



Rapid Fluvio-Thermal Erosion of a Yedoma Permafrost Cliff in the Lena River Delta

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The degradation of ice-rich permafrost deposits has the potential to release large amounts of previously freeze-locked carbon (C) and nitrogen (N) with local implications, such as affecting riverine and near-shore ecosystems, but also global impacts such as the release of greenhouse gases into the atmosphere. Here, we study the rapid erosion of the up to 27.7 m high and 1,660 m long Sobo-Sise yedoma cliff in the Lena River Delta using a remote sensing-based time-series analysis covering 53 years and calculate the mean annual sediment as well as C and N release into the Lena River. We find that the Sobo-Sise yedoma cliff, which exposes ice-rich late Pleistocene to Holocene deposits, had a mean long-term (1965–2018) erosion rate of 9.1 m yr⁻¹ with locally and temporally varying rates of up to 22.3 m yr⁻¹. These rates are among the highest measured erosion rates for permafrost coastal and river shoreline stretches. The fluvio-thermal erosion led to the release of substantial amounts of C (soil organic carbon and dissolved organic carbon) and N to the river system. On average, currently at least 5.2 × 10⁶ kg organic C and 0.4 × 10⁶ kg N were eroded annually (2015–2018) into the Lena River. The observed sediment and organic matter erosion was persistent over the observation period also due to the specific configuration of river flow direction and cliff shore orientation. Our observations highlight the importance to further study rapid fluvio-thermal erosion processes in the permafrost region, also because our study shows increasing erosion rates at Sobo-Sise Cliff in the most recent investigated time periods. The organic C and N transport from land to river and eventually to the Arctic Ocean from this and similar settings may have severe implications on the biogeochemistry and ecology of the near-shore zone of the Laptev Sea as well as for turnover and rapid release of old C and N to the atmosphere.

Keywords: permafrost, river delta, ice-rich, carbon, nitrogen, time series, Landsat

INTRODUCTION

Permafrost landscapes are sensitive to global temperature rise and may be affected by widespread degradation (Grosse et al., 2016; Biskaborn et al., 2019). Very ice-rich permafrost deposits such as the yedoma Ice Complex (Schirrmeister et al., 2013; Strauss et al., 2013, 2017) are particularly at risk to rapid, strong thaw processes and erosion. The thaw of ice-rich deposits leads to strong ground subsidence (Strozzi et al., 2018) and landscape reorganization (Morgenstern et al., 2011). Coasts and river banks in the Arctic are particularly fast changing permafrost features, which are characterized by high ground ice contents (e.g., Walker et al., 1987; Lantuit et al., 2011; Kanevskiy et al., 2016; Jones et al., 2018). These shores are affected by thermo-erosion processes induced by sea or river water combining mechanical erosion from waves, currents, and moving ice with thermal impacts of the water that is warmer than the permafrost (Are, 1983; Günther et al., 2013). The strong erosion of river banks and coastal segments leads to large sediment and organic matter inputs into Arctic river systems and the Arctic Ocean, respectively (e.g., Rachold et al., 2000; Couture et al., 2018; Rowland et al., 2018) with regional to global implications affecting the aquatic ecosystem and global C cycle (Gustafsson et al., 2011; Vonk and Gustafsson, 2013; Mann et al., 2015; Semiletov et al., 2016; Fritz et al., 2017). The C release upon permafrost thaw has been shown to have the potential to further increase atmospheric C resulting in a positive permafrost C climate feedback (Schoor et al., 2009). The deep mobilization of permafrost C caused by thermo-erosional processes along shores might add C to the atmosphere, which is not yet taken into account in current Earth System Models that include only top-down thaw (Turetsky et al., 2020). However, how much of the C released through shore erosion is re-buried on the Arctic shelf or in the deep Arctic Ocean versus is being mineralized and released into the atmosphere is still a matter of discussion (Vonk et al., 2012; Bröder et al., 2019; Grotheer et al., 2020).

Often, rapid erosion of ice-rich permafrost landscapes such as retrogressive thaw slumping (Lantz and Kokelj, 2008; Balsler et al., 2014) and river bank erosion (Walker et al., 1987; Kanevskiy et al., 2016) occurs on local spatial scales and is therefore challenging to include in permafrost C feedback models (Turetsky et al., 2020). Nevertheless, observations of increasing air and river water temperatures (van Vliet et al., 2013), decreasing sea ice cover (Stroeve and Notz, 2018), extended thaw season length (Serreze et al., 2000; Barichivich et al., 2012), and an increase in Arctic river discharge (Holmes et al., 2015; Box et al., 2019) and therefore a higher sediment transport capacity all point at a high probability for accelerated erosion rates along permafrost coasts and river shorelines in the Arctic. Several studies already indicate that acceleration of erosion is taking place over the recent decades in different regions (e.g., Günther et al., 2015; Irrgang et al., 2018; Jones et al., 2018) in response to large scale drivers such as changing sea-ice cover (Nielsen et al., 2020). It is therefore important to understand and quantify the dynamics and C mobilization of permafrost shore erosion across a range of environmental and geographical settings.

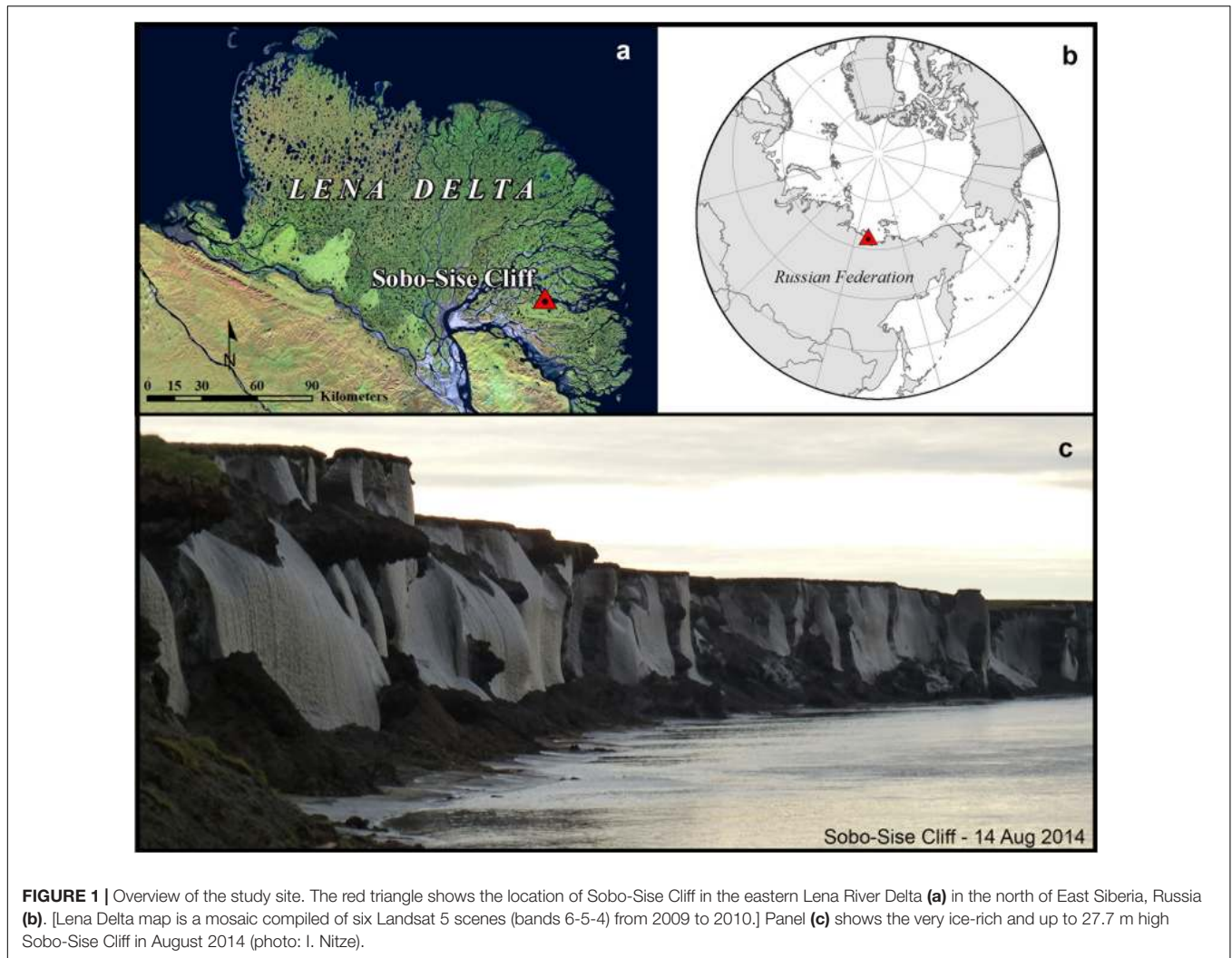
Several segments of the Arctic coast (e.g., Jorgenson and Brown, 2005; Jones et al., 2008, 2009; Lantuit et al., 2011; Ping

et al., 2011; Günther et al., 2013; Obu et al., 2017; Couture et al., 2018; Irrgang et al., 2018; Jones et al., 2018) and some individual permafrost shore sites (e.g., Walker et al., 1987; Costard et al., 2003; Günther et al., 2015; Kanevskiy et al., 2016; Stettner et al., 2018) affected by thermo-denudation and thermo-erosion have been previously studied in terms of erosion rates and organic matter flux. Here we study the fluvio-thermal erosion dynamics of a carbon-rich and fast-eroding river bank in the eastern Lena Delta, Siberia, using remote sensing image analysis and geochemical field data from sediment samples. The Sobo-Sise Cliff in the eastern Lena Delta is composed of ice-rich yedoma permafrost and strongly eroded by a main branch of the Lena River. With our study, we aim (1) to assess the cliff erosion rates of this yedoma Ice Complex exposure and (2) to quantify the sediment, C and N loss from yedoma deposits due to river bank erosion. Our study highlights the potential for continued and very rapid erosion in permafrost regions under certain environmental and geological conditions by Arctic rivers and the related environmental impacts of such a high organic matter export.

STUDY SITE

Sobo-Sise Island is located in the eastern part of the Lena River Delta at the Sardakhskaya Lena river channel (**Figure 1**), which is one of the main Lena Delta river channels with a width of around 2 km and a water discharge of $\sim 8000 \text{ m}^3 \text{ s}^{-1}$ during the summer low-water period (Fedorova et al., 2015). The Lena River is ice-covered for about 8 months (October–May) during wintertime where presumably very little or no erosion happens on the subaerial frozen deposits along the river. However, ice thickness on the Lena River is up to 2 m while the water depth at the onset of the Sardakhskaya channel can reach 22 m (Fedorova et al., 2015) and around 11 m in front of Sobo-Sise Cliff (see **Supplementary Figure S1**) allowing constant water flow underneath also in winter. The delta island Sobo-Sise is characterized by late Pleistocene, ice-rich yedoma Ice Complex deposits with Holocene cover deposits. The yedoma deposits at Sobo-Sise are affected by rapid thaw processes such as thermokarst lake dynamics and thermo-erosional valley formation (Morgenstern et al., 2011; Nitze and Grosse, 2016) as well as thaw subsidence (Chen et al., 2018). Only 19% of the area of Sobo-Sise Island consist of largely un-affected yedoma uplands, whereas the remaining area of the island consists of degraded thermokarst basins (alases) and fluvial deposits (Fuchs et al., 2018). Most of the year, snow and ice cover the landscape and water bodies, respectively. With a mean annual air temperature of -11.7°C (1998–2017) and a mean summer rainfall of 145 mm (weather station on Samoylov Island 63 km to the southwest, Boike et al., 2019), Sobo-Sise is dominated by a polar tundra climate (Peel et al., 2007).

The Sobo-Sise Cliff ($72^\circ 32' \text{ N}$, $128^\circ 17' \text{ E}$) is a 1,660 m long riverbank stretch in the central North of Sobo-Sise Island, with a northeast exposition. The vertical cliff erodes into yedoma uplands with heights up to 27.7 m above mean river water level, exposing yedoma Ice Complex and overlying Holocene deposits,



which can reach a thickness of up to two meters (Fuchs et al., 2018). Pavlova and Dorozhkina (2000) report that yedoma in the eastern part of the Lena Delta reaches to 8 to 10 m below sea level, resulting in ice-rich permafrost being exposed directly to above and underwater thermal erosion dynamics. Large syngenetic ice wedges, which dissect the entire height of the exposure, dominate the cross section exposed by erosion. The Sobo-Sise Cliff is directly exposed to the river flow of the Sardakhskaya channel, which hits the cliff at an angle of $\sim 30^\circ$. The position of the thick ice-rich deposits at this bend of the channel causes favorable conditions for erosion, removal of large sediment masses, and the formation of a steep cliff.

MATERIALS AND METHODS

Calculation of Total and Annual Erosion Rates

For the detection and analysis of recent and historic cliff retreat rates, we used remote sensing imagery from a variety

of high to medium resolution optical sensors with different spatial resolution.

Image Data Sets

For historical long-term analysis, we used Corona and Hexagon satellite images from 1965 and 1975 (**Table 1**). For the following time steps, we used Landsat data (L5, L7, L8) for the years 2000, 2005, 2010, and 2015. Unfortunately, Landsat data from before 1999 is only sparsely available over large parts of NE Siberia due to the lacking on-board memory on the platform and the region being outside of footprints of high latitude satellite receiving stations (Goward et al., 2006). Spatially very high-resolution imagery from the DigitalGlobe constellation was available for 2014 (GeoEye). Additionally, we used Planet cube-sat imagery (Planet Team, 2017) from September 2018 to finalize the long-term erosion analysis.

While Landsat data is already delivered as geometrically and radiometrically corrected data, the Corona, Hexagon and high-resolution images were orthorectified and georeferenced using the GeoEye image from 2014 as a base image (see **Table 1**) and projected to WGS84 UTM 52N. The matching between

TABLE 1 | Images used for calculation of erosion rates.

	Image dates	Ground resolution (m)	Scale used for shoreline digitization
Corona KH-4A	1965-10-01	2.5	1:2500
Hexagon KH-9	1975-07-16	10	1:5000
Landsat 7 ETM+	2000-07-29	30	1:10000
Landsat 7 ETM+	2005-07-02	30	1:10000
Landsat 5 TM	2010-09-19	30	1:10000
Landsat 8 OLI	2015-09-08	30	1:10000
GeoEye ^a	2014-08-07	0.5	1:2500
PlanetOrthoTile	2018-09-13	3.15	1:5000

^aMaster image for georeferencing. This image was not used for the calculation of erosion rates.

images for the georeferencing was done by selecting features that we considered most stable in a landscape that is affected by change and does not contain artificial human-made structures. Such features included the centers of small circular thaw ponds that largely expand radially or polygonal ponds and ice wedge junctions that are rather stable over the decadal time scales analyzed with this imagery.

Calculation of Erosion Rates

Since the Sobo-Sise Cliff is a nearly vertical cliff with up to 27.7 m height, it was sufficient to manually digitize the upper cliff line on the selected imagery at an appropriate image scale (see **Table 1**) to quantify the cliff retreat rates between several points in time. Using the oldest image in the time-series, we also digitized the boundaries between yedoma uplands and thermokarst basin based on the Corona image to determine the volume of yedoma deposits, which was eroded since 1965 and to separate it from non-yedoma deposits.

For calculating erosion rates, we used the Digital Shoreline Analysis System 5.0 (DSAS; Himmelstoss et al., 2018) in ArcGIS 10.5. DSAS calculates erosion rates between given shorelines along defined transects perpendicular to a defined baseline. In DSAS the “baseline” was defined as the straight middle line between the 1965 and 2000 cliff lines. We then cast transects at an interval of 50 m for the entire 1,660 m coastal section (resulting in 32 transects) to calculate erosion rates for individual cliff segments. We calculated erosion rates for the entire period between 1965 and 2018, as well as several sub-periods, including 1965–1975, 1975–2000, 2000–2005, 2005–2010, 2010–2015, and 2015–2018. Erosion rates are given as end point rates (EPR) of DSAS. The digitized cliff fronts for each time step as well as the erosion rates for each of the 32 segments are available on the PANGAEA data set repository (Nitze et al., 2020).

For volumetric erosion estimates, we extracted 10 m ARCTIC DEM elevation data (Mosaic v3.0 10m: tile 59_43) (Porter et al., 2018), based on representative sample areas in proximity to the cliff line for yedoma areas and the eastern and western alases as well as the river water level. We extracted elevation statistics for these training areas (see **Supplementary Figure S3**) and transformed absolute elevation to relative elevation above the mean water level. However, as we do not have any height

information for the older time steps (prior to 2000) and assume, based on topographic maps and the Corona scene, that there existed a down-slope toward the Sardakhskeya channel sometime prior to 2000, we do not calculate volumetric, C and N erosion rates prior to 2000.

Calculation of Total and Annual C and N Loss

Elemental Composition and Stock Calculations

Frozen sediment samples were taken by rope climbing across the entire vertically exposed yedoma sequence at 0.5 m intervals at three locations in close proximity and overlapping vertical profiles during an expedition to Sobo-Sise in July 2018 (Wetterich et al., 2019, 2020). In total, 60 permafrost samples were used for analyses of organic matter properties. The gravimetric segregated and pore ice content was measured in the field as difference between wet and dry weights after careful drying on a wood oven. The ice content is expressed as weight percentage (wt%, van Everdingen, 2005).

Gravimetric ice content for the samples was then converted into volumetric ice content by assuming ice saturation if the ice content is >20% (following Strauss et al., 2012) and an ice density of 0.917 g cm⁻³. For estimating the segregated and pore ice volume in %, we assumed a 3-component model of the cliff consisting of ice, a mineral component of the sediment and organic matter, and assuming component densities of 2.65 g cm⁻³ (Lide et al., 2008) and 0.25 g cm⁻³ (Adams, 1973) for the latter two components, respectively. We used this for estimating the total volume of ice and organic matter at the Sobo-Sise Cliff.

The total C and total N content of the sediment samples were measured with a Vario EL III and total organic C with a Vario Max C elemental analyzer (Elementar Analysensysteme, 2007, 2011). Results are expressed as weight percentage (wt%) of the dry sediment samples. The soil organic carbon (SOC) density, which is organic C content per soil volume unit (kg C m⁻³), was quantified following Strauss et al. (2012). Prior to conversion of the measured weight-based total organic C values to the volume-based SOC, the bulk density was estimated according to Strauss et al. (2012). More details on the bulk density calculation are presented in the **Supplementary Material** (section 2).

In addition, dissolved organic carbon (DOC) concentrations (mg L⁻¹) of segregated and pore ice were measured in 29 samples from throughout the cliff. Therefor meltwater was filtered in the field with gum-free syringes equipped with glass fiber filters (Whatman GF/F; pore size: 0.7 μm), which were rinsed with ≥20 ml of sample, and acidified with 20 μL HCl suprapur (30%) to pH < 2 in order to prevent microbial conversion. In the laboratory, samples were analyzed with a high-temperature (680°C) combustion total organic C analyzer (Shimadzu TOC-VCPH). This allows us to make a complete estimation of C loss through erosion of Sobo-Sise Cliff. For N we refer to the N loss, without differentiation between particulate and dissolved N or the form of N (e.g., organic or inorganic).

Ice Wedge Content Estimation

For the estimation of ice wedge content (IWC) we used photographs with frontal views of the Sobo-Sise Cliff taken

during field visits in summer 2014, 2015, and 2019. All analyzed photographs and front cliff sections are provided in the **Supplementary Figure S4**. In these photographs, we identified vertically continuous sediment sections and ice wedges to quantify volumetric ratios between sediment and ice wedge. We correlated the width of ice wedge sections with the width of the adjacent sediment section assuming that ice and sediment segments dissecting the entire cliff were eroded perpendicular to the polygon structure (**Figure 2**). This ice wedge to sediment column width ratio ($I:S_{width}$) was calculated according to Eq. (1).

$$I : S_{width} = Width_{ICE} / Width_{SED} \quad (1)$$

Further, assuming an ideal polygonal distribution in the landscape prior to erosion, the size of sediment columns (inner polygon) was related to adjacent ice wedges (outer polygon) (**Figure 2**). The $I:S_{width}$ was then used to calculate the volume of both inner (V_{IP} ; Eq. 2) and outer polygon (V_{OP} ; Eq. 3) and in a final step to calculate the amount of ice wedge content (IWC; Eq. 4) within the landscape.

$$V_{IP} = \left(\frac{a}{2}\right)^2 * 2\sqrt{3} * h \quad (2)$$

$$V_{OP} = \left(\left(\frac{1}{2} * \left(\frac{a}{I : S_{width}} + a\right)\right)^2 * 2\sqrt{3} * h\right) - V_{IP} \quad (3)$$

$$IWC [vol\%] = \left(\frac{V_{OP}}{V_{OP} + V_{IP}}\right) * 100 \quad (4)$$

For the diameter of the inner polygon (a) and the vertical height (h) of a polygon any reasonable realistic number for known polygon diameters and vertical heights can be chosen, since we calculate with a ratio, this does not change the final result of ice wedge content. $I:S_{width}$ for the 10 analyzed sections varied from 0.28 to 1.21 (see **Supplementary Table S1**). For each section, we calculated the ice wedge content and then calculated the average ice content over all our ten measured sections. In addition, to get an estimate for the total landscape ice volume we used Eq. (5):

$$total\ ice\ volume\ [vol\%] = IWC [vol\%] + \frac{100 - IWC [vol\%]}{100} * segregated\ and\ pore\ ice\ [vol\%] \quad (5)$$

Calculation of Annual C and N Loss

The data on ice volume and the mass of the mineral and organic components of the yedoma as well as the data on the retreat rates of the Sobo-Sise Cliff were used to estimate the amount of eroded sediments, SOC, DOC and N. According to Kanevskiy et al. (2016) we estimated the total volume (V) of frozen deposits (sediment and ice) transported annually to the river from the retreating bank using Eq. (6).

$$V[m^3 yr^{-1}] = L * H * R \quad (6)$$

L is the length of the cliff [m], H is the average height of the cliff [m above mean river water level], and R is the average retreat

rate [$m yr^{-1}$] for the cliff during the period of observation. We then calculate the flux of the dry weight of sediments (Sed_{flux}), the flux of dry weight of soil organic carbon (SOC_{flux}) and the flux of dissolved organic carbon (DOC_{flux}) lost by erosion and transported to the river (Eqs 7–9):

$$Sed_{flux}[kg yr^{-1}] = sediment\ density * V * (1 - IWC) \quad (7)$$

$$SOC_{flux}[kg yr^{-1}] = organic\ matter\ density * V * (1 - IWC) \quad (8)$$

$$DOC_{flux}[kg yr^{-1}] = DOC\ concentration * V * (1 - IWC) * ice\ content * 0.917 \quad (9)$$

The flux of the dry weight of nitrogen (N_{flux}) was calculated like SOC_{flux} following Eq. (8) by replacing organic matter density by the N density. For DOC_{flux} (Eq. 9) the segregated and pore ice volume (ice content) was considered. In addition, the volume of ice was converted in volume of water with the density difference between ice and water of 0.917. In order to get the complete DOC loss from Sobo-Sise Cliff, ice wedge DOC flux was calculated including the mean DOC concentration of $11.1\ mg\ L^{-1}$ for yedoma ice wedges based on Fritz et al. (2015), since we do not have own DOC concentrations from ice wedges at Sobo-Sise Cliff. This was combined with the total volume of ice wedges at Sobo-Sise Cliff derived from the ice wedge content estimation (see section “Ice Wedge Content Estimation”). As a result, the DOC_{flux} of Sobo-Sise Cliff is the sum of the DOC_{flux} from segregated and pore ice and the DOC_{flux} from eroded ice wedges for each investigated time period. In consequence, the total C flux from the yedoma Sobo-Sise Cliff includes the SOC, DOC from segregated and pore ice, and DOC from ice wedges. For this study, we only focus on C and N loss for the yedoma cliff. The two alas basins in the east and west of Sobo-Sise Cliff are eroding as well but were not considered here.

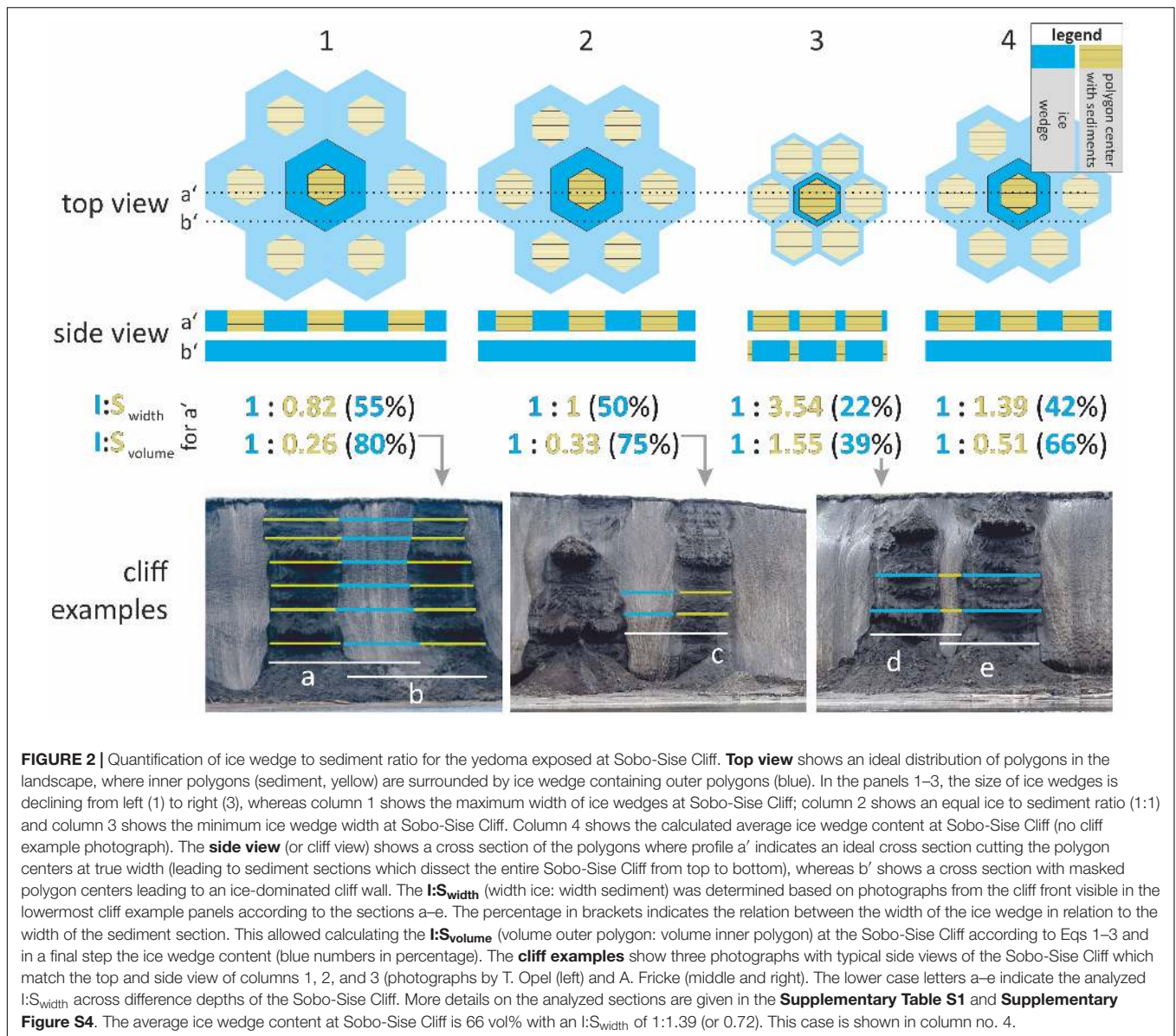
Bathymetric Survey Along the Sobo-Sise Cliff

In order to better understand how the fluvial erosion processes contribute or drive the cliff retreat a bathymetric survey was done in August 2016 during calm weather and at low speed ($<4\ km\ h^{-1}$) with a Humminbird 899cxi HD SI COMBO to detect the water depth in front of Sobo-Sise Cliff. Along four parallel and one zigzag profile lines, about 9200 depth measurements were collected. The surveyed area in the Sardakhskaya channel captured the entire cliff length and included parts of the adjacent alas basins. Bathymetric profiles covered a zone of typically 50 m width off the lower cliff base. Average water depth was $7.2 \pm 4.9\ m$.

RESULTS

Sobo-Sise Cliff Characteristics

According to the Arctic DEM, the Sobo-Sise Cliff (**Figure 3a**) has an average height of 22.3 m above mean river water



level (maximum height: 27.7 m) extending for 1,660 m from $72^{\circ}32'34$ N/ $128^{\circ}15'59$ E to $72^{\circ}32'06$ N/ $128^{\circ}18'21$ E in a concave shape along the Sardakhskaia channel in the Eastern Lena River Delta. The cliff emerged and has developed its shape from a formerly undisturbed yedoma upland gently down-sloping to the Sardakhskaia Lena River channel. Interpreting topographic maps that are based on aerial surveys from the early 1950s, the riverbank next to the yedoma hill was lower than 10 m and likely consisted of alas deposits. Over time riverbank erosion reached the yedoma upland triggering the formation of the cliff likely during the early 1970s. The cliff front is dominated by late Pleistocene ice wedges stretching the entire height of the cliff. Our ice wedge volume estimation resulted in an average of 66 vol% of wedge ice at Sobo-Sise Cliff and 34 vol% of sediments (Figure 2) with high segregated and pore ice contents. The base of

the cliff is often not visible due to sediment debris falling from the cliff onto the beach, forming piles that are rapidly removed by the river.

The sediment columns have an average segregated and pore ice content of 65.7 ± 9.0 vol%. This results in a total ice volume of 88.4 vol%. The segregated and pore ice content is included in the dry bulk density of the sediment and is therefore integrated in the calculation for mean SOC and N contents. Average dry bulk density of the Sobo-Sise Cliff sediments is 0.7 ± 0.2 g cm⁻³ resulting in average SOC and N contents of 26.2 ± 16.8 kg C m⁻³ and 2.1 ± 1.1 kg N m⁻³, respectively. If accounted for ice wedge volume (66 vol%) on Sobo-Sise Cliff, average SOC and N contents are 8.9 ± 5.7 kg C m⁻³ and 0.7 ± 0.4 kg N m⁻³. Beside SOC, DOC was analyzed from segregated and pore ice within the sediment column. Mean DOC content for Sobo-Sise Cliff is 315.5 ± 188.0 mg L⁻¹ or 0.316 kg m⁻³. For comparison,



FIGURE 3 | (a) Sobo-Sise Cliff in August 2014 with person for scale (photo: T. Opel, 11 August 2014), (b) Sobo-Sise Cliff in April 2019 with snow drifts (photo: J. Strauss, 5 April 2019), (c) cliff-top with a large crack (photo: G. Grosse, 12 August 2014), (d) cliff-top with thin snow cover (photo: J. Palmtag, 5 April 2019), (e) collapsing sediment block at the Sobo-Sise Cliff (photo: G. Grosse, 12 August 2014), (f) thermo-erosional niche and peat blocks at the base of the cliff (photo: M. Fuchs, 12 August 2014).

the average DOC content for yedoma ice wedges from Fritz et al. (2015) is 11.1 mg L^{-1} or 0.011 kg m^{-3} .

During the ice-free months (June–September), the cliff is prone to erosion mainly by flowing river water and air temperatures above freezing. Thermo-erosional niches (Figure 3f) form at the base of the cliff leading to collapse and failure of large ice and sediment blocks (Figures 3c,e).

In winter (October–May), low negative air temperatures and presence of a stable ice cover on the river prevent ground ice melt and permafrost thaw on the subaerial part of cliff, while below river ice permafrost thaw continues undercutting the cliff during the winter providing favorable conditions for block collapse in the following summer. The wind exposed cliff top is covered by a thin snow cover ($<20 \text{ cm}$; April 2019), whereas

large snow packs accumulate in windblown drifts at the cliff base (Figures 3b,d).

Dense bathymetric surveys in front of Sobo-Sise Cliff were combined with synchronous topographic survey data. The interpolated surface shows an almost immediate downward extension of the steep Sobo-Sise Cliff slope below river level (Figure 4). Within a 20 m wide shoreface zone water depths were 2.2 ± 1.2 m, while the adjacent 50 m of the cliff had mean depths of 10.3 ± 3.7 m, while maximum depths were recorded down to 19.1 m. These data suggest a narrow zone next to the cliff base, where river ice may freeze down to the bottom. Because the bathymetric survey was conducted in August several months after river ice breakup, the shoreface configuration might have changed considerably, potentially allowing for configurations during winter with a negligible bottom fast ice formation that enables direct contact of the river water flow with the cliff base and permafrost thaw throughout the winter.

Total and Annual Erosion Rates

The total erosion of the Sobo-Sise Cliff from 1965 to 2018 ranged from 322 to 679 m (Figure 5). The highest mean annual erosion rates (ER_{max}) occurred in the downstream section of the cliff with a rate of 12.8 m yr^{-1} whereas the upstream edge has a mean annual erosion rate of 6.1 m yr^{-1} . This results in an overall mean annual erosion rate (ER_{mean}) for the entire Sobo-Sise Cliff of 9.1 m yr^{-1} ranging from 4.8 to 15.7 m yr^{-1} for the mean values of the different periods. The cliff's planimetric shape changed from convex in 1965 to its present concave shape and the length of the cliff with exposed yedoma is ~ 610 m longer in 2018 than in 1965. The point of maximum erosion gradually moved upstream over time. However, ER_{mean} since 2005 are more uniform along the Sobo-Sise Cliff compared to the period before 2000.

Over the entire period from 1965 to 2018 a total area of 0.88 km^2 was eroded, of which 0.77 km^2 (87.9%) were yedoma and 0.11 km^2 (12.1%) drained lake basin. In general, erosion rates increased over time but also show short term variation (Table 2). The ER_{mean} increased considerably from 4.8 m yr^{-1} in 1965–1975 to 15.7 m yr^{-1} in 2015–2018, while the ER_{max} increased from 8.2 to 21.1 m yr^{-1} .

The recent multi-year periods 2000–2005, 2005–2010, 2010–2015, and 2015–2018 exhibited a strong variation of erosion rates. The early period from 2000 to 2005 was characterized by less erosion with ER_{mean} of 5.7 m yr^{-1} and ER_{max} of 11.8 m yr^{-1} . During 2005–2010 erosion rates accelerated to ER_{mean} of 15.1 m yr^{-1} and ER_{max} of 22.3 m yr^{-1} , the highest measured erosion rate. From 2010 to 2015 erosion rates slightly receded to ER_{mean} of 13.8 m yr^{-1} and ER_{max} of 19.3 m yr^{-1} , which still exceeded long-term mean erosion rates (1965–2018) by 51%. For the most recent period (2015–2018) ER_{mean} increased again to 15.7 m yr^{-1} .

Total and Annual Sediment, C and N Loss

Even though the total volume of ice is 88 vol% at the Sobo-Sise Cliff, large amounts of yedoma sediments have been eroded since 2000 (Table 3). The average eroded sediment flux range from 47.1 – $154.4 \times 10^6 \text{ kg yr}^{-1}$ leading to a loss of 1.7 – $5.6 \times 10^6 \text{ kg SOC yr}^{-1}$ and 0.1 – 0.4 kg N yr^{-1} for the different time periods

between 2000 and 2018. When splitting into different time periods, the highest SOC_{flux} and N_{flux} occurred between 2005 and 2010 with $5.6 \times 10^6 \text{ kg SOC yr}^{-1}$ and $0.4 \times 10^6 \text{ kg N yr}^{-1}$. However, even after 2010 the erosion rates remained on a very high level with annual erosion rates of $\sim 5 \times 10^6 \text{ kg SOC yr}^{-1}$ (Table 3). For the time period 2015–2018 in total $0.59 \times 10^6 \text{ m}^3$ of yedoma sediments (including segregated and pore ice but not including the adjacent shores along drained lake basins and the deposits below water level) were eroded into the Sardakhskaya channel, which led to a loss of $15.4 \pm 9.9 \times 10^6 \text{ kg SOC}$.

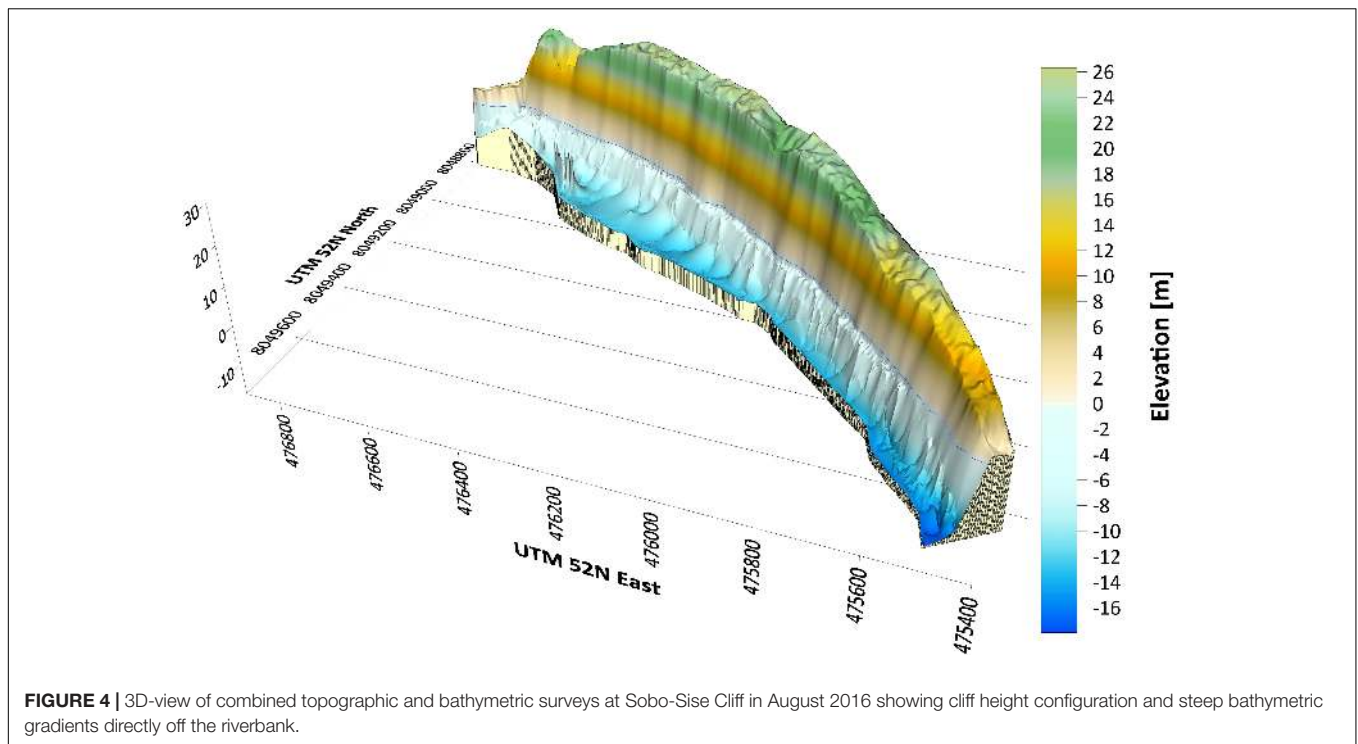
In addition to SOC and N, DOC is released from Sobo-Sise Cliff. Average DOC fluxes (including ice wedge as well as segregated and pore ice DOC) range from 13.6 to $44.8 \times 10^3 \text{ kg DOC yr}^{-1}$ for the different time periods between 2000 and 2018 (Table 3). Like sediment, SOC and N fluxes, highest DOC fluxes were estimated for the time period 2005–2010. However, compared to the SOC flux, DOC only contributes a minor amount of the total C flux of Sobo-Sise Cliff. For the four investigated time periods, DOC accounts for 0.8% of the total C lost through erosion. Nevertheless, this still accounts for a loss of $41,265 \text{ kg DOC yr}^{-1}$ in the most recent investigated time period (2015–2018). As a result, the flux ratio of DOC vs. SOC is 1:125. In total, SOC and DOC combined, $15.5 \pm 9.9 \times 10^6 \text{ kg C}$ were eroded during 2015–2018. This leads to an average C flux of $5.2 \pm 3.3 \times 10^6 \text{ kg C yr}^{-1}$ or divided by meter coastline to an average of $3,131 \text{ kg C m}^{-1} \text{ yr}^{-1}$ along Sobo-Sise Cliff. However, the overall contribution of yedoma ice wedge DOC is relatively small. For the time period 2015–2018, $11.74 \times 10^3 \text{ kg DOC}$ from ice wedges were eroded, which accounts for 9.5% of the total DOC flux and 0.08% of the total C flux from Sobo-Sise Cliff. Ice wedge DOC fluxes for the different time periods are presented in the Supplementary Table S2.

DISCUSSION

Comparison to Other Key Sites

The presence of a very ice-rich yedoma upland at a location where a major channel of the Lena Delta develops a large bend is certainly a key factor for the exceptional high erosion rates of the Sobo-Sise Cliff. These site characteristics favor high erosion rates and the local dynamics suggest that a wide range of permafrost shore erosion rates, including very high maximum rates may develop not only along wave-dominated sea coasts, but also along rivers. Maximum annual shore erosion rates of 22 m yr^{-1} on this Lena river bank exceeds all previously observed rates for Arctic rivers. Erosion dynamics along the cliff rival some of the most rapid permafrost coastal erosion rates in the Arctic and as it is the case for some sections at sea coasts (e.g., Günther et al., 2013; Irrgang et al., 2018; Jones et al., 2018), erosion of the Sobo-Sise Cliff is accelerating in the last decades.

The warm water transported by the large Lena River from further south can be considered as a contributing factor for the erosion. Water temperatures up to 15°C were measured in August 2019 (see Supplementary Figure S1). The warm and rapidly flowing water in front of the Sobo-Sise Cliff triggers the



formation of thermo-erosional niches at the base of the ice-rich cliff and continuously removes eroded debris from the base during summertime. The lack of a beach and a water depth of ~ 10 m directly in front of the cliff indicate that sediment entering the river is rapidly eroded also at depth and not accumulating. All this leads to formation of a nearly perfectly vertical cliff front where hardly any thaw slump-like features are developing at the cliff top due to the very rapid thermo-erosion of the cliff base.

For comparison, another yedoma cliff on Kurungnakh Island in the central Lena Delta has much smaller erosion rates with 4.1 to 6.9 m yr^{-1} (Stettner et al., 2018). On Kurungnakh Island, the yedoma is located on top of ice-poor fluvial sand layers (Wetterich et al., 2008), which limits erosion rates at the cliff base. Most of the erosion here currently takes place by thermodenudation, i.e., the retreat of the bluff top by thaw slumping and forming slump floors at the stratigraphic boundary between yedoma and underlying sands roughly at 17 m above river level. The Olenekskaya channel that flows around Kurungnakh Island has a low runoff rate and therefore weaker river currents compared to the Sardakhsakaya channel (Fedorova et al., 2015). For comparison, in the Olenekskaya channel in the Lena Delta, Are (1983) also reported smaller erosion rates of 1.7 – 6 m per year in up to 9 m thick deposits dissected by wedge ice.

Another key site studied for riverine yedoma cliff erosion is the Itkillik exposure on the North Slope of Alaska (Kanevskiy et al., 2016). With up to 35 m height above river water level, this 680 m long cliff retreats by 11 m yr^{-1} and erodes 70×10^6 kg yr^{-1} of sediments and $880,000$ kg C yr^{-1} . Similar to the Sobo-Sise Cliff, the Itkillik cliff is very ice-rich (total ice content of 86 vol%) and lies on a river bend (Kanevskiy et al., 2016). However, the

Sobo-Sise Cliff is more than twice the length resulting in an even higher sediment input into the river compared to the Itkillik exposure. A similar ice-rich permafrost cliff was studied by Shur et al. (2002) on the Yana River close to the village Kazachye. This 15 m high cliff had erosion rates of 6.5 m yr^{-1} from 1975 to 1990.

However, not only yedoma dominated cliff retreat leads to high erosion of sediment and C. Along the Colville River and in the Colville Delta, erosion rates of up to 3.5 m yr^{-1} (Walker et al., 1987; Payne et al., 2018) lead to a continuous transport of organic matter into delta and near-shore zones of the Colville River. Likewise, the Lena River is also eroding permafrost along its course and mobilizing sediments and organic matter. Average bank erosion on a 300 km stretch of the Lena River close to Yakutsk is 2 m yr^{-1} for the period 1967–2002 (Costard et al., 2007). At the Indigirka and Yana rivers, strong erosion of permafrost shores averages 6.5 m yr^{-1} and 2 m yr^{-1} , respectively (Shur et al., 2002). All these observations highlight the importance of major Arctic rivers for the mobilization and redistribution of formerly frozen sediments and organic matter into the fluvial and further downstream environments such as deltas and Arctic shelves (e.g., Wild et al., 2019).

Similar to the Sobo-Sise Cliff with a yedoma Ice Complex base of at least 10 m below river level, these strongly erosion-susceptible yedoma deposits extend down to 10 m below sea level on Muostakh Island (Overduin et al., 2016) and 3 m on Cape Mamontov Klyk (Schirrmeister et al., 2008). Both coastlines feature above-average erosion rates that locally may reach 39.4 and 21 m yr^{-1} during exceptional years on Muostakh Island and Cape Mamontov Klyk, respectively (Günther et al., 2013, 2015). On the Laptev Sea coast, Günther et al. (2013) studied three ice-rich coastal zones (Oyogos Yar, Buor Khaya,

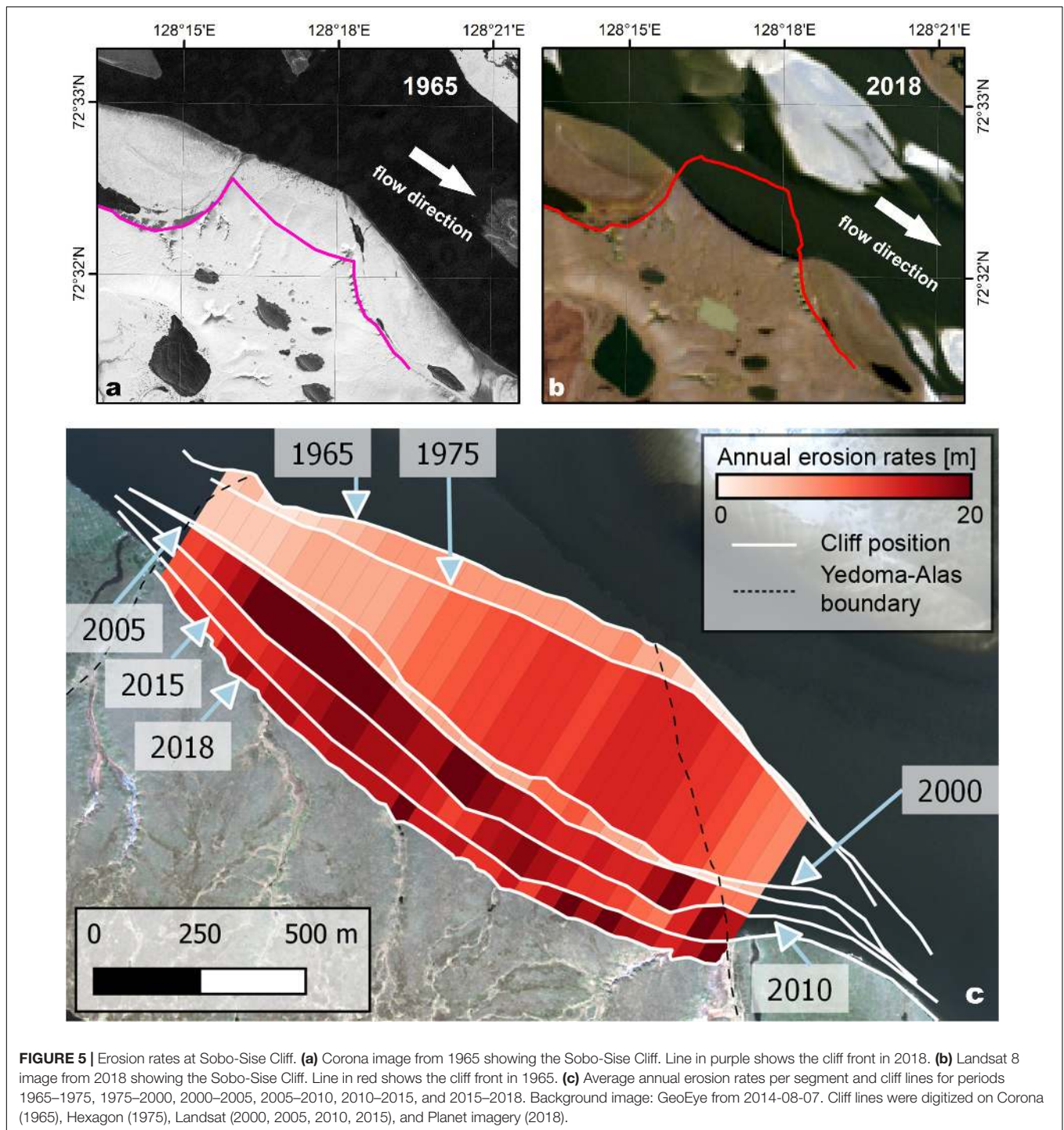


FIGURE 5 | Erosion rates at Sobo-Sise Cliff. **(a)** Corona image from 1965 showing the Sobo-Sise Cliff. Line in purple shows the cliff front in 2018. **(b)** Landsat 8 image from 2018 showing the Sobo-Sise Cliff. Line in red shows the cliff front in 1965. **(c)** Average annual erosion rates per segment and cliff lines for periods 1965–1975, 1975–2000, 2000–2005, 2005–2010, 2010–2015, and 2015–2018. Background image: GeoEye from 2014-08-07. Cliff lines were digitized on Corona (1965), Hexagon (1975), Landsat (2000, 2005, 2010, 2015), and Planet imagery (2018).

Mamontov Klyk) dominated by yedoma and alas deposits and found mean erosion rates of $0.9\text{--}2.9\text{ m yr}^{-1}$ for the yedoma dominated areas (Figure 6). Another example is the very ice-rich Alaskan Beaufort Sea coast between Drew Point and Cape Halkett with a mean erosion rate of 13.6 m yr^{-1} for the 2002–2007 period (Jones et al., 2009). Looking at a 9 km coastal segment at Drew Point only, average rates for the period 2007–2016 were even higher with 17.2 m yr^{-1} (Jones et al., 2018).

Lower rates were reported from the Yukon coast with mean erosion rates of 1.3 m yr^{-1} (Irrgang et al., 2018), Herschel Island with 0.7 m yr^{-1} (Obu et al., 2016), the Chukchi Sea coast on the northern Seward Peninsula of Alaska with 1.3 m yr^{-1} (Farquharson et al., 2018), or Barter Island (Alaskan Beaufort Sea) with 1.3 m yr^{-1} (Gibbs et al., 2019), keeping in mind that these values are long term averages of areas, where local and temporal maxima can be substantial higher (e.g., 8.1 m yr^{-1} for

TABLE 2 | Erosion rates of Sobo-Sise Cliff from 1965 to 2018.

Period	ER _{absolute} (m)	Years	ER _{mean} (m yr ⁻¹)	Scaled factor vs. long term ^a	ER _{max} (m yr ⁻¹)	Eroded total area (km ²)	Eroded yedoma area (%) ^b	Eroded total volume (10 ⁶ m ³) ^b	Eroded yedoma volume (10 ⁶ m ³) ^b	Eroded yedoma volume per year (10 ⁶ m ³ yr ⁻¹) ^b
1965–2018	321.8–679.2	53	9.13	1.00	12.83	0.878	n/a	n/a	n/a	n/a
1965–1975	6.4–80.6	10	4.76	0.52	8.24	0.082	n/a	n/a	n/a	n/a
1975–2000	110.0–371.5	25	9.24	1.01	14.84	0.406	n/a	n/a	n/a	n/a
2000–2005	11.8–58.2	5	5.73	0.63	11.82	0.046	0.043	0.972	0.961	0.192
2005–2010	49.5–116.1	5	15.12	1.66	22.26	0.145	0.141	3.165	3.150	0.630
2010–2015	47.6–96.0	5	13.79	1.51	19.32	0.121	0.118	2.648	2.636	0.527
2015–2018	34.3–63.7	3	15.67	1.72	21.21	0.078	0.078	1.743	1.743	0.581

All numbers are including ice wedges as well as segregated and pore ice, which means that 88% of all the yedoma volumes given here consist of ice. Eroded volumes are given in millions (10⁶) of m³. Eroded yedoma volumes excluding ice wedge volume are given in **Table 3**. ^aThis factor shows the ER_{mean} for the different time periods in comparison to the ER_{mean} 1965–2018. A value > 1.00 indicates that the erosion is higher for this particular period compared with the long term average. ^bWe did not estimate the volume loss for the periods 1965–2018, 1965–1975, and 1975–2000, because we have no information on the height of the yedoma cliff prior to 2000.

TABLE 3 | Sediment (Sed_{flux}), soil organic carbon (SOC_{flux}), dissolved organic carbon (DOC_{flux}) and nitrogen (N_{flux}) erosion rates.

Period	Total eroded sediment ^a (10 ⁶ m ³)	Total Sed _{flux} ^b (10 ⁶ kg)	Mean annual Sed _{flux} ^b (10 ⁶ kg yr ⁻¹)	Total SOC _{flux} (10 ⁶ kg)	Mean annual SOC _{flux} (10 ⁶ kg yr ⁻¹)	Total DOC _{flux} (10 ⁶ kg)	Mean annual DOC _{flux} (10 ³ kg yr ⁻¹)	Total eroded C (10 ⁶ kg)	Mean annual eroded C (10 ⁶ kg yr ⁻¹)	Total N _{flux} (10 ⁶ kg)	Mean annual N _{flux} (10 ⁶ kg yr ⁻¹)	Average annual C flux (SOC + DOC) per meter coast (kg C m ⁻¹ yr ⁻¹) ^c
2000–2018	2.87	2081.0	115.6	75.1 ± 48.2	4.2 ± 2.7	0.60 ± 0.33	33.5 ± 18.1	75.7 ± 48.2	4.2 ± 2.7	5.9 ± 3.2	0.3 ± 0.2	2631
2000–2005	0.32	235.5	47.1	8.5 ± 5.5	1.7 ± 1.1	0.07 ± 0.04	13.6 ± 7.4	8.6 ± 5.5	1.7 ± 1.1	0.7 ± 0.4	0.1 ± 0.1	1099
2005–2010	1.07	772.2	154.4	27.9 ± 17.9	5.6 ± 3.6	0.22 ± 0.12	44.8 ± 24.1	28.1 ± 17.9	5.6 ± 3.6	2.2 ± 1.2	0.4 ± 0.2	3546
2010–2015	0.89	646.2	129.2	23.3 ± 15.0	4.7 ± 3.0	0.19 ± 0.10	37.5 ± 20.2	23.5 ± 15.0	4.7 ± 3.0	1.8 ± 1.0	0.4 ± 0.2	2904
2015–2018	0.59	427.2	142.4	15.4 ± 9.9	5.1 ± 3.3	0.12 ± 0.07	41.3 ± 22.3	15.5 ± 9.9	5.2 ± 3.3	1.2 ± 0.6	0.4 ± 0.2	3131

All values are corrected for ice wedge volume. Total eroded sediment is given in millions (10⁶) of m³ and all other masses are given in 10⁶ kg, except mean annual DOC_{flux} is given in 10³ kg yr⁻¹. DOC fluxes include segregated and pore ice as well as ice wedge DOC flux estimations. Standard deviations are given for the eroded C and N, which are based on the variance in total organic C, dissolved organic C and total nitrogen measurements. ^aOnly accounting for the yedoma sediment. Ice wedge volume is excluded, however, segregated and pore ice is still included. The total eroded volume is presented in **Table 2**. ^bAssuming a mean dry bulk density of 0.7 g cm⁻³. This is the dry mass of all sediments (including peat) eroded from Sobo-Sise Cliff, excluding all forms of ice. ^cAn average cliff length was included in the calculation for each time period.

central parts of a bluff on Barter Island (Gibbs et al., 2019), or 22 m yr^{-1} for active slumps from 2012 to 2013 on Herschel Island (Obu et al., 2017).

However, when comparing Sobo-Sise Cliff to other locations at the coast, it is important to point out that different factors affect fluvial thermo-erosion compared to coastal erosion. The moving river water, the annual ice breakup, the spring flood and the transport of warm river water from further south certainly favor erosion of cliffs along Arctic rivers. These factors lead to a different erosion setting when compared to coastal stretches, which are affected by thermo-denudation and thermo-abrasion (e.g., Are, 1983; Overduin et al., 2014; Günther et al., 2015). In particular, the moving river water of the Sardakhskaya channel leads to fast removal of eroded debris at the cliff base, allowing continuous erosion at the base by the warm river water in summer. This prevents the flattening of the Sobo-Sise Cliff front and is likely an important factor for the continuous high erosion rates.

Annual total C fluxes from rivers into the Laptev Sea have been estimated to $6,800 \times 10^6 \text{ kg C yr}^{-1}$ (Rachold et al., 2004). Despite its distinct height, very high ground ice content, and its fast fluvial erosion, by its local nature the Sobo-Sise Cliff contributes only a small portion ($5.2 \times 10^6 \text{ kg C yr}^{-1}$, 2015–2018) of this flux. Given the 1,660 m long river shore segment studied, a shoreline-length normalized C flux yields an average of $3,131 \text{ kg C m}^{-1} \text{ yr}^{-1}$ (2015–2018), which is significantly higher than C fluxes observed in other studies (Figure 6). Compared to the yedoma coasts studied by Günther et al. (2013), Sobo-Sise Cliff is distinct by its high annual erosion rate, which leads to an overall high sediment and C loss. However, the common practice of aggregating local observations over larger coastal stretches dilutes the substantial contribution of erosional hot spots such as the Sobo-Sise Cliff to overall C fluxes. Because comparable high erosion rates sustained over several decades as we report for the Sobo-Sise Cliff are not known from other locations, our study therefore contributes to a better picture of the upper end in organic matter flux total variability when compared with the mean values summarized in Figure 6.

Sobo-Sise Cliff Will Remain a Strong Source for Mobilized C and N

Future Sobo-Sise Cliff erosion rates are difficult to predict because they depend on dynamic factors such as ice cover thickness, timing of river-ice breakup, river discharge, air and water temperatures (Walker et al., 1987). Particularly the course of the main river bed of the Sardakhskaya channel is an important factor. In case the main channel might shift slightly or a sand bank forms in front of the Sobo-Sise Cliff, riverbank erosion would suddenly decrease due to accumulation of sediment load and eroded debris, reducing niche formation and subsequent erosion at the cliff base (Grigoriev, 2007). Although a major sudden shift of the Sardakhskaya channel in near future seems unlikely, the Sobo-Sise Cliff will not form an infinite bulge on the outer side of a river bend and only retreat at the pace of the neighboring river shoreline.

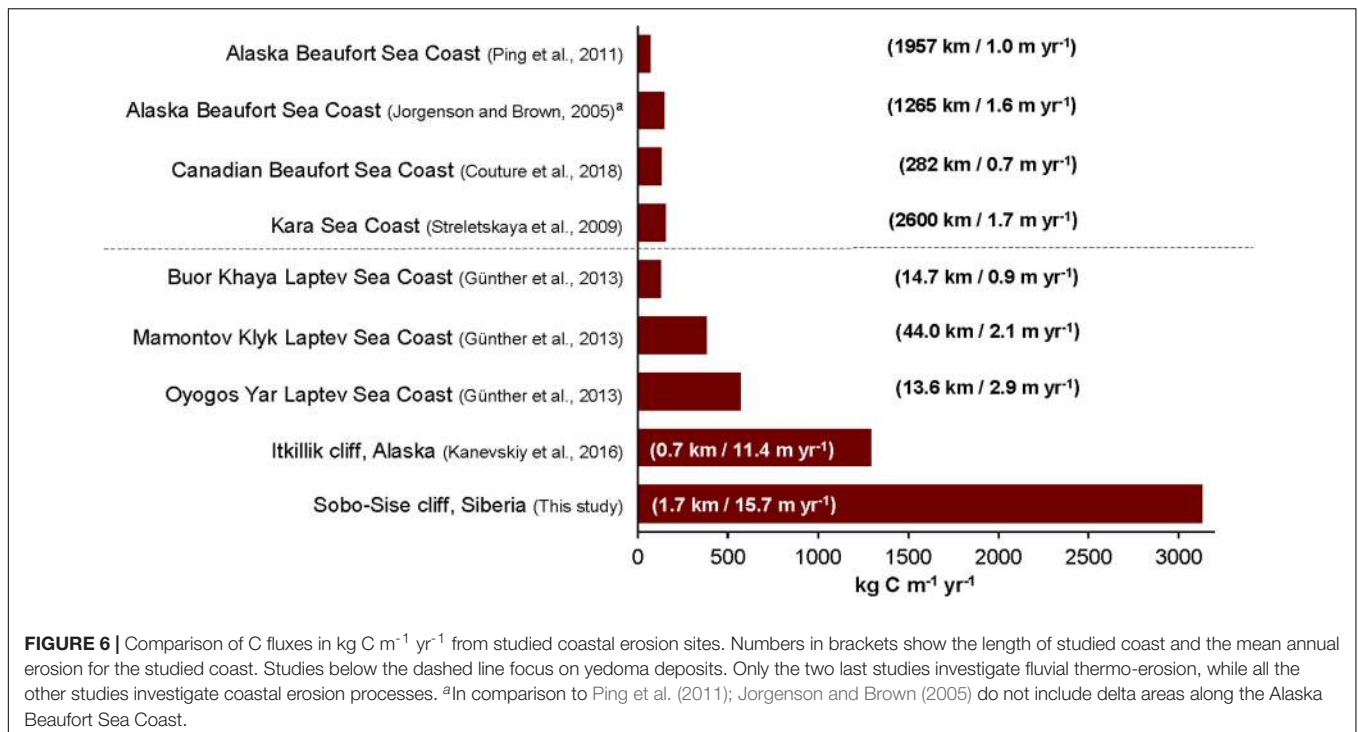
In our remote sensing time series for 1965–2018 we see an increase erosion rates at the Sobo-Sise Cliff (Table 2). Besides the increase in erosion rate, the length of the cliff itself increased by 58% from an initial length of 1,050 m in 1965 to a length of 1,660 m in 2018. The increase in length is due to continuous erosion from the initial convex shape to the concave shape of the Sobo-Sise Cliff and the progressing erosion into the crescent forms of eroded alases. With progressing erosion, the Sobo-Sise Cliff will likely further increase in length (Figure 5) and therefore would contribute an accordingly higher amount of sediment and organic C flux into the Lena River.

Our results show some variation in erosion rates of different time periods. Whereas the time period 2005–2010 is characterized by high mean erosion rates of 15.1 m yr^{-1} , the precedent period (2000–2005) showed almost three times smaller mean erosion rates of 5.7 m yr^{-1} . The reasons for this inconsistency in erosion rates are difficult to determine but likely are related to variable environmental conditions. Air temperature data (NOAA, Menne et al., 2012) from Tiksi show that from the late 1990s to the early 2000s mean monthly air temperatures were mostly below average and since then no extensive period with below-average mean monthly temperatures occurred anymore (see Supplementary Figure S5). However, air temperature is most likely not the only factor affecting Sobo-Sise Cliff erosion. Data on the Lena water level, the river discharge in the Sardakhskaya channel, duration and thickness of Lena River ice cover, or precipitation would be beneficial in determining the environmental conditions affecting Sobo-Sise Cliff erosion. Equally important, if not the major factor favoring high erosion rates, not only at the Sobo-Sise Cliff but also on outer coastal locations, is the position of the lower yedoma Ice Complex base that may have substantial variation in space due to an uneven paleo-relief prior to yedoma deposition.

We did not include erosion rates and volumes below the water surface, which makes our estimates of sediment, C, and N fluxes from the Sobo-Sise Cliff conservative. The ice-rich yedoma deposits in the eastern part of the Lena Delta are known to extend below the water surface (Pavlova and Dorozhkina, 2000) similar to the yedoma on nearby Bykovsky Peninsula (e.g., Schirrmeister et al., 2002). According to the bathymetrical survey in front of Sobo-Sise Cliff (Figure 4) mean water depth is 10.3 m. Assuming the below water deposits have a similar composition like the above-water deposits, the total volume as well as the C and N loss would increase by 46% compared to our estimates (2015–2018). However, we do not have sufficient data to characterize the deposits below the water surface nor the hydrological and erosional dynamics in the channel that would also include winter river flow, therefore this estimate remains highly speculative.

Exceptional DOC Fluxes From Ice-Rich Sobo-Sise Cliff

Up to $41.3 \times 10^3 \text{ kg DOC}$ is eroded annually (2015–2018) into the Sardakhskaya channel. Considering the length of Sobo-Sise Cliff of only 1,660 m this is a substantial amount of potentially bioavailable C, which is released into the riverine ecosystem. For comparison, Tanski et al. (2016) estimated an annual flux



of 54.9×10^3 kg yr⁻¹ from the 306 km long Yukon coast. The reasons for the very high DOC flux at Sobo-Sise Cliff are the high segregated and pore ice volume (65.7%), the dimensions of the cliff (up to 27.7 m high) and the high erosion rates. Nevertheless, the main factor is the high DOC concentration within the segregated and pore ice. Maximum DOC concentrations up to 786 mg L⁻¹ and a DOC to SOC ratio of 1:125 are among the highest observed in the Arctic. Mann et al. (2015) and Vonk et al. (2013a) reported values of 131 mg L⁻¹ and 196 mg L⁻¹ for yedoma thaw streams, respectively. Schirrmeister et al. (2017) found similar high DOC concentrations in yedoma Ice Complex outcrops and drill cores of the Buor Khaya Peninsula in the east of the Lena Delta.

Although the DOC fluxes might still be small compared to the SOC fluxes from OC-rich sediments, DOC from permafrost is chemically labile (Dou et al., 2008; Vonk et al., 2013a,b) and may directly enter local food webs (bioavailability). This means that DOC can become quickly mineralized by microbial communities and photochemical reactions (Battin et al., 2008; Vonk et al., 2013a,b; Cory et al., 2014) and returned to the atmosphere when released due to permafrost degradation (Schuur et al., 2009, 2011; Tanski et al., 2019).

Fate of Eroded Organic Matter and Implications for the Ecosystem

The high amount of organic matter eroded from the Sobo-Sise Cliff enters the Lena River and is transported downstream to the outer parts of the delta or likely to the near-shore zone of the Laptev Sea that is only about 40 km to the east. Previous studies indicated that the Lena River has an organic C flux of 0.9 Tg C yr⁻¹ coming from permafrost and peat deposits (Wild et al.,

2019). However, Wild et al. (2019) measured C concentration approximately 800 km upstream from Sobo-Sise Cliff. Therefore, it is difficult to infer on the C concentration and export at Sobo-Sise Cliff from these values. Nevertheless, this indicates that high amounts of permafrost C are transported toward the Laptev Sea.

The consequences of permafrost erosion and the subsequent organic matter input from the Lena River catchment into the Laptev Sea are not fully understood yet. Increased terrestrial C fluxes to the Arctic Ocean might affect food webs in the near-shore zone (e.g., Dunton et al., 2006, 2012; Casper et al., 2015) or lead to an increased acidification (Semiletov et al., 2016). The latter is seen in waters of the East Siberian Arctic Shelf that become more acidic due to the high CO₂ concentrations in river water input, which is affected by degradation of terrestrial organic matter. In fact, the study by Semiletov et al. (2016) stated that 57% of the terrestrial organic C in the East Siberian Shelf originates from old Pleistocene age permafrost C like the deposits from the Sobo-Sise Cliff.

However, Bröder et al. (2019) reported on slow C degradation rates of terrestrial C in the top layer of seafloor sediments, of which a majority was considered even resistant to degradation. In addition, a large amount of the permafrost C might become re-buried in marine sediments of the Laptev and East Siberian Sea shelf (Vonk et al., 2012; Vonk and Gustafsson, 2013; Bröder et al., 2019). Nevertheless, terrestrial organic C on the Laptev Sea shelf is exposed for degradation for millennia so that these areas might become a C source for ocean acidification and atmospheric export on long timescales (Bröder et al., 2018).

Also, Tanski et al. (2019) showed in an incubation study that high amounts of CO₂ are produced from thawing permafrost C in seawater, potentially leading to high CO₂ production along

permafrost coasts and near-shore zones. Another aspect raised by Sánchez-García et al. (2014) is the potential correlation of decay intensity of the organic matter in relation to active erosion of yedoma deposits, where fast erosion rates and high moisture content during summer thaw seem to favor microbial activity. In another study, Winterfeld et al. (2018) suggest that fluvial export of old (Pleistocene) C from degrading permafrost was potentially an important process for high C mobilization at the Pleistocene-Holocene transition around 11.5 kyrs BP, leading to changes in atmospheric CO₂, and indicating the significance of eroding yedoma in the Arctic. Similar findings are provided by Tesi et al. (2016) and Martens et al. (2019), who report a much higher permafrost C transport to the Laptev and Chukchi Sea at the end of the last deglaciation, indicating high permafrost C mobilization during times of rapid climate warming.

DOC from late Pleistocene permafrost as an especially bioavailable form of C due to its fraction size and molecular composition could play a crucial role for microbial use and general food web impacts. For example, Vonk et al. (2013b) showed that yedoma ice wedge meltwater can increase the degradation of organic matter due to co-metabolizing effects. At the Duvanny Yar yedoma exposure on the Kolyma river, Vonk et al. (2013a) and Mann et al. (2014) found that DOC originating from ancient yedoma deposits is highly bioavailable and therefore important to consider in a potential permafrost C feedback. The origin of organic matter and the sequestration process into ground ice plays an important role in the concentration and bioavailability of DOC (Fritz et al., 2015). Sequestration of DOC into ground ice is a complex process that is dependent on water source, freezing process, organic matter quality and even the inorganic geochemical composition of the ambient water to form ground ice. Tundra-steppe vegetation, reduced organic matter degradation rates and syngenetic permafrost aggradation were characteristic for the cold stages of the last ice age. Late Pleistocene yedoma deposits and associated ground ice are therefore the ideal reservoir for the fresh and most bioavailable DOC components, where the chemical character is preserved because of rapid incorporation into permafrost.

In summary, while we quantified the amount to C and N mobilized from the Sobo-Sise Cliff the actual fate of the eroded terrestrial material is not yet clear and further studies are needed to investigate the state (dissolved, particulate) and the lability of the C and N released from Sobo-Sise Cliff and similar shores in the Lena River and Delta. The Sobo-Sise Cliff (this study) and other permafrost-shores characterized by rapid erosion in the Lena River (Costard et al., 2003, 2007; Dupeyrat et al., 2018; Stettner et al., 2018) contribute to an enhanced fluvial transport of terrestrial organic matter and nitrogen to the Arctic Ocean. McClelland et al. (2016) estimated that the Lena River exports 0.8 Tg particulate organic C and 0.1 Tg particulate N per year into the Arctic Ocean. In addition, Holmes et al. (2012) reported of an annual DOC and total dissolved N export of 5.7 Tg and 0.2 Tg, respectively. Combining particulate and dissolved organic C fluxes results in a total export of 6.5 Tg C yr⁻¹ and 0.3 Tg N yr⁻¹ by the Lena River. Understanding the sources of this flux and the processes behind its mobilization from permafrost is key to estimate current and project future impacts of changing

river runoff regimes on biogeochemical cycling processes in permafrost rivers and the Arctic Ocean.

CONCLUSION

The 1,660 m long and 27.7 m high Sobo-Sise Cliff is one of the fastest eroding permafrost features in the Arctic, releasing a significant amount of C and N on a local scale. Erosion rates of up to 22 m per year and an average loss of 5.2×10^6 kg C (SOC and DOC) and 0.4×10^6 kg N per year (2015–2018) highlight the magnitude of biogeochemical fluxes from this individual erosional feature in the Lena River Delta. To study rapid permafrost shore erosion processes and resulting sediment fluxes at an extraordinary location such as the Sobo-Sise Cliff is extremely helpful to gain a better understanding of current and future impacts of permafrost thaw, interdependencies of permafrost erosion with runoff changes, and fluvial and marine biogeochemistry. In particular regarding the latter, our observations of current erosion dynamics at the Sobo-Sise Cliff provide an outlook of C and N fluxes from permafrost for the anticipated future accelerated coastal erosion in the course of global sea level rise, when it reaches the lower base of Ice Complex deposits in the yedoma upland along the northeast Siberian coastline.

DATA AVAILABILITY STATEMENT

The data for this study has been archived on the PANGAEA (www.pangaea.de) data set repository. Erosion rates and digitized cliff fronts are available on doi: 10.1594/PANGAEA.918507 (Nitze et al., 2020), and the geochemical results from Sobo-Sise Cliff samples are available on doi: 10.1594/PANGAEA.919470 (Wetterich et al., 2020).

AUTHOR CONTRIBUTIONS

MFu and IN contributed equally to this study. MFu, IN, JS, FG, and GG designed the study. GG, JS, and SW acquired funding. AK, SW, and MFr collected and analyzed sediment data. MFr collected and analyzed the samples for DOC. GM and SW made the bathymetric survey. IN, FG, and GG processed the remote sensing imagery. IN calculated erosion rates and cliff heights. MFu and JS calculated C and N stocks. All authors participated in field investigations. MFu wrote the original draft of the manuscript. All authors contributed in writing, reviewing, and editing the manuscript.

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