

Tropical moist *Polylepis* stands at the treeline in East Bolivia: the effect of elevation on stand microclimate, above- and below-ground structure, and regeneration

Dietrich Hertel · Karsten Wesche

Received: 24 January 2007 / Revised: 10 July 2007 / Accepted: 18 October 2007 / Published online: 29 January 2008
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Abstract We studied *Polylepis* forests along an elevational transect between 3,650 and 4,050 m a.s.l. at the treeline of the moist eastern cordillera in Bolivia to examine changes in above- and below-ground stand structure, leaf and root morphology, and regeneration in relation to stand microclimate. Field measurements and model predictions indicated relatively cold growth conditions of the *Polylepis* forests. Tree height, stem diameter, and basal area of the stands decreased markedly while stem density increased with elevation. Leaf morphology differed between the two occurring *Polylepis* species, and trees at the treeline had smaller leaves with higher specific leaf area. In contrast, fine root biomass increased from 37 g m⁻² at the lowermost stand to 234 g m⁻² at the treeline. Trees of the uppermost stand had higher specific root surface area and a much higher number of root tips per unit dry mass. Thus, root surface area and total number of root tips per unit ground area increased conspicuously from the lowermost stand to the treeline. Density of young growth inside the forest increased towards the treeline, while density in the open grassland decreased with elevation. Young growth originated from sexual reproduction at the lower forest but

was comprised exclusively of root suckers at the treeline stand. We conclude that both the marked change in carbon allocation towards the root system, as well as the changes in root morphology with elevation indicate an adaptation to reduced nutrient supply under cold conditions of these *Polylepis* stands at the treeline in E Bolivia.

Keywords Fine root biomass · Leaf morphology · Reproduction · Root morphology · Temperature

Introduction

In tropical and subtropical South America, isolated forest patches covered by trees of the genus *Polylepis* are present above the closed upper montane forest belt. Among the ca. 30 existing *Polylepis* species several are found at sites with extreme climatic conditions reaching altitudes up to 5,200 m a.s.l., and covering a wide range of high to very low rainfall (Simpson 1986; Kessler 1995; Kessler and Schmidt-Lebuhn 2005). There has been a long standing debate whether these forest represent living “fossil” remnants of warmer postglacial periods or remnants of a potentially larger natural forest belt, that was destructed by long-term human land-use (Ellenberg 1979; Goldstein et al. 1994; Kessler 1995; Kessler and Hohnwald 1998; Hensen 2002). However, stands forming at least a potential upper treeline are thus of high interest for treeline ecology.

Many studies on high elevation forest sites documented a decline in above-ground biomass and production at tree-line sites compared to forest sites at lower elevation. This is often accompanied by notable changes in leaf morphological and chemical traits. The changes in structure, morphology, and carbon cycle are predominantly influenced by the decrease in temperature with elevation (e.g.,

Communicated by M. Buckeridge.

D. Hertel (✉)
Plant Ecology, Albrecht-von-Haller Institute of Plant Sciences,
University of Göttingen, Untere Karspüle 2, 37073 Göttingen,
Germany
e-mail: dhertel@gwdg.de

K. Wesche
Institute of Biology, Geobotany and Botanical Garden,
Martin-Luther University Halle-Wittenberg,
Am Kirchtor 1, 06108 Halle/Saale, Germany

Tranquillini 1979; Oshawa 1990; Körner 1998). Recent studies confirmed that treelines are on a global scale associated with a mean seasonal root zone temperature of 6.7°C (Körner and Paulsen 2004).

Analyses of the mobile carbohydrate contents revealed higher pools at the treeline compared to trees below the treeline, which is seen as indicating carbon sink-limitation rather than carbon source-limitation as the principle factor of tree growth restriction (Hoch et al. 2002; Hoch and Körner 2003; Körner 2003a). Other authors stated that constraints in fruit production, seed dispersal and viability, and seedling establishment are crucial for forest survival and extension at the treeline (Wardle 1981; Cuevas 2000, 2003; Cierjacks et al. 2008). However, detailed studies on *Polylepis* species growing at high altitudes were rarely conducted so far, and failed to consider the root system of the trees (e.g. Rada et al. 1996; Velez et al. 1998; Rundel et al. 2003; Garcia-Nunez et al. 2004; Renison et al. 2004; Hoch and Körner 2005).

We investigated *Polylepis* stands along an elevational transect at the treeline of the eastern Andean cordillera in Bolivia. The objective of our study was to test the hypotheses that (1) elevational changes in the above-ground structure of high-Andean *Polylepis* forests are accompanied by changes in the structure of the root system as well, that (2) the *Polylepis* trees show significant adaptations in leaf and root morphology to the harsh growth conditions at the treeline, and that (3) regeneration pattern of *Polylepis* trees changes with increasing elevation.

Hence, the aims of our study were to analyse (1) how above- and below-ground structure of the *Polylepis* forests changes with elevation, (2) whether *Polylepis* trees differ in morphological and chemical leaf and root traits along the elevational transect, and (3) if constraints in regeneration of the *Polylepis* stands may contribute to treeline formation.

Material and methods

Study site location and species composition of the forests

The study was conducted in the area of Chumbre Chulumani (Department Cochabamba, Province Tiraque) at the humid eastern cordillera of Bolivia between August and September 2003. In this remote area, some largely undisturbed forest stands of *Polylepis* species grow above a closed montane forest belt. We selected three neighbouring forest stands mainly built by *Polylepis* trees along an elevational transect between 3,650 and 4,050 m a.s.l. The lowermost forest was dominated by trees of *Polylepis*

lanata (Kuntze) M. Kessler & Schmidt-Leb.¹ and had a lower second tree layer dominated by *Oreopanax macrocephalus*, *Baccharis pentlandi* and other tree species; herbaceous plants were less abundant with some ferns and *Fuchsia* plants. The two forests at higher elevation were dominated by trees of *Polylepis pepeii* B.B. Simpson. At 3,800 m a.s.l. individuals of the genus *Gynoxis* were admixed in the forest, while in the uppermost stand (at 4,050 m a.s.l.) some *Ribes* shrubs were present. In both forests, the soil was covered by mosses and herbaceous plants (genera *Luzula* and *Bomarea*). Following the definition given by Körner and Paulsen (2004), we assume that the highest forest stand (at 4,050 m a.s.l.) represents the upper treeline in this region (hence, the upper distribution limit of connected forest patches in the area). All forests grow on east-facing slopes with various inclination (Table 1).

Soil characteristics

All three *Polylepis* forests were located on boulder slopes with medium to low mineral soil contents. The mineral soil was covered by an organic layer of slightly increasing thickness with elevation (Table 1). Thickness of the mineral soil layer was ca. 8 and 14.5 cm, respectively, in the two lower forest stands, but declined to less than 2 cm in the uppermost stand. Soil chemical conditions tended to be less favourable in the two uppermost *Polylepis* stands: mineral and organic soil layers were less acidic at the lowermost site, and cation exchange capacity as well as base saturation were higher at the lowermost site (Table 1). The C:N ratio of the organic layer increased along the elevational transect, although the C:N ratio of the mineral soil showed no clear altitudinal gradient. The resin-extractable soil phosphorus fraction (as a proxy for plant-available P), on the other hand, was highest at the treeline site (Table 1).

Climatic measurements

There are no permanent weather stations in the study region, so we had to use extrapolated values from a standard model (Hijmans et al. 2005) to estimate average climatic conditions. The data are based on lapse rates and have a moderate spatial resolution (1 km²) and thus yield similar values for all three stands. However, microclimatic differences among sites were measured on a short-term base during fieldwork (4 days). Dataloggers

¹ Taxonomy of the *Polylepis* species according to Kessler and Schmidt-Lebuhn (2005).

Table 1 Topographical and soil chemical characteristics of the three *Polylepis* stands

		Plot A	Plot B	Plot C
Elevation (m a.s.l.)		3,650	3,800	4,050
Coordinates		17°16'36.6" S 65°43'09.8" W	17°16'23.6" S 65°43'27.1" W	17°16'18.5" S 65°43'54.6" W
Exposition		NE	E	SE
Inclination (°)		4	22	32
Mean thickness soil layers (cm)	Organic layer	9.8 ± 2.1	11.2 ± 1.8	13.1 ± 1.5
	Mineral soil	8.2 ± 1.7	14.5 ± 2.1	1.7 ± 0.4
pH (H ₂ O)	Organic layer	6.1 (5.8–6.3)	4.7 (4.6–5.0)	4.8 (4.7–5.3)
	Mineral soil	6.0 (5.8–6.2)	4.4 (4.1–4.6)	4.2 (4.2–4.4)
C/N (mol mol ⁻¹)	Organic layer	20.2 ± 0.6	25.4 ± 0.6	37.6 ± 3.1
	Mineral soil	16.1 ± 0.7	14.8 ± 0.2	17.9 ± 0.3
P _{resin} (μmol P g ⁻¹)	Organic layer	6.6 ± 0.5	7.1 ± 1.5	14.2 ± 2.6
	Mineral soil	3.7 ± 1.1	1.2 ± 0.1	5.6 ± 0.6
CEC (μmol _c g ⁻¹)	Organic layer	936 ± 32	740 ± 76	609 ± 55
	Mineral soil	723 ± 49	259 ± 12	496 ± 59
Base saturation (%)	Organic layer	96.2 ± 0.5	45.9 ± 3.9	54.0 ± 3.3
	Mineral soil	96.3 ± 1.1	9.3 ± 1.2	25.8 ± 2.4

Given are means and standard errors for the soil parameters; $n = 20$ for soil layer thickness, and $n = 5$ for pH, C/N, extractable phosphorus (P_{resin}), cation exchange capacity CEC, and base saturation. For pH, minimum and maximum values are presented in brackets

(*BauFeuchteLog*, Driesen & Kern, Germany) were placed inside the forests and in the open grassland east of the forest stands at the same elevation and exposition; sensors were screened against direct radiation with ventilated plastic containers. Data were recorded for two levels above the ground (+20 and +200 cm) and at -10 and -20 cm soil depth. Recording intervals were 2 min, but only mean values and maximum and minimum ranges are reported here.

Above-ground forest structure

We recorded tree height, diameter at breast height (dbh) of all tree stems, and stem density in plots of 50 m² at the uppermost site, while plot size was 100 and 300 m² in the mid-positioned and the lowermost site, respectively, to include a sufficient number of dominant tree individuals per plot. Tree height was measured using a Hagloef (Sweden) Vertex III Ultrasonic device from 12 randomly selected tree individuals in the two uppermost stands, and from 15 trees in the lowermost stand covering the two crown strata in this forest. The stem density and dbh data were used to calculate the stand basal area of the forest plots.

Leaf morphology and chemistry

Ten leaves from different branchlets were sampled from the upper sun-lit crown of six *Polylepis* trees, each stand,

resulting in a total of 60 leaves per stand. Leaves were stored in paper bags, air dried, and transported to the Göttingen laboratory within a few days. Leaf area was measured on scanned images of the rehydrated leaves (WinFolia software package, Régent Instruments Inc., Canada). Dry mass of the leaves was determined after drying for 48 h at 70°C, and then used to calculate the specific leaf area (SLA). Dried leaves were ground for analyses of leaf chemistry and carbon isotopes. Total C and N contents as well as stable carbon isotopic status ($\delta^{13}\text{C}$ values) were measured by stable-isotopic mass spectrometry. Total P concentration was determined photometrically with a staining procedure (p-yellow), after digestion of leaf tissue with 65% HNO₃ at 195°C.

Structure of the fine root system

Root samples were taken with a soil corer (3.5 cm in diameter) from the organic layer and the upper mineral soil (0–10 and 10–20 cm, respectively) at 20 randomly selected sampling locations per study plot. As described above, mean thickness of the mineral soil was only about 8 cm in the lowermost stand and less than 2 cm in the uppermost stand; hence, only in the mid-positioned stand a notable number of samples reached mineral soil depths below 10 cm. The soil samples were transferred to plastic bags and stored at ca. 4°C. The samples were transported to the laboratory at the University of Göttingen within a few days, where processing of the samples took place within 60 days. In the lab, the samples were soaked in water and cleaned

from soil residues using a sieve with a mesh size of 0.25 mm. Large root fractions (>10 mm in length) were extracted by hand with tweezers. Only fine roots of trees (roots < 2 mm in diameter) were considered for the analysis. Live (biomass) and dead rootlets (necromass) were distinguished under the stereomicroscope by color, root elasticity, and the degree of cohesion of cortex, periderm and stele. A dark cortex and stele, or a white, but non-turgid cortex, or the complete loss of the stele and cortex with only the periderm being present, were used as indicators of root death (e.g., Persson 1978; Hertel and Leuschner 2002). This fine root fraction covers the majority of living fine root mass, but represents only a minor part of the fine root necromass, because small fractions (<10 mm length) of dead roots, that may account for a large portion of the whole fine root necromass (Bauhus and Bartsch 1996; Hertel 1999), are hardly recovered. Hence, one third of the samples were subjected to an additional, more detailed analysis of finest root necromass particles (<10 mm in length) applying a method introduced by van Praag et al. (1988) and modified by Hertel (1999). After extraction of the large root fractions the residue of the sample was evenly spread on a large sheet of filter paper (730 cm²) with 36 squares marked on it. Six of the squares were randomly selected and analysed under the stereo-microscope for even the smallest dead fine root fragments. Total mass of small dead rootlets was extrapolated by regression analysis of the ratio of small rootlets to large dead roots (>10 mm in length) recorded in the same sub-sample. Fine root biomass and necromass of each sample were dried at 70°C for 48 h and weighed. The data were expressed in relation to soil volume as fine root mass density (in g l⁻¹) for the respective soil depths, and as profile fine root mass total (in m⁻²).

Root morphology and C isotopic signature

One living rootlet per soil sample was analysed for root specific surface area (SRA, in cm² g⁻¹) using a WhinRhizo (Régent Instruments Inc., Quebec, Canada) visual analysis system with a scanner in the Göttingen lab. Fine root area index (RAI, in m² root surface area per m² ground area) was calculated from SRA values and fine root biomass of a respective horizon.

The abundance of live fine root tips was determined by counting all vital tips under the stereo-microscope in the same fine root sample used for analysis of SRA. Tip abundance was expressed on a dry mass basis (n mg⁻¹) or on a stand ground area basis (number of tips per m²).

Fine root stable carbon isotopic status ($\delta^{13}\text{C}$ values) was measured by stable-isotopic mass spectrometry in the isotopic lab in Göttingen (KOSI).

Polylepis regeneration pattern

The number of juvenile plants (<2 cm dbh) was counted on 4 randomly placed subplots of 4 m² inside the forest stands; these were compared to 4 randomly chosen subplots outside. We excluded shoots emerging directly from bases of established trees. In the two lower forest stands, all seedlings were excavated and checked for obvious connections to adult trees. In the uppermost plot the sheer number of young growth rendered complete assessment impossible so we excavated only a subset (40 shoots).

Statistical analysis

Climatic data were not replicated, so we refrained from statistical analysis and present condensed raw data instead. All other data sets were analysed for distribution of normality using a Shapiro and Wilk test. Normally distributed data were tested using an ANOVA procedure followed by pair-wise comparison after Scheffé ($p < 0.05$). Other data sets were tested with one-way Kruskal-Wallis single factor analyses of variance. Significance was determined at $p < 0.05$ in all analyses. If H_0 (no significant difference among the sites or horizons) was rejected, a non-parametric Mann-Whitney two-sample test (U-test) was applied to locate the differences ($p < 0.05$). All calculations were conducted with the package SAS, version 8.01 (SAS Institute, Cary, NC, USA).

Results

Microclimate inside and outside the stands

The climate model predicted mean annual temperatures ranging from -0.1 to 15.2°C (mean minimum to maximum) with rough mean annual temperature values ranging from 8.5 at the lower to some 6.0°C at the middle stand and mean total precipitation of 700–800 mm (the uppermost stand was not differentiated due to limited spatial resolution). For September, the model predicts a mean daily thermal range between 0.2 and 16.7°C at the lower and -2.6 to 14.2°C at the uppermost stand. Our own microclimatic measurements at 2 m above ground (comparable to a standard weather hut) yielded a temperature range between -0.6 and 11.5°C at 3,850 m a.s.l. and between -2.3 and 9.5°C at 4,050 m a.s.l. Measured mean temperatures were low, ranging between 3.4°C at 3,680 m and 1.7°C at 4,050 m (Table 2). This corresponds to a mean lapse rate of 0.46 K 100 m⁻¹ along the entire elevational transect. Low lapse rate between the upper two stands

Table 2 Micro-climatic conditions inside the three *Polylepis* stands (In) and at the adjacent grasslands (Out)

Elevation (m)	3,650		3,800		4,050	
	In	Out	In	Out	In	Out
Mean air temperature +200 cm (°C)	3.4 (1.0–6.3)	3.7 (–0.6 to 11.5)	2.1 (–0.9 to 4.7)	1.8 (–1.7 to 6.1)	1.7 (–1.2 to 6.7)	1.9 (–2.3 to 9.5)
Mean air temperature +20 cm (°C)	3.4 (1.5–6.1)	n.d.	1.7 (–1.4 to 4.5)	1.5 (–4.5 to 8.6)	1.0 (–3.3 to 7.9)	2.2 (–4.2 to 17.6)
Mean soil temperature –20 cm (°C)	4.3 (4.1–4.5)	7.2 (6.1–8.6)	2.9 (2.7–3.2)	2.7 (2.4–3.1)	2.0 (1.4–3.7)	2.8 (1.4–5.2)
Mean VPD +200 cm (kPa)	0.11 (0.03–0.34)	0.15 (0.02–0.74)	0.10 (0.03–0.31)	0.09 (0.02–0.31)	0.13 (0.02–0.44)	0.15 (0.02–0.66)
Mean VPD +20 cm (kPa)	0.08 (0.04–0.23)	n.d.	0.07 (0.03–0.19)	0.08 (0.02–0.37)	0.11 (0.03–0.46)	0.19 (0.03–1.40)

Given are mean temperature and daily mean minimum and maximum values for air temperature and vapor pressure deficit (VPD) at 200 and 20 cm above ground, and soil temperature at 20 cm soil depth

suggests that air masses were not homogenous over that altitudinal range.

Conditions were warmer outside the forests, but mean temperatures did not differ much between +200 and +20 cm above ground. The range of minimum and maximum temperatures was generally larger near the ground and in the open grassland (Table 2). Both the overall highest and lowest temperatures along the transect were measured in the open grassland at 4,050 m a.s.l. Temperatures in the soil (–20 cm) were higher than in the air and this was particularly pronounced in the open grassland (Table 2). Forest soils were found to be cooler than grassland soils at the lowermost and uppermost elevation. Soil temperature in –10 cm (data not shown) had higher maximum and lower minimum values than in –20 cm, but mean temperatures in the two soil depths were almost identical with values inside the forests ranging from 4.3 to 2.0°C at 3,680 and 4,050 m a.s.l., respectively.

Mean vapor pressure deficits (VPD) were generally low at <0.2 kPa (Table 2). VPD tended to be higher outside the forests, and tended to increase with altitude, with highest values recorded at the uppermost grasslands. There was pronounced diurnal variation, extreme values ranged between 0 and 1.4 kPa.

Above-ground structure

The mean tree height of the stands decreased significantly with elevation from ca. 15 m at 3,650 m a.s.l. to some 3 m at the treeline (Fig. 1 a). This trend was even more pronounced if only *Polylepis* trees were considered: *P. lanata* formed an upper tree layer reaching a height of 30 or 32 meters, while the second stratum (10 m height) was mainly built up by *Oreopanax macrocephalus*.

Mean diameter at breast height (dbh) also decreased significantly with elevation (Fig. 1 b). In correspondence

to tree height, single individuals of *P. lanata* at 3,650 m a.s.l. showed dbh values of e.g. 112 or 152 cm, while dbh values at the treeline ranged between 3 and 23 cm.

Stem density increased linearly with elevation by a factor of six (Table 3). Due to the non-linear decrease in stem diameter with elevation, total basal area decreased from 176 m² ha^{–1} in the lowermost stand to ca. 67 m² ha^{–1} at the treeline (Table 3).

Leaf morphological and chemical status

The mean leaf area of *Polylepis* trees at the lowermost site was markedly larger than at the two uppermost stands, but this seemed to be a result of the difference between the two *Polylepis* species rather than an elevational effect. However, there was a tendency of decreasing leaf areas in *P. pepeii* with elevation (Table 4). This is confirmed by a significant reduction in specific leaf area from the middle stand towards the treeline, while *P. lanata* trees at the lowermost site had a much lower specific leaf area compared to *P. pepeii* (Table 4).

Leaf nitrogen concentrations (on a dry matter basis) increased significantly with elevation irrespective of the different species. Expressed on a leaf area unit, *P. lanata* at the lowermost site showed the highest N content, while *P. pepeii* leaves at the treeline site had significantly higher N contents per unit leaf area compared to the intermediate site (Table 4). Leaf phosphorus concentration was significantly higher in *P. lanata* compared to *P. pepeii* based on a dry matter unit as well as on a leaf area unit (Table 4). Leaf phosphorus concentration of *P. pepeii* also tended to increase with elevation, but differences were not significant (Table 4).

Analyses of the carbon isotopic status of the leaves indicated no marked differences among stands: the $\delta^{13}\text{C}$ values in *P. pepeii* were not significantly different between

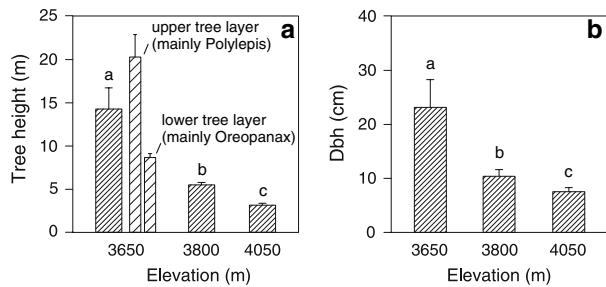


Fig. 1 Tree height and diameter at breast height (dbh) of the three *Polylepis* stands. Given are means \pm 1 SE; $n = 12$ for mean tree height and $n = 61, 75,$ and 62 for dbh values at 3,650; 3,800; and 4,050 m a.s.l., respectively. Different letters indicate significant differences between the forests. In the case of the lowermost stand the narrow bars beside indicate tree height means of the two occurring canopy strata

Table 3 Stem density and basal area of the three *Polylepis* forests

Elevation (m)	3,650	3,800	4,050
Stem density (n ha ⁻¹)	2,033	7,500	12,400
Basal area (m ² ha ⁻¹)	176.0	86.3	67.3

the two uppermost stands. *P. lanata* showed a somewhat higher (less negative) value compared to *P. pepei* (Table 4).

Below-ground structure and fine root morphology

The total mass of living fine roots (fine root biomass) of the stands increased significantly with elevation: the middle and the uppermost stand exceeded the fine root biomass of the lowermost stand by factors of 4.9 and 6.3, respectively (Fig. 2 a). This was due to a significant altitudinal increase in fine root density (biomass per soil volume) in both the organic layer and the upper mineral soil (0–10 cm; Table 5). The density of fine root necromass (dead fine

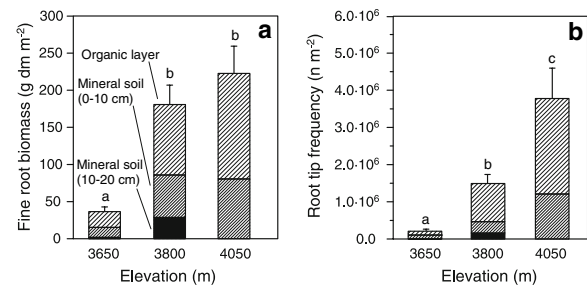


Fig. 2 Fine root biomass and root tip frequency of the *Polylepis* stands. Given are means \pm 1 SE; $n = 20$. Different letters indicate significant differences between the forests in terms of profile totals

roots) in the mineral soil also increased significantly with increasing elevation (Table 5), leading to a significantly higher standing fine root necromass at the two uppermost sites (data not shown).

Fine root morphology also changed along the altitudinal transect: the specific root tip abundance (number of tips per unit dry mass) increased linearly with elevation by a factor of ca. 2 in the organic layer (Table 5). Moreover, a significant increase in specific root tip abundance in the upper mineral soil occurred between the two *P. pepei* stands. Differences in specific fine root area SRA could also be detected among the stands: *P. lanata* fine roots at the lowermost site showed comparatively high SRA values; however, SRA increased significantly from the intermediate *P. pepei* stand towards the treeline site, where the overall maximum values were found. Fine roots of *P. pepei* at the treeline site had significantly smaller diameters compared to the two stands of *P. pepei* and *P. lanata* below the treeline.

As a result of both, changes in fine root biomass and root morphology, root tip frequency (number of root tips per unit ground area) increased conspicuously from the lowermost to the uppermost stand: there were 18 times more root tips per square metre ground area at the treeline stand compared to the lowermost site, even though the soil was much shallower in the uppermost stand (Fig. 2 b). The fine

Table 4 Structural and chemical leaf traits of the *Polylepis* trees in the three forests

Elevation (m)	3,650	3,800	4,050
Single leaf area (cm ²)	11.98 \pm 0.46 a	1.76 \pm 0.06 b	1.62 \pm 0.05 b
Specific leaf area (cm ² g dm ⁻¹)	66.0 \pm 6.1 a	112.7 \pm 1.7 b	86.0 \pm 3.1 c
Nitrogen concentration (mmol N g dm ⁻¹)	1.20 \pm 0.08 a	1.31 \pm 0.02 ab	1.50 \pm 0.07 b
(μ mol N cm ⁻²)	18.69 \pm 1.52 a	11.66 \pm 0.18 b	17.49 \pm 0.74 a
Phosphorus concentration (μ mol P g dm ⁻¹)	86.7 \pm 6.2 a	58.7 \pm 2.8 b	67.3 \pm 2.9 b
(μ mol P cm ⁻²)	1.35 \pm 0.11 a	0.52 \pm 0.02 b	0.79 \pm 0.04 b
$\delta^{13}\text{C}$ signature (‰)	-26.63 \pm 0.46 a	-26.91 \pm 0.33 a	-26.78 \pm 0.27 a

Given are means and standard errors. Different Latin letters indicate significant differences among the stands ($P < 0.05$, $n = 60$ for single leaf area, and $n = 6$ for the other parameters)

Table 5 Structural, morphological, and chemical traits of the fine root system in the three *Polylepis* forests

Elevation (m)		3,650	3,800	4,050
Fine root biomass density (g dm ⁻¹)	Organic layer	0.22 ± 0.04 a α	0.90 ± 0.10 b α	1.00 ± 0.13 b α
	Mineral soil (0–10 cm)	0.21 ± 0.04 a α	0.84 ± 0.14 b α	3.05 ± 0.52 c β
	Mineral soil (10–20 cm)	0.12 ± 0.05 a α	0.21 ± 0.07 a β	–
Fine root necromass density (g dm ⁻¹)	Organic layer	0.71 ± 0.23 a α	0.80 ± 0.14 a α	0.71 ± 0.07 a α
	Mineral soil (0–10 cm)	0.48 ± 0.04 a α	1.49 ± 0.39 a αβ	2.16 ± 0.16 b β
	Mineral soil (10–20 cm)	0.35 ± 0.08 a α	2.01 ± 0.50 a β	–
Specific fine root surface area (cm ² g dm ⁻¹)	Organic layer	361 ± 37 a α	368 ± 44 a α	584 ± 83 b α
	Mineral soil (0–10 cm)	365 ± 24 a α	251 ± 31 b β	363 ± 38 a β
	Mineral soil (10–20 cm)	365 ± 40 a α	362 ± 47 a β	–
Fine root area index (m ² m ⁻²)	Organic layer	0.67 ± 0.11 a α	3.30 ± 0.37 b α	6.63 ± 1.26 c α
	Mineral soil (0–10 cm)	0.48 ± 0.14 a α	1.21 ± 0.26 b β	2.98 ± 0.90 ab β
	Mineral soil (10–20 cm)	0.01 ± 0.01 a β	0.89 ± 0.40 b χ	–
	Profile total	1.15 ± 0.20 a	5.40 ± 0.78 b	9.61 ± 1.60 c
Specific root tip abundance (number mg dm ⁻¹)	Organic layer	5.6 ± 1.0 a α	12.0 ± 2.3 b α	25.5 ± 5.9 c α
	Mineral soil (0–10 cm)	7.7 ± 1.3 ab α	7.0 ± 1.8 a β	13.2 ± 2.5 b α
	Mineral soil (10–20 cm)	13.3 ± 11.7 a α	8.0 ± 2.0 a αβ	–
Mean fine root diameter (mm)	Organic layer	30.7 ± 1.7 a α	29.2 ± 1.5 a α	21.8 ± 0.7 b α
	Mineral soil (0–10 cm)	28.2 ± 1.0 a α	30.9 ± 2.0 a α	24.1 ± 1.5 b α
	Mineral soil (10–20 cm)	26.3 ± 3.8 a α	30.2 ± 4.1 a β	–
δ ¹³ C signature (‰)	Organic layer	–26.41 ± 0.16 a α	–28.28 ± 0.09 b α	–28.47 ± 0.11 b α
	Mineral soil (0–10 cm)	–26.19 ± 0.17 a α	–28.03 ± 0.18 b α	–28.42 ± 0.06 b α
	Profile mean	–26.33 ± 0.12 a	–28.12 ± 0.10 b	–28.45 ± 0.07 c

Given are means and standard errors. Different Latin or Greek letters indicate significant differences among the stands or among horizons, respectively ($P < 0.05$, $n = 20$ for fine root density and fine root area index, $n = 17–20$, $12–16$, and $2–7$ for specific surface area and specific root tip abundance, and fine root diameter in the organic layer, and the upper and lower mineral soil horizon, and $n = 23–40$ for the C isotopic signature)

root area index RAI (root surface area per unit ground area) increased by a factor of more than 8 from the lowermost to the uppermost stand (Table 5).

In contrast to leaf carbon isotopic signature, δ¹³C value of the fine root material decreased significantly with elevation (i.e. more negative values at higher elevation). Within a stand, fine roots in the organic layer tended to have a more negative δ¹³C value compared to those in the mineral soil (Table 5).

Young growth of the *Polylepis* stands

Density of young growth increased with altitude from a maximum of 2 shoots 4 m⁻² at the lower elevation to a maximum of 1,972 shoots at the uppermost stand (Fig. 3). Young growth at the two lower sites was not connected to adult trees, and tended to be more common in the grassland. Variance was generally large rendering differences between grasslands and forest not significant for the two lower sites. The pattern reversed at the upper plots where a mean of 1,091 shoots 4 m⁻² grew inside the stands, while

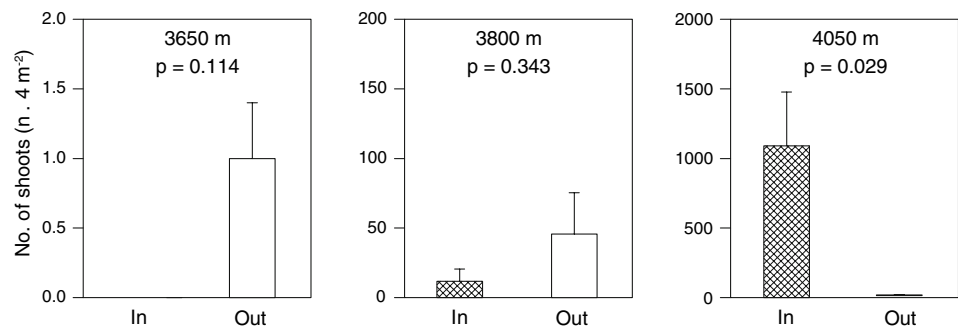
the mean density in the grasslands was 4 shoots 4 m⁻² (Fig. 3). Digging revealed that all shoots inside the upper forest stand were interconnected, and eventually could be traced back to larger roots of adult trees. Thus, *P. pepei* at the uppermost stand was capable of clonal growth by developing root suckers. Shoots in the open grassland also had long roots and may also well have originated from root suckers.

Discussion

Thermal conditions of the *Polylepis* forests at the treeline

Several thresholds were proposed for temperatures at the alpine treeline. These include the 10°C July air temperature isotherm for the temperate and boreal region (e.g. Köppen 1931) or the 3–6°C summer (growing season) air temperature isotherm at tropical or subtropical mountains (Körner 1998; Körner 2003). Recent data imply that tree-line positions are more closely associated with soil than with air

Fig. 3 Number of small shoots (<2 cm dbh) inside and outside the *Polylepis* forests. Given are means \pm 1 SE from four randomly selected subplots (4 m²) per altitude and forest or grassland site. Probability values refer to the analysis of differences between forest and adjacent grassland location



temperatures (e.g. Körner and Hoch 2006). In their global survey Körner and Paulsen (2004) found a coincidence of the treeline position with the 6.7°C growing season isotherm in the root zone (10 cm soil depth), while mean growing season temperatures in the root zone in the tropical/subtropical mountains were lower at 5 to 6°C. Accordingly, Hoch and Körner (2005) recorded a mean root zone temperature of 5.4°C in the world's highest forests formed by *Polylepis tarapacana* at the volcano Sajama in the western (dry) cordillera of Bolivia. Our own measurements at the treeline of *P. pepei* in the eastern (moist) cordillera of Bolivia point to lower values (2.0°C root zone temperature, 1.7°C air temperature), but data were collected during a short-term field campaign only. Relatively low temperatures are also reported by Kessler and Hohnwald (1998) for a very similar treeline site formed by *P. pepei* at 4,050 m a.s.l. in the eastern (moist) cordillera north of La Paz (Bolivia), where the mean soil temperature in 5 cm soil depth was also only 2.5°C. Bendix and Rafiqpoor (2001) also found relatively low mean soil temperature values (at 10 cm depth) in stands of *P. incana* and *P. pauta* at the treeline in Ecuador. In Bolivia, treelines formed by *P. tarapacana* grew at mean air temperature between 3.4 and 3.7°C (Rada et al. 2001; Garcia-Nunez et al. 2004) and values reported for the uppermost stands of *P. sericea* in Venezuela (Rada et al. 1996) are also at the lower range of temperature values given by Körner (1998) and Körner and Paulsen (2004) for the tropical regions worldwide. These records suggest that treeline stands built by *Polylepis* trees seem to occur under specifically cold conditions.

Changes in below-ground vs. above-ground carbon sequestration with altitude

The most conspicuous finding in our study was the contrasting gradients in above-ground and below-ground structure of the three *Polylepis* stands along the elevational transect. Decreasing tree height, stem diameter, and stand basal area together with increasing stem density indicated a

clear reduction of carbon allocation to above-ground tree compartments, while the size (biomass) and surface area (RAI) of the fine root system increased markedly towards the treeline.

The change in above-ground structure with elevation is a well known phenomenon. Decreasing C investment at high altitudes in above-ground tree compartments were traditionally attributed to a reduced carbon gain due to a limitation in photosynthetic activity and/or an imbalance of annual C fixation vs. respiratory losses (data reviewed by Stevens and Fox 1991; Körner 1998; Grace et al. 2002; Körner 2003a). However, several studies on C assimilation failed to give evidence that tree growth at high elevation is limited by carbon gain (e.g. Slatyer 1978; Häsler 1982; Brodersen et al. 2006; see also review by Körner 2003b). Instead, Körner and co-workers emphasized an alternative hypothesis of sink limitation due to reduced biosynthetic activity under cold growth conditions (Hoch et al. 2002; Hoch and Körner 2003; Hoch and Körner 2005; Shi et al. 2006). Specific hypotheses related to tree-stature, on the other hand, holding for damaging effects due to blowing ice, large snow charge, or winter drought are based on data from high elevation forests of the temperate, boreal, and arctic/antarctic regions (Stevens and Fox 1991; Sveinbjörnsson 2000) but cannot explain the altitudinal trends found in tropical mountains.

For *Polylepis* trees it has been shown, that above-ground growth at high elevations should not be limited by assimilate shortage: measurements by Rada et al. (1996) on *P. sericea* as well as by Rada et al. (2001) and Garcia-Nunez et al. (2004) on *P. tarapacana* indicated sufficient carbon gains even at very high altitudes leading to high concentrations of non-structural carbohydrates (NSC) in above-ground plant tissues (Hoch and Körner 2005) and high frost tolerance (Rada et al. 2001). Though some *Polylepis* species seem to form small trees only or even shrubs (Kessler 1995), there are reports of relatively high above-ground biomass accumulation and productivity of *Polylepis* stands not far below the expected climatic tree-line position (Young 1993; Fehse et al. 2002). Thus, our finding of tall *Polylepis* trees with large diameters only

400 m below the treeline in E Bolivia are in general agreement with available data, although heights of 30–32 m have never, and dbh > 100 cm have rarely been described in the literature for *Polylepis* so far (e.g., Young 1993). Marked altitudinal decreases in tree height of *Polylepis* stands were reported in a number of studies (Young 1993), and 2–4 m in height seems to be typical for *Polylepis* forests at the treeline (Rada et al. 1996, Garcia-Nunez et al. 2004, Hoch and Körner 2005).

The wealth of reports on elevational changes in above-ground stand structure is in marked contrast to the scarcity of data on changes in the root system. Luo et al. (2005) found a reduction in biomass of roots along an elevational transect in the Himalayas. However, this investigation differs from our approach by the inclusion of root fractions much larger than 2 mm in diameter and by the tree individual-based sampling procedure. The few other studies on elevational effects on the fine root system, on the other hand, also point to increasing fine root biomass with elevation on the stand area level. Hence, in moist tropical forests in Queensland (Australia) Maycock and Congdon (2000) found a 2.5 times larger fine root biomass average in stands of 600–1,000 m a.s.l. than in stands between 50 and 120 m a.s.l. A similar result was found by McGroddy and Silver (2000) for an elevational transect of tropical moist forests in Puerto Rico. However, both studies covered only lower elevation forests. More directly comparable are data from two extended altitudinal transects on forest structure and productivity at Mt. Kinabalu (Malaysia) by Kitayama and Aiba (2002): our re-analysis of that data set revealed a significant linear increase in fine root biomass along the two transects between 650 and 3,080 m a.s.l. ($r^2 = 0.74$). The results are similar to findings of Leuschner et al. (2007) for a tropical montane forest in South Ecuador, where fine root biomass increased by a factor of >4 from 1,050 to 3,060 m a.s.l. None of the described transect studies, however, reached up to the climatic treeline. In fact, we know of no study on fine root system of trees at the treeline.

Despite that, it is noteworthy that the absolute amounts in fine root biomass of the *Polylepis* stands lay far below the values of other high elevation forests. With 40–230 g m⁻² fine root biomass was markedly lower than the global average value of ca. 1,000 g m⁻² given for stands above 2,000 m a.s.l. in a recent review of the available literature data on fine root biomass values in tropical moist forests (Hertel and Leuschner 2008). Whether these relatively low amounts of fine root biomass are typical for *Polylepis* forests in the Andes remains open since there are no other comparable data on the root system of *Polylepis* forests. Nevertheless, the observed increase in fine root biomass and fine root system surface area (RAI) clearly follows the altitudinal trends of the fine root system in

tropical moist montane forests (review by Hertel and Leuschner 2008).

What possible reasons might be responsible for this increased carbon investment into the fine root system with elevation towards the treeline? According to general growth economy theory (Givnish 1984) the enlargement of the fine root system should reflect a compensation for water or nutrient deficiency. It was shown that high elevation forests can be exposed to severe drought at least periodically (James et al. 1994; Lloyd and Fastie 2002; Li et al. 2004), and those dry conditions have been reported for *Polylepis* treeline stands as well (Garcia-Nunez et al. 2004; Morales et al. 2004). However, water limitation is unlikely to be an important factor at our study transect in the humid eastern Andes cordillera. Consequently, we could not detect any significant differences in leaf $\delta^{13}\text{C}$ values among trees from different altitudes.

In fact, soils of tropical moist forests at high elevation are more often water-saturated than dry (Vitousek and Sanford 1986; Kappelle 2004). In these cold environments, this causes low pH values, oxygen depletion, and low microbial activity (Bruijnzeel and Veneklaas 1998; Silver et al. 1999), leading to a reduction in nutrient availability (Meentemeyer 1977; Vitousek and Sanford 1986). Hence, forest growth at such cold sites may be limited by the supply of nutrients, especially of nitrogen (Vitousek and Sanford 1986; Sveinbjörnsson et al. 1992; Timoney 1995; Tanner et al. 1998). Accordingly, the increase in the size of the fine root system towards the treeline might represent an adaptation to lowered nutrient availability.

Although our soil chemical analyses suggest a decrease in nutrient availability with elevation, the analyses of leaf N and P contents in our study do not seem to support the idea that the *Polylepis* trees at the treeline did experience severe nutrient deficiency. We found a moderate decrease in leaf P concentration, but a slight increase in N content (expressed on a dry weight basis) from the lowermost to the uppermost stand. This is in contrast to the finding of a weak decrease in leaf N content with elevation along an elevational transect in Tibet with relatively low N values (1.2% dm) at the treeline (Shi et al. 2006), or an even strong reduction in leaf N concentration (from 2.1 to 1.1% dm) with elevation in South Ecuadorian mountain forests (Schuldt, unpublished data). De Lillis et al. (2004), on the other hand, found an increase in leaf N content (from 1.4 to 2.1% dm) with elevation towards the treeline similar to our results. It has been stated, that high concentrations in leaf photochemical compounds might be a physiological adaptation to low temperatures (Velez et al. 1998), and thus high leaf N contents may be a consequence of physiological adaptation to cold growth conditions rather than reflecting possible nutrient limitation in the soil (Weih and Karlsson 2001). However, the N values in our study can be

judged as much lower than N contents of 3.0–3.5% found in *Polylepis quadrijuga* trees at a lower elevation of ca. 3,200 m a.s.l. in Colombia (Velez et al. 1998). Hence, it remains unclear if the N and P leaf contents may allow for identification of possible nutrient shortage in the case of our study.

Another possible explanation for the conspicuous shift in carbon allocation towards the root system with elevation could be found in a hypothesis given by Stevens and Fox (1991). These authors state that under cold growth conditions the long-distance transport of water and nutrients in trees requires a disproportional large carbon investment in the root system compared to warmer environments. Constraints in water and related nutrient uptake from soil to root and across root cells at low temperatures may have purely physical reasons (Stevens and Fox 1991; Sveinbjörnsson 2000). This line of thought would explain the obvious enlargement of the fine root system with elevation of the *Polylepis* stands towards the treeline, and also the marked decrease in shoot:root ratio illustrated by the sharp decrease in tree height and basal area. Indeed, Karlsson and Nordell (1996) and Weih and Karlsson (2001) demonstrated that N uptake of birch seedlings under cold conditions might not only be affected by low nutrient availability, but also by reduced nutrient supply due to the physiological inability of root nutrient uptake. Thus, the observed changes in the shoot:root ratio in *Polylepis* forests with elevation seem to follow the general trend described by Hertel and Leuschner (2008) of marked increase in carbon investment to the root system of tropical moist forests at high altitudes to serve for sufficient nutrient supply.

Elevational changes in leaf and root morphology

Our altitudinal transect study revealed not only marked changes in carbon sequestration but also significant differences in leaf and root morphology between the *Polylepis* trees of different elevation. Much of these adaptations in leaf morphology to high-altitude climatic conditions are well-known: leaves at higher elevations are smaller and have lower specific leaf area (SLA) values (e.g. Cordell et al. 1998; Aiba and Kitayama 1999; Li et al. 2004). These traits are interpreted as adaptations to harsh conditions at high altitudes such as increasing wind speed, low temperatures, low nutrient availability, high UV-B radiation, and, locally, high vapor pressure deficits (e.g., Box 1981; Chabot and Hicks 1982; Givnish 1987; Reich et al. 1992; Richards 1996). Not all these factors are likely to operate in the forests described here. Dry growth conditions at the treeline in the central cordillera in Bolivia at volcano Sajama are related to very low SLA values of ca. 30 cm² g⁻¹

of *Polylepis tarapacana* (Hoch and Körner 2005), while the SLAs of our *Polylepis* trees growing in the humid, eastern cordillera region ranged between 66 and 113 cm² g⁻¹. Moreover, differences between the leaf morphology data in our elevational transect are most likely biased by the presence of two different tree species. However, leaf morphology of the two upper *P. pepei* stands clearly follows the general tendency described above with lower leaf area and significantly lower SLA at the treeline site.

Fine root morphology, on the other hand, showed the opposite trend: specific surface area was highest while fine root diameter was smallest at the treeline, and specific root tip abundance increased markedly with elevation. Thus, not only increasing values of fine root biomass and fine root density but also the increase in root surface area (SRA) and root tip abundance highlights the enlargement of the absorbing parts of the fine root system towards the treeline. Like the changes in fine root biomass, the increase in SRA values and specific root tip abundance likely represents an adaptation to either decreasing nutrient availability from the soil or a physiological constraint in nutrient uptake due to low temperature conditions. Comparable data from high elevation forests in the tropics are missing, but studies on fine root morphology in temperate forests suggest that increased root surface area and root tip number result in increased nutrient uptake per unit root mass. Hence, Pregitzer et al. (1998) showed that root metabolic activity (here: respiration) was higher in fine rootlets (<0.5 mm in diameter) compared to roots of larger diameter. Pregitzer et al. (2002) and Guo et al. (2004) emphasize the meaning of small-diameter roots of different North American tree species for nutrient uptake. Small roots, on the other hand, were found to be more susceptible to root mortality under cold growth conditions (Wells and Eissenstatt 2001). Thus, an increase in size of the absorbing organs of the fine root system may also represent a compensation for a lower degree in mycorrhizal infection at the high elevation due to higher root mortality (Sveinbjörnsson 2000). We conclude, that the observed changes in fine root morphology of the *Polylepis* trees towards the treeline can be seen as an economic mechanism to save carbon costs of the increasing size of the fine root system with elevation. This is especially true comparing the two uppermost *P. pepei* stands: the observed increase in fine root area index (RAI) and root tip frequency (numbers m⁻²) of the stands was higher by far (factor > 2) than the concurrent moderate increase in fine root biomass (factor < 0.3).

Regeneration under different microclimate conditions

Regeneration pattern of the studied *Polylepis* trees changed markedly along the elevational transect in three ways: (1)

number of young growth increased with elevation, (2) seedlings were mainly found in the grassland at the lower stands, while young shoots were almost exclusively restricted to the forest interior at the uppermost forest, and (3) young plants at the two lowermost sites were actually seedlings while young growth at the uppermost site were exclusively comprised of root suckers.

Several authors describe constraints in reproduction under the harsh microclimatic conditions at the treeline. Seed rain density and seed quality may decrease with elevation leading to a decrease in seedling establishment at the treeline (Sveinbjörnsson 2000). The occurrence of strong winds as well as frost or summer drought effects hamper growth of seedling (Daniels and Veblen 2003; Piper et al. 2006; Resler 2006). Apparently, differences in microclimatic conditions in- and outside the forest stands are of pivotal importance: in a study on *Nothofagus pumilio* in South Argentina Cuevas (2000) found that seedling emergence at the treeline was much lower in the open grassland compared to the forest interior. This was attributed to differences in seed rain, but also to constraints in seedling survival under open conditions. Seeding experiments on *Polylepis* forests at the treeline in Ecuador showed that microclimatic conditions hamper seedling emergence in open grassland at the treeline compared to the forest edge or its interior (Cierjacks et al. 2007). This was related to the special microclimatic conditions near the soil surface: like in our study, mean air and soil temperatures near the ground were slightly cooler inside the uppermost forest, while air temperature in the grassland outside the treeline forest showed markedly higher maximum and lower minimum temperatures compared to the forest conditions. Similar results were found by Wang et al. (2006) for *Picea schrenkiana* growing at the treeline at Tianshan, China, where seedling recruitment was favoured at sites with relatively high monthly minimum temperatures compared to sites with colder minima. Additionally, vapor pressure deficits (VPD) at our treeline site in Bolivia reached also much more extreme values (by a factor of 3) in the grassland than inside the treeline forest (20 cm above the ground), and drought stress has also been reported to affect tree seedlings' vitality (Camarero and Gutierrez 2004; Cuevas 2003; Daniels and Veblen 2004). Thus, there is evidence that extreme events rather than mean values may be important for tree regeneration, which corresponds to the observations on *Polylepis* in our study.

Climatic conditions at the treeline also influence the adult tree and may affect fruit production and/or trigger a shift to clonal growth. The observed change in recruitment strategies towards clonal growth is in line with a range of studies from harsh environments including treeline ecotones (Holtmeier 2003). *Polylepis* spec. at the treeline in

Ecuador showed altitudinal trends in reproductive activity and flower and fruit production dropped to zero at the current treeline (Cierjacks et al. 2008). However, the decisive bottlenecks in recruitment may occur at other stages such as the transition from seedling to sapling stage rather than the establishment of seedlings (Stevens and Fox 1991; Sveinbjörnsson 2000; Hoch and Körner 2005). Thus, a range of factors can influence recruitment in a broad sense and it remains open whether the observed patterns in *Polylepis* regeneration indicate a specific mechanism of treeline formation or just a general response to harsh conditions.

Conclusions

Our study of high-Andean *Polylepis* forests in Bolivia reveals elevational changes of various ecological parameters. While some of these findings cannot be clearly rated in their importance for the formation of the *Polylepis* tree-line (e.g., the alteration of the regeneration pattern), the contrasting changes in above- versus below-ground attributes are likely to reflect systematic responses of *Polylepis* trees to deterioration in both thermal conditions and nutrient supply near the treeline. We conclude therefore that the marked shift in carbon allocation towards the fine root system as well as the significant enlargement of the surface of the fine root system (in terms of a higher surface area and a larger number of root tips) via changes in fine root morphology represent crucial adaptations of *Polylepis* trees to the harsh environmental conditions at the treeline.

Acknowledgments This study was funded by the German Research Foundation (DFG, project He 3582/1-1). We gratefully acknowledge the invaluable help of Jaime Hamel, Chochabamba, Bolivia, during the field investigations in Bolivia as well as Mechthild Stange for her skilful support of the fine root analyses. We thank Prof. Dr. Michael Kessler for his constructive comments on our manuscript.

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