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Water-Supply Paper 968-C

TOPOGRAPHIC CHARACTERISTICS OF DRAINAGE BASINS

By WALTER B. LANGBEIN AND OTHERS

Contributions to the hydrology of the United States, 1944 (Pages 125-158)



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TOPOGRAPHIC CHARACTERISTICS OF DRAINAGE BASINS

By WALTER B. LANGBEIN and others

ABSTRACT

River floods are the result of many causes, and one of the primary objectives of scientific hydrology is the segregation and evaluation of the causative factors. The climatic factor and the soil-vegetation complex are variables that exercise their principal influence on the volume of runoff. The topography of drainage basins is a sensibly permanent characteristic which influences mainly the concentration or time distribution of the discharge from a drainage basin. River systems differ in their efficiency as agencies for collecting and conducting water. In some systems, surface waters are quickly assembled, and the discharge reflects somewhat sensitively the variations of the available supply. In others, the surface drainage is longer delayed and the discharge is released slowly.

As a basis for quantitative studies of these evident differences in behavior, selected topographic features for about 340 drainage basins in the northeastern United States were studied, using Geological Survey topographic maps. The data were compiled in cooperation with the Work Projects Administration of the Federal Works Agency and included information on drainage area, length of streams, stream density, land slope, channel slope, area-altitude distribution, and area of water bodies of basins that ranged in extent from 1.64 to 7,797 square miles. Considerable effort was made to assure accuracy of the computations by appropriate checks, and the results are summarized in the table at the end of this report.

The results indicate that none of the topographic factors are unique, but each reflects a condition that also influences the others. For example, steep land slopes are generally associated with steep channel slopes and conversely. A significant variation of slope and altitude with area of basin is found, and stream density tends to vary with the land slope.

INTRODUCTION

This report presents a compilation of topographic data on drainage basins in the northeastern United States. The configuration of the earth reflects the impact of many natural forces, and it in turn exercises profound influence upon man. Most of these influences are so

basic that they have shaped life and civilization into conformity with Mountains, plains, valleys, and rivers each favor or retard them. man's search for economic stability. Within human history the first three have remained unchanged. Rivers, on the other hand, fluctuate in size from day to day and from year to year. The amplitude and frequency of these fluctuations, so significant with respect to navigation, water power, irrigation, and such riparian developments as cities and highways, are largely determined by three separate, yet interdependent features, namely climate, physiography, and the soil-The interrelation of these three features with vegetation complex. the behavior of rivers is imperfectly understood and is the subject of much investigation. This report singles out the physiography of the land for attention.

The relations between the rate, volume, and fluctuations of rivers and the topographic characteristics of the land they drain and through which they flow may be readily determined after discerning examination of the terrain and river developments, but expressing them in the quantitative terms necessary for the economic design of structures for river utilization or control requires first, topographic maps, and second, records of river flow of length adequate to define the behavior.

The stream-gaging program of the Geological Survey is Nationwide and now includes over 4,500 river-measurement stations, at which more than 65,000 station years of record were available in 1942. These records furnish an adequate source of material concerning stream behavior. The mapping program of the Geological Survey, also Nation-wide, is in general not so complete. Although about 50 percent of the country has been mapped, only States in the northeastern part have been completely covered; the scattered areas mapped in other States generally do not cover completely the areas in which stream-gaging has been carried on, so that only a small fraction of them are suitable for use in comparisons of stream-flow characteristics or river morphology.

In the northeastern and north-central States the range in topography is sufficient to furnish a basis for studying its effect on stream flow. The topographic characteristics compiled from the maps and summarized in this report can only be evaluated by a consideration of the hydrology of stream flow, the assembling of waters in a drainage system, and the hydraulic elements that regulate velocity of flow. Many stream-flow characteristics are related either directly or indirectly to topographic features. It would seem, however, that the factors most sensitive to topographic difference would be those relating to floods. In this study, therefore, particular although not exhaustive attention is given to the correlation of flood-flow characteristics with topography. This information will serve as a basis for further study of such correlations and also of other characteristics, such as volume yield, erosion, and deposition of sediments. Similarly the topographic data offers basic material for studies of river morphology, as geologic evidence suggests that a significant portion of river-channel development takes place during flood.

COOPERATION AND PERSONNEL

The cooperative project for the compilation of topographic data was undertaken in 1939 by the Works Progress Administration, which on April 25, 1939, became the Work Projects Administration under the Federal Works Agency. Their cooperation in organizing competent working groups is especially acknowledged. The Geological Survey sponsored the project and furnished technical direction, maps, and supplies. This work was carried on by W. B. Langbein, under the general direction of R. W. Davenport, chief of the Division of Water Utilization. The project at Boston, Mass., was under the supervision of H. B. Kinnison, district engineer, and his associates, particularly C. E. Knox, M. A. Benson, and B. R. Colby. The conduct of the work at Pittsburgh, Pa., was ably managed by Wm. S. Crozier, supervisor for the Works Projects Administration. Mr. Crozier died January 21, 1941, near the close of the project. H. M. Erskine, associate engineer of the Geological Survey at Pittsburgh, maintained close contact with the project there, and its continuity and efficiency may be largely credited to his competent administration.

METEOROLOGIC FACTORS AFFECTING RUNOFF

River floods are the result of many causes, and one of the primary objectives of hydrologic study is the segregation of the causative factors and the evaluation of their effects on the resultant floods under various associated conditions.

Readily apparent is the source of the water, generally an unusual amount of rainfall, which may be characterized by great intensity and in many regions may be augmented by water from melting snow. Water in excess of that which can be absorbed by the ground or evaporated into the air directly or through vegetation collects in the stream channels that drain the area. Once in the stream system, the runoff flows to the mouth through channels which, as the trunk of a network of streams, progressively increase in size as contributions are received from tributary streams.

The quantity of rainfall or snow melt, its time distribution, and the associated soil, vegetal, and climatic conditions that determine the portion of the supply that becomes direct runoff are to a large extent variable characteristics of individual storms. These variable edaphic and climatic factors are separate phases of the rainfall-runoff relation. The channel system, however, is a relatively permanent characteristic of a drainage basin. Some influential changes may take place in this system; for example, variations in seasonal vegetation along the banks may affect the hydraulic conveyance, floods may scour or deposit sediments, and old bends may be cut through and new ones created. Although the effect of these changes on local flood stages may be considerable, it is assumed that the resultant effect on total discharge from a basin will tend to be compensating and that even the cumulative effect on flood-discharge characteristics during a period as short as the usual stream-gaging record would be minor.

DIFFERENCES IN CHARACTER OF DRAINAGE BASINS

River systems differ in their efficiency as agencies for collecting and conducting water. In some systems the surface waters are quickly assembled, and the discharge therefrom reflects somewhat sensitively the variations of the available supply. In others the surface drainage is longer delayed and the discharge is released slowly. This difference is illustrated in figure 48, which shows the hydrographs of two nearby streams, each draining about 50 square miles of coastal areas of New Jersey, during a flood in June 1938. The rainfall causing these floods and the volume discharged were nearly the same for both areas. The difference in behavior illustrated by the hydrographs is normal for these two basins and may be accounted for largely by the differences in physiographic characteristics, Manasquan River having about twice the gradient of Great Egg River and about one-fifth the swamp area.

PREVIOUS STUDIES

The fact of relationship between the time distribution of discharge during a flood and the size, shape, and gradient of a drainage system is widely recognized. Few attempts have heretofore been made to determine this relationship quantitatively, probably because of the volume of labor required to evaluate the topographic factors.

Horton¹ in 1926 and again in 1932 discussed the desirability and need for a quantitative rational procedure and proposed methods for evaluating certain pertinent physiographic factors.

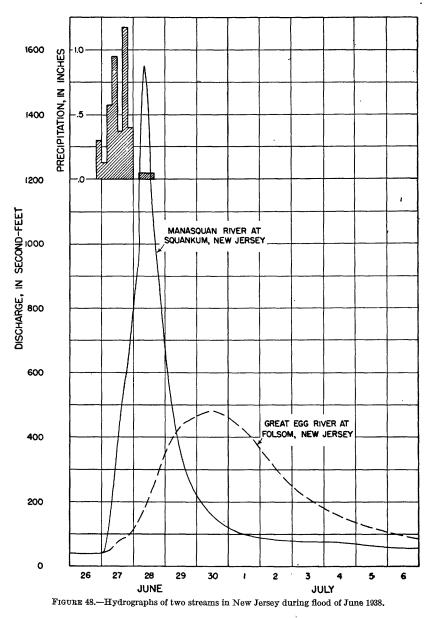
Pettis² in 1927 presented a formula to compute the maximum flood discharge, in which the five-fourths power of the average basin width was used.

Gregory and Arnold³ in 1932 developed in detail certain expressions

¹ Horton, R. E., in Jarvis, C. S., Flood flow characteristics: Am. Soc. Civil Eng. Trans., vol. 89, pp. 1681-1086, 1926; Drainage-basin characteristics: Am. Geophys. Union Trans., No. 13, pp. 350-361, 1932.

² Pettis, C. R., A new theory of river flood flow (published privately, copyrighted 1927).

³ Gregory, R. L., and Arnold, C. E., Rational runoff formulas: Am. Soc. Civil Eng. Trans., vol. 96, pp. 1038-1175, 1932.



and procedures applicable to small drainage areas for translating rainfall into rates of stream flow in terms of basin characteristics.

Bernard⁴ carried Gregory and Arnold's expressions somewhat further and presented formulas applicable to a few selected basins.

⁴ Bernard, M. M., An approach to determinate stream flow: Am. Soc. Civil Eng. Trans., vol. 100, pp. 347-395, 1935.

⁷⁴⁷⁰⁴⁹⁻⁴⁷⁻²

Sherman⁵ in 1932 presented unit hydrographs for four streams whose drainage areas and slopes differed widely. He explained how the unit hydrographs expressed these differences and suggested that basins having physical characteristics similar to the four types presented would have similar hydrographs.

McCarthy ⁶ in 1937 also used the unit hydrograph as an expression of the runoff characteristics of a drainage basin, stating further that "the agreements between graphs developed from May and November storms substantiate the contention that primarily the unit hydrograph is a function not of surface cover, which may be subject to seasonal change, but of topographic features of a watershed." From this postulate he derived a working relationship between the crest discharge and the length of base of unit hydrographs of 25 drainage basins in the Connecticut River Basin, in terms of the area, mean slope, and stream pattern, determined by inspection and expressed as one-stem basin, two-stem basin, and so on, for application to floodcontrol design.

Morgan and Hullinghorst ⁷ in 1939 stated:

The factors which determine the discharge characteristics of any watersned are innumerable, some having a major bearing on those characteristics while others are of negligible consequence. It was determined by examination of nine gaged basins having complete unit hydrograph and watershed data, and corroborated by examination of a number of gaged basins with data ranging from almost complete to fragmentary, that the discharge characteristics can be attributed principally to three fundamental, definite watershed characteristics, namely,

a. Area of the watershed in square miles,

b. Mean length of travel in miles, and

c. Mean height of watershed above outflow station in feet.

On this basis, empirical relations between these three factors were established for nine streams tributary to the Chemung River in New York.

PURPOSE AND SCOPE OF THE PRESENT STUDY

The present project was designed to provide basic material whereby investigations such as those outlined can be carried further, the range being limited, of course, to areas adequately mapped, which are mainly in the northeastern United States. Prior to this study, references to the subject were read with the view to determining which topographic factors were considered to have major influence upon

^b Sherman, L. K., The relation of hydrographs of runoff to size and character of drainage basins: Am. Geophys. Union Trans., No. 13, pp. 332-339, 1932.

⁶ McCarthy, G. T., The unit hydrograph and flood routing (unpublished manuscript presented at conference of North Atlantic Division, Corps of Engineers, U. S. Army, June 24, 1938).

⁷ Morgan, R., and Hullinghorst, D. W., Unit hydrographs for gaged and ungaged watersheds: U. S. Engineer Office, Binghamton, N. Y., July 1939. [Processed.]

discharge characteristics, and, so far as practicable, the suggestions thus obtained were incorporated in the project as proposed for cooperation to the Works Progress Administration (succeeded on April 25, 1939, by the Work Projects Administration). The compilation was based upon the topographic maps of the Geological Survey covering areas tributary to gaging stations of the Geological Survey.

In the organization of a surface-water system, and of a large part of the ground-water system as well, the drainage basin is a natural hydrologic land unit. Surface runoff is divided into drainage basins by the watersheds, and within each basin it follows a system of water courses in which the flow undergoes retardation, acceleration, or other changes that are distinctly related to the physical characteristics of that basin. Similar conditions exist with respect to all or most of the ground-water runoff. Essentially all the water within a given basin, except that which is lost by evaporation or transpiration, drains out through a common outlet or mouth.

For purposes of analysis, a major stream basin may be subdivided by considering the area tributary to the stream at any given point, for example, a gaging station, as a basin having its own characteristics. The separate characteristics of several contributory areas may then be combined to obtain the resultant for the major basin.

Geographic and topographic characteristics of drainage basins, based largely on certain horizontal and vertical dimensions, were selected for compilation and study. Geographic characteristics include water bodies, direction of stream flow, latitude, and longitude. Topographic characteristics include horizontal dimensions covering basin area, stream length, and area-distance distribution, and vertical dimensions covering land slope, tributary and principal stream slopes, and basin altitude.

In selecting basins for this study preference was given to those for which long-term stream-flow records are available and to those free from artificial regulation. In addition to areas in the northeastern States and the Ohio River Basin a few surveyed areas in Wisconsin and Kentucky were included to spread the range in geographic extent and topographic characteristics. Many basins in New York, Pennsylvania, and New Jersey that might otherwise have been included were not studied because of insufficient time.

Besides listing topographic and hydrologic data, the original records afford a gazetteer of streams and lakes. Maps were prepared showing the stream skeleton of each basin, with names of streams, length from confluence to confluence, and drainage areas and altitudes pertinent to the subdivisions. A list of lakes and ponds, giving names, locations, and approximate altitudes and areas was also prepared, much of the data for basins in New York State being based on a gazetteer by E. M. Douglas.⁸

The summarized results of the compilation, covering about 340 basins, are given in the table on pages 145–155. The original records are on file at the office of the Geological Survey in Washington, D. C., and the computations for basins in New England (except Maine) are on file also at the Boston office of the Geological Survey.

METHODS OF WORK

MAPS

Quadrangle maps on the scale 1:62,500, are the basis of this compilation, except for a few areas in New Jersey where more detailed maps were available. On these maps the gaging stations were located and the tributary basins with sub-basin divisions were outlined. Generally, each basin was divided into 50 to 75 sub-basins. Care was taken that the sub-basins crossed the streams only at confluence points. To systematize the necessary tabulations, the sub-basins were numbered in accord with the following system, which is illustrated on plate 2.

The headwater basin farthest upstream (the one farthest removed from the gaging station along the main stream) is called no. 1; the sub-basin which it joins at the first confluence point is called no. 2; the sub-basin (or intervening area along the combined channel) below the confluence of sub-basins 1 and 2 is called no. 3; the next tributary sub-basin is called no. 4, and so on. Where a large tributary stream that has been subdivided joins the main stream, the next consecutive number is assigned to the farthest upstream sub-basin of this tributary. This constitutes the lowest number on such tributary, and the sub-basins of the tributary system are then numbered in the same manner as those of the main stream, down to the sub-basin immediately above the confluence of the tributary with the main stream. The succeeding number is assigned to the sub-basin along the main river immediately below the confluence, as before. The highest number in a basin is that of the sub-basin immediately above the mouth of the main stream: it indicates the number of sub-basins into which that drainage area is divided. As shown on plate 2 the number of each sub-basin is the large integer near its center; the smaller figures represent area, length of stream, and altitude.

⁸ Douglas, E. M., Gazetteer of the lakes, ponds, and reservoirs of the State of New York: Map Information Office, Board of Surveys and Maps, 44 pp., Washington, 1926. [Processed.]

AREA OF BASINS

The total area of the basin within the watershed lines above the selected gaging station is the primary basin factor. In a humid climate the volume of discharge varies directly with the size of the tributary drainage area. Accordingly, the area in square miles was measured, not only of the main basin above the gaging station but also of a number of sub-basins (generally over 50). The size of the basins included in this compilation ranges from 1.64 to 7,797 square miles. In general, large streams were excluded because the size of sheets became unwieldy, because some contained unsurveyed areas, and because their stream-flow characteristics could best be determined by synthesis of their components.

STREAM DENSITY

The runoff from the several parts of the drainage basin is discharged by the streams, and, other factors being constant, the time required for the water to flow a given distance is directly proportional to the stream length. The stream or drainage density is the ratio between the total length of all streams within the drainage basin and the total area of the basin and is an indication of the drainage development. Accordingly, the length of all streams down to the smallest shown on the topographic maps was measured to determine the stream density and the area-distance distribution.

The number of small headwater streams shown on the topographic maps would vary with the season and the wetness of the particular year during which the survey was made, as well as with the judgment of the topographer and cartographer as to the amount of detail to be shown on the map. These circumstances introduce a measure of inconsistency in stream-density results as determined from maps.

The ratio of stream density for the basins included in this compilation, all of which are in the humid region, ranges from 0.89 to 3.37 miles per square mile and averages 1.65 miles per square mile. Other factors being equal, high drainage density indicates a more effective operation, of the agencies of stream incision. Greater incision, for example, would be associated with steep land slopes. Opportunity for incision would be greater also where most of the discharge occurs as surface runoff rather than through ground-water channels; such a condition exists in areas where the ground is sufficiently impervious to shed storm rainfall. Drainage density is greater in humid regions than in arid regions; it would approach zero in flat, sandy desert regions and would approach a maximum in steep, rocky, humid regions. The variation of stream density with the land slope is shown by the following data derived from groups of basins in New England.

| Variation of s | tream density | with lan | d slope |
|----------------|---------------|----------|---------|
|----------------|---------------|----------|---------|

| Range in stream density (miles per square mile) | Average land slope (feet per mile) |
|--|---|
| 1.00 to 1.25 | |
| 1.26 to 1.50 | 550 |
| 1.51 to 2.00 | 600 |
| 2.01 to 2.25 | |

The mean land slopes for basins in New England having drainage densities within the ranges indicated in the above data were averaged, and the results indicate that, in general, in a given region the higher drainage densities are associated with the steeper land slopes. The reciprocal of the drainage density is the average distance between streams, and half the reciprocal of drainage density is the average horizontal distance between the streams and appurtenant watershed lines, measured at right angles to the streams. Drainage density appears to be inversely related to the distance of overland flow as distinguished from flow in stream channels. However, in basins sufficiently permeable so that all rainfall can be taken directly into the soil through infiltration, the drainage density approaches zero and is associated with zero overland flow.

AREA-DISTANCE DISTRIBUTION

The concentration of runoff from drainage basins of equal size may be greatly affected by the distribution of the area with respect to distance from the gaging station or outlet. Other factors being equal, the runoff from areas close to the gaging station should reach it sooner than water from remote areas. Accordingly, a drainage basin whose tributaries are compactly organized, so that water from all parts of the basin has a comparatively short distance to travel, will discharge its runoff more quickly and reach greater flood crests than one in which the larger part of the area is remote from the gaging station or outlet. This basin characteristic is expressed in the summary table by the quantity Σal , computed by multiplying each partial area in the basin (a) by the channel distance from the midpoint of the main stem serving it, downstream to the gaging station (l). Distances along the stream channels were measured in 0.1 mile chords.

In a sense this quantity is also a measure of the volume of channel storage in the basin. For example, if under a given regime of flow the cross-sectional area of a river at a given place varies directly as the drainage area above, then the volume in any given reach would vary as the product of a coefficient by mean drainage area above the

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reach by the length of the reach. The coefficient would be a function of the stage of the regime selected, the slope of the reach, frictional resistance, and other hydraulic factors. No method is proposed for evaluating the coefficient. However, the sum of the products of mean drainage area and length of reaches for a given basin is equal to the product Σ al which was derived by both methods of computation.

The most compact drainage basin would be a glory-hole inlet, and the product Σ al for such a basin is 0.375 A^{1.50} where A represents total area; for an equilateral triangle, with reference to an outlet at one of the vertices, the product is 0.94 A^{1.50}; and for a square, with reference to a corner, it is 0.76 A^{1.50}. Figure 49 shows the results of plotting the products Σ al against the corresponding drainage area. Only enough points are shown to define the line of regression, whose equation is 0.90 A^{1.56}, or more approximately 1.2 A^{1.50}, within the range shown. Natural basins are generally less compact than any of the geometric shapes mentioned.

Additional subdivision of a basin beyond the 50 to 75 sub-basins generally used would tend to increase the value of the product Σal . However, a study of West River Basin, above Newfane, Vt., indicates that the product Σal for 20 sub-basins was 6,620, for 50 sub-basins 6,810, and for 100 sub-basins 6,860. The values given in this report may therefore be considered essentially correct limiting values.

Points on the right of the trend line (fig. 49) represent basins less compact than the average, and those on the left the more compact. The regression line therefore furnishes a standard for comparing the relative compactness of different basins.

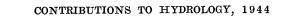
LENGTH OF BASIN

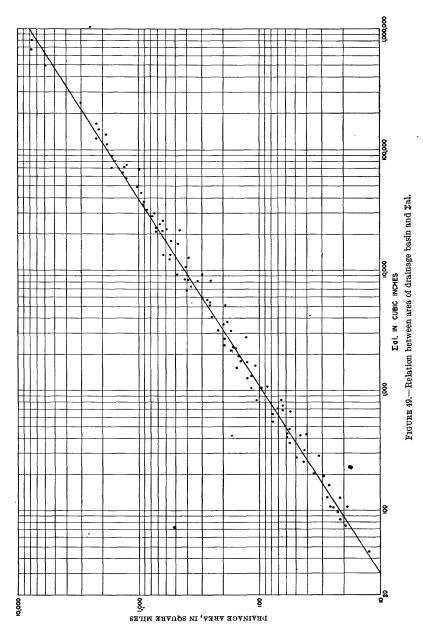
The table (pp. 145–155) lists the length of longest watercourse in each basin measured in 0.1-mile chords to the source of the most headward stream. This length, when divided by the mean velocity of flow will give the time of concentration as used in the rational formula for the computation of flood discharge.

The mean length of travel of runoff or the distance to the center of gravity of the drainage system may be found by dividing the quantity Σ al by the drainage area in square miles. This quotient is commonly identified by the symbol L_{ca} . The table also lists the length of principal streams as defined under "Channel slope."

LAND SLOPE

Rainfall or snow melt which becomes direct runoff flows over the surface of the ground or, where the surface soil is shallow and permeable, immediately beneath it over the bedrock. The average distance water travels before entering a stream channel may be





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expressed in terms of stream density. (See p. 133.) The rapidity with which the water travels to the streams likewise depends on the slope of the land.

The contours on the topographic maps provide a basis for determing the slope of the land by the intersection-line method outlined by Horton⁹ as follows:

The intersection-line method.—In order to reduce the labor of computation of slope of large areas the author has utilized the following method. An area the slope of which is to be determined is subdivided into squares of equal size by lines forming the boundaries between adjacent squares. The number of contours crossed by each subdividing line is counted and the lengths of the lines are scaled. Then the average scale-distance l' between contour crossing in the subdivision lines is

$$l' = \frac{\Sigma l}{N}$$

where N is the number of contours crossed and Σl is the total length of the subdividing lines. If α is the horizontal angle at which each of two parallel contours crosses an intersection line, then $l' \sin \alpha$ is the horizontal distance between the two contours measured normal to the contours. Contours may cross the intersection lines at all angles from 0° to 90°. The mean value of sin α for angles from 0° to 90° is

$$\int_{0}^{\frac{\pi}{2}} \frac{\sin \alpha d\alpha}{\pi} = \frac{2}{\pi} = 0.6366$$

If D is the contour interval or difference in elevation in feet, and L is the average normal horizontal distance between contours, then

$$L = 0.6366 l'$$

and the mean slope Sg of the area is

$$Sg = \frac{D}{0.6366 \frac{\Sigma l}{N}}$$
$$= 1.571 \frac{DN}{\Sigma l}$$

In applying this method it is assumed that each contour crossed represents a difference of elevation along the subdivisional line equal to the contour interval. Of course it may happen that two adjacent contours are at the same elevation and are separated by land only a little higher or lower. On an average, however, the elevations of summits or depressions between equal contours will differ from that of the adjacent contours by an amount equal to one-half the contour interval, and it can readily be seen that the average declivity between a pair of contours of equal elevation is nearly the same as if the contours were separated by the contour interval D, so that the method gives nearly correct results even where the subdivision lines cross adjoining contours of equal elevation, as in the case of summits and depressions.

[•] Horton, R. E., Drainage-basin characteristics: Am. Geophys. Union Trans., No. 13, pp. 350-361, 1932. 747049-47----3

By making the subdivision lines sufficiently frequent, the average slope of an area may be determined with whatever degree of accuracy is required.

This method has been tested by comparison of slope for the same area computed from the measured total lengths of contours, with, in general, good agreement.

In carrying out this computation, the slope along the meridian lines is computed separately from the slope along the parallels of latitude. Where there is a great difference between the land slope in the two directions, the orientation of the basin is determined by the axis of least slope. Where the east-west slope and the north-south slope are nearly the same, the line of orientation may be approximately midway between the two, or it may not be clearly defined in either direction in a cup-shaped or fan-shaped basin. Land slopes listed in the table range from 1,598 feet per mile for the upper Pemigewasset River Basin in New Hampshire to 55 feet for Great Egg River in Coastal New Jersey.

Paulsen¹⁰ found, during the flood of September 1938 in the North Atlantic States, that the infiltration index tends to increase with decrease in mean land slope. He states that "although the slope of the land itself might influence the retentive capacity of the ground, this tendency may be due to other factors related to slope, such as depth of soil cover."

CHANNEL SLOPE

Upon leaving the land the runoff enters the channel system, through which it flows in channels that increase progressively in size with the entrance of additional water. Channels in a drainage basin are classified for study as principal and tributary. The principal streams of a basin are defined as those that drain 10 percent or more of the total area of the basin; tributaries are defined as those that drain less than 10 percent of the area of the basin. The average slope of the tributaries and of the principal streams is computed separately as the quotient of the total fall divided by the corresponding total length and is reported in the summary table.

In computing the slope of the stream channels, only the largest stream in each sub-basin is considered. Thus, if a basin is divided into 75 sub-basins only 75 stream lengths and falls are measured. These stream lengths are classified as principal or tributary, and the average slope of each is computed. As only one stream in each subbasin is included in the classification, many minor headwater streams are excluded from consideration; consequently, the reported slope of the tributary stream is affected by the number of subareas into which the basin is divided. The reported mean slope of the tributary

¹⁰ Paulsen, C. G., Hurricane floods of September 1938: U. S. Geol. Survey Water-Supply Paper 867, pp. 440-441, 1940.

streams increases as the number of subareas becomes larger, thus embracing more steep minor headwater streams. This is illustrated in figure 50, which shows the result of a comprehensive study of the slope of tributaries of West River at Newfane, Vt. The asymptote resulting from that study is about 225 feet per mile, whereas the channel slope obtained with 53 subareas (see table, No. 1-354) is 200 feet per mile.

The slope information for the several drainage basins listed in the table discloses that a steep land slope is generally associated with

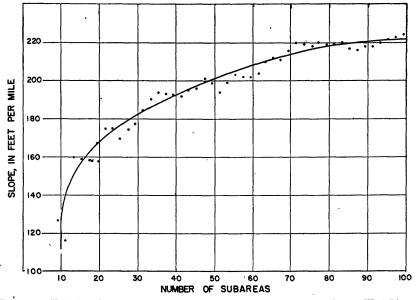


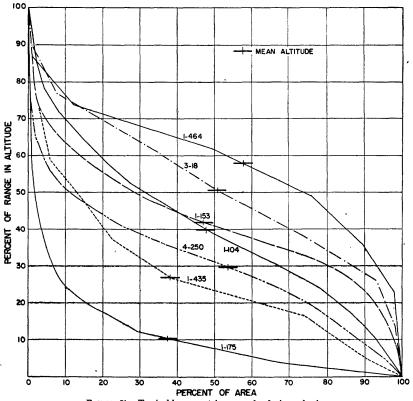
FIGURE 50.—Variation of computed slope of tributary streams with number of subareas, West River at Newfane, Vt.

steep tributary and principal channel slopes and conversely, as might be expected. There is, however, no systematic variation; moreover, according to geometric analysis by Horton,¹¹ the ratio between principal channel slope and average land slope is a measure of the horizontal angle that the lateral slope makes relative to the stream slope. A low slope ratio indicates that the laterals tend to enter the. streams at right angles, whereas the angle of inflow into the stream becomes more acute as the channel slope approaches equality with the ground slope. This slope ratio tends to decrease with increase in drainage area, but varies considerably between drainage basins of equal size.

¹¹ Horton, R. E., op. cit., p. 360.

AREA-ALTITUDE DISTRIBUTION

Another method of expressing the slope of the basin is by means of the altitude of the several parts with reference to sea level. This is best expressed through the hypsometric curve, as a graph showing the area-altitude distribution is called. Although the area-altitude distribution was derived for each basin, only the maximum, mean, and minimum altitudes as determined from the topographic maps are shown in the table. From this information, however, the areaaltitude distribution curve can be readily approximated. Figure 51 shows the hypsometric curves for several basinc plotted in terms of percent of range in altitude and percent of area above the indicated altitude. The variations are wide, but in general the mean altitude





1-104. Swift River near Roxbury, Maine.

1-153. East Branch of Pemigewasset River near Lincoln, N. H.

1-175. Lake Winnipesaukee outlet at Lakeport, N. H.

1-435. Quinnipiac River at Wallingford, Conn.

1-464. Leadmine Brook near Thomaston, Conn.

3- 18. Brokenstraw Creek at Youngsville, Pa.

4-250. Otter Creek at Center Rutland, Vt.

of a basin is located at 0.34 of the range between the minimum and maximum; thus a basin is comparable to the surface of a cone.

The area-altitude distribution curve has several applications. For example, snow surveys generally show an increase in depth of cover and water equivalent with increase in altitude; the area-altitude relation provides a means for estimating the mean depth of snow or its water equivalent over a drainage basin. Barrows¹² describes a significant variation in annual precipitation and runoff in the Connecticut River Basin with respect to altitude. The obvious variation in temperature with change in altitude is further indication of the utility of the area-altitude distribution curve.

The mean altitude of the basin above the altitude at the outlet or gaging station represents the potential head of a uniform depth of water over the basin with respect to the outlet or gaging station, and as such is a factor in determining the rate at which the waters are collected and discharged. The data in the summary table shows that, in general, the land slopes and channel slopes vary with the mean altitude of the basin above the outlet. Thus steep slopes are associated with a high altitude above the outlet, and conversely. A rough average relation between slope and mean altitude is as follows:

$$h = K_1 S_1 + K_2 S_i + K_3 S_p$$

where $K_1 = 0.31$

 K_2 ranges from 0.97 at 50 square miles to 3.0 at 1,000 square miles. K_3 ranges from 3.5 at 50 square miles to 23.4 at 1,000 square miles. S_1 =mean land slope, in feet per mile.

 S_t =slope of tributary streams, in feet per mile.

 S_p = slope of principal streams, in feet per mile.

AREA OF WATER SURFACES

The effect of storage in retaining flood runoff and prolonging its discharge until the flood in channels farther downstream has begun to subside tends to reduce flood peaks and increase the time lag between rainfall and its consequent runoff. Natural storage in lakes and ponds and artificial storage in reservoirs aids this retardation. A measure of the amount of storage available for such modification of flood discharge can be derived from the surface area of the water bodies shown on the topographic map. (See table.) The computations at Boston included the determination of swamp areas, which had been part of an earlier project carried on in 1936 in cooperation with the Works Progress Administration; this covered compilations for the Merrimac and Connecticut River Basins, both in square miles and

¹³ Barrows, H. K., Precipitation and runoff and altitude relations for Connecticut River: Am. Geophys Union Trans., 14th Ann. Meeting, pp. 396-406, 1933.

in drainage percent.¹³ The computations made at Pittsburgh did not include swamp areas.

The areas of swamps as reported would be affected by the hydrologic conditions under which the topographic surveys were made. Surveys made in spring or early summer would probably show a greater swamp area than those made in late summer or fall, and surveys in a wet year would show marked contrast with those made in a dry year. It is not known to what extent the results given in the table were affected by hydrologic conditions.

It should be pointed out that the area of water surfaces is only one measure of their effect on the time distribution of flood discharge. The position of the water bodies in the river system is also important; thus a large pond near the headwaters would affect but a small part of the runoff, whereas one of the same size farther downstream would affect a larger part of the runoff.

In addition to the effect of storage in modifying the shape of flood waves or the time distribution of runoff, the total volume of runoff may be influenced by evaporation from lakes, reservoirs, and swamps. The loss of water by evaporation from water surfaces in the northeast is about twice that from land surfaces, per unit of area. Accordingly, basins with a large proportion of water and swamp surfaces may be expected to yield less runoff than those with a small proportion.

In many of the basins listed in the table the proportion of lake and swamp areas exceeds 10 percent, and in a few, especially in New England, it approaches 20 percent; doubtless the effect on water losses is significant. The percentage of lake area is highest in New England and northern New York and generally in the glaciated portions of the areas studied.

SUMMARY OF RESULTS

The summary table that follows gives the results of measurements on topographic maps on a scale of 1:62,500. It includes about 22,000 areas covering 145,000 square miles. A total of 240,000 miles of stream length was measured, and nearly a million contours on the topographic maps were counted and translated into land and channel slopes.

Reference has already been made to general relationships between the topographic factors listed in the table. Each item is not necessarily unique, but it may reflect a condition that also influences the others, consequently other relationships between them may be found. For example, figure 52 shows that, in general, larger drainage areas are associated with flatter stream slopes; but average land slopes and

¹³ Grover, N. C., The floods of March 1936, pt. 1: U. S. Geol. Survey Water-Supply Paper 798, pp. 335-338, 1937.

mean altitudes of drainage basins above outlets or gaging stations show a tendency to increase with drainage area. The points shown on figure 52 correspond to averages of groups of drainage basins within limited ranges in size. If individual basins were plotted on figure 52, material scattering of points would result, the basins plotting on the left being relatively flatter than those on the right. The average curve therefore provides a means for comparing the slopes and alti-

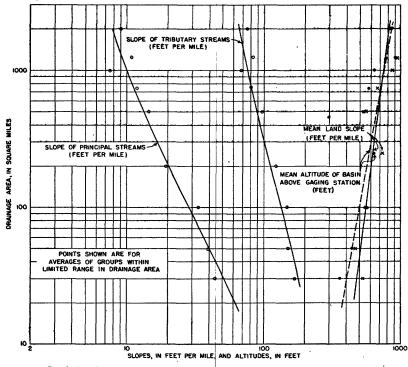


FIGURE 52.—Graph showing general variation in stream slopes and altitude in relation to size of drainage basin.

tudes of basics of different size. The divergent trends of the lines showing principal channel slope and average land slope indicate that the slope ratio of the basins analyzed tends to decrease with increase in drainage area.

A principal shortcoming of the computations of physical characteristics may be that it was not practicable to determine in detail the distribution of stream and land slopes and of lake and swamp areas within every area. Steep slopes on a few tributaries may increase the average slope considerably, yet these slopes may have little effect on flood-peak discharges. Moreover, the course of a river in a given length may be characterized either by uniform slope or by a series of pools with intervening rapids or perhaps cataracts. The velocity in a stretch of uniform slope would probably be the greater if other factors were constant, as pools have a detention storage effect and the fall at rapids or cataracts imparts but little horizontal velocity to the water. A lake on the headwaters of a stream may have no noticeable effect, whereas a lake of the same size on the main channel near the lower reaches of an otherwise flashy stream may greatly modify flood discharges. Also, steep slopes or abnormally high elevations in the part of a drainage basin upstream from a lake may affect considerably the average land slope and the mean elevation of a basin, but the lake may decrease the flood discharges so much that the outflow from the lake would differ little from that of a basin in which the slopes and elevations were much less.

Storage capacity was not computed, as topographic maps furnish no information from which channel and lake cross sections at different stages can be determined, except that they might be crudely correlated with the stream slopes.

| v pich | to tedmuN votni zgers zgw nizgd babivib | 65 31 | 818 | 67 67 | 48 | 51 | 46 46 50 45 50 45 50 45 50 45 50 45 50 45 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 50 50 50 50 50 50 50 50 50 50 50 50 | 52 62 46 | 49 59 | 48 48 48 | 39 |
|-------------------------------------|--|------------------|---|---|---|--|---|---|-----------------------------------|---|--|
| water quare | [stoT | | | | 0.09 | 1.16 | $\begin{array}{c} 7.24\\64\\ 1.58\\ 13.42\\ 3.31\\ 3.31\end{array}$ | 92.32 35.17 4.62 | 8.32 7.22 | 85.2.88 35.58 35.58 | 12.00 |
| es (s | sqmsw2 | | | | 0 | . 72 | 4.88 .58 1.32 2.57 | 8.48 25.56 2.66 | 5.38 5.17 | $\begin{array}{c} 4.95\\ 2.76\\49\\49\\23.94\end{array}$ | 10.58 |
| Area e surfac miles) | Lakes and reser- | 53.91 5.32 | 22. 79 16. 44 3. 14 | 5.54 .17 .06 | 60, | .44 | 2.36 .06 .72 .74 | 83.84 9.61 1.96 | 2.94 2.05 | 3.93 .64 2.37 11.56 | 1.42 |
| of land a hove sa level) | muminiM | 512 161 | $159 \\ 600 \\ 290 $ | 201 804 390 | 1, 020 | 620 | 460 580 560 470 | 495 270 880 | 380 430 | 340 290 65 | 35 |
| ea of | пвэМ | 860 451 | $^{440}_{1,765}$ | 959 1, 760 935 | 2,804 | 2,492 | $^{1,853}_{1,744}$ 1, 744 1, 584 1, 584 1, 258 | $^{756}_{1, 038}$ | 965 | 820 868 868 868 868 868 | 126 |
| Altitude (feet mean s | mumizeM | 2, 000 1, 463 | $\substack{1, 463\\4, 168\\4, 237\end{array}$ | 4, 116 3, 535 2, 420 | 5, 249 | 5, 249 | 5, 249 4, 810 2, 280 2, 920 2, 920 | 2,979 3,165 2,496 | 2, 722 2, 937 | $\begin{array}{c} 2,378\\ 2,280\\ 1.990\\ 755 \end{array}$ | 400 |
| sms9. | lo dtgngth of 12 lagio (miles) | 82.2 28.6 | 51. 1 73. 3 51. 6 | 62 20.8 18.4 | 18.7 | 23.9 | $\begin{array}{c} 52.4\\ 13.0\\ 15.3\\ 20.5\\ 20.5 \end{array}$ | 39.9 21.3 | 26.3 31.1 | 24.7 31.9 64 | 21.5 |
| | Longest wat willes | 75. 0 30. 9 | 41.5 81.3 41.7 | $ \begin{array}{c} 61.3 \\ 17.9 \\ 15.3 \end{array} $ | 15.3 | 23.6 | 41.8 14.7 15.0 22.5 17.0 | 34.0 72.0 23.7 | 23.9 33.1 | 26.6 34.1 20.6 51.6 | 27.5 |
| nnel (feet aile) | Indionira | 2.56 7.73 | $ \begin{array}{c} 5.55 \\ 12.88 \\ 26.71 \\ 26.71 \\ \end{array} $ | 11.95 79.57 34.5 | 117 | 76.2 | $\begin{array}{c} 21.4\\ 109.2\\ 54.9\\ 11.1\\ 33.7\\ \end{array}$ | $^{25}_{9.78}$ | 24.3 16.6 | 34.6 34.6 2.5 2.8 | 2.33 |
| Channel slope (feet per mile) | Tributary | 52.7 51.1 | 49 146 165 | 147 454 258 | 498 | 426 | 223 482 326 212 242 | 96 90.2 208 | 163 178 | 109 115 24 | 15.9 |
| (feet | A verage | 449 298 | 314 694 745 | 672 1, 044 818 | 1, 598 | 1, 478 | $1, 161 \\ 1, 004 \\ 878 \\ 603 \\ 603 \\ 764$ | 460 575 610 | $652 \\ 674$ | 511 497 416 261 | 194 |
| Land slope (feet per mile) | | 445 294 | 303 742 707 | $ \begin{array}{c} 662 \\ 662 \\ 730 \\ 730 \end{array} $ | 1,468 | 1, 363 | 1,089 870 565 565 697 | 438 522 542 | 590 627 | 505 475 372 244 | 192 |
| Land | ЕW. | 456 304 | 322 630 797 | $^{687}_{1, 088}$ | 1, 774 | 1,636 | ${ \begin{smallmatrix} 1.& 262\\ 1.& 188\\ 998\\ 654\\ 852\\ 852 \end{smallmatrix} }$ | 491 646 711 | 734 737 | 519 525 474 285 | 197 |
| (səlin | z al (cubic miles) | | 7, 010 35, 524 7, 114 | 16, 641 1, 051 622 | 824 | 2, 355 | $13, 156 \\ 1, 735 \\ 1, 735 \\ 1, 081 \\ $ | 7, 274 28, 221 669 | 1,970 2,747 | 2,270 3,154 1,168 12,390 | 1, 714 |
| Vilan Susupe | Stream de (miles per mile) | $1.18 \\ 1.10 $ | $1.08 \\ 1.42 \\ 1.79 $ | $\begin{array}{c} 1.47 \\ 1.57 \\ 2.34 \end{array}$ | 1. 45 | 1.50 | 1.62 1.62 1.62 2.05 2.05 | $ \begin{array}{c} 1.46 \\ 1.89 \\ 2.32 \end{array} $ | $1.54 \\ 1.65$ | $\begin{array}{c} 1.93\\ 2.04\\ 1.59\\ 1.50 \end{array}$ | 1.23 |
| (səli | Drainage məraups) | 867 148 | 299 872 351 | 514 95 76, 2 | 104 | 192.6 | 622 58.8 57.6 85.8 | 363 766 54.8 | 146. 2 129 | 157 170 107 405 | 124 |
| | Name of gaging station | | Vature. Passadumkeag River at Lowell, Maine Dead River at The Forks, Maine Carrabassett River near North Anson, Maine. | щõН | Maine. East Branch of Pernigewasset River near | Pemigewasset River at Woodstock, N. H. | Pemigewasset River at Plymouth, N. H. Bakers River at Wentworth, N. H. Bakers River hear Runney, N. H. Squam River at Ashland, N. H. | ROH | Warner River at Davisville, N. H. | ww.do | at Lowell, Mass. Ipswich River near Ipswich, Mass |
| | 0. V 0. | | 1-34 1-52 1-54 | 1-58. 1-104. | 1-153 | 1-153A | l-155 1-169 1-170 1-171.A | 1-175. 1-183. 1-186. | 1-188A 1-189 | $\begin{array}{c} 1-192 \\ 1-198 \\ 1-202 \\ 1-210 \end{array}$ | 1-213 |

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Summary of drainage basin topographic characteristics

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| | dəidw | Number of sreasinto asw nissd babivib | 882988 | 32522 | 865550 865550 865 | 52 88 89 89 80 80 80 80 80 80 80 80 80 80 80 80 80 | 282 282 282 282 282 282 282 282 282 282 | 80 31 50 50 50 50 50 50 50 50 50 50 50 50 50 |
|---|-------------------------------------|--|--|--|--|--|---|---|
| | water (square | IstoT | $\begin{array}{c} 1.02\\ 6.42\\ 12.36\\ 1.06\\ 26.19\\ 26.19\end{array}$ | $1.17 \\ 1.39 \\ 5.98 \\ 15.25 \\ 1.50 \\$ | $\begin{array}{c} 1.44\\ 4.52\\ 12.90\\ 1.88\\ 1.88\\ 7.54\end{array}$ | 1.00 4.52 4.10 13.16 | 25.47 1.18 3.77 2.44 | 3.11 1.08 2.83 |
| | of aces es) | eqmaw2 | 0.48 3.54 8.50 12.24 | 0.2.035 | 0 8.02 3.42 3.42 | 3.22 3.22 3.04 3.04 | 9.28 0.76 3.16 .42 | 2.28 .08 1.33 |
| | Area surf mil | Lakes and reser- voirs | $\begin{array}{c} 0.54\\ 3.86\\ 13.95\\ 13.95\end{array}$ | $ \begin{array}{c} 92 \\ 5.90 \\ 12.96 \\ 1.50 \\ \end{array} $ | 4.1.8888 4.1.8888 1.221 | 1. 30 3. 46 3. 46 10. 12 | $\begin{array}{c} 16.19\\ 1.18\\ 1.18\\ 1.64\\ 2.02\\ 2.02\end{array}$ | $\begin{array}{c} \cdot & 01 \\ 0 & \cdot & 83 \\ 0 & \cdot & 50 \\ 1 \cdot 50 \end{array}$ |
| | f land bove level) | muminiM | 100200 | 70 480 270 360 360 | 360 200 250 15 150 160 | 280 280 280 280 280 280 280 280 280 | 2220 110 110 | 1, 190 550 640 840 840 |
| | 6890 | nsəM | 124 211 213 99 | 195 746 611 495 495 | 637 500 392 620 620 | 605 614 779 696 | 553 744 536 512 512 | 2, 512 1, 712 1, 324 1, 324 1, 402 |
| | Altitude (feet mean s | mumizsM | 580 580 580 580 580 580 580 580 580 580 | $\begin{array}{c} 440\\ 1,400\\ 1,400\\ 1,400\\ 1,760\\ 760\\ \end{array}$ | 910 805 805 1, 280 1, 290 | $\substack{1,290\\1,280\\1,280\\1,280\end{array}$ | 1, 280 1, 080 860 760 | ອີກ ອີກ ເຊິ່ງ ເຊີ່ງ ເຊີ່ງ ເຊີ່ງ ເຊີ່ງ ເຊີ່ງ ເຊີ່ງ ເຊີ່ ເຊີ່ ເຊີ່ ເຊີ່ ເຊີ່ ເຊີ່ ເຊີ່ ເຊີ່ |
| | -ninq ខណៈនេ១។ | Length of ta lagio (miles) | 9.4 31.9 14.6 36.6 | 15.3 13.3 26.5 10.5 | 11.4 19.9 39 49.6 | 16.9 33 17.8 24.7 25.3 | 20.6 20.6 20.6 20.6 | 222.84 322.84 20.22.84 |
| | s) fercourse | aw tesgno.I slim) | 38.7 39.7 26.7 26.7 | 15.0 11.5 26.5 9.8 | $\begin{array}{c} 7.6\\ 27.7\\ 32.6\\ 32.6\end{array}$ | 19.6 26.5 28.3 38.1 | 65.3 9.5 16.5 16.5 | 20.2 31.3 15.2 |
| | nnel (feet nile) | Principal | $\begin{array}{c}12.8\\4.24\\3.57\\3.28\\3.28\end{array}$ | 11.6 41.5 12.3 20.0 | 46.1 21.6 15.1 25.1 13.3 | 20.7 22.0 8.97 12.2 13.7 | 12.1 37.4 26.2 24.3 24.9 | 75.8 35.9 21.8 39.2 |
| | Channel slope (feet per mile) | Tributary | 49.6 22.6 35.3 13.5 | 24.0 77.9 58.7 41.5 101 | 97.6 59.0 55.6 91.4 61.7 | 114 79.9 60.2 56.1 51.0 | 42.1 76.6 68.7 68.0 64.0 | 575 274 350 375 234 |
| | (feet | A verage | 301 236 244 176 116 | 141 371 337 307 301 | 260 260 378 344 | 334 341 415 369 369 | 328 348 262 305 305 | $^{1, 130}_{1, 295}$ |
| | Land slope (fect per mile) | 'S~'N | 257 210 223 159 105 | 129 308 279 262 262 262 | 275 213 213 312 293 | 303 330 330 330 330 330 330 330 330 330 | 261 274 237 220 271 | $1,056 \\ 1,235 \\ 693 \\ 486$ |
| 5 | Land | | 359 273 200 129 | 158 458 414 370 374 | 293 323 298 471 414 | 377 424 579 549 478 | 420 449 355 319 351 | $1,225 \\ 1,034 \\ 1,378 \\ 1,168 \\ 1,168 \\ 734 \\ $ |
| | (səlim | oiduo) le Z | 97 3,638 8,126 3,738 3,738 | 327 164 1, 861 8, 184 8, 184 | 117 966 3, 416 1, 775 6, 815 | 2, 320 2, 320 2, 362 6, 332 | 22, 477 141 617 723 738 | 968 8, 575 4, 272 172 622 |
| | Vilens Square | Stream de (miles per mile) | 2.34 1.27 1.27 1.28 | 1.36 1.36 1.37 1.57 | 1.14 1.47 1.51 1.69 1.69 | 1. 61 1. 63 1. 63 1. 55 1. 55 | 1.56 1.62 • 1.61 1.51 1.51 | 1. 72 1. 53 1. 60 |
| | area iles) | oganiar (m 91.sups) | $23.3 \\ 251 \\ 35.2 \\ 260 \\ 200 \\ 20$ | 42.2 31.3 139.2 417 26.1 | 27.8 93.3 199.8 121 401 | 76.2 169 93.8 157 331 | 711 27.7 58.7 83.5 88.6 | 88. 5 395. 5 30. 5 30. 5 80. 5 |
| | | Name of gaging station | Aberjona River at Winchester, Mass. Charles River at Charles River Village, Mass. Neponset River at Waltham, Mass. Neponset River at Norwood, Mass. Taunton River at State Farm, Mass. | Wading River near Norton, Mass Blackstone River at Worcester, Mass Blackstone River at Wortbridge, Mass Blackstone River at Woonsocket, R. 1 Quinsigamond River at North Grafton, Mass. | Mumford River at East Douglas, Mass Branch River at Forestdale, R. I. Pawturot River at Cransfon, R. I. I Willimattle River near South Coventry. Com Shetucket River near Willimantic, Com | Hop River near Columbia, Com Natchaug River at Willimantic, Conn Quinebaug River at Wurstville, Mass Quinebaug River at Quinebaug, Com Quinebaug River at Putnam, Conn | Quinebaug River at Jewett Clty, Conn Little River at Butimryille, Mass. Fire-Mile River at Killingty, Conn Moosup, River at Moosup, Conn Yantic River at Yantic, Conn | Ammonoosue River at Bethlehem Junction, N. H. Ammonoosue River near Bath, N. H. White River near Bethal, Vt. Ayers Brook at Randolph, Vt. Mascoma River at West Canaan, N. H. |
| | 'n | No. | 1-213.5. 1-216. 1-218. 1-220.4. | 1-223 1-225 1-227 1-230 | 1-232B 1-237A 1-249A 1-264 | 1–272 1–275 1–279A 1–284 | 1-289 1-294A 1-296.5 | 1-829A |

Summary of drainage basin topographic characteristics-Continued

828238 35.52 41 63535 582282 6353 5823282 $\mathbf{69}$ 35 13 238853 1.42 020200 82 1.69 2 2.69 .18 .31 .31 .31 .31 882288 3.49 1.02 48888 2.13 នុ 19 00 II H e i e,i 4.00 4.76 1.88 1.88 882028 225522 8¹⁸ 888 88 12 2668332 41 97 36 88 142 8 22 . 41 20 40 i ni 85 82 84 82 85 37 61 05 05 . 56 1. 36 9.88 04 .45 2.01 88288 33.08 **4%**82 35 1.0134 95 5 min h ന്ത്

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460 400 730 740 740 20 850 490 865 865 530 520 520 375 120 120 120 120 120 120 120 130 38 120 888 120 220 775 775 3 ŝ 285 514 246 318 318 1,6551,5171,2041,435 $1,110 \\ 1,070 \\ 988 \\ 1,038 \\ 1,108$ 918 1, 712 1, 557 1, 438 988 870 990 277 060 873 966 1, 466 1, 195 1, 192 418 20326220 132 420 _ Ľ, 900 303 900 900 900 900 900 295 859 544 166 240 890 620 839 295 166 764 560 902380 905238000 1.060 120 12841 ຕ໌ ຈີ ດ໌ ຕ໌ ດ໌ ຕໍ່ດໍ່ຕໍ່ດໍ ŝ -4 ŝ ÷. ຕໍ່ດໍ່ດໍ નંભે . ന്ന് ດໂດໂ പ്പ്പ് ¢1 c4 പ്പിന് -i 41-10-19.8 32.6 37.9 9.8 -2001 11000 21 404 10 6 41-101-1-4 10 00 00 00 6 1 3 888888 19.02 ণ 8.1.14 <u>4</u>5. 엻 82.27.28 825423 853 17. 1.282 ર્જ્સ 4 000-00 cro co co co 05070 0 8 1-10404 LQ 00 T N 1-01-400 00 CO 00 NO 03.0 IJ 20100101 23. 32.33 1.02 kg 10. 11.832.14 128823 51.30 19. ສ່ 222 প্ল 17.9 16.4 31.1 24.2 00440 0.014 21 37.2 61.3 32.3 10 00 m 00 01 0 122.3 1 00 6 9 101 ŝ ន់ខ្លន់នំនំនំ 4486 59. **** 31. ଞ୍ଚଛ୍ଷ 2 52 1835 8 8 3 10000 0 41 175 1175 193 246 50° 58.53 138 51 8275 205 245 245 218 292 89. 828624 115 157 241 227 238 238 276 8 88 811020 8 614 615 868 932 932 409 614 778 803 721 642 610 813 313 806 580 614 660 522 322 310 328 328 314 322 357 468 173 186 186 670 677 655 695 446 381 482 626 839 317 508 702 281 318 347 301 151 236862 144 282 265 278 278 278 278 33 349 527 246 246 158 185 187 222 202 520 544 530 544 530 850 88 756 215 969 969 376 370 392 376 462 619 628 628 232 232 379 537 766 884 821 899 726 404 722 678 698 476 580 539 103 88 914 022 264 262 288 258 268 912 562 211 733 666 812 751 952 274 418 890 890 919 919 204 664 236 025 74 109 344 46 838 389 017 677 998 413 216 544 352 758 758 758 749 051 പ്പിന്ന് ŝ സ്ത 6 00 1 ž **శ**ే ర్టీ సో సో ର୍ଣ୍ -Ц, 57 57 59 59 69 53 217 32 2.17 1. 72 1. 97 1. 82 1. 73 280 1.72 1.94 1.85 1.75 1.92 1.92 55052 3228 27 95 38 28 38 64 ----_____ ~----Ä ~ ci 1 308 71. 2 420 41. 8 54 50400 10 3 80 01 00 40 1 36. 18.20 88.83 80.12°. 222153 332 82 ස්සුසුසුසුව 34,158,702 197 148 148 5 83 8,8,8,7,5 74.

land, Vt. Millers River near Winchendon, Mass...... Millers River at South Royalston, Mass..... Millers River at Bruth, Mass..... Sip Pond Brook near Winchendon, Mass..... Priest Brook near Winchendon, Mass..... East Branch of Tully River near Athol, Mass. Moos Brook at Wendell Depot, Mass...... Desrfield River at Charlemout, Mass...... Desrfield River, acluding Somerset Reservoit, at Charlemout, Mass...... Desrfield River, excluding Harriman Reser-voir, at Charlemout, Mass..... Ware River at Gibbs Crossing, Mass...... Chicopee River at Bircham Bend, Mass..... Swift River at West, Mass...... Quaboag River at West Brimfeld, Mass...... Mill River at Springfield, Mass...... Westfield River at Knightville, Mass....... Westfield River near Westfield, Mass....... Westfield River, recluding Westfield Little River, near Westfield, Mass. Middle Branch of Westfield River at Goss West Briznch of Westfield River at Hunting-West Briznch of Westfield River at Hunting-Westfield Little River near Westfield, Mass.... Scantic River at Broad Brook, Com...... Farmington River near New Boston, Mass.... Farmington River, excluding Otis Reservoir, near New Boston, Mass. -----------********* -----ear Keene, N. H. Ashuelot River near Marlboro, Conn Ħ West River at Newfane, Yt. Asthuelot River near Olisum, N. H. Asthuelot River at Hinsdale, N. H. Otter Brook near Keene, N. H. Soth Branch Asthuelot River near M N. H. ₽ River near East Hartnord, zĒ Mascoma River at Mascoma, N. H. Ottauqueches River at North Earth Barth Spaga River at West Charemont, N. Black River at North Springfield, V Saxtons River near Saxtons River, V Mass. Hockanum ton, 1-363..... 1-365 1-367 1-368 1-376A 1-380A 1-380.4A 1-383.2 1-384.2 1–387 1–389 1–399 1–394 1–396 ------375.-----------------------..... -----1-340... 1-346... 1-347... 1-350... 1-399, 143. 1-406-428 928 928 928 928 928 928 928 928 -372 Ę 1-407 1-431. -371

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CONTRIBUTIONS TO HYDROLOGY, 1944

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| | A pick | io 19duuM areas inter areas inter babivided | 1288 22 28 22 | 69 | 50 50 50 50 50 50 50 50 50 50 50 50 50 5 | 52 74 53 33 49 | 234 | 45 63 105 | 53 | 49 | .49 39 | 27 51 | 40 44 |
|---|---|--|---|---|--|--|-------------------------------|-----------------------------------|--------|-------------------------------|----------------|-----------------|---|
| | water (| [ßtoT | 0.91 5.10 | 10.96 | 26.29 3.77 1.95 2.88 2.88 | . 19 1. 81 . 10 1. 24 | | | | 1.47 | .74 | 0 | -08 45 |
| | rea of surfaces (so miles) | гдтвwЗ | 0 .03 .36 | 1.72 | 9.42 2.26 0.62 .48 | 0.12 | 1 | | | 1.27 | 0 | 0 | 0.16 |
| | Area surf mil | Lakes and reser- voirs | 0.91 1.00 4.74 | 9.24 | $16.87 \\ 1.51 \\ 1.61 \\ 2.26 \\ .37 \\ .37$ | $\begin{array}{c} 10\\ 1.69\\ .10\\ .14\\ 5.85\end{array}$ | 40.97 | $3.84 \\ 18.55 \\ 13.84 \\ 13.84$ | 63 | .20 | 10 | 0 1.27 | 08 |
| | land oove level) | muminiM | 1,030 1,030 720 | 555 | 320 320 310 310 310 | $\substack{ \begin{array}{c} 410 \\ 160 \\ 420 \\ 30 \\ 1, 550 \end{array} }$ | 560 | 1, 544 700 905 | 1, 248 | 620 | 374 790 | 370 | 870 530 |
| | liftitude of land (feet above mean sea level) | пвэМ | $ \begin{array}{c} 500 \\ 500 \\ 1, 678 \\ 1, 432 \end{array} $ | 1,264 | 981 849 1, 029 685 | $\begin{smallmatrix} 1,016\\801\\870\\870\\2,193\end{smallmatrix}$ | 1,848 | 2, 231 1, 450 1, 888 | 1, 896 | 1, 700 | 1 580 | 1,650 1,269 | 1, 835 1, 649 |
| | Altitude (feet mean s | mumixsM | 1, 000 2, 300 2, 660 | 2,660 | $\begin{array}{c} 2,660\\ 1,737\\ 1,737\\ 1,070\\ 1,680\\ 1,140 \end{array}$ | $ \begin{array}{c} 1, 700\\ 1, 700\\ 1, 200\\ 1, 060\\ 5, 344 \end{array} $ | 5, 344 | 3, 865 4, 842 3, 595 | 3, 595 | 3, 816 | 2,020 3,200 | 3, 300 | 3, 143 3, 764 |
| nann | -nirq ខ៣.៩១។ | te dtans.I te lagio (miles) | 21.4 22.3 14.5 36.7 | 55.6 | 99.3 31.6 17.6 33.1 19.5 | 21.1 31.4 11 37.1 | 111.7 | 36.3 24.5 53.9 | 23.1 | 23.3 | 21.8 14 2 | 11. 5 53. 3 | 12.2 |
| | s) stercourse | aw teegno.I elim) | 17.5 23.2 12.2 43.0 | 68.7 | $\begin{array}{c} 111.2\\ 33.0\\ 19.4\\ 34.7\\ 34.7\\ 17.1\end{array}$ | 21.8 40.4 10.1 19.7 24.4 | 93.7 | 35.0 47.3 35.9 | 22.3 | 22.2 | 19.9 11 1 | 46.0 | 9.5 |
| | Channel slope (feet per mile) | Principal | $\begin{array}{c} 40.2\\ 9.42\\ 78.6\\ 13.1\end{array}$ | 8.45 | 10 17.1 17.1 31.4 33.4 | 41 17.5 72.3 32.2 13.0 | 12.37 | 32.61 16.54 35.7 | 42.5 | 52.7 | 24.36 77 | 113 17.7 | 146 80.2 |
| rerist | Cha Slope per 1 | Tributary | $\begin{array}{c} 74.6\\91.2\\165\\135\end{array}$ | 79.8 | ${ 54.2 \\ 113 \\ 85.2 \\ 88 \\ 88 \\ 153 \\ 15$ | $175 \\ 103 \\ 147 \\ 92.6 \\ 250 \\ 250 \\ 175 \\ 103$ | 127 | 159 118 124 | 183 | 314 | 163 397 | 414 181 | 479 271 |
| nunu | (feet | · 93,819V Å | 349 351 521 553 | 543 | 557 587 487 465 474 | 546 483 424 424 1,091 | 920 | $^{800}_{1,040}$ | 1, 157 | 953 | 432 673 | 699 900 | 841 |
| 0 0110 | Land slope per mile) | .8N | 286 274 486 513 | 487 | 477 480 388 388 366 | $ \begin{array}{r} 475 \\ 399 \\ 334 \\ 368 \\ 368 \\ 1,010 \\ \end{array} $ | 882 | 883 998 961 | 1, 074 | 827 | 390 544 | 533 | 719 |
| India | Land | .WЭ | 433 452 567 607 | 621 | 666 736 593 574 624 | 644 598 552 614 1, 148 | 973 | $^{684}_{1,\ 022}$ | 1, 275 | 1, 120 | 494 842 | 844 1,047 | 1,005 |
| arainaye oasin woographic characteristics | (səlim | n diduo) læ Z | 821 1, 323 405 6, 841 | 21,092 | 81, 138 3, 497 820 2, 431 710 | 822 4, 483 121 710 2, 282 | 89, 931 | 2,503 13,054 9,640 | 1, 258 | 1, 552 | 894 268 | $142 \\ 11,036$ | 210 1, 057 |
| nde oc | Ttians Sugusa | Stream de (miles per mile) | 1. 62 1. 62 1. 66 1. 91 1. 52 | 1.35 | 1.57 1.31 1.31 1.98 2.29 | 2,21 2,15 2,15 2,15 2,15 2,15 2,15 2,15 | 1.68 | 1.77 1.58 1.60 | L. 37 | 1.51 | $1.68 \\ 1.77$ | $1.79 \\ 1.68$ | 1.92 1.74 |
| | | əzanisı (] m ə1211p2) | 104.7 109 57.1 280 | 632 | $\begin{array}{c} 1,545\\ 204\\ 68.5\\ 133\\ 75.3\end{array}$ | 71.9 245.8 24.0 77.5 192 | 1,664 | 160 527 491 | 114 | 152 | 90 46.2 | 31.4 510 | 39.1 111 |
| Summary of | No. Name of gaging station | | Salmon River near East Hampton, Conn Quinntpiae River at Wallingford, Com Housatonic River at Coltsville, Mass Housatonic River near Great Barrington, Moss | Housatonic River at Falls Village, Conn | Housatonic River at Stevenson, Conn Tennile River near Gaylordsville, Conn Still River near Lanesville, Conn Biepaug River near Roybury, Conn Pomperaug River at Southbury, Conn | Naugatuck River near Thomaston, Conn Naugatuck River near Naugatuck, Conn Leadmine Brook near Thomaston, Conn Baugatuck River near Westport, Con Hudson River near Newcomb, N. Y. | Hudson River at Hadley, N. Y. | Cettar Livel Delow Lake, N. Y | | Batten Kill at Arlington, Vt. | ams, Mass | agle Bridge, N. | North Branch of Hoosic River at North Adams, Mass Walloomsac River near North Bennington, Vt. |
| | | | 1-432 1-435 1-435 1-444 | 1-446 | 1-450 1-453 1-455 1-459 1-459 | 1-462 1-463 1-464 1-466 | 1-475 | 1-485 1-486 1-488 | | 1-496 | 1-501 | 1-506. | 1-508 |

Summary of drainage basin topographic characteristics—Continued

| | | - | | | | | | | | | ~ - |
|--|---|---------------------------|----------------|--|--|--|--------|--|---|-----------------------------------|--|
| 78 66 61 | 69 69 69 69 | 49 118 49 | 49 57 | 66 54 56 | 45 | 148 33 60 7 | 55 | 53 55 56 20 2 3 3 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 | 56 55 55 | 53 19 65 | 338 |
| | | | | 3.00 | | 0 | | . 03 | 1.82 | 9.25 | |
| | | | | 2.55 | | 0 | | 0 | 1.81 | 9.24 | |
| 12.07 4.35 1.52 | .10 .65 1.11 3.09 1.06 | $^{13.90}_{3.29}$ | . 27 | 2.78 .45 .20 | .84 | $^{0.38}_{-28}$ | 0.16 | 12 23 28 | 01 10 18 07 | $ \frac{01}{56} $ | 33.08 |
| 320 770 40 | 648 55 390 196 140 | 208 360 360 | 265 258 | 2812 2812 2822 2822 2822 2822 2822 2822 | 130 | 29 113 275 192 | 52 | 294 30 72 10 | 13 10 10 10 13 13 13 13 13 13 13 13 13 13 13 13 13 | 55 48 955 | 420 |
| $1, 131 \\ 1, 821 \\ 897$ | 1, 347 930 695 624 493 | 319 456 822 | 534 687 | 654 275 778 778 | 562 | 370 229 480 587 | 378 | 732 154 137 137 265 113 | 112 84 93 93 | $125 \\ 108 \\ 79 \\ 79 \\ 1,849$ | 1, 555 |
| 3. 626 2, 780 2, 653 | 2,885 3,863 1,809 1,440 1,440 | 857 1, 496 1, 406 | 1,092 1,420 | $1, \frac{420}{710}\\643\\1, 227$ | 1, 220 | 1, 220 680 1, 122 | 1, 160 | ${\begin{array}{c} 1,205\\ 5640\\ 563\\ 561\\ 391\end{array}}$ | 308 346 212 205 205 | 199 177 131 3, 905 | 3, 905 |
| $ \begin{array}{c} 112.3 \\ 38.3 \\ 57.1 \\ \end{array} $ | 17 46.3 26.7 34.2 | 17.7 92.9 27.2 | 12.1 25.1 | 30.4 22.4 23 | 32.7 | 60.3 10.3 13.4 | 37.7 | 14.9 36.2 36.4 18.9 | 11.1 28:9 20.8 15.8 | 19 22.7 7.4 75.1 | 173.4 |
| 90.8 28.1 44.4 | 15. 9 46.7 23.9 30.6 30.6 | 11.9 59.9 31.0 | 11.3 23.9 | 35.4 18.0 18.4 23.7 | 35.1 | $ \begin{array}{c} 58.2 \\ 7.8 \\ 11.6 \\ 11.6 \end{array} $ | 33.3 | 17.6 26.6 35.6 7.47 5 9.8 | 12.3 21.8 14.7 17.1 17.6 | 16.7 16.9 7.3 62.3 | 157.9 |
| $ \begin{array}{c} 11.04 \\ 51.8 \\ 26.8 \\ 26.8 \end{array} $ | 56.4 26.8 11.2 16.1 | 9.8 7.1 14.9 | 23.9 12.3 | 15.0 19.2 31.3 | 19.6 | 10.1 24.1 112 73.7 | 23.80 | 28.46 5.80 4.15 47.3 13.6 | 8.02 6.99 6.77 6.77 | 6.05 5.46 7.2 13.3 | 7.66 |
| 73 95.7 105 | 309 144 102 63.7 63.7 | 67. 1 49. 6 88. 9 | 181 21.8 | 112 66.9 114 151 | 108 | 71.6 92.3 300 191 | 75.7 | 125 23.3 29.3 145 46.4 | 27.1 13.6 18.2 10 15.6 | 12.8 10.1 11.7 126 | 90.1 |
| 474 543 679 | 770 702 554 449 471 | 294 533 602 | 574 743 | 661 214 309 501 | 451 | 378 259 473 | 437 | 443 140 158 373 244 | 121 158 71 55 70 | 53 57 60 1, 092 | 775 |
| 497 568 612 | 820 660 412 358 | 315 473 578 | 562 648 | 596 154 305 497 | 453 | 370 262 529 | 427 | 450 154 379 250 | 132 151 72 69 | 861 53 982 61 54 | 110 |
| 441 526 765 | 709 748 636 663 663 | 266 609 632 | 580 860 | 825 302 506 | 449 | 389 255 414 | 453 | 434 131 152 365 365 240 | 114 166 70 54 71 | 53 61 59 1, 241 | 860 |
| 65, 630 3, 475 8, 756 | $\substack{9,\ 603\\1,\ 769\\26,\ 284\\3,\ 335}$ | 365 25, 065 1, 648 | 1, 307 | 2, 927 600 437 886 | 2,408 | 12, 849 109 1.84 163 | 2, 911 | $\begin{array}{c} 288\\ 2,128\\ 4,288\\ 40\\ 275\end{array}$ | 318 1, 260 474 734 693 | 1, 036 89 26, 360 | 244, 482 |
| 1.95 1.75 1.26 | $\begin{array}{c} 1.47 \\ 1.38 \\ 1.97 \\ 1.60 \\ 1.35 \\ 1.35 \end{array}$ | 1. 89 1. 73 1. 90 | 1. 70 1. 76 | 1.67 2.25 1.55 | 1.61 | 1.85 1.85 2.37 | 2.00 | 1.98 1.58 1.72 2.41 | 1.98 1.39 1.39 1.00 | $^{-96}_{1.16}$ | 1.37 |
| 1,348 261 329 | 98 386 144 711 182 | 55.4 785 116 | 29.4 118 | 160 54. 6 . 65. 3 65. 3 | 147 | 490 25.7 26.2 26.2 | 190 | 32. 8 171 258 9. 75 48. 5 | 43.4 124 56 70.5 64 | 56.3 113 22.3 783 | 3, 076 |
| Mohawk River near Little Falls, N. Y. East Canada Oreek at Dolgeville, N. Y. | Catskill Creek at Oak Hill, N.Y. Rondout Creek at Rosendale, N. Y. Walikill River near Unionville, N. Y. Walikill River at Gardiner, N. Y. Wappinger Creek near Wappinger Falls, N.Y. | | | | South Branch of Raritan River at Stanton, N.J. | | | Block Tiver near Pottersville, N. J. Block River near Pottersville, N. J. Millstone River and Blackwells Mills, N. J. Creen Brook at Plainfield, N. J. Swimming River near Red Bank, N. J. | Manasquan River at Squankum, N. J Manasquan River, N. J Codar Creek at Lanoka Harbor, N. J Batsto River at Batsto, N. J Dats Batsto River at Batsto, N. J | QSSH | Eddy, N. Y. Delaware River at Port Jervis, N. Y |
| 1-515 1-536 1-549 | 1-550 1-562 1-562 1-563 1-566 1-569 | 1-577 1-582 1-583.5 | 1-587 1-593 | 1595 1-606 1-609 1-611 | 1-614 | 1-616 1-619 1-619.5 | 1-622 | 1-623 1-624 1-625 1-625 1-632 1-632 | 1-633 1-634 1-635 1-636 1-637 | 1-639 1-640 1-641 1-644 | 1-646 |

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CONTRIBUTIONS TO HYDROLOGY, 1944

| | -dus s doidw -dus i | Number or sreasinto av nisad babivided | 52 49 | 114 | 94 | 21 | 52 52 52 53 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 54 50 50 50 50 50 50 50 50 50 50 50 50 50 | 26468 ¹ 4 | 51 51 47 | 119 48 | 382, 13 382, 13 | 51 |
|---|-------------------------------------|---|--|---|---|--|--|--|--|-------------------|--|---|
| | water (square | [stoT | | | | | | | | | | |
| | of aces es) | 20msw2 | | | | | | | | | | |
| | Area surf mil | Lakes and reser- | 36.30 .15 | . 65 | 16.21 | 3. 55 | 2. 58 2. 58 2. 58 2. 58 | 4.75 4.75 5.08 | .24 30 | .75 | 0.00 041 | 117 |
| | of land a b o v e sa level) | muminiM | 1, 335 | 955 | 874 | 874 | 645 348 410 335 560 | 404 296 1,060 210 260 | 3658 3658 | 120 | 12883 | 150 30 85 |
| | | пвэМ | 1, 911 | 1, 738 | 1, 429 | 1, 358 | 1, 644 819 883 666 666 706 | $^{689}_{1, 724}_{1, 724}_{1, 166}_{1, 166}$ | 97 275 97 | $^{873}_{1, 293}$ | 438 113 305 321 86 | 447 90 340 |
| _ | Altitude (feet mean s | mumixsM | 3, 253 | 3, 365 | 2, 654 | 2, 654 | $\begin{array}{c} 4,204\\ 1,653\\ 2,131\\ 1,600\\ 1,127\end{array}$ | $1, 248\\1, 141\\2, 320\\2, 320\\960$ | 340 660 213 | 2, 020 2, 020 | 1, 200 172 595 600 166 | 1, 100 164 640 |
| | -nirq 2m.s91 | Length of ts lagio (miles) | 201. 7 18. 3 | 60 | 54.3 | 41.1 | 46.2 23.6 27.3 10.7 | 27.6 13.7 48.3 21.4 | 20.8 36.8 26.8 | 95 16.9 | 49.3 15.8 13.4 7.5 | $^{41}_{7.2}$ |
| | s) tercourse | kw jz9gno. Jelim) | 209. 2 25. 8 | 73.3 | 34.6 | 34.6 | 8,738,87 8,738,87 2,57,88,7 2,57,89,7 2,57,89,7 2,57,99,7 2,57,70,70,70,70,70,70,70,70,70,70,70,70,70 | 28.4 13.5 99.3 23.9 | 20.5 37.5 18.4 | 79.8 | 27.8 2.7 15.7 8.2 | 38.7 7.2 20.8 |
| 87178 | nnel (feet nile) | Principal | 16.40 | 85.8 | 18.7 | 19.1 | 30.7 31.9 40.67 11.9 10.75 | 7.07 15.55 22.9 16.5 12.9 | 4.75 8.0 6.5 | 6.77 53.7 | $\begin{array}{c} 14.08\\ 20.4\\ 23.1\\ 24.6\\ 13.73\\ 13.73\end{array}$ | 12.8 11.1 20.1 |
| crer ro | Channel slope (feet per mile) | Tributary | 154 | 130 | 75.4 | 83 | 129 174 154 82 82 78. 1 | 80.3 59.4 55.3 40.5 | 22 43.7 12.7 | 70. 7 202 | 65.2 54.2 104 106 47 | 51.5 39.3 96.3 |
| crura | (feet | A Verage | 216 | 905 | 469 | 498 | 633 617 571 469 515 | 497 526 283 233 233 | 109 225 62 | 605 621 | 360 98 412 148 | 434 144 389 |
| three ci | Land slope (feet per mile) | .8N | 829 | 857 | 426 | 422 | 705 593 630 467 476 | 476 535 277 222 | 118 237 66 | 654 780 | 364 98 463 433 155 | 459 154 373 |
| pogra | Land | вW. | 1,005 | 964 | 527 | 601 | 848 641 502 472 534 | 509 515 288 288 244 244 | 99 212 57 | 540 443 | 357 96 378 378 138 | 402 130 414 |
| on ursu | (səlim əlduə) la Z | | 610, 833 1, 892 | 20, 781 | 9, 260 | 4, 857 | 5, 520 827 2, 093 138 | $^{1,611}_{7,183}$ | 810 4, 077 1, 197 | 49, 456 | 4, 277 12 335 267 76 | 5, 283 63 604 |
| nage t | vtiene eneupe | Stream de (miles per mile) | 1.43 | 1.38 | 1.27 | 1.38 | 1.41 1.58 1.59 1.99 | 1.75 1.75 1.75 | 1.60 1.18 | 2.07 1.36 | 1.68 2.144 1.23 19 19 19 19 19 19 19 19 19 19 19 19 19 | 2.26 1.75 3.03 |
| oj urun | вэтв (zəli) | əzaniar (I mənaupz) | 6, 344 142 | 593 | 519 | 290 | 222 65.1 64.4 126 31.4 | $108 \\ 36.2 \\ 322 \\ 1,280 \\ 97.4$ | 89.4 210 111 | 1, 147 42.9 | 279 6.75 33.3 31.9 31.9 19.3 | 287 14.6 52.6 |
| s unintary of aramage oaste topographic characteristics | • | Name of gaging station | Delaware River at Riegelsville, N. J. West Branch of Delaware River at Delhi, | West Branch of Delaware River at Hale Eddy, | Lackawaxen River at Hawley, Pa., (including | wallenpaupack reservor/. Lackawazen River at Hawley, Pa., (excluding Wallenpaupack Reservoir). | Neversink River at Oakland Valley, N. Y Flat Brook near Flatbrookville, N. J. McMichaels Creek at Stroudsburg, Fa Paulins Kill at Blatratown, N. J. | Pequest River at Pequest, N. J. Beaver Brook near Belvidere, N. J. Beaver Brook near Belvidere, N. J. Lehigh River at Tannery Pa. Lehigh River at Bethlehem, Pa. Tohickon Creek near Pipersville, Pa. | Assunpink Creek at Trenton, N. J. Neshaminy Creek near Langhorne, Pa North Branch of Rancocas Creek at Pember- | | Perktomen Creek at Graters Ford, Pa. Mantua Creek near Pitman, N. J. Crum Creek at Woodyn, Pa. Ridley Creek at Moylan, Pa. Oldmans Creek near Woodstown, N. J. | Brandywine Creek at Chadds Ford, Pa Salem Creek at Woodstown, N. J Big Elk Creek at Elk Mills, Md |
| • | | No. | | 1-658 | 1-664.5 | | 1-668 1-671 1-672 1-673 | 1-674 1-675 1-675 1-677 1-677 | 1-683 1-685 1-686 | 1-688 | 1-693_1-693_1-694_A_1-694_A_1-694_A_1-696_1-696_1-696_1-698_ | 1-701 1-702A 1-706 |

Summary of drainage basin topographic characteristics-Continued

| 97 198 | 67 68 68 61 61 | 210 210 41 | 451468 | 64 | 51 52 | 57 57 | 75 53 57 57 | 78 | 55 67 165 | 15 65 51 | 30 | 57 |
|--|--|---|--|---------------------------|--|---|---|---------------------------|--|--|-------------------------------------|---------------------|
| | | | | | | | | | | 0 | | |
| | | | | | | | | | | 0 | | |
| 10.63 14.23 | 23.01 3.11 2.57 2.61 2.61 2.61 | 0.2.04 | 0 2.52 .59 .37 .08 | .05 | 0.13 | 0.15 | 0.00 00.00 | .58 | 000 10. | 000 | 0 | .85 |
| 1, 152 842 | $1,175 \\ 999 \\ 880 \\ 880 \\ 920$ | 825 970 934 1, 190 | $1,060 \\ 1,010 \\ 775 \\ 560 \\ 1,225$ | 582 | 420 820 | 754 620 | 792 552 354 260 | 970 | $1,005 \\ 651 \\ 1,022 \\ 239 \\ 239$ | 970 270 155 | 270 | |
| 1,489 1,431 | 1,508 1,430 1,315 1,430 | $1, 280 \\ 1, 573 \\ 1, 579 \\ 1, 770 \\ 1$ | $\begin{array}{c} 1, 597\\ 1, 520\\ 1, 307\\ 1, 189\\ 1, 589\end{array}$ | 1, 241 | 894 1,488 | 1,477 1,056 | 1, 487 985 959 831 831 507 | 2,455 | $1, 932\\1, 633\\1, 474\\623$ | 1, 366 496 349 | 384 | |
| 2, 301 2, 740 | 2, 100 2, 120 2, 160 | 2, 133 2, 540 2, 540 2, 240 | 50000000000000000000000000000000000000 | 2,440 | 1, 813 3, 136 | 2, 1 00 2, 400 | 1, 320 1, 320 1, 320 | 4, 150 | 3,022 2,942 1,980 | 1, 880 880 600 | 560 | |
| 61. 1 124. 4 | 197 22. 71. 7 79. 4 | 40.4 69.5 1122 - 11.5 - | 25.2 433.7 50.1 50.1 | 62.3 | 42. 5 32. 4 | 29. 2 29. 2 | 85.3 37.6 63.3 41.3 | 56.6 | 13.4 37.8 57.9 | $3.2 \\ 17 \\ 21.6$ | % | 76.6 _ |
| 42. 5 128. 4 | 176.9 25.0 58.3 76.5 58.8 | 31.7 57.3 93.0 93.0 | $ \begin{array}{c} 19.2 \\ 30.6 \\ 39.3 \\ 39.3 \\ \end{array} $ | 43.4 | 36.8 31.9 | 28.1 | 78.6 80.8 35.7 | 48.7 | 17.8 33.7 111.5 63.4 | 3.5 18.6 21.7 | 6.9 | 83.2 |
| 3, 90 3, 32 | 15.40 5.16 6.99 | $\begin{array}{c} 17.45\\ 12.30\\ 7.49\\ 6.1\\ 6.1\end{array}$ | $ \begin{array}{c} 36\\ 10.30\\ 43.88\\ 22\\ 13.10\\ 13.10\\ \end{array} $ | 16.34 | 16.42 16.80 | 24. 2 19. 5 | 17.2 17.3 6.3 2 6.3 2 6.3 | 55.2 | 55.7 51.7 39.0 5.1 | 223 14.9 15.2 | 22.2 | - |
| 94.9 64.4 | 93.5 59.1 59.6 | 134 78.5 50.2 242 | 141 80.1 152 148 78.5 | III | 89.5 143 | 218 167 | 88.2 136 298 71.6 54.5 | 198 | 337 192 69. 2 | 356 78 73.4 | 80.2 | |
| 621 632 | 576 585 559 572 | 641 782 780 780 | 685 603 442 778 710 | 897 | 758 873 | 900 945 | $^{871}_{1,084}_{1,084}$ | 746 | $ \begin{array}{c} 850 \\ 1,053 \\ 1,072 \\ 438 \\ \end{array} $ | 715 524 410 | 315 | |
| 566 574 | 513 524 481 481 | 575 793 783 688 | 716 545 545 182 802 744 | 954 | 958 835 | 902 991 | $\begin{array}{c} 775\\ 1,032\\ 1,284\\ 486\\ 403\end{array}$ | 744 | 775 999 891 417 | 617 483 418 | 278 | |
| 671 710 | 626 661 653 693 | 740 768 776 | 645 683 394 755 664 | 829 | 494 921 | 897 902 | 997 811 872 872 337 | 748 | $ \begin{array}{c} 950 \\ 1,125 \\ 1,329 \\ 464 \end{array} $ | 808 567 399 | 361 | |
| 8, 829 163, 292 | $\begin{array}{c} 670, 664 \\ 1, 405 \\ 15, 249 \\ 57, 740 \\ 23, 148 \end{array}$ | $2, 961\\19, 033\\43, 398\\140, 000\\201$ | 1, 155 9, 429 4, 648 6, 191 | 12, 376 | 3, 283 5, 340 | 3,180 1,950 | 35, 640 3, 971 3, 971 19, 025 6, 368 | 7,717 | 4, 107 4, 107 165 26, 587 | 106 781 | 85 | 70, 663 |
| 1. 54 1. 44 | $ \begin{array}{c} 1.59\\ 1.53\\ 1.33\\ 1.33 \end{array} $ | 1.36 1.29 1.36 | 1.63 1.47 1.67 1.80 1.72 | 1.34 | 1.55 1.13 | $1.50 \\ 1.31$ | 1.23 1.66 1.61 1.61 | 1.36 | 1.72 1.49 1.67 2.02 | $ \begin{array}{c} 2.01 \\ 1.80 \\ 1.92 \end{array} $ | 1.56 | |
| 351 240 | 797 518 492 735 | 186 770 370 530 30.5 | 114 472 45.8 274 315 | 559 | 162 291 | 220 128 | 756 200 21.6 322 322 | 287 | 72. 4 247 30. 2 817 | 5.7 82.3 62.2 | 21.3 | 6 |
| പ് | <u> </u> | 5 7 7 7 7 | 1111 | | | | | | | | | 1, 690 |
| rsville, N. Y. | Towanda, Pa. N. Y kdale, N. Y Chenango Forks, N. Y Itaska, N. Y | Y Y N.Y | ear South Addison, N. Y. ar Campbell, N. Y | Creek at Beech Creek Sta- | East near Dalmatia, Pa of Juniata River at Wil- | Little Juniata River at Spruce Creek Pa Standing Stone Creek near Huntingdon, Pa | Raystown Branch of Juniata River at Sarton, Pa. Sherman Creek at Shermandalde, Pa Clark Creek near Carsonville, Pa Comodoquimet Creek near Hogestown, Pa Conestoga Creek at Lancaster, Pa | otomac River at Blooming- | franklin, Md Jumberland, Md Bedford Valley, Pa t Jug Bridge, near Frede | antz, Md ear Frederick, Md nerrill Drive, Washington | stia River near | 8e, N. Y. |
| at Colliersville, at Conklin, N. | ex, N. Y. ex, N. Y. Bockdale, N. J. aur Chenango I. at Itaska, N. | Owego, N. Y. Idley, N. Y. Erwins, N. Y. t Chemung, N. Y. Arkport, N. Y. | ar South Addis, r. Campbell, N. r. near Wapwall, Bloomsburg, P. Susquehanna | ek at I | ast nea I Junia | at Spruce k near Hu | of Juniata River at Shermandale, Pa. Carsonville, Pa ek near Hogestow t Lancaster, Pa | mac Ri | Franklin, Md Cumberland, Md Bedford Valley, st Jug Bridge, ne | utz, Md sar Frederick, errill Drive, | of Anacostia | Red House, N. |
| liver a | 8 J 9 L 2 | | E R S G | | reek E anch o | tiver a Creek | ch efJu at Sho ar Car Creek | | at Fran r Cun car Bec r at J | k near Sherr | | |
| unna H unna H | unna H lek at River o Rive oga R | reek n ver at ver nes Rive River | a Cree Rivel Open C Creek I Pa. | ud Ea | ND Bro | Stone | n Bran Creek sek ne umet (a Cree | anch c | Preek a ek nes eek ne V Rive d. | reek at e Cree sek at | st Bra lle. M | r Rive |
| Susquehanna River Susquehanna River | Susquehanna River Oaks Creek at Inde: Unadilla River at R Chenango River nea Tioughnioga River rea | Owego Creek near Tioga River at Lin Tioga River near H Chemung River at Canisteo River at | Tuscarora Creek n Cohocton River ne Wapwallopen Cree Fishing Creek near West Branch of Bower, Pa. | North Bald Eagle | Mahantango Creek | Little Juniata River | Raystown Branch e Sherman Creek at Clark Creek near (Conodoqumet Cre Comestoga Creek a | North Branch of P Md | Georges Creek at E Wills Creek near C Evitts Creek near C Monocacy River a rick, Md. | Owens Creek at Lantz, Md. Linganore Creek near Freder Rock Creek at Sherrill Dri D. C | Northwest Branch Colesville. Md. | Allegheny River at |
| - | | | 1111 | | 26 | 36 15 | | z | 0202 | | z | ▼ |
| 1-707 | 1-712 1-717 1-723 1-723 1-724 | 1-740 1-742 1-743 1-746 | 1-758- 1-761- 1-777- 1-778-5 1-781 | 1-790 | 1-798 | 1-802.5 1-808 | 1-810 1-818 1-818.5 1-820 1-820 | 1-851 | 1-863- 1-866- 1-867- 1-903- | 1-904- 1-905- 1-910 | 1-913 | 3-2 |

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151

.

CONTRIBUTIONS TO HYDROLOGY, 1944

| | doid w | Vurnber of sreas into w nizad babivided | 123 175 52 49 | $^{10}_{23}$ | 280 59 59 59 | $122 \\ 35 \\ 35 \\ 63 \\ 63 \\ 50 \\ 36 \\ 35 \\ 63 \\ 63 \\ 63 \\ 63 \\ 63 \\ 63$ | 49 106 275 47 | 1383555 13 | 65 59 |
|---|--|--|---|---|--|---|---|---|----------------------------------|
| | water (square | latoT | | | | 2.3 | | 4.11 | |
| | of aces as) | sqmsw2 | | | | 2.3 | | 2.18 | |
| | Area surf mile | Lakes and reser- | $\begin{array}{c} 25.53\\ 25.92\\ .16\\ .08\\ 1.25\end{array}$ | 2.71 .03 0.04 1.71 | $\begin{array}{c} 2.05 \\ 0.19 \\ 0 \\ 0 \\ 0 \end{array}$ | 0.05 - 0.03 - | $\begin{array}{c} 0.04 \\ 0.07 \\ 0.07 \end{array}$ | $\overset{1.93}{\overset{0.28}{}}$ | .03 |
| | land o v e evel) | muminiM | $ \begin{array}{c} 1,195\\ 1,220\\ 1,100 \end{array} $ | $\substack{1,\ 021\\1,\ 220\\1,\ 100\\1,\ 080\end{array}}$ | 776 950 972 972 | $^{880}_{3, 060}$ $^{865}_{1, 670}$ $^{1, 670}_{2, 175}$ | 975 828 752 860 950 | 805 940 972 870 | 952 850 |
| | Altitude of land (feet above mean sea level) | ns9M | $1,595 \\ 1,619 \\ 1,385$ | $1,\ 307\\1,\ 491\\1,\ 458\\1,\ 220\\1,\ 979$ | $^{1,\ 590}_{1,\ 575}$ | $\substack{1,\ 209\\3,\ 370\\2,\ 138\\2,\ 505}$ | $\substack{1,\ 110\\1,\ 038\\1,\ 051\\1,\ 078\\1,\ 138\end{cases}$ | $1, 254\\1, 055\\1, 132\\1, 120\\978$ | 1,072 1,101 |
| | Altitude (feet mean s | mumixeM | $\begin{array}{c} 1, 982\\ 2, 140\\ 1, 880 \end{array}$ | $\begin{array}{c} 1,880\\ 2,240\\ 2,240\\ 1,700\\ 2,949 \end{array}$ | $\substack{\textbf{2, 949}\\\textbf{2, 480}\\\textbf{1, 900}\\\textbf{1, 900}\\\textbf{1, 900}\\\textbf{2, 920}\\\textbf{1, 900}\\\textbf{2, 920}\\\textbf{2, 920}\\\textbf{2, 920}\\\textbf{2, 920}\\\textbf{3, 920}\\$ | $\begin{array}{c} 1,960\\ 4,375\\ 3,340\\ 3,213\\ 3,213\\ 3,027\\ 3,027\\ \end{array}$ | $\substack{1,\ 380\\1,\ 540\\1,\ 470\\1,\ 480\end{array}$ | $1,\ 600\\1,\ 300\\1,\ 400\\1,\ 140\\1,$ | 1,320 $1,480$ |
| | -ning 201.691 | tength of talladio (miles) | $\begin{array}{c} 184.7 \\ 259.2 \\ 45.4 \\ 35 \\ 36.8 \\ 66.8 \end{array}$ | $ \begin{array}{c} 68.3\\ 56.9\\ 33.2\\ 72.2\\ 72 \end{array} $ | 125.5 62.7 35.6 115.5 52.8 | $\begin{array}{c} 75.5\\ 27.5\\ 27.4\\ 107.5\\ 9.2\\ 9.2\end{array}$ | $\begin{array}{c} 30.8\\ 91.2\\ 69.8\\ 69.8\\ 23.4\end{array}$ | 232.2 23.8 2.5 2.5 2.5 2.5 2.5 2.5 2.5 2.5 2.5 2.5 | 21.7 47 |
| | s) tercourse | law tesgno.I slim) | $\begin{array}{c} 178.4\\ 214.7\\ 36.9\\ 58.9\\ 68.9\end{array}$ | 95.9 22.1 27.2 50.4 | 106.3 39.0 36.2 36.2 66.5 | 86.7 19.8 50.3 8.5 | 30.4 82.2 58.8 20.9 20.9 | 25.628 25.628 1.5628 25.628 | 20.9 50.8 |
| 60.0 | nnel (feet aile) | Principal | 11.2 16.8 6.6 | $ \begin{array}{c} 3.6 \\ 7.7 \\ 8.2 \\ 8.2 \\ 8.2 \\ 29.1 \\ \end{array} $ | $229 \\ 229 $ | $\begin{array}{c} 3.3\\ 13.3\\ 20.3\\ 73.59\\ 73.59\end{array}$ | $ \begin{array}{c} 6.3 \\ 3.1 \\ 3.4 \\ 2.87 \\ 10.5 \\ \end{array} $ | ${ \begin{smallmatrix} 11.52 \\ 6.21 \\ 6.64 \\ 8.66 \\ 54.4 \\ 54.4 \\ \end{smallmatrix} }$ | 7. 1 7. 45 |
| Cre1 191 | Channel slope (feet per mile) | Tributary | 59.8 101 33.6 | 34.6 97.2 87.3 67.8 111 | 97.6 64.8 155 84.7 84.7 | 43. 7 114 88. 5 84. 1 151 | 28.2 38.2 89.9 89.9 | 36. 4 39. 2 35. 3 334 2 | 54 34. 3 |
| n mu | (feet | Å verage | 467 677 339 | 323 684 764 845 845 | $680 \\ 612 \\ 612 \\ 724 \\ 1, 202 \\ 1, 265 \\ 1, 265 \\ 1$ | $1,\ 277\\520\\705\\553\\583$ | 211 158 197 197 290 | 443 214 233 233 922 | 191 345 |
| hun cu | Land slope per mile) | .8N | 442 721 311 | 308 702 777 766 575 | $663\\608\\1,179\\1,270\\1,270$ | ${ \begin{smallmatrix} 1, & 290 \\ 500 \\ 695 \\ 527 \\ 495 \\ 495 \\ \end{array} }$ | 199- 151 208 306 | 413 529 233 794 | 200 379 |
| nıßoo | Land | Е₩. | 500 633 378 | 344 665 922 622 622 | $\begin{array}{c} 704\\ 619\\ 619\\ 724\\ 1,231\\ 1,258\end{array}$ | 1, 260 541 719 586 649 | 224 166 216 190 269 | 480 225 615 233 1,098 | $180 \\ 322$ |
| usere to | (səlim olduo) la Z | | $\begin{array}{c} 495,220\\ 800,000\\ 5,626\\ 3,543\\ 21,800\end{array}$ | 48, 700 1, 700 20, 880 20, 880 | $^{111,\ 100}_{\begin{array}{c} 8,\ 230\\ 5,\ 695\\ 71,\ 600\\ 12,\ 930\end{array}}$ | 29, 520 795 85, 400 8, 373 8, 373 | $\begin{array}{c} 3,882\\ 36,418\\ 123,609\\ 18,071\\ 18,071\\ 1,026\end{array}$ | $\begin{array}{c} 10,808\\ 2,297\\ 3,608\\ 2,1115\\ 1.71\\ 1.71\end{array}$ | 1,293 11,232 |
| n afin | sdasre strafty | Stream de (miles per mile) | $ \begin{array}{c} 1.38 \\ 1.61 \\ 1.46 \end{array} $ | 1,43 1,78 1,78 1,90 1,65 1,65 | 1.75 1.75 1.75 1.64 2.07 2.07 | 2.23 2.1112 2.07211 2.07211 | 1.52 1.46 1.48 1.56 1.55 | $\begin{array}{c} 1.60\\ 1.56\\ 2.13\\ 2.03\\ 3.37\\ 3.37\end{array}$ | $1.77 \\ 1.83$ |
| n ururn | вэтв iles) | Drainage m 918102) | 5, 982 7, 671 304 233 629 | $\begin{array}{c} 1,028\\ 153\\ 321\\ 191\\ 715\end{array}$ | $1,825 \\ 390 \\ 265 \\ 1,340 \\ 384$ | 750 86. 2 1, 326 232 24. 5 | 247 899 2, 235 588 104 | 406 165 254 175 1.64 | 120 472 |
| partitude of a large out of a price is a series | | Name of gaging station | Allegheny River at Franklin, Pa. Allegheny River at Parkers Landing, Pa. Allegheny River at Parkers Landing, Pa. Berlessmann Oreek at Lynoh, Pa. Franch Creek at Lynoh, Pa. | French Creek at Utica, Pa. Mahoning Creek at Punxsutawney, Pa. Maboning Creek at Dayton, Fa. Crooked Creek at Idaho, Pa. | Kiskiminetas River at Vandergrift, Pa Blacklick Creek at Blacklick, Pa Loyalhanna Creek at New Alexandria, Pa Tygart River at Fetterman, W. Va West Fork River at Clarksburg, W. Va | West Fork River at Enterprise, W. Va Blackwater River at Davis, W. Va Progridgeny River at Connellsville, Pa Casselman River at Markleton, Pa Big Piney Run near Salisbury, Pa. | Mahoning River near Berlin Center, Ohio Mahoning River at Youngstown, Ohio Beaver Aiver at Wanpum, Pa Shenango River at Sharpsville, Pa Little Shenango River near Greenville, Pa | Slippery Rock Creek at Wurtemburg, Pa Standy Creek at Worreburg, Ohlo Standy Creek at Wayneburg, Ohlo Nimishilen Creek as North Industry, Ohlo Home Creek near New Philadelphia, Ohlo | Jerome Fork at Jeromeville, Ohio |
| | | No. | | 3-26 3-33.5 3-34 3-34 3-34.5 3-42.5 | 3-47 3-55 3-59 3-63 3-70 | 3-71 3-81 8-95 3-98 3-100 | 3-111 | 3-135 3-141 3-148.5 3-148.5 3-150 3-151.2 | 3-158 3-160 |

Summary of drainage basin topographic characteristics-Continued

| 6 1 84 82 82 | 66 112 112 104 | 69 61 74 258 | $^{47}_{58}$ | 55 53 95 180 | 61 215 37 40 | 87 108 83 83 83 | - 164 65 92 134 134 | 87 81 81 | 8 |
|--|---|--|---|--|--|--|---|--|-----------------------------|
| | | | | | | 0 | | | |
| | | | | | | 0 | | | |
| 10.00 | 4.42 0 0 0.07 | 0.10 0.04 0 | 00000 | 00.00 | 0 ⁰⁰ | 00.00 | 28822 | 0.00 | 8.96 |
| 796 785 745 | 754 702 664 638 638 | $\begin{array}{c} 1,444\\ 2,080\\ 1,522\\ 1,995\\ 1,995\\ 675\end{array}$ | 2, 185 2, 076 1, 856 925 608 | 635 674 574 704 561 | 694 569 618 601 | 910 905 879 830 830 | 708 735 715 720 667 | 584 500 516 | 928 |
| 999 958 960 | $1,\ 002\\1,\ 195\\1,\ 081\\915\\879$ | 2, 392 2, 764 2, 703 2, 703 | 3, 166 3, 079 2, 587 1, 856 | ${}^{1,581}_{1,370}$ ${}^{761}_{1,723}$ ${}^{1,723}_{1,399}$ | $1,665\\1,649\\1,378\\1,129\\1,111\\1,111$ | $\begin{array}{c} 997\\ 972\\ 972\\ 1,038\\ 1,018\\ 1,061\end{array}$ | 989 927 950 963 963 | 856 749 868 884 | 1, 036 |
| 1,280 1,280 | 1, 400 1, 240 1, 240 | 4, 300 4, 842 4, 710 4, 710 | $\begin{array}{c} 4, \ 710\\ 4, \ 524\\ 4, \ 372\\ 4, \ 839\\ 4, \ 839\end{array}$ | $\begin{array}{c} 3,400\\ 1,100\\ 3,536\\ 3,536\\ 3,536\\ 3,536\end{array}$ | 3, 095 3, 765 3, 765 3, 765 2, 500 | 1, 220 1, 050 1, 420 1, 420 1, 420 | $1,\ 420\\1,\ 220\\1,\ 343\\1,\ 343\\1,\ 343\\1,\ 343\\1,\ 343\\1,\ 342\\1,\ 342\\1,\ 343\\1,$ | 1,200 1,200 1,200 | 1, 550 |
| 46.2 10.4 28.7 | 62. 2 55. 4 53. 3 76. 9 | 63.3 75.7 75.7 75.7 75.7 75.7 51.2 51.2 | 29.1 34.8 47.4 65.6 142 | $\begin{array}{c} 61 \\ 50.9 \\ 85.5 \\ 85.2 \\ 85.2 \\ 119.7 \end{array}$ | 63 94.3 57.9 64.1 | 33.8 64.4 33.8 33.8 33.8 | 87.5 44 86.5 53.3 84.8 | 34.8 15.7 101.4 70 | 38.1 |
| 56.8 9.4 29.3 | 51.4 61.6 86.9 81.8 81.8 | 74.3 64.4 139.2 43.3 105.5 | $\begin{array}{c} 31.6\\ 21.1\\ 44.3\\ 64.8\\ 64.8\\ 152.6\end{array}$ | $\begin{array}{c} 61.9\\ 34.7\\ 79.9\\ 86.3\\ 86.3\\ 142.7\end{array}$ | 69.1 86.8 52.2 58.4 | 854.0 36.40 36.40 36.40 | 60.0 51.1 68.0 57.7 80.5 | 38.4 14.8 92.4 77.5 | 44.4 |
| 3.01 19.23 9.97 | 8.21 9797 4.97 | $\begin{array}{c} 22.76\\ 17.19\\ 10.52\\ 27.19\\ 25.15\\ \end{array}$ | $\begin{array}{c} 44.12\\ 72.76\\ 18.59\\ 35.41\\ 11.56\\ 11.56\end{array}$ | $\begin{array}{c} 24.1\\ 18.21\\ 2.28\\ 111.26\\ 7.01\end{array}$ | 18.06 13.05 7.79 7.30 | $^{2.19}_{12.04}$ | 6. 47 6. 75 5. 28 7. 32 7. 51 | 7.07 7.06 6.72 8.5 | 2.52 |
| 39.5 127 74.5 | 26.4 54.6 32.7 29.5 | 108 154 114 275 148 | 347 259 130 285 145 | 140 198 26. 5 1130 118 | 171 113 86. 1 120 120 | 19.4 11.9 17.8 18.1 | 26.6 28.7 28.7 24.2 24.2 | 59.9 64.5 21.4 | 17 |
| 417 799 660 | $^{363}_{1, 642}$ 1, 642 1, 673 742 742 | $1, 338\\1, 575\\1, 455\\1, 335\\1, 335\\1, 321\\$ | $1,\ 318\\1,\ 414\\1,\ 183\\1,\ 768\\1,\ 770\\$ | $\begin{array}{c} 2,218\\ 2,244\\ 830\\ 2,002\\ 2,077\end{array}$ | 2, 563 2, 302 2, 325 2, 339 2, 339 | 140 1100 1100 1100 | 131 146 114 96 227 | 1, 016 806 216 183 | 126 |
| 414 731 667 | $ \begin{array}{c} 358 \\ 1, 610 \\ 1, 670 \\ 532 \\ 746 \\ 746 \end{array} $ | $\substack{1,\ 332\\1,\ 606\\1,\ 474\\1,\ 349\\1,\ 302 \end{cases}$ | $^{1,373}_{1,133}$ $^{1,133}_{1,774}$ $^{1,722}_{1,722}$ | $\begin{array}{c} 2,149\\ 2,169\\ 853\\ 1,993\\ 2,061 \end{array}$ | $\begin{array}{c} 2, 511\\ 2, 268\\ 2, 268\\ 2, 077\\ 2, 496\\ 2, 496 \end{array}$ | 100 56 118 134 | 130 149 116 107 226 | 983 802 218 162 | 118 |
| 421 882 650 | 368 1, 661 1, 677 1, 677 737 737 | $1,\ 345\\1,\ 534\\1,\ 431\\1,\ 317\\1,\ 344\\1,\ 344\\1,\ 344\\$ | 1,257 1,334 1,212 1,764 1,832 | $\begin{array}{c} 2, 260\\ 2, 294\\ 2, 013\\ 2, 013\\ 2, 098\\ \end{array}$ | 2, 594 2, 343 2, 343 2, 343 2, 495 2, 113 2, 113 | 103 64 96 97 147 | 132 136 110 83 83 83 | $1,059\\811\\214\\199$ | 136 |
| $12, 196 \\ 129 \\ 1, 891$ | $\begin{array}{c} 16,414\\ 10,822\\ 35,364\\ 10,582\\ 40,234\end{array}$ | $15, 130 \\ 17, 475 \\ 91, 988 \\ 5, 597 \\ 74, 848 \\$ | $\begin{array}{c} 2,041\\ 1,569\\ 7,542\\ 11,004\\ 88,456\end{array}$ | $\begin{array}{c} 12,726\\ 4,740\\ 23,203\\ 32,960\\ 91,646\\ \end{array}$ | $\begin{array}{c} 15,246\\ 56,421\\ 5,421\\ 5,436\\ 6,708\\ 6,708 \end{array}$ | $\begin{smallmatrix} 5, & 373 \\ 9, & 778 \\ 13, & 733 \\ 3, & 533 \\ 3, & 533 \\ \end{bmatrix}$ | $\begin{array}{c} 17,090\\ 5,060\\ 9,800\\ 31,960\end{array}$ | 6, 271 551 60, 323 16, 837 | 13, 509 |
| 1.87 2.22 2.26 | 22228 2009 2009 2009 2009 2009 2009 2009 | 2.03 2.14 1.42 1.56 1.56 | 1.1.1.40 1.26 80 1.26 1.26 | $ \begin{array}{c} 1.87 \\ 2.12 \\ 1.59 \\ 2.12 \\ 2.12 \\ 2.12 \\ \end{array} $ | 22 09 22 17 1. 77 1. 64 | 1.39 1.16 1.37 1.37 1.42 | 1. 32 1. 37 1. 62 1. 65 1. 64 | 2,09 2,17 1,62 1,73 | 1.44 |
| $\frac{466}{27.5}$ | 672 386 913 944 944 | 438 540 357 315 315 | 128 150 281 281 145 | 393 267 762 226 | 389 225 150 209 209 | 255 73. 3 387 387 195 | 544 544 533 331 808 | 286 76.5 477 | 545 |
| Kilibuck Creek at Kilibuck, Ohio Mill Creek near Coshocton, Ohio Wakatomika Creek near Frazeysburg, Ohio | Licking River at Toboso, Ohio Little Kanawha River at Glenville, W. Va Hickite Kanawha River at Grantsville, W. Va Hocking River at Enterprise, Ohio | Bluestone River at Lilly, W. Va. Greenbrier River at Buckeye, W. Va. Greenbrier River at Alderson, W. Va. Gauley River at Chamden-on-Gauley, W. Va. Gauley River above Belva, W. Va. | Williams River at Dyer, W. Va. Cherry River at Feawick, W. Va. Deadow River at Nallen, W. Va. Eik River at Centralla, W. Va. | Coal River at Ashford, W. Va. Little Coal River at Madison, W. Va. River at Manswille, Ohio Guyandot River at Man, W. Va. Guyandot River at Branchland, W. Va. | Levisa Fork at Fishtrap, Ky | Scioto River at Larue, Ohio | Big Walnut Creek at Rees, Ohio Alum Creek at Columbus, Ohio Darby Creek at Darbyville, Ohio Deer Cyreek at Williamsport, Ohio | ой Ц Ц Ц Ц Ц Ц Ц Ц | Miami River at Sidney, Ohio |
| 3-161 3-163 3-166 | 3-1 67 3-1 69 3-1 70 3-175 | 3-203 3-205 3-206 3-206 3-208 2-211 | 3 -213. 3-216. 3-217. 3-220. | 3 -224 3-227 3-230 3-232 3 -233 | 3 -235.3 3 -235.7 3 -236.7 3 -238.5 3 -238.5A | 3-243 3-250.5 3-252 3-252 3-252 3-252 | 3-2 55. 3-2 56. 3-2 57. 3-2 57. 3-2 58. | 3-261.5 3-262 3-269 | 3-278 |

CONTRIBUTIONS TO HYDROLOGY, 1944

| | -duz to ruber of sub- steas into which -dus aas missd babivib | | 76 87 91 91 51 | 29 | 74 38 53 131 | 17 33 33 33 19 20 17 | 56 112 220 | 69 56 | 230 230 35 30 39 | 140 53 | 82 2 |
|---|--|------------------------------|---|---|---|---|---|--|--|-----------------------|---|
| Summary of drainage basin lopographic characteristics-Continued | water (square | IstoT | | | | | | | | | |
| | Area of surfaces (s miles) | sqmsw8 | | | | | | | | | |
| | | Lakes and reser- voirs | $ \begin{array}{c} 0.03 \\ 0.16 \\ 0 \\ 0 \end{array} $ | .04 | 2.05 .07 | | .06 1.69 | 88. | $^{16}_{-16}$ | 18,58 2,15 | 3, 14 55, 71 11, 25 |
| | Altitude of land (feet a bove mean sea level) | muminiM | 850 952 880 880 880 880 880 | 485 | 790 955 784 578 | 990 610 675 605 1,060 | 1, 1 35 655 553 | 632 800 | $^{490}_{930}$ $^{375}_{1, 195}$ | 286 736 | 1,500 226 |
| | | цвэМ | $\begin{array}{c} 1,032\\ 1,069\\ 1,079\\ 910\\ 927\end{array}$ | 845 | 1,089 988 880 | 1, 120 1, 055 986 1, 111 1, 111 1, 349 | 1,877 1,741 1,527 | 1,491 1,319 | $\substack{1,\ 408\\1,\ 576\\2,\ 076\\1,\ 650\\\end{array}$ | 1,049 $1,256$ | 1,272 2,036 1,311 |
| | | mumixeM | $\begin{smallmatrix} 1, 220\\ 1, 550\\ 1, 220\\ 1,$ | 1, 060 | 1,280 1,360 1,340 1,340 | $1,\ 280\\1,\ 860\\1,\ 920\\1,\ 800\\1,$ | 2, 565 2, 565 2, 565 | 2, 140 2, 028 | 2, 080 3, 670 2, 300 2, 300 | 2,460 | 2, 620 4, 621 3, 305 |
| | Length of prin- cipal streams (miles) | | 80.89 20.88 20.88 20.89 20.80 | 111.8 | 27.9 24.4 72.9 130.7 | 30.9 8.6 8.6 | 35 73.9 105.3 | 26.8 33.0 | 35.2 41.9 60.5 26.5 | 127.9 64.1 | 66. 2 72. 7 106. 3 |
| | Longest watercourse (miles) | | 46.5 36.1 36.1 19.7 39.6 | 85.5 | 28.8 28.1 58.5 119.5 | 39.7 39.7 33.0 33.0 7.9 | 29.4 85.1 101.9 | 22.4 29.5 | 38. 3 40. 0 39. 5 39. 5 24. 2 | 124.3 44.5 | 58.7 78.5 71.6 |
| | Channel slope (feet per mile) | [sqiənir4 | 5 11 5 56 15 29 15 29 15 7 | 5.23 | 9.18 9.34 9.34 3.96 | $\begin{array}{c} 15.4 \\ 15.05 \\ 26.38 \\ 22.5 \\ 69.2 \end{array}$ | 15.7 14.8 10.58 | 82.9 23.18 | 45.06 28.6 17.09 12.36 33.74 | 11. 29 24. 18 | 28.3 6.01 25.18 |
| | | Tributary | 12 14.8 50.4 36.3 | 23.1 | 30.2 23.4 17.4 | 19.6 78.3 112 229 | 107 78.3 98.2 | 166 | 56.1 56.9 56.9 61.2 61.2 | 39. 7 60. 6 | 47.7 81.9 48.8 |
| | Land slope (feet per mile) | A verage | 97 124 166 220 | 324 | 121 121 93 | 261 359 426 476 | 717 566 539 | 589 408 | 363 397 469 489 | 481 479 | 469 764 417 |
| | | .8N | 221 221 221 221 221 221 221 221 221 221 | 328 | 165 85 103 91 | 262 340 409 412 | 736 562 526 | $514 \\ 374$ | 322 388 388 870 870 519 | 470 | 450 819 447 |
| | | EW. | 92 117 191 174 216 | 320 | 177 112 112 95 | 260 386 295 451 560 | 697 570 557 | 641 454 | 483 404 450 634 634 | 498 500 | 495 691 378 |
| | al (cubic miles) & | | 12, 496 4, 008 9, 232 6, 070 | 21, 271 | 1, 781 1, 346 9, 138 74, 555 | 3, 510 2, 855 2, 855 88 88 | 4, 989 49, 429 74, 938 | 1,578 2,014 | 3, 497 5, 964 133, 714 8, 704 1, 052 | 68, 586 5, 176 | 10, 163 28, 408 24, 457 |
| | Stream density (miles per square mile) | | 1.2.1.33 | 1.46 | 1.16 1.40 1.38 1.55 | .65 1.71 1.94 1.94 1.06 | 1.39 1.45 1.49 | 1.52 1.53 | $ \begin{array}{c} 2.26 \\ 1.87 \\ 2.18 \\ 2.18 \\ 1.68 \\ 1.68 \\ \end{array} $ | $1.62 \\ 1.53$ | 1. 59 1. 88 1. 51 |
| | Drainage area (square miles) | | 502 195 195 315 315 | 463 | 113 89.8 299 1, 248 | $^{42}_{236}$ | 309 1, 017 1, 419 | 153 124 | 189 295 1, 876 365 85 | 973 258 | 335 722 616 |
| | Name of gaging station | | Stillwater River at Pleasant Hill, Ohio Greenville Creek near Bradford, Ohio Mad River near Springfield, Ohio Wolf Creek at Dayton. Ohio Talawanda Creek near Sevenmile, Ohio | Elkhorn Creek, at Knights Bridge, near Frank- | Cedar Creek near Cedarburg, Wis Sandusky River near Busyrus, Ohio Sandusky River near Upper Sandusky, Ohio. Sandusky River near Freemont, Ohio | Little Cuyahoga River at Akron, Ohlo Buffalo Creek at Gardenville, N. Y. Cazuga Creek near Laneaster, N. Y. Cazenovia Creek at Ebenezer, N. Y. Little Tonawanda Creek at Linden, N. Y. | Genesse River at Scio, N. Y. Genesse River at St. Helena, N. Y. Genesse River at Jones Bridge, near Mount | Canaseraga Creek near Dansville, N. Y. | East Branch of Fish Creek at Tabers, N. Y Black River near Bounville, N. Y. Black River at Wateotown, N. Y. Moose River at McKewer, N. Y Independence River at Sperryville, N. Y | | Grass River at Pyrites, N. Y Raquetta River at Prercefield, N. Y Raquetta River at Piercefield, N. Y St. Regis River at Brasher Center, N. Y |
| | No. | | 3-289 3-298 3-296 3-296 3-298A 3-300A | 3-321 | 4-50. 4-119. 4-120. | 4-134 4-139.5 4-140.3 4-140.7 4-142 | 4-143 4-144 4-146 | 4- 148. 4- 172 | 4-180. 4-197 4-201 4-204 | <u>4-227</u> 4-229 | 4-230. 4-233. 4-236. |

Summary of drainage basin topographic characteristics-Continued

822889 8 82888 442884 881 | |

| | | $ \begin{array}{c} 5.36 \\ 1.38 \\ 11.12 \end{array} $ | .31 .33 13.52 | 0 |
|--|---|--|---|---|
| | | 8. 268 8. 268 | 6. 33 | 0 |
| ¢.4.9.% 8.7888 8.2888 | .57 | . 87 . 80 . 80 . 80 . 80 | .02 7.10 0 | 0.02 |
| $1,012 \\ 880 \\ 1,630 \\ 524 \\ 524$ | 546 | 165 150 510 320 | 620 550 770 725 | 660 680 690 |
| 1, 760 1, 716 2, 434 2, 059 | 1, 905 | $^{1, 191}_{1, 617}$ | $1,486\\1,656\\1,526\\1,005\\1,025$ | 977 960 780 |
| 3, 355 3, 355 5, 112 5, 344 | 5, 344 | 4, 842 2, 727 4, 241 4, 241 | 2, 923 4, 135 3, 315 1, 420 1, 362 | $1,362 \\ 1,441 \\ 840 \\ 840 \\ 1,1,1,1,1,1,1,1,1,1,1,1,1,1,1,1,1,1,1$ |
| 27.8 25.5 21.1 54.9 | 30.5 | 49.8 35.8 71.0 | 15.7 19.8 32.7 14.6 20.7 | 20.6 3.2 3.2 |
| 19.8 40.3 40.4 | 38.3 | 42.7 25.0 33.1 74.0 | 14.0 32.2 15.2 13.3 | 23.7 91.5 3.3 |
| 38. 63 27. 49 36. 29 36. 96 | 47.2 | 24.44 24.0 34.9 9.9 | 64. 1 37. 6 19. 6 15. 27 30. 14 | 14.56 3.43 32.9 |
| 132 218 122 351 351 | 433 | 271 177 248 178 | | 136 63.3 115 |
| 712 2559 2530 2833 2590 2833 2590 2590 2590 2590 2590 2590 2590 2590 | 1, 538 | 828 814 729 | 1, 022 1, 120 578 1, 039 1, 012 | 1, 090 1 876 475 1 |
| $\begin{array}{c} 675 \\ 523 \\ 279 \\ . 211 \\ . 211 \\ 1 \end{array}$ | 1, 548 1 | , 027 690 621 621 | , 037 1 523 1 980 1 | 1,006 1 |
| 766 610 285 071 1 256 1 | 1, 524 1 | 1, 107 1, 017 949 881 | 010 010 044 1 010 1 010 1 014 | , 149 1 844 |
| 493 757 948 913 1 | 615 | 566 566 960 | 571 1 5649 1 5566 1 552 1 | $\begin{array}{c c} 1,616 \\ 34,720 \\ 5.2 \\ \end{array}$ |
| 8585258 9.1.5.1.4 | 1.94 3, | 22 27 208 25 27 208 | 28840 2,1, 2,1, | 72 1 85 34 63 |
| | | 4444 | 2.24 2.40 1.94 1.71 | |
| 132 112 247 448 | 198 | 275 186.5 307 628 | 76.1 139.0 140 77.1 77.2 | 119 699 3.6 |
| Salmon River at Chasm Palls, N. Y. Chateaugay River near Chateaugay N. Y. Grateaugay River at Perry Mills, N. Y. Chateaugay West Branch of Aussible River near Newman, N.Y. Aussible River near Aussible Forks, N. Y | East Branch of Ausable River at Ausable | Poultney I Poultney I Otter Creel | Dog River at Northfield Falls, Vt- Mad River near Moretown, Vt- Organ River at Newport, Vt- Little La Crosse River near Leon, Wis- Little La Crosse River near Leon, Wis- | Coon Creek near Stoddard, Wis Kickapoo River at Steuhen, Wis Raiston Creek at Iowa City, Iowa |
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