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February, 2014

Semantic Calibration of Digital Terrain Analysis Scale

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Available at: https://works.bepress.com/bradley_miller/10/

1 **Semantic calibration of digital terrain analysis scale**

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16 Miller, B.A. Semantic calibration of digital terrain analysis. Cartography and Geographic Information
17 Science Journal 41:166-176. doi: 10.1080/15230406.2014.883488.

18 **Abstract**

19 Digital terrain analysis (DTA) provides efficient, repeatable, and quantified metrics of
20 landscape characteristics that are important to the Earth sciences, particularly for detailed soil
21 mapping applications. However, DTA has not been field tested to the extent that traditional
22 field metrics of topography have been. Human assessment of topography synthesizes multiple
23 parameters at multiple scales to characterize a landscape, based on field experience. In order to
24 capture the analysis scale used by field scientists, this study introduces a method for calibrating
25 the analysis scale of DTA to field assessments. This method is used to calibrate land-surface
26 derivatives of relative elevation, profile curvature, and slope gradient in the context of the
27 commonly used field description of hillslope position. For a topographically diverse landscape in
28 Michigan, USA, a peak in agreement between field assessment and digital terrain analysis was
29 found at field equivalent distances of 135 m for relative elevation, 63 m for profile curvature,
30 and 9 m for slope gradient. Given the field experience of soil scientists, these calibrations of
31 DTA metrics are likely to have stronger correlations with hillslope properties and could be used
32 together to classify hillslope position consistently across large extents.

33 *Keywords:* digital terrain analysis, analysis scale, semantic calibration, land-surface derivatives

34 **Introduction**

35 Digital terrain analysis (DTA) metrics are typically scale dependent (Wood 1996a; Albani
36 et al. 2004; Hupy et al. 2004; Roecker and Thompson 2010). Therefore, when used as
37 parameters in models, attention needs to be given to using the optimal analysis scale for the
38 process being represented. Otherwise, the use of the incorrect analysis scale could result in
39 erroneous model outcomes or mistakenly disregarding important parameters (Claessens et al.
40 2005). Prior to digital techniques, cartographers utilized tacit knowledge to identify the optimal
41 analysis scale for the parameters in their mental model for creating a spatially predictive map.
42 This study calibrates the analysis scale of three land-surface derivatives to the expert
43 knowledge of hillslope position classification.

44 Processes occur at certain phenomenon scales. Analysis scale, on the other hand, is the
45 generalization that is best able to detect that phenomenon (Montello 2001). In DTA terms,
46 analysis scale is the combination of cell resolution and the number of cells incorporated in an
47 analysis neighborhood (Thompson et al. 2001; Albani et al. 2004). Experience has calibrated
48 scientists' sense of analysis scale for measuring landscape characteristics. Similar calibration of
49 analysis scale needs to be derived for digital terrain metrics. This is especially true for expert
50 knowledge-based models and models that seek to predict or be validated by human-made
51 classifications in the field (e.g. digital soil mapping). Until more field studies are conducted to
52 quantitatively determine the scale at which the processes influencing the variation of landscape
53 properties operate, utilizing the scale learned by field scientists provides the best supported

54 method for predicting metric-process relationships. Through semantic calibration, DTA can be
55 parameterized to the tacit knowledge of scientists (Dehn et al. 2001).

56 Despite the importance of analysis scale, analysis neighborhoods are often set to a 3 by
57 3 cell area for all land-surface derivatives, leaving the analysis scale dependent on cell
58 resolution (e.g. Zevenbergen and Thorne 1987; Gallant and Wilson 1996; Lapen and Martz
59 1996; MacMillan et al. 2000; Shi et al. 2009). In some cases, neighborhood size is considered to
60 be such a fundamental assumption that it is not specified in research methods (e.g. Mitášova
61 and Hofierka 1993; Joel et al. 1994; Florinsky et al. 2002). By focusing on cell resolution alone,
62 the analysis scale can be inadvertently determined by the best available resolution and
63 computational efficiency (Moore et al. 1993; Sharma et al. 2011).

64 Many studies have identified the impact of analysis scale via the effect of DEM
65 resolution on geomorphic models (e.g. Chang and Tsai 1991; Chaplot et al. 2000; Schoorl et al.
66 2000; Florinsky and Kuryakova 2000; Thompson et al. 2001; Kienzle 2004; Wu et al. 2008).
67 However, a distinction should be made between varying cell resolution and analysis
68 neighborhood size, because they are different methods of generalization. Also, cell resolution
69 can be associated with elevation accuracy (Gao 1997), especially if the DEMs are generated by
70 different methods. By using algorithms that can calculate land-surface derivatives from
71 neighborhoods larger than 3 by 3 cells, analysis scale can be tested independently of grid
72 resolution (Wood 1996b).

73 In the context of digital soil mapping, several studies have determined that analysis
74 scale affects results. However, the complex interactions between different land-surface

75 derivatives, combinations of soil properties, and value systems inherent within soil classification
76 systems, have clouded interpretation of those results. Smith et al. (2006) varied the analysis
77 scale of slope gradient, plan curvature, and profile curvature as inputs to the SoLIM soil
78 mapping model. Because analysis scale for the different parameters was not allowed to vary
79 independently, it is unclear if the determination of different optimal analysis scales for different
80 terrain types (i.e. 24-36 m for high relief, 33-48 m for gently rolling) reflects different
81 phenomenon scales for different landscapes, or a shift in the dominant parameter. In other
82 words, model performance could have been improved not by the optimal analysis scale of the
83 respective parameters changing between landscapes, but rather by the optimal parameter for
84 prediction changing between soil classes used in different landscapes. Behrens et al. (2010)
85 tested analysis scale for land-surface derivatives independently and found that the optimal
86 analysis scale varied by soil class. The variability of optimal analysis scale between soil classes
87 may be produced by the complexity and values built into the classification system, rather than
88 processes acting at different scales. Therefore, point observations for single attributes need to
89 be examined for determining optimal analysis scale. Roecker and Thompson (2010) did this and
90 concluded analysis scales between 117-189 m to be optimal for correlating point observations
91 of soil carbon, rock fragment content, and clay content at different depths with profile
92 curvature.

93 Recognizing the value of accumulated field experience, expert knowledge models seek
94 to capture tacit knowledge to improve model performance. Hillslope position is a useful, tacit
95 knowledge based, metric for many geomorphic studies; soil scientists have long used this field
96 metric extensively across the United States in their efforts to inventory soil resources. For this

97 reason, I argue that data collected by the U.S. Soil Survey has the large quantity of observation
98 points needed to calibrate DTA with field terrain analysis and to reduce the noise that is
99 inherent in human observations of continuous variables. Therefore, the purpose of this study is
100 to semantically calibrate the land-surface derivatives of slope gradient, profile curvature, and
101 relative elevation to the analysis scale used by soil scientists in determining hillslope position.
102 Because of the categorical nature of hillslope position and high variability of human
103 interpretation, a method of percent agreement is introduced for identifying the analysis scale
104 with the strongest signal. The resulting calibrated metrics can be used as parameters to a
105 variety of models, including fuzzy and discrete classifications.

106 **Methods**

107 *Relating Field Assessment of Hillslope Position to Digital Terrain Analysis*

108 Terrain characteristics for describing hillslope process zones and predicting soil
109 properties has been an important area of study in soil geography (e.g. Milne 1935; Ruhe 1960;
110 Burras and Scholtes 1987; Carter and Ciolkosz 1991; Donald et al. 1993; Stolt et al. 1993; Cerdá
111 1998; Yoo et al. 2005). Although largely qualitative, the use of hillslope position has been tuned
112 through decades of scientific study and field experience. Hillslope position is a contextual
113 metric that divides a landscape into areas where the interaction between hydrology and relief
114 affect the properties of the soil in different ways (Wysocki et al. 2000). Summit positions are
115 often flat and higher in elevation than their surrounding areas, which tends to result in more
116 infiltration than runoff, and less influence from the water table. Shoulder positions are also
117 relatively high in elevation, but their convex shape and steep slope shifts the balance to a

118 greater likelihood of runoff over infiltration. Backslopes are generally considered to be
119 transition zones, and although the slope shape is generally linear, the slope gradient is generally
120 the highest of the hillslope positions, and steeper slope gradients promote runoff. Footslope
121 positions are concave in profile curvature and lower on the slope, causing these sites to be
122 wetter and sediment accumulating positions. The toeslope is flatter and lies at the lowest
123 relative elevation along the slope. Its juxtaposition makes it the zone with the most
124 accumulation of materials transported from upslope and the most affected by the water table.

125 In the field, soil scientists synthesize the terrain characteristics of slope gradient, profile
126 curvature, and relative elevation to identify the functional zones as defined by hillslope position
127 (Figure 1). Although hillslope position is one of the most basic and widely used terrain
128 descriptions for soil geomorphology, it is primarily based on tacit knowledge without
129 quantitative definitions. Soil scientists have calibrated a mental model for identifying soil-
130 landscape patterns. This study focuses on identifying the analysis scales of land-surface
131 derivatives equivalent to the analysis scales used in the soil scientists' mental model.

132 *Semantic Calibration of Analysis Scale*

133 DTA was performed at multiple analysis scales for comparison with soil scientists' field
134 assessments. Slope gradient and profile curvature were calculated with varying neighborhood
135 sizes using the `r.param.scale` function in GRASS 6.4.2 (GRASS Development Team 2012). The
136 `r.param.scale` function calculates both slope and profile curvature by fitting a quadratic trend
137 surface using least squares (Evans 1979). The analysis scale can be expanded by including more

138 grid cells in the parameters for the polynomial and solving the least squares via a matrix (Wood
139 1996b).

140 In order to have a user controlled analysis scale for relative elevation, I developed a
141 procedure in ArcGIS 10.1 (ESRI 2012). This method for calculating relative elevation subtracted
142 the inverse elevation from the original elevation by analysis neighborhoods. The inverse
143 elevation was the elevation subtracted from the sum of the minimum and maximum elevation
144 values in the analysis neighborhood. The neighborhood size was controlled via the focal
145 statistics used to determine the minimum and maximum elevation. The resulting relative
146 elevation grid had increasing positive values above and decreasing negative values below the
147 analysis neighborhood's middle elevation (Figure 2).

148 A LiDAR-derived, 3 m resolution, elevation grid was aggregated to resolutions of 9 m
149 and 27 m to reflect the resolutions of 1/3 arc second and 1 arc second that other elevation grid
150 products are commonly available in, while still preserving cell alignment. Then, all three grids
151 were processed for the three land-surface derivatives using varied neighborhood sizes by the
152 experimental matrix in Table 1. The experimental matrix was designed to cover the full range of
153 reasonable analysis scales and to use combinations of cell resolution and neighborhood size
154 where analysis scale would align across cell resolutions.

155

156 Table 1. Experimental matrix for varying combinations of DEM cell size and neighborhood size
 157 for varying analysis scale.

	Neighborhood Size							
Distance	9 m	15 m	27 m	45 m	63 m	81 m	135 m	189 m
LiDAR (3 m)	3x3	5x5	9x9	15x15	21x21	27x27	45x45 [†]	63x63 [†]
(9 m resample)			3x3	5x5	7x7	9x9	15x15	21x21
(27 m resample)						3x3	5x5	7x7

158 [†] only used for relative elevation

159

160 Field observation points of hillslope position and slope gradient, collected by U.S.
 161 Natural Resource Conservation Service (NRCS) soil scientists, were then intersected with each
 162 of the DTA grids. The resulting match-up allowed for the comparison of DTA calculations with
 163 soil scientists' assessment in the field by location. The soil scientists recorded slope gradient as
 164 an integer, allowing for a quantitative comparison. Agreement for slope gradient was evaluated
 165 by the mean absolute difference between the field observed and the DTA calculated slope
 166 gradient in degrees. Because profile curvature and relative elevation are included in hillslope
 167 position as categorical attributes, not quantitative measures, comparison between field
 168 observation and DTA were compared on a basis of percent agreement by categorical definition
 169 (Table 2).

170 Table 2. Definition table for relating qualitative hillslope position descriptions to the DTA of
 171 profile curvature and relative elevation. Attributes highlighted in gray were used for the
 172 semantic calibration.

Hillslope Position	Profile Curvature	Relative Elevation
Summit	Linear	High
Shoulder	Convex	High
Backslope	Linear	Middle
Footslope	Concave	Low
Toeslope	Linear	Low

173 Due to the fuzzy nature of landscape elements and the subjectivity of human
174 assessment, agreement was only tested for the definitional extremes. For profile curvature,
175 agreement was evaluated for negative values corresponding to concave slope shape and
176 positive values corresponding to convex slope shapes. Relative elevation was quantified as
177 distance above (positive) or below (negative) the mid-elevation of the analysis neighborhood.
178 Agreement was evaluated as negative values corresponding with low hillslope positions and
179 positive values corresponding with high hillslope positions.

180 The level of agreement between digital and field assessment should increase as the
181 digital analysis scale is more closely aligned with the analysis scale used by the soil scientists in
182 the field. Therefore, the optimal semantic calibration was determined to be the analysis scale
183 with the highest percent agreement for the qualitative metrics. Specifically, for profile
184 curvature and relative elevation, percent agreement was calculated for each definitional
185 category separately and then summarized with the mean of those results. Using the mean of
186 the categorical results avoided over emphasizing a category that may have more observations
187 than the other. In other words, the mean percent agreement for a particular analysis scale
188 equally weights the two categories used for evaluation. For slope gradient, agreement could be
189 measured on a continuous scale. Therefore, the optimal semantic calibration for slope gradient
190 was determined by the lowest mean absolute difference between the DTA calculation and the
191 field estimate.

192 Consistent analysis scale calibration between grid resolutions provided support for the
193 determination that calibration was due to the matching of analysis scales and not to matching

194 an artifact pattern in the source DEM. Therefore, a consistent peak in percent agreement at the
195 same analysis scales was considered to be a strong signal for identifying the analysis scale used
196 by the soil scientists and that analysis scale being used consistently.

197 *Study Area*

198 This study analyzed data for Ottawa County, Michigan, on the eastern shore of Lake
199 Michigan (Figure 3). Large dunes have formed on the western edge of Ottawa County (Figure
200 4a). The central part of the county is lake plain formed beneath Glacial Lake Chicago (Figure 4b).
201 This glacial lake plain is flat with interspersed dunes of decreasing size (from west to east) that
202 have encroached from the west. The northeast and southeast portions of the area grade into a
203 hummocky terrain, more characteristic of till plains (Figure 4c). Across the entire 1,488 km²
204 study area the elevation ranges from 173 to 292 meters above sea level.

205 Ottawa County was chosen because of the availability of a LiDAR-based elevation grid
206 (2004) and a number of georeferenced observations of hillslope position, by NRCS soil
207 scientists, taken in the field. The LiDAR-based elevation grid has a 3 m resolution and was
208 provided by the Ottawa County government. Georeferenced field observations of hillslope
209 position were provided by the Grand Rapids, Michigan Soil Survey Office.

210 Recorded GPS points at observation sites were joined with a database created from the
211 paper records of field observations. Observations were made by the Grand Rapids Soil Survey
212 Office staff between 2007 and 2009. The GPS recorded observations used in this study were all
213 8-10 point transects, of typically 40-80 m point spacing, for a total of 1,068 points (Figure 3).

214 The hillslope position assessment of the soil scientists was converted into categorical attributes
215 in terms of profile curvature and relative elevation (Table 2), as interpreted from the hillslope
216 position diagram presented in Schoeneberger et al. (2012). Not all recorded observations were
217 interpretable for these two land-surface derivatives. For example, observation points that were
218 recorded simply as “flat,” could not have a relative elevation interpreted from it. From the
219 transect observation points, 966 were interpretable for profile curvature and 572 were
220 interpretable for relative elevation. Field observed values of slope gradient were available for
221 all 1,068 points. Using only points in the categories that were least likely to overlap reduced the
222 number of field observations that could be used in the calibration. Therefore, 216 points were
223 used for the evaluation of profile curvature and 406 points were used for the evaluation of
224 relative elevation.

225

226

227 **Results**

228 *Analysis Scale Calibration*

229 The calibration of all three land-surface derivatives showed a pattern of increasing
230 agreement up to a certain analysis scale and then decreasing agreement with scales coarser
231 than the optimal matching analysis scale. In describing the result details, the different
232 combinations of grid resolution and neighborhood size will be referred to by the field distance
233 equivalent for the particular analysis scale. Because all analysis scales are based on square
234 neighborhoods of square cells, it is simpler and more relatable to the field environment to use
235 the field equivalent distance for the linear dimension of the analysis scale. However, it should
236 be noted that field observations of hillslope position are likely more linear (aligned with the
237 profile of the hillslope) than the square area used by DTA.

238 Profile curvature had the highest agreement between digital calculations and field
239 observations at a field equivalent distance of 63 m (Table 3). Because the minimum possible
240 analysis scale for a grid resolution of 27 m is 81 m ($3 \text{ cells} * 27 \text{ m} = 81 \text{ m}$), the same signal could
241 not be observed with the 27 m grid. Instead, the highest percent agreement was at the smallest
242 analysis scale, which is closest to the scale signal observed in the other two resolutions. An
243 exception to this signal was the concave profile curvature agreement for the 9 and 27 m
244 resolution grids, where the highest percent agreement was at a scale of 135 m. However, for
245 the 9 m resolution grid, the percent agreement at 81 m and 63 m was only 3% less than at 135
246 m. These increases in analysis scale for the highest percent agreement could be caused by the
247 scale effect of the modifiable area unit problem (MAUP), which increases the probability of
248 categorical agreement by decreasing the variability at larger analysis scales. The percent

249 agreement peak for the combined concave and convex categories still showed a strong signal at
 250 63 m, or the closest possible analysis scale.

251

252 Table 3. Percent agreement results for profile curvature. Concave and convex 'agree' are the
 253 count of points where the digital and field assessments of slope profile shape were in
 254 agreement. Percent is the proportion of agreement out of the total number of points identified
 255 by the soil scientists to be in the respective category (105 concave and 111 convex points).[†]

Profile Curvature	3 m Grid					
	9 m	15 m	27 m	45 m	63 m	81 m
Concave agree	60	59	72	72	73	70
Convex agree	67	72	75	77	77	76
% concave agree	57.1%	56.2%	68.6%	68.6%	69.5%	66.7%
% convex agree	60.4%	64.9%	67.6%	69.4%	69.4%	68.5%
Mean % agreement	58.8%	60.5%	68.1%	69.0%	69.4%	67.6%

	9 m Grid						
	Distance	27 m	45 m	63 m	81 m	135 m	189 m
Concave agree		64	65	71	71	74	65
Convex agree		72	75	76	73	68	74
% concave agree		61.0%	61.9%	67.6%	67.6%	70.5%	61.9%
% convex agree		64.9%	67.6%	68.5%	65.8%	61.3%	66.7%
Mean % agreement		62.9%	64.7%	68.0%	66.7%	65.9%	64.3%

	27 m Grid			
	Distance	81 m	135 m	189 m
Concave agree		68	75	68
Convex agree		77	69	75
% concave agree		64.8%	71.4%	64.8%
% convex agree		69.4%	62.2%	67.6%
Mean % agreement		67.1%	66.8%	66.2%

256 [†]Analysis scale with strongest signal highlighted in dark gray; other strong agreements shown
 257 in lighter gray.

258

259 The signal for relative elevation was at 135 m (Table 4). To confirm this for the 3 m
260 resolution grid, the experimental matrix needed to be extended beyond a field equivalent
261 distance of 81 m. The signal for an optimal analysis scale at 135 m was consistent for nearly all
262 measures. The only exception was the percent agreement for relatively high elevation using the
263 27 m resolution grid. In that case, one more point was in agreement for the 81 m over the 135
264 m scale, increasing the amount of agreement by 0.5%. The analysis scale for the relative
265 elevation signal indicates relative elevation is considered contextually over a larger area in the
266 field, by mappers, as compared to the other terrain metrics.

267 The numeric field observations of slope gradient provided a quantitative determination
268 of matching analysis scales. Although it is unrealistic to expect integers recorded by human
269 observation to exactly match the rational numbers calculated by DTA, a minimal difference
270 provides a calibration of the digital method to the analysis scale used in the field. For each grid
271 resolution, the mean difference was lowest at the finest analysis scale possible (Table 5). The
272 trend of decreasing mean differences with decreasing neighborhood size suggests the optimal
273 analysis scale for digitally determining slope gradient similar to a soil scientist's characterization
274 in the field is relatively small.

275

276 Table 4. Percent agreement results for relative elevation. Low and high 'agree' are the count of
 277 points where the digital and field assessments of relative elevation were in agreement. Percent
 278 is the proportion of agreement out of the total number of points identified by the soil scientists
 279 to be in the respective category (206 low and 200 high points). †

Relative Elevation	3 m Grid								
	9 m	15 m	27 m	45 m	63 m	81 m	135 m	189 m	
Low agree	101	115	135	148	158	163	176	173	
High agree	109	117	123	136	137	140	144	140	
% low agree	49.0%	55.8%	65.5%	71.8%	76.7%	79.1%	85.4%	84.0%	
% high agree	54.5%	58.5%	61.5%	68.0%	68.5%	70.0%	72.0%	70.0%	
Mean % agreement	51.8%	57.2%	63.5%	69.9%	72.6%	74.6%	78.7%	77.0%	
				9 m Grid					
			27 m	45 m	63 m	81 m	135 m	189 m	
Low agree			125	152	159	164	177	172	
High agree			129	133	139	139	148	132	
% low agree			60.7%	73.8%	77.2%	79.6%	85.9%	83.5%	
% high agree			64.5%	66.5%	69.5%	69.5%	74.0%	66.0%	
Mean % agreement			62.6%	70.1%	73.3%	74.6%	80.0%	74.7%	
						27 m Grid			
					81 m	135 m	189 m		
Low agree					153	167	166		
High agree					145	144	138		
% low agree					74.3%	81.1%	80.6%		
% high agree					72.5%	72.0%	69.0%		
Mean % agreement					73.4%	76.5%	74.8%		

280 † Analysis scale with strongest signal highlighted in dark gray; other strong agreements shown
 281 in lighter gray.

282

283

284 Table 5. Mean absolute difference results for slope gradient.

<i>Slope (degrees)</i>	3 m Grid						
	Distance	9 m	15 m	27 m	45 m	63 m	81 m
Mean difference		3.1°	3.2°	3.4°	3.8°	4.0°	4.1°
		9 m Grid					
	Distance	27 m	45 m	63 m	81 m	135 m	189 m
	Mean difference	3.4°	3.8°	4.0°	4.1°	4.4°	4.5°
		27 m Grid					
	Distance	81 m	135 m	189 m			
	Mean difference	4.1°	4.4°	4.5°			

285

286 **Discussion**

287 The observed convergence in agreement as neighborhood size approaches an optimal
 288 analysis scale suggests 1) soil surveyors are relatively consistent in their use of analysis scale
 289 and 2) the calibration methodology was able to capture the analysis scales used by the soil
 290 scientists. By comparing a series of analysis scales for land-surface derivatives with the soil
 291 scientists’ mental model assessments, the analysis scales of land-surface derivatives that best
 292 correspond to the soil scientists’ field experience were identified. In the context of soil
 293 scientists describing hillslopes, the analysis scale used decreases from relative elevation to
 294 profile curvature to slope gradient. To many field scientists, this hierarchy of analysis scale for
 295 these components may sound intuitive, but DTA parameters are often calculated at identical
 296 analysis scales for use in environmental models. The field scientists’ use of different analysis
 297 scales for these parameters should inform quantitative modelling and thereby improve
 298 geomorphic and soil predictions.

309 Comparisons of DTA scales demonstrated the impact of analysis scale choice for each
300 land-surface derivative, particularly for capturing expert knowledge. For example, when
301 characterizing relative elevation with a 3 m resolution grid, choosing between a field equivalent
302 distance of 9 m versus 135 m had the impact of agreeing with the field scientists 52% of the
303 time or 79% of the time, respectively. Similarly, when simply characterizing profile curvature as
304 concave or convex, choosing an analysis scale of 63 m over 9 m improved agreement with
305 scientists in the field by 10%.

306 Although the method introduced would need to be applied in additional study areas to
307 determine transferability, the results of this study are supported by the observations of other
308 researchers investigating the optimal analysis scale for these land-surface derivatives in other
309 soil related contexts. The semantic calibration of profile curvature scale in this study
310 corresponds to the optimal scale Drăgut et al. (2009) determined for predicting crop yield from
311 profile curvature on an alluvial plain of the Danube River. The increasing agreement between
312 digital and field measurements of slope gradient with decreasing neighborhood size to at least
313 9 m is also consistent with the results of Shi et al. (2007). The results of Roecker and Thompson
314 (2010) indicated a coarser analysis scale for profile curvature, but the smaller quantity of
315 samples in their study may have been susceptible to issues of sampling and the scale effect of
316 MAUP.

317 The use of only categories that are better separated definitionally was successful in
318 reducing the noise in the soil scientists' mental model outcomes. The high variability of human
319 assessment greatly reduces the predictability of qualitative field metrics such as hillslope

320 position, but does not negate the ability to capture the logic in the experts' knowledge. At a
321 minimum, the DTA should be expected to produce values definitionally compatible with the
322 field observations. That is, profile curvature should be positive for convex slope shapes and
323 negative for concave slope shapes (note, some profile curvature algorithms may have those
324 definitions switched). Similarly, it would be expected that summits and shoulders would have
325 positive relative elevations with footslopes and toeslopes expected to have negative relative
326 elevations. The calibrated analysis scale at which the DTA has the highest agreement with these
327 definitional relationships provides insight to how the soil scientists conceptualized the
328 landscape in the field.

329 The similar rates in calibration agreement for matching analysis scales across grid
330 resolutions indicate that resampling to larger cell sizes did not have a major impact on the
331 analysis. Using more cells in the analysis neighborhood or aggregating to larger cell sizes and
332 then using fewer cells in the analysis neighborhood are two different routes for generalization,
333 with the potential for different outcomes. However, the effect of using one or the other
334 method of generalization did not affect the results as much as changing the realized analysis
335 scale.

336 As high resolution, digital elevation products from technologies such as LiDAR become
337 more available, smaller objects on the ground have the potential to affect DTA results.
338 Although the elevation grid used in this study was processed from the LiDAR point cloud to be
339 bare earth, the influence of either non-soil or at least man-made features are present in the
340 elevation data. This noise in the DEM could affect land-surface derivatives. Therefore, the

341 semantic calibration in this study may be influenced by not only the scale used by soil scientists,
342 but also the DTA scale that sufficiently smoothes the digital representation of terrain to
343 minimize the influence of features that would naturally be ignored by scientists in the field.

344 **Conclusions**

345 Based on data from the soil scientists in this study, this research suggests that there are
346 optimal analysis scales for relating DTA to observations made in the field. However, the optimal
347 analysis scale is likely to be different for each land-surface derivative. The method of semantic
348 calibration presented was able to identify analysis scales that optimized the agreement
349 between DTA and field observations.

350 Results from this study suggest that the optimal analysis scales for slope gradient,
351 profile curvature, and relative elevation are the field equivalent distances of 9 m, 63 m, and 135
352 m, respectively. As analysis scales tested were limited to multiples of the grid cell size, these
353 calibrated analysis scales are approximate. Application of the method introduced in this study
354 in additional areas is needed to determine the transferability of these calibrated analysis scales,
355 but similarity of these results with other studies suggests the possibility that they are not
356 unique to this study area.

357 This method of semantically calibrating analysis scales for DTA provides documentation
358 of the soil scientists' field perspective and experience in the assessment of hillslope processes.
359 Models attempting to predict landscape features traditionally identified by field scientists will
360 likely benefit from using the same analysis scales utilized by those scientists. Similarly, other

361 models using DTA as parameters will likely benefit from the lessons of analysis scale
362 accumulated by scientists in the field over centuries.

363 **Acknowledgements**

364 I am grateful to the County of Ottawa, Michigan and the NRCS Soil Survey Office in Grand
365 Rapids, Michigan for providing the data needed for this study. I thank Ashton Shortridge,
366 Randall Schaetzl, David Lusch, and Sasha Kravchenko for their advice on previous drafts. I also
367 thank the anonymous reviewers and Dr. Michael Leitner, editor of CaGIS, for their valuable
368 comments and suggestions. Support was provided by the Graduate College and the Department
369 of Geography at Michigan State University, the Soil Classifier Association of Michigan, as well as
370 by the Association of American Geographers.

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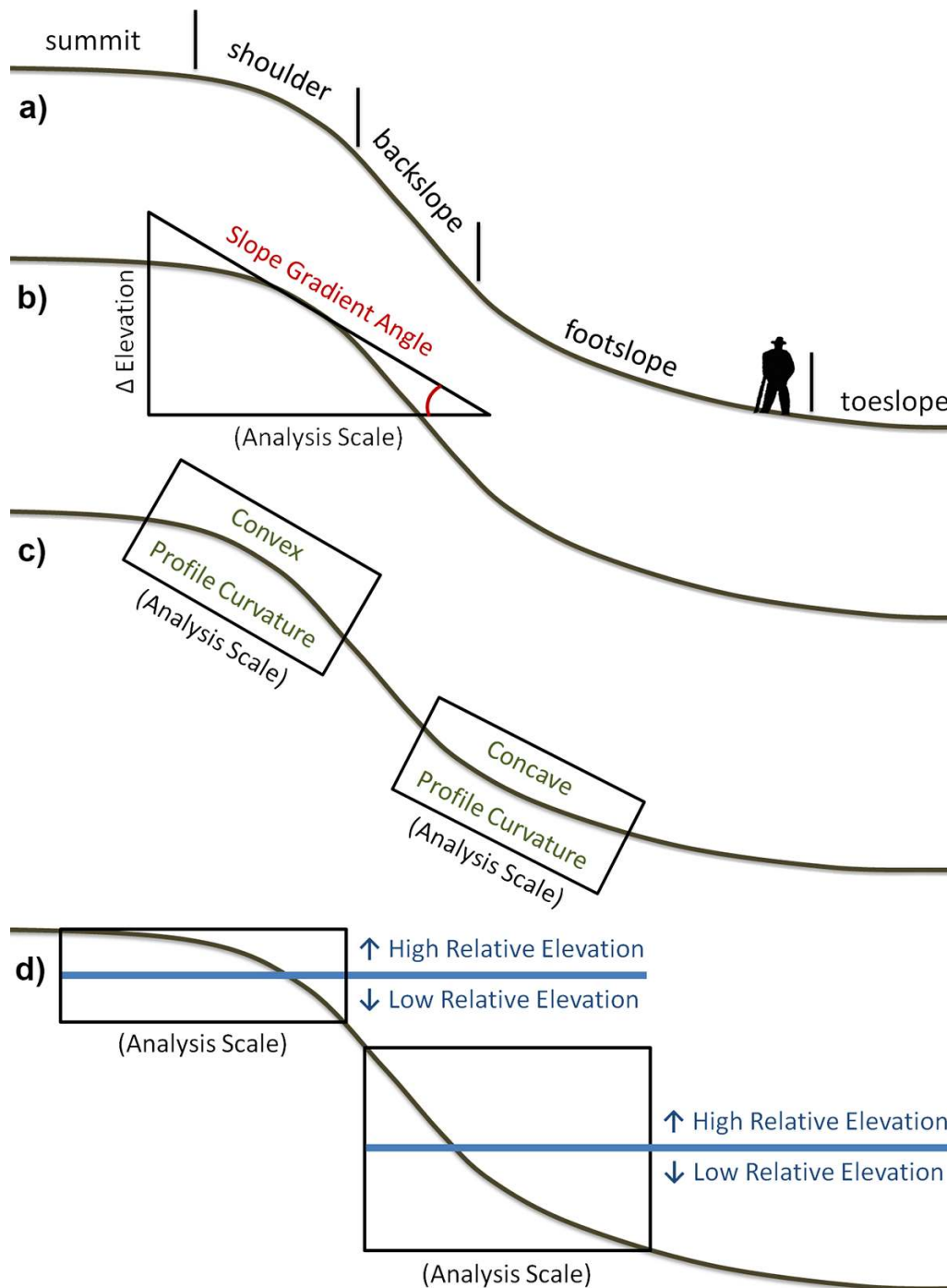
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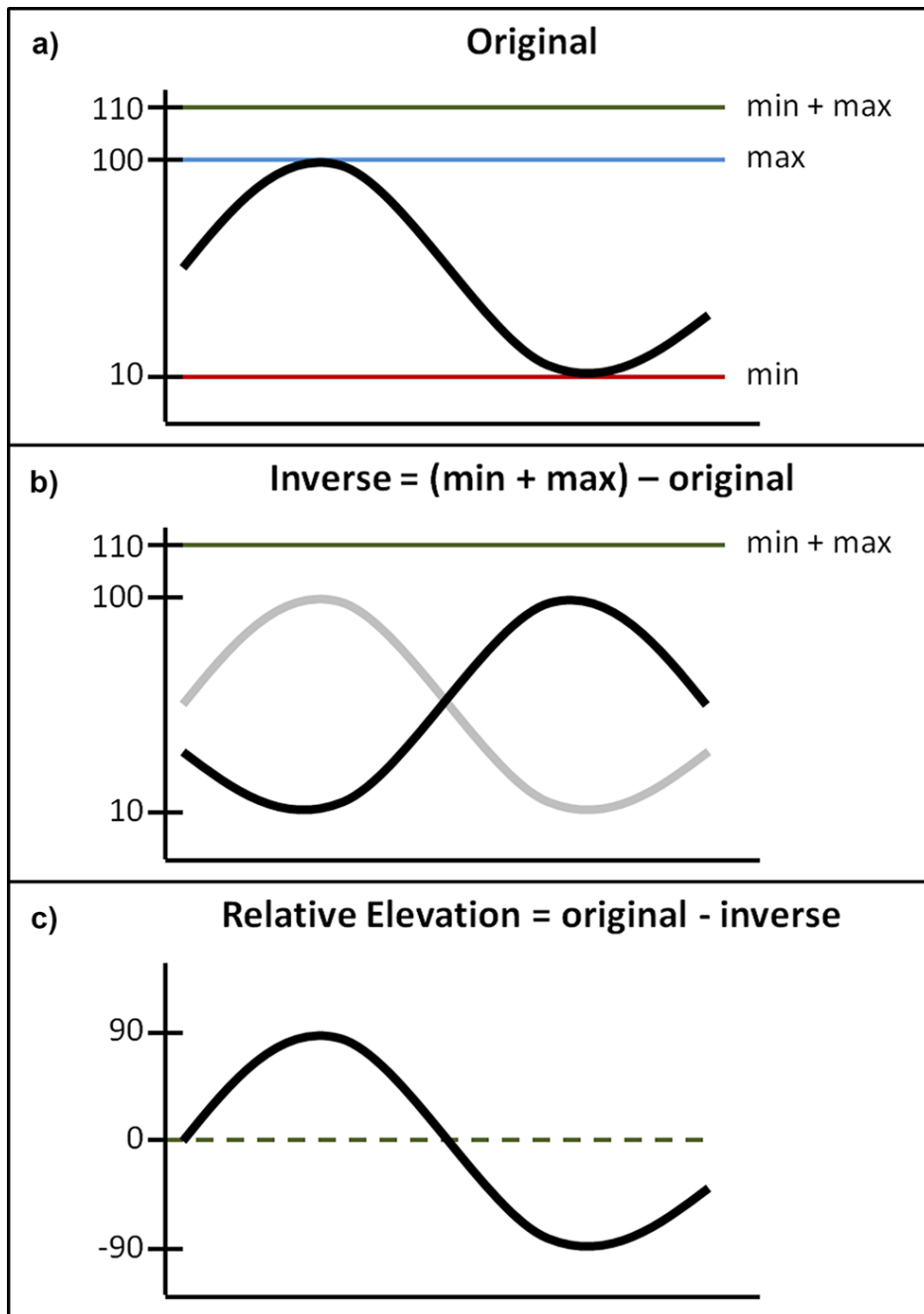
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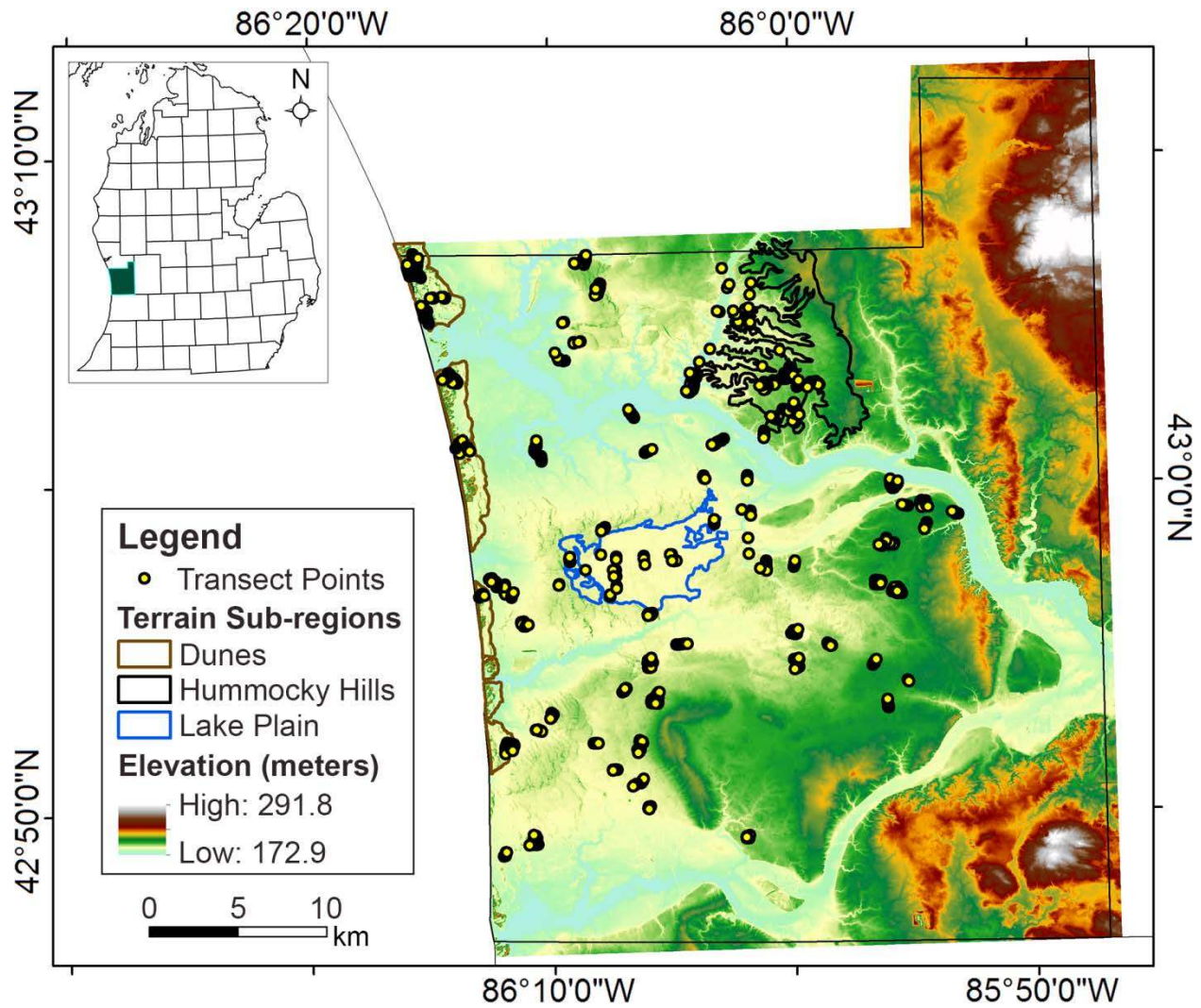
482 Figures



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484 Figure 1. In order to classify the functional zones of hillslope position, a) soil scientists in the
485 field synthesize their assessment of b) slope gradient, c) profile curvature, and d) relative
486 elevation to determine the hillslope position of a location.



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 488 Figure 2. In this research, relative elevation is calculated on a neighborhood by neighborhood
 489 basis for each cell. a) First, a reference elevation ceiling is calculated by summing the
 490 neighborhood minimum and maximum elevations. b) Then, the central cell elevation is
 491 subtracted from the elevation ceiling to calculate an inverse elevation. c) By subtracting the
 492 inverse from the original elevation, a relative elevation grid is created with the mid-point
 493 between the neighborhood minimum and maximum having a value of zero. Elevation values
 494 above the mid-point are increasingly positive. Below the elevation mid-point, values are
 495 negative and decrease with vertical distance.



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Figure 3. Relief and location of study area, Ottawa County, in the lower peninsula of Michigan, U.S. The map also includes locations of field observation transect points.



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Figure 4. Examples of terrains within study area: a) dune landscape, b) glacial lake plain landscape, and c) hummocky landscape. Photos by the author.