



On the computation of planetary boundary-layer height using the bulk Richardson number method

Y. Zhang¹, Z. Gao², D. Li³, Y. Li¹, N. Zhang⁴, X. Zhao¹, and J. Chen^{1,5}

¹International Center for Ecology, Meteorology & Environment, Jiangsu Key Laboratory of Agricultural Meteorology, College of Applied Meteorology, Nanjing University of Information Science and Technology, Nanjing 210044, China

²State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry (LAPC), Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China

³Program of Atmospheric and Oceanic Sciences, Princeton University, Princeton, NJ 08540, USA

⁴School of Atmospheric Sciences, Nanjing University, Nanjing, 210093, China

⁵Department of Geography, Michigan State University, East Lansing, MI 48824, USA

Correspondence to: Z. Gao (zgao@mail.iap.ac.cn)

Received: 15 March 2014 – Published in Geosci. Model Dev. Discuss.: 24 June 2014

Revised: 12 September 2014 – Accepted: 22 September 2014 – Published: 10 November 2014

Abstract. Experimental data from four field campaigns are used to explore the variability of the bulk Richardson number of the entire planetary boundary layer (PBL), Ri_{bc} , which is a key parameter for calculating the PBL height (PBLH) in numerical weather and climate models with the bulk Richardson number method. First, the PBLHs of three different thermally stratified boundary layers (i.e., strongly stable boundary layers, weakly stable boundary layers, and unstable boundary layers) from the four field campaigns are determined using the turbulence method, the potential temperature gradient method, the low-level jet method, and the modified parcel method. Then for each type of boundary layer, an optimal Ri_{bc} is obtained through linear fitting and statistical error minimization methods so that the bulk Richardson method with this optimal Ri_{bc} yields similar estimates of PBLHs as the methods mentioned above. We find that the optimal Ri_{bc} increases as the PBL becomes more unstable: 0.24 for strongly stable boundary layers, 0.31 for weakly stable boundary layers, and 0.39 for unstable boundary layers. Compared with previous schemes that use a single value of Ri_{bc} in calculating the PBLH for all types of boundary layers, the new values of Ri_{bc} proposed by this study yield more accurate estimates of PBLHs.

1 Introduction

The planetary boundary layer (PBL), or the atmospheric boundary layer, is the lowest part of the atmosphere that is directly influenced by earth's surface and has significant impacts on weather, climate, and the hydrologic cycle (Stull, 1988; Garratt, 1992; Seidel et al., 2010). The height of the PBL (PBLH) is typically on the order of 1 ~ 2 km but varies significantly during a diurnal cycle in response to changes in the thermal stratification of the PBL. It is an important parameter that is commonly used in modeling turbulent mixing, atmospheric dispersion, convective transport, and cloud/aerosol entrainment (Deardorff, 1972; Holtslag and Nieuwstadt, 1986; Sugiyama and Nasstrom, 1999; Seibert et al., 2000; Medeiros et al., 2005; Konor et al., 2009; Liu and Liang, 2010; Leventidou et al., 2013). As a result, accurate estimates of the PBLH under different thermal stratifications are critically needed.

The PBL is characterized by the presence of continuous turbulence, while turbulence is lacking or sporadic above the PBL. Therefore, the PBLH can be viewed as the level where continuous turbulence stops (Wang et al., 1999; Seibert et al., 2000). Using high-frequency turbulence measurements (e.g., collected from ultrasonic anemometers on aircraft), the PBLH can be readily determined. This is known as the turbulence (Tur) method. It is highly reliable, but the instruments required by this method are costly. A more economic option is to determine the PBLH through analyzing temperature and

wind profiles measured from radio soundings. In this method, the PBLs are broadly classified as strongly stable boundary layers (type I SBLs), weakly stable boundary layers (type II SBLs), or unstable boundary layers (UBLs) (Holtslag and Boville, 1993; Vogelesang and Holtslag, 1996). They are defined using the surface heat flux and the potential temperature profile, as shall be seen later.

For strongly stable boundary layers or type I SBLs, there is a strong inversion in the potential temperature profile and the PBLH is usually defined as the top of the inversion where the potential temperature gradient (PTG) first becomes smaller than a certain threshold γ_s (Bradley et al., 1993), which is chosen to be $6.5 \text{ K } 100 \text{ m}^{-1}$ following Dai et al. (2011). This is called the PTG method hereafter. For weakly stable boundary layers or type II SBLs, turbulence is generated from wind shear due to relatively high wind speed and the PBLH is defined as the height of the low-level jet (LLJ) (Melgarejo and Deardorff, 1974). This is called the LLJ method hereafter. For unstable boundary layers, UBLs, buoyancy is the dominant mechanism driving turbulence, and the PBLH is defined as the height at which a thin layer of capping inversion occurs. The PBLH of UBLs is determined first by identifying a height at which a parcel of dry air, released adiabatically from the surface, reaches equilibrium with its environment (Holzworth, 1964). This height is then corrected by another upward search for another height at which the potential temperature gradient first exceeds a threshold γ_c (Liu and Liang, 2010), which is chosen to be $0.5 \text{ K } 100 \text{ m}^{-1}$ in this study. This is called the modified parcel method hereafter.

For an atmosphere with discernible characteristics (i.e., a strongly stable potential temperature profile for the type I SBL, a strong LLJ for the type II SBL, and a capping inversion layer for the UBL), the three methods generally show good performance (e.g., Mahrt et al., 1979; Liu and Liang, 2010; Dai et al., 2011). However, for an atmosphere without these discernible characteristics, large errors can be introduced by these methods. As such, these methods are usually used in experimental studies but not in numerical models since numerical models need to determine the PBLH automatically. Instead, the bulk Richardson number (Ri_b) method is often used for numerical weather and climate models due to its reliability under a variety of atmospheric conditions (e.g., Holtslag and Boville, 1993; Jericevic and Grisogono, 2006; Richardson et al., 2013). The bulk Richardson number method assumes that the PBLH is the height at which the Ri_b reaches a threshold value (Ri_{bc} , which is called “the bulk Richardson number for the entire PBL” hereafter). The Ri_b at a certain height z is calculated with the potential temperature and wind speed at this level and those at the lower boundary of the PBL (generally the surface), as follows (Hanna, 1969):

$$Ri_b = \frac{(g/\theta_{v0})(\theta_{vz} - \theta_{v0})z}{u_z^2 + v_z^2}, \quad (1)$$

where θ_{v0} and θ_{vz} are the virtual potential temperatures at the surface and at height z , respectively, g/θ_{v0} is the buoyancy parameter, and u_z and v_z are the horizontal wind-speed components at height z . As can be seen from Eq. (1), the bulk Richardson number method is computationally cheap because it only requires low-frequency data. Nonetheless, the biggest challenge associated with the bulk Richardson number method is that the value of Ri_{bc} has to be determined as a prior known. In previous studies, the value of Ri_{bc} varies from 0.15 to 1.0 (Zilitinkevich and Baklanov, 2002; Jericevic and Grisogono, 2006; Esau and Zilitinkevich, 2010), with values of 0.25 and 0.5 most widely used (e.g., Troen and Mahrt, 1986; Holtslag and Boville, 1993). One important cause of the large variability of Ri_{bc} is the thermal stratification in the PBL. For example, Vogelesang and Holtslag (1996) reported the Ri_{bc} values of 0.16–0.22 in a nocturnal strongly stable PBL and 0.23–0.32 in a weakly stable PBL. For unstable PBLs, a value larger than 0.25 is usually needed (Zhang et al., 2011). Esau and Zilitinkevich (2010) also showed that the Ri_{bc} for nocturnal SBLs was smaller than for neutral and long-lived stable PBLs based on a large-eddy-simulation database. More recently, a linear relationship between the Ri_{bc} and the atmospheric stability parameter has been proposed and examined under stable conditions, which further suggests the impact of thermal stratification on the Ri_{bc} (Richardson et al., 2013; Basu et al., 2014).

The objective of this study is to examine the variation of Ri_{bc} with different thermal stratification conditions. To do so, a representative value of Ri_{bc} for each type of PBLs (i.e., strongly stable boundary layers, weakly stable boundary layers, and unstable boundary layers) needs to be inferred. In our study, the Tur method, the PTG method, the LLJ method, and the modified parcel method are used to determine the PBLHs from observations made in four field campaigns, which are called “observed” PBLHs. Using these “observed” PBLHs as benchmarks, the best choices of Ri_{bc} values under different stratification conditions are then inferred so that the estimates of PBLHs with the bulk Richardson number method match the “observed” PBLHs. These inferred values of Ri_{bc} are used to explore the impact of thermal stratification on the Ri_{bc} .

The study is organized in the following way: Sect. 2 describes the observational data used in this study; Sect. 3 compares estimates of PBLH from different methods that are widely used to determine the PBLH from measurements; Sect. 4 focuses on the bulk Richardson number method and describes the search for a best choice of Ri_{bc} under different stratification conditions. Section 5 concludes the paper.

2 Observational data

Observational data from four field campaigns that were conducted under different surface and atmospheric conditions are used in this study. These field campaigns are

the Litang experiment, the Atmospheric Radiation Measurement (ARM) experiment, the Surface Heat Budget of the Arctic Ocean (SHEBA) experiment, and the Cooperative Atmosphere–Surface Exchange Study (CASES) in 1999 (CASES99). Each of these four field campaigns is briefly described as follows.

The Litang site is located over a plateau meadow in the southeast of the Tibetan Plateau. The campaign provides 105 effective radio soundings of wind and temperature in three observational periods (7–16 March, 13–22 May, and 7–16 July, 2008), with a typical 6 h interval (about 00:30, 06:30, 12:30, and 18:30 LST, local standard time). The 30 min averaged wind and temperature at 3 m collected by an eddy covariance system are also used for calculating the bulk Richardson number.

The ARM experiment was carried out over a plain farmland in Shouxian, China, from 14 May to 28 December 2008. During the campaign, soundings were collected every 6 h (about 01:30, 07:30, 13:30, and 19:30 LST). Due to instrument malfunction, some data are excluded and a total of 842 radio soundings are retained. The 30 min averaged wind and temperature measured at 4 m by an eddy covariance system are also used.

The SHEBA site is located around the Canadian icebreaker *Dec Groseilliers* in the Arctic Ocean. The data set provides radio soundings from mid-October 1997 to early October 1998. During this period, rawinsondes were released 2–4 times a day (around 05:15, 11:15, 17:15, and 23:15 LST). Since the near-surface (2.5 m) data available from 29 October 1997 to 1 October 1998 at the SHEBA are hourly averages (Andreas et al., 1999; Persson et al., 2002), the surface observations and soundings do not overlap well in time. To ensure accuracy, only soundings released within 15 min around the hour were used in this study, yielding a total of 168 records.

The CASES99 is the second experiment of CASES conducted in Kansas, USA. The terrain is relatively flat (the average slope is about 0.5°). In the campaign, the National Oceanic and Atmospheric Administration (NOAA) Long-EZ and Wyoming King Air accomplished the aircraft measurements at 50 and 25 Hz sample rates, respectively, during the period 6–27 October 1999 when the PBL was primarily stable. Since the lowest flight level was restricted (e.g., for security reasons), only 35 effective aircraft soundings are used in our study. The 5 min averaged near-surface (3 m) wind and temperature data recorded at the no. 16 flux tower in CASES99 (www.eol.ucar.edu/projects/cases99) are also used. The surface observations and soundings in CASES99 overlap well in time, but their horizontal positions slightly differ due to the movement of aircraft. Due to the fact that most of the sounding data from CASES99 were collected under strongly stable conditions and data under other conditions were too limited, in this study only soundings under strongly stable conditions (i.e., in type I SBLs) are used; except in Fig. 1 where one weakly stable boundary layer case from

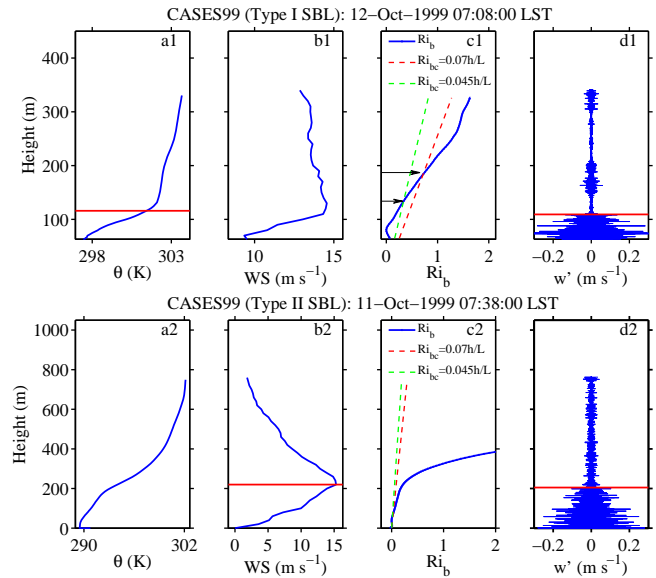


Figure 1. Examples of vertical profiles of the type I SBL (upper panels) and the type II SBL (lower panels) from CASES99 aircraft measurements: (a) potential temperature (K); (b) horizontal wind speed (m s^{-1}); (c) bulk Richardson number Ri_b and Ri_{bc} ; (d) w' perturbation (m s^{-1}). The red solid lines on (a1) and (b2) denote the PBLH calculated by the PTG and LLJ methods, respectively, and those on (d) denote the PBLH determined by the Tur method. The black arrows on (c1) denote the PBLHs determined by the bulk Richardson number method with Ri_{bc} from Eq. (7).

CASES99 is presented in order to compare the LLJ method to the Tur method.

In postprocessing, a 20 m moving-window average is used for all the soundings from all the sites (except the turbulence measurements by aircraft in CASES99) to remove the measurement noise.

3 PBLHs determined from observational data

As mentioned in the introduction, the PBLs during a typical diurnal cycle are categorized into three types: type I SBLs (i.e., strongly stable boundary layers at night), type II SBLs (i.e., weakly stable boundary layers at early morning/night), and UBLs (i.e., unstable boundary layers during the daytime). The PTG method, the LLJ method, and the modified parcel method are usually used to determine the PBLH for type I SBLs, type II SBLs, and UBLs, respectively. Based on previous studies (e.g., Holtslag and Boville, 1993; Vogelezang and Holtslag, 1996), they are classified using the surface heat flux H and the potential temperature θ profile:

$$\begin{cases} H \geq \delta & \text{for UBLs,} \\ H < \delta \text{ and } d^2\theta/dz^2 < 0 & \text{for type I SBLs,} \\ H < \delta \text{ and } d^2\theta/dz^2 \geq 0 & \text{for type II SBLs,} \end{cases} \quad (2)$$

where δ is the minimum H for unstable conditions, which, in practice, is specified as a small positive value instead of zero (Liu and Liang, 2010). Due to different thermodynamic properties of land and ice, the value of δ is specified as 1 W m^{-2} over land and 0.5 W m^{-2} over ice through trial and error. Under stable conditions (i.e., $H < \delta$), the PBLs are further classified into type I SBLs and type II SBLs according to $d^2\theta/dz^2$. For type I SBLs, the PTG decreases with height and the inversion near the surface is relatively strong, so there is always a sudden decrease of PTG at the PBL top (e.g., see Fig. 1a1). As such, the derivative of PTG with respect to z should be negative; that is, $d^2\theta/dz^2 < 0$. For type II SBLs, the PTG increases with height and the inversion is relatively weak. No sudden change of PTG is seen at the PBL top (e.g., see Fig. 1a2) and thus $d^2\theta/dz^2 \geq 0$. In this study, $d^2\theta/dz^2$ is calculated between 40 and 200 m; the selection of 40 m as the lower boundary is to avoid near-surface variability caused by landscape heterogeneity.

Note that cases with $-\delta < H < \delta$ (i.e., under near-neutral conditions) are typically treated as type II SBL cases according to our classification. This is because stable stratification usually prevails above the boundary layer and wind shear is the only source of turbulence under near-neutral conditions. Both of these features are similar to those of a stable boundary layer and, as a result, the near-neutral cases are treated as SBL cases (Serbert et al., 2000). It appears there might be an abrupt change in the calculation of PBLH at $H \approx \delta$ if different values of Ri_{bc} are used for SBLs and UBLs, which is the aim of this study. However, we note that changes of Ri_{bc} at $H \approx \delta$ from SBLs to UBLs have little effect on the PBL height determination, because the Ri_b increases drastically with height at the PBL top under near-neutral conditions and using Ri_{bc} for either SBLs or UBLs gives reasonable estimates of PBLH (Supplementary Fig. S1).

For any of the three types of PBLs, the Tur method is the most direct and accurate approach for the PBLH estimation because it measures the turbulence intensity directly. Figure 1 shows vertical profiles of potential temperature, mean wind velocity, bulk Richardson number, and wind velocity perturbations from CASES99 for a type I SBL (a1–d1) and a type II SBL (a2–d2). The wind velocity perturbations (u' , v' , w'), or turbulence intensities, are obtained by removing the slowly varying part of the corresponding winds (u , v , w) through a high-pass wavelet filter (Wang et al., 1999; Wang and Wang, 2004). In the Tur method, continuous wavelet transform is applied to the absolute magnitude of turbulent fluctuations of each velocity component. The PBLH is automatically determined to be the level at which the absolute magnitude of these velocity fluctuations shows the most rapid decrease with height (Dai et al., 2011, 2014). The PBLHs determined by u' , v' , w' are then averaged using the absolute magnitude of the reciprocal velocity fluctuations as weights. As can be seen in Fig. 1d1 and 1d2, the PBLHs determined by the Tur method are denoted by the red solid lines.

Figure 1 further shows the PBLHs determined with the PTG method (see the red solid line on a1) and the LLJ method (see the red solid line on b2) for type I and type II SBLs, respectively. It is clear that the estimates of PBLHs with these two methods are comparable to the PBLHs determined from the Tur method, suggesting that the PTG method and the LLJ method work well for type I and type II SBLs, respectively.

Figure 2 shows the sounding profiles taken from Litang on 9 July 2008 and the PBLHs estimated by the PTG, LLJ, and modified parcel methods for the three different PBLs, respectively. At midnight (00:35 LST), the PBL was very stable due to radiative cooling from the surface and is classified as a type I SBL. According to the PTG method, the PBLH was found at the top of the strong inversion (125 m; see Fig. 2a). In the early morning (06:35 LST), the surface temperature increased and thus the inversion near the surface became weak; the low-level wind speed increased rapidly and formed a LLJ. The PBL is classified as a type II SBL. With the LLJ method, the PBLH was determined at the height of the maximum wind (260 m; see Fig. 2b). As the surface heating continued, a super-adiabatic layer in which the potential temperature decreased with height formed near the surface and a UBL was developed by midday (12:45 LST). With the modified parcel method, the PBLH is estimated to be 1654 m (see Fig. 2c). Consequently, it can be concluded that the three methods mentioned above are useful for a PBL with discernible characteristics (Figs. 1, 2).

However, for a PBL without these discernible characteristics, these methods may introduce large biases (see Fig. 3 and e.g., Russell et al., 1974; Martin et al., 1988; Balsley et al., 2006; Meillier et al., 2008). For type I SBLs, when the underlying inversion is not strong, it will be difficult to determine the PBLH by the PTG method due to the fact that the maximum PTG can be less than the threshold γ_s (Fig. 3b1). For type II SBLs, when there is no clear wind-speed maximum or when multiple maxima exist, the LLJ method will have difficulties in determining the PBLH. For example, there were two maxima in the wind profile (at 160 and 400 m, see Fig. 3c2). If the PBLH is simply determined as the height where the first maximum occurs, the PBLH would be 160 m. Combining information from the Ri_b profile (Fig. 3d2), a more reasonable estimate of the PBLH should be 400 m instead of 160 m since the Ri_b profile undergoes a significant transition at 400 m. For UBLs similar complex situations may occur. The results of the modified parcel method with a specified PTG threshold may be subjective since the threshold may depend on the vertical resolution and data precision (Beyrich, 1997; Joffre et al., 2001). For example, there are two PTG maxima at 900 and 2000 m (see Fig. 3b3) due to the sharp drop of relative humidity at these two heights. A more accurate estimate of the PBLH should be 900 m when combining the information from the Ri_b profile (Fig. 3d3), while it might have been determined to be 2000 m by the modified parcel method if $\gamma_c = 0.5 \text{ K (100 m)}^{-1}$ was used.

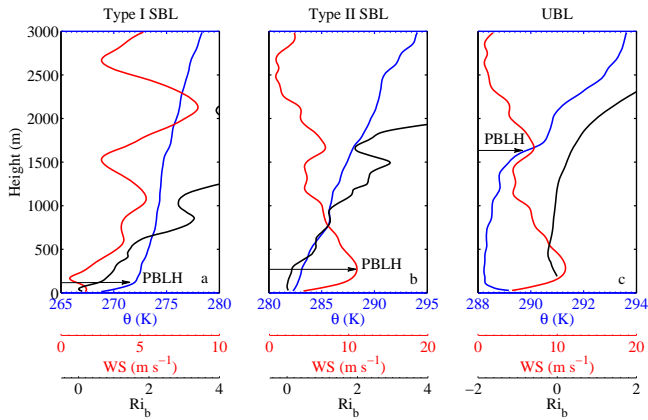


Figure 2. Typical profiles of potential temperature (blue), wind speed (red), and Ri_b (black) for different types of boundary layers: (a) type I SBL, (b) type II SBL, and (c) UBL. The indicated PBLHs in (a–c) are calculated by the PTG, LLJ, and modified parcel methods, respectively. The observations in (a–c) are from Litang on 8 July 2008 16:35 UTC (00:35 LST), 8 July 2008 22:45 UTC (06:45 LST), and 9 July 2008 04:45 UTC (12:45 LST), respectively.

Although these special cases do not always exist, they limit the applications of the three methods. The accuracy of the determined PBLH can be improved with additional information, as have been demonstrated before. The following sums up the procedures that are used in this study for estimating PBLH by using these four methods. First, whenever turbulence measurements are available, the Tur method is used to determine the PBLH. Second, for type I SBL cases with a relatively weak inversion (the local PTG maximum is $< 6.5 \text{ K} (100 \text{ m})^{-1}$ between 40 and 200 m), if there is a LLJ the case is reclassified to a type II SBL, if not the case is removed. Third, type II SBL cases without clear wind-speed maximum are removed. Fourth, when there are multiple wind maxima for a type II SBL or multiple PTG maxima for a UBL, the information from the Ri_b profile is combined to determine the PBLH. With these procedures, the PBLHs obtained by using these methods are treated as “observed” PBLH hereafter. The observed PBLH and the bulk stability parameter (PBLH/ L , where L is the surface Obukhov length) for these four field experiments are provided in Table 1.

4 The bulk Richardson number method and the Ri_{bc}

The PTG method, the LLJ method, and the modified parcel method are usually used to determine the PBLH in observational data. However, they do not work well when the PBL has no distinct features that are required by these methods. Instead, the bulk Richardson number method with a prescribed Ri_{bc} is often used in numerical methods to automatically determine the PBLH. For example, in the non-local PBL scheme of the Community Climate Model ver-

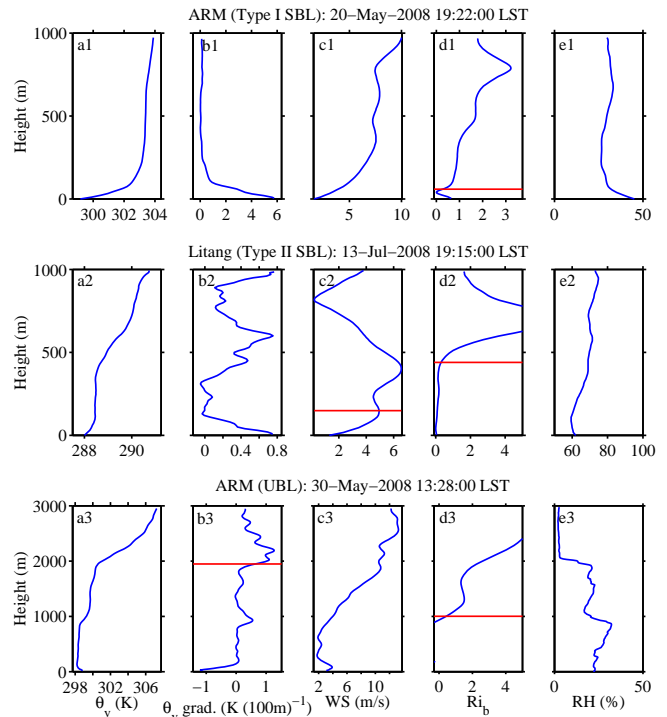


Figure 3. Examples of vertical profiles in type I SBLs (upper panels), type II SBLs (middle panels), and UBLs (lower panels): (a) potential temperature (K), (b) potential temperature gradient ($\text{K } 100 \text{ m}^{-1}$), (c) horizontal wind speed (m s^{-1}), (d) bulk Richardson number Ri_b , and (e) relative humidity (%). The red solid lines on (b3), (c2), and (d1–d3) denote the PBLH determined by the modified parcel, LLJ, and bulk Richardson number methods, respectively.

sion 2 (CCM2), Eq. (1) is applied to estimate the PBLH with $Ri_{bc} = 0.5$. The computation starts by calculating the Ri_b between the surface and subsequent higher levels of the model. Once Ri_b exceeds Ri_{bc} , the PBLH is derived by a linear interpolation between the level with $Ri_b > Ri_{bc}$ and the level below.

To avoid overestimating the shear production in Eq. (1) for relatively high wind speeds (i.e., in type II SBL) and to account for turbulence generated by surface friction under neutral conditions, Vogelesang and Holtslag (1996) proposed an updated formulation, which is employed in the Community Atmosphere Model version 4 (CAM4), written as

$$Ri_b = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - v_s)^2 + 100u_*^2}, \quad (3)$$

where z_s is the height of the lower boundary for the PBL (generally the top of the atmospheric surface layer), θ_{vs} is the virtual potential temperature at the height z_s , u_s and v_s are the wind-speed components at z_s . z_s is often taken as 20, 40, or 80 m for SBLs (Vogelesang and Holtslag, 1996) and taken as 0.1 PBLH (Troen and Mahrt, 1986) or the height of super-adiabatic layer (z_{SAL}) where the potential temperature first

Table 1. The “observed” PBLH and the stability parameter at four observational sites.

| PBL types | PBLH (m) | | | | PBLH / L | | | |
|-------------------------|------------|----------|------------|-----------|---------------|------------|---------------|---------------|
| | Litang | CASES99 | ARM | SHEBA | Litang | CASES99 | ARM | SHEBA |
| Type I SBL | 45 ~ 265 | 25 ~ 157 | 54 ~ 593 | 42 ~ 414 | 0.12 ~ 323.0 | 1.5 ~ 94.2 | 0.22 ~ 327.2 | 0.4 ~ 38.3 |
| Type II SBL ($H < 0$) | 68 ~ 543 | \ | 131 ~ 670 | 97 ~ 312 | 0.64 ~ 74.8 | \ | 0.36 ~ 113.1 | 0.1 ~ 21.3 |
| Type II SBL ($H > 0$) | 357 ~ 678 | \ | 152 ~ 879 | 138 ~ 414 | -33.4 ~ -0.32 | \ | -34.1 ~ -0.2 | -55.1 ~ -0.01 |
| UBL | 315 ~ 2594 | \ | 293 ~ 1693 | 121 ~ 981 | -866.4 ~ -4.3 | \ | -350.9 ~ -1.3 | -342 ~ -0.03 |

reaches its local minimum for UBLs. In our study, $z_s = 40$ or 80 m is used under stable conditions while $z_s = 0.1$ PBLH or z_{SAL} is used under unstable conditions. The term $100u_*^2$ makes Eq. (3) more applicable for the near-neutral condition, which is classified as a type II SBL in our study (Seibert et al., 2000).

Under unstable conditions, the virtual potential temperature at the lower boundary θ_{vs} is replaced by θ'_{vs} (Troen and Mahrt, 1986; Holtslag et al., 1995):

$$\theta'_{vs} = \theta_{vs} + b_s \frac{\overline{(w'\theta'_v)}_0}{w_m}, \quad (4)$$

where $b_s = 8.5$, $\overline{(w'\theta'_v)}_0$ is the virtual heat flux at the surface, and w_m is a turbulent velocity scale:

$$w_m = \left(u_*^3 + 0.6w_*^3\right)^{1/3}, \quad (5)$$

and

$$w_* = \left[\frac{g}{\theta_{v0}} \overline{(w'\theta'_v)}_0 h\right]^{1/3} \quad (6)$$

is the convective velocity scale. The second term on the right-hand side of Eq. (4) represents a temperature excess, which is a measure of the strength of convective thermals.

In this study, the virtual potential temperature is estimated as the potential temperature in the calculation because the former can lead to significant fluctuations in the estimated PBLH due to inaccurate humidity measurements (Liu and Liang, 2010).

After Ri_b is computed from Eqs. (3)–(6), the PBLH can be determined as the height where the Ri_b exceeds Ri_{bc} . In our study, instead of calculating the PBLH using a prescribed Ri_{bc} , we infer a representative Ri_{bc} for each type of PBLs using the “observed” PBLH (see Sect. 3) and examine the variation of the inferred Ri_{bc} with thermal stratification. It is pointed out here that our methodology is different from that of Richardson et al. (2013), who proposed a stability dependent Ri_{bc} for SBLs:

$$Ri_{bc} = \alpha \frac{\text{PBLH}}{L}, \quad (7)$$

where PBLH / L is a bulk stability parameter and L is the surface Obukhov length. α is a proportionality constant, which

depends on surface characteristics and/or atmospheric conditions. It varies between 0.03 and 0.21 with suggested values of 0.045 and 0.07 (Richardson et al., 2013; Basu et al., 2014). As shown in Fig. 1c1 and c2, in the type I SBL case, a relatively reliable PBLH (133 m) was calculated with $\alpha = 0.045$, but an overestimation (184 m) occurs when $\alpha = 0.07$. While in the type II SBL case both α values (0.045 and 0.07) yield too small estimates of PBLH, because the two values are determined by idealized stable large-eddy-simulation data sets (Richardson et al., 2013) and observational data sets under weakly and moderately stable conditions (Basu et al., 2014), respectively. In addition, Eq. (7) is only applicable for SBLs but not UBLs. As such, instead of adopting this equation, we inferred a representative Ri_{bc} value for each type of PBL in our study.

Because each profile provides a Ri_{bc} value, a representative Ri_{bc} at each experimental site is determined by fitting a linear relationship between the numerator and the denominator of Eq. (3) at the PBLH, as will be shown in Sect. 4.1, or using statistical error minimization methods, as will be shown in Sect. 4.2.

4.1 Representative Ri_{bc} from the linear fitting method

The representative Ri_{bc} values for type I SBLs are shown in Fig. 4. The soundings are taken from Litang, CASES99, ARM, and SHEBA, with the height z_s of 40 (left) and 80 m (right). Note that with $z_s = 80$ m, only cases with a PBLH ≥ 80 m are used. Except for CASES99, the fitted Ri_{bc} values at each site are about 0.25. The difference in Ri_{bc} when different values for z_s (40 or 80 m) are used is small. However, the results are slightly more consistent with $z_s = 40$ m compared to $z_s = 80$ m, as can be seen from the higher correlation coefficients at ARM and CASES99. The value of Ri_{bc} for type I SBLs from CASES99 aircraft measurements is 0.20–0.21, which is smaller than the values determined from radio soundings at other experimental sites. This may be because the depth of the nocturnal inversion is generally thicker than the depth of the turbulent layer (Mahrt et al., 1979; Andre and Mahrt, 1982). Therefore, the PBLH determined by the Tur method is smaller than that determined by the PTG method. Note the Tur method is always used to determine the PBLH for CASES99 since turbulence measurements are available at this site.

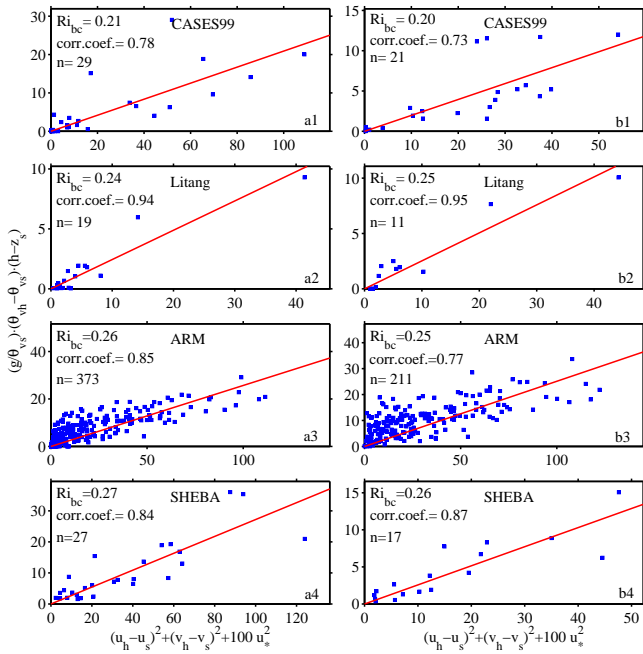


Figure 4. Linear fitting method inferred Ri_{bc} for type I SBLs, with $z_s = 40$ m (left) and $z_s = 80$ m (right). The red solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

The inferred Ri_{bc} values for type II SBLs are shown in Fig. 5. Compared to the results in Fig. 4, the correlation coefficients in Fig. 5 are smaller, indicating that the PBLH is more difficult to determine for weakly stable boundary layers, which is consistent with previous studies (e.g., Esau and Zilitinkevich, 2010). The correlation coefficients indicate that the agreement with $z_s = 80$ m is slightly better than with $z_s = 40$ m. In particular, the inferred Ri_{bc} is sensitive to the height z_s in the SHEBA data. It changes from 0.21 to 0.29 as the height z_s changes from 40 to 80 m. The main cause of the large variation of Ri_{bc} is because the LLJs above the ice surface in SHEBA are considerably strong (up to 20 m s^{-1}) and the vertical wind-speed gradients are large, so the denominator in Eq. (3) decreases more rapidly with the height z_s than the numerator, which leads to an increase in the Ri_{bc} value when z_s increases from 40 to 80 m. However, the Ri_{bc} over land varies little with z_s (Fig. 5), which is consistent with the findings of Vogelesang and Holtslag (1996) using the Cabauw data.

For UBLs, the height z_s is chosen to be 0.1 PBLH (left) and z_{SAL} (right) in Fig. 6. As can be seen, the correlation coefficients are smaller than 0.4 at all sites, implying large variability in the Ri_{bc} inferred from each sounding. The representative value of Ri_{bc} is larger than 0.25 and varies from 0.28 to 0.34. However, it appears that the PBLH estimated by the bulk Richardson number method seems to be less sensitive to Ri_{bc} under unstable conditions. The estimates of PBLH using the bulk Richardson number method with $Ri_{bc} = 0.25$ or 0.5 are both in good agreement with the “observed” PBLH at the

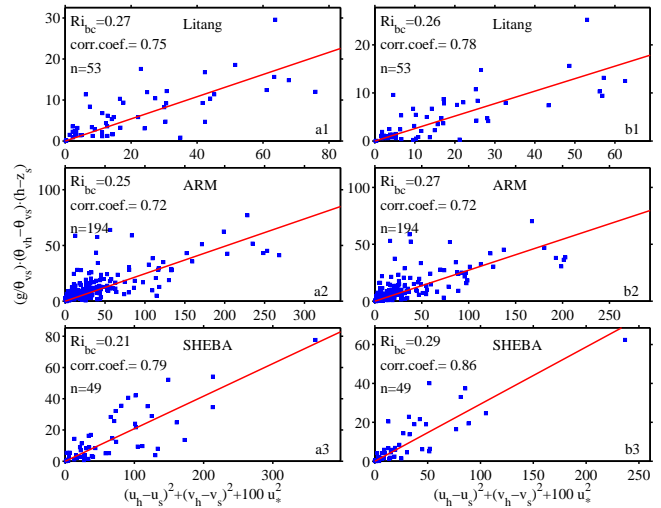


Figure 5. Linear fitting method inferred Ri_{bc} for type II SBLs, with $z_s = 40$ m (left) and $z_s = 80$ m (right). The red solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

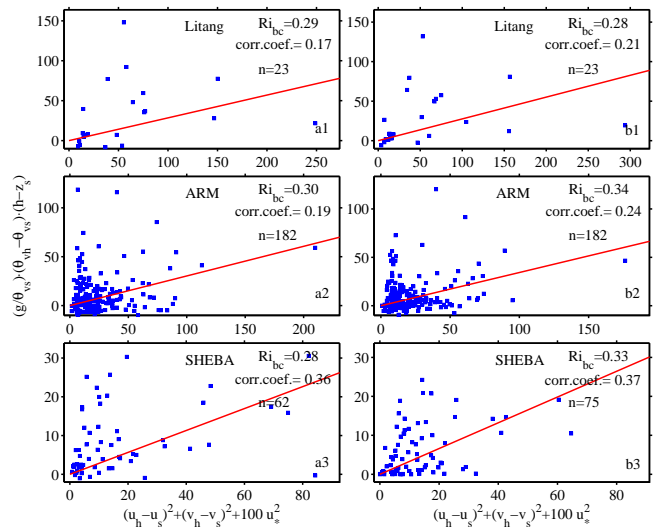


Figure 6. Linear fitting method inferred Ri_{bc} for UBLs, with $z_s = 0.1$ PBLH (left) and $z_s = z_{SAL}$ (right). The red solid lines are the best linear fittings and their slopes represent the values of Ri_{bc} .

three sites (Fig. 7). This is also in agreement with some previous studies (Troen and Mahrt, 1986). Therefore, the bulk Richardson number method is still reliable in estimating the PBLH of UBLs, despite the inferred Ri_{bc} showing large variability.

4.2 Representative Ri_{bc} from the error minimization method

It is seen that the linear fitting method yields small correlation coefficients under unstable conditions. Under stable conditions, the linear fitting method also has some

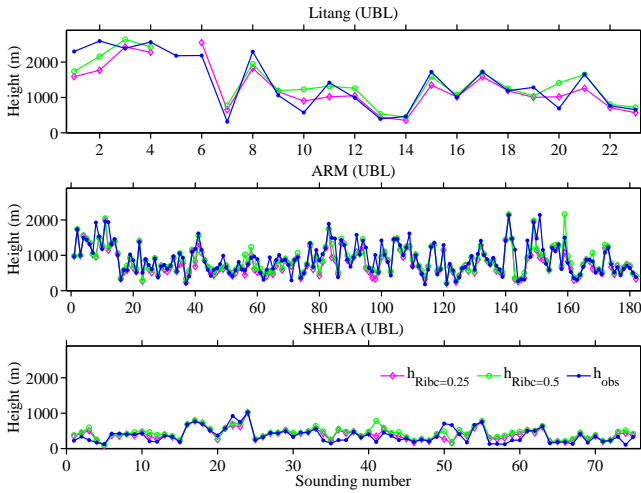


Figure 7. Comparisons of the heights of UBL at different sites determined by the bulk Richardson number method with $Ri_{bc} = 0.25$ (diamond) and 0.5 (circle), and the observed PBLHs (point).

disadvantages. For example, the inferred value of Ri_{bc} and the correlation coefficients highly depend on the larger value points, while the impact of the smaller value points is reduced (see e.g., Fig. 4a2). Therefore, we apply error minimization methods in this section to determine the optimal Ri_{bc} . The values of Ri_{bc} between 0.1–0.4 in stable conditions and 0.2–0.5 in unstable conditions are first used to calculate the PBLH; then, three statistical measures are used to examine the accuracy of the estimated PBLH (Gao et al., 2004):

$$\text{bias} = \frac{\sum_{i=1}^n |h_{Ri_{bc}} - h_{obs}|}{n}, \quad (8)$$

$$\text{SEE} = \sqrt{\frac{\sum_{i=1}^n |h_{Ri_{bc}} - h_{obs}|^2}{n-2}}, \quad (9)$$

$$\text{NSEE} = \sqrt{\frac{\sum_{i=1}^n (h_{Ri_{bc}} - h_{obs})^2}{\sum_{i=1}^n (h_{obs})^2}}, \quad (10)$$

where $h_{Ri_{bc}}$ is the estimated PBLH by the bulk Richardson number method, and h_{obs} represents the observed PBLH (i.e., calculated using the Tur, PTG, LLJ, or modified parcel method). Bias, SEE, and NSEE are the absolute bias, standard error, and normalized standard error of $h_{Ri_{bc}}$ against h_{obs} , respectively, and n is the sampling number. Optimal values of Ri_{bc} can be determined based on the minimum bias, SEE, and NSEE. Note that the optimal Ri_{bc} determined based on the minimum bias, or the minimum SEE/NSEE can be different, however, the minimum SEE and the minimum NSEE always yield the same optimal Ri_{bc} . In this study, minimum SEE and NSEE are used as the final criterion for the optimal

Ri_{bc} . To compare the error minimization method with the linear fitting method, the correlation coefficients between h_{obs} and $h_{Ri_{bc}}$ are also presented.

The correlation coefficient, bias, SEE, and NSEE with different values of Ri_{bc} for type I SBLs are shown in Fig. 8 when $z_s = 40$ (top panels) and 80 m (bottom panels). Quadratic curves are fitted to these data and then the maximum or minimum of the fitted quadratic curves are obtained, which are used to select the optimal Ri_{bc} for each site. The weighted averages based on the sampling number at the four sites are treated as the representative optimal Ri_{bc} across the four sites (see the black dashed lines in Fig. 8) and the error bars depict the range of the optimal Ri_{bc} across the four sites (Fig. 8). The variability of the optimal Ri_{bc} values for different sites is probably caused by the diversity of surface characteristics (e.g., surface roughness). Compared to the results with $z_s = 80$ m, the error bars are smaller and thus the optimal Ri_{bc} across different sites are more concentrated with $z_s = 40$ m. Furthermore, the maximum correlation coefficient is larger and the minimum bias, SEE, and NSEE are smaller with $z_s = 40$ m.

Compared to type I SBLs, the correlation coefficients are smaller and errors are larger for type II SBLs (Fig. 9), again indicating that the PBLH for weakly stable boundary layers is more difficult to determine. However, the maximum correlation coefficient, minimum bias, SEE and NSEE, and the range of optimal Ri_{bc} show smaller differences between different values of z_s (40 or 80 m). Compared to the results of the linear fitting method, the values of Ri_{bc} are generally larger for each site, which is understandable given that the scatter distribution is mostly above the fitted lines in Fig. 5, especially at ARM and SHEBA. The optimal Ri_{bc} based on minimum SEE and NSEE for type II SBLs is 0.30–0.31. The result is consistent with the value (i.e., 0.3) from Melgarejo and Dardorff (1974).

For UBLs, Fig. 10 shows that the maximum correlation coefficient is larger, the minimum bias, SEE, and NSEE are smaller, and the values of optimal Ri_{bc} for each site are more concentrated with $z_s = z_{SAL}$ (bottom panels) compared to $z_s = 0.1$ PBLH (top panels). Therefore, z_{SAL} is more appropriate as the lower boundary height in estimating the PBLH under unstable conditions. The minimum SEE and NSEE indicate that the optimal Ri_{bc} is 0.39 under unstable conditions. The results with $z_s = 40$ or 80 m are also examined but not shown here. The maximum correlation coefficient and minimum bias, SEE, and NSEE are close to those with $z_s = 0.1$ PBLH, but the values of the optimal Ri_{bc} are more scattered across different sites.

Through the above statistical error minimization methods, the optimal Ri_{bc} for different stratifications and sites with different choices of z_s are summarized in Table 2. It appears that the optimal Ri_{bc} value increases when the PBL stability decreases (i.e., as the PBL becomes more unstable). The optimal Ri_{bc} value is 0.24 ($z_s = 40$ m) or 0.23 ($z_s = 80$ m) for type I SBLs, and 0.30 ($z_s = 40$ m) or 0.31 ($z_s = 80$ m)

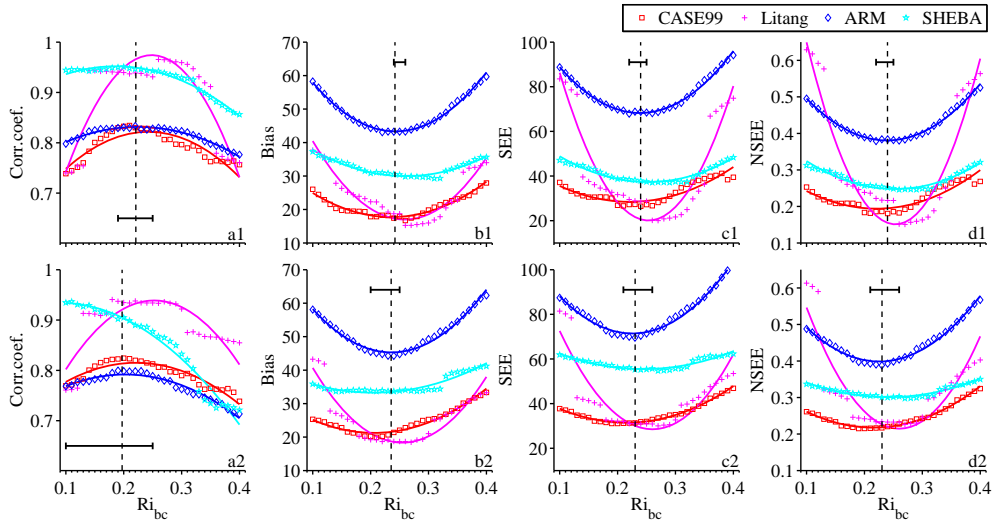


Figure 8. Comparison between estimated PBLH using the bulk Richardson number method with $z_s = 40$ m (upper panels) and $z_s = 80$ m (lower panels) and observed PBLHs for type I SBLs. The correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), CASES99 (square), ARM Shouxian (diamond), and SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all four sites, and the error bars indicate the range of Ri_{bc} across the four sites.

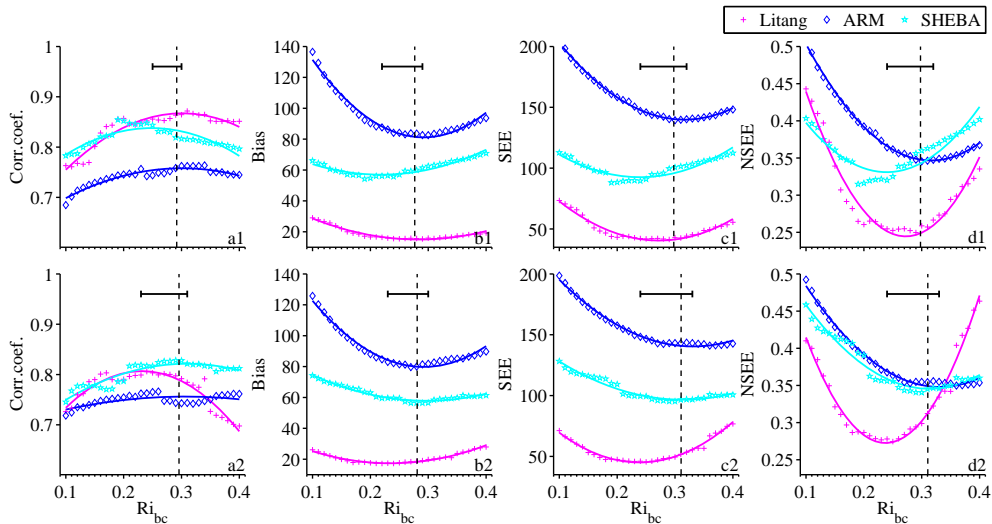


Figure 9. Comparison between estimated PBLHs using the bulk Richardson number method with $z_s = 40$ m (upper panels) and $z_s = 80$ m (lower panels) and observed PBLHs for type II SBLs. The correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of Ri_{bc} across the three sites.

for type II SBLs. For UBLs, the optimal Ri_{bc} value falls between 0.33 and 0.39, depending on the choice of z_s . To be exact, the best choices of Ri_{bc} suggested by this study are 0.24 ($z_s = 40$ m), 0.31 ($z_s = 80$ m), and 0.39 ($z_s = z_{SAL}$) for type I SBLs, type II SBLs, and UBLs, respectively. Note that z_s is recommended to be 80 m for type II SBLs, given that the surface layer is usually thicker for type II SBLs than for type I SBLs.

4.3 Impacts of thermal stratification on Ri_{bc}

With the above analyses, the best choices of Ri_{bc} are inferred under different thermal stratification conditions. Hence, the traditional way of determining the PBLH using a single value of Ri_{bc} without considering the dependence of Ri_{bc} on thermal stratification (e.g., Troen and Mahrt, 1986) needs to be revised. For example, the Yonsei University

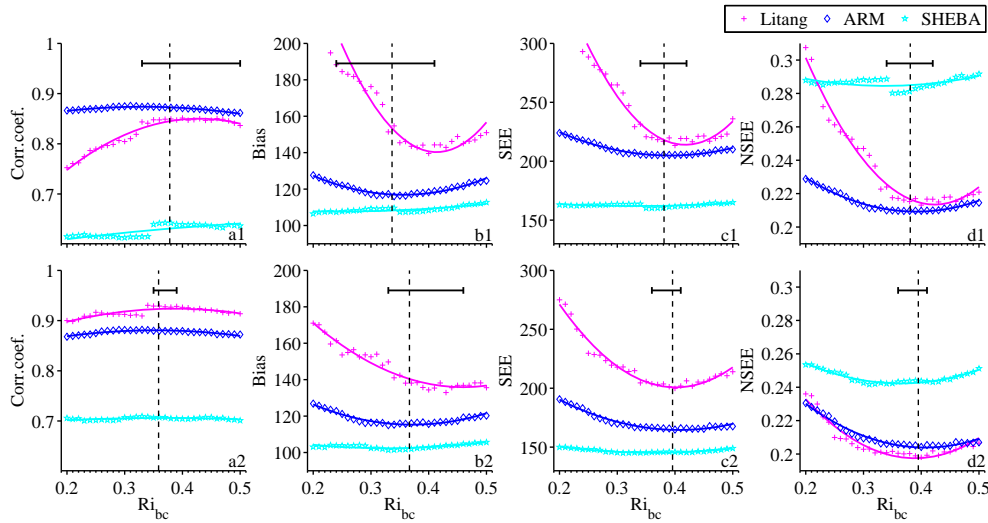


Figure 10. Comparison between estimated PBLHs using the bulk Richardson number method with $z_s = 0.1$ PBLH (upper panels) and $z_s = z_{SAL}$ (lower panels) and observed PBLHs for UBLs. The correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of Ri_{bc} across the three sites.

Table 2. Inferred bulk Richardson number of the entire PBL, Ri_{bc} , for different types of PBLs and sites, with different values of z_s . n refers to the sample number. * indicates the best choice.

| Regime | z_s (m) | Litang | | CASES99 | | ARM | | SHEBA | | Total | |
|---------|-----------|-----------|-----|-----------|-----|-----------|-----|-----------|-----|-----------|-----|
| | | Ri_{bc} | n | Ri_{bc} | n | Ri_{bc} | n | Ri_{bc} | n | Ri_{bc} | n |
| Type I | 40 | 0.25 | 19 | 0.22 | 29 | 0.24 | 373 | 0.25 | 27 | 0.24* | 448 |
| | 80 | 0.26 | 11 | 0.21 | 21 | 0.23 | 211 | 0.24 | 17 | 0.23 | 261 |
| Type II | 40 | 0.27 | 53 | \ | \ | 0.32 | 194 | 0.24 | 49 | 0.30 | 296 |
| | 80 | 0.24 | 53 | \ | \ | 0.33 | 194 | 0.31 | 49 | 0.31* | 296 |
| UBL | 40 | 0.41 | 23 | \ | \ | 0.36 | 182 | 0.20 | 75 | 0.33 | 280 |
| | 80 | 0.41 | 23 | \ | \ | 0.38 | 182 | 0.32 | 75 | 0.39 | 280 |
| | 0.1 PBLH | 0.42 | 23 | \ | \ | 0.39 | 182 | 0.34 | 62 | 0.38 | 267 |
| | z_{SAL} | 0.39 | 23 | \ | \ | 0.41 | 182 | 0.36 | 75 | 0.39* | 280 |

(YSU) PBL scheme in the Weather Research and Forecasting (WRF) model assumes $Ri_{bc} = 0.25$ over land (Hong, 2010), while $Ri_{bc} = 0.5$ is used in the Holtslag and Boville (HB) boundary-layer scheme in CCM2 (Holtslag and Boville, 1993). To examine the impact of thermal stratification on Ri_{bc} , we obtained a single representative Ri_{bc} for all stratification conditions with the same sounding data from Litang, ARM, and SHEBA sites, assuming the lower boundary heights z_s of 40 and 80 m, and z_{SAL} for type I SBLs, type II SBLs, and UBLs, respectively. According to the minimum SEE and NSEE, the optimal choice of Ri_{bc} for all PBL types is 0.33 (Fig. 11), which is close to that used in CAM4 ($Ri_{bc} = 0.3$; Neale et al., 2010). In Fig. 12, the errors that occur when a single value of Ri_{bc} is used ($Ri_{bc} = 0.33$ determined by our study, $Ri_{bc} = 0.25$ in WRF–YSU, and $Ri_{bc} = 0.5$ in

CCM2–HB) are presented, as compared to the errors with a new scheme that uses $Ri_{bc} = 0.24, 0.31,$ and 0.39 for type I SBLs, type II SBLs, and UBLs, respectively. It is found that the new scheme with variable Ri_{bc} is more reliable in estimating PBLH, suggesting that the impact of atmospheric stability or thermal stratification on Ri_{bc} is significant and that the variation of Ri_{bc} with atmospheric stability should be taken into account when estimating the PBLH using the bulk Richardson number method.

To further investigate the improvements in estimating PBLHs with the new, variable Ri_{bc} values, simulations using CAM4 are conducted at the ARM site, with the default (i.e., 0.3) and the new, variable Ri_{bc} values used to estimate PBLHs. Figure 13 shows a comparison between the observed and the CAM4-simulated PBLHs with the default and new

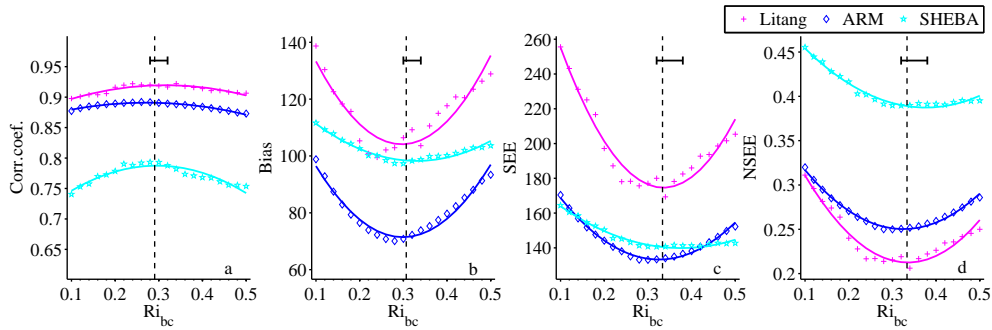


Figure 11. Comparison between estimated PBLH using the bulk Richardson number method and observed PBLHs for all types of PBLs. The correlation coefficient (a), bias (b), standard error (c), and normalized standard error (d) are shown. The sounding data are taken from Litang (plus sign), ARM Shouxian (diamond), and SHEBA (pentacle). The curved lines are obtained by quadratic curve-fitting, the black vertical dashed lines indicate a representative Ri_{bc} for all three sites, and the error bars indicate the range of Ri_{bc} across the three sites.

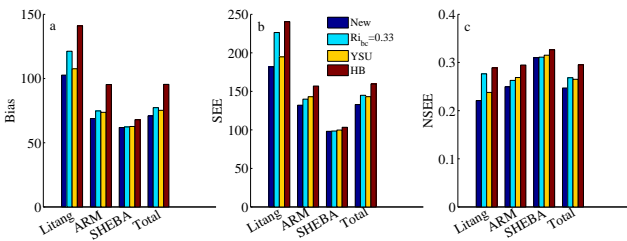


Figure 12. Comparisons between observed and estimated PBLHs with a single $Ri_{bc} = 0.33$ for all PBL conditions, with $Ri_{bc} = 0.25$ as in the YSU scheme, with $Ri_{bc} = 0.5$ as in the HB scheme, and with the new, variable values ($Ri_{bc} = 0.24, 0.31,$ and 0.39 for type I SBLs, type II SBLs, and UBLs, respectively): (a) bias, (b) standard error, (c) normalized standard error.

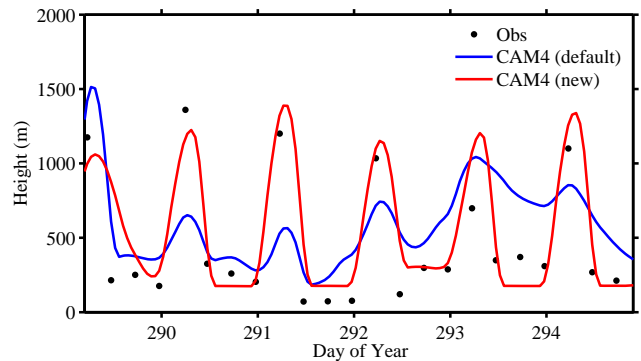


Figure 13. Comparison of observed and simulated PBLHs using CAM4 with the default and new Ri_{bc} values during the period 16–21 October 2008 at the ARM site.

5 Conclusions

Ri_{bc} values over a 6-day period. It can be seen that the simulated PBLHs with the new Ri_{bc} values have a more pronounced diurnal cycle, which are also closer to the observations. Over the whole observational period, results indicate that the bias, SEE, and NSEE are 270.1 m, 379.3 m, and 0.75 with the new, variable Ri_{bc} values, respectively, and 306.2 m, 417.5 m, 0.83 with the default Ri_{bc} value, respectively. Again, these results indicate that the impact of thermal stratification on Ri_{bc} should be considered in calculating PBLH with the bulk Richardson number method and that the new Ri_{bc} values determined in this study improve model results in real applications. It is pointed out here that there are still large biases in the CAM4-simulated PBLH even with the new Ri_{bc} values, which are probably related to the biases in the model physics and parameterizations (e.g., parameterizations of land–atmosphere interactions and boundary-layer turbulence). Unraveling how biases in the model physics and parameterizations affect the PBLH is nevertheless out of the scope of this study.

The PBLH is an important parameter in boundary-layer research and accurate estimates of the PBLH are vital for many environmental applications. In this study, we investigated several methods for computing the PBLH under different stratification conditions. The Tur method is considered as the most accurate approach for any atmospheric stratification due to its direct measurement of turbulence intensity. However, such a method is expensive and thus cannot be widely applied. However, determining of the PBLH with radio soundings through the PTG, LLJ, and modified parcel methods is more affordable. These methods usually work well when the PBL has certain unique features but may fail under special conditions (e.g., a weak underlying inversion for strongly stable boundary layers, multiple wind maxima for weakly stable boundary layers, and no clear maximum of vertical gradient of potential temperature for unstable boundary layers). With corrections made for these special cases, we used the Tur, PTG, LLJ, and modified parcel methods to determine PBLHs from Litang, ARM Shouxian, SHEBA,

and CASES99 field experiments. The estimated PBLHs using these methods are treated as observed PBLHs.

The bulk Richardson number method is more commonly used in numerical models due to its reliability for all atmospheric stratification conditions, which requires a specified value of the bulk Richardson number for the entire PBL, or Ri_{bc} . In many numerical models, the Ri_{bc} is specified as one single value (e.g., 0.25 for WRF–YSU, 0.5 for CCM2–HB, 0.3 for CAM4) and hence its dependence on the thermal stratification is ignored. This study infers a representative Ri_{bc} for each stratification condition from observed PBLHs using linear fitting and statistical error minimization approaches. Results indicate that the best choices for Ri_{bc} are 0.24, 0.31, and 0.39 for strongly stable boundary layers (type I SBLs), weakly stable boundary layers (type II SBLs), and unstable boundary layers (UBLs), respectively. Both offline and online evaluations show that the new and variable Ri_{bc} values proposed in this study yield more reliable estimates of the PBLH, suggesting that the variation of Ri_{bc} should be considered when computing the PBLH with the bulk Richardson number method.

The Supplement related to this article is available online at doi:10.5194/gmd-7-2599-2014-supplement.

Acknowledgements. This study is supported by the China Meteorological Administration under grant GYHY201006024, the National Program on Key Basic Research Project of China (973) under grants 2012CB417203 and 2011CB403501, the National Natural Science Foundation of China under grant 41275022, and the CAS Strategic Priority Research Program, grant 5 XDA05110101. We are grateful to two anonymous reviewers for their careful review and valuable comments, which led to a substantial improvement of this manuscript.

Edited by: A. Colette

References

- Andre, J. C. and Mahrt, L.: The nocturnal surface inversion and influence of clear-air radiative cooling, *J. Atmos. Sci.*, 39, 864–878, 1982.
- Andreas, E. L., Fairall, C. W., Guest, P. S., and Persson, P. O. G.: An overview of the SHEBA atmospheric surface flux program, in: Preprints, Fifth Conference On Polar Meteorology and Oceanography, 10–15 January 1999, Dallas, TX, American Meteorological Society, Boston, 411–416, 1999.
- Balsley, B. B., Frehlich, R. G., Jensen, M. L., and Meillier, Y.: High-resolution in situ profiling through the stable boundary layer: Examination of the SBL top in terms of minimum shear, maximum stratification, and turbulence decrease, *J. Atmos. Sci.*, 63, 1291–1307, 2006.
- Basu, S., Holtslag, A. A. M., Caporaso, L., Riccio, A., and Steenveeld, G. J.: Observational support for the stability dependence of the bulk Richardson number across the stable boundary layer, *Bound.-Lay. Meteorol.*, 150, 515–523, doi:10.1007/s10546-013-9878-y, 2014.
- Beyrich, F.: Mixing height estimation from sodar data—a critical discussion, *Atmos. Environ.*, 31, 3941–3954, 1997.
- Bradley, R. S., Keimig, F. T., and Diaz, H. F.: Recent changes in the North American Arctic boundary layer in winter, *J. Geophys. Res.*, 98, 8851–8858, doi:10.1029/93JD00311, 1993.
- Dai, C., Wang, Q., Kalogiros, J. A., Lenschow, D. H., Gao, Z., and Zhou, M.: Determining boundary-layer height from aircraft measurements, *Bound.-Lay. Meteorol.*, 152, 277–302, doi:10.1007/s10546-014-9929-z, 2014.
- Dai, C. Y., Gao, Z. Q., Wang, Q., and Cheng, G.: Analysis of atmospheric boundary layer height characteristics over the Arctic Ocean using the aircraft and GPS soundings, *Atmospheric and Oceanic Science Letters*, 4, 124–130, 2011.
- Deardorff, J. W.: Parameterization of the planetary boundary layer for use in general circulation models, *Mon. Weather Rev.*, 100, 93–106, 1972.
- Esau, I. and Zilitinkevich, S.: On the role of the planetary boundary layer depth in the climate system, *Adv. Sci. Res.*, 4, 63–69, doi:10.5194/asr-4-63-2010, 2010.
- Gao, Z., Chae, N., Kim, J., Hong, J., Choi, T., and Lee, H.: Modeling of surface energy partitioning, surface temperature, and soil wetness in the Tibetan prairie using the Simple Biosphere Model 2 (SiB2), *J. Geophys. Res.*, 109, D06102, doi:10.1029/2003JD004089, 2004.
- Garratt, J. R.: *The atmospheric boundary layer*, Cambridge University Press, 316 pp., 1992.
- Hanna, S. R.: The thickness of the planetary boundary layer, *Atmos. Environ.*, 3, 519–536, 1969.
- Holtslag, A. A. M. and Boville, B. A.: Local versus nonlocal boundary-layer diffusion in a global climate model, *J. Climate*, 6, 1825–1842, 1993.
- Holtslag, A. A. M. and Nieuwstadt, F. T. M.: Scaling the atmospheric boundary layer, *Bound.-Lay. Meteorol.*, 36, 201–209, 1986.
- Holtslag, A. A. M., van Meijgaard, E., and de Rooy, W. C.: A comparison of boundary layer diffusion schemes in unstable conditions over land, *Bound.-Lay. Meteorol.*, 76, 69–95, 1995.
- Holzworth, C. G.: Estimates of mean maximum mixing depths in the contiguous United States, *Mon. Weather Rev.*, 92, 235–242, 1964.
- Hong, S.-Y.: A new stable boundary-layer mixing scheme and its impact on the simulated East Asia summer monsoon, *Q. J. Roy. Meteor. Soc.*, 136, 1481–1496, 2010.
- Jericevic, A. and Grisogono, B.: The critical bulk Richardson number in urban areas: verification and application in a numerical weather prediction model, *Tellus A*, 58, 19–27, 2006.
- Joffre, S. M., Kangas, M., Heikinheimo, M., and Kitaigorodskii, S. A.: Variability of the stable and unstable atmospheric boundary-layer height and its scales over a boreal forest, *Bound.-Lay. Meteorol.*, 99, 429–450, 2001.
- Konor, C. S., Boezio, G. C., Mechoso, C. R., and Arakawa, A.: Parameterization of PBL processes in an atmospheric general circulation model: Description and preliminary assessment, *Mon. Weather Rev.*, 137, 1061–1082, 2009.

- Leventidou, E., Zanis, P., Balis, D., Giannakaki, E., Pytharoulis, I., and Amiridis, V.: Factors affecting the comparisons of planetary boundary layer height retrievals from CALIPSO, ECMWF and radiosondes over Thessaloniki, Greece, *Atmos. Environ.*, 74, 360–366, 2013.
- Liu, S. and Liang, X. Z.: Observed diurnal cycle climatology of planetary boundary layer height, *J. Climate*, 23, 5790–5809, doi:10.1175/2010JCLI3552.1, 2010.
- Mahrt, L., Heald, R. C., Lenschow, D. H., Stankov, B. B., and Troen, I.: An observational study of the nocturnal boundary layer, *Bound.-Lay. Meteorol.*, 17, 247–264, 1979.
- Martin, C. L., Fitzjarrald, D., Garstang, M., Oliveira, A. P., Greco, S., and Browell, E.: Structure and growth of the mixing layer over the Amazonian rain forest, *J. Geophys. Res.*, 93, 1361–1375, 1988.
- Medeiros, B., Hall, A., and Stevens, B.: What controls the climatological depth of the PBL?, *J. Climate*, 18, 2877–2892, doi:10.1175/JCLI3417, 2005.
- Meillier, Y. P., Froehlich, R. G., Jones, R. M., and Balsley, B. B.: Modulation of small-scale turbulence by ducted gravity waves in the nocturnal boundary layer, *J. Atmos. Sci.*, 65, 1414–1427, 2008.
- Melgarejo, J. W. and Deardorff, J. W.: Stability functions for the boundary-layer resistance laws based upon observed boundary-layer heights, *J. Atmos. Sci.*, 31, 1324–1333, 1974.
- Neale, R. B., Richter, J. H., Conley, A. J., Park, S., Lauritzen, P. H., Gettelman, A., Williamson, D. L., Rasch, P. J., Vavrus, S. J., Taylor, M. A., Collins, W. D., Zhang, M., and Lin, S.-J.: Description of the NCAR Community Atmosphere Model (CAM 4.0), NCAR/TN-485+STR NCAR, Boulder, CO, 224 pp., 2010.
- Persson, P. O. G., Fairall, C. W., Andreas, E. L., Guest, P. S., and Perovich, D. K.: Measurements near the atmospheric surface flux group tower at SHEBA: Near-surface conditions and surface energy budget, *J. Geophys. Res.*, 107, 8045, doi:10.1029/2000JC000705, 2002.
- Richardson, H., Basu, S., and Holtslag, A. A. M.: Improving stable boundary-layer height estimation using a stability-dependent critical bulk Richardson number, *Bound.-Lay. Meteorol.*, 148, 93–109, doi:10.1007/s10546-013-9812-3, 2013.
- Russell, P. B., Uthe, E. E., Ludwig, F. L., and Shaw, N. A.: A comparison of atmospheric structure as observed with monostatic acoustic sounder and lidar techniques, *J. Geophys. Res.*, 79, 5555–5566, 1974.
- Seibert, P., Beyrich, F., Gryning, S. E., Joffre, S., Rasmussen, A., and Tercier, P.: Review and intercomparison of operational methods for the determination of the mixing height, *Atmos. Environ.*, 34, 1001–1027, 2000.
- Seidel, D. J., Ao, C. O., and Li, K.: Estimating climatological planetary boundary layer heights from radiosonde observations: Comparison of methods and uncertainty analysis, *J. Geophys. Res.*, 115, D16113, doi:10.1029/2009JD013680, 2010.
- Stull, R. B.: An introduction to boundary layer meteorology, Kluwer Academic, 666 pp., 1988.
- Sugiyama, G. and Nasstrom, J. S.: Methods for Determining the Height of the Atmospheric Boundary Layer, UCRL-ID-133200, Lawrence Livermore National Laboratory Report, February 1999.
- Troen, I. and Mahrt, L.: A simple model of the planetary boundary layer: Sensitivity to surface evaporation, *Bound.-Lay. Meteorol.*, 37, 129–148, 1986.
- Vogelezang, D. H. P. and Holtslag, A. A. M.: Evaluation and model impacts of alternative boundary-layer height formulations, *Bound.-Lay. Meteorol.*, 81, 245–269, doi:10.1007/BF02430331, 1996.
- Wang, Q. and Wang, S.: Turbulent and thermodynamic structure of the autumnal arctic boundary layer due to embedded clouds, *Bound.-Lay. Meteorol.*, 113, 225–247, 2004.
- Wang, Q., Lenschow, D. H., Pan, L., Schillawski, R. D., Kok, G. L., Prevot, A. S. H., Laursen, K., Russell, L. M., Bandy, A. R., Thornton, D. C., and Suhre, K.: Characteristics of the marine boundary layers during two Lagrangian measurement periods: 2. Turbulence structure, *J. Geophys. Res.*, 104, 21767–21784, 1999.
- Zilitinkevich, S. S. and Baklanov, A.: Calculation of the height of stable boundary layers in practical applications, *Bound.-Lay. Meteorol.*, 105, 389–409, 2002.
- Zhang, J. A., Rogers, R. F., Nolan, D. S., and Marks, F. D.: On the Characteristic Height Scales of the Hurricane Boundary Layer, *Mon. Weather Rev.*, 139, 2523–2535, 2011.