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Resource availability and disturbance shape maximum tree height across the Amazon

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1 **Abstract**

2 The factors shaping the distribution of giant tropical trees are poorly understood, despite its importance as a link
3 between evolutionary biology and ecosystem biogeochemistry. The recent discovery of clusters of trees over 80
4 metres tall in the Guiana Shield region of the Amazon rainforest challenges the current understanding of the factors
5 controlling the growth and survival of giant trees. The new discovery led us to revisit the question: what determines
6 the distribution of the tallest trees of the Amazon?

7 Here, we used high-resolution airborne LiDAR (Light Detection and Ranging) surveys to measure canopy height

8 across 282,750 ha of primary old-growth and secondary forests throughout the entire Brazilian Amazon to investigate
9 the relationship between the occurrence of giant trees and the environmental factors that influence their growth and
10 survival. Our results suggest that the factors controlling where trees grow extremely tall are distinct from those
11 controlling their longevity. Trees grow taller in areas with high soil clay content ($> 42\%$), lower radiation (< 130
12 clear days per year) and wind speeds, avoiding alluvial areas (elevations higher than 40 m a.s.l), and with an optimal
13 precipitation range of 1,500 to 2,500 mm yr⁻¹. We then used an envelope model to determine the environmental
14 conditions that support the very tallest trees (i.e. over 70 m height). We found that, as opposed to the myriad of
15 interacting factors that control the maximum height at a large scale, wind speed had by far the largest influence on the
16 distribution of these sentinel trees, and explained 67% of the probability of finding trees over 70 m in the Brazilian
17 Amazon forest.

18 The high-resolution pan-Amazon LiDAR data showed that environmental variables that drive growth in height are
19 fundamentally different from environmental variables that support their survival. While precipitation and temperature
20 seem to have lower importance for their survival than expected from previous studies, changes in wind and radiation
21 regimes could reshape our forested biomes. This should be carefully considered by policy-makers when identifying
22 important hotspots for the conservation of biodiversity in the Amazon.

23 **Introduction**

24 The Amazon is the largest rain forest on Earth, covering 5.5 million square kilometres, and stor-
25 ing about 17% of all vegetation carbon. Ecologists have long taken an interest in comparing the
26 structure and composition of rain forests across the tropics (Yang *et al.* 2016), and have reached a
27 consensus that the Amazon supports shorter trees, and therefore stores a lower amount of carbon
28 per hectare, than the forests of tropical Africa and Asia (Cao & Woodward 1998; Feldpausch *et al.*
29 2012). However, the recent discovery of giant trees - up to 88 m tall - in the Amazon basin (Gor-
30 gens *et al.* 2019) challenges this paradigm and poses new questions about the drivers causing the
31 spatial distribution of tall trees in the Amazon.

32 Previous studies have debated the factors which govern Amazon tree growth and have particularly

33 focused on productivity drivers related to the wet and dry seasons (Huete *et al.* 2006; Morton
34 *et al.* 2014). This paper's findings inform this important question but extend the investigation
35 beyond these factors to include the influence of 18 climatic and other environmental conditions on
36 achieving greatest tree height.

37 Tree height is fundamentally linked to growth, survival, and reproduction strategies, and is ulti-
38 mately related to the ability to pre-empt light resources and disperse diaspores (Díaz *et al.* 2016).
39 Xylem conduit diameter and total path length resistance to water flow increase with canopy height,
40 making water transport to higher leaves more difficult (Koch *et al.* 2004; Givnish *et al.* 2014). To
41 counteract this difficulty, taller species have higher xylem hydraulic conductivity but are more vul-
42 nerable to xylem embolism (Liu *et al.* 2019). Across species, higher wood density, and stomata
43 closure in response to water deficit are often positively related to embolism resistance (Bennett *et*
44 *al.* 2015; McDowell & Allen 2015; Greenwood *et al.* 2017). Height growth is partly governed by
45 small-scale factors such as water availability, temperature, rooting depth, and soil type (Anderegg
46 *et al.* 2016; McDowell & Allen 2015; Coomes *et al.* 2006; Niklas 2007), with precipitation and
47 potential evapotranspiration consistently reported as key factors determining plant height across
48 biomes (Moles *et al.* 2009; Larjavaara 2013; Rueda *et al.* 2016).

49 Forest giants are disproportionately vulnerable to disturbances and thus their conservation requires
50 particular attention (Pennisi 2019; Yanoviak *et al.* 2019; Stovall *et al.* 2019; Enquist *et al.* 2020).
51 To reach such immense sizes, trees must fulfill at least three conditions: they must (1) have an
52 evolutionary design that is capable of transporting water to great heights and overcome highly
53 negative water potentials to deliver that water toward tissues in the upper canopy; (2) inhabit an
54 area with optimal environmental conditions (such as climate, soil properties, and water) that meet
55 species-specific requirements (Simard *et al.* 2018; Scheffer *et al.* 2018) and (3) grow in regions
56 with a low frequency of natural or anthropogenic disturbance events (Larjavaara 2013; Linden-
57 mayer & Laurance 2016; Scheffer *et al.* 2018; Enquist *et al.* 2020). Resource availability (e.g.

58 sunlight, nutrients, CO₂, and water) controls a tree's ability to produce biomass through photosyn-
59 thesis. Natural disturbances (e.g. windthrow, drought, or lightning) and history of anthropogenic
60 actions (e.g. selective logging, forest fragmentation) increase the likelihood of mortality and limit
61 the time available to trees to grow taller (Bennett *et al.* 2015; Powers *et al.* 2020; Yanoviak *et*
62 *al.* 2019; Almeida *et al.* 2019). Tall trees are likely to have developed strategies for surviving
63 diseases and pathogens (van Gelder *et al.* 2006; Aleixo *et al.* 2019) as well as climatic fluctu-
64 ations (Sakschewski *et al.* 2016) and resisting wind damage (Jagels *et al.* 2018). However, the
65 question of how resource supply and disturbances interact to determine canopy height across the
66 Amazon has not been fully explored.

67 The sheer size of the Amazon, its environmental heterogeneity and species diversity, pose chal-
68 lenges and practical difficulties to understand general ecological relationships and biogeographical
69 patterns (Tuomisto *et al.* 2019). Forest plots provide many valuable insights to investigate the
70 influences of the environment on tree height but they can only represent a minuscule fraction of
71 the total forest area (Chave *et al.* 2020). Currently, a network of 5,351 forest inventory plots
72 established across the Brazilian Amazon, of known and published sites recently compiled by (Te-
73 jada *et al.* 2019), represents only 0.0013% of the total forest area in this region. In addition, the
74 plot distribution is spatially clustered in close proximity to major roads or large rivers (Stropp *et*
75 *al.* 2020), implying a spatial distribution bias (Marvin *et al.* 2014) since about 42% of the total
76 Brazilian Amazon lies over 50 km from the nearest forest inventory plots (Tejada *et al.* 2019). Re-
77 mote sensing can remove sampling biases and uncertainty about ecological patterns (Schimel *et al.*
78 2015) and provides large datasets with which to uncover the environment controls of forest struc-
79 ture (Asner *et al.* 2010). In particular airborne LiDAR (Light Detection and Ranging) generates
80 valuable high-resolution 3D information of forest canopy structure (Görgens *et al.* 2016; Coomes
81 *et al.* 2017), and can be used as an intermediary to integrate field data with satellite sources (Asner
82 2009; Bae *et al.* 2019).

83 The “Improving Biomass Estimation Methods for the Amazon” (EBA) project performed high-
84 resolution LiDAR flights over 3 years, totaling more than 800 transects across mature and sec-
85 ondary forests in the Brazilian Amazon (for more information about the EBA project see the
86 Method section). The transects were randomly distributed considering both spatial location of the
87 start point and flight direction, allowing us to conduct statistical design-based models of tree height
88 since there is a growing consensus that ecosystem traits like tree height can be better measured with
89 LiDAR than other methods (Valbuena et al. 2020). A total of 282,750 ha were covered, 0.183% of
90 the Brazilian Amazon, which is 100 times the area of available permanent plot networks (Tejada
91 *et al.* 2019). This unprecedented dataset has led to remarkable discoveries in Amazon (Gorgens *et*
92 *al.* 2019; Pereira *et al.* 2019; Santos *et al.* 2019; Almeida *et al.* 2019). In this study, we employed
93 it to contribute to our understanding of how resources and disturbances shape the maximum height
94 distribution across the Brazilian Amazon. We conducted an extensive analysis relating remotely
95 sensed environmental variables to the maximum height recorded in the transects. We concluded
96 that drivers of height development are fundamentally different from those influencing the survival
97 of tree giants. Thus changes in wind and light availability shape their distribution as much as
98 precipitation and temperature, altogether shaping the demographics and composition of forested
99 biomes.

100 **Material and methods**

101 Between 2016 and 2018, an airborne mission (held by National Institute for Space Research -
102 INPE and funded by Amazon Fund) collected airborne LiDAR data from 906 transects of 375
103 ha (12.5 x 0.3 km) each, randomly spread across primary and secondary forests defined by the
104 PRODES database - layer mask of primary old-growth forests (**PRODES, INPE, 2016**) and by the
105 TerraClass database - a layer mask of secondary forest (**TerraClass, INPE, 2014**).

106 The LiDAR sensor was the Trimble Harrier 68i (Trimble, California, USA) aboard a Cessna 206

107 aircraft. The average pulse density was set at 4 pulses m^{-2} , the field of view equal to 30° , and
108 flying altitude of 600 m. The Global Navigation Satellite System (GNSS) collected data on a dual-
109 frequency receiver (L1/L2). The pulse footprint was set to be below 30 cm, based on a divergence
110 angle between 0.1 and 0.3 mrad. Horizontal accuracy was controlled to be under 1 m, and the
111 vertical accuracy to be under 0.5 m.

112 Details about LiDAR parameterization, processing, and the EBA project characteristics can be
113 consulted in (Gorgens *et al.* 2019). Briefly, each transect was processed by identifying the re-
114 turns backscattered from the ground and interpolating a 1m spatial resolution digital terrain model
115 (DTM) from them. Then, the DTM was employed to calculate the heights above ground from the
116 returns backscattered from the vegetation (Görgens *et al.* 2016). The uppermost vegetation heights
117 were then employed to compute a canopy height model CHM at the same spatial resolution as
118 the DTM. The height of the tallest tree per transect was identified from the CHM using a local
119 maximum moving window algorithm (Dalponte & Coomes 2016). All transects were finally ma-
120 nually inspected to exclude non-trees maximum derived from artifacts, ensuring that all the largest
121 heights indeed depicted a tall tree.

122 **Environmental variables**

123 To investigate drivers influencing the spatial distribution of giant trees, we initially considered a
124 total of 18 environmental variables: (1) fraction of absorbed photosynthetically active radiation
125 (FAPAR; in %); (2) elevation above sea level (Elevation; in m); (3) the component of the horizon-
126 tal wind towards east, i.e. zonal velocity (u-speed ; in m s^{-1}); (4) the component of the horizontal
127 wind towards north, i.e. meridional velocity (v-speed ; in m s^{-1}); (5) the number of days not affected
128 by cloud cover (clear days; in days yr^{-1}); (6) the number of days with precipitation above 20 mm
129 (days $> 20\text{mm}$; in days yr^{-1}); (7) the number of months with precipitation below 100 mm (months
130 $< 100\text{mm}$; in months yr^{-1}); (8) lightning frequency (flashes rate); (9) annual precipitation (in

131 mm); (10) potential evapotranspiration (in mm); (11) coefficient of variation of precipitation (pre-
132 cipitation seasonality; in %); (12) amount of precipitation on the wettest month (precip. wettest; in
133 mm); (13) amount of precipitation on the driest month (precip. driest; in mm); (14) mean annual
134 temperature (in °C); (15) standard deviation of temperature (temp. seasonality; in °C); (16) annual
135 maximum temperature (in °C); (17) soil clay content (in %); and (18) soil water content (in %).

136 The FAPAR was derived from land surface reflectance product calibrated and corrected from the
137 National Oceanic and Atmospheric Administration's (NOAA) Advanced Very High-Resolution
138 Radiometer (AVHRR), which is a consistent time-series dataset spanning from the mid-1980s to
139 present and suitable for climate studies (Tao *et al.* 2016). FAPAR is a primary vegetation variable
140 controlling the photosynthetic activity of plants and is considered an essential climate variable (Ma-
141 son *et al.* 2010). The algorithm to create this layer relies on artificial neural networks calibrated
142 with the MODIS FAPAR dataset and validated using a set of globally-distributed sites. The inputs
143 to generate the FAPAR were 1) the surface reflectance from NOAA-AVHRR which is measured
144 in two wavelengths (red, 580–680 nm, and near-infrared, 725–1000 nm); 2) a reference dataset
145 from MODIS FAPAR to calibrate the NOAA-AVHRR FAPAR; and 3) a land cover map classi-
146 fication, used to stratify the outputs. Five land cover classes were included: evergreen broadleaf
147 forest, deciduous broadleaf forest, needle leaf forest, shrubland, croplands and grasslands, and
148 non-vegetated (Claverie *et al.* 2016).

149 The elevation was computed based on the third version of the Shuttle Radar Topography Mission
150 (SRTM) provided by National Aeronautics and Space Administration Jet Propulsion Lab (NASA
151 JPL) (Farr *et al.* 2007; Liu *et al.* 2014). The SRTM mission was launched on Space Shuttle Endeav-
152 or on 11th February 2000 and collected data during ten days of operations, using two synthetic
153 aperture radars: NASA's C band system (5.6 cm wavelength) and an X band system by DLR (3.1
154 cm). The C-band digital elevation model (DEM) used in this study is now available at 30-m spati-
155 al resolution from 60° north latitude and 56° south latitude, covering 80% of Earth's land surface.

156 C-band partially penetrates the vegetation canopy, with depth varying with vegetation structure.
157 Since Amazonian vegetation is dense throughout, for the purposes of this study the C-band DEM
158 is assumed to vary consistently with topography across the region.

159 We used the maximum daily mean wind speeds over the last 5 years from the fifth major glo-
160 bal reanalysis (ERA5) produced by the European Centre for Medium-Range Weather Forecasts
161 (ECMWF). The reanalysis combined model data with observations from across the world into a
162 globally complete and consistent dataset (Olauson 2018). The products from a reanalysis include
163 many variables such as wind speeds, temperature, and atmospheric pressure. They were produced
164 on reduced Gaussian grids, by using a different number of grid points along different latitudes
165 and thus keeping the grid point separation in metres approximately constant. ERA5 has an hourly
166 resolution and spans from 1950 to near real-time. Two wind velocities were considered: u-speed
167 which is the zonal velocity (i.e. the component of the horizontal wind towards east), and v-speed
168 which is the meridional velocity (i.e. the component of the horizontal wind towards north). These
169 products are used extensively for modelling wind power both in academia and industry (Olauson
170 2018; Albergel *et al.* 2019; Ramon *et al.* 2019).

171 The number of clear days was computed based on Moderate Resolution Imaging Spectroradio-
172 meter (MODIS) surface reflectance products. MODIS products provide an estimate of the surface
173 spectral reflectance as it would be measured at ground level in the absence of atmospheric scat-
174 tering or absorption (Kang *et al.* 2005; Bisht & Bras 2010). MODIS operates onboard Terra and
175 Aqua satellites. Terra satellite has a 10:30 am equator over-passing time, and the $\pm 55^\circ$ scanning
176 pattern at 705 km altitude achieves a 2,330 km swath that provides global coverage every one to
177 two days (Ruhoff *et al.* 2012). We used the Terra MOD09GA Version 6 product, which provides
178 an estimate of the surface spectral reflectance of MODIS, corrected for atmospheric conditions
179 such as gases, aerosols, and Rayleigh scattering.

180 Temperature and precipitation were obtained from the WorldClim database of bioclimatic varia-

181 bles, which are derived from weather station data compiled for the 1950–2000 period ([Hijmans *et*](#)
182 [al. 2005](#); [Fick & Hijmans 2017](#)). The main source of data was the Global Historical Climatolo-
183 gy Network (GHCN), complemented with other global, national, regional, and local data sources,
184 which were added if they were further than 5 km away from stations already included in the GH-
185 CN. After removing stations with errors, the final database consisted of precipitation records from
186 47,554 locations, mean temperature from 24,542 locations, and minimum and maximum tem-
187 peratures from 14,835 locations. To interpolate the weather station data, latitude, longitude, and
188 elevation were used as independent variables.

189 The lightning frequency was provided by Lightning Imaging Sensor (LIS) instrument onboard
190 the Tropical Rainfall Measuring Mission provided by NASA Earth Observing System Data and
191 Information System (EOSDIS) Global Hydrology Resource Center. The LIS was launched in No-
192 vember 1997 into a precessing orbit inclination of 35° at an altitude of 350 km and was powered off
193 in April 2015. The LIS datasets were collected during 16 years (1998–2013) and they are available
194 at 0.1° spatial resolution (approx. 11 km in the equator). The LIS provided the basis for the develop-
195 ment of a comprehensive global thunderstorm and lightning climatology to detect the distribution
196 and variability of total lightning occurring in the Earth. This information is used for severe storm
197 detection and analysis, and also for lightning-atmosphere interaction studies ([Albrecht *et al.* 2016](#)).

198 The potential evapotranspiration was provided by the TerraClimate dataset, a global monthly cli-
199 mate and water balance for terrestrial surfaces spanning 1958–2015. The layer used climatically ai-
200 ded interpolation (bilinear interpolation of temporal anomalies), combining high-spatial-resolution
201 climatological normals from WorldClim with Climate Research Unit (CRU) Ts4.0 and the Japa-
202 nese 55-year Reanalysis (JRA-55) data. The CRU Ts4.0 provides monthly average maximum and
203 minimum temperature, vapor pressure, and cumulative precipitation from 1901–2015. The JRA-
204 55 is the longest-running (1958–present) observing-system and provides spatially and temporally
205 complete data for mean temperature, vapor pressure, wind speed, downward shortwave flux at the

206 surface, and accumulated monthly precipitation. The Reference Evapotranspiration was calculated
207 using the Penman-Monteith approach ([Abatzoglou *et al.* 2018](#)).

208 The number of months per year with precipitation below 100 mm and the number of days per year
209 with precipitation above 20 mm was computed based on the Climate Hazards Group InfraRed
210 Precipitation with Station data (CHIRPS) dataset. CHIRPS incorporated 0.05° resolution satellite
211 imagery with in-situ station data to create gridded rainfall time series for trend analysis and seasonal
212 drought monitoring ([Funk *et al.* 2015](#)). The CHIRPS process involves three main components:
213 1) the Climate Hazards group Precipitation climatology (CHPclim), 2) the satellite-only Climate
214 Hazards group Infrared Precipitation (CHIRP), and 3) the station blending procedure that produces
215 the CHIRPS. Two sets of monthly historical long-term means were used to create the CHPclim.
216 The first set was a collection of 27,453 monthly stations obtained from the Agromet Group of the
217 Food and Agriculture Organization of the United Nations (FAO). The second set of 20,591 stations
218 was taken from version two of the Global Historical Climate Network (GHCN). The CHIRP
219 relies on two global thermal infrared archives that are: the 1981–2008 Globally Gridded Center Satellite
220 (GriSat) produced by NOAA’s National Climate Data and the 2000-present dataset NOAA
221 Climate Prediction Center. The CHIRPS station datasets were obtained from the GHCN monthly,
222 GHCN daily, Global Summary of the Day, Global Telecommunication System and Southern
223 African Science Service Centre for Climate Change and Adaptive Land Management. The station
224 blending procedure that produces CHIRPS is a modified inverse distance weighting algorithm.
225 Daily CHIRPS are then produced for the globe by using daily Cold Cloud Duration (CCD) data
226 to identify non-precipitating days. Whenever the daily CCD is zero, precipitation is assumed to be
227 zero.

228 Edaphic variables were obtained from The OpenLandMap produced by the OpenGeoHub Foundation
229 and contributing organizations. Soil texture is connected with soil granulometry or the composition
230 of the particle sizes (clay, silt, and sand), typically measured as volume percentages. The

231 clay content (fine particles $< 2 \mu\text{m}$) and water content layers, both with a spatial resolution of 250
232 m, were created based on machine learning predictions from a global compilation of soil profiles
233 and samples (Arsanjani *et al.* 2014).

234 To help visualization of the regional-level, we divided the Brazilian Amazon into eight regions,
235 according to the classification of (Morrone 2014): I - Para; II - Xingu-Tapajos; III - Roraima; IV -
236 Guianan Lowlands; V - Madeira; VI - Yungas; VII - Pantepui; VIII - Imeri. This regionalization is
237 based on biogeographic analyses of terrestrial plant and animal taxa of the Neotropical region and
238 seeks to provide universality, objectivity, and stability, such that it can be applied when describing
239 distributional areas of particular taxa or comparing different biogeographic analyses.

240 **Random Forest and Maximum Entropy**

241 To better understand the environmental requirements for development in tree height, we employed
242 Random Forest modelling and marginal plots to observe the relative variable importance. Among
243 the initial 18 environmental variables, two of them (precipitation on driest month and months $<$
244 100mm) were excluded due to high correlation (> 0.80) to other independent variables (Table 1).
245 Using the coordinates of the tallest tree within each lidar transect, we extracted the values from the
246 variable layers. Tree height was then modeled against the factors using a random forest algorithm,
247 which recursively computes classification and regression trees (CART) from random subsets, a
248 k-fold ($k = 15$) cross-validation method, and 500 as the number of CART. The number of variables
249 randomly sampled as candidates at each split was set to 10. The adjusted model was evaluated
250 considering the mean absolute error (MAE), root mean squared error (RMSE), and coefficient of
251 determination (R^2) of cross-validated predicted versus observed values. To assess the overall rela-
252 tive variable importance we used the mean decrease in Gini importance, which evaluates at each
253 split in each tree, how much each variable contributes to decreasing the weighted impurity (i.e.,
254 variance in the case of regression trees). The resulting Random Forest model was finally imple-

255 mented using the environmental variables to deliver a map of estimated heights of tallest trees
256 across the Amazon. Then we focused on the tallest trees only - those over 70 m in height - to
257 determine the conditions which allow them to survive. We employed a maximum entropy en-
258 velope approach (MaxEnt) commonly applied to modelling species geographic distributions with
259 presence-only data and indicate better discrimination of suitable versus unsuitable areas for the
260 species (Phillips *et al.* 2006). The variable importance of the MaxEnt model was used to indicate
261 the most relevant characteristics sustaining extreme height individuals and the potential location
262 for new occurrence. The observations higher than 75 m were filtered out and used to adjust an
263 envelope model based on maximum entropy. In its optimization routine, the algorithm tracked
264 how much the model gain was improved when small changes were made to each coefficient value
265 associated with a particular variable. Each variable was then ranked based on the proportion of
266 all contributions. The resulting MaxEnt model was finally implemented using the environmental
267 variables to deliver a map of probability of occurrence for trees taller than 70 m across the Amazon.

268 **Results**

269 Trees exceeded 50 m in height across many parts of the Brazilian Amazon, but trees over 80 m were
270 only observed in the eastern Amazon (micro-region III, Roraima Province; Fig. 1). To examine
271 why the trees grow taller in some regions and determine the environmental variables modulating
272 height pattern in the Amazon, we predicted maximum tree height as a function of environmental
273 variables using a Random Forest approach. The number of clear days, clay content in the soil,
274 elevation and mean annual precipitation were found to be the strongest drivers of maximum tree
275 height, while the average monthly temperature and soil water content were weak predictors (Ta-
276 ble 2). The Random Forest model obtained MAE = 3.62 m, RMSE = 4.92 m, and $R^2 = 0.735$. A
277 resulting map of Random Forest model predicted maximum tree height shows that occurrence
278 is highest in eastern Amazon (Fig. 2), with the tallest trees more specifically achieving greatest

279 heights in the northeastern part of Roraima (III), in Pantepui (VII) and in the confluence of Ma-
280 deira (V) and Xingu-Tapajos (II). Since low values of FAPAR are related to degraded forests and
281 anthropogenic regions, we performed the same analysis after excluding areas with FAPAR values
282 under 80%, which resulted in the elimination of 133 transects. Similar spatial distributions for ma-
283 ximum tree height persisted similarly after removing these potential anthropogenic effects (Fig. 3),
284 demonstrating that the underlying patterns we report are naturally driven by the environmental fac-
285 tors.

286 The marginal plot obtained for each environmental variable in the random forest model, allows
287 us to interpret its influence on the height of tall trees directly in the units that correspond to each
288 (Fig. 4). Lines close to horizontal indicate a given environmental factor having little effect on the
289 height of tall trees. The number of clear days was the strongest predictor of maximum height
290 (Table 2). The shape of this relationship resembles a step function (Fig. 4), in which regions with
291 the number of clear days below 130 days per year support tall trees, above this level, we observe
292 an abrupt decline in maximum height. Elevation was also a key predictor of tree height, with low-
293 lying forests growing 7 m lower than trees in terrains above 40 m above sea level. An increase
294 in soil clay content from 20% to 40% translated into a 7 m increase in maximum height. Our
295 results also demonstrate mean annual precipitation as a key factor for trees to grow taller, with
296 a tolerance curve peaking at around 2,300 mm yr⁻¹ as optimal annual precipitation across the
297 Brazilian Amazon. In comparison to these areas, we observe a 4 m decline in maximum tree
298 height in regions with annual precipitation below 1,500 mm yr⁻¹ or above 3,000 mm yr⁻¹. From the
299 intermediate importance variables, we highlight the zonal velocity (u-speed) and FPAR influencing
300 height variation in ranges around 6 m.

301 The results of the MaxEnt approach are focusing on the survival of trees taller than 70 m in height
302 (Fig. 5). The extraordinarily tall trees had a unique niche, characterized by a much smaller set
303 of environmental variables from those which drove the large-scale patterns of maximum height.

304 The maximum entropy model shows that the niche is dominated mostly by wind speed (relative
305 importance of 67.7 %). The second most important driver of tall tree survival was the elevation
306 above sea level (relative importance of 12.3 %). It is worthwhile noting that relative importance
307 values reflect the proportion of all contributions to explain the presence of the tallest trees. The
308 resulting map of predicted occurrence of the tallest trees in the Amazon from the MaxEnt model
309 shows that the probability of maximum tree height occurrence is highest in northeastern Amazon
310 (Fig. 6), more specifically in the Roraima (III) and Guianan Lowlands (IV).

311 Discussion

312 The locations of the tall trees in the eastern and southern Amazon coincide with forests that have
313 a high basal area predicted by statistical modelling of permanent plot data (Malhi *et al.* 2006).
314 The basal area generally declines with increasing dry season length, for regions with dry seasons
315 lasting four months or longer. Young soils nearer the Andes, as well as the sedimented and flooded
316 lowlands, are richer in nutrients, thereby supporting fast-growing, low wood density species with
317 high turnover rates and, as a result, the trees do not reach extremely large sizes (Marra *et al.* 2014;
318 Quesada *et al.* 2011; Phillips *et al.* 2004). The species *Dinizia excelsa* (Ducke), for example,
319 has been reported as the tallest trees in the Amazon reaching 88 m in height in the region of
320 Roraima (Gorgens *et al.* 2019), and also has been reported as the highest average species-level
321 wood density in the Amazon of 0.94 g cm⁻³ (Fauset *et al.* 2015) with a large contribution to the
322 total forest biomass.

323 Many physiological and structural traits in the Amazon have strong phylogenetic associations with
324 effects on tree growth and mortality (Baker *et al.* 2004; Fyllas *et al.* 2012; Patiño *et al.* 2012).
325 Forests of the western Amazon are more homogeneous in composition at the family (Myristi-
326 caceae, Arecaceae, Moraceae) and species levels (Condit 2002; Pitman *et al.* 2001), while species
327 from eastern Amazonian have broadly different patterns of family-level composition being dom-

328 inated by the Sapotaceae, Chrysobalanaceae, Fabaceae and Lecythidaceae (Chave *et al.* 2006).
329 Wood density is driven by shifts in tree species composition (Terborgh & Andresen 1998) and tends
330 to peak in the slow-growing forests on infertile soils in eastern Amazon and the Guyanas (Malhi
331 *et al.* 2006). Soil physical properties in combination with limited nutrient supply in eastern Ama-
332 zon favour slow-growing species and increases species that invest their resources in structures
333 that can support taller and bigger trees with a long lifespan (Malhi *et al.* 2004; Quesada *et al.*
334 2009). Temperature and dry season precipitation effects on the structure and wood density are
335 more ambiguous (Quesada *et al.* 2012), although species with higher wood density are better able
336 to resist drought-induced embolism (Hacke *et al.* 2001) and therefore tolerate longer periods of
337 high vapor water deficit and evaporative demand (McDowell *et al.* 2018). This myriad of environ-
338 mental variables with confounding effects on species composition, as well as on their physiological
339 and structural traits, play a crucial role in the tree lifespan and the size of trees (Muller-Landau
340 2004).

341 **Conditions supporting tall trees**

342 In our study, the low wind speed was determined as the single most important predictor of the
343 occurrence of the tallest trees in the Brazilian Amazon. The fact that trees adapt to their wind en-
344 vironment and are shorter in windy locations has been widely observed in temperate regions (Bon-
345 nesoeur *et al.* 2016; Telewski 2006). We can see a similar effect across the Amazon, with trees over
346 70 m tall having a 50-75% likelihood of surviving in the calmest areas but a sharply decreasing
347 probability with stronger winds. This agrees with previous findings that disturbance rates are far
348 higher in the Western Amazon (Espírito-Santo *et al.* 2014) and may demonstrate how significant
349 the role of wind is in shaping the niche for extraordinarily tall trees. The importance of wind speed
350 was also apparent in the Random Forest model which showed a 9 m reduction in the estimated
351 tree height from the calmest to the windiest areas (Figure 2). The zonal velocity (i.e. the eastward

352 component), which is the prevailing wind direction in the region, drives this pattern. Interestingly,
353 our data showed that the lightning rate was only weakly related to maximum forest height patterns
354 in both the Random Forests and MaxEnt models. Despite being relevant to the death of individual
355 trees (Marra *et al.* 2014; Bonnesoeur *et al.* 2016; Niklas 1998) and being the key factor causing
356 tree deaths in tropical forests of Panama (Yanoviak *et al.* 2019), lightning and storms do not seem
357 to impact the potential dominant tree of a region, nor to limit the survival of the tallest trees, in
358 light of our results.

359 A balance between tree structural strength and wind shearing forces contributes to set an upper
360 limit to tree height development (Klein *et al.* 2015). The wind has a direct effect on tree height,
361 since trees adapt their growth rates to their local wind environment, although the scale of this effect
362 is unknown (Telewski 2006; Bonnesoeur *et al.* 2016). Extreme wind speeds, often associated
363 with convective storms in the tropics, can also snap or uproot trees. Large-scale wind patterns
364 in the Amazon are dominated by the easterly trade winds. Wind damage is most common from
365 September to February (Negrón-Juárez *et al.* 2017) and taller trees have higher rates of mortality
366 in wind storms (Rifai *et al.* 2016). Remote sensing analyses have shown that disturbance rates are
367 much higher in the western Amazon compared to the east (Espírito-Santo *et al.* 2014).

368 A decrease in cloud-free days goes together with an increase in solar radiation (Barkhordarian
369 *et al.* 2019), which, along with changes in the Vapor Pressure Deficit, or atmospheric dryness,
370 drive changes in the physiological function of trees (Williams *et al.* 2012; Nunes *et al.* 2019).
371 The increase in diffuse radiation led by cloudy conditions induces an increase in photosynthetic
372 activity (Gu 2003). Tree responses to direct solar radiation are dependent on the species and
373 developmental stage, with physiological and structural changes to maximize either growth or sur-
374 vival (Wright *et al.* 2004; Nunes *et al.* 2019; Poorter & Bongers 2006). As the traits of individual
375 trees are at least conserved at the species level, additional variation is determined by the local en-
376 vironment (Fyllas *et al.* 2009). As trees grow taller, increasing leaf water stress due to gravity and

377 path length resistance may limit leaf expansion and photosynthesis for further height growth (Koch
378 *et al.* 2004). Tall trees have direct exposure to sunlight and high temperatures lead to higher stom-
379 atal control to avoid excessive water loss (Drake *et al.* 2018; Rowland *et al.* 2015).

380 Elevation was also a key predictor of tree height, with low-lying forests growing potentially less
381 than trees in terrains over 40 m a.s.l. (Fig. 4). The topographic gradient is likely to be related to the
382 likelihood of flooding in the low elevation transects on the lowlands. Rivers erode the *terra firme*
383 terraces and create floodplains of variable sizes dating to the Miocene, with terrace–floodplain ele-
384 vation differences decreasing eastwards from the Andes (Hamilton *et al.* 2007). Shifts in multiple
385 canopy chemical traits between the terrace and floodplain forests in the Amazon are paralleled by
386 species turnover, which reveals the micro-topography effects on the growth-defense trade-off in
387 Amazonian forests, and its associated processes of nutrient mobilization and deposition (Asner *et*
388 *al.* 2015). The species and trait shifts with topographical variation in the Amazon also confers
389 an adaptive drought resistance, with species from the plateaus more susceptible to prolonged peri-
390 ods with lower soil water content, and, therefore, investing in higher hydraulic safety with higher
391 wood density, lower mean vessel hydraulic diameter, lower mean vessel area and smaller stem
392 cross-sectional sapwood area than species in valley forests (Cosme *et al.* 2017).

393 An increase in soil clay content also translated into an increase in maximum height. Clay content
394 is usually highest on flat terrain (Laurance *et al.* 1999) decreasing from about 75% to 5% when
395 moving from the plateau areas to the valleys (Ferraz *et al.* 1998; Toledo *et al.* 2016). Previous
396 studies also indicated the presence of clayey soil in the plateau areas of the Amazon (Broedel *et*
397 *al.* 2017; Cerri & Volkoff 1987; Marques *et al.* 2002; Marques *et al.* 2004; Marques *et al.* 2015).
398 Our results suggest that 1) if the clayey soils of our study occur in the plateau areas with lower
399 soil water content, a shift of species associated with the plateaus favoured species with higher
400 hydraulic safety, otherwise 2) access to structured soils seems to be essential for trees to grow
401 taller. A previous study showed an increase in wood density from stands on sandy soils in valleys

402 to clayey soils on plateaus at a local scale in Central Amazon, and lower tree mortality rates in
403 clayey soils (Toledo *et al.* 2016). These patterns were primarily driven by soil moisture - correlated
404 to depth to water table - causing shifts in tree community composition (Schietti *et al.* 2013), and
405 favouring higher hydraulic safety in the lower soil moisture areas of the plateaus (Toledo *et al.*
406 2016; Cosme *et al.* 2017). We suggest that the structured soils allow trees to obtain an additional
407 volume of water during the dry season towards eastern Amazon, where soils tend to be richer in
408 clay compared to central and western Amazon (Fisher *et al.* 2008; Hodnett *et al.* 1997). The
409 dimorphic root systems associated with structured, clayey soils can redistribute water from deep
410 layers to the soil surface during periods of drought (Broedel *et al.* 2017).

411 Chemical and physical properties of soils across the Amazon Basin tend to correlate with variations
412 in and type of parent material, and exhibit an east-west soil age gradient (Quesada *et al.* 2011).
413 This edaphic variation across geological formations has strong influences on the floristic, structural
414 and demographic patterns in the Amazon (Quesada *et al.* 2012; ter Steege *et al.* 2006), with abrupt
415 changes in species composition following changes in soil properties and topography (Phillips *et*
416 *al.* 2003; Higgins *et al.* 2011). These patterns reflect more than a simple east-west gradient, due to
417 a complex history of deposition and erosion dating to the Miocene (Higgins *et al.* 2011). Despite
418 the clear heterogeneity caused by abrupt edaphic variation, two main gradients explain 24% of the
419 total variation in tree community composition: one from the Guiana Shield to the southwestern
420 Amazon, congruent with variation in soil fertility and its effects on tree wood density and seed
421 mass, and another gradient from Colombia to the southeastern Amazon related to the length of
422 the dry season (ter Steege *et al.* 2006). These gradients have distinctions in terms of their most
423 abundant genera and occurrence of the Fabaceae family, which contains most of the large trees and
424 grow successfully in low-dynamics environments such as the Guiana Shields. Higher occurrence
425 of the Fabaceae in these low-fertility soils may occur due to the ability to fix nitrogen in the soil
426 and ectomycorrhizal association (Webb & Sprent 2002; Sprent 2009).

427 Our results also demonstrate mean annual precipitation as a key factor for trees to grow taller. A tol-
428 erance curve associated the height of tall trees with precipitation, peaking at 2,300 mm yr⁻¹ as op-
429 timal, but also showing that areas too dry or too wet may both inhibit the growth of tall trees. Thus,
430 we observed 4 m decline in maximum tree height in regions with annual precipitations below 1,500
431 mm yr⁻¹ or above 3,000 mm yr⁻¹. The availability of soil water depends on both precipitation and
432 evapotranspiration, and our results suggest that below 1,500 mm yr⁻¹ evapotranspiration may ex-
433 ceed precipitation in the Amazon (Scheffer *et al.* 2018), and mortality by the hydraulic failure may
434 occur for trees near their maximum height (McDowell *et al.* 2008). Mean annual precipitation
435 above 2,300 mm year⁻¹ may be related to exceeding water, and the combination of high precip-
436 itation and poorly drained soils may result in anaerobic conditions with negative effects on tree
437 growth and survival (Quesada *et al.* 2009). Furthermore, higher precipitation tends to be related to
438 the occurrence of storms and stronger winds with increases in tree mortality (Aleixo *et al.* 2019).

439 Temperature and precipitation are key variables modulating the composition of species in the north-
440 western to southeastern seasonality gradient (ter Steege *et al.* 2006). The mean precipitation in the
441 Brazilian Amazon varies from less than 2,000 mm year⁻¹ (in the south, east, and extreme north) to
442 more than 3,000 mm year⁻¹ (in the northwest) (Liebmann & Marengo 2001). The annual convective
443 movement of the inter-tropical convergence zone results in distinct wet and dry seasons (Marengo
444 & Nobre 2001). However, the dry season in the Amazon basin varies from virtually nonexistent
445 to periods reaching up to seven consecutive months with less than 100 mm month⁻¹ of rain (Som-
446 broek 2001). A global analysis provided evidence for the control of water availability over forest
447 canopy height around the world, but the predictability between wet/dry indicates the involvement
448 of additional limiting factors as temperature or radiation (Klein *et al.* 2015).

449 **Conclusion**

450 Plant size distributions can be understood as the demographic consequence of size-dependent vari-
451 ation in growth and mortality in old-growth forests, and the mortality of large trees is independent
452 of resource availability and competition (Coomes *et al.* 2003). Understanding the spatial dis-
453 tribution of maximum tree height in tropical forests and how it is associated with environmental
454 conditions and tree functional traits is of fundamental importance. Emergent trees that reach their
455 maximum height are responsible for a significant amount of the transpired water flux and the above-
456 ground carbon storage. Trees which reach these extraordinary heights are rare and only a small
457 proportion of species have the necessary adaptations to achieve this. However, these adaptations are
458 not sufficient alone, and maximum tree height is strongly influenced by environmental conditions.
459 We found that, across the Brazilian Amazon, the most important conditions were a lower number
460 of clear sky days (reducing stress from direct sunlight), and soil clay content (improving water
461 retention). Our second analysis emphasized the importance of disturbance, showing that the tallest
462 trees are only found in places with low wind speed, allowing trees to grow for centuries without
463 substantial damage.

464 Current climate models differ in their predictions of large-scale changes in wind patterns, although
465 warmer temperatures will mean that the air can hold more moisture, which will likely make con-
466 vective storms more intense. Whatever the change in environmental conditions, it is likely to occur
467 faster than trees can adapt. Our results showed that precipitation and temperature have a lower
468 importance than expected from previous studies. Nevertheless, changes in the precipitation and
469 radiation regimes (strongly linked to the number of cloudy days) could reshape our forest biomes.
470 Ultimately, the association between environmental conditions and mechanisms of natural selec-
471 tion, where some traits have some advantages in comparison to others influencing the survival of
472 the most adaptable, are key to understanding the complexity of this process in a changing climate.

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489 **Author contributions**

490 EBG, TJ, DC, MK, NH, MHN, JO conceived of the idea. EBG, MA, GS, FSP, AZM developed
491 the analysis and performed the computations. EBG, MHN, TJ, MK, DV, RV, NH, CRR, RC, DAA,
492 JR, BG, JO verified the results, interpreted the results, and wrote the manuscript.

493 **Competing interests**

494 The authors have no conflicts to declare.

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815 **Tables**

	fa- par	srtm	us- peed	vspeed	clear- Days	days20	light- ning	month100	pan- nual	pdri- est	pet	psea- son	pwettest	tan- nual	tsea- son	tmax	clay- Con- tent	water- Content
fapar	1	0.14	0.1	0.15	0.08	0.01	0.17	-0.21	0.12	0.09	-	-	-	-	-	-	0.01	0.11
											0.28	0.18	0.06	0.17	0.06	0.03		
srtm	0.14	1	0.44	0.27	0.6	-	-0.03	0.45	-	-	0.28	0.5	-0.1	-	0.51	0.41	-0.18	0.29
						0.28			0.41	0.53				0.74				
uspeed	0.1	0.44	1	0.67	0.73	-	0.33	0.13	-	-	-	0.32	-	-	0.42	0.5	-0.56	0.28
						0.25			0.19	0.36	0.05		0.13	0.48				
vspeed	0.15	0.27	0.67	1	0.56	-	0.23	-0.03	-	-	-	0.17	-	-	0.33	0.31	-0.42	0.43
						0.42			0.11	0.34	0.46		0.24	0.49				
clear- Days	0.08	0.6	0.73	0.56	1	-	0.36	0.41	-	-	0.2	0.54	-	-0.5	0.53	0.78	-0.48	0.39
						0.47			0.37	0.62			0.11					
days20	0.01	-	-	-	-0.47	1	-0.1	-0.52	0.64	0.6	0.03	-	0.4	0.3	-	-0.3	0.02	-0.31
		0.28	0.25	0.42								0.46			0.22			
light- ning	0.17	-	0.33	0.23	0.36	-0.1	1	-0.2	0.2	0.04	-	-	-	0.07	-	0.34	-0.29	0.01
		0.03									0.15	0.13	0.02		0.18			
month100	-	0.45	0.13	-	0.41	-	-0.2	1	-	-	0.63	0.89	0.04	-	0.4	0.5	0.09	0.18
		0.21		0.03		0.52			0.77	0.82				0.24				
pannual	0.12	-	-	-	-0.37	0.64	0.2	-0.77	1	0.75	-	-	0.46	0.31	-	-	-0.1	-0.2
		0.41	0.19	0.11							0.32	0.73			0.38	0.29		
pdriest	0.09	-	-	-	-0.62	0.6	0.04	-0.82	0.75	1	-	-	0.05	0.45	-	-0.6	0.03	-0.35
		0.53	0.36	0.34							0.36	0.91			0.43			
pet	-	0.28	-	-	0.2	0.03	-0.15	0.63	-	-	1	0.58	0.4	0.08	0.25	0.48	0	-0.11
		0.28	0.05	0.46					0.32	0.36								
pseason	-	0.5	0.32	0.17	0.54	-	-0.13	0.89	-	-	0.58	1	0.16	-	0.43	0.6	-0.04	0.24
		0.18				0.46			0.73	0.91			0.38					
pwettest	-	-0.1	-	-	-0.11	0.4	-0.02	0.04	0.46	0.05	0.4	0.16	1	0.2	-	0.19	0.06	-0.13
		0.06	0.13	0.24										0.12				
tannual	-	-	-	-	-0.5	0.3	0.07	-0.24	0.31	0.45	0.08	-	0.2	1	-0.4	-	0.2	-0.33
		0.17	0.74	0.48	0.49						0.38					0.12		
tseason	-	0.51	0.42	0.33	0.53	-	-0.18	0.4	-	-	0.25	0.43	-	-0.4	1	0.42	-0.23	0.21
		0.06				0.22			0.38	0.43			0.12					
tmax	-	0.41	0.5	0.31	0.78	-0.3	0.34	0.5	-	-0.6	0.48	0.6	0.19	-	0.42	1	-0.33	0.21
		0.03							0.29					0.12				
clay- Content	0.01	-	-	-	-0.48	0.02	-0.29	0.09	-0.1	0.03	0	-	0.06	0.2	-	-	1	-0.13
		0.18	0.56	0.42								0.04			0.23	0.33		
water- Content	0.11	0.29	0.28	0.43	0.39	-	0.01	0.18	-0.2	-	-	0.24	-	-	0.21	0.21	-0.13	1
						0.31				0.35	0.11		0.13	0.33				

Table 1: Correlation between the environmental variables.

Layer	Definition	Related to	Unit	Source	Importance
clearDays	number of clear days per year	energy balance - water balance - radiation	days	MODIS	14.7
clayContent	fraction of clay content	soil structure - physical properties - water availability	%	Open-LandMap	13.8
topography	elevation above sea level	distance to water - flooding zones - soil	m	SRTM	11.2
pannual	average annual precipitation	precipitation - precipitation intensity - precipitation distribution	mm	WorldClim	8.9
pseason	precipitation seasonality	precipitation - precipitation intensity - precipitation distribution	mm	WorldClim	6.9
fapar	fraction of absorbed photosynthetically active radiation	radiation - vegetation health - anthropic regions - soil exposure	%	NOAA AVHRR	6.3
pwettest	precipitation of the wettest month	precipitation - precipitation intensity - precipitation distribution	mm	WorldClim	5.8
uspeed	zonal speed (W-E)	storms - convective winds	m/s	ECM-RWF	5.6
days20	days with precipitation higher then 20 mm	storms - convective winds	days	CHIRPS	5.5
pet	potential evapotranspiration	energy balance - water balance - radiation - vegetation health - anthropic regions - soil exposure	mm	TerraClimate	5.2
tseason	temperature seasonality	temperature - temperature distribution	C	WorldClim	4.6
tmax	maximum temperatura	storms - convective winds	C	WorldClim	4.2
vspeed	meridional speed (N-S)	storms - convective winds	m/s	ECM-RWF	3.6
lightning	lightning rate	storms - convective winds	flashes rate	LIS TRMM	3.5
tannual	daily average annual temperature	temperature - temperature distribution	C	WorldClim	0.3
water-Content	fraction of water content	soil structure - physical properties - water availability	%	Open-LandMap	0
month100	month with precipitation below 100 mm	precipitation - precipitation intensity - precipitation distribution	months	CHIRPS	Removed by high correlation
pdriest	precipitation of the driest month	precipitation - precipitation intensity - precipitation distribution	mm	WorldClim	Removed by high correlation

Table 2: Variables used to estimate maximum height distribution and evaluate its distribution, ranked by variable importance results in the Random Forest model

816 **Figure Captions**

Figure 1. Maps of the Brazilian Amazon showing the location of trees > 50 m, > 60 m, > 70 m, and > 80 m in height. Black circles indicate the presence of a tree above the height thresholds. Background color indicates the biogeographical subdivisions proposed by (Morrone 2014): I - Para; II - Xingu-Tapajos; III - Roraima; IV - Guianan Lowlands; V - Madeira; VI - Yungas; VII - Pantepui; VIII - Imeri.

Figure 2. Maximum height estimation based on remote sensing variables estimated by Random Forest method. Black lines indicate the biogeographical subdivisions: I - Para; II - Xingu-Tapajos; III - Roraima; IV - Guianan Lowlands; V - Madeira; VI - Yungas; VII - Pantepui; VIII - Imeri.

Figure 3. Maximum tree height distribution when FAPAR values under 80% were excluded from our analysis.

Figure 4. Marginal plot for each variable considering the Random Forest model for maximum height estimation

Figure 5. Marginal plot for each variable considering the Maximum Entropy model for niche determination

Figure 6. Occurrence of giant trees (black dots) and niche capability to support the development of tall trees (probability of tall tree occurrence). Black lines indicate the biogeographical subdivisions: I - Para; II - Xingu-Tapajos; III - Roraima; IV - Guianan Lowlands; V - Madeira; VI - Yungas; VII - Pantepui; VIII - Imeri.

817 **Figures**

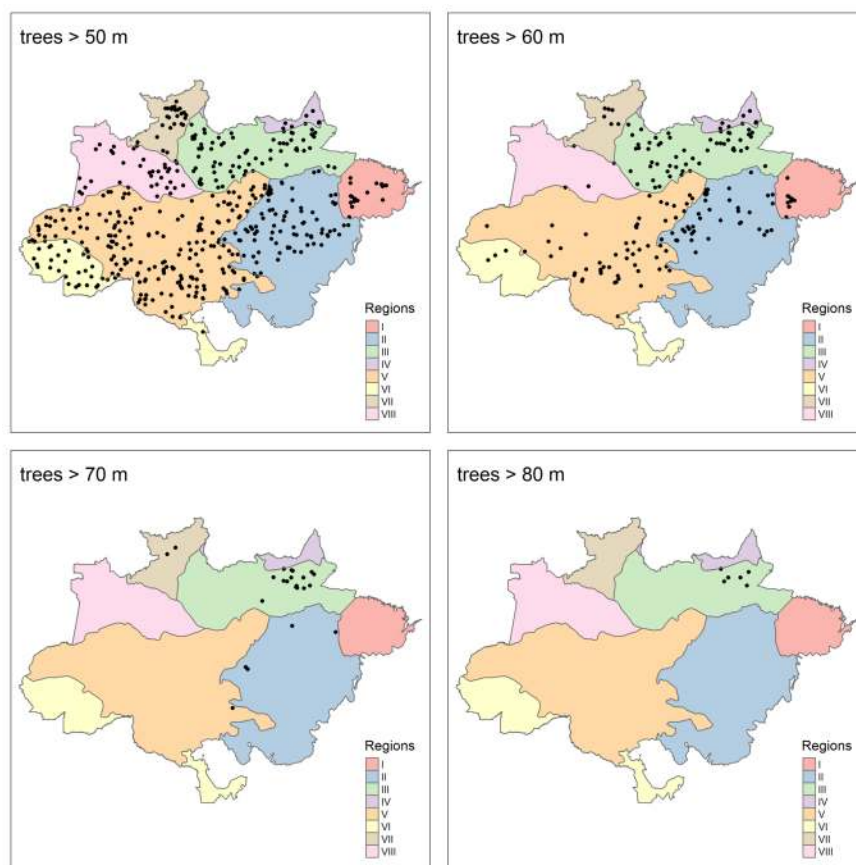


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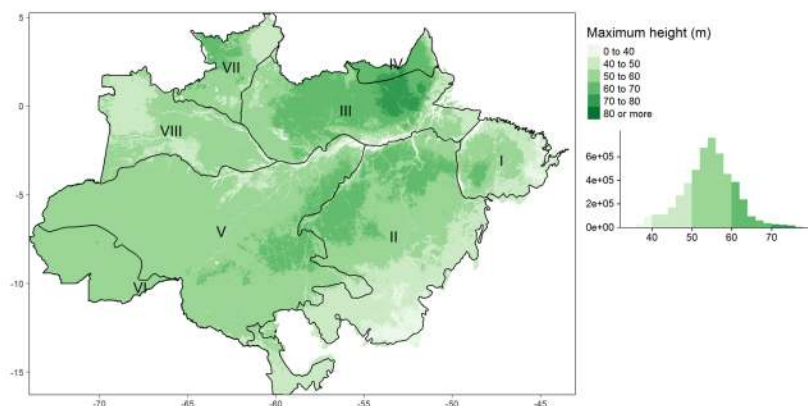


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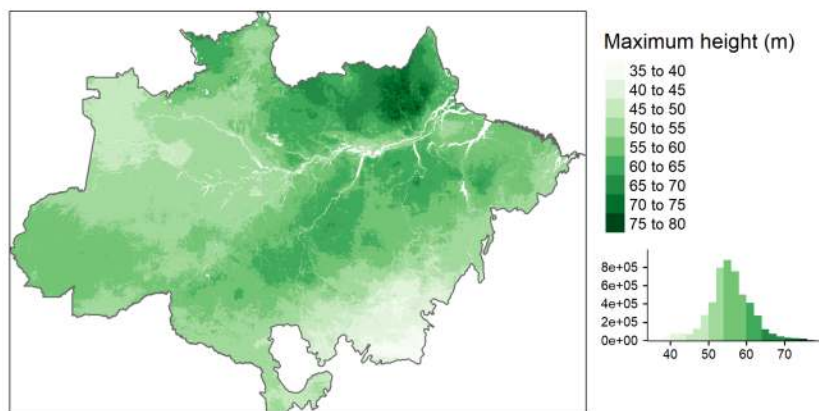


Figure 3: Maximum tree height distribution when FAPAR values under 80% were excluded from our analysis.

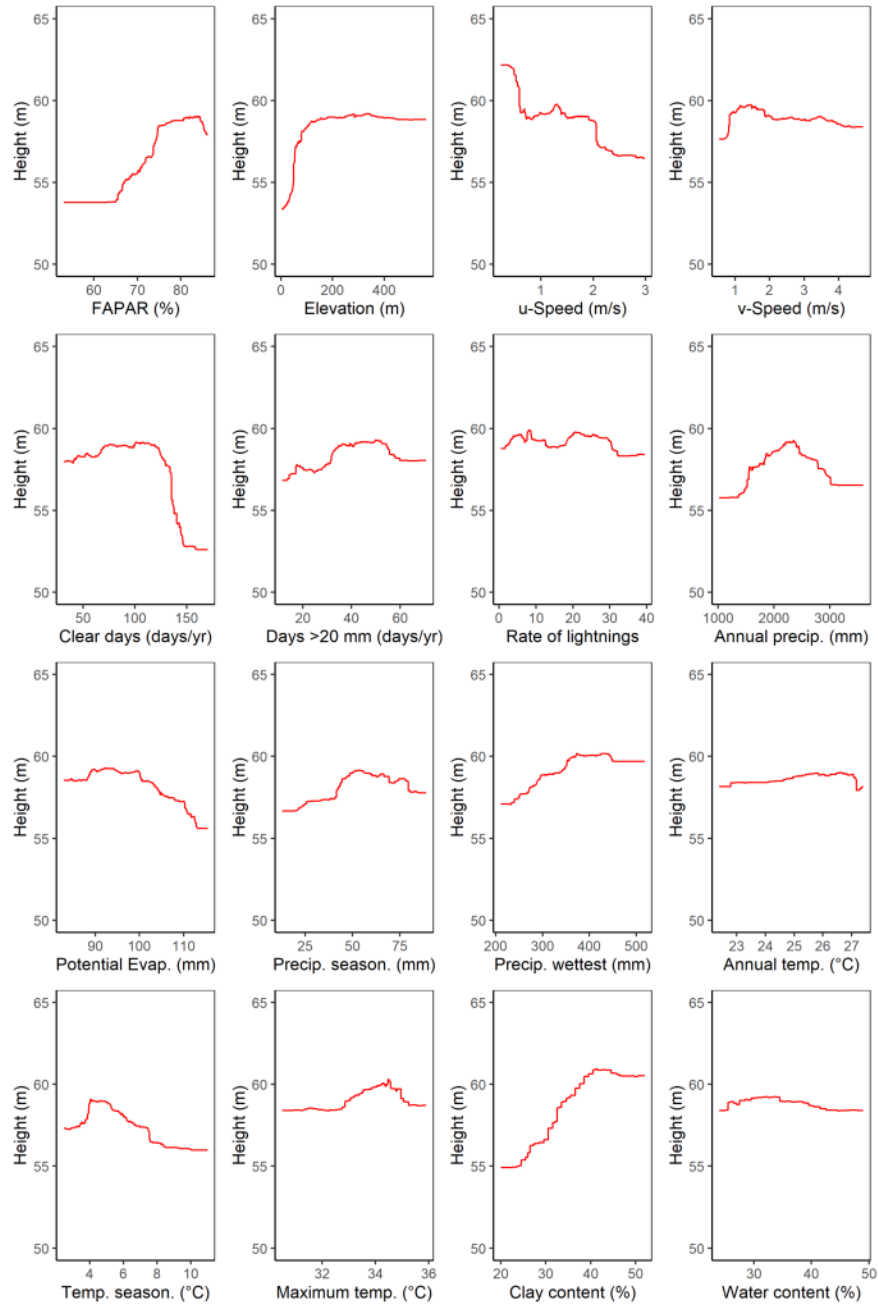


Figure 4: Marginal plot for each variable considering the Random Forest model for maximum height estimation

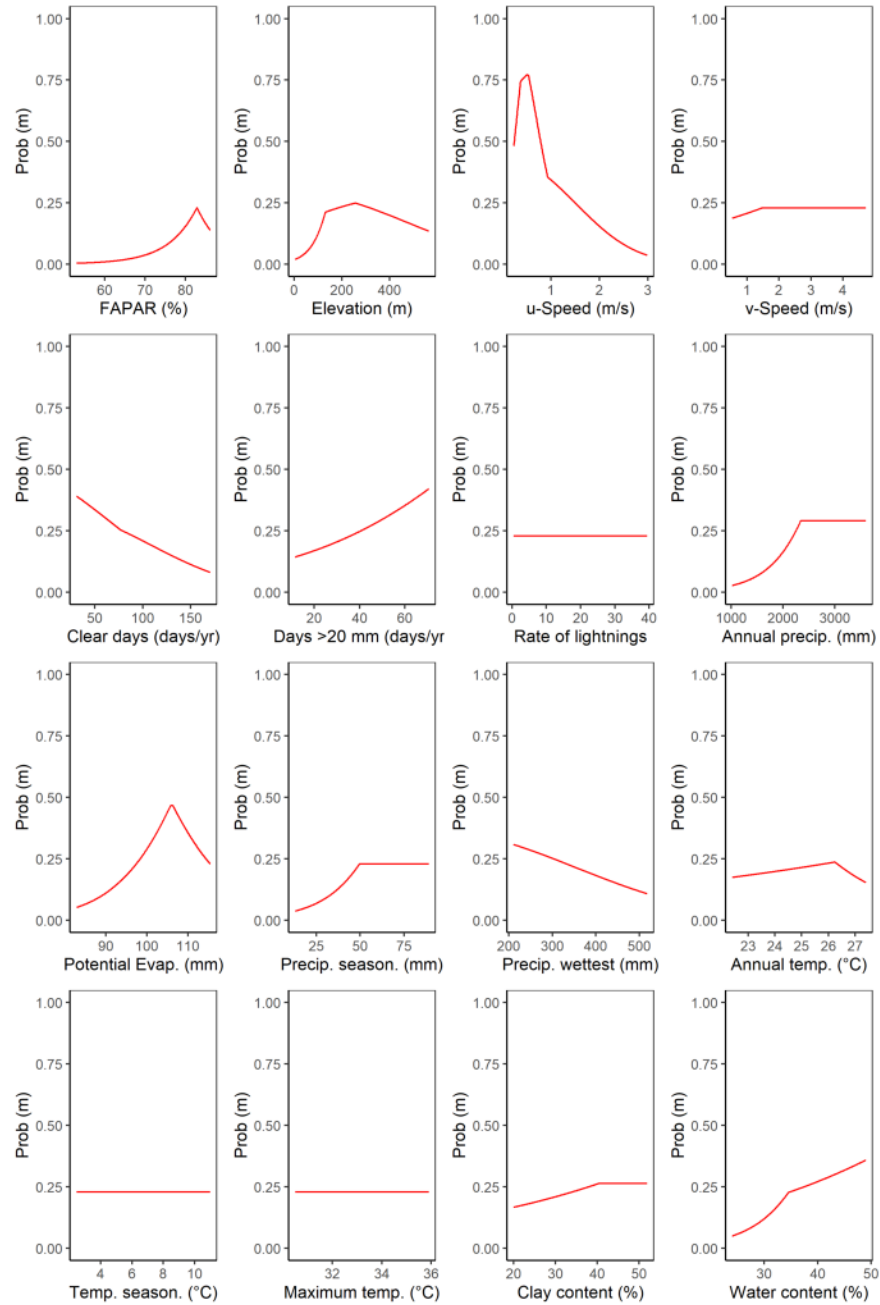


Figure 5: Marginal plot for each variable considering the Maximum Entropy model for niche determination

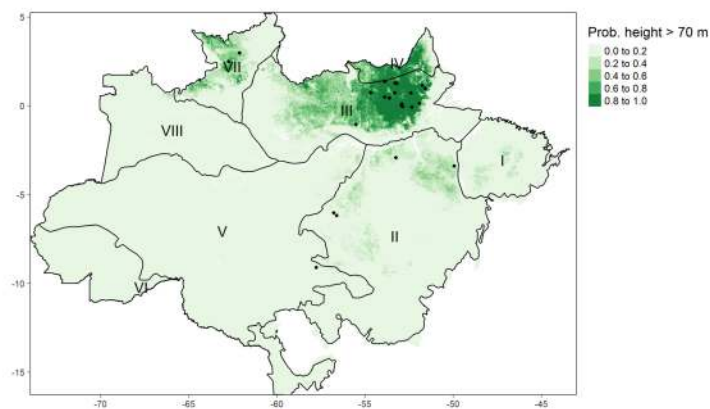


Figure 6: Occurrence of giant trees (black dots) and niche capability to support the development of tall trees (probability of tall tree occurrence). Black lines indicate the biogeographical subdivisions: I - Para; II - Xingu-Tapajos; III - Roraima; IV - Guianan Lowlands; V - Madeira; VI - Yungas; VII - Pantepui; VIII - Imeri.