

Secondary organic aerosol reduced by mixture of atmospheric vapours

DOI:

[10.1038/s41586-018-0871-y](https://doi.org/10.1038/s41586-018-0871-y)

Document Version

Accepted author manuscript

[Link to publication record in Manchester Research Explorer](#)

Citation for published version (APA):

Mcfiggans, G., Mentel, T. F., Wildt, J., Pullinen, I., Kang, S., Kleist, E., Schmitt, S., Springer, M., Tillman, R., Wu, C., Zhao, D., Hallquist, M., Faxon, C., Breton, M. L., Hallquist, Å. M., Simpson, D., Bergström, R., Jenkin, M. E., Ehn, M., ... Kiendler-Scharr, A. (2019). Secondary organic aerosol reduced by mixture of atmospheric vapours. *Nature*, 565, 587–593. <https://doi.org/10.1038/s41586-018-0871-y>

Published in:

Nature

Citing this paper

Please note that where the full-text provided on Manchester Research Explorer is the Author Accepted Manuscript or Proof version this may differ from the final Published version. If citing, it is advised that you check and use the publisher's definitive version.

General rights

Copyright and moral rights for the publications made accessible in the Research Explorer are retained by the authors and/or other copyright owners and it is a condition of accessing publications that users recognise and abide by the legal requirements associated with these rights.

Takedown policy

If you believe that this document breaches copyright please refer to the University of Manchester's Takedown Procedures [<http://man.ac.uk/04Y6Bo>] or contact uml.scholarlycommunications@manchester.ac.uk providing relevant details, so we can investigate your claim.



Secondary organic aerosol reduced by mixture of atmospheric vapours

Gordon McFiggans¹, Thomas F. Mentel², Jürgen Wildt^{2,3}, Iida Pullinen^{2,*}, Sungah Kang², Einhard Kleist³, Sebastian Schmitt^{2,#}, Monika Springer², Ralf Tillmann², Cheng Wu^{2,+}, Defeng Zhao^{5,2}, Mattias Hallquist⁶, Cameron Faxon⁶, Michael Le Breton^{1,6}, Åsa M. Hallquist⁷, David Simpson^{8,12}, Robert Bergström^{6,8,13}, Michael E. Jenkin⁹, Mikael Ehn¹⁰, Joel A. Thornton¹¹, M. Rami Alfarra^{1,14}, Thomas J. Bannan¹, Carl J. Percival^{1,°}, Michael Priestley¹, David Topping^{1,14}, Astrid Kiendler-Scharr^{2,4}

1. University of Manchester, School of Earth and Environmental Sciences, Oxford Road, Manchester, M13 9PL, UK

2. Institut für Energie- und Klimaforschung, IEK-8, Forschungszentrum Jülich, 52425 Jülich, Germany

3. Institut for Bio- and Geosciences, IBG-2, Forschungszentrum Jülich, Jülich 52425, Germany

4. I. Physikalisches Institut, Universität zu Köln, 50937 Köln, Germany

5. Department of Atmospheric and Oceanic Sciences & Institute of Atmospheric Sciences, Fudan University, Shanghai, China

6. Atmospheric Science, Department of Chemistry and Molecular Biology, University of Gothenburg, SE-41296 Gothenburg, Sweden

7. IVL Swedish Environmental Research Institute, PO Box 5302, SE-400 14 Gothenburg, Sweden

8. Department of Earth, Space and Environment, Chalmers University of Technology, 41296 Gothenburg, Sweden

9. Atmospheric Chemistry Services, Okehampton, Devon, EX20 4QB, UK

10. Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of Helsinki, P.O. Box 64, 00014, Helsinki, Finland

11. Department of Atmospheric Sciences, University of Washington, Seattle, WA98195, USA

12. EMEP MSC-W, Norwegian Meteorological Institute, Oslo, Norway

13. Swedish Meteorological and Hydrological Institute, 60176 Norrköping, Sweden

14. National Centre for Atmospheric Science (NCAS), Oxford Road, Manchester, M13 9PL, UK

* now at: Department of Applied Physics, University of Eastern Finland, 70211 Kuopio, Finland

now at: TSI GmbH, 52068 Aachen, Germany

+ now at: Stockholm University, Department of Environmental Science & Analytic Chemistry, SE-10691 Stockholm, Sweden

° now at: Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA 91109, USA

Secondary organic aerosol contributes to the atmospheric particle burden with implications for air quality and climate. Biogenic volatile organic compounds emitted from plants are important secondary organic aerosol precursors with isoprene emissions dominating globally. However, its yield of particle mass from oxidation is generally modest compared to that of other terpenoids. Here we show that isoprene, carbon monoxide and methane can suppress the instantaneous mass and the overall mass yield derived from monoterpenes in mixtures. We find that isoprene scavenges hydroxyl radicals preventing reaction with monoterpenes and the resulting isoprene peroxy radicals scavenge highly oxygenated monoterpene products. These effects reduced the yield of low-volatility products that would otherwise form secondary organic aerosol. Global model calculations indicate that oxidant and product scavenging can operate effectively in the real atmosphere. Highly reactive, modest aerosol yield compounds are not necessarily net producers and their oxidation can suppress both particle number and mass.

Introduction

The secondary organic fraction of fine aerosol is substantial^{1,2} and contributes significantly to the fine particulate matter (PM) burden³ and aerosol impacts on climate⁴. Conventionally, Secondary Organic Aerosol (SOA) particle mass is considered to be formed independently from the condensable oxidation products of each contributing volatile organic compound (VOC)⁵. The ambient atmosphere is a complex mixture of biogenic and anthropogenic VOC, a limited number of which are normally considered to control the oxidant concentrations in air quality or climate models^{4,5}. Isoprene has been found to dominate plant VOC emissions globally^{6,7}, though it has been found only moderately effective at forming SOA particle mass^{8,9,10}. Nevertheless, isoprene oxidation tracers have been observed in particles in both chamber experiments^{11,12} and the atmosphere^{11,13,14,15} showing that they condense from the gas phase. Oxidation of the monoterpenes (MT), including α -pinene, and sesquiterpenes is generally more efficient at producing SOA particle mass^{16,17,18}. When oxidation of a VOC contributing to SOA is explicitly considered in models, some interaction between the oxidant field and the SOA formation process is implicit⁵, though recently observed behaviour is yet to be described. Two such recent findings have motivated the current work. First, it has been shown that C₅ isoprene molecules suppress the nucleation of SOA particles in the oxidation of the C₁₀ MT in mixtures of plant emissions by successfully competing for the hydroxyl radical, OH^{19,20,21}. Second, recent direct observation of highly oxygenated organic molecules (HOM) formed from MT oxidation²², and elucidation of the auto-oxidation mechanisms by which they are formed, have implicated them in the formation of new SOA particles. Our work moves beyond the suppression of particle nucleation by isoprene using the new mechanistic understanding to explain SOA mass and yield suppression in MT-containing mixtures. These results highlight a need for more realistic consideration of SOA formation in the atmosphere analogous to the treatment of ozone formation, where interactions between the mechanistic pathways involving peroxy radicals are recognised to be essential²³.

90 **Use of SOA yields in mixtures**

91 Ostensibly, the concept of a SOA yield is straightforward and unambiguous²⁴⁻²⁶. It has
92 been widely used to interpret the potential of precursor molecules to produce
93 particulate mass, most extensively from smog chamber studies and usually with the
94 goal of quantifying the formation of ambient atmospheric particles. Yield is normally
95 defined in terms of the particulate mass condensed for a given mass of gaseous parent
96 VOC consumed. The process of determining SOA yields presents substantial practical
97 challenges. Notwithstanding measurement uncertainties and artefacts (for example,
98 wall losses and those associated with the accurate determination of the mass of semi-
99 volatile material²⁷⁻³¹) there are numerous reasons why SOA yields may vary
100 according to the conditions under which they are measured. This paper does not aim
101 to provide a critical evaluation of the determination of SOA yields and their
102 extrapolation to atmospheric conditions and the reader is referred to a number of
103 recent publications for a discussion of challenges and the state-of-the-science
104 concerning the interpretation of SOA formation from chamber experiments^{22,27,29-32}.
105 However, there are conceptual aspects often implicit in the treatment of SOA
106 formation either in chamber experiments, real atmospheric mixtures, or models of the
107 atmosphere that provide the context for the current study.

108 First, the concept of yield frequently includes some assumption of equilibrium. An
109 interpretation of particle mass in terms of the partitioning of components of known
110 volatility by absorptive partitioning invokes an equilibrium assumption. In reality,
111 mass takes a finite time to transfer between the continuous and dispersed phases^{33,34}.
112 Second, a parent VOC will continue to react, as will its reaction products, provided
113 they have reaction partners. Under atmospheric, or simulated atmospheric, conditions
114 the reaction partners are normally oxidants, such as the hydroxyl radical OH (the
115 main focus in this paper). This means that the distribution of vapour phase
116 components that is available to transfer to the condensed phase by condensational
117 growth or nucleation is continuously evolving through photochemical reactions.
118 Third, a yield integrates across all oxidation products formed from the parent
119 hydrocarbon, summing the fraction of these products that will partition to the particle
120 phase from whatever stage of oxidation under the chosen reaction conditions. We
121 further would like to note that the yields themselves can be dependent on the oxidant
122 concentration^{e.g.31,32}.

123 This combination of characteristics raises first order challenges when considering
124 SOA formation in the real atmosphere which comprises a complex mixture of organic
125 vapours of widely varying volatilities, from numerous sources of anthropogenic and
126 biogenic origin. This paper is concerned with the challenge surrounding the
127 introduction of oxidants into the atmospheric mixture and the reactivity of the
128 evolving components in the mixture – specifically, the yield of condensed organic
129 particulate material of any of the single component vapours in the mixture when it is
130 consumed by the oxidant. Any of the vapours may react with the available oxidant, so
131 from the perspective of any other vapour molecule, the oxidant has been “scavenged”
132 and is unavailable as a reaction partner. Our experimental systems illustrate the
133 implications on SOA formation of mixtures of precursors. We demonstrate how SOA
134 yield in mixtures is dependent on whether the reactivity of one of the SOA forming
135 compounds controls the oxidant and contrast conditions in the laboratory and the
136 atmosphere. It is not only the oxidants that can be removed from the system by
137 components of a mixture. The products from oxidation that can react to form the final
138 condensable SOA precursors may alternatively be scavenged by other reactive

intermediates. This process can form the basis for the reduction in yield of the individual components in a mixture, as shown by our results below and detailed in the supplement sections.

Experiments to SOA formation in mixtures

In the Jülich Plant Atmosphere Chamber (JPAC; suppl. section 1) we find that the presence of isoprene substantially suppresses the SOA mass formed from oxidation of α -pinene, the most abundant MT in the atmosphere (Figure 1A). Furthermore, we show that this reduction of SOA mass is not trivially due to the lower amount of MT consumed in the presence of isoprene. Rather the presence of isoprene actually suppresses the SOA mass yield from the amount of VOC consumed (Figure 1B). Isoprene thereby limits not only new particle formation¹⁹ (hence exhibiting control over particle number concentration) but also the growth of pre-existing particles and thus their size. This has significant consequences for the abundance of PM mass formed in the mixture and the number of particles that may grow to sizes where they can act as cloud seeds. Our findings are surprising and unexpected in the context of existing conceptions of SOA mass yields (suppl. section 2). As shown in Figure 1, the presence of isoprene substantially suppresses the SOA mass formation from α -pinene oxidation by about 60% and the SOA yield by 40%. Figure S2 shows even greater suppression in greater isoprene excess, with both mass and yield reduced in the presence of both neutral and acidic seed. This contrasts sharply with current understanding, since isoprene oxidation should contribute substantially to SOA production when isoprene concentrations are high.

Below we describe the two effects contributing to SOA suppression - oxidant scavenging and product scavenging. This demonstrates the importance of accounting for the reactivity of the parent VOC in atmospheric mixtures and hence their turnover (suppl. section 2) as well as interactions between their reaction intermediates and products. Whilst our experiments illustrate this requirement predominantly with reference to the photochemical isoprene – α -pinene system with OH as the major oxidant, it should be noted that in all mixed systems investigated, OH (oxidant) scavenging and product scavenging were found to reduce both the SOA particle mass and the SOA yield. Note that during daytime, both, ozonolysis and OH oxidation of α -pinene, contribute to SOA, as in our experiments, however the contribution via OH reactions here was much greater than 90%. According to our mechanistic understanding presented below, product scavenging will be similarly effective for OH and O₃ oxidation.

Oxidant and product scavenging

Oxidant scavenging is demonstrated by the decrease in SOA mass and yield of the mixture with increasing isoprene fraction as well as by the decrease of the α -pinene yield itself (Figure 1B and suppl. section 3). The contribution of isoprene to SOA mass in the presence of neutral ammonium sulfate aerosols is small and α -pinene oxidation products make up the overwhelming fraction of SOA in mixtures. Isoprene increasingly efficiently competes with α -pinene for the available OH as the reaction system is stepped from high SOA yield pure α -pinene to low SOA yield pure isoprene (Figure S2). With respect to α -pinene, the isoprene “scavenges” the OH¹⁹, i.e. the steady state OH concentration is lowered by the introduction of isoprene for a given

OH source strength, therefore less OH can react with α -pinene, resulting in less SOA mass being formed. In addition, the SOA yield itself is dependent on $[\text{OH}]^{31,32}$.

By uniquely being able to control the OH at the same concentration with and without isoprene addition by readjusting its source strength, we were able to reveal that the SOA mass (Figure 1A), and by implication the α -pinene SOA yield (Figure 1B, 3rd bar), decreased with increasing isoprene in the mixture (suppl. section 4). Strikingly, the α -pinene SOA yield was reduced by 30% relative to the pure α -pinene yield, even though the OH scavenging effect has been removed ($\Delta\text{isoprene}/\Delta\alpha\text{-pinene} \approx 1 \pm 0.2$ [ppb/ppb] = 0.5 ± 0.1 [$\mu\text{g}/\mu\text{g}$]). Moreover, by re-adjusting $[\text{OH}]$, we held the α -pinene sub-system of the mixture at the same stage of oxidation as in the pure case. This results in the condensing oxidation products having the same “chemical age” in both cases, thereby avoiding comparison of chemical systems of different chemical evolution.

In Figure 2 we show the actual α -pinene SOA yields (y_{act}) in the mixture normalized to the α -pinene SOA yield $y_{\text{AP}} = 17\%$ for pure α -pinene in absence of isoprene. The y_{act} were calculated from the observed SOA mass and the actual consumption of α -pinene. Depending on the isoprene to α -pinene ratio, SOA mass produced from the same amount of consumed α -pinene in the presence of isoprene can be reduced by more than half compared to the amount formed when oxidised alone (Figure 2). Clearly, smaller amounts of SOA forming products from α -pinene oxidation are formed in the presence of isoprene – we denote this effect “product scavenging”.

The product scavenging effect (and the OH scavenging effect) on SOA formation is somewhat masked in the presence of the acidic seed aerosol with which isoprene has a significant yield^{12,35-37} (suppl. section 4, Figure S3, Figure 1B, 6th bar). Replacing isoprene by CO shows, that the acidity of the seed particles did not affect the behaviour of α -pinene (Figure S3), but indeed enhanced the isoprene contribution (Figure S4). This underlines the importance of liquid phase processes for isoprene, but also clearly shows that both scavenging effects are general phenomena in the gas phase.

SOA yield suppression also occurs when isoprene is mixed with β -pinene and when CO or CH_4 were mixed with α -pinene (suppl. section 5, Figures S5-S7, Figure 4). This generality ensures that the influence of the relative reactivity of components in mixtures on the SOA formation potential has much more profound implications than a simple reduction in yield in each of the binary mixtures. Moreover, because of potential interactions with additional SOA contributors, the yields in anything beyond the binary mixtures cannot be simply added or in any way predicted without specific mechanistic insight.

The key to a mechanistic understanding of the product scavenging is the formation of HOM with the peculiarity that HOM peroxy radicals also dimerise in the gas phase³⁸ (suppl. section 6). It can be inferred from their structure and O/C ratios that HOM and even more so their dimers, are low to extremely low volatility organic compounds^{22,39} and will contribute substantially to SOA formation and initial growth^{22,39}.

In Figure 3A we present direct observation of the product scavenging of α -pinene HOM by comparing mass spectra measured in OH reaction systems for α -pinene-only and for an α -pinene/isoprene mixture. For α -pinene-only we observed about equal signal intensities in the ranges of α -pinene HOM monomers with 10 C-atoms and of HOM dimers with 17-20 C-atoms. In the presence of isoprene, α -pinene dimers are suppressed by a factor of about 3. Comparable dimer suppression by a factor of 2 is observed for CO (Figure 4A). HOM dimers have been observed in the boreal ambient atmosphere, lower during daytime than at night-time, but at lower fractions than in

our α -pinene-only experiments^{40,41}. Since product scavenging will be occurring, particularly during daytime, these observations are consistent with our findings. The strength of product scavenging is estimated by comparing two sets of experiments where the [OH] was tuned, either by varying the OH sink through adding different amounts of isoprene to the reaction system or by varying the OH production rate through modifying the photolysis rate $J(\text{O}^1\text{D})$ for O_3 (Figure 3B). With increasing isoprene, the dimers are more strongly suppressed than with decreasing $J(\text{O}^1\text{D})$ as a result of increased scavenging of α -pinene HOM peroxy radicals by isoprene related RO_2 and HO_2 radicals⁴². This leads to formation of less HOM, mainly by diminishing α -pinene HOM self-dimerization. Suppression of the HOM dimers alone accounts for 27% reduction in HOM mass (Figure 3C) and explains most of the 35% SOA mass reduction by the product scavenging effect shown in Figure 1, underlining the pivotal role of HOM dimers in new particle and SOA formation^{22,39}. The involvement of HOM does not preclude the involvement of multigenerational oxidation as such (cf. suppl. section 6). Indeed, the non-linear dependence of the total HOM on the turnover in Figure 3B requires at least a second oxidation step and explains the dependence of the α -pinene SOA yield on [OH], contributing to the OH scavenging effect. The mechanism of product scavenging via dimer suppression by short chained peroxy radicals or HO_2 is general for peroxy radical chemistry and is not limited to the presence of isoprene in the mixture. It is effective in mixtures of α -pinene with CO via HO_2 (Figure 4) and with CH_4 via CH_3O_2 radicals (Figures S7). CO and CH_4 are extreme cases, each with essentially no SOA yield. In both cases, the scavenging effects are solely related to gas-phase chemistry and independent of aerosol composition: organic or inorganic neutral or inorganic acidic particles. We therefore conclude that product scavenging effect is a general effect in mixtures of VOC.

Atmospheric implications

Our new findings in JPAC would hold in atmospheric concentration regimes (suppl. section 7). However, in the atmosphere, the OH concentration is not solely controlled by the isoprene / α -pinene ratio as in the JPAC chamber. Simulations using the EMEP MSC-W model^{43,44} (suppl. section 8) show that the scavenging effects by isoprene under realistic atmospheric conditions lowers the OH concentration compared to the hypothetical "no isoprene" case and affects the SOA formation. However, even at small isoprene SOA yields, the gains of isoprene SOA and MT- O_3 -SOA offset the reduction from OH scavenging (Figure 5, noIso, OHS). When a parameterisation of the product scavenging constrained to our measurements (Figure 2) is included in the global model simulations, we find that photochemically generated MT-OH-SOA is strongly suppressed, in areas with sufficiently high MT and isoprene emissions (Figure 5, PS). The dramatic effects on MT-OH-SOA show the potential for product scavenging to reduce aerosol loadings substantially in many regions of the world, even more so if one considers that product scavenging similarly could also affect e.g. the MT- O_3 -SOA (Figure 5, PSox). It is not possible to be so sure about the magnitude of OH scavenging from all potential reactants, since more measurement constraint is needed to include mechanistic descriptions in such systems.

Extending our results from mixtures of isoprene, CH_4 , and CO with α -pinene or β -pinene to the reactive atmospheric mixture of vapours, any of the reaction products from any of the components in the mixture may react together, such that from the perspective of any of the others, they have been "scavenged" from the mixture. Similarly, from the perspective of the formation of organic particulate material, some of the condensable oxidation products that are direct precursors may be "scavenged",

in all cases reducing the ambient yield. Spatial distribution of short chain organic compounds and specific composition of the reaction mixture will determine the strength of the OH and the product scavenging effects. Whilst long lived compounds, like CH₄, may act ubiquitously in cleaner regions of the earth, compounds with strong anthropogenic sources, like CO, may not exert their full scavenging potential because, when high in concentration, they are always accompanied by high concentrations of other pollutants, like NO_x. While oxidant scavenging will take place under nearly all conditions, product scavenging across SOA precursors is a complex function RO₂-terminating agents, such as peroxy radicals of VOC with moderate or no SOA yield, HO₂, and NO_x. Product scavenging and “non-additivity” of SOA yields will thus vary from region to region, accordingly. Without further experimentally constrained parameterisation of the interactions between all relevant components in realistic atmosphere, it is not possible to predict the magnitude and geographical distribution of suppression of SOA formation. Inclusion of explicit interactions of such complexity in global simulations is practically impossible. It is only possible to say for the considered systems that the SOA yields and net SOA mass must be reduced compared to linear addition of yields, even if the control of oxidant were realistically included.

Outlook and recommendations

Our findings demonstrate that SOA precursors with modest yields suppress SOA forming products of higher SOA yield precursors substantially. Our work directly challenges two widespread treatments of SOA formation and suggests an alternative mechanistically-based solution. A first obvious recommendation following this work is that the decoupling of the modulation of oxidant fields from SOA precursor turnover is clearly discouraged. We have further demonstrated that simple linear addition of SOA mass from the individual yields of components in a VOC mixture will likely lead to a substantial overestimation of SOA production. Measuring SOA yields with individual compounds leads to insight into the mechanisms of SOA production but the results do not reflect the conditions of the real environment. Such data should therefore be used with caution when modelling aerosol formation. In the general case, the abundance of HOM, HOM-RO₂ and RO₂ products from potential SOA precursors as well as from volatile compounds not producing SOA mass should be considered when predicting the mixture’s yield. Our results highlight a need for more realistic consideration of SOA formation in the atmosphere analogous to the treatment of ozone formation, where interactions between the mechanistic pathways involving peroxy radicals are recognised to be essential²³. It may be possible to produce lumped classes of precursors according to their likelihood of increasing or decreasing a mixture’s yield and this could be a focus of further investigation. It is likely that both background oxidant concentrations and VOC emissions (and hence OH reactivity) will change in the future. Without a reasonable representation of SOA yields in different atmospheric VOC mixtures it will not be possible to achieve predictive capability for the SOA contribution to particulate matter.

330 References

- 331
- 332 1. Hallquist, M. *et al.* The formation, properties and impact of secondary organic
- 333 aerosol: current and emerging issues. *Atmos. Chem. Phys.* **9**, 5155-5236 (2009).
- 334
- 335 2. Jimenez, J. L. *et al.* Evolution of Organic Aerosols in the Atmosphere. *Science*,
- 336 **326**, 1525-1529 (2009)
- 337
- 338 3. Goldstein, A. H. & I. E. Galbally, Known and Unexplored Organic Constituents in
- 339 the Earth's Atmosphere, *Environ. Sci. & Technol.* **41**, 1514-1521 (2007).
- 340
- 341 4. Spracklen, D.V. *et al.* Aerosol mass spectrometer constraint on the global
- 342 secondary organic aerosol budget. *Atmos. Chem. Phys.* **11**, 12109-12136 (2011).
- 343
- 344 5. Kanakidou, M. *et al.* Organic aerosol and global climate modelling: a review.
- 345 *Atmos. Chem. Phys.* **5**, 1053-1123 (2005).
- 346
- 347 6. Guenther, A. *et al.* Estimates of global terrestrial isoprene emissions using
- 348 MEGAN (Model of Emissions of Gases and Aerosols from Nature). *Atmos. Chem.*
- 349 *Phys.* **6**, 3181-3210 (2006).
- 350
- 351 7. Guenther, A. *et al.* The Model of Emissions of Gases and Aerosols from Nature
- 352 version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic
- 353 emissions. *Geosci. Model Dev.* **5**, 1471-1492, (2012).
- 354
- 355 8. Carlton, A.G., Wiedinmyer, C., & Kroll, J.H. A review of Secondary Organic
- 356 Aerosol (SOA) formation from isoprene. *Atmos. Chem. Phys.* **9**, 4987-5005 (2009).
- 357
- 358 9. Clark, C.H. *et al.* Temperature Effects on Secondary Organic Aerosol (SOA) from
- 359 the Dark Ozonolysis and Photo - Oxidation of Isoprene. *Environ. Sci. & Techn.* **50**,
- 360 5564-5571 (2016).
- 361
- 362 10. Liu, J. *et al.* Efficient Isoprene Secondary Organic Aerosol Formation from a
- 363 Non - IEPOX Pathway. *Environ. Sci. & Technol.* **50**, 9872-9880, (2016).
- 364
- 365 11. Edney, E.O. *et al.* Formation of 2-methyl tetrols and 2-methylglyceric acid in
- 366 secondary organic aerosol from laboratory irradiated isoprene/NOX/SO2/air mixtures
- 367 and their detection in ambient PM2.5 samples collected in the eastern United States.
- 368 *Atmos. Environ.* **39**, 5281-5289 (2005).
- 369
- 370 12. Surratt, J.D. *et al.* Reactive intermediates revealed in secondary organic aerosol
- 371 formation from isoprene. *Proc. Natl. Acad. Sci. USA* **107**, 6640-6645 (2010).
- 372
- 373 13. Claeys, M. *et al.* Formation of secondary organic aerosols from isoprene and its
- 374 gas-phase oxidation products through reaction with hydrogen peroxide. *Atmos.*
- 375 *Environ.* **38**, 4093-4098 (2004).
- 376
- 377 14. Robinson, N.H. *et al.* Evidence for a significant proportion of Secondary Organic
- 378 Aerosol from isoprene above a maritime tropical forest. *Atmos. Chem. Phys.* **11**,
- 379 1039-1050 (2011).

380
381 15. Xu L. *et al.* Effects of anthropogenic emissions on aerosol formation from
382 isoprene and monoterpenes in the southeastern United States. *Proc. Natl. Acad. Sci.*
383 *USA* **112**, 37-42 (2015).
384
385 16. Lee, A. *et al.* Gas-phase products and secondary aerosol yields from the
386 ozonolysis of ten different terpenes, *J. Geophys. Res.* **111**, D07302 (2006).
387
388 17. Lee, A., *et al.* Gas-phase products and secondary aerosol yields from the
389 photooxidation of 16 different terpenes, *J. Geophys. Res.* **111**, D17305 (2006).
390
391 18. Ng, N. L. *et al.* Effect of NO_x level on secondary organic aerosol (SOA)
392 formation from the photooxidation of terpenes. *Atmos. Chem. Phys.* **7**, 5159-5174
393 (2007).
394
395 19. Kiendler-Scharr, A. *et al.* New particle formation in forests inhibited by isoprene
396 emissions. *Nature* **461**, 381-384 (2009).
397
398 20. Kanawade, V.P. *et al.* Isoprene suppression of new particle formation in a mixed
399 deciduous forest. *Atmos. Chem. Phys.* **11**, 6013-6027 (2011).
400
401 21. Lee, SH *et al.* Isoprene suppression of new particle formation: Potential
402 mechanisms and implications. *J. Geophys. Res. Atmos.* **121**, 14,621-14,635 (2016).
403
404 22. Ehn, M. *et al.* A large source of low-volatility secondary organic aerosol. *Nature*
405 **506**, 476-479 (2014).
406
407 23. Jenkin, M.E., Derwent, R.G. & Wallington, T.J. Photochemical ozone creation
408 potentials for volatile organic compounds: Rationalization and estimation. *Atmos.*
409 *Environ.* **163**, 128-137 (2017).
410
411 24. Odum, J. R. *et al.* Gas/particle partitioning and secondary organic aerosol yields.
412 *Environ. Sci. & Tech.* **30**, 2580-2585 (1996).
413
414 25. Hoffmann, T. *et al.* Formation of organic aerosols from the oxidation of biogenic
415 hydrocarbons. *J. Atmos. Chem.* **26**, 189-222 (1997).
416
417 26. Seinfeld, J. H. & Pankow, J. F. Organic atmospheric particulate material. *Ann.*
418 *Rev. Phys. Chem.* **54**, 121-140 (2003).
419
420 27. Matsunaga, A. & Ziemann, P. J. Gas-Wall Partitioning of Organic Compounds in
421 a Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield
422 Measurements. *Aerosol Sci. Tech.* **44**, 881-892 (2010).
423
424 28. Zhang, X. *et al.* Influence of vapor wall loss in laboratory chambers on yields of
425 secondary organic aerosol. *Proc. Natl. Acad. Sci. USA* **111**, 5802-5807 (2014).
426
427 29. Zhang, X. *et al.* Vapor wall deposition in Teflon chambers. *Atmos. Chem. Phys.*
428 **15**, 4197- 4214 (2015).
429

430 30. Krechmer, J. E., Pagonis, D., Ziemann, P. J. & Jimenez, J. L. Quantification of
 431 Gas-Wall Partitioning in Teflon Environmental Chambers Using Rapid Bursts of
 432 Low-Volatility Oxidized Species Generated in Situ. *Environ. Sci. Tech.* **50**, 5757-
 433 5765 (2016).
 434
 435 31. Sarrafzadeh, M. *et al.* Impact of NO_x and OH on secondary organic aerosol
 436 formation from β -pinene photooxidation. *Atmos. Chem. Phys.* **16**, 11237-11248
 437 (2016).
 438
 439 32. Eddingsaas, N. C. *et al.* Alpha-pinene photooxidation under controlled chemical
 440 conditions - Part 2: SOA yield and composition in low- and high-NO_x environments.
 441 *Atmos. Chem. Phys.* **12**, 7413-7427, (2012).
 442
 443 33. Zhang, X., Pandis, S. N. & Seinfeld, J. H. Diffusion-Limited Versus Quasi-
 444 Equilibrium Aerosol Growth. *Aerosol Sci. Tech.* **46**, 874-885 (2012).
 445
 446 34. O'Meara, S., Topping, D. O. & McFiggans, G. The rate of equilibration of viscous
 447 aerosol particles. *Atmos. Chem. Phys.* **16**, 5299-5313 (2016).
 448
 449 35. Surratt, J. D. *et al.* Effect of acidity on secondary organic aerosol formation from
 450 isoprene. *Environ. Sci. & Technol.* **41**, 5363-5369 (2007).
 451
 452 36. Gaston, C. J. *et al.* Reactive Uptake of an Isoprene-Derived Epoxydiol to
 453 Submicron Aerosol Particles. *Environ. Sci. & Technol.* **48**, 11178-11186 (2014).
 454
 455 37. Riva, M. *et al.* Effect of Organic Coatings, Humidity and Aerosol Acidity on
 456 Multiphase Chemistry of Isoprene Epoxydiols. *Environ. Sci. & Technol.* **50**, 5580-5588
 457 (2016).
 458
 459 38. Berndt, T. *et al.* Accretion Product Formation from Self- and Cross-Reactions of
 460 RO₂ Radicals in the Atmosphere. *Angew. Chem. Int. Edit.* **57**, 3820-3824 (2018).
 461
 462 39. Tröstl, J. *et al.* The role of low-volatility organic compounds in initial particle
 463 growth in the atmosphere. *Nature* **533**, 527-531 (2016).
 464
 465 40. Mohr, C. *et al.* Ambient observations of dimers from terpene oxidation in the gas
 466 phase: Implications for new particle formation and growth. *Geophys. Res. Lett.* **44**,
 467 2958-2966 (2017).
 468
 469 41. Yan, C. *et al.* Source characterization of highly oxidized multifunctional
 470 compounds in a boreal forest environment using positive matrix factorization. *Atmos.*
 471 *Chem. Phys.* **16**, 12715-12731 (2016).
 472
 473 42. Wennberg, P. O. *et al.* Gas-Phase Reactions of Isoprene and Its Major Oxidation
 474 Products. *Chem. Reviews* **118**, 3337-3390 (2018).
 475
 476 43. Simpson, D. *et al.* The EMEP MSC-W chemical transport model - technical
 477 description. *Atmos. Chem. Phys.* **12**, 7825-7865 (2012).
 478

44. Stadtler, S. *et al.*, Ozone impacts of gas-aerosol uptake in global chemistry-transport models. *Atmos. Chem. Phys.* **18**, 3147-3171 (2018)

Supplementary Information is available in the online version of the paper

Author contributions:

G.M., T.F.M., J.W. edited the manuscript and the supplement
G.M., T.F.M., J.W., A.K.S., M.H., D.S. M.E.J. conceptualized and planned the study, and conducted data interpretation
J.W. I.P., S.K., E.K., S.S., M.S., R.T., C.W., D.Z., C.F., M.L.B., Å.M.H., M.R.A., T.J.B., C.J.P., M.P., D.T conducted data collection and analysis
D.S., R.B., M.E.J. contributed the global model calculations
J.T., M.E., Å.M.H., M.H. provided specific inputs to manuscript and supplement
All co-authors discussed the results and commented the manuscript and the

Acknowledgements

The EMEP modelling work has been partially funded by EMEP under UNECE. Computer time for EMEP model runs was supported by the Research Council of Norway through the NOTUR project EMEP (NN2890K) for CPU, and NorStore project European Monitoring and Evaluation Programme (NS9005K) for storage of data. The research presented is a contribution to the Swedish strategic research area Modelling the Regional and Global Earth system, MERGE. This work was supported by Formas (grant numbers 214-2010-1756, 942-2015-1537); the Swedish Research Council (grant number 2014-5332) and the European Research Council (Starting grant no 638703, "COALA"). ÅMH acknowledge Formas (214-2013-1430) and Vinnova, Sweden's Innovation Agency (2013-03058), including support for her research stay at FZJ. Participation of the Manchester group was facilitated by the UK Natural Environment Research Council (NERC) funded CCN-Vol project (NE/L007827/1) and underpinning National Centre for Atmospheric Science (NCAS) funding.

Author Information Reprints and permissions information is available at <http://www.nature.com/>. The authors declare no competing interests. Correspondence and requests for data should be addressed to T.F.M (t.mentel@fz-juelich.de).

Data availability statements

All data used are shown as Figures or Tables in the manuscript or in the supplemental information. Raw data are available from the corresponding author on reasonable request.

Figure Captions

Figure 1: Reduced SOA mass and yield of α -pinene by product scavenging and OH scavenging by isoprene.

The top panel depicts symbolically the transition of the reaction system. **a:** Neutral seed particles: isoprene (red triangles) lowered [OH] (black diamonds) and the corresponding consumption of α -pinene (blue squares), reducing the organic mass dramatically. The organic mass was normalized to sulfate mass (organics/SO₄²⁻, green circles). The newly discovered product scavenging (red) in the presence of isoprene is separated from OH scavenging (orange) by re-adjustment of [OH]. Because of product scavenging only 60% of the organic mass is formed despite the same α -pinene consumption. (The acidic case is depicted in Figure S3.) Vertical error bars for $\Delta\alpha$ -pinene and Δ isoprene are s.d. for averaging 14-15 samples, horizontal bars show the averaging intervals. For OH we estimate an overall error of $\pm 33\%$.

b: The SOA yields in presence of neutral ammonium sulfate (blue) and acidic ammonium bisulfate (red) seed particles were calculated at times B, C, and, D as *organic mass/ α -pinene consumption* (acidic seed particles, see suppl. section 3). Not only the SOA mass (indicated by the green spheres in the middle) but also its yield decreased strongly (bar plots). B shows the SOA yield for α -pinene alone, while C shows the summed effect of the new product scavenging and OH scavenging in the α -pinene & isoprene reaction system. For the neutral seeds D demonstrates the magnitude of the product scavenging effect on the SOA yield, as [OH] and α -pinene consumption were re-adjusted to the same levels as before isoprene addition. On acidic seed aerosol (red) isoprene makes a substantial contribution to SOA, which compensates for some of the OH scavenging effect and most of the product scavenging effect. The errors reflect the estimated overall 15% uncertainty for an individual experiment.

Figure 2: The reduction of the SOA yield of α -pinene by isoprene as a function of the isoprene/ α -pinene consumption (Δ isoprene/ $\Delta\alpha$ -pinene).

The actual SOA yield of α -pinene ($y_{\text{act}} = \Delta\text{organic mass} / \Delta\alpha\text{-pinene}$) in α -pinene & isoprene mixtures was normalized to the average SOA yield obtained with α -pinene-only (y_{AP}) at same conditions for neutral (blue) and acidic seeds (red). The grey area indicates the $\pm 20\%$ band for the scattering of the α -pinene-only yields around $y_{\text{AP}} = 0.17$ which includes the 15% error in the individual SOA mass yields (vertical bars) and the variability of the experimental conditions for the reaction system with α -pinene-only. Horizontal error bars show the uncertainties of the isoprene/ α -pinene ratio ($\pm 10\%$) derived from the measurement errors of the compounds.

The dashed line is the fit to the results for the neutral seeds that was implemented in the global model calculations to parametrize product scavenging:

$$y_{\text{act}} = y_{\text{AP}} \cdot (0.48 + 0.52 \cdot \exp[-1.53 \cdot (\Delta\text{isoprene}/\Delta\alpha\text{-pinene})])$$

Figure 3: HOM monomer/dimer distribution in the presence and absence of isoprene illustrating the product scavenging effect.

a: comparison of two mass spectra taken in experiments with (orange) and without isoprene (blue). In the presence of isoprene the dimer/monomer drops to about 1/3, indicating that the formation of these extremely low vapour pressure compounds is efficiently suppressed in the presence of isoprene. **b:** Both the total HOM (monomers+dimers, circles) and HOM dimers (squares, molecular mass > 370 Da) decrease with decreasing α -pinene turnover. Here [OH] was varied in two different ways: i) by decreasing the OH production rate, i.e. by decreasing the photolysis of ozone ($J(O^1D)$), blue) and ii) by adding increasing amounts of isoprene (red). In the presence of isoprene HOM decrease more strongly, mainly because the product scavenging is suppressing the α -pinene HOM dimers. The total HOM from the two different experiments were normalized to 1 for better comparison. **c:** Estimate of the product scavenging effect on α -pinene HOM dimers by isoprene at turnover $5.0 \times 10^7 \text{ cm}^{-3} \text{ s}^{-1}$ based on molecular mass weighted signals in the mass spectra (details in Figure S8). Data for α -pinene-only were interpolated from the nearest $J(O^1D)$ data (arrows in **b**). The reduction in total HOM mass is 33%, to which dimers (dark green) contribute 26% and monomers 7% (light green). Error bars show the s.d. of the averages over 31 data points. The suppression of α -pinene HOM dimers of 26% explains most of the 35% product scavenging derived from reduced SOA formation in Figure 1. The balance is a result of the reduced yield of HOM monomers and/or shifts in their volatility spectrum.

Figure 4: Suppression of α -pinene SOA in presence of CO illustrating the generality of the product scavenging effect.

a: comparison of three mass spectra taken in experiments with (orange [CO]=10 ppm, red [CO]=57 ppm) and without CO (blue). In the presence of [CO] > 10 ppm the dimer/monomer drops by more than 1/2, indicating that the formation of the dimers is efficiently suppressed in the presence of CO. Because of the high [CO] overall HOM are suppressed by factor of 4 and 5 respectively. **b:** The product scavenging effect reduces the α -pinene yield in the reaction system α -pinene & CO & OH in the presence of ammonium sulfate seeds. The actual SOA yield, calculated as actual organic mass/actual α -pinene consumption, decreases by more than 50%, when 40 ppm CO was added to the reaction system. CO concentrations were calculated from known inlet concentration, the flow rate through the RC and the reaction rate of CO with OH. Light colored data points show steady state averages, whereby horizontal error bars show the averaging interval; s.d. of the averages are of the same size or smaller than the symbols.

Note, panels **a** and **b** were not taken from the same experiment: experiments to show HOM in the gas-phase and to show SOA suppression cannot be performed in optimal fashion under the same conditions. To measure the unperturbed HOM-MS, particle formation must be avoided. In contrast, mass formation and its suppression must be determined in presence of seed surface as HOM losses at the walls of the chamber have to be minimized. Therefore, the experiment to measure the HOM-MS was performed without seed aerosol whereas ammonium sulfate particles were added to measure the suppression of mass formation by CO. Otherwise the conditions of the experiments were similar.

614
615 **Figure 5:** Atmospheric implications of product scavenging and OH scavenging.
616 Global model calculations showing that OH scavenging (OHS) and product
617 scavenging (PS) in presence of isoprene reduce the photochemically derived MT-OH-
618 SOA mass (red) significantly compared to the reference cases NoIso. For the
619 calculation a 4% iSOA yield for acidic aerosols was applied (sum of the brownish
620 bars). OHS is inherently considered in models as long as biogenic SOA (BSOA)
621 formation is directly linked to the oxidant fields, but the new PS effect reduces the
622 MT-OH-SOA even further. To illustrate the maximum potential PS effect it was also
623 applied to ozone derived MT-O₃-SOA (orange, PSOx). In this case the iSOA gain for
624 1% iSOA yield on neutral aerosols (dark brown) would not be able to compensate for
625 the loss of MT-SOA. We show the global average and results for regions with high
626 and low isoprene emissions and high and low population density. The map shows
627 isoprene emissions, green shaded in [mg m⁻²] and the areas of the selected regions.
628 All contributions are normalised to MT-OH-SOA of the respective NoIso case, when
629 biogenic isoprene emissions were omitted in the model (suppl. section 8).









