

The SCEC Southern California Reference Three-Dimensional Seismic Velocity Model Version 2

by Harold Magistrale, Steven Day, Robert W. Clayton, and Robert Graves

Abstract We describe Version 2 of the three-dimensional (3D) seismic velocity model of southern California developed by the Southern California Earthquake Center and designed to serve as a reference model for multidisciplinary research activities in the area. The model consists of detailed, rule-based representations of the major southern California basins (Los Angeles basin, Ventura basin, San Gabriel Valley, San Fernando Valley, Chino basin, San Bernardino Valley, and the Salton Trough), embedded in a 3D crust over a variable depth Moho. Outside of the basins, the model crust is based on regional tomographic results. The model Moho is represented by a surface with the depths determined by the receiver function technique. Shallow basin sediment velocities are constrained by geotechnical data. The model is implemented in a computer code that generates any specified 3D mesh of seismic velocity and density values. This parameterization is convenient to store, transfer, and update as new information and verification results become available.

Introduction

The dense population and active tectonics of southern California necessitate extensive seismic hazard evaluations that include precise earthquake location determinations, path and site effect studies, and strong ground motion simulations. These studies require a realistic three-dimensional (3D) seismic velocity model defined on spatial scales appropriate for each application. Here we describe a 3D seismic velocity model for southern California assembled from geological and geophysical data and designed to serve as a reference model for multidisciplinary research activities in the area.

A velocity model, to be widely useful, must integrate data from multiple disciplines, including seismic imaging, geologic mapping, and geotechnical investigations, in order to capture the wide range of spatial scales that are important for both basic research and earthquake hazard applications. Consider, for example, the problem of deterministic 3D modeling of long period (>1 sec) strong ground motion in southern California. Regional seismic tomography provides 3D seismic velocity information with resolution on the order of tens of kilometers (Magistrale *et al.*, 1992; Zhou, 1994; Hauksson, 2000). This resolution is useful for modeling the propagation of long-period seismic waves in crystalline basement rocks outside of the sedimentary basins, where wavelengths are long and velocity variations are relatively small. In the basins, however, much higher spatial resolution is required: basin depths are typically less than 10 km, and seismic velocities vary dramatically. In the low-velocity basin sediments, 1-sec *S* waves have wavelengths ranging from

a few kilometers in the deep basins, down to only a few hundred meters in the shallow basin layers. Also, important amplification and interference effects are likely to be localized near the basin edges, which therefore need to be well resolved. Geologic mapping, geotechnical investigations, and borehole velocity logs can provide the necessary high spatial resolution.

The Southern California Earthquake Center (SCEC) has supported an effort to develop a standard 3D reference model for southern California. The designation “reference model” is meant to emphasize the following characteristics. (1) The model incorporates contributions from multiple types of data. (2) It represents a standard agreed to among a large number of researchers working in southern California, against which anomalies (in, e.g., seismic travel times, waveforms, and amplitudes; gravity; and borehole data) can be identified, quantified, and compared. (3) The model description is reviewed and maintained by SCEC and made widely available, and its periodic revisions are documented and tracked by version number. (4) By integrating a large, diverse body of both seismic and nonseismic data, the reference model provides a starting model for application of perturbative approaches to the 3D inversion of seismic travel time and waveform data. A prototype reference model (Magistrale *et al.*, 1996; we will refer to it as Version 0) has been widely used for simulating ground motions from past earthquakes (e.g., Wald and Graves, 1998) as well as for estimating basin effects from potential future earthquakes (e.g., Olsen *et al.*, 1996). The need for a single standard reference

model motivated the much more comprehensive model development reported here.

Version 2 of the SCEC reference model consists of detailed, rule-based representations of the major southern California basins (Fig. 1) embedded in a 3D crust over a variable depth Moho. The model includes the populated Los Angeles area basins (Los Angeles basin, Ventura basin, San Gabriel Valley, San Fernando Valley, Chino basin, and San Bernardino Valley), and the Salton Trough. The basins are parameterized as a set of objects (constructed from geological, geophysical, and geotechnical data) and rules implemented in a computer code that generates any specified 3D mesh of seismic velocity and density values. This parameterization is convenient to store, transfer, and update as new information and verification results become available. It allows any distribution of velocities; for example, fast-over-slow velocities are easily modeled. A fine spatial resolution is achieved by the use of geologic information to constrain the locations and ages of structural and stratigraphic boundaries. Outside of the basins, the model crust is based on regional tomographic results. The model Moho is represented by a surface with the depths determined by the receiver function technique.

Several studies in this special volume (Field, 2000; Lee and Anderson, 2000; Olsen, 2000; Steidl, 2000) use Version 1 of the SCEC reference model (Magistrale *et al.*, 1998). Version 1 contains the Los Angeles area basins in a 1D crust over a constant depth Moho. The Version 1 model improved the Version 0 model of Magistrale *et al.* (1996) by adding the Ventura basin, Chino basin, and San Bernardino Valley, and revising the San Fernando Valley. Version 2 is an ad-

vance over Version 1 in that it includes the Salton Trough, a 3D distribution of crustal velocities outside of the basins, a 3D Moho, and detailed shallow basin velocities from geotechnical logs. The ground-motion simulations reported in this volume (Olsen, 2000) focused on the Los Angeles area basins and imposed a V_S lower bound of 1 km/sec (due to computational limitations), so the conclusions based on those simulations would be little affected by the Version 2 modifications. Basin depth effects on ground motion reported in this volume (Field, 2000; Lee and Anderson, 2000; Steidl, 2000) use the depth to the 2.5 km/sec V_S isovelocity surface to define basin depth. In the Los Angeles area basins that isovelocity surface is the same in the Version 1 and Version 2 models.

Model Construction

Reference Surfaces and Rule Definition

In the model sedimentary basins, V_P is determined by the application of empirical rules to interpolate properties from the model objects, and density and V_S are derived from V_P . Outside and below the basins, V_P and V_S are assigned by interpolation from the regional tomographic results of Hauksson (2000). Within the basins, V_P and V_S in the top 300 m are constrained by geotechnical borehole seismic velocity data. Where V_P and V_S are independently specified, the density is derived from V_P .

There exists a great deal of information about the age and depth of the sediments in the Los Angeles area basins from oil and water exploration activities and other geologic

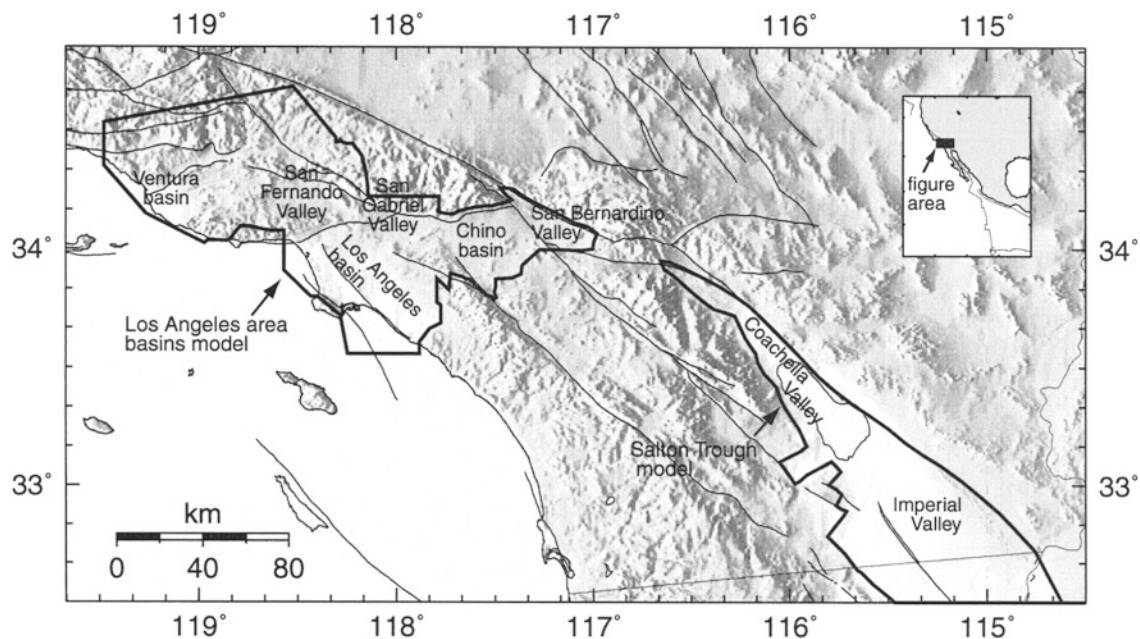


Figure 1. Location map of southern California showing the extent of the basin models (heavy black lines) and basin names. Light black lines are faults. Inset shows location of figure area; western North America is shaded.

studies (Fig. 2). From this information, we define reference surfaces (objects) of known depth and age in the detailed portion of the model representing the sedimentary basins. We examine structural cross sections and maps to define widespread, well-defined reference surfaces representing stratigraphic horizons, sediment-basement contacts, and faults (many of the surfaces are in multiple pieces). The maps and cross sections are digitized, and the reference surfaces are carefully interpolated and resampled on regular grids with a spacing of 100 to 300 meters. Uplift of each reference surface is estimated, or sometimes has been explicitly mapped (e.g., Wright, 1991).

Faust (1951) examined well surveys from North America and determined an empirical relation between sediment age, depth, and P -wave seismic velocity:

$$V_P = k(da)^{1/6} \quad (1)$$

where V_P is P -wave velocity, d is the maximum depth of burial of the sediments, a is the sediment age, and k is a constant. The one-sixth power reflects the tendency of sediments to compact as they are buried and to indurate as they age (Dobrin, 1976). Age at any point in a basin can be interpolated from the reference surfaces. The constant k is calibrated for each reference surface by comparison to oil well sonic logs and seismic refraction surveys. At each point of interest within a basin (defined by a latitude, longitude, and depth) for which the velocity is desired: (1) The age and k of the point are interpolated by comparing the point depth to the depths, ages, and k values of the reference surfaces at the same latitude and longitude. (2) The maximum depth of burial is found by correcting the current depth by any known amount of uplift. (3) V_P is determined from the Faust equation. (4) Other physical parameters are derived: density is found from V_P using the relation of Nafe and Drake (1960); density is used to find Poisson's ratio with the relation of Ludwig *et al.* (1970); V_S is calculated from the P -wave velocity and Poisson's ratio.

The seismic velocity structure of the Salton Trough has been characterized by several seismic refraction studies (Fuis *et al.*, 1982, 1984; Mooney and McMechan, 1982; Parsons and McCarthy, 1996). Thus, instead of constructing reference surfaces from sediment stratigraphy information, the Salton Trough is modeled by digitizing V_P cross sections derived from the seismic refraction lines (Fig. 2) and converting the cross sections into isovelocity surfaces. At a point of interest, V_P is interpolated from the isovelocity surfaces, and the other properties are derived from V_P as in step 4 mentioned previously.

Los Angeles Basin and San Gabriel Valley

Wright (1991), in an extensive summary, presents structure-contour maps of two widespread sedimentary stratigraphic horizons: the base of the Repetto Formation, about 4.5 Ma; and the base of the Mohnian Stage, about 14 Ma. Age control of the stratigraphic horizons is from microfossils

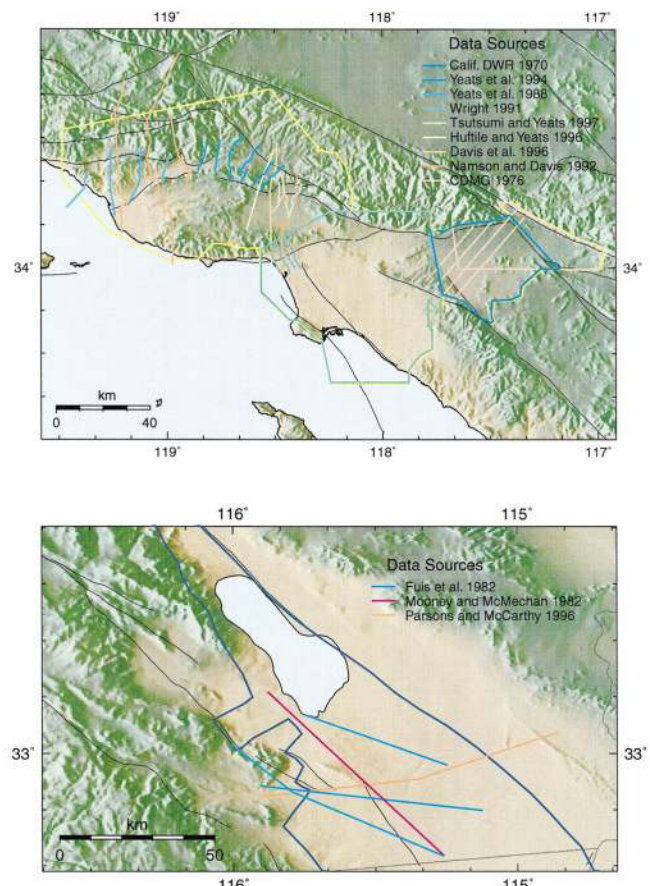


Figure 2. Sources of the information used to construct the basin model reference surfaces for the Los Angeles area (top panel) and Salton Trough (lower panel).

(e.g., Blake, 1991). Wright (1991) also presents a contour map of the amount of uplift during the Pasadenan deformation (3.5 Ma to present); we use this information to correct current sediment depths to depth of maximum burial. McCulloh (1960) and Yerkes *et al.* (1965) show a structure-contour map of the top of crystalline basement rocks inferred mainly from gravity data. The age we use for this horizon is not the rock age, but rather an early Miocene age (20 Ma) that just predates the development of major basement relief and so dates the base of the sediment fill. The age and distribution of material at the ground surface is indicated on California Division of Mines and Geology (CDMG) geologic maps (Jennings, 1962; Rogers, 1965, 1967; Jennings and Strand, 1969).

The Santa Monica area within the Los Angeles basin is of particular interest to strong-motion modelers because of the unexpectedly high damage to the area from the Northridge earthquake (e.g., Gao *et al.*, 1997). Wright (1991) shows four detailed cross sections that we use to refine the Mohnian, Repetto, and basement surfaces in that area.

We calibrate the model by adjusting the constant k in the Faust relation (equation 1) to match seven oil well sonic

logs (Fig. 3) in the Los Angeles basin and the San Gabriel Valley (Brocher *et al.*, 1998). In the Los Angeles basin, $k = 197$; in the San Gabriel Valley, $k = 218$. The sonic logs indicate a V_P inversion within the sediments of the San Gabriel Valley. The inversion starts at a constant fraction (0.6) of the depth to the Mohnian reference surface and reaches a constant ~ 1250 m/sec about 400 m deeper. The inversion is modeled by subtracting the 1250 m/sec from the calculated velocities, tapering the subtraction over the top 400 m of the inversion.

This version of the Los Angeles basin and the San Gabriel Valley differs from Version 0 in the different values of k calibrated from the oil well sonic logs, the San Gabriel Valley velocity inversion, and the Santa Monica area details. The current Los Angeles basin and San Gabriel Valley in Version 2 are the same as in Version 1, except for the geotechnical constraints described subsequently.

San Fernando Valley and Ventura Basin

The San Fernando Valley and the Ventura basin share similar stratigraphy and so are considered together. Yeats *et al.* (1988, 1994), Namson and Davis (1992), Huftile and Yeats (1996), Davis *et al.* (1996), and Tsutsumi and Yeats (1999) present structural cross sections of the San Fernando Valley and the Ventura basin from which we define a total of 12 reference surfaces in 57 pieces. The lateral extent of the reference surfaces at the Earth's surface is from a CDMG geologic map (Jennings and Strand, 1969).

The 11 reference surfaces in the Ventura basin have ages of 0.5, 0.975, 1.5, 2.3, 5.0, 24, 37, 47, 67, 75, and 100 Ma; lacking independent calibration, we set $k = 180$ for all those surfaces to produce model velocities in the deepest sediments approaching the velocities of the surrounding basement rock. In the San Fernando Valley, the seven reference surfaces have ages of 2.0, 2.3, 5.0, 37, 67, 75, and 100 Ma. Four oil well sonic logs (Fig. 3) are available in the San Fernando Valley (Brocher *et al.*, 1998); from these we determine a different k for each reference surface ($k = 189, 189, 160, 180, 123, 180, 180$, respectively). We correct current sediment depth to maximum depth of burial by calculating the average depth of each reference surface and, because the strata are deformed largely by relatively recent (< 1 Ma, e.g., Huftile and Yeats, 1995) activity, assume any depth above the average depth was formerly at least as deep as the average. If the current depth is below the average depth, the current depth is used as the maximum depth of burial.

This version of the San Fernando Valley supplants the Version 0 model. It uses entirely new reference surfaces, and new k values calibrated to oil well sonic logs in the valley. The Version 0 model did not include the Ventura basin. The Version 2 San Fernando Valley and Ventura basin are the same as in Version 1, except for the geotechnical constraints described subsequently.

San Bernardino and Chino Basins

The Chino and San Bernardino basins are shallow (generally < 1 km deep) basins filled mostly with terrestrial sed-

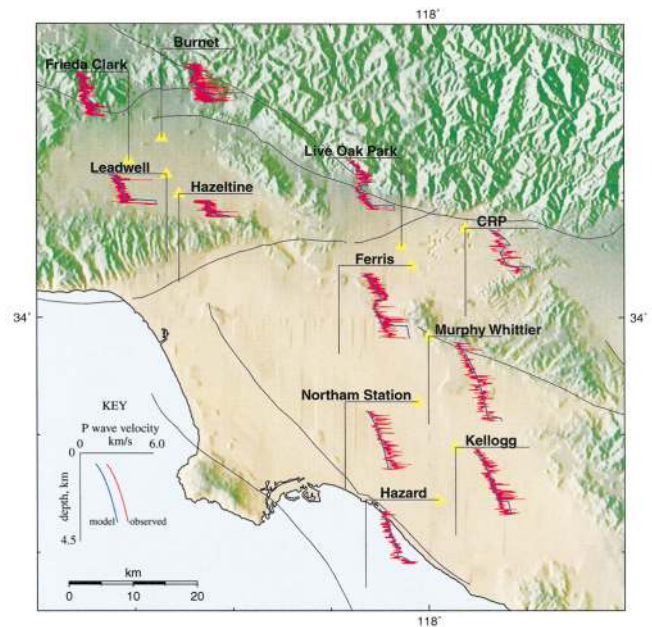


Figure 3. Oil well sonic logs (red) from Brocher *et al.* (1998) in the Los Angeles basin, San Gabriel Valley, and San Fernando Valley used to calibrate the model (blue). Yellow triangles indicate oil well locations.

iments. We use structural cross sections and maps of the depth to the base of water-bearing strata from Department of Water Resources (1970) and Fife *et al.* (1976) to define three reference surfaces: a 14.5 Ma Mohnian and a 6.0 Ma Miocene (both limited to the westernmost portion of the Chino basin), and the base of the water bearing strata. The age and distribution of material at the ground surface is from CDMG geologic maps (Rogers, 1965; 1967).

Hadley and Combs (1974) obtained a seismic refraction profile in San Bernardino basin. We note that the top of their 2.9 km/sec V_P layer corresponds to the base of the water bearing strata, and we interpret the top of that 2.9 km/sec layer to correspond to the top of weathered crystalline basement rock. Below the 2.9 km/sec layer, Hadley and Combs (1974) defined a 5.3 km/sec layer that we interpret to represent hard rock, and we define a hard rock reference surface at a constant depth below the weathered basement surface to mark the bottom of the basin. We compare model velocity profiles to the seismic refraction profile and calibrate the model by adjusting the nominal ages of the weathered and hard basement surfaces (while keeping k fixed at 180) to match the refraction results. The final ages are 6.0 and 16.5 Ma, respectively.

Frankel (1993) combined the Hadley and Combs (1974) refraction profile and water well logs to develop a model of the San Bernardino basin to use in ground-motion simulations. That model used the base of water bearing strata in the well logs and the top of the 5.3 km/sec refraction profile layer to define the top of the basement, and thus is dominated

by a deep basement trough at the refraction profile site. The current basin model differs greatly by having a relatively flat bottom because of our identification of the top of the 2.9 km/sec layer as the base of the water-bearing strata. The Version 0 model did not include the San Bernardino and Chino basins; these basins in Version 2 are the same as in Version 1, except for the geotechnical constraints described subsequently.

Geotechnical Constraints

It is desirable to have well-constrained, detailed shallow properties (<300 m depth) in the model because (1) shear waves in low-velocity, shallow sediments at frequencies relevant to engineered structures have wavelengths of a few hundred m or less, and so the model requires definition at that scale; (2) since shallow S -wave impedance has an especially strong role in determining ground-motion amplification (Boore *et al.*, 1993; Anderson *et al.*, 1996; Day, 1996) and because V_p/V_s can be highly variable in unconsolidated, possibly saturated shallow material, it is important to have direct V_s estimates in the near surface rather than relying on the empirical V_p/V_s relation used elsewhere in the model; and (3) some long-period modeling efforts (Graves, 1995) suggest that near-surface, small-scale structures significantly influence the amplitude and duration of recorded waveforms.

W. Silva (personal comm., 1999) compiled V_p and V_s measured for geotechnical studies in several hundred boreholes by various organizations. The boreholes are tens to hundreds of meters deep, and borehole coverage (Fig. 4) is

dense in the Los Angeles area basins. These data provide direct constraints for the shallow parts of the model basins.

To incorporate the geotechnical data into the velocity model we must interpolate point measurements (borehole velocities at a given depth) over the model area while also preserving the geologic boundaries separating different surface geology types. The latter is important because different surface geologies can have different ground-motion responses (e.g., Silva *et al.*, 1999). We use the detailed NEHRP category (Martin, 1994) classification map developed by the California Division of Mines and Geology (CDMG) (Wills *et al.*, 2000) (Fig. 4). The map plots, with uniform statewide coverage, surface regions classified by NEHRP site category based on the average shear-wave velocity in the top 30 m; the site category regions are outlined by geologic boundaries. The CDMG added the intermediate site categories BC, CD, and DE to the original NEHRP site categories B, C, and D that are present in southern California. We calculate average V_p and V_s velocity-depth profiles for each site category by finding, for a given depth, the log normal mean value of the velocities of every borehole in the site category. The mean profiles are determined to maximum depths (25 to 150 m) controlled by the number of samples at each depth; at least three samples are required. Interestingly, the mean V_s profiles of two site types (CD and D) vary by basin for depths below 30 m (Fig. 5), with V_s lower in the Los Angeles basin than in the San Gabriel and San Fernando Valleys. The difference may be because the Los Angeles basin is further from the sediment sources north and

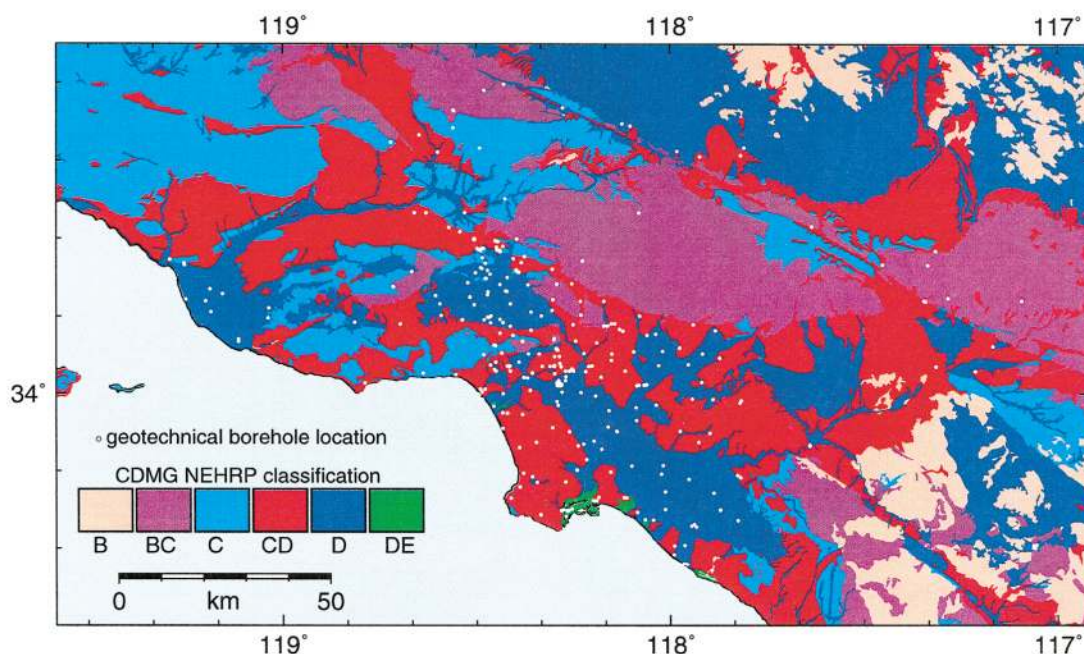


Figure 4. Geotechnical borehole locations (white circles; W. Silva, personal comm.) and the CDMG NEHRP site classifications (Wills *et al.*, 2000) for the Los Angeles region.

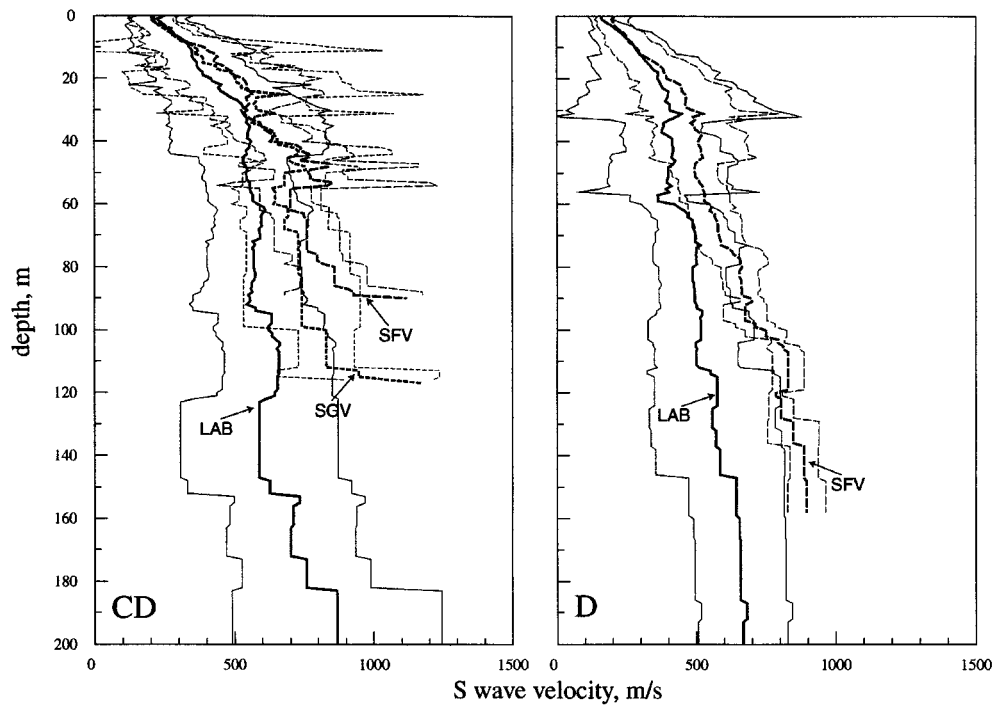


Figure 5. Mean (unsmoothed log-normal) V_S profiles (thick lines) of site categories CD and D for the Los Angeles basin (LAB, solid), San Fernando Valley (SFV, long dashes), and San Gabriel Valley (SGV, short dashes). Note differences below 30 m depth. Thin lines are $\pm 1 \sigma$; the number of samples varies from 3 to 88 at different depths.

east of the basins than the two valleys, and so receives finer grained, seismically slower sediments. We use basin-specific mean profiles (defined by finding the mean velocities of the boreholes of each site type within each basin) for site types CD and D.

Separate V_P and V_S mean profiles for all the site categories are used; the V_S profiles are smoothed by eye to remove minor velocity inversions that result from the averaging process. V_P/V_S values derived from the (unsmoothed) mean profiles are about 1.7 to 2.5 in site category C, about 1.7 to 3.5 in site categories CD and D (except for in the San Gabriel Valley, where category D V_P/V_S reaches about 5.5), and up to about 9.5 in category DE (Fig. 6). Site category BC had too few V_P data to calculate V_P/V_S .

The velocity at a specific shallow point is found by (1) looking up the site category the point is in; (2) looking up nearby (<5 km distance) boreholes in the same site category with data at the same depth as the point; and (3) assigning the velocity as a weighted combination of the appropriate mean profile and nearby boreholes. If there are no nearby boreholes the velocity from that site type mean profile is used. This allows reasonable velocity values to be assigned to the areas where geotechnical data are sparse. If the point is within 50 m of a borehole, the velocity from that borehole is used, so the original borehole data can be recovered. Between 50 m and 2 km (2 km and 5 km) the boreholes and generic profile are weighted by 2/3 and 1/3 (1/3 and 2/3),

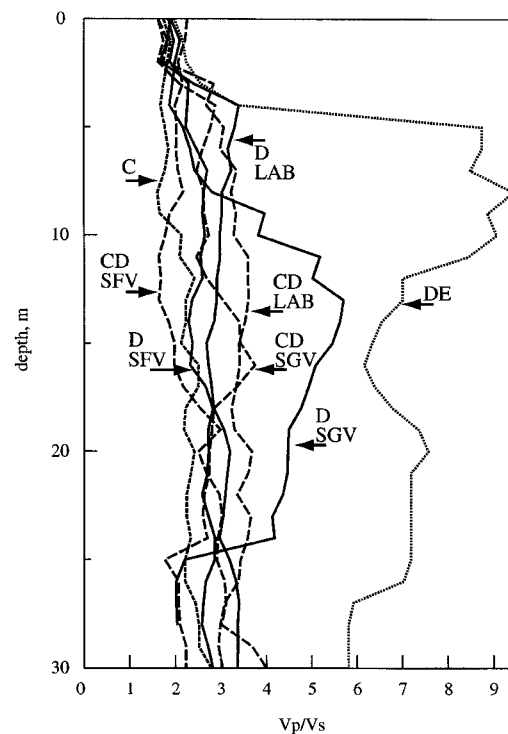


Figure 6. V_P/V_S for site categories C (short dashes), CD (long dashes), D (solid), and DE (dots). Abbreviations as in Figure 5.

respectively. The weighting scheme preserves impedance contrasts between site category regions while allowing smooth variations within a site category region. Below the maximum depth of the mean profiles, V_S is extrapolated using velocity-depth gradients from the deepest boreholes (Silva *et al.*, 1999; W. Silva, personal comm., 2000) and compared to V_S calculated by the rule-based scheme. The extrapolated V_S is used until it reaches a value equal to the rule-based value, typically between 100 and 300 m depth; below 300 m, the rule-based V_S is always used. V_P is determined by a weighted sum of the bottom of the generic profiles and the rule-based velocities, so that the two smoothly merge at 200 m depth. That depth was determined by comparing the borehole profiles with trial model predictions. The geotechnical constraints are implemented in the Los Angeles area basins (Los Angeles basin, Ventura basin, San Gabriel Valley, San Fernando Valley, Chino basin, and San Bernardino Valley) but not elsewhere in the model. These constraints are not present in any previous versions of the reference model.

Salton Trough

The Salton Trough has been subject to seismic refraction studies by Fuis *et al.* (1982), Fuis and Kohler (1984), Mooney and McMechan (1982), and Parsons and McCarthy (1996). They present V_P cross sections (Fig. 2) derived from the refraction results. Following Magistrale (1999), we digitize the cross sections and define five isovelocity reference surfaces of 2.5, 5.5, 6.0, 6.5, and 7.0 km/sec. We fix the surface V_P at 1.8 km/sec, consistent with all the refraction studies. Parsons and McCarthy (1996) determine a 22 km deep Moho, with an upper mantle V_P of about 7.7 km/sec; we use those values. Lower crustal V_P determinations vary from 6.9 km/sec (Parsons and McCarthy, 1996) to about 7.5 km/sec (Fuis *et al.*, 1982). We fix the V_P just above the Moho to 7.1 km/sec, producing a gentle gradient between the 7.0 km/sec surface and the 22 km deep Moho, and a sharp jump to the upper mantle 7.7 km/sec V_P .

The seismic lines constraining the model are in the Imperial Valley (the southern part of the Salton Trough). Areas outside of the Imperial Valley, but still within the Trough (e.g., the Coachella Valley) are assigned V_P from a velocity-depth profile determined within the Imperial Valley. The lack of control on the Coachella Valley sediment velocities and basement configuration mean that this part of the model is relatively crude. The Salton Trough is not present in any previous versions of the model.

Crustal Model

The seismic velocity model requires realistic velocities to represent the rocks outside of the rule-based basin models. Tomographic results (Hauksson, 2000) from the inversion of local earthquake travel times recorded on the southern California seismic network provide 3D basement rock V_P and V_S values. Those velocities are determined on a 15 km \times 15 km grid covering southern California at 9 depths (1.0,

4.0, 6.0, 10.0, 15.0, 17.0, 22.0, 31.0, and 33.0 km). V_P and V_S at a point outside of the basins is determined by interpolation of the velocities of the eight tomography grid nodes surrounding the point.

The tomographic calculations used an initial 1D model based on Hadley and Kanamori (1977) that did not include the model Los Angeles area basins. The compatibility of travel times upon the insertion of the basins into the tomographic background is tested by replacing the final tomography values at grid nodes within the basins with velocities from the basin models, and performing an additional iteration of the travel-time inversion. Around the basin edges, the results show the influence of the basin low velocities smeared over 4 to 5 grid lengths into the surrounding crust as concentric bands of slightly (<4%) higher and lower velocities relative to the velocities found in the original tomography results. Away from the basins, the velocity differences are <1%. Because the smearing is an artifact of the tomography method, and the changes elsewhere are small, we conclude that the basins model is compatible with regional travel-time data.

The incorporation of tomographically determined velocities outside the basins is new to Version 2. The Version 0 and 1 models use a 1D, depth-dependent crustal model from Hadley and Kanamori (1977).

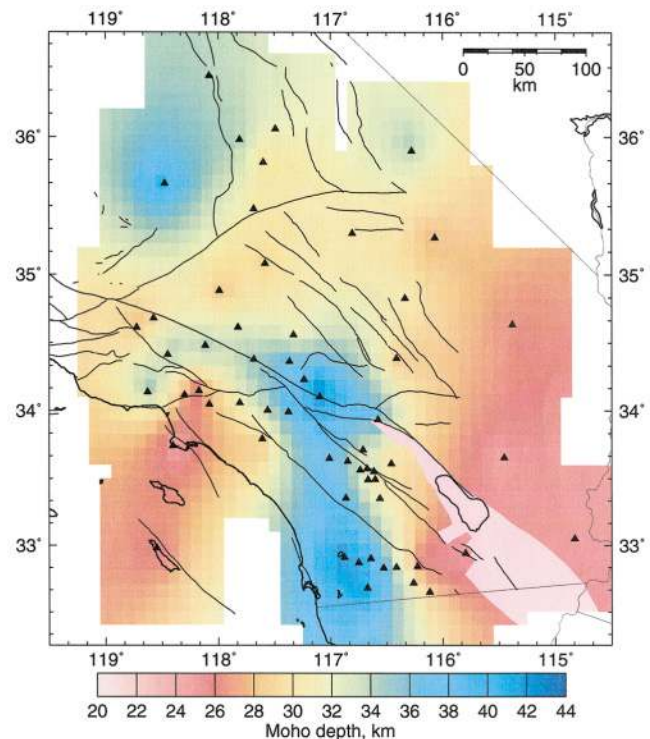


Figure 7. Model Moho surface. Moho depths are found using the receiver-function technique and regional tomography crustal velocities except in the Salton Trough where the Moho depth is fixed at 22 km. Triangles indicate broadband station locations.

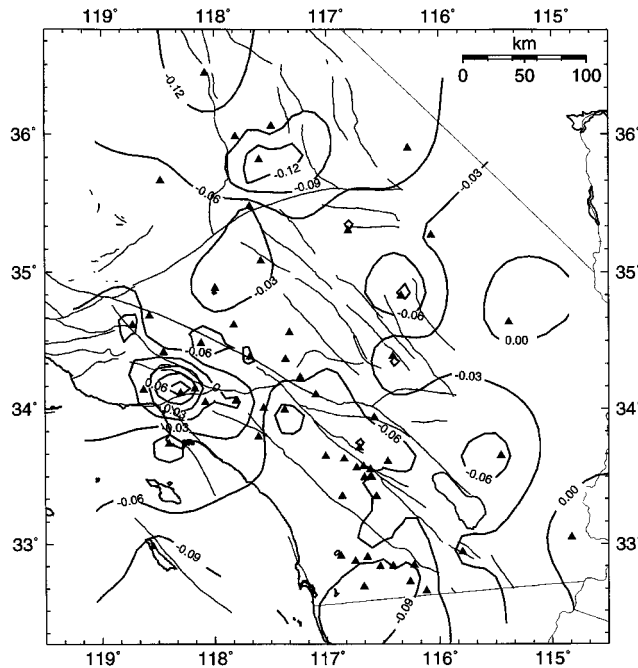


Figure 8. Contours of V_P/V_S differences between the tomographic crustal velocities of Hauksson (2000) and multiple converted phase-stacking technique of Zhu and Kanamori (2000). Note that the tomographic V_P/V_S are generally lower (negative differences) than the stacking technique values. Triangles indicate broadband station locations. Contour interval is 0.03.

Moho

A well characterized, variable depth Moho is a desirable element of the model not only to accurately model the regional distribution of seismic velocities, but also to constrain models of the tectonic evolution of the region. We determine the crustal thickness of southern California using P_s minus P times measured from receiver functions of teleseismic events recorded at broadband stations, and crustal velocities determined by tomographic analysis of local earthquake travel times.

We use the receiver functions determined at ~ 70 sites by Zhu and Kanamori (2000) and the 3D crustal V_P and V_S regional tomography models of Hauksson (2000). For each recording site, we construct V_S and V_P crustal velocity profiles from the 3D tomography results. These crustal velocities, described previously, are used to produce a self-consistent model. Moho depths are obtained from the P_s minus P times by adjusting the thickness of the lower crust to match the differential travel times.

The receiver function point determinations of Moho depth are converted (via minimum curvature gridding) into a surface and resampled onto a $6'$ by $6'$ grid (Fig. 7). Under the Salton Trough, the Moho depth is fixed at 22 km depth. A point of interest is tested to determine if it is above or below the Moho: if below, a V_P of 7.8 km/sec is assigned

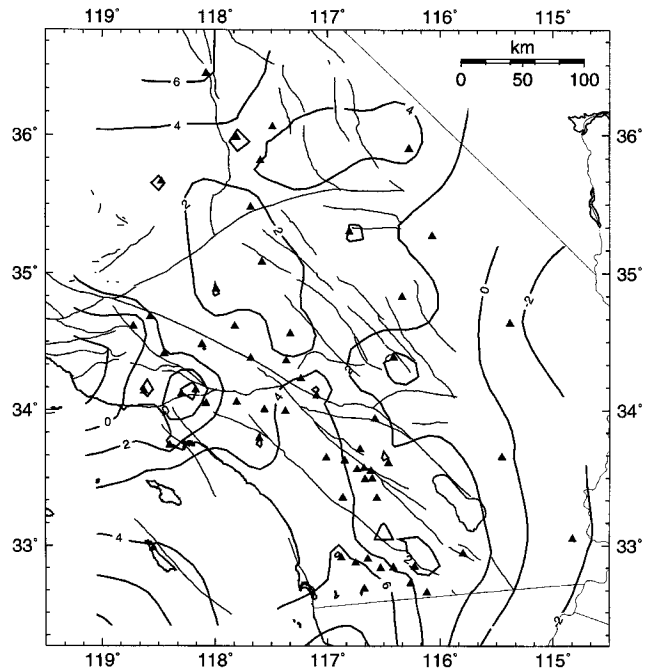


Figure 9. Contours of the differences between the receiver function Moho depths of the current work and those of Zhu and Kanamori (2000). Note that the current work depths are generally deeper (positive differences). Triangles indicate broadband station locations. Contour interval is 2 km.

(based on Hadley and Kanamori, 1977), except in the Salton Trough, where an upper mantle V_P of 7.7 km/sec is used (Parsons and McCarthy, 1996). Future model versions will incorporate 3D upper mantle velocities.

Crustal thickness found from receiver functions is a strong function of V_P/V_S . We estimate the uncertainties in the Moho depth determination by comparing different V_P/V_S models and the Moho depths calculated from those models. Zhu and Kanamori (2000) calculate crustal thickness and a vertically integrated V_P/V_S with a receiver function stacking technique that exploits the converted phase multiples to constrain the tradeoff between crustal thickness and V_P/V_S . Hauksson (2000) and Zhou (1994) performed tomographic inversions of local earthquake travel times to determine V_P and V_S structure. The vertically integrated crustal V_P/V_S of the two tomographic studies are similar, and both tend to be lower than the V_P/V_S found by stacking (Fig. 8). The Moho depths found here using the tomographic V_P and V_P/V_S are typically 2 to 4 km deeper than the Moho depths found by Zhu and Kanamori (2000) (Fig. 9); by comparison, the Moho depth errors they estimate at all sites average to ± 0.9 km.

Model Application and Availability

The southern California reference seismic velocity model (Figs. 10 and 11) is suitable for a variety of applications. As a reasonable description of crustal properties, the

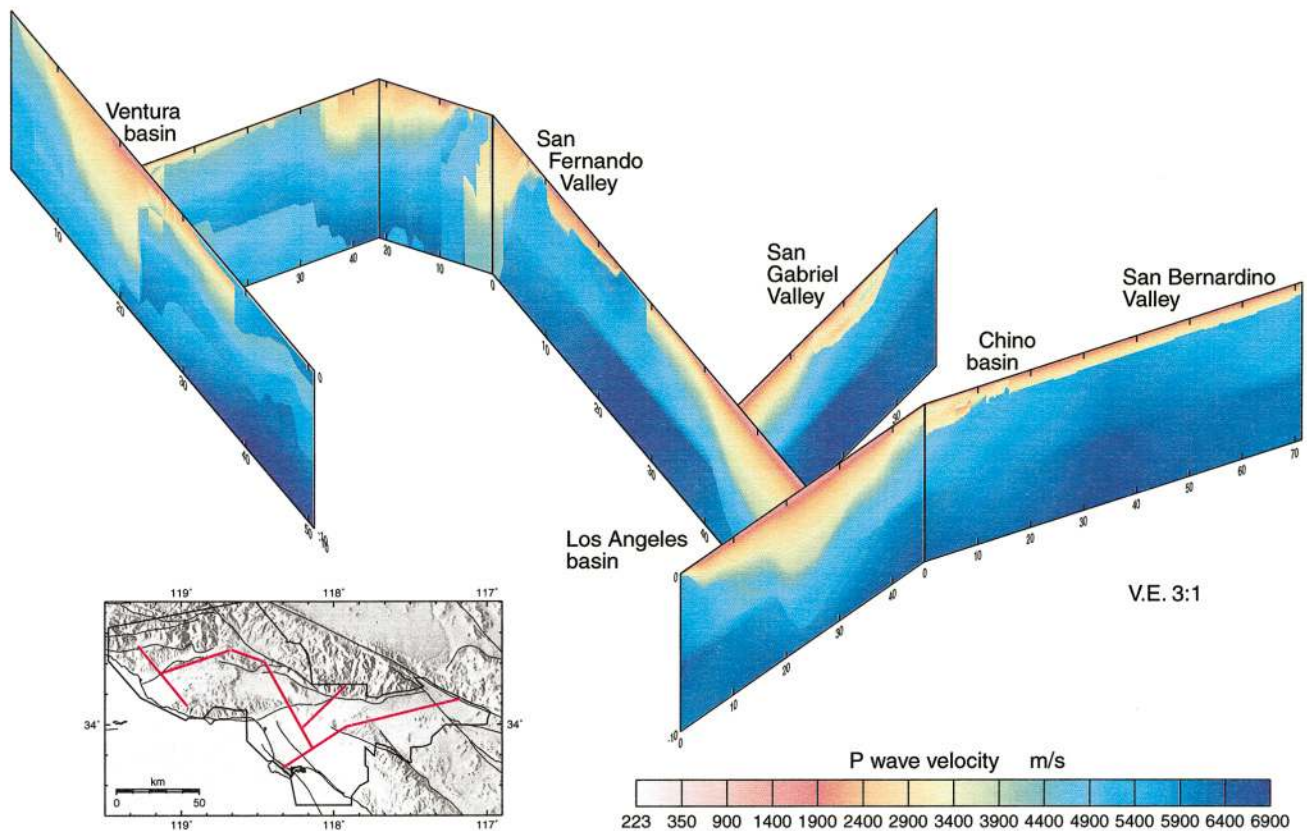


Figure 10. Fence diagram of P -wave velocity in and around the Los Angeles region basins. Cross section locations shown as red lines in lower left panel. Seismic velocities in the basins are controlled by rule-based models within the basins, and by tomographic results outside the basins. See Field *et al.* (2000) for an S -wave velocity fence diagram.

Version 0 and 1 models have been used to model basin effects on ground motions from past (e.g., Day *et al.*, 1994) and potential future earthquakes (e.g., Olsen *et al.*, 1996), to incorporate basin structure into earthquake ground-motion attenuation relations (Lee and Anderson, 2000; Steidl, 2000; Field, 2000), and to perform 3D source inversion for southern California earthquakes (Liu and Archuleta, 1999).

The reference model may be appropriate for use as a starting model in perturbation studies, such as linearized inversions of travel times for crustal velocities (e.g., Magistrale, 1999) or of seismic waveforms for crustal structure. SCEC has supported investigations to verify and improve the model, for example, testing the ability of the model to produce synthetic waveforms that match observations of recent, well-recorded earthquakes, such as Landers and Northridge (Graves *et al.*, 1999; Olsen, 2000), and testing how well the model densities can fit gravity observations (Roy and Clayton, 1999).

Version 2 of the standard three-dimensional seismic velocity model for southern California is available as a FORTRAN source code and associated files on the SCEC Data Center website at <http://www.scecdc.scec.org>.

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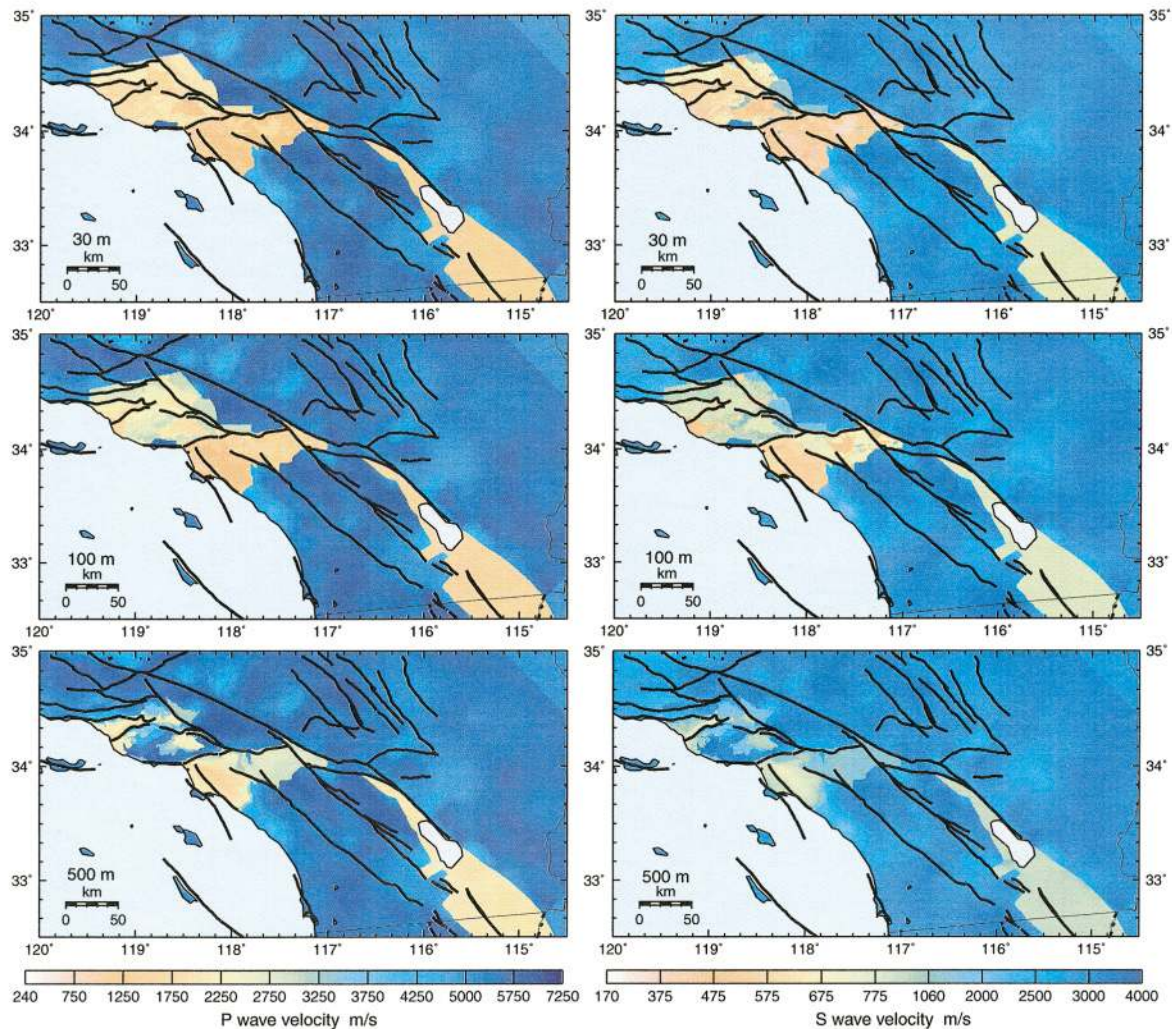


Figure 11. Depth slices through the model. Left column is P -wave velocity, right column is S -wave velocity (scales differ). The velocities are controlled by the rule-based basin models within the basins, and by tomographic results outside the basins. Depth of each slice is indicated on each panel.

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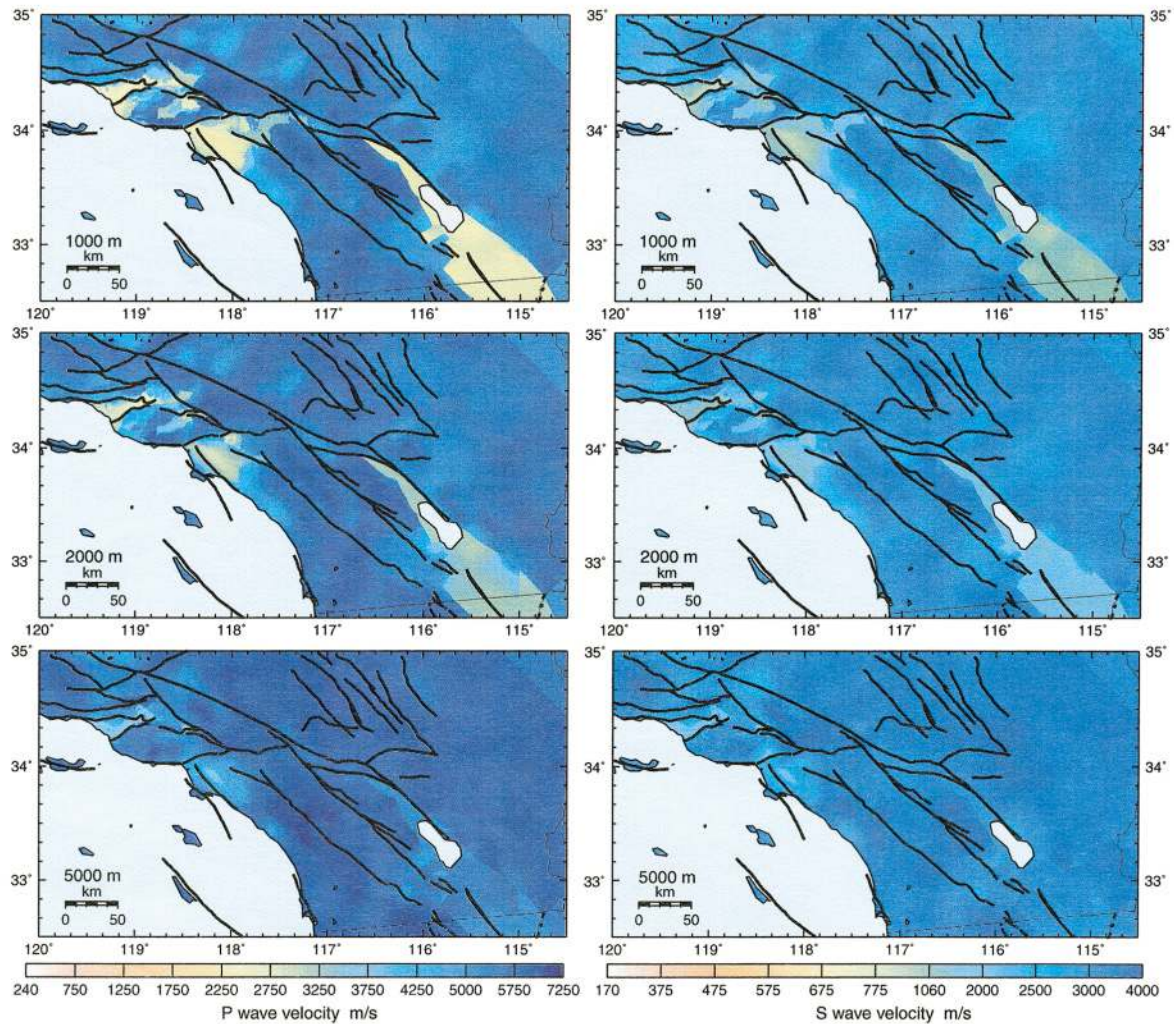


Figure 11. (Continued).

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Department of Geological Sciences
San Diego State University
San Diego, CA 92182-1020
(H. M., S. D.)

Seismological Laboratory 252-21
California Institute of Technology
Pasadena, CA 91125
(R. W. C.)

URS Greiner Woodward Clyde
566 El Dorado Street
Pasadena, CA 91101
(R. G.)

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