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1 Predicting groundwater arsenic contamination in Southeast Asia

from surface parameters 2

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18 Arsenic contamination of groundwater resources threatens the health of 19 millions of people worldwide, particularly in the densely populated river deltas 20 of Southeast Asia. Although many arsenic-affected areas have been identified in 21 recent years, a systematic evaluation of vulnerable areas remains to be carried 22 out. Here we present maps pinpointing areas at risk of groundwater arsenic concentrations exceeding 10 µgL⁻¹. These maps were produced by combining 23 geological and surface soil parameters in a logistic regression model, calibrated 24 25 with 1756 aggregated and geo-referenced groundwater data points from the 26 Bengal, Red River and Mekong deltas. We show that Holocene deltaic and 27 organic-rich surface sediments are key indicators for arsenic risk areas and that the combination of surface parameters is a successful approach to predict 28 29 groundwater arsenic contamination. Predictions are in good agreement with the known spatial distribution of arsenic contamination and further indicate 30 31 elevated risks in Sumatra and Myanmar where no groundwater studies exist.

33 More than 100 million people worldwide ingest excessive amounts of arsenic (As) 34 through drinking water contaminated from natural geogenic sources. Many Asian countries, in particular, are known to be affected by high groundwater As 35 concentrations as a result of chemically reducing aguifer conditions: Bangladesh¹⁻¹⁰, 36 India^{3,11,12}, China^{13,14}, Nepal¹⁵, Cambodia¹⁶⁻¹⁸ and Vietnam^{17,19,20}. However, since As 37 analysis is expensive and time-consuming, groundwater resources of many regions 38 39 still remain to be tested. Therefore, maps pinpointing areas vulnerable to As contamination can guide households at risk of arsenic contamination, as well as 40 scientists and policy-makers to initiate early mitigation measures and protect the 41 42 populations from chronic As poisoning.

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44 Though the exact chemical conditions and reactions leading to As mobilization are still 45 under debate, it is generally accepted that microbial and/or chemical reductive dissolution of As-bearing iron minerals in the aguifer sediments^{1,4,21} is the main cause 46 47 for the release of As. Reducing conditions are often associated with the presence of natural (bio)degradable organic carbon embedded in sediments^{9,11,22-25}. Other 48 identified key characteristics of contaminated areas are rapidly buried, young 49 (Holocene) sediments and low hydraulic gradients in flat and low-lying areas^{8,18,26,27}. 50 51 Ideally an As prediction model for groundwater should be based on parameters that 52 indicate the key characteristics mentioned above in three dimensions. However, in the absence of a 3-dimensional spatially continuous database of aquifer conditions to 53 54 depth, globally and regionally available (two dimensional) surface parameters can be used as indicators for As-enrichment in underlying aquifers^{28,29}. 55

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In the past, several geostatistical interpolation methods (e.g. kriging) have been used
 to predict elevated As in groundwater on a regional scale³⁰⁻³². However, for predictions

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59 of areas where no groundwater quality data exist, interpolation methods are not applicable³² and models based on logistic regression are more appropriate^{33,34}. In an 60 expert-based statistical model to delineate areas at risk of groundwater As 61 62 contamination on a coarse global scale, we found that geological information was of crucial importance²⁹. Here we focus on an in-depth assessment of depositional 63 64 environments in Southeast Asia. We use a logistic regression approach based on 65 relationships between sedimentary information, soil maps and measured groundwater As data of Bangladesh, Cambodia and Vietnam, to assess the relative importance of 66 the different surface proxies in these countries. We apply these relationships to set up 67 68 prediction maps of As contamination in Southeast Asia including Indonesia (Sumatra) 69 and other countries where groundwater quality data is scarce (Myanmar and Thailand). 70 Furthermore, we verify the predicted risk in South Sumatra, where the groundwater 71 has not previously been tested for the presence of As.

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73 ARSENIC PREDICTION MODEL

74 Our model is based on three assumptions. First, sedimentary depositional environments are characterized by a unique combination of chemical, physical, and 75 biological properties³⁵ and can serve as indicators (proxies) for chemical and physical 76 77 conditions of the aguifers beneath the surface. Second, soil properties are proxies for 78 present and past drainage conditions and they are also indicators of recent 79 depositional environments. Third, soil textures, for example clay and silt are proxies 80 for the chemical maturity of the sediments, where clay is more mature than silt. An important factor in the development of soil textures is topography³⁶, which allows the 81 82 delineation of areas where the model is applicable.

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84 GIS-datasets were established from digital elevation data, countrywide geological 85 maps and global soil data (FAO), which were converted to a raster format using ArcGIS (ver. 9.2). An overview of GIS data used in this study is provided in Table S1 86 87 (see Supplementary Information). Because each geological map applied a different 88 classification terminology, we created an uniformly classified geological map for all 89 regions (Figure 1). Although Bangladesh does geographically not belong to Southeast Asia, it was included in the model because of the large number of available data 90 91 points⁵. Statistical relations between As concentrations and 30 parameters related to 92 soil properties, geology, climate, and hydrology (Table S2) were initially evaluated by 93 stepwise regression. The six parameters exhibiting a significance >95% and two 94 additional soil parameters were employed in the final model (see Methods and Table 95 1). Since young geological deposits and As groundwater contamination are rarely 96 observed in areas with steeper slopes, and groundwater As concentration data are only available for regions with a flat topography (slope $<0.1^{\circ} \approx 0.17\%$), areas with 97 98 slopes >0.1° were excluded from the model (see Supplementary Figure S1).

99

100 The prediction required three steps: i) Aggregation and binary-coding of measured As 101 concentrations to reduce spatial heterogeneities (dependent variables), ii) Logistic 102 regression to obtain weighting coefficients of independent variables (see Methods 103 section and corresponding results in Table 1), and, iii) Calculation of the probability of As contamination based on the threshold value of 10 µgL⁻¹. The spatial datasets 104 105 considered as independent variables for the model are topographic data to delineate 106 the model area, sedimentary depositional environments as a proxy for aquifer 107 conditions, and soil variables as a proxy for drainage and chemical maturity of 108 sediments (see the Methods section).

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110 SURFACE PARAMETERS CONTRIBUTING TO THE MODEL

111 The weighting factors (λ) and significance of the eight independent variables retained 112 in the final model are listed in Table 1. In general, variables describing sedimentary 113 depositional environments have a larger contribution to the model than soil variables, 114 presumably because geological variables come closest to describing As 115 contamination in the aguifer itself. The presence of young deltaic deposits (λ =1.65) is 116 a particularly significant indicator for As-contaminated aguifers. In Southeast Asia. 117 delta initiation and progradation occurred simultaneously with the Holocene Climate Optimum³⁷. Therefore, delta progradation resulted in the burial of organic matter at a 118 119 high rate. The presence of relatively fresh organic carbon provides favourable 120 conditions to establish reducing environments, which may lead to As enrichment in 121 groundwater. Logistic regression confirms that organic-rich deposits (λ =1.11) play an 122 important part in the model. Recent alluvial deposits (λ =0.55) are also indicative for 123 elevated As concentrations in groundwaters (Table 1). Of the soil parameters, medium 124 textured soils seem to be indicative of As-bearing aquifers evolved from rapid 125 accumulation of young (Holocene) sediments. In contrast, the negative weighting 126 coefficient (λ = -0.95) for floodplain deposits implies that these fine-grained deposits 127 could overlie aguifers low in dissolved As (compare Figures 1 and 3). Fine-grained 128 deposits (high clay content) thereby point to low-energy depositional environments 129 with condensed sediments where Holocene aquifers are rare and where groundwater 130 is likely drawn from older aquifers.

131

Pre-Holocene deposits, other Holocene deposits and tidal deposits (Figure 1) were found to be statistically insignificant (p >0.05) and they were excluded from the model in the first stepwise regression (see Table S2 and Table 1). Tidal sediments are

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generally associated with aquifers abundant in sulfate that may be microbially reduced
to sulfide and re-precipitate As^{38,39}. They might serve as proxies for low-risk areas, but
our As data in such aquifers were possibly too few to see such a relation.

138

139 INFERRING DEPOSITIONAL ENVIRONMENTS AT DEPTH FROM THE SURFACE

As mentioned in the introduction, our model is based on two- dimensional data (i.e. 140 141 surface maps). Nevertheless, geological information (sedimentary depositional 142 environments) inherently contains a three-dimensional component. The recent sedimentary history of major Southeast Asian basins is characterized by delta 143 144 initiation, which occurred on a global scale at about 8500-6500 years BP and was principally controlled by the deceleration of sea-level rise⁴⁰. Delta progradation 145 146 resulted in the unconformable deposition of thick late Pleistocene-Holocene sediments on Pleistocene and older sediments, e.g. as incised-valley-fill deposits^{8,18}. As-147 148 contaminated aguifers are mainly present in these Holocene aguifers, whereas deeper Pliocene-Pleistocene aquifers are, to a large extent, free of As^{7,8,26}. The boundary 149 150 between Pleistocene and Holocene sediments is not located at a constant depth⁸ and 151 this can lead to misclassifications, since our model inherently assumes that the 152 underlying aguifer belongs to the same sedimentary depositional environment as on 153 the surface.

154

Two situations exist where the environment at the surface does not reflect the geology at depth, i) when the As measurements used in the model were obtained from tubewells tapping deep (Pleistocene) aquifers because shallow (Holocene) aquifers are not present, or shallow groundwater is too saline for consumption (e.g. in coastal regions), and ii) when Holocene sediments were deposited at sedimentation rates to small to form a usable aquifer. In both situations Holocene depositional environments

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are present at the surface and high-risk areas are indicated, although measured As
concentrations are low (false positive cases). Even though the probability maps would
be improved by including the thickness of Holocene sediments, the absence of
country-wide three-dimensional geological data rules this out. Furthermore, the
complexity of aquifer heterogeneity at a local scale makes it inevitable that
misclassifications occur.

167

168 **PROBABILITY MAPS**

169 Predicted areas at risk of As contamination agree well with known spatial 170 contamination patterns, a finding which is supported by the results of the model 171 classification indicating the performance of model prediction (Figure 2) and the 172 Hosmer-Lemeshow goodness-of-fit statistics (Table S3). An absolute average 173 deviation of 7.3% is found between expected and modelled probabilities of As being ≤ 10 or $>10 \mu g L^{-1}$. The model is further characterised by the receiver operating 174 175 characteristics curve area of ~0.7 (Figure S2), which is a good result considering that 176 neither depths of analyzed groundwater wells, nor aquifer hydrological data form part 177 of the model.

178

The probability maps for As concentrations exceeding 10 µgL⁻¹ are presented in 179 Figure 3 and supplementary Figure S3 (probability map of whole Southeast Asia). The 180 highest probabilities (0.7-0.8) for As contamination (>10 μ gL⁻¹) are found in the south-181 182 central part of Bangladesh, with a value of 0.5-0.6 in the north-eastern Sylhet basin 183 (Figure 3a). The probability of finding contaminated wells in the Red River delta 184 reaches a value of 0.7 (Figure 3e). The sedimentary depositional environments 185 present along the Mekong river differ from the deltaic environments of Bangladesh 186 and the Red River delta in that organic-rich deposits are found at close distance of the

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modern Mekong and Bassac river courses, surrounded by extensive floodplain areas
free of As¹⁸ (compare Figures 1 and 3). The probability map for the Mekong delta
shows values of up to 0.6 at close distance to the modern Mekong and Bassac river
courses and adjacent swampy marshes (Figure 3c). In addition, the large floodplain of
Lake Tonle Sap with organic-rich sediments is a risk area with probabilities ranging
between 0.4-0.6.

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194 To compare our predicted low-risk (probability ≤ 0.4) and high-risk areas (probability 195 >0.4) (Figures 3b and 3d) with measured As contamination, a misclassification 196 analysis was performed based on the 1756 aggregated and binary-coded (≤10 or >10 µgL⁻¹) As measurement data for the three deltas. We report 71% (1023 aggregated 197 198 points), 59% (107 a.p.), and 75% (100 a.p.) of correctly classified cases for 199 Bangladesh, Red River and Mekong deltas respectively (average 70%). In our study, 200 the Red River and Mekong deltas have 6 and 9% false negative cases. In Bangladesh false negative cases (13%) were found specifically in wells lying at a close distance to 201 202 rivers. The number of false negatives is outnumbered by false positives in all the three 203 deltas with 17, 35 and 16% for Bangladesh, the Red River and Mekong deltas, 204 respectively. Errors in the prediction can arise from the commonly reported well-to-well 205 variability where wells with low As levels are often present at close distance from wells high in arsenic^{5,6,18,19}, as well as from uncertainties in measured As (estimated at 25%) 206 leading to misclassifications of concentrations being close to the threshold of 10 µgL⁻¹. 207 208 However, we interpret these misclassifications to be mainly an effect of modelling 209 three-dimensional processes based on two-dimensional data.

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211 Apart from the three deltas discussed above, our Southeast Asia probability map

212 (supplementary Figure S3) also highlights risk areas that are largely unknown or

213 unreported, particularly in Sumatra, Myanmar and Thailand (Figure 3f).

214

215 VERIFICATION OF PREDICTIONS FOR SUMATRA

According to the modelled probability, an area of about 100,000 km² at Sumatra's east 216 217 coast is prone to high risk of As contamination (probability >0.4) (see Figure 4a and 218 supplementary Figure S3). To validate the Sumatra prediction map, 97 groundwater 219 samples were collected in 2007 in the Province of South Sumatra in the vicinity of 220 Palembang (see Methods section). This area was chosen because it is at the border 221 of a low- and high-risk area and not previously studied. Since there is no large 222 variation in geological and topographical features along Sumatra's east coast, the 223 study area is representative of the whole high risk area. Figure 4b shows As concentrations ($\leq 10 \ \mu g L^{-1}$ and $> 10 \ \mu g L^{-1}$) measured in the groundwater, imposed on 224 225 the binary probability map. The classification results for Sumatra (63% correctly 226 classified, 36% false positive and 1% false negative) are comparable to those of the 227 Bangladesh, Red River, and Mekong deltas. In total, 94% of the 50 tubewells located in the low-risk area have As levels below 10 µgL⁻¹. Of the 12 contaminated wells (As 228 >10 µgL⁻¹) 75% are positioned in the high-risk area. However, in the high-risk area the 229 230 contaminated wells are clearly outnumbered by uncontaminated groundwater 231 measurements (9 contaminated wells vs. 38 uncontaminated wells), for reasons 232 explained below.

233

On average, both high and low risk areas in Sumatra are characterized by high DOC, NH₄⁺, HCO₃⁻, PO₄³⁻, and Fe(II) and low SO₄²⁻ concentrations (Supplementary Table S5). At least two-thirds of the sampled groundwaters are reducing in nature, especially

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those located in the high-risk area. Since the chemical conditions of the aquifers would
permit the reductive dissolution of As, other explanations for the overall low As
concentrations must be considered.

240

241 To gain a greater insight into the characteristics of the aquifer, the local geology was examined laterally and in depth. The studied area is geologically young (Tertiary-242 Quaternary) (Figure 4c). The sediment contains deposits from peat swamp forests that 243 244 developed during the Holocene era (past 5,000 to 10,000 years) and unconformably overlie older sediments⁴¹. The outline of the high-risk As-contamination zone mainly 245 246 follows the outline of these deposits. The peat deposits are usually found ranging from 4-8 meters although depths of up to 24 metres have been reported⁴². However, the 247 248 depths of sampled tubewells in Sumatra average 46 meters, which implies that most 249 of them tap groundwaters from aquifers below the Holocene peat deposits. This 250 shows that the prediction map is a useful tool for the identification of areas at risk of 251 As contamination, but that understanding the local geology as a function of depth is of 252 vital importance for specific areas.

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254 ARSENIC CONTAMINATION IN THAILAND AND MYANMAR

255 The probability map in Figure 3f shows that the Chao Phraya basin in central Thailand 256 and the Irrawaddy delta in Myanmar have a risk of elevated As being present in 257 groundwater, whereas the Sittang basin (Myanmar) has a lower risk, and the Salween 258 basin (Myanmar) has virtually no risk at all. The problem of As contamination in the Irrawaddy delta is partly known⁴³ although its spatial extent has not been investigated 259 260 to date, which is particularly worrying considering the size of the area at risk. In the 261 Chao Phraya basin in central Thailand, a groundwater survey was undertaken in 2001 using As field test kits to test wells with minimum depths of 80 m⁴⁴. The range of 262

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263 measured As concentrations in the 37 tested wells was <1 μ gL⁻¹ to 100 μ gL⁻¹ with an 264 average of 11 μ gL⁻¹. These results correlate with our maps that predict a low to 265 moderate As contamination. Indeed, North-South geological profiles across the 266 basin⁴⁵ indicate maximal depths of 20 m for the pre-Holocene incisions implying that 267 sampled tubewells draw water from Pleistocene or older aquifers.

268

269 In contrast to the slow sedimentation rates of the Sittang and Chao Phraya deltas, the 270 Irrawaddy delta received massive amounts of sediments during the Cenozoic era⁴⁶. The Irrawaddy River⁴⁷ still has a 10 times larger sediment load than the Chao Phraya 271 River⁴⁵. In 2002 the Departments of Medical Research and Health (Lower Myanmar), 272 273 financially supported by UNICEF, conducted a groundwater sampling campaign in the Irrawaddy division⁴³. In total, 99 groundwater samples (90 shallow tubewells and 9 274 275 dug wells) were collected in 25 villages, and As was quatified by atomic absorption spectroscopy. It was reported that 67% of the sampled wells had As levels >50 μ gL⁻¹. 276 277 These results show that the risk of As contamination in the Irrawaddy delta, as 278 indicated by the probability map, should be taken seriously and that there is an urgent 279 need to test shallow tubewells in other townships or districts in the high risk area. 280

281 ASSESSING ARSENIC ELSEWHERE

In the study presented here, we identified regions in Southeast Asia, based on As
prediction maps, where tubewells should be tested for elevated As concentrations
(>10 µgL⁻¹). Although the use of such maps in other scientific investigations (e.g.
climate research) is a common practice, the prediction of As (and other groundwater
conditions) is still an emerging technique in the field of natural groundwater
contaminants. Our As model differs from earlier models in its ability to predict
contamination in areas of unknown groundwater quality and on a sub-continental

- 11-

289 scale. The strength of the prediction lies rather in the combination of surface 290 parameters than in the individual parameters. As an example, the Sittang basin and 291 the vicinity of the lower Mekong river branches are both characterized by alluvial 292 deposits, but only the latter has a modelled high-risk because of the contribution of soil 293 properties. A limitation to be considered is that shallow, As-bearing groundwater can 294 only be expected where Holocene aguifers are present. Where this is not the case, 295 high risk areas may indicate the presence of reducing aquifers, but As concentrations 296 in groundwater could be lower than predicted.

297

298 Our approach provides a blueprint for further modelling and mapping of As-tainted 299 aquifers in and outside of Southeast Asia. The probability maps can further be 300 improved when data with higher spatial resolution or in three dimensions becomes 301 available, although we emphasize that it will not be possible to account for the local 302 heterogeneities of aguifers. The presented prediction maps are a valuable and 303 resource-saving tool that can serve both scientists and policy-makers to initiate early 304 mitigation measures in order to protect the people from As-related health problems as 305 well as to efficiently guide water resources management.

306

307 Methods

Geological maps of Bangladesh, Cambodia, Thailand and Vietnam were available in
digital format (Supplementary Table S1). Maps of Myanmar and Sumatra were
digitized for this study. Geological maps of Malaysia and Laos were not available. The
sedimentary depositional environments employed in the model as independent
variables are deltaic deposits, organic-rich sediments (e.g. sediments deposited in
marshy environments), floodplains, alluvial deposits, tidal deposits, other Holocene
sediments and pre-Holocene sediments. Soil variables are percentages of silt, clay

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and sand in both the topsoil (0-30 cm) and subsoil (30-100 cm), and coarse, medium
and fine soil textures.

317

318 We compiled >4600 data points of groundwater As-concentrations from Bangladesh 319 (median well depth 35 m), the Mekong delta (Cambodia and Southern Vietnam, m.w.d. 320 39 m) and the Red River delta (Northern Vietnam, m.w.d. 30 m). The data originates from BGS and DPHE (Bangladesh, n=3534)⁵, from Buschmann et al. (Mekong delta, 321 n=352)^{18,20}, and from Berg et al. (Red River delta, n=720)^{17,25,48} and was used without 322 a restriction in well depths. To test the model, 97 tubewell were randomly sampled for 323 this study in Sumatra at about 5-10 km intervals at an average sampling density of 1 324 sample per 54 km². This study area is positioned at a latitude of 2°872S to 3°911S 325 and a longitude of 103°949E to 104°993E (sampling area 85 km by 65 km). 326 327 Procedures of sampling and analysis were carried out as described in Buschmann et al.¹⁸. Concentrations of As and additional parameters measured in Sumatra 328 329 groundwater samples are provided in Table S5 (Supplementary Information). 330 Arsenic concentrations are point measurements within vertical depths of the wells, 331 332 while the other variables have coarser spatial resolution being generally greater than 333 30 arc seconds. Point data of measured As-concentrations were therefore aggregated using the geometric mean to a resolution of one point per pixel with a size of 5 arc 334 335 minutes (~9.3 km at the equator), which is the pixel size of the global soil data (FAO). Aggregation resulted in a decreased dataset of 1756 pixel-based data points 336 337 (Bangladesh 1443, Red River Delta 180, and Mekong delta 133 points). The 338 aggregated point-data were binary-coded using the WHO guideline value for As in drinking water (10 μ gL⁻¹) as a threshold. We acknowledge that several countries apply 339 a guideline of 50 µgL⁻¹, but adopting this threshold would result in significantly less 340

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data points (992) for the calibration of the model. The binary variable of whether As
concentrations exceed the WHO threshold (1) or not (0) was used as the dependent
variable in this study.

344

Logistic regression was used to determine the weighting of the independent variables. This is a common statistical method used in environmental research and allows for the concurrent use of continuous and categorical variables^{33,34}. The parameter P denotes the probability of As concentrations exceeding the WHO threshold. Logistic regression models log(odds), which is defined as the ratio of the probability that an event occurs to the probability that it fails to occur log(P/(1–P)), as a linear combination of independent variables⁴⁹.

$$352 \quad \log(odds) = C + \sum_{i=1}^{n} \lambda_i X_i \tag{1}$$

353 where C is the intercept of regression, X_i are independent variables, and λ_i are the weighting coefficients that were obtained using the maximum likelihood procedure⁴⁹. 354 355 Exponential values of coefficients, Wald statistics, and p-values (see Table 1) indicate 356 the importance of each variable. Statistically insignificant independent variables were 357 excluded from the model during each of the subsequent regression steps (Table S2). 358 The threshold for maintaining a variable in the model was determined by the 95% 359 significance level (p < 0.05). The silt contents in the subsoil and medium textured soils 360 were kept in the model because of their good spatial match with known contaminated 361 areas, which is supported by the presence of silty sands at the surface of regions exhibiting contaminated aguifers in West Bengal⁵⁰, and by reported elevated 362 363 groundwater As concentrations in aguifers capped with fine surface material (silt and clav) in Bangladesh^{10,27}. 364

According to the calculated odds, the probability (P) of having an As concentration above 10 μ gL⁻¹ was calculated as follows:

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$$P = \frac{\exp(C + \sum_{i=1}^{n} \lambda_i X_i)}{1 + \exp(C + \sum_{i=1}^{n} \lambda_i X_i)}$$
(2)

In addition, we tested how successful the model predicted the number of
contaminated cases for probability intervals of 0.1 (Figure 2 and Supplementary
Information).

372

373 Misclassification occurs when either a point with As concentration >10 μ gL⁻¹ falls in an

374 uncontaminated area (false negative) or an As concentration <10 μgL⁻¹ falls in a

375 contaminated area (false positive). Based on the model classification results (Figure 2

and Table S4), a probability threshold of 0.4 resulted in a minimum misclassification

377 rate. The binary risk maps was hence categorized into low-risk areas (probability <0.4)

378 and high risk areas (probability >0.4).

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521 Competing financial interests statement

522 The authors declare that they have no competing financial interests.

523 Figure captions

Figure 1. Uniformly classified geological map of Southeast Asia. It indicates
seven different sedimentary depositional environments in the mapped countries of
Bangladesh, Myanmar, Thailand, Cambodia, Vietnam, and Sumatra (Indonesia).

Figure 2. Model classification results. The graph shows the sensitivity (true
positives) and specificity (true negatives) of the model for different probability cutoff
values. The full classification table is provided in the Supplementary Information
(Table S4). A probability threshold of 0.4 was applied to delineate low- and high-risk
areas in the binary risk maps shown in Figures 3b, 3d, and 4b (see Methods).

533

Figure 3. Modelled probability of As concentrations exceeding 10 µgL⁻¹ under 534 535 reducing aquifer conditions. a, Continuous probability map of Bangladesh. b, Binary 536 map of Bangladesh (probability threshold 0.4) indicating high- and low-risk areas 537 overlain by aggregated As concentrations. Areas where groundwater is mainly drawn 538 from Pleistocene aquifers are sketched in brown. c, Continuous probability map of the 539 Mekong delta (Cambodia and Vietnam). d, Binary risk map of the Mekong delta 540 overlain by aggregated As concentrations. e, Continuous probability map of the Red 541 River delta (Vietnam). f, Continuous probability map of the Irrawaddy delta (Myanmar) 542 and Chao Phraya basin (Thailand).

543

544 Figure 4. Maps of the model verification study area in Southeast Sumatra. a,

545 Map of whole Sumatra depicting the probability of groundwater As concentrations 546 exceeding 10 μ g L⁻¹. The corresponding colour code is given in Figure 3. **b**, Binary risk 547 map (probability cuttoff 0.4) and As concentrations measured for this study in the 548 vicinity of Palembang (South Sumatra). Swampy areas (high-risk area) are scarcely

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- 549 populated and the number of sampled tubewells in this area is therefore limited. **c**,
- 550 Geological map (source: see Supplementary Table S1).

552 Table 1. Results of logistic regression analysis. Weighting coefficients of the 553 independent variables (λ) were used to calculate probabilities of As contamination. 554 Wald values (%) indicate the relative importance of the variables, and p-values the 555 absolute significance, where a value <0.05 indicates a significance of at least 95%. 556 Variables that were not statistically significant (based on a 95% level) were excluded 557 from the model (Table S2), with exception of medium textured soils and silt in subsoil 558 (see Methods). Excluded variables are: tidal deposits, other Holocene deposits, pre-559 Holocene deposits (see Figure 1), coarse textured soils, soil sand and clay contents, 560 and climate.

561

	λ	Wald	p-value
Deltaic deposits	1.65	71.55	<0.001
Organic-rich deposits	1.11	72.20	<0.001
Alluvial deposits	0.55	12.51	<0.001
Floodplain deposits	- 0.95	2.32	0.010
Soil variables Medium textured soils	0.60	7.80	0.128
Fine textured soils	0.24	1.00	0.005
Silt in subsoil Silt in topsoil	0.10	6.62	0.317
	- 0.09	6.28	0.012
Regression constant	- 1.54	20.67	<0.001
	Organic-rich deposits Alluvial deposits Floodplain deposits Medium textured soils Fine textured soils Silt in subsoil Silt in topsoil	Deltaic deposits1.65Organic-rich deposits1.11Alluvial deposits0.55Floodplain deposits- 0.95Medium textured soils0.60Fine textured soils0.24Silt in subsoil0.10Silt in topsoil- 0.09	Deltaic deposits 1.65 71.55 Organic-rich deposits 1.11 72.20 Alluvial deposits 0.55 12.51 Floodplain deposits -0.95 2.32 Medium textured soils 0.60 7.80 Fine textured soils 0.24 1.00 Silt in subsoil 0.10 6.62 Silt in topsoil -0.09 6.28

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