1	Photolytic Hazes in the Atmosphere of 51 Eri b
2	K. Zahnle
3	NASA Ames Research Center, Moffett Field, CA 94035
4	Kevin.J.Zahnle@NASA.gov
5	M. S. Marley
6	NASA Ames Research Center, Moffett Field, CA 94035
7	Mark.S.Marley@NASA.gov
8	and
9	C. V. Morley
10	Department of Astronomy and Astrophysics, University of California, Santa Cruz, CA 95064
11	cmorley@ucolick.org
12	and
13	J. I. Moses
14	Space Science Institute, 4750 Walnut Street, Suite 205, Boulder, CO 80301
15	jmoses@spacescience.org
16	ABSTRACT

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We use a 1D model to address photochemistry and possible haze formation in the irradiated warm Jupiter 51 Eridani b. The intended focus was to be carbon, but sulfur photochemistry turns out to be important. The case for organic photochemical hazes is intriguing but falls short of being compelling. If they form, they are likeliest to do so if vertical mixing in 51 Eri b is weaker than in Jupiter, and they would be found below the regions where methane and water are photolyzed. The more novel result is that photochemistry turns H_2S into elemental sulfur, here treated as S_8 . In the cooler models, S_8 is predicted to condense in optically thick clouds of solid sulfur particles, whilst in the warmer models S_8 remains a vapor along with several other sulfur allotropes that are both visually striking and potentially observable. For 51 Eri b, the division between models with and without condensed sulfur is at an effective temperature of

700 K, which is within error its actual effective temperature; the local temperature where sulfur condenses is between 280 and 320 K. The sulfur photochemistry we have discussed is quite general and ought to be found in a wide variety of worlds over a broad temperature range, both colder and hotter than the 650-750 K range studied here, and we show that products of sulfur photochemistry will be nearly as abundant on planets where the UV irradiation is orders of magnitude weaker than it is on 51 Eri b.

¹⁸ Subject headings: planetary systems — stars: individual(51 Eri b)

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1. Introduction

The star 51 Eridani is a pre-main-sequence F dwarf that is only 20 million years old. Directimaging observations with GPI (Gemini Planet Imager) reveal that the star is orbited by a selfradiant young Jupiter, designated 51 Eri b, that emits with an effective temperature on the order of $T_{\rm eff} = 700 \pm 50$ K (Macintosh et al. 2015). Thermal evolution models predict that a 20 Myr old jovian planet with 51 Eri b's luminosity will have mass ~ $2M_{\rm Jup}$ and radius ~ $1R_{\rm Jup}$ (Macintosh et al. 2015).

Comparison by Macintosh et al. (2015) of the available spectral and photometric data to 26 spectral models reveal that while the planet shows methane in absorption, methane is depleted 27 compared to thermochemical equilibrium. Carbon monoxide is therefore expected to be abundant 28 but available data do not yet constrain it. Spectral matching with radiative transfer models also 29 strongly suggest that clouds, possibly patchy, are present in the atmosphere (Macintosh et al. 2015). 30 However, the planet is cool enough that silicate clouds if present would be confined to levels deep 31 beneath the photosphere and thus unlikely to affect what can be seen. Clouds of salts like Na₂S 32 and NaCl are possible, but even these would be expected to be confined to levels beneath the 33 photosphere by the low temperature of the planet (Morley et al. 2012). 34

In this study we use a 1D chemical kinetics model to ask whether, and under what conditions. 35 photochemical hazes are likely to form in the atmosphere of 51 Eri b and perhaps be the agent 36 responsible for the observed particulate opacity. We consider two candidates, one familiar, the 37 other more novel. The familiar candidate is an organic photochemical haze loosely analogous to 38 the hazes seen over Titan, Pluto, or Beijing. Such hazes have been proposed by many workers, but 39 to date the case for them has been inconclusive (Moses 2014). We will find here that a reasonable 40 case for a photochemical organic haze in 51 Eri b can be made, but we do not follow the chain of 41 polymerization reactions to molecules big enough and refractory enough that we can prove that 42 condensates actually form. The novel candidate is sulfur. With sulfur we can follow a much shorter 43 chain of polymerization to the point where sulfur condenses. We will show here that a good case 44 can be made for the presence of photochemical sulfur clouds in the atmosphere of 51 Eri b. 45

⁴⁶ This paper begins with a brief review of some related previous work. We next reprise our own

⁴⁷ model. In section 4 we present results for models that span the parameter space in which 51 Eri ⁴⁸ b probably resides. We will find that for some of these parameters organic hazes might form, and ⁴⁹ for some parameters sulfur clouds will form, for some parameters both might form, and for some ⁵⁰ parameters neither kind of haze is likely to form. The important role of sulfur raises the issue that ⁵¹ much of the sulfur chemistry is very poorly known. In section 5 we perform a series of sensitivity ⁵² tests to examine how the model responds to alternative assumptions about sulfur's photochemistry.

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2. Previous Models

The possibility that photochemical organic hazes might be important in irradiated brown dwarfs was first raised by Griffith et al. (1998). It remains an open question.

The first exoplanet photochemical models showed that small hydrocarbons would not condense 56 in the solar composition atmospheres of hot Jupiters (Liang et al. 2003, 2004). Line et al. (2010, 57 2011) confirmed this result for hot Jupiters. They predicted the flowering of a rich disequilibrium 58 non-methane hydrocarbon (NMHC) photochemistry in the cooler (~ 800 K) and presumptively 59 metal-rich warm Neptune GJ 436b, but stopped short of concluding that the chemistry would 60 necessarily lead to smogs. Moses and coworkers (Moses et al. 2011; Visscher and Moses 2011; Moses 61 et al. 2013a,b; Moses 2014) extended this model to bigger molecules, concluding that "complex 62 hydrocarbons and nitriles might produce high-altitude photochemical hazes" (Moses 2014). On the 63 other hand, as Moses (2014) also points out, methane has not yet been seen in GJ 436b. 64

There are several other models of exoplanet thermochemistry and photochemistry that have 65 been used to address a variety of hydrogen-rich exoplanets, from Jupiters to Neptunes to super-66 Earths, but none of them go as far as predicting the photochemical production of organic hazes. 67 Venot et al. (2012) examined C-N-O photochemistry on HD 189733b and HD 209458b; Kopparapu 68 et al. (2012) explored the effect of the C/O ratio on the hot Jupiter WASP-12b; Venot et al. (2013)69 used high-temperature UV cross sections to study the effect of CO₂ photolysis on the warm Neptune 70 GJ 436b; Hu and Seager (2014) addressed temperature and elemental abundances in super-Earths 71 and mini-Neptunes, with application to GJ 1214b, HD 97658b, and 55 Cnc e; Agúndez et al. 72 (2014b) added tidal heating and metallicity variations to GJ 436b; Venot et al. (2014) looked at 73 temperature, metallicity, UV flux, tidal heating, and atmospheric mixing in warm Neptunes, with 74 application to GJ 3470b and GJ 436b; Miguel and Kaltenegger (2014) took into account stellar type 75 and orbital distance; Miguel et al. (2015) focused on Lyman α irradiation of GJ 436b and other 76 warm Neptunes; Koskinen et al. (2013) and Lavvas et al. (2014) addressed ion chemistry; Agúndez 77 et al. (2012, 2014a) used a 2D model to address the horizontal quenching that occurs when winds 78 carry hot air to cold places; and Benneke (2015) combined photochemistry with retrievals from 79 exoplanet transit spectra to mine for C/O ratios in several planets. 80

Two recent models do include heavier organic molecules (Rimmer and Helling 2016; Venot et al. 2015). Rimmer and Helling (2016) compile an extensive reaction network that includes both

neutral and ion chemistry; they pay particular attention to the formation of prebiotic molecules 83 like glycine, but they do not yet address photochemical hazes. Venot et al. (2015) have expanded 84 their reaction network to include selected hydrocarbons with as many as eight carbon atoms. A 85 plus is that their reaction network has been tested against combustion experiments. On the other 86 hand, it should be borne in mind that complex models of complex systems often achieve empirical 87 agreement by cancellation of errors, and that things can go awry when the model is applied to new 88 conditions. Venot et al. (2015) compute that cyclohexadiene (cC_6H_8 , an obscure but reasonably 89 stable molecule) is a major photochemical product in 500 K stratospheres, exceeding even acetylene 90 (C_2H_2) and CO in abundance. Although Venot et al. (2015) do not mention photochemical hazes, 91 it is obvious that cyclohexadiene is well along the path to building a heavy smog. However, the 92 stated pathway for cC_6H_8 formation goes through 93

$$C_2H_2 + C_2H_2 \rightarrow nC_4H_3 + H, \tag{R60r}$$

⁹⁴ a very endothermic reaction that we will encounter again in section 4.1.1 when we discuss its reverse. ⁹⁵ We estimate that the rate for R60r is $k_{60r} = 3 \times 10^{-13} e^{-33000/T} \text{ cm}^3/\text{s}$, which at 500 K is very close ⁹⁶ to never. It is hard to imagine how a reaction with such a huge activation energy could actually ⁹⁷ be a major factor in a planetary atmosphere.

We have used our own code to address photochemistry and thermochemistry in giant planets 98 and brown dwarfs Zahnle et al. (1995, 2009); Zahnle and Marley (2014). Early versions of this 99 code (2011 and earlier) had some issues with the implementation of thermochemical equilibrium 100 that were corrected after consultations with Channon Visscher. Miller-Ricci Kempton et al. (2012) 101 and Morley et al. (2013, 2015) used the corrected code to address photochemistry in the warm 102 $(T_{\rm eff}\approx 550$ K) super earth GJ 1214b and similar planets. They suggested that hazes should form 103 when reduced organic radicals like CH_3 (building blocks of bigger organic molecules) were more 104 abundant than OH. If so, NMHCs can be abundant enough that organic hazes show potential to 105 provide a viable alternative to clouds of other condensible substances such as Na_2S . However, as 106 with GJ 436b, methane has not been seen in GJ 1214b. 107

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3. Model Details

We use a vanilla 1D kinetics code to simulate atmospheric photochemistry. Such codes parameterize vertical transport as a diffusive process with an "eddy diffusion coefficient," denoted K_{zz} $[cm^2/s]$. Volume mixing ratios f_i of species *i* are obtained by solving continuity

$$N\frac{\partial f_i}{\partial t} = P_i - L_i N f_i - \frac{\partial \phi_i}{\partial z} \tag{1}$$

112 and diffusion

$$\phi_i = b_{ia} f_i \left(\frac{m_a g}{kT} - \frac{m_i g}{kT} \right) - \left(b_{ia} + K_{zz} N \right) \frac{\partial f_i}{\partial z} \tag{2}$$

equations for each species. In these equations N is the total number density (cm⁻³); $P_i - L_i N f_i$ represent chemical production and loss terms, respectively; ϕ_i is the upward flux; b_{ia} , the binary diffusion coefficient between i and the background atmosphere a, describes true molecular diffusion; and m_a and m_i are the molecular masses of a and i.

For the base model we use 481 forward chemical reactions and 42 photolysis reactions for 78 117 chemical species made from H, C, O, N, and S. We supplement these with 12 additional reactions 118 and two additional species for sensitivity tests. Every forward chemical reaction (e.g., $CO + OH \rightarrow$ 119 $CO_2 + H$) is balanced by the corresponding reverse reaction (e.g., $CO_2 + H \rightarrow CO + OH$) at a 120 rate determined by thermodynamic equilibrium. We have not included reverses of the photolysis 121 reactions; that is, we include reactions such as $H_2O + h\nu \rightarrow H + OH$, but we do not include 122 $H + OH \rightarrow H_2O + h\nu$ because radiative recombination of small molecules is typically slow, and our 123 chemical system does not include large molecules for which radiative attachment can be important 124 (Vuitton et al. 2012). 125

Organic photochemistry begins with photolysis of methane. Methane fragments can react with 126 each other to make more complicated organic molecules. Non-methane hydrocarbons (NMHCs) 127 with unsaturated bonds are in turn prone to polymerizing to form chains, rings, PAHs (polycyclic 128 aromatic hydrocarbons), and soots (disorganized agglomerations of PAHs and sheets of PAHs). In 129 this study we truncate NMHC chemistry at C_2H_n , with the exception of C_4H_2 . How we handle 130 C_4H_2 as a proxy for polymerization is discussed in detail in section 4.1.1 below. The more abundant 131 NMHC species in this model are C_2H_2 , C_2H_4 , C_2H_6 , C_4H_2 , H_2CO , CH_3OH , and HCN. The total 132 NMHC abundance is assessed as the total number of carbon atoms in the NMHCs and reported in 133 several figures below. 134

Sulfur photochemistry is the important new thing here. Sulfur photochemistry begins with 135 photolysis of, or chemical attack on, H_2S . Sulfur can be successively oxidized by OH (from H_2O 136 photolysis) to SO, SO₂, and SO₃ or H_2SO_4 . Sulfuric acid (H_2SO_4) is a major aerosol on Venus and 137 Earth worth looking for generally. Sulfur can also react with hydrocarbons to make CS, CS_2 , and 138 OCS. All three were abundant in the wake of the impacts of Comet Shoemaker-Levy 9 into Jupiter 139 in 1994 (Harrington et al. 2004). Finally, sulfur can polymerize, condense, and precipitate as the 140 element. The S_2 molecule was seen as a strong signature in the SL9 plumes (Moses et al. 1995; 141 Zahnle et al. 1995) and it has been seen in volcanic plumes over Io (Spencer et al. 2000). There is 142 strong circumstantial evidence in sulfur's isotopic record in Archean sediments that precipitation 143 of elemental sulfur was commonplace in the anoxic atmosphere of early Earth (Pavlov and Kasting 144 2002). Here we use a simplified system consisting of S, S_2 , S_3 , S_4 , and S_8 . As there is considerable 145 uncertainty in sulfur's reactions, we have listed our choices for key reactions in Table 1. Most of 146 the key reaction rates will be varied — and in one case, created — in sensitivity studies in section 147 5 below. All small sulfur-bearing molecules are rather easily photolysed but the sulfur rings -148 here gathered together under the master ring S_8 — are more stable to UV (Young et al. 1983; 149 Kasting et al. 1989; Yung et al. 2009). Thus, as we shall see, there is a strong tendency for sulfur 150 to polymerize to S_8 under UV radiation. 151

The background atmosphere is assumed to be 84% H₂ and 16% He. The relative abundances of

C, N, O, and S are presumed solar and to scale as a group according to metallicity; scavenging of O 153 and S by silicates and chalcophiles is taken into account (Lodders and Fegley 2006). For simplicity 154 we assume solar metallicity in the base models (the star 51 Eridani itself is very slightly subsolar, 155 [Fe/H] = -0.027). We consider one set of models with metallicity that is a Jupiter-like $3 \times$ solar. 156 It is not immediately obvious that higher metallicity always favors haze formation, despite the 157 greater abundance of haze-forming elements. Indeed, in atmospheres where CH₄ is less abundant 158 than CO, raising metallicity reduces the CH₄/CO ratio, and hence can make organic haze formation 159 less favorable. Here we will find that raising the metallicity from solar to $3 \times$ solar in 51 Eri b has 160 a negative effect on NMHC formation. 161

¹⁶² 51 Eridani is a bright star that was observed decades ago by the International Ultraviolet ¹⁶³ Explorer (IUE). We use the observed UV spectrum for $115 < \lambda < 198$ nm, the range of wavelengths ¹⁶⁴ for which data are available. For $\lambda > 198$ nm we use a standard stellar model photosphere for an ¹⁶⁵ FOIV star of radius $1.6R_{\odot}$, which makes the star's luminosity appropriate to 51 Eridani itself. We ¹⁶⁶ note in passing that the UV irradiation of 51 Eri b is about twice what it is at Earth today, or ¹⁶⁷ about $200 \times$ what it is at Titan.

An important simplification is that we treat vertical mixing by an eddy diffusion parameter K_{zz} that does not vary with height. What K_{zz} should be in a stratified atmosphere like that of 51 Eri b is not well-constrained (Freytag et al. 2010). Values ranging from 10³ cm²/s at the top of the troposphere to $10^6 - 10^7$ cm²/s at the top of the stratosphere seem to be useful for Jupiter (Moses et al. 2005), and values as high as 10^{10} cm²/s have been suggested for hot Jupiters. Here we consider $10^5 \leq K_{zz} \leq 10^{10}$ cm²/s.

We set surface gravity to $g = 32 \text{ m/s}^2$ in the nominal model. To test the response of the model to different gravities we consider $g = 56 \text{ m/s}^2$ as a variant. These bracket what is expected for 51 Eri b; $g = 32 \text{ m/s}^2$ is not better than $g = 56 \text{ m/s}^2$. The higher gravity models are cooler at a given pressure and thus are more favorable to CH₄ and to sulfur condensation.

The pressure-temperature profile is computed by a radiative-convective equilibrium model 178 assuming a cloud-free atmosphere. In the troposphere these assumptions produce a relatively cool 179 model. Unlike the thermal structure of the troposphere, which is governed by the planet's own 180 luminosity, temperatures at very high altitude depend also on heating by the star. Here we simply 181 extend an isothermal atmosphere to altitudes above the top of the radiative-convective model. 182 This is an important limitation on our models: we don't know the temperature well enough to 183 categorically state that sulfur does or does not condense in 51 Eri b. The temperature structure of 184 a sulfurous atmosphere is a big enough topic that it is best deferred to a future study. 185

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4. Results

¹⁸⁷ We begin with a particular model that illustrates the general features of 51 Eri b photochem-¹⁸⁸ istry. We then look at how the models respond to parameter variations. 189



Fig. 1.— Photochemistry in a nominal 51 Eri b model ($T_{\rm eff} = 700$ K, g = 32 m s⁻², solar metallicity, cloud-free atmosphere, $K_{zz} = 10^7$ cm²s⁻¹). The top and bottom rows differ in how C₄H₂ is treated. How C₄H₂ is treated has little effect on the more abundant molecules. *Left.* Carbon and oxygen. In the top panel, "C₄H₂" is treated as the gateway to C₂H₂ polymerization. Where "C₄H₂" is more abundant than acetylene (C₂H₂), our chemical scheme has broken down. In the bottom panel, C₄H₂ is chemically recycled. *Right.* Sulfur shows a rich photochemistry that tends to build toward the relatively photolytically stable S₈ molecule. This particular model is about 5 K too warm for S₈ to condense. Abundances of SO, CS, and S in the upper stratosphere will be smaller than shown here if sulfur condenses. Note that S₄ is abundant at the interface between H₂S and S₈.

¹⁹⁰ The particular model documented in Figure 1, which we call the nominal model, assumes an

effective temperature $T_{\rm eff} = 700$ K, an eddy diffusivity of $K_{zz} = 10^7$ cm²s⁻¹, constant gravity 191 $g = 32 \text{ m/s}^2$, solar metallicity m = 1, and a cloud-free atmosphere. Figure 1 plots volume mixing 192 ratios of selected carbon-, oxygen-, and sulfur- bearing species as a function of altitude (pressure). 193 For carbon and oxygen we plot CO and CH_4 , the major oxidized photochemical product CO_2 , the 194 reduced photochemical products acetylene (C_2H_2) and C_4H_2 , the bleaching agents OH and O_2 , and 195 atomic H. For sulfur we plot most of the species that are abundant, although CS_2 and SO_2 are not 196 labeled and S_3 , which is coincident with S_4 but less abundant in these models, is omitted entirely 197 for clarity. We do not plot H_2O (the most abundant molecule other than H_2), atomic O, other 198 hydrocarbons, nor any N-bearing species. 199

Figure 1 illustrates the vertical structure of chemical products. The top of the atmosphere is 200 relatively oxidized by OH from H_2O photolysis, but it is also where CH_4 is photolyzed by Lyman 201 α , and so the top is also the primary source of small hydrocarbon radicals. Reactions with OH 202 are the chief competition to hydrocarbon polymerization because the CO bond once formed is 203 effectively unbreakable in the haze-forming region. Thus NMHC production is possible only when 204 OH is suppressed. OH is controlled by reaction with H_2 to reconstitute H_2O , or with CO to make 205 CO_2 ; this is why CO_2 is always a major photochemical product in all 51 Eri b models. Conditions 206 are more reduced at greater depth. 207

4.1.1. Alternative carbon polymerizations

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It is self-evident that hydrocarbon polymerization can ramify without any known limit, espe-209 cially in the presence of nitrogen and a little oxygen. In the bigger picture this is obviously a good 210 thing, but our modeling effort cannot ramify without limit. We must either be able to show that 211 abundances go to zero for molecules with more than a few carbon atoms, or we must artificially 212 truncate the system. If the atmosphere is sufficiently oxidized, the first option is workable. The 213 system will stop at CO_2 without much of interest happening — this has historically been the bane 214 of terrestrial prebiotic atmospheric chemistry models (Abelson 1966; Pinto et al. 1980). But here 215 we are dealing with H₂-rich atmospheres and it is not obvious a priori that the chemistry converges. 216

In this study we truncate the system at C_4H_2 , the first molecule to form as the product of two C₂H_n molecules. The state of the art in exoplanets takes the chemistry up to C₈H_n (Moses 2014; Venot et al. 2015; Rimmer and Helling 2016), but only a tiny fraction of all possible C_mH_n (m \leq 8) can be taken into account, and the combinatorial nature of the chemistry rapidly approaches or exceeds the limit of what can be done with a detailed chemical kinetics model. Further progress requires working with a limited number of generic or representative species. We consider two extreme assumptions that might bound the problem.

In one set of numerical experiments we treat C_4H_2 as a bucket in which polymerizing carbon accumulates, rather than as an actual chemical species. The only loss is the reverse of the formation ²²⁶ reaction,

$$C_2H + C_2H_2 \to C_4H_2 + H. \tag{R57}$$

The underlying idea is that C_4H_2 is destined to grow into ever larger $C_mH_nN_xO_yS_z$ molecules by the addition of free radicals. When used in this way, we will from here forward put quotes on " C_4H_2 " to indicate that we are treating it as a representative species rather than as the real C_4H_2 molecule. This is the case documented by the upper left-hand panel of Figure 1 and in most other spaghetti plots in this paper.

In the other set of numerical experiments we add three chemical reactions with H to crack C_4H_2 : first an addition,

$$C_4H_2 + H + M \to C_4H_3 + M \tag{R58}$$

234 followed either by H-abstraction

$$C_4H_3 + H \to C_4H_2 + H_2 \tag{R59}$$

235 or by fission

$$C_4H_3 + H \to C_2H_2 + C_2H_2.$$
 (R60)

Reaction R58 is a fast reaction that has been studied both theoretically and experimentally (Eiteneer and Frenklach 2003; Klippenstein and Miller 2005); we use rates for k_{58} from the latter. The other two reactions are inventions. For R59, we assume that $k_{59} = 5 \times 10^{-11} \exp(-500/T) \text{ cm}^3/\text{s}$, which is not unusual for an H-abstraction, if perhaps a bit fast. For R60, the unusual reverse reaction R60r discussed above with respect to cyclohexadiene suggests that there ought to be a considerable activation barrier and a rather small collision factor to the reverse reaction to account for the special geometry that would seem required. We assume that

$$k_{60} = 5 \times 10^{-11} \exp\left(-2000/T\right) \text{ cm}^3/\text{s.}$$
 (3)

The lower left-hand panel of Figure 1 shows that adding reactions R58-R60 to the network reduces the peak abundance of C_4H_2 and restricts the molecule to the photochemical region. Not shown is that if k_{60} is reduced by a factor of 30, the C_4H_2 altitude profile reverts to the " C_4H_2 " profile seen in the upper left panel of Figure 1.

We note that neither R59 nor R60 are likely to be important in reality. Much more likely is that the reaction with H will be another addition (Harding et al. 2007) and the hydrocarbon will continue to grow,

$$C_4H_3 + H + M \to C_4H_4 + M, \tag{R61}$$

with no natural truncation point in the photochemical region where C-bearing radicals are also abundant; that is, additions and ramifications will continue, and there is no obvious end to this. From this perspective " C_4H_2 " is a gateway species. At greater depth in a hydrogen-rich atmosphere, hydrogenation will probably focus on the unsaturated carbon bonds until what is left is an alkane or alkanes, and in the end the alkanes will be hydrogenated to CH₄ and H₂, completing the cycle.

In most figures that follow we will show " C_4H_2 " profiles computed with the high C_4H_2 because these are more interesting to look at.

4.1.2. Sulfur photochemistry and sulfur condensation

The righthand panels of Figure 1 line up the sulfur chemistry with the carbon and oxygen chemistry in the nominal model. Several things stand out. The first is that H_2S — sulfur's stable form in the abyss — barely makes it past the tropopause. Although H_2S is susceptible to UV photolysis, that is not what is happening here. Rather, H_2S is being destroyed by atomic H flowing down from the high altitude photochemical source region,

$$H_2S + H \to HS + H_2. \tag{R23}$$

²⁶³ The HS radical reacts quickly with H to free S,

$$\mathrm{HS} + \mathrm{H} \to \mathrm{S} + \mathrm{H}_2,\tag{R9}$$

and atomic S reacts with HS to make S_2 ,

$$\mathrm{HS} + \mathrm{S} \to \mathrm{S}_2 + \mathrm{H},\tag{R8}$$

and the polymerization of sulfur has begun, which is the second thing to stand out: S_8 is very abundant, generally at a lower altitude than the NMHCs and under more reduced conditions.

The high predicted abundance of S_8 suggests that it might condense. Sulfur vapor is compli-267 cated by the presence of several allotropes. Our first simplification is to lump S_6 and S_7 together 268 with the more abundant S_8 . Lyons (2008) gives simple curve fits to many allotropes above the liq-269 uid, and then describes a scheme for extrapolating these to lower temperatures above solid sulfur. 270 A complication is that the vapor pressure curves given by Lyons (2008) are discontinuous by nearly 271 a factor of two at sulfur's melting point ($T_m = 398$ K). We use a blended approximation in which 272 the vapor pressure over the solid is extended to higher temperature until it intersects the reported 273 vapor pressure over the liquid, 274

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$$p_v(S_8) = \exp(20 - 11800/T)$$
 $T < 413 \text{ K}$
 $p_v(S_8) = \exp(9.6 - 7510/T)$ $T > 413 \text{ K}$ (4)

where the vapor pressure is in bars. In Figure 1, the S₈ mixing ratio is $\sim 2 \times 10^{-6}$ for atmospheric pressure levels between 100 µbars and 10 mbars. At these partial pressures, $2 \times 10^{-10} < p(S_8) < 2 \times 10^{-8}$, sulfur's condensation temperature is between 280 and 310 K. The uncertainty in Eq 4 in this temperature range is probably less than a factor of two (the coldest datum is at ~ 310 K), which is insignificant compared to the uncertainty in the temperature in our models. For context, the corresponding condensation temperatures for water are between 170 and 200 K at the same altitudes. At higher metallicity both condensation temperatures are $\sim 20 \log_{10}(m)$ K higher.

The vapor pressure of S_2 over solid or liquid sulfur is tiny (Lyons 2008),

$$p_v(S_2) = \exp(27 - 18500/T)$$
 $T < 413 \text{ K}$

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$$p_v(S_2) = \exp(16.1 - 14000/T)$$
 $T > 413$ K. (5)

All of our models of 51 Eri b predict more S_2 than would be consistent with the presence of condensed sulfur. Evidently S_2 (and S_3 and S_4 as well) would be drawn down to negligible amounts where S_8 condenses.

Saturation mixing ratios of S_8 and S_2 over solid sulfur are plotted on Figure 1. A third thing stands out: S_8 in 51 Eri b is very close to its condensation point. In this particular model S_8 does not condense, but if the model were a few degrees cooler it would condense. If S_8 condenses, we can presume that there would be much less S_2 , SO, CS, and S above the clouds than is shown here.

4.2. Dependence on vertical mixing

In this study vertical mixing is a free parameter. Figure 2 shows what happens when K_{zz} is 293 made much bigger or much smaller. These are high "C₄H₂" models. Strong vertical mixing (K_{zz} = 294 $10^9 \text{ cm}^2 \text{s}^{-1}$, top panels) creates a more oxidized environment at the top of the atmosphere that is 295 unfavorable to NMHC growth. In particular, "C₄H₂" is all but wiped out. Weak vertical mixing 296 $(K_{zz} = 10^5 \text{ cm}^2 \text{s}^{-1})$, bottom panels) is more favorable to NMHCs, especially at lower altitudes 297 that are too deep for oxidants to reach when the mixing is weak. This is somewhat obscured by 298 our plotting volume mixing ratios in Figure 2, which exaggerates the apparent importance of trace 290 species at high altitudes, and understates the importance of anomalies at $K_{zz} = 10^5 \text{ cm}^2 \text{s}^{-1}$. In 300 fact $K_{zz} = 10^5 \text{ cm}^2 \text{s}^{-1}$ is more conducive to hydrocarbon polymerization than is $K_{zz} = 10^7 \text{ cm}^2 \text{s}^{-1}$. 301

The effects of changing K_{zz} on sulfur are parallel to those on carbon but more exaggerated. 302 Strong vertical mixing (Figure 2, upper right-hand panel) enables H₂S to get higher before it gets 303 destroyed, which creates a more favorable environment for S_2 , which becomes rather abundant. If 304 S_8 does not condense, eddy mixing also lifts it to high altitudes where it is photolyzed and oxidized 305 to SO and SO_2 or reduced to CS. Weak vertical mixing (Figure 2, lower right-hand panel) squeezes 306 the sulfur photoproducts into a relatively thin region below the homopause and above the H_2S 307 destruction horizon at 30 mbars; the high molecular weight of S_8 prevents sulfur getting very high, 308 which markedly depletes the top of the atmosphere in all sulfur species even if sulfur does not 309 condense. 310

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4.3. Dependence on effective temperature

Figure 3 shows what happens when the effective temperature of the planet is raised or lowered by 50 K. These are high " C_4H_2 " models. The cooler atmosphere is clearly more oxidized. In carbon this is seen in the higher abundance of CO₂ and the lower abundances of C_2H_2 and " C_4H_2 ," in sulfur it is seen in higher abundance of SO and SO₂ and the disappearance of CS. The primary oxidant is OH from H₂O photolysis. The most important sink on OH is the temperature-sensitive



Fig. 2.— The effect of K_{zz} on carbon and sulfur photochemistry in our nominal 51 Eri b model. Mixing is $100 \times$ stronger (top row) and $100 \times$ weaker (bottom row) than in Figure 1 (these are both high "C₄H₂" models to be compared to the top panels of Figure 1). High vertical mixing creates a more oxidized environment at the top of the atmosphere that is less favorable to S₈ and very unfavorable to NMHC growth. Strong vertical mixing is more favorable to S₂, SO, and SO₂, less favorable to S₈. Weak vertical mixing produces a more reduced atmosphere that is more favorable to NMHCs and to S₈, which forms abundantly in deeper, warmer regions.

 $_{317}$ reaction with H_2

$$H_2 + OH \rightarrow H_2O + H_2$$



Fig. 3.— The effect of temperature on carbon and sulfur photochemistries. These are 50 K hotter (top) or colder (bottom) than the corresponding model high C_4H_2 model in Fig 1. Note that the $T_{eff} = 650$ K (lower panels) models are more oxidized and less favorable to NMHCs. Note that S_8 condenses in the cooler models but does not condense in the warmer.

which puts H₂O back together. The high abundance of H₂ in a solar composition gas ensures that the reaction with H₂ is the leading sink on OH for T > 200. It is only where T < 200 K that the temperature-insensitive reaction with CO,

$$\rm CO + OH \rightarrow CO_2 + H$$

becomes more important, but we do not encounter temperatures this low in 51 Eri b models. The reaction with H_2 becomes much slower as the temperature drops and consequently the OH ³²³ abundance becomes much larger as the temperature drops. In turn the higher OH abundance ³²⁴ promotes CO₂ formation and inhibits NMHC growth. Both trends are clearly seen in the lower ³²⁵ panels of Figure 3.

The other effect of temperature on sulfur is the obvious one that condensation becomes more likely in the cooler models. Sulfur readily condenses in the cooler $T_{\rm eff} = 650$ K model at around 3 mbar (Figure 3, lower right-hand panel). This is also the altitude where organic hazes would form if any do, if the proxy "C₄H₂" is a useful guide.

330

4.4. Overview of carbon chemistry

Figure 4 gives an overview of carbon photochemistry for solar composition models over the phase space of different T_{eff} , g, and K_{zz} pertinent to 51 Eri b. We consider temperatures of $T_{\text{eff}} = 750$ K and $T_{\text{eff}} = 650$ K in addition to the nominal model with $T_{\text{eff}} = 700$ K, and we consider a gravity of g = 56 m/s² in addition to the nominal g = 32 m/s². Figure 4 is restricted to solar metallicity and the UV radiation observed by *IUE*. We vary K_{zz} between 10⁵ cm²/s and 10¹⁰ cm²/s for all variants of T_{eff} and g.

Figure 4 plots the quenched disequilibrium CO and CH_4 mixing ratios and it plots the peak mixing ratios reached by the major photochemical products. 51 Eri b is near the boundary between CO-dominated and CH_4 -dominated atmospheres (in equilibrium, the carbon would almost entirely be in CH_4). Both gases are abundant in all models, although CO is more abundant in most of them. In general, CH_4 is most abundant when K_{zz} is small, or the gas cooler, or the gravity higher (Zahnle and Marley 2014). None of the cases are truly methane-rich.

Smaller values of K_{zz} are more favorable to photochemical NMHC formation and high values of 343 K_{zz} are very unfavorable. The apparently lower NMHC production at low values of K_{zz} is illusory, 344 a consequence of plotting peak mixing ratios in Figure 4. The peak occurs at higher pressure at 345 $K_{zz} = 10^5 \text{ cm}^2 \text{s}^{-1}$ than at $K_{zz} = 10^7 \text{ cm}^2 \text{s}^{-1}$, so that NMHC densities at $K_{zz} = 10^5 \text{ cm}^2 \text{s}^{-1}$ 346 are actually higher. Some of the trend with K_{zz} can be ascribed to the CH₄/CO ratio, but the 347 trend is even stronger in CO_2 , which suggests that the weaker mixing is also acting to isolate and 348 preserve the photochemical products. On the other hand, the relative dearth of NMHCs in the 349 cooler $T_{\rm eff} = 650$ K models is a real feature caused by the strong temperature dependence of the 350 $H_2 + OH \rightarrow H_2O + H$ reaction that holds OH in check. 351

352

4.5. Overview of sulfur chemistry

Figure 5 presents the corresponding overview of sulfur photochemistry. Here we count sulfur atoms, so that S_2 is counted doubly and S_8 is counted eight-fold. The symbols are not like pie charts. They do not show how sulfur is apportioned at any one height. Rather, they show each



Fig. 4.— Carbon photochemistry in some possible 51 Eri b's of solar metallicity subject to the observed *IUE* UV flux from 51 Eri a. Peak volume mixing ratios are plotted in proportion to the areas of the disks. CO (black) and CH₄ (green) are quenched disequilibrium abundances welling up from below. The photochemical products CO₂ (blue) and NMHCs (red, chiefly C_2H_2) are maxima found at higher altitudes where photochemistry is king.

category at its peak abundance, which in most cases are at different heights. In this way we see that, for example at $K_{zz} = 10^7$, almost all the sulfur transitions from H₂S to S₈ at higher altitudes, or that at $K_{zz} = 10^{10}$, almost all the sulfur that started in H₂S is found in S₂ higher up and then still higher up it is found as S.

Several trends are evident in Figure 5. One is that H_2S is quantitatively converted to elemental sulfur. For weaker vertical mixing the sulfur will pool in S_8 . The OCS molecule will be abundant. Strong vertical mixing favors S_2 and S. As with carbon, the cooler atmospheres are more strongly oxidized, but with sulfur the more strongly oxidized species are more prevalent when the vertical



Fig. 5.— Overview of sulfur photochemical products. Sulfur is grouped into relatively oxidized species (SO and SO₂), carbonized species (mostly OCS), and two allotropes of elemental sulfur (S₂ and S₈). The circles and semicircles represent maximum mixing ratios as the areas of the implicit disks. The outer ring is the mixing ratio of H₂S in the deep atmosphere. In cooler models S₈ is predicted to condense; in these half of the S₈ is colored green.

mixing is stronger because when mixing is weak S_8 settles out, as was seen in the bottom-right panel of Figure 2. About half the models predict that sulfur condenses in clouds.

366

4.6. Metallicity

Figure 6 illustrates the effect of higher metallicity m = 3 in the $g = 32 \text{ m/s}^2$ models as a function of K_{zz} for both carbon and sulfur chemistry. In this figure the carbon and sulfur mixing ratios are plotted to the same scale to facilitate cross-comparison, but as a consequence sulfur's circles are rather small. In order to see both S₂ and S, these are plotted as quarter circles. As is well-known, higher metallicity strongly favors CO and CO_2 over CH_4 . Higher metallicity has little effect on sulfur speciation because (i) all of its major products are metal-rich and (ii) its most abundant product, S_8 , is the metal-richest.



Fig. 6.— Higher metallicity (m = 3). Symbols and colors have the same meaning as in Figures 4 and 5 above. On this figure we plot sulfur and NMHC mixing ratios to the same scale to facilitate direct comparison.

4.7. Optical depths

It is possible that sulfur will be optically thick when it condenses, and it is possible that several of its optically-active allotropes will be visible when it does not. Here we estimate the optical depths of sulfur clouds and of sulfur vapor through the S_4 di-radical, and we give an optimistic estimate of the opacity from organic hazes.

374

For the sulfur clouds we gather all the S_8 above the condensation height into either 1 or 10 μ m diameter particles, a size range that seems appropriate for condensation clouds. We assume effective particle densities of 1.5 g/cm³. Both optical depths are shown in Figure 7 as gold and green disks, respectively. For S_4 we show the optical depth at 500 nm (red disks), which is near the center of its strong broadband visible light absorption.

For the organic hazes, we consider two cases. The first (black disks) is based on NMHC mixing ratios: we presume that 10% of the NMHCs (chiefly C_2H_2) go into haze particles at altitudes where the total mixing ratio of NMHC's exceeds 1 ppmv. The second case (gray disks) is based on the



Fig. 7.— Optical depths at a glance. We plot $1 - e^{-\tau}$ rather than τ itself to give a better graphical sense of how much light is blocked. The outer circle represents the incident light. The black and gray disks show two upper bounds on clouds of 100 nm diameter organic particles. Black is deduced from mixing ratios in the photochemical zone where acetylene peaks, while gray refers to number densities of "C₄H₂," which typically peaks deeper in the atmosphere. Gold shows clouds of 1 micron diameter sulfur particles and green shows clouds of 10 micron diameter sulfur particles. Cherry red disks show the optical depth of S₄ vapor at 500 nm. Where sulfur condenses, S₄ should condense too, and so it is not shown. Results are for solar composition; τ for sulfur scales with metallicity but τ of organic hazes may not.

computed number densities of " C_4H_2 " when treated as a portal through which every carbon that passes ultimately gets incorporated in a haze particle. For both cases we assume that organic particles are 100 nm diameter and of effective density 0.6 g/cm³.

The results of the exercise are presented in Figure 7 for the same range of solar composition models discussed above. Sulfur is in the top hemisphere of each circle and carbon in the bottom

half. There appears to be considerable potential for sulfur to be optically significant in 51 Eri b. 392 This can be as sulfur clouds if 51 Eri b is a cool object, or as vapor if sulfur does not condense. The 393 sulfur clouds can be optically thick at solar metallicity, and they could be significantly thicker on 394 planets because sulfur optical depths will scale linearly with metallicity. The sulfur vapors can also 395 be important, especially the chains. E.g., S_4 absorbs strongly at 500 nm, and longer chains absorb 396 to 750-850 nm (Meyer 1976) (the rings, which confer invisibility, typically absorb $\lambda < 330$ nm). It is 397 well known that liquid sulfur when heated turns from light yellow to dark red as S_8 rings decompose 398 into a soup of chains and rings (the depth of red depends on impurities, especially hydrocarbons 399 (Moses and Nash 1991)). We might expect similar behavior in 51 Eri b as S_8 rings thermochemically 400 decompose between 10 and 100 mbars. On the other hand sulfanes, alkane analogs with the general 401 formula HS_nH , may be the intermediaries between S_8 and H_2S ; like the rings, sulfanes typically 402 absorb $\lambda < 330$ nm (Meyer 1976). 403

There is also some potential for organic hazes to be important, especially where K_{zz} is small, but this potential is model dependent. At high altitudes where CH₄ and H₂O are photolyzed, optical depths near unity (black disks) are achievable only if conversion of acetylene into PAHs is highly efficient, which seems unlikely. Lower altitudes that coincide with the more reduced S_n-H₂S photochemistry are more promising, but interpreting "C₄H₂" as a bucket full of particles is a leap that future work could prove baseless. A difference from sulfur is that we do not expect that modestly higher metallicity will lead to more organic haze.

411

5. Sensitivity of the results to model uncertainties

We have found that most of our models predict that S_8 is a major product of sulfur photolysis 412 (Figure 5). We have also found that NMHC formation is sensitive to sulfur photochemistry. We 413 have discussed truncation of hydrocarbon chemistry at C_4H_2 above. Here we perform a series of 414 tests to determine how sensitive the model is to other uncertain or unknown factors. These are (i) 415 different amounts of stellar ultraviolet radiation; (ii) different rates of S_8 photolysis; (iii) different 416 estimates of H_2S thermolysis and recombination; (iv) different rates of sulfur polymerization; and 417 (v) unknown chemical reactions that would compromise S_8 's stability. The latter proves the matter 418 of most concern. 419

420

5.1. Sensitivity to UV

In Figure 8 we have explored the sensitivity of the nominal model to reduced levels of UV radiation. With UV irradiation at 10% that in the nominal model, the general pattern of the photochemistry is similar to that in the nominal model. Chief differences are that there is less CO_2 and C_2H_2 , H_2S reaches higher altitudes before it is destroyed, and there is a modest shift away from S_8 as the chief product. Even when the UV is reduced to 0.1% that of the nominal model, there



Fig. 8.— 51 Eridani b models with reduced and greatly reduced UV irradiation. *Left.* UV irradiation is 10% that in the nominal model. *Right.* UV irradiation is reduced to 0.1% that in the nominal model. Photochemical CO_2 and C_2H_2 nearly disappear, but a rich sulfur photochemistry remains.

are enough photons for H_2S to be fully consumed and a complete suite of sulfur photochemical products is generated. It is only when UV irradiation is reduced by another factor of ten that most of the H_2S survives and the sulfur photochemistry becomes photon-limited.

429

5.2. Sensitivity to S_8 photolysis

We have used Young et al. (1983)'s method for estimating S_8 's photolysis rate. Young suggested that the first UV photon absorbed cleaves the ring. The resulting linear S_8 molecule can either be put back into the form of a ring by a collision, or it can be broken into two pieces (here both S_4) by absorbing a visible light photon. We assume an absorption cross section of 3×10^{-18} cm² to visible light ($\lambda < 850$ nm, Meyer 1976). The effective photolysis rate is

$$P(S_8) = P(S_{8,r}) \frac{N_c P(S_{8,l})}{N_c P(S_{8,l}) + \nu_c},$$
(6)

where $P(S_{8,r})$ and $P(S_{8,l})$ are the photolysis rates of the ring and linear S_8 molecules, respectively; N_c is the number of collisions required to close the ring; $\sigma_c = 3 \times 10^{-15}$ cm² is the collision cross section of a molecule; and $\nu_c = N \sigma_c \bar{\nu}$ is the collision frequency in terms of the mean thermal speed $\bar{\nu}$. In the nominal model we take $N_c = 1$. For the sensitivity test (Figure 9) we take $N_c = 30$. The chief consequence of higher S_8 photolysis is that catalytic sulfur is more abundant and NMHC

440 yield is reduced.



Fig. 9.— Two sensitivity tests to compare to Figure 1. Left. Enhancing the efficiency of S_8 photolysis at low pressures by raising N_c in Eq 6 also generates more S-bearing free radicals that reduce the yield of C_2H_2 and other NMHCs at the top of the atmosphere. Right. This model uses the faster rate k'_{22r} for the spin-forbidden insertion reaction $H_2 + S(^{3}P) + M \rightarrow H_2S + M$. The model is quite sensitive to this. With the faster rate H_2S survives to much higher altitudes than in the nominal model (Figure 1) and S_8 is less abundant and restricted to higher levels in the atmosphere. Photolysis of other sulfur-containing small molecules generates S-bearing free radicals that reduce the yield of C_2H_2 and other NMHCs.

441

5.3. Sensitivity to H₂S recombination

Rates of many of the chemical reactions that involve sulfur are poorly known. In particular, 442 a major source of model pathology is the 3-body recombination of H_2S , either from HS and H, 443 or from H_2 and S. There is limited information on H_2S recombination, but the reverse process, 444 thermolysis of H_2S , is industrially important and has been the subject of several experiments that 445 elude easy consensus (Bowman and Dodge 1977; Roth et al. 1982; Tesner et al. 1990; Woiki and 446 Roth 1994, 1995a; Olschewski et al. 1994; Shiina et al. 1996, 1998; Karan et al. 1999). Measured 447 rates from high temperature (1800 < T < 3500 K) shock tube experiments in Ar (Bowman and 448 Dodge 1977; Woiki and Roth 1994, 1995a; Olschewski et al. 1994; Shiina et al. 1996, 1998) differ 449 among themselves by an order of magnitude; it is not clear why. Moreover, lower temperature 450 (800 < T < 1400 K) flow reactor experiments in N₂ (Tesner et al. 1990; Karan et al. 1999) imply 451 rates that are 100-300 times higher than extrapolation of the shock tube data predict. 452

453 It was at first presumed that the dominant decomposition channel was

$$H_2S + M \to HS + H + M \tag{R21}$$

 $_{454}$ as with H_2O , because the alternative

$$H_2S + M \to H_2 + S(^{3}P) + M, \qquad (R22)$$

although much less endothermic, is spin forbidden (Roth et al. 1982). But parallel shock tube experiments by Woiki and Roth (1994), who monitored $S(^{3}P)$ production, and Olschewski et al. (1994), who monitored H₂S disappearance, gave a consistent picture of the spin-forbidden path being dominant. The straightforward, thermodynamically self-consistent reverse reaction

$$H_2 + S(^{3}P) + M \rightarrow H_2S + M$$
(R22r)

459 was therefore predicted to be fast at low temperatures. The possibility that the interesting reaction

⁴⁶⁰ R22r might be fast motivated follow-up experiments by Woiki and Roth (1995a) and Shiina et al.

(1996, 1998) to directly determine the reaction rate between $S(^{3}P)$ and H_{2} . Shiina et al. (1998)

found that, for T > 900 K, R22r is negligible compared to the competing abstraction reaction

$$H_2 + S(^{3}P) \to HS + H. \tag{R9r}$$

Shiina et al. (1998) do not dispute that R22 is the more important thermolysis channel for H_2S , but they change the extrapolation to low temperatures to take into account the considerable energy barrier that they computed,

$$k_{22} = 8.9 \times 10^{-7} \left(T/300 \right)^{-2.61} \exp\left(-44640/T \right).$$
(7)

⁴⁶⁶ The rate we use for R22r in our standard models is the reverse of k_{22} ,

$$k_{22r} = 1.4 \times 10^{-31} \, (T/300)^{-1.9} \exp\left(-8140/T\right),\tag{8}$$

which is far below the upper bound determined by Shiina et al. (1998) and very slow (but not negligible) at low temperatures.

For the sensitivity test (Figure 9, right-hand panel) we use a parallel pair of rates that are consistent both with the higher thermolysis rates reported by Olschewski et al. (1994) and Woiki and Roth (1994) and with the lower activation energy estimated by Olschewski et al. (1994):

$$k'_{22} = 8.9 \times 10^{-7} \, (T/300)^{-2.61} \exp\left(-38800/T\right) \tag{9}$$

472 with reverse

$$k'_{22r} = 1.4 \times 10^{-31} \left(T/300 \right)^{-1.9} \exp\left(-2300/T \right).$$
(10)

The rate k'_{22r} is comparable to the upper bound reported by Shiina et al. (1998). Although much slower than the rate that Shiina et al. (1998) had hoped to see, k'_{22r} is fast enough to affect our results significantly (Figure 9). With k'_{22r} , H₂S reaches altitudes 3 scale heights above where it gets to with k_{22r} .

477 We do not attempt to take into account the flow reactor data. These experiments suggest thermolysis rates that are orders of magnitude faster than either k_{22} or k'_{22} at 1000 K, and therefore 478 the recombination reactions must also be. However, we were unable to reproduce Karan et al. 479 (1999)'s argument that the different reported rates can be brought into agreement. We favor the 480 shock tube data because the flow reactor system is more complicated (more reactions need to be 481 taken into account) and less straightforwardly interpreted. E.g., what Karan et al. (1999) actually 482 measured is whether the system has had time enough to reach thermochemical equilibrium, which 483 isn't quite the same thing as determining a particular reaction rate. In the end, we think that the 484 slower rates k_{22r} and k'_{22r} are more plausible given the extensive molecular rearrangements that 485 must occur if an unlikely-looking reaction like R22r is to take place 486

487

5.4. Sensitivity to S_n polymerization

Our nominal sulfur polymerization scheme is mostly encompassed by reactions R2-R8 in Table 1. We have kept the system simple because the reactions and rates are very uncertain. Our rates are similar to those used elsewhere (e.g., Moses et al. 2002; Yung et al. 2009) and are not inconsistent with the few experimental reports (Fair and Thrush 1969; Langford and Oldershaw 1972, 1973; Nicholas et al. 1979). For the sensitivity tests we raise [lower] the rates of

$$S + S_3 + M \to S_4 + M \tag{R4}$$

493 and

$$S_2 + S_2 + M \to S_4 + M \tag{R5}$$

⁴⁹⁴ by a factor of 10, and raise [lower] the rate of

$$S_4 + S_4 + M \to S_8 + M \tag{R7}$$

⁴⁹⁵ by a factor of 100. These two cases of faster and slower polymerization are illustrated in left- and ⁴⁹⁶ right-hand panels of Figure 10, respectively. The figure shows that our model is not very sensitive ⁴⁹⁷ to reasonable uncertainties in the sulfur polymerization rate.

498

5.5. Sensitivity to unknown mechanisms of S_8 chemical destruction

 $_{499}$ Other than photolysis, the main sink of S_8 in our basic model is thermal destruction

$$S_8 + M \to S_4 + S_4 + M. \tag{R7r}$$

This is predicted to be rather fast, because ΔH is a relatively modest 150 kJ/mol, there is a considerable gain in entropy, and the rate for the forward reaction R7 is probably fast. We know of



Fig. 10.— Sensitivity of results to rates of S_n polymerization (to be compared to Figure 1). Left. Faster polymerization decreases the abundances of S and S₂ without noticeably changing S₈, because in the nominal case most of the sulfur was already pooling in S₈. Right. Slower polymerization increases the abundances of S and S₂, but not enough in this model to noticeably affect the S₈ abundance. The greater abundance of S-bearing radicals causes the acetylene (C₂H₂) yield to shrink.

⁵⁰² no reported kinetic data regarding S_8 's reactions. Yet S_8 should be reactive because reactions with ⁵⁰³ important free radicals should be significantly exothermic. To test the sensitivity of our model ⁵⁰⁴ to these unknown reactions, we need to invent both the reactions and the products. The most ⁵⁰⁵ abundant free radical by far is H, which makes reactions with H the likeliest to be important.

At high pressures we might expect a 3-body reaction to unmake the ring into a quasi-linear HS₈ radical, which would then be followed by reactions that either return S₈ or cleave the S₈ chain. We have not pursued this strategy here because (i) we would have to invent many species and many rates and (ii) our rate for R7r is pretty fast at high pressures.

At low pressures any plausible reaction would have to cleave the chain in two places. The invented reaction that adds the least new complexity to our model is

$$S_8 + H \to HS_4 + S_4, \tag{R51}$$

because we need to add only one invented species, HS₄. We estimate a standard heat of formation of 110 kJ/mol and standard entropy of 330 J/mol/K by analogy to HS₂ (Benson 1978). The invented reaction R51 is therefore substantially exothermic but undoubtedly faces a considerable energy barrier. We consider a slow rate

$$k_{\rm slow} = 3 \times 10^{-12} \exp\left(-5000/T\right) \tag{11}$$

516 and a fast rate

$$k_{\text{fast}} = 1 \times 10^{-11} \exp\left(-2500/T\right).$$
 (12)

For the temperature dependence we use Pauling's rule of thumb (Pauling 1970, p. 568) that the activation barrier of a radical-molecule reaction is about 8% of the bond energy of the bond to be broken. The slow rate presumes that the two S-S bonds are additive.

We then need a set of reactions for HS_4 . One category of reaction will be the H-abstraction reactions with H, OH, and some other radicals, such as

$$H + HS_4 \to H_2 + S_4. \tag{13}$$

These will probably have small activation barriers. At high altitudes S_4 will be promptly photolysed by visible light, so that the sulfur chain is quickly broken down. The other representative category will be molecular rearrangements that reconstitute H_2S , such as

$$H + HS_4 \to H_2S + S_3. \tag{14}$$

Reactions of this type face considerable activation barriers but can be important at depths where downwelling S_8 is converted back to H_2S . For these we use the same 2500 K activation barrier that we used for breaking the S-S bond.

Our expectation had been that deep thermal recycling of S_8 would be much sped up by the 528 new chemistry, but this is not really evident in Figure 11. Rather, the greater impact of the new 529 chemistry is to convert S_8 in the upper atmosphere into other more active species, and finally to 530 atomize it. The more abundant S-containing radicals catalyze the oxidation of organics. On the 531 other hand our chemical schemes do not encompass the speculative possibility that sulfur might also 532 catalyze carbon polymerization. In summary, what we don't know about sulfur chemistry appears 533 to have relatively little impact on whether S_8 forms, but there appears to be a strong impact on 534 carbon chemistry. If sulfur does not condense, the fast rate k_{fast} for S₈ destruction does not bode 535 well for organic hazes. Prospects for organics then become better in cooler atmospheres because 536 sulfur condensation would deplete S-containing radicals above the cloudtops. 537

538

6. Discussion

Photochemical hazes are widespread in the solar system but they are not yet established as fact on any actual exoplanet. Observations do not go much beyond showing that many exoplanetary spectra require a broad-band opacity resembling that of clouds. What these clouds might be made of has been a problem for theory (Morley et al. 2015), but as many substances can condense, it is reasonable to expect that there are many kinds of cloud.

Our purpose when we began this study was to make a case for organic hazes on the particular planet 51 Eri b. The idea was to use C_4H_2 , the first product of acetylene polymerization, as a proxy



Fig. 11.— Chemical sensitivity tests. Top. With the slow rate k_{slow} (Eq 11) for H+S₈, the nominal model is little altered, although there is notably more atomic S at high altitudes. Bottom. With the fast rate k_{fast} (Eq 12), the model looks rather different, with S₈ eliminated above 100 µbars and much increased abundances of photochemically active S-containing radicals and molecules. The overall character of the upper atmosphere is more oxidized, acetylene is much reduced, and the proxy "C₄H₂" is nearly wiped out.

for further polymerization: every " C_4H_2 " formed was assumed to eventually become incorporated into a particle. The quotes on " C_4H_2 " indicate that we are no longer talking about C_4H_2 the molecule but instead about everything downstream from it. But even in those cases where we have clearly tipped the scales to favor " C_4H_2 ," it only becomes more abundant than C_2H_2 at depths well below the primary photochemical region, where conditions are more reducing. It doesn't help our case that when we treat C_4H_2 as an actual molecule subject to cracking, we find that there isn't all that much of it there. In summary, the case for soots is intriguing but falls short of being compelling.

⁵⁵⁴ On the other hand we have rediscovered the importance of sulfur. Sulfur can have a dispropor-⁵⁵⁵tionate influence on photochemistry because most S-bearing species are relatively easily photolyzed ⁵⁵⁶by UV photons with $\lambda > 200$ nm, which 51 Eri a, an F star, emits copiously. Thus sulfur pho-⁵⁵⁷tochemistry becomes a major source of free radicals that can catalyze other chemistries. One ⁵⁵⁸consequence of sulfur catalysis is a tendency to drive carbon away from the disequilibrium NMHCs ⁵⁵⁹and toward the stronger bonds of CO and CO₂.

More interesting is that sulfur itself can be the photochemical cloud that we are looking for. 560 We find that for a wide range of conditions the major photochemical product of sulfur in a planet 561 like 51 Eri b is the ring molecule S_8 , which typically forms at ~ 10 mbars and extends up to 100 562 μ bars. The overall sulfur cycle is simple: H₂S flows up and S₈ and H₂ flow down. In the cooler 563 half of our models sulfur condenses to make a photochemical haze that, depending on particle 564 size, can be optically thick, while in the warmer half of the models sulfur remains in the vapor 565 phase. The sulfur vapor itself might also be optically important, especially at the interface between 566 abyssal H_2S and S_8 , where the latter thermally decomposes into a wide range of optically active 567 molecules that are eventually hydrogenated to recombine H_2S . The sulfur photochemistry we have 568 discussed in this paper is quite general and ought to be found in a wide variety of worlds over a 569 broad temperature range, both much cooler and much hotter than the 650-750 K range studied 570 here, and will be present on planets where the UV irradiation is very weak. Sulfur clouds should 571 be found in many of these. Whether 51 Eri b itself is cold enough for sulfur to condense cannot 572 be answered until radiative transfer models incorporate sulfur vapors and sulfur clouds, which is a 573 project beyond the scope of this paper, or until the yellow clouds are seen. 574

575

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REFERENCES

Abelson, P. (1966). Chemical events on the primitive Earth. <u>Proc. Natl. Acad. Sci. USA 55</u>, 1365-1372.

- Agúndez, M., Venot, O., Iro, N., Selsis, F., Hersant, F., Hébrard, E., Dobrijevic, M., (2012). The impact of atmospheric circulation on the chemistry of the hot Jupiter HD 209458b. <u>Astron.</u> Astrophys. 548, A73.
- Agúndez, M., Parmentier, V., Venot, O., Hersant, F., Selsis, F., (2014). Pseudo 2D chemical model
 of hot-Jupiter atmospheres: application to HD 209458b and HD 189733b. <u>Astron. Astrophys.</u>
 564, A73.
- Agúndez, M., Venot, O., Selsis, F., Iro, N., (2014). The Puzzling Chemical Composition of GJ 436b's Atmosphere: Influence of Tidal Heating on the Chemistry. Astrophys. J. 781, 781.
- Anders, E., Grevasse, N., (1989). Abundances of the elements: Meteoritic and solar. <u>Geochim.</u>
 Cosmochim. Acta 53, 197-214.
- Atkinson, R., Baulch, D.L., Cox, R.A., Crowley, J.N., Hampson, R.F., Hynes, R.G., Jenkin, M.E.,
 Rossi, M.J., Troe, J. (2004), Evaluated kinetic and photochemical data for atmospheric
 chemistry: Volume I gas phase reactions of Ox, HOx, NOx and SOx species. Atmos. *Chem. Phys.* 4, 1461-1738.
- Benneke, B., (2015). Strict Upper Limits on the Carbon-to-Oxygen Ratios of Eight Hot Jupiters
 from Self-Consistent Atmospheric Retrieval. ArXiv e-print 1504.07655.
- Benson, S.W. (1978). Thermochemistry and kinetics of sulfur-containing molecules and radicals.
 Chem. Rev. 78, 23-35. DOI: 10.1021/cr60311a003.
- Bowman, C.T., Dodge, L.G., (1977). Kinetics of the Thermal Decom- position of Hydrogen Sulfide
 Behind Shock Waves. 16th Symp. (Int.) Combust. 16, 971.
- Darwin, D.C., Moore, C.B. (1995). Reaction rate constants (295 K) for ³CH2 with H2S, SO2,
 and NO2: upper bounds for rate constants with less reactive partners. J. Phys. Chem. 99,
 13467-13470.
- DeMore, W.B., Sander, S.P., Golden, D.M., Hampson, R.F., Kurylo, M.J., Howard, C.J., Ravishankara, A.R., Kolb, C.E., Molina, M.J. (1997). <u>Chemical kinetics and photochemical data</u>
 for use in stratospheric modeling. Evaluation number 12. JPL Publication 97-4.
- ⁶¹¹ Du, S.Y., Francisco, J.S., Shepler, B.C., Peterson, K.A. (2008). Determination of the rate con ⁶¹² stant for sulfur recombination by quasiclassical trajectory calculations. <u>J. Chem. Phys. 128</u>,
 ⁶¹³ 204306.
- Eiteneer, B., Frenklach, M., (2003). Experimental and Modeling Study of Shock-Tube Oxidation
 of Acetylene. Int. J. Chem. Kinet. 35, 391-414.
- Fair, R.W., Thrush, B.A. (1969). Mechanism of S2 chemiluminescence in the reaction of hydrogen atoms with hydrogen sulphide. Trans. Faraday Soc. 65, 1208-1218.

- Freytag, B., Allard, F., Ludwig, H.-G., Homeier, D., Steffen, M. (2010). The role of convection,
 overshoot, and gravity waves for the transport of dust in M dwarf and brown dwarf atmospheres. Astron. Astrophys. 513, A19 (14 pp.).
- Griffith, C.A., Yelle, R.V., Marley, M.S. (1998). The Dusty Atmosphere of the Brown Dwarf Gliese
 229B. Science 282, 2063-2065.
- Hansen, N., Klippenstein, S.J., Taatjes, C.A., Miller, J.A., Wang, J., Cool, T.A., Yang, B., Yang, R.,
 Wei, L., Huang, C., Wang, J., Qi, F., Law, M.E., Westmoreland, P.R., (2006). Identification
 and Chemistry of C4H3 and C4H5 Isomers in Fuel-Rich Flames. J. Phys. Chem. A, 110,
 3670-3678.
- Harding, L.B., Klippenstein, S.J., Georgievskii, Y., (2007). On the combination reactions of hydro gen atoms with resonance-stabilized hydrocarbon radicals. J. Phys. Chem. A 111, 3789-3801.
- Harrington, J, de Pater, I., Brecht, S.H., Deming, D., Meadows, V., Zahnle, K.J., Nicholson, P.
 (2004). Lessons from Shoemaker-Levy 9 about Jupiter and Planetary Impacts. In <u>Jupiter</u>:
 <u>The Planet, Satellites and Magnetosphere</u>, F. Bagenol, T. Dowling, W. McKinnon, eds.
 Cambridge Univ. Press. pp. 158-184.
- Hills, A.J., Cicerone, R.J., Calvert, J.G., Birks, J.W. (1987). Kinetics of the reactions of S2 with
 O, O2, O3, N2O, NO, and NO2. J. Phys. Chem. 91, 1199-1204.
- Hu, R., Seager, S., (2014). Photochemistry in Terrestrial Exoplanet Atmospheres. III. Photochemistry and Thermochemistry in Thick Atmospheres on Super Earths and Mini Neptunes.
 Astrophys. J. 784, 63.
- Hubeny, I., Burrows, A. (2007). A systematic study of departures from chemical equilibrium in the
 atmospheres of substellar mass objects. Astrophys. J. 669, 1248-1261.
- Karan, K., Mehrotra, A.K., Leo A. Behie, L.A., (1999). On reaction kinetics for the thermal
 decomposition of hydrogen sulfide. Amer. Inst. Chem. Eng. J. 45, 383-389.
- Kasting, J.F., Zahnle, K.J., Pinto, J., Young, A. (1989). Sulfur, ultraviolet radiation, and the early
 evolution of life. Origins of Life 19, 95-108.
- Klippenstein, S.J., Miller, J.A., (2005). The addition of hydrogen atoms to diacetylene and the
 heats of formation of i-C4H3 and n-C4H3. J. Phys. Chem. A 109, 4285-4295.
- ⁶⁴⁶ Kopparapu, R.K., Kasting, J.F., Zahnle, K.J., (2012)Kopparapu, R. k. and Kasting, J. F. and
 ⁶⁴⁷ Zahnle, K. J., A Photochemical Model for the Carbon-rich Planet WASP-12b. <u>Astrophys.</u>
 ⁶⁴⁸ J. 745, 77.
- Koskinen, T.T., Harris, M.J., Yelle, R.V., Lavvas, P., (2013). The escape of heavy atoms from the
 ionosphere of HD 209458b. I. A photochemical-dynamical model of the thermosphere. <u>Icarus</u>
 226, 1678-1694.

- ⁶⁵² Kurbanov, M.A., Mamedov, Kh.F. (1995). The role of the reaction $CO + SH \rightarrow COS + H$ in ⁶⁵³ hydrogen formation in the course of interaction between CO and H2S. <u>Kinet. Catal. 36</u>, ⁶⁵⁴ 455-457.
- Langford, R.B., Oldershaw, G.A., (1973). Mechanism of Sulfur Formation in the Flash Photolysis
 of Carbonyl Sulphide. J. Chem. Soc. Faraday Trans. 69, 1389.
- Langford, R.B., Oldershaw, G.A., (1972). Flash Photolysis of H₂S. J. Chem. Soc. Faraday Trans. 658 68, 1550.
- Lavvas, P., Koskinen, T., Yelle, R.V., (2014). Electron Densities and Alkali Atoms in Exoplanet Atmospheres. Astrophys. J. 796, (15 pp).
- Lee, J.H., Stief, L.J., Timmons, R.B. (1977). Absolute Rate Parameters for the Reaction of Atomic
 Hydrogen with Carbonyl Sulfide and Ethylene Episulfide, J. Chem. Phys. 67, 1705-1714.
- Liang, M.C., Parkinson, C.D., Lee, A.Y.-T., Yung, Y.L., Seager, S. (2003). Source of atomic hydrogen in the atmosphere of HD 209458b. Astrophys. J. 596, L247-L250.
- Liang, M.C., Seager, S., Parkinson, C.D., Lee, A.Y.-T., Yung, Y.L. (2003). On the Insignificance
 of Photochemical Hydrocarbon Aerosols in the Atmospheres of Close-in Extrasolar Giant
 Planets Astrophys. J. 605, L61-L64.
- Line, M.R., Liang, M.-C., Yung, Y.L. (2010). High-Temperature Photochemistry in the Atmosphere of HD 189733b. Astrophys. J. 717, 496-502.
- Line, M.R., Vasisht, G., Chen, P., Angerhausen, D., Yung, Y.L. (2011). Thermochemical and
 Photochemical Kinetics in Cooler Hydrogen-dominated Extrasolar Planets: A Methanepoor GJ436b? Astrophys. J. 738, article id. 32, 14 pp.
- Lodders, K., Fegley, B., (2006). In <u>Chemistry of Low Mass Substellar Objects.</u> Ed. J.W. Mason (Berlin: Springer), pp. 1-31.
- ⁶⁷⁵ Lu, C.W., Wu, Y.J., Lee, Y.P., Zhu, R.S., Lin, M.C. (2004). Experimental and theoretical investi-⁶⁷⁶ gations of rate coefficients of the reaction $S(^{3}P) + O2$ in the temperature range 298-878 K. ⁶⁷⁷ J. Chem. Phys. 121, 8271-8278
- ⁶⁷⁸ Lu, C.W., Wu, Y.J., Lee, Y.P., Zhu, R.S., Lin, M.C. (2006). Experimental and theoretical investiga-⁶⁷⁹ tion of rate coefficients of the reaction $S(^{3}P) + OCS$ in the temperature range of 298-985 K. ⁶⁸⁰ J. Chem. Phys. 125, 164329.
- Lyons, J.R. (2008). An estimate of the equilibrium speciation of sulfur vapor over solid sulfur and implications for planetary atmospheres. J. Sulfur Chem. 29, 269-279.
- Macintosh, B., Graham, J.R., Barman, T., De Rosa, R.J., Konopacky, Q., Marley, M.S., Marois,
 C., Nielsen, E.L., Pueyo, L., Rajan, A., and 72 coauthors (2015). Discovery and spectroscopy
 of the young jovian planet 51 Eri b with the Gemini Planet Imager. Science 350, 64-67.

- Marley, M.S., Fortney, J.J., Seager, S., Barman, T. (2007). An Imposing and Comprehensive Title.
 In Protostars and Planets V. Eds. B. Reipurth, D. Jewitt, and K. Keil. Univ. Ariz. Press,
 Tucson, pp. 733–747.
- Meyer, B., (1976). Elemental sulfur. Chemical Reviews 76, 367-387.
- Miguel, Y., Kaltenegger, L., (2014). Exploring Atmospheres of Hot Mini-Neptunes and Extrasolar
 Giant Planets Orbiting Different Stars with Application to HD 97658b, WASP-12b, CoRoT 2b, XO-1b, and HD 189733b, Astrophys. J. 780, 166.
- Miguel, Y., Kaltenegger, L., Linsky, J.L., Rugheimer, S., (2015) The effect of Lyman α radiation on
 mini-Neptune atmospheres around M stars: application to GJ 436b. Mon. Not. Roy. Astron.
 Soc. 446, 345-353.
- Millar, T.J., Farquhar, P.R.A., Willacy, K., (1997). The UMIST database for astrochemistry 1995.
 Astron. Astrophys. Suppl. Ser. 121, 139-185.
- Miller-Ricci Kempton, E., Zahnle, K., Fortney, J.J. (2011). The Atmospheric Chemistry of GJ 1214b: Photochemistry and Clouds. Astrophys. J. 745, 3 (13pp).
- Morley, C.V., Fortney, J.J., Visscher, C., Marley, M.S., Sauman, D., Legett, S.K. (2012). Neglected
 clouds in T and Y dwarf atmospheres. Astrophys. J. 756, article id. 172, 17 pp.
- Morley, C.V., Fortney, J.J., Kempton, E.M.R., Marley, M.S., Visscher, C., Zahnle, K.J. (2013).
 Quantitatively Assessing the Role of Clouds in the Transmission Spectrum of GJ 1214b.
 Astrophys. J. 775, article id. 33, 13 pp.
- Morley, C.V., Fortney, J.J., Kempton, E.M.R., Marley, M.S., Visscher, C., Zahnle, K.J. (2013).
 Quantitatively Assessing the Role of Clouds in the Transmission Spectrum of GJ 1214b.
 Astrophys. J. 775, article id. 33, 13 pp.
- Moses, J.I., Nash, D.B., (1991). Phase transformations and the spectral reflectance of solid sulfur:
 can metastable sulfur allotropes exist on Io? Icarus 89, 277-304.
- Moses, J.I. Allen, M., Gladstone, G.R. (1995). Post-SL9 sulfur photochemistry on Jupiter. <u>Geophys.</u>
 Res. Lett. 22, 1597-1600.
- Moses, J.I., Zolotov, M.Y., Fegley, B. (2002). Photochemistry of a volcanically driven atmosphere
 on Io: Sulfur and oxygen species from a Pele-type eruption. Icarus 156, 76-106.
- Moses, J.I., Fouchet, F., B'ezard, B., Gladstone, G.R., Lellouch, E., Feuchtgruber, H. (2005).
 Photochemistry and diffusion in Jupiter's stratosphere: Constraints from ISO observations and comparisons with other giant planets. J. Geophys. Res. 110, E08001 (45 pp.).
- Moses, J.I., Visscher, C., Fortney, J.J., Showman, A.P., Lewis, N.K., Griffith, C.A., Klippenstein,
 S.J., Shabram, M., Friedson, A.J., Marley, M.S., Freedman, R.S. (2011). Disequilibrium

- Carbon, Oxygen, and Nitrogen Chemistry in the Atmospheres of HD 189733b and HD
 209458b. Astrophys. J. 737, article id. 15, 37 pp.
- Moses, J.I., Madhusudhan, N., Visscher, C., Freedman, R.S. (2013). Chemical Consequences of
 the C/O Ratio on Hot Jupiters: Examples from WASP-12b, CoRoT-2b, XO-1b, and HD
 189733b. Astrophys. J. 763, article id. 25, 26 pp.
- Moses, J.I., Line, M.R., Visscher, C., Richardson, M.R., Nettelmann, N., Fortney, J.J., Barman,
 T.S., Stevenson, K.B., Madhusudhan, N. (2013). Compositional Diversity in the Atmospheres of Hot Neptunes, with Application to GJ 436b. <u>Astrophys. J. 777</u>, article id. 34, 23
 pp.
- Moses, J.I. (2014). Chemical Kinetics on Extrasolar Planets. <u>Phil. Trans. Roy. Soc. A 372</u>, 20130073.
 DOI: 10.1098/rsta.2013.0073.
- Nicholas, J.E., Amodio, C.A., Baker, M.J. (1979). Kinetics and Mechanism of the Decomposition
 of H₂S, CH₂SH and (CH₃)₂S in a Radio-frequency Pulse Discharge <u>J. Chem. Soc. Faraday</u>
 Trans. 75, 1868.
- Olschewski, H.A., J. Troe, Wagner, H.Gg., (1994). UV absorption study of the thermal decomposition reaction $H_2S \rightarrow H_2 + S(^{3}P)$. J. Phys. Chem. 98, 12964-12967.
- Oya, M., Shiina, H., Tsuchiya, K., Matsui, H. (1994). Thermal decomposition of COS. <u>Bull. Chem.</u>
 Soc. Japan 67, 2311-2313.
- ⁷³⁷ Pauling, L. (1970). General Chemistry. Dover.
- ⁷³⁸ Pavlov, A.A., Kasting, J.F., (2002). Astrobiology
- Pen, J., Hu, X., Marshall, P. (1999). Experimental and *ab initio* investigations of the kinetics of
 the reaction of H atoms with H₂S. J. Phys. Chem. A, 103, 5307-5311.
- Perrin, D., Richard, C., Martin, R. (1988). Etude cinetique de la reaction thermique du pentene-2
 cis vers 500°C. III Influence de H₂S. J. Chim. Phys. 85, 185-192.
- Pinto, J., Gladstone, R., Yung, Y., (1980). Photochemical production of formaldehyde in the Earth's
 primitive atmosphere. Science 210, 183-185.
- Rimmer, P.B., Helling, C., (2016). A Chemical Kinetics Network for Lightning and Life in Planetary
 Atmospheres. Astrophys. J. Supp., *accepted.* arXiv:1510.07052.
- Roth, P., Lohr, R., Braner, U., (1982). Thermal Decomposition of Hydrogen Sulfide at Low Con centrations. Combust. Flame 45, 273.
- Sander, S.P., Friedl, R.R., Ravishankara, A.R., Golden, D.M., Kolb, C.E., Kurylo, M.J., Huie,
 R.E., Orkin, V.L., Molina, M.J., Moortgat, G.K., Finlayson-Pitts, B.J. (2003). Chemical

- Kinetics and Photochemical Data for Use in Atmospheric Studies. Evaluation Number 14.
 JPL Publication 02-25.
- Schofield, K. (1973). Evaluated chemical kinetic rate constants for various gas phase reactions. <u>J.</u>
 Phys. Chem. Ref. Data 2, 25-84.
- Shiina, H., Oya, M., Yamashita, K., Miyoshi, A., Matsui, H. (1996). Kinetic studies on the pyrolysis
 of H₂S. Phys. Chem. 100, 2136-2140.
- ⁷⁵⁷ Shiina, H., Miyoshi, A., Matsui, H. (1998). Investigation on the insertion channel in the $S(^{3}P) + H_{2}$ reaction. J. Phys. Chem. A 102, 3556 3559.
- Shum, L.G.S., Benson, S.W., (1985). The pyrolysis of dimethyl sulfide, kinetics and mechanism.
 Int. J. Chem. Kinet. 17, 749.
- Singleton, D.L., Cvetanovic, R.J. (1988). Evaluated chemical kinetic data for the reactions of atomic
 oxygen O(³P) with sulfur containing compounds. J. Phys. Chem. Ref. Data 17, 1377-1399.
- Spencer, J.R., Jessup, K.L., McGrath, M.A., Gilda E. Ballester, G.E., Roger Yelle, R.V. (2000).
 Discovery of gaseous S₂ in Io's Pele plume. Science 288, 1208-1210.
- Tesner, P.A., Nemirovskii, M.S., Motyl, D.N. (1990). Kinetics of the thermal decomposition of
 hydrogen sulfide at 600-1200°C. Kinet. Catal. 31, 1081-1083.
- Tsuchiya, K., Yamashita, K., Miyoshi, A., Matsui, H. (1996). Studies on the reactions of atomic sulfur (³P) with H₂, D₂, CH₄, C₂H₆, C₃H₈, n-C₄H₁₀, and i-C₄H₁₀. J. Phys. Chem. 100, 17202-17206.
- Tsuchiya, K., Kamiya, K., Matsui, H., (1997). Studies on the Oxidation Mechanism of H₂S Based
 on Direct Examination of the Key Reaction. Int. J. Chem. Kinet. 29, 57.
- Venot, O., Hébrard, E., Agúndez, M., Dobrijevic, M., Selsis, F., Hersant, F., Iro, N., Bounaceur,
 R., (2012). A chemical model for the atmosphere of hot Jupiters. <u>Astron. Astrophys. 546</u>,
 A43.
- Venot, O., Fray, N., Bénilan, Y., Gazeau, M.-C., Hébrard, E., Larcher, G., Schwell, M., Dobrijevic,
 M., Selsis, F., (2013). High-temperature measurements of VUV-absorption cross sections of
 CO₂ and their application to exoplanets. Astron. Astrophys. 551, A131.
- ⁷⁷⁸ Venot, O., Agúndez, M., Selsis, F., Tessenyi, M., Iro, N., (2014). The atmospheric chemistry of the ⁷⁷⁹ warm Neptune GJ 3470b: Influence of metallicity and temperature on the CH_4/CO ratio. ⁷⁸⁰ Astron. Astrophys. 562, A51.
- Venot, O., Hébrard, E., Agúndez, M., Decin, L., Bounaceur, R., (2015). New chemical scheme for
 studying carbon-rich exoplanet atmospheres. Astron. Astrophys. 577, A33.

- Visscher, C.W., Lodders, K., Fegley, B. (2006). Atmospheric chemistry in giant planets, brown
 dwarfs, and low-mass dwarf stars. ii. Sulfur and phosphorus. Astrophys. J. 648, 1181-1195.
- Visscher, C.W., Moses, J.I. (2011). Quenching of carbon monoxide and methane in the atmospheres
 of cool brown dwarfs and hot Jupiters. Astrophys. J. 738, article id. 72, 12 pp.
- Visscher, C.W. (2012). Chemical timescales in the atmospheres of highly eccentric exoplanets.
 Astrophys. J. 757, article id. 5, 8 pp.
- Vuitton, V., Yelle, R.V., Lavvas, P., Klippenstein, S.J. (2012). Rapid association reactions at low
 pressure: impact on the formation of hydrocarbons on Titan. <u>Astrophys. J. 744</u>, article id.
 11, 7 pp.
- Woiki, D.; Roth, P. (1994). Kinetics of the high-temperature H_2S decomposition. J. Phys. Chem. 98, 12958-12963.
- Woiki, D., Roth, P., (1995a). A shock tube study of the reaction $S + H_2 = SH + H$ in pyrolysis and photolysis systems. Int. J. Chem. Kinet. 27, 547-553.
- Woiki, D., Roth, P. (1995b). Oxidation of S and SO by O₂ in high-temperature pyrolysis and
 photolysis reaction systems. Int. J. Chem. Kinet. 27, 5-71.
- ⁷⁹⁸ Young A.T. (1983). Venus cloud microphysics. Icarus 56, 568.
- Yung, Y.L., Liang, M.C., Jiang, X., Shia, R.L., Lee, C., Bézard B., Marcq, E., (2009). Evidence
 for carbonyl sulfide (OCS) conversion to CO in the lower atmosphere of Venus. J. Geophys.
 Res. 114, E00B34.
- Zahnle, K.J, Mac Low, M.-M., Lodders, K., B. Fegley, B. (1995). Sulfur chemistry in the wake of
 Comet Shoemaker-Levy 9. Geophys. Res. Lett. 22, 1593-1596.
- Zahnle, K.J., Marley, M.S., Freedman, R.S., Lodders, K., Fortney, J.J. (2009). Atmospheric sulfur
 chemistry on hot Jupiters. Astrophys. J. Lett. 701, L20-L24.
- Zahnle, K.J., Marley, M.S. (2014). Methane, Carbon Monoxide, and Ammonia in Brown Dwarfs
 and Self-Luminous Giant Planets. Astrophys. J. 797, article id. 41, 19 pp.

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	Table 1				
	Reactants		Products	Rate $[cm^3s^{-1}]$ or $[cm^6s^{-1}]$	Reference
R1	$S + S + M^*$	\rightarrow	$S_2 + M$	$2.0 \times 10^{-33} e^{206/T}$	Du et al. (2008)
	S + S	\rightarrow	S_2	$2.3 \times 10^{-14} e^{415/T}$	Du et al. (2008)
R2	$S + S_2 + M$	\rightarrow	$S_3 + M$	$1.0 \times 10^{-30} (T/298)^{-2.0}$	assumed
	$S + S_2$	\rightarrow	S_3	5.0×10^{-11}	assumed
R3	$S + S_3$	\rightarrow	$S_2 + S_2$	4.0×10^{-11}	assumed
R4	$S + S_3 + M$	\rightarrow	$S_4 + M$	$1.0 \times 10^{-30} (T/298)^{-2.00}$	assumed, varied $10 \times$
	$S + S_3$	\rightarrow	S_4	5.0×10^{-11}	assumed
R5	$S_2 + S_2 + M$	\rightarrow	$S_4 + M$	$1.0 \times 10^{-30} (T/298)^{-2.00}$	varied $10 \times$, see note
	$S_2 + S_2$	\rightarrow	S_4	3.0×10^{-11}	assumed
R6	$S + S_4$	\rightarrow	$S_2 + S_3$	$4.0 \times 10^{-11} e^{-500/T}$	Moses et al. (1995)
$\mathbf{R7}$	$S_4 + S_4 + M$	\rightarrow	$S_8 + M$	$7.0 \times 10^{-30} (T/298)^{-2.00}$	assumed, varied $100 \times$
	$S_4 + S_4$	\rightarrow	S_8	7.0×10^{-11}	assumed
$\mathbf{R8}$	S + HS	\rightarrow	$S_2 + H$	1.0×10^{-11}	assumed, see note
$\mathbf{R9}$	H + HS	\rightarrow	$S + H_2$	$3.0 \times 10^{-11} (T/298)^{0.7}$	reverse of R9r
R9r	$S + H_2$	\rightarrow	H + HS	$5.3 \times 10^{-10} (T/298)^{0.95} e^{-9920/T}$	see note
R10	HS + HS	\rightarrow	$S_2 + H_2$	$1.3 \times 10^{-11} e^{-20600/T}$	like $2OH \rightarrow H_2 + O_2$
R11	$H + S_3$	\rightarrow	$HS + S_2$	$5.0 \times 10^{-11} e^{-500/T}$	like $\rm H\!+\!O_3\rightarrow OH\!+\!O_2$
R12	$H + S_4$	\rightarrow	$HS + S_3$	$5.0 \times 10^{-11} e^{-500/T}$	like $H + S_3$
R13	O + HS	\rightarrow	OH + S	$1.7 \times 10^{-11} \left(T/298\right)^{0.67} e^{-956/T}$	Schofield (1973)
R14	HS + OH	\rightarrow	$H_2O + S$	$4.0 \times 10^{-12} e^{-240/T}$	inspired by R23
R15	S + CH	\rightarrow	HS + C	$1.7 \times 10^{-11} \left(T/298\right)^{0.50} e^{-4000/T}$	Millar et al. (1997)
R16	S + NH	\rightarrow	HS + N	$1.7 \times 10^{-11} (T/298)^{0.50} e^{-4000/T}$	Millar et al. (1997)
R17	$NH_2 + HS$	\rightarrow	$NH_3 + S$	$5.0 \times 10^{-12} e^{-500/T}$	Moses et al. (1995)
R18	$HS + CH_2$	\rightarrow	$S + CH_3$	$4.0 \times 10^{-12} e^{-500/T}$	Moses et al. (1995)
R19	$HS + CH_3$	\rightarrow	$S+ CH_4$	$4.0 \times 10^{-11} e^{-500/T}$	Shum & Benson (1985)
R20	S + HCO	\rightarrow	HS + CO	6.0×10^{-11}	Moses et al. (1995)
R21	H + HS + M	\rightarrow	$H_2S + M$	$1.4 \times 10^{-31} (T/298)^{-2.5} e^{+500/T}$	see note
	H + HS	\rightarrow	H_2S	1.0×10^{-10}	assumed
R22	$S + H_2 + M$	\rightarrow	$H_2S + M$	$1.4 \times 10^{-31} (T/298)^{-1.9} e^{-8140/T}$	see text
	$S + H_2 + M$	\rightarrow	$H_2S + M$	$1.4 \times 10^{-31} (T/298)^{-1.9} e^{-2300/T}$	alternate rate, see text
	$S + H_2$	\rightarrow	H_2S	1.0×10^{-11}	assumed
R23	$H + H_2S$	\rightarrow	$\mathrm{HS} + \mathrm{H}_2$	$3.7 \times 10^{-12} (T/298)^{1.94} e^{-455/T}$	Pen et al. (1999)
R24	$H_2S + S$	\rightarrow	HS + HS	$1.4 \times 10^{-10} e^{-3720/T}$	Shiina et al. (1996)
R25	$O + H_2S$	\rightarrow	HS+ OH	$9.2 \times 10^{-12} e^{-1800/T}$	DeMore et al. (1997)
R26	$OH + H_2S$	\rightarrow	$H_2O + HS$	$6.1 \times 10^{-12} e^{-81/T}$	Atkinson et al. (2004)
R27	HS + HCO	\rightarrow	$H_2S + CO$	5.0×10^{-11}	like R20
R28	$CH_2 + H_2S$	\rightarrow	$CH_3 + HS$	$2.5 \times 10^{-11} e^{-750/T}$	Darwin&Moore (1995)
R29	$H_2S + CH_3$	\rightarrow	$\mathrm{HS} + \mathrm{CH}_4$	$2.1 \times 10^{-13} e^{-1160/T}$	Perrin et al. (1988)
R30	O + HS	\rightarrow	$SO + \overline{H}$	7.0×10^{-11}	Sander et al. (2003)
R31	S + OH	\rightarrow	H + SO	6.6×10^{-11}	DeMore et al. (1997)
R32	$O + S_2$	\rightarrow	SO + S	1.1×10^{-11}	Hills et al. (1987)
R33	$S+O_2$	\rightarrow	SO + O	$1.5 \times 10^{-13} (T/298)^{2.11} e^{-730/T}$	Lu et al. (2004)
R34	$S_3 + O$	\rightarrow	$S_2 + SO$	$2.0 \times 10^{-11} e^{-500/T}$	Moses et al. (1995)

	Reactants		Products	Rate $[cm^3s^{-1}]$ or $[cm^6s^{-1}]$	Reference
R35	$S_4 + O$	\rightarrow	$S_3 + SO$	$2.0 \times 10^{-11} e^{-500/T}$	Moses et al. (1995)
R36	S + CO + M	\rightarrow	OCS + M	$3.6 \times 10^{-34} (T/298)^{-0.57}$	see note
	S + CO	\rightarrow	OCS	3.0×10^{-14}	see note
R37	O + OCS	\rightarrow	$\rm CO + SO$	$7.8 \times 10^{-11} e^{-2620/T}$	Singleton and Cve-
					tanovic (1988)
R38	HS + CO	\rightarrow	OCS + H	$4.2 \times 10^{-14} e^{-7660/T}$	Kurbanov et al. (1995)
R39	OCS + S	\rightarrow	$\rm CO + S_2$	$1.5 \times 10^{-13} \left(T/298\right)^{2.57} e^{-1180/T}$	Lu et al. (2006)
R40	O + OCS	\rightarrow	$S + CO_2$	$8.3 \times 10^{-11} e^{-5530/T}$	Singleton and Cve-
					tanovic (1988)
R41	OCS + OH	\rightarrow	$CO_2 + HS$	$1.1 \times 10^{-13} e^{-1200/T}$	Atkinson et al. (2004)
R42	S + HCO	\rightarrow	OCS + H	6.0×10^{-11}	Moses et al. (1995)
R43	$\rm CO + S_3$	\rightarrow	$S_2 + OCS$	$1.0 \times 10^{-11} e^{-10000/T}$	see note
R44	O + CS	\rightarrow	$\rm CO + S$	$2.7 \times 10^{-10} e^{-760/T}$	Atkinson et al. (2004)
R45	S + CH	\rightarrow	CS + H	2.0×10^{-11}	assumed
R46	OH + CS	\rightarrow	OCS + H	2.0×10^{-13}	assumed
R47	$CH_2 + S$	\rightarrow	$CS + H_2$	2.0×10^{-11}	assumed
R48	$H + S_4 + M$	\rightarrow	$HS_4 + M$	7.0×10^{-30}	assumed
	$H + S_4$	\rightarrow	HS_4	7.0×10^{-11}	assumed
R49	$H + S_8$	\rightarrow	$HS_4 + S_4$	$3.0 \times 10^{-12} e^{-5000/T}$	assumed, see text
	$H + S_8$	\rightarrow	$HS_4 + S_4$	$3.0 \times 10^{-11} e^{-2500/T}$	alternate rate
R50	$H + HS_4$	\rightarrow	$H_2 + S_4$	$1.0 \times 10^{-10} e^{-500/T}$	assumed, see text
R51	$H + HS_4$	\rightarrow	$H_2S + S_3$	$1.0 \times 10^{-10} e^{-2500/T}$	assumed, see text
R52	$OH + HS_4$	\rightarrow	$H_2O + S_4$	$3.0 \times 10^{-11} e^{-500/T}$	assumed, see text
R53	$NH_2 + HS_4$	\rightarrow	$NH_3 + S_4$	$3.0 \times 10^{-11} e^{-500/T}$	assumed, see text
R54	$HS + HS_4$	\rightarrow	$H_2S + S_4$	$3.0 \times 10^{-11} e^{-2500/T}$	assumed, see text
R55	$HS_4 + HS_4$	\rightarrow	$H_2 + S_8$	$3.0 \times 10^{-12} e^{-2500/T}$	assumed, see text
R56	$S + HS_4$	\rightarrow	$HS + S_4$	$3.0 \times 10^{-11} e^{-1000/T}$	assumed, see text
R57	$\mathrm{C_2H} + \mathrm{C_2H_2}$	\rightarrow	$C_4H_2 + H$	$1.25 \times 10^{-10} \left(T/298\right)^{0.24} e^{37.3/T}$	Eiteneer & Frenklach (2003)
R58	$C_4H_2 + H + M$	\rightarrow	$C_4H_3 + M$	$5.9 \times 10^{-25} \left(T/298\right)^{-8.9} e^{-1260/T}$	Klippenstein & Miller (2005)
	$C_4H_2 + H + M$	\rightarrow	$C_4H_3 + M$	$5.2 \times 10^{-11} \left(T/298\right)^{1.2} e^{-882/T}$	Klippenstein & Miller (2005)
R59	$C_4H_3 + H$	\rightarrow	$C_4H_2 + H_2$	$5 \times 10^{-11} e^{-500/T}$	see text
R60	$C_4H_3 + H$	\rightarrow	$\mathrm{C_2H_2} + \mathrm{C_2H_2}$	$5 \times 10^{-11} e^{-2000/T}$	see text
R61	$C_4H_3 + H$	\rightarrow	C_4H_4	1.8×10^{-10}	Harding et al. (2007)

* M refers to the background atmosphere, principally H_2 and H_2 ; units of density $[cm^{-3}]$.

R2-R8. These assumed rates are generally consistent with those of Moses et al. (2002); Yung et al. (2009).

R5. Reported rates are 1×10^{-29} (Langford and Oldershaw 1973), 2.2×10^{-29} (Nicholas et al. 1979).

R8. Reported rates are 4×10^{-11} (Schofield 1973), $< 5 \times 10^{-12}$ (Nicholas et al. 1979).

R9r. A blend of Woiki and Roth (1995a) and Shiina et al. (1998).

R21. This is the reverse of Shiina et al. (1998) upper bound on $H_2S + M \rightarrow H + HS + M$.

R36. These are reverses of Oya et al. (1994) rate for OCS + M \rightarrow CO + S + M and of

Schofield (1973) high pressure limit $OCS \rightarrow CO + S$.

R43. Exothermic, but the analogous $CO + O_3 \rightarrow CO_2 + O_2$ has an upper limit at 298 K of 4×10^{-25} . R45-R47. There is little information re reactions of CS.

R48-R61. HS_4 and C_4H_3 are invoked to create sinks on S_8 and C_4H_2 , respectively.