# THERMAL EXPANSION OF SINGLE-CRYSTAL MANGANOSITE

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Thermal expansion of manganosite MnO at temperatures between  $20^{\circ}\text{C}$  and  $850^{\circ}\text{C}$  has been carefully determined by a dilatometric technique. The theory of thermal expansion by Grüneisen is improved and the acquired data are analyzed to derive the harmonic and anharmonic parameters of MnO. Grüneisen's parameter  $\gamma$  is 1.56, Debye temperature is 441 K, and pressure derivative of bulk modulus is 7. The volume expansion coefficient is  $34.3 \times 10^{-6}/\text{K}$  at  $20^{\circ}\text{C}$  and  $47.1 \times 10^{-6}/\text{K}$  at  $850^{\circ}\text{C}$ ; almost the same in magnitude to those of MgO and FeO.

#### 1. Introduction

We have been studying the thermal properties of the materials composing the earth's mantle and their related compounds. Subsequent to the latest measurement on tephroite Mn<sub>2</sub>SiO<sub>4</sub> (OKAJIMA et al., 1978), the thermal expansion of manganosite MnO, the manganese analogue to periclase MgO and wüstite FeO, was determined up to 850°C, to obtain a better understanding of the anharmonic properties of these materials.

The useful information on the anharmonic property of a solid is derived by an appropriate analysis of accurate thermal expansion data (WACHTMAN et al., 1962; Suzuki, 1975). The usual practice is the application of Grüneisen's theory of thermal expansion with some approximation. In this paper, an improved expression of thermal strain, in which the second order term is exactly included, is proposed and comparison of the analyzed result with that by the previous method is made on the manganosite data.

## 2. Sample and Experimental Procedure

A single-crystal manganosite grown by a flame fusion technique was afforded by Nakazumi Crystal Co. This specimen is glossy and black, but neither crack nor flaw is found penetrating into the interior of the specimen.

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Table 1. Observed and calculated values of thermal expansion and its coefficient for manganosite MnO.

Temp. $\frac{t}{{}^{\circ}C}$	Observed expansion $\frac{Y_{\rm L}}{10^{-2}}$	Calculated expansion and its coefficient				
		Linear		Volume		
		$\frac{Y_{\rm L}}{10^{-2}}$	$\frac{\alpha_{ m L}}{10^{-6}{ m K}^{-1}}$	$\frac{Y_{\rm V}}{10^{-2}}$	$\frac{\alpha_{\rm V}}{10^{-6}{ m K}^{-1}}$	
-200 -150 -100 -50 0		-0.1950 -0.1685 -0.1272 -0.0777 -0.0234	3.3 7.1 9.3 10.5 11.2	-0.5837 -0.5046 -0.3812 -0.2330 -0.0702	9.8 21.2 27.7 31.4 33.7	
20 25 50 75	0.0 0.0045 0.0328 0.0630	0.0008 0.0050 0.0340	11.4 11.5 11.7	-0.0023 0.0149 0.1021	34.3 34.5 35.2	
100 125	0.0942 0.1249	0.0937	12.1	0.2814	36.3	
150 175	0.1564 0.1870	0.1551	12.4	0.4661	37.3	
200 225	00 0.2187	0.2180	12.7	0.6556	38.1	
250 275	0.2818 0.3138	0.2823	12.9	0.8492	38.8	
300	0.3474 0.3807	0.3477	13.2	1.047	39.5	
325 350	0.4143 0.4467	0.4142	13.4	1.248	40.1	
375 400 425	0.4821 0.5160	0.4819	13.6	1.453	40.7	
450 475	0.5503 0.5863	0.5507	13.8	1.661	41.4	
500	0.6199	0.6205	14.0	1.873	42.0	
525 550 575	0.6567 0.6918 0.7275	0.6916	14.2	2.089	42.7	
600 625	0.7634 0.8006	0.7637	14.4	2.309	43.3	
650 675	0.8376 0.8743	0.8371	14.7	2.532	44.0	
700 725	0.9112 0.9495	0.9117	14.9	2.760	44.7	
750 775	0.9884 1.027	0.9876	15.2	2.992	45.5	
800 825	1.064 1.104	1.065	15.4	3.229	46.3	
850	1.144	1.144	15.7	3.470	47.1	
900 950 1,000 1,050 1,100 1,150		1.224 1.306 1.389 1.474 1.561 1.651	16.0 16.3 16.6 17.0 17.4 17.8	3.716 3.968 4.225 4.488 4.758 5.034	48.0 48.9 49.9 51.0 52.1 53.3	
1,200 1,250 1,300 1,350 1,400		1.742 1.836 1.932 2.032 2.134	18.2 18.7 19.2 19.8 20.4	5.318 5.609 5.910 6.219 6.540	54.6 56.0 57.6 59.3 61.2	

This crystal was shaped into a cylinder along  $\{1\ 0\ 0\}$  axis about 7 mm in height, and both ends of the cylinder were polished to be parallel. Bulk density determined from the size and mass of the specimen is  $5.373\pm0.011$  g/cm³ at  $20^{\circ}$ C.

The experimental technique and apparatus to measure thermal expansion were reported in previous papers (Suzuki, 1975; Okajima et al., 1978). By using a 10:10:1 mixture of  $CO_2$ , Ar, and  $H_2$  gasses, the oxygen partial pressure of the specimen environment was brought to the order of  $10^{-16}$  bar at 850°C. The specimen after the measurement did not show any evidence of alteration. The acquired data of the thermal expansion are shown in the second column of Table 1.

## 3. Theoretical Representation of Thermal Expansion

In the previous work, the following theoretical expression of thermal expansion was used for the analysis of data (GRÜNEISEN, 1926; SUZUKI, 1975);

$$y_{v}(T) = (V(T) - V(0))/V(0)$$

$$= E(\theta, T)/(Q_{0} - kE(\theta, T))$$
(1)

where  $y_v(T)$  is relative volume expansion at temperature T referred to 0 K, V is the volume,  $E(\theta, T)$  is the thermal energy as a function of T, and  $\theta$  is the Debye temperature  $\theta$ .  $Q_0(\gg kE)$  is such a quantity that determines the magnitude of thermal expansion, and is defined by

$$Q_0 = B_0 \cdot V_0 / \gamma$$

where  $\gamma$  is Grüneisen's parameter,  $B_0$  and  $V_0$  are bulk modulus and volume at 0 K. k is also an anharmonic parameter related to the pressure derivative of bulk modulus  $dB_0/dP$  at 0 K;

$$k = (dB_0/dP - 1)/2$$
.

The thermal expansion in Eq. (1) referred to 0 K, is converted to that referred to a reference temperature  $T_{\rm r}$ ,

$$Y_{v}(T) = y_{v}(T)/a_{v} + (1 - a_{v})/a_{v}$$

$$= E(\theta, T)/a_{v}(Q_{0} - kE(\theta, T)) + (1 - a_{v})/a_{v}$$
(2)

where  $a_v$  is the ratio of volumes at  $T_r$  and 0 K, i.e.,  $a_v = V(T_r)/V(0)$ . The volume expansion is related to linear expansion defined by  $Y_L(T) = (L(T) - L(T_r))/L(T_r)$ , where L(T) is a linear dimension of the crystal; thus,

$$Y_{v}(T) = 3Y_{L}(T) + 3Y_{L}^{2}(T) + Y_{L}^{3}(T)$$
. (3)

Then Eq. (2) can also be represented in terms of linear expansion. The approximate expression which have been employed (WACHTMAN et al., 1962) is

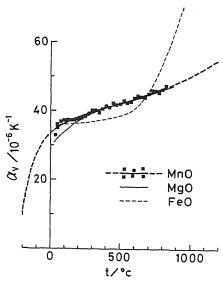


Fig. 1. Observed and calculated values of volume expansion coefficient  $\alpha_V$  of manganosite MnO.  $\alpha_V$  is extrapolated over wider range of temperature by Eq. (6).  $\alpha_V$  of periclase MgO (SUZUKI, 1975) and of polycrystalline wüstite Fe<sub>1-z</sub>O (Carter, 1959) are compared.

based on the first term alone of Eq. (3); so that,

$$Y_{L}(T) = E(\theta, T)/3a_{L}(Q_{0} - kE(\theta, T)) + (1 - a_{L})/a_{L}$$
 (4)

where  $a_{\rm L}$  is the ratio of linear dimensions at  $T_{\rm r}$  and 0 K, i.e.,  $a_{\rm L} = L(T_{\rm r})/L(0) = a_{\rm v}^{1/3}$ . In fitting Eq. (2) or Eq. (4) to the experimental data,  $a_{\rm v}$  or  $a_{\rm L}$  is an additional parameter to be determined.

The procedure in the derivation of Eq. (1) is as follows; 1) the Taylor expansion of the potential term in the Mie-Grüneisen equation of state is terminated at the first order, and 2) the additional second order term is approximated from the first order term by way of iteration (see Suzuki, 1975, for detail). Because of the approximate treatment of the second order term of  $y_v(T)$  in Eq. (1), the constants  $Q_0$ ,  $\theta$ , and k to be determined may be systematically biased. Further, the conversion of Eq. (2) to Eq. (4) involves the approximation  $Y_v(T) = 3Y_L(T)$ , which would also lead to the numerical difference of  $Q_0$ ,  $\theta$  and k depending on the way of analysis, namely, fitting by linear expansion or by volume expansion.

Yet, the approximations in deriving Eqs. (1) and (4) have scarecely given rise to any problem, since the accuracy of thermal expansion measurement has not been sufficiently high. If, however, we are to derive more information from more accurate data, the exact treatment of Grüneisen's theory and

an accurate conversion of volume and linear expansions should be made. When we take account of the contribution of the second order term of Grüneisen's theory exactly, we have to solve the quadratic equation

$$ky_{v}-y_{v}^{2}+E(\theta,T)/Q_{0}=0$$

whose solution is

$$y_{v}(T) = (1 - \sqrt{1 - 4kE(\theta, T)/Q_0})/2k$$
 (5)

This equation has the same parameters as Eq. (1). Volume expansion referred to a reference temperature  $T_{\rm r}$  is thus represented by

$$Y_{v}(T) = (1 + 2k - \sqrt{1 - 4kE(\theta, T)/Q_0})/2ka_{v} - 1$$
 (6)

By using the exact relation;  $1+Y_v(T)=(1+Y_L(T))^3$ , we have

$$Y_{\rm L}(T) = (1 + 1/2k - \sqrt{1 - 4kE(\theta, T)/Q_0/2k})^{1/3}/a_{\rm L} - 1.$$
 (7)

Equation (7) is certainly more complicated but more rigorous than Eq. (4) given by Wachtman *et al.* (1962). The complicated expression of the equation is not, however, a matter of any difficulty with the practical data analysis with a computer.

## 4. Results of Data Analysis

Calculations based on the least squares method using Eqs. (2), (4), (6), and (7) were made with the aid of an iteration technique to determine the numerical parameters;  $Q_0$ ,  $\theta$ , k, and a's from the present data of manganosite. These parameters having been determined, their probable errors, and standard deviations are listed in Table 2. The standard deviations are equally small of the order of  $10^{-5}$  (the last column in Table 2), indicating that the previous Eqs. (2) and (4) are as good as the proposed new Eqs. (6) and (7) in fitting the experimental data. However, the value of k determined by volume expansion (Eq. (2)) and by linear expansion (Eq. (4)) are significantly different, while the same value of k is derived in the proposed Eqs. (6) and (7) as

Table 2. Comparison of expansion parameters in Eqs. (2), (4), (6), and (7) for manganosite MnO.

	$\frac{Q_0}{10^6 \text{ J/mol}}$	$\frac{\theta}{K}$	k	$\frac{a^{1)}-1}{10^{-2}}$	$\frac{\sigma^{2)}}{10^{-4}}$
Eq. (2)	1.403±0.008	359±18	4.20±0.08	$0.661 \pm 0.014$	0.19
Eq. (4)	$1.403 \pm 0.008$	$359 \pm 18$	$3.88 \pm 0.08$	$0.220 \pm 0.005$	0.06
Eq. (6)	$1.343 \pm 0.005$	$441 \pm 13$	$3.03 \pm 0.04$	$0.608 \pm 0.010$	0.19
Eq. (7)	$1.343 \pm 0.005$	$441 \pm 13$	$3.03 \pm 0.04$	$0.202 \pm 0.003$	0.06

Reference temperature 20°C,  $a=a_{\rm V}$  for Eqs. (2) and (6),  $a=a_{\rm L}$  for Eqs. (4) and (7).

 $<sup>\</sup>sigma = \sum (Y_{\text{obs}} - \hat{Y}_{\text{calc}})^2/(n-q)^{1/2}$ , n, number of input data; q, number of parameters to be determined.

a matter of course. Further, the probable errors of  $Q_0$ ,  $\theta$ , k, and a obtained by Eqs. (6) and (7) are certainly less than those by Eqs. (2) and (4). Therefore, it is concluded that Eq. (6) or Eq. (7) is definitely better than Eqs. (2) and (4). The systematic deviation of  $Q_0$ ,  $\theta$ , k, and, a by the application of Eqs. (2) and (4) has been clarified as shown in Table 2. It is stressed that the second order term of strain should be taken into account in order to achieve a reliable analysis of data.

In Table 1, the observed linear expansion  $Y_L$  of manganosite is compared with that calculated by Eq. (7) with the parameters listed in Table 2. The volume expansion  $Y_v$  and expansion coefficient  $\alpha$  are calculated over a wider range of temperature by Eq. (6) and are shown in Table 1.

## 5. Discussion

The volume expansion coefficient of MnO is compared with those of periclase (Suzuki, 1975) and wüstite (Carter, 1959) in Fig. 1. These oxides have almost the same magnitude of expansion coefficients in the temperature range of measurement. A noticeable feature in wüstite is the steep increase of the expansion coefficient above 600°C. There are several possible causes: 1) wüstite is not stable below 560°C (Carter, 1959), 2) the change of stoichiometry (vacancy concentration) with temperature and partial pressure of oxygen (Darken and Gurry, 1946), 3) the changes of structure and of the associated defects in Carter's polycrystalline specimen, and/or 4) the effect of magnetic transition (Néel temperature  $T_N = -84$ °C; Bozorth et al., 1972). The second and third possibilities are most likely, but the discussion should be reserved until the specified data on the single-crystalline sample is worked out. The last effect may also be present in MnO ( $T_N = -153$ °C; Bozorth et al., 1972), but the transition temperature is too low to affect significantly the properties of MnO at high temperature.

The similarity of thermal expansion coefficients in MgO, FeO and MnO suggests that rock-salt oxides relevant to the earth's mantle show essentially the same thermal expansion coefficient; 35 at room temperature and 50 at  $1,000^{\circ}$ C in the unit of  $10^{-6}/K$ . Above the Debye temperatures, these oxides show an expansion coefficient about a quarter smaller than those of the typical ionic crystals such as alkali-halides with rock-salt structure (Kirby et al., 1972). This fact is understood by comparing the valence product  $(2 \times 2 = 4)$  of the divalent compounds and that  $(1 \times 1 = 1)$  of monovalent alkali-halides.

The harmonic and anharmonic parameters from the present expansion measurement are compared with those derived from other sources. Debye temperature calculated by the elastic wave velocities of MnO (Sumino, personal communication, 1978) is  $\theta = 537$  K. The characteristic temperature

441 K derived from Eq. (6) or Eq. (7) is smaller by 100 K than the acoustic Debye temperature. The source of this discrepancy is not yet understood. It is noted that the characteristic temperature 359 K derived by an application of the previous Eq. (2) or Eq. (4) is too small. This may be another reason for supporting the use of Eqs. (6) and (7) instead of Eqs. (2) and (4). The value of  $k = (dB_0/dP - 1)/2$  gives the pressure derivative of bulk modulus  $dB_0/dP =$ 7.0, which is not an unreasonable magnitude when compared with  $\partial B/\partial P$  of ordinary solid materials. The lattice parameter of manganosite at 25°C is given as 4.4448 Å (Robie et al., 1966). By using the values of a from Eq. (6) or Eq. (7) in Table 2, the lattice constant and molar volume of manganosite at 0 K are given as  $4.4358 \pm 0.0005$  Å and  $13.140 \pm 0.005$  cm<sup>3</sup>/mol. By combining the values of  $Q_0 = 1.34 \times 10^6$  J/mol and  $B_0 \cdot V_0 = 2.095 \times 10^6$  J/mol (approximated from the room temperature data), we get Grüneisen's parameter This  $\gamma$  is theoretically different from the thermal  $\gamma(\cong B \cdot V/Q_0) \cong 1.56.$ Grüneisen's parameter  $\gamma_{th}$ , but the numerical difference may be very small, a few percent at most (Suzuki and Kumazawa, 1978).

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