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Corresponding Author	Family Name	Bindi
	Particle	
	Given Name	D.
	Suffix	
	Division	Helmholtz Centre Potsdam GFZ
	Organization	German Research Centre for Geosciences
	Address	Helmholtzstraße 7, Potsdam, 14467 , Germany
	Email	bindi@mi.ingv.it

Author	Family Name	Massa
	Particle	
	Given Name	M.
	Suffix	
	Division	
	Organization	Istituto Nazionale di Geofisica e Vulcanologia
	Address	via Bassini 15, Milano, 20133 , Italy
	Email	

Author	Family Name	Luzi
	Particle	
	Given Name	L.
	Suffix	
	Division	
	Organization	Istituto Nazionale di Geofisica e Vulcanologia
	Address	via Bassini 15, Milano, 20133 , Italy
	Email	

Author	Family Name	Ameri
	Particle	
	Given Name	G.
	Suffix	
	Division	
	Organization	FUGRO-Geoter
	Address	Auriol, 13390 , France
	Email	

Author	Family Name	Pacor
	Particle	
	Given Name	F.
	Suffix	

	Division	
	Organization	Istituto Nazionale di Geofisica e Vulcanologia
	Address	via Bassini 15, Milano, 20133 , Italy
	Email	
Author	Family Name	Puglia
	Particle	
	Given Name	R.
	Suffix	
	Division	
	Organization	Istituto Nazionale di Geofisica e Vulcanologia
	Address	via Bassini 15, Milano, 20133 , Italy
	Email	
Author	Family Name	Augliera
	Particle	
	Given Name	P.
	Suffix	
	Division	
	Organization	Istituto Nazionale di Geofisica e Vulcanologia
	Address	via Bassini 15, Milano, 20133 , Italy
	Email	
Schedule	Received	22 December 2012
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Keywords (separated by '-')	Ground motion prediction equation - Europe - RESORCE data set	
Footnote Information	Electronic supplementary material The online version of this article (doi:10.1007/s10518-013-9525-5) contains supplementary material, which is available to authorized users.	

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~~Pan-European~~ ground-motion prediction equations for the average horizontal component of PGA, PGV, and 5 %-damped PSA at spectral periods up to 3.0 s using the RESORCE dataset

D. Bindi · M. Massa · L. Luzi · G. Ameri · F. Pacor ·
R. Puglia · P. Augliera

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D. Bindi (✉)
Helmholtz Centre Potsdam GFZ, German Research Centre for Geosciences,
Helmholtzstraße 7, 14467 Potsdam, Germany
e-mail: bindi@mi.ingv.it

M. Massa · L. Luzi · F. Pacor · R. Puglia · P. Augliera
Istituto Nazionale di Geofisica e Vulcanologia, via Bassini 15, 20133 Milano, Italy

G. Ameri
FUGRO-Geoter, 13390 Auriol, France

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 21 aftershock classification) or to the non-linear site ~~amplifications~~. The proposed GMPEs have
 22 lower median values than global models at short periods and large distances, while are con-
 23 sistent with global models at long periods ($T > 1$) s. Consistency is found with two regional
 24 models developed for Turkey and Italy, as the considered dataset is dominated by waveforms
 25 recorded in these regions.

26 **Keywords** Ground motion prediction equation · Europe · RESORCE data set

27 1 Introduction

28 In the framework of the FP7 EU project SHARE (<http://www.share-eu.org/>), Yener et al.
 29 (2010) compiled a new databank for hazard studies in Europe. The SHARE databank is a
 30 collection of records and metadata contained in previously compiled databases with a ranking
 31 of preferred metadata in case of data overlaps and no waveform processing. The SHARE
 32 databank was then exploited by the SIGMA project (<http://projet-sigma.com/organisation.html>),
 33 in order to improve the seismic hazard assessment in France. A new databank (hereafter
 34 referred to as RESORCE) has been compiled, from the subset of European data included in
 35 SHARE after metadata revision and uniform re-processing of waveforms (Akkar et al. 2013).

36 The RESORCE databank has been used by different authors to derive a set of ground-
 37 motion models following different approaches (Douglas et al. 2013). In this article, we
 38 first describe the data selection applied to RESORCE for deriving ~~Pan-European~~ GMPEs
 39 and we introduce the functional form, along with the explanatory variables considered
 40 in the regressions. The results, provided in terms of coefficients of the equations and
 41 associated 95 % confidence levels, are discussed through a comparison with previously
 42 derived global and regional models and analyzing the residual distributions. A compre-
 43 hensive comparison among the different models derived from RESORCE is presented
 44 in Douglas et al. (2013).

45 2 Flat file compilation from the RESORCE data bank

46 The RESORCE databank (Akkar et al. 2013) originally includes 5,882 waveforms from
 47 1814 earthquakes occurred in Europe and Middle East from 1967 to 2011 in the magnitude
 48 range from 2.8 to 7.8 (the largest magnitude is relevant to the 1969/02/28 02:40:31 Portu-
 49 gal, offshore single-recorded earthquake). In this study no additional analyses are performed
 50 to identify poor quality data and a preliminary selection is made in order to exclude: (i)
 51 unprocessed records, (ii) data lacking the three components of ground-motion and (iii) earth-
 52 quakes for which moment magnitude is not provided. Events with unreliable magnitude are
 53 also disregarded.

54 The following criteria are adopted to select the records for the regression:

- 55 – *Range of validity.* Given the recent interest in considering small magnitude earth-
 56 quakes for assessing the hazard in several regions of Europe ([http://projet-sigma.com/](http://projet-sigma.com/ScientificObjectives.html)
 57 [ScientificObjectives.html](http://projet-sigma.com/ScientificObjectives.html)), records from events with moment magnitudes larger than or
 58 equal to four are considered. Hypocentral depths are lower than 35 km and distances
 59 (Joyner-Boore, R_{jb} , or epicentral R_{epi}) shorter than 300 km. The epicentral distance, R_{epi} ,
 60 is used to approximate R_{jb} when the latter is unspecified, but only when $M \leq 5$ and

- 61 $R_{\text{epi}} \geq 10$ km. For larger magnitudes and smaller epicentral distances, records without
 62 R_{jb} are disregarded.
- 63 – *Waveform selection.* Only records filtered with low-pass corner frequency larger than
 64 or equal to 20 Hz and, for each period T , only recordings filtered with high-pass corner
 65 frequency $f_{\text{hp}} \leq 1/(1.25 T)$ are considered.
- 66 – *Sampling.* Single recorded earthquakes are not selected.
- 67 Two different subsets of data are then compiled, based on the information available about
 68 site classification and style of faulting:
- 69 – *DS-EC8 dataset.* It is composed by waveforms recorded by stations characterized by
 70 EC8 site classes (CEN 2004), which are based on the average shear-wave velocity of the
 71 uppermost 30 m, V_{s30} (class A: $V_{\text{s30}} \geq 800$ m/s, class B: $360 \leq V_{\text{s30}} \leq 800$ m/s, class
 72 C: $180 \leq V_{\text{s30}} \leq 360$ m/s, class D: $V_{\text{s30}} \leq 180$ m/s and class E: 5–20 m of C or D-type
 73 alluvium underlain by stiffer material with $V_{\text{s30}} \geq 800$ m/s).
- 74 The RESORCE site categories are determined either from shear wave velocity profiles or
 75 inferred by surface geology and, among them, only classes from A to D are accounted
 76 for, as only few stations (less than 5) are classified as class E. Waveforms from events
 77 with unspecified focal mechanism are included in DS-EC8. This data set contains 2,126
 78 recordings from 365 earthquakes and 697 stations.
- 79 – *DS-VS30 dataset.* It is composed by waveforms recorded by stations with measured V_{s30}
 80 and relative to events with known focal mechanism. It contains 1,224 recordings relevant
 81 to 255 earthquakes recorded by 345 stations.

82 Figure 1 shows the magnitude-distance distribution of the two datasets at $T = 0.1$ s. Each
 83 individual magnitude-distance entity is shown in the left panel, where different colors indicate
 84 different networks, whereas the right panel shows the hit counts, computed over a grid where
 85 the distance is discretized into 30 bins, equally spaced over a logarithmic scale from 1 to
 86 300 km, and the magnitude range is discretized into 0.15 unit intervals. The bulk of the
 87 selected data sets is represented by Italian and Turkish data, that mainly cover the magnitude
 88 range from 4 to 6.5 and the distance range from 10 to 200 km. The data set includes sixteen
 89 events with magnitude larger than 6.5, seven of which having magnitude larger than 7.

90 Figure 2 shows the distribution of site categories and focal mechanisms for the two datasets.
 91 The main features of the DS-EC8 are shown in the upper left panel, which indicates that
 92 categories A, B and C characterize the Italian dataset, while Turkish data have been mainly
 93 recorded by stations belonging to soil category B and C. When data characterized by shear
 94 wave velocity profiles are selected (DS-VS30, lower left panel) there is a strong reduction of
 95 class A and a relative increase of Turkish data, as the majority of European networks lack of
 96 detailed site characterization.

97 The style of faulting distribution is shown in Fig. 2 (top right panel for DS-EC8 and bottom
 98 right panel for DS-VS30). Normal and strike-slip mechanisms, that are the bulk of the Italian
 99 and Turkish datasets, are the most represented. The relative proportion of waveforms from
 100 normal and strike-slip style of faulting does not change in the two datasets, although there is
 101 a considerable reduction of Italian normal events in DS-VS30.

102 Since we consider only the spectral ordinates within the usable frequency band, which is
 103 function of the low-cut frequency, the number of selected recordings varies with period and,
 104 in particular, decreases with increasing periods, as detailed in Fig. 3 for the DS-EC8 dataset.
 105 The number of selected recordings, stations and earthquakes for $T = 0.1$ s are 2,126, 697 and
 106 365, respectively. At 3 s the recordings decrease relevantly being 1,460, from 580 stations
 107 and 226 earthquakes. Figure 3 (right) displays the plot of the low cut filter corner in function
 108 of the earthquake magnitude for the available data set, which indicates that some of the small

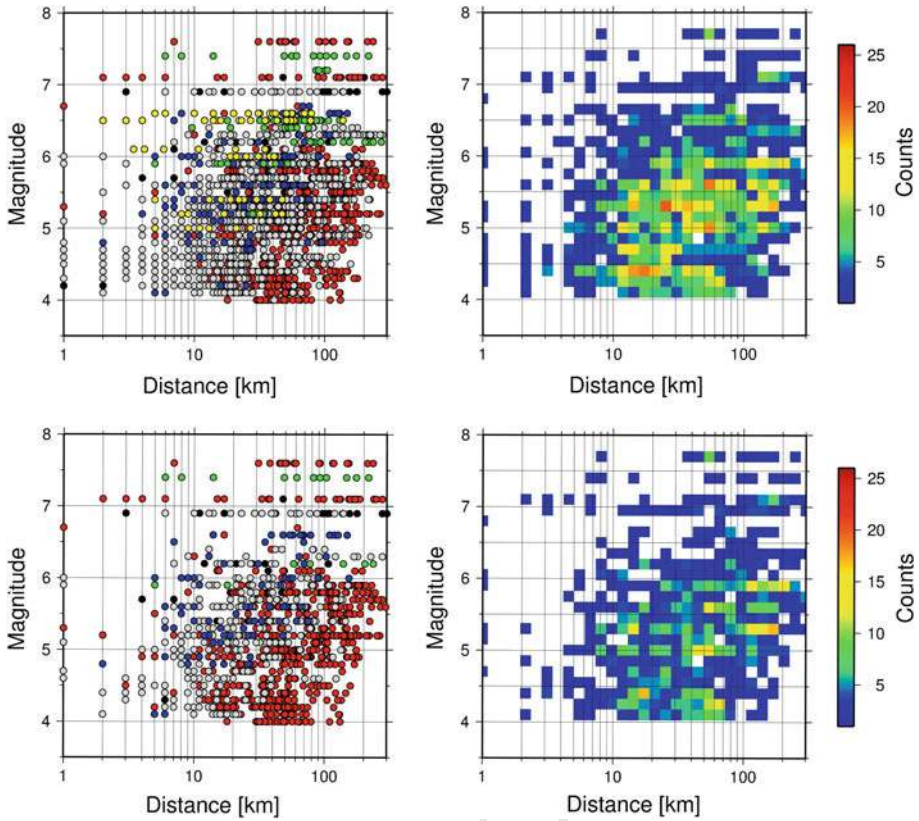


Fig. 1 *Right panels*: Magnitude versus distance scatter plot for DS-EC8 (*top*) and DS-Vs30 (*bottom*) at $T = 0.1$ s. The records are colour coded accordingly to the network: *red* (Turkey); *gray* (Italy); *blue* (Greece); *green* (Iran); *yellow* (Iceland); *black* (other countries). *Left panels*: hit counts computed for the data distribution shown in the *left panels*, discretizing the distance range (1–300 km) into 30 equally spaced bins over a logarithmic scale and considering 0.15 magnitude unit intervals

109 magnitude events have low filter corners and, therefore, noise at long periods could still be
110 present in the waveforms after processing.

111 3 Functional form and regression

112 The GMPEs are derived considering a parametric model based on the following functional
113 form (e.g. [Boore and Atkinson 2008](#); [Akkar and Cagnan 2010](#); [Bindi et al. 2011a](#))

$$114 \log_{10} Y = e_1 + F_D(R, M) + F_M(M) + F_S + F_{sof} \quad (1)$$

115 where the distance F_D and magnitude F_M functions are given by:

$$116 F_D(R, M) = [c_1 + c_2 (M - M_{ref})] \log_{10} \left(\sqrt{R^2 + h^2/R_{ref}} \right) - c_3 \left(\sqrt{R^2 + h^2} - R_{ref} \right) \quad (2)$$

117

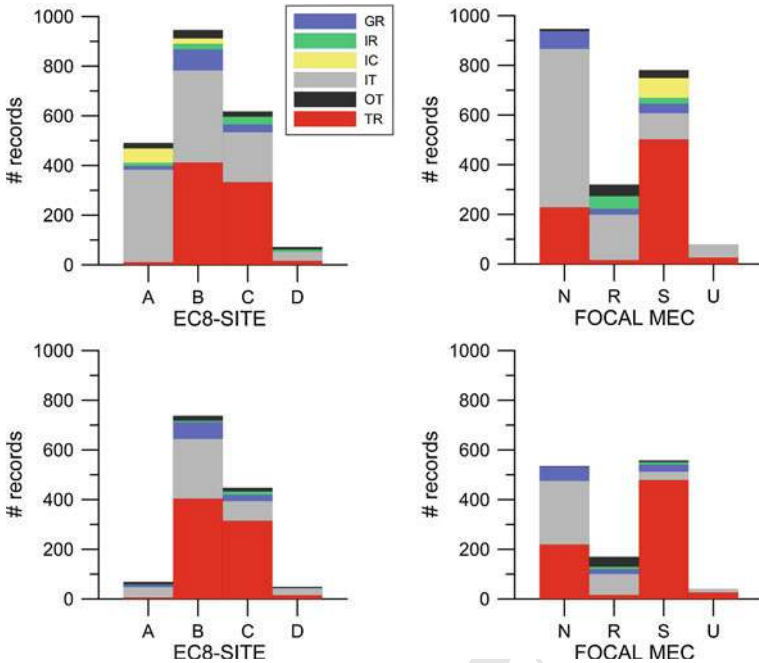


Fig. 2 EC8 site categories (*right*) and style-of-faulting (*left*) distributions for the DS-EC8 (*top*) and DS-Vs30 (*bottom*) datasets at $T = 0.1$ s. Different colours indicate different countries

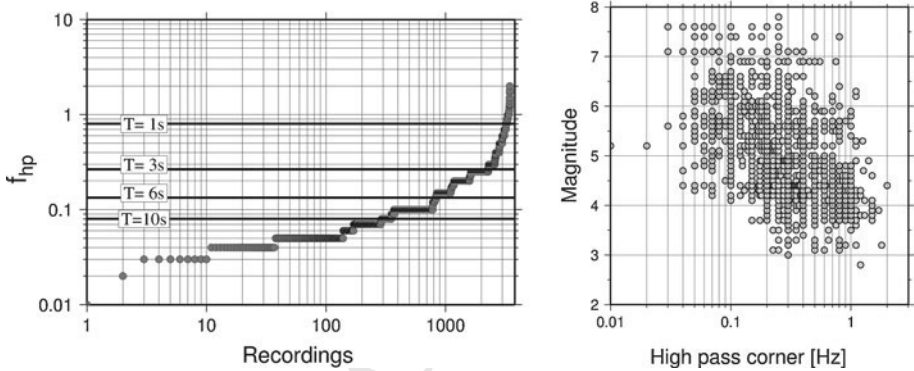


Fig. 3 *Left*: Number of recordings (abscissa) as function of the high-pass corner frequency (ordinate). The selections, considering $f = 1/(1.25 T)$, at $T = 1, 3, 6$ and 10 s, are indicated. *Right*: Magnitude versus corner frequency scatter plot, considering the high-pass filter applied to the original data set

118

$$F_M(M) = \begin{cases} b_1 (M - M_h) + b_2 (M - M_h)^2 & \text{for } M \leq M_h \\ b_3 (M - M_h) & \text{otherwise} \end{cases} \quad (3)$$

119

120 Preliminary analysis about the data scaling with magnitude and distance confirmed the suit-
 121 ability of the selected functional form to describe the dependences of the ground-motion
 122 parameters on the explanatory variables. Following [Bommer and Akkar \(2012\)](#), the regres-
 123 sions are performed considering both a point-source and an extended-source measure of the

124 source-to-site distance R , namely the Joyner and Boore distance R_{JB} and the hypocentral dis-
 125 tance R_{hypo} . Additional explanatory variables related to the source model (e.g. hanging/foot
 126 walls effect; depth to the top of the rupture; etc.) or other measure for the source-to-station
 127 distance (e.g. distance from the rupture) are not considered because of the lack of information
 128 in RESORCE.

129 The functional form F_S in Eq. (1) represents the site amplification. The model derived in
 130 this article includes only a linear site amplification term although nonlinear site effects are
 131 expect to be important for strong shaking at soil sites, that is for large and close earthquakes
 132 recorded at site with low V_{s30} values (e.g. classes C and D of EC8). Unfortunately these
 133 conditions are not well sampled in RESORCE (Figs. 1 and 2). We show in the following
 134 that nonlinear site effects, if present, do not significantly bias the median predictions of the
 135 model.

136 Regarding the linear site amplification term, we consider two models, depending on the
 137 dataset (DS-EC8 or DS-VS30). The first model is $F_S = s_j C_j$, for $j = 1, \dots, 4$, where s_j are the
 138 coefficients to be determined through the regression analysis, while C_j are dummy variables
 139 used to denote the four considered EC8 site classes (A–D). The regression for the EC8 model
 140 is performed constraining to zero the coefficient for class A (reference site class). In the
 141 second model, the site effects are expressed in terms of V_{S30} as $F_S = \gamma \log_{10}(V_{S30}/V_{ref})$
 142 where $V_{ref} = 800$ m/s and γ is to be determined through the regression.

143 The functional form F_{sof} in Eq. (1) represents the style of faulting correction and it is
 144 given by $F_{sof} = f_j E_j$, for $j = 1, \dots, 4$, where f_j are the coefficients to be determined during
 145 the analysis and E_j are dummy variables used to denote the different fault classes: nor-
 146 mal (N), reverse (R), strike-slip (S) and unspecified (U). Since earthquakes with unknown
 147 style of faulting are not included in DS-VS30 the class U is considered only for DS-
 148 EC8. The reference style of faulting conditions (i.e. parameters constrained to zero in the
 149 regressions) are class U for DS-EC8 and the average over the three classes N, R, S for
 150 DS-VS30.

151 After trial regressions, the variables M_{ref} , M_h , R_{ref} (Eqs. 2 and 3) have been fixed to
 152 5.5, 6.75 and 1 km, respectively. Coefficients c_3 and b_3 are constrained to be non-negative.
 153 As response variable Y , the geometric mean of the horizontal components for peak ground
 154 acceleration (PGA in cm/s^2) and velocity (PGV in cm/s) are considered, along with 5%
 155 damped pseudo-spectral acceleration (PSA in cm/s^2) computed over 27 periods in the range
 156 0.02–3 s. The regressions are performed applying a random effect approach (Abrahamson
 157 and Youngs 1992), that allows to determine the components of the standard deviation of
 158 the regression (commonly referred to as sigma, σ), namely the between-events (τ) and the
 159 within-event (ϕ) components, as well as the site-to-site component (ϕ_{S2S}) (e.g., Bindi et al.
 160 2009, 2011b). For the definition of the components of variability, see Al Atik et al. (2010).
 161 Finally, for each period, the standard error of the distribution of the coefficients is obtained
 162 through a bootstrap analysis (Efron and Tibshirani 1994) considering 30 different bootstrap
 163 replications of the original data set, being each replication composed by the same number of
 164 data as the original set but randomly selected with repetitions.

165 4 Results

166 The regression coefficients for the two datasets, and the relevant 95% confidence intervals,
 167 are shown in Tables 1 and 2, for Joyner-Boore distance, and in Tables 3 and 4, for hypocentral
 168 distance. The tables including the 95% confidence intervals are reported in the Appendix
 169 (Tables 6 and 7 for R_{JB} ; Tables 8 and 9 for R_{hypo}).

Table 1 Coefficients for the model derived in this study (see Eqs. 1–3) for R_{IB} and EC8 ground categories

T[sec]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	3.47806	3.58006	3.78163	3.7926	3.77838	3.69276	3.6761	3.66966	3.59721	3.55671	3.50177	3.45717
c1	-1.37519	-1.43327	-1.46134	-1.41441	-1.29344	-1.18195	-1.16549	-1.1752	-1.14479	-1.1452	-1.1308	-1.11631
c2	0.218095	0.238839	0.225844	0.208667	0.16355	0.119101	0.102609	0.099164	0.095008	0.094317	0.100456	0.101994
h	5.90684	5.79394	6.62019	6.89248	6.71735	5.78659	5.45192	5.40732	5.02064	5.08066	4.95777	4.69877
c3	0.00071	0.000685	0.001176	0.001602	0.002029	0.002123	0.001654	0.001248	0.000919	0.000673	0.000583	0.000509
b1	-0.02683	-0.05688	-0.04305	-0.05845	-0.03586	0.067202	0.129716	0.145499	0.168179	0.173884	0.190813	0.203522
b2	-0.0726	-0.06373	-0.04979	-0.06443	-0.09154	-0.09151	-0.09751	-0.10488	-0.11422	-0.12015	-0.12318	-0.12608
b3	0	0	0	0	0.085537	0.145251	0.135986	0.135159	0.149582	0.151849	0.130847	0.122339
ClassA	0	0	0	0	0	0	0	0	0	0	0	0
ClassB	0.134904	0.133973	0.139714	0.155236	0.158937	0.138968	0.126737	0.113881	0.109638	0.110223	0.108079	0.108783
ClassC	0.226827	0.218136	0.206862	0.210168	0.199726	0.216584	0.249141	0.259274	0.274211	0.280836	0.298022	0.305295
ClassD	0.213357	0.176183	0.145621	0.156052	0.186495	0.1995	0.229736	0.252504	0.282686	0.301657	0.34708	0.370989
sofN	-0.02809	-0.03866	-0.03889	-0.01955	-0.02056	0.018953	0.023563	0.018438	0.012675	0.02215	0.017165	0.016712
sofR	0.077532	0.060308	0.07126	0.084246	0.074269	0.133352	0.143428	0.138662	0.122472	0.129181	0.115968	0.114252
sofS	-0.02064	-0.0334	-0.02736	-0.02283	-0.02673	0.026665	0.039234	0.043489	0.036662	0.046123	0.044778	0.049822
sofU	0	0	0	0	0	0	0	0	0	0	0	0
τ	0.182533	0.18063	0.194176	0.181926	0.18138	0.177903	0.178211	0.184254	0.184085	0.191734	0.19969	0.200063
φ	0.278823	0.289652	0.296609	0.306918	0.305998	0.300131	0.300652	0.295463	0.295192	0.292878	0.291096	0.29164
φ _{S2S}	0.208393	0.220859	0.235714	0.244969	0.241833	0.219913	0.200662	0.193285	0.187569	0.180758	0.182941	0.175988
σ	0.333258	0.341358	0.354515	0.356785	0.355716	0.348896	0.349501	0.348207	0.347887	0.350056	0.353006	0.353665

Table 1 continued

T[see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV
e1	3.38799	3.34381	3.25802	3.16899	3.14649	2.89515	2.76366	2.63662	2.6215	2.46318	2.3968	3.45078
c1	-1.1047	-1.11609	-1.10907	-1.08714	-1.09387	-1.03042	-1.01437	-1.04838	-1.0543	-1.07308	-1.05706	-1.36061
c2	0.104529	0.099889	0.119754	0.1117879	0.114285	0.136666	0.1441	0.180838	0.181367	0.226407	0.248126	0.215873
h	4.54643	4.64017	4.63849	4.50481	4.53118	4.53208	4.61172	5.39607	5.56772	6.23491	6.7674	6.14717
c3	0.000249	0	0	0	0	0	0	0	0	0	0	0.000733
b1	0.242603	0.280922	0.291242	0.311362	0.359324	0.393471	0.432513	0.434162	0.458752	0.475305	0.48108	-0.02087
b2	-0.12601	-0.12461	-0.12126	-0.12373	-0.11774	-0.11544	-0.1043	-0.0963	-0.09558	-0.07881	-0.07197	-0.07224
b3	0.095965	0.092048	0.032748	0.052576	0.044584	0	0	0	0	0	0	0
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.106929	0.102965	0.097481	0.087057	0.086496	0.092091	0.103385	0.107251	0.099358	0.105913	0.127642	0.137715
Class C	0.321296	0.331801	0.341281	0.342803	0.34521	0.345292	0.342842	0.333706	0.329709	0.312454	0.318684	0.233048
Class D	0.4440581	0.503562	0.542709	0.581633	0.590175	0.618805	0.653192	0.618956	0.604177	0.577657	0.597588	0.214227
SoFN	0.013695	0.024399	0.024483	0.042376	0.053679	0.087972	0.123393	0.161886	0.139794	0.125695	0.052424	-0.03228
SoFR	0.100223	0.092189	0.078739	0.091254	0.091382	0.119863	0.165217	0.193198	0.167929	0.153396	0.047119	0.073678
SoFS	0.042018	0.049609	0.049226	0.068452	0.067455	0.100768	0.143638	0.201695	0.185814	0.173281	0.116645	-0.01943
SoFU	0	0	0	0	0	0	0	0	0	0	0	0
τ	0.207756	0.208828	0.211136	0.220213	0.221524	0.222493	0.218105	0.212905	0.22224	0.223041	0.236576	0.180904
ϕ	0.289459	0.290952	0.294168	0.293618	0.295365	0.296657	0.303878	0.31036	0.309638	0.310755	0.302186	0.276335
ϕ	0.176453	0.178954	0.18031	0.194549	0.196091	0.196817	0.19849	0.201126	0.202676	0.20708	0.21241	0.206288
σ	0.356299	0.358137	0.362096	0.367022	0.369206	0.370822	0.374047	0.376367	0.381138	0.382513	0.383777	0.330284

Acceleration is in (cm/s²), velocity in (cm/s). The symbols τ , ϕ , ϕ_{S2S} , and σ stand for between-events, within-event, site-to-site and total standard deviations. Sof indicates the Style of Faulting (*N* normal, *R* reverse, *S*:strike slip, *U* unknown). EC8 ground categories are indicated as Class A, B, C, D

Table 2 Coefficients for the model derived in this study (see Eqs. 1–3) for R_{fB} and V_{s30} classification

T[see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	3.37053	3.43922	3.59651	3.68638	3.68632	3.68262	3.64314	3.63985	3.5748	3.53006	3.43387	3.40554
c1	-1.26358	-1.31025	-1.29051	-1.28178	-1.17697	-1.10301	-1.08527	-1.10591	-1.09955	-1.09538	-1.06586	-1.05767
c2	0.220527	0.244676	0.231878	0.219406	0.182662	0.133154	0.115603	0.108276	0.103083	0.101111	0.109066	0.112197
h	5.20082	4.91669	5.35922	6.12146	5.74154	5.31998	5.13455	5.12846	4.90557	4.95386	4.6599	4.43205
c3	0.001118	0.001092	0.001821	0.002114	0.00254	0.002421	0.001964	0.001499	0.001049	0.000851	0.000868	0.000789
b1	-0.08906	-0.11692	-0.08501	-0.11355	-0.09287	0.010086	0.02994	0.03919	0.052103	0.045846	0.060084	0.088319
b2	-0.09162	-0.07838	-0.057	-0.07533	-0.10243	-0.10518	-0.12717	-0.13858	-0.15139	-0.16209	-0.1659	-0.16411
b3	0	0	0	0	0.073904	0.150461	0.178899	0.189682	0.216011	0.224827	0.197716	0.15475
γ	-0.29402	-0.24177	-0.20763	-0.17324	-0.20249	-0.29123	-0.35443	-0.39306	-0.45391	-0.49206	-0.56446	-0.5962
sofN	-0.03924	-0.03772	-0.04594	-0.03805	-0.02673	-0.03265	-0.03384	-0.03725	-0.02791	-0.02563	-0.01866	-0.01742
sofR	0.081052	0.079778	0.087497	0.08471	0.067844	0.075977	0.074982	0.076701	0.06979	0.072567	0.064599	0.060283
sofS	-0.04182	-0.04206	-0.04155	-0.04666	-0.04111	-0.04332	-0.04114	-0.03946	-0.04188	-0.04694	-0.04594	-0.04286
τ	0.15867	0.154621	0.172785	0.169691	0.152902	0.150055	0.151209	0.157946	0.165436	0.157728	0.173005	0.18082
ϕ	0.282356	0.291143	0.291499	0.301967	0.305804	0.300109	0.302419	0.297402	0.294395	0.296992	0.291868	0.289957
ϕ_{S2S}	0.183959	0.187409	0.199913	0.208178	0.212124	0.190469	0.187037	0.174118	0.175848	0.169883	0.164162	0.16509
σ	0.323885	0.329654	0.33886	0.346379	0.3419	0.335532	0.338114	0.336741	0.337694	0.336278	0.33929	0.341717

Table 2 continued

Tl[sec]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV
e1	3.30442	3.23882	3.1537	3.13481	3.12474	2.89841	2.84727	2.68016	2.60171	2.39067	2.25399	3.32819
c1	-1.05014	-1.05021	-1.04654	-1.04612	-1.0527	-0.97383	-0.98339	-0.98308	-0.97922	-0.97753	-0.94037	-1.2398
c2	0.121734	0.114674	0.129522	0.114536	0.103471	0.104898	0.109072	0.164027	0.163344	0.211831	0.227241	0.21732
h	4.21657	4.17127	4.20016	4.48003	4.41613	4.25821	4.56697	4.68008	4.58186	5.39517	5.74173	5.26486
c3	0.000487	0.000159	0	0	0	0	0	0	0	0	0	0.001186
b1	0.120182	0.166933	0.193817	0.247547	0.306569	0.349119	0.384546	0.343663	0.331747	0.357514	0.385526	-0.0855
b2	-0.16333	-0.16111	-0.15655	-0.15382	-0.14756	-0.14948	-0.13987	-0.13593	-0.14828	-0.12254	-0.11145	-0.09256
b3	0.117576	0.112005	0.051729	0.081575	0.092837	0.108209	0.098737	0	0	0	0	0
γ	-0.66782	-0.73839	-0.79408	-0.8217	-0.82658	-0.84505	-0.8232	-0.77866	-0.76924	-0.76961	-0.73207	-0.3019
SoFn	-0.00049	0.011203	0.016526	0.016449	0.026307	0.025234	0.018674	0.011371	0.005535	0.008735	0.022989	-0.03977
SoFR	0.044921	0.028151	0.020352	0.021242	0.018604	0.022362	0.023089	0.016688	0.019857	0.023314	-0.02066	0.077525
SoFS	-0.04443	-0.03935	-0.03688	-0.03769	-0.04491	-0.0476	-0.04176	-0.02806	-0.02539	-0.03205	-0.00233	-0.03776
τ	0.182233	0.189396	0.189074	0.191986	0.195026	0.181782	0.177752	0.163242	0.164958	0.17028	0.176546	0.149977
ϕ	0.292223	0.289307	0.288815	0.293264	0.297907	0.306676	0.316312	0.326484	0.329916	0.320626	0.314165	0.282398
ϕ	0.175634	0.168617	0.16817	0.183719	0.200775	0.209625	0.218569	0.221367	0.22535	0.210193	0.207247	0.165611
σ	0.344388	0.345788	0.3452	0.350517	0.356067	0.356504	0.362835	0.36502	0.368857	0.363037	0.360373	0.319753

Acceleration is in (cm/s²), velocity in (cm/s). The symbols τ , ϕ , ϕ_{25} , and σ stand for between-events, within-event, site-to-site and total standard deviations. Sof indicates the Style of Faulting (*N* normal, *R* reverse, *S* strike slip). The site coefficient is indicated as γ

Table 3 Coefficients for the model derived in this study (see Eqs. 1–3) for R_{HYP0} and EC8 ground categories

T[see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	4.42044	4.54992	4.73285	4.67503	4.56965	4.45017	4.45593	4.47171	4.38799	4.37609	4.33372	4.29359
c1	-1.77754	-1.8546	-1.87822	-1.79917	-1.61405	-1.46501	-1.44342	-1.46016	-1.41842	-1.42843	-1.42503	-1.41465
c2	0.147715	0.165968	0.157048	0.151808	0.105601	0.056755	0.032061	0.025927	0.02215	0.016902	0.025903	0.028368
h	7.06428	6.98227	8.1337	8.38098	7.49625	6.27222	5.4804	5.50316	4.76952	4.81974	5.10961	4.95519
c3	0	0	0	0.000548	0.001183	0.001431	0.000982	0.000554	0.000269	0	0	0
b1	0.147874	0.124402	0.138028	0.098832	0.125747	0.236642	0.313239	0.332549	0.355357	0.368987	0.379142	0.38941
b2	-0.06621	-0.0566	-0.04079	-0.05694	-0.0835	-0.08346	-0.08972	-0.09722	-0.10604	-0.11196	-0.11515	-0.11815
b3	0.29709	0.260601	0.27609	0.322027	0.464456	0.542025	0.555789	0.551296	0.543724	0.547881	0.511833	0.495459
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.14111	0.14035	0.145543	0.158622	0.162534	0.143446	0.133443	0.121637	0.118062	0.119481	0.117659	0.118871
Class C	0.225339	0.21701	0.206101	0.208849	0.197589	0.213637	0.244854	0.254554	0.268087	0.275041	0.291964	0.29887
Class D	0.187033	0.146507	0.115846	0.125428	0.158161	0.170195	0.202162	0.226009	0.258058	0.275672	0.321124	0.344584
sofN	-0.06531	-0.06538	-0.05129	-0.03749	-0.04709	-0.02145	-0.03049	-0.04227	-0.05667	-0.05327	-0.06251	-0.06474
sofR	0.091732	0.088098	0.113143	0.120065	0.098046	0.139454	0.132769	0.119803	0.092863	0.09198	0.073772	0.069449
sofS	-0.05613	-0.05767	-0.03762	-0.0369	-0.05061	-0.01246	-0.01516	-0.01923	-0.03496	-0.03219	-0.03929	-0.03741
sofU	0	0	0	0	0	0	0	0	0	0	0	0
τ	0.197407	0.204345	0.208843	0.19539	0.193856	0.191231	0.192222	0.199096	0.199491	0.207716	0.216313	0.225415
φ	0.287767	0.297881	0.304438	0.31332	0.310861	0.306652	0.308241	0.304125	0.304728	0.302796	0.30138	0.300553
φ _{S2S}	0.216309	0.222929	0.242821	0.251339	0.247987	0.226544	0.214042	0.207111	0.201784	0.194828	0.197633	0.198934
σ	0.348969	0.361234	0.369185	0.369252	0.366353	0.361392	0.363266	0.363499	0.36422	0.367194	0.370974	0.375691

Table 3 continued

T[see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV _y
e1	4.23915	4.19696	4.11453	4.03249	4.0114	3.68402	3.53587	3.46588	3.4691	3.28384	3.2647	4.36693
c1	-1.40603	-1.41297	-1.40429	-1.38977	-1.39543	-1.30231	-1.27351	-1.36102	-1.38111	-1.38977	-1.39974	-1.75212
c2	0.02698	0.020876	0.038146	0.037094	0.034061	0.069535	0.082246	0.137018	0.137878	0.188643	0.216533	0.150507
h	4.63597	4.29377	4.01059	3.97812	4.09668	3.7329	4.07408	6.0971	6.53917	7.04011	8.33921	7.32192
c3	0	0	0	0	0	0	0	0	0	0	0	0
b1	0.430341	0.470648	0.481962	0.504043	0.550001	0.544404	0.570581	0.524014	0.551312	0.547984	0.552993	0.144291
b2	-0.11928	-0.1181	-0.11674	-0.11665	-0.11086	-0.11362	-0.10376	-0.10109	-0.09877	-0.08423	-0.07134	-0.06608
b3	0.475308	0.460014	0.393948	0.400442	0.386023	0.282169	0.24976	0.046975	0	0	0	0.284211
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.117717	0.115734	0.110981	0.103765	0.103026	0.108865	0.119032	0.123814	0.115091	0.124833	0.143969	0.143778
Class C	0.314097	0.325887	0.334461	0.334934	0.336196	0.337519	0.33311	0.323505	0.320404	0.306133	0.315187	0.231064
Class D	0.412316	0.477053	0.51753	0.559004	0.566463	0.592894	0.626267	0.60053	0.586654	0.548523	0.559213	0.187402
SoFN	-0.07608	-0.07496	-0.08163	-0.06429	-0.05717	-0.03466	-0.01067	-0.00297	-0.0238	-0.05066	-0.14667	-0.07175
SoFR	0.045871	0.028575	0.008429	0.019498	0.014893	0.029824	0.060267	0.058459	0.034964	0.003435	-0.12866	0.084958
SoFS	-0.05488	-0.05564	-0.06343	-0.04562	-0.05139	-0.02508	0.007386	0.039471	0.02527	0.007396	-0.06757	-0.0571
SoFU	0	0	0	0	0	0	0	0	0	0	0	0
τ	0.234484	0.246498	0.249844	0.261433	0.274446	0.26531	0.269363	0.27539	0.277179	0.278908	0.283885	0.195249
ϕ	0.299514	0.301897	0.305995	0.30722	0.309616	0.311777	0.316539	0.323622	0.325724	0.327756	0.320266	0.284622
ϕ_{S2S}	0.208675	0.212696	0.224068	0.240384	0.244465	0.244067	0.236824	0.257636	0.259839	0.263531	0.267078	0.213455
σ	0.380383	0.389747	0.395038	0.403399	0.413742	0.409383	0.415637	0.424936	0.427696	0.430364	0.427973	0.345155

Acceleration is in (cm/s²), velocity in (cm/s). The symbols τ , ϕ , ϕ_{S2S} , and σ stand for between-events, within-event, site-to-site and total standard deviations

Table 4 Coefficients for the model derived in this study (see Eqs. 1–3) for R_{HYPO} and Vs30 classification

T[sec]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	4.3397	4.46839	4.5724	4.55255	4.51119	4.49571	4.49224	4.51726	4.46559	4.46834	4.3715	4.34198
c1	-1.60402	-1.68536	-1.63863	-1.57947	-1.4471	-1.37039	-1.36679	-1.40078	-1.40973	-1.42893	-1.40655	-1.39751
c2	0.103401	0.126703	0.123954	0.125609	0.08461	0.038536	0.012937	0.00198	0.000489	-0.0091	0.00101	0.004238
h	4.47852	4.58063	5.12096	5.67511	4.8248	4.56965	3.94802	4.26816	4.39978	4.6039	4.60254	4.43045
c3	2.63E-05	0	0.000722	0.001239	0.001692	0.001586	0.001059	0.000565	5.97E-05	0	0	0
b1	0.230422	0.205651	0.226272	0.167382	0.194714	0.289627	0.321065	0.336096	0.346351	0.353351	0.35717	0.384532
b2	-0.06654	-0.05281	-0.0298	-0.05091	-0.07845	-0.08155	-0.10418	-0.11526	-0.12711	-0.13778	-0.14277	-0.14092
b3	0.363906	0.323734	0.311109	0.348968	0.448903	0.533244	0.596455	0.612107	0.600314	0.621323	0.589127	0.543301
γ	-0.286524	-0.232462	-0.195629	-0.168432	-0.194539	-0.270912	-0.323355	-0.363199	-0.430464	-0.467397	-0.531694	-0.555531
sofN	-0.04692	-0.04517	-0.05321	-0.04704	-0.03631	-0.03868	-0.03658	-0.03807	-0.02853	-0.02616	-0.01928	-0.01758
sofR	0.115063	0.114597	0.121653	0.119021	0.102481	0.107555	0.103236	0.104818	0.095509	0.097198	0.090202	0.086012
sofS	-0.06814	-0.06943	-0.06845	-0.07198	-0.06617	-0.06888	-0.06666	-0.06675	-0.06697	-0.07104	-0.07092	-0.06843
τ	0.154538	0.158402	0.169775	0.165148	0.145533	0.144701	0.156869	0.165195	0.164907	0.165146	0.181401	0.189686
ϕ	0.290986	0.298261	0.302117	0.310963	0.310621	0.308845	0.313737	0.311052	0.310509	0.310959	0.306033	0.304174
ϕ_{S2S}	0.18825	0.192664	0.203229	0.212643	0.216313	0.20204	0.199484	0.186722	0.180734	0.182064	0.176797	0.178065
σ	0.329477	0.337714	0.346552	0.352097	0.343023	0.341063	0.350769	0.352197	0.351583	0.352092	0.355756	0.358473

Table 4 continued

T[see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA _y	PGV _y
e1	-1.37164	4.14832	4.09246	4.08324	4.07207	3.77954	3.69447	3.45408	3.38901	3.06601	2.89391	4.27391
c1	17.7384	-1.37169	-1.37736	-1.38649	-1.38735	-1.27343	-1.26477	-1.27364	-1.28283	-1.23427	-1.16461	-1.57821
c2	0.216704	0.002264	0.008956	-0.00453	-0.01855	-0.01377	-0.00337	0.083746	0.086724	0.150146	0.162354	0.108218
h	886.652	3.00978	3.15727	3.4537	3.3163	3.04976	3.65482	4.59988	4.95285	4.45511	4.62321	4.82743
c3	0.05606	0	0	0	0	0	0	0	0	0	0	9.64E-05
b1	-0.25583	0.466754	0.510102	0.567727	0.631338	0.650829	0.6746	0.563304	0.548353	0.54175	0.590765	0.217109
b2	-0.12108	-0.13807	-0.13263	-0.12724	-0.12124	-0.12901	-0.11908	-0.1178	-0.12957	-0.1037	-0.08533	-0.06826
b3	0	0.498126	0.437529	0.45811	0.474982	0.488244	0.461122	0.184126	0.171017	0.009303	0.034058	0.352976
γ	-0.457888	-0.698998	-0.757522	-0.786632	-0.791438	-0.803656	-0.780198	-0.749008	-0.744073	-0.744468	-0.693999	-0.293242
SoFN	0.022367	0.010003	0.015018	0.01638	0.026396	0.024922	0.019123	0.011676	0.004993	0.006027	0.018621	-0.04721
SoFR	0.125552	0.054388	0.045865	0.044224	0.041137	0.038329	0.038697	0.029249	0.033587	0.030508	-0.01898	0.110979
SoFS	-0.14793	-0.06439	-0.06088	-0.0606	-0.06753	-0.06325	-0.05782	-0.04092	-0.03858	-0.03653	0.000361	-0.06376
τ	0.259955	0.20181	0.211664	0.225279	0.238973	0.212162	0.208441	0.203238	0.205751	0.190711	0.183363	0.145783
ϕ	0.397088	0.30827	0.30855	0.313873	0.318631	0.324083	0.33425	0.342873	0.347114	0.339373	0.326297	0.291566
ϕ_{S2S}	0.189183	0.264361	0.208994	0.225906	0.246861	0.245588	0.24415	0.256308	0.26183	0.242015	0.22865	0.186662
σ	0.474611	0.368453	0.374172	0.386351	0.398289	0.387354	0.393917	0.398582	0.403511	0.389288	0.374289	0.325981

Acceleration is in (cm/s²), velocity in (cm/s). The symbols τ , ϕ , ϕ_{S2S} , and σ stand for between-events, within-event, site-to-site and total standard deviations

170 4.1 Influence of the dataset

171 Figures 4a, b show the comparison between ground-motion predictions obtained with the DS-
 172 EC8 or the DS-VS30 datasets. In particular, Fig. 4a exemplifies the comparison for class B of
 173 EC8, which is sampled by stations belonging to different countries (Fig. 2) and considering
 174 $V_{S30} = 580$ m/s. No significant differences can be appreciated both for the median and the
 175 total standard deviation.

176 Figure 4b shows the dependence on period of the total (σ), between-events (τ), within-
 177 event (ϕ) and site-to-site (ϕ_{S2S}) standard deviations, for the two data sets. For DS-EC8, sigma
 178 increases from 0.33 at 0.02 s to 0.38 at 3 s, while only a modest reduction of sigma is observed
 179 when DS-VS30 is considered (sigma ranges from 0.32 to 0.36 at 0.02 and 3 s, respectively),
 180 although DS-VS30 includes waveforms from stations characterized by measured V_{S30} . To
 181 investigate the reasons of such limited improvement, Fig. 4b also compares the other compo-
 182 nents of variability. The between-events component τ , obtained for DS-VS30, is smaller than
 183 0.2 over the analyzed period range, while, for DS-EC8, τ is lower than 0.2 only for periods
 184 shorter than 0.5 s. The within-event component ϕ is significantly larger than τ and similar for
 185 the two data sets, except for periods longer than 1 s, where ϕ associated to DS-EC8 is smaller
 186 than the one associated to DS-VS30. This feature can be ascribed to the site-to-site standard
 187 deviation (ϕ_{S2S}). In fact, the ϕ_{S2S} evaluated using a continuous function of V_{S30} is smaller
 188 than the one obtained from EC8 site categories in the short periods range (< 1 s) while, for
 189 periods longer than 1 s, it assumes larger values. Since stations belonging to EC8 classes
 190 C and D usually exhibit large amplification above 1 s, the increase of ϕ_{S2S} at long periods
 191 suggests that V_{S30} is not a good proxy to capture the site effects for soft sites included in
 192 RESORCE.

193 4.2 Influence of distance metrics

194 In Fig. 5 the influence of the use of hypocentral or Joyner-Boore distances is shown for Mw
 195 7.5 and 4, and considering EC8 class A. The two magnitudes correspond to the limits of
 196 applicability of the model. The comparison is performed for PGA and PSA at 1 s. The two
 197 models predict similar values for low magnitudes ($M_w = 4$) at all distances. Since the point
 198 source approximation can be applied, for small magnitudes R_{JB} is similar to the epicentral
 199 distance and the difference between R_{JB} and R_{hypo} is related to the focal depth. For short
 200 distances and large magnitudes the model based on R_{hypo} predicts larger values than the one
 201 based on R_{JB} , as the difference in the definition of the two metrics, can cause R_{JB} close to
 202 zero, even for considerable epicentral (and hence hypocentral) distances.

203 In the following, since the models based on DS-EC8 or DS-VS30 are similar in terms of
 204 median predictions and standard deviation, we limit the discussion to the DS-EC8 dataset
 205 and the Joyner-Boore distance.

206 4.3 Coefficients of GMPEs

207 Figure 6 shows that nearly all coefficients have a significant dependence on period. The
 208 coefficient b_3 , controlling the magnitude-dependence over the hinge-magnitude, is positive
 209 in the period range 0.15–1.5 s, although these values are not significantly different from zero
 210 at a 5% significance level. The coefficient c_3 , relative to the linear attenuation with distance
 211 is significantly different from zero, at a 5% significance level only for $0.04 < T < 0.4$ s.
 212 The pseudo-depth parameter h varies from 4.5 to 6.9 km, with an average of 5.5 km.

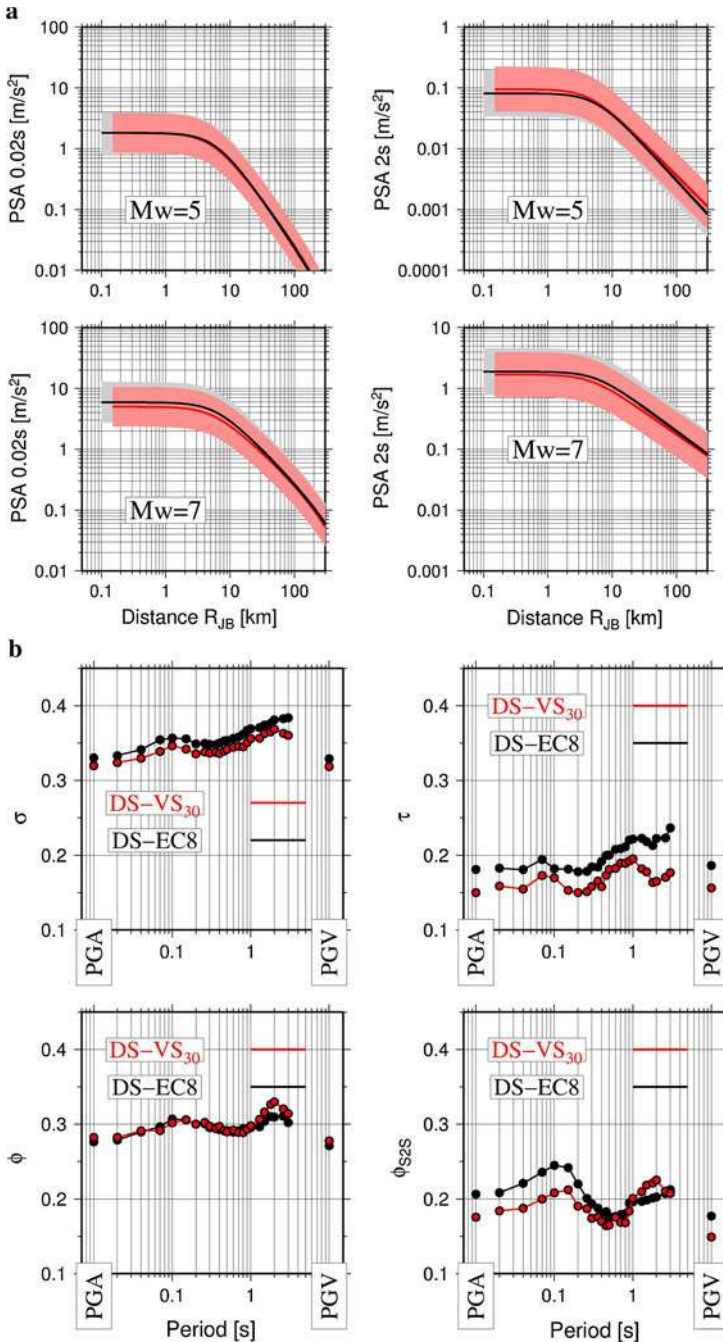
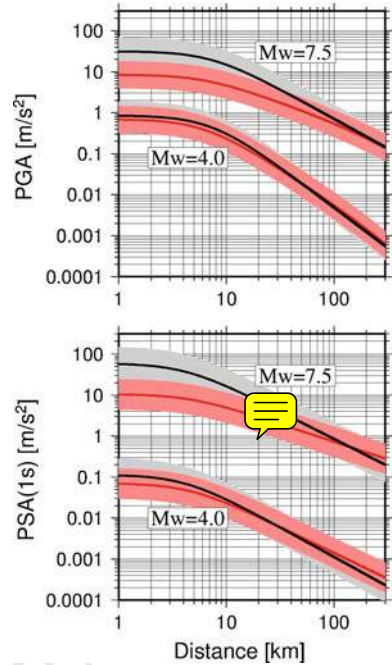


Fig. 4 **a** Comparisons between the ground motion predictions obtained with the two datasets (red color is DS-Vs₃₀ and grey color is DS-EC8). Left panel T = 0.02 s, right panel T = 2 s. The curves are relative to normal fault and Ec8 class B (Vs₃₀ = 580 m/s). **b**. Standard deviations obtained for the two datasets (top left: total sigma; top right: between-events standard deviation; bottom left: within-event standard deviation; bottom right: site-to-site standard deviation)

Fig. 5 Comparison between models based on Joyner-Boore (red) or hypocentral distances (grey scale) for PGA (upper panel) and PSA at $T = 1$ s (lower panel) for DS-EC8 dataset, considering class C and strike slip faulting



213 The site coefficients are shown in Fig. 7. Site category B (i.e. stiff sites) amplifies the entire
 214 period range with values between 0.10 and 0.15 log10 units; class C shows an almost constant
 215 amplification of about 0.21, for periods up to 0.2 s, and a peak of amplification of about 0.34
 216 at 1 s; class D has a relevant amplification (up to 0.62) for periods longer than 1 s. Class D
 217 is represented by 31 stations but only 5 are characterized by at least 5 recordings, namely:
 218 Bevagna (BVG, Italian station with 7 records analyzed in this study at 1 s); Colfiorito (CLF,
 219 Italy, 14 records at 1 s); Norcia (NOR, Italy, 5 records at 1 s); Rieti (RTI, Italy, 7 records);
 220 Ambarli (ATS, Turkey, 6 records). Therefore, the above mentioned stations strongly control
 221 the amplification coefficient of the entire class (Bindi et al. 2011b). In fact, the amplification
 222 peak of class D (Fig. 7) reflects the long period (> 1 s) site amplifications found by Rovelli
 223 et al. (2001) for CLF; Bindi et al. (2011c) for station NOR; Foti et al. (2011), for RTI; Luzi
 224 et al. (2005) for BVG. Figure 7 also shows the style of faulting coefficients, indicating that
 225 reverse faulting causes amplitudes higher than strike-slip and normal faulting, at short periods
 226 (0–1s), while normal and strike slip coefficients show similar trend.

227 Regarding the propagation of errors from data to solutions, Fig. 8 (top panel) exemplifies
 228 the unit covariance matrix (Menke 1989), computed for $T = 0.1$ s. The parameters most
 229 affected by the amplification of error from data to solutions are e_1 , h and b_3 (i.e., those
 230 having the largest diagonal elements), that also show significant trade-offs with the other
 231 parameters. These results are in agreement with the data distribution shown in Fig. 1, relative
 232 to the sparse sampling at short distances (controlling e_1 and h) and large magnitudes
 233 (controlling b_3). Figure 6 (top panel) illustrates the trade-off of coefficients e_1 (C_{1j} elements)
 234 and h (C_{4j} elements). The off-set coefficient e_1 shows positive trade-off with pseudo-depth
 235 h ($C_{14} = 0.597$ at $T = 0.1$ s), with b_3 ($C_{16} = 0.036$) and negative trade-off with c_1 ($C_{12} =$
 236 -0.0597), which, in turn, shows a negative trade-off with the pseudo-depth parameter
 237 ($C_{42} = -0.434$).

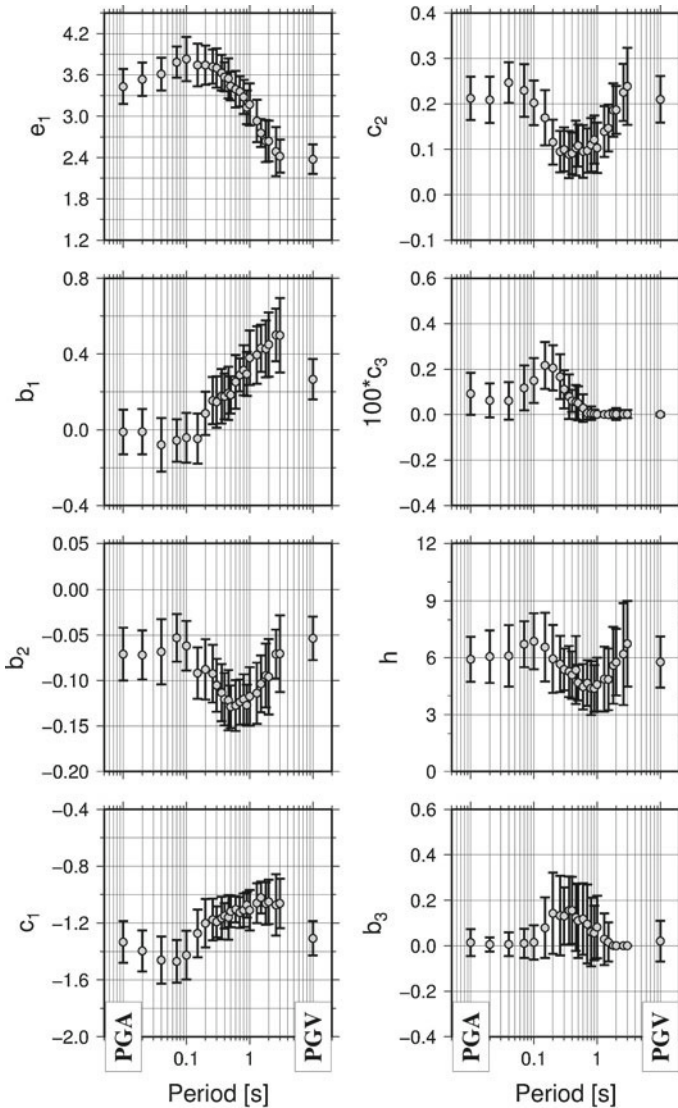


Fig. 6 Mean and 95% confidence interval versus period of the coefficients relevant to the model described in Eqs. 1–3 (DS-EC8 data set). PGA is reported at $T = 0.01$ s (first point the left), while the Peak Ground Velocity (PGV) at $T = 10$ s (last point on the right)

238 The constraint applied to the site amplification for class A removes the trade-off between
 239 the offset coefficient e_1 and the site coefficients while a weak trade-off among the style
 240 of faulting coefficients and e_1 still persists. Finally, all the entries of the covariance matrix
 241 show a weak-dependence on period (Fig. 8 middle and bottom panels) although C_{14} and
 242 C_{44} , have a sharp increase for periods longer than 1 s. At long periods, the trade-off between
 243 the pseudo-depth parameter h and the off-set parameter e_1 (C_{14}) increases and h is poorly
 244 constrained (C_{44}). An explanation for the increase of this trade-off could be the decreasing
 245 number of recordings due to the filter selection (Fig. 3).

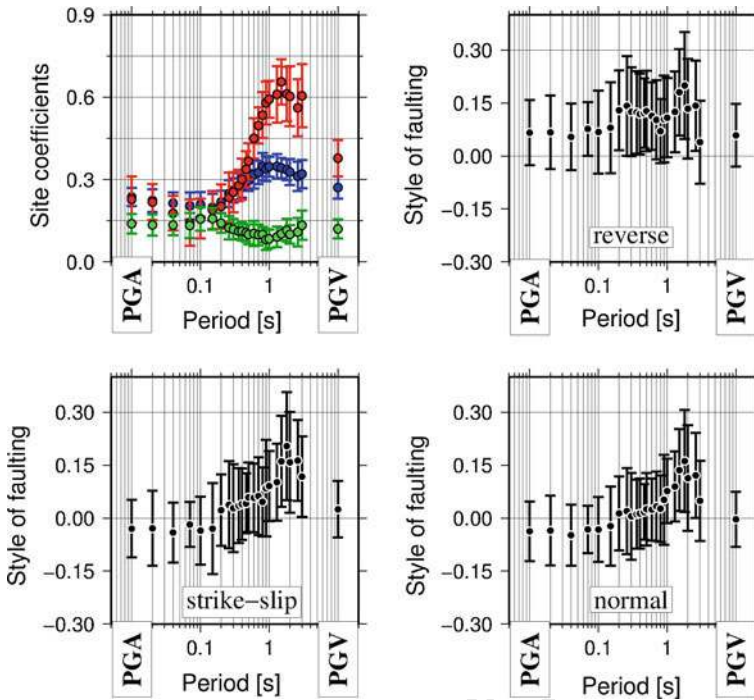


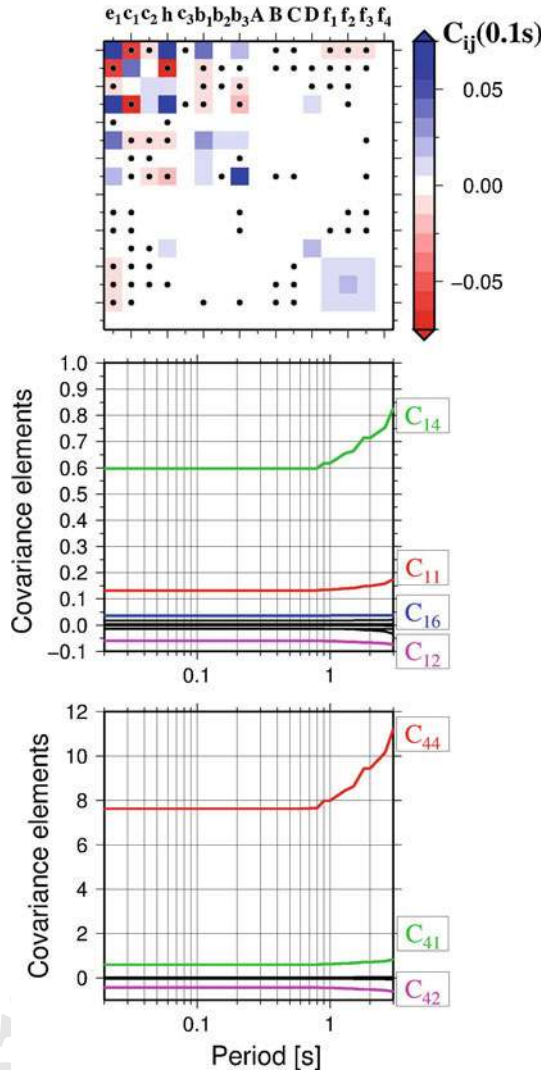
Fig. 7 Top Left: Site coefficients for EC8 classes B through D versus period (green: class B; blue: class C; red: class D). Top Right: Style-of-faulting coefficients versus period for reverse faulting. Bottom: style of faulting for strike slip (left) and normal (right) faulting. Peak Ground Acceleration (PGA) is reported at $T = 0.01$ s (first point the left), while the Peak Ground Velocity (PGV) at $T = 10$ s (last point on the right)

246 **5 Ground-motion variability**

247 We carry out the residuals analysis with the aim of investigating the possible origins of
 248 uncertainties not captured by the GMPEs, and to check for any dependence on the primary
 249 explanatory variables. Figure 9 shows, at fixed periods of 0.02 s (left panels) and 2.0 s (right
 250 panels): (i) the residuals (log10 of the observed—predicted values), (ii) the between-events
 251 errors in function of magnitude (second row), (iii) the site-to-site errors in function of the
 252 EC8 ground categories (third row) and (iv) the record-to-record errors (that is, the residuals
 253 corrected for the between-events and site-to-site errors) in function of R_{JB} distance (fourth
 254 row).

255 To investigate the influence of the earthquake type on the predicted ground-motion, the
 256 total residuals (Fig. 9, first row) relative to aftershocks and mainshocks are displayed with
 257 different colors. Since it is difficult to find objective criteria to separate mainshocks and
 258 aftershocks in the European datasets (see also Douglas and Halldórsson 2010), we apply
 259 the Gardner and Knopoff (1974) approach, and we exploit the information available in the
 260 metadata. The histograms of the residuals for aftershocks and mainshocks are shown in Fig. 10
 261 (top panels), indicating that the distributions are almost unbiased for both type of earthquakes,
 262 and the normal distributions that best fit the histograms have nearly the same sigma (e.g.,
 263 0.369 and 0.364 for mainshocks and aftershocks, respectively, at $T = 2$ s). Considering the
 264 between-events errors distribution (bottom panels of Fig. 10), a slightly smaller standard

Fig. 8 *Top panel:* unit covariance matrix computed for the final model at $T = 0.1$ s (blue: positive entries, red: negative entries; black dots also indicate negative entries). *Middle panel:* entries in the columns of the covariance matrix relevant to the parameter e_1 (first row or first column) against period. *Bottom panel:* entries in the columns of the covariance matrix relevant to the parameter h (fourth row or fourth column)



265 deviation is observed for the mainshock distribution with respect to the aftershock one (0.169
 266 against 0.181 at 2s; 0.130 against 0.154 at 0.02s) together with slightly larger bias for
 267 mainshocks (0.0332 against 0.0147 at 2s). The small differences in the statistical parameters
 268 of the normal distributions shown in Fig. 10 suggest that neglecting the classification of the
 269 earthquake type in the predictive model, introduces a negligible bias in the residuals, with a
 270 weak tendency of underestimating the ground-motion at long periods for mainshocks.

271 The plot of the between-events errors in function of magnitude (Fig. 9, second row), shows
 272 an increase of dispersion for small magnitude and long periods, although the low dispersion
 273 at large magnitude could be an apparent effect related to the poor sampling. The large positive
 274 between-events error at $T = 2$ s is relevant to Mw 6.2, 2004 Baladeh (Iran) earthquake (e.g.,
 275 Tatar et al. 2007; Ghasemi et al. 2008), whose processing over long periods should be revised.
 276 The top panel of Fig. 11 presents the between-events standard deviation τ computed grouping

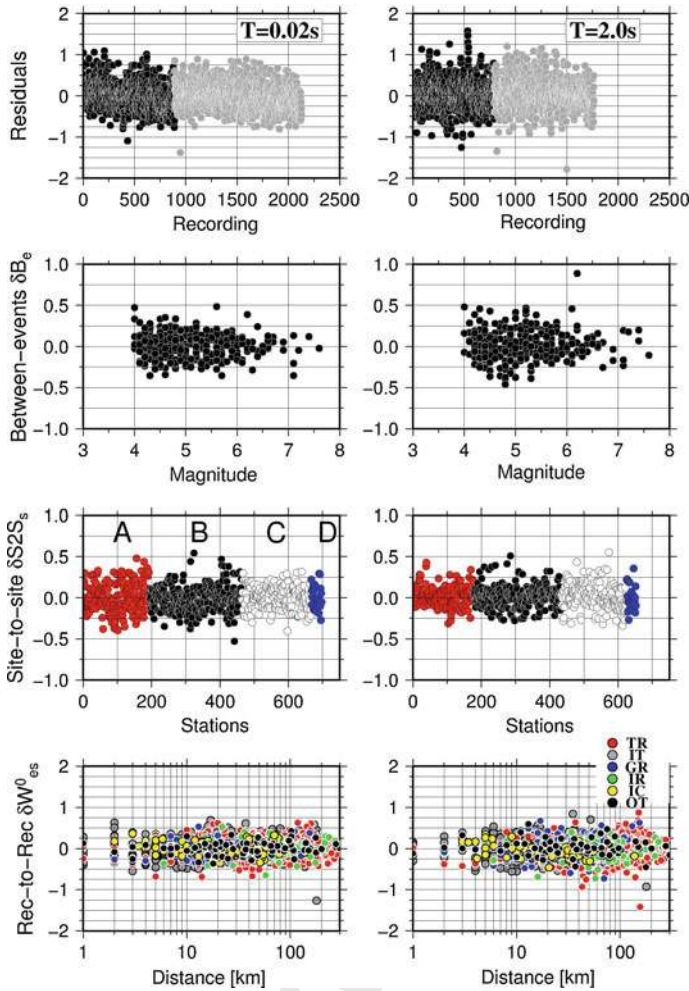


Fig. 9 *Left panels are $T = 0.02\text{ s}$ and right panels are $T = 2.0\text{ s}$. First row: distribution of residuals (computed as \log_{10} of observations over predictions), where the *black* and *gray dots* indicate mainshocks and aftershocks, respectively; *second row*: between-events residuals in function of magnitude; *third row*: site-to-site residuals (*red*: class A *black*: class B *white*: class C *blue*: class D); *fourth row*: record-to-record residuals (the colours indicate different recording networks)*

277 the earthquakes into three magnitude ranges ($M \leq 5$; $5 < M \leq 6$; $M > 6$), without
 278 considering the Baladeh earthquake. While the standard deviations are almost the same for
 279 short period ($T < 0.2\text{ s}$), at longer periods ($T > 1\text{ s}$) τ is significantly larger for small
 280 magnitudes ($M < 5$), with respect to the well sampled magnitude range $5 < M < 6$. The
 281 dependence of τ on magnitude suggests that a heteroscedastic model, including a magnitude
 282 dependent sigma, should be considered in deriving GMPEs, although it would be necessary
 283 an increase of sampling of large earthquakes.

284 The site-to-site errors are shown according to their EC8 site classes (Fig. 9, third row). A
 285 larger dispersion of class A at short periods and of classes B and C at long periods is observed.
 286 This trend is analyzed in detail in the middle panel of Fig. 11, where the standard deviation

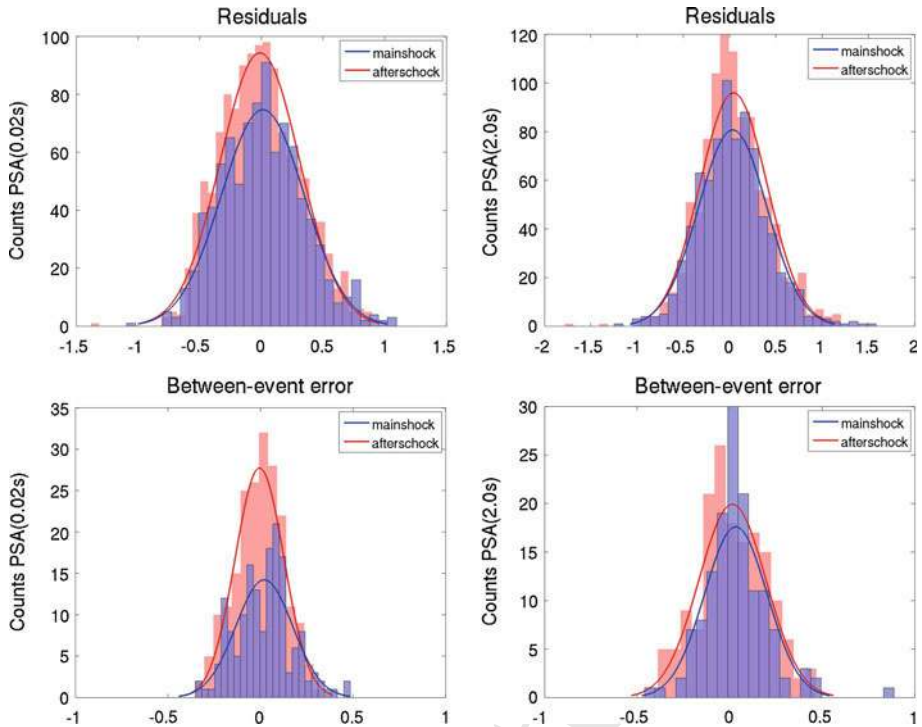


Fig. 10 *Top*: Histograms of the total residual at $T=0.02$ s (*left*) and at $T=2$ s (*right*) and best fit Gaussian models. *Bottom*: the same as in the top panels but for the between-events residuals. *Red bars and curves*: aftershocks; *blue bars and curves*: mainshocks. Please note the different scales on the axes

287 ϕ_{S2S} of the site-to-site error is computed separately for the four different site categories. For
 288 all classes, ϕ_{S2S} assumes the largest value around 0.1 s and the minimum around 0.5 s. For
 289 $T < 0.2$ s, the largest ϕ_{S2S} are those relevant to classes A and B while, for $T > 1$ s, the largest
 290 values are obtained for classes C and D.

291 The record-to-record errors are shown in the bottom panels of Fig. 9, grouped according
 292 to the recording network. The histograms computed for each country are shown in Fig. 12.
 293 For $T = 0.2$ s, the distributions for the different networks have zero mean and similar standard
 294 deviations. For $T = 2$ s, small differences among the mean record-to-record error for different
 295 networks are observed. Greece and Turkey shows the largest standard deviation (0.26 and
 296 0.24, respectively), while the standard deviation for Iran, Italy and Iceland are 0.23, 0.20 and
 297 0.15, respectively.

298 To investigate the presence of possible biases in the residual distributions due to non linear
 299 site effects not modeled in the GMPEs, Fig. 13 shows the record-to-record distribution of
 300 errors for PGA and PGV versus V_{s30} (top panels) and the observed peak values (bottom panels).
 301 The error distributions do not show any trend with V_{s30} , while a weak positive trend is
 302 observed with respect to the input values. This trend is opposite to the one expected in
 303 case of non-linear site effects. Moreover, the trend is the same for soft ($V_{s30} < 360$ m/s) and
 304 stiff/rock ($V_{s30} > 360$ m/s) sites. We conclude that non-linear site effects are not leaving a
 305 significant imprint in the residuals with respect to the predictions from a model including only
 306 a linear site term.

Fig. 11 *Top*: Between-events standard deviation τ compute for three different magnitude classes; *Bottom*: site-to-site standard deviation ϕ_{S2S} computed for the 4 considered EC8 site classes

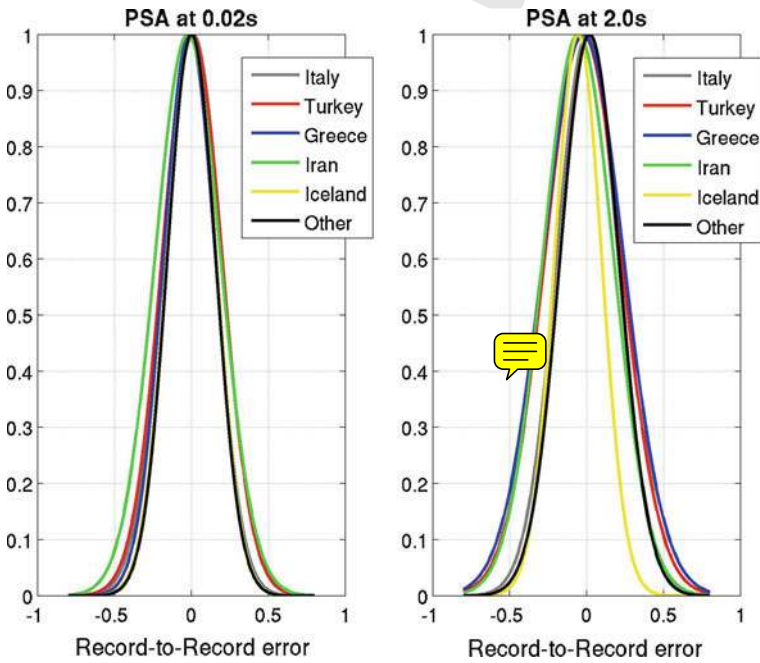
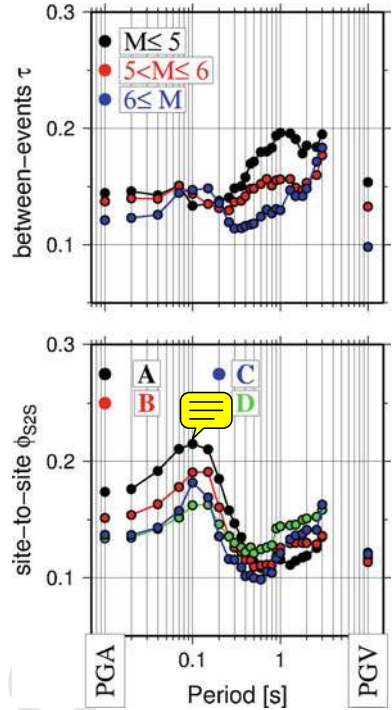


Fig. 12 Record-to-record residuals at $T = 0.02\text{ s}$ (left) and at $T = 2\text{ s}$ (right). The colors are relative to the different networks as indicated in the legend

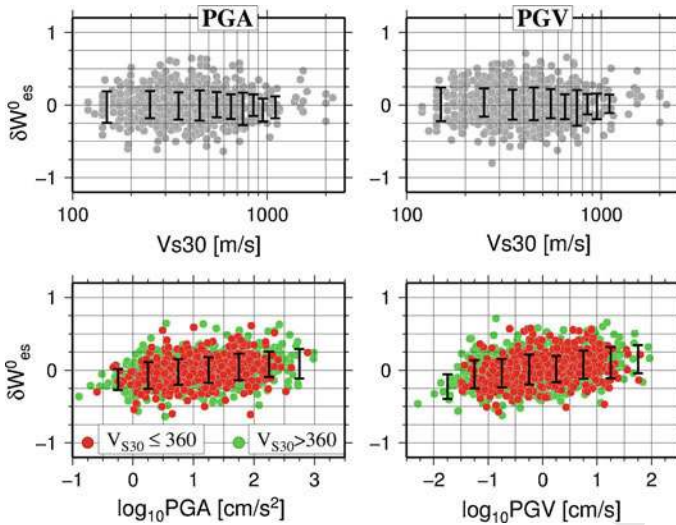


Fig. 13 Top: Record-to-record δW_{se}^0 errors versus V_{s30} for PGA (left) and PGV (right). Mean \pm one standard deviation values (vertical bars) are computed over velocity bins 100 m/s wide except than the first vertical bar, computed over the range $V_{s30} < 200$ m/s², and the last over the range $V_{s30} > 1000$ m/s². Bottom: Record-to-record δW_{se}^0 errors versus \log_{10} PGA (left) and \log_{10} PGV (right). Red and green filled circles correspond to recordings at site with $V_{s30} \leq 360$ m/s² and $V_{s30} > 360$ m/s², respectively

307 6 Comparison with NGA and regional models

308 We compare the equations derived in this study to both global (Akkar and Bommer 2010,
 309 AB10; Cauzzi and Faccioli 2008, CF08; NGA equations by Boore and Atkinson 2008, BA08;
 310 Campbell and Bozorgnia 2008, CB08) and regional GMPEs (Bindi et al. 2011a, ITA10,
 311 developed for Italy; Akkar and Cagnan 2010, AC10, developed for Turkey; Danciu and
 312 Tselentis 2007, DT07, developed for Greece). The main characteristics of these GMPEs are
 313 listed in Table 5. For the BA08 model, the correction factors for small magnitudes proposed
 314 by Atkinson and Boore (2011) are applied. As the GMPEs are based on different distance
 315 metrics, the hypocentral distances, used by CF08, have been converted into R_{JB} using the
 316 relationships proposed by Scherbaum et al. (2004), while the closest distance to the rupture
 317 plane, R_{rup} in CB08, has been estimated with the empirical relationship of Kaklamanos et
 318 al. (2011). The depth to the top of the co-seismic rupture plane, Z_{TOR} , and the depth to the
 319 2.5 km/s shear-wave velocity horizon, typically referred to as basin or sediment depth, $Z_{2.5}$,
 320 used in CB08, have been estimated with Kaklamanos et al. (2011). The comparison has been
 321 made using a vertical strike slip fault, in order to neglect the hanging wall effect present in
 322 CB08.

323 Figure 14 displays the total sigma, as well as its different components, for the considered
 324 GMPEs. In general the total sigma's associated to the NGA models are the smallest, the
 325 current European model (AB10) and the Greek model (DT07) have intermediate values,
 326 while the GMPE developed for Italy (ITA10), Turkey (AC10) and, as consequence, this
 327 study have very similar trends and the largest values.

328 CF08, developed on a global data set, shows large sigma's at short periods and decreasing
 329 sigma's at long periods, comparable to the model developed for Europe. The main difference
 330 among the models is due to the between-events sigma, which is higher for two regional

Table 5 Main features of the global and regional GMPEs used for the comparisons with the GMPEs derived in this study

GMPE code	M range	Distance range [km]	Style of faulting	Site	Period range [s]
AB10 Pan-EU	5–7.6	R _{jb} 0–99	N, R, SS	V _{s30} < 360, 360 < V _{s30} < 760, V _{s30} > 760	0.05–3.0
BA08 NGA	4.2–7.9	R _{jb} 0–200	N, R, SS, U	Function of V _{s30} including non linear effects	0.01–10.0
CF08 global	5–7.2	R _{hypo} 6–150	N, R, SS	Both EC8 A, B, C, D and Function of V _{s30}	0.05–20.0
CB08 NGA	4.0 to 7.5–8.5 (depending on SQF)	R _{rup} 0–200	Function including hanging wall effects	Function of V _{s30} including non linear effects and basin response term	0.01–10.0
DT07 Greece	4.5–6.9	R _{epi} 0–136	N, R, SS	Rock stiff soft	0.1–4.0
ITA10 Italy	4–6.9	R _{jb} 0–200	N, R, SS, U	EC8 A, B, C, D, E	0.04–2.0
AC10 Turkey	5–7.6	R _{jb} 0–200	N, R, SS,	Function of V _{s30} including non linear effects	0.03–2.0
This study Pan-EU	4–7.6	R _{jb} 0–300	N, R, SS, U	EC8 A, B, C, D	0.02–3.0

331 GMPEs (ITA10, AC10), the GMPE developed in this study and the CF08 for period shorter
332 than 0.2s.

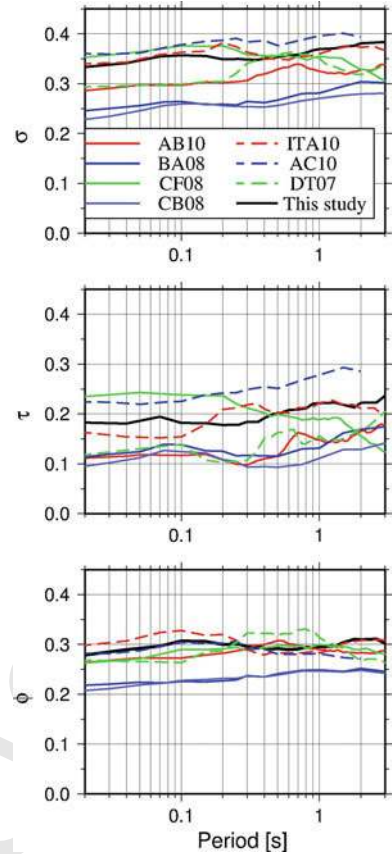
333 Finally, the within-event sigmas obtained from the NGA data set are the lowest (CB08
334 and BA08), while other GMPEs, based on regional or pan-European data sets (ITA10, AC10,
335 DT07, this study) have higher and comparable values.

336 In Figs. 15 and 16, the comparisons among GMPEs are carried out for sites with
337 V_{s30} = 800 m/s and for strike slip style of faulting. Figure 15a shows the median pre-
338 dictions of the global GMPEs in function of magnitude, for PGA and for PSA at
339 T = 2.0s. The major discrepancies among models are found for PGA, in particular at
340 distances of 100km, where the NGA models and AB10 predict higher median val-
341 ues than the present study or CF08. At T = 2.0s, the effect of the magnitude sat-
342 uration model adopted for this study is evident above the magnitude threshold of
343 6.5.

344 Figure 15b shows the median predictions of the global GMPEs in function of distance, for
345 PGA and for SA at T = 2.0s. The major discrepancies among models are observed for PGA,
346 regardless the magnitude. In particular, it is observed that this study and CF08 have similar
347 PGA attenuation with distance, while NGA models and AB10 predict lower attenuation with
348 distance. At long periods the major discrepancies are found in the near source and for low
349 magnitudes.

350 Figure 16a shows the median predictions of the regional GMPEs in function of mag-
351 nitude, for PGA and pseudo spectral acceleration at T = 2.0s. Major differences are evi-
352 dent at long periods and for small magnitudes and at large distances for both PGA
353 and T = 2.0s. In particular, the GMPE with the largest difference is the one devel-

Fig. 14 *Top panel:* total standard deviation; *middle panel:* between-events τ standard deviation; *lower panel:* within event ϕ standard deviation of the considered GMPEs. *Continuous lines* are for global or pan-European models; *dashed lines* for regional models, as indicated in the legend of the *top panel*. See Tables 6 and 7 for the explanation of acronyms



354 oped for Greece (DT07). Figure 16b shows the median predictions of the regional
 355 GMPEs as function of distance, for PGA and PSA at $T=2.0$ s. The major discrepan-
 356 cies among models are observed for PGA for small magnitudes and long periods. In
 357 particular, at $T=2.0$ s the median predictions for Turkey and Italy are lower than the
 358 prediction of this study. The GMPE developed for Greece (DT07) shows systematically
 359 higher median predictions at distances larger than 10km, especially at small mag-
 360 nitudes.

361 7 Discussion

362 After this study, the following issues can be evidenced:

363 7.1 Range of applicability of the GMPEs

364 The GMPEs are strictly usable in their range of applicability (magnitude M_w in the range
 365 4–7.4; distances R_{JB} smaller than 300 km; periods in the range 0.02–3 s).

366 To this end Fig. 17 (left panel) shows the total standard deviation along with its between-
 367 and within-components, **are** computed for PSA up to 10s, as included in RESORCE data
 368 set. The sharp drop in the different components casts some doubts on the reliability of the

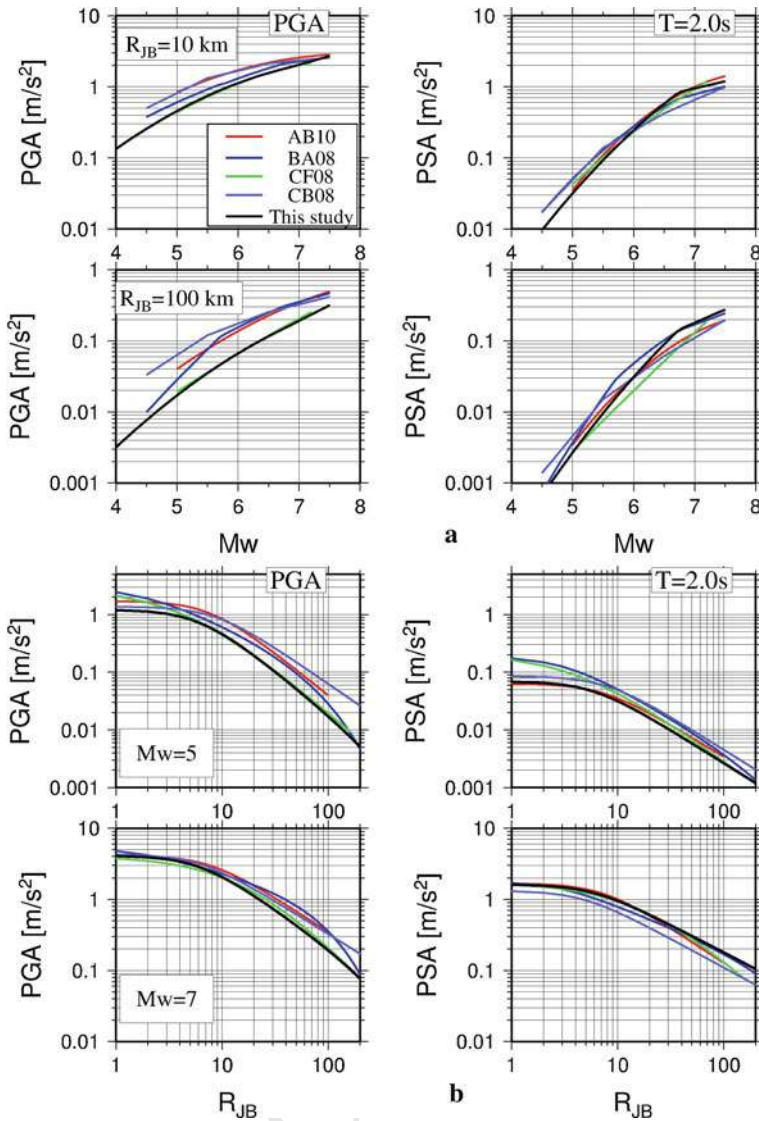


Fig. 15 Global GMPEs: **a** Pseudo spectral acceleration ordinates PSA in function of magnitude; *upper panel* PGA (left) and $T = 0.2\text{ s}$ (right) at a distance of 10 km; *lower panel* PGA (left) and $T = 0.2\text{ s}$ (right) at a distance of 100 km; **b** pseudo spectral acceleration ordinates in function of distance; *upper panel* PGA (left) and $T = 0.2\text{ s}$ (right) for an event of $M_w = 5.0$; *lower panel* PGA (left) and $T = 0.2\text{ s}$ (right) for an event of $M_w = 7$. The predictions are computed for $V_{s30} = 800\text{ m/s}^2$. See Tables 6 and 7 for the explanation of acronyms

369 variability captured by the model at periods longer than 3 s. This is confirmed by the increase
 370 of the regression coefficient instability at periods longer than 1 s (Figs. 4, 6).

371 These results could be related both to the reduction in the number of the considered
 372 recordings at long period (see Fig. 2, right) and to the processing scheme which was not
 373 optimized for long periods.

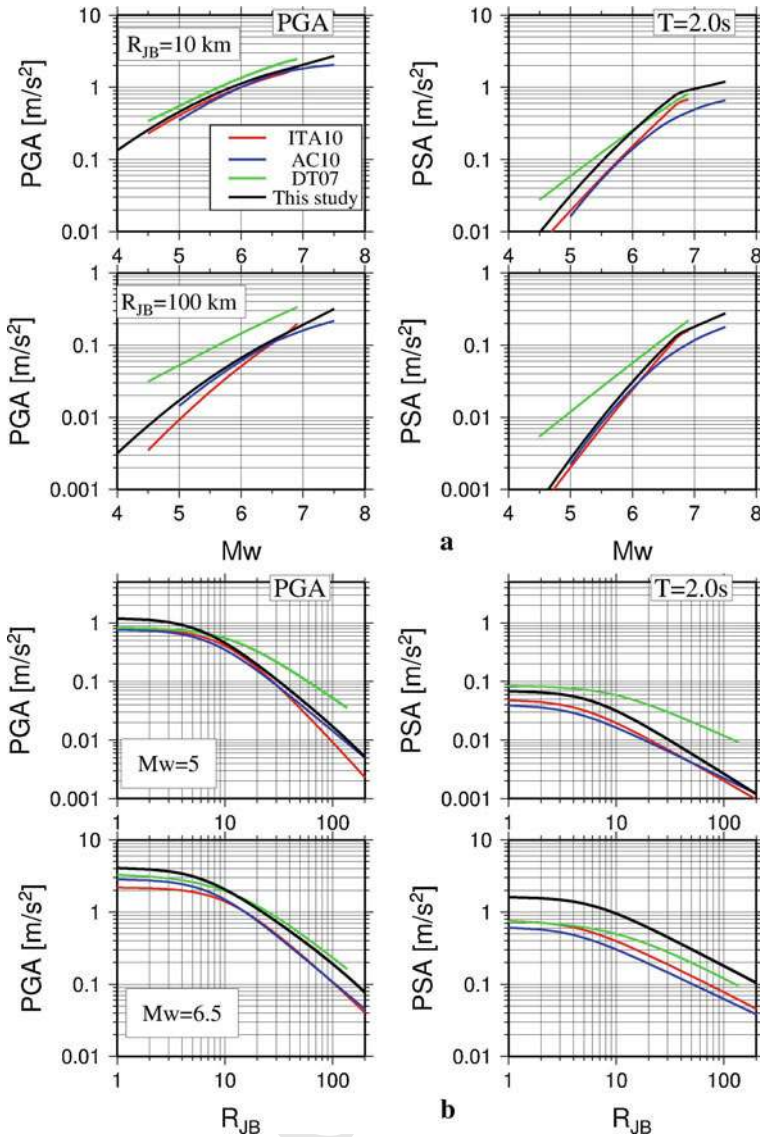


Fig. 16 Regional GMPEs: **a** Pseudo spectral acceleration ordinates in function of magnitude; *upper panel* PGA (left) and $T = 0.2$ s (right) at a distance of 10 km; *lower panel* PGA (left) and $T = 0.2$ s (right) at a distance of 100 km; **b** Pseudo spectral acceleration spectral ordinates in function of distance; *upper panel* PGA (left) and $T = 0.2$ s (right) for an event of $M_w = 5.0$. *Lower panel* PGA (left) and $T = 0.2$ s (right) for an event of $M_w = 7$. See Tables 6 and 7 for the explanation of acronyms

374 7.2 Mainshock /aftershocks

375 The distributions of the residuals for aftershocks and mainshocks are almost unbiased and
 376 with similar sigma (Fig. 10), therefore the features of mainshocks and aftershocks can-
 377 not be captured by the empirical ground-motion equations developed for Europe. This can
 378 be attributed to the tectonic complexity in Europe, where events belonging to the same

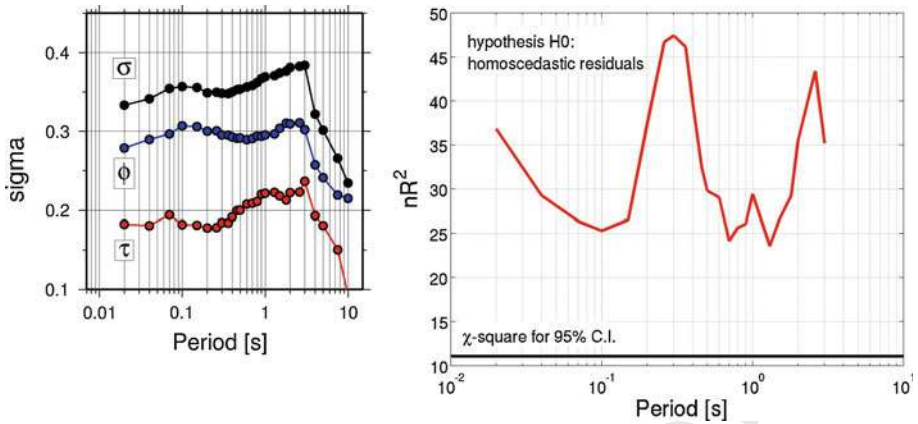


Fig. 17 Left: Standard deviations (*black*) along with the between-events (*blue*) and within-event (*red*) components computed up to 10 s. Right: Result of the White test applied for testing the null hypothesis of homoscedastic residuals. Red curve: nR^2 coefficient of the regression performed over the squared residuals versus the explanatory variable (magnitude and distance), their squares and cross product. Black line: the 95 % confidence value for a χ^2 distribution with 6 degree of freedoms indicated that the null hypothesis can be rejected at this level of confidence

379 seismic sequence frequently occur on adjacent faults (e.g. [Umbria-Marche 1997](#); [L'Aquila](#)
380 [2009](#)).

381 7.3 Site characterization and nonlinear site effects

382 Soil category A (rock) dominates the site-to-site variability for short-periods while for long
383 periods the difference among the classes is less pronounced, with soil sites (classes C and D)
384 showing larger site-to-site variability (Fig. 11). The long period amplifications ($T > 1$ s)
385 for classes C or D can be related to both one-dimensional (1D) resonant effects or 2D–3D basin
386 effects (e.g. [Rovelli et al. 2001](#); [Bindi et al. 2009b, 2011c](#)), which cannot be captured by any
387 simple site model, based either on the V_s30 or on classes identified by V_s intervals (e.g.
388 [Luzi et al. 2011](#)).

389 Non-linear effects are expect to be important for strong shaking at soil sites, that is for large
390 and close earthquakes recorded at sites with low V_s30 values (e.g. classes C and D of EC8).
391 These conditions are not adequately sampled in RESORCE (Figs. 1, 2). The characteristics
392 of the dataset, in terms of magnitude, distance and site characteristics, do not evidence non-
393 linear site effects, as shown by the residual distribution in Fig. 13. Therefore a non-linear site
394 model cannot be calibrated with the data set used to derive the GMPEs (see also [Akkar and](#)
395 [Bommer 2007a,b](#)). Nonlinear models can be only calibrated on external data sets, introducing
396 potential bias in the predictions.

397 In conclusion, given the RESORCE data set, we preferred to derive a ground-motion model
398 which includes a linear site term. For application where non-linear effects are expected,
399 the predictions of these GMPEs should be corrected by factors calibrated on the specific
400 characteristics of the investigated sites.

401 7.4 SIGMA

402 The sigma values obtained in this study are of the same order of the sigma's obtained by
403 the regional models for Turkey and Italy, that mostly contribute to RESORCE ([Bindi et al.](#)

2011a; Akkar and Cagnan 2010). These values are larger than sigma's obtained for global or Pan-European GMPEs, which are generally derived for magnitudes larger than 5 (i.e. BA08, CF08, AB10).

The between-events sigma (Fig. 14) of this study is within the values obtained by Akkar and Cagnan (2010, Turkish dataset), and those obtained by Bindi et al. (2011a, Italian dataset) and is generally large for small magnitude events and long periods (Fig. 11). This effect can be ascribed to the large variability of the ground-motion at low magnitudes and/or to the low-cut corner filters applied to the small magnitude events, which have probably introduced low-frequency noise in the analysis. The between-events sigma could also be affected by the conversion into Mw from other magnitude scales.

Figure 17 (right panel) shows the result of a White test (1980) applied to the residual distribution, to test the null hypothesis of homoscedastic residuals. This test is based on the regression of the squared residuals versus the explanatory variables (magnitude and distance in our test), their square and their cross-product. The comparison of the nR^2 coefficient of the regression (red curve in Fig. 17, right) with the critical value for a chi-squared distribution (black line) allows us to reject the null hypothesis of homoscedastic residuals for all periods at 95% confidence interval. The result of this preliminary test confirms that it is worth to investigate in future the possibility of integrating RESORCE with selected strong earthquakes to constrain a heteroscedastic models over the magnitude range from 4 to 8.

8 Conclusions

We derived a set of GMPEs from a ~~Pan-European data set~~, RESORCE, compiled in the framework of the project SIGMA (<http://projet-sigma.com/organisation.html>), ~~to improve the seismic hazard assessment in France.~~

This data set includes the most recent (up to 2011) strong motion data for Europe and Middle East. From the dataset, we extracted a ~~dataset~~, named DS-EC8, containing recordings from magnitude and distance ranges wide enough to satisfy a large spectrum of applications: moment magnitudes larger than or equal to 4, hypocentral depths lower than 35 km and Joyner-Boore (R_{JB}) or epicentral (R_{epi}) distances lower than 300 km. The sites are categorized according to the EC8 ground categories (from class A to D), while the style-of-faulting is accounted for as four categories (normal, reverse, strike-slip or unspecified).

A subset of DS-EC8, named DS-VS30, containing only waveforms characterized by known style of faulting and recorded by stations having a measured shear wave velocity profiles, has been used to test of the accuracy of the mean prediction and the variability associated to DS-EC8.

A parametric model has been adopted for the regression of both datasets, following Boore and Atkinson (2008), Bindi et al. (2011a) and Akkar and Cagnan (2010). The predictions associated to the two datasets have similar variability, although the sigma relative to DS-VS30 is slightly smaller than the one relative to DS-EC8 (Fig. 4b). The differences are mainly ascribed to a reduction of the between-events standard deviation, due to the specification of focal mechanisms, and of the site-to-site sigma, due to the knowledge of the subsoil profile, especially at short periods. Both models are reliable, although the one derived from the DS-EC8 dataset has a wider range of applicability, as style of faulting can be unspecified and sites classification can be inferred from surface geology. The median predictions of

450 this study are in agreement with the results of models derived from global data sets at long
451 periods. At short periods (PGA), the model derived in this study shows a better agreement
452 with the [Cauzzi and Faccioli \(2008\)](#) one, while it shows a different rate of attenuation with
453 distance with respect to the considered NGA models. The sigma obtained in this study is
454 larger than sigma's obtained for global or Pan-European GMPEs, which are generally derived
455 for magnitudes larger than 5 (i.e. BA08, CF08, AB10). There is a good agreement with the
456 equations derived for Italy and Turkey, in terms of median and standard deviation at short
457 periods, although the GMPEs derived in this study predict larger ground shaking at long
458 periods.

459 Considering the sigma values of the model derived in this study, we suggest a revision
460 of the metadata relevant to earthquakes with magnitude < 5 , and particular attention
461 should be paid to the conversion from local to moment magnitude for small events,
462 since the between-events sigma for these earthquakes largely contributes to the total variability
463 at long periods. A revision of the high-pass corner frequencies for these earthquakes
464 is also suggested. The evaluation of GMPEs at periods longer than 3 s requires an
465 increase of the number of large-magnitude events, that can be achieved including also earthquakes
466 occurred outside Europe. Moreover, in order to better capture the ground motion
467 variability at different magnitudes, the implementation of heteroscedastic models might be
468 explored.

469 Finally, the reduction of the epistemic uncertainty affecting the sites might be reached by
470 considering more sophisticated site amplification functions, including for example nonlinear
471 models and basin depth, although there is still a strong limitation of geotechnical information
472 regarding the recording sites, as few data providers in Europe promoted site characterization
473 programs.

474 **Acknowledgments** The coefficients of the GMPEs derived in this study and their 95% confidence intervals
475 are available in the Electronic Supplements of this article, along with some Matlab scripts to compute predictions
476 for different selections of the explanatory variable. Comments from an anonymous Reviewer and from
477 C. Cauzzi triggered significant improvements in the article. The Editor J. Douglas is also acknowledged. The
478 research activities of D. Bindi have been founded through the REAKT (Strategies and Tools for Real Time
479 Earthquake Risk Reduction) EU-FP7 project.

480 9 Appendix

481 See Tables [6](#), [7](#), [8](#) and [9](#).

Table 6 95 % Confidence intervals for the model derived in this study (see Eqs. 1–3) for R_{fB} and EC8 ground categories

T [see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	0.12391	0.121527	0.115083	0.163395	0.157876	0.145339	0.126843	0.145158	0.131924	0.108779	0.148158	0.106598
c1	0.073593	0.084836	0.076821	0.087507	0.086066	0.087099	0.075175	0.056234	0.076841	0.045285	0.080007	0.056646
c2	0.025963	0.022801	0.029608	0.025116	0.031075	0.02521	0.023317	0.024694	0.026571	0.023454	0.030767	0.023685
h	0.705677	0.826334	0.617483	0.754097	0.922982	0.914605	0.7553	0.561224	0.685506	0.637961	0.913656	0.480418
c3	0.000379	0.000421	0.000501	0.000505	0.000528	0.000308	0.000504	0.00043	0.000499	0.000248	0.000383	0.000381
b1	0.061038	0.072102	0.056655	0.067203	0.06722	0.058176	0.06342	0.069368	0.07324	0.061945	0.07168	0.050007
b2	0.013753	0.018276	0.013364	0.013868	0.014396	0.016944	0.016195	0.014831	0.015996	0.014721	0.0169	0.011982
b3	0.015539	0.026582	0.032912	0.038649	0.068145	0.091046	0.089351	0.064411	0.076461	0.075277	0.075031	0.083733
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.019884	0.019101	0.023396	0.021834	0.02041	0.018293	0.021751	0.018888	0.014545	0.017174	0.016245	0.018516
ClassC	0.021135	0.019827	0.025217	0.023074	0.024861	0.020882	0.025004	0.022629	0.024476	0.020305	0.019206	0.019526
Class D	0.033404	0.032331	0.043163	0.037048	0.035284	0.041073	0.046337	0.037456	0.029156	0.038351	0.033783	0.033809
sofN	0.050451	0.044131	0.034368	0.046779	0.057222	0.053451	0.06246	0.06278	0.050022	0.050633	0.040136	0.052285
sofR	0.053122	0.048082	0.038819	0.059826	0.065816	0.057652	0.072096	0.064313	0.059711	0.058684	0.049947	0.058154
sofS	0.05393	0.043475	0.032421	0.049164	0.065702	0.051469	0.0633086	0.063357	0.054041	0.05164	0.043349	0.050162
sofU	0	0	0	0	0	0	0	0	0	0	0	0

Table 6 continued

T [sec]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV
e1	0.135883	0.111665	0.125455	0.125496	0.155883	0.156283	0.102505	0.155931	0.151116	0.182144	0.122927	0.130382
c1	0.046427	0.05541	0.060255	0.045397	0.072641	0.071594	0.056977	0.076385	0.07979	0.109924	0.088425	0.075018
c2	0.029473	0.024827	0.030003	0.027033	0.028322	0.028474	0.026231	0.029861	0.026756	0.031878	0.043311	0.024405
h	0.614371	0.613053	0.723977	0.633978	0.718858	0.875666	0.82935	1.03247	0.896828	1.36986	1.15451	0.604569
c3	0.000311	0.000139	0.000153	9.03E-05	5.26E-05	3.97E-05	7.66E-06	8.94E-05	0.000132	9.26E-05	8.97E-05	0.00047
b1	0.072167	0.044728	0.067567	0.059175	0.073389	0.077649	0.063371	0.076473	0.086585	0.07048	0.100127	0.059753
b2	0.014247	0.01267	0.014035	0.011545	0.016435	0.017151	0.01591	0.017941	0.021173	0.013728	0.021513	0.014809
b3	0.08016	0.08867	0.077122	0.061423	0.070072	0.05788	0.043681	0.002306	0	0	0	0.030284
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.023451	0.01943	0.014079	0.019225	0.017356	0.019117	0.011715	0.021086	0.01732	0.026648	0.027186	0.018097
Class C	0.026519	0.021883	0.025502	0.023149	0.018456	0.023132	0.018751	0.019497	0.024489	0.026961	0.026467	0.016873
Class D	0.037533	0.033704	0.042822	0.039871	0.034117	0.052518	0.042387	0.050683	0.05548	0.053011	0.058549	0.043388
SoFN	0.044883	0.050114	0.048332	0.064846	0.046582	0.050787	0.059159	0.073939	0.076347	0.061466	0.057963	0.04305
SoFR	0.048823	0.057269	0.046387	0.06263	0.057384	0.058538	0.062408	0.077215	0.072066	0.065347	0.059997	0.04705
SoFS	0.049879	0.056254	0.049868	0.069784	0.050601	0.055853	0.06573	0.078287	0.07267	0.058471	0.058345	0.04176
SoFu	0	0	0	0	0	0	0	0	0	0	0	0

Acceleration is in (cm/s²), velocity in (cm/s). Values equal to zero correspond to coefficients constrained in the regression

Table 7 95 % Confidence intervals for the model derived in this study (see Eqs. 1–3) for R_{fB} and V_{s30} classification

T[see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	0.119376	0.165741	0.114871	0.138983	0.114853	0.15357	0.130826	0.133852	0.119588	0.164146	0.140609	0.147126
c1	0.071763	0.089741	0.067479	0.072709	0.057429	0.089002	0.0773	0.0879	0.067489	0.092358	0.078861	0.070891
c2	0.031129	0.034681	0.030171	0.034202	0.035005	0.034895	0.028872	0.026112	0.024534	0.030809	0.033135	0.038603
h	0.763209	0.75251	0.737684	0.740023	0.682503	1.01786	0.823344	0.804603	0.672935	0.794876	0.723385	0.717549
c3	0.000494	0.00059	0.000405	0.00051	0.000388	0.000521	0.000514	0.000655	0.000404	0.000571	0.000574	0.000452
b1	0.072438	0.085452	0.083027	0.082115	0.079822	0.089221	0.061706	0.069364	0.077476	0.084642	0.090658	0.09876
b2	0.01431	0.018165	0.018655	0.019665	0.016307	0.020788	0.019248	0.01606	0.019264	0.0185	0.018696	0.022122
b3	0.039825	0.037486	0.04596	0.070542	0.099805	0.088819	0.104139	0.072424	0.072665	0.1110506	0.096252	0.111751
γ	0.052584	0.045789	0.03917	0.050205	0.03688	0.054365	0.060873	0.045314	0.063349	0.051652	0.058377	0.046624
sofN	0.01437	0.013604	0.013656	0.015221	0.014105	0.010393	0.013403	0.014455	0.01333	0.014434	0.013395	0.013092
sofR	0.021365	0.015231	0.020062	0.017379	0.02124	0.018135	0.021546	0.022927	0.016435	0.023006	0.022461	0.019377
sofS	0.01647	0.01487	0.012361	0.016317	0.016003	0.016366	0.016502	0.017352	0.011701	0.018001	0.014508	0.015469

Table 7 continued

T [see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGY
e1	0.136069	0.161653	0.143452	0.148138	0.13107	0.120361	0.176144	0.196548	0.178821	0.181898	0.175894	0.145007
c1	0.065181	0.083973	0.07617	0.067917	0.069687	0.071903	0.085837	0.096141	0.098345	0.097969	0.091141	0.080703
c2	0.032502	0.035902	0.031267	0.035507	0.027695	0.032235	0.040639	0.048326	0.041525	0.04314	0.048851	0.03229
h	0.722435	0.809068	0.863448	0.942758	0.493574	0.737819	1.09324	1.18635	1.21633	1.97771	1.5161	0.604155
c3	0.000381	0.000298	0.000302	0.000181	9.07E-05	0.000163	0.000124	8.91E-05	0.000256	0.000214	0.000319	0.000508
b1	0.090228	0.088144	0.08326	0.087647	0.081742	0.091426	0.112675	0.111454	0.10073	0.106514	0.133328	0.08088
b2	0.020646	0.018802	0.023769	0.017214	0.021989	0.026164	0.026381	0.02855	0.027099	0.022866	0.035092	0.019505
b3	0.090883	0.09918	0.081987	0.078417	0.070708	0.101311	0.1110192	0.067413	0.04185	0.035124	0.03977	0.039424
γ	0.043885	0.061556	0.050754	0.061085	0.047408	0.069918	0.057408	0.059191	0.062944	0.068649	0.075872	0.045541
SoFN	0.013892	0.011203	0.010746	0.017231	0.012809	0.014081	0.018581	0.019061	0.01663	0.015954	0.016157	0.01447
SoFR	0.018023	0.017466	0.022276	0.024694	0.02027	0.019541	0.027255	0.024237	0.02499	0.028409	0.023208	0.018719
SoFS	0.015531	0.01588	0.021322	0.019235	0.017704	0.013794	0.018027	0.016931	0.020535	0.022017	0.019899	0.01403

Acceleration is in (cm/s²), velocity in (cm/s)

Table 8 95 % Confidence intervals for the model derived in this study (see Eqs. 1–3) for R_{HYPO} and EC8 ground categories

T[see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	0.162223	0.154254	0.189283	0.171375	0.169115	0.17316	0.199215	0.219545	0.1617	0.162699	0.132846	0.141996
c1	0.101039	0.092358	0.097128	0.095642	0.102693	0.108807	0.093717	0.1112268	0.078306	0.091217	0.08547	0.060794
c2	0.033926	0.037746	0.037693	0.042037	0.033336	0.035483	0.040852	0.033752	0.029748	0.033652	0.028229	0.028434
h	0.962928	1.19234	1.1525	1.27163	1.24611	1.37899	1.60572	1.50953	1.00863	1.01162	1.30428	0.973498
c3	0.000116	0.000147	0.000286	0.000394	0.000571	0.00058	0.000573	0.000478	0.000297	0.000242	0.000103	0.000019
b1	0.070048	0.075749	0.088331	0.087252	0.07528	0.065611	0.095364	0.076053	0.086192	0.072464	0.075332	0.067039
b2	0.011773	0.012351	0.016414	0.01577	0.015674	0.017542	0.017092	0.014208	0.017884	0.013348	0.016165	0.015296
b3	0.107028	0.103664	0.117345	0.139338	0.121656	0.143074	0.098256	0.119121	0.118327	0.127049	0.109	0.125675
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.021473	0.020169	0.022509	0.020515	0.020449	0.019537	0.016488	0.015888	0.021704	0.019351	0.02195	0.027305
Class C	0.019609	0.024995	0.023025	0.025756	0.020895	0.02229	0.019367	0.018069	0.018154	0.025808	0.023742	0.020612
Class D	0.036033	0.041074	0.045499	0.042577	0.041605	0.045418	0.033054	0.038452	0.039912	0.045533	0.029208	0.039484
sofN	0.042147	0.039597	0.041724	0.044908	0.047293	0.044762	0.056859	0.056113	0.038272	0.052119	0.039565	0.059416
sofR	0.04405	0.043664	0.04608	0.045385	0.055023	0.047656	0.069293	0.054435	0.044668	0.056132	0.04171	0.065653
sofS	0.041382	0.037388	0.041507	0.045641	0.046228	0.047956	0.058056	0.053163	0.045456	0.054731	0.035614	0.061234
sofU	0	0	0	0	0	0	0	0	0	0	0	0

Table 8 continued

T [see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV
e1	0.176433	0.148246	0.136189	0.151171	0.145034	0.158629	0.146203	0.148056	0.200347	0.220246	0.229088	0.104644
c1	0.087692	0.090535	0.081746	0.078289	0.098783	0.068711	0.095043	0.099082	0.112949	0.1111766	0.136501	0.07744
c2	0.036387	0.027564	0.031472	0.030833	0.026075	0.039745	0.032644	0.021522	0.039154	0.038436	0.036221	0.031948
h	1.17002	1.15677	2.0927	1.27159	1.819	0.978392	1.85639	1.82087	1.84098	1.89477	1.83951	0.93246
c3	7.32E-05	0	0	1.55E-06	0	0	0	0	0	2.29E-05	0	0.000121
b1	0.086479	0.056097	0.079706	0.078264	0.073053	0.093396	0.070858	0.062862	0.08394	0.103439	0.096826	0.076066
b2	0.015014	0.013019	0.014594	0.01926	0.016978	0.016364	0.014635	0.017022	0.019467	0.020413	0.020473	0.015053
b3	0.100437	0.090815	0.092353	0.106859	0.084317	0.118405	0.102441	0.064018	0.079199	0.011837	0.002244	0.089606
Class A	0	0	0	0	0	0	0	0	0	0	0	0
Class B	0.018676	0.018839	0.013379	0.019991	0.019596	0.022271	0.016666	0.020877	0.019221	0.021426	0.028038	0.016506
Class C	0.023959	0.024281	0.021598	0.01854	0.024414	0.023715	0.022187	0.025692	0.020558	0.02693	0.024545	0.02008
Class D	0.049813	0.040629	0.038765	0.032032	0.040218	0.044745	0.053531	0.055989	0.040841	0.053971	0.07465	0.041802
SoFN	0.063185	0.049799	0.066682	0.048557	0.046751	0.058559	0.053649	0.060967	0.080818	0.074659	0.074224	0.045222
SoFR	0.066195	0.050842	0.068955	0.058334	0.054271	0.062446	0.061971	0.061758	0.091634	0.071649	0.07204	0.049115
SoFS	0.06809	0.052927	0.064046	0.049255	0.045955	0.058149	0.052172	0.061587	0.085699	0.077932	0.069574	0.044288
SoFu	0	0	0	0	0	0	0	0	0	0	0	0

Acceleration is in (cm/s²), velocity in (cm/s). Values equal to zero correspond to coefficients constrained in the regression

Table 9 95 % Confidence intervals for the model derived in this study (see Eqs. 1–3) for R_{HYPO} and $Vs30$ classification

T[see]	0.02	0.04	0.07	0.1	0.15	0.2	0.26	0.3	0.36	0.4	0.46	0.5
e1	0.148614	0.221493	0.161911	0.127908	0.155409	0.23013	0.228819	0.224915	0.165211	0.147006	0.199166	0.223988
c1	0.077069	0.104547	0.079365	0.077602	0.084072	0.113165	0.123219	0.116923	0.093036	0.078328	0.094014	0.103695
c2	0.048909	0.059291	0.044418	0.037208	0.04363	0.050617	0.05433	0.050143	0.040449	0.038115	0.051277	0.042569
h	1.06658	1.00828	1.12928	1.00474	1.48942	1.34527	1.1677	1.59307	1.12573	1.36232	1.04717	1.46143
c3	0.000316	0.000141	0.000454	0.000468	0.000485	0.00077	0.000753	0.000546	0.000274	0.000134	0.00031	0.000237
b1	0.107778	0.120828	0.099091	0.085023	0.103087	0.116569	0.133753	0.109063	0.098902	0.094054	0.110216	0.126226
b2	0.01887	0.023244	0.01927	0.01913	0.021987	0.022348	0.024811	0.022083	0.02013	0.017234	0.018042	0.024789
b3	0.148945	0.191562	0.147567	0.117771	0.143702	0.163601	0.161578	0.12711	0.168939	0.127675	0.11769	0.115833
γ	5.56E-02	5.53E-02	5.44E-02	5.54E-02	4.38E-02	6.07E-02	4.94E-02	5.36E-02	7.12E-02	5.16E-02	5.97E-02	5.43E-02
sofN	0.012945	0.015325	0.015092	0.010938	0.01641	0.015702	0.013105	0.014415	0.01489	0.012312	0.015086	0.012935
sofR	0.020097	0.01527	0.020411	0.014855	0.019031	0.022561	0.019818	0.025483	0.024511	0.019294	0.021172	0.020085
sofS	0.015127	0.01885	0.013671	0.01232	0.01542	0.015604	0.01503	0.018095	0.017524	0.016167	0.019834	0.017467

Table 9 continued

T [see]	0.6	0.7	0.9	1	1.3	1.5	1.8	2	2.6	3	PGA	PGV
e1	0.163636	0.179788	0.257854	0.18107	0.205358	0.2181	0.202476	0.187422	0.208332	0.259683	0.22777	0.154951
c1	0.115852	0.105718	0.102813	0.094651	0.118656	0.129953	0.11866	0.091345	0.103871	0.152804	0.136084	0.085984
c2	0.04157	0.041201	0.060975	0.038254	0.033847	0.047836	0.04956	0.038437	0.048528	0.060871	0.05732	0.033303
h	1.53667	1.36415	1.51506	1.63559	1.78867	1.9574	2.33189	2.03347	2.24876	2.38718	2.38106	1.06636
c3	2.59E-04	0.000247	1.81E-05	1.02E-06	0	0	0	0	0	6.94E-05	0.000177	0.000347
b1	0.060884	0.097142	0.128807	0.099575	0.061382	0.089216	0.107353	0.114807	0.11234	0.142646	0.166586	0.085348
b2	0.021587	0.018132	0.018901	0.020257	0.014423	0.018383	0.022548	0.031173	0.02943	0.033805	0.038534	0.01764
b3	0.159759	0.107094	0.164912	0.119502	0.110014	0.143337	0.174913	0.129941	0.130879	0.097421	0.096057	0.093406
γ	4.67E-02	4.04E-02	6.95E-02	5.67E-02	5.68E-02	6.54E-02	5.22E-02	6.68E-02	6.17E-02	6.44E-02	4.72E-02	4.55E-02
SoIN	0.015261	0.017573	0.016388	0.017726	0.017304	0.020619	0.016116	0.018223	0.021319	0.016337	0.018286	0.014969
SoIR	0.020816	0.024844	0.027932	0.020622	0.015056	0.025939	0.024476	0.024936	0.030811	0.028956	0.03236	0.022779
SoIS	0.015042	0.015076	0.02162	0.014413	0.016785	0.015322	0.022126	0.021087	0.020439	0.021276	0.025577	0.015203

Acceleration is in (cm/s²), velocity in (cm/s)

482 **References**

- 483 Abrahamson NA, Youngs RR (1992) A stable algorithm for regression analyses using the random effects
484 model. *Bull Seismol Soc Am* 82(1):505–510
- 485 Akkar S, Bommer JJ (2007a) Prediction of elastic displacement response spectra in Europe and the Middle
486 East. *Earthq Eng Struct Dyn* 36:1275–1301. doi:10.1002/eqe.679
- 487 Akkar S, Bommer JJ (2007b) Empirical prediction equations for peak ground velocity derived from strong-
488 motion records from Europe and the Middle East. *Bull Seismol Soc Am* 97(2):511–530. doi:10.1785/
489 0120060141
- 490 Akkar S, Bommer JJ (2010) Empirical equations for the prediction of PGA, PGV, and spectral accelerations
491 in Europe, the mediterranean region, and the middle east. *Seismol Res Lett* 81(2):195–206. doi:10.1785/
492 gssrl.81.2.195
- 493 Akkar S, Cagnan Z (2010) A local ground-motion predictive model for Turkey, and its comparison with
494 other regional and global ground-motion models. *Bull Seismol Soc Am* 100(6):2978–2995. doi:10.1785/
495 0120090367
- 496 Akkar S, Sandikkaya MA, Şenyurt M, Azari AS, Ay BÖ (2013) Reference database for seismic ground-motion
497 in Europe (RESORCE), *Bull Earthq Eng*. [Submitted to this issue](#).
- 498 Al Atik L, Abrahamson NA, Bommer JJ, Scherbaum F, Cotton F, Kuehn N (2010) The variability of ground-
499 motion prediction models and its components. *Seismol Res Lett* 81(5):794–801. doi:10.1785/gssrl.81.5.
500 794
- 501 Atkinson G, Boore D (2011) Modifications to existing ground-motion prediction equations in light of new
502 data. *Bull Seismol Soc Am* 101:1121–1135
- 503 Bindi D, Luzi L, Pacor F (2009) Interevent and interstation variability computed for the Italian Accelerometric
504 Archive (ITACA). *Bull Seismol Soc Am* 99(4):2471–2488. doi:10.1785/0120080209
- 505 Bindi D, Parolai S, Cara F, Di Giulio G, Ferretti G, Luzi L, Monachesi G, Pacor F, Rovelli A (2009b) Site
506 amplifications observed in the Gubbio Basin, Central Italy: hints for lateral propagation effects. *Bull Seism
507 Soc Am* 99:741–760. doi:10.1785/0120080238
- 508 Bindi D, Pacor F, Luzi L, Puglia R, Massa M, Ameri G, Paolucci R (2011a) Ground-motion prediction
509 equations derived from the Italian strong motion database. *Bull Earthq Eng* 9(6):1899–1920. doi:10.1007/
510 s10518-011-9313-z
- 511 Bindi D, Luzi L, Pacor F, Paolucci R (2011b) Identification of accelerometric stations in ITACA with distinctive
512 features in their seismic response. *Bull Earthq Eng* 9:1921–1939. doi:10.1007/s10518-011-9271-5
- 513 Bindi D, Luzi L, Parolai S, Di Giacomo D, Monachesi G (2011c) Site effects observed in alluvial basins: the
514 case of Norcia (Central Italy). *Bull Earthq Eng* 9:1941–1959
- 515 Bommer JJ, Akkar S (2012) Consistent source-to-site distance metrics in ground-motion prediction equations
516 and seismic source models for PSHA. *Earthq Spectra* 28(1):1–15
- 517 Boore DM, Atkinson GM (2008) Ground-motion prediction equations for the average horizontal component of
518 PGA, PGV, and 5%-damped PSA at spectral periods between 0.01 s and 10.0 s. *Earthq Spectra* 24:99–138
- 519 Campbell KW, Bozorgnia Y (2008) NGA ground-motion model for the geometric mean horizontal component
520 of PGA, PGV, PGD and 5% damped linear elastic response spectra for periods ranging from 0.01 to 10 s.
521 *Earthq Spectra* 24(1):139–171
- 522 Cauzzi C, Faccioli E (2008) Broadband (0.05 to 20 s) prediction of displacement response spectra based on
523 worldwide digital records. *J Seismol* 12(4):453–475. doi:10.1007/s10950-008-9098-y
- 524 Comité Européen de Normalisation (CEN) (2004) Eurocode 8: design of structures for earthquake resistance—
525 Part 1: general rules, seismic actions and rules for buildings. Comité Européen de Normalisation, Brussels
- 526 Danciu L, Tselentis G-A (2007) Engineering ground-motion parameters attenuation relationships for Greece.
527 *Bull Seismol Soc Am* 97(1B):162–183. doi:10.1785/0120040087
- 528 Douglas J, Halldórsson B (2010) On the use of aftershocks when deriving ground-motion prediction equations.
529 In: Proceedings of the 9th U.S. national and 10th Canadian conference on earthquake engineering, paper
530 no. 220
- 531 Douglas J, Akkar S, Ameri G, Bard P-Y, Bindi D, Bommer JJ, Bora SS, Cotton F, Derras B, Hermkes M, Kuehn
532 NM, Luzi L, Massa M, Pacor F, Riggelsen C, Sandikkaya MA, Scherbaum F, Stafford PJ, Traversa P (2013)
533 Comparisons among the five ground-motion models 1 developed using RESORCE for the prediction of
534 response spectral accelerations due to earthquakes in Europe and the Middle East, *Bulletin of Earthquake
535 Engineering*, this issue
- 536 Efron B, Tibshirani RJ (1994) *An introduction to the bootstrap*. Chapman & Hall/CRC, Boca Raton, Florida.
537 ISBN 978-0412042317, 456 pp
- 538 Foti S, Parolai S, Bergamo P, Di Giulio G, Maraschini M, Milana G, Picozzi M, Puglia R (2011) Surface wave
539 surveys for seismic site characterization of accelerometric stations in ITACA. *Bull Earthq Eng* 9:1797–1820

- 540 Gardner JK, Knopoff L (1974) Is the sequence of earthquakes in Southern California, with aftershocks removed,
541 Poissonian? *Bull Seis Soc Am* 64:1363–1367
- 542 Ghasemi H, Zare M, Fukushima Y (2008) On the scattering in normal distribution of the peak ground accel-
543 eration residuals in Alborz region, 14th World Conference on Earthquake Engineering, October 12–17,
544 Beijing, China
- 545 Kakkalamos J, Baise LG, Boore DM (2011) Estimating unknown input parameters when implementing the
546 NGA ground-motion prediction equations in engineering practice. *Earthq Spectra* 27(4):1219. doi:10.1193/
547 1.3650372
- 548 Luzi L, Bindi D, Franceschina G, Pacor F, Castro RR (2005) Geotechnical site characterisation in the Umbria-
549 Marche area and evaluation of earthquake site-response. *Pure Appl Geophys* 162:2133–2161. doi:10.1007/
550 s00024-005-2707-6
- 551 Luzi L, Puglia R, Pacor F, Gallipoli MR, Bindi D, Mucciarelli M (2011) Proposal for a soil classification
552 based on parameters alternative or complementary to V_s , 30. *Bull Earthq Eng* 9(6):1877–1898. doi:10.
553 1007/s10518-011-9274-2
- 554 Menke W (1989) Geophysical data analysis: discrete inverse theory. In: Dmowska R, Holton JR (eds) *IntGeo-*
555 *phys series*, vol 45. Academic Press, New York, p 289
- 556 Rovelli A, Scognamiglio L, Marra F, Caserta A (2001) Egediffracted 1-sec surface waves observed in a
557 small-size intramountainbasin (Colfiorito, central Italy). *Bull Seismol Soc Am* 91:1851–1866
- 558 Scherbaum F, Schmedes J, Cotton F (2004) On the conversion of source-to-site distance measures for extended
559 earthquake source models. *Bull Seismol Soc Am* 94(3):1053–1069
- 560 Tatar M, Jackson J, Hatzfeld D, Bergman E (2007) The 2004 May 28 Baladeh earthquake (Mw 6.2) in the
561 Alborz, Iran: overthrusting the South Caspian Basin margin, partitioning of oblique convergence and the
562 seismic hazard of Tehran. *Geophys J Int* 170:249–261
- 563 White H (1980) A Heteroskedasticity-consistent covariance matrix estimator and a direct test for heteroskedas-
564 ticity. *Econometrica* 48(4):817–838
- 565 Yenier E, Sandikkaya MA, Akkar S (2010) Report on the fundamental features of the extended strong-motion
566 databank prepared for the SHARE project. Report WP4—strong ground-motion modeling. Project SHARE
567 <http://www.share-eu.org/node/73>