO₃ is exploited to evaluate the MOZART-4 model performance for different air mass types (fresh plumes, free troposphere and stratospheric-contaminated air masses).

Based on model simulations of O₃ and HNO₃ tagged by source type and region, we find that the anthropogenic pollution from the Northern Hemisphere is the dominant source of O₃ and HNO₃ in the Arctic at pressures greater than 400 hPa, and that the stratospheric influence is the principal contribu-

tion at pressures less than 400 hPa. During the summer, intense Russian fire emissions contribute some amount to the tropospheric columns of both gases over the American sector of the Arctic. North American fire emissions (California and Canada) also show an important impact on tropospheric ozone in the Arctic boundary layer.

Additional analysis of tropospheric O₃ measurements from ground-based FTIR and from the IASI satellite sounder made at the Eureka (Canada) and Thule (Greenland) polar sites during POLARCAT has been performed using the tagged contributions. It demonstrates the capability of these instruments for observing pollution at northern high latitudes. Differences between contributions from the sources to the tropospheric columns as measured by FTIR and IASI are discussed in terms of vertical sensitivity associated with these instruments. The first analysis of O₃ tropospheric columns observed by the IASI satellite instrument over the Arctic is also provided. Despite its limited vertical sensitivity in the lowermost atmospheric layers, we demonstrate that IASI is capable of detecting low-altitude pollution transported into the Arctic with some limitations.

1 Introduction

As a part of the international POLARCAT (Polar Study using Aircraft, Remote Sensing, Surface Measurements and Models, of Climate, Chemistry, Aerosols, and Transport) program of field observations of Arctic atmospheric composition during the 2007–2008 International Polar Year, the Arctic Research of the Composition of the Troposphere from Aircraft and Satellites (ARCTAS) and Aerosol, Radiation, and Cloud Processes affecting Arctic Climate (ARCPAC) missions were conducted during spring and summer 2008 to better understand factors driving change in the atmospheric composition and climate of the Arctic (Jacob et al., 2010; Brock et al., 2011). Cool surface temperature and stable conditions characteristic of the Arctic atmosphere facilitate low-altitude isentropic transport into this region, making it a major receptor for mid-latitude pollution (Klonecki et al., 2003; Stohl, 2006). This process produces the so-called “Arctic Haze”, a recurring phenomenon observed every winter and spring, that results from long-range transport of aerosols and persistent pollutants such as mercury and ozone (O_3). Previous studies based on Arctic measurements from ground stations, aircraft, sondes and satellites, which provide observations of several trace species, show that the Arctic troposphere is strongly influenced by emissions from Europe, North America and Asia. They particularly showed that transport from anthropogenic source regions occurs through different processes at different altitudes: near-surface pollution is dominated by Northern European and North American sources, while Asian sources are the dominant contributors at higher altitudes (e.g., Stohl, 2006; Shindell et al., 2008; Fisher et al., 2010; Tilmes et al., 2011). They also highlight the strong impact of boreal forest fires in North America, Asia and Eastern Europe on the Arctic atmospheric pollution (e.g., Generoso et al., 2007; Leung et al., 2007; Turquety et al., 2007; Stohl et al., 2007; Fisher et al., 2010; Tilmes et al., 2011). However, despite the major interest in Arctic pollution for air quality and climate concerns, there remains considerable uncertainty concerning the source-receptor relationships between mid-latitude and Arctic atmospheric pollution throughout the troposphere (Shindell et al., 2008).

The ARCTAS and ARCPAC aircraft campaigns, which cover periods of several weeks in spring and summer, provide a large and suitable observational dataset for studying the long-range transport of pollution into the Arctic. The spring period offers a good opportunity to observe Asian influence, which is less understood than transport from Europe. This period is also favorable in terms of photochemistry controlling ozone production (Jacob et al., 2010) and of stratosphere-troposphere exchange affecting tropospheric ozone and nitric acid (HNO_3) in the Arctic. A primary goal of the summer deployment of ARCTAS was to better characterize boreal fire emissions, the chemical evolution within the fire plumes, and the regional scale impact of these forest fires. Another research theme of ARCTAS was to better un-

derstand the chemical processes controlling nitrogen oxides (NO_x) in the Arctic (Jacob et al., 2010). NO_x plays an essential role in the processes that control the ozone abundance in the lower atmosphere. HNO_3 is one of the principal reservoir species for the nitrogen oxides. Ozone and nitric acid can either be produced in mid-latitude source regions or during transport into the Arctic from the NO_x precursors. Ozone production is limited by oxidation of NO_x to reservoirs, such as PAN (peroxyacetylnitrate) and HNO_3 . Correlations between NO_x , NO_y (the sum of all reactive nitrogen species) and O_3 have frequently been used as a probe of chemistry and transport in the lower atmosphere (Murphy et al., 1993; Trainer et al., 1993; Ridley, et al., 1994; Singh et al., 1996; Fischer et al., 2000; Zellweger et al., 2003). Strong correlations between HNO_3 and O_3 were established in the lower stratosphere/upper troposphere and have been used previously to characterize air masses in the Arctic and midlatitude regions (Bregman et al., 1995; Talbot et al., 1997; Schneider et al., 1999; Neuman et al., 2001; Popp et al., 2009), to validate satellite HNO_3 profiles (Irie et al., 2006; Popp et al., 2009), or to infer the efficiency of ozone production (Cooper et al., 2011). Study of the correlation of HNO_3 and O_3 in the lower/middle troposphere is less common, despite its importance in quantifying the dependence of ozone on nitrogen compounds. The DC-8 aircraft from the spring and summer ARCTAS mission, as well as the concurrent NOAA WP-3D aircraft for the ARCPAC campaign (Brock et al., 2011), provide relevant information for this purpose.

Another central motivation for the ARCTAS campaign was to augment and interpret satellite observations of atmospheric composition (Jacob et al., 2010). Here, we show that the high spatiotemporal coverage of the Infrared Atmospheric Sounding Interferometer (IASI) instrument, launched onboard the polar orbiting MetOp satellite in October 2006, provides an excellent platform for observing the long-range transport of mid-latitude pollution into the Arctic. The comparison of satellite retrievals with aircraft data and chemistry transport model simulations provides a useful evaluation of the capability of satellites to measure Arctic atmospheric composition.

We present in this paper an analysis of the sources and the transport of Arctic pollution of O_3 and HNO_3 during spring and summer 2008 using a combination of the MOZART-4 chemical transport model, in situ measurements from aircraft (ARCTAS and ARCPAC campaigns) and remote measurements from ground-based Fourier Transform InfraRed (FTIR) and IASI spectrometers. First, a brief description of the model and measurements is given, followed by an evaluation of the model using observations of O_3 and HNO_3 during the campaigns in Sect. 3. In Sect. 4, the source attribution and their partitioning derived from the “tagging method” are presented. In Sect. 5, tropospheric observations of O_3 from FTIR instruments, at the Thule (Greenland) and Eureka (Canada) polar sites, and from the IASI satellite during the aircraft campaigns are discussed in comparison with

MOZART-4 results. We also use IASI observations to further investigate the capability of the instrument to observe the sources of pollution and the transport into the Arctic. Conclusions are given in Sect. 6.

2 Model and measurements

2.1 MOZART-4 model description

For this paper, model simulations were performed with the MOZART-4 global 3-D chemical transport model (Emmons et al., 2010a). It was driven by offline meteorological fields from the National Centers for Environmental Prediction (NCEP) Global Forecast System (GFS), with a horizontal resolution of $1.4^\circ \times 1.4^\circ$ (approximately 140 km) with 64 levels from the surface up to 2 hPa. MOZART-4 was run with its standard chemical mechanism, including 97 species and approximately 200 reactions (Emmons et al., 2010a). As described in detail in Emmons et al. (2010a), MOZART-4 does not have a complete chemistry in the stratosphere. For specifying long-lived species in the stratosphere (including O_3 , CO, CH_4 , HNO_3 , NO_x), it uses zonal mean climatological values above 50 hPa, coming from MOZART-3 simulations (Kinnison et al., 2007), and then it constrains the mixing ratios by relaxation toward the climatology down to the tropopause. MOZART-4 is also not complete for the troposphere since it does not include halogen chemistry. Our model simulations, covering the period April 2008–July 2008, have been initialized by model simulations at $2.8^\circ \times 2.8^\circ$ starting July 2007 with 42 levels in the vertical. MOZART-4 simulations of O_3 and related tracers have been previously compared to numerous in situ observations and used to track the intercontinental transport of pollution (e.g., Emmons et al., 2010b; Pfister et al., 2006, 2008).

The anthropogenic emissions used here are taken from the 2006 inventory of Zhang et al. (2009). This emission inventory is built upon the INTEX-B inventory with additional data from other inventories to obtain global emissions (see <http://www.cgrrer.uiowa.edu/arctas/emission.html> for more information). These emissions are constant in time, with no monthly variation. The emissions for VOCs are only available as total VOCs, so speciation to the MOZART-4 VOCs was based on the VOC speciation of the RETRO emissions inventory, as in Lamarque et al. (2010). The fire emissions were calculated using the global Fire Inventory from NCAR (FINN) version 1 (Wiedinmyer et al., 2011). Emissions for individual fires, based on daily MODIS fire counts, were calculated and then gridded to the simulation resolution (Wiedinmyer et al., 2006, 2011). Aircraft emissions of NO, CO and SO_2 from general aviation and military traffic for 1999 are also taken into account (Emmons et al., 2010a). Other natural emissions, for example, NO from soil and lightning are also included in the standard MOZART-4 scheme. Soil NO emissions are a combination of interac-

tive natural emissions (Yienger and Levy, 1995) and fertilizer use (Bouwman et al., 2002). The lightning parameterization is based on cloud-top height (Emmons et al., 2010a). In MOZART-4, NO_x is emitted as NO and the partitioning between NO and NO_2 is calculated explicitly in the chemistry.

Global and regional model NO_x emissions for the 1–19 April 2008 and the 18 June–13 July periods are summarized in Table 1. Model NO_x emissions using these inventories for these two periods are shown in Fig. 1. The highest NO_x surface emissions for April 2008 are due to particularly intense biomass burning over Southeast Asia and Southern Russia, where values of $4 \times 10^{-10} \text{ kg } (NO_x) \text{ m}^{-2} \text{ s}^{-1}$ are introduced in the model. For summer 2008, some fires are observed over Southern Russia, over California and over central Canada.

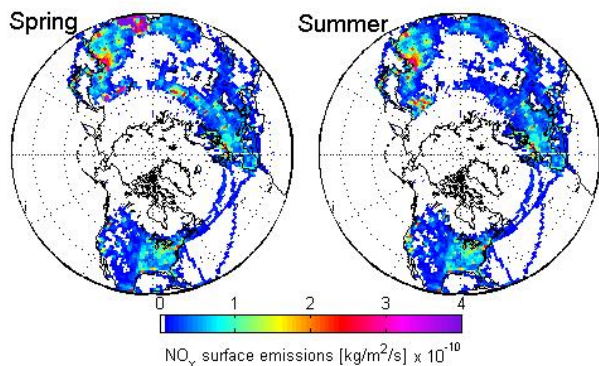
2.2 Aircraft, ground-based FTIR and satellite data

The ARCTAS mission acquired measurements using three research aircrafts, including the NASA DC-8 with a detailed chemical payload (Jacob et al., 2010). The spring ARCTAS-A deployment took place from 1 to 19 April 2008. The DC-8 was based in Fairbanks, Alaska, but made two overnight flights to Thule, Greenland and Iqaluit, Nunavut. The summer ARCTAS-B deployment took place from 26 June to 13 July 2008. The DC-8 had a two-night stay in Thule to overfly Summit, Greenland and focused on sampling North American biomass burning outflow. The summer ARCTAS-CARB deployment (18–24 June) was focused on California air quality, targeting wildfires in Northern California and urban and agricultural pollution. The ARCPAC airborne field experiment, which took place in Alaska from 1 to 23 April 2008, was coordinated with the POLARCAT activity and focused on the factors driving the radiative forcing by aerosols and tropospheric ozone in the Arctic (Brock et al., 2011). It used a NOAA WP-3D aircraft, with extensive in situ instrumentation for aerosols, ozone and related species and with flights coordinated with the spring ARCTAS aircraft mission concurrently based in Fairbanks.

Here, we use the O_3 and HNO_3 data from the NASA DC-8 aircraft to study the long-range transport of pollution from mid-latitudes into the Arctic. We consider all flights from the three deployments (ARCTAS-A, ARCTAS-B and ARCTAS-CARB) of the ARCTAS mission, covering two periods: the spring period (1–19 April 2008) and the summer period (18 June–13 July). We also consider data from all flights of the NOAA WP-3D ARCPAC mission (1–23 April 2008) during the spring period. Measurements of O_3 and HNO_3 were made respectively by NO chemiluminescence and chemical ionization mass spectrometry (CIMS) aboard the DC-8 and the WP-3D aircrafts (Weinheimer et al., 1994; Ryerson et al., 1998; Crouse et al., 2009; Huey et al., 2004; Brock et al., 2011 and references therein). The observations have an estimated uncertainty of 0.05 ppbv + 4 % for O_3 (Ryerson et al., 1998), of $25 \% \pm 0.15$ ppbv for WP-3D HNO_3 (Huey et al.,

Table 1. Global and regional NO_x emission sources used in MOZART-4 simulations for the ARCTAS spring (1–19 April) and summer (18 June–13 July 2008) campaign periods.

Sources	NO _x emissions (Tg)	
	Spring period (19 days)	Summer period (26 days)
Anthropogenic Total	3.73	5.11
North America (180–45° W, 20–90° N)	0.76	1.05
Europe (15° W–60° E, 30–90° N)	1.22	1.67
Asia (60–180° E, 0–90° N)	1.22	1.68
Biomass burning Total	1.66	1.14
USA (130–60° W, 25–50° N)	0.021	0.051
Canada (170–50° W, 50–90° N)	0.0013	0.050
North Asia (30–180° E, 40–70° N)	0.62	0.49
South Asia (60–180° E, 0–40° N)	0.92	0.017
Lightning Total [Tg-N]	0.29	0.41
Other sources (Soil emission + aircraft)	0.60	0.94
TOTAL	6.28	7.60

**Fig. 1.** Model NO_x surface emission fluxes during the ARCTAS spring (1–19 April 2008) and the ARCTAS summer (18 June–13 July 2008) campaign periods.

2004), and of 30 % ± 30 pptv for DC-8 HNO₃ (Crouse et al., 2006). As described in Brock et al. (2011), instrument comparisons were performed during separate coordinated flights with the NASA DC-8 and the NOAA WP-3D at the same altitude. Results from the comparison can be found at: http://www-air.larc.nasa.gov/TAbMEP2_polarcat.html. The preliminary POLARCAT O₃ assessment report estimates a deduced relative bias between the DC-8 and WP-3D instruments as the sum of: 0.904 (ppb) + 0.0480*DC-8 O₃ (ppb) with the DC-8 taken as an arbitrary reference. The comparison between DC-8 and WP-3D instruments for HNO₃ can be found at: <http://www-air.larc.nasa.gov/cgi-bin/ic2008r>. The ambient levels reported by the DC-8 instrument during the comparison period were within the reported imprecision of

the WP-3D instrument making it difficult to assess any bias for HNO₃ between these instruments.

For comparisons between MOZART-4 model simulations and aircraft observations, the MOZART-4 model results from 6-h window averages are sampled along the flight track at the same location, altitude and time of the observations, making use of 1-min average aircraft measurements.

Additional independent comparisons of O₃ model results with ground-based FTIR solar absorption measurements are also performed for two sites: Eureka (80° N, 86° W) and Thule (77° N, 69° W). These measurements were made with Bruker 125 HR FTIR spectrometers, operated in the framework of the Network for the Detection of Atmospheric Composition Change (NDACC) (Kurylo, 1991; Kurylo and Zander, 2000). The retrieval of trace gas profiles from the FTIR spectra are performed using the SFIT2 algorithm, a solar transmittance and retrieval model based on the optimal estimation method (OEM), and a specific set of microwindows (Batchelor et al., 2009; Hannigan et al., 2009; Lindemaier et al., 2010). The typical total error on the retrieved O₃ total columns is ~10 % (Batchelor et al., 2009). During the aircraft campaign periods, measurements at Thule were only available from 1 to 15 April, while a larger set of measurements at Eureka were made coinciding with the spring and summer campaigns. Typical FTIR averaging kernels, which express the sensitivity of the retrieval to the abundance of the target species within the layers throughout the atmospheric column, are shown in Fig. 2 (left) for the retrieved tropospheric layers and for the tropospheric and total columns. The analysis of the averaging kernels indicates that most of the retrieved information comes from the measurements.

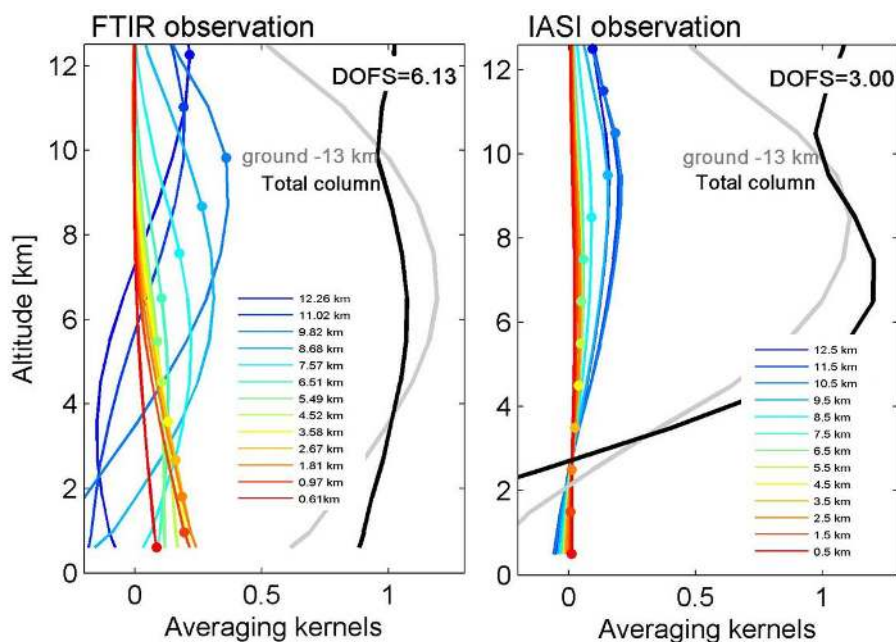


Fig. 2. Typical FTIR (left) and IASI (right) O_3 averaging kernels in partial column units ($\text{molecules cm}^{-2}/\text{molecules cm}^{-2}$) for different retrieved layers for O_3 at Eureka (10 July).

They reveal a DOFS (Degrees Of Freedom for Signal, which represents the pieces of independent information contained in the measurement) ranging from about 5.8 to 6.5 over the spring and the summer campaigns, with a maximum sensitivity around 6 km.

O_3 profiles from the IASI infrared sounder are also used here for further comparison with MOZART-4 results. IASI measures the thermal infrared emission of the Earth-atmosphere system between 645 and 2760 cm^{-1} with a field of view of 2×2 circular pixels on the ground, each of 12 km diameter at nadir. The IASI measurements are taken at nadir every 50 km along the track of the satellite and they are also taken across-track over a swath width of 2200 km. IASI offers in its standard observing mode a global coverage twice a day with overpass times at 09:30 and 21:30 mean local solar time. IASI measures concentrations of several gas-phase species including O_3 , HNO_3 , CO , and detects fires and associated smoke plumes (Clerbaux et al., 2009; Wespes et al., 2009; Coheur et al., 2009; Turquety et al., 2009; Pommier et al., 2010). O_3 profiles are retrieved with the FORLI- O_3 (Fast Operational/Optimal Retrievals on Layers for IASI) processing chain set up by ULB/LATMOS groups (Clerbaux et al., 2009; Dufour et al., 2011). It provides profiles on 39 layers from the surface up to 39 km, relying on a fast radiative transfer and retrieval methodology, based on the Optimal Estimation Method (Rodgers, 2000), which has been developed initially for the retrieval of HNO_3 from IASI (Wespes et al., 2009). As described in Scannell et al. (2011), the a priori information (a priori profile and a priori covariance matrix) used for the retrieval of O_3 is built

from the Logan/Labow/McPeters climatology (McPeters et al., 2007) which is a combination of data from the Stratospheric Aerosol and Gas Experiment II (SAGE II; 1988–2001), the Microwave Limb Sounder (MLS; 1991–1999) and data from balloon sondes (1988–2002) to represent the best approximation of the atmospheric state. Note that a single O_3 a priori profile and variance-covariance matrix are used. The FORLI- O_3 product is currently under validation, including polar regions (Scannell et al., 2011; Dufour et al., 2011; Anton et al., 2011). The total error on the retrieval ranges between $\sim 5\%$ (Northern latitudes) and $\sim 15\%$ (equatorial latitudes) for the surface-300 hPa column. Only daytime O_3 IASI observations with a good spectral fit (RMS of the spectral residual lower than $4 \times 10^{-8} \text{ W cm}^{-2} \text{ sr cm}^{-1}$) have been used in this study. The reason for only considering the daytime IASI observations relies on a better vertical sensitivity to the troposphere associated with a higher surface temperature and a higher thermal contrast (Clerbaux et al., 2009; Boynard et al., 2009). The cloud information from the operational processing is further used for filtering the observations, and only scenes with cloud coverage below 25% are analyzed (Clerbaux et al., 2009). An example of typical FORLI- O_3 averaging kernel functions for polar observation with a DOFS of 3 is represented on the right panel of Fig. 2. The averaging kernels show a maximum sensitivity at around 8 km with a sharp decrease of sensitivity down to the surface, which is inherent to nadir thermal IR sounding in cases of low surface temperature and low thermal contrast. This decrease of sensitivity, below around 4 km, indicates that the retrieved information principally comes from the a priori at

these altitudes. Taken globally, the DOFS ranges from ~ 2.5 in cold polar regions to ~ 3.5 at mid-latitudes and up to ~ 4.5 in hot tropical region, with a maximum of sensitivity peaking around 6–8 km altitude for almost all situations. For comparisons with MOZART-4 model simulations, the IASI retrievals are averaged over the MOZART grid and over the 6-h window of the MOZART-4 outputs.

When comparing model profiles with trace gas retrievals from remote sounders, the high resolution modeled layers must be smoothed using the averaging kernels (**A**) to be meaningfully compared to the observations. The derived modeled profile is here called $x_{\text{Model Smoothed}}$. In our case, the MOZART-4 profiles were firstly vertically interpolated to the pressure levels of the a priori profiles (x_a) used in the retrieval algorithms. Then the smoothing of the model profiles (x_{Model}) to the lower vertical resolution of the observations (ground-based FTIR or IASI) was performed following (Rodgers, 2000):

$$x_{\text{Model Smoothed}} = x_a + \mathbf{A}(x_{\text{Model}} - x_a) \quad (1)$$

In order to take into account the specific scene of each observation, the averaging kernels of the different observations contained in each MOZART grid have been considered to smooth the gridded MOZART profile.

HNO₃ measurements from FTIR and IASI instruments could not be used here due to a weak sensitivity of these instruments to this compound in the troposphere (Vigouroux et al., 2006; Wespes et al., 2009).

3 O₃ and HNO₃ observations and model evaluation along the aircraft flight tracks

Observed and simulated volume mixing ratio (VMR) of O₃ and HNO₃ along the flight tracks of ARCPAC, ARCTAS-A, -CARB and -B are shown in Fig. 3. The flight altitudes extend from 0 to 7 km for ARCPAC and from 0 to 12 km for ARCTAS. The highest values observed north of 50° N both for O₃ (red hotspots above 120 ppbv in Fig. 3) and HNO₃ (above 0.3 ppbv), principally during ARCTAS, correspond to observations made above 8 km, and show influences from stratospheric air masses. The stratospheric air masses can be diagnosed as [O₃]/[CO] > 1.25 (Hudman et al., 2007). Excluding these, we observe, along the ARCPAC flights, concentrations ranging from 20 to 90 ppbv for O₃, while they range from 0.01 to 0.12 ppbv for HNO₃. For ARCTAS, O₃ measurements range from 20 to 120 ppbv and HNO₃ measurements reach 14 ppb. The maxima for HNO₃ are observed over California during ARCTAS-CARB and they can be attributed to fires and urban pollution, as further discussed in Sect. 4.1. These observations are broadly consistent with previous measurements in urban plumes (e.g., Emmons et al., 2000; Bey et al., 2001; Neuman et al., 2006).

Figure 4 presents the vertical distribution, averaged over 0.5 km layers, of the ARCPAC and the ARCTAS aircraft ob-

servations, along with the model values interpolated to the flight tracks. The differences relative to the observations, expressed as a percentage, are shown on the right panels for each campaign. Vertical distributions of observations and model results show an increase of O₃ mixing ratios from around 50 ppbv at the surface to 90 ppbv in the upper troposphere. It is followed by a sharp increase above 8 km for ARCTAS-A and above 11 km for ARCTAS-B, reflecting the influence of stratospheric air. The latter is also observed in the vertical distributions of HNO₃. During the ARCTAS-CARB and the ARCTAS-B campaigns, a maximum for HNO₃ near the surface, followed by a decrease of VMR with altitude to around 5 km is observed. The highest values near the surface are principally representative of fresh plumes from anthropogenic pollution in addition to fires from California, for the ARCTAS-CARB, and from Canada, for the ARCTAS-B aircraft observations.

The aircraft data set shows the largest variability in the upper troposphere – lower stratosphere during ARCTAS-A and ARCTAS-B and in the boundary layer for the HNO₃ measurements during ARCTAS-CARB and ARCTAS-B (horizontal bars in Fig. 4). We also find a larger relative variability associated with the HNO₃ observations (50–200 %) than those associated with the O₃ observations (0–60 %) consistent with the much shorter lifetime of HNO₃ and to a larger relative uncertainty in the HNO₃ observations (Sect. 2.2). Relative to the aircraft measurements, the model underestimates O₃ and HNO₃ concentrations by 5–15 % and 3–35 % for overall campaign averages, respectively. The relative differences between observations and model are larger for HNO₃ (around –80 to 150 %) than for O₃ (–40 to 50 %), possibly explained by the larger variability associated with the HNO₃ observations that cannot be captured with the relatively coarse model resolution. The high correlation coefficients (*R*) between model and observations, principally for the ARCTAS data sets probably illustrate the large vertical gradient sampled during the aircraft flight tracks and the performance of the model for reproducing the variability of observations in space and time. The lower value of *R* for ARCPAC data, particularly for HNO₃, could result from the lower vertical gradient sampled during the ARCPAC flights, from the long-distance displacement and smearing of plumes from sources into the Alaskan Arctic, inducing larger model error (Fisher et al., 2010; Rastigejev et al., 2009; Tilmes et al., 2011) and from the large relative uncertainty in the HNO₃ observations at concentrations below 0.1 ppbv.

Figure 5 shows a comparison of O₃ and HNO₃ aircraft data with MOZART-4 for data representative of the altitude and latitude sampled during the ARCTAS campaign, with the flight altitude plotted on the right axis. In general, the MOZART-4 model reproduces the spatial and temporal variability of observations well. However, some discrepancies can be observed. Extremely low values of O₃ in the boundary layer in the Arctic (e.g., flights on 8 and 17 April, Fig. 5) are not reproduced in the model. A likely explanation is the

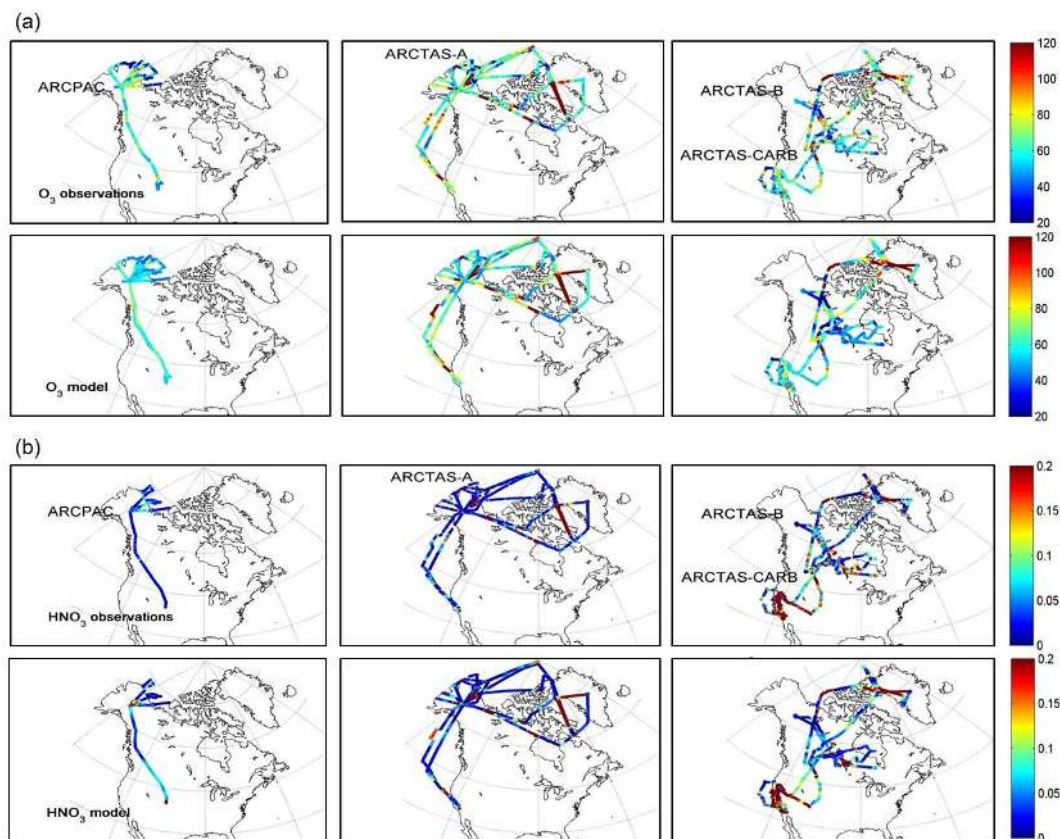


Fig. 3. (a) O₃ and (b) HNO₃ mixing ratios (ppb) observed during the ARCPAC (3 to 23 April 2008), the ARCTAS-A (1 to 19 April 2008) and the ARCTAS summer (18 June to 13 July 2008) campaigns (ARCTAS-CARB and ARCTAS-B) (top panels) compared to model values (bottom panels) sampled along the flight track at the same time, location and altitude of the observations. The flight tracks extend from 0 to 7 km for ARCPAC, from 0 to 9 km for ARCTAS-CARB and from 0 to 12 km for ARCTAS-A and -B.

missing treatment of the halogen chemistry in MOZART-4, in particular, the bromine chemistry that is responsible for the extreme ozone depletion events that occur within the Arctic boundary layer. Underestimation of HNO₃ within the Arctic boundary layer is also simultaneously observed with the overestimation of O₃. This probably results from the misrepresentation of the wet deposition processes in the Arctic boundary layer in MOZART-4. The fact that MOZART-4 systematically and simultaneously overestimates O₃ and underestimates HNO₃ within the Arctic boundary layer in April can also be observed in Fig. 4 (lowest 0.5 km). Values of around 100 ppbv or higher observed for O₃ and higher than 0.5 ppbv for HNO₃ along the highest altitude flights (e.g., flights on 8, 9 and 17 April, Fig. 5), related to stratospheric air masses, are largely underestimated in the model. This probably results from the coarse horizontal and vertical resolution of the model results, which dilutes the stratospheric influence, and from the constraint of O₃ and HNO₃ and other long-lived species to a climatology in the stratosphere in MOZART-4 (see Sect. 2.1). For HNO₃, other discrepancies are observed near the surface (between the ground and 2 km)

for flights on 9, 12 and 17 April (see Fig. 5) at Barrow, with a significant underestimation in the model. This suggests that low-altitude transport or subsidence of HNO₃-rich air masses into this region is not well simulated. It is also worth pointing out that most of the variations in both O₃ and HNO₃ are correlated with the altitude, except for the flight on 13 July near Trinidad Head. This site, which is representative of fresh urban and wildfire plumes in northern California, shows variations that are anti-correlated with the altitude, suggesting fast oxidation of NO_x to O₃ and HNO₃ in source regions.

As another measure of the model performance, the ratio O₃/HNO₃, can also be compared with that from the aircraft observations. Figure 6a and b show the O₃ vs. HNO₃ plots for both the aircraft observations and the MOZART-4 results, with the points colored by the altitude of the observations. For ARCPAC and ARCTAS-A (Fig. 6a, left), the [O₃]/[HNO₃] ratios show, on average, values of 321 for the stratospheric-influenced air masses, diagnosed by [O₃]/[CO] ratio larger than 1.25 (Hudman et al., 2007), while much higher values are observed in the troposphere (4×10^3). Identically, the ARCTAS-B measurements show a

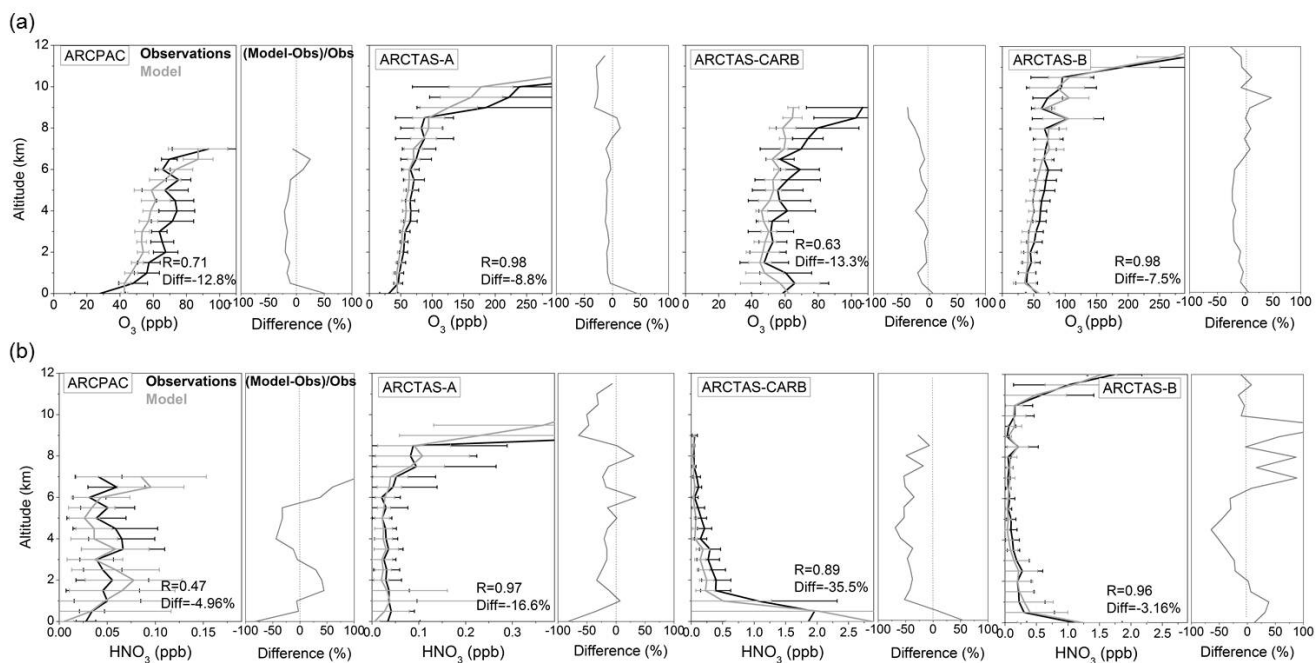


Fig. 4. Vertical distribution of O₃ (a) and HNO₃ (b) ARCPAC and ARCTAS observations averaged over 0.5 km altitude, compared to the MOZART-4 model values (from 6-h averages). Horizontal bars represent the variability associated with the observations (3σ, where σ is the standard deviation). The differences between the observations and the model are shown on the right panels for each campaign. The correlation coefficient and median of the differences are also indicated.

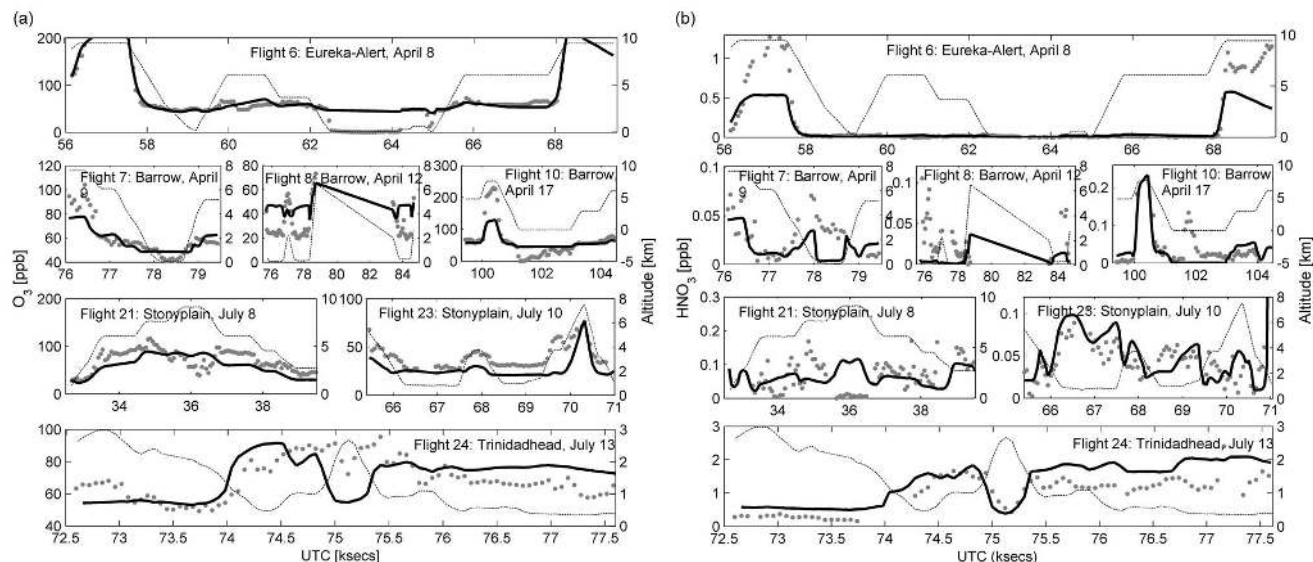


Fig. 5. (a) O₃ and (b) HNO₃ observations along the track of seven flights representative of the altitude and the latitude sampled during ARCTAS (dark grey dots). The MOZART-4 results (from 6-h averages) are interpolated to the flight path (from 1-min averages) (black line). The flight altitude (thin grey line) is plotted against the right axis.

high [O₃]/[HNO₃] ratio (5×10^3) in the free troposphere (not shown), while a ratio of 322 is measured in higher-altitude stratospheric air. However, below 2 km, averaged ratios of 453 and 652, respectively observed during ARCTAS-CARB

and ARCTAS-B, with values as low as 8, indicate weaker ozone production despite the rapid removal of HNO₃ by deposition in the boundary layer. These particular lower values characterize fresh air masses from North American urban and

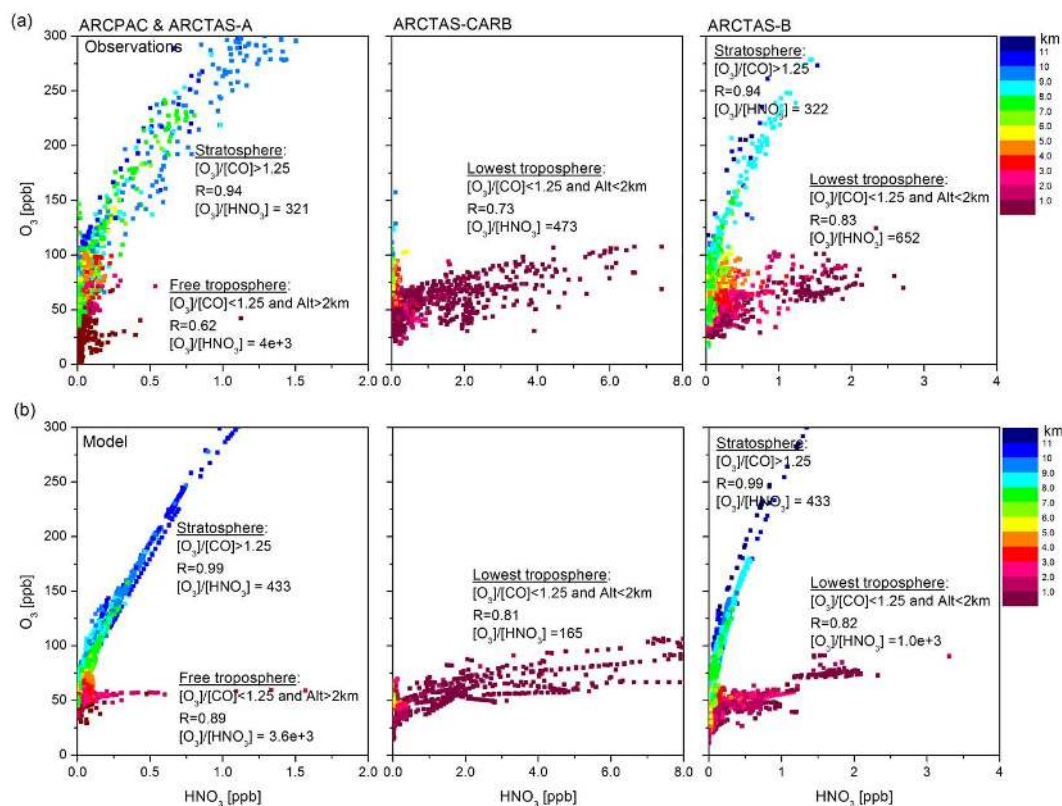


Fig. 6. O₃ vs. HNO₃ from ARCPAC and ARCTAS-A (left), ARCTAS-CARB (middle) and ARCTAS-B (right) observations (a) compared to the MOZART-4 results (b), with points colored by altitude (in km) of observations.

fire emissions (Californian for ARCTAS-CARB and Canadian for ARCTAS-B) corresponding to the highest concentrations of HNO₃ illustrated in Fig. 3. This is consistent with earlier studies which point to different HNO₃ vs. O₃ relationships in the troposphere and in the stratosphere (Bregman et al., 1995; Talbot et al., 1997; Schneider et al., 1999; Neuman et al., 2001; Popp et al., 2009) and which show lower ozone production for air masses with fresh emissions from biomass burning and urban sources than in the free troposphere (Shon et al., 2008). For each specific air mass (fresh pollution, free troposphere or stratosphere-contaminated air mass) and for each data set, we find that O₃/HNO₃ ratios and the scatter values are of comparable magnitude between the aircraft observations and the model results. Ratios as low as 8 below 2 km are well reproduced by the model. The largest differences between the ratios from observations and from model are observed for ARCTAS-CARB. They are driven by a too high formation of HNO₃ in the lowermost layer coinciding with an underestimation of O₃, as indicated in the Figs. 4 and 5. Some discrepancies in the model are also identified for the stratosphere-contaminated air masses: the ARCTAS-A observations show a relatively broad O₃/HNO₃ relationship in comparison with the observations during ARCTAS-B. This presumably reflects different dynamical and photochemical

histories of the air masses sampled by the aircraft during the spring, when the stratosphere is dynamically active, and during the summer; characterized by a more uniform photochemical history associated with summertime easterly flow and very little stratospheric wave activity. The differences in the compactness between the modeled and the observed stratospheric O₃/HNO₃ relationship can be attributed to the missing stratospheric chemistry and the use of climatological values in the stratosphere (see Sect. 2.1).

4 Source attribution of O₃ and HNO₃ Arctic pollution during spring and summer 2008

In this section, we use MOZART-4 model simulations to diagnose the contributions from the individual sources to the O₃ and HNO₃ profiles. As described in Lamarque et al. (2005) and Pfister et al. (2006, 2008), the amount of O₃ or HNO₃ produced from the different NO_x emission sources is quantified by tagging the NO emissions. The tagged NO (from the emissions shown in Fig. 1) is traced through the hydrocarbon and CO oxidation and through all the odd nitrogen species (HNO₃, PAN, N₂O₅, organic nitrates, etc.) to account for recycling of NO_x. The tagged O₃ and HNO₃ result from the photolysis and oxidation of NO₂. As explained in

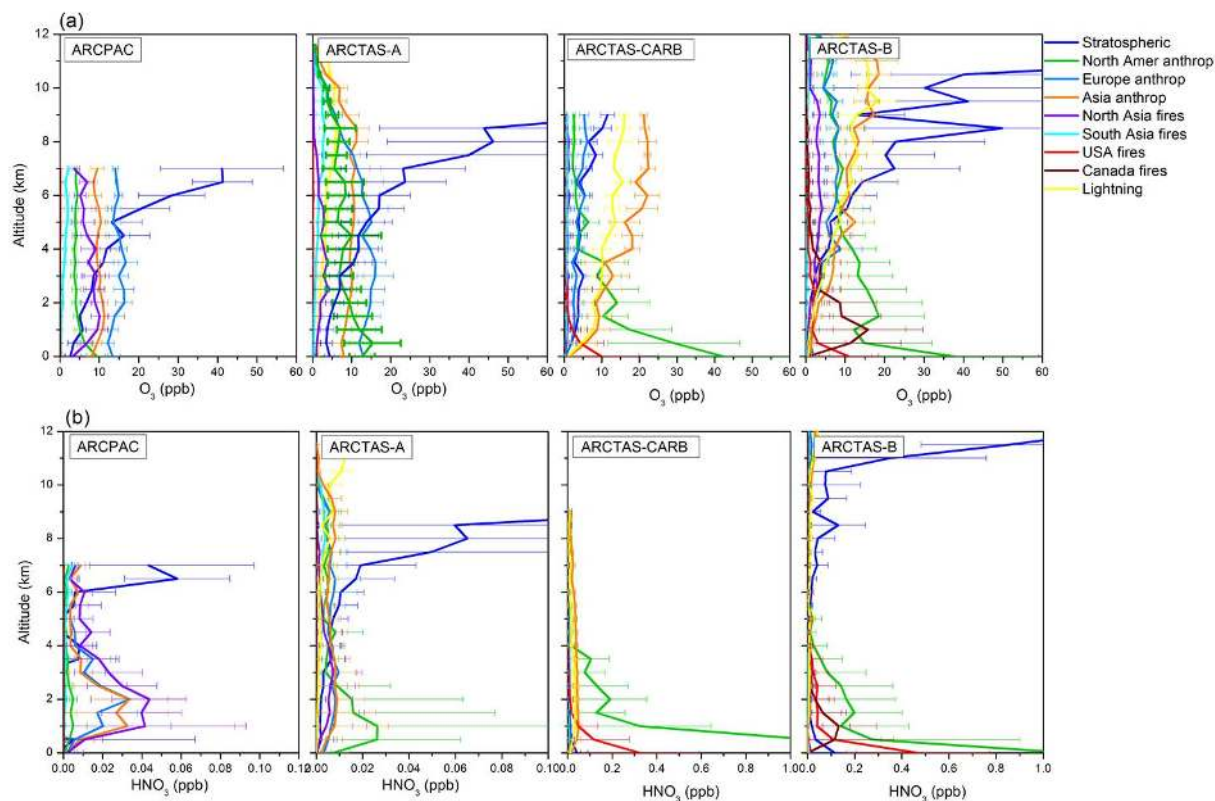


Fig. 7. Vertical distributions of O_3 (a) and HNO_3 (b) source contributions from MOZART-4 simulations along the flight tracks of the ARCPAC, ARCTAS spring and summer deployments. Data are averaged over 0.5 km layers and tagged by source type and region. Horizontal bars represent the variability associated with each tagged contribution (3σ , where σ is the standard deviation).

Emmons et al. (2010b, 2011), this tagging technique is additive. The sum of the tagged species resulting from each NO_x source tagged separately reproduces within a few percent the tagged species from the total NO_x emissions.

4.1 Partitioning along the aircraft flight tracks

Figure 7 presents the vertical distributions of O_3 and HNO_3 concentrations from the MOZART-4 simulations along the ARCPAC and ARCTAS flight tracks for different source types and regions: stratospheric, anthropogenic (North American, European and Asian), biomass burning (North Asian, South Asian, American and Canadian) and lightning-produced NO_x . For both the ARCPAC and ARCTAS campaigns, the mean concentrations of O_3 are dominated by European anthropogenic emissions at low altitude. They are followed by North American and Asian anthropogenic contributions for the ARCTAS-A measurements, while the ARCPAC measurements contain large contributions from Asian anthropogenic emissions and from North Asian fires. The latter, which reflects intense Russian fires, is much weaker along the ARCTAS-A than along the ARCPAC flight tracks (Fisher et al., 2010), while the North American anthropogenic contribution is much larger along the ARCTAS-A

than along the ARCPAC flight tracks. This reflects the different sampling strategies between the two campaigns (Jacob et al., 2010; Brock et al., 2011). Transport from the stratosphere, which is responsible for a large enhancement of tropospheric ozone in regions of subsidence (Cooper et al., 2002; Liang et al., 2009), is the predominant source above 5 km. It becomes even larger than the sum of the tropospheric contributions above 8 km. During ARCTAS-A, the mean concentrations of HNO_3 are dominated at low altitude by the North American anthropogenic contribution, while they are dominated by the Russian fires contribution during ARCPAC. Transport of anthropogenic HNO_3 from Europe and from Asia into the Arctic is lower in comparison with that of ozone. This can be explained by a shorter lifetime of HNO_3 in the lower layers (a few hours) compared to that of O_3 (a few days). Above 5 km, transport from the stratosphere dominates the HNO_3 distributions.

During the ARCTAS-CARB and the ARCTAS-B campaigns, the stratospheric influence on tropospheric ozone and nitric acid is less pronounced than during the spring. This is due to weaker stratosphere-troposphere exchanges during the summer but also to the lower latitudes sampled during ARCTAS-CARB. For both O_3 and HNO_3 , the anthropogenic

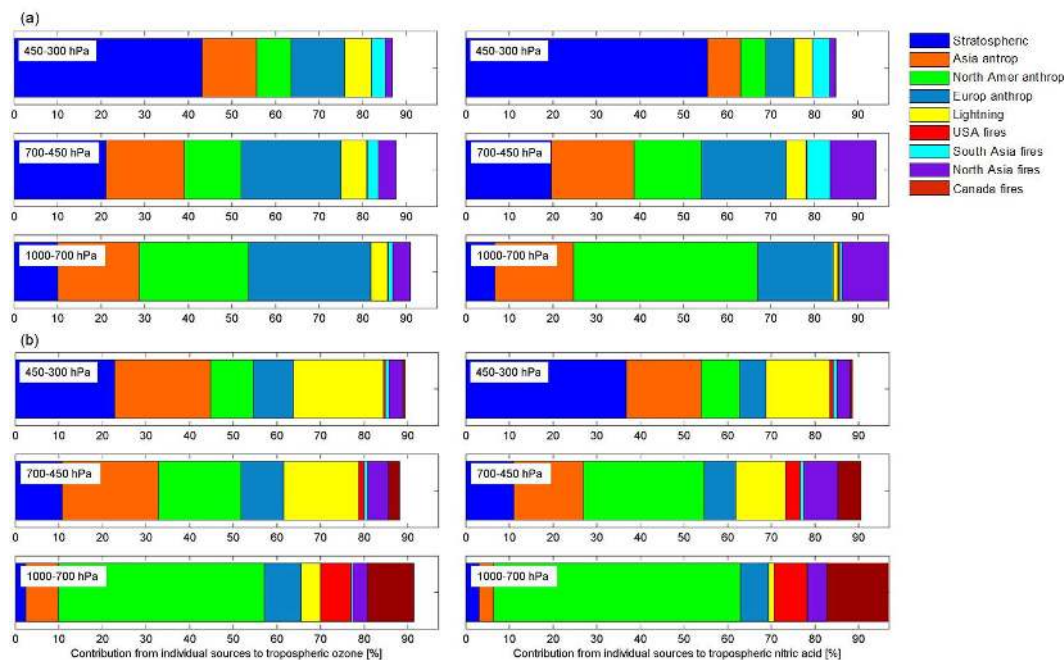


Fig. 8. Partitioning of principal sources of O_3 (left) and HNO_3 (right) in altitude bands of ground–700 hPa, 700–450 hPa and 450–300 hPa along the ARCTAS spring (a) and summer (b) flight tracks (from 18 June to 13 July), as simulated by MOZART-4, expressed in percentage (%) of the total.

contribution from North America shows a strong maximum in the lowest layers. Contribution from intense North American fire emissions (California and/or Canada) also shows an important impact on tropospheric ozone in the boundary layer. Evidence for enhanced ozone in boreal fires plumes during the ARCTAS-B campaign has already been reported in Alvarado et al. (2010). We also observe, especially for O_3 , an increasing contribution in the summer months from lightning-produced NO_x and from Asian anthropogenic pollution with altitude. The latter is in agreement with the fact that trans-Pacific transport of Asian ozone pollution mainly takes place in the free troposphere where the ozone lifetime is longer (several months) and the winds are strong (Liang et al., 2004; Price et al., 2004). Note also that, based on the model bias relative to the aircraft observations (see Sect. 3), less confidence should be attributed to the quantitative partitioning of the smallest contributors to tropospheric O_3 and HNO_3 .

For both O_3 and HNO_3 during ARCTAS-A/B, North American anthropogenic pollution sources dominate the variability in the lower troposphere indicating direct low-altitude transport into the Arctic. The ARCPAC variability is principally determined by the Russian fires emissions and the European anthropogenic pollution. Transport from the stratosphere dominates the variability in the higher altitudes. We also found that the model captures the vertical gradient in the variability (see Sect. 3, Fig. 4) but it underestimates the variability range. When separating data by anthropogenic,

stratospheric or fires influence, we found that the underestimation of O_3 and HNO_3 concentrations in the lower/middle altitudes in the model (see Sect. 3, Fig. 4) is principally dominated by an underrepresentation of local pollution sources (North American anthropogenic pollution, Californian and Canadian fires) during ARCTAS, while it is attributed to an underrepresentation of Russian fire emissions, Asian and European anthropogenic pollution during ARCPAC. At higher altitudes, for the ARCPAC, ARCTAS-A/B observations, the underestimation is related to an underrepresentation of the stratospheric influence. The underrepresentation of both modeled concentrations and variability are possibly due to a mix of uncertainties introduced by the coarse grid of the model leading to too much diffusion, problems in the transport of fine-scale plumes in the model, uncertainties in the emissions, possible bias in the NO_y partitioning, incomplete stratospheric chemistry of MOZART-4 and interpolation of the MOZART-4 results (from 6-h averages) to the flight path (1-min averages).

In order to further illustrate the tagged contributions to tropospheric O_3 and HNO_3 , Fig. 8 represents the partitioning by source type and region over three altitude ranges during the ARCTAS spring and summer campaigns, as simulated by MOZART-4. Differences between total simulated O_3 or HNO_3 (100%) and the sum of tagged contributions are principally due to NO_x emissions from soils, but also from aircraft and from ships. An interesting feature in Fig. 8 lies in the observation of increasing relative contribution, during

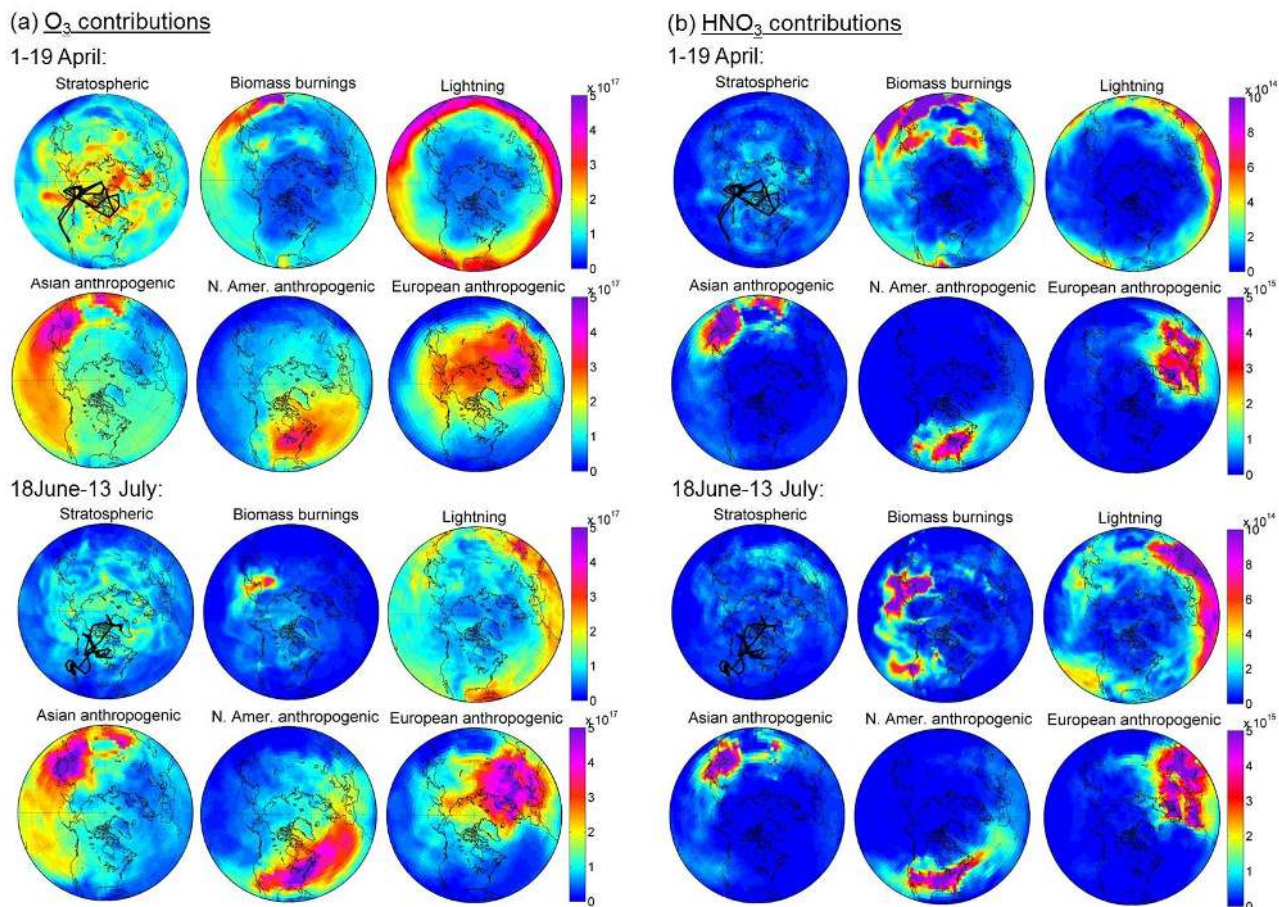


Fig. 9. Principal contributions to O_3 (a) and to HNO_3 (b) columns (molecules cm^{-2}) from the ground to 300 hPa during spring (from 1 to 19 April) and summer (from 18 June to 13 July) ARCTAS campaigns, as simulated by MOZART-4. The flight tracks corresponding to these periods are superimposed in black on the stratospheric contribution plots. Note that the color scales for anthropogenic contributions to HNO_3 are different than for the other sources.

the summer, from Asian anthropogenic pollution with altitude for HNO_3 . The partitioning from that source is larger than from other anthropogenic individual sources (American and European) at higher altitudes, suggesting that this source tends to favor HNO_3 in the upper troposphere where it is longer lived (several weeks to a month), making it a candidate for the transport of pollution (Miyazaki et al., 2003; Neuman et al., 2006). It is also worth pointing out that, both during the summer and the spring, the relative partitioning of O_3 and of HNO_3 are similar along the flight tracks of the aircraft missions, reflecting the chemical link between the two compounds.

4.2 Partitioning over the Arctic (north of 60°N)

Figure 9 shows simulated spring and summer O_3 and HNO_3 columns from the ground to 300 hPa, of each tagged tracer over the Northern latitudes. Asian, European and North American anthropogenic pollution are of comparable

magnitude (reaching 5×10^{17} and 5×10^{15} molecules cm^{-2} respectively for O_3 and for HNO_3), but with distinct geographical signatures. Sources in Europe dominate in the European sector of the Arctic over Scandinavia (4×10^{17} molecules cm^{-2} on average for O_3) reflecting northward transport, and in the Asian sector of the Arctic (3×10^{17} molecules cm^{-2} on average for O_3) through a westerly trans-Siberian transport pathway. Pollution from North America contributes only a small amount to the Arctic pollution (around 1.5×10^{17} molecules cm^{-2} on average for O_3). According to the ARCPAC observations, the Asian and the European influences dominate the pollution over Alaska. Relative to anthropogenic O_3 , which shows horizontally smoothed distributions, anthropogenic HNO_3 presents a strong horizontal gradient overall the troposphere. This difference is probably due to washout or to the uptake of HNO_3 on dust or ice (Querol et al., 2008; Zhang et al., 2008), making it shorter-lived than O_3 and therefore, less available for transport into the Arctic.

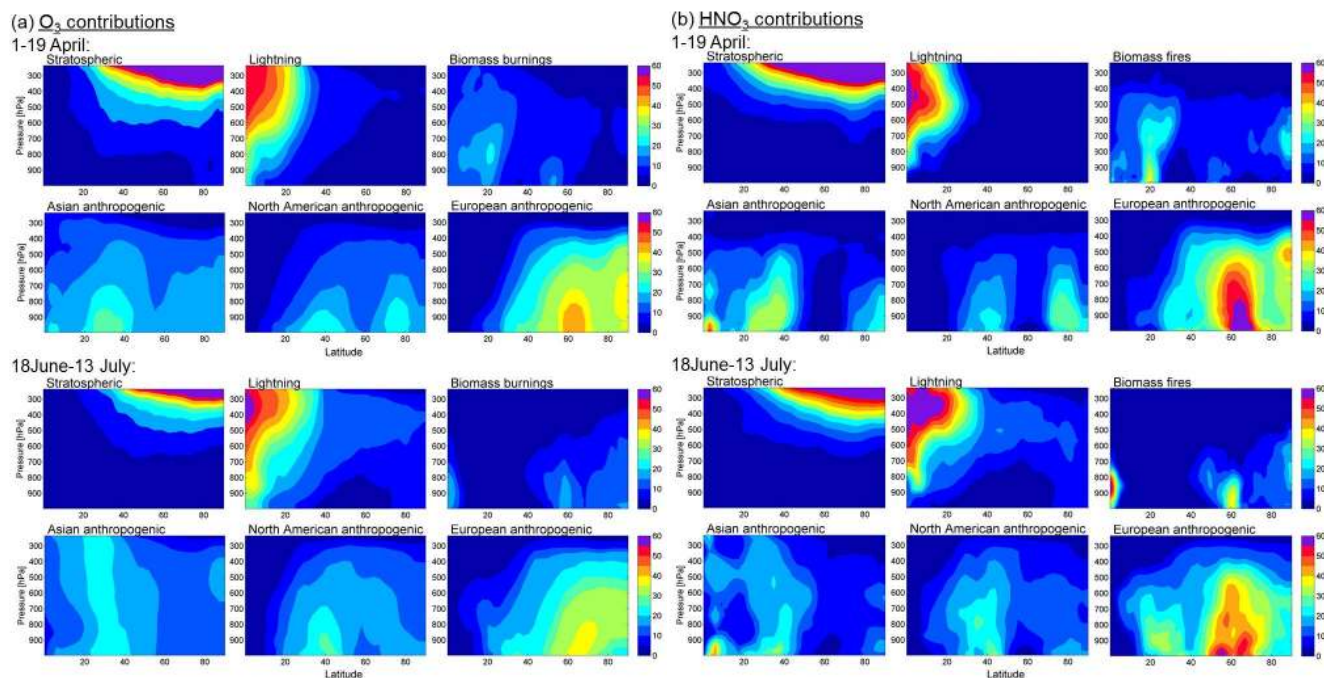


Fig. 10. Vertical and latitudinal distributions of the relative contributions (%) to O_3 (a) and HNO_3 (b) concentrations during the spring (from 1 to 19 April) and summer (from 18 June to 13 July) ARCTAS campaigns, as simulated by MOZART-4.

As a result, the anthropogenic emissions contribute only slightly to the Arctic background concentrations of HNO_3 (5×10^{14} molecules cm^{-2}). Even though they were particularly intense during the spring campaigns, the Southeast Asian fires show only a small influence on the Arctic due to the low latitude of the emissions. During the summer, Russian fire emissions greatly contribute to the tropospheric columns over Eastern Siberia (3×10^{17} molecules cm^{-2} for O_3 and 1.5×10^{15} molecules cm^{-2} for HNO_3) with some contributions over the American sector of the Arctic. Emissions from Californian fires are also observed. Finally, lightning which principally occurs at the equatorial belt also has a non-negligible contribution to the tropospheric columns of O_3 (1×10^{17} molecules cm^{-2}) and HNO_3 (1×10^{15} molecules cm^{-2}) in the Arctic principally during the summer.

Figure 10 presents the vertical and latitudinal distributions of the most important contributors to O_3 (Fig. 10a) and HNO_3 (Fig. 10b) for the spring and the summer periods, as simulated by MOZART-4. Consistent with studies of CO during ARCTAS (Fisher et al., 2010; Tilmes et al., 2011), anthropogenic pollution dominates concentrations over the Northern Hemisphere between the surface and 400 hPa, through lifting to the middle troposphere and transport towards higher latitudes. The European contribution particularly dominates throughout the Arctic, reaching up to 45 % and 60 % of the total abundance of O_3 and HNO_3 , respectively. The Asian emissions are dominant below $40^\circ N$ with a contribution of 35 % for O_3 and HNO_3 . They reach

up to 20 % of the Arctic background via uplift and transport associated with warm conveyor belts (WCBs) over China (Stohl, 2006). The much smaller North American emissions, which also involve uplift in WCBs, make a limited contribution to the Arctic concentrations of 20 % between the surface and 700 hPa. The influence of the biomass burning emissions on the Arctic composition, with a contribution of around 15 % on average for O_3 and HNO_3 , results principally from the Siberian fires (maxima at $55\text{--}60^\circ N$) occurring during spring and summer. The large fires in the Southeast Asian region in April (maxima at $20\text{--}30^\circ N$) have a minimal impact on the Arctic abundances. Their contribution is smaller than the Asian anthropogenic contribution, even in the upper troposphere where it results from transport with WCBs (Bey et al., 2001). Subsiding stratospheric air masses dominate concentrations at pressure less than 400 hPa into the Arctic, where they exceed 60 % of O_3 or HNO_3 Arctic abundance. During April, transport from the stratosphere shows a contribution of around 10 % down to the boundary layer. Contributions from lightning to the Arctic composition during April are relatively small, even at high altitude. During summer, they reach up to 20 % of the O_3 or HNO_3 Arctic background concentrations at 400–500 hPa with some descent (10 %) down to the surface. It is also worth noting that, as previously discussed, the principal differences between the distributions of O_3 and HNO_3 lie in a deeper horizontal and vertical gradients for HNO_3 concentrations.

Moreover, it is interesting to indicate that the partitioning based on the aircraft track flights (Fig. 8) is generally

Table 2. Contribution, expressed in percentage (%) of the total, of the principal sources of O₃ and HNO₃ in altitude bands of ground–700 hPa, 700–450 hPa and 450–300 hPa, based on the aircraft sampling (>60° N) and for the entire Arctic (>60° N) for the spring and the summer phases of the ARCTAS campaigns, as simulated by MOZART-4.

Contributions (%)	Spring period				Summer period			
	O ₃		HNO ₃		O ₃		HNO ₃	
Anthropogenic								
North America 1000–700 hPa	19	24	21	35	16	30	14	29
700–450 hPa	14	13	4	16	15	18	12	11
450–300 hPa	8	8	7	8	12	10	8	4.5
Europe	37	29	3	21	34	27	39	29
	31	23	5	21	30	16	33	13
	15	13	14	8	19	12	15	4.5
Asia	18	19	0	22	13	13	10	10
	15	21	0	23	15	14	10	9
	10	14	6	11	12	14	6	4
Biomass burning Total	4.5	5.5	9	14	14	8	17	14
	5	6	9	14	9	11	12	23
	3	5	4.5	9	5	6	4	3
Stratospheric	11	9.5	2.5	4	3	4	2	2
	19	19	13	17	10	16	12	21
	48	43	49	49	28	37	41	56
Lightning	3	3	3	1	8	6	7	5
	0.1	5	0.1	2	0.3	15	0.3	13
	0.006	6	0	4	0.007	13	0.002	10

Left: for the entire Arctic (>60° N)

Right: based on aircraft sampling (>60° N)

different from that for entire Arctic (Fig. 10), except for middle/upper tropospheric O₃, which is longer-lived and better mixed than HNO₃. Table 2 summarizes the contribution of the principal sources based on the aircraft sampling and for the entire Arctic, for the spring and the summer phase of the ARCTAS campaigns. The aircraft sampling patterns are observed to influence the partitioning analysis, with an overrepresentation of the North American anthropogenic contribution (~10 % for O₃ and ~15 % for HNO₃ on average) and an underrepresentation of the European anthropogenic contribution (~10 % for O₃ and ~17 % for HNO₃ on average) in the lower/middle troposphere, in comparison with partitioning from the entire Arctic. The overestimation of the biomass burning contribution to lower/middle tropospheric HNO₃, based on aircraft sampling patterns, reflects the influence from intense Asian fires in addition to the impact of Canadian fires during the summer. The lightning contribution over all the troposphere and the stratospheric contribution in the middle/upper troposphere for the summer period are also influenced by the aircraft sampling patterns.

5 Tropospheric O₃ observed by FTIR and IASI during ARCTAS

In this section, we test the ability of the FTIR and IASI instruments to observe the long-range transport of pollution into the Arctic by investigating the daily variability of O₃ measurements from IASI and from FTIR at Eureka (80° N, 86° W) and at Thule (77° N, 69° W) in comparison with the MOZART-4 results previously evaluated in Sect. 3. Based on an IASI sensitivity analysis, we consider, in this section, tropospheric columns between ground and 300 hPa both to limit as much as possible the stratospheric and the tropopause height variation influence and to contain the altitude range of maximum sensitivity in the troposphere (see Sect. 2.2). We also make a further use of the tagged tracers to assist in the interpretation of the O₃ variability observed by these instruments.

5.1 FTIR observations of Arctic pollution

Figure 11 (top) shows the comparison of model results with observations of daily mean O₃ columns from ground to 300 hPa, from FTIR observations at Eureka and at Thule during spring and summer 2008. For the purpose of the

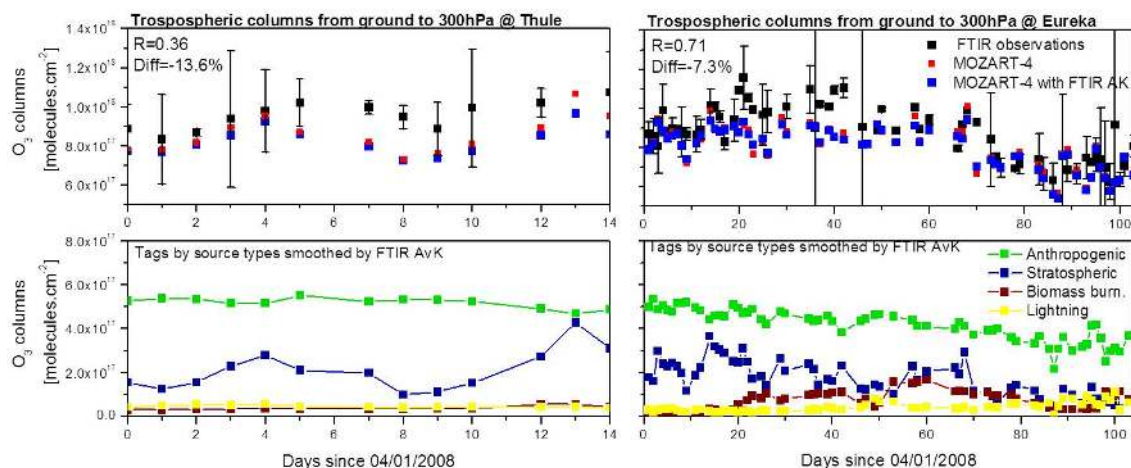


Fig. 11. Top: daily mean O₃ columns from ground to 300 hPa at Thule (left) for 1 to 15 April 2008 and at Eureka (right) for 1 April to 13 July 2008. The ground-based FTIR measurements (black) are compared to raw model values (red) and smoothed (blue) with the FTIR averaging kernels. Vertical bars represent the daily variability associated with the tropospheric columns (3σ , where σ is the daily standard deviation). The correlation coefficient and median of the differences between the observations and the model (calculated as $[(\text{model}-\text{obs.})/\text{obs.}]\cdot 100\%$) are also indicated. Bottom: contributions of principal sources to O₃ tropospheric columns over Thule (left) for 1 to 15 April 2008 and Eureka (right) for 1 April to 13 July 2008, smoothed with the FTIR averaging kernels.

comparison, MOZART-4 results have been smoothed by the FTIR averaging kernels (see Sect. 2.2, Fig. 2). Comparison of measurements with model results shows good agreement within 1×10^{17} molecules cm^{-2} , with a negative bias within -15% on average. This can be explained by the underestimation of model results of around 5–15% as discussed in Sect. 3, by the retrieval total errors (Sect. 2.2) and by the daily variation associated with the FTIR observations (vertical bars in Fig. 11).

Similarly to the partitioning analysis along the flight tracks (Sect. 4.1), MOZART-4 simulations have been used to simulate the FTIR tropospheric O₃ contributions from the principal sources: stratospheric, anthropogenic, biomass burning and lightning (Fig. 11, bottom). Following the formalism of Rodgers (2000) (Eq. (1), Sect. 2.2), it can be shown that:

$$\mathbf{x}_{\text{Model Smoothed}} = \mathbf{x}_a + \mathbf{A}[(\sum \mathbf{x}_{\text{Tracer}}) - \mathbf{x}_a] \quad (2)$$

where $\mathbf{x}_{\text{Tracer}}$ represents the profile from each tagged contribution. This expression can be separated into two components: one representing the contributions from each tracer smoothed by the averaging kernels, calculated as $\mathbf{A}(\mathbf{x}_{\text{Tracer}})$, and the second, the a priori contribution, representing the contribution from the a priori to the retrieved tropospheric columns, due to the limited vertical sensitivity of the instrument. It is calculated as $\mathbf{x}_a - \mathbf{A}(\mathbf{x}_a)$.

The observations at Thule and Eureka are dominated by the anthropogenic contribution, followed by the stratospheric influence (Fig. 11, bottom). At Thule, the anthropogenic pollution represents, on average, 54% of total concentration (23% from North America, 18% from Europe and 12% from Asia), the stratospheric fluxes count for 21% and light-

ning for 5%. Some contribution from biomass burning is also likely affecting the FTIR measurements to a weak extent (less than 4%). The small fraction from the a priori information (only 1.2%) reflects a high vertical sensitivity throughout the troposphere. The differences between the total simulation and the sum of the tagged contributions (around 15% on average) can be attributed to the contribution from soils and aircraft emissions. The day-to-day variability observed at Thule can mainly be attributed to the day-to-day variability associated with the stratospheric flux of ozone, which shows an amplitude of 3×10^{17} molecules cm^{-2} .

At Eureka, the contributions are identified as follows: 52% from the anthropogenic sources (14% from North America, 23% from Europe and 14% from Asia), 20% from the stratospheric flux, smaller influence from biomass fires (9%) and from lightning (6%) and a negligible a priori contribution (1%). The stratospheric and the anthropogenic sources have the most variability, with amplitudes of 3×10^{17} molecules cm^{-2} and 2×10^{17} molecules cm^{-2} respectively. The anthropogenic source amplitude over the campaign periods shows up as a decrease in the summer while the stratospheric source is strongly variable on a daily basis, though being also much less in summer as compared to spring. As discussed in Sect. 4.2, the contributions from lightning to the Arctic composition are small during April, but they are as large as 18% during the summer. The most elevated concentrations from fires over May and June with contributions reaching up to 20% principally result from the Siberian emissions.

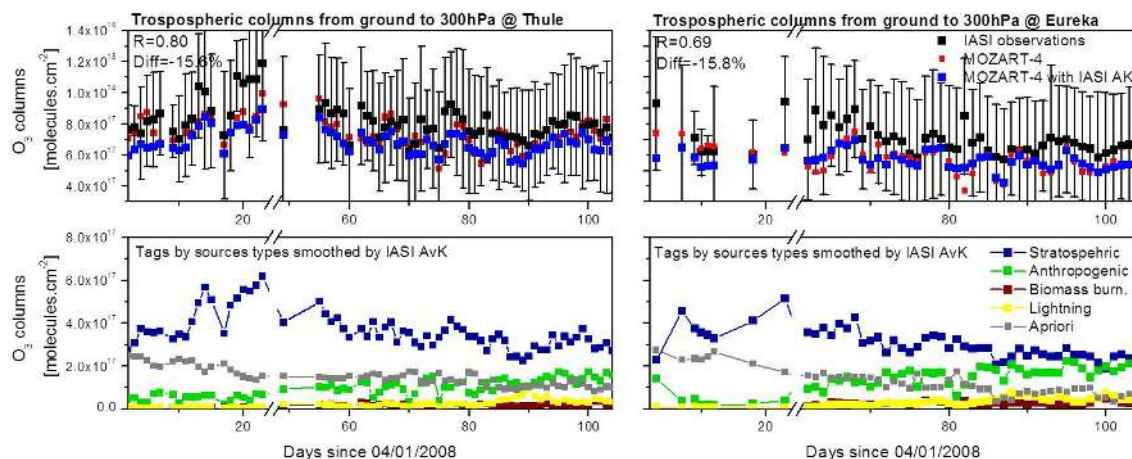


Fig. 12. Same as Fig. 11 but for comparison with IASI satellite measurements within the $1.4^{\circ} \times 1.4^{\circ}$ MOZART-4 grid around Thule (left) or Eureka (right).

5.2 IASI observations of Arctic pollution

Figure 12 is similar to Fig. 11, but for the IASI observations extracted within the MOZART grid box around Thule or Eureka. Here, MOZART-4 results have been smoothed by the IASI averaging kernels (see Sect. 2.2, Fig. 2). Comparison of the measurements with model results shows an agreement within $1\text{--}2 \times 10^{17}$ molecules cm^{-2} on average, which corresponds to a negative average bias of around -15% . It probably reflects, on one hand, the tendency of IASI retrievals to overestimate the tropospheric ozone columns by $\sim 10\%$ in the Northern latitudes (Boynard et al., 2009), and, on the other hand, the underestimation of the concentrations from MOZART-4 in these latitudes (see Sect. 3). It also can be largely explained by the IASI daily variations (3σ values of $\sim 6 \times 10^{17}$ molecules cm^{-2} on average).

Contrary to the FTIR observations, the observations from IASI at Thule and Eureka are dominated by the stratospheric contribution. This results from a higher sensitivity of IASI in the higher layers where the stratospheric influence represents the largest contributions (Figs. 7 and 10). This also makes the underrepresentation of the stratospheric influence in the model (reaching 25%, as previously discussed when comparing aircraft data and model results in Sect. 3) largely responsible for the differences between the IASI observations and the model results. The IASI measurements are also more affected by the a priori information than the ground-based FTIR measurements, because of a lower vertical sensitivity of IASI in the lowest layers. As a consequence, IASI detects smaller contributions from anthropogenic pollution and fires than the FTIR instruments.

At Thule, the stratospheric and the anthropogenic contributions represent, on average, 53% and 14% of the total concentration (with 5.5% from North America, 4.5% from Europe and 4.0% from Asia) respectively. The a priori infor-

mation contributes importantly with 22% over the entire troposphere. The contribution from fires, mainly from Siberia, and from lightning is significantly weaker (less than 3.5% each on average). Contributions of 6% from soils and aircraft emissions were calculated. The day-to-day variability can mainly be attributed to the variability of the ozone stratospheric contribution, which shows large amplitudes up to 3×10^{17} molecules cm^{-2} , as indicated by the MOZART-4 simulations.

At Eureka, the stratospheric influence is again the most important (46.1% on average) decreasing over the summer. The relative contribution from the anthropogenic sources accounts for 17% (5.5% from North America, 6.5% from Europe and 5.0% from Asia). The a priori contribution represents a fraction of 19%. Influences from fires and from lightning contribute to the Eureka background concentration with a fraction smaller than 4.5% each. Both the stratospheric and the anthropogenic sources have the maximum variability, with amplitudes of 3×10^{17} molecules cm^{-2} and 2×10^{17} molecules cm^{-2} respectively. It is also worth pointing out that IASI retrievals at Thule and at Eureka are characterized by a sharp decrease of the a priori contribution between April and July observations, with contributions ranging from 3×10^{17} molecules cm^{-2} to 5×10^{16} molecules cm^{-2} . The larger contribution during the spring reflects a lack of sensitivity probably due to both smaller thermal signal above ice-covered surfaces and smaller thermal contrast in the cold boundary layers.

In order to further test the ability of IASI to observe the transport and the variability over the scale of the Arctic, Fig. 13 shows the mean O_3 columns from ground to 300 hPa during the ARCTAS spring and summer periods as observed by IASI and compared to MOZART-4 model values. The horizontal distributions of the Northern latitude concentrations show similar patterns between the observations and the

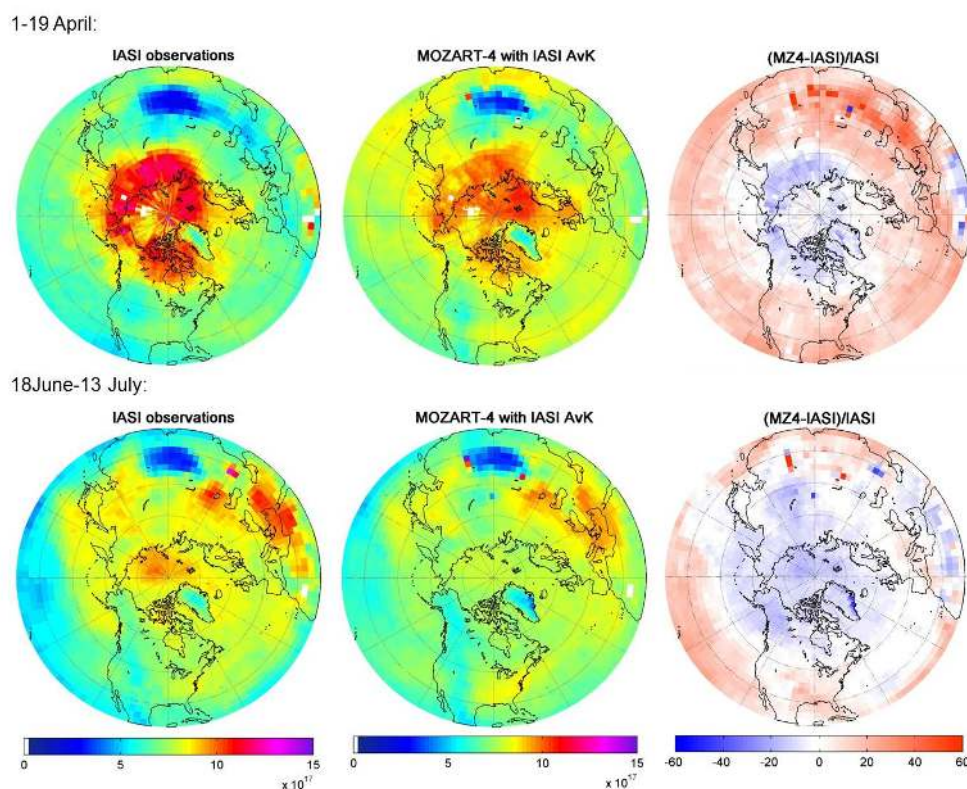


Fig. 13. Mean O_3 columns ($\times 10^{17}$ molecules cm^{-2}) from ground to 300 hPa during the spring (from 1 to 19 April) and summer (from 18 June to 13 July) campaign periods, observed by the IASI satellite instrument and simulated by MOZART-4, smoothed according to the averaging kernels of the IASI observations. The percent differences between the observations and the model results (calculated as $[(\text{model} - \text{obs.})/\text{obs.}] \cdot 100\%$) are shown on the right panel.

simulations. During the spring, both IASI and MOZART-4 show highest concentrations over the Arctic, while during the summer, the highest values are observed above south-west Asia. Observations and model datasets are reasonably well correlated ($R = 0.79$ during the spring and $R = 0.88$ during the summer campaign periods) and the mean relative differences ($4.1 \pm 39.4\%$ and $-4.4 \pm 29.9\%$, respectively for the two periods) indicates the absence of a systematic bias between the model and the observations. The high differences (averaged negative bias of around 20%) observed over the Arctic can be explained by the underestimation of the stratospheric contribution in the model (as previously discussed for observations at Eureka and at Thule), while the averaged positive bias (around 20%) measured at low latitudes could be mostly attributed to the tendency of MOZART-4 to overestimate the O_3 concentrations in the Northern Tropics (Emmons et al., 2010a) and to larger total retrieval errors on the IASI tropospheric columns, principally found at the intertropical latitudes ($\sim 15\%$ on average, Sect. 2.2). These larger total errors are possibly due to the stronger impact of the water vapor lines on O_3 retrievals at these latitudes (Hurtmans et al., 2011). Dilution on the model spatial scale (~ 140 km) compared to IASI spatial scale (~ 12 km at nadir)

could also explain a part of the differences. These differences lead to a stronger latitudinal gradient from IASI observations in comparison to that from the model.

Figure 14 shows O_3 columns from ground to 300 hPa as observed by IASI and modeled by MOZART-4 during two specific episodes of pollution transport across the Arctic: 14–16 April during the spring campaign and 2–5 July during the summer campaign. Principal contributions (stratospheric, anthropogenic, fires and lightning) as simulated by the model and smoothed using the IASI averaging kernels are also shown for the purpose of our analysis. Both IASI and MOZART-4 show elevated columns over Northern latitudes. Based on modeled source contributions, this enhancement reflects principally a strong stratospheric influence during the spring, while it mainly results from a mix of anthropogenic pollution from Europe, Asia and North America during the summer. Similar to the analysis presented in Sect. 4.2., the highest values for anthropogenic pollution (up to 7×10^{17} molecules cm^{-2}) are observed over the European sector of the Arctic during the two periods (a contribution of $\sim 50\%$ to the tropospheric column), followed by maxima of 5×10^{17} molecules cm^{-2} over the Asian and the American sectors (see Sect. 4.2). Large anthropogenic plumes from

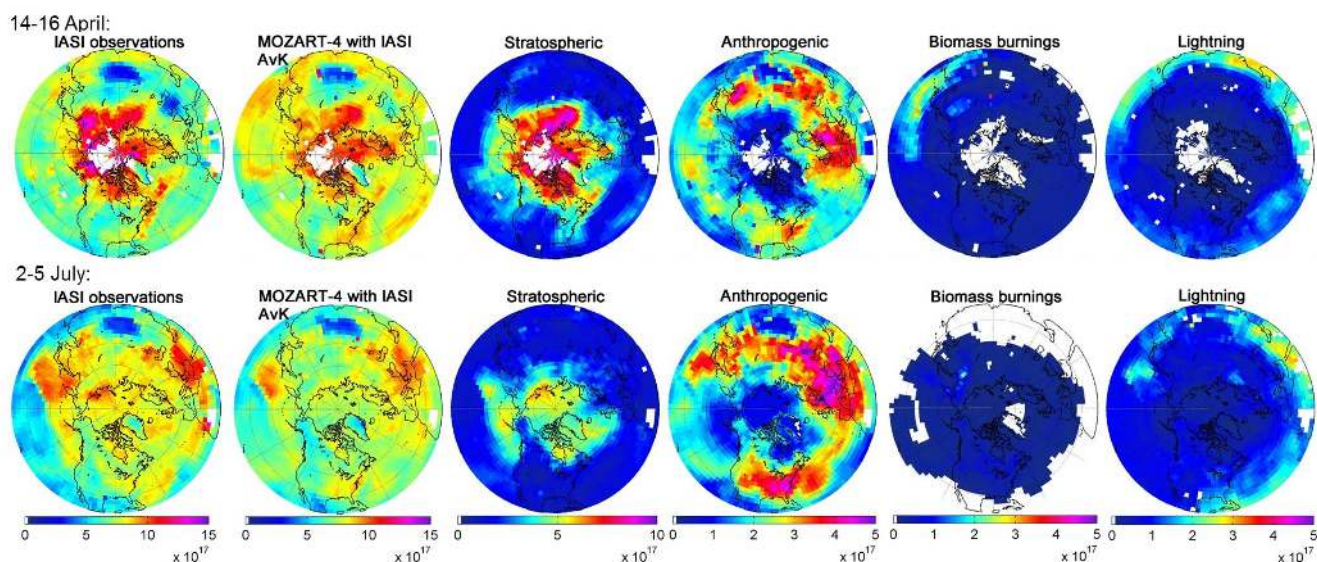


Fig. 14. Mean O_3 columns (molecules cm^{-2}) from ground–300 hPa, over 14–16 April and 2–5 July periods observed by the IASI satellite instrument and simulated by MOZART-4 smoothed according to the averaging kernels of IASI observations. Contributions of principal sources to O_3 tropospheric columns as modeled in MOZART-4 and smoothed with the IASI averaging kernels are also shown.

Europe to Asia through a westerly trans-Siberian transport pathway, from eastern Asia across the Pacific and from eastern USA across the Atlantic are clearly shown in the distributions. This example illustrates the ability of IASI to observe anthropogenic outflow transported over a long distance and reaching the Arctic border. However, evidence of enhanced ozone from a mix of anthropogenic pollution is not clearly detected from the anthropogenic contribution smoothed using the IASI averaging kernels. This results from a sharp decrease of the IASI vertical sensitivity in the middle/lower troposphere above these cold ice-covered surfaces even during the summer time. This lack of vertical sensitivity over the Arctic leads to a collar region of enhanced anthropogenic ozone circling the colder Arctic, in contrast to the non-smoothed MOZART-4 anthropogenic tag (see Sect. 4.2, Fig. 9). This collar region is characterized by DOFS larger than 3.2 while colder Arctic surfaces show lower DOFS. This threshold DOFS value of ~ 3.2 indicates the minimum sensitivity required for the detection of anthropogenic ozone into the Arctic. Principally during April, evidence for biomass burning emissions from eastern Asia (reaching concentration of 3×10^{17} molecules cm^{-2} , $\sim 30\%$ of tropospheric columns) with some transport to higher latitudes are also detected in the IASI measurements. Finally, some contributions from lightning to tropospheric O_3 as observed by IASI are also detected in higher latitudes principally during the summer with values of $\sim 1 \times 10^{17}$ molecules cm^{-2} (reaching 15% of the tropospheric columns).

6 Conclusions

O_3 and HNO_3 observations from the NASA ARCTAS and NOAA ARCPAC aircraft campaigns provide information of the sources of pollution into the Arctic during spring and summer 2008.

The MOZART-4 model generally reproduces the variability of observations in space and time for both O_3 and HNO_3 , especially during ARCTAS. During ARCPAC, the long-distance displacement and smearing of plumes from sources into the Alaskan Arctic seems to introduce larger model error. The model underestimates both O_3 and HNO_3 concentrations by 5–15% and 3–35%, respectively. For the ARCTAS measurements, the underestimation of O_3 and HNO_3 concentrations is attributed to an underrepresentation of both anthropogenic sources and stratospheric influence. During ARCPAC, the underestimation is tied to an underrepresentation of anthropogenic sources and Russian fires emissions. The O_3/HNO_3 ratio characterizing different air mass types (fresh plumes, free troposphere and stratosphere-contaminated air masses) are well reproduced in the model. The largest differences between observations and model are observed in the freshly polluted plumes of the ARCTAS-CARB and the ARCTAS-B campaigns, pointing to the probable impact of the coarse model resolution and the associated diffusion on O_3 production in polluted regions. In summary, the underestimation of concentrations in the model could result from a mix of uncertainties in emission inventories, underestimation of the stratospheric influence, transport errors in the model introduced by too much diffusion due to the coarse resolution of the model and interpolation

of the MOZART-4 results (from 6-h averages) to the flight path (1-min averages). In the spring Arctic boundary layer, the misrepresentation of extremely low values of O_3 is attributed to the missing treatment of the halogen chemistry in MOZART-4.

Over the scale of the Arctic, we find that anthropogenic pollution, especially from Europe, dominates concentrations over the Arctic below 400 hPa (up to 45 % and 60 % of total abundance of O_3 and HNO_3 respectively). This principally reflects lifting to the middle troposphere with northward and westerly trans-Siberian transport pathways. The stratospheric influence dominates concentrations above 400 hPa in the Arctic (~60 %). The Asian emissions and the much smaller North American emissions contributes weakly up to 20 % of the Arctic background between the surface and 500 hPa. Important contributions from Siberian fires reaching the American sector of the Arctic and from Californian fires are also clearly observed. We also found that the principal differences between the distributions of O_3 and HNO_3 lie in a deeper horizontal and vertical gradient for HNO_3 concentrations, because of its shorter lifetime and consequently weaker mixing of HNO_3 than of O_3 throughout the troposphere.

Finally, through comparisons with MOZART-4 total and tagged tracers results, we demonstrate that ground-based FTIR and IASI satellite instruments capture the temporal variations of Northern Hemisphere pollution during the ARCTAS campaigns, with events in the free troposphere, but the satellite observations are less sensitive in the lowermost layers. We show that the O_3 observations from FTIR instruments at Thule and Eureka are dominated by stratospheric and anthropogenic influences (20 % and 55 %, respectively) while the observations from the IASI satellite principally reflect the stratospheric contribution (~50 %) due to the weaker IASI sensitivity in the lowest layers. As a result, we also find that the IASI measurements are characterized by a stronger contribution from the a priori information (~20 % for IASI vs. ~1 % for FTIR) and a smaller contribution from anthropogenic pollution and fires (~15 % for IASI vs. ~55 % for FTIR and <5 % for IASI vs. 10 % FTIR, respectively) than FTIR measurements.

For the extratropical Northern Hemisphere, IASI measurements capture well the mean spatial structure of tropospheric ozone during the spring and the summer 2008, with a good correlation and a good mean relative difference between observations and model. The relative difference of 20 % between IASI observations and model results over all the Arctic further illustrates the underestimation of the stratospheric fluxes in the model. We also demonstrated, for the first time, the capability of the IASI satellite to detect O_3 polluted plumes from fires or anthropogenic emissions during spring and summer 2008, over long distances to the Arctic, but not over the colder surface of the Arctic because of a lack of sensitivity above snow and ice.

Transport of O_3 into high latitudes can strongly impact the radiative forcing of climate and the oxidizing capacity of the troposphere. We demonstrate in this study the capability of combining global model and remote sensing instruments for better observing and quantifying the long-range transport of pollution into the Arctic. It opens perspectives for further studies in exploring the trend in tropospheric ozone pollution and its impact on chemistry and climate. Furthermore, work is in progress to investigate the relevant information for tropospheric HNO_3 from the IASI measurements which would be useful in improving our understanding of the HO_x - NO_y - O_3 system from a more global perspective.

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