

The methanesulfonic acid (MSA) record in a Svalbard ice core

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ABSTRACT. Svalbard ice cores have not yet been fully exploited for studies of climate and environmental conditions. In one recently drilled ice core from Lomonosovfonna, we have studied the methanesulfonic acid (MSA) records in relation to temperature and sea ice. Under the present climatic conditions, MSA appears to be negatively correlated with the sea-ice conditions in the Barents Sea, and positively correlated with the instrumental temperature record from Svalbard. However, prior to about 1920 the MSA concentrations were about twice as high, despite the more extensive sea-ice coverage. After exploring different possibilities, we suggest that MSA concentrations were higher in the 19th century than in the 20th century due to increased primary production, in response to increased vertical stability of the sea surface layers, caused by increased meltwater production from the more extensive sea-ice cover. Thus, the MSA record from Lomonosovfonna probably both is a measure of the regional sea-ice variability on the multi-decadal scale and reflects locally favorable conditions for marine biogenic dimethyl sulfide (DMS) production on the sub-decadal scale.

1. INTRODUCTION

In ice-core studies, methanesulfonic acid (MSA, $\text{CH}_3\text{SO}_3\text{H}$) is used as a proxy for marine biogenic productivity (e.g. Saigne and Legrand, 1987). MSA is a product of marine biogenic dimethyl sulfide (DMS, $(\text{CH}_3)_2\text{S}$) emission and does not appear to have any source other than DMS. In polar waters, MSA production is strongly influenced by sea-ice coverage. Under present conditions in the Arctic Ocean, the highest DMS concentrations are found in open water at the ice edge, and the lowest concentrations occur beneath heavy pack ice (Leck and Persson, 1996). Thus, MSA has been suggested as a potential sea-ice proxy in ice-core studies, and many attempts have been made to link the present sea-ice conditions with the MSA record from high-altitude ice cores from the large ice sheets in Antarctica (Welch and others, 1993; Pasteur and others, 1995; Meyerson and others, 2002; Curran and others, 2003) and Greenland (Whung and others, 1994; Legrand and others, 1997). However, not much is known about MSA in ice cores from smaller ice caps at lower elevations where there could be a more direct link with the source areas than at high-altitude sites. In our previous work with a Svalbard ice core from the ice field Lomonosovfonna (Fig. 1), we have suggested that MSA is related to sea ice and sea surface temperatures (SSTs) in the nearby Barents Sea between 1920 and 1997 (O'Dwyer and others, 2000). At that time, we had not analyzed the complete core. In this paper, we discuss the MSA records from the entire 121 m deep core (Fig. 1), spanning the period from approximately AD 1200 to AD 1997.

2. PHYSICAL SETTINGS

Svalbard is positioned in the zone that includes both the polar ocean and atmospheric fronts between the Arctic Ocean, Nordic Seas and Barents Sea (Fig. 1). The Norwegian Current brings warm water towards the west coast of the

archipelago and into the Barents Sea, and mild air is transported from lower latitudes by southerly to south-westerly winds. This creates a very mild climate over Svalbard, considering its northern position (Hisdal, 1998).

Temperature measurements in Svalbard started in 1911, and by combining several records from the vicinity of Longyearbyen at Isfjorden (Fig. 1) a continuous record has been created (Nordli and others, 1996). The most characteristic feature is the abrupt warming following a temperature minimum in 1917 that lasted until the late 1930s. After a cooler period culminating in the late 1960s, there has been a significant increase in temperature, but Svalbard is still somewhat colder than it was in the 1930s.

Sea ice has a strong effect on Svalbard climate. In a recent study using Arctic temperature records and gridded sea-ice data, a strong link was found between the local climate in Svalbard and the ice edge (Benestad and others, 2002). This connection was particularly strong after 1950, when the quality of sea-ice data was good. The sea-ice reconstructions around Svalbard based on records from whaling and sealing ships compiled by Vinje (1999, 2001) form a valuable climatic archive. His work has shown that sea-ice variability in this area has been large and that in general the period before 1920 was characterized by heavy sea ice interrupted by periods with more modest sea-ice conditions.

In the Barents Sea, the phytoplankton *Phaeocystis pouchetii* dominates DMS emissions (Lancelot and Wassmann, 1994). A major bloom in the Barents Sea usually starts in late April, triggered by a combination of increases in insolation, surface water stability and nutrient input from the melting sea ice. Thus, the bloom follows the retreating sea-ice edge, and can be seen clearly on satellite images (Engelsen and others, 2002). A second bloom occurs somewhat later, in the first half of May, in the permanently ice-free southwestern part of the Barents Sea. In general, there is a positive relationship between plankton growth and warm years. In addition, plankton growth is affected by many other factors such as the position of storm tracks and



Fig. 1. Svalbard with the drill location on Lomonosovfonna. The location of the Austfonna ice core mentioned in the text is also indicated.

the storm intensity. There is also an annual variability, which is partly related to large-scale changes such as the North Atlantic Oscillation and the El Niño–Southern Oscillation (Sakshaug and Walsh, 2000).

3. DESCRIPTIONS OF THE ICE CORE AND ANALYTICAL METHODS

The 121 m deep Lomonosovfonna ice core was retrieved in April 1997 from the highest ice field in Spitsbergen (1230 m a.s.l.) and, according to ice radar measurements, extended to within a few meters of the bed. Dating with a glaciological flow model (Nye, 1963) using the 1963 ^{137}Cs peak (Pinglot and others, 1999) and the Laki (Iceland) 1783 eruption as reference horizons suggests that the core contains at least 800 years of climate and environmental information (Kekonen and others, 2005).

The ice core was subsampled in 5 cm pieces to detect seasonal signals for the best possible dating. The analysis has involved major ions, methylsulfonate (CH_3SO_3^-), acidity (H^+), $\delta^{18}\text{O}$ and δD . More details on the analytical methods can be found in Jauhiainen and others (1999) and Isaksson and others (2001). Although the original ice-core signals have probably been altered to some degree by melt, our analysis suggests that annual or pseudo-annual signals are preserved (Pohjola and others, 2002a), permitting the development of an annual timescale back to at least 1920 using ions (Isaksson and others, 2001), and to the early 18th century using $\delta^{18}\text{O}$ (Pohjola and others, 2002b).

To evaluate the significant trends and cycles in the data, we have used the statistical tool Significant Zero Crossings of Derivatives (SiZer) (Chaudhuri and Marron, 1999; Godtliobsen and others, 2003) on the original data smoothed at different levels. SiZer is a graphical tool that discriminates quantitatively between features that represent real underlying structure in the dataset, and features attributable to noise or sampling variability. In SiZer, significant features are found at different scales (e.g. at different levels of data smoothing).

Table 1. The MSA concentration data in the Lomonosovfonna ice core for different time periods

	1800–99	1900–97	Whole core
Number of samples	278	325	1128
Mean concentration ($\mu\text{g L}^{-1}$)	11.72	7.43	9.33
Median concentration ($\mu\text{g L}^{-1}$)	7.91	4.34	6.70
Minimum concentration* ($\mu\text{g L}^{-1}$)	0.94	0.39	0.39
Maximum concentration ($\mu\text{g L}^{-1}$)	120.9	196.6	196.6
MSA/nss SO_4^{2-} (%)	19.5	4.89	13.7
MSA/(MSA+nss SO_4^{2-}) [†] (%)	22.6	4.45	18.5

*Minimum value is the smallest detectable value.

[†]Ratios are calculated using $\mu\text{g L}^{-1}$.

Color coding is used to visualize the statistical significance of the data trends at various smoothing timescales. This is controlled by the size of the smoothing window (bandwidth, h) and location (x) of the signal. For each scale and location of the signal, SiZer tests whether the smoothed series has a derivative significantly different from zero.

4. RESULTS AND DISCUSSION

The mean MSA concentration for the whole Lomonosovfonna ice core is $9.33 \mu\text{g L}^{-1}$ (determination limit (DL) = $0.39 \mu\text{g L}^{-1}$ and $n = 1128$). For the period 1900–97 the mean concentration is $7.43 \mu\text{g L}^{-1}$ ($n = 325$), and for 1800–99 it is $11.72 \mu\text{g L}^{-1}$ ($n = 278$). The MSA concentrations are shown in Table 1. The complete MSA record from Lomonosovfonna is shown in Figure 2 together with the results from the SiZer analysis.

The MSA fractions (MSA/MSA+non-sea-salt (nss) SO_4^{2-}) range from 4% to 23% (Table 1), which are similar to those found in Greenland (Whung and others, 1994; Legrand and others, 1997) and from aerosol sampling in the atmospheric boundary layer in the ocean west of Svalbard (Leck and Persson, 1996) and in Ny-Ålesund (Teinilä and others, 2003). This implies a strong influence of high-latitude sources for the MSA under present conditions.

4.1. Seasonality of MSA in recent Svalbard snow and air

In aerosol studies from Ny-Ålesund, MSA seems to peak in mid-summer (Hara and others, 1997), most likely due to the phytoplankton bloom west of Svalbard. Furthermore, the marine air masses from the west have higher concentrations of MSA and sea salts than the colder, drier Barents Sea air, and during some winter months MSA was found in air associated with coarser particles, suggesting transport from middle latitudes (Hara and others, 1997). In another study, no MSA was found above the detection limit during the polar night (Teinilä and others, 2003).

Snow-pit studies performed during two winters on Lomonosovfonna show that high concentrations of MSA begin to appear in the snow in April (K. Virkkunen, unpublished information) at the same time as the start of the algae bloom. Prevailing easterly winds suggest that conditions in the Barents Sea should have a strong influence on the amount of MSA deposited on Lomonosovfonna.

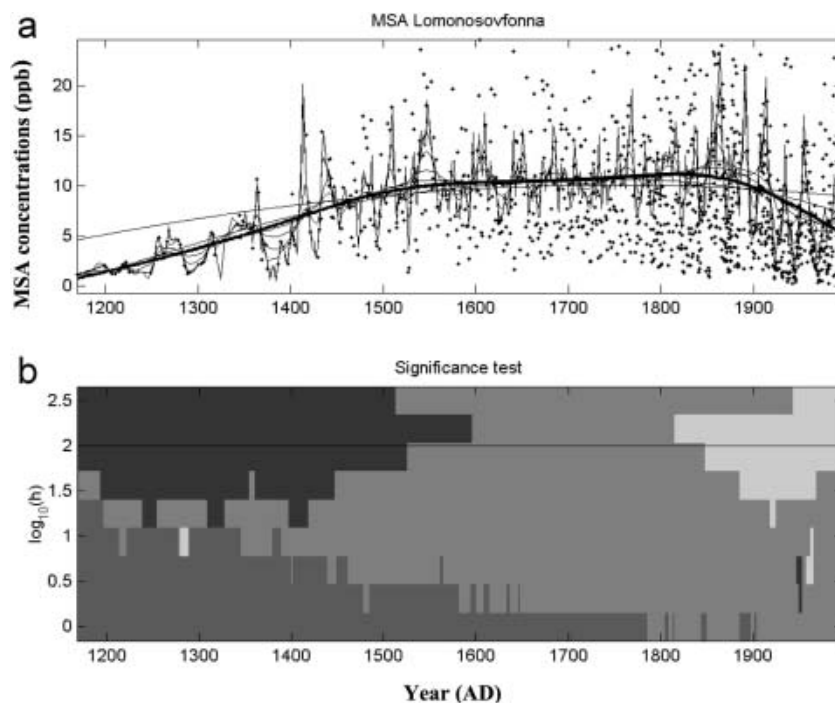


Fig. 2. (a) The MSA concentration record (in ng g^{-1}) from Lomonosovfonna, AD 1200–1997, with various smoothings. The dots shows all the data points. The thick line corresponds to the 100 year smoothing time window, indicated by a line in (b). (b) The significance of the trends at different smoothing scales (y axis) shown in a greyscale for the 90% level. h is the time window in which the smoothing operation is performed. The SiZer analysis suggests a significant increase in MSA concentrations from the beginning of the record until about AD 1600 and a decrease from about AD 1800 to the present. Color coding: black: increasing trend; dark grey: not enough data; medium grey: no trend; light grey: decreasing trend.

4.2. The MSA record after 1900

As already mentioned, our previous work has shown that the MSA and the sea-ice records over the 1920–97 period are closely related (O'Dwyer and others, 2000). MSA concentrations appear to be higher for warm years with reduced ice cover in the Barents Sea, which is probably related directly to the availability of a productive source, where less sea ice and warmer water promotes more DMS production and thus more MSA. A greater influx of Atlantic water in low-sea-ice years may be responsible for this. Analysis by O'Dwyer and others (2000) suggests that MSA concentrations in Lomonosovfonna snow correlate better with Barents Sea SST and ice cover than with conditions west of Svalbard. The same study also hypothesized that MSA might indirectly reflect larger-scale changes in the region via Atlantic Ocean input driven by the positive state of the Arctic Oscillation increasing the inflow of warm Atlantic water to the Barents Sea (Dickson and others, 2000). Thus, previous results suggested that MSA variations may be used as a proxy for past local climate changes on a decadal scale.

There is a close relationship between nssSO_4^{2-} and MSA concentrations after about 1920 ($R = 0.7$) (Kekonen and others, 2005) (Fig. 3), whereas between 1900 and 1920 there is no correlation.

4.3. The MSA record before 1900

The MSA record shows that the highest concentrations occurred during the mid-19th century (Fig. 2). The SiZer analysis suggests that there was a significant increase in MSA concentrations from the beginning of the record to about AD 1600 and a significant decrease after about AD 1800 (Fig. 2). The decadal MSA and nssSO_4^{2-} concentration profiles show relatively weak correlations during the 19th

century ($R = 0.6$) (Fig. 3), but their ratio is different compared with the post-1920 period. MSA/ nssSO_4^{2-} ratios are 19.5% during the 19th century, and 2.5% from 1920 to 1997 (Fig. 4a). During the 17th and 18th centuries, there are no correlations between MSA and nssSO_4^{2-} (Fig. 4b). During the 15th and 16th centuries, weak correlations are evident, but the reduced correlations are probably due to post-depositional movement rather than climatic reasons because other ions are also weakly correlated with MSA (Kekonen and others, 2005).

The deepest 10 m of the ice core, dated to the period AD \sim 1200–1400, shows low MSA concentration and has a very different character than the rest of the core. All the ions except ammonium have anomalously low concentrations. Possible explanations for this include migration of ions from

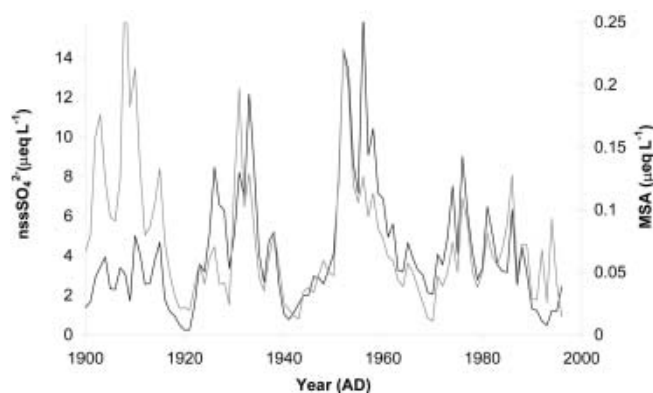


Fig. 3. The relationship between nssSO_4^{2-} (black line) and MSA (grey line) concentrations, 1920–97.

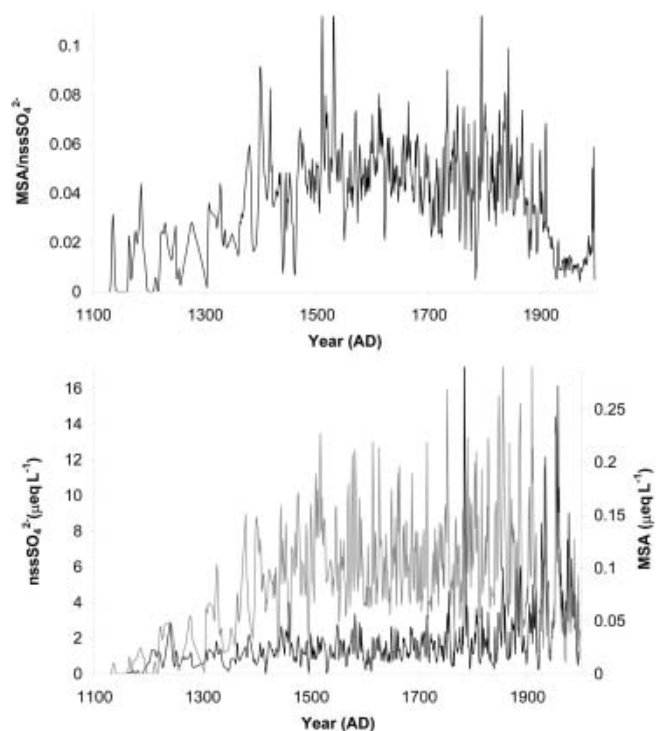


Fig. 4. (a) The MSA/nssSO₄²⁻ ratio over the whole time period in the ice core. (b) The nssSO₄²⁻ (black line) and MSA (grey line) concentrations over the whole core length.

the deepest and warmest ice to the bed, or loss of ions in the original snowpack caused by warm conditions prior to about 1300. Calculations of likely flow rates to the bed and snow-pit analyses over recent seasons suggest that loss of ions by runoff from the original snowpack is the most likely explanation (Kekonen and others, 2005). However, we want to emphasize that this only applies to the lowermost part of the ice core and does not affect the trends we are discussing in this paper.

Because the sea ice was more extensive before 1920 (Vinje, 2001) (Fig. 5), the negative MSA–sea-ice relationship after 1920 does not apply before 1920. Despite the higher sea-ice concentrations, the MSA concentrations were more than twice as high before 1920 as after 1920. Since this is not true for all the ion species (Kekonen and others, 2005), it is not purely an effect of more vigorous circulation, as has been suggested for the atmosphere over Greenland during the Little Ice Age (O'Brien and others, 1995; Kreutz and others, 1997), nor can it be an effect of changes in accumulation rate, since our accumulation record does not suggest any such increase (Pohjola and others, 2002b).

There could be many possible reasons for the concentration pattern we find, and below we discuss the ones we find most likely.

Source change

Because sea-ice conditions were more severe during the 19th than during the 20th century (Fig. 5), it is possible that there was a stronger contribution from a DMS source in mid-latitudes with higher production than in Arctic waters. The decrease in MSA over the last 200 years is in agreement with the high-resolution MSA record from the Greenland interior (Whung and others, 1994; Legrand and others, 1997) (Fig. 6), suggesting that there may have been a

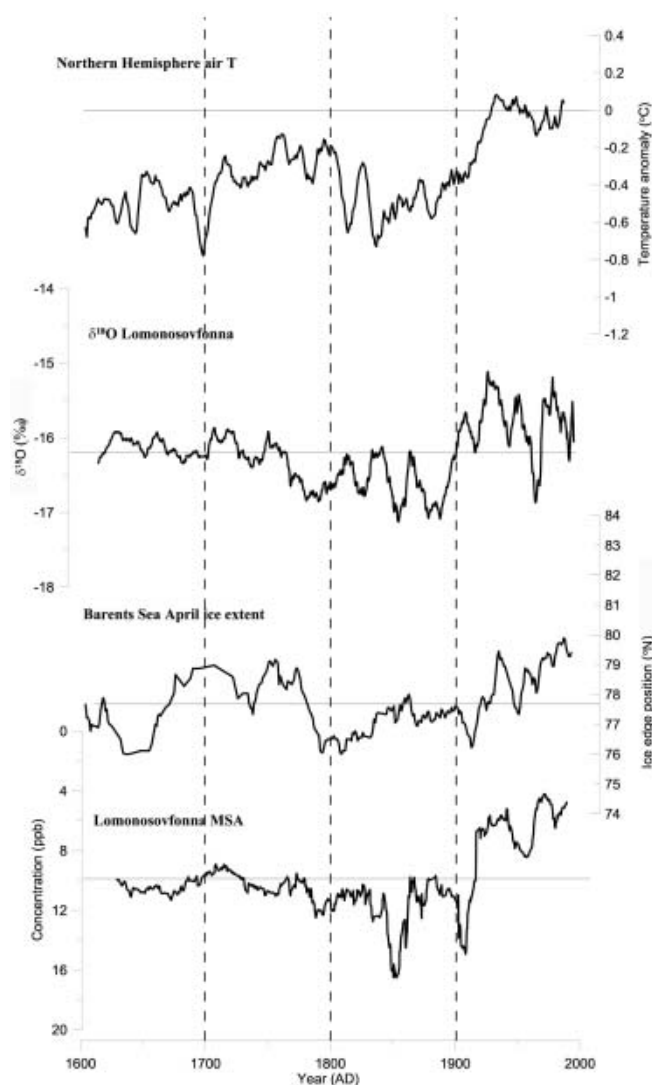


Fig. 5. Comparisons of the $\delta^{18}\text{O}$ (Isaksson and others, 2003) and MSA records from Svalbard, sea ice from the Barents Sea (Vinje, 1999) and Northern Hemisphere air-temperature anomaly record (Jones and others, 1998). All records are smoothed to an equivalent of a 10 year running mean.

common source and/or similar mechanisms at work on Svalbard as well. Because the Greenland sites are at high elevations, it is likely that the source is partly local (the Greenland Sea) and partly distant (mid-latitudes) (Legrand and others, 1997 and references therein). In this context, it is of interest to note that Fischer (2001) found a 30% enhanced sea-salt concentration during the 19th century in ice cores from North Greenland. Recent weather patterns suggest that the Pacific is an important source region for North Greenland, and Fischer (2001) proposes that the sea-salt increase during the 19th century was related to an enhancement of sea-salt export from the Pacific. Also, Kahl and others (1997) concluded, based on trajectory studies, that long-range transport might be responsible for increased MSA aerosols in the air. As already mentioned, the present conditions indicate the highest MSA concentrations in the air at Ny-Ålesund with westerly winds (Hara and others, 1997). The similarities between Greenland and Svalbard data may also reflect large-scale changes in marine productivity or in the oxidative capacity of the atmosphere (Legrand and others, 1997).

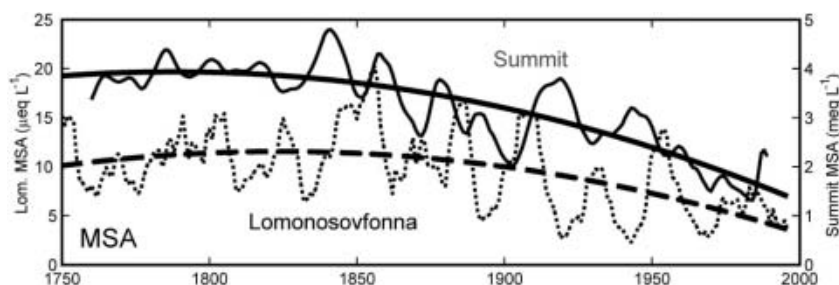


Fig. 6. The smoothed MSA records from Lomonosovfonna (stippled) and Summit (black), Greenland (Legrand and others, 1997). Both records show a decrease from the mid-19th century to the present, suggesting that this might be a regional signal.

Sea ice

As already mentioned, the highest concentrations of DMS are found in open water at the ice edge, and the lowest concentrations occur beneath heavy pack ice under present conditions in the Arctic Ocean (Leck and Persson, 1996). Thick ice cover prevents biogenic production, but the increased meltwater release from the sea ice during spring and summer increases the stability of the water column. This holds the phytoplankton in the euphotic zone, and, in combination with the favorable light conditions, creates a highly productive environment (e.g. Sakshaug and Walsh, 2000). Therefore, it is likely that an increased length of the ice edge would cause increased biomass production. In addition, it is possible that, given the generally colder temperatures, the sea ice was also thicker. This would lead to increased production of meltwater and thus favorable conditions for the phytoplankton.

Storminess

Under the present conditions, there is a positive relationship between storm activity and primary production in Atlantic waters due to increased vertical mixing (Sakshaug and Walsh, 2000 and references therein). If this was true for the past, the more stormy conditions that supposedly prevailed during the 19th century (Lamb, 1995) may have increased DMS production and may have also increased aerosol incorporation. However, since the sea-salt records do not show such increased concentrations during this period (Kekonen and others, 2005), it seems more plausible that the MSA increase is due to increased primary productivity. However, present atmospheric aerosol measurements of MSA at Ny-Ålesund suggest that MSA is not affected by increased cyclonic activity in the same way as sea salts, and that high concentrations of MSA are associated with warm, humid storm conditions (Hara and others, 1997). Cyclonic storm conditions increase the coarse fraction of sea salt but not MSA, and the abundance of MSA in the dark winter months is due to transport from mid-latitudes (Hara and others, 1997).

Seasonality

A change of storm season could directly affect the incorporation of DMS into the atmosphere. More prevalent storms in the spring/summer season during the 19th century but more frequent winter storms in the 20th century would explain the higher MSA concentrations in the 19th-century ice. However, one can assume that this would also affect other ions, and there is no such indication (Kekonen and others, 2005). Another scenario could be a longer emission

season for DMS due to favorable conditions, i.e. melting sea ice. With the data available, it is difficult to find such evidence in our cores.

Change in branching ratio

Colder air temperatures favor MSA oxidation which changes the branching ratio (Seinfeld and Pandis, 1998). Since the air temperatures were lower in the 19th century than after 1920, this may explain the higher MSA concentrations in the ice-core data before 1920. MSA/nssSO₄²⁻ ratios are clearly different after 1920 (2.47%) than during the 19th century (19.5%) (Fig. 4a). However, increases in nssSO₄²⁻ concentrations after 1920 are also affected by anthropogenic sources.

After considering the above possibilities, we believe that the most plausible explanations for the high MSA concentrations during the 19th century are a change of source and/or more favorable growing conditions for the DMS-producing algae in a more extensive sea-ice environment. Vinje (2001) suggested that the April sea-ice extent has decreased by 33% since 1864. That decrease coincides with the high MSA concentrations in the ice cores (Fig. 5) and it is therefore tempting to believe that sea ice and MSA are connected.

4.4. Discussion

As mentioned in the introduction, the relationship between MSA and sea ice has been discussed in several papers concerned with recent conditions in Antarctica. When comparing MSA records from different ice cores, it is evident that the transfer function between DMS production and the MSA in ice cores is not straightforward. Most recently, Curran and others (2003) found a clear positive correlation between the MSA record from Law Dome and local sea-ice extent. Welch and others (1993) found a positive relationship between MSA and sea-ice extent in the Ross Sea, while Pasteur and others (1995) found the opposite effect when comparing an ice core from the Antarctic Peninsula with the Weddell Sea ice. On the glacial-to-interglacial scale, the MSA was higher during glacial times in the Vostok ice core (Legrand and others, 1988). This has been attributed to more favorable conditions for DMS-producing *Phaeocystis pouchetii* (Gibson and others, 1989) with extensive sea ice. In contrast, ice cores from Greenland suggest lower MSA during glacial times (Hansson and Saltzman, 1993), possibly related to sea-ice cover over local DMS-producing sources.

The δ¹⁸O records from both Lomonosovfonna and Austfonna suggest that the 19th century was the coldest

during the past 600 years (Isaksson and others, 2003), as is also the case for other proxy records of air temperatures (Jones and others, 1998) (Fig. 5). A comparison of the ice-core and sea-ice records suggests that sea-ice extent and Austfonna $\delta^{18}\text{O}$ are related over the past 400 years (Isaksson and others, 2003). This may reflect the position of the storm tracks and their direct influence on the relatively low-altitude Austfonna. Lomonosovfonna may be less sensitive to such changes and may primarily record atmospheric changes due to its higher elevation. Thus, it appears that lower-altitude ice cores might better reveal a sea-ice connection, although currently there are relatively few such records available. In a study of sulfate and MSA in air and snow at two sites on Greenland, the variations of the MSA concentrations did not differ between a coastal and an inland site (Jaffrezo and others, 1994). Svalbard ice cores might, together with some of the other lower-lying Arctic ice caps such as Zemlya Frantsa Iosifa and Severnaya Zemlya, provide a link between MSA and sea ice.

There is variability on a 20–25 year scale in the MSA data from about 1800 to the present. Such variations have also been described from Greenland ice cores, and no correspondence with known climatic oscillations has been found there (Whung and others, 1994). At this point, we do not have any good suggestion to explain this 20–25 year variability.

Clearly, more knowledge about the whole DMS–MSA oxidative process, the incorporation of MSA into air–snow, and the depositional and post-depositional process is required (cf. Pasteur and Mulvaney, 1999).

5. CONCLUSIONS

The exploration of MSA records from one Svalbard ice core suggests that sea-ice extent has a dominant effect on MSA concentrations in the snow/ice. The ice core suggests lower MSA concentrations during the 20th century than during the 19th century. Between 1920 and 1997, there is a negative correlation between sea ice, SST in the Barents Sea and MSA on sub-decadal scales in the Lomonosovfonna snow, but our data suggest that, before 1920, sea ice affected MSA in a positive direction, i.e. periods with large sea-ice extent coincide with high MSA content in Svalbard snow. The same general pattern is apparent in the available MSA data from Greenland spanning the same time period, suggesting that MSA reflects processes on a regional scale. However post-depositional processes modify the local records, making detailed comparisons impossible.

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