

STUDIES OF THE MASS BUDGET OF ARCTIC PACK-ICE FLOES*

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ABSTRACT. The processes of melting, freezing, precipitation, evaporation and condensation were considered in the computation of the mass budget of an ice floe in the northern Chukchi Sea. It was found that condensation and evaporation contribute negligible amounts, and that melting accounted for the largest loss from the upper surface of the floe. From an original mass of 270 g. cm.^{-2} , a unit ice column lost 9.2 g. cm.^{-2} by melting of snow and 35.2 g. cm.^{-2} by melting of ice during the melt season. During the freezing period 5.8 g. cm.^{-2} were regained by the accumulation of snow and frost. Measurements at the bottom of the ice indicate an estimated net loss of 5 g. cm.^{-2} over a seven-month study. A comparison of six summers at three stations shows no obvious correlation of ablation and latitude, and an average loss of ice at the surface of 40 g. cm.^{-2} .

RÉSUMÉ. *Études du bilan de masse de floes du pack arctique.* Dans l'établissement du bilan de masse d'un floe de la Mer Chukchi septentrionale, on a considéré les processus de fusion, gel, précipitation, évaporation et condensation. Il s'est avéré que les influences dues à la condensation et évaporation étaient négligeables, et que la plus grande perte était due à la fonte de la surface du floe. A partir d'une masse de 270 g cm^{-2} , un cylindre unitaire de glace perdit $9,2 \text{ g cm}^{-2}$ par fonte de neige et $35,2 \text{ g cm}^{-2}$ par fonte de glace pendant la saison de fonte. Pendant la saison froide, $5,8 \text{ g cm}^{-2}$ s'accumulèrent par chute de neige et gel. Des mesures au bas du floe montrent une ablation nette d'environ 5 g cm^{-2} pour la période d'observation de sept mois. Les observations effectuées pendant six périodes estivales en trois stations ne montrent aucune corrélation entre l'ablation et la latitude, et une perte moyenne de glace à la surface de 40 g cm^{-2} .

ZUSAMMENFASSUNG. *Untersuchung des Massenhaushaltes arktischer Packeis-schollen.* Unter Berücksichtigung beobachteter Werte von Ablation, Eisneubildung, Niederschlag, Verdunstung und Kondensation wird der Massenhaushalt einer Eisscholle in der nördlichen Tschuktschen-See berechnet. Es zeigt sich, dass Verdunstung und Kondensation nur unerhebliche Beiträge liefern und dass der grösste Massenverlust durch Abschmelzung an der Oberfläche eintritt. Sie reduziert die Ausgangsmasse einer Einheitsssäule von 270 g cm^{-2} um $9,2 \text{ g cm}^{-2}$ durch Schnee-Ablation und um weitere $35,2 \text{ g cm}^{-2}$ durch Eis-Ablation in der Abschmelzperiode. In der Frostperiode beträgt die Akkumulation durch Schnee und Rauhreif $5,8 \text{ g cm}^{-2}$. Messungen an der Unterseite der Eisscholle ergaben einen Nettoverlust von schätzungsweise 5 g cm^{-2} während eines 7-monatigen Beobachtungszeitraumes. Ein Vergleich von sechs Ablationsperioden an drei Stationen zeigt keine eindeutige Breitenabhängigkeit. Der durchschnittliche Eisverlust an der Oberfläche beträgt 40 g cm^{-2} .

INTRODUCTION

The present report is mainly a study of the mass budget of U.S. drifting station "Charlie". Data from the preceding station "Alpha" and the succeeding station ARLIS II (Arctic Research Laboratory Ice Station) are used for comparisons.

Station "Alpha", which was used during the I.G.Y., was set up on drifting pack ice during April 1957 and abandoned in November 1958. During the 19 months of occupation the station drifted in the area bounded by lat. 79° to 86.5° N. and long. 110° to 176.5° W.

Station "Charlie" (also known as "Alpha II") was established on drifting pack ice during April 1959 and was withdrawn during January 1960. During this period it drifted in an area bounded by lat. 76° to 78.2° N. and long. 159.5° to 174.4° W.

ARLIS II is on an ice island. It was occupied in May 1961 at lat. $76^\circ \text{ N.},$ long. 155° W. and had, up to May 1964, drifted with the Beaufort Sea gyral to lat. $86^\circ \text{ N.},$ long. 28° W. Continuous studies of the sea ice adjacent to the island have been hampered by contamination and frequent ice movement. Only during the summer of 1963, a small floe of clean pack ice 2.5 m. thick was accessible for measurements of ablation.

The budget described here pertains to the mass of a column of unit cross-sectional area and a length equal to the mean thickness of the floe. By assuming densities (Untersteiner, 1961) of 0.9 g. cm.^{-3} for the ice, errors smaller than those from other sources are introduced.

The mass changes considered in the computation of the mass budget at station "Charlie"

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were melting, freezing, precipitation, evaporation and condensation. Internal melting was considered to have no net effect.

METHODS

The mass of the snow cover was measured just before the onset of the melt season by more than four hundred measurements of the snow depth.

The amount of ice melt at the upper surface of the floe was determined from several sets of white wooden ablation stakes set in the ice approximately 1 m. deep at the time the snow disappeared. Pond stakes were set in the lower areas where melt water had saturated the snow.

The accretion or ablation of the lower surface of the floe was determined from thickness measurements made by drill holes and by thickness gauges devised by Untersteiner and Badgley (1958) (Fig. 1).

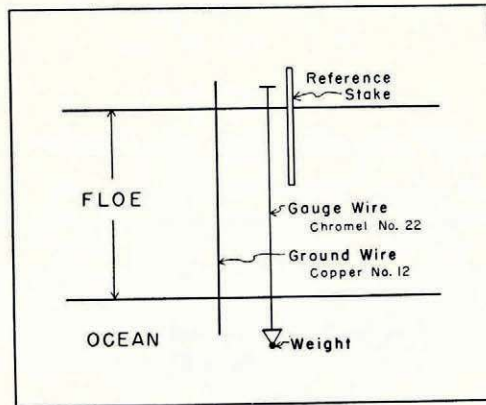


Fig. 1. Thickness gauge to measure ice bottom accretion or ablation

Observations of evaporation, condensation and/or sublimation were carried out by weighing two clear "Plexiglass" pans which were filled with wet snow or disintegrated ice and placed with the pan tops even with the surface. The pans had an area of 500 cm.², were 1.5 and 2.0 cm. deep and had walls approximately 2 mm. thick. (Pans of different depths showed no systematic difference in the mass changes of their contents.)

During the melt season (June to mid-August) measurements of the ablation stakes and weighing of the evaporation pans were made twice daily, which also permitted calculation of diurnal changes. After the termination of the melt season the stakes and the thickness gauges were measured several times each month.

A water level recorder which can be used to determine changes in ice thickness was in operation at station "Charlie". A description of the use of this instrument for such purposes has been given by Untersteiner and Badgley (1958). The rise of the floe of 37.5 cm. observed between 21 June and 31 August corresponds to a surface ablation of 41.7 cm., in almost perfect agreement with the stake measurements.

RESULTS

Upper surface

The measurements of snow depth preceding the melt season ranged from 0 to 122 cm.; the average was 28 cm. With a density of 0.33 g. cm.⁻², as found at station "Alpha"

(Untersteiner, 1961), this amounts to 9.2 g. cm.^{-2} . At the end of the melt season the amount of the winter snow cover remaining on the floe was estimated to be less than 1 per cent.

Solid precipitation (snow and frost) during the melt season amounted to 2.82 cm. when melted, all of which for purposes of mass budget calculations was assumed to have melted. Table I gives monthly totals of precipitation measurements by U.S. Weather Bureau personnel on station "Charlie". From the end of the melt season to the end of the year snow and frost accumulation amounted to 5.8 g. (Table II).

TABLE I. STATION "CHARLIE", MONTHLY PRECIPITATION (cm. of water)

Type	June	July ¹⁹⁵⁹	August	September	Total June-September
Rain, drizzle	0.15	0.18	0.43	0.13	0.89
Snow	0.08	1.93	trace	0.81	2.82
Total precipitation	0.23	2.11	0.43	0.94	3.71

TABLE II. HEIGHTS OF SNOW AND FROST DEPOSIT (cm.)

Period 1959	Increment	Total
15-31 August	0.7	0.7
25 September	2.1	2.8
28 October	6.0	8.8
27 November	5.3	14.1
31 December	3.6	17.7

Net accumulation (density 0.33 g. cm.^{-3}) = 5.8 g. cm.^{-2} .

Snow melt and formation of ponds began in June. The ablation values during 5-day periods in non-ponded areas as computed from two sets of ablation stakes are given in Table III. A "weighted total" is given because stake set II incompletely spans the melt season.

TABLE III. 5-DAILY ICE ABLATION OF NON-PONDED AREAS (cm.),
20 JUNE-20 AUGUST 1959

Period ending 1959	Set I	Number of stakes	Set II	Number of stakes
26 June	6.9	4 to 6*	—	0
1 July	6.6	6	—	0
6 July	1.1†	6	—	0
11 July	1.3	6	—	0
16 July	2.5†	6	—	0
21 July	3.3	6 to 11*	—	0
26 July	4.1	11	3.0 (3 days)	10
31 July	1.4†	11	2.7	10
5 August	4.7	11	5.4	10
10 August	3.7	11	3.8	10
15 August	2.5†	8	2.7	10
20 August	0.0	8	0.0	10
Total	38.1		17.6	
	Weighted total		39.2 cm.	
	Mass, using 0.9 g. cm.^{-3}		35.3 g. cm.^{-2}	

* Number of stakes increased.

† Snow or frost during interval.

TABLE IV. POND ABLATION (cm.), MINIMUM VALUES

<i>Period</i>	<i>Pond 1*</i>	<i>Pond 2†</i>	<i>Pond 3</i>
1959			
25 June–18 July	41.8	33.4	39.7
18 July–11 August	—	28.0	46.6
11 August–10 September	—	—	11.9
Total melt	—	61.4	98.2
Pond depth at end of season		37.1	58.9
Net mass lost using 0.9 g. cm. ⁻³		18.2 g. cm. ⁻²	29.5 g. cm. ⁻²
Weighted net loss for all ponds (at least)		28 g. cm. ⁻² .	

* Pond 1 measurements were discontinued after it filled with drifting snow.

† Pond 2 measurements were discontinued after the surface froze.

The “weighted net loss for all ponds” given in Table IV was computed by using the three intervals shown in Table IV. Pond 2 was about twice the area of pond 1 but only one-seventh the area of pond 3, and therefore the “net loss” included areal weighting. The difference in the amount of melt in any interval between ponds was due in the main part to differences in pond depth and ice type and to a lesser part to errors of measurement.

Diurnal differences in melt rate were observed. The data analysed by half-day periods is shown in Table V. The pond melting rate was smoother by almost a factor of two.

TABLE V. MEAN DIURNAL ABLATION RATES (mm.)

<i>Period</i>	<i>Location</i>	<i>Day</i>	<i>Night</i>	<i>24-hr. rate</i>
1959				
25 June–23 July	3 ponds	9.8	5.0	14.8
22 June–14 August	Stake set I	7.8	2.0	9.8
23 July–14 August	Stake set II	7.8	1.7	9.5

Lower surface

Ice thickness measurements were made at two sites by drilling; these were helpful not only to determine changes of thickness but also for the study of “false bottoms” and under-ice ponds. A “false bottom” is a thin layer of ice which forms in summer underneath the floe where fresh water lies between the salt water and the ice; this is described by Untersteiner and Badgley (1958). Measurements of the thickness of a “false bottom” are accurate to 1 cm.

The results obtained by repeated drilling in a small pond with natural drainage are shown in Figure 2a. The formation of two fresh-water under-ice ponds indicates that there must have been two periods of pond drainage in this area prior to the initiation of measurements. The floe continued to thin after the pond froze over, because of the supply of heat from the fresh-water pond and from short-wave radiation passing through the clear pond ice and water.

At another site near an artificial drainage hole the under-ice pond had vanished between 19 and 30 August (Fig. 2b). All drilling at this site was near an ablation stake which served as a datum plane. The step in the false bottom which occurred between 19 and 21 July was probably caused by the escape of some of the under-floe pond water.

Bottom ablation and accumulation measurements by thickness gauges were initiated in September (Table VI); three of these gauges had been installed through under-ice ponds. Ablation continued at the “false bottom” of the floe well into the winter season (shown by gauges 2 and 3 in December). The large variation of the net thickness change was caused by differences in ice thickness, snow depth and by the presence or absence of surface and under-ice ponds, as well as by the variable amount of water within the floe.

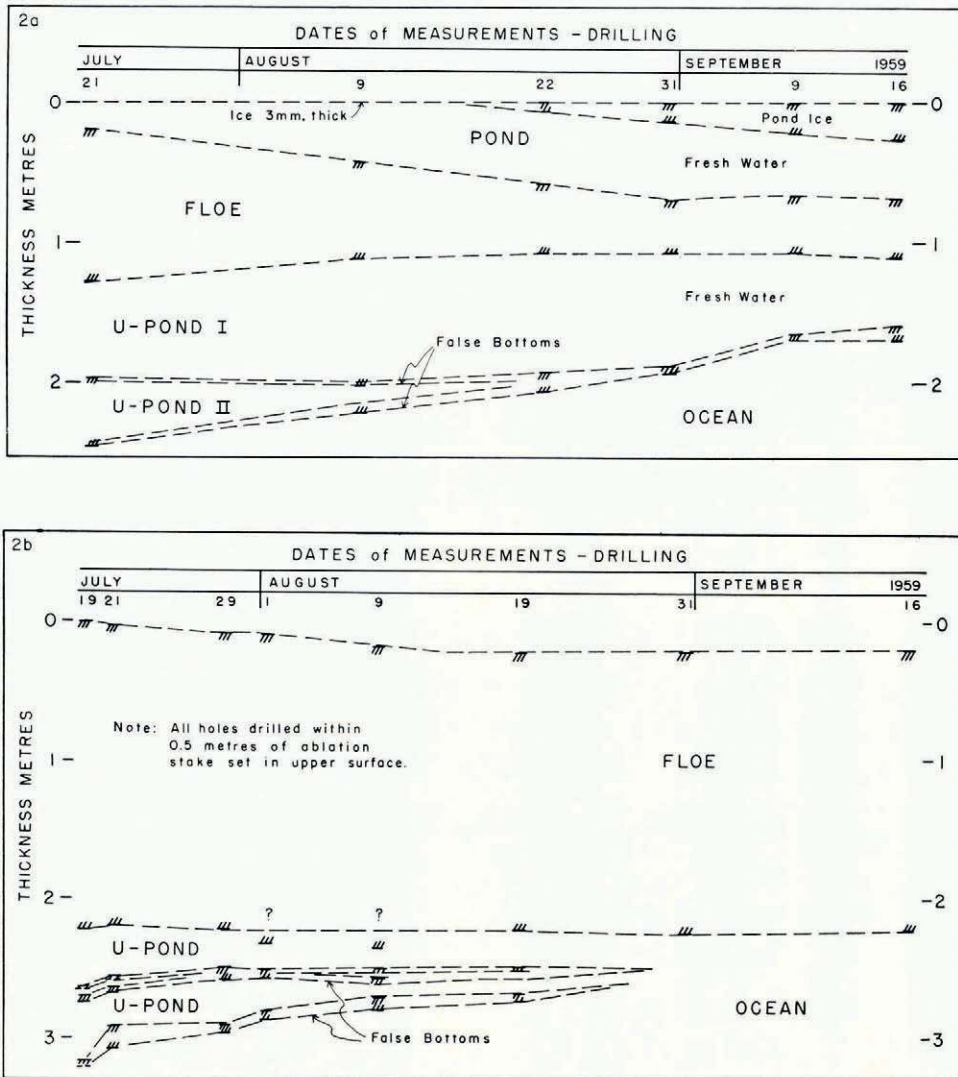


Fig. 2. (a) Time cross-section of floe at a site with both surface pond and under-floe pond near natural drain hole. (b) Time cross-section of floe near artificial drain hole

DISCUSSION

Measurements of the initial ice thickness of the floe at six points within the study site area ranged from 1.5 to 5.2 m., with an average of 2.9 m. The mass of a unit ice column, including 9.2 g. cm.⁻² of snow, was estimated as 270 g. cm.⁻² at the start of the ablation season.

The amounts of evaporation and condensation were so nearly in balance and were so small (mg. cm.⁻² day⁻¹) that these processes have been neglected in the analysis.

The mass loss at the upper surface of the floe (Tables I, III and IV) was calculated as a weighted mean of areas free of and covered by ponds, reduced by the amount of refrozen pond water. The largest source of error in the pond ablation measurements was caused by the

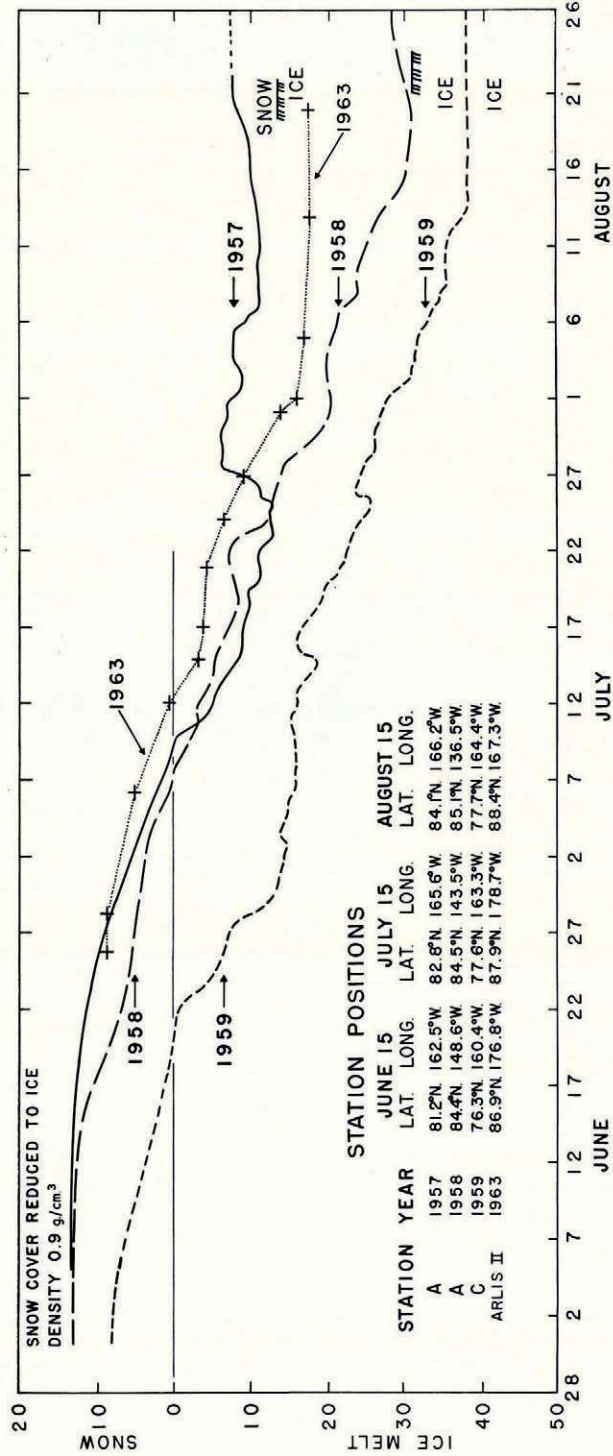


Fig. 3. Summer snow and ice ablation; non-ponded areas

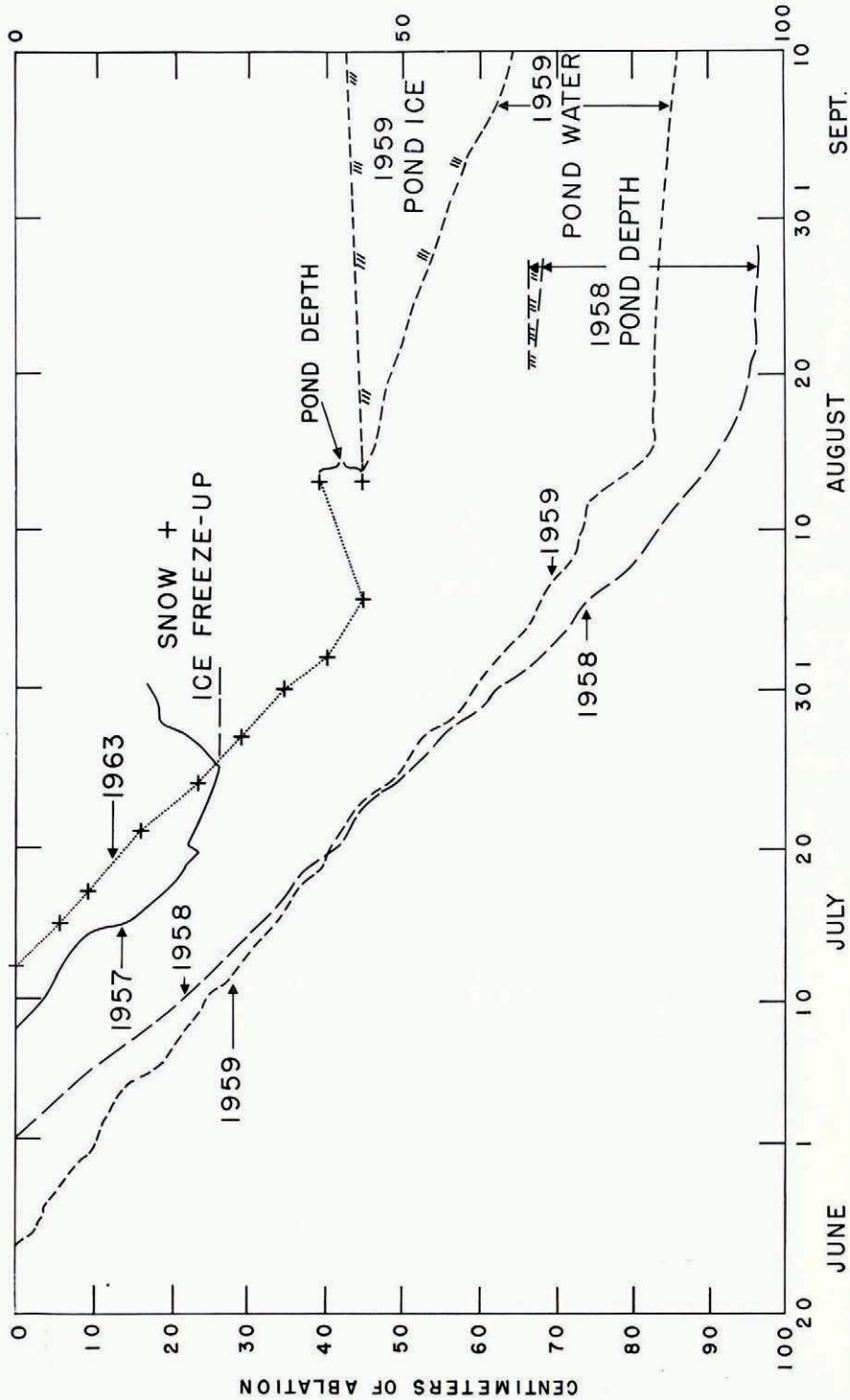


Fig. 4. Summer ice ablation; ponded areas

ablation stakes melting loose. An analysis of the pond ablation measurements gave a minimum value for the melt in the ponds. The actual loss was probably slightly greater.

The amount of ice gained or lost from the bottom of the floe was not well determined owing to the unknown fraction of the total surface which had under-floe ponding, but the best estimate from drilling is that under-floe ponds covered approximately one-half the floe bottom. The under-floe ponds, no doubt, protect the floe bottom from ablation; in fact some accretion can be expected (Untersteiner and Badgley, 1958). From the drilling and thickness-gauge measurements it is estimated that, while more than 9 g. of ice was lost by ablation, there was an average accretion of 6.3 g. (average of gauges 1, 4, 5 and 6; Table VI). The data of Table VI alone show a near balance in ablation and accretion, but the installation dates extend over a five-week period beginning after the end of the surface melt season. In the absence of the insulating effect of under-floe ponds, a larger bottom ablation would probably have been measured and also a larger accretion. A summary of the mass balance for station "Charlie" is presented in Table VII. From June through December approximately 20 per cent of the original mass was lost.

TABLE VI. BOTTOM ABLATION AND ACCRETION (in cm.), AS DETERMINED BY THICKNESS GAUGES

Date began	Gauge No. 1			Gauge No. 2*			Gauge No. 3†		
	Thickness	Ablation	Accretion	Thickness	Ablation	Accretion	Thickness	Ablation	Accretion
	2 September			15 September			28 August		
	450.0	—	—	252.7	—	—	221.0	—	—
30 September	442.5	7.5	—	251.3	1.4	—	204.9	16.1	—
29 October	442.2	0.3	—	249.0	2.3	—	199.7	5.2	—
29 November	442.6	—	0.4	248.8	0.2	—	200.6	—	0.9
21 December	443.2	—	0.6	247.5	1.3	—	198.3	2.3	—
Total		7.8	1.0		5.2	—		23.6	0.9
Net thickness change	-6.8			-5.2			-22.7		

Date began	Gauge No. 4			Gauge No. 5‡			Gauge No. 6		
	Thickness	Ablation	Accretion	Thickness	Ablation	Accretion	Thickness	Ablation	Accretion
	28 September			30 September			3 October		
	307.0	—	—	268.6	—	—	198.1	—	—
30 September	307.2	—	0.2	274.4	—	5.8	—	—	—
29 October	305.0	2.2	—	276.4	—	2.0	195.6	2.5	—
29 November	305.5	—	0.5	284.3	—	7.9	193.6	2.0	—
21 December	309.8	—	4.3	294.2	—	9.9	204.4	—	10.8
Total		2.2	5.0		—	25.6		4.5	10.8
Net thickness change	2.8			25.6			6.3		

* Under-floe pond 81 cm. thick.

† Under-floe pond 28 cm. thick; also surface pond with 26 cm. water and 11 cm. ice.

‡ Surface pond with 3 cm. water and 31 cm. ice.

TABLE VII. STATION "CHARLIE", MASS BUDGET FOR JUNE-DECEMBER 1959 (g. cm.⁻²)

	Gain/Loss	Balance
Initial mass	270.0	—
Loss of initial snow cover	-9.2	260.8
Solid precipitation during melt season, all of which melted	+2.8	263.6
	-2.8	260.8
Surface ice melt	-33.8	227.0
Snow and frost	+5.8	232.8
Bottom ablation	-10.0	222.8
Bottom accretion	+6.3	228.1

In Figures 3 and 4 melt season curves of four summers at three stations are presented. Furthermore, a summary of the net losses for the summers of 1957, 1958, 1959, 1961, 1962 and 1963, which have been adjusted according to estimates of pond area and remaining pond water, show rather wide variations (Table VIII). Apparently, the effect of latitude is outweighed by the variations of weather conditions during individual summers.

The 70 cm. melt estimated for 1961 is based on the measurement of the ablation of approximately 1 m. of contaminated ice. Similarly, for 1962 the measurement of the melt of contaminated ice was approximately 30 cm.

TABLE VIII. COMPARISON OF LOSS DUE TO MELTING OF THE UPPER SURFACE OF PACK ICE (g. cm.⁻²)

Station/Year	Net loss	Latitude
"Alpha" 1957	32	80-85° N.
"Alpha" 1958	53	84-86° N.
"Charlie" 1959	46	76-78° N.
ARLIS II 1961	(70)*	73-76° N.
ARLIS II 1962	(20)*	81-83.4° N.
ARLIS II 1963	28	86-88.5° N.

* Estimated (see text).

For energy budget calculations it must be considered that the total mass of ice melted is greater than the net loss by the amount of water retained in surface ponds. It may vary between 5 and 25 per cent of the total melt. An average of 40 g. cm.⁻² (Table VIII) appears to be a reasonable value of the average net loss of mass from the upper surface of old clean pack ice in the Arctic Ocean.

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