

Variable impact of late-Quaternary megafaunal extinction in causing ecological state shifts in North and South America

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Loss of megafauna, an aspect of defaunation, can precipitate many ecological changes over short time scales. We examine whether megafauna loss can also explain features of lasting ecological state shifts that occurred as the Pleistocene gave way to the Holocene. We compare ecological impacts of late-Quaternary megafauna extinction in five American regions: southwestern Patagonia, the Pampas, northeastern United States, northwestern United States, and Berin-gia. We find that major ecological state shifts were consistent with expectations of defaunation in North American sites but not in South American ones. The differential responses highlight two factors necessary for defaunation to trigger lasting ecological state shifts discernable in the fossil record: (i) lost megafauna need to have been effective ecosystem engineers, like proboscideans; and (ii) historical contingencies must have provided the ecosystem with plant species likely to respond to megafaunal loss. These findings help in identifying modern ecosystems that are most at risk for disappearing should current pressures on the ecosystems' large animals continue and highlight the critical role of both individual species ecologies and ecosystem context in predicting the lasting impacts of defaunation currently underway.

megafauna | extinction | Quaternary | North America | South America

Defaunation is occurring at a rapid pace presently (1–3). Losses are particularly severe for megafauna (1) (considered here as animals with an average body size ≥ 44 kg), whose removal can trigger the following: changes in vegetation structure and species composition; reductions in environmental heterogeneity, species richness, evenness, seed dispersal, nutrient cycling and distribution, and ecosystem services; coextinction of dependent species; and increases in disease-transmitting organisms (1, 4–14) and fire frequency and/or intensity (15–17).

Most work on defaunation has been in contemporary ecosystems. Much less is known about how it manifests over millennial time scales. A natural experiment to assess lasting effects of megafauna loss is provided by the extinctions of late-Quaternary megafauna in the Americas, part of global-scale ecological state shift (18), during which about half of the world's large-bodied mammal species (19, 20) disappeared. In North America, ~60 megafaunal species died out, with the youngest occurrences of dated species typically falling between ~13,000 and 11,000 y ago (19). In South America, ~66 species were lost over a longer time span (21–23).

With a few important exceptions (6, 17, 24–29), the major changes in vegetation and mammalian community structure that accompanied Quaternary extinctions have been interpreted as responses to changing climate (17–19, 21, 23, 25–27, 29–35). Here, we build on recent work of paleoecologists (17, 25, 28, 29, 32, 36) and ecologists (1, 3–7, 9, 10, 15, 16, 37) who have been asking instead: Are the observed biotic responses consistent with

megafauna loss, and if so, what does this loss imply for the future of ecosystems at risk for losing their megafauna today?

Approach

The late-Quaternary impact of losing 70–80% of the megafauna genera in the Americas (19) would be expected to trigger biotic transitions that would be recognizable in the fossil record in at least two respects. First, vegetation should change noticeably, consistent with ecological release from browsing, grazing, and trampling; such changes should be apparent in fossil-pollen time series (17, 25, 29) and possibly also in charcoal (fire-frequency) records (15–17, 36). Second, surviving mammal communities should demonstrate changes in species composition, richness, and evenness (32). Testing for such impacts requires fossil sites with the following: (i) good records of extinct megafauna; (ii) proximal and temporally overlapping records of vegetation change and fire frequency; (iii) associated records of species that allow assessment of mammalian diversity; (iv) robust dating of megafauna population crashes and extinction and vegetation, fire, and diversity changes;

Significance

Removing megafauna from contemporary ecosystems changes vegetation and small mammal communities over ecological time scales. We show that similar dynamics seem to operate over millennial time scales but only if the megafaunal loss includes ecosystem engineers in settings that also contain plant species susceptible to ecological release. Under such conditions, megafauna extinction can initiate changes that quickly lead to new, lasting ecological states. This implies that should some megafauna currently at risk for extinction actually become extinct, their characteristic ecosystems—for example, mosaics of savannah and forest—would also disappear, rapidly transforming into novel systems with respect to what is considered normal in today's world.

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and (v) a fossil record that temporally samples before and after the extinctions.

Here, we examine evidence from five regions where paleontological records are adequate in at least most of these respects: southwestern Patagonia, the Pampas, northeastern United States, northwestern United States, and the Alaska/Yukon area (Beringia) (Fig. 1). Each region featured different megafauna and/or vegetation, and each provides relatively rich datasets within constrained biogeographic settings (see [Supporting Information](#) for the South American sites). If defaunation triggers ecological state shifts universally, in all areas, we would expect to see as megafauna go extinct: (i) increases in understory and forest plants susceptible to ecological release in the absence of herbivory, trampling, and related pressures, potentially with increased fire frequency; and (ii) predictable changes in surviving communities of mammal species, such as adjustments in geographic range, species density, and diversity. Disentangling defaunation impacts from those triggered by end-Pleistocene climate change requires observing a discordance between the timing of defaunation and climate change proxies with respect to the hypothesized biotic responses, and/or by observing biotic indicators (for example, of vegetation change or fire frequency) that make more sense as the result of a defaunation rather than from a climatic trigger. The relative contribution of human activity and climate change in causing the late-Quaternary defaunation events is still debated (19). However, our goal here is simply to understand whether or not the chronology of megafauna extinction, changes in vegetation and fire records, and changes in mammal species diversity support or reject the expectations of ecological changes triggered by megafauna extinction, regardless of the ultimate cause of defaunation. We emphasize that, given the limitations of the fossil record, only the coarsest effects of defaunation—such as vegetation and diversity changes we examine here—can be recognized by our approach.

Vegetational Changes

South America—Southwestern Patagonia. Recent work in the Última Esperanza region in Chile has produced a detailed chronology of colonization of the region by humans and megafauna, megafauna extinction, and vegetation, fire, and climatic changes for the period 19,000 y ago to 5,000 y ago (23) (Figs. S1–S4). There, *Lama gracilis* and *Vicugna vicugna* [these two taxa may be synonymous (38)] became locally extinct about 15,000 y ago. Humans arrived in the area by 13,750 y ago. The extinction of megacarnivores (*Smilodon* and *Panthera*), ~1,100 and 1,700 y

later, respectively, occurred just after the Antarctic Cold Reversal, during the beginning of climatic warming that apparently initiated a shift from cold grassland vegetation to *Nothofagus* forest (23, 39) (Fig. S4). *Hippidion saldiasi* and a llama (*Lama* cf. *L. owenii* of past publications) disappeared ~2,400 y later, as transition to *Nothofagus* continued. At about the same time, fire frequency also increased, consistent with warming temperatures and increased fuel provided by the encroachment of trees. It may be that extinction of the equid and llama also helped to increase fire fuels (Fig. S4) and promote expansion of the *Nothofagus* forests, which had arrived in the region at least 500 y earlier (Figs. S2–S4). By ~11,000 y ago, *Nothofagus* forests were firmly established, and it is only after that (most likely near ~10,100 y ago) that *Mylodon darwini* goes extinct. The overall pattern of herbivore extinction (Fig. S4) is equivocal with respect to defaunation triggering some of the observed vegetation and fire changes (23). Although loss of *Hippidion* and *L. cf. L. owenii* potentially corresponds with increasing forest cover and fire frequency, these changes also could be explained entirely by contemporaneous climate changes in the region. The extinction of *Mylodon* does not correlate with any major vegetation or fire changes and comes only after *Nothofagus* forest is well established. Dung samples of *Mylodon* confirm that it primarily grazed in open landscapes (40) (e.g., specializing on grasses), suggesting its loss would have little impact on woody vegetation (23).

South America—Pampas. The Pampas region (here considered bounded by 40° S, 60' W; 35° S, 65' W; 25° S, 55' W; 28° S, 45' W; following ref. 41) (Fig. S5) is located between tropical Brazil and the cold dry Southern Cone. The region contains ecotones that shifted as late-Quaternary climate changed (42) such that the timing of the Pleistocene-to-Holocene vegetation transitions varied across the study area over ~2,000 y, beginning about 12,000 y ago (Fig. 2) (43–45). Although different parts of the region featured locally distinctive vegetation regimes (see [Supporting Information](#) for more detail), in general, the end-Pleistocene Pampas was a C3-dominated grassland steppe. In the Holocene, C4 vegetation increased, as did seasonal variation in proportions of C3 versus C4 grasses (43–45). Overall, the increase in C4 grasses (46) and other vegetational changes ([Supporting Information](#)) are consistent with late-Quaternary climatic warming driving a transition to warmer, more humid conditions, as also indicated by sedimentology (47), malacological remains (48), and paleobotanical data (43–45, 49).

Most likely well after the vegetation changed (Fig. 2), proboscideans (*Notiomastodon*), horses (*Equus* and *Hippidion*), a llama (*Palaeolama*), and meridiungulates (*Toxodon* and *Macrauchenia*) disappear from the record. Xenarthrans (ground sloths and glyptodonts) go extinct even later (Fig. 2). Based on morphological analyses and isotopic data from the Pampas region, *Notiomastodon* and *Macrauchenia* were mixed feeders that consumed both C3 shrubs and C4 grasses; *Equus* and *Toxodon* were grazers, and *Hippidion* consumed primarily C3 vegetation (50). Generally, mammalian herbivores prefer C3 grasses when they are available (46), and in Uruguay today, grazing promotes the dominance of C4 grasses where both C3 and C4 grasses are present, although the effect is variable; C3 and C4 grasses both are heavily used but in different seasons (51). These considerations—in addition to the fact that the grazers and the xenarthrans likely became extinct only well after the vegetation transition began (Fig. 2)—suggest that climate change, not defaunation, was the primary cause of the C3 to C4 transition.

There is no evidence that loss of *Notiomastodon* significantly impacted the amount of shrubby vegetation in the dry southern pampas (Fig. S5), which actually retracted in range across the Pleistocene–Holocene transition (43, 48). However, many undated *Notiomastodon* specimens (52) come from northern Uruguay and southern Brazil (Fig. S5), where locally gallery forests along rivers and tropical woodlands expanded somewhat at the

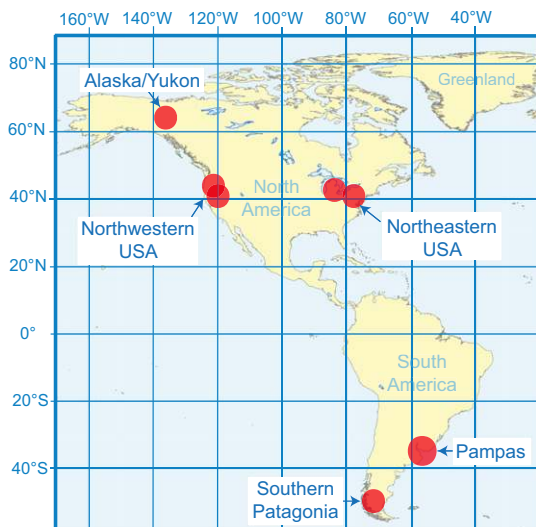


Fig. 1. Areas examined for impacts of late-Quaternary defaunation.

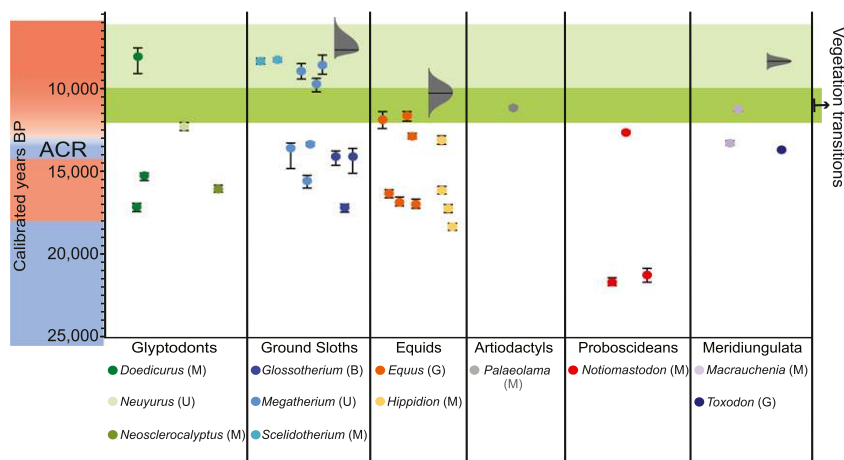


Fig. 2. Chronology of megafaunal loss, vegetation, and climate change in the Pampas. The diagram shows robust calibrated radiocarbon dates on bone collagen or tooth enamel (the error bars span the 2- σ probability distribution and, for many dates, are too small to appear on the chart). A transition (darker green shading) from an ecosystem dominated by C3 (Pooid) grasses to C4 (Panicoid) grasses begins ~12,000 y ago and is asynchronous across the Pampas (see [Supporting Information](#) for details). The red and blue shading on the left illustrates cool (blue) and warm (red) times. ACR, Antarctic Cold Reversal, which overlaps the northern hemisphere's Younger Dryas to some extent but is somewhat earlier. Gray distributions indicating the probability of timing of extinction for megafauna were calculated using the GRIWM best-estimate method (59); also see ref. 68. for R-code. GRIWM's best estimate of the extinction time and its 95% confidence band are younger than 5,000 y ago for glyptodonts (estimated probable timing of extinction: 1,903 calibrated years BP; range: 4,126 to -256 calibrated years BP; the negative number indicates the probability extends into the future) and *Notiomastodon* (estimated probable time of extinction: 2,415 calibrated years BP; range: 3,001–1,838 calibrated years BP). Clearly, these anomalously young extinction estimates come from a paucity of radiocarbon dates, with those few available being widely spaced in time. Such results do not provide evidence for persistence of megafauna well into the Holocene; rather, they emphasize that more radiocarbon dates are needed. Dietary preferences of extinct megafauna are indicated by G (grazer), B (browser), M (mixed feeder), and U (unknown) (50, 69).

end of the Pleistocene (44, 45). More *Notiomastodon* dates from these areas are needed to assess whether a defaunation signal would emerge. Available data are insufficient to evaluate fire history in this region.

North America—Northeastern United States. Previous work links megafaunal population crashes with vegetation response and increased fire at Appleman Lake, IN (17, 29); Silver Lake, OH (25); and in southeastern New York (Otisville, Binnewater Pond, Pawelski Farm) (36), using percentage decline of the dung-spore fungus *Sporormiella* (17, 25, 29, 53, 54) in palynological records as a proxy for local decline of megafauna, with recognition that the taphonomy of *Sporormiella* is not yet fully understood (24, 53, 55). Vegetation and fire response is consistent with climate change interacting with defaunation (17, 29), but less so with climate change on its own (17, 25, 29). Of interest is the increase in taxa such as *Fraxinus nigra* (ash), *Ostrya/Carpinus* (hornbeam/ironwood), and *Quercus* (oak) in the Indiana and Ohio sites, and *Alnus* (alder), *Betula* (birch), and *Quercus* in the New York sites during the Younger Dryas, a temporary return to cool conditions that presumably would not have favored their expansion. At the same time, charcoal frequencies rose. Increasing percentages of ash, hornbeam/ironwood, oak, alder, and birch and a concomitant increase in fire fits the model of defaunation well, because contemporary ecosystems exhibit such increases in woody understory and tree-forming plants when megafauna, especially elephants, are removed (3, 11, 56, 57), and the increase in combustible woody fuel promotes fires (15–17, 36). The proboscideans *Mammuthus* (mammoth) and *Mammuthus americanus* (mastodons) occupied the region, as did *Cervalces* (stag-moose), *Megalonyx* (ground sloth), *Ovibos moschatus* (shrub ox), *Castoroides ohioensis* (giant beaver), *Platygonus compressus* (peccary), and *Bison priscus* (bison) (58). These megafauna remained on the landscape for >2,000 y after the *Sporormiella* decline, which occurred by 13.7 ka at Appleman Lake, 13.9 ka at Silver Lake, and between ~13–14 ka at the New York sites. Summed probability distributions of radiocarbon dates from extinct megafauna in the northeast compiled by ref. 58.

suggest significant decline in their abundance by 12,600 y ago, with last records by 11,700 y ago. Confidence intervals calculated by the Gaussian-resampled, inverse-weighted McNerny et al. (GRIWM) method (59) from the data in ref. 58. indicate high probability of extinction by ~11,000 y ago. Thus, if the *Sporormiella* decline does indicate significant decrease in megafauna abundance, this decline also indicates that noticeable impacts of defaunation can occur from reducing animal densities even in the absence of final extinction. Also near 11 ka, *Quercus* pollen (oak) and charcoal abundance increase dramatically, an expected consequence of both megafaunal extinction and warming climate.

North America—Northwestern United States. We used composite data for the biogeographic region covering northern California into the Pacific Northwest. We correlated palynological and charcoal records from Twin Lake (34) in California's Klamath Mountains and Mumbo Lake in the Trinity Mountains (35) with mammal data from nearby Potter Creek (60) and Samwell Caves (32, 61), within 100 km to the southeast. Regional timing of extinction is constrained by megafauna dates from Oregon's Willamette Valley (62) plus the Manis Mastodon (63), all within ~800 km north of the California cave sites, and *Sporormiella* declines documented from lake sediments in California's central Sierra Nevada, about 800 km to the south of the caves (53, 54). Corresponding pollen profiles from the *Sporormiella* sites have not been published. For proxies of regional vegetation and sea surface temperature, we used information from Ocean Drilling Program (ODP) site 1019 (64).

These data document that by 14,000 y ago, megafauna near Twin Lake and Mumbo Lake included *Mammuthus* (mammoth) and *Mammuthus* (mastodon), *Bison* (bison), *Platygonus* (peccary), *Equus* (horse), and the sloths *Megalonyx*, *Paramylodon*, and *Nothrotheriops* (32, 60, 61). The youngest radiocarbon dates and associated confidence intervals indicate that these taxa were likely extinct by ~12,400 y ago (Fig. 3). Unlike the northeastern United States, where *Sporormiella* declines long before final extinction of megafauna, in the northwestern United States, the

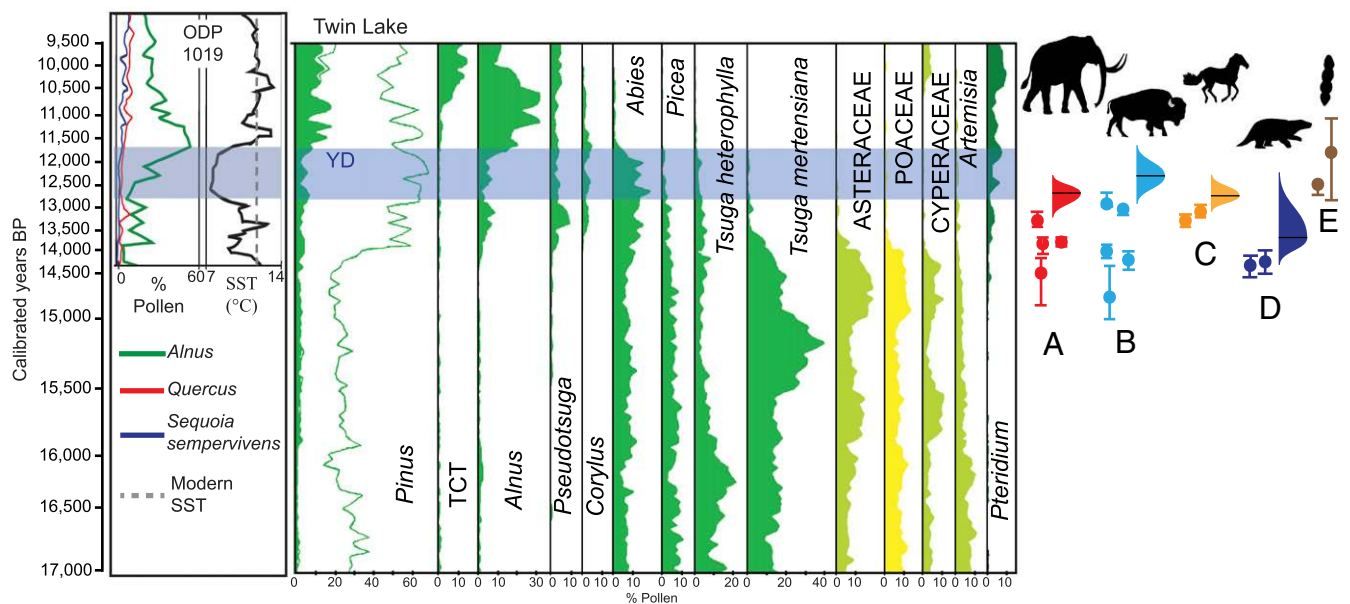


Fig. 3. Chronology of vegetation change in northern California compared with sea surface temperature (SST) and regional extinction of megafauna. (Left) SST and pollen-percentage curves are from ODP site 1019 (64). (Center) Twin Lake pollen diagram reproduced with permission from ref. 34. TCT, undifferentiated pollen from Taxodiaceae, Cupressaceae, and Taxaceae. Blue shading indicates Younger Dryas (YD). Megafauna dates are from ref. 62, except for the oldest proboscidean date from ref. 63. (Right) Proboscideans (*Mammuthus* and *Mammot*) (A), *Bison* (B); *Equus* (C), *Paramylodon* (D), and *Sporormiella* (E) from Mono Lake on the left and Exchequer Meadow on the right (53, 54). Probability distributions for extinction timing at right estimated using GRIWM method (59); see Fig. 2 for explanation. The dating of the *Sporormiella* decline is not amenable to estimating extinction timing with the GRIWM method because of the nature of the data, so here we simply report the 95% uncertainty in the dates used to constrain the ages of the youngest occurrences at each of two separate lake localities.

youngest radiocarbon dates on bones and the *Sporormiella* decline appear generally coeval (Fig. 3).

The Twin Lake record documents that *Alnus* (alder), a tree-forming shrub that spreads rapidly in the absence of browsing and trampling by ungulates (57), arrived in the region about 14,000 y ago, as the climate became favorable for its growth and when megafauna were still present. *Alnus* abundance declines slightly at the beginning of the Younger Dryas (~12,900 y ago), consistent with the expected response to cooler conditions, but then begins to increase in the midst of the Younger Dryas cooling (~12,000 y ago), at odds with expectations of an exclusively climate-driven change but consistent with predicted response to defaunation (Fig. 3). This is in marked contrast to the persistence of other plant taxa that reflect cool conditions all of the way to the end of the Younger Dryas: *Abies* (fir), *Picea* (spruce), and *Tsuga heterophylla* (western hemlock); these disappear at ~11,600 y ago. At the same time (from about 12.6–11.8 ka), charcoal abundances indicate that fire frequency increased (34), consistent with expectations of decreased large-herbivore populations.

At Mumbo Lake, both *Alnus* and *Quercus* (oaks) were present by ~13,500 y ago. Beginning ~12,000 y ago, an increase in evergreen and deciduous oaks and in *Alnus*, and an inferred transition to a more closed, diverse forest with more burnable biomass, fits expectations of megafaunal loss, although previously those changes have been attributed entirely to climatic drivers (35).

Climate proxy and vegetation data from ODP site 1019 suggest that regionally vegetation changed in the mid-Younger Dryas in ways that are inconsistent with expectations from climate change alone (64). Especially suggestive of defaunation is the dramatic rise in *Alnus* pollen, which precedes, by about 600 y, the warming of local sea surface temperatures (Fig. 3).

North America—Alaska and the Yukon (Beringia). Past work has pointed out that end-Pleistocene vegetation changes in Beringia are consistent with defaunation (26, 27, 33) (summarized in

Supporting Information). The ecological transformation from productive forests and grasslands to tundra has been interpreted (26, 27) to result from the loss of mammoths plus human-caused reduction in the densities of megafaunal species that survived. With fewer megafauna on the landscape and proboscideans absent, the reduction of trampling, browsing, and grazing pressure by big animals is postulated to have stimulated an increase in woody and leafy plants, which set off an ecological cascade by promoting accumulation of surface leaf litter that insulated soil, reduced summer soil temperatures, caused formation of permafrost, and favored growth of mosses and shrubs at the expense of grasses and other nutritious fodder for herbivores. Experimental studies support this idea (26, 27, 65).

Mammal Communities

Of regions discussed in this paper, only northern California has an adequate published record, from Samwell Cave (32), to test for local impacts of defaunation on mammalian community structure. Samwell Cave is near the two pollen sites discussed in *North America—Northwestern United States*. There, coeval with the regional signal for megafaunal extinction, a lasting decline in both richness and evenness (a metric describing to what degree abundances of individuals in different taxa are equal) in the small

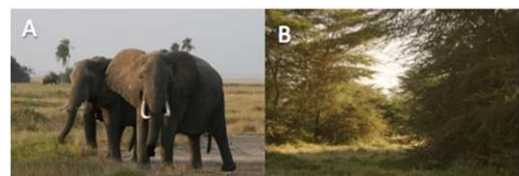


Fig. 4. Amboseli National Park, Kenya. Typical vegetation is shown where elephants roam freely (A) compared with a nearby area where elephants are excluded (B).

mammal community has been well documented (32). At the continental scale and within biogeographic provinces, most mammalian communities in the lower 48 United States also became less species-rich during the time when megafauna extinction was accelerating (30, 31) (also see *Supporting Information*). Such diversity losses at the local to continental scales are consistent with megafaunal extinction triggering adjustments of small-mammal density and distribution.

Discussion

Pinpointing a particular taxon's extinction precisely is not possible given vagaries of the fossil record; therefore, we have expressed extinction timing in terms of confidence bands (Figs. 2 and 3 and Fig. S4; also see ref. 23). Wide bands result from a paucity of dates, a limiting factor in our approach. Also, the fossil record is likely to preserve only the most dramatic, obvious signs of defaunation. This is both a strength and a weakness of our study. The strength is that if the fossil record does reveal effects of losing megafauna, those effects are severe and the implication that they have for ongoing defaunation are robust. The weakness is that subtler but perhaps equally important effects of defaunation, such as changes in nutrient cycling or in environmental heterogeneity, may not be revealed by our methods. With those caveats, the data available warrant some tentative conclusions.

In all regions examined, end-Pleistocene climate change clearly played a role in reshaping ecosystems, as is widely recognized (17–19, 21, 23, 25–27, 29–35). However, in all three North American regions, a common aftermath of losing herbivore megafauna was an apparent increase in plants that form understory and deciduous forest; this was not observed for either of the South American regions. These differences may illustrate an important point about potential effects of megafauna loss—the extent of ecological change depends on: (i) the ecological roles of the deleted megafauna; and (ii) the chance abiotic constraints—including climate and historical contingency—that provide opportunities for marked transformation of vegetation. In the North American sites, a critical ecosystem engineer seems to have been proboscideans; only after mammoths and/or mastodons disappeared or declined substantially did denser understory and deciduous forests began to flourish, much as is evident in African savannah-grasslands today (Fig. 4).

In southern Patagonia, forest-forming vegetation—*Nothofagus*—was there, but proboscideans were not. This finding may explain why, in this study area, the signal for defaunation is at best equivocal for extinction of an equid and a llama and absent for deletion of the mylodont sloth. The lost taxa were not ecosystem engineers in the sense of knocking over trees, tearing off their branches, and trampling and eating new shoots. Instead, the lost taxa inhabited open grasslands almost exclusively; only when the grasslands disappeared, as a result of climatic changes that allowed *Nothofagus* forest to become widespread, did those megafauna disappear from the region.

In the Pampas a forest-ecosystem engineer, *Notiomastodon*, was present (and lost), but in most of the region that contains dated *Notiomastodon* fossils, either the soil and climatic regimes remained unsuitable for the establishment of forests, or tree species capable of thriving in that environment never dispersed there, so even after *Notiomastodon* disappeared, major restructuring of the vegetational component of the ecosystem did not result. Further work is needed to see whether loss of *Notiomastodon* might help explain local early-Holocene expansion of forests in the northern, more tropical pampas (northernmost Uruguay and southeastern Brazil).

Conclusions

Our data provide a test of whether many different kinds of megafauna—such as proboscideans, giant ground sloths, glyptodonts, toxodonts, equids, and llamas—cause similar, lasting ecological changes when they go extinct. We find that the clear and lasting defaunation effects are variable, depending on the megafauna species and the particular ecosystem from which it was deleted. In general, it seems extinction of megafauna will likely trigger permanent ecological state shifts that manifest as substantially modified vegetation and mammalian community structure if (i) the lost megafauna are significant ecosystem engineers; and (ii) the ecosystem contains plant species susceptible to ecological release with declining herbivore pressure. The North American systems we studied, which contained proboscideans, showed particularly strong response to defaunation. The two South American areas we studied did not exhibit as strong a defaunation signal, either because megafauna that acted as major forest-ecosystem engineers were absent (southwestern Patagonia, which lacked proboscideans) or because soil and climatic limitations prevented the plant taxa capable of forming dense forest to flourish, or the requisite plants never dispersed there (the Pampas).

As has been previously recognized (17, 25, 29), the ecosystems that are most prone to shift into new regimes are those that are simultaneously impacted by both climate change, which sets the stage for previously minor components of vegetation to flourish, and defaunation, which further promotes major vegetational transformations. Both conditions are present in contemporary ecosystems with megafauna, as they were during the Pleistocene–Holocene transition.

Our study also provides insights about which modern-day ecosystems are most in danger of disappearing as a result of contemporary defaunation. Risks of irreversible change may be highest in mixed mosaics of grassland and forest that still contain the largest megafauna, which tend to be ecosystem engineers. For example, large swaths of savannah in East and South Africa, where elephants are in danger of extinction within the next two decades from intense poaching, would be likely to transform radically (Fig. 4). However, because our study demonstrates that not all megafaunal species play equal roles in maintaining ecosystem structure and function, it highlights the necessity to thoroughly understand the ecological role of each species before making predictions about how its removal will or will not cascade through an ecosystem to trigger irreversible changes.

Methods

Following refs. 21 and 66, we assume that the actual extinction of a given species was more recent than its youngest dated specimens, so we calculated estimates of the temporal intervals over which taxa became regionally extinct using the GRIWM method (59) (Figs. 2 and 3). GRIWM provides an estimate of the true time of extinction by extending the observed stratigraphic range by the average gap size (but up-weighting younger gaps to accommodate nonrandom preservation and recovery of fossils) within the known stratigraphic range. GRIWM accommodates the uncertainties in the radiometric dates by providing a 95% confidence band around the estimated extinction time. All dates mentioned in the text were vetted from published literature and are expressed in calendar years before present calibrated using Calib version 7.0.2 (67). For South American megafauna dates, only radiocarbon dates considered robust (with a score of ≥ 11 , following ref. 21) were used.

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