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The Capability of Satellite Borne Remote Sensors to Measure Stratospheric Trace Constituents

Volume II: Ozone and Aerosol Related Missions

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ABSTRACT

This document is Volume II of a three volume report issued as MITRE/METREK Technical Report, MTR-7519. The three volumes cover the following principal subjects:

Volume I contains a synthesis of the results of two previous MITRE/METREK studies {1,2} and an update of the information contained in them. The update was made during the Summer and Fall of 1977. These studies deal with a comprehensive review of stratospheric trace constituent measurement requirements. The scope of the study was restricted to those constituents which fall into the general category of "air pollutants."

Volume II separates stratospheric trace constitutent measurement requirements into two somewhat overlapping areas. In the first area, it is assumed that the only problem of interest is ozone; its chemistry chain, environmental effects and measurement requirements. In like manner, in the second area it is assumed that the only problem of interest is stratospheric aerosols; their chemistry, effects and measurement requirements.

Volume III contains material of a supportive nature not considered to be of sufficient importance to be included in the other two volumes. This material is of two types:

- Information and numerical evaluations used in the development of mission evaluations for stratospheric trace constituent measurement.
- Various spatial and temporal distributions for those stratospheric trace species having sufficient measurements available to warrant their presentation.

The reader is advised to note that the results and conclusions presented here are based on the specific combination of remote sensors, Shuttle orbits and analysis values selected to exemplify the technique presented. Although these sensors and orbits are typical, extension of the study to include all available sensors and many orbits, or to another specific small combination could result in different results and conclusions.



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LIST OF CHEMICAL SYMBOLS*

Symbol ·	Name
A	Argon
A1 ⁺⁺⁺⁺	Aluminum ion
nA1203	Aluminum oxide aerosol
Br	Atomic bromine
Br	Bromide ion aerosol
BrO	Bromine oxide
Ca ⁺⁺	Calcium ion aerosol
CBr ₄	Tetrabromomethane (carbon tetrabromide)
CC1 ₂ =CHC1	Trichloroethylene
cc1 ₄	Tetrachloromethane (carbon tetrachloride)
CFC12 ⁺	Dichlorofluoromethane radical
CFC1 ₃	Trichlorofluomethane (F-11)
CF ₂ C1 ⁺	Chlorodifluoromethane radical
CF2C1CFC12	Trichlorotrifluoroethane (F-113)
CF ₂ Cl ₂	Dichlorodifluoromethane (F-12)
CHC1F ₂	Chlorodifluoromethane (F-22)
CHC1 ₂ F	Dichlorofluoromethane
CHC1 ₃	Trichloromethane (chloroform)
CH ₂ :CHC1	Vinyl chloride
CH ₂ Cl ₂	Dichloromethane (methyl dichloride)

^{*}Common name given in parentheses where appropriate. Unless specifically stated, species is assumed to be in gaseous state.

Symbol	<u>Name</u>
CH ₂ O	Methanal (formaldehyde)
сн ₃	Methyl radical
CH ₃ Br	Bromomethane (methyl bromide)
CH3CC13	Trichloroethane (methyl chloroform)
CH ₃ C1	Chloromethane (methyl chloride)
сн ₃ о	Methyl oxy radical
CH302	Methyl peroxy radical
(CH ₃) ₂ S	Methyl sulfide
CH ₄	Methane
CO	Carbon monoxide
COCI	Carbonyl monochloride
cos	Carbonyl sulfide
co ₂	Carbon dioxide
nCO ₂	Carbon dioxide in cluster formation; quasi aerosol
cs ₂	Carbon disulfide
C2H4Cl2	Dichloroethane
с ₂ н ₅ с1	Chloroethane (ethyl chloride)
C H x y	Non-methane hydrocarbons (NMHC)
Cl	Atomic chlorine
C1 ⁻	Chloride ion aerosol
C1 ₂ C:CC1 ₂	Tetrachloroethene (perchloroethylene)
C1FCO ·	Fluoroformyl chloride

Symbol	Name
C10	Chlorine monoxide
Clono ₂	Chlorine nitrate
ClO ₂	Chlorine dioxide
C10 x	"Odd" chlorine
Cu ⁺⁺	Copper ion aerosol
F ₂ CO	Carbonyl fluoride
Fe ⁺⁺ or Fe ⁺⁺⁺	Iron ion aerosol
Н	Atomic hydrogen
HBr	Hydrogen bromide
HC1	Hydrogen chloride
HF	Hydrogen fluoride
HNO 2	Nitrous acid
HNO ₃	Nitric acid
nHNO 3	Nitric acid aerosol
HO or OH	Hydroxyl
но 2	Hydroperoxyl .
нѕо_	Bisulfite radical
H ₂	Molecular hydrogen
н ₂ 0	Water vapor
nH ₂ 0	Liquid water or ice (as aerosol or in cluster formation)
H O 2 2	Hydrogen peroxide

Symbol	Name
H ₂ S	Hydrogen sulfide
H ₂ SO ₄	Sulfuric acid
H ₂ SO ₄ nH ₂ O	Sulfuric acid aerosol
H C O	Unspecified organic compound
ī .	Iodide ion aerosol
κ* .	Potassium ion aerosol
М	Unspecified third body
Mg	Magnesium aerosol
Mn ⁺⁺ or Mn ⁺⁺⁺	Manganese ion aerosol
N	Atomic nitrogen
N ₂	Molecular nitrogen
^{nN} 2	Molecular nitrogen in cluster formation; quasi aerosol
NH 3	Ammonia
NH ₄ ⁺	Ammonium ion
nNH ₄	Ammonium ion aerosol
NH ₄ HSO ₄ • nH ₂ O	Ammonium bisulfate aerosol
(NH ₄) ₂ SO ₄	Ammonium sulfate aerosol
(NH ₄) ₂ s ₂ 0 ₈	Ammonium peroxydisulfate aerosol
NO	Nitric oxide
nNO	Nitric oxide in cluster formation; quasi aerosol
NO .	Nitrogen dioxide

Symbol	Name
NO ₂	Nitrite ion aerosol
NO ₃	Nitrogen trioxide
NO ₃	Nitrate ion aerosol
NO x	"Odd" nitrogen (nitrogen oxides)
N ₂ O	Nitrous oxide
N ₂ O ₅	Nitrogen pentoxide
Na ⁺	Sodium ion aerosol
0	Atomic oxygen, unspecified
o(¹ _D)	Atomic oxygen, excited state ¹ D
o(³ P)	Atomic oxygen, normal state
o(¹ s)	Atomic oxygen, excited state ¹ S
o ₂	Molecular oxygen
0 ₂ (¹ Δ)	Molecular oxygen, excited state $\overset{1}{\Delta}$
03	Ozone
sf ₆	Sulfur hexafluoride
so ₂ .	Sulfur dioxide
nS0 ₂	Sulfur dioxide in cluster formation; quasi aerosol
so ₃	Sulfur trioxide
so ₄ =	Sulfate ion aerosol
Si ⁺⁺⁺⁺	Silicon ion aerosol

1.0 INTRODUCTION AND CONCLUSIONS

1.1 General Objectives

Planned Shuttle missions provide another opportunity for

NASA to evaluate and utilize stratospheric remote sensors. This

unique platform will allow larger payloads than have been possible

while permitting the man-machine interface which has been lacking.

The key goals of any remote sensing experiment are the optimization

of sensors to the mission and the timely communication of the results

to the intended users.

In recent studies [1,2] for NASA/LaRC, MITRE/METREK has examined the sensor/constitutent/platform problems for the case of multi-purpose space missions. In the first of these, an assessment of the capabilities of specific NASA remote sensing systems to provide appropriate measurements of stratospheric parameters was made. This study emphasized the roles of the aerosol, the nitrogen oxide/ozone chemistry cycle and the chlorine/ozone chemistry cycle in the stratsophere. It also evaluated the capabilities of six specific instruments to provide required measurements of stratospheric constituents.

In the second study a more comprehensive view of all stratospheric trace constituents was taken. This included:

- development of a prioritized list of requirements for stratospheric trace constituent measurement;
- a comprehensive summary of present knowledge of stratospheric trace constituents;
- development of a detailed structured constituent/instrument/mission evaluation technique; and

 application of the technique to a specific set of instrument/ orbit combinations.

Since the completion of the two original studies a need has been recognized to synthesize the previous studies and combine this synthesis with additional updated information to produce a single document having the following principal objectives:

- providing the scientific community with a concise view of the current status of knowledge of stratospheric trace constituents and generating the impetus for frank and indepth discussions of future requirements;
- providing the instrument development community with an information set which would guide them in selecting optimum directions for new instrument development based on the combination of scientific needs and instrument capabilities.

Volume I of this document presents the results of this synthesis and updating.

In the current volume the fact is recognized that for a variety of reasons concern may focus on one particular aspect of the stratosphere such as the ozone cycle or the role of aerosols. The reasons for this specific concern may be political, scientific, economic or involve public concern over current issues such as supersonic aircraft or chlorofluoromethanes. These basic nontechnical concerns may result in limiting the scope of any stratospheric measurement program. Two such logical limitations would involve studies restricted to:

- the ozone chemistry cycle
- the role of aerosols in the stratosphere

It is these two limited measurement programs which are of concern in the present study. Each significant component of the

chemistry chains for both ozone and aerosols is identified and prioritized. The approach to this prioritization will be similar to that used in Volume I for the overall stratospheric study. This will yield two separate prioritized lists of constituents with some species appearing on both lists. From this point the analysis will proceed as in the previous study. The same instrument and orbit combinations will be used but evaluations will be limited to those species involved in either the ozone or aerosol chains. Two sets of results will be produced one giving the optimum instrument/orbit combinations for ozone chemistry and the other the optimum instrument/orbit combinations for measuring aerosols.

In most areas covered by this study, considerable effort has already been expended by many groups, not only within NASA but also among other government agencies, the private sector and in the two previous MITRE studies. MITRE's principle role in the present study was to integrate and reconcile these sometimes disparate sources and to provide informed opinions in the areas where either no data existed or a concensus was absent. The following subsections summarize the major sections of this report.

1.1.1. The "Natural" and "Perturbed" Stratosphere

The purpose of this section (2.0) is threefold. The first is to provide a readily available short summary of the general characteristics of the stratosphere. The temperature regime and circulation are discussed in terms of the general dynamic processes to illustrate

the formation of the unperturbed stratosphere. This leads to a summary of the stratospheric constituents and their role in atmospheric chemistry. The later two parts of the section present summaries of ozone chemistry and the role of stratospheric aerosols. This section is intended only as a supporting base of information for understanding the various topics covered later.

1.1.2 User Requirements for Stratospheric Measurements

In volume I of this report the results of an extensive user requirements study are presented. That presentation discussed some general features of NASA's interaction with users of stratospheric data and presented the two major examples (ultraviolent radiation* and climate) of pressing atmospheric pollution problems which demand of NASA a careful and effective program of development. The emphasis was placed on who uses the data and how they use it in order to develop the specifics of the measurement requirements. In the present section (3.0) this work is summarized with emphasis placed on the distinction in requirements to satify users concerned with ultraviolet radiation (UV) and those concerned with climate. The general assumption is made that ozone chemistry primarily affects UV while aerosols primarily affect climate.

^{*}Biologists divide the ultraviolet spectrum into three wavelength regions: UV-A: 0.32 to 0.4 μm ; UV-B: 0.28 to 0.32 μm ; UV-C: less than 0.28 μm . Unless specifically stated the term UV when used in this report refers to all three regions.

1.1.3 Science Requirements

In this section (4.0) the analysis of user needs and general measurement requirements developed in the previous setion (3.0) are used in combination with present knowledge of stratospheric chemistry and with the results of many other recent studies to develop two sets of prioritized scientific requirements for stratospheric trace constituent measurements. The first set focuses on ozone chemistry and the second on aerosols.

1.1.4 Mission Evaluations

This section (5.0) presents the results of the application of a method for the evaluation of various stratospheric species measurement missions. The method itself was developed during previous MITRE effort [2] and is presented in detail as Appendix A in both Volumes I and II of this report. Much of the supporting data used in these evaluations has been assembled in several of the appendices of Volume III of this report. Separate evaluations are made for ozone and aerosol applications. The evaluations include all possible combinations of three orbits (a 30° shuttle type, a 56° shuttle type and a polar type) with one or more of four remote instruments,

- Limb IR Monitor for the Stratosphere (LIMS)
- Stratospheric Aerosol and Gas Experiment (SAGE)
- Correlation Interferometer for the Measurement of Atmospheric Trace Species (CIMATS)
- Halogen Occultation Experiment (HALOE)

1.2 Conclusions

Many of the generalized conclusions contained in Volume I of this report apply equally well to specific ozone or aerosol missions. The following sections will present the more significant of these conclusions along with specific points relating to the ozone and aerosol mission evaluations.

1.2.1 Current and Projected Measurement Capability

Analysis of the material presented here and in Volume I indicates three key conclusions:

- (1) The performance of current remote stratospheric sensors, in some cases, compares quite well with identified measurement requirements. Their ability to measure other species has not been demonstrated. A number of in-situ methods also exist with comparable sensitivity and accuracy but whose measurements are of a limited utility, given their spatial and temporal sampling characteristics.
- (2) None of the current, in-situ methods have the capability to satisfy the requirements for global monitoring and the temporal constraints derived from the users needs portion of the study.
- (3) Existing, non-remote techniques will continue to play an important role in stratospheric investigations for both corroboration of remotely collected data and in the evolutionary development of future remote sensors.

All of the measurment techniques discussed have their strengths and weaknesses. The <u>in-situ</u> methods are extremely sensitive and accurate but suffer from limited coverage and local contamination problems. Remote sensing techniques offer wide area coverage and relatively long mission lifetimes. Their disadvantages lie in the

reduced senstivity to low concentration levels and the requirements for auxiliary data to invert the integrated path measurements which most utilize. Indeed, the masses of data which must be processed in order to yield the desired information is at least a temporary disadvantage of remote sensing methods. The development of better models and improved data handling techniques is expected to minimize these problems.

The general features of remote sensors of the stratosphere aboard a satellite platform reveal two key features:

- (1) nadir-viewing instrumentation provides superior performance in the areas of horizontal resolution and measurement time per orbit
- (2) limb-viewing instrumentation provides superior sensitivity, and vertical resolution

In most other areas, the two basic monitoring methods are equally capable. The science requirements include the need for vertical profiles and data of fairly high quality. Limb-viewing instrumentation appears to satisfy these needs but provides limited temporal sampling for solar occultation when certain orbits are used. As a result, instrumentation of the limb emission type represents the optimum choice. In general, this type of instrument has the potential of satisfying scientific requirements for vertical profiles as well as those for spatial and temporal sampling.

Orbital considerations emerge as a key element in the applicability of various sensor systems to specific measurement roles. Sunsynchronous orbits provide optimum coverage for nadir-viewing, thermal source sensors and limb-viewing emission source sensors. High angle non-sunsynchronous

orbits are preferred to nadir-viewing reflected solar source or limbviewing solar occultation sensors, if geographical coverage is to be maximized.

1.2.2 Requirements for Stratospheric Measurements

Material utilized in the selection of requirements for stratospheric monitoring has been derived from the user needs survey as
well as the detailed investigation of data needed for a better
understanding of stratospheric chemistry. In addition, a review of
current measurement methods examined the quality of data currently
available for a variety of gases of interest. The proposed accuracy
requirements reflect improvements, where required, over current
limitations.

In many cases no specific requirements have been expressed for spatial or temporal sampling. In view of the generally infrequent and localized nature of current measurements, any satellite monitoring system will represent an improvement in these categories. It is anticipated that the need will exist for global coverage at a rate which provides data on diurnal and seasonal variations as well as longer term trends.

Based on the measurement requirements expressed by various categories of users and numerous other specific studies, prioritized lists
of properties and species have been developed for ozone and aerosol
oriented missions (Tables 4-II and 4-III respectively). Both lists
give the highest priority for measurement of three basic properties:

- stratospheric temperature
- solar irradiance
- earth radiance

In addition the ozone oriented list stresses measurements in the following categories:

- pure oxygen forms, 0
- hydrogen oxides, HO_v
- nitrogen oxides, NO_x
- chlorine oxides, ClO_x

The aerosol related list places the highest priority on measurements of:

- total aerosols
- total sulfate aerosols
- sulfuric acid aerosols
- ammonium sulfate, and
- major sulfate aerosol precursor gases

It must be remembered that these lists have been developed on a purely scientific basis, without regard to present knowledge or potential measurment capability. Later in the report application of the evaluation methodology presented indicates that many of the high priority properties and species do not receive the highest priority for planned satellite missions since their distributions are much more understood than most of the other important stratospheric species. Those species which show high priority for satellite missions are typically the

components of the basic reactions involving the direct production or depletion of ozone or the formation of sulfuric acid or ammonium sulfate aerosols.

As our understanding of the stratosphere matures, various constituents will receive more or less emphasis with respect to sampling and data quality. While these lists are presently current, changes should be anticipated, particularly when measurements exceed the current minimum requirements.

It should be noted that these requirements have been generated independently of any instrument considerations. Therefore, this material represents a set of performance goals for contact or remote sensors placed on airborne, orbiting, or terrestrial platforms. In the case of those species not yet measured, airborne measurements should receive considerable attention in order to establish background levels and to corroborate proposed remote sensing techniques.

1.2.3 Selected Instrument/Orbit Evaluations

Within the constraints imposed by the sensor complement examined and the choice of three orbits selected, the various sensor-orbit combinations were evaluated for each species of interest. For stratospheric study, the limb-scanners scored significantly higher than either the nadir-viewing or the solar occultation class of instruments. This is attributable to the direct vertical profiles which the limb-scanners provide.

Among the three orbits investigated, the polar orbit does not automatically give the best coverage. This tends to be a function of instrument type. For limb emission instruments, the higher the inclination angle the greater the global coverage. However, the poorest latitudinal coverage of all the combinations examined is obtained in the case of solar occultation from polar orbits.

It must be emphasized that the present evaluation was performed for a limited number of instruments and orbits. The methodology is sufficiently flexible to allow new instruments to be included in subsequent analyses of this type. If any of the instruments considered should prove incapable of all the measurements for which they are credited, their relative standing in a later analysis would suffer proportionately.

Evaluation of the various instrument/orbit combinations for an ozone related mission does not yield as clear a distinction among the combinations as in the evaluation of the total stratospheric program done in Volume I. The results (Table 5-XXII) indicate some preference for the three instrument mission containing the LIMS, CIMATS and HALOE. In addition a two instrument mission which contains LIMS and CIMATS scores favorably. The reasons for the superiority of these missions lies primarily in the number of species measured and secondarily in the relative importance of those measurements. For all of the eleven ozone mission combinations examined, the 56° orbit is seen to be capable of satisfying the greatest number

of scientific requirements. This is due to the preponderance of solar occultation measurements in these ozone missions.

In contrast with the ozone results, orbit selection for combinations of instruments for aerosol missions does not consistently show the 56° orbit to score the highest. The polar orbit tends to prevail in combinations containing the LIMS instrument while the 56° orbit is superior when it is absent. In no case is the 30° orbit shown to be superior.

As with the ozone results, aerosol missions tend to score higher based on the number of species measured with the LIMS, CIMATS, HALOE combination and the LIMS, CIMATS combination prevailing. The relative distinction between the various combinations is somewhat clearer than in the ozone case. Those instruments (i.e., LIMS and CIMATS) measuring the high priority sulfate precursor gases, ammonia and water vapor, give the best evaluation results.

2.0 THE "NATURAL" AND "PERTURBED" STRATOSPHERE

Many recent studies refer to the stratosphere in one of two ways, namely, "natural" or "perturbed". In general the exact meaning of these terms is assumed to be known rather than defined. In this report the term "natural" refers to the long term average characteristics of the stratosphere which existed before any significant anthropogenic activities. The term "perturbed" is used when refering to any significant change in the natural state whether temporary (months to years) or quasi-permanent. Stratospheric perturbations can be caused in several ways:

- natural disturbances such as volcanic eruptions
- direct anthropogenic disturbances such as chlorofluorocarbon releases, fossil fuel burning or high altitude aircraft operation
- indirect anthropogenic activities such as massive changes in land use or the ordinary growth and progress of the human race

The purpose of the following subsections is to acquaint the reader with the object of satellite remote sensor measurements—the stratosphere. The temperature regime and circulation are discussed in terms of the general dynamic processes to illustrate the formation of the unperturbed stratosphere. This leads to a summary of the stratospheric constituents and their role in atmospheric chemistry.

The two major reasons for observing or monitoring the stratosphere are to gain a more complete understanding of the subject and to be able to predict changes in the environment. Inadvertent modifications of the atmosphere by pollutants can have far-reaching effects upon man's activities. Chemical and physical processes, in terms of both ozone destruction and aerosol formation, will be summarized below to provide a background for later discussions contained in this report.

2.1 The "Natural" Stratosphere

2.1.1 Physical Formation of the Tropopause

The earth's atmosphere may be divided into layers which are characterized by the average vertical temperature structure. The three layers closest to the earth's surface are:

- Troposphere,
- · Stratosphere, and
- Mesosphere.

Averaged over reasonably long periods of time, the temperature of the troposphere decreases regularly with altitude. At an elevation that varies systematically with latitude and seasons, the temperature becomes isothermal. This property defines the tropopause, which lies between 8 and 16 km. The stratosphere is the region above the tropopause and below the stratopause. In this region, the temperature is typically constant or increasing with altitude. This increase is reversed at an altitude of about 45 to 50 km—the stratopause. The region immediately above the stratopause is the mesophere.

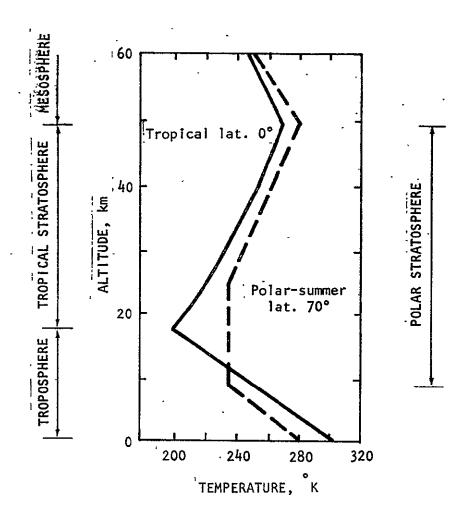


FIGURE 2-1
SAMPLE TEMPERATURE PROFILES IN TROPICAL AND POLAR ZONES [3]
(Troposphere, stratosphere, stratopause, and mesosphere defined in terms of vertical temperature profiles)

The vertical distribution of temperature in the tropical and the polar zones is shown in Figure 2-1 [3]. The two temperature profiles of Figure 2-1 show substantial differences between polar and tropical regions.

The special properties of the stratosphere—its temperature inversion and the resulting slow vertical mixing—are a consequence of the presence of 0_3 , which is formed rapidly in the upper stratosphere. The formation of 0_3 occurs at an altitude of 30 to 50 km by the photolysis of 0_2 , producing 0, which in turn recombines with 0_2 to form 0_3 . Some of the physical reasons behind the temperature inversions at the tropopause are discussed below.

If heat from the ground were the only source of energy in the atmosphere, the vertical temperature at a given location would decrease monotonically with altitude. In contrast, measurement of the vertical temperature profiles shows that beyond the tropopause, to a height of about 50 km, the temperature increases. At this height, the stratopause, the temperature undergoes an inversion and again starts to decrease.

One to three percent of the incoming solar radiation is absorbed by the 03 layer in the stratosphere. The absorbed energy heats adjacent layers. The model now contains two sources of energy in the atmosphere, one at the surface and the other at an altitude of about 30 to 50 km. From this simplified picture, it is evident that a temperature inversion should occur at a height between the two

sources. The region where the inversion occurs defines the tropopause, which lies between 8 and 16 km depending on the season, latitude, and synoptic weather situation.

2.1.2 Transport Phenomena

Clouds, rain, and thunderstorms are strong evidence for the considerable vertical motion characteristic of the troposphere. In thunderstorms vertical velocities, which are generally 10 cm/sec in normal latitude cyclones and anticyclones, may reach 10 to 20 m/sec. In the stratosphere, however, the temperature increases with height providing an equilibrium condition. For this reason, the vertical motions rarely exceed a few centimeters per second and are often much smaller. In other words, an air parcel moves up or down more slowly in the stratosphere than it does in the troposphere. This is not true for horizontal motions in the stratosphere which are significantly more rapid than the vertical motions. Typical horizontal wind velocities in the stratosphere are of the order of 1 to 100 m/sec, whereas vertical velocities are in the range of 10⁻⁴ to 10⁻¹ m/sec.

The overall structure of the wind field in the stratosphere has been investigated and shows a complicated latitudinal and seasonal dependence [5,6]. In general, there are some correlations between the meridional (N-S) and vertical wind fields at different times of the year [4]. No correlation seems to exist between the rapid zonal (E-W) circulation and vertical wind data.

In summary, because of the slow vertical mixing, the contaminants which are introduced into the stratosphere at a particular altitude will remain near that altitude for periods as long as several years [7]. This long residence time allows the contaminants to take part actively in the chemical and radiative processes of the stratosphere. In the case where a contaminant is capable of entering a catalytic process which would lead to the destruction of an important stratospheric constituent such as ozone, the consequences are of great importance and must be thoroughly investigated.

2.1.3 Stratospheric Constituents

The constituents of the stratosphere may be separated into four categories. These are:

- · Major gaseous constituents,
- · Minor gaseous constituents,
- Trace gaseous constituents, and
- Aerosols.

The major atmospheric constituents are N_2 , O_2 , A and CO_2 . The accepted value for N_2 concentration is 78.08 percent by volume of dry air. Recent oxygen measurements show a concentration of 20.95 percent by volume when corrected to dry air conditions [4]. Argon has a stratospheric background concentration of 0.93 percent and carbon dioxide of 0.03 percent at about 20 km.

. The minor constituents, such as 0_3 , H_2O , CH_4 , etc.; have concentrations of a few parts per million in the stratosphere. Table 2-I

summarizes some of the minor constituents at 20 km that are known to be important in stratospheric chemistry and circulation. The number of important trace gaseous constituents known to play a significant role in stratosphere chemistry is very large and growing rapidly. A complete development of the characteristics of these constituents is given in later sections.

Besides these gaseous chemical constituents, a layer of particles several kilometers thick, exists in the stratosphere. This layer, called the "Junge layer," is located several kilometers above the tropopause. The Junge, or sulfate layer, has a particle density of two to ten times that exhibited above and below this layer. The particle size is predominately in the 0.1 to 1.0 µm radius range. The particle distribution shows a decreasing concentration with increasing size. The particles consist mainly of sulfuric acid or sulfate solutions and are probably in a supercooled liquid state. The characteristics of this aerosol layer and of all stratospheric aerosols will be covered in detail in later sections.

2.2 The "Perturbed" Stratosphere

As stated previously perturbations in the stratosphere may be caused by either natural or anthropogenic changes. Natural changes generally take the form of injection of gaseous, liquid or solid materials through the tropopause into the stratosphere. Such changes may be either abrupt as in the case of volcanic eruption or quasicontinuous as in the case of the daily worldwide occurrence of thunderstorms. However, in this later case there is some question as to

TABLE 2-I
MINOR STRATOSPHERIC CONSTITUENTS

Species	Concentration at 20 km	Variability	Importance
03	6 ppmv	Factor of two or more diurnal, season, latitude and height.	UV-shield, . radiative heating and cooling of strato- sphere.
Н ₂ 0	3 ppmv	With latitude, season, and altitude.	Radiative balance, clouds, particle formation, 0 chemistry.
CH ₄	1 ppmv	Decreases with height above tropo- pause.	Chemical source of OH. Possible sink of C1, indicator of tropopause interchange.
н ₂	0.55 ppmv	Increases to a maximum of 0.8 ppmv at 28 km and decreases to 0.4 at 50 km.	0 ₃ -chemistry.
N ₂ 0	0.1 ppmv	Decreases with altitude, sea-son, and latitude.	Source of stratospheric
CO	0.05 ppmv	May decrease above tropopause, but actual profile and variations are unknown.	Indicator of troposphere-stratosphere exchange. By-product of CH ₄ chemistry.

whether thunderstorms should be considered a perturbation or merely one part of the natural stratosphere. For the purposes of this report they will be treated as perturbations.

Injection of extraterrestrial material downward through the mesopause seems to occur but the characteristics of this mechanisms are poorly understood. At present the total possible mass of such injections is thought to be small and would represent a negligible optical perturbation [66].

Anthropogenically caused changes may take the form of direct release of material in the stratosphere as with aircraft and rocket exhaust or upward diffusion and transport through the tropopause of material released into the troposphere: Table 2-II lists the major categories of materials currently known to cause significant stratospheric perturbations along with the sources of each. The table includes both natural and anthropogenic materials.

The ultimate stratospheric effects caused by these perturbations belong to two chains, the UV chain and the climate chain. The direct channels for production of these effects are three:

- (1) Changes in UV transmission resulting from changes in ozone concentration.
- (2) Changes in the radiative heat balance caused by formation of aerosols.
- (3) Changes in the radiative heat balance caused directly by some gases such as carbon dioxide.

Since 0_3 concentration controls the amount of UV-B radiation that reaches the surface of the earth, a reduction in 0_3 concentration will

TABLE 2-II PRINCIPAL SOURCES OF STRATOSPHERIC PERTURBATIONS

TYPE OF PERTURBATION	TRANSPORTED FROM TROPOSPHERE	RELEASED BY AIRCRAFT IN-SITU
Liquid water or ice Particles Gases	√* √**	✓`
Carbon Dioxide Water Vapor Nitrous Oxide Other Nitrogen Oxides Ammonia	✓ ✓ ✓	✓ ✓
Hydrogen Sulfide Sulfur Dioxide Carbon Monoxide and Hydrocarbons Chlorine Compounds including Chlorofluoromethanes	√*** √*** ✓	✓

^{*} Major mechanism is thunderstorms and cumulonimbus clouds.

** Major mechanism is volcanic eruption.

Through ordinary transport and volcanic eruption.

increase the amount of this radiation, which has been shown to cause skin cancer and other biological effects [42]. The increase in aerosol concentrations (besides increasing the potential for hetrogeneous reactions whose effects are not well understood at present) will perturb the radiation balance of the earth's atmosphere and may lead to climatic changes, affecting sunshine, temperature, and precipitation. In addition to these, CO₂ and H₂O vapor introduced into the stratosphere by aircraft or Space Shuttle vehicles may increase the greenhouse effect and lead to stratospheric watering, which would perturb the natural circulation of the stratosphere.

Table 2-III shows which of the three channels for producing UV or radioactive heat balance changes are associated with the various categories of stratospheric perturbations. In the following subsections the mechanisms for ozone change and aerosol production are discussed in more detail. Direct radioative heat balance changes are not discussed since they are beyond the scope of this report.

2.3 Stratospheric Chemistry

The stratosphere contains many different kinds of reactive chemical species. Any one of these species can react with a number of others, or be generated by a variety of other reactions which are sometimes very complex and quite indirect.

As related to stratospheric chemistry in general, three types of reactions may be distinguished. These are:

· Photochemical reactions,

TABLE 2-III

PRINCIPAL CHANGES ASSOCIATED WITH STRATOSPHERIC PERTURBATIONS

TYPE	CHANGES IN	CHANGES RADIATIVE HEAT	
OF PERTURBATION	OZONE AND UV TRANSMISSION	VIA AEROSOLS	DIRECT
Liquid water or ice	√ .	✓	✓
Particles · .	·	✓	
Gases			ļ
Carbon Dioxide			✓
Water Vapor	✓	. ~	√
Nitrous Oxide	✓		
Other Nitrogen Oxides	✓ .		_
Ammonia	✓	✓	1
Hydrogen Sulfide		✓	
Sulfur Dioxide	` ∀	· ·	
Carbon Monoxide and Hydrocarbons	✓		
Chlorine Compounds including Chlorofluoromethanes	√		

- · Homogeneous reactions, and
- · Heterogeneous. reactions.

Photochemical reactions involve the interaction of electromagnetic radiation of of varying wavelengths with constituents of the stratosphere. Photochemical interactions are the only known source of stratospheric ozone production.

Homogeneous reactions are those reactions in which both the reactant species and the products are in a gaseous phase. If in these reactions a "third body" is needed to carry off energy to prevent dissociation of the product, that third body is a gas molecule.

Heterogeneous reactions are those reactions in which a particle, solid or liquid, interacts with gaseous species. The interaction may be catalytic, or the particle itself may take part in the reaction.

In the following subsections a more detailed discussion of the chemistry of ozone and aerosols will be presented.

2.3.1 Ozone Chemistry

As stated above many reactive trace gases from both natural and anthropogenic sources are tansported into the stratosphere or released in situ. Although the mass of the stratosphere is an order of magnitude less than that of the troposphere, the rate of exchange with the troposphere is slow and there is no significant washout by precipitation. For these reasons relatively minor perturbations in the stratospheric balance can have significant long term effects.

The most important trace constituent in the stratosphere is ozone. Any possibility of decreasing or even increasing the ozone content of the stratosphere by altering the natural chemical balance must be viewed as potentially dangerous. Two major applied problems exist in current studies of stratospheric ozone. These are the effect of aircraft and rocket exhausts released in the stratosphere and the transport of chlorofluoromethanes from the troposphere.

Other relatively less significant problems also exist as was shown previous in Table 2-II. In regard to ozone chemistry, perhaps key among these others are anthropogenic releases of nitrous oxide, ammonia and various organic chemicals.

An assessment of the effects of these stratospheric perturbations requires an understanding of the natural chemistry and atmospheric dynamics of the ozone balance. This is the basic problem. The chemistry which originally was thought to be relatively simple, is now known to be quite complex and involving many species. In fact, except for the inert gas Argon, every gas currently known to exist in the stratosphere can be related either directly or indirectly with either ozone chemistry, aerosol chemistry or both.

In order to describe the chemistry of ozone formation and destruction the various reactions will be discussed in the following sequence.

- pure oxygen reactions
- hydrogen-oxygen reactions

- methane and hydrocarbon reactions
- nitrogen reactions
- chlorine reactions
- chlorine-nitrogen reactions
- other chemicals

In discussing the various ozone chemistry chains, the various gases can be divided into three types, reactive, inactive and passive. The term reactive will refer to those gases which readily and vigorously participate in reactions leading to the increase or decrease of ozone. Examples of such are atomic oxygen, hydroxyl, nitric oxide and chlorine monoxide.

Inactive gases are those which when formed during some ozone related process are removed completely from further participation in ozone chemistry. The prime example of such is carbon dioxide usually formed from the reaction of carbon monoxide and hydroxyl. There is no known mechanism for the reentry of carbon dioxide into any stratospheric ozone chemistry chain.

Passive gases are less reactive gases formed from one or more reactive components which temporarily prevent the reactive components from vigorous participation in ozone chemistry. An example of such is hydrogen chloride gas. While in this state the chlorine atom is prevented from reacting with ozone. Since the rate of formation of hydrogen chloride gas by reaction of atomic chlorine with methane or HO is one or two orders of magnitude greater than the rate of

destruction of hydrogen chloride gas by hydroxyl or atomic oxygen the hydrogen chloride gas serves as a quasi sink for chlorine atoms.

Such gases are referred to as passive (i.e., less reactive).

2.3.1.1 <u>Pure Oxygen Reactions</u>. The pure oxygen reactions in ozone chemistry refer to those in which only some forms of oxygen participate. These reactions may involve solar ultraviolet light or the presence of a third body (typically N₂ or another O₂ molecule). These reactions may be termed natural since they occur irregardless of any stratospheric perturbations. The principal pure oxygen reaction are:

$$0_2 + h\nu \ (\lambda < 0.242 \ \mu m) \longrightarrow 0 + 0$$
 (1)

$$0 + O_2 + M \longrightarrow O_3 + M$$
 (2)

$$0_3 + hv (\lambda : 0.45 - 0.65 \mu m) \longrightarrow 0 + 0_2$$
 (3a)

$$0_3 + hv (\lambda : 0.31-0.34 \mu m) \longrightarrow 0_2 (^1\Delta) + 0 (^3P)$$
 (3b)

$$o_3 + hv (\lambda < 0.31 \mu m) \longrightarrow o (^1D) + o_2$$
 (3c)

$$0_3 + 0 \longrightarrow 0_2 + 0_2 \qquad . \tag{4}$$

$$0 + 0 + M \longrightarrow 0_2 + M \tag{5}$$

Reaction (1) is concentrated in the upper stratosphere and provides the initial source of atomic oxygen. The reactions (2) through (4) occur at lower altitudes in the stratosphere and reach their maxima in the same altitude area as the ozone maximum. Reaction (5) typically occurs in the mesosphere and can be neglected in the stratosphere since reaction (2) predominates [75].

2.3.1.2 Hydrogen-Oxygen Reactions. Theoretically reactions involving hydrogen, oxygen and/or hydrogen oxides (HO_X) can occur via direct two body collision, three body collisions involving a neutral third body (M) or by photochemistry. The two body collision is the predominant one in the stratosphere with photochemical processes occurring in the mesosphere.

The principal reactions involving a single hydrogen atom are:

$$H + O_2 + M \longrightarrow HO_2 + M \tag{6}$$

$$H + O_3 \longrightarrow HO + O_2 \qquad (7)$$

$$HO_2 + O \longrightarrow O_2 + HO \tag{8}$$

$$HO + O \xrightarrow{} H + O_2$$
 (9)

$$HO_2 + O_3 \longrightarrow HO + 2O_2$$
 (10a)

$$HO + O_3 \xrightarrow{\cdot} HO_2 + O_2 \tag{10b}$$

The first three of the above reactions principally occur in the upper stratosphere in the 35 km and above region. The atomic hydrogen involved initially in equation (6) is transported downward from higher levels in the mesosphere. These three reactions provide a mechanism for the production of hydroperoxyl and hydroxyl in the upper stratosphere. In the lower stratosphere (20 to 40 km) these two radicals which may be transported downward from the upper stratosphere or produced from other in situ reactions act as removal mechanisms for ozone and atomic oxygen (equations 7 through 10b).

Direct production and depletion of hydroxyl in the lower stratosphere depends on the two principal reactions:

$$0 (^{1}D) + H_{2}O \longrightarrow HO + HO$$
 (11)

$$HO + HO_{2} + H_{2}O + O_{2}$$
 (12)

Thus the presence or absence of water vapor in the lower stratosphere enters as a significant factor in the ozone balance.

The principal reactions which involve the production and destruction of molecular hydrogen, ${\rm H_2}$, are:

$$H + HO_2 \longrightarrow H_2 + O_2 \tag{13}$$

$$H_2 + O(^1D) \longrightarrow H + HO$$
 (14)

Reaction 13 is important in the mesosphere and the hydrogen molecules produced may be transported downward into the stratosphere where they react with atomic oxygen to produce additional hydroxyls.

The extent to which hydrogen peroxide enters into the hydrogen/
oxygen chemistry cycle is poorly understood and requires more experimentation [75]. The basic reaction for formation of hydrogen peroxide
involves the direct combination of 2 hydroperoxyl radicals:

$$HO_2 + HO_2 \longrightarrow H_2O_2 + O_2$$
 (15)

Hydrogen peroxide is destroyed either through photodissociation,

$$H_2O_2 + h\nu \longrightarrow HO + HO$$
 (16)

or by reaction with hydroxyl,

$$HO + H_2O_2 \longrightarrow H_2O + HO_2$$
 (17)

This entire chain of reactions should be the subject of further investigation.

2.3.1.3 <u>Carbon Monoxide and Hydrocarbon Reactions</u>. The role of carbon monoxide in the stratospheric ozone balance is indirect but nevertheless very important. One principal reaction dominates,

$$HO + CO \longrightarrow CO_2 + H \tag{18}$$

The hydrogen atoms released are thus available to react with molecular oxygen to produce the highly reactive hydroperoxyl as shown in reaction (6). In the upper troposphere and lower stratosphere reactions (17) and (6) combine to provide the principal pathway for conversion of hydroxyl to hydroperoxyl. It is therefore important to know the concentration and vertical distribution of carbon monoxide in the lower stratosphere. In the middle stratosphere the hydroxyl shows preference for reaction with ozone rather than carbon monoxide while in the upper stratosphere the atomic oxygen-hydroxyl reaction is favored.

The principal hydrocarbon involved in ozone chemistry is methane. Methane, which is transported upward from the troposphere, is photodissociated in the mesosphere and dissociated by oxidation in the stratosphere. The two principal reactions are with excited atomic oxygen, 0 (1D) or with hydroxyl,

$$- 0 (^{1}D) + CH_{4} \longrightarrow CH_{3} + HO$$
 —(19)

$$HO + CH_4 \longrightarrow CH_3 + H_2O$$
 (20)

In the stratosphere the methyl radicals which are produced in the above reactions have a preference for reaction with molecular oxygen in a three body manner,

$$CH_3 + O_2 + M \longrightarrow CH_3O_2 + M$$
 (21)

The methyl peroxy radicals produced usually enter into the nitrogen chemistry chain in a variety of ways as will be discussed later.

These nitrogen chemistry reactions can yield formaldehyde which added to any formaldehyde transported from the troposphere can photodecompose yielding either atomic or molecular hydrogen,

$$CH_{2}O + h\nu \longrightarrow CHO + H$$
 (22)

$$CH_2O + h\nu \longrightarrow CO + H_2 \tag{23}$$

2.3.1.4 Nitrogen, Hydrogen and Oxygen Reactions

Although the nitrogen oxides chemistry cycle is extremely complex with many possible reactions, a simple description of the essential reactions indicates the basic cycle. Nitric oxide may be released into the stratosphere by high flying aircraft. However, the predominent source is through the reaction,

$$O(^{1}D) + N_{2}O \longrightarrow 2NO$$
 (24)

 ${
m N}_2{
m O}$ is the principal nitrogen oxide transported upward from the troposphere where it is formed at the earth's surface through biological processes.

Once the NO is formed an equilibrium is established between NO and NO, through a series of reactions. The three principal ones are,

$$NO + O_3 - NO_2 + O_2$$
 (25)

$$NO_2 + O \longrightarrow NO + O_2 \tag{26}$$

$$NO_2 + h_V (\lambda < 0.398 \ \mu\text{m}) \longrightarrow NO + O$$
 (27)

Reaction (25) is the same as that occurring in urban areas during smog production. Of the two reactions for contaction of NO₂ reaction (26), as expected, occurs principally higher in the stratosphere where more atomic oxygen is available and reaction (27) occurs at lower altitudes. Due to the need for UV in reaction (27) the overall net effect of the three reactions is an increase in NO₂ at night and NO in the daytime.

At present the only know sink mechanism for stratospheric NO $_{\rm X}$ is conversion to one of the nitrogen oxide acids. This is followed by some type of aerosol formation and eventual settling into the troposphere where rainout can occur. The principal NO $_{\rm X}$ conversion reactions involve interaction with HO $_{\rm X}$ and a third body,

$$NO + HO_2 + M \longrightarrow HNO_3 + M$$
 (28)

$$NO_2 + HO + M \longrightarrow HNO_3 + M$$
 (29)

$$NO + HO + M \longrightarrow HNO_2 + M \tag{30}$$

The conversion of NO $_{
m x}$ to acids is also possible through a wide variety of reactions with organic compounds. Typical of these are reactions with methyl oxy or peroxy radicals,

$$CH_3O + NO_2 \longrightarrow CH_2O + HNO_2$$
 (31)

$$CH_3O_2 + NO_2 \longrightarrow CH_2O + HNO_3$$
 (32)

As noted earlier reactions of this type form a link between methane or hydrocarbon chemistry and NO chemistry.

2.3.1.5 Reactions Involving Chlorine

The potential for significant stratospheric ozone depletion through reactions with odd chlorine has been clearly indicated by many groups including the National Academy of Sciences [67,68]. The basic source of stratospheric odd chlorine is through photodissociation of chlorine containing compounds transported upward from the troposphere. These compounds may have natural origins such as some of the CCl₄, CH₃Cl and HCL or they may be completely anthropogenic as in the case of CFCl₃(F-11), CF₂Cl₂ (F-12) and the other chlorofluorocarbons. At the present time it is felt that the most prevalent of these compounds in the atmosphere is CH₃Cl since its natural concentration is relatively high [75]. However the most controversial compounds are the chlorofluorocarbons due to the enormous economic impact associated with their utilization. Considerable doubt still remains over the question of the possibility of tropospheric-sinks for these compounds [114].

The basic mechanism for production of atomic chlorine in the stratosphere is photolytic with the principal reaction at the present time being,

$$CH_3C1 + h\nu (\lambda < 0.22 \mu m) \longrightarrow CH_3 + C1$$
 (33)

The amount of atomic chlorine produced in this reaction is probably equal to that produced in the three following reactions [75].

$$CC1_4 + h\nu \longrightarrow CC1_3 + C1$$
 (34)

$$CFCl_3 + h\nu \ (\lambda < 0.226 \ \mu m) \longrightarrow CFCl_2^+ + C1$$
 (35)

$$CF_2CL_2 + h\nu \ (\lambda < 0.215 \ \mu m) \longrightarrow CF_2C1^+ + C1$$
 (36)

The dissociation of HCl is also of lesser magnitude,

HC1 + hv (
$$\lambda$$
< 0.22 μ m) \longrightarrow H + C1

Thus methyl chloride is the most important overall source of stratospheric chlorine at the present.

Once formed atomic chlorine rapidly enters into catalytic chain decomposition of ozone in a two step reaction,

$$c1 + o_3 \longrightarrow c10 + o_2 \tag{38}$$

$$c10 + 0 \longrightarrow c1 + 0_{2} \tag{39}$$

The atomic chlorine may also be regenerated through reactions such as,

$$c10 + NO \longrightarrow c1 + NO_2 \qquad (40)$$

$$C10 + hv \longrightarrow C1 + 0 \tag{41}$$

The principal sink for stratospheric chlorine is conversion into inactive hydrogen chloride followed by downward transport to the troposphere and rainout. The conversion into HCl generally occurs by reaction of chlorine monoxide with HO or possibly methane. For example,

$$C1 + HO_2 \longrightarrow HC1 + O_2 \tag{42}$$

$$c1 + cH_4 \longrightarrow Hc1 + cH_3$$
 (43)

Recent reports [67,68] indicate that chlorine nitrate may serve as a passive chemical (temporary sink) for both Cl and NO $_{\rm x}$. This is due to its apparent non-reaction with ozone and slow reaction with atomic oxygen. This slowly occurring cycle has three principal reactions [68],

$$c10 + NO_2 + M \longrightarrow c10NO_2 + M \tag{44}$$

$$clono_2 + hv \longrightarrow clo + No_2 . (45)$$

$$clono_2 + o \longrightarrow clo + No_3$$
 (46)

The role of chlorine nitrate in preventing ozone destruction has been studied [67,68] and found to reduce ozone depletion significantly. It concluded by the National Academy of Sciences [68] that, "unless there is some presently unknown process that quickly returns ClONO2 into active ClO_x and NO_x species, the effect of chlorine nitrate formation is to decrease the projected ozone reductions by about a factor of 1.85 compared with the values calculated for the CFMs without this reaction."

2.3.1.6 Other Halogen Reactions

The possibility that bromine or fluorine compounds analogous to the chlorine compounds discussed in the previous section could play a role in ozone destruction has been studied [68]. Diametrically opposed results were reported for the two classes of chemicals.

The reactions for production of atomic bromine are very rapid. However the sink reactions for production of HBr are largely ineffective in producing this acid. Thus, an equivalent quantity of Br is much more effective than Cl in catalytic ozone destruction. It is fortunate that at the present time natural and anthropogenic releases of bromine compounds such as the fumigant CH₃Br appear to be an order of magnitude less than chlorine compounds.

In contrast to the above it appears that although stratospheric production of atomic fluorine and catalytic destruction of ozone is possible, the sink reactions for production of HF are extremely rapid and dominate this cycle. Thus $FO_{\mathbf{x}}$ is unlikely to be a major factor in ozone chemistry.

2.3.2 Aerosol Chemistry

Of equal importance to stratospheric processes as the ozone chemistry cycle is the aerosol system which exists in the stratosphere. An aerosol is by definition [115], "a colloidal system in which the dispersed phase is composed of either solid or liquid particles, and in which the dispersion medium is some gas, usually air." This aerosol which varies significantly over time and space greatly influences stratospheric radiative transfer in both the incoming solar and outgoing infrared ranges. Changes in the nominal opacity of the stratosphere, which is about 2 percent [4], effect the passage of sunlight and heat to an extent which could result in variations in global mean temperature, general circulation and precipitation patterns. These climatic impacts are connected to a whole range of social, political and economic interdependence amongnations [100].

Stratospheric aerosols have their origins from both natural and anthropogenic sources near the earth's surface and possibly somewhat from extraterrestial sources. In recent years the presence of aircraft and rockets in the stratosphere has also made a contribution.

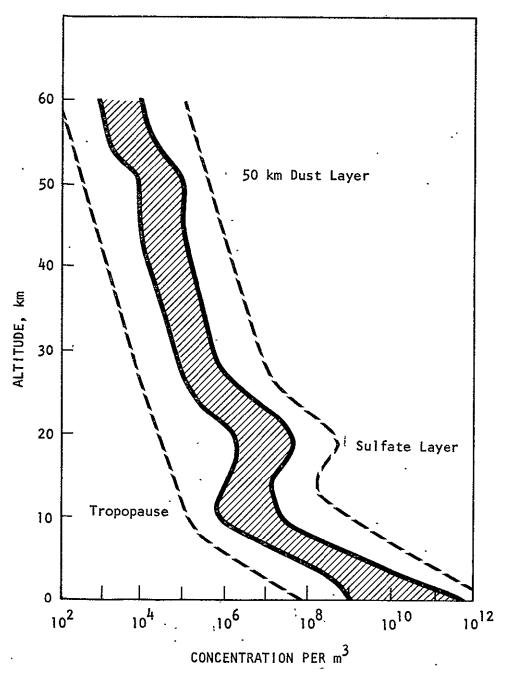
Information on the global distribution of aerosols which varies considerably in time and space is given in volume III of this report.

In the present discussion, only a brief summary of the vertical distribution is covered as a background for presentation of stratospheric aerosol chemistry.

In general the concentration of stratospheric aerosols shows a decrease with height. A depiction of this distribution is given in figure 2-2. The increase in concentration at an altitude of about 20km is called the Junge or sulfate layer. These aerosols are thought to be formed principally through the oxidation of hydrogen sulfide or sulfur dioxide which has been transported upward from the troposphere. The concentration of the aerosols can increase by as much as a factor of 10 to 100 after a volcanic eruption [66]. Study efforts during the Climatic Impact Assessment Program [4] have shown that the concentration of aerosols tends to decrease with time and the size distribution shifts toward smaller sizes between major volcanic eruptions.

At the 50km level another maximum is observed sporadically. These aerosols may be of extra-terrestial origin. In any event, their concentration is quite small and has a negligible effect on the total stratospheric system.

As stated above the principal mechanism for stratospheric aerosol formation appears to be oxidation of sulfur type gases and conversion to sulfuric acid or sulfate aerosols. In the following



MOST FREQUENT CONCENTRATIONS

---- EXTREME VALUES

FIGURE 2-2
VERTICAL DISTRIBUTION OF AEROSOLS [66]

sections the discussion of aerosol chemistry will be divided into the sulfuric acid/sulfate cycle and other known mechanisms.

2.3.2.1 <u>Sulfate Aerosol Chemistry</u>. The mechanisms for conversion of sulfur dioxide to sulfuric acid and sulfates are not well understood even for tropospheric situations [116]. The most important tropospheric sulfuric and sulfate formations identified to date [116] are shown in table 2-IV. Of the five mechanisms shown only the first two are likely to occur in the stratosphere. The lack of liquid water in the stratosphere makes mechanisms 3 and 4 unlikely. In addition, except for the possibility of carbon particles in aircraft or rocket exhaust, mechanism 5 can be neglected.

However, the probability of occurrence of mechanisms 1 and 2 in the stratosphere is increased by the ambient conditions, principally the availability of atomic oxygen, HO_X and NO_X. The basic mechanisms for direct photooxidation involves the availability of atomic_oxygen from a photochemical reaction such as (1) or (3a, b or c),

$$so_2 + o + M \longrightarrow so_3 + M \tag{47}$$

which is followed by,

$$so_3 + nH_2O \longrightarrow H_2SO_4 \cdot (n-1)H_2O$$
 (48)

If HO enters the reaction rather than atomic oxygen the procedure can be as follows,

$$SO_2 + HO_2 + h\nu (\lambda:0.24 - 0.34 \mu m) \longrightarrow SO_3 + HO$$
 (49) which then proceeds as in reaction (48).

TABLE 2-IV

TROPOSPHERIC MECHANISMS FOR CONVERTING SULFUR DIOXIDE
TO
SULFURIC ACID AND SULFATES[116]

	MECHANISM	FACTORS CONTROLLING REACTION
1.	Direct photo-oxidation	Sulfur dioxide concentration, sunlight intensity
2.	Indirect photo-oxidation	Sulfur dioxide concentration, organic oxidant concentration, HO or NO concentration
3.	Air oxidation in liquid droplets	Ammonia concentration, liquid water
4.	Catalyzed oxidation in liquid droplets	Concentration of heavy metal (iron, world on, manganese) ions, liquid water
5.	Catalyzed oxidation on dry surfaces	Carbon/particle concentration

Laboratory studies [117] have indicated the importance of ammonia in the growth of stratospheric aerosols. These studies concluded that a significant portion of stratospheric aerosols are composed of ammonium sulfate, ammonium bisulfate or possibly ammonium persulfate. A mechanism for the conversion to sulfate could be

$$NH_3 + H_2SO_4 \cdot nH_2O \longrightarrow NH_4HSO_4 \cdot nH_2O$$
 (50)

followed by

$$NH_3 + NH_4HSO_4 \cdot nH_2O \longrightarrow (NH_4)_2SO_4 \cdot nH_2O$$
 (51)

As soon as the ammonium sulfate salt is formed, it provides a medium for rapid catalytic oxidation of more SO_2 ,

$$2SO_2 + 2H_2O + O_2 \xrightarrow{NH_4 \text{ salt}} 2H_2SO_4$$
 (52)

In summary the three main mechanisms for sulfate aerosol production are,

- Photolysis of 0_3 to produce 0 atoms which oxidize 0.50^{-1} into sulfuric acid,
- · Neutralization of sulfuric acid by ammonia and
- Rapid oxidation of additional SO₂ to sulfate through the .
 catalytic action of ammonium ions

2.3.2.2 Other Aerosol Mechanisms

Some of the non-sulfate mechanisms for production of stratospheric aerosols wee indicated briefly in the introduction to this section. A listing of these and others indicates six possible pathways,

- . Direct transport of particles from the troposphere
- Anthropogenic release in situ
- · Conversion of non-sulfur stratospheric trace gases
- Extraterrestial matter
- Non-thunderstorm associated water or ice clouds
- Molecular clusters

Each of the above could contribute to the aerosol concentration in the stratosphere but none has shown any evidence of approaching the magnitude and scope of the sulfuric acid/sulfate aerosols.

The presence of stratospheric aerosols which have been transported upward from the troposphere is well substantiated by several factors. The first is the great increase in aerosols after volcanic activity. These eruptions probably contribute most of the basalt (aluminum, calcium, magnesium) type species which have been detected. The second factor indicating tropospheric to stratospheric transport is the relatively continuous decrease in concentration of aerosols with height and a third factor is the marked bimodal size distribution of the aerosols. It has been concluded from CIAP studies [4] that aerosols formed from gases in the lower stratosphere have a much larger mean size than those of tropospheric origin.

Anthropogenic releases of aerosols in the stratosphere occur from aircraft and rocket exhausts. Results of the Department of Transportation's High Altitude Pollution Program study [113] show

that current and projected aircraft operations in the lower stratosphere emit significant amounts of water, soot [carbon particles],
lubricating oils and trace elements such as the heavy metals found
in many crude oils. In addition to these releases, operation of
propulsion systems such as the Space Shuttle booster could release a
different set of chemicals such as aluminum oxides and hydrogen
chloride.

In addition to the formation of sulfuric acid aerosols from sulfur dioxide and other sulfur gases, a lesser amount of aerosols appear to be formed during the NO $_{\rm x}$ and ClO $_{\rm x}$ cycles discussed previously in the ozone chemistry section. These aerosols serve as a sink for NO $_{\rm x}$ and chlorine in ozone chemistry. The main species appear to be nitric acid and hydrochloric acid. Other possible species include nitrous acid, nitrate and nitrite salts, chloride salts, hydrobromic and hydrofluoric acids, and their related halogen salts.

Non-thunderstorm associated water clouds are rare in the stratosphere. Two general types occur occasionally [115,118] from natural
causes while aircraft contrails which transform into cirrus clouds
have an anthropogenic origin. The first natural type called nacreous
clouds occurs in polar regions at altitudes in the 21 to 30 Km range.
They are presumed to be composed of supercooled water or ice. The
second natural type called noctilucent clouds also occurs in polar
regions at altitudes between 75 and 90km. These clouds may be

composed of extra-terrestial dust possibly in combination with ice.

Neither of these natural clouds has a significant impact on the stratospheric aerosol system. However, it has been shown that upper tropospheric and stratospheric aircraft contrails can cause a measurable change in radiative transfer.

During the CIAP study period [4] the possibility or speculation was indicated that certain gas molecules showed an affinity which caused small numbers of them to cluster together. These clusters are presumed to be much too small to be considered true aerosols but they should exhibit some non-gaseous characteristics. Among the gases speculated to cluster are sulfur dioxide, nitric oxide, molecular nitrogen, water vapor and carbon dioxide. The significance of clustering has not been theorized. This subject would require extensive investigation which does not appear to be warranted at present.

3.0 USER REQUIREMENTS FOR OZONE AND AEROSOL RELATED STRATOSPHERIC MEASUREMENTS

In volume I of this report an extensive discussion of user requirements for stratospheric measurements is presented. This discussion will be summarized here with emphasis placed on the distinction between the ozone problem which has changes in UV transmission as its principal effect and the aerosol problem which effects radiative transfer and hence global climate.

3.1 Identification of the Users

The uses of both ozone and aerosol related data may be grouped into three major categories, namely,

- understanding stratospheric chemistry and physics
- the effects of stratospheric change on the biosphere
- · monitoring for regulatory activities and long-term trends.

The first function shown focuses on analysis and understanding of the details of the various stratospheric cycles and is obviously the most demanding in terms of the magnitude and scope of information required.

The second function focuses primarily on the results of stratospheric change namely changes in UV transmission and climate and how
they affect the earth's environment. Emphasis is not placed on the
particular cause of the change but rather on what results from the
change. Except for a few special cases such as direct intrusion of
ozone into high flying aircraft, in situ particulars are not of
concern to these users.

The third function of monitoring and regulation is of interest to a wide spectrum of users extending from governmental agencies to manufacturers or users of known or potentially hazardous materials. For example, an understandably large number of industrial organizations are known to be supporting studies of the chlorofluorocarbon/ozone problem. These users, both governmental and industrial, are concerned not only with the details of how ozone and aerosols alter the UV and radiative transfer but also with the effects of such alteration. The end goal of both groups is the formulation of equitable control and regulation when indicated.

3.2 Measurement Requirements for UV Studies

The predominant interest in the field of UV is in the interaction of biological systems with UV. As a result, it appears that this field is dominated by users far removed from effective utilization of observations which describe the state or variability of those features of the stratosphere which control UV transmission. However, at the same time, considerable interest has developed among scientific and regulatory groups in measurements of particular properties and species which will help clarify the physical and chemical processes which control the UV environment.

Specific measurement requirements are not generally clear since the largest potential group of users is not interested in detailed physics or chemistry of the stratosphere but rather with responses of biological systems to changes in the environment. However, some general comment can be made. First, it is clear that a topic of primary concern is the intensity and wavelength distribution of ultraviolet radiation at the Earth's surface. Inference of this data from spacecraft measurements provides a unique opportunity to supplement present observations and provide more comprehensive coverage in space and time.

Secondary studies would include determination of the variability of radiation features, studies of the influence of polluting gases on the atmospheric transmission in the UV spectrum and data which relates the UV environment to biological variability.

Direct NASA contact with those studying the subject will begin the communication cycle so necessary if experiments are to be developed which satisfy these users.

3.3 Measurement Requirements for Climate Studies

Climate effects are much more pervasive than those defined previously for UV. In the case of UV, the chain of concern is traceable from the stratosphere directly to the well-defined set of users, both direct and indirect. For climate, the end point of such a consequence chain is much more diffuse. For this reason, the user community will be restricted to the primary users of remotely sensed data with the understanding that concern with climate is almost limitless.

The primary users of climatic data consist of those who are studying the physical processes of climate either through modeling or other approaches and those whose concern is monitoring the climate for indications of significant change. The total system which comprises the Earth's climate is extremely complex and highly interrelated. Stratospheric aerosols play only a small (but by no means insignificant) role in the total climate. However, in many cases perturbations in one or more of these smaller components of the climate system can have far reaching consequences throughout the world. With the full realization of the complexity of climate, it is still desirable to separate the system into various component processes,

- radiation
- water clouds
- surface
- atmospheric

Within these processes a set of parameters may be identified as being required for adequate characterization and monitoring of climate and in particular climate change. NASA [119] has proposed such a set of 40 parameters covering the full span of climatic component processes.

Those parameters involving atmospheric composition include:

- solar ultraviolet flux
- stratospheric aerosol optical depth
- tropospheric aerosol optical depth
- ozone
- stratospheric H₂0
- nitrous oxide and nitrogen oxides
- carbon dioxide

- chlorofluoromethanes
- methane

As can be seen in the above list many of the parameters also play key roles in the ozone chemistry cycle. Their roles in climate generally focus on their contribution to the radiative heat balance of the atmosphere. However, as stated earlier in this section this discussion is limited to aerosol related processes. The role of stratospheric gases in relation to climate has been discussed in Volume I.

Table 3-I shows NASA's proposed measurement requirements for stratospheric aerosol optical depth. In conjunction with the presentation of these requirements they indicate the need for satellite remote sensing as a key part of stratospheric aerosol monitoring.

The WMO-ICSU [47] has developed a set of requirements for the Global Atmospheric Research Program (GARP). Those requirements pertaining to atmospheric aerosol processes are shown in table 3-II. It is noteworthy that under the monitoring section the variables of concern are not only the number and mass of aerosols but also the concentration of precursor gases. It is unfortunate however that this and other studies do not provide sufficient detail to present the specifics of these precursor monitoring requirements. This again can be understood from the fact that most users are concerned with the effects of stratospheric aerosols rather than the detail chemical and physical processes.

TABLE 3-I

NASA'S PROPOSED MEASUREMENT REQUIREMENTS
FOR
STRATOSPHERIC AEROSOL OPTICAL DEPTH[119]

Desired accuracy	0.002
Spatial resolution	·
North-South	· 250 km
East-West	1000 km
Vertical	3 km
Temporal resolution	1 month

TABLE 3-II ROSOL PROCESSES-SUMMARY OF TENTATIVE OBSERVATIONAL REQUIREMENTS[47]

-I.	STUI	OY OF PROCESSES		
	a)	Radiative effects of aerosols.	•	
•		Required aerosol parameter for troposphere and stratosphere	Observational require- ment and accuracy	
	-	Size distribution		
		dn dr in cm -4 STP	. 5%	
		Vertical profile of size distribution	5% Required vertical reso- lution generally 0.5 to 1.0 kilometer	
		Real refractive index of bulk material n	1% over the range $1.0 \le n \le 2$	
_	•	Imaginary part of the re- fractive index k	10% over the range 0.001 < k < 0.1	
-		Bulk density δ of aerosol particles, in g cm ⁻³	5% over the range $1.0 \le \delta < 3.0$	
		Solubility of aerosol particles and/or growth characteristic with relative humidity	Use of 3 to 4 typical growth curves	
		For necessary data to calculate energy balance of the atmosphere		
	b)	Aerosol cloud interaction	Cannot be specified at this time.	
tr.	MON	ITORING Space	Time Accuracy	
		Variables to be monitored Resolution	Resolution	
	1) 2) 3) (4)	Total number concentration Concentration of optically important particles Total mass concentration Concentration of gaseous precursors about 20 baseline stations distributed over the globe	daily} 5%	

4.0 SCIENCE REQUIREMENTS

A number of recent major study groups (CIAP, GARP, NAS, etc.) and many smaller ones have addressed the general question of man's interaction with and impact on the stratosphere. A few of these groups have studied the entire stratospheric problem but in general the studies focus on one or more specific aspects of the stratospheric such as ozone, nitrogen oxides, chlorofluorocarbons or aerosols. In this section the results of these efforts are summarized and used to develop two sets of scientific requirements for stratospheric trace constituent measurements. The first set is developed on the assumption that user concern is only with the depletion of stratosphere ozone. The second assumes that user concern is only with stratospheric aerosols and associated chemistry.

In Volume I a listing is presented of those stratospheric measurements which would be desirable. The listing is presented without regard to involvement of any species or property in any specific chemical chain such as the ozone chain, the aerosol chain or others. Table 4-I presents the properties and gases from this list and indicates in capsule form the major or typical reasons why the measurement would be appropriate in a specific study of ozone or aerosols. If the term no direct involvement (NDI) is applied to any measurement it does not imply that no chain of reactions or events exists which would link the measurement with ozone or aerosol chemistry. It is intended to indicate any possible link is at least relatively remote.

TABLE 4-I CAPSULE SUMMARY OF THE PARTICIPATION OF VARIOUS PROPERTIES AND GASES IN OZONE AND AEROSOL CHEMISTRY

1			
	PARTICIPATION		
: PROPERTY OR GAS	OZONE CHEMISTRY	AEROSOL CHEMISTRY	
Stratospheric Temperature	Changes with changes in ozone chemistry cycles	Change with aerosol changes	
Solar Irradiance (inc UV)	UV-B increases with decreasing ozone	Radiation transmission affected by aerosol concentration	
Earth Radiance	Many species associated with ozone absorb IR	Increase in aerosols create "greenhouse"	
, Water Vapor	Involved in production of H ₂ and OH in upper stratosphere	Involved in H ₂ SO ₄ ·nH ₂ O formation	
Ozone, 0 ₃	a priori	Involved in oxidation of SO_2 and H_2S , through O and HO_2	
Aerosols	NDI*	a priori	
Carbon Dioxide, CO ₂	Sink for CO through reaction with HO	NDI, except possible CO ₂ cluster formation	
Hydroxy1, HO	Involved in 0x, NOx, ClOx and CH4 chemistry	Involved in SO chemistry	
Atomic Oxygen, O(3P) or O(1D)	Involved in O _x , HO _x , NO _x , and ClO _x chemistry	Involved in basic oxidation of SO ₂	
Ammonia, NH ₃	Transport from troposphere and reaction with HO leads to NO _x chain	Neutralizes H ₂ SO ₄ to produce (NH ₄) ₂ and related compounds	
Nitric Oxide, NO	NO chemistry	NDI	
Nitrogen Dioxide, NO 2	NO chemistry	NDI	
Atomic Chlorine, C1	Clo chemistry, catalytic ozone depletion	Formation of HCl and metal chlorides	
Chlorine Monoxide, C10	ClO _x chemistry	Formation of HC1 and metal chlorides	
Hydrogen, H ₂ or H	$0_{_{ m X}}$ and ${ m HO}_{_{ m X}}$ chemistry, HCl equilibrium	Indirectly through HCl and metal chlorides	
Hydroperoxyl, HO ₂	O _x , HO _x , NO _x and C1 chemistry	Oxidation of SO ₂	
Nitrous Oxide, N ₂ O	NO _x chemistry	NDI.	
Nitrogen Pentoxide, N205	NO _x chemistry	NDI .	
Nitric Acid Vapor, HNO ₃	NO _X chemistry	HNO ₃ aerosols and nitrates	
Chlorine Nitrate, C10NO ₂	Reduces 03 depletion; quasi sink for Cl and NO in lower stratosphere	NDI	
Carbon Monoxide, CO	Involved in HO chemistry	NDI	
Methane, CH ₄	Reacts with O to produce HO and CH ₃ radicals; also reacts with Cl to produce HCl and CH ₃	NDI	

	PARTICIPATION	
PROPERTY OR GAS	OZONE CHEMISTRY	AEROSOL CHEMISTRY
Hydrogen Chloride Gas, HCl Trichlorofluoromethane, F-11, CFCl ₃	C1 chemistry C1 chemistry, reacts with O and UV	Chloride aerosol formation
Dichlorodifluoromethane, F-12, CF ₂ Cl ₂	C1 chemistry, reacts with 0 and UV	NDI
, Sulfur Dioxide, SO ₂	Depletes 0, HO, HO ₂ etc., by oxidation to SO ₃	Principal source for H ₂ SO ₄ and sulfates
Tetrachloromethane, CCl ₄	C1 chemistry, reacts with 0 and UV	NDI
Chloromethane, CH ₃ Cl	C1 chemistry, reacts with O and UV	NDI
Dichloromethane, CH ₂ Cl ₂	C1 chemistry, reacts with O and UV	NDI
Trichloromethane, CHCl3	Cl chemistry, reacts with O and UV	NDI
Formaldehyde, CH ₂ O	HO chemistry, reacts with UV to form H	NDI
Chlorodifluoromethane, F-22, CHC1F ₂	Cl chemistry, reacts with O and UV	NDI
Dichlorofluoromethane, F-21, CHCl ₂ F	Cl chemistry, reacts with O and UV	NDI
Methyl Bromide, CH ₃ Br	Br chemistry	NDI
Trichloroethylene, CCl2 CHC1	Cl chemistry, reacts with 0 and UV '	NDI
Methylchloroform, CH ₃ CCl ₃	Cl chemistry, reacts with O and UV	NDI .
Trichlorotrifluoroethane, F-113, CF2C1CFC12	Cl chemistry, reacts with O and UV	NDI .
Carbonyl Fluoride, F ₂ C0	Halogen chemistry	NDI
Fluoroformyl Chloride, C1FCO	Halogen chemistry	NDI
Tetrabromomethane, CBr ₄	Br chemistry	NDI .
Methyl Peroxy Radical, CH ₃ O ₂	CH ₄ and NO _x chemistry	NDI .
Methyloxy Radical, CH ₃ 0	CH ₄ and NO _x chemistry	NDI .
Chlorodifluoromethane Radical, CF ₂ C1 ⁺	Halogen chemistry	NDI
Dichlorofluoromethane Radical, CFC1+	Halogen chemistry	NDI .
Chlorine Dioxide, ClO ₂	Clo _x chemistry	NDI .
Methyl Sulfide, (CH ₃) ₂ S	NDI	Sulfur and SO _x chemistry
Carbonyl Sulfide, COS	NDI .	Sulfur and SO _x chemistry

TABLE 4-I (CONCLUDED)

	PARTICIPATION		
PROPERTY OR GAS	OZONE CHEMISTRY	Aerosol chémístry	
Carbon Disulfide, CS ₂	NDI ·	Sulfur and SO _x chemistry	
Dichloroethane, C ₂ H ₄ Cl ₂	C1 chemistry	NDI .	
Chloroethane, C ₂ H ₅ Cl	Cl chemistry	NDI · ·	
Carbonyl Monochloride, COC1	C1 chemistry .	NDI	
Tetrachloroethene, Cl ₂ C:CCl ₂	.Cl chemistry	NDI	
Vinyl Chloride, CH ₂ :CHC1	C1 chemistry	NDI , .	
Hydrogen Sulfide, H ₂ S	NDI	Sulfur and SO _x chemistry	
Hydrogen Fluoride, HF	Halogen chemistry	HF aerosol formation	
Hydrogen Bromide, HBr	Halogen chemistry .	HBr and bromide aerosol formation	
Hydrogen Peroxide, H ₂ O ₂	·HO _x chemistry	HO _x /SO _x chemistry	
Ammonium Ion, NH ₄ Sulfur Hexafluoride, SF ₆ Sulfur Trioxide, SO ₃	Possible involvement with NH3/OH reactions	Directly involved in (NH ₄) ₂ SO ₄ production	
Sulfur Hexafluoride, SF ₆	NDI, tracer in dispersion studies	NDI, tracer in dispersion studies	
Sulfur Trioxide, SO3	Product of HO _x /SO ₂ reaction	Precursor of H ₂ SO ₄ serosol	
Sulfur Hexafluoride, SF ₆ Sulfur Trioxide, SO ₃ Bisulfite Radical, HSO ₃ Nitrogen Trioxide, NO ₃	Involved in possible depletion of HO	Possible precursor in H ₂ SO ₄ and SO ₄ . aerosol formation	
Nitrogen Trioxide, NO3	NO chemistry	Precursor for HNO3 and NO3 aerosols	
Bromine Oxide, BrO	Br chemistry, analogous to C10	Precursor for HBr and Br aerosols	
Atomic Bromine, Br	Br chemistry	Precursor for HBr and Br aerosols	
Atomic Oxygen, O(¹ S).	Possibly involved in O chemistry	NDI	
Molecular Oxygen, O ₂ (¹ Δ)	Possible involved in O _x chemistry	NDI	
Non-Methane Hydrocarbon, C H	Possible reactions analogous to CH ₄	NDI	
Various Organics, H _x C _y O _z	Possible reactions analogous to CH ₂ O	NDI	

Specific aerosols have not been included in the table since they are a priori involved in aerosol chemistry and as a group have no direct involvement in ozone chemistry. This is due to the quasi irreversable nature of aerosol formation. The only known sinks for aerosols are gravitational settling into the troposphere or possible washout in stratospheric thunderclouds.

4.1 Development of Scientific Criteria

The scientific criteria developed for stratospheric pollution measurements must have as their basis the major objectives of the entire stratospheric program. These objectives may be primary or secondary depending upon the nature of their interaction with man and his environment. The primary objectives are:

- Monitoring climatic changes caused by changes in the concentrations of the various stratospheric trace constituents, particularly aerosols; and,
- Monitoring changes in ultraviolet received at the earth's surface as a result of changes in the concentrations of the various stratospheric trace constituents, particularly ozone.

The secondary objectives may be considered as indirect objectives of the entire program. These are:

- Increased understanding of the chemistry and physics of the stratosphere and its constituents; and,
- Increased understanding of the meteorology and hydrodynamics of the stratosphere.

Obviously, there is considerable overlap between the primary and secondary objectives, since the latter have a much broader scope which includes the former. The following sections present discussions supporting the prioritization of ozone and aerosol related measurements into the various groupings shown.

4.2. Prioritization of Ozone Related Measurements

The list of stratospheric ozone related measurements has been presented in four groups which are considered to be of descending order of importance in terms of the absolute need for the measurement without regard to present knowledge or measurement capability. However, it must be emphasized at this point that none of these groups is considered unimportant. The groupings merely show the degree of importance, and relative placement within a group has no significance.

The rationale for placement of a required measurement in any one of the categories is given below:

Ozone Group 1. This group contains those properties and species which are considered to be directly related to changes in the untraviolet flux. For example, ozone is directly related to the major absorption of ultraviolet while the freon compounds are not.

Ozone Group 2. In this group are listed the components of the basic reactions involved in the direct production or depletion of the ozone concentration in the stratosphere. These species participate in the principal chemical equations which directly involve ozone. These equations are given below for each of the significant chemistry

chains. The numbers in parentheses are the equation numbers as used previously in section 2.

Pure oxygen reactions:

$$0 + 0_2 + M \longrightarrow 0_3 + M \tag{2}$$

$$0_3 + h\nu (\lambda:0:45-0.65 \mu m) \longrightarrow 0 + 0_2$$
 (3a)

$$0_3 + hv (\lambda:0.31-0.34 \mu m) \longrightarrow 0_2(^1\Delta) + 0 (^3P)$$
 (3b)

$$0_3 + h\nu \ (\lambda < 0.31 \ \mu m) \longrightarrow 0 \ (^1D) + 0_2$$
 (3c)

$$0_3 + 0 \longrightarrow 0_2 + 0_2$$
 (4)

Hydrogen-oxygen reactions:

$$H + O_3 \longrightarrow HO + O_2 \tag{7}$$

$$HO_2 + O_3 - HO + 2O_2$$
 (10a)

$$HO + O_3 \longrightarrow HO_2 + O_2 \tag{10b}$$

Nitrogen-oxygen reactions:

$$O_3 + NO \longrightarrow NO_2 + O_2 \tag{25}$$

Chlorine-oxygen reactions:

$$c1 + o_3 \longrightarrow c10 + o_2$$
 (38)

Ozone Group 3. This group contains those species considered to be the most important ones in the indirect chemistry chains; that is, those which result in the production or depletion of the major species discussed under Group 2.

Ozone Group 4. This group contains those species considered to be involved in a lesser but nor unimportant way on the indirect chemistry chains discussed above.

Table 4-II presents the list of ozone related stratospheric measurements which should be made or would be of scientific interest. The measurements are grouped according to the criteria discussed above and placement was made after analyzing all available references that discuss the importance of the various species. The list also shows the major references supporting the selection of the measurement and its placement in the appropriate group. A number of other references [49, 57-61, 74-103, 113-120] were also consulted during preparation of the list.

4.3 Prioritization of Aerosol Related Measurements

The list of stratospheric aerosol related measurements has been presented in six groups which are considered to be of descending order of importance in terms of the absolute need for the measurement without regard to present knowledge or measurement capability. However, it must be emphasized at this point that <u>none</u> of these groups is considered totally unimportant. The groupings merely show the degree of importance, and relative placement within a group has no significance.

The rationale for placement of a required measurement in any one of the categories is given below:

Aerosol Group 1. This group contains those properties and species which are considered to be directly related to changes in climate and/or the stratospheric aerosol content. The group contains both stratospheric properties such as temperature and radiative flux and species such as total aerosols and sulfate aerosols.

TABLE 4-II
PRIORITIZED LIST OF DESIRED STRATOSPHERIC OZONE RELATED MEASUREMENTS

NAME OF SPECIES/PROPERTY AND					LAM	OR RE	FEREN(CES WI	IERE C	ITED	·				7
SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73	1
OZONE GROUP 1, MEASUREMENTS OF PROPERTIES AND SPECIES DIRECTLY ASSO- CLATED WITH UV CHANGE															
Temperature .	✓	√	✓	✓	✓	✓				✓		V	V	✓	
Solar Irradiance (including UV)	V	✓	√	✓	√	√		V	V	√	√	√	V	✓	1.
Earth Radiance	1	√	✓		1	✓				✓			√],
Ozone, O ₃	V	✓	✓	1	~	✓		,	₹	V	V	√	✓	√	
OZONE GROUP 2, COMPONENTS OF THE BASIC REACTIONS INVOLVED IN THE DIRECT PRODUCTION OR DEPLETION OF OZONE					,										ľ
Hydroxy1, HO	✓	✓	√	<u> </u>	V	√	✓	· 🗸	√	✓ '		√	√	V	
Atomic Oxygen, O(³ P)	✓		✓	 √	√	✓	√	✓	V	√		√	√	√	
Atomic Oxygen, O(¹ D)	✓		✓	 √	√	1	√	√	√	√		√	√	√	
Nitric Oxide, NO	✓	√	√	√	✓	√		✓	✓	√	✓	√	√	√	
Nitrogen Dioxide, NO ₂	✓.	√	✓	√	✓	✓		√	√ :	√	✓	¥	√	√	
Atomic Chlorine, Cl	\ \		✓	√	✓			√	✓	√		√	√	✓	
Chlorine Monoxide, ClÓ	✓	✓	√	4	√	√		√	✓	√		✓	√.	V	
Hydrogen, H ₂ or H	✓	✓	√	V		✓		√	V	√ .		√		√	
Hydroperoxyl, HO ₂	✓		√;	✓		✓	✓	√	√	√		√		· •/	
,			,							-		•		,	

TABLE 4-II (Continued)

NAME OF SPECIES/PROPERTY		 			LAM	OR REI	ERENC	es wh	ERE CI	TED		-		
AND SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
OZONE GROUP 3, MAJOR COMPONENTS OF THE BASIC REACTIONS INDIRECTLY INVOLVED IN THE PRODUCTION OR DEPLETION OF OZONE								, ,		,			÷	
Water Vapor, H ₂ O	✓	√	✓	-	✓	√.	✓	,	√	√		V	V	v
Ammonia, NH ₃	·	√			√	√	✓		.*	•		√		
Nitrous Oxide, N ₂ O .	✓	✓	√	✓	✓		✓	✓	✓	✓	✓	✓	√.	✓
Nitrogen Pentoxide, N ₂ O ₅	√		√	✓	✓	✓	1	√					✓	✓
Nitric Acid Vapor, HNO3	✓	✓	√	V	✓	✓		V	√	✓	✓	V	✓	✓
Chlorine Nitrate, Clono	ļ		<u> </u>					V	. 🗸			✓		✓
Carbon Monoxide, CO		✓	✓	✓		✓				✓		✓	√]	✓
Methane, CH ₄	✓	✓	✓	✓	✓	✓		✓	V	√	√	✓	V	✓
Hydrogen Chloride Gas, HCl	✓	✓	✓	✓	✓			✓	✓	✓	,	. 🗸	✓.	✓
Trichlorofluoromethane, F-11, CFC1	✓	✓	✓	✓	✓			✓	✓	✓		\	V	✓
Dichlorodifluoromethane, F-12, CF ₂ Cl ₂	✓	✓	✓	✓	✓	'		✓	✓	✓		-/	✓	✓
Sulfur Dioxide, SO ₂	V		✓		✓	✓	√			1		· 🗸	✓	
OZONE GROUP 4, OTHER SIGNIFICANT COMPONENTS OF THE OZONE CHEMISTRY CHAINS					١.									
Carbon Dioxide, CO ₂	✓	✓	√					✓		√	_:	✓	√	✓
Tetrachloromethane, CCl ₄ , (carbon tetrachloride)	✓		:	~				√	~	~		✓		. ~

TABLE 4-II (Continued)

NAME OF SPECIES/PROPERTY AND					LAM	OR REI	FERENC	es Wi	ERE C	ITED				
· SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
OZONE GROUP 4 (Continued)						!		Ì			:		:	
Chloromethane, CH ₃ Cl, (Methyl Chloride)		;			✓								٠	
Dichloromethane, CH ₂ Cl ₂ , (Methyl Dichloride)	:	1	:	+	✓	+	-				<u>.</u> .		::	-
Trichloromethane, CHCl ₃ (Chloroform)	<u>.</u>			:	V	1	1					ŗ.	'	;
Methanal, CH ₂ O, (Formaldehyde)	<u>:</u>	}	√	,	ļ ·	!	4.		1 :	:	1	· 	<u>_</u>	
Chlorodifluoromethane, F-22, CHC1F2		1	. '	-		+		✓	√.	-;		;	İ	-
Dichlorofluoromethane, F-21, CHC1 ₂ F			1	-	- .	:		✓			1:	ļ ;	:	
Bromomethane, CH3BR, (Methyl Bromide)	√.			:	_		1.1.	✓	√	√	1 :	₹.	; ;-	
Trichloroethylene, CC1 ₂ =CHC1		-	;	;					✓	4	.: .	<u>'</u> -	+	✓
Methylchloroform, CH ₃ CC1 ₃				<u>:</u>				_	√.	√.	-			1
Trichlorotrifluoroethane, F-113, CF2C1CFC12	-	†				-		٠.	✓		1, -		1	-
Carbonyl Fluoride, F ₂ CO		;				i	-		V	√,	, ;.	√.	1	√.
Fluoroformyl Chloride, C1FCO		+		-		.4.	-	/	✓.	4			,	1
Tetrabromomethane, CBr ₄	✓				-	-	_			,	. :		1	
Methyl Peroxy Radical, CH302	✓	':		-	-								. :	
Methyl Oxy Radical, CH30	✓	,					.	-	-		-		:	
Chlorodifluoromethane Radical, CF ₂ Cl ⁺	✓	.		-						V				1 L
Dichlorofluoromethane Radical, CFC12+									-	✓		-	1	1
		<u> </u>				1						1		,

NAME OF SPECIES/PROPERTY AND	<u>, </u>				MAJ	OR RE	FEREN	CES WI	ERE C	ITED				
SYMBOI.	39	40.	62	63	64	65	66	67	68	69	70	7,1	72	T :
OZONE GROUP 4 (Continued)			· 1.			:	,			;				┪
Chlorine Dioxide, ClO ₂	✓				1 1	,								
Olchloroethane, C2H4Cl2				1	<u>'</u>									
Sthyl Chloride, C ₂ H ₅ Cl	+			. 1.	:	ì.			_	į				
Carbonyl Monochloride, COCl	1 :		,	_	<u>.</u>	,	,							
etrachloroethene, CL ₂ C:CCl ₂	1.				,			-					,	
inyl Chloride, CH ₂ :CHCl		4	,			,			·	i			'	
ydrogen Fluoride, HF			√		wie	_	,		* /	<i>'</i>		√	1	
ydrogen Bromide, HBr	!	;	*	, ,				_			1	v. √		
ydrogen Peroxide; H ₂ O ₂	√	✓	√			√		√.		ļ	*	√		
mmonium Ion, NH ₄		 ✓		. i	_	√	√		,	1		Ť	. 1	
ulfur Trioxide, SO ₃	\ \times \	· ·			-	'		;	-	,		√		
isulfite Radical, HSO3		· •	'		· ;			•		- 1		V √		
itrogen Trioxide, NO ₃	V	;		, i	`	✓	-4-		_	,		v. √	 √	
romine Oxide, BrO	` ;						,		٠,			√		
tomic Bromine, Br	_						Ţ		_			√		
tomic Oxygen, O(¹ S)	<u>:</u>		\	_			_		-	- 		·		
xygen, 0 ₂ (¹ Δ)					!									٠,
Aygen, 02(1)	-	- [٠, ا		ł			-			. [. 1	

NAME OF SPECIES/PROPERTY					MAJ	OR REI	FERENC	es Wi	ere c	ITED				
AND SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
ZONE GROUP 4 (Concluded)					,				,					
on-Methane Hydrocarbons, C _x H _y	- ,			,	+ ,				1	1	٠	V		
rious Organics, H _C C _y O _z	✓			د		-			! ->-		•	-	•	
		;			1					,	i.	; ;		
]	;	;					;	;		į		
•					,	,			- 3		!	i		
		,		i	ŧ ;		:	1	`,	, i		;	,	
S OF		,		;		·		:		٠.	:	;		
ORIGINAL PACE IS					4					<u>.</u>	٠,	,		
2 T 2					ľ									
AGE JALI												٠.		
7 8				,										
										, '	٠.			
							}				,	<u> </u>		
												:		
				<u> </u>	ļ 									

Aerosol Group 2. In this group are listed the gaseous precursor components which are involved in the basic reactions for production of either sulfuric acid or ammonium sulfate aerosols. These reactions were shown in equations (47) through (52) in Section 2.3.2.1.

Aerosol Group 3. This group contains the major specific aerosols both in the sulfate and non-sulfate category. The latter includes aerosols such as nitric acid and hydrochloric acid.

Aerosol Group 4. This group contains all other aerosols which may be of significance in the stratosphere.

Aerosol Group 5. This group contains the precursor gases for non-sulfate aerosols along with the gases indirectly associated with sulfate aerosols.

Aerosol Group 6. This group contains those species suspected of being capable of forming molecular clusters as discussed in Section 2.3.2.1.

Table 4-III presents the list of aerosol related stratospheric measurements which should be made or would be of scientific interest. The measurements are grouped according to the criteria discussed above and placement was made after analyzing all available references that discuss the importance of the various species. The list also shows the major references supporting the selection of the measurement and its placement in the appropriate group. A number of other references [49, 57-61, 74-103, 113-120] were also consulted during preparation of the list.

NAME OF SPECIES/PROPERTY					MAJ	OR REI	ERENC	es Wii	ere c	(TED				
AND SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
AEROSOL GROUP 1A, DIRECT MEASUREMENTS OF CLIMATIC CHANGE AND ULTRA-VIOLET CHANGE	 					,								
Temperature	✓	✓	✓	✓	√	√,				✓		✓	✓	√.
Solar Irradiance (including UV)	✓	✓	 ✓	✓	✓	✓		√.	. ~	√	✓	V	✓	√
Earth Radiance	✓	✓	✓		√.	✓				✓			✓	
AEROSOL GROUP 1B, MAJOR AEROSOLS														
Total Aerosols		✓	✓	√	√.	✓	√.	Ì				.√	√.	✓
Total Sulfate Aerosols	V	 ✓	 √.			✓	√				. '	V	ŀ	
Sulfuric Acid Aerosol, H2SO4.nH2O	✓	✓	✓			√	✓				,	✓		
Ammonium Sulfate, (NH ₄) ₂ SO ₄		✓	V			✓	V					✓		
AEROSOL GROUP 2, MAJOR SULFATE AEROSOL PRECURSOR GASES													ļ	
Cultim Diagrido SO Int 20	_/		~		✓	V	V				,	V	V	
Sulfur Dioxide, SO ₂ Sulfur Trioxide, SO ₃ Atomic Oxygen, O	\ \		*				✓			,,		✓		
Atomic Oxygen, O	1		√	✓	√	V	✓	✓	√	V		√	✓	✓
Water Vapor, H ₂ O	V	1	1		✓	√	✓		√	✓		✓	✓	✓
Water Vapor, H ₂ O Hydroxyl, HO	√	 √	√		✓	√	√	V	√	✓		.~	✓	.√.
Hydroperoxy1, HO,	1		V	✓.		√.	✓	√.	V	✓		V		✓
			. !				<u> </u>	<u> </u>				<u> </u>		

TABLE 4-III

NAME OF SPECIES/PROPERTY AND					MAJ	OR RE	PERENC	CES WI	ERE C	ITED .				
SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
AEROSOL GROUP 2 Concluded									;	,			:	. 1
Ammonia, NH ₃			√		√.	✓	√.		;			3		
Ammonium Ion, NH ₄ +		✓			;	✓	✓			<u> </u>	,			
AEROSOL GROUP 3, MAJOR SPECIFIC AEROSOLS]			,						
Nitric Acid Aerosol, nHNO3		✓	✓					,						
Chloride Aerosol, Cl										}				•
Nitrate Aerosol, NO ₃		✓	√.			√.	√.				ļ ,			
Ammonium Ion Aerosol, nNH ₄ .		✓	✓		'	✓	√.							
Ammonium Peroxydisulfate, (NH ₄) ₂ S ₂ O ₈			₃⁄.			√.								
AEROSOL GROUP 4, OTHER SPECIFIC AEROSOLS			·											
													! ,	
Aluminum Oxide Aerosol, nAl ₂ O ₃	√				,	•						√		
Nitrite Aerosol, NO ₂		√					✓						[
Liquid Water or Ice, nH20	√.	√.	√		٠									
Aluminum Ion, Al		✓						,		l '				
Bromide Ion, Br		✓	√		,		√.							
Calcium Ion, Ca		✓					✓						·	
Copper Ion, Cu++		√.	~					,						
]			!] .						

NAME OF SPECIES/PROPERTY					MAJ	or rei	ERENC	es Will	ere Ci	ITED	المحدث والمساور			
. AND SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
AEROSOL GROUP 4 Concluded	'		.	!	1 .		ţ				'			
· Iodide Ion, I		✓	✓		,				•		<u> </u>	1		
Iron Ion, Fe or Fe	;		V	:	ï				,					1
Magnesium, Mg		√												;
Manganese Ion, Mn or Mn H			~			·			,					.;
Potassium Ion, K [†]		√.	✓		,				1					
Silicon Ion, Si		√.	✓				,	<u> </u>						,
Sodium Ion, Na		1/2	√.				×							
AEROSOL GROUP 5, OTHER PRECURSOR GASES							4	,	,					:
Ozone, 03 Atomic Chlorine, Cl	✓	√.	√.	~	√.	~	:	√.	~	~	.	~	*	V
Atomic Chlorine, Cl	√.		V	✓	✓	'		✓	V	✓		√.	✓	✓
Chlorine Monoxide, ClO	✓	√	\	✓	✓	.~		√.	Y	✓		✓.	~	. v .
Chlorine Monoxide, ClO Hydrogen, H ₂ or H Nitric Acid Vapor, HNO ₃	√		✓_	✓		✓_		✓	✓	Y	'			_√.
Nitric Acid Vapor, HNO	√.	√.	√	√.	¥	✓		· √.	✓	✓	√.	\	✓.	4
Hydrogen Chloride Gas, HCl	√.	✓	✓.	✓.	√.			✓.	✓	✓		V	Y	Y
Hydrogen Sulfide, H ₂ S				.	√.	√.	\ \		1			✓	,	ļ. i
Methyl Sulfide, (CH ₃) ₂ S		;										.√.		,
					<u> </u>	<u></u>	1	1 :	<u> L </u>	<u></u>	<u> </u>	<u>Ļ</u>	1	

TABLE 4-III (Concluded)

NAME OF SPECIES/PROPERTY AND	L I				L'AM	OR RE	FERENC	CES WH	ERE C	ITED	- <u></u>			
SYMBOL	39	40	62	63	64	65	66	67	68	69	70	71	72	73
AEROSOL GROUP 5 Concluded	•	ļ					,		-					
Carbonyl Sulfide, COS		-					<u>'</u>				٠			1
Carbon Disulfide, CS ₂									<u>'</u>					√.
Hydrogen Fluoride, HF	ļ ,	, •	ļ,	✓					✓	✓		¥	✓	V
Hydrogen Bromide, HBr	•		1 1 1	✓.	ļ				<u> </u>			√	ľ	√.
Hydrogen Peroxide, H ₂ O ₂	✓	1	✓	<u>'</u>		.√.		✓		,		√.	4 .	×
Bisulfite Radical, HSO3					·		√.		ι			√		
Nitrogen Trioxide, NO ₃	✓		ŀ			.√.					, ;	√	V	
Bromine Oxide, BrO		'										√		√.
Atomic Bromine, Br											·	√		₹
AEROSOL GROUP 6, QUASI AEROSOLS		:				:	,							
Sulfur Dioxide, nSO2, (in cluster formation		√.	√								,			
'Nitric Oxide, nNO, (" " "			√.									,	1	
Nitrogen, nN ₂).		✓							1				
Liquid Water or Ice, nH20 (""""""""""""""""""""""""""""""""""""			.∠		· }							2		
Carbon Dioxide, nCO ₂ (" " "			✓			}								
														٠
•														

5.0 MISSION EVALUATIONS

This section presents the results of the application of a method for the evaluation of various stratospheric species measurement missions. The method was developed previously [2] and is presented in detail as Appendix A in both Volumes I and II of this report. In the current application the method has been revised and updated to evaluate the same set of missions and instruments from two standpoints,

- the monitoring of stratospheric ozone and related processes,
- the monitoring of stratospheric aerosols and related precursor gaseous processes.

5.1 Evaluation of Specific Missions

The missions selected for evaluation were:

- A Shuttle-type mission with a 30° inclination and a fourto six-month duration,
- A Shuttle-type mission with a 56° inclination and a fourto six-month duration,
- A polar-type mission with a one- to two-year duration.

 Several instruments under development were evaluated for each of these missions. The instruments evaluated are shown in Table 5-I along with the generic type of each of the species evaluated.

Tables 5-II through 5-XVII show the results of these evaluations for each species/instrument/mission combination. Included with each parameter value for the three missions is the performance used to determine the value. The values (V) shown in each table for present, required and mission capability are taken from the value matrices presented in Volume III of this report. The values represent the

TABLE 5-I
STRATOSPHERIC INSTRUMENTS AND SPECIES EVALUATED
FOR
OZONE AND AEROSOL MISSIONS

INSTRUMENT	GENERIC TYPE	SPECIES	OZONE MISSION	AEROSOL MISSION
LIMS*	Limb scanning	co.	V	-
	•	03 .	✓	. ✓
	•	н ₂ о	✓	✓
	,	NO ₂	√	
		HNO 3	✓ .	. 🗸
. SAGE	Solar occultation	03	✓	✓
		Aerosols		✓
CIMATS	Solar occultation	н ₂ о	V	√
•		CH ₄	√	•
		N ₂ O	V	, ,
-		NH ₃	· •	√
	,	CO	 	
HALOE	Solar occultation	HF	√	√
		СН ₄ .	√ ✓	
		· HC1	√ . √	· • •
		NÓ	√	

^{*} Instrument descriptions are given in Section 6.2 of Volume I.

TABLE 5-II EVALUATION OF CARBON DIOXIDE, CO2, LIMS WITH 80° AZIMUTH SCAN

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	0.1	8	0.8	9	0.8	5 0.5 90°	8 0.8 140°	10 1.0 170°
Duration of Program	0.3	8	2.4	8	2.4	5 1.5 4-6 mos	5 1.5 4-6 mos	7 2.1 1-2 yrs
Diurnal Coverage	0.1	8	0.8	8	0.8	10' 1.0 Full	10 1.0 Full	9 1.0 Part D&N
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	0.2	10	2.0	. 10	2.0	10 2.0 Full	10 2.0 Full	10 2.0 Full
Vertical Profile Resolution	0.2	8	1.6	9	1.8	10 2.0 <1Km	10 2.0 <1Km	10 2.0 <1Km
Longitude	0.1	8	0.8	8	0.8	10 1.0 Full	10 1.0 Full	10 1.0 Full
	1.0		8.4	, 	8.6	8.0	8.3	9.1
Total Value			8		9	.8	. 8	9
Incremental Gain Over Present					<1	<1	<1	1



LEGEND:

V = Value

VXWF = Value x weighting factor

TABLE 5-III

EVALUATION OF OZONE, LIMS WITH 80° AZIMUTH SCAN

Parameter	. WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.25	10	2.5	10	2.5	5 1.25 90°	6 1.5 140°	10 2.5 170°
Duration of Program	.25	7	1.75	10	2.5	4 1.0 4-6 mos	4 1.0 4-6 mos	6 · 1.5 1-2 yrs
Diurnal Coverage	.15	2	.3	8	1.2	10 1.5 Full	10 1.5 Full	8 1.2 Part D&N
Launch Time	0	10	0	10	0	10 0 .	10 0	10 0 .
Vertical Profile Coverage	.1	7	.7	10	1	10 1.0 Full	10 1.0 Full	10 1.0 Full
Vertical Profile Resolution	.15	5	.75	10	1.5	10 1.5 <1Km	10 1.5 <1Km	10 1.5 <1Km·
Longitude	.1	10	1	10	1	10 1.0 Full	10 1.0 Full	10 1.0 Full
	1.0		7.0		9.7	7,25	7.5	8.7
Total Value			7		10	7	8	9
Incremental Gain Over Present				,	3 · · · · · · · · · · · · · · · · · · ·	<1	1	2

LEGEND:

V = Value

ORIGIN JUALITY

TABLE 5-IV EVALUATION OF OZONE, SAGE, SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle · 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.25	10	2,5	10	2.5	4 1.0 90° sparse at extremes	7 1.75 150° spars at extreme	
Duration of Program	.25	7	1.75	10	2.5	4 1.0 4-6 mos	4 1.0 4-6 mos	6 1.5 1-2 yrs
Diurnal Coverage	.15	2	.3	8	1.2	2 0.3 Part Day 2 points	2 0.3 Part Day 2 points	2 0.3 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 O	10 0
Vertical Profile Coverage	.1`	7	.7	10	1	10 1.0 Full	10 1.0 Full	10 1.0 Full
Vertical Profile Resolution	.15	5	.75	10	1.5	7 1.05 ~10 points	7 1.05 ~10 points	7 1.05 ~10 points
Longitude	,	10	1.	10	1	10 1.0 Ful1	10 1.0 Full	10 1.0 Full
	1.0		7.0	· •••••••	9.7	5.35	6.1	4.85
Total Value		•	7		10	. 5	, 6	5 `
Incremental Gain Over Present					3	<1	<1	<1

LEGEND:

V = Value

VXWF = Value x weighting factor

TABLE 5-V EVALUATION OF WATER VAPOR, H2O, CIMATS SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle .56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.3	6	1.8	9	2.7	6 1.8 90° sparse at extremes	8 2.4 150° sparse at extremes	
Duration of Program	.2		1.0	9	1.8	6 1.2 4-6 mos	6 1.2 4-6 mos	9 1.8 1-2 yrs
Diurnal Coverage	.1	7	0.7	8	0.8	2 0.2 Part Day 2 points	2 0.2 Part Day 2 points	2 0.2 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.15	5	0.75	. 10	1.5	10 1.5 Full ·	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution	.15	7 .	1.05	10	1.5	5 0.75 ~20 points	5 0.75 ~20 points	5 0.75 ~20 points
Longitude	.1	0	0	8	8.0	10 0.1 Full	10 0.1 Full	10 0.1 Full
-	1.0		5.3		9.1	5.55	6,15	4.35
Total Value ' · ·	,		5		. 9	6	. 6 ·	4
Incremental Gain Over Present		,			4	1	1	<1

LEGEND: $\overline{V} = Value$ VXWF = Value x weighting factor

TABLE 5-VI EVALUATION OF WATER VAPOR, H₂O; LIMS WITH 80° AZIMUTH SCAN

 Parameter	WF 0-1		sent ledge VXWF		uired oility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.3	6 :	1.8	9	2.7	7 2.1 90°	9 2.7 140°	10 3.0 170
Duration of Program	. 2	5	1.0	9	1.8	6 1.2 4-6 mos	6 1.2 4-6 mos	9 1.8 1-2 yrs
Diurnal Coverage	.1	7	0.7	8	0.8	10 1.0 Full	10 1.0 Full	9 0.9 Part D&N
Launch Time	٠ ٥	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.15	5	0.75	10	1.5	10 1.5 Full	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution ·	.15	7	1.05	10	1.5	10 1.5 <1Km	10 1.5 <1Km	10 1.5 <1Km
Longitude	1	. 0	0	8	0.8	10 1.0 Full	10 1.0 Full	10 1.0 Full
	1.0		5.3	-	9.1	8.3	8.9	9.7
Total Value			5		9	8	9 .	10
Incremental Gain Over Present		,			4	3	· 4	5



LEGEND: $\overline{V} = Value$ VXWF = Value x weighting factor

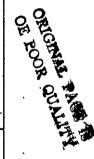


TABLE 5-VII EVALUATION OF AEROSOLS, SAGE SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.15	9	1.35	10	1.5	4 0.6 90° sparse at extremes	7 1.05 150° spars at extreme	
Duration of Program	.15	8	1.2	9	1.35	7 1.05 4-6 mos	7 1.05 4-6 mos	9 1.35 1-2 yrs
Diurnal Coverage	.05	9	0.45	9	0.45	6 .3 Part Day 2 points	6 .3 . Part Day 2 points	6 .3 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.25	. 8	2.0	10	2.5	10 2.5 Full	10 2.5 Full	10 2.5 Full
Vertical Profile Resolution	.15	7	1.05	10	1.5	7 1.05 ~10 points	7 1.05 ~10 points	7 1.05 ~10 points
Longitude	.25	6	1.5	10	2.5	10 2.5 Full	10 2.5 Full	10 2.5 Full
	1.0		7.55	-	9.8	8.00	8.45	7.7
Total Value			8	•	10	8	. 8.	8
Incremental Gain Over Present					2	<1	<1	<1

LEGEND:

 $\overline{V} = Value$ VXWF = Value x weighting factor

TABLE 5-VIII EVALUATION OF AMMONIA, NH3, CIMATS SOLAR OCCULTATION

Parameter	WF	Pres V	ent VXWF	, Requ	ired VXWF	Shuttle V VXWF	Shuttle V VXWF	Sun-Sync V VXWF
Latitude	. 2	0	0	7	1.4	7 1.4 90° sparse at extremes	8 1.6 150° sparse at extremes	
Duration of Program	.1 .	0	0	6	0.6	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.9 1-2 yrs
Diurnal Coverage	.15	0	O .	6	0.9	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points
Launch Time	Ò	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.25	0	0	7	1.75	10 2.5 Full	10 2.5 Full	10 2.5 Full
Vertical Profile Resolution	.25	0	0 .	7	1.75	9 2.25 ~20 points	9 2.25 ~20 points	9 2.25. ~20 points
Longitude .	.05	0	0	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
,	1.0		0,		6,8	8.05	8.25	6.75
Total Value			0		7	8	8 ·.	· 7 _.
Incremental Gain Over Present	,				7	8		7

 $\frac{\text{LEGEND:}}{V = Value}$



TABLE 5-IX EVALUATION OF NITROGEN DIOXIDE, NO2, LIMS WITH 80° AZIMUTH SCAN .

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF	
Latitude' .	.15	4	•6	10	1.5	7 1.05 90°	9 1.35 140°	10 1.5 170°	
Duration of Program	.15	5	.75	9	1.35	7 1.05 4-6 mos	7 1.05 4-6 mos	9 1.35 1-2 yrs	٠.
Diurnal Coverage	.35	5	1.75 ·	9	3,15	10 3.5 Full	10 3.5 Full	8 2.8 Part D&N	•
Launch Time	. Ò	10	0	10	0 ·	10 0	10 0	10 0	
Vertical Profile Coverage	.15	6	0.9	10	1.5	10 1.5 Full	10 1.5 Full	10 1.5 Full	
Vertical Profile Resolution	.15	4	0.6	10	1.5	10 1.5 ⋜1Km	10 1.5 ≂1Km	10 1.5 ≈1Km	•
Longitude	.05	0	0	8	0.4	10 .5 Full	10 .5 Full	10 .5 Full	,
	1.0		4.6		9.4	9.1	9.4	9.15	
Total Value			5		9	9	9 .	, 9	,
Incremental Gain Over Present			•		4	· 4	4	4 .	

LEGEND: V = ValueVXWF = Value x weighting factor

TABLE 5-X EVALUATION OF NITRIC ACID. VAPOR, NHO $_3$, LIMS WITH 80 $^{\circ}$ AZIMUTH SCAN

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF	
Latitude	.3	5	1.9	10	3.0	7 2.1 90°	9 2.7 140°	10 3.0 170°	
Duration of Program	.25	3	.75	9	2,25	7 1.75 4-6 mos	7 1.75 4-6 mos	9 2.25 1-2 yrs	,
Diurnal Coverage	.1	7	. 7	8	.8	10 1.0 Full	10 1.0 Full	8 0.8 Part D&N	
Launch Time	0	10	0	10	0	10 0	10 0	10 0	
Vertical Profile Coverage	.15	7	1.05	10	1.5	10 1.5 Full	10 1.5 Full	10 1.5 Full	
Vertical Profile Resolution	. 1.	8	.8	10	1.0	10 1.0 ₹1Km	10 1.0 ₹1Km	10 1.5 ≂1Km	
Longitude	.1	0	0	8	.8	10 1.0 Full	10 1.0 Full	10 1.0 Full	
	1.0		4.8		9,35	8.35	8.95	9.55	•
Total Value			5.		9	8	9	10	
Incremental Gain Over Present					4	3	4	5	

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LEGEND: ·

V = Value VXWF = Value x weighting factor D & N = Day & Night

TABLE 5-XI EVALUATION OF HYDROGEN CHLORIDE GAS, HC1, HALOE SOLAR OCCULTATION

Parameter .	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.35	4	1.4	9	3.15	6 2.1 90° sparse at extremes	8 2.8 150° sparse at extremes	
Duration of Program	•1	5	.5	8	.8	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.8 1-2 yrs
Diurnal Coverage	.1	0	0	7	.7	3 0.3 Part Day 2 points	3 0.3 Part Day 2 points	3 0.3 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	2	6	1.2	9	1.8	9 1.8 10-40Km	9 1.8 10-40Km	9 1.8 10-40Km
Vertical Profile Resolution	.2	7	1.4	9	1.8	9 1.8 2Km	9 1.8 2Km	9 1.8 2Km
Longitude 	.05	0	0	8	.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
	1.0		4.5		8.65	7.35	8.0	5,2
Total Value			5		9	7 .	. 8	5
Incremental Gain Over Present					4	2	.3	, <1

 $\frac{\text{LEGEND}}{V = Value}$

EVALUATION OF METHANE, CH4, CIMATS SOLAR OCCULTATION

Parameter	WF 0−1		sent , ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync · Noon V VXWF
Latitude	. 4	0	0	8	3.2	7 2.8 90° sparse at extremes	8 3.2 150° spars at extreme	
Duration of Program	.1	0	0	6	0.6	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.9 1-2 yrs
Diurnal Coverage	.15	0.	0	6	0.9	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points
Launch Time	. 0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.15	6	0.9	8	1.2	10 1.5 Full	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution	.15	3	0.45	9	1.35	9 1.35 ~ 20 points	9 1.35 ~20 points	9 1:35 ~20 points
Longitude	.05	, 0	0	8	0,.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
	1.0		1.35		7.65	7.55	7.95	4.85
Total . Value			1		8	8	8	5
Incremental Gain Over Present					7	7	7	4

LEGEND: $\overline{V = Value}$. VXWF = Value x weighting factor



TABLE 5-XIII EVALUATION OF METHANE, CH4, HALOE SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.4	· 0	0	8.	3.2	7 2.8 90° sparse at extremes	8 3.2 150° sparse at extremes	
Duration of Program	.1	0	0	6	0.6	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.9 1-2 yrs
Diurnal Coverage	.15	0	0	6	0.9	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 Q ·
Vertical Profile Coverage	.15	6	0:9	8	1.2	10 1.5 Full	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution	.15	3	0.45	9	1.35	9 1.35 2Km	9 1.35 2Km	9 1.35 2Km
Longitude	.05	0	0	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 Full .
·	1.0		1.35		7.65	7.55	7.95	. 4.85
Total Valué			1		8	8	8	. 5
Incremental Gain Over Present			•	•	, 7	7	· 7	4

LEGEND: $\overline{V} = Value$ VXWF = Value x weighting factor

EVALUATION OF NITROUS OXIDE, n_2° 0, CIMATS SOLAR OCCULTATION

Parameter	WF 01		esent vledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.25	4	1.0	10	2.5	7 1.75 90° sparse at extremes	9 2.25 150° spars at extreme	
Duration of Program	.15	5	0.75	9	1.35	7 1.05 4-6 mos	7 1.05 4-6 mos	9 1.35 1-2 yrs
Diurnal Coverage	.1	8	0.8	. 8	0.8	1 0.1 Part Day 2 points	1 0.1 Part Day 2 points	1 0.1 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.15	6	.0.9	10	1.5	10 1.5 Full	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution	.15	4	0.6	10	1.5	9 1.35 ~20 points	9 1.35 ~20 points	9 1.35 ~20 points
Longitude	.05	0	0	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
	1.0		4.05		8.05	6.25	6.75	4.8
Total Value			4		8	6	7	5.
Incremental Gain Over Present				-	4	2	3	1

 $\frac{LECEND}{V = Value}$

TABLE 5-XV

EVALUATION OF CARBON MONOXIDE, CO, CIMATS SOLAR OCCULTATION

Parameter	WF. 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude 	.4	0	0	8	3.2	7 2.8 90° sparse at extremes	8 3.2 150° sparse at extremes	
Duration of Program	. 1.	0	0	6.	0.6	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.9 1-2 yrs
Diurnal Coverage	15	0	0	6	0.9	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points
Launch Time	0	10	0	10	0 .	10 0	10 0 .	10 0
Vertical Profile Coverage	.15	5	.75	9	1.35	10 1.5 Full	10 1.5 Full	10 1.5 Full
Vertical Profile Resolution	15	3	.45	9	1.35	9. 1.35 ~ 20 points		9 1.35 ~20 points
Longitude	.05	0	0	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
	1.0		1.2	• •••••	7.8	7.55	7 . 95	4.85
Total Value	•		1		8.	8	8	5 .
Incremental Gain Over Present	•				7	7	7	4

LEGEND:

 $\overline{V} = Value$

TABLE 5-XVI EVALUATION OF HYDROGEN FLUORIDE, HF, HALOE SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		ired ility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude	.2	0	0	7	1.4	7 1.4 90° sparse at extremes	8 1.6 150° spar at extrem	
Duration of Program	.1	0	0	6	0.6	8 0.8 4-6 mos	8 0.8 4-6 mos	9 0.9 1-2 yrs
Diurnal Coverage	.15	0	0	6	0.9	4 0.6 Part Day 2 points	4 0.6 Part Day 2 points	- 1
Launch Time	0 ,	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.25	0	0	7	1.75	10 2.5 Full	10 2.5 Full	10 2.5 Full
Vertical Profile Resolution	.25	0	0	7	1.75	9 2.25 2Km	9 2.25 2Km	9 2.25 2Km
Longitude	.05	0	· o	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 Full
	1.0	•	0		6.8	8.05	8.25	6.75
Total Value		0		7		8	8	7
Incremental Gain Over Present				7		8	8	7

LEGEND:

V = Value

VXWF = Value x weighting factor



TABLE 5-XVII EVALUATION OF NITRIC OXIDE, NO, HALOE SOLAR OCCULTATION

Parameter	WF 0-1		sent ledge VXWF		uired bility VXWF	Shuttle 30° V VXWF	Shuttle 56° V VXWF	Sun-Sync Noon V VXWF
Latitude .	.25	4	1.0	10	2.5	6 1.5 90° sparse at extremes	9 2.25 150° sparse at extremes	0 0 ∿5°
Duration of Program	.2	8	1.6	9	1.8	5 1.0 4-6 mos	5 1.0 4-6 mos	9 1.8 1-2 yrs
Diurnal Coverage	.3	5	1.5	9	2.7	3 0.9 Part Day 2 points	3 0.9 Part Day 2 points	3 0.9 Part Day 2 points
Launch Time	0	10	0	10	0	10 0	10 0	10 0
Vertical Profile Coverage	.1	7	0.7	10	1.0	10 1.0 Full	10 1.0 . Full	10 1.0 Full
Vertical Profile Resolution	.1	7	0.7	, 10	1.0	8 0.8 2Km	8 0,•8 2Km	.8 0.8 2Km
Longitude	.05	0	0	8	0.4	10 0.5 Full	10 0.5 Full	10 0.5 . Full
	1.0		5.5		9.4	5.7	6.5	5.0
Total Value			6		9	6	7	5
Incremental Gain Over Present				•	3	<1	. 1	<1

LEGEND:

V = Value

VXWF = Value x weighting factor

relative value on a scale of 0 to 10 (low to high) for the stated performance where 0 indicates no capability and 10 indicates perfect capability. The weighting functions show the value of one parameter relative to the others under study. The product of the value and its corresponding weighting function (VXWF) yields the desired weighted value for each parameter. The sum of the weighted values for each parameter yields the total relative value for each pollutant (see Appendix A for full explanation).

5.1.1 Ozone Related Single Instrument Missions

In Table 5-XVIII the incremental gains have been summarized to show the totals for each instrument/orbit/species combination which could be used in an ozone related mission. The results have been weighted by the weighting factors for the various ozone related pollutant groups. These weights adjust to individual pollutant values to account for the different priority groups into which they were placed in Section 4.1.1 (Table 4-II). The factors assigned to the different ozone priority groups are as follows:

- Group 1 Measurements of properties and species 1.0 directly associated with UV change
- Group 2 Components of the basic reactions involved in the direct production or
 depletion of ozone
- Group 3 Major components of the basic reactions 0.8 indirectly involved in the production or depletion of ozone
- Group 4 Other significant components of the ozone .0.6 chemistry chains

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TABLE 5-XVIII

OF POOR QUALITY SUMMARY OF INCREMENTAL GAINS FOR EACH OZONE RELATED SPECIES/INSTRUMENT/ORBIT COMBINATION

				• .*				
	OZONE	GROUP		WEIGHTED	WEIGHTED INCREMENTAL GAIN			
SPECIES	PRIORITY GROUP	WEIGHTING FACTOR	INSTRUMENT	REQUIRED GAIN		56° ORBIT,	POLAR ORBIT	
· 0 ₃	1	1.0	LIMS	3	<1	1	2	
	· · 1	1.0	SAGE	. 3	<1	<1	<1	
NO	2	0.9	HALOE	3	<1	.<1	< 1	
NO ₂	2	0.9	LIMS	4.	. 4	4	4	
H ₂ 0	. 3	0.8	CIMATS	3	<1	<1	<1	
	3	0.8	LIMŠ	3 .	. 2	3	4 .	
NH ₃	3	.0.8.	CIMATS	6	6	6	. 6	
HNO ₃	3	0.8	LIMS	3 ·	2	3	4	
HCl	3	0.8	HALOE	3	2	Ž.	<1	
CH ₄	3	0.8	CIMATS	6	6	6	3	
	3	0.8	HALOE	6	. 6	6	. 3	
N ₂ O	3	0.8	CIMATS	3	2	2	<1	
СО	3	0.8	CIMATS	6	6	6	3	
co ₂	4	0.6	LIMS	<1	<1	<1	<1	
HF.		0.6	HALOE	. 4	5	5	4	
			Σ.	1 44-45			. ,	
,			2=	44-45				

^{*}Uncertainty due to use of values <1. The total equals the sum of the best values for each species.

The incremental gain totals for each instrument/orbit combination which could be used in an ozone related mission are summarized in Table 7-XIX. It is obvious that those combinations showing the highest gains exhibit two prominent characteristics,

- (1) the instrument measures a larger number of species, and/or
- (2) most of the species measured represent those for which little data now exist; this allows large incremental gains for any successful measurement.

5.1.2 Aerosol Related Single Instrument Missions

In Table 5-XX the incremental gains have been summarized to show the totals for each instrument/orbit/species combination which could be used in an aerosol related mission. The results have been weighted by the weighing factors for the various aerosol related pollutant groups. These weights adjust the individual pollutant values to account for the different priority groups into which they were placed in Section 4.1.2 (Table 4-III). The factors assigned to the different aerosol priority groups are as follows:

Group 1A - Direct measurements of climatic change and ultraviolet change	1.0
Group 1B - Major aerosols	. 1.0
Group 2 - Major sulfate aerosol precursor gases	1.0
Group 3 - Major specific aerosols	0.8
Group 4 - Other specific aerosols	0.6
Group 5 - Other precursor gases	0.5
Group 6 - Quasi aerosols (clusters)	0.3

TABLE 5-XIX

SUMMARY OF INCREMENTAL GAINS FOR EACH OZONE RELATED INSTRUMENT/ORBIT COMBINATION

		WEIGHTED	WEIGHTED INCREMENTAL GAIN				
INSTRUMENT	SPECIES MEASURED	REQUIRED GAIN	30° ORBIT	56° ORBIT	POLAR ORBIT		
LIMS	co ₂ , o ₃ ,	13-14*	8-10*	11-12*	· 14-15 [*]		
	H ₂ O, NO ₂						
SAGE	03	3	<1	<1	<1		
CIMATS (solar occultation)	H ₂ O, CH ₄ , N ₂ O, NH ₃ ,	24	20-21*	20-21*	12-14*		
HALOE	CO HF, CH ₄ HC1, NO	16	13-14*	13-14*	7-9*		
				- 	Taryana de la composition della composition dell		
		t					

^{*}Uncertainty due to use of values of <1

SUMMARY OF INCREMENTAL GAINS FOR EACH AEROSOL RELATED SPECIES/INSTRUMENT/ORBIT COMBINATION

TABLE 5-XX

	OZONE GROUP PRIORITY WEIGHTING			WEIGHTED REQUIRED	WEIGHTED INCREMENTAL GAIN			
SPECIES	GROUP	FACTOR	INSTRUMENT	GAIN	30° ORBIT	56° ORBIT	POLAR ORBIT	
iotal aerosols	1	1.0	SAGE	2	<1	<1	<1	
H ₂ 0 -	2	1.0	CIMATS	4.	1	1	<1	
	2	1.0	LIMS	4	3	4	5	
NH ₃	2	1.0	CIMATS	7 .	8	8	7	
³ .	5	0.5	LIMS	2	<1	<1	1	
	5	0.5	SAGE	2	<1 .	<1	<1	
-1NO3	5	0.5	LIMS	2	2	2	3	
HC1	. 5	0.5	HALOE	2	1	2	<1	
I F	5	0.5	HALOE	4	4	4 .	4	
			\sum	= 23*				
							,	
		٠						
					1			
					1			

^{*}The total equals the sum of the best values for each species.

The incremental gain totals for each instrument/orbit combination which could be used in an aerosol related mission are summarized in Table XXI. Inspection of the data shows that the highest gains are produced by those instruments which,

- (1) measure a larger number of species, particularly those species in high priority groups, and/or
- (2) measure high priority species for which little data now exists; thus allows large incremental gains for any successful measurement.

Thus the CIMATS instrument, which has the potential for measuring the high priority but previously unmeasured ammonia, scores very well as opposed to SAGE which measures relatively well known ozone and total aerosol distributions.

5.2 Evaluation of Multiple Species or Instrument Missions

5.2.1 Ozone Related Missions

Table 5-XXII shows the summary of incremental gains resulting when various combinations of two, three, or four instruments are flown on the same ozone oriented mission. These values are obtained by adding the individual contributions of each species/instrument except in those cases where two instruments measure the same species. In this later case, the value is determined by using the better value for each parameter between the instruments involved.

Inspection of the results reemphasizes some previous intuitive knowledge and also presents some new concepts. In the former category are such results as:

SUMMARY OF INCREMENTAL GAINS FOR EACH AEROSOL RELATED - ORIGINAL PAGE IS INSTRUMENT/ORBIT COMBINATION

TABLE 5-XXI

OF POOR QUALITY

	SPECIES	WEIGHTED	WEIGHTED INCREMENTAL GAIN				
INSTRUMENT	MEASURED	REQUIRED GAIN	30° ORBIT	56° ORBIT	POLAR ORBIT		
LIMS	H ₂ O, O ₃ ,	8	* · 5–.6 [*]	6-7*	9		
SAGE	Aerosols,	4	<1	<1	<1		
CIMATS (solar occul- tation)	н ₂ о, мн ₃	11	9	9	7-8 [*]		
HALOE	HC1, HF	6	-		4-5*		

^{*}Uncertainty due to use of values <1

TABLE 5-XXII

SUMMARY OF INCREMENTAL GAINS RESULTING

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FROM VARIOUS INSTRUMENT COMBINATIONS USED IN OZONE ORIENTED MISSIONS

		WEIGHTED	WIIGHTED INCREMENTAL GAIN		
INSTRUMENTS	SPECIES MEASURED	REQUIRED GAIN	30° ORBIT	56° ORBIT	POLAR ORBIT
Four Instruments LIMS, SAGE, CIMATS, HALOE	O ₃ , NO, NO ₂ , H ₂ O, NH ₃ , HNO ₃ , HCl, CH ₄ , N ₂ O, CO, CO ₂ , HF	44-45*	36–37	38-40	31–33
Three Instruments LIMS, CIMATS, HALOE	O ₃ , NO, NO ₂ , H ₂ O, NH ₃ , HNO ₃ , HC1, CH ₄ , N ₂ O, CO, CO ₂ , HF	44–45	. 3637	38~40	31–33
LIMS, SAĜE, CIMATS	о ₃ , мо ₂ , н ₂ о, мн ₃ , нмо ₃ , сн ₄ , м ₂ о, со, со ₂	34–35	29-30	31-32	2628
SAGE, CIMATS, HALOE	03, NO, H20, NH3, HC1, CH4, N20, CO, HF	37	28-29	28-29	18-19
LIMS, SAGE, HALOE	o ₃ , no, no ₂ , H ₂ o, Hno ₃ , HC1, CH ₄ , CO ₂ , HF	29-30	22–23	24-26	22-23
Two Instruments LIMS, CIMATS	о ₃ , мо ₂ , н ₂ о, мн ₃ , нмо ₃ , сн ₄ , м ₂ о, со, со ₂	34–35	29–30	3132	26-28
CIMATS, HALOE	NO, H ₂ O, NH ₃ , HC1, CH ₄ , N ₂ O, CO, HF	34	27-28	27–28	17-19
SAGE, CIMATS	03, H20, NH3, CH4, N20, CO	27	20–21	20-21	13-14
LIMS, HALOE	o ₃ , no, no ₂ , H ₂ o, Hno ₃ , HC1, CH ₄ , CO ₂ , HF	29-30	22–23	24-26	22-23
LIMS, SAGE	о ₃ , мо ₂ , н ₂ о, нмо ₃ , со ₂	13-14	8-10	11-12	14-15
SAGE, HALOE	о ₃ , мо, нс1, сн ₄ , нғ	19	13-15	13–15	8-9

^{*}Uncertainty due to use of values <1

- the more individual species and/or instruments involved the greater the value
- Solar occultation-type instruments give poor global coverage in polar orbits
- Limb-looking instruments give excellent global coverage in polar orbits

The principal conclusion in the later category is that the highest potential for gain in value lies in the measurement of those species in ozone priority group 3 which play important roles in stratospheric processes but whose characteristics and spatial/temporal distribution are poorly known. Thus those instruments which measure species such as ammonia, methane and carbon monoxide score relatively high. These factors consistently lead to higher results for instrument/orbit combinations involving LIMS, CIMATS and HALOE as opposed to the SAGE instrument which measures only the reasonably well understood ozone.

In terms of orbit selection it is clearly shown in every combination that the 56° Shuttle type orbit is superior to the 30° Shuttle type orbit and the polar orbit.

5.2.2 Aerosol Related Missions

Table XXIII shows the summary of incremental gains resulting when various combinations of two, three or four instruments are flown on the same aerosol oriented mission. As was the case with ozone missions these results show that combinations which measure more species score relatively higher, particularly if the species is high priority and relatively unmeasured.

TABLE XXIII

SUMMARY OF INCREMENTAL GAINS RESULTING FROM VARIOUS INSTRUMENT COMBINATIONS USED IN AEROSOL ORIENTED MISSIONS

	SPECIES MEASURED	WEIGHTED REQUIRED GAIN	WEIGHTED INCREMENTAL GAIN		
INSTRUMENTS			30° ORBIŢ	56° ORBIT	POLAR ORBIT
Four Instruments LIMS, SAGE, CIMATS, HALOE	Aerosols, H ₂ O, NH ₃ , O ₃ , HNO ₃ , HCl, HF	23	18-20 [*]	20-22	20–22
Three Instruments LIMS, CIMATS, HALOE	н ₂ 0, мн ₃ , 0 ₃ , - нмо ₃ , нс1, нғ	21	18–19	2021	20-21
LIMS, SAGE, CIMATS	Aerosols, H ₂ O,	. 17	13-15	15-17	16-17
SAGE, CIMATS, HALOE	Aerosols, H ₂ O, NH ₃ , O ₃ , HCl, HF	21	14-16	15–17	12–14
LIMS, SAGE, HALOE	Aerosols, H ₂ O, O ₃ , HNO ₃ , HC1, HF	16	10-12	12–14	13-15
Two Instruments					
LIMS, CIMATS	н ₂ о, мн ₃ , о ₃ , нмо ₃	15	13-14	14-15	. 16
CIMATS, HALOE	H ₂ O, NH ₃ , HC1, HF	17	14	· 15	11-13
SAGE, CIMATS ·	Aerosols, H ₂ O,	. 15	9–11	9-11	8~9
LIMS, HALOE	н ₂ о, о ₃ , ноо ₃ , нс1, нғ	14	10-11	12–13	1,3-14
LIMS, SAGE	Aerosols, H ₂ O, O ₃ , HNO ₃	10 	5–7	6-8, ,	9–10
SAGE, HALOE	Aerosols, 0 ₃ , HC1, HF	10	5–7	6-8	5-6

^{*}Uncertainty due to use of values <1

In contrast with the ozone results, orbit selection for combinations of instruments for aerosol missions does not consistently show the 56° orbit to score highest. The polar orbit tends to prevail in combinations containing the limb-emission LIMS instrument while the 56° orbit is superior when it is absent. In no case is the 30° orbit shown to be superior.

APPENDIX A

MISSION EVALUATION METHODOLOGY

A.1 INTRODUCTION

In order to properly determine how well any selected stratospheric species measurement mission improves on present knowledge of
the characteristics and spatial/temporal distribution of the species,
a method is presented that evaluates a selected mission in terms of
the present status of stratospheric knowledge of the species of
interest and the required level of knowledge (as expressed by the
scientific user community). The method has also been inverted and
used to select the mission that is most effective.

The selection of an optimum mission involves not only the evaluation of orbital characteristics but also the selection of those species to be measured that provide the optimum incremental improvement from present knowledge to required knowledge. Thus, two factors are involved:

- (1) Prioritization of pollutants based on a combination of present knowledge and required knowledge.
- (2) Selection of the "optimum" mission (orbit plus instrument) based on present measurement knowledge and required knowledge.

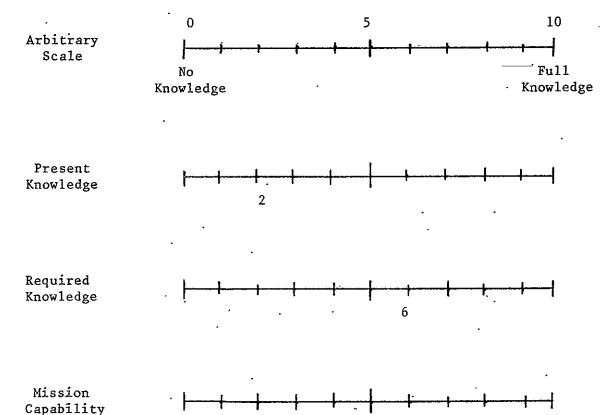
The following sections will be limited to a discussion of the "optimum" mission selection for a single species. The prioritization of species based on requirements was discussed in Section 5.2. Incorporation of these priorities into the evaluation methodology will be discussed later.

This evaluation technique can be applied specifically to orbit evaluation, instrument evaluation, or both by selection of the appropriate parameters.

A.2 DEVELOPMENT OF THE METHOD

A.2.1 Approach to the Ranking and Evaluation

For each stratosphere species of interest one may assign a ranking or value in terms of an arbitrary scale of, say, 0 to 10 based on a comparison of either: (1) the present knowledge of the species distribution, (2) the required knowledge of the species distribution, or (3) the projected measurement capability of a specific mission with the total possible four-dimensional knowledge. For a typical species this may be exemplified as follows:



. 4 . A 6

В

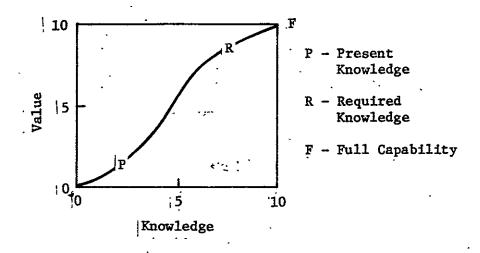
8

C

The key to assessing the value of a particular mission lies in comparing the mission capability with the incremental improvement between present knowledge and required knowledge. In the example illustrated above, the present level of knowledge has been given an arbitrary rating of 2 and the required knowledge an arbitrary rating of 6. It is important to note that the required knowledge level is not always set at the maximum. This may be for two reasons. On the one hand, a full capability of 10 may provide the user with much more data than he needs or could ever make use of. On the other hand, the present level of knowledge may be so low that the user would require only a small increase in knowledge to achieve a significant improvement in understanding the chemistry and distribution of the pollutant. Requirements should be set at the level that best equals the capability of the user community to assimilate the data measured.

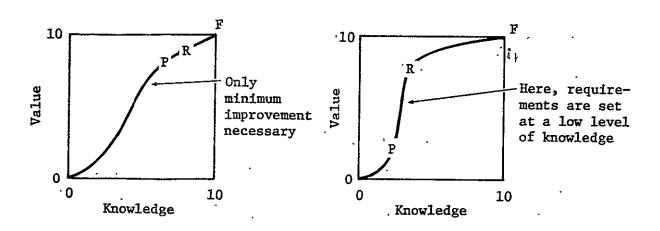
Thus, in the given example, the critical area for gain lies between the present knowledge and the required knowledge. Therefore, system C is not automatically much better than system B. However, each (B and C) is significantly better than system A.

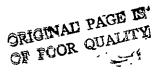
In order to indicate this in a more powerful way, the ranking scheme may be presented in a slightly different manner:



Here we see a sharp rise in value between present and required knowledge and little gain thereafter. Present knowledge is assigned a value at or near zero and required knowledge is assigned a value approaching 10 but allowing some small value for additional knowledge up to full.

In other cases the present knowledge may be such that it commands a high value in relation to full capability leaving little room for improvement. Conversely, the current requirements may be such that they can be fulfilled with only a minimum additional capability.





This type of evaluation has been used previously in a variety of system evaluations [108-112]. These reports give the details of the application of the method to both real cases and illustrative examples. The evaluation method makes use of value judgments of experts, either individually or by consensus, to provide information where "hard" data are unavailable. The objective is to make use of as much information as is available to the system. Much of this information is derived from the experience of experts associated with the system being evaluated. It is the objective of the evaluation to extract this information and check for its validity and utility. Critical areas can be identified where further gathering of information would be most effective. The success of the method depends on two critical factors:

- Availability of expert opinions or facts on the subject either directly or through adequate documentation.
- A thorough understanding of the structure and utilization of the evaluation procedure.

A logical sequence of steps in the application of the evaluation method is shown in Figure A-1. The first step is to identify the appropriate evaluation parameters. These parameters when measured will provide the information needed to describe and adequately evaluate the candidate species, instruments, and orbits. The selection of the parameters must be made independent of any particular knowledge of instruments or orbits.

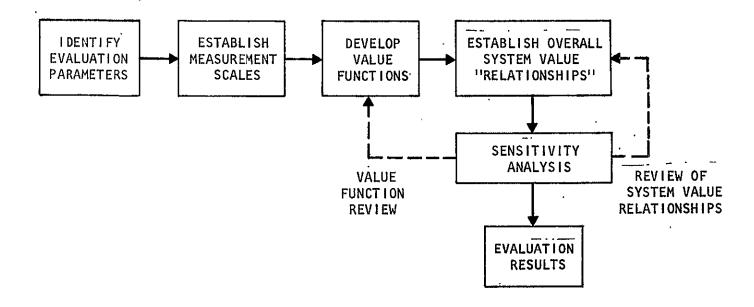


FIGURE A-1
EVALUATION TECHNIQUE DEVELOPMENT AND VALIDATION

Once the parameters are identified, measurement scales must be established for each parameter. The ranges of the technical parameter measurement values can be based either on established facts (which are generally unavailable) or expert judgments. The analytical formulation of the technique begins with the development of the value functions. The value function and its graphic representation, the value judgment curve, are the basic inputs of the method. The value function relates points on the parameter measurement scale to a value scale that ranges between zero for no value to the user and some arbitrary positive number for maximum value to the user. (Ten was selected as maximum in this study.)

The first step in developing a typical value judgment function is to establish the maximum and minimum points for each of the evaluation parameters. Additional points between the parameter maximum and minimum points are defined and each assigned a value to the user.

Identification of all break points is very valuable in this procedure. These points are then plotted on a value judgment scale to indicate the nature of the actual relationship. In most cases the judgment curves should have the following characteristics,

- Smooth variation over the entire range
- Zero slope at the origin
- An asymptotic approach to zero or the maximum for large values of the parameters
- Flexibility so that special cases are easily incorporated

These characteristics are best represented by the family of hyperbolic tangent curves characterized by the scale factors α and n. Then,

$$V = \tanh(\alpha x^n)$$
 or $V = 1-\tanh(\alpha x^n)$

where, V = value to the user; x = parameter value; α determines at what point a change in parameter value begins to have a significant effect on the value to the user and n determines the slope of the change. In order for value to user to increase with increasing parameter change n must be greater than 1. While the hyperbolic tangent curve is used in most cases, it should be noted that other types of value functions can be used. These may in some cases be step functions or binary functions.

The next phase in the formulation of the technique is to develop the overall system value relationship. This is accomplished by establishing the relative importance of each of the parameters through weighting functions. The initial step in developing these functions is to designate each parameter as a factor or a term. A parameter is designated as a factor if it is of such paramount importance that if the value to the user is zero for that parameter, the entire system is considered valueless. If a parameter is not of the same level of criticality as a factor, it is designated a term. A term is related to the other parameters through an additive relationship.

The second step in establishing the relative importance of the performance parameters is to assign weights to each parameter designated a term; where the sum of these weights is equal to unity. Various methods can be used to assign the weights. For example, the Delphi technique developed by the RAND Corporation has been used to reach a consensus within a group of experts as to the weights which should be assigned. Another method is to assign an initial set of weights and evaluate them against candidate species whose characteristics and relative importance are known. Refinement of the weights is then made based on the results. However, there is no substitute for the participation of experts in the field, either actually or by proxy.

The relationship among all parameters, including terms and factors, is then established, taking the general form of the following equation:

$$V = \prod_{j=1}^{n} \left(F_{j}(x_{j}) \right) \sum_{i=1}^{m} \left(A_{i}G_{i}(x_{i}) \right)$$

where

$$\sum_{i=1}^{n} A_{i} = 1$$

V = value

 A_{i} = weight

F; = value function (factor)

-G; = value function (term)

x_i,x_i = parameter measurement

This equation is termed a value set and can be used to evaluate for example all candidate instruments and/or orbits for a single stratos-pheric species.

A total system value can be calculated by combining all the individual value sets for the various species into one equation such as,

Total System Value =
$$v_1 v_2 \left(w_3 v_3 + \cdots + w_8 v_8 \right)$$

where

 V_1, V_2 are individual value sets which are factors $V_3 \cdots V_8$ are individual value sets which are terms $W_3 \cdots W_8$ are term weighting functions where $W_3 + \cdots + W_8 = 1$

A sensitivity analysis can be performed on all value sets and value functions if desired. The analysis should indicate which evaluation parameters are most critical to the system value. In addition this analysis may also indicate if the various weighting functions or value set algorithms should be modified.

This technique is of high utility for decision making. However, it is a tool for use in decision making and not a decision maker itself. The ultimate decisions should be made by the experts in the field who have benefited from the logical presentation of available information by means of this structured technique.

A.2.2 Application of the Method to Stratospheric Species Measurement

The evaluation method discussed in the previous section was used in the development of the evaluation techniques applied to stratospheric species measurement. However, two basic changes were made in its present application:

- (1) Incremental values were used in place of smoothly varying value functions
- (2) Two-dimensional value functions were used for each measurement parameter

The first change was indicated by the minimal amount of information available about most species of interest. The second change was made because the quality and quantity of the various measurements were considered to be an important part of the value function development. In a sense, these may be considered as weighting factors on each measurement parameter. In the actual application, these were combined into a common parameter called the data status.

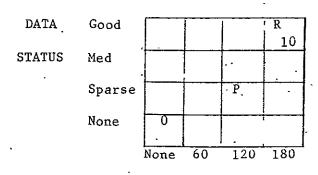
The parameters considered to be of sufficient importance to be included in stratospheric species . . . analysis are:

- Latitude coverage
- Duration of the mission or measurement program
- Diurnal coverage
- Launch date
- Vertical coverage

- Vertical resolution, and
- Longitude coverage

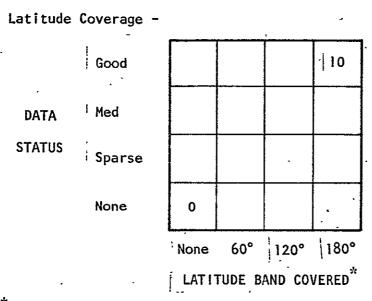
Each of the above parameters must be analyzed and values assigned to the various performance levels from zero to full capability. The measurement scales selected for each parameter are shown in Figures A-2 through A-5.

For each matrix shown, values must be selected for each incremental improvement from no capability for both the parameter and the status of the data up to full capability for both. The general approach is first to determine the level of present knowledge and the required level of knowledge for each species. These levels are then assigned appropriate values from 0 to 10 and the levels beyond and in between these levels are given other appropriate values based upon the present and required knowledge. For example, for the case of latitude coverage for nitric acid vapor, it is known from Section 5.3 and supporting information that nitric acid has been measured in the stratosphere over various latitudes that cover approximately 120°. However, the quantity of data available is very small. Thus the value matrix for nitric acid versus latitude becomes:



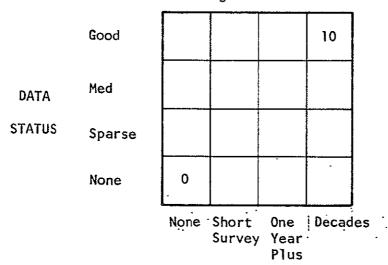
LATITUDE BAND COVERED

Nitric acid vapor,



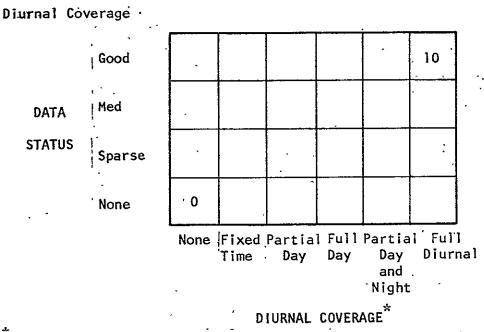
^{*}Includes nadir coverage plus any additional coverage due to orientation of instrument.





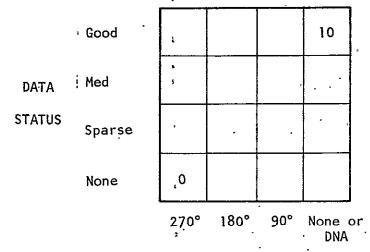
DURATION OF MEASUREMENT PROGRAM

FIGURE A-2
PARAMETERIZATION OF LATITUDE COVERAGE AND PROGRAM DURATION



Based on both orbit and instrument capability.

Launch Date or Beginning of Experiment -

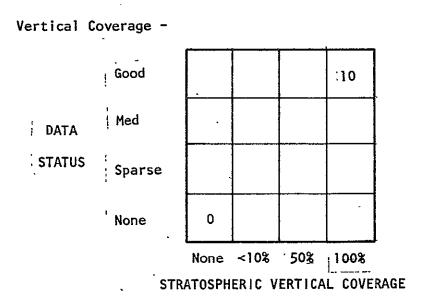


SEASONAL PHASE DEVIATION*

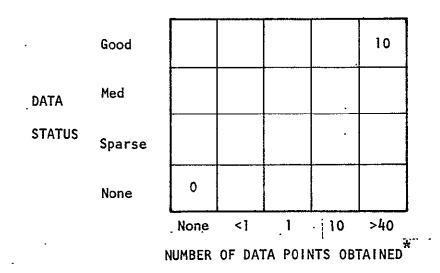
* 90° - Launch is one season prior to desired season
180° - Launch is two seasons prior to desired season
270° - Launch is three seasons prior to desired season
DNA - Launch time not important therefore does not apply

FIGURE A-3

PARAMETERIZATION OF DIURNAL COVERAGE AND TIME OF LAUNCH



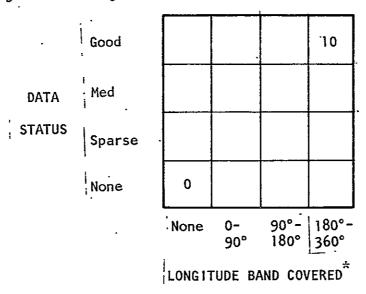
Vertical Resolution -



Note: <1 data point refers to column density through entire atmosphere which provides only part of a stratospheric data point.

FIGURE A-4
PARAMETERIZATION OF VERTICAL COVERAGE AND VERTICAL RESOLUTION

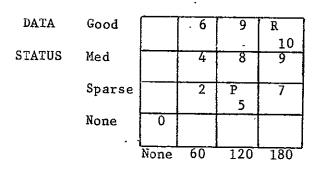




^{*}It is assumed that all orbits being considered for stratospheric pollution missions automatically provide good longitudinal coverage. Therefore mission capability is automatically raised from present knowledge to full capability.

FIGURE A-5
PARAMETERIZATION OF LONGITUDINAL COVERAGE

where the P indicates the present knowledge. Since nitric acid vapor is considered to be one of the very important members of the NO $_{\rm x}$ chemistry chain, requirements (R) have been set at full capability. Values from 0 to 10 are then assigned to each of the matrix areas yielding:



Nitric acid vapor, HNO.

LATITUDE BAND COVERED

These value matrices were prepared for all species prioritized into Groups 1 and 2 plus those in Groups 3 and 4 for which satellite-borne remote sensing instruments either exist or are under development.

The matrices are presented in Volume III of this report.

A.2.3 Weighting Factors

In order to determine the extent (in terms of value) to which each orbit and/or instrument under consideration raises the present knowledge of the species distribution up to or beyond the required knowledge, the capability of the mission for each parameter (i.e., latitude coverage, vertical coverage, etc.) must be known. The values corresponding to the capabilities for each parameter are then combined into the value set for each species which provides a construction.

of how the entire orbit/instrument improves on present knowledge and how it compares with other orbit/instrument missions. However, as indicated in Section A.2.1, simple combination of such values assumes that all of the parameters are of equal importance. This is definitely not true. For any given species some of the parameters are of much greater interest to the user community than others. Thus weighting factors must be assigned for each measurement parameter. For example, in general the latitudinal distribution of stratospheric species is considered to be more important than the longitudinal distribution. Thus, it is more valuable to measure the latitudinal distribution before the longitudinal distribution if both cannot be measured simultaneously. However, if the latitudinal distribution is already well known then the primary value lies in extending knowledge to include the longitudinal distribution.

For most stratospheric species distributions the desirable progression from "no knowledge" to "full knowledge" would be:

- (1) No data
- (2) a. Fixed point data exist (one latitude, longitude, altitude, and time.)
 - Fixed point column burden data exist (one latitude, longitude, and time.)
- (3) Fixed point vertical profile
- (4). Latitude coverage

- (5) Seasonal coverage*
- (6) Diurnal coverage*
- (7) Longitude coverage
- (8) Long time coverage (years or decades)

Thus weighting factors must be assigned to each parameter for each species based on present and required knowledge and the logical progression of desired knowledge given above. High weights should be given to those parameters that would yield the best improvement from present to required knowledge and smaller weights to the other parameters.

The various values for each parameter (adjusted by the weighting functions) are combined to yield the total value for the mission under study. Each mission value is then compared with the value of the present knowledge and the required knowledge. The mission that provides the largest improvement from present knowledge to required knowledge should be considered the "optimum" system. If any mission achieves a value beyond the required knowledge level, the mission value should be truncated at the required knowledge level since this is the goal for each pollutant. However, if several missions achieve approximately equal values then this additional benefit should be acknowledged.

In some cases the mission may show only a small improvement over present knowledge or in fact none at all. Thus, the incremental

^{*} For a few specific species diurnal coverage may be more important than seasonal coverage and possibly latitude coverage.

gain in value over the value of present knowledge would be zero. However, this in no way implies that the entire mission under evaluation has no value at all. At the present state of the art of remote sensing of the stratosphere any successful mission would have value in terms of engineering, technological, and scientific advances. The value derived from the present evaluation only indicates that the mission would not significantly advance our knowledge of the mean stratospheric distribution of the species measured. For this reason, no mission will be given an absolute zero in the actual application of this method. Such cases will be indicated as less than one. In order to evaluate a multiple pollutant or multiple instrument mission the value of each individual orbit/instrument is added to give the total value. In the case where several instruments measure the same pollutant the highest capability for each parameter is used to determine the contributing value. However, in the case of a multi-species mission, simple addition of the individual species values assumes that all are of equal importance. As was discussed in Section 5.2 and again at the beginning of this section, the species have been prioritized. These priorities must be taken into account when comparing the values of different species. This is accomplished by applying weighting factors. These factors have been assigned to the different species groups as follows:

Group la - Direct measurements of climatic 1.0 change and ultraviolet change

Group 1b - Species directly associated with 1.0 changes in climate and/or ultra-Group 2 - Important species associated with 0.9 two or more chemistry chains Group 3 - Components of the basic reactions 0.9 involved in the direct production or depletion of ozone Group 4 - Components of the basic reactions 0.8 indirectly involved in the production or depletion of ozone 0.6 Group 5 - Other significant components of the chemistry chains 0.6 Group 6 - Specific aerosols

The rationale for selecting these factors is as follows. On a scale of 0 to 1 a factor of 1 was given to Groups 1a and 1b since no distinction in importance could be identified. Group 2 rates almost as high due to the fact that the species are involved in more than one major chemistry chain. The Group 3 species are considered to be primary from both the NO and C1 chemistry chains. All of these species are directly related to the ozone generation and destruction reactions. Thus, the weighting remains high. Group 4 species are considered to be secondary in the sense that they are primarily involved in the production of the primary species listed in Group 3. The Groups 5 and 6 species, although very important in stratospheric chemistry, cannot be considered as important as the species in the previous groups. In the actual evaluation an initial set of weights was postulated. This set was exercised against a small set of

species for which relative importance was known with some confidence. From this the final revised set of weights was determined.

The combined values for present and required knowledge for all pollutants for which value matrices were generated are given in Volume III of this report. The combined values also include the parameter weighting functions and the rationale for the selection of each. It should be mentioned, that for the particular stratospheric species and missions considered here, all final values are rounded off to the nearest integer since this is considered to be the maximum preciseness that can be justified by the accuracy of the input values.

APPENDIX B

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NOTE: For the convenience of the user, the same set of references is presented in Volumes I, II and III of this report. Therefore, in any one volume, all references are not cited in the text.

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