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*Published in:*  
Journal of Fluid Mechanics

*Link to article, DOI:*  
[10.1017/jfm.2022.92](https://doi.org/10.1017/jfm.2022.92)

*Publication date:*  
2022

*Document Version*  
Peer reviewed version

[Link back to DTU Orbit](#)

*Citation (APA):*  
Li, Y., Larsen, B. E., & Fuhrman, D. R. (2022). Reynolds stress turbulence modelling of surf zone breaking waves. *Journal of Fluid Mechanics*, 937, [A7]. <https://doi.org/10.1017/jfm.2022.92>

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# 1 Reynolds stress turbulence modelling of surf zone 2 breaking waves

3 Yuzhu Li<sup>1,2</sup>†, Bjarke Eltard Larsen<sup>1</sup> and David R. Fuhrman<sup>1</sup>

4 <sup>1</sup>Technical University of Denmark, Department of Mechanical Engineering, Section for Fluid Mechanics,  
5 Coastal and Maritime Engineering, 2800 Kgs. Lyngby, Denmark

6 <sup>2</sup>National University of Singapore, Department of Civil and Environmental Engineering, Singapore 117576

7 (Received xx; revised xx; accepted xx)

8 Computational fluid dynamics is increasingly used to investigate the inherently complicated  
9 phenomenon of wave breaking. To date, however, no single model has proved capable of  
10 accurately simulating the breaking process across the entirety of the surf zone for both  
11 spilling and plunging breakers. The present study newly considers the Reynolds stress- $\omega$   
12 turbulence closure model for this purpose, where  $\omega$  is the specific dissipation rate. Novel  
13 stability analysis proves that, unlike two-equation closures (at least in their standard forms),  
14 the stress- $\omega$  model is neutrally stable in the idealized potential flow region beneath surface  
15 waves. It thus naturally avoids unphysical exponential growth of turbulence prior to breaking,  
16 which has plagued numerous prior studies. The analysis is confirmed through simulation of  
17 a progressive surface wave train. The stress- $\omega$  model is then applied to simulate a turbulent  
18 wave boundary layer, demonstrating superior accuracy relative to a two-equation model,  
19 especially during flow deceleration. Finally, the stress- $\omega$  model is employed to simulate  
20 spilling and plunging breaking waves, with seemingly unprecedented accuracy. Specifically,  
21 the present work marks the first time that a single turbulence closure model collectively:  
22 (1) avoids turbulence over-production prior to breaking, (2) accurately predicts the breaking  
23 point, (3) provides reasonable evolution of turbulent normal stresses, while also (4) yielding  
24 accurate evolution of undertow velocity structure and magnitude across the surf zone, for  
25 both spilling and plunging cases. Differences in the predicted Reynolds shear stresses (hence  
26 flow resistance) are identified as key to the improved inner surf zone performance, relative  
27 to a state-of-the-art two-equation model.

28 **Key words:**

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## 29 1. Introduction

30 Breaking water waves feature a rather amazing variety of fluid mechanics, ranging from nearly  
31 potential flow prior to breaking, to unsteady turbulent boundary layers at the sea bed, to a  
32 turbulent jet flow e.g. during the initial plunge, to a highly complicated and turbulent multi-  
33 phase (air and water) flow throughout the surf zone. Over the past decades, significant efforts

† Email address for correspondence: pearl.li@nus.edu.sg

34 have been made to better understand the breaking wave process through both experimental  
35 and numerical means.

36 A large number of experimental studies have been performed, with focus on e.g. the  
37 breaking onset location, turbulence characteristics, as well as the undertow velocity field in  
38 the surf zone, which is especially important in nearshore sediment transport processes. The  
39 surf zone is the part of the shoreface from the most seaward wave breaking point to the most  
40 landward broken wave (Van Rijn 1993). The surf zone can be divided into two sub-regions,  
41 i.e. the outer and the inner surf zone. For spilling breakers, there has not been a specific  
42 definition of the threshold between two sub-regions. It can be considered that the outer surf  
43 zone extends from the breaking point up to the part with rapid changes in wave shape, and the  
44 inner surf zone consists of the breaking bores with slow changes in wave shape. For plunging  
45 breakers, the splash point (where the water pushed upwards by the plunging jet hits the water  
46 again) is often used to mark the start of the inner surf zone. When breaking waves propagate  
47 to the shore, a return flow (known as undertow) beneath the wave trough is generated to  
48 compensate the amount of water waves that is transported shoreward. The undertow velocity  
49 is generally strongest in the surf zone (Svendsen 1984). Most of the experimental studies have  
50 been performed in relatively small scale facilities (e.g. Nadaoka *et al.* 1989; Chang & Liu  
51 1999; Ting & Kirby 1994, 1996; Stansby & Feng 2005; De Serio & Mossa 2006; Lara *et al.*  
52 2006). Among these, the spilling and plunging breaking wave experiments of Ting & Kirby  
53 (1994, 1996) have been most often used for validating numerical models. Spilling breaking  
54 is a rather gentle breaking at the wave crest and is followed by a gradual dissipation of energy  
55 over the surf zone, while plunging breaking is more violent with the crest curling over and  
56 plunging into the surface as a turbulent jet flow. Recently, several large-scale experimental  
57 studies involving breaking waves over a fixed barred bed profile (e.g. Scott *et al.* 2005; van  
58 der A *et al.* 2017; van der Zanden *et al.* 2018, 2019) have likewise been performed, with  
59 detailed measurements provided for the flow and turbulence fields throughout the surf zone,  
60 as well as in the near bed bottom boundary layer region (van der Zanden *et al.* 2018).

61 With the continual increase in computer power, computational fluid dynamics (CFD)  
62 modelling has been increasingly utilized as an alternative means of studying breaking waves,  
63 due to its cheaper cost and faster set-up compared to conventional laboratory tests. CFD can  
64 also, in principal, overcome scale effects and operation disturbances that exist in laboratory  
65 experiments. CFD simulations on breaking waves have typically been conducted based  
66 on Reynolds-averaged Navier Stokes (RANS) equations, coupled with various turbulence  
67 closure models (e.g. Lin & Liu 1998; Bradford 2000; Chella *et al.* 2015; Derakhti *et al.*  
68 2016*a,b*; Lupieri & Contento 2015; Brown *et al.* 2016; Devolder *et al.* 2018; Liu *et al.*  
69 2020). Additionally, large eddy simulation models (LES, e.g. Christensen & Deigaard 2001;  
70 Christensen 2006; Zhou *et al.* 2017) have also been employed to study wave breaking  
71 processes, as have models based on so-called smoothed particle hydrodynamics (SPH, e.g.  
72 Shao 2006; Shadloo *et al.* 2015; Wei *et al.* 2018; Lowe *et al.* 2019). In recently years,  
73 some high-fidelity direct numerical simulation (DNS) studies have been made on breaking  
74 waves with focus on air-entrainment and bubble statistics (e.g. three-dimensional simulations  
75 of Deike *et al.* 2016; Wang *et al.* 2016; Chan *et al.* 2021), which have built largely upon  
76 previous two-dimensional simulations solving the Navier-Stokes equations (e.g. Iafrati 2009,  
77 2011). These high-fidelity simulations are at small length scales and are not yet practically  
78 applicable to surf zone breaking waves due to computational time and costs. Among those  
79 various approaches, RANS models have been those most widely used for surf zone breaking  
80 wave modelling, as they are the most computationally affordable.

81 Regarding RANS two-equation models, the pioneering work of Lin & Liu (1998) applied a  
82 nonlinear  $k-\varepsilon$  model for simulating breaking waves ( $k$  is the turbulent kinetic energy density,  
83 and  $\varepsilon$  is the dissipation rate). Their simulations showed a pronounced over-production of

84 turbulence at their most offshore point of comparison (near the breakpoint). This is similar to  
85 other more recent works (e.g. Brown *et al.* 2016; Derakhti *et al.* 2016*a,b*; Devolder *et al.* 2018;  
86 Liu *et al.* 2020) using other two-equation models such as  $k-\omega$  and  $k-\omega$  shear stress transport  
87 (SST) models ( $\omega$  being the specific dissipation rate). Several of the simulations mentioned  
88 just above even clearly demonstrate turbulence levels prior to breaking that are similar  
89 in magnitude to those within the surf zone, which obviously defies physical explanation  
90 as well as measurements. Hsu *et al.* (2002) also identified that the  $k-\varepsilon$  turbulence model  
91 tended to predict unrealistically high turbulence in regions that were supposed to contain  
92 low turbulence levels during their long-time simulations. They suspected that this problem  
93 was due to convection and diffusion mechanisms. To combat this issue they have used an  
94 empirical damping coefficient to reduce the eddy viscosity in such regions.

95 The persistent problem of over-production of turbulence in the potential flow region  
96 beneath (non-breaking) surface waves in RANS turbulence closure models has only recently  
97 been fully explained and analyzed. Building on the proof of conditional instability of the  
98  $k-\omega$  closure model of Mayer & Madsen (2000), Larsen & Fuhrman (2018) proved that nearly  
99 all two-equation models in wide use (several  $k-\omega$  and  $k-\varepsilon$  variants) are (asymptotically)  
100 unconditionally unstable in such regions. (An exception is the realizable  $k-\varepsilon$  model of Shih  
101 *et al.* (1995), which was proved to be conditionally unstable in such regions by Fuhrman & Li  
102 2020). Larsen & Fuhrman (2018) devised a simple and general method for formally stabilizing  
103 two-equation models, based on a reformulation of the eddy viscosity. Their “stabilized”  $k-\omega$   
104 model was tested on small-scale spilling waves over a constant slope in Larsen & Fuhrman  
105 (2018), and on full-scale plunging waves over a breaker bar in Larsen *et al.* (2020). These  
106 works have collectively shown that the stabilized  $k-\omega$  model leads to marked improvement in  
107 the predicted turbulence, undertow velocity profiles, and the bottom boundary layer dynamics  
108 in the pre-breaking region and outer surf zones, likely to be of considerable importance for  
109 e.g. breaking wave hydrodynamics and cross-shore sediment transport predictions. However,  
110 even the best of the models considered in Larsen & Fuhrman (2018) and Fuhrman & Li  
111 (2020) were still rather inaccurate in the inner surf zone (i.e. closer to the shoreline), thus  
112 seemingly requiring yet more advanced methods of achieving turbulence closure. To date,  
113 no single turbulence closure model has demonstrated the ability to accurately simulate the  
114 entirety of the breaking process, from shoaling to the inner surf zone, including accurate  
115 prediction of the undertow velocity structure and magnitude, for both spilling and plunging  
116 breaking waves.

117 NASA’s CFD Vision 2030 Study white paper (Slotnick *et al.* 2014) identifies advanced  
118 turbulence modelling based on Reynolds stress models (RSMs) as a priority in the coming  
119 decades. Motivated by this, and especially the persistent shortcomings encountered with  
120 two-equation turbulence closure models noted above, the present study will consider novel  
121 applications of a Reynolds stress turbulence model for the simulation of breaking waves.  
122 Specifically, we will consider applications of the stress- $\omega$  model proposed by Wilcox (2006),  
123 which has not been utilized previously for this purpose. Unlike two-equation models, RSMs  
124 (e.g. Wilcox 2006; Launder *et al.* 1975) simulate all components of the Reynolds stress  
125 tensor with their own respective transport equation, eliminating the need to resort to a  
126 Boussinesq eddy viscosity approximation. RSMs are therefore theoretically superior to  
127 their two-equation counterparts, while still maintaining reasonable computational efficiency,  
128 compared to turbulence-resolving methods such as DNS and LES. Comparing to two-  
129 equation RANS models, RSMs must provide closure for a larger number of terms, which  
130 can present a challenge. In the present work, the closure terms and coefficients provided in  
131 Wilcox (2006) will be adopted.

132 To the authors’ knowledge, the study of Brown *et al.* (2016) has been the only one to  
133 have attempted application of a RSM to study breaking waves, in their case utilizing the

134 Launder-Reece-Rodi (LRR) stress- $\varepsilon$  model (Launder *et al.* 1975). However, they found a  
135 significant over-estimation of the turbulent kinetic energy for spilling breakers both pre-  
136 and post-breaking, which was even more pronounced than found with several of their two-  
137 equation closures. Their results suggest that RSMs may share the same problem of instability  
138 in the nearly potential flow region beneath surface waves, leading to unphysical exponential  
139 growth of turbulence. The formal stability of RSMs in the potential flow region beneath  
140 non-breaking surface waves is an open question, which will be definitively answered by the  
141 present work. We further aim to establish the ability of the stress- $\omega$  model to accurately  
142 simulate coastal fluid mechanics problems involving breaking waves.

143 The present work is organized as follows: We begin by conducting a novel stability analysis  
144 of the Wilcox (2006) stress- $\omega$  model in a region of idealized potential flow beneath surface  
145 waves (Section 2). We will prove that this model is formally neutrally stable in such regions,  
146 and therefore ought not give rise to unphysical exponential growth of turbulence. The stress-  
147  $\omega$  model (with buoyancy production terms included, as derived in Appendix A) will then  
148 be tested in CFD simulations throughout Section 3. Here the formal stability analysis will  
149 be directly verified through simulations of a progressive surface wave train (Section 3.1).  
150 We then move from the surface to the sea bed, and consider CFD simulations of a turbulent  
151 wave boundary layer, with comparisons made against a two-equation  $k$ - $\omega$  model (Section  
152 3.2). We finally test the performance of the stress- $\omega$  model in simulations involving both  
153 the spilling (Section 3.3) and plunging (Section 3.4) breaking wave cases of Ting & Kirby  
154 (1994, 1996), with direct comparison made against the best of the  $k$ - $\omega$  models devised by  
155 Larsen & Fuhrman (2018). The present breaking wave results are discussed relative to those  
156 of prior CFD studies in Section 4, before drawing conclusions in Section 5.

157 Although it is not the primary focus of the present work, for completeness, we similarly  
158 analyze the LRR stress- $\varepsilon$  model for stability in Appendix B. Similar to the stress- $\omega$  model, we  
159 prove that the stress- $\varepsilon$  model is likewise neutrally stable in the potential flow region beneath  
160 non-breaking surface waves. This has also been confirmed through testing with surface wave  
161 trains, as noted there. The likely explanation of the LRR stress- $\varepsilon$  model significantly over-  
162 predicting turbulence prior to breaking in the work of Brown *et al.* (2016) is also provided  
163 there.

## 164 2. Stability analysis of the Wilcox (2006) stress- $\omega$ turbulence model in the 165 potential flow region beneath waves

### 166 2.1. Turbulence closure model

167 While computational power has improved immensely in recent decades, for many fluid  
168 mechanics problems, it is still not practically feasible to resolve the small scales required  
169 for either DNS or LES simulations. Rather, it is often necessary in practice to work with a  
170 Reynolds-averaged description of the flow, with the effects of turbulence on the mean flow  
171 accounted for with the aid of a turbulence closure model. For this purpose, the present study  
172 will focus on the Wilcox (2006) stress- $\omega$  model (where  $\omega$  is again the specific dissipation  
173 rate of turbulence). This model, in a form suitable for a two-phase (water-air) fluid mixture,

175 consists of the following stress-transport equations:

$$\begin{aligned}
 & \underbrace{\frac{\partial \bar{\rho} \tau_{ij}}{\partial t}}_{\text{Time variation}} + \underbrace{\bar{u}_k \frac{\partial \bar{\rho} \tau_{ij}}{\partial x_k}}_{\text{Convection}} = - \underbrace{\bar{\rho} P_{ij}}_{\text{Production}} + \underbrace{\frac{2}{3} \bar{\rho} \beta^* \omega k \delta_{ij}}_{\text{Dissipation}} - \underbrace{\bar{\rho} \Pi_{ij}}_{\text{Pressure-strain}} \\
 & + \underbrace{\bar{\rho} \alpha_b^* \frac{k}{\omega} N_{ij}}_{\text{Buoyancy production}} + \underbrace{\frac{\partial}{\partial x_k} \left[ \bar{\rho} \left( \nu + \sigma^* \frac{k}{\omega} \right) \frac{\partial \tau_{ij}}{\partial x_k} \right]}_{\text{Diffusion}}
 \end{aligned} \tag{2.1}$$

178 combined with a separate transport equation for the specific rate of dissipation  $\omega$ :

$$\begin{aligned}
 & \underbrace{\frac{\partial \bar{\rho} \omega}{\partial t}}_{\text{Time variation}} + \underbrace{\bar{u}_j \frac{\partial \bar{\rho} \omega}{\partial x_j}}_{\text{Convection}} = \underbrace{\bar{\rho} \alpha \frac{\omega}{k} \tau_{ij} \frac{\partial \bar{u}_i}{\partial x_j}}_{\text{Production}} - \underbrace{\bar{\rho} \beta \omega^2}_{\text{Dissipation}} + \underbrace{\sigma_d \frac{\bar{\rho}}{\omega} \frac{\partial k}{\partial x_j} \frac{\partial \omega}{\partial x_j}}_{\text{Cross-diffusion}} \\
 & + \underbrace{\frac{\partial}{\partial x_k} \left[ \bar{\rho} \left( \nu + \sigma \frac{k}{\omega} \right) \frac{\partial \omega}{\partial x_k} \right]}_{\text{Diffusion}}
 \end{aligned} \tag{2.2}$$

180 In the above  $x_j$  are the Cartesian coordinates,  $\bar{u}_j$  are the mean (Reynolds-averaged) com-  
 181 ponents of the velocity,  $g_j$  is gravitational acceleration,  $\delta_{ij}$  is the Kronecker delta,  $\nu$  is the  
 182 kinematic fluid viscosity,  $\bar{\rho}$  is the fluid density, and  $t$  is time. The specific Reynolds stress  
 183 tensor is defined as:

$$184 \quad \tau_{ij} = -\overline{u'_i u'_j} \tag{2.3}$$

185 where a prime superscript denotes turbulent fluctuations and the overbar denotes Reynolds  
 186 averaging. The turbulent kinetic energy (per unit mass) is thus:

$$187 \quad k = \frac{1}{2} \tau_{kk} \tag{2.4}$$

188 Buoyancy production (as derived in Appendix A) is included with terms proportional to the  
 189 Brunt-Väisälä frequency tensor:

$$190 \quad N_{ij} = \frac{1}{\rho_0} \left( g_i \frac{\partial \bar{\rho}}{\partial x_j} + g_j \frac{\partial \bar{\rho}}{\partial x_i} \right) \tag{2.5}$$

191 where  $\rho_0$  is the constant reference density of the fluid.

193 The pressure-strain correlation is:

$$\begin{aligned}
 & \Pi_{ij} = \beta^* C_1 \omega \left( \tau_{ij} + \frac{2}{3} k \delta_{ij} \right) - \hat{\alpha} (P_{ij} - \frac{2}{3} P \delta_{ij}) \\
 & - \hat{\beta} (D_{ij} - \frac{2}{3} P \delta_{ij}) - \hat{\gamma} k (S_{ij} - \frac{1}{3} S_{kk} \delta_{ij})
 \end{aligned} \tag{2.6}$$

195 where

$$196 \quad P_{ij} = \tau_{im} \frac{\partial \bar{u}_j}{\partial x_m} + \tau_{jm} \frac{\partial \bar{u}_i}{\partial x_m} \tag{2.7}$$

$$198 \quad D_{ij} = \tau_{im} \frac{\partial \bar{u}_m}{\partial x_j} + \tau_{jm} \frac{\partial \bar{u}_m}{\partial x_i} \tag{2.8}$$

$$200 \quad S_{ij} = \frac{1}{2} \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \tag{2.9}$$

$$201 \quad P = \frac{1}{2} P_{kk} \quad (2.10)$$

202 The model closure coefficients, taken directly from Wilcox (2006), are defined as follows:

$$203 \quad \begin{aligned} C_1 = 1.8, & \quad C_2 = 10/19, & \hat{\alpha} = (8 + C_2)/11, & \hat{\beta} = (8C_2 - 2)/11 \\ \hat{\gamma} = (60C_2 - 4)/55, & \alpha = 0.52, & \beta^* = 0.09, & \beta_0 = 0.0708 \\ \beta = \beta_0 f_\beta, & \sigma = 0.5, & \sigma^* = 0.6, & \sigma_{d0} = 0.125 \end{aligned} \quad (2.11)$$

$$204 \quad \sigma_d = \begin{cases} 0, & \frac{\partial k}{\partial x_j} \frac{\partial \omega}{\partial x_j} \leq 0 \\ \sigma_{d0}, & \frac{\partial k}{\partial x_j} \frac{\partial \omega}{\partial x_j} > 0 \end{cases} \quad (2.12)$$

$$206 \quad f_\beta = \frac{1 + 85\chi_\omega}{1 + 100\chi_\omega}, \quad \chi_\omega = \left| \frac{\Omega_{ij}\Omega_{jk}\hat{\delta}_{ki}}{(\beta^*\omega)^3} \right|, \quad \hat{S}_{ki} = S_{ki} - \frac{1}{2} \frac{\partial \bar{u}_m}{\partial x_m} \delta_{ki} \quad (2.13)$$

207 with  $\alpha_b^* = 1.36$  (following Larsen & Fuhrman 2018, see also Appendix A). Unless explicitly  
208 stated otherwise, this value is fixed in what follows. A detailed description of the closure  
209 evolution from third-order turbulence correlations to second-order can be found in Wilcox  
210 (2006, p. 41–43) and Launder *et al.* (1975, their Section 3).

211 Compared to the LRR (Launder *et al.* 1975) stress- $\varepsilon$  model, the  $\omega$ -based stress-transport  
212 model formulated above reduces the complexity of the diffusion term and the pressure-strain  
213 relation considerably. Moreover, since the  $\omega$  equation yields better near-wall behaviour,  
214 the pressure-strain relation does not require an artificial wall-reflection term. (As discussed  
215 by Parneix *et al.* 1998, the LRR wall-reflection term is more to mitigate a deficiency in  
216 the  $\varepsilon$  equation than to correctly or physically represent the pressure-echo process.) We  
217 therefore adopt the Wilcox (2006) stress- $\omega$  model as our primary focus for both analysis and  
218 applications in what follows.

## 219 2.2. Stability analysis

220 As shown and explained by Mayer & Madsen (2000), Larsen & Fuhrman (2018), and  
221 Fuhrman & Li (2020) (see also Section 7.6 of Sumer & Fuhrman 2020), standard two-  
222 equation turbulence closure models can result in turbulence over-production in the potential  
223 flow core region beneath surface waves. This is due to their inherent instability in such  
224 regions, leading to non-physical exponential growth of the turbulent kinetic energy and  
225 eddy viscosity. Computational results of Brown *et al.* (2016), who used the LRR stress- $\varepsilon$   
226 turbulence model to simulate breaking waves, demonstrated seemingly similar turbulence  
227 over-production prior to incipient wave breaking. This suggests that RSMs may share the  
228 same inherent instability in nearly potential flow regions having finite strain. It is therefore  
229 of interest to extend the analysis of Larsen & Fuhrman (2018) to consider the formal asymptotic  
230 stability of Reynolds stress models. In what follows in the main text we will formally analyze  
231 the Wilcox (2006) stress- $\omega$  model. Similar analysis (and findings) of the LRR stress- $\varepsilon$  model  
232 is provided in Appendix B for completeness.

233 Consider now an incompressible fluid region having constant density beneath a small  
234 amplitude plane surface wave train propagating in the horizontal  $x_1 = x$  direction, where the  
235 turbulence model described above is active. We will assume the mean flow is described by  
236 linear potential flow (Stokes first-order) wave theory, with velocity fields:

$$237 \quad \bar{u}_1 = u = \frac{H\sigma_w}{2} \frac{\cosh(k_w y)}{\sinh(k_w h)} \cos(k_w x - \sigma_w t) \quad (2.14)$$

238

239

$$\bar{u}_2 = v = \frac{H\sigma_w}{2} \frac{\sinh(k_w y)}{\sinh(k_w h)} \sin(k_w x - \sigma_w t) \quad (2.15)$$

240

where the vertical  $x_2 = y$  axis is placed at the bed,  $\sigma_w$  is the angular wave frequency,  $k_w$  is the wave number,  $h$  is the water depth, and  $H$  is the wave height.

242

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245

246

Following Mayer & Madsen (2000), Larsen & Fuhrman (2018) and Fuhrman & Li (2020) diffusive and convective terms will be neglected in the analysis, which is reasonable in the potential flow region. Meanwhile, the buoyancy production term goes to zero in the region beneath surface waves where the density is again assumed constant. From the assumptions stated above, (2.1) and (2.2) simplify to the following system of seven governing equations:

247

$$\frac{\partial \tau_{ij}}{\partial t} = -P_{ij} + \frac{2}{3}\beta^* \omega k \delta_{ij} - \Pi_{ij} \quad (2.16)$$

248

249

$$\frac{\partial \omega}{\partial t} = \alpha \frac{\omega}{k} \tau_{ij} \frac{\partial \bar{u}_i}{\partial x_j} - \beta \omega^2 \quad (2.17)$$

250

251

252

253

254

We may simplify the governing equations yet further by (1) assuming that the turbulence field under consideration has equivalent normal stresses (such that  $\tau_{11} = \tau_{22} = \tau_{33}$ ), (2) accounting for both assumed zero mean flow ( $\bar{u}_3 = w = 0$ ) and uniformity ( $\partial/\partial x_3 = 0$ ) in the transverse  $x_3 = z$  direction, and (3) invoking local continuity  $\partial \bar{u}_i/\partial x_i = 0$ . Equations (2.16) and (2.17) then reduce considerably to the following system of three ODEs:

255

$$\frac{\partial k}{\partial t} = 2\tau_{12} S_{12} - \beta^* \omega k \quad (2.18)$$

256

257

$$\frac{\partial \tau_{12}}{\partial t} = \left( \frac{4}{3} - \frac{4}{3}\hat{\alpha} - \frac{4}{3}\hat{\beta} + \hat{\gamma} \right) k S_{12} - C_1 \beta^* \omega \tau_{12} \quad (2.19)$$

258

259

$$\frac{\partial \omega}{\partial t} = 2\alpha \frac{\omega}{k} \tau_{12} S_{12} - \beta \omega^2 \quad (2.20)$$

260

261

262

263

where (2.18) stems from the trace of (2.16). Notice that even in this reduced form the resulting Reynolds stress model differs fundamentally from a simpler  $k$ - $\omega$  turbulence model (see Larsen & Fuhrman 2018), with the Reynolds shear stress  $\tau_{12}$  governed by its own equation.

264

265

266

For analysis purposes, it turns out to be convenient to introduce a dimensionless utility variable  $\Psi = k/\tau_{12}$ . Combining (2.18) and (2.19), while also invoking  $\Psi$  into the  $\omega$  equation (2.20) then leads to:

267

$$\frac{\partial \Psi}{\partial t} = \underbrace{\left( \frac{4}{3}\hat{\alpha} + \frac{4}{3}\hat{\beta} - \hat{\gamma} - \frac{4}{3} \right)}_{-8/15} \Psi^2 S_{12} + (C_1 - 1)\beta^* \Psi \omega + 2S_{12} \quad (2.21)$$

268

269

$$\frac{\partial \omega}{\partial t} = -\beta \omega^2 + 2\alpha \frac{\omega}{\Psi} S_{12} \quad (2.22)$$

270

271

272

273

274

From inspection of (2.21) and (2.22) it is clear that, for any reasonable initial conditions i.e. with  $\tau_{12}$  (hence  $\Psi$ ) and  $S_{12}$  having the same sign, both  $\Psi$  and  $\omega$  will evolve asymptotically towards equilibrium values such that their respective time derivatives are zero. A brief mathematical analysis follows. Setting both (2.21) and (2.22) to zero, and solving for  $\Psi$  and  $\omega$  (discarding the unphysical solution with  $\omega = 0$ ) leads to the following asymptotic values



275 (so called fixed points):

$$276 \quad \Psi_{\infty} = \pm \sqrt{6 \cdot \frac{(1 - C_1)\alpha\beta^* - \beta}{\beta(4\hat{\alpha} + 4\hat{\beta} - 3\hat{\gamma} - 4)}} \approx \pm 2.394 \quad (2.23)$$

277

$$278 \quad \frac{\omega_{\infty}}{S_{12}} = \pm \alpha \sqrt{\frac{2}{3} \cdot \frac{4 - 4\hat{\alpha} - 4\hat{\beta} + 3\hat{\gamma}}{\beta^2 + (C_1 - 1)\alpha\beta\beta^*}} \approx \pm 6.135 \quad (2.24)$$

279 where the closure coefficients have been invoked to arrive at the constants. For positive  $S_{12}$ ,  
 280 the fixed point is  $(\Psi_{\infty}, \omega_{\infty}) = (2.394, 6.135S_{12})$ , while for negative  $S_{12}$ , the fixed point is  
 281  $(\Psi_{\infty}, \omega_{\infty}) = (-2.394, -6.135S_{12})$ .

282 Now let us check for formal stability of the fixed points based on the eigenvalues of the  
 283 Jacobian matrix for (2.21)–(2.22) which is defined by

$$284 \quad J = \begin{bmatrix} \frac{\partial}{\partial \Psi} \left( \frac{\partial \Psi}{\partial t} \right) & \frac{\partial}{\partial \omega} \left( \frac{\partial \Psi}{\partial t} \right) \\ \frac{\partial}{\partial \Psi} \left( \frac{\partial \omega}{\partial t} \right) & \frac{\partial}{\partial \omega} \left( \frac{\partial \omega}{\partial t} \right) \end{bmatrix} \quad (2.25)$$

285 After invoking the right-hand sides of (2.21)–(2.22) in the above, in addition to the model  
 286 closure coefficients, this becomes:

$$287 \quad J = \begin{bmatrix} -1.067S_{12}\Psi + 0.072\omega & 0.072\omega \\ -\frac{1.04S_{12}\omega}{\Psi^2} & -0.1416\omega + \frac{1.04S_{12}}{\Psi} \end{bmatrix} \quad (2.26)$$

288 By linearizing about (i.e. inserting) the fixed points  $(\Psi_{\infty}, \omega_{\infty})$ , the eigenvalues of  $J$  are  
 289 found to be  $(-1.99, -0.558)|S_{12}|$ . As these are negative, the fixed points correspond to  
 290 stable nodes (Strogatz 2018). This is also visually demonstrated for the positive quadrant by  
 291 the dimensionless stream plot of  $(1/|S_{12}|\partial\Psi/\partial t, 1/(S_{12}|S_{12}|\partial\omega/\partial t))$  in figure 1, depicting  
 292 evolution to a single point in the  $\omega/|S_{12}|$ - $\Psi$  plane, there indicated by the filled circle. The  
 293 plot with  $\Psi$  and  $S_{12}$  both having negative sign is symmetric to that shown in figure 1. This  
 294 behaviour has been confirmed through numerous numerical simulations of (2.18)–(2.20),  
 295 examples of which (with initial conditions for  $S_{12}$  and  $\tau_{12}$  having both positive and negative  
 296 values) are shown in figure 2. The asymptotic constants found above are likewise consistent  
 297 with figure 1.

298 Inserting the asymptotic values  $\Psi_{\infty}$  and  $\omega_{\infty}$  back into (2.18) and (2.19) and simplifying  
 299 then leads to linearized equations of the form

$$300 \quad \frac{1}{k} \frac{\partial k}{\partial t} = \frac{1}{\tau_{12}} \frac{\partial \tau_{12}}{\partial t} = \Gamma_{\infty} \quad (2.27)$$

301 where

$$302 \quad \Gamma_{\infty} = (\beta - \alpha\beta^*) \sqrt{\frac{2}{3} \cdot \frac{4 - 4\hat{\alpha} - 4\hat{\beta} + 3\hat{\gamma}}{\beta^2 + (C_1 - 1)\alpha\beta\beta^*}} \cdot |S_{12}| \approx 0.2831 \cdot |S_{12}| \quad (2.28)$$

303 defines the asymptotic exponential growth rate of both  $k$  and  $\tau_{12}$ .

304 It is seen from (2.28) that the exponential growth rate is expressed in terms of the strain-rate

$$305 \quad S_{12} = \frac{1}{2} \left( \frac{\partial u}{\partial y} + \frac{\partial v}{\partial x} \right) \quad (2.29)$$

306 which has been treated as fixed above at some unknown value for the sake of keeping the  
 307 analysis tractable. Note that this is entirely consistent with the prior analysis of the  $k$ - $\omega$

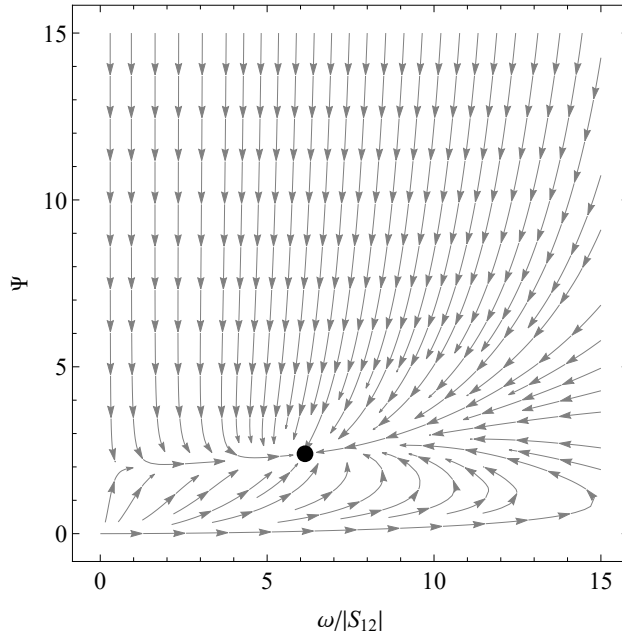


Figure 1: Dimensionless stream plot of  $\left(\frac{1}{|S_{12}|} \frac{\partial \Psi}{\partial t}, \frac{1}{S_{12}|S_{12}|} \frac{\partial \omega}{\partial t}\right)$  depicting the evolution of  $\Psi$  and  $\omega/|S_{12}|$  to a single point indicated by the filled circle.

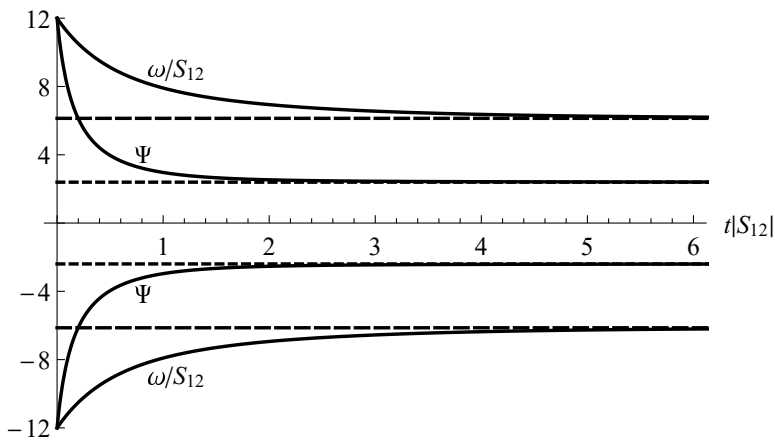


Figure 2: Simulated development (full lines) and predicted asymptotic values (dashed lines) of  $\Psi$  and  $\omega/S_{12}$  based on ODEs of (2.18)–(2.20) for the stress- $\omega$  closure model.  $S_{12}$  and  $\tau_{12}$  are provided with both positive and negative initial conditions.

308 model (and several other two-equation turbulence models) made by Larsen & Fuhrman  
 309 (2018), who similarly assumed their variable  $p_0 = 2S_{ij}S_{ij}$  to be fixed. This was interpreted in  
 310 practice e.g. as a period- and depth-averaged value beneath the considered surface wave field.  
 311 Adopting a similar approach, we therefore insert (2.14) and (2.15) into (2.29) and period

312 average. This leads to the rather trivial, but contextually important, result that

$$\begin{aligned}
 \langle S_{12} \rangle &= \frac{1}{T} \int_0^T S_{12} dt \\
 &= \frac{1}{T} \int_0^T \frac{1}{2} H k_w \sigma_\omega \cos(\sigma_\omega t - k_w x) \operatorname{csch}(h k_w) \sinh(k_w y) dt \\
 &= 0
 \end{aligned}
 \tag{2.30}$$

314 (where  $\langle \cdot \rangle$  indicates period-averaging), such that the exponential growth rate will, in fact, be  
 315 simply (on average) zero.

316 This thus proves that, under the simplifying assumptions made above, the Wilcox (2006)  
 317 stress- $\omega$  turbulence model is neutrally stable in the potential flow region beneath small  
 318 amplitude progressive waves. We find similarly for the LRR stress- $\varepsilon$  Reynolds stress model,  
 319 the details of which are again provided in Appendix B. These results are in contrast to the  
 320 authors' original expectations, based on the computational results of Brown *et al.* (2016).  
 321 The reason for this discrepancy is likewise explained in Appendix B. The present results are  
 322 also in stark contrast to similar analysis made for several two-equation models, most of which  
 323 have been proved to be either unconditionally unstable (Larsen & Fuhrman 2018) or (in the  
 324 special case of the realizable  $k$ - $\varepsilon$  model) conditionally unstable (Fuhrman & Li 2020), under  
 325 the same assumptions as considered here.

326 For the interested reader, an alternative analysis based on eigenvalues of the Jacobian matrix  
 327 for the governing equations (2.18)–(2.20), linearized about the fixed points, is presented in  
 328 Appendix C. The alternative analysis confirms the asymptotic growth rate found in (2.28),  
 329 and hence the finding of neutral stability above.

### 330 2.3. Comparison with analysis of two-equation models

331 Given the fundamental differences in the formal stability of Reynolds stress turbulence  
 332 models compared to their two-equation counterparts, it seems worthwhile to briefly revisit  
 333 the prior analysis of these simpler models to pinpoint precisely where these differences arise.  
 334 For this purpose, consider the  $k$  equation in (2.18), where the turbulence production term  
 335 corresponds to

$$336 \quad P_k = 2\tau_{12}S_{12}, \tag{2.31}$$

337 the form of which is theoretically based. With a Reynolds stress closure model,  $\tau_{12}$  is free to  
 338 evolve naturally based on its own transport equation (2.19). Conversely, with two-equation  
 339 closure models it is instead conventionally based on the Boussinesq approximation

$$340 \quad \tau_{ij} = 2\nu_t S_{ij} - \frac{2}{3} k \delta_{ij} \tag{2.32}$$

341 where  $\nu_t$  is the kinematic eddy viscosity. For the conditions specifically analyzed in Section  
 342 2.2, (2.32) leads to the Reynolds stress  $\tau_{12} = 2\nu_t S_{12}$ , such that the turbulence production  
 343 term becomes

$$344 \quad P_k = p_0 \nu_t, \quad p_0 = 4S_{12}S_{12} \tag{2.33}$$

345 i.e. proportional to  $p_0$  rather than simply  $S_{12}$ . Similarly, in their analysis of standard two-  
 346 equation models, Larsen & Fuhrman (2018) showed that they inevitably lead to asymptotic  
 347 values of  $\omega_\infty$  and  $\Gamma_\infty$  that are both proportional to  $\sqrt{p_0}$ , rather than  $S_{12}$ . Critically in the  
 348 present context, in the potential flow region beneath surface waves  $\langle p_0 \rangle$  is finite (Mayer &  
 349 Madsen 2000; Larsen & Fuhrman 2018), rather than zero as is the case for  $\langle S_{12} \rangle$ , see (2.30).

350 Thus, this clarifies that it is the Boussinesq approximation of the Reynolds shear stress  
 351 in two-equation turbulence closure models that is responsible for their formal instability in

352 the potential flow region beneath surface waves. Notably, this finding lends credence to the  
 353 approach adopted by Larsen & Fuhrman (2018), who utilized a re-formulated eddy viscosity  
 354 (to include an additional stress-limiting feature) in order to formally stabilize such closures.

### 355 3. CFD simulations with the Wilcox (2006) stress- $\omega$ model

356 This section will present a series of CFD simulations, where the Wilcox (2006) stress- $\omega$   
 357 model is used as turbulence closure for a numerical model solving incompressible Reynolds-  
 358 averaged Navier-Stokes (RANS) equations. The selected simulations will build towards the  
 359 ultimate aim of accurately simulating breaking surface waves with significantly improved  
 360 accuracy compared to existing two-equation closures. Specifically, Section 3.1 will consider  
 361 simulation of a simple progressive non-breaking wave train, as a direct test of the model's  
 362 stability in the potential flow core region (as analyzed in the preceding section). We will  
 363 then focus on simulation of the turbulent wave boundary layer in Section 3.2, of fundamental  
 364 interest beneath both non-breaking and breaking waves. This section will finally culminate  
 365 with CFD simulations of both spilling (Section 3.3) and plunging (Section 3.4) breaking  
 366 waves. All simulations in the present work have been carried out within the OpenFOAM®  
 367 v1812 framework. Free surface simulations utilize the waves2FOAM toolbox (Jacobsen *et al.*  
 368 2012) for wave initiation or generation and absorption.

369 The free surface is modelled using the volume of fluid (VOF) method, and the phases  
 370 in terms of the two fluids (i.e. air and water) are tracked by a scalar field  $\gamma$ , where  $\gamma = 0$   
 371 denotes pure air and  $\gamma = 1$  denotes pure water. Any intermediate  $\gamma$  value between 0 and 1  
 372 represents a fluid mixture. The  $\gamma$  field is governed by the advection equation (see also Sumer  
 373 & Fuhrman 2020, p. 558):

$$374 \quad \frac{\partial \gamma}{\partial t} + \frac{\partial(\bar{u}_i \gamma)}{\partial x_i} + \frac{\partial[\bar{u}_i^r \gamma(1 - \gamma)]}{\partial x_i} = 0 \quad (3.1)$$

375 where  $\bar{u}_i^r$  is a relative velocity for interface compression according to Berberović *et al.* (2009).  
 376 Any fluid property (represented by  $\Phi$ ) is calculated by:

$$377 \quad \Phi = \gamma \Phi_{\text{water}} + (1 - \gamma) \Phi_{\text{air}} \quad (3.2)$$

378 i.e. fluid properties are weighted linearly based on the local value of  $\gamma$ . For modelling  
 379 the free-surface of breaking waves with strong turbulence, Brocchini & Peregrine (2001)  
 380 and Brocchini (2002) also proposed an approach using averaged equations (i.e. mass and  
 381 momentum conservation equations along with an equation for the turbulent kinetic energy),  
 382 with boundary conditions obtained through integration across the two-phase surface layer.  
 383 This may provide a useful alternative for modelling the disturbed free-surface of breaking  
 384 waves, though this approach will not be pursued here.

#### 385 3.1. Simulating a progressive wave train

386 The stability analysis in Section 2.2 demonstrates that the Wilcox (2006) stress- $\omega$  model  
 387 is neutrally stable in the ideal potential flow region beneath surface waves. This is again  
 388 in contrast to our original suspicions, since the Reynolds-stress CFD simulations of Brown  
 389 *et al.* (2016) demonstrated turbulence over-production prior to breaking. As an initial test  
 390 to confirm our stability analysis, we therefore conduct CFD simulations involving the  
 391 simple propagation of a theoretically (based on potential flow theory) steady wave train.  
 392 For comparative purposes, two simulations will be considered, having buoyancy production  
 393 either on ( $\alpha_b^* = 1.36$ , as indicated in Section 2.1) or off ( $\alpha_b^* = 0$ ). The reason for this  
 394 comparison is to elucidate any effects of the buoyancy production term (which will cause a

395 sink of turbulence near the air-water interface), since it was not considered in the stability  
 396 analysis for reasons of simplicity.

397 Following Larsen & Fuhrman (2018), we adopt the wave properties associated with the  
 398 incident wave from the spilling breaker experiments of Ting & Kirby (1994) for the present  
 399 simulations, corresponding to period  $T = 2$  s and wave height  $H = 0.125$  m on a constant  
 400 water depth  $h = 0.4$  m. The numerically exact stream function wave (potential flow) solution  
 401 of Fenton (1988) (as implemented by Jacobsen *et al.* 2012), yields the dimensionless depth  
 402  $k_w h = 0.664$  and steepness  $k_w H = 0.207$ . This wave solution is set as the initial conditions  
 403 on a domain spanning a single wave length with periodic left and right boundaries. An  
 404 initially small turbulence field is set with  $\tau_{11} = \tau_{22} = \tau_{33} = -\tau_{12} = -1.33 \times 10^{-6} \text{ m}^2 \text{ s}^{-2}$ , such  
 405 that the initial turbulent kinetic energy  $k_0$  is  $2.0 \times 10^{-6} \text{ m}^2 \text{ s}^{-2}$ . The setup utilized (including  
 406 mesh, discretization schemes, and multi-phase flow solver) is adopted directly from Larsen  
 407 & Fuhrman (2018), who performed similar tests utilizing two-equation ( $k$ - $\omega$ ) turbulence  
 408 models. Specifically, the maximum Courant number is set to  $Co = 0.05$ , and a diffusive  
 409 balance scheme as discussed in Larsen *et al.* (2019) is adopted. The bottom boundary is  
 410 modelled as a slip wall, to mimic potential flow as much as possible.

411 Figure 3(a,b) depicts time series of the dimensionless surface elevation as well as the  
 412 period- and depth-averaged (note that  $[\cdot]$  herein indicates depth-averaging) turbulence level,  
 413 respectively, over a simulated duration of  $100T$ . It is seen in figure 3(a) that the wave  
 414 propagates with nearly constant form in both cases (the two results for the free surface  
 415 elevations are indistinguishable). It is seen from figure 3(b) that the case with  $\alpha_b^* = 0$  results  
 416 in a growth rate in the turbulent kinetic energy that may indeed be reasonably characterized as  
 417 zero. This result is consistent with our simplified analysis of this problem in Section 2.2, again  
 418 predicting that the model is neutrally stable. Minor deviations (e.g. the initial slow decay  
 419 and later rise of  $[\langle k/k_0 \rangle]$ ) are relatively insignificant, and are likely due to terms neglected  
 420 in the analysis and/or from accumulation of small numerical errors, which may cause the  
 421 solution to deviate from the ideal potential flow solution over extended times. It is likewise  
 422 seen from figure 3(b) that the buoyancy production term being active instead leads to a decay  
 423 in turbulence levels. This is clearly due to the additional sink in turbulence caused by this  
 424 term near the air-water interface, which was not considered in the formal stability analysis.  
 425 Hence, both simulations largely confirm our analysis, that the Wilcox (2006) stress- $\omega$  model  
 426 is indeed stable in the ideal potential flow core region beneath non-breaking surface waves.  
 427 Note that both results presented in figure 3 differ considerably from the simulation using the  
 428 Wilcox (2006)  $k$ - $\omega$  closure model in its standard form, as presented in figure 4(a) of Larsen  
 429 & Fuhrman (2018), which resulted in immediate exponential growth of the eddy viscosity  
 430 (hence turbulence) and eventual wave decay, due to this model's inherent instability, as shown  
 431 and discussed therein.

### 432 3.2. Simulating the oscillatory turbulent wave boundary layer

433 We will now turn our attention to the performance of the Wilcox (2006) stress- $\omega$  model in  
 434 the bottom boundary layer region beneath waves, an area of special importance beneath both  
 435 non-breaking and breaking waves. (Recall that this region was neglected in the previous wave  
 436 train simulations due to the use of a slip condition at the sea bed.) For this purpose, we will  
 437 consider the experiments of Jensen *et al.* (1989) conducted in a full-scale oscillating tunnel  
 438 facility. We will specifically consider their Test 13, involving the boundary layer beneath a  
 439 sinusoidally varying free stream flow (having velocity magnitude  $U_{0m} = 2.0$  m/s and period  
 440  $T = 9.72$  s) yielding a Reynolds number  $Re = aU_{0m}/\nu \approx 6 \times 10^6$ , where  $a = U_{0m}/\sigma_w$  and  
 441  $\nu = 1.14 \times 10^{-6} \text{ m}^2/\text{s}$ . The bottom wall is rough, with Nikuradse's equivalent roughness  
 442  $k_s = 0.84$  mm. A model height of 0.145 m corresponding to half of the physical tunnel  
 443 height (0.29 m) in Jensen *et al.* (1989) is used, hence only the bottom boundary layer is

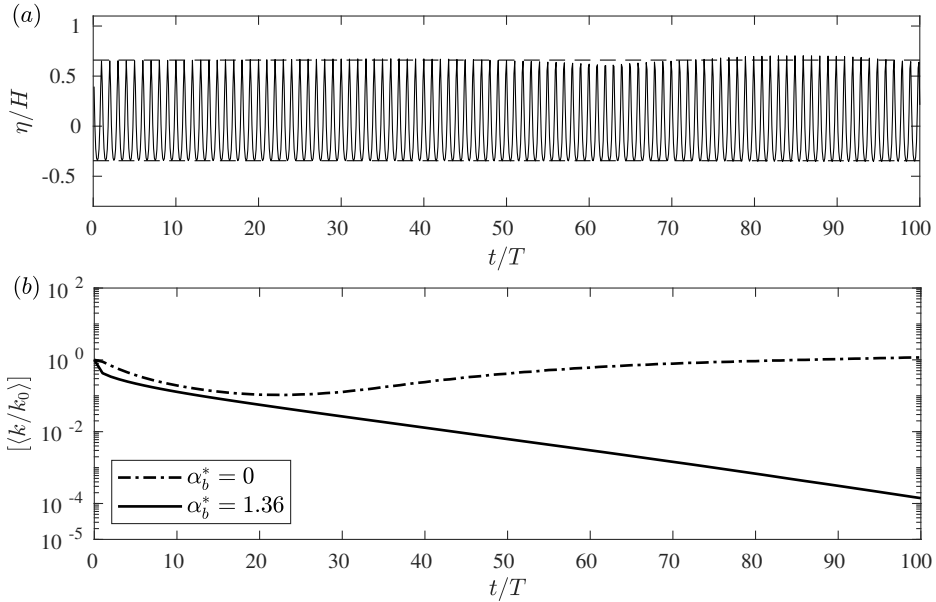


Figure 3: Computed (a) surface elevation time series and (b) the time- and depth-averaged turbulence level for the progressive waves with the Wilcox (2006) stress- $\omega$  model, with buoyancy production term both off ( $\alpha_b^* = 0$ ) and on ( $\alpha_b^* = 1.36$ ).

444 simulated. The top boundary is treated as a frictionless (slip) lid. The bottom boundary is  
 445 set as a no slip wall, where the  $\omega$  wall function with a viscous-inertial sublayer blending  
 446 method (Menter & Esch 2001; Popovac & Hanjalic 2007) is applied, combined with a zero  
 447 normal gradient condition for the Reynolds stress. The first cell center near the bottom wall  
 448 lies at  $y_c/k_s = 0.5$ . An oscillatory body force is applied to drive the flow until an equilibrium  
 449 (periodic in time) state is reached and comparisons are made.

450 Computed and experimental results are compared in figure 4 at four phases during the  
 451 oscillation cycle:  $\sigma_{wt} = 0^\circ$  (free stream flow reversal),  $45^\circ$  (flow acceleration due to a  
 452 favorable pressure gradient),  $90^\circ$  (peak free stream flow) and  $135^\circ$  (flow deceleration due  
 453 to an adverse pressure gradient). Results are shown for the dimensionless mean flow  $\bar{u}/U_{0m}$   
 454 (figure 4a); the turbulent kinetic energy density  $k/U_{0m}^2$  (figure 4b), which for the experiments  
 455 of Jensen *et al.* (1989) has been empirically approximated from (Justesen 1991):

$$456 \quad k = -0.65(\tau_{11} + \tau_{22}); \quad (3.3)$$

457 as well as the Reynolds stress components:  $-\tau_{11}/U_{0m}^2$ ,  $-\tau_{22}/U_{0m}^2$ , and  $\tau_{12}/U_{0m}^2$  (figure 4c,d,  
 458 and e, respectively). Results computed utilizing both the Wilcox (2006) stress- $\omega$  and  $k$ - $\omega$   
 459 models are shown, such that those of the Reynolds stress model (the primary focus of the  
 460 present work) may be compared directly with a simpler two-equation model. Note that for  
 461 the  $k$ - $\omega$  model, the Reynolds stress components are obtained directly from the Boussinesq  
 462 approximation (2.32).

463 From figure 4(a) it is seen that the computed mean flow velocities from both models are  
 464 largely similar, and in good agreement with the experiments. The most notable difference is  
 465 the slight reduction (and increased accuracy) in the mean flow computed with the stress- $\omega$   
 466 model at phase  $\sigma_{wt} = 135^\circ$  (i.e. during adverse pressure and flow deceleration), relative to  
 467 the  $k$ - $\omega$  model. This difference will be explained immediately below. It is seen in figure 4(b)  
 468 that the stress- $\omega$  model obviously improves the accuracy of the turbulence kinetic energy

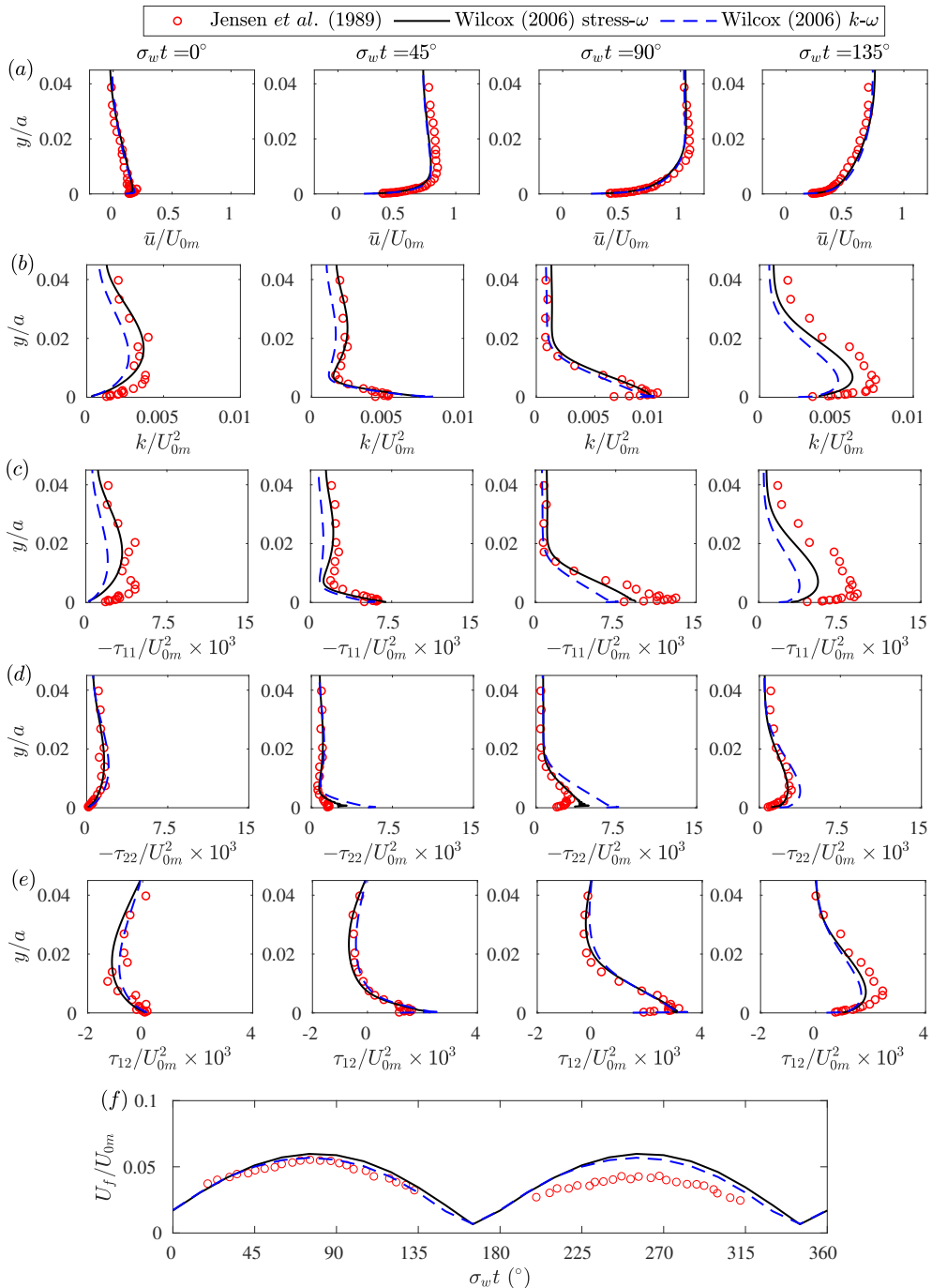


Figure 4: Comparison of computed and measured vertical profiles for (a)  $\bar{u}/U_{0m}$ , (b)  $k/U_{0m}^2$ , (c)  $-\tau_{11}/U_{0m}^2$ , (d)  $-\tau_{22}/U_{0m}^2$ , (e)  $\tau_{12}/U_{0m}^2$ , and (f)  $U_f/U_{0m}$  for the oscillatory wave boundary layer case of Jensen *et al.* (1989, their Test 13). The depicted CFD simulations utilize both the Wilcox (2006) stress- $\omega$  and  $k$ - $\omega$  turbulence models.

469  $k$ , relative to the  $k$ - $\omega$  model, especially at phase  $\sigma_w t = 135^\circ$ . Note that Sumer & Fuhrman  
 470 (2020) have similarly documented relatively poor performance of the Wilcox (2006)  $k$ - $\omega$   
 471 model in simulating the deceleration stage of the wave boundary layer (see their figures  
 472 5.90–5.92), and this is a well-known deficiency with two-equation models in general (see  
 473 e.g. Justesen 1991, for similar finding with a  $k$ - $\varepsilon$  closure model). From inspection of the  
 474 results just discussed in figure 4(a), it is clear that the over-prediction of  $\bar{u}$  seen there with  
 475 the  $k$ - $\omega$  model is associated with its under-prediction of  $k$  at this phase i.e. that the  $k$ - $\omega$   
 476 model does not extract enough energy from the mean flow during the flow deceleration  
 477 stage. Since the form of the turbulence production term in the  $k$  equation ( $\tau_{ij}\partial\bar{u}_i/\partial x_j$  which  
 478 simplifies to  $\tau_{12}\partial\bar{u}/\partial y$  in the present horizontally-uniform case) is theoretical (hence exact  
 479 if its determination is free of error), it is then clear that this shortcoming must be due to  
 480 inaccuracy of  $\tau_{12}$  from the Boussinesq approximation (2.32).

481 The individual Reynolds stress component profiles at each stage are presented in figure  
 482 4(c–e). It is seen that the stress- $\omega$  model captures both the dynamics of the turbulent normal  
 483 and shear stress components with better accuracy compared to the  $k$ - $\omega$  model, although  $\tau_{11}$   
 484 and  $\tau_{12}$  at  $\sigma_w t = 135^\circ$  are still slightly under-predicted in the near bottom region. It is seen  
 485 in figure 4(c,d) that  $\tau_{11}$  and  $\tau_{22}$  predicted by the  $k$ - $\omega$  model (blue dashed lines) are identical  
 486 and deviate from the experimental measurements. This is simply because application of the  
 487 Boussinesq approximation (2.32) for the present case leads simply to:

$$488 \quad \tau_{11} = \tau_{22} = -\frac{2}{3}k, \quad \tau_{12} = \nu_t \frac{\partial\bar{u}}{\partial y} \quad (3.4)$$

489 the former of which is well-known to be incorrect, even in the simpler case of a steady  
 490 horizontally uniform turbulent boundary layer flow, see e.g. Chapter 3 of Sumer & Fuhrman  
 491 (2020). In line with the discussion above, it is notable that  $\tau_{12}$  (figure 4e) is indeed under-  
 492 predicted by the  $k$ - $\omega$  model at  $\sigma_w t = 135^\circ$ . Overall, the Wilcox (2006) stress- $\omega$  model is  
 493 demonstrated to be superior to the  $k$ - $\omega$  model in simulating the turbulence dynamics for the  
 494 oscillatory wave boundary layer flows, as measured by Jensen *et al.* (1989).

495 The measured and modelled friction velocity  $U_f$  is likewise presented in figure 4(f). In the  
 496 experiment of Jensen *et al.* (1989), the friction velocity was determined by fitting straight  
 497 lines to the logarithmic-layer portion of the mean velocity distribution (see Sumer & Fuhrman  
 498 2020, Section 5.4.1). It is noted that the difference in the measurements for two half cycles are  
 499 quite obvious, and are due to apparent asymmetries that occurred in the experiment, which  
 500 are avoided in the numerical simulations. It is seen that both stress- $\omega$  and  $k$ - $\omega$  model results  
 501 match the friction velocity closely for the first half cycle. The friction velocity simulated with  
 502 the stress- $\omega$  model is identical to that with the  $k$ - $\omega$  model in the flow acceleration stage, while  
 503 being slightly larger than the  $k$ - $\omega$  model in the peak and deceleration stages. As the difference  
 504 in the measurements over the two half cycles is larger than that of the two numerical results,  
 505 both numerical model results are considered acceptable.

### 506 *3.3. Simulating spilling breaking waves*

507 The preceding preliminary simulations have demonstrated potential advantages of using a  
 508 stress- $\omega$  model (rather than a traditional two-equation  $k$ - $\omega$  turbulence closure) for applica-  
 509 tions relevant to non-breaking waves, ranging from the free surface (the progressive wave  
 510 train) to the sea bottom (the turbulent wave boundary layer). Let us now apply the model  
 511 to simulate breaking wave hydrodynamics, the primary aim of the present paper. For this  
 512 purpose, we will first consider the spilling breaking wave experiment of Ting & Kirby  
 513 (1994, 1996), to be followed by their plunging breaking wave experiment in the following  
 514 sub-section.

515 The numerical set-up for simulation of the experiments of Ting & Kirby (1994, 1996)



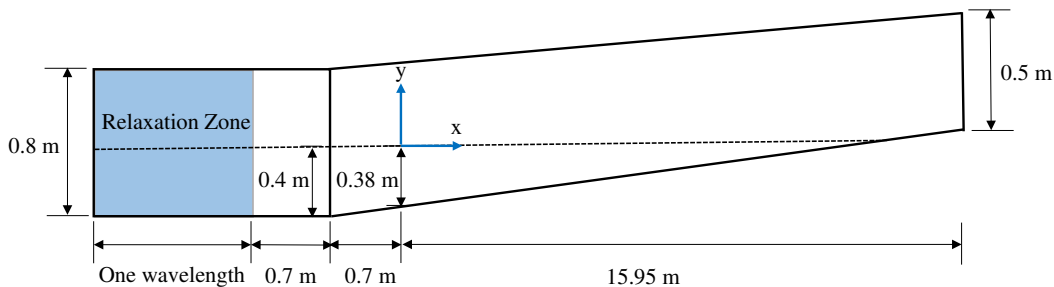


Figure 5: Computational domain set-up for plunging and spilling breaker cases corresponding to Ting & Kirby (1994)'s breaking wave experiments.

Case	$T$ (s)	$H$ (m)	$h_0$ (m)	$x_b$ (m)	$k_w H$	$k_w h$	$\xi_0$
Spilling	2	0.125	0.4	6.400	0.208	0.664	0.20
Plunging	5	0.128	0.4	7.795	0.076	0.238	0.60

Table 1: Wave properties for breaking wave simulations. In the above  $x_b$  is the measured breaking point in Ting & Kirby (1994), and  $\xi_0$  is the surf similarity parameter.

516 is shown in figure 5, where a  $\tan \beta = 1/35$  constant slope is connected to a region having  
 517 constant still water depth  $h_0 = 0.4$  m. The origin is placed at the same water depth ( $h = 0.38$   
 518 m) as in the experiments, for consistency. A relaxation zone (Jacobsen *et al.* 2012) of one  
 519 wave length is set at the inlet for wave generation, which also serves to absorb any reflected  
 520 waves. A no slip condition along with standard smooth bed wall functions are employed  
 521 as the bottom boundary conditions, since in the experiments of Ting & Kirby (1994) and  
 522 Ting & Kirby (1996) a roughness value was not explicitly indicated. The computational  
 523 mesh utilized is identical to that used previously by Larsen & Fuhrman (2018). Dimensional  
 524 and dimensionless wave properties utilized for the simulation of both spilling and plunging  
 525 breaking wave cases are indicated in table 1, where a numerically exact stream function  
 526 (potential flow) theory is used for specification of the generated wave at the inlet. In table 1  
 527  $x_b$  denotes the position of incipient breaking and

$$528 \quad \xi_0 = \frac{\tan \beta}{\sqrt{H_0/L_0}} \quad (3.5)$$

529 is the surf similarity parameter, where  $L_0 = gT^2/(2\pi)$  is the deep-water wave length and

$$530 \quad H_0 = H \sqrt{\tanh(k_w h) \left( 1 + \frac{2k_w h}{\sinh(2k_w h)} \right)} \quad (3.6)$$

531 is the deep-water wave height, calculated according to linear wave theory. The breaking wave  
 532 simulations are initially run for  $50T$  to reach equilibrium, followed by a subsequent  $50T$   
 533 which is utilized for period-averaging purposes. The simulated spilling breaking case with  
 534 the stress- $\omega$  model required approximately 12 days to run in parallel on 16 processors on the  
 535 supercomputing cluster at the Technical University of Denmark (DTU). Note that the total  
 536 computational time using the stress- $\omega$  model is approximately 15% more than that using the  
 537  $k$ - $\omega$  model.

538 To elucidate differences between the Wilcox (2006) stress- $\omega$  and two-equation  $k$ - $\omega$   
 539 turbulence closure models, simulations utilizing a stabilized version of the Wilcox (2006)  $k$ - $\omega$

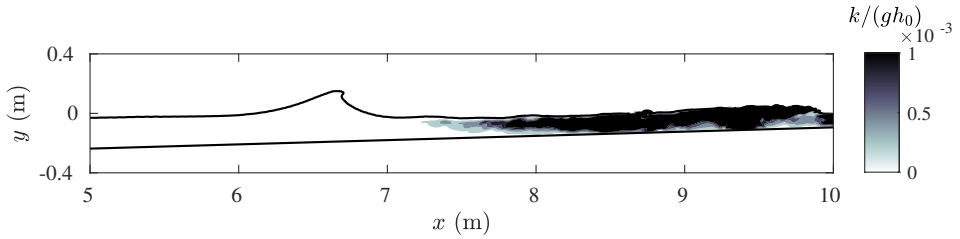


Figure 6: Snapshot of the spilling breaker turbulent kinetic energy simulated with the Wilcox (2006) stress- $\omega$  model at  $t/T = 100$ .

540 model, as proposed by Larsen & Fuhrman (2018, with stress-limiter coefficients  $\lambda_1 = 0.2$  and  
 541  $\lambda_2 = 0.05$ , as suggested there and in their notation), will also be considered for comparison.  
 542 This model will hereafter be called the LF18  $k$ - $\omega$  model. Note that results based on the LF18  
 543  $k$ - $\omega$  model have been re-simulated for presentation herein, to ensure full consistency with the  
 544 stress- $\omega$  results. This ensures that any effects associated e.g. with the specific OpenFOAM  
 545 software version or boundary treatment are fully controlled for. (Such effects have not been  
 546 found to be very significant, but this accounts for subtle differences in the results presented  
 547 herein compared to those originally presented in Larsen & Fuhrman 2018).

548 To begin our investigation, figure 6 depicts a snapshot of the spilling breaker turbulent  
 549 kinetic energy (here presented dimensionless as  $k/(gh_0)$  where  $h_0 = 0.4$  m is the constant still  
 550 water depth prior to the slope) simulated with the Wilcox (2006) stress- $\omega$  model at  $t/T = 100$ .  
 551 It is observed that there is no sign of turbulence over-production prior to breaking, indicating  
 552 that the Wilcox (2006) stress- $\omega$  model is indeed stable i.e. free of unphysical exponential  
 553 growth of turbulence in nearly potential flow regions. This is once again consistent with  
 554 our analysis of this model (Section 2) as well as our previous CFD simulations involving a  
 555 progressive wave train (Section 3.1). The present result is in stark contrast to those stemming  
 556 from two-equation models (both  $k$ - $\omega$  and  $k$ - $\epsilon$  variants) in their standard forms, see e.g. Brown  
 557 *et al.* (2016), Larsen & Fuhrman (2018, their figure 6a,b), Larsen *et al.* (2020) and Fuhrman  
 558 & Li (2020, their figure 7a).

559 Figure 7 shows the surface elevation envelopes for the spilling breaker simulations, where  
 560  $\langle \eta \rangle$  is the period-averaged mean water level (over the final  $50T$ ), and  $\eta_{max}$  and  $\eta_{min}$  are  
 561 respectively the averaged maximum and minimum surface elevations. Results from both the  
 562 stress- $\omega$  and LF18  $k$ - $\omega$  models are shown separately. The grey shaded regions depict plus and  
 563 minus one standard deviation, hence indicating the degree of wave-to-wave variability. Good  
 564 agreement is observed in figure 7(a) between the simulation with Wilcox (2006) stress- $\omega$   
 565 model and the measurements of Ting & Kirby (1994). The predicted breaking point (where  
 566  $\eta_{max} - \langle \eta \rangle$  is the highest) is consistent with the experimental measurement. The surface  
 567 elevation envelopes predicted by the LF18  $k$ - $\omega$  model are also similarly in line with the  
 568 experimental measurement (figure 7b), consistent with previous demonstrations.

569 Figure 8(a-d) compares the computed phase-averaged surface elevations with the exper-  
 570 imental measurements of Ting & Kirby (1994) at four post-breaking cross-shore locations,  
 571 where  $\bar{\eta}$  denotes the phase-averaged surface elevation and  $\langle \eta \rangle$  denotes the period-averaged  
 572 surface elevation. Additionally, the two model results are compared even further onshore  
 573 ( $x = 9.725$  m) in figure 8(e), for completeness. (Although the phase-averaged surface  
 574 elevations from the experiments were not directly reported at this position, velocity and  
 575 turbulence profiles were, to be presented in what follows.) It is seen that the numerical  
 576 predictions with both turbulence models are generally in line with the experimental data  
 577 for the three positions furthest offshore (figure 8a-c). Further onshore, the stress- $\omega$  model  
 578 maintains this accuracy. However, it is seen in figure 8(d,e) that the wave front computed

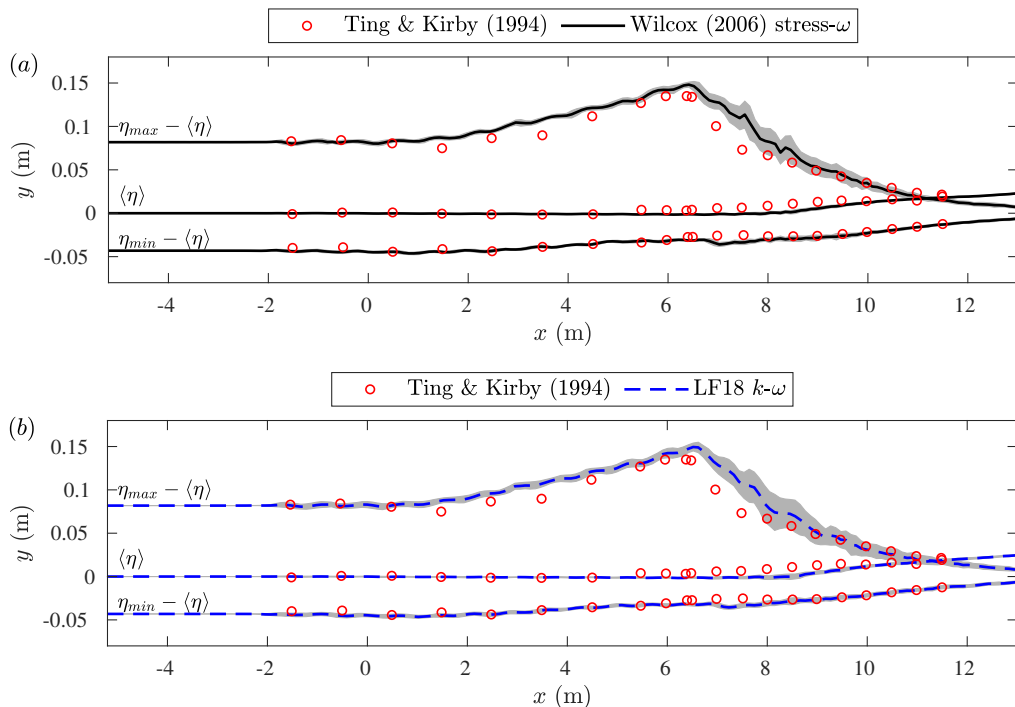


Figure 7: Period-averaged surface elevation envelopes for the spilling breaker simulated with (a) the Wilcox (2006) stress- $\omega$  model and (b) the LF18  $k$ - $\omega$  model, comparing to the experimental measurement of Ting & Kirby (1994). Grey shaded areas are the plus and minus one standard deviation.

579 with the LF18  $k$ - $\omega$  model is well ahead of what was measured. This was also noticed by  
 580 Larsen & Fuhrman (2018), indicating that the breaking bore travels too rapidly in the inner  
 581 surf zone. Larsen & Fuhrman (2018, see their figure 10) showed clearly that this problem was  
 582 due to the conventional stress-limiter on the eddy viscosity (controlled by the  $\lambda_1$  coefficient  
 583 in their notation) within the Wilcox (2006)  $k$ - $\omega$  model. Simulations where this feature was  
 584 on ( $\lambda_1 > 0$ ) resulted in significantly improved results (in terms of undertow velocity and  
 585 turbulence profiles) in the outer surf zone, but at the expense of reduced accuracy in the inner  
 586 surf zone. The stress- $\omega$  model, on the other hand, breaks free of the eddy viscosity concept  
 587 altogether, and hence avoids this issue entirely.

588 Let us now turn our attention to the turbulence quantities beneath the spilling breaking  
 589 waves. Ting & Kirby (1994, 1996) have reported results for  $\sqrt{\langle k \rangle}$ ,  $\langle \sqrt{-\tau_{11}} \rangle$  and  $\langle \tau_{22} \rangle / \langle \tau_{11} \rangle$  at  
 590 each measurement position. Although the measurements for  $\langle \tau_{22} \rangle$  were not directly reported,  
 591 they can be obtained from their reported  $\sqrt{\langle k \rangle}$  and  $\langle \tau_{22} \rangle / \langle \tau_{11} \rangle$  values. In Ting & Kirby (1994),  
 592 because the transverse velocity component was not measured,  $k$  was estimated empirically  
 593 by

$$594 \quad \langle k \rangle = \frac{1.33}{2} (\langle \tau_{11} \rangle + \langle \tau_{22} \rangle), \quad (3.7)$$

595 which is also utilized for the experimental  $k$  values presented in what follows. For the LF18  
 596  $k$ - $\omega$  model, the Reynolds stress components are again obtained directly from the Boussinesq  
 597 approximation (2.32).

598 Figures 9–10 compare specific period-averaged Reynolds normal stresses (non-  
 599 dimensionalized  $-\tau_{11} = \overline{u'u'}$  and  $-\tau_{22} = \overline{v'v'}$  period-averaged over the final simulated 50T;

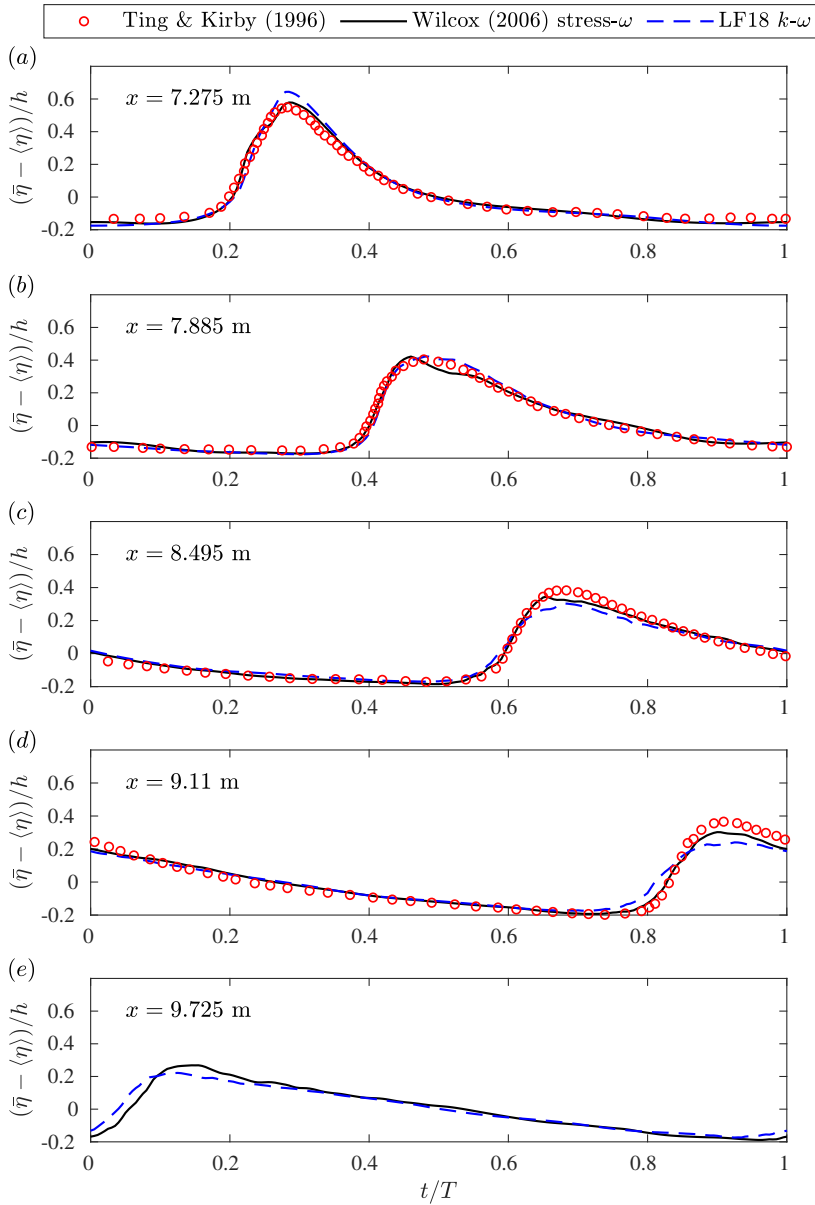


Figure 8: Phase-averaged surface elevation for the spilling breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations. The subfigures (a, b, c) are in the outer surf zone while (d, e) are in the inner surf zone.

600 results are similarly period-averaged in several forthcoming figures) profiles at each of the  
 601 measured cross-shore positions. Figure 11 similarly presents a comparison of computed and  
 602 measured period-averaged turbulent kinetic energy density  $k$  profiles. From these figures,  
 603 it can be surmised that both the Wilcox (2006) stress- $\omega$  model and the LF18  $k$ - $\omega$  model  
 604 predict Reynolds normal stress components that are reasonably, though not perfectly, in  
 605 line with the measurements. It is noted that the stress- $\omega$  model predicts streamwise normal  
 606 stresses ( $\tau_{11}$ ) significantly better than vertical ones ( $\tau_{22}$ ). This might be attributed to the

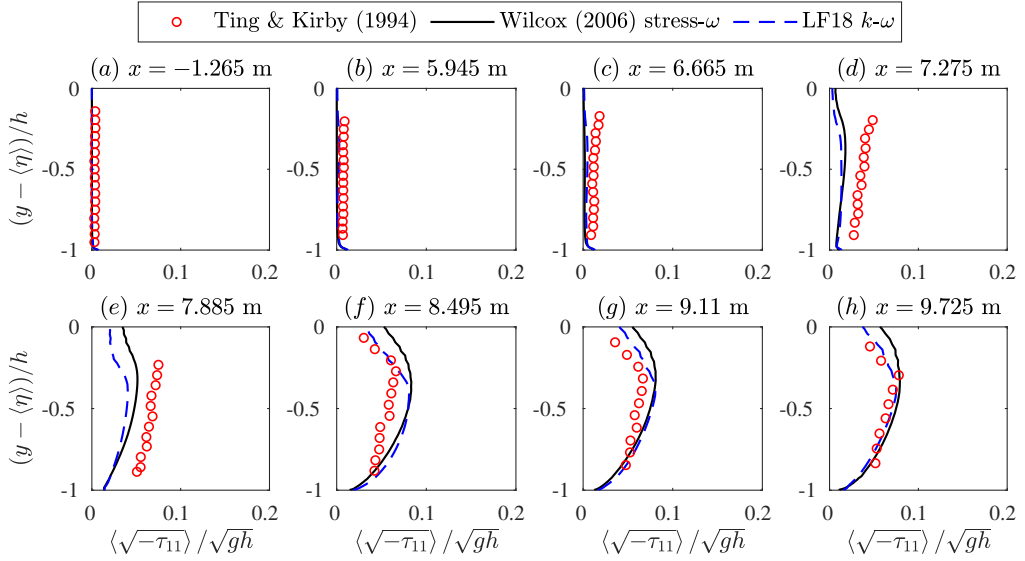


Figure 9: Period-averaged Reynolds normal stress  $-\tau_{11}$  for the spilling breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations. The subfigures (a, b) are in the pre-breaking region while (c, d, e) are in the outer surf zone and (f, g, h) are in the inner surf zone.

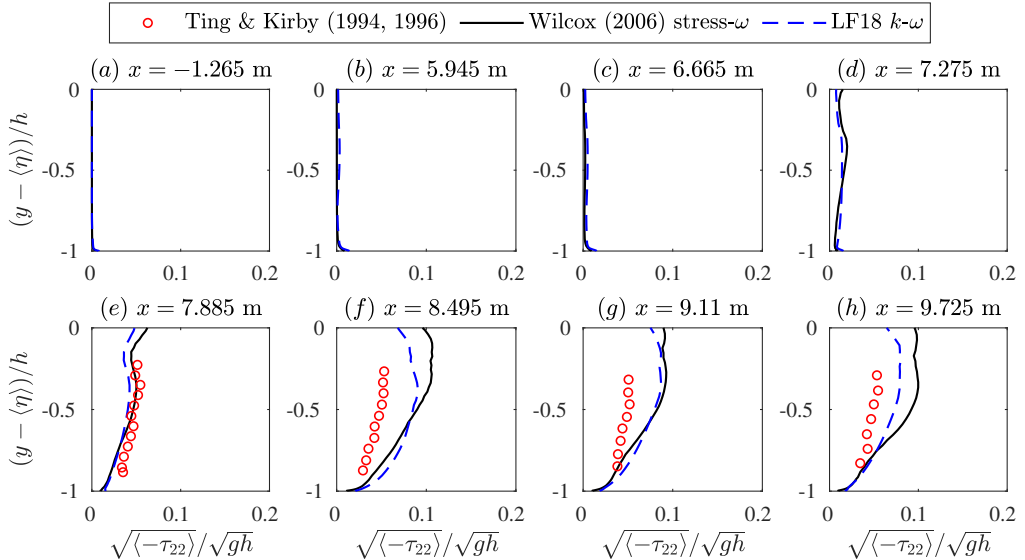


Figure 10: Period-averaged Reynolds normal stress  $-\tau_{22}$  for the spilling breaker from the experimental measurement of Ting & Kirby (1994, 1996) and the present simulations.

607 simple formulation of pressure-strain closure in the Wilcox (2006) stress- $\omega$  model, as the  
 608 streamwise normal stresses ( $\tau_{11}$ ) are dominated by the production term  $P_{11}$  while  $\tau_{22}$  is  
 609 mainly driven by the pressure-strain correlation  $\Pi_{22}$ . It is seen in figure 11(d-h) that there  
 610 is also a tendency for the LF18  $k$ - $\omega$  model to predict more accurate turbulence near the  
 611 free surface, where the stress- $\omega$  model predicts slightly higher turbulence than the  $k$ - $\omega$   
 612 model. This can also be attributed to the standard Wilcox (2006) stress-limiting feature in

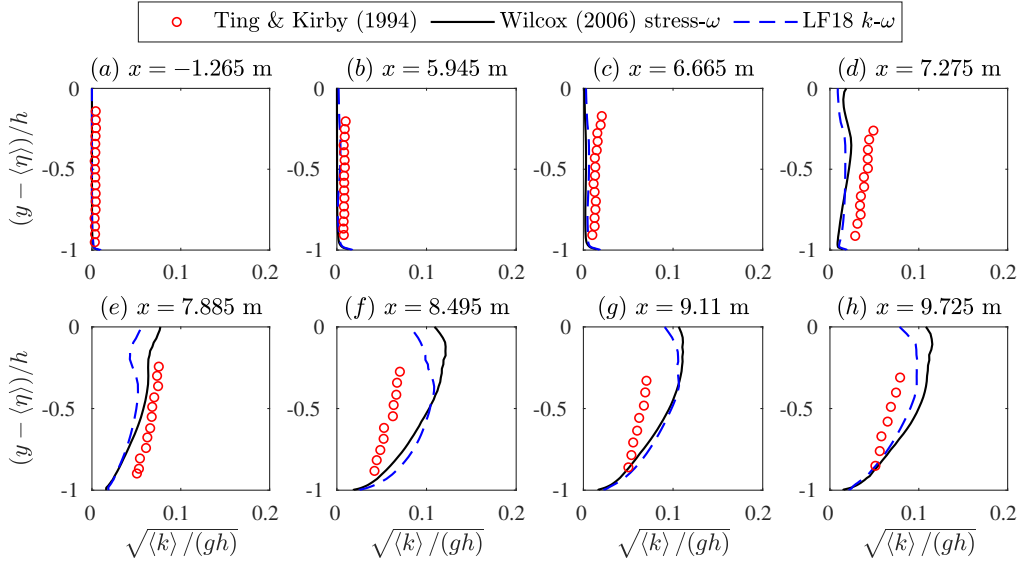


Figure 11: Period-averaged turbulent kinetic energy  $k$  for the spilling breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations.

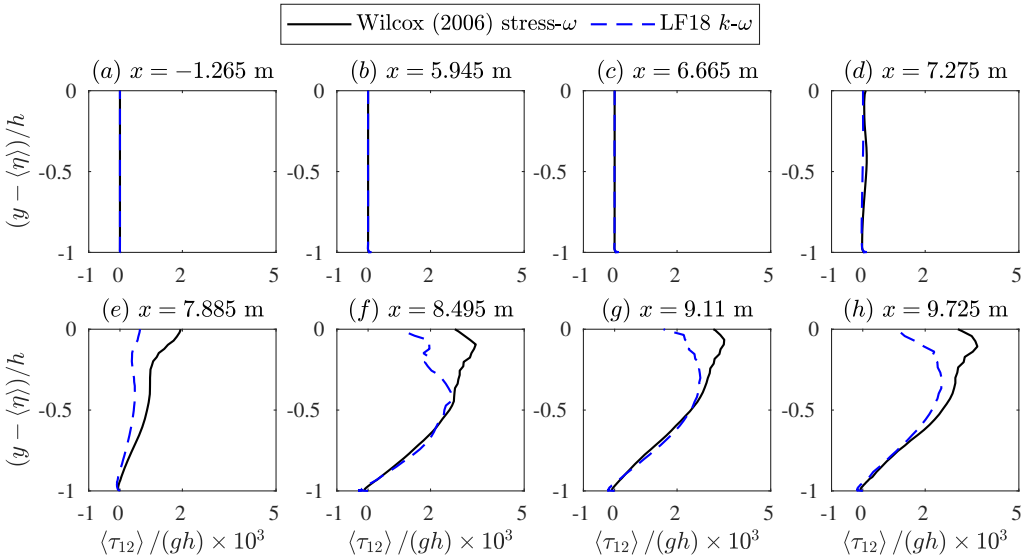


Figure 12: Period-averaged specific Reynolds shear stress  $\tau_{12}$  for the spilling breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations.

613 the  $k$ - $\omega$  model, as shown through systematic testing by Larsen & Fuhrman (2018, compare  
614 e.g. Cases 3 and 5 in their figure 12).

615 Let us now similarly investigate the computed Reynolds shear stresses  $\tau_{12} = -\overline{u'v'}$ , which  
616 can be expected to play a much more important role in terms of flow resistance than the  
617 turbulent normal stresses. Figure 12 compares the period-averaged  $\tau_{12}$  (again over the final  
618 simulated  $50T$ ) profiles from both models at all eight measurement positions considered  
619 previously. Note that this quantity was not reported by Ting & Kirby (1994), thus we are  
620 not able to compare directly with their measurements; nevertheless, important differences

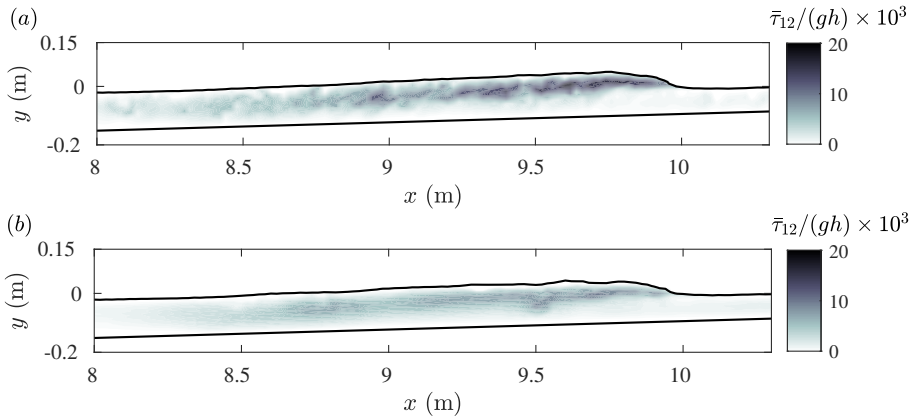


Figure 13: Phase-averaged  $\tau_{12}$  at  $t/T = 0.08$  for the spilling breaker case computed with the (a) Wilcox (2006) stress- $\omega$  and (b) LF18  $k$ - $\omega$  model. Results are scaled using the depth  $h = 0.102$  m at  $x = 9.725$  m.

621 between the two models will be revealed. It is seen from figure 12(a–d) that neither model  
 622 predicts significant Reynolds shear stress prior to breaking (as should be expected) or in  
 623 the outer surf zone. However, further shoreward the Reynolds shear stress predicted with  
 624 the Wilcox (2006) stress- $\omega$  model is significantly larger than with the LF18  $k$ - $\omega$  model,  
 625 particularly in the upper part of the water column i.e. near the surface. These differences  
 626 can also be seen directly in figure 13, which compares (phase-averaged) snapshots of the  
 627 specific Reynolds shear stress ( $\tau_{12}$ ) field beneath breaking bores computed with both models  
 628 in the inner surf zone. The instant shown has been selected such that the surface breaking  
 629 wave front is approximately at the inner-most measurement position ( $x = 9.725$  m). The  
 630 increased Reynolds shear stresses with the stress- $\omega$  model will in turn increase flow resistance  
 631 in the upper part of the water column. Although we again cannot compare directly with  
 632 measurements of  $\tau_{12}$  in the present case, it is now evident that it is this increased flow  
 633 resistance that is responsible for slowing the propagation of the breaking wave front in  
 634 the inner surf zone, bringing the resulting (phase-averaged) surface elevation time series  
 635 computed with the stress- $\omega$  model in line with that measured (see again e.g. figure 8d). In  
 636 Larsen & Fuhrman (2018), the flow resistance was represented through the eddy viscosity  
 637  $\nu_t$ , as shown in their figure 14. A higher eddy viscosity in the upper part of the flow extracts  
 638 more energy from the mean flow, which reduces the mean flow velocities. However, the  
 639 stress- $\omega$  model does not utilize the eddy viscosity assumption. Therefore, we compare the  
 640 flow resistance between two turbulence models through  $P_k$ , as given in (2.31) and (2.33),  
 641 which represents the rate at which kinetic energy is transferred from the mean flow to the  
 642 turbulence (Wilcox 2006, p. 109). For the stress- $\omega$  model, the turbulence shear production  
 643 is in the form of  $\tau_{12}S_{12}$  which is seen to be the rate at which work is done by the mean shear  
 644 strain rate against the Reynolds shear stress. Therefore,  $P_k$  is an indicator of flow resistance  
 645 that is induced by the Reynolds shear stress  $\tau_{12}$ . For the two-equation model  $P_k$  is calculated  
 646 based on  $\nu_t$ , as is presented in (2.33). As shown in figure 14 in the upper part of the flow  
 647 (right beneath the breaking bore), the shear production of turbulence with the stress- $\omega$  model  
 648 is larger than with LF18  $k$ - $\omega$  model, indicating higher flow resistance near the broken wave  
 649 surface with the stress- $\omega$  model. The related effects on the period-averaged undertow velocity  
 650 profiles will be considered in the next paragraph.

651 As hinted immediately above, figure 15 compares computed and measured period-averaged  
 652 undertow velocity profiles. It is seen that the stabilized LF18  $k$ - $\omega$  model provides accurate

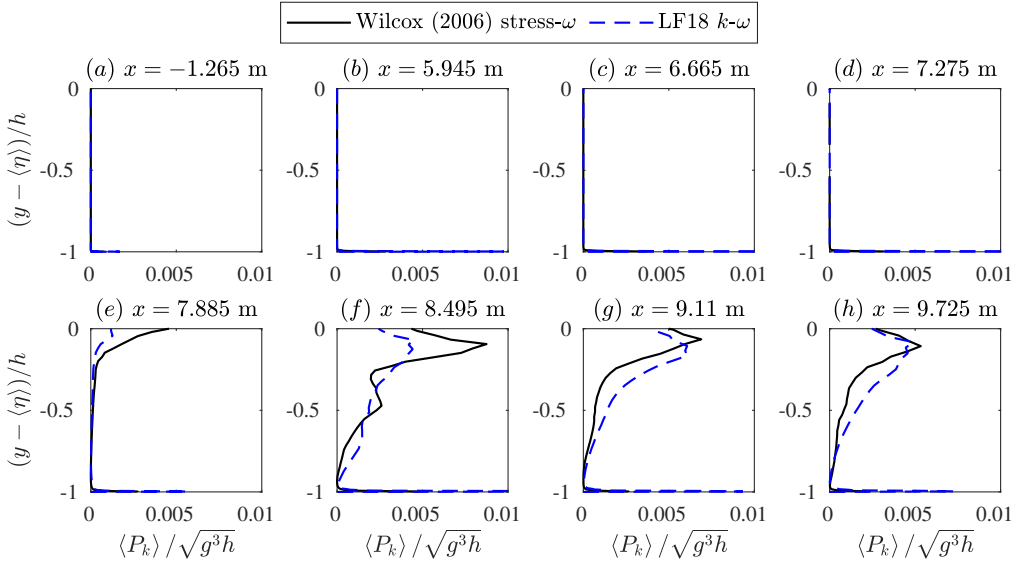


Figure 14: Period-averaged  $P_k$  for the spilling breaker from the present simulations.

653 undertow velocity profiles before wave breaking and in the outer surf zone (figure 15a–e),  
 654 generally consistent with the earlier findings of Larsen & Fuhrman (2018). Once reaching the  
 655 inner surf zone (figure 15f–h), however, the LF18  $k-\omega$  model yields exaggerated undertow  
 656 velocities. In contrast, the stress- $\omega$  model maintains consistent accuracy in the computed  
 657 undertow velocity profile across the entirety of the measured surf zone, resulting in a  
 658 significant increase in accuracy. These differences seem clearly linked to the increased  
 659 flow resistance near the surface shown in figure 12(e–h) and figure 13, and the related  
 660 increased accuracy of the breaking bore propagation evident from figure 8(d). As the  
 661 Reynolds shear stress in two-equation turbulence closure models is computed based on  
 662 the Boussinesq approximation, it seems clear that this classical assumption utilized within  
 663 two-equation models (even in their stabilized form) fails to yield the correct evolution of the  
 664 flow resistance in the inherently complicated inner surf zone, which further leads to locally  
 665 inaccurate undertow predictions.

666 The accurate prediction of undertow velocities is of major importance in the fluid  
 667 mechanics of the surf zone, as they are important drivers of fluid, pollutants, and sediment  
 668 transport in nearshore coastal regions. Despite this importance, the problem of inaccurate  
 669 undertow velocity profiles has consistently plagued RANS CFD simulations of breaking  
 670 waves over the past two decades. The present results utilizing the Wilcox (2006) stress- $\omega$   
 671 model are novel, in that they represent the first time that consistent quantitative accuracy in  
 672 the computed undertow has been maintained throughout the entirety of the nearshore wave  
 673 breaking process i.e. during shoaling (prior to breaking), to the outer surf zone, and all the  
 674 way into the inner surf zone. Other RANS models (typically using two-equation turbulence  
 675 closure) yield incorrect undertow structure prior to breaking and in the outer surf zone (e.g.  
 676 Lin & Liu 1998; Brown *et al.* 2016; Devolder *et al.* 2018; Liu *et al.* 2020) or exaggerated  
 677 undertow in the inner surf zone (e.g. Jacobsen *et al.* 2012; Larsen & Fuhrman 2018; Larsen  
 678 *et al.* 2020), or both. A detailed discussion on the results and problems in previous works  
 679 will be presented in Section 4.



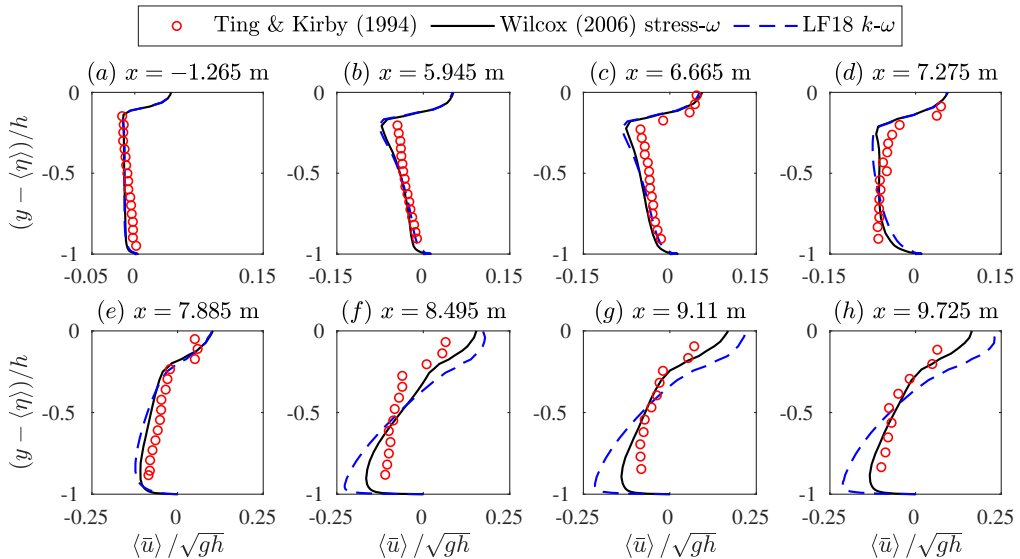


Figure 15: Period-averaged undertow velocity profiles for the spilling breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations.

680

### 3.4. Simulating plunging breaking waves

681 We will now employ the Wilcox (2006) stress- $\omega$  model to simulate the plunging breaking  
 682 wave experiments of Ting & Kirby (1994, 1996). For these simulations the numerical set-up  
 683 and protocol is identical to that used for spilling breakers in Section 3.3, with the wave  
 684 parameters as indicated in table 1. The simulation of the plunging breaking waves required  
 685 approximately 25 days to run in parallel on 16 processors on the supercomputing cluster at  
 686 DTU. As before, comparison will be made with the LF18  $k$ - $\omega$  model (Larsen & Fuhrman  
 687 2018), which again represents a stabilized form of the basic model presented by Wilcox  
 688 (2006). As much of the story to follow bears similarity to that in Section 3.3, it will be told  
 689 with far greater brevity in the present sub-section.

690 Figure 16 depicts a snapshot of the dimensionless turbulence field  $k/(\omega\nu)$  for the plunging  
 691 breaking case, computed with the stress- $\omega$  model at a time instant of  $t/T = 50.825$ , similar  
 692 to figure 6. This time instant has been chosen, as it corresponds to wave over-turning just  
 693 prior to the subsequent plunge. Similar to our findings in the spilling breaking case, there  
 694 is no turbulence over-production prior wave breaking. This should by now be expected as  
 695 we have definitively established that the stress- $\omega$  model is stable in nearly potential flow  
 696 regions beneath surface waves. It can be noted that this plunging case is not nearly as prone  
 697 to significant turbulence over-production prior to breaking as the spilling case, because the  
 698 unstable growth rate is much smaller due to a small value of  $[\langle p_0 \rangle]$ , as discussed by Larsen  
 699 & Fuhrman (2018).

700 Figure 17 compares the surface elevation envelopes from the model simulations with  
 701 the experimental measurements, similar to figure 7. A reasonable match is again achieved.  
 702 Both the Wilcox (2006) stress- $\omega$  and LF18  $k$ - $\omega$  models predict the breaking point, and  
 703 subsequent wave decay, reasonably. The set-up in the mean water level is likewise similarly  
 704 well predicted. It is noted that right after the breaking point (at  $x \approx 8$  m), the maximum surface  
 705 elevation predicted with both numerical models has small deviations from the experimental  
 706 measurement (with the stress- $\omega$  model result being slightly closer to the measurement). This  
 707 deviation could be due to the plunging jet splashing down and causing turbulent mixture

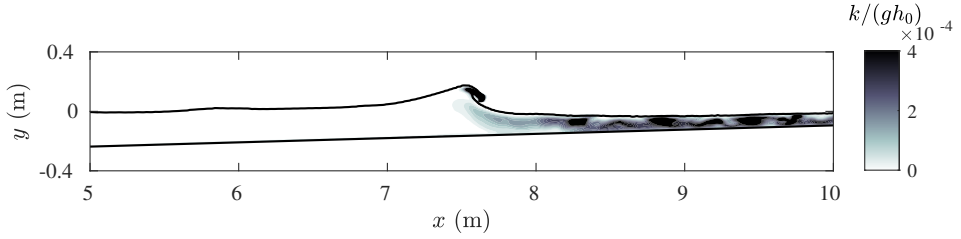


Figure 16: Snapshot of the plunging breaker turbulent kinetic energy simulated with the Wilcox (2006) stress- $\omega$  model at  $t/T = 100$ .

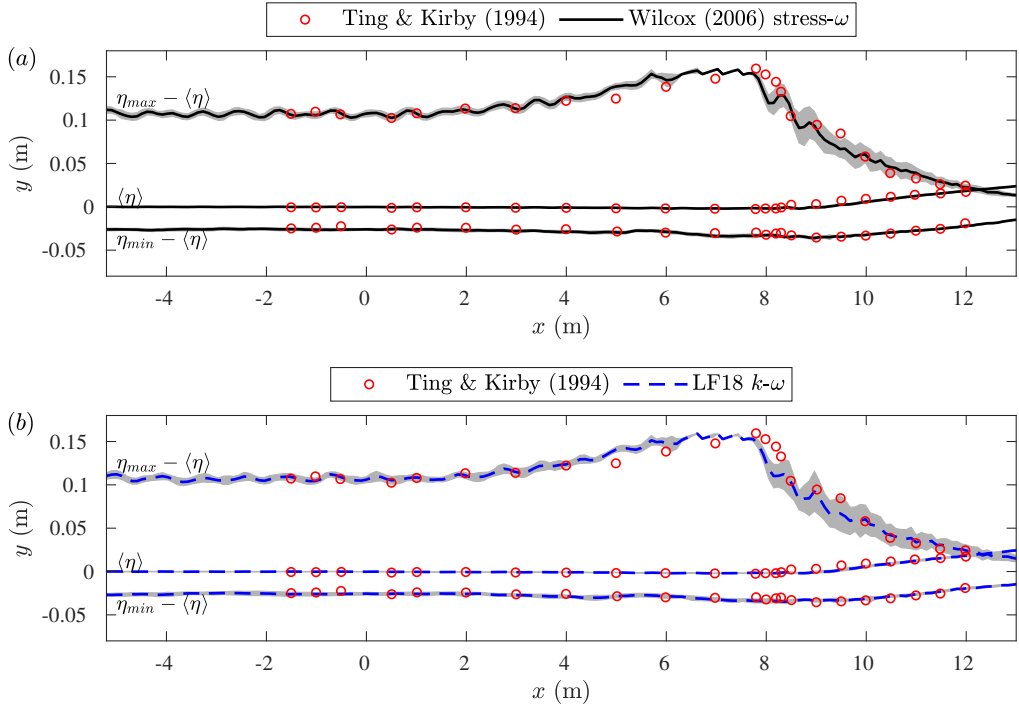


Figure 17: Surface elevation envelopes for the plunging breaker simulated with (a) the present Wilcox (2006) stress- $\omega$  model and (b) the LF18  $k$ - $\omega$  model, comparing to the experimental measurement of Ting & Kirby (1994). Grey shaded areas are the plus and minus one standard deviation.

708 of the surface layer (as discussed in Brocchini 2002) which makes accurate modelling  
 709 challenging. However, our numerical models are able to show reasonable consistency with  
 710 the measurements, with minor deviations in the splash region. Comparison of computed  
 711 and measured phase-averaged time series of the surface elevation at several measurement  
 712 positions are additionally provided in figure 18. Interestingly, apart from the deviations near  
 713 the crest in figure 18(c-e) with the  $k$ - $\omega$  model, the computed wave front in the present  
 714 plunging case does not propagate noticeably faster with the  $k$ - $\omega$  model in the inner surf zone.  
 715 This differs from our findings in the spilling case, see figure 8(d,e), and will be explained  
 716 later in this sub-section.

717 Computed and measured (when available) period-averaged (over the final  $50T$ , as before)  
 718 turbulent normal stress profiles are compared in figure 19 (for  $-\tau_{11}$ ) and figure 20 (for  
 719  $-\tau_{22}$ ), with profiles for the turbulent kinetic energy density  $k$  similarly presented in figure

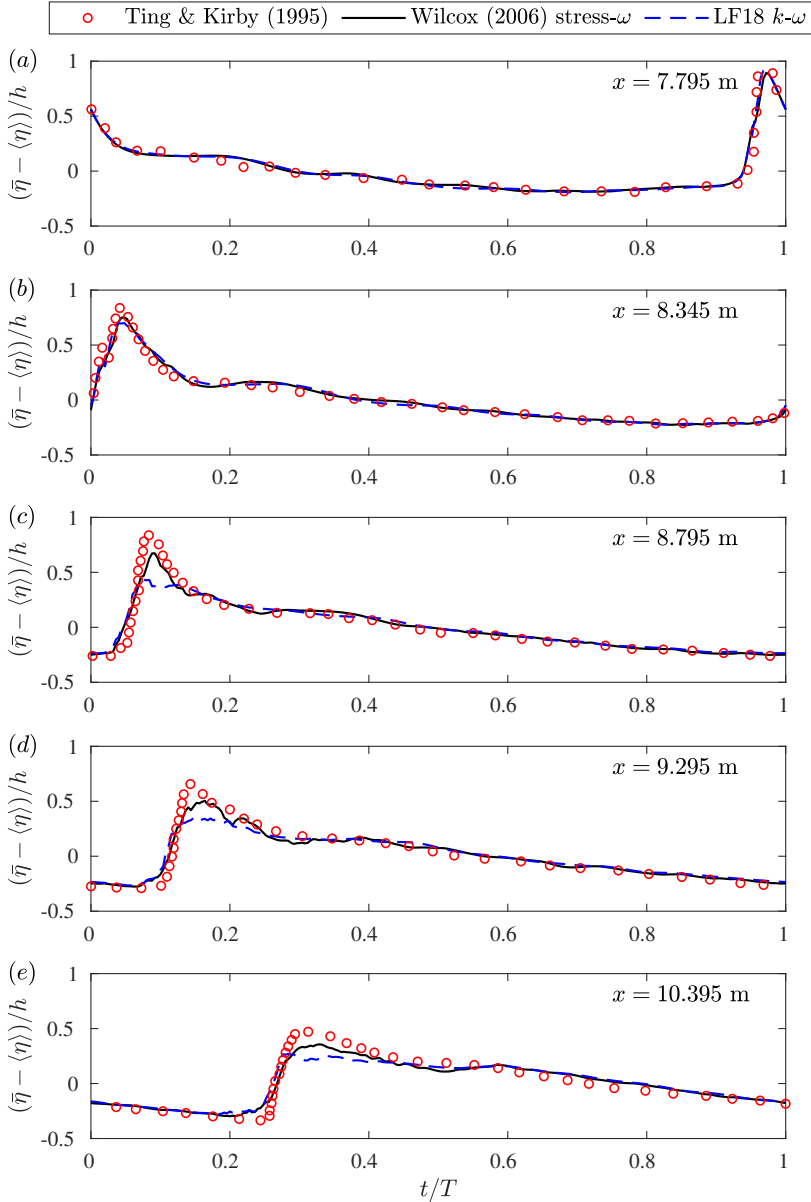


Figure 18: Phase-averaged surface elevation for the plunging breaker from the experimental measurement of Ting & Kirby (1995) and the present simulations. The subfigures (a, b, c) are in the outer surf zone while (d, e) are in the inner surf zone.

720 **21.** In the experiments  $k$  was again estimated from (3.7). As Ting & Kirby (1994) did not  
 721 provide measurement data for  $\langle \tau_{22} \rangle$  or  $\langle \tau_{22} \rangle / \langle \tau_{11} \rangle$  for their plunging case, only model results  
 722 are shown in figure 20. From these figures it is seen that the two models seem to provide  
 723 comparable accuracy for the turbulent normal stresses, similar to what was shown in our prior  
 724 simulations involving spilling breaking waves. The results for  $k$  are somewhat more accurate  
 725 with the stress- $\omega$  model specifically at  $x = 9.795$  m (figure 21g), though this increased  
 726 accuracy is not consistent throughout the surf zone as a whole. The overall prediction for  $k$

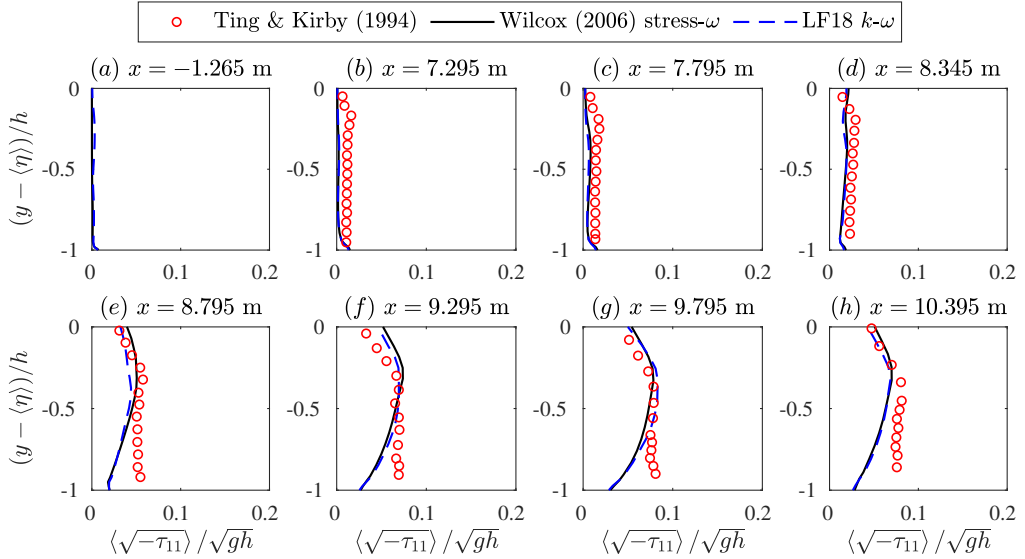


Figure 19: Period-averaged specific Reynolds normal stresses  $-\tau_{11}$  profiles for the plunging breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations. The subfigures (a,b) are in the pre-breaking region while (c,d,e) are in the outer surf zone and (f,g,h) are in the inner surf zone.

727 in the inner surf zone with both turbulence models are larger than the measurement with a  
 728 maximum factor of two (figure 21 g). The reason remains uncertain to the authors. However,  
 729 it is worthwhile to mention that the experimental study of Scott *et al.* (2005) presented  
 730  $k$  profiles post-processed with three different turbulence separation methods, with results  
 731 varying by up to a factor of two to six from one another. The vertical gradient of their largest  
 732 prediction is much higher than that of the lowest prediction (as shown in their figure 5).  
 733 Therefore, the difference between our numerical results and the measurement of Ting &  
 734 Kirby (1994) might still be considered reasonable.

735 Figure 22 presents the computed phase-averaged  $\tau_{12}$  field in the surf zone for the plunging  
 736 case with both models, in a fashion similar to figure 13. The phase plotted has been selected  
 737 to capture the propagation of the breaking wave front in the inner surf zone. Similar to our  
 738 findings in the spilling case, it is clearly seen that the stress- $\omega$  model predicts turbulent shear  
 739 stresses that are significantly larger in the inner surf zone than that with the  $k$ - $\omega$  model. It  
 740 can thus be expected to result in increased flow resistance in this region. From comparison  
 741 of figures 22 and 13 it is also seen that the increased turbulent shear stresses in the plunging  
 742 case are spread more uniformly throughout the water column than in the spilling case, where  
 743 they were more concentrated near the surface. This is likely due to the more violent surf  
 744 zone initiated by the plunging breaking, and thus also explains why the breaking surface  
 745 front propagates at approximately the same speed in the inner surf zone with both models in  
 746 the present case (see again figure 18). The flow resistance indicated by  $P_k$  for the plunging  
 747 breaker is likewise presented in figure 23. It is clearly seen that in the upper part of the  
 748 flow,  $P_k$  predicted with the stress- $\omega$  model is much larger than that predicted with the LF18  
 749  $k$ - $\omega$  model, indicating higher flow resistance and therefore smaller magnitude of mean flow  
 750 velocity with the stress- $\omega$  model.

751 Figure 24 finally compares computed and measured undertow velocity profiles. It is seen  
 752 that before wave breaking (figure 24a-c), the numerical simulations with both turbulence  
 753 models are almost identical, and are in line with the experimental measurement. This is as

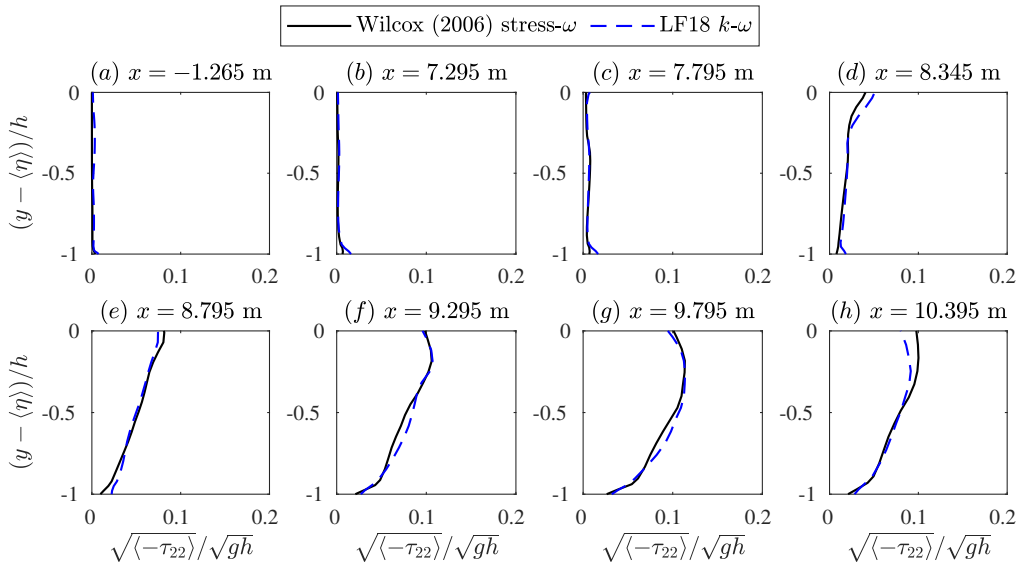


Figure 20: Period-averaged specific Reynolds normal stresses  $-\tau_{22}$  profiles for the plunging breaker from the present simulations.

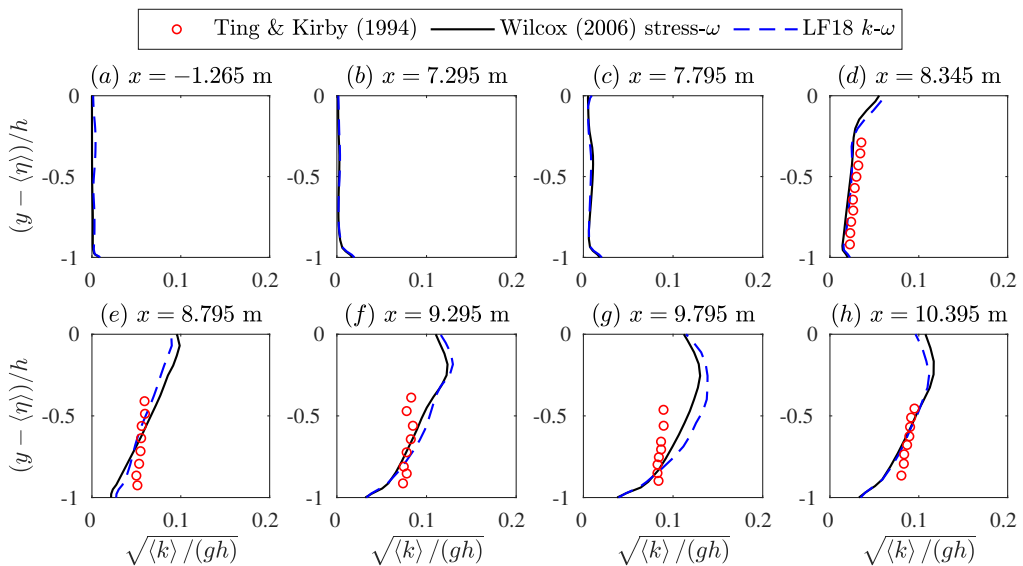


Figure 21: Period-averaged turbulent kinetic energy  $k$  profiles for the plunging breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations.

754 expected, since both model variants considered herein are formally stable in the potential flow  
 755 regions beneath surface waves, hence the choice of turbulence model has little impact prior to  
 756 breaking. Results are also similar in the outer surf zone, as seen in figure 24(d,e). Much more  
 757 significant differences become apparent once the inner surf zone is reached, as seen in figure  
 758 24(f-h). Consistent with the previously considered spilling breaking case (figure 15), in the  
 759 inner surf zone the LF18  $k-\omega$  model results in undertow velocity profiles that are much larger  
 760 than were measured. The LF18 turbulence model similarly yielded over-predicted undertow  
 761 velocities in the simulation of large-scale plunging breakers made by Larsen *et al.* (2020).

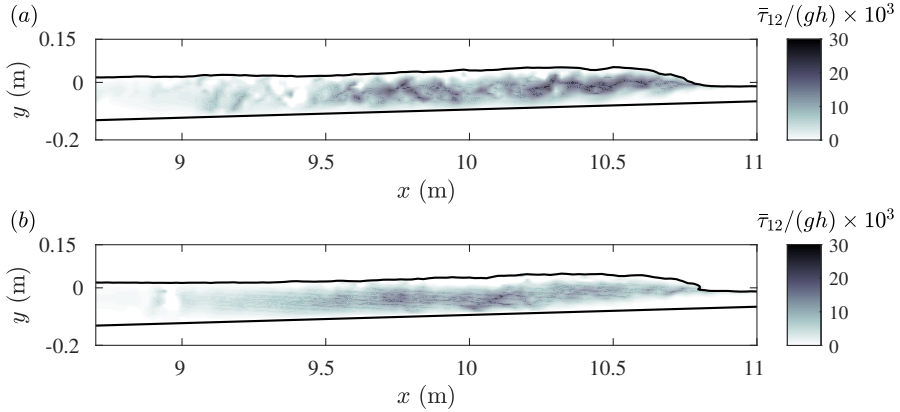


Figure 22: Phase-averaged  $\tau_{12}$  at  $t/T = 0.30$  for the plunging breaker case computed with the (a) Wilcox (2006) stress- $\omega$  and (b) LF18  $k$ - $\omega$  model. Results are scaled using the depth  $h = 0.083$  m at  $x = 10.395$  m.

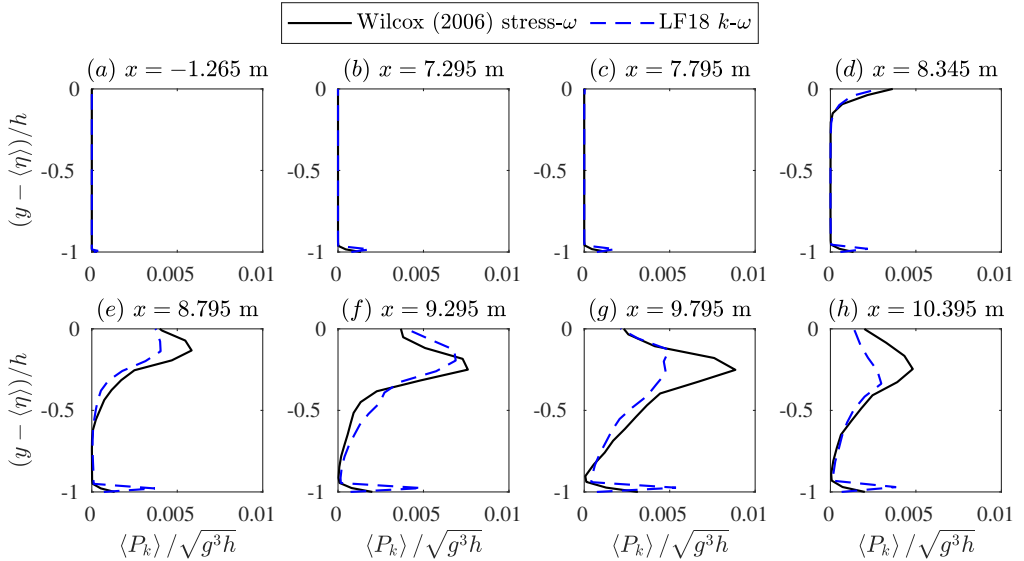


Figure 23: Period-averaged  $P_k$  for the plunging breaker from the present simulations.

762 It is thus now evident that this is a consistent shortcoming with this model, which stems  
 763 from the inclusion of the traditional stress-limiter within the Wilcox (2006)  $k$ - $\omega$  model (see  
 764 again the comparisons made by Larsen & Fuhrman 2018, with this feature switched on and  
 765 off). The stress- $\omega$  model, on the other hand, reduces this exaggeration considerably, though  
 766 not completely. The undertow profiles predicted with this model in the inner surf zone are  
 767 much more uniform, having a similar structure to what has been measured. The reduction in  
 768 the undertow magnitudes computed with the stress- $\omega$  model is consistent with the increased  
 769 flow resistance in the inner surf zone, as illustrated in figures 22 and 23.

770 Though a substantial improvement of the predicted undertow in the inner surf zone is  
 771 seen with the stress- $\omega$  model, there are still some disagreements between the stress- $\omega$  model  
 772 prediction and the experimental measurement for the plunging breaker (figure 24f-h). The  
 773 reasons are, as yet, uncertain to the authors. One possible reason could be the simplistic

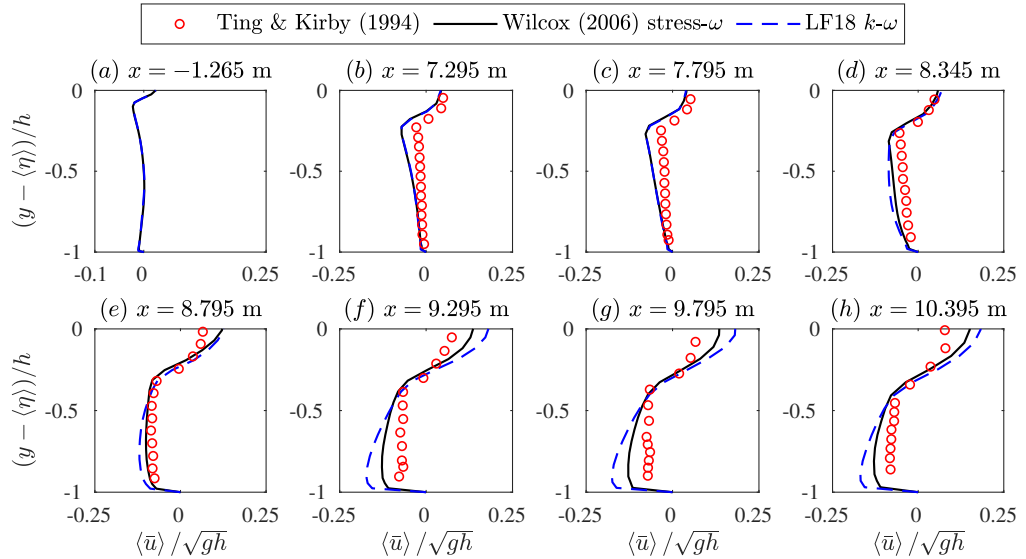


Figure 24: Period-averaged undertow velocity profiles for the plunging breaker from the experimental measurement of Ting & Kirby (1994) and the present simulations.

774 formulation of the pressure-strain terms in the Wilcox (2006) stress- $\omega$  model, for which more  
 775 complex closures for pressure-strain terms could potentially make further improvement for  
 776 the undertow predictions in the complicated inner surf zone of plunging breakers. Another  
 777 possible reason could be air-bubbles entrained in the plunging surf zone. The present study  
 778 has not specifically employed a model for the air bubbles/pockets. The bubble-mass cascade  
 779 phenomena (Chan *et al.* 2021) and the bubble break-up in the inner surf zone may further  
 780 increase the flow resistance. These may be interesting to investigate in future work.

#### 781 4. Discussion

782 The present work represents the first time that such accurate prediction of the breaking point,  
 783 turbulence characteristics, and evolution of the undertow structure from pre-breaking to the  
 784 inner surf zone has been achieved with a single turbulence closure model for both the spilling  
 785 and plunging breaking cases of Ting & Kirby (1994, 1996), which have widely served as the  
 786 basis for validating breaking wave models over the past two decades. In what follows, we  
 787 provide a discussion of the results and problems encountered in previous studies which have  
 788 attempted to model breaking waves with comparable CFD models. Such results mainly fall  
 789 into three categories:

790 (i) **Over-production of turbulence prior to breaking, especially in the spilling case.**  
 791 Models in this category become polluted due to unphysical turbulence over-production during  
 792 the shoaling process i.e. before the wave breaking process even starts, and therefore cannot  
 793 claim to have modelled the processes leading up to and including the surf zone correctly.  
 794 Results in this category typically stem from two-equation closure models in their standard  
 795 forms, as the analysis of Larsen & Fuhrman (2018) has proved that these are (asymptotically)  
 796 unconditionally unstable in nearly potential flow regions beneath non-breaking surface waves.  
 797 Numerous examples include wave breaking simulations using standard formulations of the  
 798  $k$ - $\varepsilon$  turbulence model (e.g. Lin & Liu 1998; Bradford 2000; Xie 2013; Brown *et al.* 2016;  
 799 Derakhti *et al.* 2016b). Results using the “non-stabilized” (standard) variant of the realizable  
 800  $k$ - $\varepsilon$  model to simulate breaking waves by Fuhrman & Li (2020) also fall into this category.

801 The same problem is also evident in SPH simulations coupled with a  $k$ - $\varepsilon$  model (e.g. Shao  
 802 2006). Likewise, results from  $k$ - $\omega$  or  $k$ - $\omega$  SST closure models have also demonstrated the  
 803 same turbulence over-production problem (e.g. Brown *et al.* 2016; Devolder *et al.* 2018; Liu  
 804 *et al.* 2020). Results using “non-stabilized” variants of the  $k$ - $\omega$  model by Larsen & Fuhrman  
 805 (2018), i.e. those having  $\lambda_2 = 0$  (their notation), would similarly fall into this category. In  
 806 some other works employing standard two-equation models, the undertow velocity profiles  
 807 and turbulence were simply not presented. For example, Lupieri & Contento (2015) utilized  
 808 the  $k$ - $\omega$  SST model, but did not present undertow and turbulent kinetic energy predictions.  
 809 However, the phase-averaged surface elevations for both spilling and plunging breakers near  
 810 the breaking point were significantly under predicted, which would be consistent with a  
 811 polluted pre-breaking region, causing the simulated waves to decay prematurely. Similarly,  
 812 Chella *et al.* (2015) utilized a standard  $k$ - $\omega$  model to simulate breaking waves, but did not  
 813 present detailed predictions of either undertow or turbulence. As turbulence models of this  
 814 type were shown by Larsen & Fuhrman (2018) to be formally unstable, combined with  
 815 the numerous simulations with similar models leading to turbulence over-production noted  
 816 above, there would seem to be little doubt as to the inherent instability in the nearly potential  
 817 flow leading up to wave breaking in their model. It is worth mentioning that some recent  
 818 notable works using two-equation turbulence closure models have attempted to improve the  
 819 accuracy of breaking wave modelling by focusing on the air-water interface region near the  
 820 surface. For example, Devolder *et al.* (2018) added buoyancy production terms to the  $k$ - $\omega$  and  
 821  $k$ - $\omega$  SST models, to account for density gradients near the air-water interface. Additionally,  
 822 Liu *et al.* (2020) applied a free surface jump condition to the  $k$ - $\omega$  SST model, while also  
 823 separately considering a variant incorporating buoyancy production as in Devolder *et al.*  
 824 (2018), to simulate the experiments of Ting & Kirby (1994, 1996). Their works showed that  
 825 such features could improve prediction of the breaking point relative to standard models  
 826 without these features. However, over-production of turbulence prior to breaking still clearly  
 827 persists in these models, which is especially apparent in the spilling case, as can clearly  
 828 be seen e.g. in figure 17 of Liu *et al.* (2020). This is also clear from results of Larsen &  
 829 Fuhrman (2018) and Larsen *et al.* (2020) using “non-stabilized” models, but where buoyancy  
 830 production was still included, as also discussed by Fuhrman & Larsen (2020). These results  
 831 thus collectively indicate that, while inclusion of buoyancy production will cause a local sink  
 832 of turbulent kinetic energy near the free surface (and thus may be beneficial), it does little  
 833 to stabilize two-equation turbulence models (and hence avoid turbulence over-production) in  
 834 the nearly potential flow core region prior to breaking as a whole.

835 (ii) **Turbulence over-production eliminated prior to breaking, but undertow poorly**  
 836 **predicted in the inner or outer surf zone.** This category consists of CFD simulations  
 837 using turbulence models which avoid turbulence over-production prior to breaking, but  
 838 typically yield poor undertow velocity structure and/or magnitude in either the outer or  
 839 the inner surf zone e.g. Mayer & Madsen (2000), Jacobsen *et al.* (2012, 2014), Larsen &  
 840 Fuhrman (2018), and Fuhrman & Li (2020). This category can be further sub-divided into  
 841 those turbulence closure models which incorporate a conventional stress-limiter on the eddy  
 842 viscosity (corresponding to  $\lambda_1 > 0$  in the notation of Larsen & Fuhrman 2018), and those  
 843 which do not (corresponding to  $\lambda_1 = 0$ , again in their notation). Mayer & Madsen (2000) made  
 844 the first attempt to control the instability inherent in the standard  $k$ - $\omega$  model through ad-hoc  
 845 modification of the production terms (i.e. the production terms in the  $k$  and  $\omega$  equations were  
 846 modified to be based on the rotation-rate tensor instead of the strain-rate tensor). Jacobsen  
 847 *et al.* (2012, 2014) adopted the modification of Mayer & Madsen (2000), such that the  
 848 turbulence over-production in the potential flow region prior to breaking was avoided, while  
 849 also incorporating a conventional stress-limiter on the eddy viscosity. The resulting model  
 850 yielded reasonable prediction of the undertow velocity structure in the spilling breaking case



851 of Ting & Kirby (1994), both prior to breaking and in the outer surf zone, but unfortunately  
852 resulted in exaggerated undertow magnitudes (by approximately a factor of two) in the  
853 inner surf zone. Larsen & Fuhrman (2018) discussed theoretical inconsistencies with the  
854 modification proposed by Mayer & Madsen (2000) (namely, that it leaves the Reynolds  
855 stress tensor doubly-defined) and instead formally stabilized two-equation models through  
856 re-formulation of the eddy viscosity. Fuhrman & Li (2020) adopted a similar approach and  
857 stabilized the realizable  $k$ - $\varepsilon$  model. The “stabilized” model results of Larsen & Fuhrman  
858 (2018) with the conventional stress-limiter on ( $\lambda_1 > 0$ ) were qualitatively similar to those  
859 of Jacobsen *et al.* (2012), with undertow velocity profiles quite accurate prior to breaking  
860 and in the outer surf zone, but exaggerated in the inner surf zone. Larsen & Fuhrman  
861 (2018) additionally conducted simulations where their “stabilized” closure models had the  
862 conventional stress-limiter switched off ( $\lambda_1 = 0$  in their notation). This variant produced  
863 quite accurate undertow profiles in the inner surf zone, but at the expense of grossly over-  
864 predicted turbulence levels and erroneous undertow structure in the outer surf zone. From this  
865 comparison, it seems clear that the classical Boussinesq approximation utilized within two-  
866 equation turbulence closure models (even with advanced features, such as stress-limiters)  
867 is not capable of yielding the correct evolution of the flow resistance beneath breaking  
868 waves over the entirety of the surf zone, even in the relatively calm conditions associated  
869 with spilling breaking. Experience with “stabilized” closure models in the CFD simulation of  
870 plunging breaking waves (Larsen *et al.* 2020; Sumer & Fuhrman 2020) has likewise produced  
871 results that are generally consistent with those described above. As such, while the models  
872 cited above avoid over-production of turbulence in potential flow regions prior to the onset  
873 of breaking, none can reasonably claim to have accurately simulated the breaking process  
874 (including accurate evolution of the undertow velocity profile) across the entirety of the surf  
875 zone in either the spilling or plunging cases of Ting & Kirby (1994, 1996).

876 (iii) **Results simulated with other CFD approach such as LES and SPH with a sub-**  
877 **grid scale turbulence model.** We finally discuss results from a third category, consisting  
878 of models not working within the confines of Reynolds-averaged Navier-Stokes equations.  
879 Watanabe & Saeki (1999) applied LES with a sub-grid scale model to simulate breaking  
880 waves. However, their model was only qualitatively validated. Christensen (2006) simulated  
881 both spilling and plunging breaking wave experiments of Ting & Kirby (1994, 1996) with  
882 LES and two different sub-grid scale models, one in terms of the Smagorinsky model, and the  
883 other based on the  $k$ -equation. However, compared to the present results, the breaking points  
884 were not accurately captured and the turbulence levels were in general too high compared  
885 to experiments of Ting & Kirby (1994, 1996). Zhou *et al.* (2017) also conducted LES with  
886 a Lagrangian dynamic sub-grid closure model. Their model over-predicted the turbulent  
887 intensity especially near the surface. The undertow velocities were more or less similar to the  
888 work of Jacobsen *et al.* (2014) which have been classified into category (ii) above. Makris  
889 *et al.* (2016) applied an SPH approach with a Smagorinsky-type sub-particle scale approach,  
890 which is similar to the LES concept. Their study on a weakly plunging breaker showed clear  
891 underestimation of the ensemble-averaged surface elevation at the incipient breaking region  
892 compared to the experiment of Stansby & Feng (2005). Lowe *et al.* (2019) also conducted an  
893 SPH simulation for breaking waves, and it was found that the turbulent kinetic energy was  
894 over-predicted with this approach, even with no sub-particle scale turbulence closure models  
895 included. This over-prediction was even greater with inclusion of a sub-grid scale model.  
896 They specifically highlighted the need for further improvement in sub-grid scale turbulence  
897 models within the surf zone.

898 In contrast to those models discussed above, the present study marks the first time that the  
899 Wilcox (2006) stress- $\omega$  Reynolds stress model has been utilized to simulate the multiphase  
900 wave breaking process. As can be seen from the results presented and discussed above, this

901 approach solves several of the problems which have consistently plagued other comparable  
902 models of breaking waves over the past two decades. Most notably, the present results  
903 have demonstrated, for both spilling and plunging breaking cases: (1) no turbulence over-  
904 production prior to breaking, (2) accurate prediction of the breaking point, (3) reasonable  
905 (though certainly not perfect) evolution of turbulence quantities within the surf zone, and (4)  
906 undertow velocity profile structure and magnitudes for the spilling breaker are in line with  
907 measurements from pre-breaking regions all the way to the outer and inner surf zones, while  
908 for the plunging breaker the undertow results are largely improved comparing to the best  
909 two-equation model of Larsen & Fuhrman (2018), especially in the inner surf zone. Hence,  
910 the present model seemingly provides the most accurate and consistent (for both spilling  
911 and plunging cases) description of the turbulent wave breaking process achieved with CFD  
912 models to date.

913 Indeed, many of the issues faced by the comparable CFD models discussed above are  
914 rather naturally avoided with the stress- $\omega$  turbulence closure. As proved in Section 2, this  
915 model is formally (neutrally) stable in the potential flow region beneath non-breaking surface  
916 waves. Hence, it avoids (without any modification) the over-production of turbulence prior  
917 to breaking plaguing the standard two-equation models in category (i) above. Following  
918 Devolder *et al.* (2018) and Larsen & Fuhrman (2018), we have additionally added buoyancy  
919 production to this model, such that these benefits are likewise retained. Finally, the stress- $\omega$   
920 model breaks free of the Boussinesq approximation, and hence the eddy viscosity concept  
921 (and associated complications such as stress-limiters) altogether. Rather, the Reynolds stress  
922 is allowed to evolve according to its own governing equation, resulting in a model that is both  
923 theoretically superior, and more capable of predicting the dynamic variations in the flow  
924 resistance that arise between the outer and inner surf zones. This freedom seems to solve the  
925 problem consistently encountered by the models falling into category (ii) above, where users  
926 were seemingly faced with having to choose between accurate undertow profiles in either  
927 the outer or inner surf zone. It is finally worth mentioning that, by still working within the  
928 confines of Reynolds-averaged equations, the stress- $\omega$  model additionally avoids the practical  
929 resolution issues that are commonly faced and raised in LES applications, while also avoiding  
930 any need for sub-grid scale modelling, as described in relation to category (iii) above. It would  
931 thus seemingly offer an attractive compromise that has been under-utilized to date, providing  
932 a turbulence model that is dynamic enough to handle the inherently complicated surf zone  
933 at reasonable computational expense.

## 934 5. Conclusions

935 The present work has considered the Reynolds stress- $\omega$  model of Wilcox (2006), as a  
936 new candidate for providing turbulence closure in the CFD simulation of breaking waves  
937 with Reynolds-averaged Navier-Stokes equations. We have first conducted novel stability  
938 analysis of this model, formally proving that it is neutrally stable in the potential flow region  
939 beneath non-breaking surface waves. Unlike simpler two-equation models in their standard  
940 forms (see Larsen & Fuhrman 2018), this model should therefore not lead to unphysical  
941 exponential growth of turbulence during the shoaling process leading up to incipient breaking.  
942 Comparison with prior analysis of two-equation models has also definitively shown that their  
943 instability arises as a result of the widely-utilized Boussinesq approximation. The stability  
944 of the stress- $\omega$  model in potential flow regions has been directly confirmed through CFD  
945 simulations involving a progressive surface wave train.

946 As coastal waves (both breaking and non-breaking) also involve a wave boundary layer  
947 at the sea bottom, the stress- $\omega$  model has subsequently been applied to simulate unsteady  
948 oscillatory turbulent wave boundary layer flow, as measured by Jensen *et al.* (1989). The

949 computational results are generally in line with those measured, with notable improvement  
950 over two-equation turbulence closures apparent in the deceleration stage, where e.g. the  $k-\omega$   
951 turbulence model fails to accurately capture the turbulence kinetic energy and the Reynolds  
952 shear stress distribution. The stress- $\omega$  model has improved the accuracy for predicting the  
953 anisotropic Reynolds normal stress and Reynolds shear stress components within the wave  
954 boundary layer.

955 This work has culminated with CFD simulations employing the stress- $\omega$  turbulence closure  
956 model in the simulation of both the spilling and plunging breaking wave cases of Ting &  
957 Kirby (1994, 1996). Surface elevation envelopes, turbulence characteristics and undertow  
958 velocity profiles have been predicted with consistent accuracy maintained from pre-breaking  
959 all the way into the inner surf zone in both cases. Comparison with the stabilized  $k-\omega$  model  
960 of Larsen & Fuhrman (2018) demonstrates that both models predict Reynolds normal stresses  
961 (and turbulent kinetic energy) that are reasonably in line with measurements. Both models  
962 likewise predict similar undertow velocity profiles prior to breaking and in the outer surf zone.  
963 In the inner surf zone, however, the Larsen & Fuhrman (2018)  $k-\omega$  model predicts undertow  
964 velocity profiles that are exaggerated by approximately a factor of two in magnitude relative  
965 to measurements, a feature that has similarly plagued several other two-equation turbulence  
966 closure models in the literature. The stress- $\omega$  model, on the other hand, generally results in  
967 undertow velocity profiles that are reasonably accurate (both in the uniformity of structure  
968 and magnitude) throughout the surf zone. These differences have been shown to stem from  
969 predictions in the Reynolds shear stresses within the inner surf zone, which are significantly  
970 larger with the stress- $\omega$  model (near the surface in the spilling case, more distributed across  
971 the depth in the plunging case). This in turn results in greater flow resistance in the inner surf  
972 zone. Based on a survey of previous CFD simulations of breaking waves in the literature,  
973 we conclude that the stress- $\omega$  model considered herein is seemingly the first demonstrating  
974 the collective ability to: (1) naturally avoid turbulence over-production prior to breaking,  
975 (2) accurately predict the breaking point, (3) provide reasonable evolution of turbulent  
976 normal stresses across the surf zone, while also providing (4) accurate undertow structure  
977 and magnitude from pre-breaking regions all the way to the outer and inner surf zones for  
978 the spilling breaking waves, and improvement for the plunging breaking waves compared to  
979 previous numerical simulations. It may therefore be useful for other studies involving various  
980 aspects of breaking waves, as it seems to have been under-utilized in the literature to date.  
981 The authors are freely releasing their source code implemented in the OpenFOAM framework,  
982 to hopefully help make such applications more accessible, as described in more detail in the  
983 next section.

984 While the present work has focused primarily on analysis and applications of the Wilcox  
985 (2006) stress- $\omega$  model, a stability analysis of the Launder *et al.* (1975) (LRR) Reynolds  
986 stress- $\varepsilon$  turbulence closure model in the potential flow region beneath non-breaking waves  
987 is also novelly considered in Appendix B, for completeness. Similar to our findings for the  
988 stress- $\omega$  model, the stress- $\varepsilon$  model is likewise proved to be neutrally stable. This has similarly  
989 been confirmed through CFD simulation of a propagating wave train, similar to Section 3.1.  
990 The likely explanation of the turbulence over-production experienced by Brown *et al.* (2016)  
991 is also provided there.

## 992 **Availability of source codes**

993 The source code implemented and utilized in the present work is publicly available at: <https://github.com/LiYZPearl/ReynoldsStressTurbulenceModels>. This includes our im-  
994 plementations of all turbulence models utilized within, namely the Wilcox (2006) stress- $\omega$   
995 and  $k-\omega$  models, for use in both single- and two-phase flow simulations (including buoyancy

997 production terms). In case of the  $k$ - $\omega$  model, the two-phase flow implementation also includes  
 998 stabilization of the model as described in Larsen & Fuhrman (2018), deemed the LF18 model  
 999 within. The OpenFOAM set-ups for the simulations presented herein of the turbulent wave  
 1000 boundary layer, as well as both spilling and plunging breaking wave cases, are likewise  
 1001 provided as tutorials.

## 1002 **Acknowledgement**

1003 The first author acknowledges financial support from the European Union's Horizon 2020  
 1004 research and innovation program, Marie Skłodowska-Curie Grant No. 713683 (COFUND-  
 1005 fellowsDTU, H. C. Ørsted Postdoc project SUBSEA: SimULating Breaking waves and  
 1006 SEDiment trANsport with stabilized turbulence models). The third and the last authors  
 1007 acknowledge financial support from the Independent Research Fund Denmark (project  
 1008 SWASH: Simulating WAVE Surfzone Hydrodynamics and sea bed morphology, Grant No.  
 1009 8022- 00137B). This support is greatly appreciated.

## 1010 **Declaration of interests**

1011 The authors report no conflict of interest.

## 1012 **Appendix A. Buoyancy production term for the Wilcox (2006) stress- $\omega$ model**

1013 In this appendix we will derive the buoyancy production term for use in the Wilcox (2006)  
 1014 stress- $\omega$  turbulence closure model equation (2.1). The derivation of the buoyancy production  
 1015 term starts from the exact form given in Burchard (2002, p. 18):

$$1016 \quad B_{ij} = \frac{1}{\rho_0} \left( g_i \overline{u'_j \rho'} + g_j \overline{u'_i \rho'} \right) \quad (\text{A } 1)$$

1017 Following Burchard (2002, p. 37), the correlation between the fluctuating velocity and density  
 1018 can be written as

$$1019 \quad \overline{u'_j \rho'} = -\alpha_b^* \frac{k}{\omega} \frac{\partial \bar{\rho}}{\partial x_j} \quad (\text{A } 2)$$

1020 where  $k/\omega$  here effectively plays the role of the eddy viscosity. Invoking (A 2) within (A 1),  
 1021 the buoyancy production term becomes:

$$1022 \quad B_{ij} = -\alpha_b^* \frac{k}{\omega} N_{ij} \quad (\text{A } 3)$$

1023 where  $N_{ij}$  is from (2.5). This matches the term seen within (2.1).

1024 Note that taking half the trace of  $B_{ij}$  above leads to:

$$1025 \quad B_k = -\frac{1}{2} B_{ii} = \alpha_b^* \frac{k}{\omega} N^2, \quad N^2 = \frac{g_i}{\rho_0} \frac{\partial \bar{\rho}}{\partial x_i}. \quad (\text{A } 4)$$

1026 This matches the buoyancy production term utilized in the  $k$ - $\omega$  turbulence closure model by  
 1027 Larsen & Fuhrman (2018). They showed that requiring the steady-state Richardson number  
 1028 to be smaller than 0.25 (Schumann & Gerz 1995; Burchard 2002) corresponds to a minimum  
 1029 value  $\alpha_b^* = 1.36$ . This value has similarly been adopted within the Wilcox (2006) stress- $\omega$   
 1030 model (which did not originally include a buoyancy production term) in the present work.

1031 **Appendix B. Stability analysis of the Launder *et al.* (1975) stress- $\varepsilon$  model**

1032 The stress- $\varepsilon$  closure model of Launder *et al.* (1975) (called the LRR stress- $\varepsilon$  model in the  
1033 present work), with additional buoyancy production terms (derived similar to above), may  
1035 be written in full as:

$$\begin{aligned}
 & \underbrace{\frac{\partial \bar{\rho} \tau_{ij}}{\partial t}}_{\text{Time variation}} + \underbrace{\bar{u}_k \frac{\partial \bar{\rho} \tau_{ij}}{\partial x_k}}_{\text{Convection}} = - \underbrace{\bar{\rho} P_{ij}}_{\text{Production}} + \underbrace{\frac{2}{3} \bar{\rho} \varepsilon \delta_{ij}}_{\text{Dissipation}} - \underbrace{\bar{\rho} \Pi_{ij}}_{\text{Pressure-strain}} + \underbrace{\bar{\rho} \frac{C_\mu}{P_r} \frac{k^2}{\varepsilon} N_{ij}}_{\text{Buoyancy production}} \\
 & \underbrace{- C_s \frac{\partial}{\partial x_k} \left[ \frac{\bar{\rho} k}{\varepsilon} \left( \tau_{im} \frac{\partial \tau_{jk}}{\partial x_m} + \tau_{jm} \frac{\partial \tau_{ik}}{\partial x_m} + \tau_{km} \frac{\partial \tau_{ij}}{\partial x_m} \right) \right]}_{\text{Diffusion}}
 \end{aligned} \tag{B 1}$$

1038

$$\begin{aligned}
 & \underbrace{\frac{\partial \bar{\rho} \varepsilon}{\partial t}}_{\text{Time variation}} + \underbrace{\bar{u}_j \frac{\partial \bar{\rho} \varepsilon}{\partial x_j}}_{\text{Convection}} = \underbrace{\bar{\rho} C_{1\varepsilon} \frac{\varepsilon}{k} \tau_{ij} \frac{\partial \bar{u}_i}{\partial x_j}}_{\text{Production}} - \underbrace{\bar{\rho} C_{2\varepsilon} \frac{\varepsilon^2}{k}}_{\text{Dissipation}} - \underbrace{\bar{\rho} C_{1\varepsilon} C_{3\varepsilon} C_\mu \frac{1}{P_r} \frac{\varepsilon}{k} N^2}_{\text{Buoyancy production}} \\
 & \underbrace{- C_\varepsilon \frac{\partial}{\partial x_k} \left[ \frac{\bar{\rho} k}{\varepsilon} \tau_{km} \frac{\partial \varepsilon}{\partial x_m} \right]}_{\text{Diffusion}}
 \end{aligned} \tag{B 2}$$

1039

1041 where

$$\begin{aligned}
 & \Pi_{ij} = C_1 \frac{\varepsilon}{k} (\tau_{ij} + \frac{2}{3} k \delta_{ij}) - \hat{\alpha} (P_{ij} - \frac{2}{3} P \delta_{ij}) - \hat{\beta} (D_{ij} - \frac{2}{3} P \delta_{ij}) \\
 & - \hat{\gamma} k (S_{ij} - \frac{1}{3} S_{kk} \delta_{ij}) + \left[ 0.125 \frac{\varepsilon}{k} (\tau_{ij} + \frac{2}{3} k \delta_{ij}) - 0.015 (P_{ij} - D_{ij}) \right] \frac{k^{3/2}}{\varepsilon n}.
 \end{aligned} \tag{B 3}$$

1043 The last term on the right-hand side of (B 3) is the LRR stress- $\varepsilon$  wall-reflection term, where  
1044  $n$  is the distance normal to the surface. In the above  $\varepsilon$  is the turbulence dissipation rate, and  
1045  $P_{ij}$  and  $D_{ij}$  are respectively defined in (2.7) and (2.8). The closure coefficients are (Gibson  
1046 & Launder 1978):

$$\begin{aligned}
 & C_\mu = 0.09, \quad C_1 = 1.8, \quad C_2 = 0.60 \\
 & \hat{\alpha} = (8 + C_2)/11, \quad \hat{\beta} = (8C_2 - 2)/11, \quad \hat{\gamma} = (60C_2 - 4)/55 \\
 & C_s = 0.11, \quad C_\varepsilon = 0.18, \quad C_{1\varepsilon} = 1.44 \\
 & C_{2\varepsilon} = 1.92, \quad C_{3\varepsilon} = -0.33, \quad P_r = 0.85
 \end{aligned} \tag{B 4}$$

1048 with  $C_{3\varepsilon} = -0.33$  and (the Prandtl number)  $P_r = 0.85$  adopted from the standard  $k$ - $\varepsilon$  closure  
1049 model.

1050 Similar to the Wilcox (2006) stress- $\omega$  model, the governing equations for the LRR stress- $\varepsilon$   
1051 model defined in (B 1) and (B 2) can be simplified for stability analysis purposes in the 2D  
1052 potential flow region beneath propagating surface water waves. An additional assumption is  
1053 made that the term for the near-wall effect in the pressure-strain correlation is negligible.  
1054 This is justifiable in the potential flow region above the bottom boundary layer. Following  
1055 the derivation in Section 2.2, the analogous resulting simplified  $k$ ,  $\tau_{12}$  and  $\varepsilon$  equations for  
1056 the LRR stress- $\varepsilon$  model are:

$$\frac{\partial k}{\partial t} = 2\tau_{12} S_{12} - \varepsilon \tag{B 5}$$

1057

1058

1059

$$\frac{\partial \tau_{12}}{\partial t} = \left( \frac{4}{3} - \frac{4}{3} \hat{\alpha} - \frac{4}{3} \hat{\beta} + \hat{\gamma} \right) k S_{12} - C_1 \frac{\varepsilon}{k} \tau_{12} \quad (\text{B } 6)$$

1060

$$\frac{\partial \varepsilon}{\partial t} = 2C_{1\varepsilon} \frac{\varepsilon}{k} \tau_{12} S_{12} - C_{2\varepsilon} \frac{\varepsilon^2}{k} \quad (\text{B } 7)$$

1061

1062

1063

To perform a stability analysis on the three-equation system above, it turns to be convenient to utilize two utility variables, namely  $\Psi = k/\tau_{12}$  and  $\Xi = \varepsilon/\tau_{12}$ . The equations for these quantities work out to be:

1064

$$\frac{\partial \Psi}{\partial t} = \frac{\partial(k/\tau_{12})}{\partial t} = \underbrace{\left( \frac{4}{3} \hat{\alpha} + \frac{4}{3} \hat{\beta} - \hat{\gamma} - \frac{4}{3} \right)}_{-8/15} \Psi^2 S_{12} + (C_1 - 1) \Xi + 2S_{12} \quad (\text{B } 8)$$

1065

$$\frac{\partial \Xi}{\partial t} = \frac{\partial(\varepsilon/\tau_{12})}{\partial t} = \underbrace{\left( \frac{4}{3} \hat{\alpha} + \frac{4}{3} \hat{\beta} - \hat{\gamma} - \frac{4}{3} \right)}_{-8/15} \Psi \Xi S_{12} + (C_1 - C_{2\varepsilon}) \frac{\Xi^2}{\Psi} + 2C_{1\varepsilon} \frac{\Xi}{\Psi} S_{12} \quad (\text{B } 9)$$

1066

1067

Setting both (B 8) and (B 9) equal to zero, their (constant) asymptotic equilibrium values can be found as:

1068

$$\Psi_{\infty} = \pm \sqrt{6 \cdot \frac{C_1 + C_{1\varepsilon} - C_1 C_{1\varepsilon} - C_{2\varepsilon}}{(C_{2\varepsilon} - 1)(4\hat{\alpha} + 4\hat{\beta} - 3\hat{\gamma} - 4)}} \approx \pm 2.277 \quad (\text{B } 10)$$

1069

1070

$$\frac{\Xi_{\infty}}{S_{12}} = \frac{2(C_{1\varepsilon} - 1)}{C_{2\varepsilon} - 1} \approx 0.957 \quad (\text{B } 11)$$

1071

1072

1073

Thus the fix points for the nonlinear ODEs (B 8)–(B 9) are  $(\Psi_{\infty}, \Xi_{\infty}) = (\pm 2.277, 0.957 S_{12})$ . To check for formal stability of these two fixed points, the Jacobian matrix for (B 8)–(B 9) is defined as

1074

$$J = \begin{bmatrix} \frac{\partial}{\partial \Psi} \left( \frac{\partial \Psi}{\partial t} \right) & \frac{\partial}{\partial \Xi} \left( \frac{\partial \Psi}{\partial t} \right) \\ \frac{\partial}{\partial \Psi} \left( \frac{\partial \Xi}{\partial t} \right) & \frac{\partial}{\partial \Xi} \left( \frac{\partial \Xi}{\partial t} \right) \end{bmatrix} \quad (\text{B } 12)$$

1075

1076

After invoking the right-hand sides of (B 8)–(B 9) in the above, in addition to the model closure coefficients, this becomes:

1077

$$J = \begin{bmatrix} -1.067 S_{12} \Psi & 0.8 \\ \frac{0.12 \Xi^2}{\Psi^2} - 0.533 S_{12} \Xi - \frac{2.88 S_{12} \Xi}{\Psi^2} & -0.533 \Psi S_{12} + \frac{2.88 S_{12}}{\Psi} - \frac{0.24 \Xi}{\Psi} \end{bmatrix} \quad (\text{B } 13)$$

1078

1079

1080

After linearizing about the fixed points  $(\Psi_{\infty}, \Xi_{\infty})$ , the eigenvalues of  $J$  are found to be  $(-2.012, -0.4663)|S_{12}|$ . As these are negative, the fixed points correspond to stable nodes, similar to what was found for the stress- $\omega$  model.

1081

1082

Now inserting these fixed points  $(\Psi_{\infty}, \Xi_{\infty})$  back into (B 5) and (B 6) and simplifying, then leads to the following linearized equation for the exponential growth rate for  $k$ :

1083

$$\Gamma_{\infty} = \frac{1}{k} \frac{\partial k}{\partial t} = \frac{2S_{12} - \Xi_{\infty}}{\Psi_{\infty}} \quad (\text{B } 14)$$

1084 Substituting the closure coefficients finally yields:

$$1085 \quad \Gamma_{\infty} = (C_{2\varepsilon} - C_{1\varepsilon}) \sqrt{\frac{2}{3} \cdot \frac{(4\hat{\alpha} + 4\hat{\beta} - 3\hat{\gamma} - 4)}{(C_{2\varepsilon} - 1)(C_1 + C_{\varepsilon} - C_1 C_{1\varepsilon} - C_{2\varepsilon})}} \cdot |S_{12}| \approx 0.458 \cdot |S_{12}| \quad (\text{B } 15)$$

1086 As discussed in Section 2.2, since  $\langle S_{12} \rangle = 0$  in the idealized potential flow region beneath  
 1087 propagating waves, then  $\Gamma_{\infty}$  will (on average) likewise be zero. Therefore, this proves that,  
 1088 similar to the Wilcox (2006) stress- $\omega$  model, the LRR stress- $\varepsilon$  model is neutrally stable in  
 1089 the ideal potential flow region beneath non-breaking surface waves.

1090 While the model analyzed above has not been the main focus of the present work, for  
 1091 the sake of completeness the progressive wave train simulations from Section 3.1 have also  
 1092 been repeated using the LRR stress- $\varepsilon$  model, maintaining the same schemes and settings as  
 1093 before (maximum Courant number  $Co = 0.05$ ). The results for the free surface elevations are  
 1094 presented in figure 25(a, simulated with buoyancy production terms on) and figure 25(b, with  
 1095 buoyancy production terms off). They are unsurprisingly identical and similar to those from  
 1096 the Wilcox (2006) stress- $\omega$  model (figure 3a). The period- and depth-averaged  $k/k_0$  time  
 1097 series are presented in figure 25(e). The black solid line (with buoyancy production terms  
 1098 on) has an immediate decrease of  $k$ , while the black dashed line (with buoyancy production  
 1099 terms off) has a zero growth of  $k$  in the early stage and then decreases at the same rate  
 1100 as the solid line. Both simulations are stable, confirming our analysis. It is noted that the  
 1101 simulations with the Wilcox (2006) stress- $\omega$  model and LRR stress- $\varepsilon$  model are essentially  
 1102 consistent with buoyancy production terms on (comparing figures 3 and 25 in the black  
 1103 solid lines). Conversely, the wave trains simulated with buoyancy production terms off are  
 1104 different in the growth rate, i.e. the Wilcox (2006) stress- $\omega$  model has a zero growth rate  
 1105 in general (figure 3b, black dashed-dotted line), while the LRR stress- $\varepsilon$  model has a zero  
 1106 growth in the beginning and a decreasing  $k$  later on (figure 25e, black dashed line). This  
 1107 slight difference may due to the wall-reflection term in the LRR stress- $\varepsilon$  model which could  
 1108 be interesting to investigate in detail in future work. These results, combined with those  
 1109 in the main text, thus demonstrate that RSM models (both stress- $\omega$  and stress- $\varepsilon$  variants)  
 1110 are generally (neutrally) stable in the idealized potential flow region beneath non-breaking  
 1111 surface waves. They should therefore not be expected to suffer from the problem of unphysical  
 1112 over-production (exponential growth) of turbulence in potential flow regions prior to wave  
 1113 breaking, common to many two-equation turbulence closure models in their standard forms,  
 1114 as shown by Larsen & Fuhrman (2018).

1115 A final remaining open question (which we shall now attempt to close) is: Why then did  
 1116 Brown *et al.* (2016) experience pronounced over-production of turbulence prior to spilling  
 1117 breaking in their CFD simulation using the LRR stress- $\varepsilon$  model? In this context, it is important  
 1118 to emphasize that for the analysis (predicting neutral stability) above to hold in practice, a  
 1119 CFD model must maintain the nearly potential flow region beneath a surface wave with  
 1120 sufficient accuracy such that  $\langle S_{12} \rangle \approx 0$ . If this is not the case, since  $\Gamma_{\infty} \sim |S_{12}|$  in (B 15), then  
 1121 our analysis suggests RSMs may, in fact, still be prone to unphysical exponential growth of  
 1122 turbulence beneath non-breaking waves, if they do not solve the flow with sufficient accuracy.  
 1123 We hypothesize that this is precisely what has occurred in the simulation of Brown *et al.*  
 1124 (2016) mentioned just above. Note that Brown *et al.* (2016) utilized a significantly larger  
 1125 maximum Courant number ( $Co = 0.2$ , hence numerical time step) than considered herein  
 1126 (the present results have uniformly used  $Co = 0.05$ ). Moreover, Larsen *et al.* (2019) have  
 1127 specifically demonstrated that such large Courant numbers can indeed lead to pronounced  
 1128 inaccuracies in the resulting flow kinematics (hence  $S_{12}$ ), even beneath computed free surfaces  
 1129 that may otherwise appear reasonable. To test this hypothesis, we will repeat our simulation

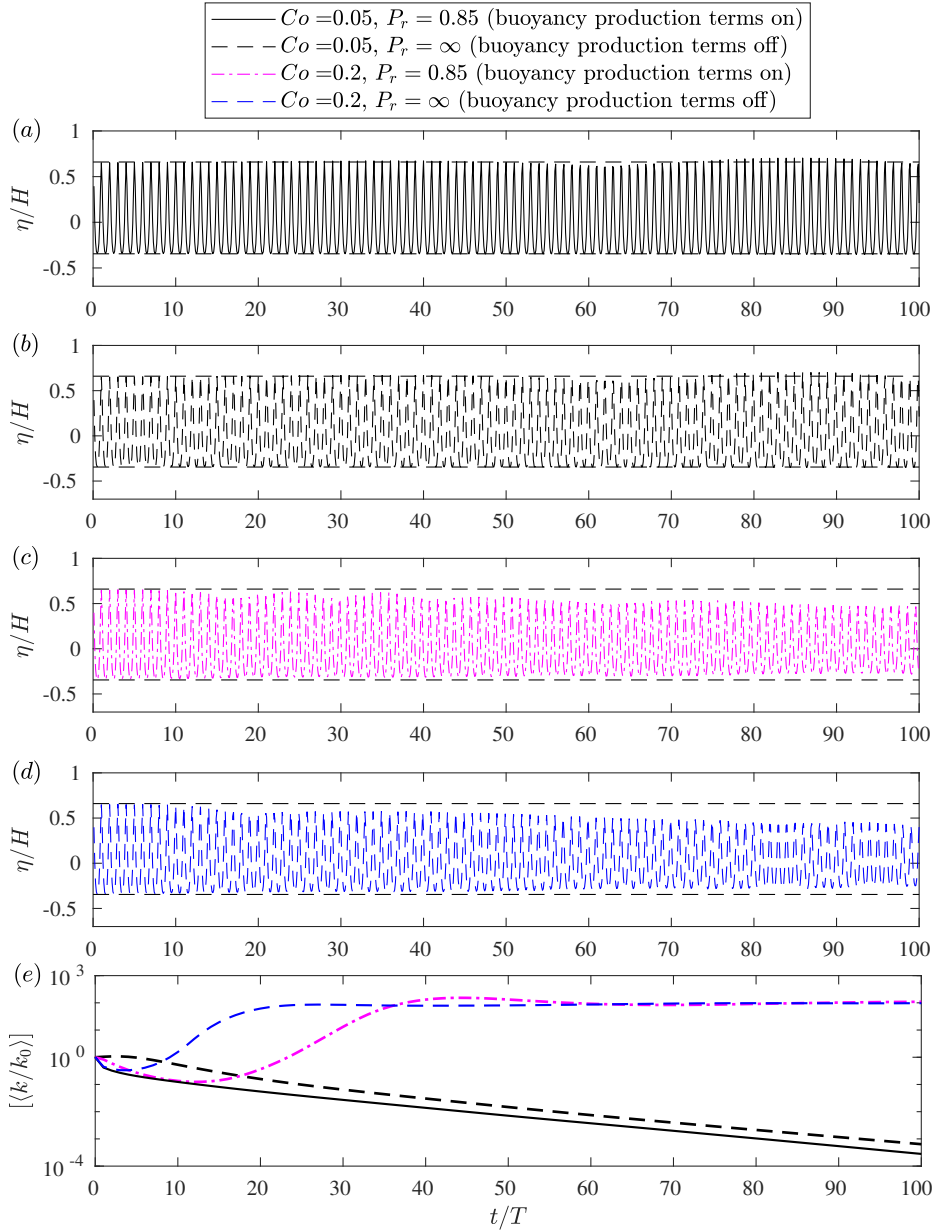


Figure 25: Computed time series of (a)–(d) surface elevations and (e) depth- and period-averaged turbulent kinetic energy beneath wave trains simulated with the LRR stress- $\varepsilon$  model. The results depicted as blue dashed lines in (d) and (e), with  $Co = 0.20$  and without buoyancy production terms, are chosen to match most closely those used by Brown *et al.* (2016).

1130 of the wave train above, but now with  $Co = 0.2$ , while also switching schemes to those stated  
 1131 by Brown *et al.* (2016). We consider two otherwise-identical simulations, having buoyancy  
 1132 production terms both on ( $P_r = 0.85$ , as before) and off ( $P_r = \infty$ , as in Brown *et al.* 2016).  
 1133 These results are respectively also shown as the pink (dashed-dotted) and blue (dashed)  
 1134 lines in figure 25. For the case believed to most-resemble the setup used by Brown *et al.*



1135 (2016) (blue dashed lines in figure 25*d,e*) it is seen that, due to accumulated numerical  
 1136 errors in the velocity kinematics, the turbulence indeed begins to grow exponentially already  
 1137 by  $t/T = 10$ . By  $t/T = 20$  the turbulence has reached several hundred times the initial  
 1138 level, becoming large enough to cause unphysical decay of the wave train. A similar (but  
 1139 delayed) process occurs for the case with buoyancy production terms on (figure 25*c,e*, pink  
 1140 dashed-dotted lines). Based on these results, it seems clear that the over-production of pre-  
 1141 breaking turbulence experienced by Brown *et al.* (2016) with the LRR stress- $\varepsilon$  model can  
 1142 be attributed to numerical inaccuracies in the velocity kinematics (hence  $S_{12}$ ) during their  
 1143 simulated shoaling stage. These inaccuracies can be attributed to the larger Courant number  
 1144 used, in accordance with what has been shown previously (there without a turbulence model)  
 1145 by Larsen *et al.* (2019). Because buoyancy production terms create a sink of turbulence in  
 1146 the air-water interface region, their inclusion may delay the onset of this problem, but will  
 1147 not eliminate it. Similar issues could be expected with the stress- $\omega$  model if accurate velocity  
 1148 kinematics are not maintained in nearly-potential flow regions beneath surface waves, since  
 1149 similarly  $\Gamma_\infty \sim |S_{12}|$  in (2.28), though the predicted growth rate would be smaller due to the  
 1150 lower coefficient in front of  $|S_{12}|$ .

### 1151 **Appendix C. Alternative stability analysis of the stress- $\omega$ model using eigenvalues**

1152 During the peer review process of the present paper, it became apparent that the stability of  
 1153 the turbulence closure models could be equivalently analyzed based on eigenvalues of the  
 1154 Jacobian matrix, after linearizing about the fixed points. We will hence briefly outline this  
 1155 procedure in what follows for the stress- $\omega$  closure model.

1156 The Jacobian matrix for the simplified stress- $\omega$  model governing equations in (2.18)–(2.20)  
 1157 is defined by:

$$1158 \quad J = \begin{bmatrix} \frac{\partial}{\partial k} \left( \frac{\partial k}{\partial t} \right) & \frac{\partial}{\partial \tau_{12}} \left( \frac{\partial k}{\partial t} \right) & \frac{\partial}{\partial \omega} \left( \frac{\partial k}{\partial t} \right) \\ \frac{\partial}{\partial k} \left( \frac{\partial \tau_{12}}{\partial t} \right) & \frac{\partial}{\partial \tau_{12}} \left( \frac{\partial \tau_{12}}{\partial t} \right) & \frac{\partial}{\partial \omega} \left( \frac{\partial \tau_{12}}{\partial t} \right) \\ \frac{\partial}{\partial k} \left( \frac{\partial \omega}{\partial t} \right) & \frac{\partial}{\partial \tau_{12}} \left( \frac{\partial \omega}{\partial t} \right) & \frac{\partial}{\partial \omega} \left( \frac{\partial \omega}{\partial t} \right) \end{bmatrix} \quad (C 1)$$

1159 After invoking the right-hand sides of (2.18)–(2.20) in the above, in addition to the model  
 1160 closure coefficients, this becomes:

$$1161 \quad J = \begin{bmatrix} -0.09\omega & 2S_{12} & -0.09k \\ 0.5\bar{3}S_{12} & -0.162\omega & -0.162\tau_{12} \\ -\frac{1.04S_{12}\tau_{12}\omega}{k^2} & \frac{1.04S_{12}\omega}{k} & \frac{1.04S_{12}\tau_{12}}{k} - 0.1416\omega \end{bmatrix} \quad (C 2)$$

1162 Further invoking  $k = \Psi\tau_{12}$  and linearizing about (i.e. inserting) the fixed points from (2.23)–  
 1163 (2.24), the eigenvalues of  $J$  are found to be:  $(-1.675, -0.5891, 0.2831)|S_{12}|$ . It is seen that  
 1164 the critical (third) eigenvalue matches precisely the asymptotic growth rate  $\Gamma_\infty$  from (2.28),  
 1165 confirming our analysis in the main text.

1166 Although we will not present full details for the sake of brevity, we have also conducted  
 1167 an analogous stability analysis of the LRR stress- $\varepsilon$  model equations defined in (B 5)–(B 7).  
 1168 Should the interested reader wish to repeat said analysis, we find that the eigenvalues of  
 1169 the Jacobian matrix, after linearizing about the fixed points for this system, correspond to:  
 1170  $(-1.555, -0.008060, 0.4583)|S_{12}|$ . It is again seen that the critical (third) eigenvalue matches  
 1171 precisely the growth rate  $\Gamma_\infty$  from (B 15).

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