

Relative contributions of crust and mantle to the origin of the Bijli Rhyolite in a palaeoproterozoic bimodal volcanic sequence (*Dongargarh Group*), central India

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New mineralogical, bulk chemical and oxygen isotope data on the Palaeoproterozoic Bijli Rhyolite, the basal unit of a bimodal volcanic sequence (*Dongargarh Group*) in central India, and one of the most voluminous silicic volcanic expressions in the Indian Shield, are presented. The Bijli Rhyolite can be recognized as a poorly sorted pyroclastic deposit, and comprises of phenocrystic K-feldspar + albite \pm anorthoclase set in fine-grained micro-fragmental matrix of quartz-feldspar-sericite-chlorite-iron-oxide \pm calcite. The rocks are largely metaluminous with high SiO₂, Na₂O + K₂O, Fe/Mg, Ga/Al, Zr, Ta, Sn, Y, REE and low CaO, Ba, Sr contents; the composition points to an 'A-type granite' melt. The rocks show negative Cs-, Sr-, Eu- and Ti- anomalies with incompatible element concentrations 2–3 times more than the upper continental crust (UCC). LREE is high (La/Yb \sim 20) and HREE 20–30 times chondritic. $\delta^{18}\text{O}_{\text{whole-rock}}$ varies between 4.4 and 7.8‰ (mean $5.87 \pm 1.26\%$).

The Bijli melt is neither formed by fractionation of a basaltic magma, nor does it represent a fractionated crustal melt. It is shown that the mantle-derived high temperature basaltic komatiitic melts/high Mg basalts triggered crustal melting, and interacted predominantly with deep crust compositionally similar to the Average Archaean Granulite (AAG), and a shallower crustal component with low CaO and Al₂O₃ to give rise to the hybrid Bijli melts. Geochemical mass balance suggests that \sim 30% partial melting of AAG under anhydrous condition, instead of the upper continental crust (UCC) including the Amgaon granitoid gneiss reported from the area, better matches the trace element concentrations in the rocks. The similar Ta/Th of the rhyolites (0.060) and average granulite (0.065) vs. UCC (0.13) also support a deep crustal protolith. Variable contributions of crust and mantle, and action of hydrothermal fluid are attributed for the spread in $\delta^{18}\text{O}_{\text{whole-rock}}$ values. The fast eruption of high temperature (\sim 900°C) rhyolitic melts suggests a rapid drop in pressure of melting related to decompression in an extensional setting.

1. Introduction

The petrogenesis of large volumes of rhyolites in association with basalts, but without rocks of intermediate composition is of considerable interest in igneous petrology for decades (e.g., Yoder 1973; Hildreth 1981; Cox 1988; Peccerillo *et al* 2003).

The unlikelihood of a wholly basaltic source for many voluminous potassic rhyolites is recognized time and again (e.g., Coleman and Walker 1992; Annen and Sparks 2002). The emerging idea based on isotopic, theoretical and experimental studies considers the important role of crust and/or juvenile contribution from the mantle in the generation

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of rhyolite to high-silica rhyolitic ($\text{SiO}_2 \geq 74$ wt%) melts (e.g., Landenberger and Collins 1996; Patiño Douce 1999). The discovery of compositional similarities of the older Yellowstone silicic tuff to the interstitial glass of the Columbia River Plateau basalts (Nash *et al* 1996) is significant in this context. Evidence of chemical interaction between mafic magma and lower crustal material is well preserved in lower-crustal xenoliths (Downes 1993), and in exposed sections of the lower crust (e.g., the Ivrea Verbano zone, Sinigoi *et al* 1991). Yet, some workers believe that the traditional idea of basaltic magma fractionation is an effective mechanism for the origin of silicic melts (e.g., Garland *et al* 1995; Ayalew *et al* 2002). The other competing mechanism considers extensive fractional crystallization of the mantle-derived basaltic melt accompanied by assimilation of crustal components (e.g., Musselwhite *et al* 1989; Peccerillo *et al* 2003). In short, the nature of the involvement of both crust and mantle is where most controversy lies. The question is important, because if the only mantle contribution is heat, then the continental crust is largely recycled, but if mantle material is supplied, then it involves crustal growth. The issue perhaps assumes further significance if the silicic magmatism took place at or near 2.5 Ga, known to be a significant continental crust-building period. The petrogenetic study of the Bijli Rhyolite in the central Indian Craton, one of the most voluminous expressions of silicic volcanism in the Indian Precambrian Shield, is thus of topical relevance and would help in understanding the origin of similar rocks elsewhere.

This paper presents:

- new mineralogical and bulk chemical (major and trace elements) data of the Bijli Rhyolite,
- evaluation of its geochemical characteristics and
- important constraints for the genesis of such high silica melts including relative roles of crust and mantle and the nature of the protoliths.

Our recently published oxygen isotope data (Sensarma *et al* 2004) are also used for the purpose. The regional significance of this silicic volcanic activity is discussed at the end of the paper.

2. Geological setting

The Precambrian terrain in the central Indian Shield is divided into Northern Crustal Province (NCP) and the Southern Crustal Province (SCP) separated by a ENE-WSW trending belt referred to as the Central Indian Tectonic Zone (CITZ) (Roy *et al* 2000 and references therein). The SCP

comprises a collage of N-S trending low-grade volcanic-sedimentary litho-stratigraphic units including the Dongargarh volcanic-sedimentary succession (*Dongargarh Group*) (figure 1). This group of rocks occurs in a linear belt in parts of Bhandara district of Maharashtra and Rajnandgaon district of Chattisgarh, and is southerly extended in the Bastar district of Chattisgarh. The Palaeoproterozoic (2.2–2.5 Ga) Dongargarh Group is weakly folded into a synclinal structure (Sitagota Syncline), and is intruded by a post-tectonic granite pluton (*Dongargarh Granite*) (Sarkar 1957–58; Sensarma 2001; Sensarma and Mukhopadhyay 2003) (table 1). Both felsic and mafic volcanic rocks are abundantly present in a bimodal distribution within the Dongargarh Group as demonstrated in the SiO_2 vs. Mg-number plot (figure 2), and are broadly contemporaneous to each other (Sensarma 2001). The silicic volcanic unit (Bijli Rhyolite) occurs at the base of the Dongargarh Group, and represents the first major magmatic activity in the belt. The overlying mafic volcanic rocks (Munro-type basaltic komatiite/high Mg basalts – tholeiite assemblages) of the Pitepani and Sitagota volcanic units are products of pulsatory eruptions of mantle plume origin (Sensarma 2001; Sensarma and Mukhopadhyay 2003). The Dongargarh Group of rocks is affected by low-grade green-schist facies metamorphism.

The Bijli Rhyolite is a ~ 4.5 km thick and voluminous (> 8000 cubic km) particle-enriched fragmental deposit of pyroclastic origin, and principally exposed on the western limb of the Sitagota Syncline. It comprises several pyroclastic facies including ignimbrite, rheo-ignimbrite, breccia and breccio-conglomerate deposits (Sensarma 2001; Sensarma and Mukhopadhyay 2003). The rocks have variable concentration of crystals, vitric particles (shards, collapsed pumice, glassy fragments) and lithic fragments. The rocks show a N-S trending foliation with steep dip ($\sim 60^\circ$) parallel to the regional schistosity. The primary flow layers are recognized on the weathered surface at a few places as rib like structures that cut across the regional foliation at a low angle (15° – 20°). The Bijli Rhyolite yielded Rb-Sr whole rock isochron ages of 2180 ± 25 Ma (ISR = 0.70571 ± 0.0015) (Sarkar *et al* 1981) and 2503 ± 35 Ma (ISR = 0.7035 ± 0.0017) (Krishnamurthy *et al* 1990). Recent Rb-Sr whole rock isochron ages determined by Dr. K Gopalan and his associates at N.G.R.I., Hyderabad, (personal communication) from the rhyolite samples collected by the present authors suggest an age of 2089 ± 29 Ma. The Rb-Sr whole rock isochron ages for the intrusive Dongargarh Granite is 2270 ± 90 Ma (Sarkar *et al* 1981) and 2465 ± 22 Ma (ISR = 0.7010 ± 0.0042) (Krishnamurthy *et al*

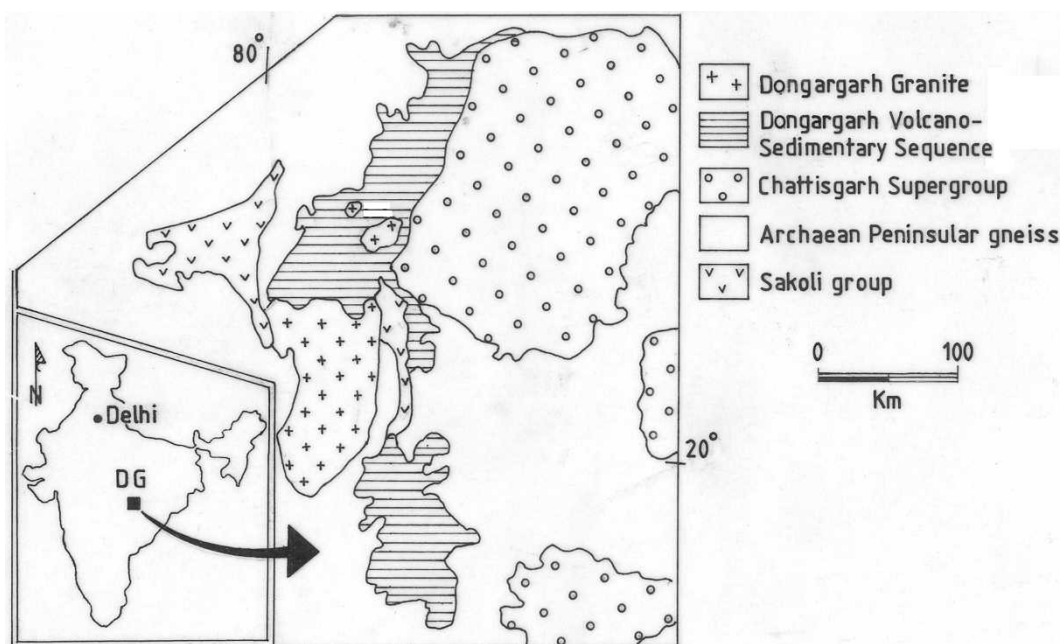


Figure 1. Location of the Dongargarh volcano-sedimentary sequence (Dongargarh Group) in the central Indian Craton. Map of India is shown in the inset.

Table 1. *Litho-stratigraphic succession in the Dongargarh Belt, central India (Sensarma 2001; Sensarma and Mukhopadhyay 2003). The principal lithology of the units is shown within corresponding paranthesis.*

Bijepar Formation (low dipping arkosic sediments)	
————— Unconformity —————	
Dongargarh Granite (granite, granophyre, aplite)	
————— Intrusive contact —————	
Mangikhuta Volcanics (andesitic lava)	} Dongargarh Group
Karutola Formation (quartz arenite ± phyllite)	
Sitagota Volcanics (basaltic komatiite/high-Mg basalts-basalts)	
Chandsuraj Formation (pebbly volcanogenic wacke)	
Pitepani Volcanic (andesitic basalts-basalts)	
Bijli Rhyolite (silicic pyroclastic rock) (with Halbitola Rhyolitic breccio-conglomerate Member)	
Basement ? ?	

1990). The geochemical investigations of the Bijli Rhyolite include those by Sarkar *et al* (1994); Neogi *et al* (1996); Gangopadhyay and Roy (1997); Sensarma (2001). The petrogenetic aspects are discussed in Divakara Rao *et al* (2000).

The Munro-type komatiitic rocks are believed to occur at or near the Archaean-Proterozoic transition (Arndt 1994). Similarly, the presence of near contemporaneous Munro-type basaltic komatiitic melts to the Bijli Rhyolite in the area, and the Rb/Sr whole rock age of the Bijli Rhyolite suggest a palaeoproterozoic age (~ 2.5 Ga) for the Bijli Rhyolite.

3. Analytical procedures and data base

3.1 Analytical procedures

Aphanitic to ultrafine cherty samples with/without sparse crystals and free from effects of alteration, amygdules, veins and fractures have been given preference for chemical analysis, as these samples are likely to represent approximate bulk liquid composition. Most of the samples so chosen are of welded to extremely welded ignimbrite variety. The samples were collected near Salekasa (21°10': 80°29'), Makkatola (21°15': 86°26½'), and further south near Deori in the Survey of India sheet no.

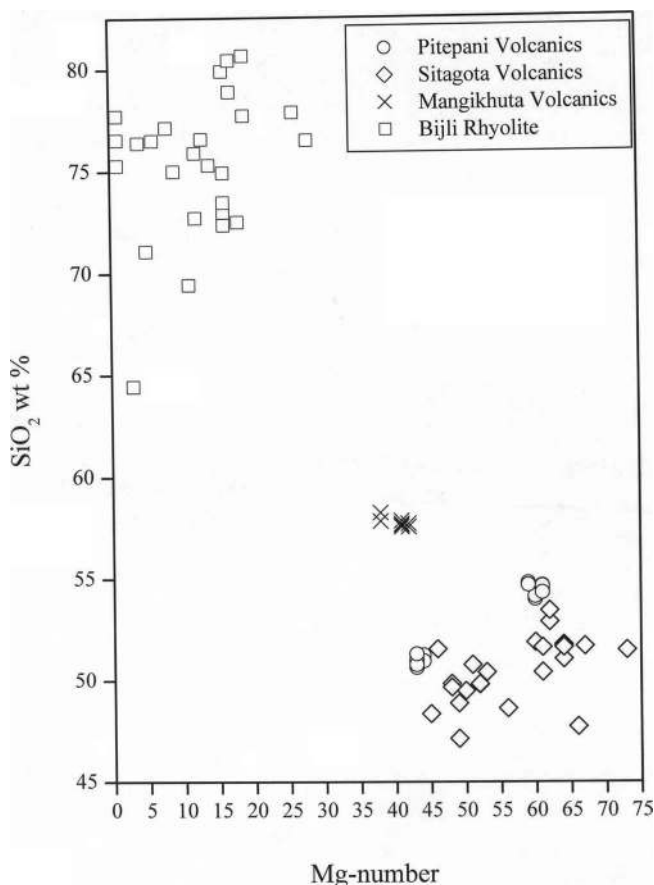


Figure 2. SiO_2 -Mg number plot showing bimodal character of the volcanic units in the Dongargarh Group.

64C/8. The analytical work were mostly carried out at the Institut für Mineralogie und Geochemie, Universität Zu Köln, Germany and the Activation Laboratory Limited, USA. Some additional analyses were done at the School of Environmental Sciences, Jawharlal Nehru University (JNU), New Delhi and Geological Survey of India (GSI), Calcutta.

At the Universität Zu Köln, major and selected trace elements concentrations (Rb, Sr, Zr, Ba, Y, Ni, Cr, Co, Sc, V, Zn) were determined by a Philips PW 2400 XRF machine using fused glass discs. For trace elements, INAA method was used. The powdered samples were irradiated in the carousel of the TRIGA reactor of the Universität Mainz, and subsequently counted on large Ge-detectors at the Institut für Mineralogie und Geochemie, Universität Zu Köln. Based on extensive comparison with standard rocks, the accuracy of the XRF data is estimated to be below 3%. For INAA, accuracies are indicated in the respective table. Details of analytical procedures followed, and the inter-laboratory comparison on accuracy of the XRF and INAA methods applied here are available in Sensarma *et al* (2002).

Some major and trace element data were generated at the Activation Laboratory Limited, USA,

by XRF and ICP-MS methods using international standards, through the courtesy of Professor P H Reitan of the University at Buffalo. At JNU, major and trace element analyses were performed on a Labtam ICP-AES using international and in house standards. Silica was analyzed by colorimeter following the method of Shapiro and Brannock (1962). REE determination was done by column separation, as described in detail by Giritharan and Rajamani (1998). At GSI, major element concentrations of few samples were determined by XRF method with fused pallets using both in house and international standards. Data generated in these laboratories are in good agreement within the limits of uncertainty of the methods (10–15%).

Electron microprobe analyses (EPMA) of feldspar were performed with a JEOL electron microprobe (JXA-8900) at the same institute in Köln. An acceleration voltage of 20 kV and a beam current of 20 nA were used. A set of natural minerals (for Si, Ca, Na, K) and synthetic oxides (for Mg, Fe, Al, Ti) were used for standardization.

The bulk analyses are supplemented with oxygen isotope analyses of whole rock and mineral separates (phenocrysts) of selected samples (Sensarma *et al* 2004). The analyses were done at the Universität Bonn, Germany on 8–10 aliquotes of whole rock powders using a modified (Clayton and Mayeda 1963) type extraction line. Purified F2 gas (Asprey 1976) was used to liberate oxygen, which was then converted to CO_2 and measured on SIRA-9 mass spectrometer (VG Instruments). For mineral analyses, quartz and feldspar phenocrysts were handpicked and cleaned in warm dilute HF to remove any glass, carbonate and other impurities. These analyses were also replicated.

3.2 Data base

EPMA data of feldspar is presented in table 2. The XRF and ICP data for major and selected trace elements are shown in table 3. Selected trace element data (ICP and INAA) are presented in table 4. Oxygen isotope data for both whole rock and phenocrysts (quartz and feldspar) are listed in table 5.

In order to develop a unified petrogenetic model for such a huge volume of magmatic deposit, we have assembled major and trace element data ($n = 85$) from the literature (e.g., Neogi *et al* 1996; Gangopadhyay and Roy 1997; Divakara Rao *et al* 2000; Sensarma *et al* 2002) representing different parts of the unit including the Kotri belt in Bastar district (e.g., Shivkumar *et al* 2003). The published data are in reasonable agreement with the new data presented here.

4. Petrography and mineralogy

Patchy alteration, devitrification, low-grade metamorphism and deformation have influenced the petrographic character of the rocks. Yet, the Bijli Rhyolite can be recognized to be ill-sorted pyroclastic deposit in which crystals, vitric clasts (pumice, shards, collapsed vitric fragments) and lithic particles occur in a fine-grained fragmental to microfragmental matrix of quartz-feldspar-sericite-chlorite-iron-oxide \pm calcite assemblage. Crystal-to-crystal-vitric tuff is the dominant petrographic variety. The crystals are quartz and feldspar, the latter being mostly K-feldspar. Framework quartz is anhedral to euhedral with frequent embayed margins. Clots of biotite, veinlets of quartz and calcite are occasionally present. Zircon, Fe-oxide particles are common accessories. Quartz-k-feldspar-plagioclase-orthopyroxene-clinopyroxene-ilmenite-apatite represents the normative mineralogy. Lithic fragments are mostly of cognate type, devitrified and altered. Vitroclastic texture is common, with development of eutaxitic fabric in the welded varieties. The petrographic details of the rocks are discussed elsewhere (Sensarma 2001; Sensarma and Mukhopadhyay (in preparation)).

More than 70% of the phenocrysts in the Bijli Rhyolite are feldspars. Microprobe analyses of feldspars carried out on a number of samples show feldspar phenocrysts to be mostly of alkali feldspar, often with albite-orthoclase solid solutions or intergrowths. The composition of feldspar is presented in table 2. K-feldspar ($\text{Or}_{100} - \text{Or}_{71.7}\text{Ab}_{28.2}$) and albite ($\text{Or}_{0.4}\text{Ab}_{99.4} - \text{Or}_{5.0}\text{Ab}_{94.3}$) constitute the main population, with subordinate anorthoclase ($\text{Or}_{40.8}\text{Ab}_{59.2} - \text{Or}_{14.7}\text{Ab}_{85.2}$). Abundance of K-rich feldspar in the rocks (table 2) suggests K-rich character for the bulk liquid. Potassic rim around albite is also present.

5. Geochemistry

5.1 Evaluation of post-crystallization alteration

Alteration, low-grade metamorphism, devitrification, deformation have probably affected the chemistry of the original rocks. It is, therefore, important to evaluate the extent to which the chemical data of the Bijli Rhyolite represent the primary chemistry of the melts. Values of LOI, considered to be indicators of aqueous alteration of volcanic rocks, though low ($< 2 \text{ wt}\%$) in the analyzed samples, grossly maintain a positive relation with similarly low Cs concentration (figure 3). This may be attributed to preferential Cs mobilization during interaction with aqueous fluids (e.g.,

Rudnick and Presper 1990). The relatively immobile refractory lithophile major elements like Al, Ti and to some extent Ca in the rocks, however, show systematic patterns in different variation diagrams (see later). K/Rb ratio (200–300) of the rocks is not far from typical igneous rock values (~ 250) (Rudnick *et al* 1985). K/Ba and Rb/Sr ratios show generally consistent pattern when plotted against silica (figures 4 and 5). Some rocks (e.g., Shivkumar *et al* 2003), however, have consistently higher Rb/Sr values by about 3 times from rest of the data irrespective of silica concentration, probably implying original ratios. Five samples have higher but inconsistent K/Ba values with respect to other samples signifying Ba mobility for these samples (figure 5). Th/U ratios in the rocks are not much variable. U generally shows greater mobility compared to Th. Th/U ratios in many samples (~ 4) (figure 5) are close to that CI-chondrites (3.6), and consequently do not indicate any preferred mobilization of U compared to Th. On the other hand, a set of samples (e.g., Shivkumar *et al* 2003) show comparatively lower Th/U (< 2.5) (figure 5), but with no significant scatter on Th-U plot. These suggest generally unaltered characters of the suite, though not completely free from mobility of selected trace elements (e.g., Cs, Ba) during action of aqueous fluids. In addition, primitive upper mantle (PUM)-normalized incompatible element and chondrite-normalized REE patterns are regular implying that their concentrations and ratios in the samples, except those of Cs, reflect original magma characteristics. Since high field strength elements (HFSE) like Ti, Th, Nb, Ta, Zr, Hf, Y and REE generally remain immobile during various secondary processes, our petrogenetic interpretations mostly rely on these elements.

5.2 Classification

On a total Alkali-Silica diagram (Le Bas *et al* 1986) (figure 6), the samples mostly plot in the rhyolite field and are generally subalkaline in character. A few samples from Kotri area in the far southern part of the Dongargarh belt (e.g., Shivkumar *et al* 2003), however, plot as trachyte to trachy-andesite. The rocks (see in table 3) predominantly have very high silica content ($\text{SiO}_2 \text{ wt}\% > 74 \text{ wt}\%$), and thus can be termed high-silica rhyolite (cf., Pankhurst and Rapela 1995). These are high-K rocks ($\text{K}_2\text{O} \sim 4-4.5 \text{ wt}\%$) with total alkalis ($\text{Na}_2\text{O} + \text{K}_2\text{O}$) up to 8–10 wt%, but reaching up to $\sim 12 \text{ wt}\%$ especially in some samples from the Kotri area. The rocks are generally metaluminous (Sensarma 2001; this work), though some samples show peraluminous chemistry (Gangopadhyay and Roy 1997; Divakara Rao *et al* 2000; Shivkumar *et al* 2003).

Table 2. *Microprobe data of feldspar.*

[wt%]	1	2	3	4	5	6	1*	2*	4*	6*	1***	2***	3***	4***	5***	6***	7***	8***	9***	10***	11***	
SiO ₂	68.34	64.64	65.49	65.25	65.21	64.7	68.14	68.11	67.13	65.46	63.01	63.82	63.36	64.45	64.66	63.99	69.16	68.53	67.73	66.22	68.20	
TiO ₂	-	-	-	-	-	0.02	0.09	-	0.05	0.14	0.05	-	-	-	0.16	-	0.05	-	0.05	-	-	
Al ₂ O ₃	18.66	17.50	17.94	17.52	18.18	18.10	19.67	19.02	19.69	19.64	18.89	18.02	18.67	17.78	18.14	17.86	19.07	18.95	19.25	19.91	19.66	
FeO	-	-	-	-	0.04	0.03	0.13	0.08	0.29	0.71	0.62	0.06	0.41	0.18	0.16	-	-	0.03	0.32	0.54	0.07	
MnO	-	-	-	-	0.03	0.01	0.00	-	-	-	-	0.01	0.01	-	-	0.01	0.02	-	-	-	0.02	
MgO	-	-	-	-	0.03	-	0.08	0.03	0.10	0.06	0.12	-	-	-	-	-	0.06	-	0.06	0.05	0.00	
CaO	0.27	-	-	-	0.63	0	0.04	0.03	0.02	-	-	0.02	0.01	-	-	0.01	0.05	0.08	0.07	0.05	0.06	
Na ₂ O	11.71	0.05	0.28	0.11	2.86	0.56	11.23	10.13	10.90	6.55	0.31	0.11	0.26	0.22	0.41	0.34	11.89	11.63	11.72	10.95	11.78	
K ₂ O	0.2	17.05	16.83	16.95	12.38	16.00	0.97	2.65	0.89	6.85	15.75	16.49	16.02	16.47	16.24	16.27	0.1	0.12	0.37	1.17	0.3	
Total	99.18	99.24	100.54	99.83	99.36	99.42	100.35	100.05	99.07	99.41	98.75	98.53	98.74	99.10	99.77	98.48	100.40	99.34	99.57	98.89	100.09	
Total cation : on the basis of 24 oxygen																						
Si	9.040	9.064	9.048	9.083	8.994	9.017	8.938	9.006	8.916	8.862	8.867	8.999	8.916	9.035	8.995	9.019	9.027	9.036	8.951	8.849	8.948	
Ti	-	-	-	-	-	0.002	0.009	-	0.005	0.015	0.006	-	-	-	0.017	-	0.005	-	0.005	-	-	
Al	2.909	2.892	2.922	2.875	2.955	2.973	3.041	2.965	3.082	3.135	3.133	2.995	3.097	2.937	2.974	2.967	2.934	2.945	2.998	3.137	3.041	
Fe	-	-	-	-	0.005	0.003	0.014	0.009	0.032	0.08	0.073	0.007	0.048	0.022	0.018	-	-	-	0.004	0.035	0.06	0.010
Mn	-	-	-	-	0.004	0.001	0.000	-	-	-	-	0.001	0.000	-	-	0.002	0.002	-	-	-	0.000	
Mg	-	-	-	-	0.007	-	0.015	0.006	0.020	0.012	0.025	-	-	-	-	-	0.012	-	0.011	0.01	0.001	
Ca	0.038	-	-	-	0.092	0.000	0.006	0.004	0.003	-	-	0.004	0.001	-	-	0.001	0.008	0.011	0.01	0.008	0.010	
Na	3.003	0.015	0.075	0.031	0.765	0.152	2.856	2.597	2.807	1.720	0.086	0.031	0.072	0.060	0.112	0.093	3.010	2.972	3.004	2.838	2.997	
K	0.033	3.051	2.966	3.011	2.179	2.844	0.162	0.447	0.151	1.184	2.828	2.966	2.877	2.945	2.882	2.926	0.017	0.019	0.062	0.199	0.050	
Total	15.023	15.022	15.011	15.000	15.001	14.992	15.041	15.034	15.016	15.008	15.018	15.003	15.011	14.999	14.998	15.008	15.015	14.987	15.076	15.101	15.057	

Table 2. (Continued)

	Central crystal**						Rim around the crystal**			
	1	2	3	4	5	6	1	2	3	4
SiO ₂	67.67	68.62	67.42	68.37	68.71	68.89	64.78	64.05	63.31	64.35
TiO ₂	-	-	-	-	-	-	-	0.07	0.07	0.11
Al ₂ O ₃	19.48	19.23	19.47	19.34	18.85	19.41	17.88	18.12	18.06	17.93
FeO	-	-	-	-	-	-	-	0.08	-	0.02
MnO	-	-	-	-	-	-	-	-	0.00	-
MgO	-	-	-	-	-	0.02	-	-	-	-
CaO	0.09	0.11	0.25	0.08	0.11	0.26	0.14	-	0.09	0.06
Na ₂ O	12.32	12.48	11.91	12.06	12.01	11.79	0.15	0.24	0.27	0.28
K ₂ O	0.08	0.04	0.06	0.04	0.07	0.10	16.72	16.38	16.33	16.52
Total	99.64	100.48	99.11	99.89	99.75	100.47	99.67	98.94	98.13	99.27
Total cation : on the basis of 24 oxygen										
Si	8.929	8.974	8.934	8.98	9.033	8.990	9.033	8.990	8.967	9.008
Ti	-	-	-	-	-	-	-	0.008	0.008	0.011
Al	3.03	2.965	3.041	2.994	2.921	2.986	2.938	2.997	3.015	2.958
Fe	-	-	-	-	-	-	-	0.009	-	0.002
Mn	-	-	-	-	-	-	-	-	-	-
Mg	-	-	-	-	-	0.003	-	-	-	-
Ca	0.013	0.016	0.036	0.010	0.016	0.036	0.020	-	0.014	0.009
Na	3.153	3.166	3.061	3.07	3.06	2.983	0.042	0.067	0.075	0.077
K	0.014	0.006	0.009	0.006	0.012	0.017	2.974	2.933	2.951	2.95
Total	15.139	15.127	15.081	15.060	15.042	15.015	15.007	15.004	15.030	15.015

Analytical condition: 15 kv, pd 15 micron, 20 nA.

Analyzed sample nos. DM 13, DM 7*, DM 10**, DM 3***

Table 3. Major and selected trace element data for the Biglii Rhyolite.

Sample [wt%]	DM 12	DM 15	DM 16	DM 2	DM 21	DM 3	DM 3U	DM 4	DM 5	DM 7	DM 7U	DM 9	C16+	C13+	160+	C03+
SiO ₂	76.58	76.40	69.44	72.37	76.52	72.99	73.47	75.00	79.90	80.43	80.65	78.88	64.44	76.55	72.51	62.41
TiO ₂	0.23	0.21	0.46	0.36	0.22	0.36	0.36	0.29	0.24	0.24	0.19	0.19	0.64	0.22	0.36	0.53
Al ₂ O ₃	11.72	11.92	13.25	12.55	11.76	11.84	12.02	11.69	11.49	11.34	10.87	11.81	12.72	11.41	12.15	13.45
FeOT	1.88	1.68	4.94	3.69	2.17	3.47	3.35	2.8	0.58	0.44	0.31	0.26	8.98	2.37	4.85	8.42
MnO	0.03	0.03	0.06	0.04	0.04	0.04	0.03	0.03	0.01	0	0	0	0.09	0.01	0.04	0.08
MgO	0.16	0.04	0.35	0.39	0.08	0.38	0.36	0.15	0.06	0.05	0.04	0.03	2.11	0.15	0.61	2.06
CaO	0.63	0.42	0.99	0.79	0.62	1.36	1.17	0.79	0.13	0.08	0.13	0.28	2.87	0.8	0.89	3.96
Na ₂ O	3.42	3.73	3.12	4.1	4.14	2.38	2.47	2.2	4.07	3.97	3.59	4.52	3.19	2.87	3.04	1.88
K ₂ O	3.41	4.69	4.21	3.83	3.12	4.8	4.78	5.11	2.82	2.92	3.61	3.28	2.29	4.03	4.15	3.98
P ₂ O ₅	0.03	0.02	0.06	0.06	0.02	0.06	0.06	0.04	0.02	0.02	0.02	0.03	0.27	0.01	0.08	0.28
LOI	1.42	0.55	1.79	1.24	1.28	1.78	1.67	1.33	0.63	0.33	0.44	0.48	2.68	0.92	1.03	3.78
Total	99.51	99.69	98.67	99.42	99.97	99.46	99.74	99.43	99.95	99.82	99.85	99.76	100.28	99.34	99.71	100.83
[ppm]																
Ba	821	1046	975	718	851	860	860	805	596	616	928	893				
Rb	116	139	164	127	77	206	204	170	101	106	113	98				
Sr	54	48	70	75	72	93	86	29	37	41	42	48				
Ga	24	24	32	24	24	26	25	24	28	25	22	20				
Zr	367	409	358	405	435	379	389	542	422	418	346	337				
Y	47	47	52	60	42	55	57	57	48	51	31	49				
Ni	-	-	-	-	-	-	-	-	-	-	-	-				
Cr	8	8	7	7	6	8	10	11	9	9	6	8				
V	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.				
Zn	64	43	105	156	109	62	54	27	27	27	26	17				
Rb/Sr	2.15	2.90	2.34	1.69	1.07	2.22	2.37	5.86	2.73	2.59	2.69	2.04				
K/Rb	244	280	213	250	336	193	195	250	232	229	265	278				
Rb/Ba	0.14	0.13	0.17	0.18	0.09	0.24	0.24	0.21	0.17	0.17	0.12	0.11				
Ga/Al	3.87	3.80	4.56	3.61	3.86	4.15	3.93	3.88	4.60	4.17	3.82	3.20				
P ₂ O ₅ /TiO ₂	0.13	0.10	0.13	0.17	0.09	0.17	0.17	0.14	0.08	0.08	0.11	0.16				
Mg #	0.13	0.04	0.11	0.16	0.06	0.16	0.16	0.09	0.16	0.17	0.19	0.17	0.30	0.10	0.18	0.30

Mg#: Mg-number.

Data collected by XRF methods at Universität zu Köln, Germany, and ⁺ Geological Survey of India, Calcutta.

Table 3. (Continued)

Sample [wt%]	M30*	M35*	M36*	M38*	M17*	M29*	10C#	51#	120A#	101A#	138B#	150B#	211#	90-67#	99#	141#	106A#
SiO ₂	75.30	75.40	75.16	74.52	75.68	74.66	76.49	71.09	72.72	75.9	75.29	77.71	77.86	77.72	77.14	75.3	74.92
TiO ₂	0.21	0.19	0.20	0.18	0.2	0.20	0.26	0.36	0.39	0.27	0.3	0.27	0.25	0.18	0.08	0.2	0.26
Al ₂ O ₃	12.1	11.5	11.30	11.4	11.3	11.40	11.36	11.55	11.83	11.4	11.61	11.14	10.82	12.12	12.4	11.38	11.47
FeOT	2.03	1.07	1.96	2.44	0.16	2.74	1.42	3.89	3.79	2.88	3.37	2.67	1.74	1.75	1.97	2.83	2.65
MnO	0.02	0.01	0.03	0.04	-	0.04	0.03	0.05	0.04	0.05	0.06	0.02	0.03	0.02	0.04	0.04	0.04
MgO	0.12	-	-	-	-	-	0.31	0.12	0.3	0.23	0.22	0.35	0.34	0.11	0.09	0.26	0.29
CaO	0.27	0.15	0.48	0.75	0.7	1.13	1.98	1.68	1.61	1.28	1.33	0.11	1.26	0.27	0.28	0.35	0.78
Na ₂ O	1.9	2.51	3.86	3.19	3.24	3.00	2.32	2.5	2.37	2.98	2.6	2.75	3.45	5.17	3.62	2.62	1.17
K ₂ O	4.85	4.94	2.96	3.46	3.68	3.94	3.8	4.72	4.92	4.06	4.14	3.76	2.7	3.06	4.7	4.88	4.66
P ₂ O ₅	0.04	0.04	0.04	0.04	0.03	0.06	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.01	0.02	0.01	0.01
LOI	2.3	3.18	3.21	2.21	3.4	2.00	2.12	2.19	2.18	1.67	2.01	1.11	2.13	0.41	0.59	0.89	1.74
Total	99.14	98.99	99.20	98.23	98.39	99.17	100.1	98.17	100.2	100.7	101.0	99.91	100.60	100.82	100.93	98.76	97.98

[ppm]	611	1472	796	921	1108	1065	730	770	900	650	770	640	730	740	440	380	530
Ba	-	-	-	-	-	-	150	215	190	136	155	127	106	75	180	177	217
Rb	29.3	54.2	63.2	68.1	53.8	71.6	80	60	70	60	60	40	140	50	50	50	40
Sr	-	-	-	-	-	-	14	19	15	18	17	15	15	24	19	18	14
Ga	393	408	431	353	397	420	270	400	400	430	350	210	270	320	120	430	470
Y	65.8	52.1	57.3	59.7	66.7	63.5	34	56	58	50	46	28	28	34	34	58	52
Ni	9.23	9.39	14	13.7	8.78	16.1	<10	<10	<10	<10	<10	<10	<10	<10	<10	<10	<10
Cr	35.3	42.9	31	42.4	42	30.7	17	<0.5	2.3	5.8	6.8	<0.5	43	2.5	6	4.5	<0.5
V	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	16	<2	8	<2	<2	4	20	<2	<2	<2	8
Zn	-	-	-	-	-	-	120	110	85	110	110	95	85	130	75	150	90
Rb/Sr	-	-	-	-	-	-	1.88	3.58	2.71	2.27	2.58	3.18	0.76	1.50	3.60	3.54	5.43
K/Rb	-	-	-	-	-	-	210	182	215	248	222	246	211	339	217	229	178
Rb/Ba	-	-	-	-	-	-	0.21	0.28	0.21	0.21	0.20	0.20	0.15	0.10	0.41	0.47	0.41
Ga/Al	-	-	-	-	-	-	2.33	3.11	2.40	2.98	2.77	2.54	2.62	3.74	2.89	2.99	2.31
P ₂ O ₅ /TiO ₂	0.19	0.21	0.20	0.22	0.15	0.30	0.08	0.06	0.05	0.07	0.07	0.07	0.08	0.06	0.25	0.05	-
Mg #	0.10	0.10	-	-	-	-	0.28	0.05	0.12	0.12	0.10	0.19	0.26	0.10	0.08	0.14	0.16

Mg#: Mg-number.
Data collected by ICP methods at * JNU, New Delhi and # Activation Laboratory, USA.

Table 4. Trace element data for the Bijli Rhyolite.

Sample [ppm]	10C	51	120A	101A	138B	150B	211	90-67	99	141	106A	M36*	M30*	M35*	M29*	M17*	M38*	DM 2	Error (%)
Sc	3	4.6	5.1	2.5	3	3.3	0.2	1.6	3.2	1.5	2.7							5.4	±3
Cr	17	< 0.5	2.3	5.8	6.8	< 0.5	43	2.5	6	4.5	< 0.5	30.9	35.3	43	30.7	42	42.4	50	±5
Mn																		200	±8
Co	41	61	76	57	69	64	63	100	84	38	34							5.0	±5
Ni	< 10	< 10	< 10	< 10	< 10	< 10	< 10	< 10	< 10	< 10	< 10	14	9.23	9.39	16.1	8.78	13.7	< 10	
Cu	10	5	25	< 5	10	10	25	< 5	< 5	10	5								
Zn	120	110	85	110	110	95	85	130	75	150	90							170	±5
Ga	14	19	15	18	17	15	15	15	19	18	20							13	±15
As	1	< 1	1	< 1	< 1	< 1	1	< 1	1	2	< 1							1.20	±25
Se	< 0.5	< 0.6	< 0.7	< 0.5	< 0.6	< 0.6	< 0.5	< 0.5	< 0.5	< 0.7	< 0.5								
Br	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5	< 0.5								
Rb	150	215	190	136	155	127	106	75	180	177	217							130	±5
Sr	80	60	70	60	60	40	140	50	50	50	40	63.2	29.3	54.2	71.6	53.8	68.1	80	±20
Y	34	56	58	50	46	28	28	34	34	58	52	57.3	65.8	52.1	63.5	66.7	59.7		
Zr	270	400	400	430	350	210	270	320	120	430	470	431	393	408	420	397	353	330	±10
Nb	20	29	31	23	24	18	19	22	21	32	23							14.4	±12
Mo	< 5	< 8	< 8	< 6	< 5	< 2	< 5	< 4	2	< 1	< 6							< 1	
Ag	< 0.1	< 0.1	< 0.1	0.1	< 0.1	< 0.1	0.2	0.1	< 0.1	< 0.1	< 0.1								
Sn	< 5	< 5	< 5	< 5	< 5	< 5	< 5	< 5	< 5	< 5	7								
Sb	0.2	0.2	0.2	0.4	0.4	0.3	0.2	< 0.1	0.5	0.6	0.4							0.3	±20
Cs	2.1	1.7	1.1	1	1.6	1.1	1.2	< 0.2	1.6	2.3	2.6							2.0	±5
Ba	730	770	900	650	770	640	730	740	440	380	530	796	611	1472	1065	1108		755	±10
La	65.9	86.3	105	87.1	68.5	76	53.1	50	31.9	115	94.7	98.09	60.86		5.75			80	±3

Table 4. (Continued)

Sample [ppm]	10C	51	120A	101A	138B	150B	211	90-67	99	141	106A	M36*	M30*	M35*	M29*	M17*	M38*	DM 2	Error (%)
Ce	110	154	172	152	119	131	94	90	62	190	160		182	118	10.9			200	±8
Pr																		26	±30
Nd	42	60	58	60	45	50	35	44	30	71	65		65.04	49.91	3.41			78	±10
Sm	7.1	9.8	9.8	9.9	7.7	7.6	5.6	7.6	6.1	11	10		11.48	9.07	0.93			11.0	±5
Eu	1.05	1.5	1.26	1.61	1.41	1.16	0.77	1.2	0.66	0.92	1.18		1.41	1.56	0.17			1.3	±5
Gd													10.37	8.21	0.82			21	±10
Tb	1.3	1.8	1.5	1.6	1.4	0.9	0.9	1.2	1.2	1.7	1.7							1.4	±8
Dy													9.62	7.24	0.78			7.4	±10
Ho																		2	±8
Er													9.01	6.39	0.75				
Tm																			
Yb	3.19	5.57	4.68	5.51	3.97	2.33	2.41	3.75	3.5	5.02	5.3		5.64	4.54	0.48			0.60	±15
Lu	0.47	0.86	0.68	0.74	0.58	0.36	0.4	0.52	0.5	0.8	0.78		0.88	0.70	0.08			5.6	±8
Hf	6.3	8.7	8.5	10	8.4	5	6.3	8.6	4.1	10	10							0.88	±10
Ta	1.5	2.4	2.9	2.1	2.2	2	1.5	2.9	3.4	2.1	1.8							10.8	±5
W	370	550	620	440	570	490	480	860	670	320	290							2.5	±5
Ir	<1	<1	<1	<1	<1	<1	<1	<1	<1	<1	<1							<2	
Au	<2	<2	<2	4	<2	<2	2	4	<2	<2	<2							0.003	±30
Hg	<1	<1	<1	<1	<1	<1	<1	<1	<1	<1	<1								
Th	21	35	35	24	20	19	18	13	21	29	27							39.0	±5
U	5.3	8	7.8	5.6	4.8	2.6	5.2	3.6	4.4	8.1	5.6							9.0	±10

DM2: INAA data (done at the Universität zu Köln, Germany).
 Others: ICP data at Activation Laboratory, USA and * JNU, New Delhi.

Table 5. *O-isotope data for the Bijli Rhyolite.*

Sample no.	DM 21	DM 7	DM 9	DM 3
$\delta^{18}\text{O}$ (whole rock)	7.8 ± 0.1	6.1 ± 0.1	5.2 ± 0.1	4.4 ± 0.05
$\delta^{18}\text{O}$ (quartz)	7.35 ± 0.05	–	–	4.5 ± 0.1
$\delta^{18}\text{O}$ (feldspar)	6.6 ± 0.1	–	–	4.2 ± 0.1

The Bijli rocks have lower Al_2O_3 , MgO , CaO , Sr , and higher FeO , Zn and HFSE (e.g., Zr , Nb , Ta , Hf , Ga) and REE concentrations (tables 3 and 4) in comparison to the fractionated I-type granites (felsic I-type granite of Whalen *et al* 1987) with similar SiO_2 contents. High SiO_2 , $\text{Na}_2\text{O} + \text{K}_2\text{O}$, Fe/Mg , Zr , Ta , Ga , Sn , Y , and REE (except Eu) contents and low CaO , Ba , Sr of the Bijli samples rather indicate similarity with a ‘A-type granite’ (cf. Collins *et al* 1982; Whalen *et al* 1987; Eby 1990). High Ga/Al (average 3.71) of the Bijli rocks is also similar to A-type granites (see Whalen *et al* 1987; Eby 1990).

5.3 Major elements

The most striking feature of the Bijli Rhyolite samples is their high silica content and narrow range of major element compositions (table 3). The silica content varies from 64 wt% to 80.65 wt% with an average $\text{SiO}_2 \sim 75$ wt%. The different major oxides are plotted against SiO_2 wt% (figure 7). The plots show that TiO_2 wt%, FeO wt% and MgO wt% decrease with increasing SiO_2 wt%. The data in MgO diagram plot along a curved trend, whereas TiO_2 , FeO_t trends are practically linear. The scatter of data is larger for Al_2O_3 and CaO plots; the average Al_2O_3 content (~ 12 wt%) significantly less than in calc-alkaline rhyolite (> 14 wt%) (Philpotts 1990). A close study of Al_2O_3 plot, however, shows decrease in Al_2O_3 contents with increasing silica, but along two sub-parallel trends with an increase in Al_2O_3 contents at $\text{SiO}_2 \sim 76$ wt%. A similar increase in CaO concentrations in the suite can also be detected at similar silica level ($\text{SiO}_2 \sim 76$ wt%). The K_2O content in the rocks is much higher than that of calc-alkaline rhyolites (Hess 1989, p 224). The $\text{K}_2\text{O}/\text{Na}_2\text{O}$ ratio of the rocks is 1.43, with K_2O and Na_2O showing mutually negative relationship (figure 8). It is important to note here that concentrations of K_2O and Na_2O are clearly higher in Kotri samples than rest of the data, which is consistent with the presence of more alkali rich lava like trachyte and trachy-andesite there. The rise in Al_2O_3 and CaO contents at a particular silica value may be of significance, as discussed later.

5.4 Trace elements

Unlike major elements, trace element distributions in the rocks exhibit larger variations (tables 3 & 4) (this work; Neogi *et al* 1996; Divakara Rao *et al* 2000; Sensarma 2001; Sensarma *et al* 2002). The rocks are enriched in incompatible elements like K , Ba , Rb , Th , U , Ta , Zr , and LREE, and depleted in compatible elements like Cr (~ 8 ppm), Sc (< 10 ppm) and Mn compared to the upper continental crust (UCC). When plotted against silica, certain incompatible element ratio like K/Rb (200–400) and Rb/Sr (1.5 – 6) in the suite varies within a small range (figure 4). K/Ba ratios maintain constant value at 100 (figure 5), though a couple of samples show large scatter associated with occasional Ba mobilization during secondary processes, as discussed earlier. PUM (Primitive Upper Mantle of McDonough and Sun 1995) normalized plots for the incompatible elements (figure 9) show a pattern similar to that of Upper Continental Crust (UCC) of Taylor and McLennan (1985), but display overall enrichment by a factor of 2 to 3 compared to the UCC. There are distinct negative Cs -, Sr -, Eu - and Ti -anomalies; Sr -anomaly being the most prominent. The negative Cs -anomaly is responsible for the very high Rb/Cs ratios (> 100) of the rhyolites compared to Rb/Cs (30) of the UCC. Thorium (Th) and Uranium (U) concentrations in the rocks are high (about 10 times) compared to the UCC. Th/U ratios in the rocks range between 1 and 7 against a crustal Th/U ratio (~ 4) (figure 5). The samples maintain crustal ratios of Nb/Ta (11) (figure 10), and Zr/Hf (35), but have low Ta/Th (≤ 0.06 vs. 0.13) compared to the UCC. Zr/Y ratios in the Kotri samples vary by about two times (6–10) (figure 10). The REE abundances in the rocks are high. Chondrite normalized Rare Earth Elements (REE) patterns (figure 11) show that the samples have very steep LREE pattern with high La/Sm (~ 6.5) and La/Yb (~ 20) and very pronounced negative Eu -anomalies. HREE patterns show slightly negative trend ($\text{Gd}/\text{Yb} \sim 3$) within the limits of analytical error with concentrations more than 20–30 times chondritic. The REE patterns are not parallel to one another, rather sometimes cross each other. Such crossing of REE patterns from presumably co-genetic lava is unique to high-silica rhyolites (Hess, 1989).

5.5 Oxygen isotopes

Oxygen isotope data for selected samples of the Bijli rhyolite rocks are presented in table 5. Whole-rock data show an average value of $5.87 \pm 1.26\text{‰}$ comparable to mantle value ($5.7 \pm 0.3\text{‰}$) (e.g., Rollinson 1993), though with a larger spread. Since large shifts in $\delta^{18}\text{O}$ value can occur in vol-

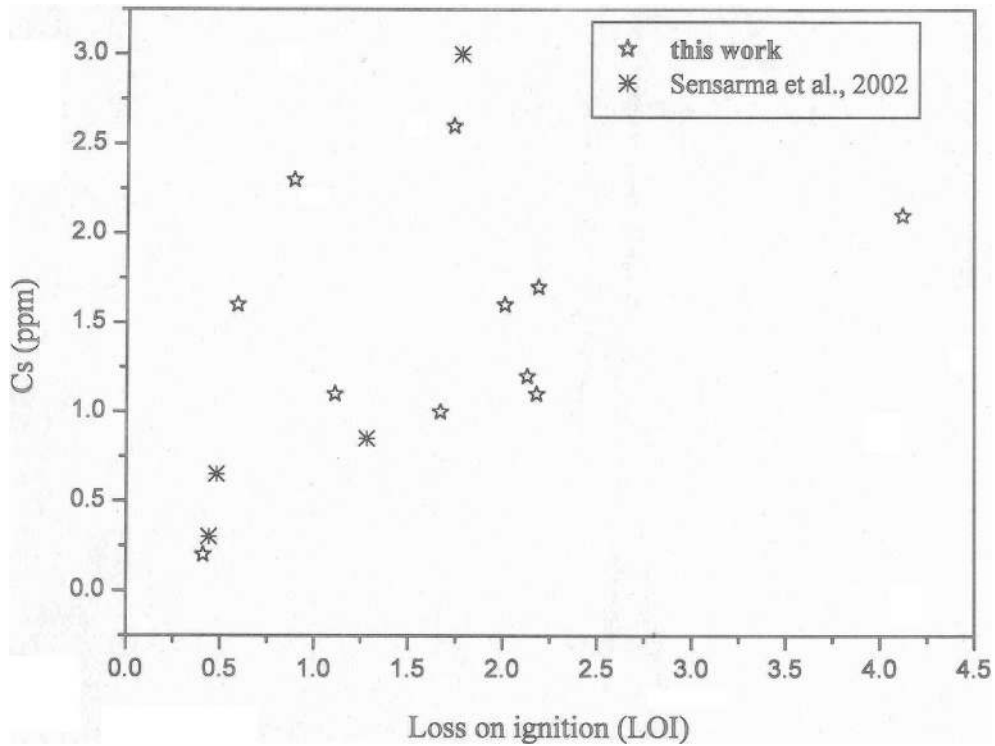


Figure 3. LOI (loss on ignition) vs. Cs (ppm) plot showing a general rise in Cs concentrations with increase in LOI in the Bijli Rhyolite.

canic rocks due to low temperature alteration processes, we also analyzed grains of quartz and K-feldspar phenocrysts from the same samples, which are likely to remain unaffected by alteration processes (Taylor 1980). We note that the $\delta^{18}\text{O}$ value for whole rock and quartz separates for DM3 ($4.4 \pm 0.05\text{‰}$ vs. $4.5 \pm 0.1\text{‰}$) and DM21 ($7.8 \pm 0.1\text{‰}$ vs. $7.35 \pm 0.05\text{‰}$) (table 5) are comparable.

This further confirms the generally unaltered character of our samples. However, as quartz must always be heavier than the corresponding whole rock value, slightly enriched DM21_{whole-rock} value relative to its quartz may indicate some degree of alteration. $\delta^{18}\text{O}_{\text{K-feldspar}}$ in DM3 ($4.2 \pm 0.1\text{‰}$) and DM21 ($6.6 \pm 0.1\text{‰}$) is lower than that of both quartz and whole rock (table 5), though the quartz-K-feldspar fractionation between 0.8 (DM21) and 0.3 (DM3) is very close to values typical of fresh unaltered magmatic rocks.

6. Thermometry

Temperatures calculated from oxygen isotope fractionations (e.g., Hoffbauer *et al* 1994) suggest 880°C for DM21, a value close to solidus temperature of dry rhyolitic magmas (Nekvasil 1992); and a value of 1550°C for DM3, which is, however, by far too high. This might be connected to the unusual oxygen isotopic composition of

DM3. The whole-rock value of 4.4‰ for this sample is too low even for any mantle derived melt (Taylor and Sheppard 1986), and thus could be indicative of fluid-rock exchange with a meteoric component. Any substantial exchange, however, would drive the feldspar isotopic composition to effect larger quartz-K-feldspar fractionation. The low values reported here therefore indeed represent low ^{18}O magma, and at the same time a strong indication of participation of deeper crustal component/mantle in the genesis of the Bijli Rhyolite.

The Zr solubility in our melts (~ 400 ppm) with a low cation ratio $((\text{Na} + \text{K} + 2\text{Ca})/(\text{Si} \times \text{Al}) = 0.03)$ (cf., Watson and Harrison 1983) indicate a temperature range of $860^\circ\text{--}930^\circ\text{C}$. The P_2O_5 - (Harrison and Watson 1984) and Mg-thermometers (see in Johannes and Holtz 1996) yield a similarly high temperature ($900^\circ\text{--}950^\circ\text{C}$ and $850^\circ\text{--}900^\circ\text{C}$ respectively) for our rocks. One sample (C03) with higher P_2O_5 content (0.28 wt%) at SiO_2 (64.41 wt%) indicates even a higher temperature ($\sim 950^\circ\text{C}$). The high temperature ($\sim 900^\circ\text{C}$) for the Bijli melts as obtained by different thermometers is primary, because such a temperature is not attainable during low grade metamorphism of the green schist facies condition. Also, it is independent of pressure of melting or melts H_2O contents. The presence of anorthoclase in the rocks also argues for high temperature Bijli melts.

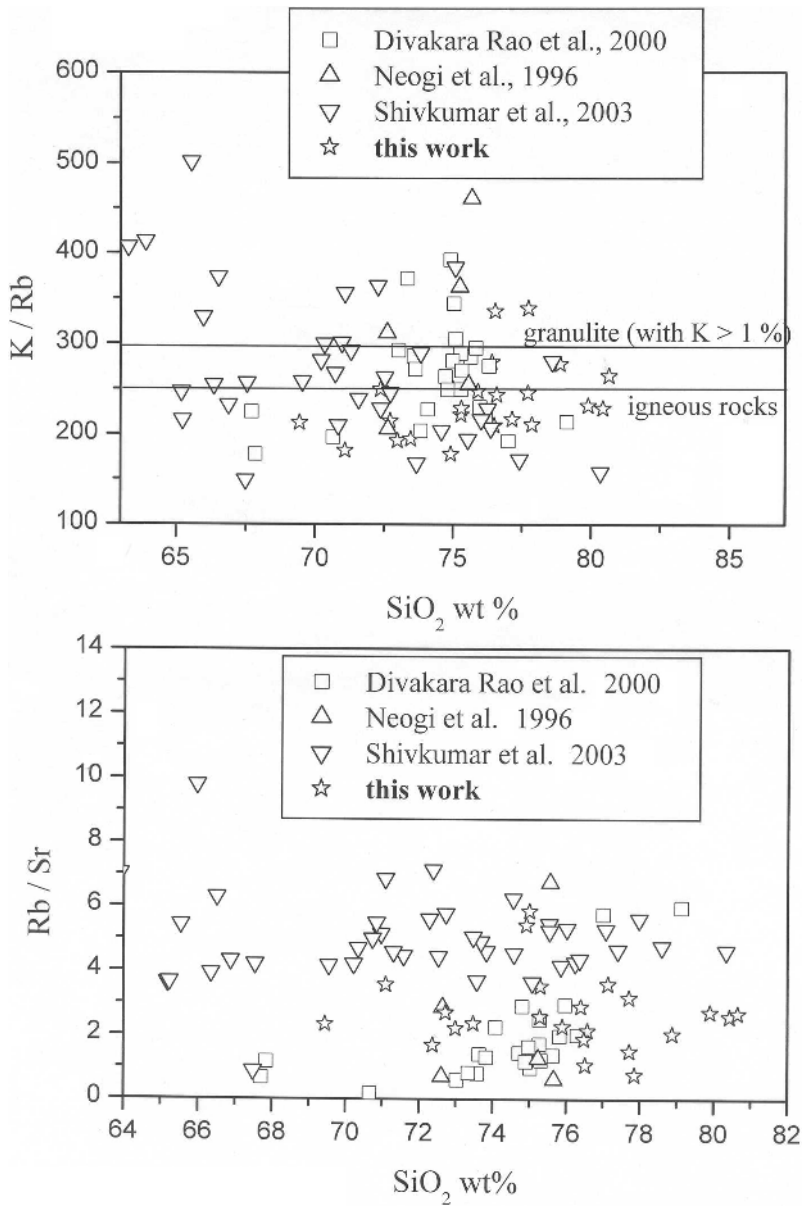


Figure 4. SiO_2 wt% vs. K/Rb and Rb/Sr plots. K/Rb (200–300) of the rocks are in the range of average igneous rocks. Rb/Sr of the rocks vary by two to three times. Symbols are described in the legend. Note higher Rb/Sr values for the Kotri samples.

7. Petrogenesis

7.1 Fractional crystallization of basaltic magma

The trends in SiO_2 -FeO, SiO_2 -MgO and SiO_2 -TiO₂ plots show decrease in FeO, MgO and TiO₂ contents with silica in the bulk liquid (figure 7), which may indicate some sort of crystal-liquid fractionation process, e.g., fractional crystallization, partial melting and/or contamination. To test the fractional crystallization hypothesis, we adopted the 'lever rule' principle of Cox *et al* (1987), where it is possible to calculate graphically the way in which a liquid composition changes as a particular mineral

is removed. On SiO_2 -Al₂O₃ plot (figure 12), fractionations of plagioclase-pyroxene seem possible, as the line joining pyroxene-plagioclase intersect the bulk compositional trend of the Bijli rocks. Plot of pyroxene along the bulk compositional trend in CaO-SiO₂ plot (figure 12) also indicates involvement of pyroxene in the evolution of the Bijli melts. However, the proportions of pyroxene fractionation as per these plots do not tally with the negligible amount required by the MgO-CaO plot (figure 12). So, the plots are not internally consistent to signify any role of pyroxene fractionation in the evolution of the Bijli melts. A similar conclusion can be drawn for ilmenite fractionation from FeO-TiO₂

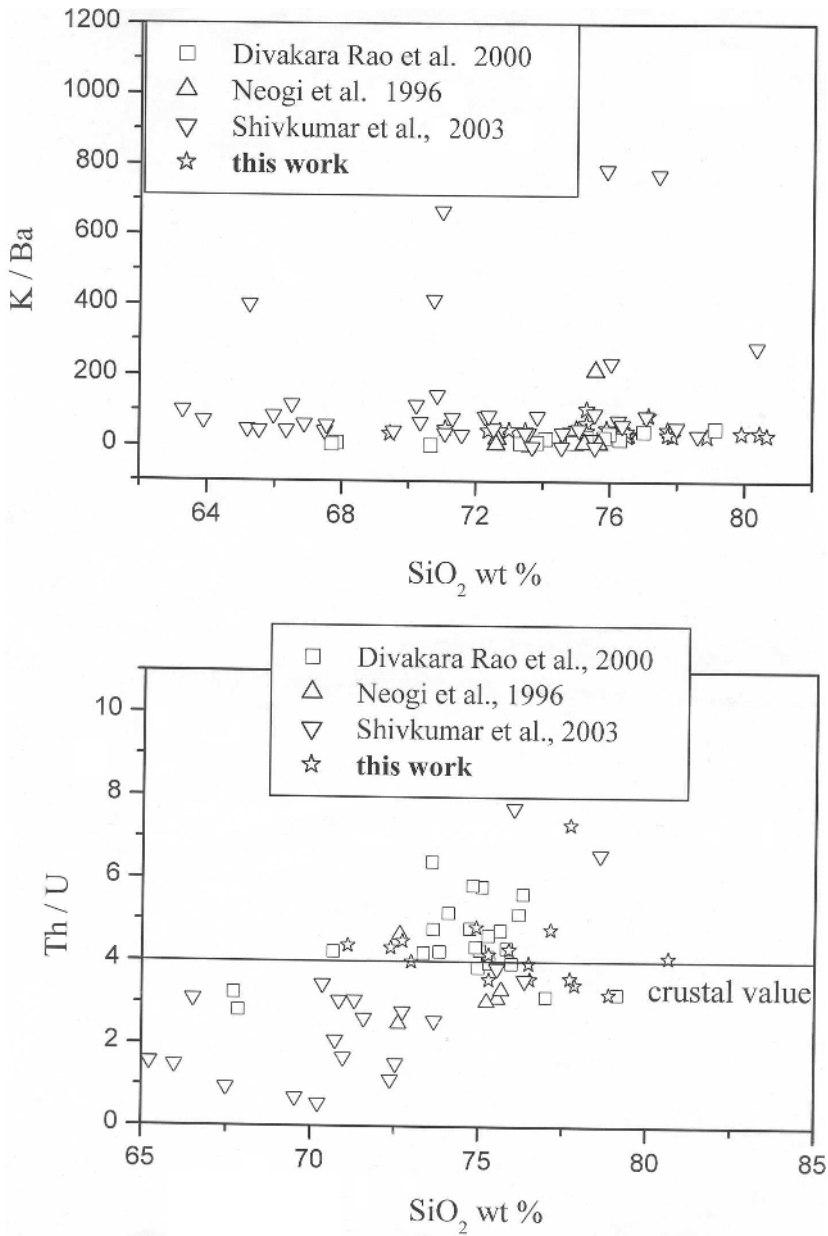


Figure 5. SiO₂ wt% vs. K/Ba and Th/U plots. Note the constant K/Ba (100) for the Bijli samples. Few samples with higher K/Ba values indicate Ba mobilization. Th/U for many samples are at or near crustal value (4). Symbols are described in the legend. Note low Th/U values for the Kotri samples.

and CaO-TiO₂ plots (figure 12). In the basaltic magma fractionation model, a large amount of basalt is needed, and it would produce only a small amount of derivative silicic melt contrary to what we have in Dongargarh. Also, continuous ranges of differentiated product of a basaltic magma like andesite, dacite, rhyodacite do not occur in spatial proximity to the rhyolite in the area. Rather, the Bijli Rhyolite and overlying basalts show clear presence of the 'Daly gap' (figure 2). We therefore do not see the possibility of a basaltic magma fractionation to be solely responsible for generation of the Bijli melts.

However, it is difficult to explain elevated concentrations of certain compatible elements (e.g., Co, Cr) in few rhyolite samples with high silica (SiO₂ > 75 wt%) and incompatible trace element (LILE, HFSE, LREE) contents (tables 3 and 4) without participation of mantle-derived melts. At the same time, the Bijli suite includes rocks with SiO₂ contents of 65 wt% or less, that are not easily explained as pure crustal melts. The involvement of a mantle derived magma in the origin of the Bijli melts therefore still remains valid, but could be better known once we understand the role of crust in this scenario.

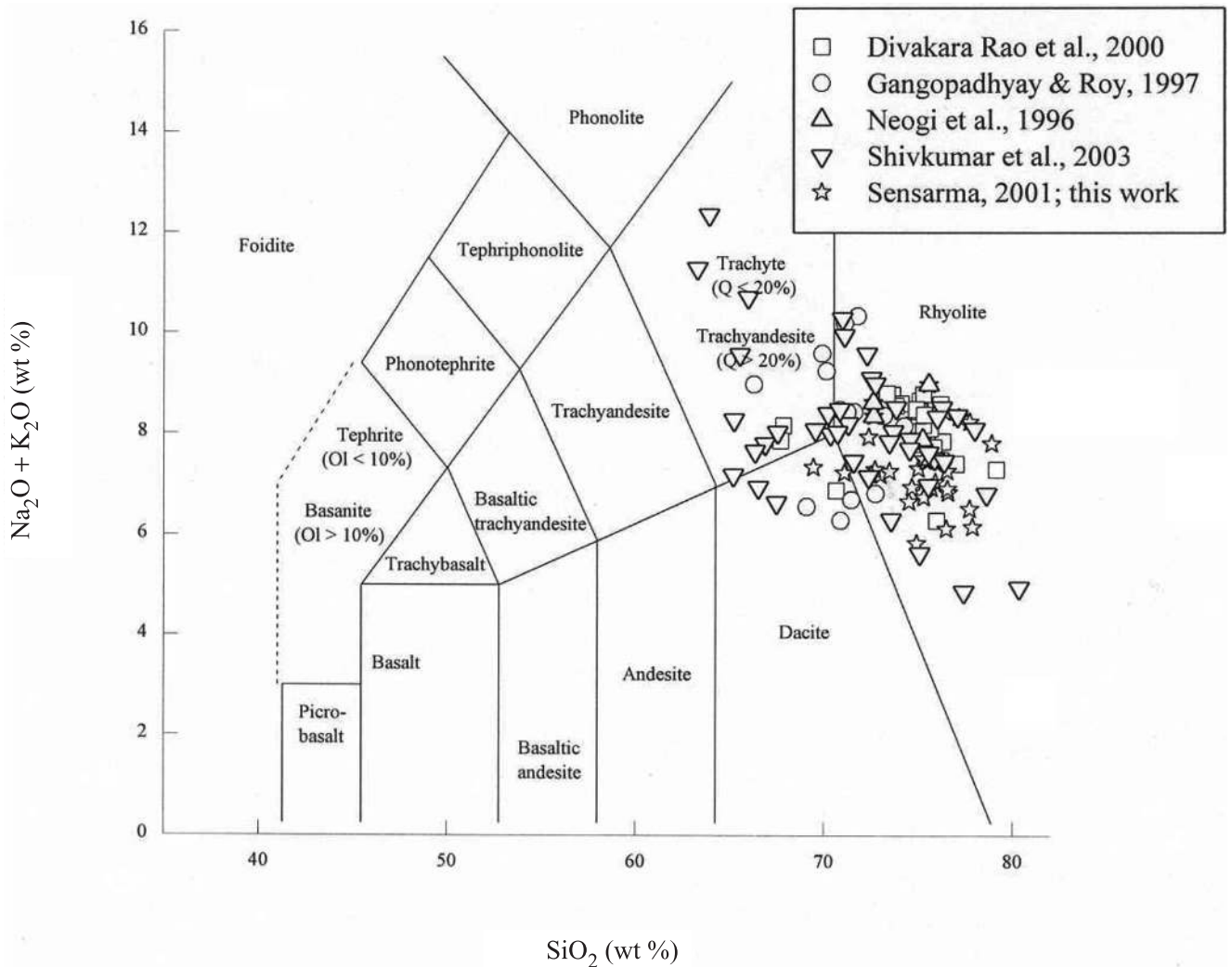


Figure 6. Plot of the Bijli Rhyolite samples on a total alkali-silica (TAS) diagram (after Le Bas *et al* 1986). Different symbols indicate different data sources, as shown in the legend.

7.2 Crustal origin

The possibility of crustal melting in the origin of the Bijli Rhyolite is certainly attractive as the presence of the silica gap is often considered as evidence in support of a partial melting model (Bellieni *et al* 1986). Comparable $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of the Bijli Rhyolite, 0.70305 ± 0.0017 (Krishnamurthy *et al* 1990) and 0.7057 ± 0.0015 (Sarkar *et al* 1981), to that of the crustal value (~ 0.703) at ~ 2.5 Ga (see Philpotts, 1990, p 434) may indicate a crustal origin for these rocks. The high alkali contents and the similarity in trace element distribution patterns of the samples and that of UCC (figure 9) also probably point to a crustal derivation. Crustal anatexis requires anomalously high temperatures within the continental crust through localized heat source, and/or by decompression of heated rock during crustal thinning (Thompson 1999). The potential heat source could be the high temperature mantle-

derived melts (cf. Clemens and Vielzeuf 1987; Huppert and Sparks 1988), which erupted subsequently as basaltic komatiite/high-Mg basalts in the Dongargarh succession (see above). It is not clear whether the fractionation of crustal partial melts was important in the evolution of this high silica melts, as suggested for the Chon Aike rhyolites in Patagonia (Pankhurst and Rapela 1995). We evaluate below the role of fractional crystallization of crustal partial melt, and then examine the possible crustal source(s) and degree of melting that could have given rise to the Bijli melts.

7.2a Fractionation of a crustal melt

Feldspar

On $\text{SiO}_2\text{-Al}_2\text{O}_3$ and MgO-CaO diagrams (figure 12), the join between potential fractionating phases from a silicic melt, such as K-feldspar and plagi-

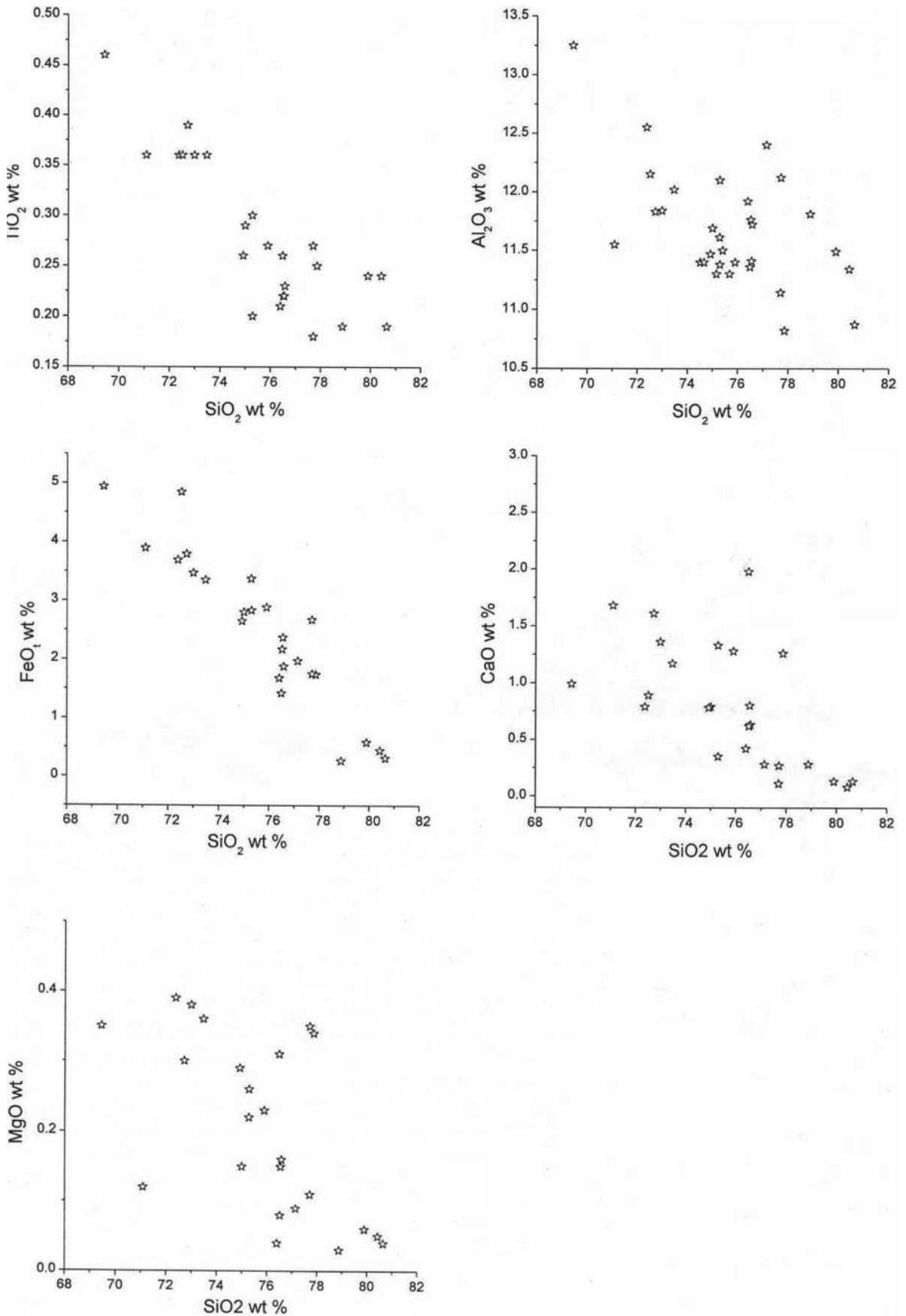


Figure 7. The Bijli rocks show decrease in major oxide contents with increasing SiO₂ wt%. The trends in MgO, FeO_t, TiO₂ plots are linear. Al₂O₃ plot shows two sub-parallel trends with rise in Al₂O₃ and CaO contents at SiO₂ ~ 76 wt%.

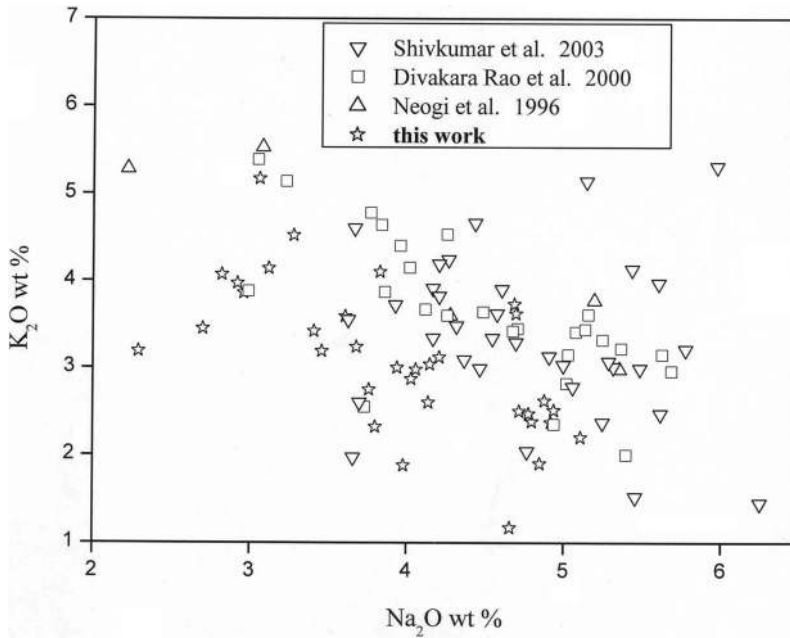


Figure 8. Na_2O - K_2O plot for the Bijli Rhyolite samples. Na_2O and K_2O show mutually negative relationship.

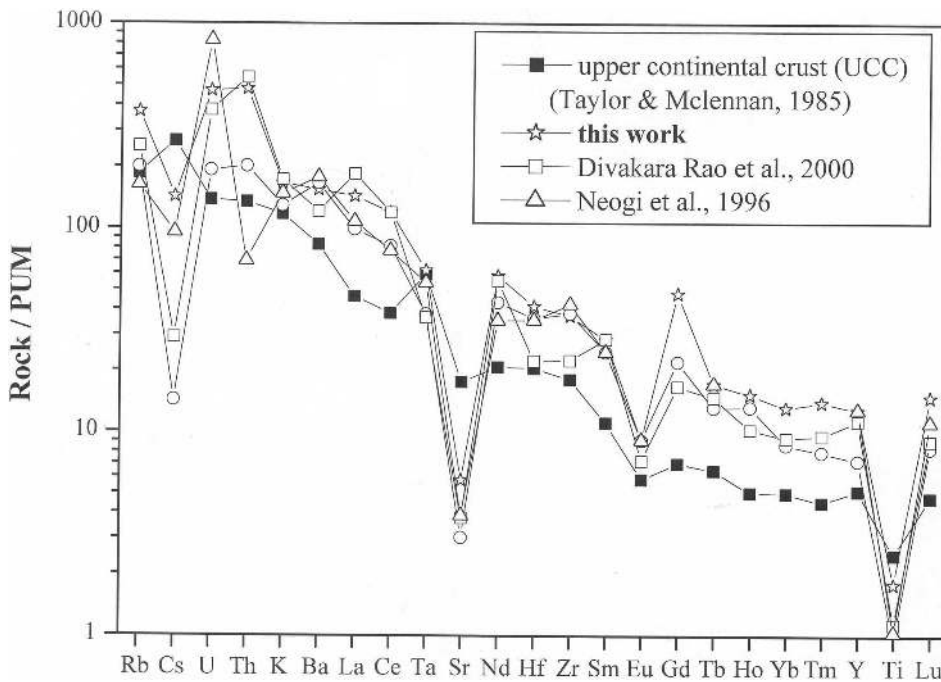


Figure 9. Primitive upper mantle (PUM)-normalized incompatible elements plot of the Bijli Rhyolite. PUM-normalized incompatible element concentrations of upper continental crust (UCC) are shown for comparison. PUM values are taken from McDonough and Sun (1995). Symbols are described in the legend.

clase, intersect the bulk compositional trend. From the proportion of intercepts of the join with the bulk compositional trend, plagioclase appears to be the main fractionating phase. On the other hand, in Al_2O_3 - Na_2O and SiO_2 - CaO diagrams, plagioclase and K-feldspar not only plot on the bulk compo-

sitional trend, but their plots coincide with some bulk compositions (figure 12). This may be suggestive of feldspar-melt relationship, but does not argue for fractionation crystallization of feldspars, as removal of feldspars would drive the bulk composition away from it. Further arguments against K-

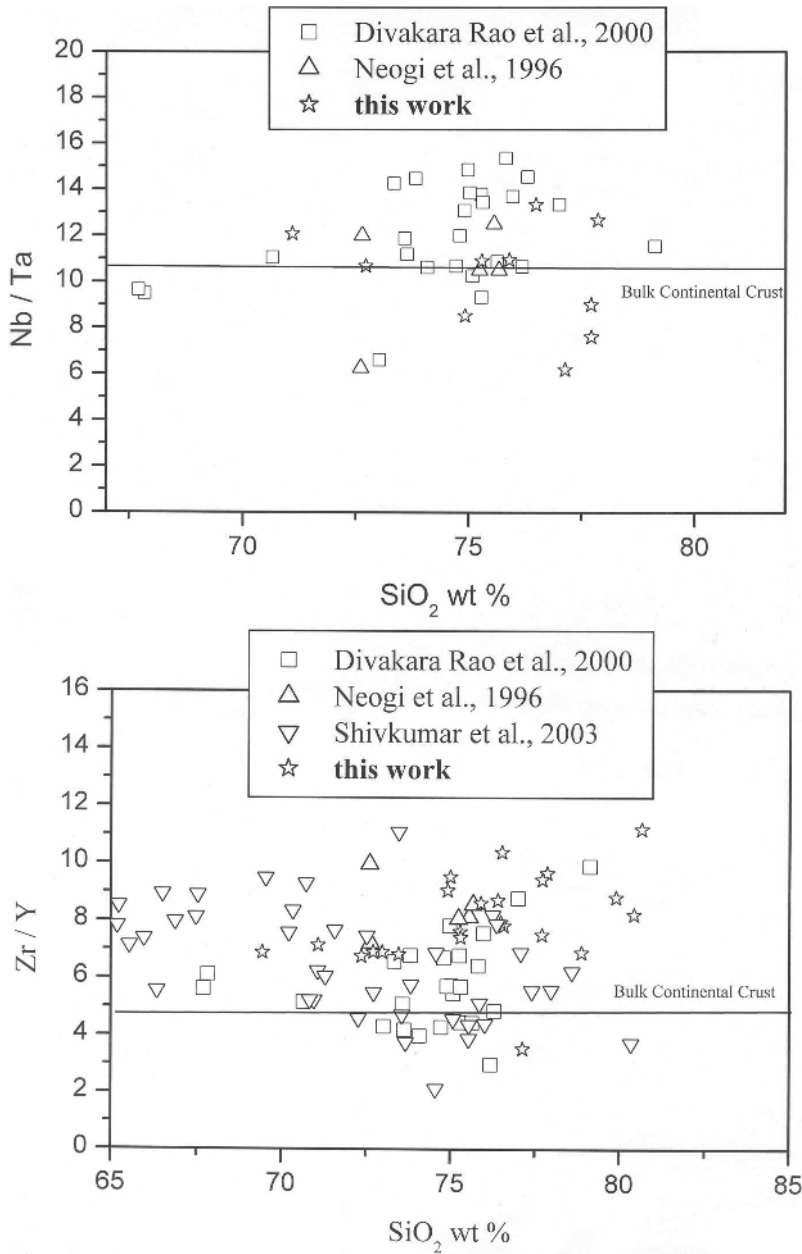


Figure 10. Nb/Ta for the Bijli Rhyolite show near crustal values, but Zr/Y are little higher than the bulk continental crust (BCC). Different symbols for different data sources, as shown in the legend.

feldspar fractional crystallization come from trace elements data. As already discussed, K-feldspar is the principal phenocrystic phase in our rocks, and its removal by fractionation would have considerably decreased K₂O, Ba and Rb concentrations and consequently decrease Rb/Sr ratios with increasing silica. But the near constant K₂O contents (4–5 wt%) with little variation in Ba contents and no mutual negative relationship between Rb/Sr and SiO₂ wt% (tables 3 and 4, figure 4) do not suggest K-feldspar fractionation. K-feldspar rather occurs as a near-liquidus phase early in the crystallization history as also observed in many ‘A-type gran-

ite’ (Landenberger and Collins 1996). Also, if negative Sr- and Eu- anomalies (figures 9 and 11) were indicative of plagioclase fractionation as argued by Divakara Rao *et al* (2000), Rb/Sr ratio would have significantly increased with increasing silica, which is also not observed in our samples (figure 4). Simultaneous rise in Al₂O₃ and CaO concentrations in the melt at a particular silica level (SiO₂ ~ 76 wt%) (figure 7) rather indicate addition of plagioclase. Negative Sr- and Eu-anomalies thus could be indicative of retention of plagioclase in the residue during partial melting, as discussed below.

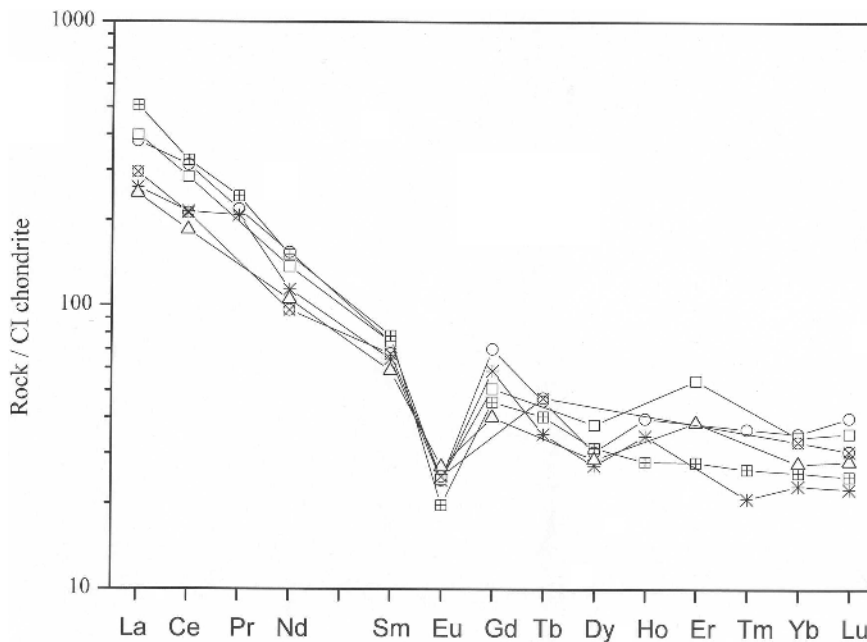


Figure 11. CI-chondrite normalized Rare Earth Elements (REE) plot for the Bijli Rhyolite samples.

Accessory minerals

Most felsic magmas are saturated with accessory minerals (REE-, Y-, Th-, U-bearing), which play an important role in controlling the geochemistry of the crustal melts (Watson 1988). Magmatic epidote has high to moderate REE contents. Fractionation of epidote may account for part of the strong negative Eu-anomalies in our samples. Allanite contains Th ($\text{ThO}_2 \sim 1 \text{ wt}\%$), but hardly any U ($\text{UO}_2 \sim 0.0\text{--}0.2 \text{ wt}\%$) (Bea 1996). High concentrations of Th and U in our samples thus possibly preclude any role for allanite. Low Th/U ratios particularly in Kotri samples (e.g., Shivkumar *et al* 2003) may suggest fractionation of phases like Th-orthosilicate. The concentrations of Th and U are highly variable in zircon, even among grains from the same thin section (Bea 1996). So low Th/U values may also suggest zircon fractionation. The scattering in Zr/Y ratios may also indicate zircon fractionation. Near crustal Nb/Ta ratios in the rocks (figure 10) do not indicate fractionation of phases like niobotantalates. The role of LREE-rich accessory minerals such as monazite is not clear given the smooth LREE patterns in the samples (figure 11). Strong negative Ti anomalies may, however, indicate titanomagnetite fractionation, rather than sphene, because sphene fractionation would have resulted in Y depletion, which is not observed in the samples (figure 9). In summary, fractionation crystallization of feldspars and accessory phases were not important in the evolution of the Bijli melts, except for zircon and Th-orthosilicate.

7.2b Nature of source rock

There is a lot of controversy about the possible nature of the crustal protolith for the generation of high silica rhyolite. Here, we evaluate the possibility of two types of crustal sources: (a) the possible basement granitoid gneissic rocks (Amgaon Gneiss) in the area, and (b) a deeper crustal source compositionally similar to the Average Archaean Granulite (AAG) of Rudnick and Presper (1990). Sarkar *et al* (1994), on the basis of mass balance computations, have suggested extensive partial fusion (60–68%) of the granitoid basement for the generation of the Bijli rhyolites. Sedimentary rocks and its metamorphic equivalents are not suitable as source, because experimental studies reveal that the granitic melts produced by partial melting of this type of source are strongly peraluminous (Johannes and Holtz 1996; Patiño Douce 1999), whereas the Bijli rhyolites are dominantly metaluminous in composition.

7.2c Comparison with experimental melts

The normative compositions of the first melts produced by dry partial melting of charnockite at higher temperatures and pressure ($\sim 950^\circ\text{C}$, 15 kbar) (Litvinovsky *et al* 2000), and H_2O -undersaturated basaltic andesite to dacite melting at $\sim 1000^\circ\text{C}$ at 1 atmosphere (Brugger *et al* 2003) are compared with the Bijli melts on a Q-Ab-Or space (figure 13). The plots of the Bijli rhyolites apparently show scatter. Nevertheless, the plot brings out some useful information.

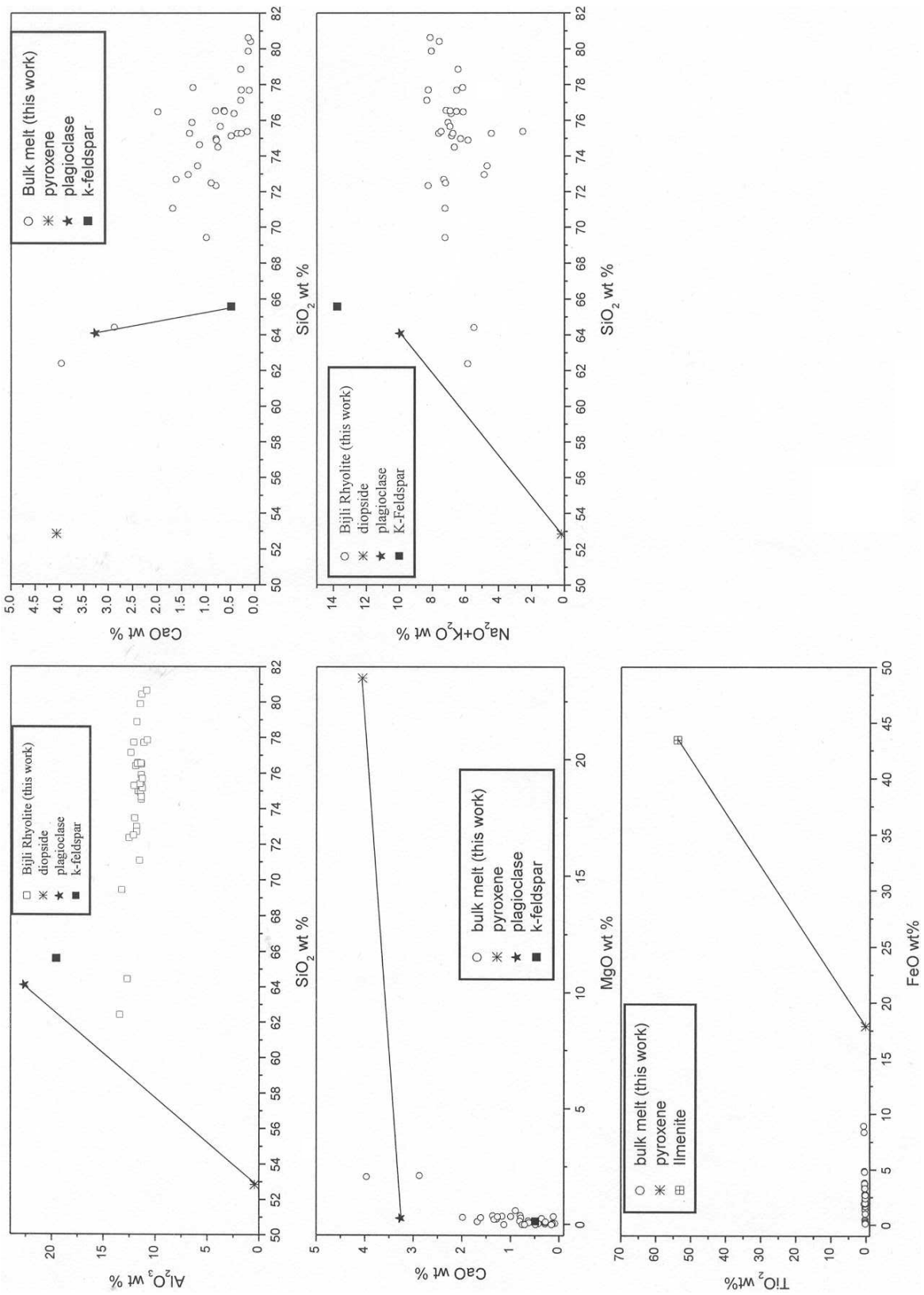


Figure 12. Figure showing relation between potential fractionating phases (joined by line) and bulk liquid trend. Pyroxene-plagioclase assemblage intersects the bulk liquid trend in different proportions in Al₂O₃ wt%-SiO₂ wt%, CaO wt%-MgO wt% and (Na₂O + K₂O) wt%-SiO₂ wt% plots. CaO wt%-MgO wt% shows compositional similarity of K-feldspar to many bulk samples. In FeO wt%-TiO₂ wt%, pyroxene-ilmenite assemblage does not intersect the bulk trend. Plagioclase intersection is not in similar proportions.

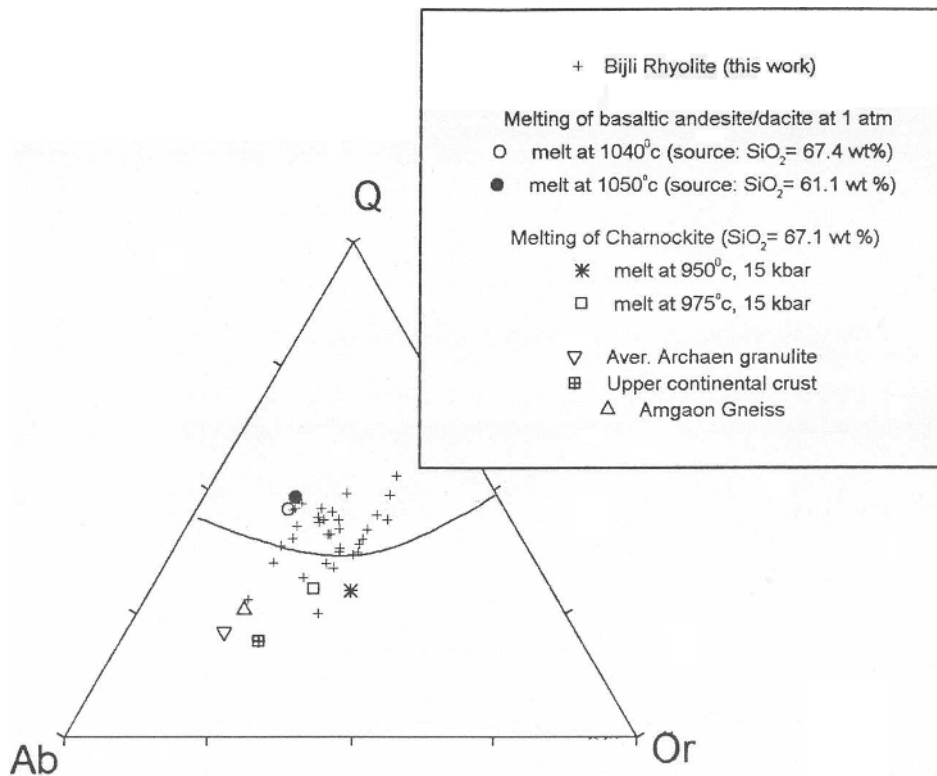


Figure 13. Plot of Bijli Rhyolite bulk compositions, experimental melts and crustal lithologies including the minimum in the Quartz (Q)–Albite (Ab)–Orthoclase (Or) system at $P_{\text{H}_2\text{O}} = 1 \text{ atm}$. **Data sources:** upper continental crust (UCC) (Taylor and McLennan 1985), average Archaean Granulite (AAG) (Rudnick and Presper 1990), Amgaon Gneiss (average) (Sarkar *et al* 1994), partial charnockite melts (Litvinovsky *et al* 2000), and partial basaltic andesite/dacite melts (Brugger *et al* 2003).

- There are samples which are compositionally not far from the normative composition of the first partial melt ($\text{Ab}_{35.32} \text{Or}_{35.03} \text{Q}_{29.65}$ with $\text{SiO}_2 \sim 72 \text{ wt}\%$) of charnockite melting.
- There is a continuous change in bulk rock compositions from partial melts of charnockite at high pressure (15 kbar) to eutectic melt composition ($\text{Ab}_{38.06} \text{Or}_{15.98} \text{Q}_{45.97}$) of basaltic andesite/dacite melting (Brugger *et al* 2003) with concomitant rise in Ab and Qtz components.
- There are melts with similar levels of normative quartz as in (b), but elevated normative orthoclase at the expense of albite ($\text{Ab}_{30.21} \text{Or}_{32.43} \text{Q}_{37.36} - \text{Ab}_{23.24} \text{Or}_{32.02} \text{Q}_{44.74}$).

The Bijli samples compositionally similar to the type (a) might indicate that some deeper crustal source of charnockitic composition might be directly involved in the formation of silicic melts in the area. The bulk Amgaon granitoid composition (Sarkar *et al* 1994) is not much different from that of the bulk charnockite composition of Litvinovsky *et al* (2000), except for higher Al_2O_3 (17 wt%) and slightly lower CaO and MgO con-

tents. The Amgaon rocks could thus be an equally potential source as suggested by Sarkar *et al* (1994). Irrespective of the crustal source, however, moderate negative K- and Eu-, and strong negative Sr-anomalies in the incompatible elements plot (figure 9) indicate the presence of both K-feldspar and plagioclase in the residue, which may have retained the elements resulting in negative anomalies and also much of Al_2O_3 .

The increase in Ab and Qtz in our melts with a drop in pressure is consistent with experimental findings at decreasing crustal pressure (Patiño Douce and Beard 1995; Johannes and Holtz 1996). The change in melt compositions (type b) from 15 kbar pressure (crustal depth) to one atmosphere within a small temperature range of about 100°C (figure 13) can be attributed to a rapid drop in pressure leading to melt equilibration with surface conditions. This leads to suggest that the high temperature Bijli melts may have erupted soon after their generation that inhibited fractionation, consistent with our findings in section 7.2a. The high magmatic temperatures, in turn, lead to the enrichment in HFSE, characteristic of A-type granitic melts including the Bijli Rhyolite, owing to increased solubility of accessory

phases (cf., Collins *et al* 1982). Low H₂O activity ($a_{\text{H}_2\text{O}}$) with decreasing pressure at shallow crustal melting ($P \leq 4$ kbar) causing the melt to be enriched in normative orthoclase at the expense of albite (type c) (Conrad *et al* 1988, Holtz and Johannes 1991) is also consistent with the melting regime considered here. So, there could as well be some contributions of shallower crustal components including the Amgaon Gneiss, if it is the basement to the Bijli Rhyolite (Sarkar 1957–58; Sarkar *et al* 1994). H₂O-undersaturated conditions, however, commonly prevail during partial melting in the middle and lower crust, and basalt-induced crustal melting is conducive to rapid generation of melts (Huppert and Sparks 1988; Annen and Sparks 2002).

7.2d Trace element modeling

An attempt has been made to further constrain the nature of the sources by selected trace element modeling. Models of different extent of melting of two different crustal sources are considered here. Major element data for the Amgaon granitoid rocks are not far from that of the upper continental crust (UCC) of Taylor and McLennan (1985). Since the trace element data for the Amgaon rocks are not available, and the upper continental crust (UCC) of Taylor and McLennan (1985) is believed to be compositionally similar to the Precambrian basement granitoids, the UCC is taken as a potential crustal source, as a first approximation, for modeling.

Melt segregation is a complex process to model in detail (Spiegelman and Kenyon 1992). The silicic melts have a higher permeability threshold, and thus probably behave according to the batch melting equation (Shaw 1970) as follows:

$$C_L/C_0 = 1/F + D - FD \quad (1)$$

where F = fraction of partial melting, D = bulk distribution coefficient, C_0 = concentration of element in original source, C_L = concentration of element in final melt

The normative minerals of the respective source compositions are considered as participating phases for the distribution of trace elements into the melt fraction. It is always a problem to choose the correct mineral/melt partition coefficients for individual elements in different phases. Besides, comparative data of partition coefficients in silicic systems are scarce, and partition coefficients for minerals from high-silica rhyolites differ substantially from those of less silicic magmas. Nevertheless, partition coefficient data for individual elements in highly silicic melts (e.g., Mahood

Table 6. Partition co-efficient data used for calculations.

	qtz	k-feld	plag	cpx	opx	ilm	Apt
K	0.013	0.00	0.263	0.037	0.002	0.00	0.00
Sc	0.01	0.02	0.06	50	18	5.90	0.00
Ti	0.04	0.00	0.00	0.00	0.00	0.00	0.10
Sr	0.00	7.30	13	0.516	0.009	0.00	0.00
Rb	0.015	1.80	0.06	0.032	0.003	0.00	0.00
Th	0.01	0.02	0.05	0.13	0.15	7.50	0.00
La	0.012	0.07	0.34	0.87	0.84	7.10	14.5
Ce	0.017	0.02	0.24	1.40	0.99	7.80	21.1
Sm	0.017	0.02	0.15	3.80	1.70	6.90	46
Nd	0.024	0.04	0.19	2.40	1.30	7.60	32.8
Eu	0.073	6.50	7.90	3.30	0.81	2.50	26
Yb	0.025	0.03	0.10	4.0	2.4	4.1	15.4
Lu	0.024	0.04	0.13	3.80	2.40	3.60	13.8
Ba	0.06	24	3.3	0.131	0.003	0.00	0.00
Ta	0.008	0.01	0.03	0.13	0.20	106	0.00
U	0.04	0.04	0.13	0.00	0.17	3.20	0.00
Hf	0.057	0.04	0.29	0.56	0.22	3.10	0.10
Cr	0.00	0.00	0.00	0.00	0.00	3.00	0.00

qtz = quartz, k-felds = k-feldspar, plag = plagioclase, cpx = clino-pyroxene, opx = ortho-pyroxene, ilm = ilmenite, Apt = apatite.

Data taken from Mahood and Hildreth (1983), Nash and Crecraft (1985). Rb, Sr, K, Ba for cpx and opx from Arth (1976).

and Hildreth 1983; Nash and Crecraft 1985) are used to calculate the bulk distribution coefficient for each element for the Bijli melts. The partition coefficient data of Nash and Crecraft (1985) have some advantages over those of Mahood and Hildreth (1983), for no partition data are available in the latter for plagioclase, clino- and orthopyroxenes separately for the same eruptive unit. REE partition coefficients of Nash and Crecraft (1985) are, however, very similar to that reported by Mahood and Hildreth (1983) and Arth (1976). With these limitations, the same set of partition coefficient values (table 6) are used for all the calculations irrespective of the sources considered here.

Since the modal abundance of REE, Th, U, Y-bearing accessory minerals like Zircon, monazite, Th-orthosilicate in AAG and UCC are not known, these could not be taken as participating phases. Moreover, consideration of these phases would have complicated the situation as they do not generally obey Henry's law during melt-solid partitioning, their concentrations in partial melts not being ruled by crystal-melt distribution coefficients but by solubility relations and reaction kinetics (Rapp and Watson 1986).

7.2e Results

Geochemical mass balance (table 7) shows that trace element data fits better with 25–35% partial melting of AAG rather than a source of UCC composition of Taylor and McLennan (1985), particularly in terms of Sr, Rb, Th, La, Nd, Lu, Sc and Hf. For some incompatible elements, e.g., Ba, K, Ti, Ta, U, Sm, Yb, however, the extent of melting are more uncertain, which could partly be attributed to compositional heterogeneity of crustal protolith including the depth of melting in the crust.

Given these uncertainties, partial melting (~ 30%) of deeper crustal rocks similar to the AAG may better explain the data. The occurrence of charnockite patches, lenticular bodies and enclaves within the granitic rocks of the Archaean complex in south and north Bastar (Bhate and Krishna Rao 1981) also supports the model. Similar Ta/Th ratios of the Bijli Rhyolite (0.060) and of average granulite (0.065), but higher values for UCC (0.13) (Sensarma *et al* 2002) also points to a deeper crustal source. However, limited role of the UCC including the Amgaon Gneiss cannot be ruled out, as some trace elements fit quite well with those of UCC values (table 7).

LILE depletion, particularly depletion of U compared to Th has been generally assumed to be a common feature of granulites (Rudnick *et al* 1985). We do not consistently find U depletion in our rocks. So, the deeper crustal source in Dongargarh was probably melt-undepleted in character. Melt-undepleted character of granulites, though not very common, is also reported from the southern Indian granulites (Weaver 1980; Janardhan *et al* 1983).

8. Crust-mantle interaction

It is in this scenario that the question of whether the komatiitic basalts provided only heat for crustal melting, or also interacted chemically with it, needs to be addressed. This assumes significance as recent experimental studies confirm reaction of basaltic magmas with quartzofeldspathic rocks yielding silicic magmas (Patiño Douce 1999 and references therein).

8.1 Constraints from major elements

Interaction of mantle-derived melts and crustal rocks in terms of major elements should involve alkali elements, Al_2O_3 , CaO, FeO, MgO and TiO_2 (Patiño Douce 1999). We plot the major element compositions of crustal melts produced by experimental dehydration melting of biotite gneiss and quartz-amphibolites at 3–15 kbar (Patiño Douce and Beard 1995), and the Bijli rocks as a ratio

between two variables versus the sum of the variables. The Bijli melts yield a tight trend that passes close to, but not exactly within the fields of crustal melts (figure 14). It is interesting that the basaltic komatiite/high-Mg basalt samples reported from the Dongargarh belt (Sensarma 2001) plot at one end of this trend. The Bijli rocks have two distinct sub-trends:

- constant $(\text{Na}_2\text{O} + \text{K}_2\text{O})/(\text{FeO} + \text{MgO} + \text{TiO}_2)$ and $\text{Al}_2\text{O}_3/(\text{FeO} + \text{MgO} + \text{TiO}_2)$, but wide range in $\text{Na}_2\text{O} + \text{K}_2\text{O} + \text{FeO} + \text{MgO} + \text{TiO}_2$ and $\text{Al}_2\text{O}_3 + \text{FeO} + \text{MgO} + \text{TiO}_2$, and
- wide range in $(\text{Na}_2\text{O} + \text{K}_2\text{O})/(\text{FeO} + \text{MgO} + \text{TiO}_2)$ and $\text{Al}_2\text{O}_3/(\text{FeO} + \text{MgO} + \text{TiO}_2)$, but constant $\text{Na}_2\text{O} + \text{K}_2\text{O} + \text{FeO} + \text{MgO} + \text{TiO}_2$ and $\text{Al}_2\text{O}_3 + \text{FeO} + \text{MgO} + \text{TiO}_2$.

The Bijli samples with constant $(\text{Na}_2\text{O} + \text{K}_2\text{O})/(\text{FeO} + \text{MgO} + \text{TiO}_2)$ and $\text{Al}_2\text{O}_3/(\text{FeO} + \text{MgO} + \text{TiO}_2)$ have similar values to that of the basaltic komatiites, but slightly lower $\text{Al}_2\text{O}_3/(\text{FeO} + \text{MgO} + \text{TiO}_2)$ compared to the amphibolite considered here (figure 14). Wide range in $\text{Na}_2\text{O} + \text{K}_2\text{O} + \text{FeO} + \text{MgO} + \text{TiO}_2$ and $\text{Al}_2\text{O}_3 + \text{FeO} + \text{MgO} + \text{TiO}_2$ in the same set of samples can be attributed to variable contributions of crust and mantle-derived basaltic komatiite melts into the Bijli liquid with mantle contribution largely in terms of Fe, Mg and Ti. Both AAG and the Amgaon Gneiss plot very close to these samples implying their involvement. The second type of samples, on the other hand, calls for predominant participation of shallow crustal components (≤ 4 kbar) comparatively richer in Na_2O , K_2O , Al_2O_3 , with limited mantle input. We are not certain about the identity of this crustal component in Dongargarh, but low Ca, $\text{CaO} + \text{Al}_2\text{O}_3$ (~ 12) and $\text{CaO}/\text{Al}_2\text{O}_3$ (0.02) of the rocks do not argue for the Amgaon Gneiss, which has higher CaO and Al_2O_3 contents and $\text{CaO}/\text{Al}_2\text{O}_3$ ratios. The compositions of these samples are rather analogous to the experimental partial melt compositions with a calc-alkaline granite protolith at 4 kbar pressure (Patiño Douce 1997). So, variable extents of interaction between mantle-derived basaltic komatiite melts and shallow to deeper crustal components, not only shallow hybridization of basalts and calc-alkaline rocks as suggested for the origin of the rhyolites in some basaltic plateaus (Paraná and Karoo) (Patiño Douce 1999), may have given rise to the hybrid Bijli melts.

8.2 Further constraints from oxygen isotope

Igneous rocks with O-isotope values in the range of +5 to +8‰, as observed in our samples (table 5), are generally considered to have been derived from

Table 7. Partial melting calculations for selected trace elements (in ppm) of the Bijli Rhyolite.

	Melting of AAG				Melting of UCC				Bijli Rhyolite (aver)
	AAG (Proportion of phases (%))	F = 25% Cl (calculated)	F = 30% Cl (calculated)	F = 35% Cl (calculated)	UCC (Proportion of phases (%))	F = 25% Cl (calculated)	F = 30% Cl (calculated)	F = 35% Cl (calculated)	
K	qtz = 16.34	53320	47382	42633	qtz = 16.1	81632	72488	65187	33271
Sc	k-feld = 13.41	2.76	2.91	3.08	k-feld = 20.11	2.85	3.00	3.17	3.19
Ti	plag = 47.48	13385	11206	9637	plag = 46.86	11744	9831	8454	1720
Sr	cpx = 6.21	57.46	60.80	64.54	cpx = 5.83	58.89	62.35	66.24	66.92
Rb	opx = 15.46	125	116	107.83	opx = 10.16	205	194	185	145.5
Th	ill = 0.6	24.37	21.45	19.15	ill = 0.45	34.05	29.73	26.38	23.4
La	Apt = 0.5	67.90	65.05	62.44	Apt = 0.5	53	51	48	70.3
Ce		107	103	100		110	105	100	135.3
Sm		6.44	6.38	6.33		5.63	5.54	5.45	8.89
Nd		50.25	49.12	48.05		38.33	37.15	36.05	52.12
Eu		0.32	0.34	0.36		0.21	0.22	0.23	1.24
Yb		2.62	2.58	2.55		3.05	2.97	2.90	4.31
Lu		0.54	0.54	0.53		0.45	0.43	0.42	0.64
Ba		191	201	212		109	115	122	800
Ta		1.82	1.78	1.75		3.43	3.31	3.19	6.96
U		11.79	10.43	9.36		8.48	7.47	6.67	5.67
Hf		8.7	8.0	7		7.20	6.58	6.06	8.66
Cr		770.4	649.4	561		134.55	113.10	97.55	26.16

qtz = quartz, k-feld = k-feldspar, plag = plagioclase, cpx = clinopyroxene, opx = orthopyroxene, ill = ilmenite, Apt = apatite.
 AAG: Average Archean Granulite (Rudnick and Presper 1990).
 UCC: Upper Continental Crust (Taylor and McLennan 1985).

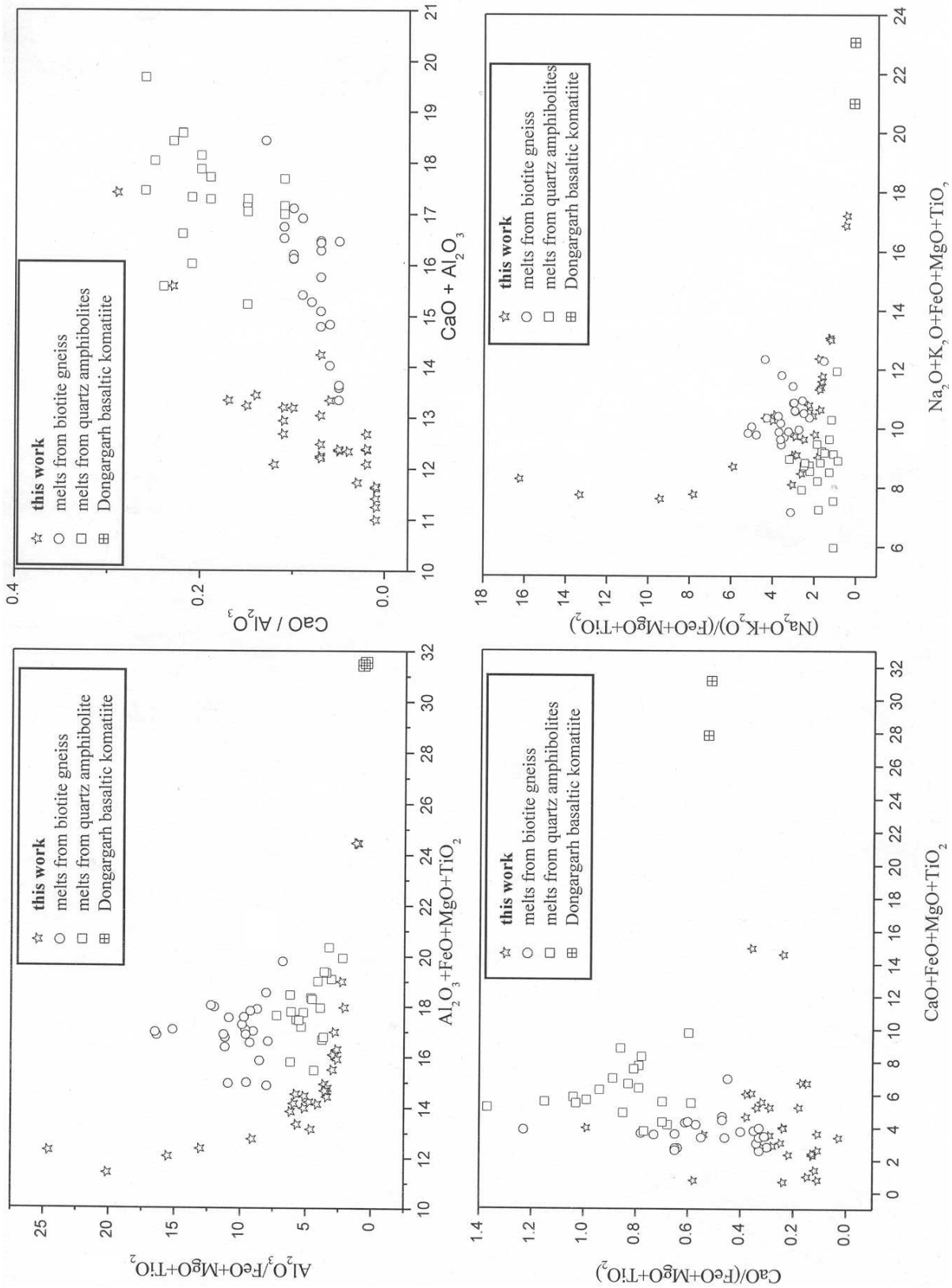


Figure 14. Figure showing plots of the Bijli rocks, experimental crustal melts (Patiño Douce and Beard 1995) and Dongargarh basaltic komatiite (Sensarma 2001) as a ratio between two variables versus the sum of the variables of selected major oxides. The Bijli samples show tight trends with basaltic komatiite as one end member, and passes close to partial crustal melts compositions. $CaO/FeO + MgO + TiO_2$ vs. $CaO + Al_2O_3$ and $CaO/(Al_2O_3 + Al_2O_3)$ vs. $CaO + Al_2O_3$ plots indicate very low Ca and Al_2O_3 contents of the samples.

a source including mafic to intermediate lower crust (granulites) ($8‰ \pm 1.8‰$) and mantle derived materials, but not upper crustal rocks with higher $\delta^{18}\text{O}$ value ($> 13‰$) (Kempton and Harmon 1992). Average K/Rb (300) in our rocks (figure 4) may reflect the presence of some mafic component, as average compositions of common igneous rocks must include rocks of basaltic composition (cf., Rudnick and Presper 1990). Our oxygen isotope data is therefore evaluated in terms of interaction of granulitic crust and mantle-derived melt. Since oxygen constitutes approximately half the mass of the rock, crust-mantle interaction in terms of oxygen isotope should approximate simple mixing process. For example, addition of about 15% of mantle-derived material ($\delta^{18}\text{O}_{\text{mantle}} = +5.7‰$) to a lower crust ($\delta^{18}\text{O}_{\text{lowercrust}} = +8‰$) may give rise to a melt with $\delta^{18}\text{O}$ of about 7.65‰, similar to DM21 value. On the other hand, only 10% crustal contribution would be required to achieve a value of 6.1‰, similar to DM7. So, variable contribution of deeper crust and mantle could thus be responsible for the spread in $\delta^{18}\text{O}$ values in our samples. On the other hand, a range in $\delta^{18}\text{O}$ values from 4.4 to 7.6‰, and low initial $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of 0.7031 and 0.7057 in the samples may also be explained in terms of interaction of a mafic source rock and hydrothermal fluid. Some recent rhyolites from Iceland show such oxygen and Sr isotope signatures and have been explained by involvement of hydrothermally altered oceanic rocks. Irrespective of the possibilities, isotopic signatures in our samples well argue for contributions of mafic components in the origin of the Bijli Rhyolite.

The nature of these chemical interactions is complex, and may involve assimilation, magma mixing, crystal accumulation (Hildreth and Moor bath 1988). It is not the purpose of this paper to elucidate the details, nor the relative importance of all these processes. Nevertheless, the present contribution highlights that crust-mantle interaction was an important process, wherein komatiitic basalts did not only supply heat to partially melt the crust, but contributed material to generate the hybrid silicic liquid. So, basaltic magmas are involved in the origin of the Bijli Rhyolite both physically and chemically. The recent discovery of a Siliceous High Magnesian Basalt (SHMB), which indicates basaltic komatiite and Bijli Rhyolite interaction and is occurring in spatial proximity to the Bijli Rhyolite (Sensarma *et al* 2002), strongly supports the role of crust-mantle interaction in the origin of the Bijli melts. Similar suggestion is also made for producing the Karoo rhyolites (Cox 1988) and the large-volume silicic volcanic centers in the western USA (Johnson 1991). It may happen that fluid introduction partly occurred as a consequence of brittle fracturing of the crust during the period

between basaltic komatiite underplating and melting to produce the rhyolitic magmas.

9. Regional implication

Absence of hydrous phases and low Al_2O_3 content ($\sim 12 \text{ wt}\%$), elevated HFSE, LILE and LREE contents in our rocks reflect its formation in an essentially anhydrous system (e.g., Collins *et al* 1982). An independent study by Bergantz (1989) also emphasized that liquid composition obtained under dry partial melting at $\sim 10 \text{ kbar}$ (crustal depth) more commonly matches the natural melt composition compared to that in saturated conditions. Low aH_2O also requires higher temperature ($\sim 950^\circ\text{C}$) during partial melting (Clemens and Vielzeuf 1987; Johannes and Holtz 1996), which fits better for our samples (see above). Presence of anorthoclase in the rocks also indicates high temperature for the melts. This temperature was attainable in the given geologic situation in Dongargarh where successive high Mg basalts of mantle plume origin overlie the Bijli Rhyolite (Sensarma 2001), as discussed earlier.

The high silica rhyolites in the Dongargarh belt therefore manifest the onset of crustal growth at or near the end Archaean-Palaeoproterozoic in the central Indian Craton, where crustal mobilization at various depth as well as material input from deep mantle were responsible to build the new crust. The mantle-derived basaltic komatiite supplied heat and triggered melting at middle to lower crustal level, which subsequently involved shallower crust as well. The rapid eruption of the high temperature rhyolitic melts not only inhibited fractional crystallization, but is also indicative of a rapid drop in pressure related to decompression in an extensional setting within the crust. The Amgaon Gneiss may or may not be the basement on which the Bijli Rhyolite was laid down, but is not the only source component to the high silica rhyolite melts, as suggested by Sarkar *et al* (1994). Based on field and petrographic studies, we earlier suggested near contemporaneous relation between the Bijli Rhyolite and the immediately overlying Pitepani Volcanics in the area (Sensarma and Mukhopadhyay 2003). The common process of crust-mantle interaction in the origin of the Bijli Rhyolite as well as the SHMB-type lava in the Pitepani Volcanics (Sensarma *et al* 1999, 2002) surely indicates their contemporaneity, and need not be placed across an unconformity (Sarkar 1957; Sarkar *et al* 1994).

10. Concluding remarks

The following conclusions can be drawn from this study:

- The Bijli Rhyolite is rhyolite to high silica rhyolite, and show close similarity to “A-type granite”.
- It is neither a product of fractionation of a basaltic magma, nor is related to fractional crystallization of partial crustal melts. It is a product of crust-mantle interaction where crustal mobilization took place at various crustal depths. Deep crust compositionally similar to the average Archaean granulite of Rudnick and Presper (1990), shallower crust with high alkali, but low CaO, Al₂O₃ contents, the Amgaon Gneiss and the basaltic komatiitic melts/high-Mg basaltic melts had variable contributions in the origin of the hybrid Bijli melts. About 30% partial melting of average Archaean granulite compared to the Amgaon Gneiss better explain the data.
- The Amgaon granitoid gneiss of Sarkar *et al* (1994) may or may not be the basement to the Bijli Rhyolite in the area, but cannot be the sole source for the Bijli melt as advocated by Sarkar *et al* (1994).

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