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GAS TRANSFER VELOCITIES

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## The Influence of Bubble Plumes on Air-Seawater Gas Transfer Velocities

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### Abstract

Air-sea gas exchange is an important process in the geochemical cycling of carbon dioxide ( $\text{CO}_2$ ). The air-sea flux of  $\text{CO}_2$  is determined in part by the physical forcing functions, which are parameterized in terms of the air-sea transfer velocity,  $k_L$ . Past studies have attempted to correlate  $k_L$  with wind speed,  $U$  (e.g., Liss and Merlivat, 1986; Wanninkhof, 1992). Because strong winds occur in ocean regions thought to be important sources or sinks of  $\text{CO}_2$ , accurate knowledge of  $k_L$  at high  $U$  is important in estimating the global air-sea flux of  $\text{CO}_2$ . Better understanding of the physical processes affecting gas transfer at large  $U$  will increase the accuracy in estimating  $k_L$  in ocean regions with high  $\text{CO}_2$  fluxes. Increased accuracy in estimating  $k_L$  will increase the accuracy in calculating the net global air-sea  $\text{CO}_2$  flux and provide more accurate boundary and initial conditions for global ocean carbon cycle models.

High wind speeds are associated with the presence of whitecaps, which can increase the gas flux by generating turbulence, disrupting surface films, and creating bubble plumes. Bubble plumes will create additional turbulence, prolong the surface disruption, and transfer gas to or from individual bubbles while they are beneath the surface. These turbulence and bubble processes very effectively promote gas transfer. Because of this, Monahan and Spillane (1984) postulated that breaking waves, if present, will dominate non-whitecap related gas exchange. Under this assumption,  $k_L$  will increase linearly with increasing fractional area whitecap coverage,  $W_c$ . In support of this, Asher et al. (1992) found  $k_L$  measured in a whitecap simulation tank (WST) was linearly correlated with bubble plume coverage,  $B_c$  (the laboratory analog of  $W_c$ ). Further evidence supporting the large effect of breaking waves on air-sea gas fluxes is given by the results of Wallace and Wyrick (1992) and Farmer et al. (1993). However, it is not definitively known how the presence of breaking waves and bubble plumes affect the dependence of  $k_L$  on Schmidt number,  $Sc$ , and aqueous-phase solubility,  $\alpha$ . Knowledge of this dependence is necessary to determine the best method for parameterizing  $k_L$  in the presence of breaking waves.

The effects of whitecaps on air-water gas exchange could be studied in detail if  $k_L$  could be measured for reproducible breaking waves for different gases. In the research described here,  $k_L$  values for invasion and evasion of  $\text{CO}_2$ , oxygen ( $\text{O}_2$ ), helium ( $\text{He}$ ), sulfur hexafluoride ( $\text{SF}_6$ ), and dimethyl sulfide (DMS) through clean and surfactant-influenced (SI) water surfaces were measured in a tipping-bucket WST. Reproducible, simulated breaking waves were generated in the WST by releasing a known volume of water vertically onto the tank water surface from a tipping bucket. The bubble populations in the WST are similar to bubble populations measured in the ocean (Asher and Farley, 1995) and the WST can generate  $B_c$  values of up to 0.77%. The WST and the characteristics of the bubble plumes generated in it are described in detail in Asher and Farley (1995).

All measurements were made in seawater that was filtered and sterilized using a flow-through ultraviolet sterilizer. This procedure removed particles, helped prevent biologically produced bubbles, and reduced water surface contamination. Prior to the start of each experiment, accumulated surface contaminants were removed by vacuuming the water surface. For comparison, SI water surfaces were generated by adding 1 ppmv of the soluble surfactant Triton X-100 to the WST. Salinity in the WST was measured by refractometry to be 30 *psu*.

The measurements of  $k_L$  for evasion through a cleaned water surface,  $k_{L,E}(C)$ , show that  $k_{L,E}(C)$  is linearly related to  $B_c$ . Furthermore, the data show  $k_{L,E}(C)$  is very sensitive to small changes in  $B_c$  even for the relatively soluble gases  $\text{CO}_2$  and DMS. For these two gases, an increase in  $B_c$  from 0% to 0.77% increases  $k_{L,E}(C)$  by a factor of seven for  $\text{CO}_2$  and three for DMS. Similarly,  $k_L$  for invasion through a cleaned water surface,  $k_{L,I}(C)$ , was found to be linearly related with  $B_c$ . Invasive transfer velocities also were very sensitive to small changes in  $B_c$  with  $k_{L,I}(C)$  for  $\text{CO}_2$  showing an increase with increasing  $B_c$  similar to that observed for  $k_{L,E}(C)$ .

Bubble-mediated gas transfer models indicate that the gas flux due to bubbles decreases as  $\alpha$  increases (Memery and Merlivat 1985; Woolf and Thorpe, 1992), suggesting that bubble plumes generated by breaking waves will be of little importance in the air-sea exchange of  $\text{CO}_2$ . However, the WST measurements show that the simulated breaking waves are very effective at increasing both  $k_{L,E}(C)$  and  $k_{L,I}(C)$  for  $\text{CO}_2$ . The sensitivity of  $k_{L,I}(C)$  and  $k_{L,E}(C)$  for soluble gases to small changes in  $B_C$  demonstrates that bubble transfer processes are not the only mechanisms by which breaking waves increase the gas flux. The turbulence generated by the plunging water and rising bubbles must make a significant contribution to the overall  $k_L$  measured in the WST. In agreement with Kitaigorodskii (1984), extending this conclusion to oceanic conditions suggests that the turbulence generated by oceanic breaking waves could be an important air-sea transfer pathway even for a relatively soluble gas such as  $\text{CO}_2$ .

Models of bubble-mediated gas transfer suggest that if the bubble gas flux is a significant fraction of the net gas flux,  $k_{L,I}(C)$  measured at a particular  $B_C$  will be larger than  $k_{L,E}(C)$  measured for the same gas at the same  $B_C$  (Memery and Merlivat, 1985; Woolf and Thorpe, 1992). Furthermore, this invasion-evasion asymmetry will be a function of  $\alpha$  with the asymmetry increasing as  $\alpha$  decreases. The WST data support this hypothesis, the ratio of  $k_{L,I}(C)$  to  $k_{L,E}(C)$ ,  $R_{IE}$ , calculated for transfer of  $\text{SF}_6$ ,  $\text{O}_2$ , and  $\text{CO}_2$  decreased as  $\alpha$  increased when bubbles were present in the WST. In the absence of bubbles,  $R_{IE}$  was equal to unity, and no significant invasion-evasion asymmetry was observed.

Transfer velocities measured for invasion and evasion through SI water surfaces,  $k_{L,E}(SI)$  and  $k_{L,I}(SI)$ , respectively, were also found to be linearly correlated with  $B_C$ . The presence of the surfactant caused a significant reduction in both  $k_{L,E}(SI)$  and  $k_{L,I}(SI)$  compared to  $k_{L,E}(C)$  or  $k_{L,I}(C)$ , respectively, for the range of  $B_C$  studied here, the maximum reduction in  $k_L$  occurred at  $B_C=0\%$ . Although the effect of the surfactant is largest when the water surface is not disrupted by breaking waves, the data show that soluble surface active compounds can decrease  $k_L$  even in the presence of breaking waves.  $R_{IE}$  calculated using  $k_{L,I}(SI)$  and  $k_{L,E}(SI)$  for  $\text{SF}_6$  and  $\text{O}_2$  shows that invasion-evasion asymmetry seen in the cleaned-surface case is also found in the SI data. Because this asymmetry is caused by bubble-mediated transfer processes, the similarity in the behavior of  $R_{IE}$  for transfer at cleaned and SI surfaces suggests that the presence of a soluble surfactant does not drastically decrease the importance of the bubble-mediated gas flux.

Using the results from the gas transfer measurements, an empirical parameterization has been developed that can be used to estimate  $k_L$  in the WST from  $B_C$ ,  $Sc$ , and  $\alpha$ . Based on the modelling studies of Memery and Merlivat (1985) and Keeling (1993),  $k_L$  is partitioned into a component due to mechanically generated turbulence,  $k_M$ , a component due to bubble plume turbulence,  $k_T$ , and a component due to bubble-mediated transfer,  $k_B$ . For conditions where the gas is far from equilibrium, this expression has the form

$$k_L = (A_M + B_C(A_T - A_M))Sc^{-n} + B_C \left( \frac{a_1}{\alpha} + b_1 \alpha^{-m_1} Sc^{-n'_1} \right) \quad (1)$$

where  $A_M$  and  $A_T$  are constants determined by the mechanically generated and whitecap-generated turbulence, respectively, and  $n=1/2$  for transfer through a cleaned surface or  $n=2/3$  for transfer through an SI surface. The constants  $a_1$  and  $b_1$  are functions of water surface cleanliness and flux direction (i.e., gas invasion or evasion) and the exponents  $m_1$  and  $n'_1$  are functions of water surface cleanliness. The set of coefficients  $A_M$ ,  $A_T$ ,  $n$ ,  $a_1$ ,  $b_1$ ,  $m_1$ , and  $n'_1$  were determined by nonlinear optimization of Equation 1 to the data for  $k_{L,E}(C)$ ,  $k_{L,I}(C)$ , or  $k_{L,E}(SI)$ .

Comparison of  $k_{L,E}(C)$  estimated using Equation 1 and  $B_C$ ,  $Sc$ , and  $\alpha$  with  $k_{L,E}(C)$  measured in the WST showed the overall accuracy of Equation 1 was  $\pm 10\%$  for  $B_C$  in the range 0% to 0.77% for the gases studied. Equation 1 was able to predict  $k_{L,I}(C)$  with an overall accuracy of  $\pm 5\%$  for the same range of  $B_C$ . Finally,  $k_{L,E}(SI)$  could be estimated with an accuracy of  $\pm 10\%$  using Equation 1. The accuracy of the model-predicted  $k_L$  values is within the experimental uncertainty of the direct measurements of  $k_L$  in the WST for all three cases. This shows that Equation 1 does an excellent job of describing the functional dependence of  $k_{L,E}(C)$ ,  $k_{L,I}(C)$ , and  $k_{L,E}(SI)$  on  $B_C$ ,  $Sc$ , and  $\alpha$  for the conditions present in the WST.

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## References

- Asher, W. E., P. J. Farley, R. Wanninkhof, E. C. Monahan, and T. S. Bates, 1992. "Laboratory and Field Measurements Concerning the Correlation of Fractional Area Foam Coverage with Air/Sea Gas Transport" In *Precipitation Scavenging and Atmosphere-Surface Exchange, Volume 2-The Semonin Volume: Atmosphere-Surface Exchange Processes*, S. E. Schwartz and W. G. N. Slinn, eds., Hemisphere, Washington D.C.: 815-828.
- Asher, W. E., and P. J. Farley, 1995. "Phase-Doppler Anemometer Measurement of Bubble Concentrations in Laboratory-Simulated Breaking Waves" Accepted in *J. Geophys. Res., Oceans*.
- Farmer, D. M., C. L. McNeil, and B. D. Johnson, 1993. "Evidence for the Importance of Bubbles in Increasing Air-Sea Gas Flux" *Nature* 361: 620-623.
- Keeling, R. F., 1993. "On the Role of Large Bubbles in Air-Sea Gas Exchange and Supersaturation in the Ocean" *J. Mar. Res.* 51: 237-271.
- Kitaigorodskii, S. A., 1984. "On the Fluid Dynamical Theory of Turbulent Gas Transfer Across an Air-Sea Interface in the Presence of Breaking Wind-Waves" *J. Phys. Ocean.* 14: 960-972.
- Liss, P. S., and L. Merlivat, 1986. "Air-Sea Gas Exchange Rates: Introduction and Synthesis" In *The Role of Air-Sea Exchange in Geochemical Cycling*, P. Buat-Menard, ed., D. Reidel, New York/New York: 113-127.
- Memery, L., and L. Merlivat, 1985. "Modeling of the Gas Flux Through Bubbles at the Air-Water Interface" *Tellus* 37B: 272-285.
- Monahan, E. C., and M. C. Spillane, 1984. "The Role of Whitecaps in Air-Sea Gas Exchange" In *Gas Transfer at Water Surfaces*, G. H. Jirka and W. Brutsaert, eds., D. Reidel, Hingham/Massachusetts: 495-504.
- Wallace, D. W. R., and C. D. Wirrick, 1992. "Large Air-Sea Gas Fluxes Associated with Breaking Waves" *Nature* 356: 694-696.
- Wanninkhof, R., 1992. "Relationship Between Wind Speed and Gas Exchange Over the Ocean" *J. Geophys. Res.*, 97C: 7373-7382.
- Woolf, D. K., and S. Thorpe, 1992. "Bubbles and the Air-Sea Exchange of Gases in Near-Saturation Conditions" *J. Mar. Res.* 49: 435-466.